

Reversible and irreversible gas-particle partitioning of dicarbonyl compounds observed in the real atmosphere

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Abstract. Glyoxal and methylglyoxal are vital carbonyl compounds in the atmosphere and play substantial roles in radical cycling and ozone formation. The partitioning process of glyoxal and methylglyoxal between the gas and particle phase via reversible and irreversible pathways could efficiently contribute to secondary organic aerosol (SOA) formation. However, the relative importance of two partitioning pathways still remains elusive, especially in the real atmosphere. In this study, we launched five field observations in different seasons and simultaneously measured glyoxal and methylglyoxal in the gas and particle phase. The field-measured gas-particle partitioning coefficients were 5–7 magnitudes higher than the theoretical ones, indicating the significant roles of reversible and irreversible pathways in the partitioning process. The particulate concentration of dicarbonyls and product distribution via the two pathways were further investigated using a box model coupled with the corresponding kinetic mechanisms. We recommended the irreversible reactive uptake coefficient γ for glyoxal and methylglyoxal in different seasons in the real atmosphere, and the average value of 8.0×10^{-3} for glyoxal and 2.0×10^{-3} for methylglyoxal best represented the loss of gaseous dicarbonyls by irreversible gas-particle partitioning processes. Compared to the reversible pathways, the irreversible pathways played a dominant role, with a proportion of more than 90% in the gas-particle partitioning process in the real atmosphere and the proportion was significantly influenced by relative humidity and inorganic components in aerosols. However, the reversible pathways were also substantial, especially in winter, with a proportion of more than 10%. ~~The partitioning processes of dicarbonyls in reversible and irreversible pathways jointly contributed to more than 25% of SOA formation in the real atmosphere. These two pathways of dicarbonyls jointly contributed to more than 25% of SOAs in the real atmosphere.~~ To our knowledge, this study ~~was~~is the first to systemically examine both reversible and irreversible pathways in the ambient atmosphere, ~~strives~~to narrow the gap between model simulations and field-measured gas-particle partitioning coefficients, and ~~revealed~~reveal the importance of gas-particle processes for dicarbonyls in SOA formation.

1 Introduction

Glyoxal and methylglyoxal, the simplest α -dicarbonyls, are recognized as being of great importance in atmospheric chemistry

29 due to their unique physicochemical properties. The α -dicarbonyl functionality leads to higher water solubility and reactivity
30 of dicarbonyls than expected, as the α -dicarbonyl functionality is hydrophilic and contributes to hydrate formation. The hydrate
31 form of carbonyls is less volatile and more water-soluble than the unhydrated form (EPA., 2012), owing to the strong effect of
32 the two hydrogen-bonding groups in the hydrated form (Elrod et al., 2021). Moreover, hydrates can easily participate in
33 continuous radical reactions with higher reactivity by H-abstraction to form higher-molecular-weight oligomers (Michailoudi
34 et al., 2021). ~~The α -dicarbonyl functionality increases their water solubility and reactivity more than expected.~~ The traditional
35 opinion is that methylglyoxal is less reactive compared to glyoxal due to its unreactive methyl substitution, while a very recent
36 study noted that methylglyoxal could be more reactive under an atmospheric-relevant concentration (Li et al., 2021). Overall,
37 both of them play crucial roles in radiation balance, air quality, brown carbon formation, and SOA formation (Laskin et al.,
38 2015; Qiu et al., 2020). Moreover, as major carcinogenic and genotoxic compounds, dicarbonyls can cause serious damage to
39 human health. They have relatively limited primary sources, except for biomass burning and biofuel combustion (Zarzana et
40 al., 2017; Zarzana et al., 2018), compared to secondary formation that occurs with photooxidation of both biogenic volatile
41 organic compounds (VOCs), such as isoprene, and anthropogenic VOCs, such as aromatic hydrocarbons (Lv et al., 2019).
42 Considering the atmospheric sink, glyoxal and methylglyoxal can be lost in the gas phase by self-photolysis, oxidation by
43 active radicals (such as OH radicals, NO₃ radicals) and wet/dry deposition; photochemical reactions, oxidation by OH radicals,
44 and dry deposition; however, there is still a missing sink for the two dicarbonyls (Volkamer et al., 2007), that's the gas-particle
45 partitioning process, which would be fully discussed in this study.
46 Gas-particle partitioning was recently found to be the most important removal pathway for both glyoxal and methylglyoxal,
47 especially in regions like Beijing with high particulate matter (PM) pollution that provides sufficient aerosol surface area.
48 Although ~~with they have~~ relatively high vapor pressure, glyoxal and methylglyoxal can efficiently partition into the particle
49 phase due to their α -dicarbonyl functionality. The surface-adsorbed dicarbonyls could alter the properties of the particle's
50 surfaces and the organic surface films could act as a kinetic barrier to gas-aerosol mass transport and thereby influence particle
51 equilibration and water/gas uptake (Donaldson and Vaida, 2006). Upon physical adsorption, besides desorption or reaction at
52 the surface, dicarbonyls could undergo solvation and incorporation into the bulk liquid, and then they could go through
53 diffusion and chemical reactions in the bulk phase. The product may return into surfaces and gas phase, or stay in the bulk
54 phase (Donaldson, 2006 #252)(Paul et al., 2011). Moreover, chemical reactions occurred at the surface or in the bulk phase
55 could in turn accelerate the physical adsorption and greatly contribute to the formation and growth of atmospheric particulate
56 matter. Whereas, as it is difficult to distinguish the surface reactions and bulk reactions in field observations, we regard both
57 of them as particle-phase reactions in this study. ~~Upon physical adsorption to the particle phase, they can undergo various~~
58 ~~chemical reactions and subsequently form larger molecular weight products retained in the condensed phase. Moreover,~~
59 ~~chemical reactions can in turn accelerate the physical adsorption and greatly contribute to the formation and growth of~~

60 ~~atmospheric particulate matter~~. The chemical reactions occurring in the gas-particle partitioning processes can be divided into
61 reversible pathways, including reversible hydration and self-oligomerization, and irreversible pathways, which can be driven
62 by oxidative compounds. These processes can also efficiently explain observed aerosols properties – including relatively high
63 oxygenation levels, compositions ~~like-such as~~ organic acids and oligomers, and higher light absorption – that cannot be
64 explained by traditional absorptive models of gas-particle partitioning (Pankow, 1994; Pankow and James, 1994; Odum et al.,
65 1996).

66 Many laboratory and model studies have made a great effort to investigate the reversible and irreversible pathways of
67 dicarbonyls to further understand their gas-particle partitioning mechanisms and reveal their contribution to SOA formation.
68 Fu et al. (2008) found that the modeled SOA concentrations were largely increased when accounting for irreversible uptake of
69 dicarbonyls in the GEOS-Chem model. Considering the surface-controlled reactive uptake of dicarbonyls into the CMAQ
70 model, the aerosol uptake of dicarbonyls accounted for more than 45% of total SOA in the eastern US (Ying et al., 2015);
71 similarly, the contribution of glyoxal and methylglyoxal to SOA formation in China was 14% to 25% and 23% to 28%,
72 respectively (Hu et al., 2017). Although reversible and irreversible pathways of dicarbonyls have been separately investigated
73 in previous studies, solely incorporating just one pathway into models could lead to a large discrepancy between model results
74 and observational data, highlighting the importance of comprehensively considering both reversible and irreversible pathways
75 when quantifying the gas-particle partitioning process of dicarbonyls (Li et al., 2014; Hu et al., 2017; Ling et al., 2020).

76 Despite increasing interest in dicarbonyls and their gas-particle partitioning processes, the detailed chemical mechanisms of
77 two partitioning pathways remain poorly understood. First, previous studies have exposed seed particles to high concentration
78 levels of dicarbonyl vapors, from hundred ppb to ppm levels, or used bulk samples; thus, their applicability to the real
79 atmosphere requires further validation. Second, prior studies always used one constant coefficient to present all heterogeneous
80 processes occurring on the aerosol, which neglects the influencing factors in real atmospheric partitioning processes. Further
81 studies have shown that the two pathways in the gas-particle partitioning process for glyoxal and methylglyoxal are rather
82 complex, and their relative contribution to the partitioning process can be influenced by many factors such as relative humidity
83 (Curry et al., 2018; Shen et al., 2018), particle acidity (Liggio et al., 2005b; Shi et al., 2020), and particle organic/inorganic
84 components (Kampf et al., 2013). However, there persist controversies in the specific partitioning mechanisms of glyoxal and
85 methylglyoxal, especially conflicting views on their role in SOA formation, which urgently warrants further investigation.

