

Quantification of carbon monoxide emissions from African cities using TROPOMI

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Abstract. Carbon monoxide (CO) is an air pollutant that plays an important role in atmospheric chemistry and is mostly emitted by forest fires and incomplete combustion in for example road transport, residential heating, and industry. As CO is co-emitted with fossil fuel CO₂ combustion emissions, it can be used as a proxy for CO₂. Following the Paris agreement, there is a need for independent verification of reported activity-based bottom-up CO₂ emissions through atmospheric measurements.

5 CO can be observed daily at global scale with the TROPOMI satellite instrument with daily global coverage at a resolution down to 5.5x7 km². To take advantage of this unique TROPOMI dataset, we develop a cross-sectional flux-based emission quantification method that can be applied to quantify emissions from a large number of cities, without relying on computationally expensive inversions. We focus on Africa as a region with quickly growing cities and large uncertainties in current emission estimates. We use a full year of high-resolution WRF-simulations over three cities to evaluate and optimize the performance of
10 our cross-sectional flux emission quantification method and show its reliability down to emission rates of 0.1 Tg CO yr⁻¹. Comparison of the TROPOMI-based emission estimates to the DACCIWA and EDGAR bottom-up inventories shows CO emission rates in northern Africa are underestimated in EDGAR, suggesting overestimated combustion efficiencies. We see the opposite when comparing TROPOMI to the DACCIWA inventory in South Africa and Côte d'Ivoire, where CO emission factors appear to be overestimated. Over Lagos and Kano (Nigeria) we find that potential errors in the spatial disaggregation of national emis-
15 sions cause errors in DACCIWA and EDGAR, respectively. Finally, we show that our computationally-efficient quantification method combined with the daily TROPOMI observations can identify a weekend effect in the road transport-dominated CO emissions from Cairo and Algiers.

1 Introduction

Carbon monoxide (CO) is an air pollutant that is mostly emitted by anthropogenic sources. It is a product of incomplete combustion in for example road transport, residential heating, industry and forest fires (Zhong et al., 2017). CO is a precursor of ozone, and because it reacts with the hydroxyl radical (OH) its presence effectively increases the atmospheric lifetime of methane (Daniel and Solomon, 1998; Jacob, 1999; Wuebbles and Hayhoe, 2002). The concentration of CO is therefore im-

portant for climate modelling. Furthermore, as many processes that emit CO also emit carbon dioxide (CO₂), knowledge of CO emission rates can provide additional information about CO₂ emissions (Wu et al., 2022; Park et al., 2021). The Inter-governmental Panel on Climate Change (IPCC) identified a need for independent verification of the reported greenhouse gas emissions through measurements (IPCC, 2019). From space this is challenging for CO₂ as its long atmospheric residence time results in high background concentrations, making it hard to detect emissions. For this reason, measuring the 'short lived' CO can be a useful alternative (Silva et al., 2013; MacDonald et al., 2023). We present a method to quantify CO emission rates over cities in Africa using TROPOMI satellite observations.

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As transport (23%) and residential heating (35%) are key contributors to total anthropogenic CO emissions (Zhong et al., 2017), cities are an important source of CO. In Africa, the contributions of transport and residential heating are estimated at 27% and 38% of anthropogenic CO emissions respectively in the DICE-Africa inventory (Marais and Wiedinmyer, 2016) and 17% and 72% in the DACCIWA inventory (Keita et al., 2021). The importance of these two sectors is further confirmed by a large number of ground-based measurements specifically aimed at traffic (Diab et al., 2005; Lindén et al., 2008; Zakari et al., 2020; Doumbia et al., 2021) and domestic heating (Havens et al., 2018; Kansime et al., 2022; Saleh et al., 2023) that show CO concentrations in African cities exceeding air quality guidelines by the World Health Organisation. Urbanisation scenarios predict a growth in both the number of mega-cities and their populations, leading to larger emission rates and increased health risks. Africa is predicted to have a large urbanisation rate in the coming years. Hoornweg and Pope (2017) predict the continent to house five of the ten largest cities by 2100, compared to one of ten today. Africa is also a region for which relatively large uncertainties are present in emission inventories as only a few are dedicated to the region (Keita et al., 2021). Current emission inventories are based on so called bottom-up methods where emissions are estimated by combining activity data (e.g. national fuel consumption statistics) with emission factors and spatially distribute the emission estimates using proxies like population density (Janssens-Maenhout et al., 2019). These bottom-up methods are also used to report country-level greenhouse emission estimates to the United Nations Framework Convention on Climate Change (UNFCCC). However, lack of detailed data results in large uncertainties (Macknick, 2011; Cai et al., 2019; Oda et al., 2019).

Independently of bottom-up methods, emissions can also be estimated by top-down methods where atmospheric concentrations are measured and used to infer the corresponding emission rates. Multiple studies have investigated urban CO emissions using ground-based measurements (Badarinath et al., 2007; McKain et al., 2012; Bi et al., 2022). Many studies have also shown the capability of satellite measurements for this specific task (Borsdorff et al., 2020; Tian et al., 2022a; Plant et al., 2022; Wu et al., 2022). For CO, the TROPospheric Monitoring Instrument (TROPOMI) on ESA's Sentinel 5 precursor satellite is of particular interest (Veefkind et al., 2012). It was launched in 2017 and provides daily global coverage with a resolution of 5.5 by 7 km², which makes it suited to investigate city emissions worldwide.

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An advantage of polar-orbiting satellites is their ability to monitor the entire globe. However, most satellite-based studies of CO so far have focused either on regional inversions (Yumimoto et al., 2014; Qu et al., 2022) or on trends in concentrations

(Lama et al., 2019; Park et al., 2021; Hedelius et al., 2021), while only a few studies have tried to quantify emissions from individual cities or point sources (Dekker et al., 2017; Borsdorff et al., 2020). These urban emission quantifications use atmospheric inversions, which require computationally expensive high-resolution simulations with Chemical Transport Models (CTMs). Although inversions are able to get relatively accurate emission estimates, they are difficult to apply to a large number of sources. To take full advantage of the TROPOMI data, we adjust the mass balance Cross-Sectional Flux (CSF) method, originally developed for high-resolution point-source quantifications, to be used with TROPOMI data over urban areas. After evaluating the method using atmospheric transport simulations, we use it to estimate emissions from the largest cities in Africa.

2 Data & Methods

This section describes the different data products used in development of the Cross-Sectional Flux method and the simulations that were used to calibrate the model and evaluate its performance using an Observing System Simulation Experiment (OSSE). An OSSE is an experiment where a model or method is applied to synthetic data to evaluate the benefit of using this data and/or method. Which for this work means evaluating whether the CSF can be used to correctly estimate emissions from TROPOMI-like synthetic data. Figure 1 shows the roles of the different data products that are used and further described in Section 2.1 to 2.6. In addition, in Section 2.6 we show that the CSF method can be successfully applied to simulated data.

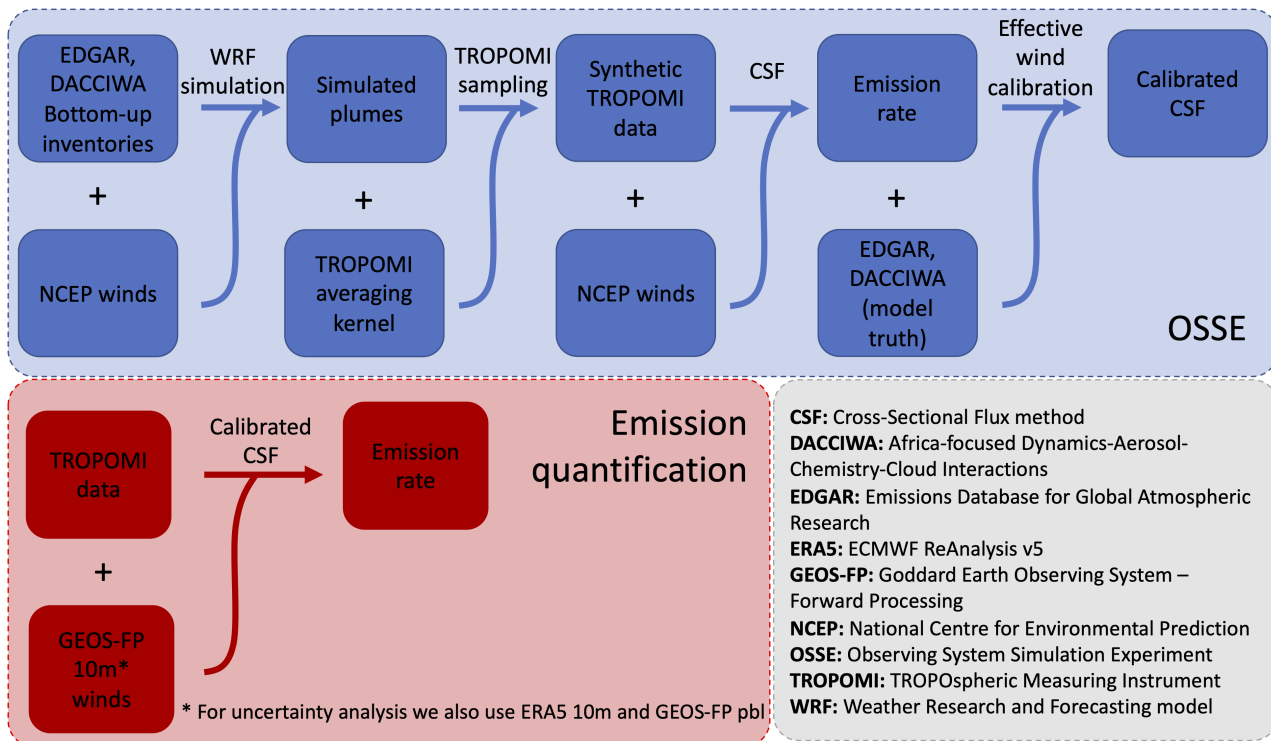


Figure 1. Schematic description of the use of the different data products within the OSSE and the subsequent emission quantification using TROPOMI data. The data products are discussed in the Section 2.1-2.6. First, the CSF is applied to simulated synthetic plumes in order to determine appropriate values for the various parameters used in our method. Second, an effective wind is calibrated by using the known emission rates of the simulated plumes following the procedure by Varon et al. (2018). The CSF, now calibrated on the synthetic plumes, is subsequently applied to satellite data to estimate emission rates of African cities.