86 In this study, five field observations were launched over urban Beijing in four seasons, and glyoxal and methylglyoxal in the
87 gas and particle phase were simultaneously measured. Beijing, as the political center of China, is the most prosperous city with
88 numerous key environmental issues. Chen et al. (2021) found that the average concentration of dicarbonyls in Beijing is lowest
89 among ~~the~~ key regions ~~that have with~~ relatively higher PM_{2.5} concentrations, indicating there is a more efficient partitioning
90 process of dicarbonyls. Thus, it is more environmentally significant to discuss the gas-particle partitioning processes in urban

91 Beijing. These processes are divided into ~~two reversible~~ pathways: ~~reversible~~ and irreversible pathways, which is based on the
92 reversibility of chemical reaction of dicarbonyls occurred on condensed phase (Galloway et al., 2008; Ervens and Volkamer,
93 2010; Kampf et al., 2013; Ling et al., 2020). On the basis of field-measured data, we could estimate the product distribution,
94 main influencing factors, and relative importance of the two gas-particle partitioning pathways for glyoxal and methylglyoxal
95 in the real atmosphere.

96 2 MATERIALS AND METHOD

97 2.1 Field sampling ~~and analysis~~

98 We performed field observations on the roof of a six-story teaching building (26 m above the ground) on the Peking University
99 campus (39.992°N, 116.304°E) in northwest urban Beijing. The field observations in this study were launched during four
100 different seasons from 2019 to 2021.

101 Gaseous carbonyls were collected by adsorption reactions in a 2,4-dinitrophenyl hydrazine (DNPH) cartridge (Sep-Pak; Waters
102 Corporation). The air samples were first passed through an ozone scrubber (Sep-Pak; Waters Corporation) to eliminate
103 interference by ozone and then trapped in the DNPH cartridge. To prevent deliquescence of the potassium iodide in the ozone
104 scrubber, the air samples were mixed with ultrapure nitrogen before pumped into the sampling tubing. Air samples were
105 continuously collected every 3 h in daytime and 9 h in nighttime. The total flow rate was 0.8 L·min⁻¹.

106 Particulate carbonyls were collected by a four-channel ambient particles sampler (TH-16A, Wuhan Tianhong) with Teflon
107 filter and quartz filter (47 mm, Whatman). The Teflon filter was used to measure the mass concentration of collected PM_{2.5}
108 and water-soluble inorganic compounds (Na⁺, NH₄⁺, K⁺, Mg²⁺, Ca²⁺, Cl⁻, NO₃⁻, and SO₄²⁻). ~~and~~ The quartz filters were used
109 for carbonyl analysis. The flow rate was set at 16.7 L·min⁻¹ and particle samples were continuously collected every 12 h daily.

110 Detailed information about field sampling and analysis were provided in the previous studies (Rao et al., 2016; Qian et al.,
111 2019). To estimate the positive artifacts by adsorption of gas-phase dicarbonyls onto the filter (Hart and Pankow, 1994; Mader
112 and Pankow, 2001; Liggio, 2004; Odabasi and Seyfioglu, 2005), throughout our previous field observations, we placed a
113 backup quartz filter after the particle sampling quartz filter using an independent filter holder. The sampling filters would
114 collect the particles and adsorbed gaseous dicarbonyls, while the backup filter would only collect gaseous dicarbonyls. And
115 the ratio of measured dicarbonyls in second filter to that in the first were lower than 20%, which was equal to the previous
116 study (Shen et al., 2018). And the particulate concentrations of dicarbonyls used in this study were already corrected by the
117 possible adsorption artifacts.

118 ~~(Hart and Pankow, 1994; Mader and Pankow, 2001; Liggio, 2004; Odabasi and Seyfioglu, 2005)(Shen et al., 2018)~~

119 The meteorological station was co-located at our sampling site and provided meteorological parameters. ~~Common~~ Common

120 trace gases, ~~such as~~ NO/NO₂, SO₂, CO, and O₃, were detected online by Thermo 42i, 43i, 48i, and 49i analyzers,
121 respectively. A TEOM 1400A analyzer was applied to measure the mass concentrations of PM_{2.5} and PM₁₀, the results of which
122 were consistent with the PM_{2.5} weighing results (Fig. S1). The time resolution for all of the above data was 1 min. Detailed
123 information about these five observations is shown in Table S1.

124 2.2 Sample extraction and analysis

125 The gaseous carbonyl samples were eluted with acetonitrile (HPLC/GC-MS grade) at a flow rate of less than 3 mL/min (higher
126 flow rates can result in reduced recovery). And the particulate carbonyl samples on quartz filter were eluted with acidic DNPH
127 solutions in the flask and then were shaken for 3 h at 4 °C with a rotation rate of 180 rpm in an oscillator (Shanghai Zhicheng
128 ZWY 103D). The derived solutions were placed in darkness for 12-24 h to ensure complete derivatization, and then they were
129 analyzed by high-performance liquid chromatography-ultraviolet (HPLC-UV) for separation and detection. Carbonyls were
130 separated effectively with each other (Fig. S2) in this method. They were calibrated using a mixing standard solution with a
131 concentration range of 0.1–10 μM, and the linearity was indicated by a correlation of determination (r^2) of at least 0.999. The
132 detailed analysis method was presented in the previous study (Wang et al., 2009).

133 The Teflon samples were also extracted by deionized water using an ultrasonic bath for 30 min at room temperature. The
134 extracted solutions were analyzed by ion chromatography (IC Integriion and Dionex ICS 2000, USA) to measure the water-
135 soluble inorganic compounds (Na⁺, NH₄⁺, K⁺, Mg²⁺, Ca²⁺, Cl⁻, NO₃⁻, and SO₄²⁻) and low-molecular-weight organic acids
136 (formate, acetate, and oxalate) in aerosols.

137 2.3 Quality assurance / quality control

138 As carbonyl compounds are ubiquitous in environmental media, following measurements were conducted during the sample
139 collection, pretreatment and analysis to ensure the accuracy of results: (1) Before sampling, flow calibration and airtightness
140 tests of sampling devices were conducted, and flow difference were less than 10%; (2) After sampling, the gas-phase samples
141 were resealed by its end cap and plug, and stored in the provided pouch under cool environment (<4°C); the particle-phase
142 samples were stored in the sealed boxes wrapped by pre-baked aluminum foils under freezing environment (<-18°C), both
143 gas-phase and particle-phase samples were extracted and analyzed within a week; (3) The extraction processes were conducted
144 in fume hood with glassware, which was rinsed with acetonitrile for at least three times; (4) A calibration run was performed
145 each day to determine the response factor of the detector and recalibration was performed if the relative deviation of the RF
146 was beyond 5%.

147 Blank samples were collected every three days and then were stored and extracted by the same procedure as that for ambient
148 samples. The blank gas-phase samples were collected by placing blank DNPH cartridges near the gas inlet for the same

149 duration without artificial pumping. And the blank particle-phase samples were collected by placing blank quartz filter on the
 150 PM_{2.5} inlet with flow rate of 0 L/min. All data used in this study were all calibrated by blanks.
 151 The limit of detection (LOD) of two methods was 50 pptv for gaseous carbonyls and 1 ng·m⁻³ for particulate carbonyls, which
 152 is similar to the previous literature (Shen et al., 2018). Sample amount to limit of detection ratios were significantly higher
 153 than 1.0 for both gas- and particle-phase samples, indicating that the sensitivity of the methods was sufficient to analyze the
 154 samples.
 155 Additional field-sampling were launched to estimate the sampling efficiency during the collection. Two blank DNPH cartridges
 156 were connected in tandem to assess the sampling efficiency of gas-phase carbonyls. The sampling efficiency of the cartridges
 157 were the ratios of dicarbonyl concentrations in the first cartridge to the total concentrations in the two cartridges and the results
 158 were more than 95% for both glyoxal and methylglyoxal. Similarly, a backup Teflon filter were placed after the particle
 159 sampling Teflon filter using an independent filter holder to estimate the particle collection efficiency. Both Teflon filters were
 160 weighed by a semimicro balance (Sartorius, Germany) to obtain the mass concentration of collected particles. The mass
 161 concentrations of particles collected on the backup filter were closed to zero, indicating that the sampling efficiency of particle
 162 were more than 99%.
 163 Moreover, recovery tests were also conducted using two methods - adding standard solution and repeated extraction. We added
 164 the standard solution at three spiked levels of 0.025, 0.25 and 2.5 µg (namely 50 µL of 0.5 µg·mL⁻¹, 5 µg·mL⁻¹ and 50 µg·mL⁻¹
 165 1 analytical standards) into blank DNPH cartridges and blank quartz filter to determine the carbonyl lost during the extraction
 166 and derivation. And then the cartridges and filters were extracted in the same way as the ambient samples. Each group were
 167 set with five parallel. The recoveries were ranged from 88% to 96% for gas-phase method and ranged from 85% to 96% for
 168 particle-phase method. Moreover, we also estimated the recovery efficiencies by repeated extraction and the recoveries were
 169 the ratios of dicarbonyl concentrations in the first extraction to the total concentrations in the two extractions. The results
 170 ranged from 92.8% to 99.9% for gas-phase method and ranged from 90% to 99.9% for particle-phase method.

171 **2.3 Estimation of effective partitioning coefficient**

172 To estimate the effective partitioning process of gas-phase carbonyls to the particle phase, we could use Pankow's absorptive
 173 partitioning theory for the gas-organic phase (Eqs. (1), (2)) (Odum et al., 1996) and Henry's law for the gas-liquid phase (Eq.
 174 (3)):

$$175 K_p^f = \frac{C_p}{C_g \times TSP} \quad (1)$$

$$176 K_p^t = \frac{RTf_{om}}{10^6 MW_{OM} C_{PL}^0} \quad (2)$$

$$177 \text{eff } K_H = 10^3 \frac{c_p}{c_g \times M \times ALWC / \rho_{\text{water}}} \quad (3)$$

178 In Eq. (1), K_p^f ($\text{m}^3 \cdot \mu\text{g}^{-1}$) is the field-measured gas-particle partitioning coefficient; C_p ($\mu\text{g} \cdot \text{m}^{-3}$) is the concentrations of
 179 dicarbonyls in the particle phase, which is derived from the analysis of extracts, including monomers and their reversibly
 180 formed products (the product distribution is discussed in Section 3.2); and C_g ($\mu\text{g} \cdot \text{m}^{-3}$) ~~are~~ is the concentrations of dicarbonyls
 181 in the ~~particle and~~ gas phase, ~~respectively~~; and TSP ($\mu\text{g} \cdot \text{m}^{-3}$) is the mass concentration of suspended particles (mass
 182 concentrations of PM_{2.5} were used in this study). In Eq. (2), K_p^t ($\text{m}^3 \cdot \mu\text{g}^{-1}$) are the theoretical gas-particle partitioning
 183 coefficients determined by Pankow's absorptive model, f_{om} is the absorbing fraction of total particulate mass, MW_{OM} ($\text{g} \cdot \text{mol}^{-1}$)
 184 is the mean molecular weight of the organic phase, and ζ is the activity coefficient of target compounds. In the estimation
 185 of K_p^t in this study, f_{om} and ζ are unity and $MW_{OM} = 200 \text{ g} \cdot \text{mol}^{-1}$, as used in previous studies (Healy et al., 2008; Williams et
 186 al., 2010; Xie et al., 2014; Shen et al., 2018), and p_L^0 (Pa) is the supercooled vapor pressure of compounds as a pure liquid at
 187 temperature T , which is calculated by the extended aerosol inorganic model (E-AIM,
 188 http://www.aim.env.uea.ac.uk/aim/ddbst/pcalc_main.php) (Clegg et al., 1998). The possible uncertainty in K_p^t calculation were
 189 fully discussed in Supporting Information. In Eq. (3), $\text{eff } K_H$ ($\text{M} \cdot \text{atm}^{-1}$) is the field-derived effective Henry's law coefficient;
 190 c_p ($\mu\text{g} \cdot \text{m}^{-3}$) and c_g (atm) are particle- and gas-phase concentrations of carbonyls, respectively; ALWC ($\mu\text{g} \cdot \text{m}^{-3}$) is the aerosol
 191 liquid water content calculated by the thermodynamic model ISORROPIA-II(forward model, metastable state), the results of
 192 which are comparable to the actual measured contents confirmed by previous studies (Guo et al., 2015).

193 The irreversible reactive uptake coefficient γ could efficiently describe the irreversible pathways of the gas-particle partitioning
 194 process of dicarbonyls driven by OH radicals. We could estimate the reactive uptake coefficients γ based on the effective
 195 Henry's constant via theory calculation (Hanson et al., 1994; Curry et al., 2018) and then calculate the effective uptake rate
 196 $k_{\text{eff, uptake}}$, following Eqs. (4)–(7):

$$197 \frac{1}{\gamma} = \frac{1}{\alpha} + \frac{v}{4RT \text{ eff } K_H \sqrt{k^1 D_{\text{aq}}}} \times \frac{1}{[\text{coth}q - 1/q]} \quad (4)$$

$$198 v = \sqrt{\frac{8RT}{\pi M_X}} \quad (5)$$

$$199 q = R_p / l = R_p / \sqrt{\frac{D_{\text{aq}}}{k^1}} \quad (6)$$

$$200 k_{\text{eff, uptake}} = \frac{1}{4} v \times \gamma \times A_{\text{surf}} \quad (7)$$

201 where γ is the dimensionless uptake coefficient, v ($\text{m} \cdot \text{s}^{-1}$) is the gas-phase thermal velocity of glyoxal/methylglyoxal, D_{aq}
 202 ($\text{m}^2 \cdot \text{s}^{-1}$) is the diffusion coefficient in the liquid phase, α is dimensionless mass accommodation coefficient, $\text{eff } K_H$ ($\text{M} \cdot \text{atm}^{-1}$)
 203 is the effective Henry's law constant calculated by field-measured data in Table 2, R is the universal gas constant, k^1 (s^{-1}) is the
 204 first-order aqueous loss rate, M_X ($\text{kg} \cdot \text{mol}^{-1}$) is the average molar mass of gas-phase dicarbonyls, q is the parameter for
 205 measuring in-particle diffusion limitations, R_p (m) is the particle radius, l (m) is the diffusion reactive length, $k_{\text{eff, uptake}}$ (s^{-1}) is
 206 the effective uptake rate, and A_{surf} ($\text{m}^2 \cdot \text{m}^{-3}$) is the aerosol surface area density. This formulation is based on the effective Henry's

law constant under high RH conditions (RH > 40%). Moreover, the formulation describes the reactive uptake due to irreversible multiple-phase loss processes in the presence of OH. The uncertainty in the γ calculation is mainly attributed to the uncertainty in OH concentration, which was 3×10^{-12} M on average and varied from 5.5×10^{-14} to 8×10^{-12} M (Herrmann et al., 2010).

3 RESULT AND DISCUSSION

3.1 Observation results and partitioning coefficients calculation

3.1.1 Dicarbonyls in the gas and particle phase

We launched five field observations in different seasons. Table S1 details the information about the field observations, including observation periods, sample volume, and meteorological parameters. We totally collected 387 gas-phase samples and 130 particle-phase samples in four seasons. In these samples, carbonyls were simultaneously measured in both gas phase and particle phase. Ten carbonyls were measured in the gas phase, including formaldehyde, acetaldehyde, acetone, propionaldehyde, methacrolein, butyraldehyde, methyl vinyl ketone, benzaldehyde, glyoxal, and methylglyoxal; and six carbonyls were measured in the particle phase, including formaldehyde, acetaldehyde, acetone, propionaldehyde, glyoxal, and methylglyoxal. In this study, we mainly discuss the gas-particle partitioning processes of glyoxal and methylglyoxal because of their significant roles in atmospheric chemistry.

Figure 1 and Table 1 show the temporal characteristics and seasonal variation of glyoxal and methylglyoxal, respectively. Gaseous dicarbonyls showed obvious seasonal variation. Concentrations in summer (0.99 ± 0.59 ppbv) were generally much higher than in other seasons, followed by autumn and spring, and the concentrations in winter were the lowest. This seasonal variation could be partly attributed to the higher temperature and more intensive radiation in summer, which could greatly enhance the secondary formation of gaseous carbonyls via photochemical reactions. The diurnal variation in the dicarbonyls during summer support this interpretation of the data; gas-phase dicarbonyls exhibited obviously diurnal variations in summer, whereas this variation was irregular in other seasons (Fig. S32). The concentration levels of gaseous dicarbonyl in summer rapidly increased after sunrise, remained relatively high during the daytime (12:00–14:00), and then decreased at dusk. Although methylglyoxal has a shorter lifetime compare to glyoxal (glyoxal 2.9 h vs. MG-methylglyoxal 1.6 h) (Fu et al., 2008), its gas-phase concentration levels were generally higher than those of glyoxal, consistent with previous studies (Rao et al., 2016; Mitsuishi et al., 2018; Qian et al., 2019), mainly due to the relatively larger production from isoprene and acetone for methylglyoxal.

The concentrations of particulate dicarbonyls were an order of magnitude smaller than the gaseous concentrations using the unit of nanogram per cubic meter of air (ng/m^3 air). The average particulate glyoxal and methylglyoxal were 19.37 and 11.24 ng/m^3 , respectively, which were slightly higher than previously reported values (Zhu et al., 2018; Shen et al., 2018; Qian et al.,

236 2019; Cui et al., 2021). Dicarbonyls measured in the particle phase also showed obvious seasonal variation. The particulate
237 concentrations of the two dicarbonyls in winter ($43.38 \pm 32.42 \text{ ng/m}^3 \text{ air}$) were 2–2.3 times higher than those in other seasons,
238 suggesting that the dicarbonyls were more favored into the particle phase in winter. Moreover, particulate dicarbonyls in
239 different seasons exhibited the same diurnal variation (Fig. S32). The particulate concentrations of dicarbonyls in daytime
240 were generally higher than those in nighttime, especially in winter.