2.1 TROPOMI carbon monoxide data product

TROPOMI provides total column carbon monoxide concentrations with daily global coverage at 13:30 local time using the shortwave-infrared band (SWIR) at 2305–2385 nm (Veeffkind et al., 2012). From the spectral signal, the CO concentration is inferred using the shortwave-infrared CO retrieval (SICOR) algorithm (Borsdorff et al., 2018). We use three years of data (2019–2021) from the operational data product (Landgraf et al., 2018). To assure high quality data, all pixels with a TROPOMI quality flag below 0.7 are removed, leaving data that are cloud-free or only have low altitude clouds. The CO concentration over cloud-free water surfaces is difficult to retrieve, due to the low intensity of reflected light, therefore, we only use observations with a quality flag equal to 0.7 (low altitude clouds) over water. The resulting dataset shows good agreement with ground-based measurements, with a mean difference per station of $2.45 \pm 3.38\%$ to the unscaled Total Carbon Column Observing Network (TCCON, (Wunch et al., 2011)) columns and $6.5 \pm 3.54\%$ to the Infrared Working Group of the Network for the Detection of Atmospheric Composition Change (NDACC-IRWG) measurement stations (Sha et al., 2021).

2.2 EDGAR and DACCIWA bottom-up inventories

85 We use two different bottom-up inventories to compare the TROPOMI emission estimates with: the Emissions Database
for Global Atmospheric Research (EDGAR) version 5 (Oreggioni et al., 2021) and the Africa-focused Dynamics-Aerosol-
Chemistry-Cloud Interactions in West-Africa (DACCIWA) inventory (Keita et al., 2021). These inventories are also used in
our atmospheric transport simulations (Section 2.3) to simulate TROPOMI observations. Both inventories provide yearly grid-
90 ded emission rates at 0.1° resolution up to 2015. Due to its global scope, the EDGAR inventory relies mostly on international
statistics and spatial proxies combined with national data while its emission factors are based on IPCC methodology for green-
house gases (Eggleston et al., 2006) and the EMEP/EEA emission inventory guidebook for air pollutants (Nielsen, 2013). The
DACCIWA inventory provides emission rates over the African continent, ranging from -35 to 38° latitude and -25.5 to 63.5°
longitude. It uses similar international data but is supplemented by local measurements of emission factors and data from local
authorities (Keita et al., 2021). As the DACCIWA inventory characterizes emission from fewer (sub)sectors than EDGAR, we
95 merge different sectors in EDGAR to match those used in the DACCIWA inventory to make them intercomparable. When
reporting urban emissions from EDGAR and DACCIWA, we sum emissions over the pixel closest to the city center, its 8
neighbours and all directly attached pixels where the population density exceeds the surroundings by 1.8 standard deviation.
Changing the city mask to $0.3 \times 0.3^\circ$ or $0.7 \times 0.7^\circ$ boxes changes the emissions by 10-20% for 17 out of 29 cities in EDGAR
and 16 out of 29 cities in DACCIWA. Although larger deviations up to 50% in densely populated areas like South Africa are
100 observed, the observed patterns discussed in Section 3 are valid for these masks as well.

2.3 WRF chemical transport model

To test and calibrate our emission quantification approach we apply our CSF method to simulated TROPOMI data for three
urban areas. We use the Weather Research and Forecasting (WRF) chemical transport model version 4.1 (Powers et al., 2017),
to simulate column CO mixing ratios over Cairo (Egypt), Bamako (Mali) and Lagos (Nigeria) for 2019, using December 2018
105 as spin-up month. These three African cities form a diverse set, with Cairo next to the Nile river, Bamako at the boundary of the
Sahara desert and Lagos at the coast of the Atlantic Ocean. We simulate CO as an inert tracer and drive the simulations with
meteorological fields from the National Centre for Environmental Prediction (NCEP, 2000). All simulations have three-layer
nested domains where the outer domain covers $2673 \times 2673 \text{ km}^2$ at a resolution of 27 km; the middle and inner domain cover
 $891 \times 891 \text{ km}^2$ at 9 km resolution and $315 \times 315 \text{ km}^2$ at 3 km resolution respectively (Fig. 2). Initial and 6-hourly boundary con-
110 ditions to capture the background CO are taken from the Copernicus Atmosphere Monitoring Service (CAMS) at $0.25^\circ \times 0.25^\circ$
resolution, (Inness et al., 2015). The resulting background is scaled to match the mean background observed by TROPOMI
over the full year. We use emissions from the global Emissions Database for Global Atmospheric Research (EDGAR) version
5 and the Africa-focused Dynamics-Aerosol-Chemistry-Cloud Interactions in West-Africa (DACCIWA) inventory distributed
across the vertical model levels according to the sector specific vertical profiles provided by Bieser et al. (2011). Typical injec-
115 tion heights for CO emissions from transport and the residential sector are 0-20 m, while emissions from industry are typically
injected into the atmosphere at 100-200 m (Bieser et al., 2011). City-specific hourly, daily and monthly temporal profiles for

each emission sector are taken from Guevara et al. (2021). To maintain flexibility over model output the different sectors in the emission inventories (19 for EDGAR and 6 for DACCIWA) are simulated separately. We sample the model output (at the TROPOMI overpass time) to facilitate comparison to the TROPOMI carbon monoxide data as discussed in detail in Section 120 2.6.

While EDGAR and DACCIWA only include the primary production of CO, the concentrations observed by TROPOMI include CO from secondary production as well. CO is produced by oxidation of volatile organic compounds (VOC) with methane as main contributor (Rozante et al., 2017). Mixing ratios of non-methane volatile organic compounds (NMVOC) observed in urban locations are typically of the order of $10^{-3} - 10^{-2}[\text{NMVOC}]/[\text{CO}]$ (Von Schneidemesser et al., 2010). 125 Dekker et al. (2019) showed that chemical production of CO by methane and NMVOC over cities only contribute 4% to the total CO signal, justifying the simulation of CO as an inert tracer in our approach. Due to the 10 year atmospheric lifetime of methane, its contribution to CO production will result in a uniform concentration (Park et al., 2013), that is subtracted with the background. NMVOC have lifetimes of 0.6-10 days (Guo et al., 2007) that are much shorter than the lifetime of CH₄, but due to their low urban mixing ratios (~ 1%) their effect on the estimated emission rate is much smaller than the reported 130 uncertainty of the CSF. This is consistent with the observation that the emission estimates of individual transects (that span a timescale of up to ~ 10 hours) are stable and do not increase with increasing distance from the city (Section 2.4).

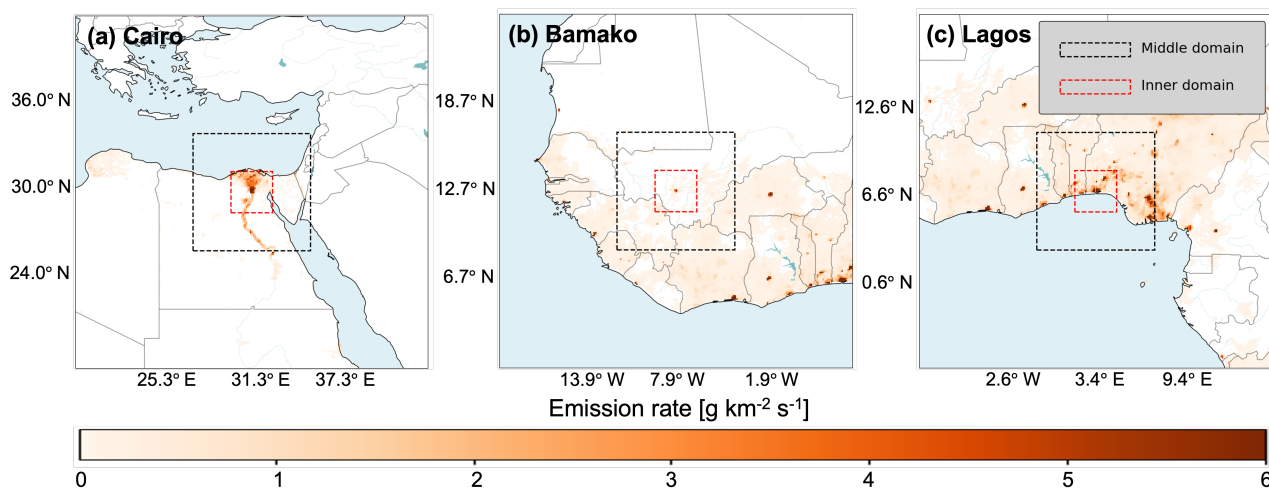


Figure 2. Domain setup of the WRF simulations over Cairo (Egypt), Bamako (Mali) and Lagos (Nigeria). The inner domain (red) spans $315 \times 315 \text{ km}^2$ around the city at 3 km resolution. The middle (black) and outer (full figure) domain cover $893 \times 893 \text{ km}^2$ and $2673 \times 2673 \text{ km}^2$ at 9 and 27 km resolution respectively. All panels show the emission rates from the DACCIWA inventory.

2.4 Cross-Sectional Flux method

The Cross-Sectional Flux method (CSF) has been shown to be an effective way to quantify emission rates of plumes observed by satellites (Varon et al., 2018, 2020; Sadavarte et al., 2021b; Tian et al., 2022b). It is based on the continuity equation which

135 relates the flux through a closed surface to the associated emission rate:

$$Q = \oint U_{\perp} \Delta\Omega dA, \quad (1)$$

where Q [kg s^{-1}] is the emission rate, U_{\perp} [m s^{-1}] is the wind speed perpendicular to the closed surface, $\Delta\Omega$ [kg m^{-3}] is the enhancement at the closed surface and dA [m^2] is a surface element. As illustrated in Fig. 3A, the plumes have a distinct direction as they move with the wind, they are very directional and it suffices to integrate over perpendicular transects that
140 cover the entire plume-width (Fig. 3B). Equation 1 can then be rewritten as:

$$Q = \int U_{\perp}(x, y) \Delta\Omega(x, y) dy, \quad (2)$$

with x, y coordinates along and perpendicular to the plume respectively as in Varon et al. (2018). Assuming a constant emission rate, transects at different distances downwind of the source should yield the same emission quantification and can be averaged to make the method more robust.

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We then optimize the implementation of the CSF on TROPOMI data for city-like sources. Figure 3A shows a CO-plume observed by TROPOMI over Cairo on April 7th 2019. As expected, the plume follows the 10-m wind direction given by NASA/GMAO GEOS-FP reanalysis data (Molod et al., 2012). We start by determining the city center, for our purposes defined as the location at the center of the urban emissions and therefore best representing the origin of the city's total emission.
150 The location of the center is determined by taking the weighted average position of the pixels in DACCIWA that are part of the city mask introduced in Section 2.2. Weights of the pixels are equal to their emission rate. By determining the city location using the emission inventory, we ensure that we are comparing similar regions when we compare our satellite-based emission estimates with the emission inventories in Section 3. To make sure the entire plume is downwind, we start the transects of our CSF 0.1° upwind of this city center. As the wind direction is an important source of uncertainty, the downwind direction
155 can not be solely based on the GEOS-FP reanalysis wind data. Instead, following Sadavarte et al. (2021a), we infer the wind direction from the satellite observations by selecting the direction of the highest mean downwind concentration within 90° of the reanalysis wind direction. To do so, we calculate the mean downwind concentration over 180 boxes (0.1° width and 0.4° length) rotated at 1° intervals, and pick the direction with the highest downwind concentration. In the absence of a clear plume, this method would create a positive bias as it would select the highest enhancement in the noise. Therefore, we use
160 the reanalysis wind direction if the mean enhancement does not exceed 5 ppb. To calculate CO enhancements, we subtract a background calculated as a mean upwind concentration over a $0.4^{\circ} \times 0.4^{\circ}$ square starting 0.3° upwind from the city center (Fig. 3B). If this box contains fewer than five valid pixels, we extend it symmetrically with two arcs of a circle of 10° , 20° , 45° up to 60° until there are at least five TROPOMI pixels in the background region (Fig. 3C). We use extension in an arc-like fashion rather than increasing the size of the square to be able to get estimate background values for coastal cities. Retrievals over
165 water are only possible if there are clouds present, hence increasing the size of the background square upwind could result in background pixels far away from the city that are not representative for the local background.