241 3.1.2 Gas-particle partitioning coefficient

242 Dicarbonyls could partition between gas and aerosol phases or the liquid phase, following Pankow's absorptive partitioning
243 theory or Henry's law, respectively, as listed in Table 2. Both gas-particle partitioning coefficient (K_p^f) and effective Henry's
244 law coefficient (K_H^f) were calculated on the basis of field-measured data and were in the range of 10^{-4} – $10^{-2} \text{ m}^3 \cdot \mu\text{g}^{-1}$ and 10^6 –
245 $10^8 \text{ M} \cdot \text{atm}^{-1}$, respectively. The partitioning coefficient values of the two dicarbonyls exhibited the same seasonal variation, as
246 winter and spring > autumn > summer. A higher aerosol concentration accompanied by higher aerosol surface area
247 concentration and lower relative humidity resulted in a higher partitioning coefficient in winter and spring, when heavy
248 pollution and sandstorms always occurred. In the case of temperature variation varied from 265.53 K to 310.75 K in different
249 seasons, lower temperature promoted the gas-to-particle partitioning processes as K_p^f values for the dicarbonyls and
250 temperature showed negative correlation with significant difference ($p < 0.001$) (Fig. S43). Moreover, The K_p^f and K_H^f values
251 of glyoxal were always higher than those of methylglyoxal, implying the former was more likely to partition to the particle
252 phase; this could be attributed to their different structures. Glyoxal were more soluble and reactive because of the adjacent
253 electron-poor aldehydic carbons, whereas methylglyoxal was more stable due to the reduced electron-deficient ketone moiety
254 (Kroll et al., 2005).

255 Both K_p^f and K_H^f were relatively closed to those found in previous field-measured studies (Shen et al., 2018; Qian et al., 2019;
256 Cui et al., 2021). However, compared with the theoretical partitioning coefficients K_p^t calculated by Pankow's absorptive
257 theory, K_p^f values were approximately 5–7 orders of magnitudes higher than the corresponding K_p^t values. ~~The influencing
258 factors in Pankow's absorptive model, like the activity coefficient ζ or absorbing fraction f_{om} , could not explain this great
259 difference between the field-measured values and the theoretical ones. The underestimation of gas-particle partitioning
260 coefficients can be attributed to the misidentification of condensed phase species produced in heterogeneous chemical reactions.~~
261 Similarly, K_H^f values were approximately 2–5 orders of magnitudes higher than the theoretical Henry's law coefficient K_H^t
262 calculated in pure water, which could be attributed to salting effects in wet aerosol (Figure S5). K_p^f/K_p^t values in this study
263 were close to but slightly higher than the values published in previous literature (Table S2) and the discrepancy between field-
264 measured partitioning coefficients and the theoretical ones was fully discussed in Supporting Information. The discrepancy
265 could be explained by the complex components in aerosol liquid water. Figure S4 presents the Setschenow plot of dicarbonyls

266 ~~versus aqueous sulfate, nitrate, and ammonia (SNA) concentration in aerosol. The negative salting constant indicated the~~
267 ~~“salting in” effects, which could result in exponential solubility, for both glyoxal and methylglyoxal in the real atmosphere.~~
268 ~~Moreover, both K_p^f and K_H^f of dicarbonyls were more than one magnitude higher than the reported laboratory partitioning~~
269 ~~coefficient values from chamber experiments (Healy et al., 2008; Healy et al., 2009), indicating that the real atmosphere is~~
270 ~~more favorable for the partitioning of gaseous dicarbonyls to the particle phase. Actual atmospheric environment conditions~~
271 ~~and complex particle compositions, such as higher ionic strength, could greatly affect the partitioning process and the chemical~~
272 ~~reactions in the aerosols.~~

273 To narrow the large discrepancy between field-measured partitioning coefficients and theoretical ones, we needed to further
274 investigate the mechanism and product distribution of chemical reactions occurring in the aerosols during the partitioning
275 processes. The products of the reversible and irreversible pathways mostly have lower saturated vapor pressure, and thus
276 leading to higher partitioning coefficients compared to monomer dicarbonyls. Take glyoxal for example, the effective
277 saturation vapor pressures of the product set in reversible pathways are $\sim 10^{-5}$ Torr in the real atmosphere (Shen et al., 2018).
278 And the products of the irreversible pathways had much lower vapor pressure values than those of reversible pathways. for
279 example, the vapor pressure of oxalic acids and ammonium oxalates are $\sim 10^{-5}$ Torr (Saxena and Hildemann, 1996) and $5.18 \times$
280 10^{-8} Torr (EPA, 2011), respectively, and those of glyoxal trimer dihydrates are $\sim 10^{-11}$ Torr at 20 °C (SPARC, 2003), indicating
281 the irreversible pathways make larger contributions to the underestimation of partitioning processes of dicarbonyls. For
282 example, the calculated vapor pressures of the products of glyoxal hydration and dimerization are, respectively, 5 and 10 orders
283 of magnitudes less than that of glyoxal monomer (Hastings et al., 2005). Moreover, the products of the irreversible pathway,
284 such as organic acids produced in radical chemistry, also have lower vapor pressure and efficiently contribute to the
285 underestimation of partitioning coefficients. The following sections further discuss the mechanism and product distribution of
286 reversible and irreversible pathways to explain the partitioning process of dicarbonyls.

287 3.2 Reversible pathways

288 Gas-particle partitioning of dicarbonyls via reversible pathways mainly consists of hydration and self-oligomerization. Since
289 glyoxal and methylglyoxal had high water solubility and reactivity, they could easily dissolve into aerosol liquid water, and
290 then form hydrates and oligomers. Hemiacetal/acetal formation (Loeffler et al., 2006) and aldol condensation (Haan et al.,
291 2009) are the most thermodynamically favored oligomer reactions for glyoxal hydrates and methylglyoxal hydrates,
292 respectively. The proposed mechanism for the reversible formation of glyoxal and methylglyoxal in aerosols is shown in Fig.
293 S5S6. By adding excess derivatization agent (like 2,4-dinitrophenylhydrazone in this study), dicarbonyls as well as their
294 reversibly formed products are efficiently transformed into dicarbonyl-bis-2,4-dinitrophenylhydrazone, which are quantified
295 as monomers by means of analysis techniques (Kampf et al., 2013). Moreover, Healy et al. (2008) have confirmed that

296 derivatization agent was found to efficiently dissolve a trimeric glyoxal standard and convert the resulting monomers to oxime
297 derivatives, and oligomers could not be detected in the extracts of filter samples by GC-MS analysis, also indicating the use
298 of excess derivatization agent could efficiently convert the hydrates and oligomers back to the monomeric species by removing
299 dicarbonyl monomers from the extract as soon as they are formed. Both dissolves dicarbonyl monomers and reversibly formed
300 production are efficiently transformed into carbonyl-bis-2,4-dinitrophenylhydrazone, which was quantified by means of
301 HPLC-UV in this study. The concentrations of dissolved dicarbonyl monomers were estimated using Henry's law coefficients,
302 which is used to determine the physical solubility of carbonyls (e.g., $K_H=5 \text{ M}\cdot\text{atm}^{-1}$ for glyoxal) (Schweitzer et al., 1998). The
303 results were negligible compared to the concentrations of carbonyls in hydrate and oligomer forms. Thus, the concentrations
304 of particle-phase dicarbonyl in reversible partitioning pathways~~Overall, since the products of the reversible pathway, including~~
305 ~~hydrates and oligomers, are thermodynamically unstable and could easily revert to their original monomer form during~~
306 ~~extraction and analysis were close to the measured concentration of carbonyls by HPLC-UV.~~

307 As glyoxal and methylglyoxal have similar trend under different conditions, we focused on the total concentration of the two
308 dicarbonyls in the following discussion. As shown in Fig. 2a, the particulate concentration of dicarbonyls via a reversible
309 pathway was strongly dependent on RH. It increased significantly when RH increased from <10% to 60%, as dicarbonyls were
310 more favorable to dissolve into hygroscopic aerosols during their growth (Mitsubishi et al., 2018; Xu et al., 2020). However,
311 from 60% to 80% RH, it exhibited the opposite trend and decreased with increasing RH, as higher water concentrations at
312 elevated RH levels may dilute the monomer concentration in the particle phase and hinder oligomerization reactions (Healy et
313 al., 2009), and the product distribution of the reversible formation could also well explain this phenomenon. The results
314 exhibited similar partten to a previous study, in which the partitioning of glyoxal and methylglyoxal gradually increased as
315 RH increased to 40%, peaked sharply around 50%, and subsequently decreased as RH increased towards 80%(Healy et al.,
316 2009). Moreover, ionic strength could also influence the reversible partitioning process as it is closely related to aerosol liquid
317 water and RH conditions. The presence of inorganic ions could catalyze and participate in oligomerization reactions via salting
318 effects (Sareen et al., 2010; Mcneill, 2015). Whereas, increasing viscosity of particles with increasing ionic strength could
319 slow down all particle-phase reactions, and the reversible nucleophilic addition of inorganic ions (e.g., sulfate ions) at carbonyl
320 carbons deactivates the molecule for further oligomerization (Kampf et al., 2013).~~Moreover, under high RH conditions, the~~
321 ~~particulate concentration of dicarbonyls via a reversible pathway had a strong and positive dependence on particle acidity (pH).~~
322 ~~The product distribution of the reversible formation could well explain this phenomenon.~~