After determining the initial direction of the plume, we need to better capture the shape of the plume to draw transects perpendicular to the plume. The shape of the plume is determined in two steps. First, we select all pixels in a downwind box (0.3° width, 0.8° length) that exceed the mean concentration in the surrounding $3^\circ \times 3^\circ$ area by more than 1.8 standard deviations, these pixels are referred to as the spline mask. We then fit a 2D-spline (0.8° length) through the resulting spline mask. If, due to a lack of signal or missing pixels, the spline mask contains fewer than 3 pixels, a spline fit is unlikely to capture the true plume shape and we use a straight line in the (optimized) wind direction instead (Fig. 3C). The transects (0.4° width) are drawn perpendicular to the spline, separated by 0.04° . The transects have a larger width than the box used to determine the spline mask to ensure the transects cover the entire plume width. All pixels overlapping with the transects are used in the emission quantification. We also include days where no clear plume is visible to avoid systematic overestimation of the average emission rate. Using Eq. 2, an emission estimate can be derived for every transect. We stop drawing transects when the emission rate estimates of two consecutive transects are more than one standard deviation below the mean estimate of the earlier transects, indicating the end of the plume. Transects with less than 70% pixel coverage are removed from the estimate as they will not have a complete integral, resulting in underestimated emissions.

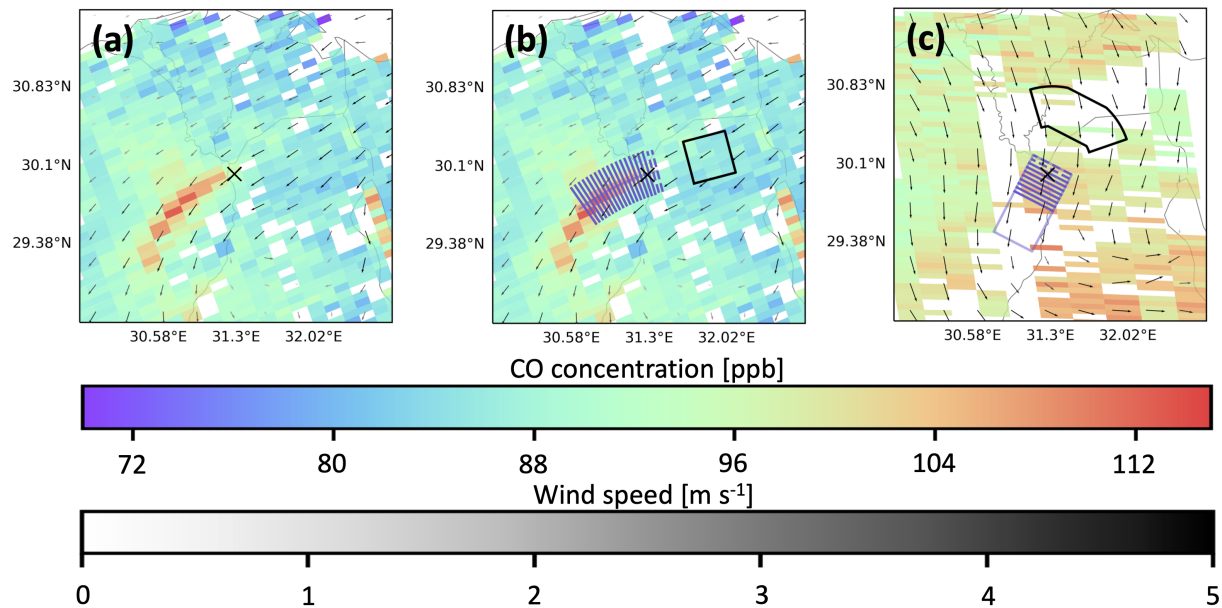


Figure 3. Example of how the cross-sectional flux transects perpendicular to the plume are drawn. (a) TROPOMI data over Cairo on April 7th 2019. The city center is shown with a cross and taken from the DACCIWA emission inventory. The arrows show direction and magnitude of GEOS-FP 10-m winds at their native $0.25^\circ \times 0.3125^\circ$ resolution (Molod et al., 2012). (b) Pixels downwind of the city that surpass the regional background by more than 1.8σ form a plume mask through which a 2D spline is fitted (grey line). The transects used for quantification (purple lines) are drawn perpendicular to the spline fit starting 0.1° upwind of the city center. The first two transects are dashed to reflect that they are not used for the emission quantification. The background is estimated over the black $0.4^\circ \times 0.4^\circ$ box upwind. (c) TROPOMI observation over Cairo on March 27th 2020. Due to a lack of coverage there are insufficient pixels to generate a reliable spline mask. A rectangular box (grey) is therefore used to draw transects instead. The basic background is extended symmetrically with circle arcs to compensate for a lack of coverage upwind.

Contrary to studies using high-resolution satellites (Varon et al., 2018, 2020), the plumes observed with TROPOMI cover distances over which there can be significant fluctuations in wind speed and direction. We therefore use the wind speed at each transect instead of a single wind speed for the entire plume. The wind speed at the transects is calculated in two steps. First, a wind speed is calculated for each TROPOMI pixel by interpolation of the reanalysis wind product. Second, the wind speed for each transect is determined by taking the average wind speed of the overlapping TROPOMI pixels, weighted by the length of the overlap. Similar to trends observed in (Sadavarte et al., 2021b), the first two transects are found to have roughly 30% lower emissions than the transects further away, which have a stable mean emission rate. This pattern is consistent across the cities investigated. One reason is that the early plume only captures part of the city's emissions, another explanation is that the associated pixels might see a partial-pixel absorption saturation effect (Pandey et al., 2019). Incorporating the first two

190 transects would result in an average underestimation of emissions by 8%. We therefore remove the first two transects from the emission estimation.

On the spatial scale relevant to plumes observed by TROPOMI, there can be contamination of the city signal by carbon monoxide produced by open fires (e.g. agricultural fires or wildfires). CO enhancements caused by open fires can result in overestimation of either the background or the downwind urban enhancement depending on their location. To avoid this, days with considerable CO contributions from open fires have been removed from our estimates. These days were selected based on the fire emission data from the Global Fire Assimilation System fire emission (GFAS) inventory (Kaiser et al., 2012) that is based on satellite measurements of fire radiative power. Days with cumulative fire emissions over 57 Mg per hour (equivalent to 0.5 Tg yr⁻¹) within 1.5° from the city center are removed. Additionally, days with strong burning events closer to the city (23 Mg hr⁻¹ within a 0.75° radius) are removed as well (Appendix B). Although the change in emission rate by this filtering is limited for most cities, the filter can change estimated emission rates by up to 47%, as seen in Lusaka (Zambia).

2.5 Uncertainty analysis

To estimate the uncertainty of the estimated emission rates, we compile an ensemble of emission estimates for each city. We generate the ensemble by varying parameters of the quantification method such as the wind database used. For example, we vary parameters such as the number of transects and the distance of the background box. To incorporate the uncertainty on the wind data, we use our method with GEOS-FP 10 meter altitude winds, GEOS-FP planetary boundary layer averaged winds (Molod et al., 2012) as well as 10 meter altitude winds from the ERA5 product, provided by the European Centre for Medium-Range Weather Forecasts (ECMWF; (Hersbach et al., 2020)). A complete list of the varied parameters and their ranges can be found in appendix A. For each city, the spread in the resulting ensemble is reported as uncertainty.

2.6 Calibration and validation

210 This section describes the application of the CSF to simulated CO column mixing ratios. The simulations are used to determine parameter settings (e.g. spline length and transect width) and to calibrate an effective wind (Varon et al., 2018) for TROPOMI-sized pixels. In addition, the simulations are used to evaluate how well the CSF can quantify emission rates of simulated plumes.

As simulated (and TROPOMI observed) plumes stay within the inner domain, only the inner domain is used to test the performance of the CSF. A set of synthetic TROPOMI observations is created by sampling the simulation output over the TROPOMI footprints, applying its averaging kernel, selecting pixels based on quality value as discussed in Section 2.1, and adding Gaussian noise with a standard deviation equal to the reported uncertainty of the respective TROPOMI pixel. The TROPOMI quality value filtering ensures relatively clear sky observations with good surface sensitivity. We also calculate "idealized" pressure weighted columns, which assume a uniform vertical sensitivity (flat averaging kernel), over the TROPOMI footprints without taking into account whether there is a valid TROPOMI observation as a first check to see whether the CSF can reproduce the emissions used as model input.

We first test the validity of the CSF method using the idealized columns with 10-m winds output by the WRF simulation. The WRF winds are directly responsible for transport within the simulation, and can therefore be considered as the true wind fields behind the modeled concentrations. Parameters like the number of transects and distance of the background region are tuned to get optimal quantification estimates on the simulated data, such that the fitted splines capture the observed curvature of the plumes and the background is not affected by the urban emissions. A list of the different parameters and their values can be found in Appendix A. While the true wind field varies with altitude, the CSF method requires just a single (2D) wind field that is representative for the transport of the plume. We use the simulations to calibrate the CSF by introducing an effective wind speed that replaces the wind speed in Eq. 2 following the procedure by (Varon et al., 2018). The effective wind speed is the wind that best captures the transport of the plume. It is a parametrization of the true wind speed to account for the effects of turbulence and variation in vertical wind speed and injection height. As the emission rates in the WRF simulations are known, the effective wind can be calculated explicitly for every orbit for each of the simulated cities. Figure 4 shows the relation between the effective wind (U_{eff}) and the WRF 10-m winds U_{10} averaged over the plume; the fitted linear relation is:

$$U_{eff} = a_{10}U_{10} + b_{10}, \quad (3)$$

with $a_{10} = 1.43$ and $b_{10} = -0.92 \text{ m s}^{-1}$ ($R^2 = 0.82$). U_{10} is the wind speed at the time of overpass at 10 meter altitude. We determine the effective wind relationship separately for the planetary boundary layer averaged winds, which tend to be higher than the surface winds. The resulting calibration gives $a_{PBL} = 0.98$ and $b_{PBL} = -0.20$ ($R^2 = 0.62$). While the absolute value of the PBL winds is closer to the effective wind speeds, using the U_{10} winds captures more of the variability.

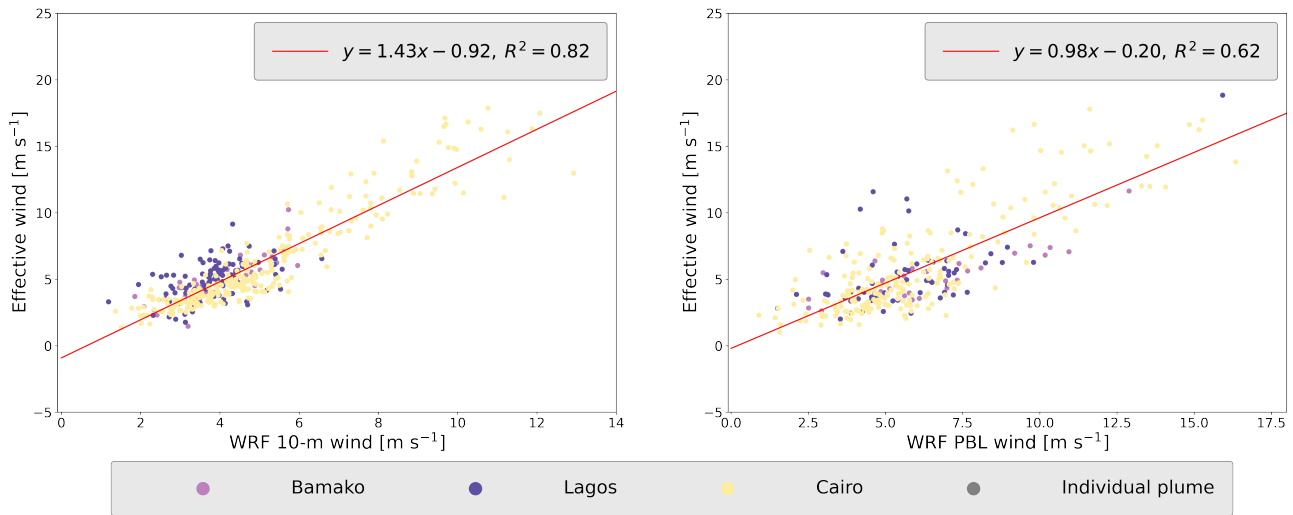


Figure 4. Determination of the relation between the effective wind and both the wind speed at 10 meter altitude (left) and the planetary boundary layer averaged winds (right). The effective wind corrects for the effects of turbulence, injection height and variation in the vertical wind profile. The simulated plumes used in the calibration cover a full year over Cairo, Bamako and Lagos.

240 After determination of the effective wind on plumes with idealized pressure profiles, we test the performance of the CSF on more realistically sampled plumes which include the TROPOMI quality filtering and averaging kernel sensitivities as described in Section 2.1 to see whether the CSF can correctly quantify emission rates from synthetic observations with quality filtering and non-uniform vertical sensitivity. To test the method's sensitivity, we perform an additional effective wind calibration on these data. The resulting linear fit ($a = 1.4$, $b = -0.85$, $R^2 = 0.23$) yields similar results and shows the filtering has limited
245 impact on the calibration, while the lower R^2 value reflects the larger variation in estimated emission rates. At the same time we test the lower limit to which we can trust the resulting emission estimates, as smaller enhancements are more difficult to distinguish from the background. As the modeled output concentration from the WRF simulations without chemistry scales linearly with the magnitude of the input emissions, emissions from the different sectors provided by the bottom-up inventories can be scaled up and down without having to rerun the chemical transport model. This allows us to simulate plumes from
250 cities with different emission rates with limited effect on the simulated background through scaling of the emission sector most concentrated in the considered urban area. We use this to determine the lower limit to which our method can be trusted. Figure 5 shows a comparison between the simulated input and the retrieved emission rates for the three simulated cities. The results suggest that the CSF is able to reproduce input of the WRF simulations when using one year of data for cities with emission rates larger than 0.1 Tg yr^{-1} .

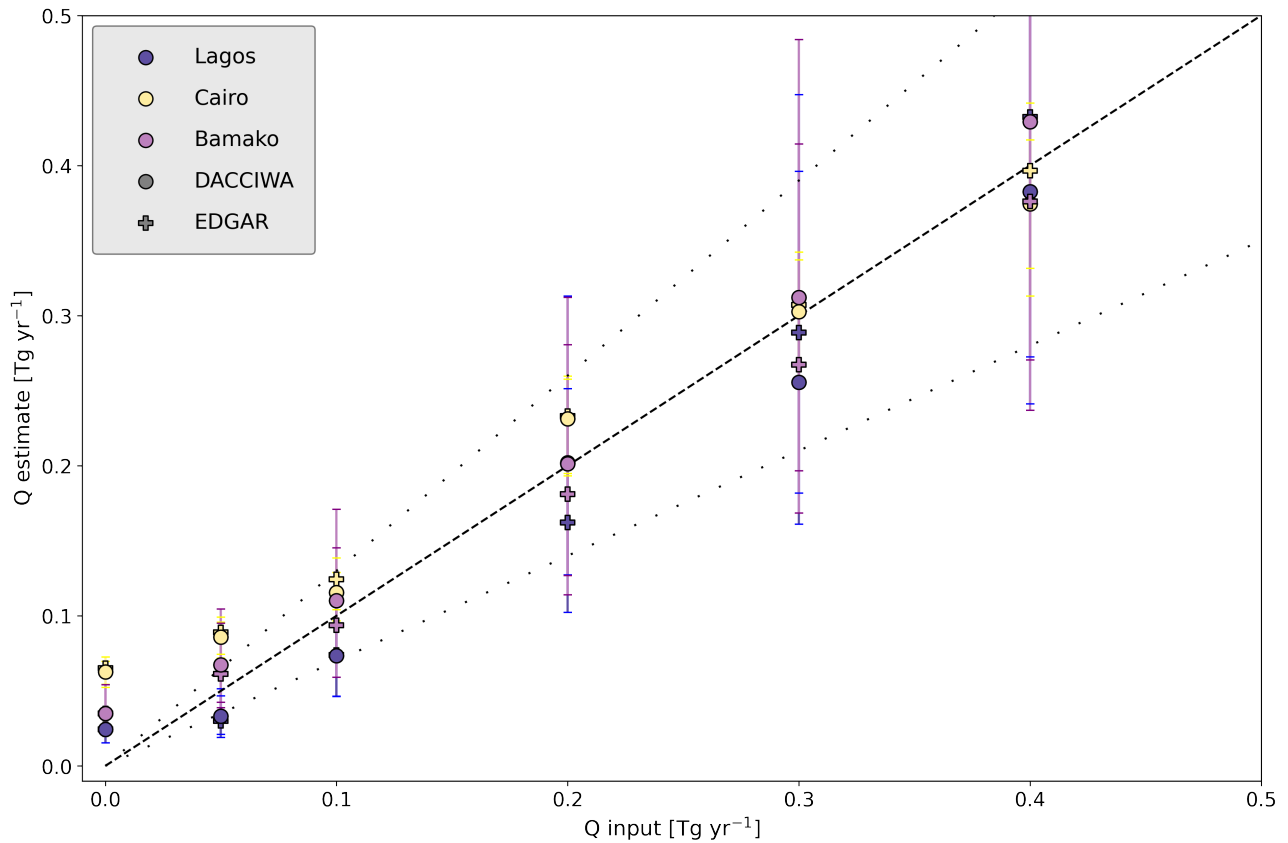


Figure 5. Emission quantification by the CSF on simulated plumes. The plumes are simulated with the WRF-model for the year 2019 over Lagos (Nigeria), Cairo (Egypt) and Bamako (Mali). The simulations either used the EDGAR global bottom-up inventory or the DACCWA inventory. The dotted lines show a 30% deviation from the (dashed) 1:1 line.

255 To quantify TROPOMI plumes over all major cities in Africa we will use the NASA/GMAO GEOS-FP wind fields (Molod et al., 2012) rather than the WRF simulated wind fields which are available only for the 3 cities selected for evaluation and calibration. For each TROPOMI pixel, we spatially interpolate the GEOS-FP wind field, which has a $0.25^\circ \times 0.3125^\circ$ spatial resolution and a 1-hour time resolution. The NCEP winds that drive the WRF simulated wind fields have a coarser time resolution of 6 hours. Figure 6 shows emission estimates of simulated data using the GEOS-FP wind fields instead of the
 260 WRF fields to mimic uncertainties in the wind fields. Individual days are shown as colored dots, while the mean over the full year is shown as a colored line to represent the average estimate. The uncertainty of the average is determined as discussed in section 2.5 and shown as a shaded area. The true emission is shown as a black dotted line, and lies within the uncertainty of the estimate for both Cairo and Bamako. For Lagos the emission rate is underestimated when using the WRF simulations as the NCEP wind fields that drive the simulations are higher than both the GEOS-FP and ERA5 wind products specifically over

265 Lagos by about 60%. The difference between the wind products might be caused by the fact that Lagos lies in the West-African monsoon region where transport has been shown to be difficult to model (Liu et al., 2014).