323 To roughly estimate the product distribution of the reversible pathway in the real atmosphere, we simplified reaction
324 mechanisms and calculated the product distribution on the basis of the kinetic mechanisms listed in Table S3 using a 0-D box
325 model with a steady-state approach,~~the equilibrium constant reported in previous literature (Table S2).~~ Generally, more
326 dicarbonyls existed in oligomer forms ~~(83.5% for glyoxal and 80.8% for methylglyoxal)~~ than in hydrate forms ~~(16.3% for~~

glyoxal and 20.8% for methylglyoxal) in the reversible formation. Moreover, their distribution exhibited obvious seasonal variations. Summer had the highest proportion of hydrate forms (52.80% for GL and 35.10% for MG), while winter had the highest proportion of oligomer forms (86.15% for GL and 86.31% for MG). Detailed information is shown in Table S43. The seasonal variation could be attributed to the RH in different seasons – relatively high in summer and low in winter. As shown in Fig. 2b, the product distribution of the reversible formation has a strong dependence on RH. The proportion of dicarbonyls in hydrate forms increased with increasing RH and could reach more than 75% in high RH, while the proportion of dicarbonyls in oligomer forms exhibited the opposite trend. Hydrates play a dominant role in dilute solutions under high RH conditions with a relatively high aerosol liquid water concentration, which might hinder oligomer formation. And large quantities of oligomers, including dimers and trimers, would form until the aerosol liquid concentration became greater than 1 M (Liggio et al., 2005b) when RH decreased. However, the product distribution here was simulated based on the bulk-phase mechanisms and higher ionic strength in aerosol phase would influence reaction equilibria and rate constant (Ervens and Volkamer, 2010; McNeill, 2015). The lack of quantitative reaction rate in aerosol phase could contribute more uncertainties to the simulation, whereas, the RH dependence of product distribution and the order of magnitude of estimated K_p values were close to those in aerosol-phase and the roughly simulation could help to understand the reversible partitioning pathways of dicarbonyls. Combined with the vapor pressure of dominant products published in previous studies (Hastings et al., 2005; Axson et al., 2010), their gas-particle partitioning coefficient can be roughly estimated and can effectively fit the field-measured values, as shown in Fig. 2c. The estimated gas-particle partitioning coefficients in this study are five orders of magnitude higher than the theoretical ones but still 1–2 orders of magnitude lower than the field-measured coefficients, especially in winter. The difference between the estimated partitioning coefficients and the field-measured ones suggests that the current understanding of the equilibrium in reversible formations cannot reasonably explain the gas-particle partitioning processes of dicarbonyls. There still exist extra pathways of reversible formation. Cross-oligomerization of glyoxal and methylglyoxal is nonnegligible and could form similar molecular structure products and contribute to SOA yield (Schwier et al., 2010). Esterification and amination of diols also occur in aerosol liquid water but are negligible compared to hydration and polymerization (Zhao et al., 2006). However, these reactions are not further discussed here. The hydrates and oligomers mentioned above are the dominant forms of glyoxal/methylglyoxal in the particle phase, while the higher molecular oligomers up to nonamer could also exist with a relatively smaller but still significant fraction at equilibrium. Although the reactions are thermodynamically reversible, upon evaporation of the aerosol liquid water, the oligomer formation is faster than the evaporation of dehydrated dicarbonyls, and the dicarbonyl evaporation is limited (Liggio et al., 2005b; Loeffler et al., 2006). This results in relatively stable oligomers and yielding SOA. Moreover, other nucleophilic species may also form oligomers with glyoxal and methylglyoxal and effectively prevent their evaporation. Besides reversible pathways, higher carbon number products with lower volatility were mainly formed through irreversible pathways, such as like radical reactions (e.g., OH radicals), which are fully discussed in the

358 next section.

359 3.3 Irreversible pathways

360 3.3.1 Irreversible pathways driven by hydroxyl radicals

361 Reactive uptake driven by hydroxyl radicals (OH) is the dominant process for glyoxal and methylglyoxal in their irreversible
362 gas-particle partitioning pathways. Compared to other irreversible pathways, ~~such as like~~ imidazole formation,
363 glyoxal/methylglyoxal + OH chemistry occurs on much shorter timescales (Teich et al., 2016). The reaction is the initial step
364 for most radical-based chemistry of glyoxal/methylglyoxal and has been proven to be an important source of SOA in both
365 cloud/fog droplets and wet aerosols (Tan et al., 2012; Lim et al., 2013), producing low-volatility products such as organic acids,
366 large multifunctional humic-like substances, and oligomers. The proposed mechanism for the irreversible pathway of glyoxal
367 and methylglyoxal driven by hydroxyl radicals in aerosols is shown in Fig. S76. The OH radicals in aerosol liquid water are
368 mainly from the direct uptake of gas-phase OH radicals with a Henry's law constant of 30 M/atm (Faust and Allen, 1993) and
369 Fenton reactions, and Fenton reactions are closely related to hydrogen peroxide, iron ions, and manganese ions in the particle
370 phase. The sources of OH radicals are one of the major uncertainties in SOA formation (Ervens et al., 2014).

371 ~~The irreversible reactive uptake coefficient γ could efficiently describe the irreversible pathway of the gas-particle partitioning~~
372 ~~process of dicarbonyls driven by OH radicals. We could estimate the reactive uptake coefficients γ based on the effective~~
373 ~~Henry's constant via theory calculation (Hanson et al., 1994; Curry et al., 2018) and then calculate the effective uptake rate~~
374 ~~$k_{\text{eff, uptake}}$, following Eqs. (4)–(7):~~

$$375 \frac{1}{\gamma} = \frac{1}{\alpha} + \frac{v}{4RTH^*} \sqrt{k^1 D_{\text{aq}}} \times \frac{1}{[\text{eothq}^{-1}/q]} \quad (4)$$

$$376 v = \sqrt{\frac{8RT}{\pi M_x}} \quad (5)$$

$$377 q = R_p / l - R_p / \sqrt{\frac{D_{\text{aq}}}{k^1}} \quad (6)$$

$$378 k_{\text{eff, uptake}} = \frac{1}{4} v \times \gamma \times A_{\text{surf}} \quad (7)$$

379 where γ is the dimensionless uptake coefficient, v ($\text{m}\cdot\text{s}^{-1}$) is the gas phase thermal velocity of glyoxal/methylglyoxal, D_{aq}
380 ($\text{m}^2\cdot\text{s}^{-1}$) is the diffusion coefficient in the liquid phase, α is the dimensionless mass accommodation coefficient, H^* ($\text{M}\cdot\text{atm}^{-1}$)
381 is the effective Henry's law constant calculated by field-measured data in Table 2, R is the universal gas constant, k^1 (s^{-1}) is the
382 first order aqueous loss rate, M_x ($\text{kg}\cdot\text{mol}^{-1}$) is the average molar mass of gas phase dicarbonyls, q is the parameter for
383 measuring in particle diffusion limitations, R_p (m) is the particle radius, l (m) is the diffusion reactive length, $k_{\text{eff, uptake}}$ (s^{-1}) is
384 the effective uptake rate, and A_{surf} ($\text{m}^2\cdot\text{m}^{-3}$) is the aerosol surface area density. This formulation is based on the effective Henry's
385 law constant under high RH conditions ($\text{RH} > 40\%$). Moreover, the formulation describes the reactive uptake due to irreversible