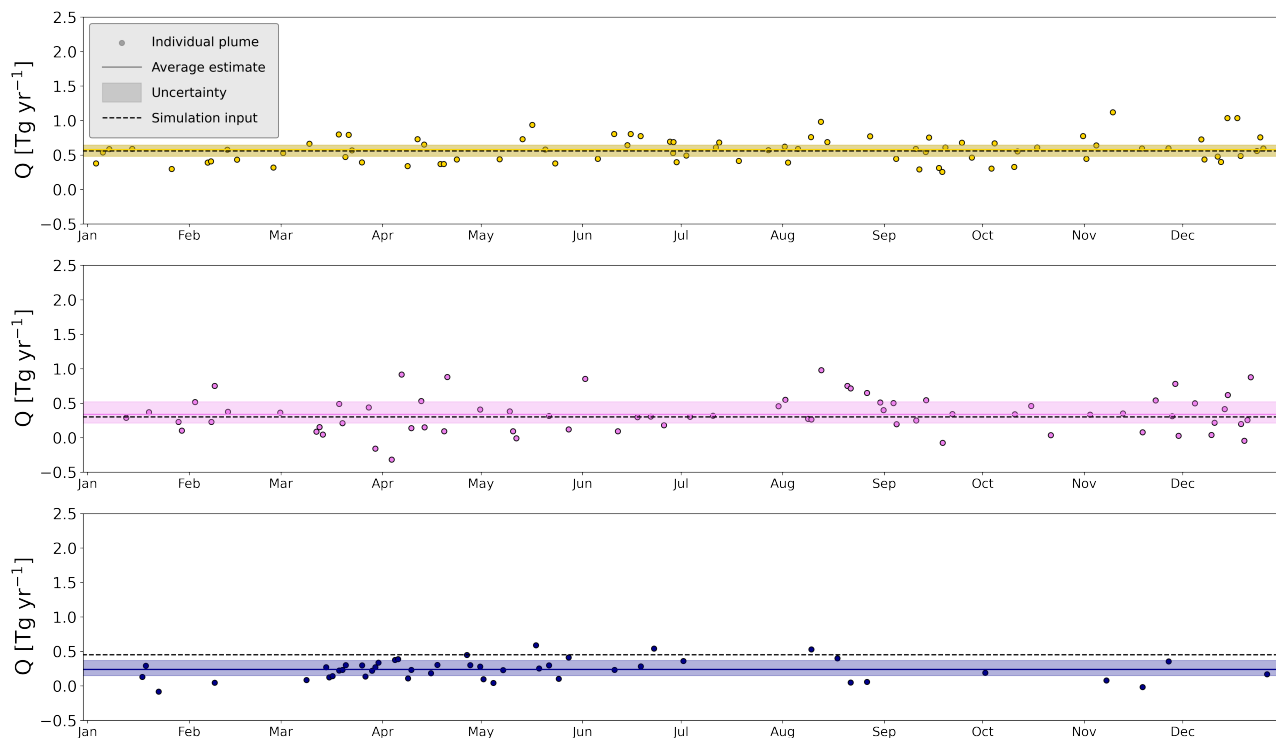


Figure 6. To check the validity of the CSF method for quantification of city emissions we apply the method to simulated plumes sampled as TROPOMI would see them. The dots show CSF emission estimates for individual days over Cairo, Bamako and Lagos respectively. The dark colored line shows the annual CSF mean with the uncertainty based on the emission ensemble shown by the shaded area. The simulation emission input, black dotted line, lies within the uncertainty of the mean CSF emission estimate for Cairo and Bamako showing that the CSF can successfully quantify these urban emissions. For Lagos the emissions are underestimated as the NCEP winds used to drive the simulation are much higher than both the GEOS-FP and ERA5 wind products.

3 Results & Discussion

270 After verification of the validity and calibration of the method, we apply it to 29 of the largest cities in Africa. These cities are chosen based on their population or because they are emitting above the CSF's quantification threshold in the DACCIWA inventory. Figure 7 shows the results of our TROPOMI quantification and a comparison with the DACCIWA and EDGAR inventories. This data is also included in the appendix in Table D1. As discussed in Section 2.2, we used different sizes for the city masks applied to the bottom-up inventories to ensure a fair comparison to the satellite-based emission estimates and found the choice of city mask did not impact our conclusions.

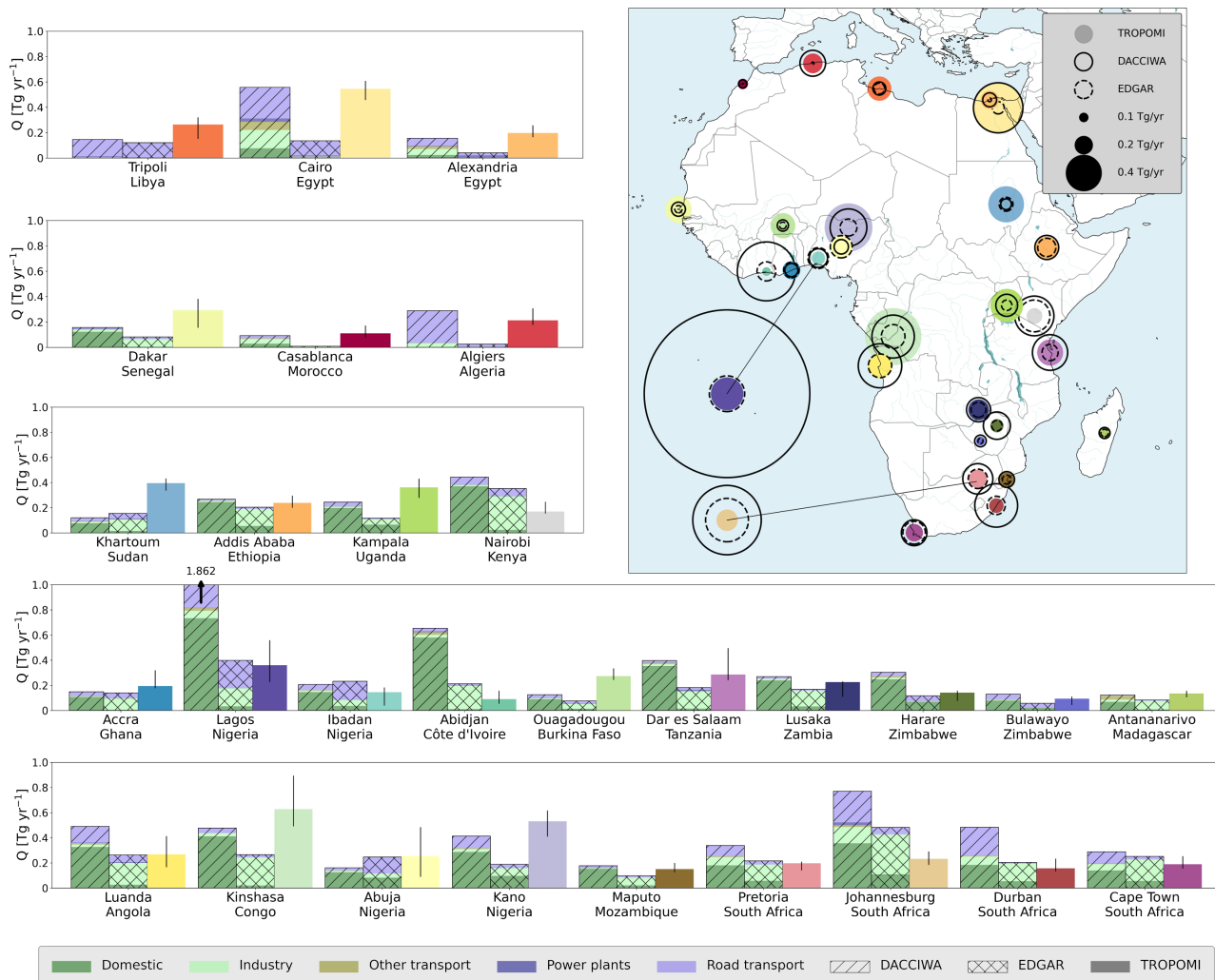


Figure 7. CSF emission quantifications for the largest African cities. Comparison between TROPOMI emission estimates averaged for 2019-2021 (shown as colored circles) and the DACCIWA and EDGAR v5 emission inventories for 2015 shown by the black (dashed) rings. The emission strength is indicated by the size of the circles or rings. The same comparison is made in bar plots where the first two bars show the emission rates from DACCIWA and EDGAR respectively, including the sectoral breakdown. The third bar gives the corresponding TROPOMI estimate where the uncertainty is given by the range of the ensemble. The cities are ordered by geographical location. The emission estimate for Lagos in DACCIWA extends beyond the figure boundary.

On average we find TROPOMI emissions of 0.25 Tg yr^{-1} per city, compared to 0.35 Tg yr^{-1} in DACCIWA and 0.18 Tg yr^{-1} in EDGAR. Except for Abuja (Nigeria) and Khartoum (Sudan), the DACCIWA emission estimates are consistently higher than the EDGAR estimates. Additionally, the two inventories disagree on the sectoral breakdown of the emission estimates with the domestic sector contributing 59% to the total emission rate in DACCIWA while EDGAR attributes 54% of total emissions to the industry sector. For 10 cities TROPOMI and DACCIWA agree within the TROPOMI uncertainty, that is the case for 9 cities

in EDGAR. For 16 cities, the TROPOMI estimates are closer to DACCIWA than to EDGAR. The largest differences between
280 TROPOMI and DACCIWA are found for Abidjan (627%) and Lagos (417%) while estimates for Cairo (2%) and Antananarivo
(9%) agree best. To explain the differences between TROPOMI and the inventories we will now focus on some specific areas.

Northern Africa

Two cities that stand out in Fig. 7 are Algiers and Casablanca. Unlike DACCIWA, EDGAR does not include any major
285 emission sources around these cities, even though they both have populations of 4.2 million. EDGAR also appears to largely
underestimate the emissions of the two Egyptian cities that were investigated, Cairo and Alexandria, while the DACCIWA
emissions for these cities agree well with our TROPOMI estimates. As we can not directly obtain the underlying emission
factors and activity data that are used in EDGAR and DACCIWA, we compare the TROPOMI CO emission rates to the
corresponding CO₂ emission rates in EDGAR as shown in Fig. 8. The CO₂ emission rates are also included in Table D2 in
290 the appendix. Cairo, Alexandria, Casablanca and Algiers clearly deviate from the other cities. Their much higher values for
 $CO_{TROPOMI}/CO_{EDGAR}$ correspond to lower values for CO/CO₂ in EDGAR, indicating that not the activity data but the CO
emission factors for these cities are underestimated in EDGAR. This is further confirmed by the higher CO/CO₂ values in
DACCIWA and the fact that the absolute CO₂ emission rates for these cities agree well between the two inventories. The
underestimate in CO may point at an overestimated combustion efficiency used in the compilation of the EDGAR emissions
295 for this region. Similar observations over Cairo were made by MacDonald et al. (2023) when comparing measured CO and
CO₂ concentrations.

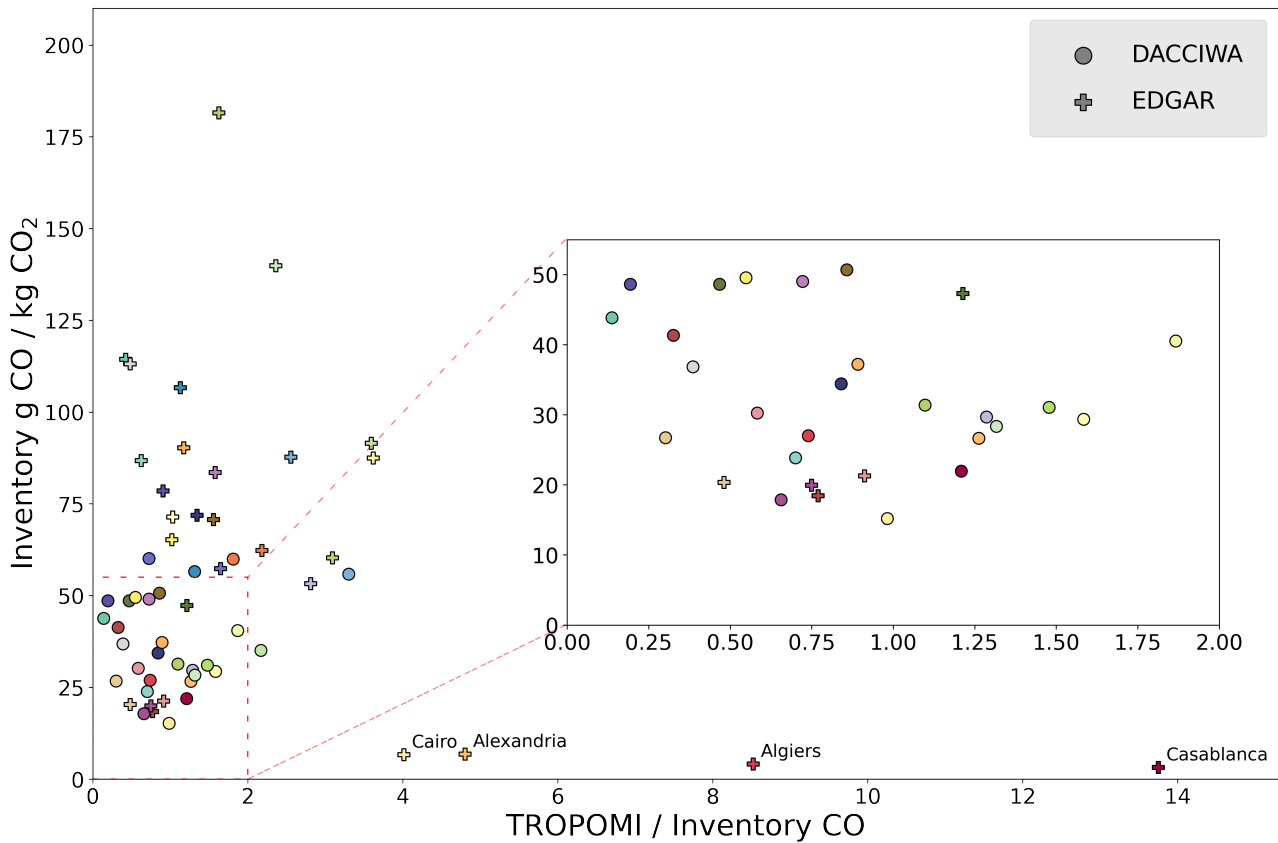


Figure 8. Comparison between inventory combustion completeness and TROPOMI to inventory CO estimates. Each marker represents a single city. The CO₂ values for both inventories include both fossil fuel and biofuel combustion emissions. As power plants hardly emit any CO per kg of emitted CO₂ due to their high combustion efficiency, the contributions of this sector are removed from the CO₂ values. Cairo, Alexandria, Algiers and Casablanca have very low CO emission rates in EDGAR compared to TROPOMI and compared to EDGAR CO₂ emissions, which indicates that EDGAR largely overestimates the combustion efficiency for these cities. The four cities in EDGAR with CO/CO₂ values around 20 are all cities in South Africa showing lower CO-emission rates than the other African cities.

South Africa

In South Africa we find closer agreement between EDGAR and TROPOMI than in northern Africa. However, the emission rates for the 4 considered cities in DACCIIWA are on average 2.4 times higher than those based on TROPOMI (Fig. 7). The emission ratios from Fig. 8 show that the South African cities stand out from the other cities in EDGAR, as they have relatively low CO/CO₂ emission ratios, suggesting high average combustion efficiencies. This does not hold for DACCIIWA, where the South African cities have CO/CO₂ emission ratios comparable to other cities. This indicates that the CO emission factors for South Africa are overestimated in DACCIIWA, and these cities have higher combustion efficiencies more in line with EDGAR.

305 Nigeria

The four investigated cities in Nigeria show varying results when comparing TROPOMI to the inventories, but the two cities that stand out are Lagos and Kano. In Lagos we estimate emissions of 0.36 (0.23-0.56) Tg yr⁻¹, that are consistent with EDGAR, but DACCIWA has emissions that are 5.2 times higher, a difference which is much larger than the uncertainty in wind data discussed in Section 2.6. For Kano, in contrast to Lagos, we observe an emission rate of 0.53 (0.41-0.62) Tg yr⁻¹ which is consistent with DACCIWA but more than twice the EDGAR estimate (0.19 Tg yr⁻¹). The CO/CO₂ ratios of the inventories agree within 50% but the differences are caused by the activity data. Figure 9 shows that the CO₂ emission rates in DACCIWA for Lagos, Kano and Ibadan are respectively 8.1, 3.8 and 3.3 times higher than in EDGAR, this data is also available in appendix in Table D2. Comparing Nigeria's national CO₂ budget there is a 24% difference between the inventories (EDGAR 530 Tg yr⁻¹ to DACCIWA 700 Tg yr⁻¹) but the larger regional discrepancies (over 700% for Lagos' CO₂ emissions) suggest differences in spatial allocation as well.

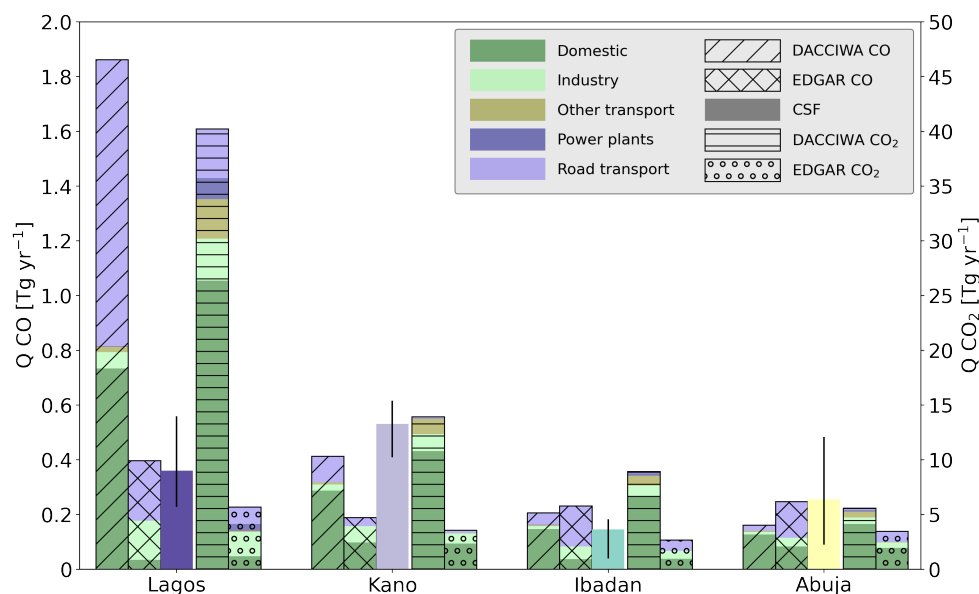


Figure 9. Comparison between the TROPOMI CO emission estimates and EDGAR and DACCIWA CO and CO₂ for four cities in Nigeria with the same color scheme as Fig. 7. The differences between the two inventories in CO₂ emission rates indicate a different spatial allocation - based on gridded activity data- of the national totals.

Côte d'Ivoire

Abidjan in Côte d'Ivoire has the largest relative discrepancy between DACCIWA (0.65 Tg yr⁻¹) and TROPOMI (0.1 (0.05-0.16) Tg yr⁻¹). In DACCIWA, the domestic sector contributes for 89% of the city's emissions and Abidjan is the city with one of the highest CO/CO₂ values of all investigated cities. In EDGAR, the domestic CO/CO₂ ratio for Abidjan is four times lower, which would indicate a four times lower emission rate. This would bring the DACCIWA emission rate much closer to

the TROPOMI observed emission, indicating that, similar to South Africa, DACCIWA may overestimate the city's domestic sector's CO emission factor.

325 **Libya**

Tripoli, the capital of Libya, stands out as its CO emissions in both inventories are almost exclusively (90+%) due to road transport. The TROPOMI estimate for this city of 0.26 (0.15-0.32) Tg yr⁻¹ is 1.7 and 2.2 times higher than DACCIWA and EDGAR respectively. The difference can be partly explained by considering the domestic and industry sectors. In both emission inventories the CO/CO₂ ratios for these sectors are four to five times lower than the mean of the other cities and two
330 to three times lower than the next lowest city (excluding Egypt, Morocco and Algeria). This implies these sectors in Tripoli have high combustion efficiency compared to other cities. However, based on the TROPOMI estimate, both inventories seem to underestimate the emission factor for Tripoli specifically for the non-road transport sectors.

Temporal emission patterns

335 With the three-year TROPOMI dataset, we can also investigate the temporal variability of emissions. Earlier studies, focusing on concentration trends rather than emission estimates, have found that CO concentrations over Cairo are lower on Fridays, which is the day off in the Islamic world (Rey-Pommier et al., 2022). This "weekend effect" has also been observed for nitrogen dioxide (NO₂) and ozone (O₃), which like CO in Cairo are dominated by transport emissions (Beirle et al., 2003; Khoder, 2009; Stavrakou et al., 2020). Combined with the fact that both emission inventories agree on road transport as the main
340 contributor to emissions in Cairo, lower CO emissions are indeed expected on Fridays when there is less commuter traffic. Figure 10 indeed shows a 32% drop in emissions on Fridays over Cairo. A similar reduction in emissions can be seen over Algiers, which can also be attributed to reduced road traffic. Similar significant patterns were not seen for the other cities that tend to have relatively lower contributions from road traffic.