386 ~~multiple phase loss processes in the presence of OH. The uncertainty in the γ calculation is mainly attributed to the uncertainty~~
387 ~~in OH concentration, which was 3×10^{-12} M on average and varied from 5.5×10^{-14} to 8×10^{-12} M (Herrmann et al., 2010). The~~
388 calculated γ and $k_{\text{eff, uptake}}$ values for different seasons are listed in Table 3. The reactive uptake coefficients of glyoxal were in
389 the range 10^{-4} – 10^{-2} , and the average value of 8.0×10^{-3} in this study was closed to the ones representing the loss of glyoxal by
390 surface uptake during the KORUS-AQ campaign in a very recent studies (Kim et al., 2022). And the value slightly exceeded
391 the one commonly used in model simulations ($\gamma = 2.9 \times 10^{-3}$), which was based on an experimental study for $(\text{NH}_4)_2\text{SO}_4$ aerosols
392 at 55% RH (Liggio et al., 2005a), and also far outweighs the uptake coefficients of glyoxal on clean and acidic gas-aged
393 mineral particles ($\gamma = 10^{-6}$ – 10^{-4}) (Shen et al., 2016), implying that a real atmospheric aerosol provides a far more reactive
394 interface for physiochemical processes than that of mineral particles. Moreover, uptake coefficients for methylglyoxal were
395 with an average value of 2.0×10^{-3} and were higher than those reported in other experimental studies, which varied from 10^{-6}
396 to 10^{-3} (Curry et al., 2018; De Haan et al., 2018). On the one hand, conflicting with previous experimental results (Waxman et
397 al., 2015), methylglyoxal exhibited an unexpected salting-in effect in real atmosphere due to much more complex compositions
398 and higher ionic strength in ambient particles, which was also reported in other observational studies (Shen et al., 2018; Cui
399 et al., 2021). And the higher Henry's law coefficient values in Eq.4 could lead to higher uptake coefficient values. One the
400 other hand, a recent study also provided direct experimental evidence to confirm that methylglyoxal is more reactive and have
401 larger uptake coefficients on seed particles under atmospherically relevant concentrations (Li et al., 2021). ~~Moreover, the~~
402 ~~reactive uptake coefficients of methylglyoxal were slightly lower than those for glyoxal, with an average value of 2.0×10^{-3} .~~
403 ~~Conflicting with previous experimental results, methylglyoxal exhibited unexpected salting-in effects in real atmospheric~~
404 ~~particles and had much higher uptake coefficients, which could be attributed to the increased reactivity of methylglyoxal with~~
405 ~~a high uptake coefficient under an atmospheric relevant concentration (Li et al., 2021).~~ The γ values for both glyoxal and
406 methylglyoxal exhibited similar seasonal variations, which were lowest in summer and reached their highest in winter. This
407 seasonal variation could be attributed to RH variation and particle composition. Moreover, the effective uptake rate ($k_{\text{eff, uptake}}$),
408 which is regarded as a pseudo-first-order reaction rate, is a net result of competition between reversible and irreversible
409 processes, and it varied from 10^{-4} s^{-1} to 10^{-5} s^{-1} in the real atmosphere in this study. As shown in Fig. 3a, the negative dependence
410 of $k_{\text{eff, uptake}}$ on RH also confirmed that the irreversible uptake of dicarbonyls could be inhibited in high RH conditions. What's
411 more, as we can see in Figure 3b, the irreversible uptake increased exponentially with increasing SNA (SNA: Sulfate, Nitrate
412 and Ammonia) concentrations, mainly because that higher SNA concentrations always occurred in lower RH conditions with
413 lower aerosol liquid water content (Figure S8) and the irreversible uptake was promoted by combined efforts of RH effects
414 and ion effects. And the relatively low SNA concentration under high RH conditions also attenuated the irreversible uptake as
415 the weakening of ion effects (Figure 3b). Whereas, for a given RH, uptake coefficients γ for both glyoxal and methylglyoxal
416 showed a weak dependence on the ratio of SNA (Sulfate:Ammonia and Sulfate:Nitrate) with significant scatter (Fig. S8S9).

417 Moreover, it was worth noting that under extremely low RH (<40%), the aerosol was not completely deliquescent, and the
418 uptake coefficients based on Henry's law could not explain the irreversible pathways. Previous research indicated that the
419 irreversible uptake of dicarbonyls could still occur under a low RH condition (Liggio et al., 2005a; De Haan et al., 2018), and
420 that these uptake values were generally lower due to the inefficient reactive uptake process onto the crystallized aerosols.

421 3.3.2 Reactive uptake of dicarbonyl compounds

422 We could not directly measure the particulate concentration of dicarbonyls via an irreversible pathway, as the dicarbonyls
423 irreversibly reacted with oxidative radicals on aerosols. To quantitatively evaluate the contribution of the irreversible pathway
424 of dicarbonyls, we calculated their average concentration based on Eqs. (S4S3)–(S5S7) in the Supplement with the calculated
425 γ values in this study. The samples estimated here were collected under high RH conditions (RH > 40%) because of the
426 calculation limitation of irreversible uptake coefficients. Although the products of irreversible pathways could not be directly
427 detected in particle phase and didn't directly contribute to the increase of particulate dicarbonyls, the irreversible pathways
428 could contribute to the decrease of gaseous dicarbonyls and well explain the overestimation of modeled dicarbonyl mixing
429 ratios, which is about 3-6 times higher than the observed ones (Volkamer et al., 2007; Ling et al., 2020).

430 The total particulate concentration of glyoxal and methylglyoxal via irreversible pathway varied from several to more than
431 100 nanograms per microgram PM_{2.5} (ng/ μ g PM_{2.5}), and it was strongly dependent on RH, as shown in Fig. 3c, which generally
432 decreased with increasing RH. Concentrated inorganic solutions and relatively higher ionic strength in aerosol water under
433 low RH conditions could jointly contribute to the hydration of dicarbonyls, the products of which could easily participate into
434 the following irreversible radical reactions via H-abstraction. Since the irreversible uptake coefficients γ of dicarbonyls tended
435 to decrease with increasing RH due to the "salting-in" effects, both glyoxal and methylglyoxal in this study exhibited high
436 solubility in low RH conditions, where the aerosol water was coupled with concentrated inorganic solutions and relatively
437 high ionic strength. Moreover, the total particulate concentrations of dicarbonyls were positively dependent on particle acidity
438 under high RH conditions, while there were no obvious correlations under low RH conditions.

439 To further discuss the product distribution of the reaction of glyoxal/methylglyoxal with hydroxyl radicals, we used the kinetic
440 mechanisms of glyoxal/methylglyoxal + OH chemistry proposed by Lim et al.(2013) on the basis of a 0-D box model with a
441 steady-state approach. The average OH radical concentration setting in the modeling was 3.2×10^{-12} M, which is based on the
442 hypothesis of the Henry equilibrium of OH radicals between the gas and particle phase (Sander, 2015; Shen et al., 2018).
443 Oxalate can be considered as a tracer for this aqueous chemistry, since it does not have any other significant chemical sources.
444 Oxalate was detected in the particle-phase samples by ion chromatography. The modeling results of oxalate concentration
445 agreed well with the measured values, and their deviations were in the considered range (Fig. S107). Meanwhile, we can
446 estimate the distribution of major products in irreversible glyoxal/methylglyoxal-OH radical chemistry under different RH

447 conditions, as illustrated in Figure 3d. Generally, oxalate is the major product in wet aerosols, contributing ~60%, and its
448 proportion increases with increasing RH. Besides oxalate, oligomers also play significant roles in glyoxal/methylglyoxal-OH
449 radical chemistry with a contribution of ~30%, and their proportion is maximum under relatively low RH conditions. The RH
450 dependence of the product distribution could mainly be attributed to the particulate concentration of glyoxal/methylglyoxal,
451 which significantly affects the OH radical chemistry. With relatively high carbonyl concentrations (0.1–10 M) in aerosol liquid
452 water, self-reactions of organic molecules become more favorable, resulting in new carbon–carbon bonds and high molecular
453 weight oligomers via radical–radical chemistry (Lim et al., 2013). Moreover, besides OH radical chemistry, reaction with
454 sulfate and ammonium also contribute to the oligomer formation and irreversible uptake of gaseous dicarbonyls (Ortiz-
455 Montalvo et al., 2014; Lin et al., 2015; Lim et al., 2016). The oligomer proportion could be more than 30% in concentrated
456 carbonyl solutions (~0.1 M) and only account for 1% in diluted solutions (~0.01 M).

457 **3.4 Relative importance of two partitioning pathways**

458 Table 4 summarizes the particulate concentration of glyoxal and methylglyoxal via reversible and irreversible pathways in
459 different seasons. The average particulate concentrations of glyoxal (0.43 ng/μg in the reversible pathway and 24.26 ng/μg in
460 the irreversible pathway) were generally higher than those of methylglyoxal (0.25 ng/μg in the reversible pathway and 16.53
461 ng/μg in the irreversible pathway), mainly due to the relatively higher water solubility and reactivity of glyoxal. Comparing
462 two gas-particle partitioning processes, the irreversible pathway played extremely dominant roles and generally accounted for
463 96.7% and 95.0% for glyoxal and methylglyoxal, respectively. The proportion of the irreversible pathway varied from 90% to
464 99.9% and reached its highest in summer for glyoxal (98.8%) and in autumn for methylglyoxal (99.2%), while it was minimum
465 in winter (92.9% for glyoxal and 92.8% for methylglyoxal). Overall, the irreversible pathway played a dominant role in the
466 gas-particle partitioning process for both glyoxal and methylglyoxal in the real atmosphere, while the reversible pathway was
467 also substantial and nonnegligible, especially in winter, with an proportion of ~10%. Furthermore, as discussed above, the
468 particulate concentrations of dicarbonyls and their relative importance were influenced by environmental factors such as
469 relative humidity and particle composition, which could jointly influence both the reversible and irreversible pathways of
470 dicarbonyls. As shown in Figure 4, the proportion of irreversible pathways in the gas-particle partitioning process for
471 dicarbonyls increased with aqueous SNA concentrations, and reached maximum when SNA concentrations were more than 100
472 M under low RH conditions. Moreover, higher organic concentrations in aerosol may lead to an OH-limit environment,
473 hindering the irreversible pathways driven by radicals and influencing the relative importance of the two pathways (Waxman
474 et al., 2013; Ervens et al., 2014). But the OH limitations are still exclusive due to the uncertainties in the sources of OH in
475 aerosol particles (Herrmann et al., 2010).
476 Comprehensively considering the contribution of both reversible pathways and irreversible pathways occurred in gas-particle