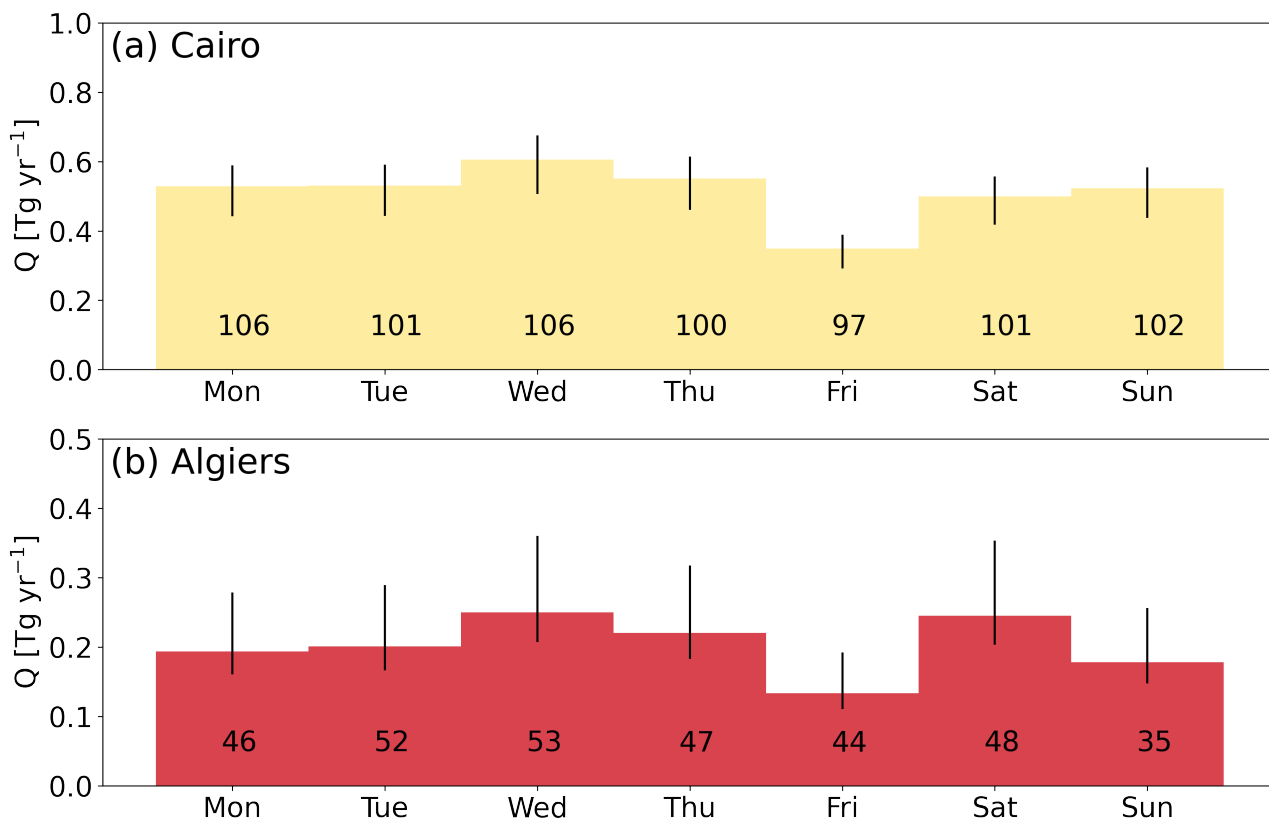


Figure 10. TROPOMI emission estimates over Cairo (panel a) and Algiers (panel b) for different days of the week averaged over 2019-2021. The numbers in the bars show the number of days the averages are based on. Both cities show lower emissions on Friday consistent with road transport being the main contributor to urban emissions and Friday being the standard day off in Islamic countries.

4 Conclusions

345 We adapted and calibrated the computationally efficient Cross-Sectional Flux (CSF) method to quantify urban carbon monoxide emission rates from major cities in Africa using TROPOMI data. We determined optimal values for the parameters of the CSF by applying the method to a full-year of simulated WRF plumes over three distinctly different African cities (Cairo, Lagos, and Bamako), such that the transects drawn best match the shape and curvature of the simulated plumes. These simulations were also used to calibrate the CSF's effective wind speed relationship for TROPOMI data. By applying the calibrated

350 CSF to the simulated data with known emission rates, we found that we can quantify urban CO emissions down to 0.1 Tg yr⁻¹ within 30% uncertainty. After calibration, we applied our CSF method to TROPOMI observations of 29 of Africa's most populated/emitting cities. We focus on Africa as there are relatively few dedicated emission inventories for the continent and large uncertainties in emission rates are expected.

355 We compared our TROPOMI-based emission estimates with the global EDGAR emission inventory and the Africa-focused
DACCIWA inventory. There are substantial differences between urban CO emissions from both inventories. We did not find
vastly better average agreement of either inventory with TROPOMI. DACCIWA is closer to the TROPOMI estimate for 16
out of 29 cities. For 10 cities, the DACCIWA and TROPOMI estimates agree within the uncertainty of the TROPOMI-based
360 estimate, but there are also cities with large significant differences of over 600%. Compared to EDGAR, we find 9 cities agree
within the uncertainty and we similarly find cities with large discrepancies.

We then evaluated our results for different regions. In northern Africa, TROPOMI observes higher emission rates than shown
in EDGAR for cities in Egypt, Algeria and Morocco. The EDGAR CO to CO₂ emission ratios for these four cities are relatively
low, implying that the mismatch with TROPOMI may originate from the emission factors, which implies that EDGAR would
365 overestimate the average combustion efficiency in these cities. In South-Africa, the TROPOMI estimates agree with EDGAR
but the DACCIWA estimates are high in comparison. EDGAR shows lower CO to CO₂ ratios and hence higher combustion
efficiency for South Africa compared to other parts of Africa, implying those ratios may be too high in DACCIWA. Similarly,
DACCIWA appears to underestimate the combustion efficiency in Abidjan (Côte d'Ivoire). For Tripoli (Libya), both invento-
ries estimate lower emission rates than the estimate based on TROPOMI. Specifically the domestic and industry sectors show
370 particularly high combustion efficiencies compared to other cities in both inventories, which can explain part of the discrepancy.

We also found some discrepancies that can be attributed to the activity data used by the inventories. We found a factor
~4 lower emissions based on TROPOMI for Lagos (Nigeria) than estimated by DACCIWA. The associated DACCIWA CO₂
emissions are eight times larger than in EDGAR which can be partly a difference in activity data but also suggests a mismatch
375 related to the spatial distribution of emissions. For Kano (Nigeria) DACCIWA estimates CO₂ emissions that are four times
larger than EDGAR. Here, the TROPOMI estimate agrees better with DACCIWA and the activity rate, which corresponds to
CO₂ emission, in EDGAR seems to be an underestimation.

The large TROPOMI data volume enables identification of temporal emission patterns. Over Cairo and Algiers we find
380 significantly lower emission rates on Fridays - the local rest day - compared to other days of the week. The ability to recognise
such patterns builds confidence and shows the strength of TROPOMI's daily global coverage combined with a computationally
efficient method like the CSF method developed here.

Code and data availability. TROPOMI CO data are publicly available at <https://ftp.sron.nl/open-access-data-2/TROPOMI/tropomi/co/>.
GEOS-FP wind data can be downloaded at https://gmao.gsfc.nasa.gov/GMAO_products/. ERA5 wind data are available at
385 <https://cds.climate.copernicus.eu>. WRF-Chem code is available at <https://github.com/wrf-model/WRF/releases>, in this work version
4.1.5 was used. EDGAR v5 CO data is available at https://edgar.jrc.ec.europa.eu/dataset_ap50. EDGAR v5 CO₂ data is available at
https://edgar.jrc.ec.europa.eu/dataset_ghg50. DACCIWA CO and CO₂ data are available at <https://eccad.aeris-data.fr/> at the Emissions of

atmospheric Compounds and Compilation of Ancillary Data (ECCAD) system after creation of a login account. GPW v4 gridded population density is available at <https://sedac.ciesin.columbia.edu/data/set/gpw-v4-population-density-rev11>. Open fire emissions from GFAS are available at <https://atmosphere.copernicus.eu/global-fire-emissions>.

Appendix A: TROPOMI-based CO uncertainty

The CSF method as applied in this work has many parameters that were calibrated on simulated plumes. In order to determine the uncertainty of the estimated emission rate we have created an ensemble of emission estimates by varying these parameters. The members of the ensemble and the ranges over which they were varied are shown in Table A1. The wind databases are the three wind products as described in section 2.5 and are responsible for a mean uncertainty of -19% and + 12%. The standard deviation threshold for spline pixels is the number of standard deviations a pixel has to be above the background concentration in order to be considered part of the plume. Pixels identified as part of the plume are only used for fitting the spline shape. The number of cross-sections is the total number of transects drawn perpendicular to the direction of the plume. As described in section 2.4 not all transects are taken into account in the quantification. The number of cross-sections is mostly a measure for the line density, or the distance between consecutive transects as the transects are evenly spaced over the full length of the spline. The minimum pixel coverage is the minimum fraction of a line that needs to be covered by pixels and is a balance between retaining enough days with a valid estimate and not underestimating emissions. A lower limit of 50% coverage per line will retain a lot of lines, and thus, more days with an estimate, however, the cross-sections will potentially miss large parts of the plume. The distance of the background box is important as one would like the box to be close to the city to get a good representation of the local background. However, it must not overlap with any urban emissions to have a clean background. As explained in section 2.4 the transects start upwind of the city center to capture the full city plume. We vary the distance between the first transect and the city center for our uncertainty estimate. As a last member of the ensemble we use the spread in emission estimates of the individual transects. We include the means of the transects with the lowest and highest 50% emission rates in the ensemble.

Table A1. Variables used in the uncertainty analysis and the ranges over which they were varied. The resulting ensemble spreads are reported as uncertainty.

Parameter	Domain	Default
Wind database	GEOS-fp 10m, ERA5 10m & GEOS-fp pbl	GEOS-fp 10m
Standard deviation threshold for spline pixels	{1.2 - 2.4}	1.8
Number of transects	{15 - 25}	20
Minimum pixel coverage per transect	{50% - 90%}	70%
Distance of background box	{0.2° - 0.4°}	0.3°
Varying upwind distance first transect	{0° - 0.2°}	0.1°
Transects used for estimate	{Lowest 50% - highest 50%}	All

410 **Appendix B: TROPOMI data filtering**

Although the CSF has been shown to reproduce simulated emission rates (Fig. 5), it can not be applied to every single overpass of TROPOMI. For example, days with a lot of missing pixels in the TROPOMI data can lead to underestimation of the emission rates. To prevent a positive bias, it is important to not only use days with strong, clearly visible plumes. With the filters chosen in this work, over 400 days are accepted as quantifiable over the 3-year period studied for most non-coastal cities and coastal cities
415 with predominantly inland winds (e.g. Cairo has 713 estimates, Johannesburg 427 and Khartoum 570). With TROPOMI's daily overpasses this means that we estimate emissions on roughly 40% of all days. Regions with fewer estimates tend to be coastal. For example, we only have estimates for 160 days over Lagos and 113 for Dakar because of limited TROPOMI coverage over water. An additional reason for a small number of valid estimates lies in the occurrence of open fires; for example, 224 orbits (42%) are removed from our estimate over Lusaka (Zambia) due to fires within 1.5° of the city center and stronger fires within
420 0.75° .

The filters employed are shown in Table B1.

Table B1. Filtering applied to the data to ensure correct application of the CSF method.

Description	Value
Per pixel	
Quality flag TROPOMI (land)	≥ 0.7
Quality flag TROPOMI (water)	$= 0.7$
Per transect	
Misalignment between plume and wind direction	$< 45^\circ$
Minimum pixel coverage	$> 70\%$
Per plume	
Downwind coverage in a $0.3^\circ \times 0.8^\circ$ box	$> 60\%$
Effective wind speed	$> 2\text{m s}^{-1}$
Maximum concentration outside the plume	$< 200\text{ppb}$ (1.5° radius)
Number of transects used for the estimate	> 3
Fraction estimate transect 8-20 to transect 3-7. High emission estimates of the far away lines tend to indicate interference of different sources	$< 2.5\text{x}$
Second derivative of spline scaled to 1° pixel size. This represents the dimensionless curvature of the fitted spline.	< 0.05
Mismatch between the first transect and the starting pixel of the plume	$< 0.35^\circ$
Fire emission from Global Fire Assimilation System (GFAS) database, (Kaiser et al., 2012)	$< 23\text{Mg hr}^{-1}$ (0.75° radius)
	$< 57\text{Mg hr}^{-1}$ (1.5° radius)

Appendix C: CSF estimates

Table C1. Emission estimates for the studied African cities by applying the CSF method to the TROPOMI CO product (2019-2021) as well as the corresponding emission rates according to the DACCIWA (2015) and EDGAR (2015) inventories. All emission rates are in Tg yr⁻¹. Population is taken from the Center for International Earth Science Information Network CIESIN (2018)

City	Country	Population	TROPOMI estimate	Lower limit	Upper limit	DACCIWA	EDGAR
Algiers	Algeria	4.2M	0.213	0.177	0.307	0.288	0.025
Luanda	Angola	5.1M	0.268	0.166	0.413	0.489	0.263
Ouagadougou	Burkina Faso	2.8M	0.273	0.243	0.335	0.126	0.076
Kinshasa	Congo	7.4M	0.628	0.49	0.894	0.477	0.266
Abidjan	Côte d'Ivoire	7.1M	0.09	0.054	0.157	0.654	0.266
Alexandria	Egypt	3.3M	0.197	0.164	0.257	0.156	0.041
Cairo	Egypt	16.7M	0.546	0.456	0.608	0.556	0.136
Addis Ababa	Ethiopia	4.4M	0.239	0.2	0.295	0.268	0.204
Accra	Ghana	3.5M	0.194	0.176	0.318	0.148	0.172
Nairobi	Kenya	4.8M	0.17	0.153	0.247	0.441	0.353
Tripoli	Libya	1.2M	0.264	0.151	0.32	0.146	0.121
Antananarivo	Madagascar	3.0M	0.135	0.104	0.156	0.123	0.083
Casablanca	Morocco	4.2M	0.11	0.078	0.171	0.091	0.008
Maputo	Mozambique	2.5M	0.151	0.125	0.199	0.176	0.097
Abuja	Nigeria	2.6M	0.255	0.089	0.483	0.161	0.247
Ibadan	Nigeria	2.2M	0.145	0.039	0.183	0.207	0.232
Kano	Nigeria	5.3M	0.531	0.409	0.616	0.413	0.189
Lagos	Nigeria	10.9M	0.36	0.227	0.558	1.862	0.397
Dakar	Senegal	4.2M	0.293	0.154	0.381	0.157	0.081
Cape Town	South Africa	4.1M	0.189	0.155	0.252	0.288	0.252
Durban	South Africa	3.3M	0.157	0.132	0.233	0.482	0.204
Johannesburg	South Africa	9.1M	0.232	0.184	0.291	0.77	0.482
Pretoria	South Africa	6.4M	0.197	0.139	0.209	0.338	0.216
Khartoum	Sudan	3.1M	0.396	0.336	0.432	0.12	0.155
Dar es Salaam	Tanzania	5.4M	0.286	0.241	0.496	0.396	0.181
Kampala	Uganda	4.3M	0.362	0.279	0.431	0.245	0.117
Lusaka	Zambia	2.4M	0.226	0.108	0.23	0.269	0.168
Bulawayo	Zimbabwe	0.8M	0.094	0.043	0.111	0.13	0.057
Harare	Zimbabwe	2.7M	0.142	0.076	0.157	0.304	0.117

Appendix D: Inventory emission rates

Table D1. Sectoral breakdown of the CO emission rates for the studied African cities according to the DACCIWA and EDGAR inventory. All emission rates are in Gg yr⁻¹.

City	DACCIWA					EDGAR				
	Domestic	Industry	Other	Power	Road	Domestic	Industry	Other	Power	Road
Algiers	0.3	33.7	0.6	0.9	252.2	1.3	2.7	0.2	0.4	20.0
Luanda	326.1	19.4	12.0	3.4	129.0	25.4	174.0	0.3	6.5	56.1
Ouagadougou	90.3	8.0	0.0	0.8	26.4	13.2	39.8	0.1	2.8	20.0
Kinshasa	411.1	25.4	0.0	0.0	40.4	20.1	222.3	0.1	0.1	23.5
Abidjan	579.9	21.4	20.4	7.0	25.5	10.8	183.6	0.3	1.6	15.7
Alexandria	24.3	46.7	20.7	4.4	59.6	3.4	1.8	0.3	15.8	19.9
Cairo	75.8	145.5	64.6	26.1	244.3	3.5	6.4	1.1	9.3	115.9
Addis Ababa	245.3	8.5	1.8	0.0	12.3	55.3	134.8	0.2	0.0	13.9
Accra	108.2	3.2	3.9	0.0	32.4	5.0	90.2	1.0	0.5	43.2
Nairobi	370.8	11.7	0.2	2.3	59.1	23.5	266.9	0.3	2.1	60.0
Tripoli	3.1	4.2	0.0	6.6	131.6	1.9	0.1	0.1	0.0	118.9
Antananarivo	68.8	17.0	30.2	0.1	6.5	7.5	68.3	0.1	0.0	7.3
Casablanca	29.2	37.1	4.6	0.5	20.0	1.5	3.9	0.0	0.2	2.2
Maputo	153.1	5.3	0.1	3.1	14.7	21.0	62.7	0.0	0.3	13.1
Abuja	126.2	10.2	3.6	0.1	21.0	82.6	32.1	0.0	0.0	131.9
Ibadan	145.8	11.7	4.1	0.3	44.7	36.2	46.4	0.0	0.0	149.1
Kano	286.3	23.0	8.1	0.0	95.5	96.9	60.6	0.0	0.0	31.0
Lagos	733.5	59.0	20.8	2.4	1046.3	33.2	143.5	0.1	0.7	219.8
Dakar	123.3	17.4	0.1	6.1	10.0	2.7	62.7	0.2	3.2	12.6
Cape Town	139.9	50.7	5.7	1.4	90.5	56.9	170.1	0.7	0.0	23.8
Durban	183.8	66.6	7.5	0.0	224.4	53.3	139.4	0.3	0.0	10.9
Johannesburg	355.8	128.9	14.5	22.4	247.9	108.8	317.5	0.1	0.0	55.1
Pretoria	180.4	65.4	7.4	0.9	83.4	57.3	129.0	0.1	0.4	29.3
Khartoum	77.2	10.8	3.8	1.6	26.2	16.3	90.0	0.1	5.6	42.6
Dar es Salaam	354.2	15.0	9.6	1.7	15.8	15.1	138.1	0.1	10.4	17.7
Kampala	198.8	10.5	0.1	1.2	34.2	65.6	42.5	0.0	0.0	8.8
Lusaka	239.0	11.7	1.2	0.0	16.7	31.7	126.3	0.0	0.0	9.8
Bulawayo	79.1	4.3	3.5	0.2	42.4	21.5	2.0	0.0	0.2	33.0
Harare	248.2	13.3	10.8	0.2	31.5	69.8	4.5	0.1	0.4	41.7

Table D2. Sectoral breakdown of the CO₂ emission rates for the studied African cities according to the DACCIWA and EDGAR inventory. All emission rates are in Tg yr⁻¹.

City	DACCIWA					EDGAR				
	Domestic	Industry	Other	Power	Road	Domestic	Industry	Other	Power	Road
Algiers	2.0526	2.3547	0.3026	0.7006	5.9577	2.2141	2.3123	0.1231	0.0446	1.4081
Luanda	4.8853	1.4104	2.2871	0.0	1.2816	1.8331	1.4681	0.1276	1.6398	0.6036
Ouagadougou	2.1571	0.5752	0.701	0.4603	0.1553	0.3469	0.2106	0.1102	0.4868	0.1633
Kinshasa	9.1058	2.4811	4.9228	0.0	0.3205	0.4757	1.1183	0.0897	0.0245	0.2184
Abidjan	5.6168	1.2141	7.6905	4.1632	0.3916	0.3994	1.1479	0.1115	1.1956	0.1948
Alexandria	1.7076	3.1371	0.3293	0.0	0.6755	0.8145	4.1211	0.0639	23.3842	0.9927
Cairo	9.7364	17.8872	3.5146	17.7906	5.4195	3.1328	11.4975	0.7648	13.7938	4.9941
Addis Ababa	5.2452	0.4889	1.0817	0.0	0.3851	1.2733	0.5438	0.2141	0.0005	0.2297
Accra	1.172	0.3122	0.7803	0.6024	0.3533	0.231	0.6034	0.4059	0.3562	0.3729
Nairobi	5.3718	0.9503	4.7633	0.429	0.8817	0.6676	1.6001	0.2878	0.4414	0.5657
Tripoli	0.3477	0.8711	0.3068	7.4683	0.9077	0.252	0.6718	0.022	0.0	0.9969
Antananarivo	1.7812	0.9208	1.0695	0.0798	0.1474	0.205	0.144	0.046	0.0195	0.0622
Casablanca	1.4466	2.1949	0.0158	4.0377	0.4885	0.7635	1.2911	0.0514	1.2078	0.3914
Maputo	2.3379	0.2568	0.6895	0.0	0.1877	0.4888	0.7266	0.0355	0.0795	0.1202
Abuja	4.1275	0.6013	0.5368	0.1024	0.2142	1.9347	0.4962	0.0413	0.0	0.9858
Ibadan	6.678	0.9728	0.8684	0.2404	0.1572	0.944	0.7214	0.0006	0.0006	1.0063
Kano	10.773	1.5693	1.401	0.0311	0.1601	2.3444	0.9346	0.0206	0.0008	0.2502
Lagos	26.3529	3.8388	3.5988	1.9078	4.5136	1.1902	2.2007	0.0889	0.6383	1.5765
Dakar	1.5405	1.01	0.9282	1.9502	0.3954	0.1379	0.4475	0.1115	1.2811	0.2292
Cape Town	3.1332	10.8663	1.2788	1.0896	0.823	1.759	9.0686	0.247	0.0009	1.5695
Durban	2.3207	8.0484	0.543	0.1116	0.7436	1.5599	8.5055	0.0943	0.1335	0.9045
Johannesburg	5.5358	19.1984	2.2775	23.3095	1.7699	3.1791	17.0838	0.1623	4.9948	3.2495
Pretoria	2.2686	7.8676	0.5308	3.0576	0.5046	1.6006	6.6839	0.1254	3.4728	1.7434
Khartoum	0.6315	0.9569	0.3614	0.1673	0.198	0.5049	0.6555	0.101	1.4848	0.5062
Dar es Salaam	5.582	1.0448	1.4222	1.073	0.0259	0.3639	1.5205	0.0908	1.3461	0.1923
Kampala	5.0997	0.8774	1.6523	0.5054	0.2586	1.4997	0.329	0.009	0.0	0.1042
Lusaka	3.8478	1.6338	2.2394	0.0	0.0951	0.7361	1.4979	0.0327	0.0	0.0699
Bulawayo	1.8723	0.1759	0.0062	0.2274	0.1068	0.5132	0.2041	0.0134	0.2151	0.263
Harare	5.522	0.5187	0.0183	0.2416	0.193	1.6508	0.3461	0.059	0.4645	0.4179

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