477 ~~partitioning processes could benefit the two pathways in partitioning processes could be conducive to~~ ambient dicarbonyls
478 simulations. Ling et al. (2020) found that the observation and simulation of the gas-phase concentration level of dicarbonyls
479 could reach reasonable agreement when the irreversible uptake and reversible partitioning were incorporated into the model,
480 as these jointly contribute ~62% to the sink of dicarbonyls. Moreover, the contribution of gas-particle partitioning processes
481 of dicarbonyls to SOA formation were higher as the two partitioning pathways were jointly considered. In this study, gas-
482 particle partitioning processes of dicarbonyls accounted for a relatively large proportion of total particle mass (PM_{2.5}), on the
483 average of ~5% considering both reversible and irreversible gas-particle partitioning pathways. Since a large fraction of PM_{2.5}
484 mass in Beijing consists of SOAs (~30%) (Huang et al., 2014), we could roughly estimate the contribution of gas-particle
485 partitioning processes of dicarbonyls to SOA yields (by mass). There were approximately 25% SOAs formed from glyoxal
486 and methylglyoxal in this study. However, the particulate dicarbonyls calculated here only contained simple reversible
487 pathways and irreversible pathways driven by OH radicals. More complicated chemical processes, ~~such as like~~ NO₃ radical
488 chemistry, were not considered, which still resulted in the underestimation of their contribution to SOA formation.

489 4 Conclusions

490 We simultaneously measured glyoxal and methylglyoxal concentration in the gas and particle phase in different seasons over
491 urban Beijing. Based on field-measured data, the field-derived gas-particle partitioning coefficients were calculated and found
492 to be 5–7 magnitudes higher than the theoretical values. Such a large discrepancy provides field evidence that the gas-particle
493 partitioning process does not occur by physical absorption alone but also results from the combined and simultaneous effects
494 of reversible and irreversible pathways. Hydration and oligomerization occurred in the reversible pathway, producing
495 compounds with lower volatility in the condensed phase, and the irreversible pathway could accelerate the uptake of gaseous
496 dicarbonyls. The two pathways jointly contributed to the underestimation of gas-particle partitioning of dicarbonyls.

497 This study systemically considers both reversible and irreversible pathways in the ambient atmosphere for the first time.
498 Compared to the reversible pathways, the irreversible pathways play a dominant role in the gas-particle partitioning process
499 for dicarbonyls, accounting for ~90% of this process. We recommend the irreversible reactive uptake coefficient for glyoxal
500 and methylglyoxal in different seasons in the real atmosphere. The values we calculated here are higher than those used in
501 model simulations to date, especially for methylglyoxal which exhibits an unexpected salting-in effect under an atmospheric-
502 relevant concentration. We expect the application of these parameterizations will increase the calculated contribution of
503 irreversible uptake of dicarbonyls to SOA formation and narrow the gap between model predictions and field measurements
504 of ambient dicarbonyl concentrations. Moreover, relative humidity and inorganic particle compositions are defined as the most
505 important factor influencing particulate concentration and product distribution of dicarbonyls via both reversible and

506 irreversible pathways, implying the significance of considering different RH conditions in dicarbonyl SOA simulations.

507 Furthermore, we note that there may be other potential explanations for the increase in particle mass caused by dicarbonyls
508 ~~particulate concentrations~~ and the uncertainty in the gas-particle partitioning process, including physical adsorption, reversible
509 pathways and irreversible pathways. Physical adsorption of dicarbonyls could be enhanced by water-soluble organics and
510 mineral dust. Other reversible pathways, like such as adducts formed from glyoxal with inorganic species, ~~like sulfate and~~
511 ~~ammonia~~, could also promote the gas-particle partitioning process. Irreversible pathways driven by other oxidants, like such
512 as NO₃ radicals, can also perform a substantial role. Shen et al. (2016) found that glyoxal could irreversibly produce formic
513 acid, glycolic acid, and oligomers on particles without illumination or extra oxidants. Besides gas-particle partitioning,
514 particulate dicarbonyls formed via the heterogeneous reaction of VOCs could contribute to the uncertainty in partitioning
515 research. Dong et al. (2021) recently revealed that aqueous photooxidation of toluene could yield glyoxal and methylglyoxal
516 via a ring-cleavage process. Overall, the real gas-particle partitioning process of glyoxal and methylglyoxal is more
517 complicated and their contribution to SOA formation is still indistinct; thus, more laboratory experiments and field
518 measurements are urgently needed to improve our understanding of the gas-particle partitioning process for glyoxal and
519 methylglyoxal.

520

521 *Data availability.* The data are accessible by contacting the corresponding author (zmchen@pku.edu.cn).

522

523 *Author contributions.* In the framework of the five field measurements in different seasons, ZC and JH designed the study, and
524 JH performed all carbonyl measurements used in this study, analyzed the data, and wrote the paper. ZC helped interpret the
525 results, guided the writing, and modified the manuscript. XQ and PD contributed to the methods of sampling and analyzing
526 gas- and particle-phase carbonyls. All authors discussed the results and contributed to the final paper.

527

528 *Competing interests.* The authors declare that they have no conflict of interest.

529

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533 **Reference**

- 534 SPARC performs automated reasoning in chemistry., <http://ibmlc2.chem.uga.edu/sparc/index.cfm>., 2003.
- 535 Axson, J. L., Takahashi, K., De Haan, D. O., and Vaida, V.: Gas-phase water-mediated equilibrium between methylglyoxal
536 and its geminal diol, *Proc Natl Acad Sci U S A*, 107, 6687-6692, 10.1073/pnas.0912121107, 2010.
- 537 Chen, X., Zhang, Y., Zhao, J., Liu, Y., Shen, C., Wu, L., Wang, X., Fan, Q., Zhou, S., and Hang, J.: Regional modeling of
538 secondary organic aerosol formation over eastern China: The impact of uptake coefficients of dicarbonyls and semivolatile
539 process of primary organic aerosol, *Science of the Total Environment*, 793, 148176, 10.1016/j.scitotenv.2021.148176, 2021.
- 540 Clegg, S. L., Brimblecombe, P., and Wexler, A. S.: Thermodynamic Model of the System $H^+-NH_4^+-SO_4^{2-}-NO_3^- -H_2O$
541 at Tropospheric Temperatures, *Journal of Physical Chemistry A*, 102, 2137-2154, 1998.
- 542 Cui, J., Sun, M., Wang, L., Guo, J., Xie, G., Zhang, J., and Zhang, R.: Gas-particle partitioning of carbonyls and its
543 influencing factors in the urban atmosphere of Zhengzhou, China, *Science of the Total Environment*, 751, 142027,
544 10.1016/j.scitotenv.2020.142027, 2021.
- 545 Curry, L. A., Tsui, W. G., and McNeill, V. F.: Technical note: Updated parameterization of the reactive uptake of glyoxal
546 and methylglyoxal by atmospheric aerosols and cloud droplets, *Atmospheric Chemistry and Physics*, 18, 9823-9830,
547 10.5194/acp-18-9823-2018, 2018.
- 548 De Haan, D. O., Jimenez, N. G., de Loera, A., Cazaunau, M., Gratien, A., Pangui, E., and Doussin, J. F.: Methylglyoxal
549 Uptake Coefficients on Aqueous Aerosol Surfaces, *Journal of Physical Chemistry A*, 122, 4854-4860,
550 10.1021/acs.jpca.8b00533, 2018.
- 551 Donaldson, D. J. and Vaida, V.: The Influence of Organic Films at the Air–Aqueous Boundary on Atmospheric Processes,
552 *Chemical Reviews*, 106, 1445-1461, 2006.
- 553 Dong, P., Chen, Z., Qin, X., and Gong, Y.: Water Significantly Changes the Ring-Cleavage Process During Aqueous
554 Photooxidation of Toluene, *Environmental Science & Technology*, 55, 16316-16325, 10.1021/acs.est.1c04770, 2021.
- 555 Elrod, M. J., Sedlak, J. A., and Ren, H.: Accurate Computational Model for the Hydration Extent of Atmospherically
556 Relevant Carbonyls on Aqueous Atmospheric Particles, *ACS Earth and Space Chemistry*, 5, 348-355,
557 10.1021/acsearthspacechem.0c00322, 2021.
- 558 EPA: Estimation Programs Interface (EPI) Suite TM for Microsoft®Windows, v4.1, Environmental Protection Agency
559 (EPA), Washington, DC., 24, 2011.
- 560 EPA., U.: Estimation Programs Interface Suite for Microsoft Windows, v 4.11. , United States Environmental Protection
561 Agency: Washington, DC, USA, <https://www.epa.gov/tsca-screeningtools/epi-suitetm-estimation-program-interface>., 2012.
- 562 Ervens, B. and Volkamer, R.: Glyoxal processing by aerosol multiphase chemistry: towards a kinetic modeling framework
563 of secondary organic aerosol formation in aqueous particles, *Atmospheric Chemistry and Physics*, 10, 8219-8244, 10.5194/acp-
564 10-8219-2010, 2010.
- 565 Ervens, B., Sorooshian, A., Lim, Y. B., and Turpin, B. J.: Key parameters controlling OH-initiated formation of secondary
566 organic aerosol in the aqueous phase (aqSOA), *Journal of Geophysical Research: Atmospheres*, 119, 3997-4016,
567 10.1002/2013jd021021, 2014.
- 568 Faust, B. C. and Allen, J. M.: Aqueous-phase photochemical formation of hydroxyl radical in authentic cloudwaters and
569 fogwaters, *Environmental Science & Technology*, 27, 113-122, 1993.
- 570 Fu, T.-M., Jacob, D. J., Wittrock, F., Burrows, J. P., Vrekoussis, M., and Henze, D. K.: Global budgets of atmospheric
571 glyoxal and methylglyoxal, and implications for formation of secondary organic aerosols, *Journal of Geophysical Research*,
572 113, 10.1029/2007jd009505, 2008.
- 573 Guo, H., Xu, L., Bougiatioti, A., Cerully, K. M., and Weber, R. J.: Fine-particle water and pH in the southeastern United
574 States, *Atmospheric Chemistry and Physics*, 15, 5211-5228, 2015.
- 575 Haan, D. O. D., Corrigan, A. L., Smith, K. W., Stroik, D. R., Turley, J. J., Lee, F. E., Tolbert, M. A., Jimenez, J. L.,
576 Cordova, K. E., and Ferrell, G. R.: Secondary organic aerosol-forming reactions of glyoxal with amino acids, *Environmental
577 Science & Technology*, 43, 2818, 2009.

578 Hanson, D. R., Ravishankara, A. R., and Solomon, S.: Heterogeneous reactions in sulfuric acid aerosols: A framework
579 for model calculations, *Journal of Geophysical Research Atmospheres*, 99, 3615-3629, 1994.

580 Hart, K. M. and Pankow, J. F.: High-Volume Air Sampler for Particle and Gas Sampling.2.Use of Backup Filters To
581 Correct for the Adsorption of Gas-Phase Polycyclic Aromatic Hydrocarbons to the Front Filter, *Environmental Science &*
582 *Technology*, 28, 655-661, 1994.

583 Hastings, W. P., Koehler, C. A., Bailey, E. L., and Haan, D. O. D.: Secondary organic aerosol formation by glyoxal
584 hydration and oligomer formation: humidity effects and equilibrium shifts during analysis, *Environmental Science &*
585 *Technology*, 39, 8728-8735, 2005.

586 Healy, R. M., Temime, B., Kuprovskite, K., and Wenger, J. C.: Effect of relative humidity on gas/particle partitioning
587 and aerosol mass yield in the photooxidation of p-xylene, *Environmental Science & Technology*, 43, 1884-1889, 2009.

588 Healy, R. M., Wenger, J. C., Metzger, A., Duplissy, J., Kalberer, M., and Dommen, J.: Gas/particle partitioning of
589 carbonyls in the photooxidation of isoprene and 1,3,5-trimethylbenzene, *Atmospheric Chemistry and Physics*, 8, 2008.

590 Herrmann, H., Hoffmann, D., Schaefer, T., Brauer, P., and Tilgner, A.: Tropospheric aqueous-phase free-radical chemistry:
591 radical sources, spectra, reaction kinetics and prediction tools, *Chemphyschem*, 11, 3796-3822, 10.1002/cphc.201000533,
592 2010.

593 Hu, J., Wang, P., Ying, Q., Zhang, H., Chen, J., Ge, X., Li, X., Jiang, J., Wang, S., Zhang, J., Zhao, Y., and Zhang, Y.:
594 Modeling biogenic and anthropogenic secondary organic aerosol in China, *Atmospheric Chemistry and Physics*, 17, 77-92,
595 10.5194/acp-17-77-2017, 2017.

596 Huang, R. J., Zhang, Y., Bozzetti, C., Ho, K. F., Cao, J. J., Han, Y., Daellenbach, K. R., Slowik, J. G., Platt, S. M.,
597 Canonaco, F., Zotter, P., Wolf, R., Pieber, S. M., Brun, E. A., Crippa, M., Ciarelli, G., Piazzalunga, A., Schwikowski, M.,
598 Abbazade, G., Schnelle-Kreis, J., Zimmermann, R., An, Z., Szidat, S., Baltensperger, U., El Haddad, I., and Prevot, A. S.:
599 High secondary aerosol contribution to particulate pollution during haze events in China, *Nature*, 514, 218-222,
600 10.1038/nature13774, 2014.

601 Ip, H., Huang, X., and Jian, Z. Y.: Effective Henry's law constants of glyoxal, glyoxylic acid, and glycolic acid,
602 *Geophysical Research Letters*, 36, 2009.

603 Kampf, C. J., Waxman, E. M., Slowik, J. G., Dommen, J., Pfaffenberger, L., Praplan, A. P., Prevot, A. S., Baltensperger,
604 U., Hoffmann, T., and Volkamer, R.: Effective Henry's law partitioning and the salting constant of glyoxal in aerosols
605 containing sulfate, *Environmental Science & Technology*, 47, 4236-4244, 10.1021/es400083d, 2013.

606 Kim, D., Cho, C., Jeong, S., Lee, S., Nault, B. A., Campuzano-Jost, P., Day, D. A., Schroder, J. C., Jimenez, J. L., Volkamer,
607 R., Blake, D. R., Wisthaler, A., Fried, A., DiGangi, J. P., Diskin, G. S., Pusede, S. E., Hall, S. R., Ullmann, K., Huey, L. G.,
608 Tanner, D. J., Dibb, J., Knute, C. J., and Min, K.-E.: Field observational constraints on the controllers in glyoxal (CHOCHO)
609 reactive uptake to aerosol, *Atmospheric Chemistry and Physics*, 22, 805-821, 10.5194/acp-22-805-2022, 2022.

610 Kroll, J. H., Ng, N. L., Murphy, S. M., Varutbangkul, V., Flagan, R. C., and Seinfeld, J. H.: Chamber studies of secondary
611 organic aerosol growth by reactive uptake of simple carbonyl compounds, *Journal of Geophysical Research*, 110,
612 10.1029/2005jd006004, 2005.

613 Laskin, A., Laskin, J., and Nizkorodov, S. A.: Chemistry of atmospheric brown carbon, *Chem Rev*, 115, 4335-4382,
614 10.1021/cr5006167, 2015.

615 Li, X., Rohrer, F., Brauers, T., Hofzumahaus, A., Lu, K., Shao, M., Zhang, Y. H., and Wahner, A.: Modeling of HCHO
616 and CHOCHO at a semi-rural site in southern China during the PRIDE-PRD2006 campaign, *Atmospheric Chemistry and*
617 *Physics*, 14, 22(2014-11-21), 14, 2014.

618 Li, Y., Ji, Y., Zhao, J., Wang, Y., Shi, Q., Peng, J., Wang, Y., Wang, C., Zhang, F., Wang, Y., Seinfeld, J. H., and Zhang,
619 R.: Unexpected Oligomerization of Small alpha-Dicarbonyls for Secondary Organic Aerosol and Brown Carbon Formation,
620 *Environmental Science & Technology*, 55, 4430-4439, 10.1021/acs.est.0c08066, 2021.

621 Liggio, J.: Uptake of carbonyls to atmospheric particulate matter: Ambient measurements and laboratory studies, York
622 University (Canada). 2004.

623 Liggio, J., Li, S. M., and McLaren, R.: Reactive uptake of glyoxal by particulate matter, *Journal of Geophysical Research*

624 Atmospheres, 110, 2005a.

625 Liggio, J., Shao-Meng, L. I., and McLaren, R.: Heterogeneous Reactions of Glyoxal on Particulate Matter: Identification
626 of Acetals and Sulfate Esters, *Environmental Science & Technology*, 39, 1532-1541, 2005b.

627 Lim, Y. B., Tan, Y., and Turpin, B. J.: Chemical insights, explicit chemistry, and yields of secondary organic aerosol from
628 OH radical oxidation of methylglyoxal and glyoxal in the aqueous phase, *Atmospheric Chemistry and Physics*, 13, 8651-8667,
629 10.5194/acp-13-8651-2013, 2013.

630 Lim, Y. B., Kim, H., Kim, J. Y., and Turpin, B. J.: Photochemical organonitrate formation in wet aerosols, *Atmospheric
631 Chemistry and Physics*, 16, 12631-12647, 10.5194/acp-16-12631-2016, 2016.

632 Lin, P., Laskin, J., Nizkorodov, S. A., and Laskin, A.: Revealing Brown Carbon Chromophores Produced in Reactions of
633 Methylglyoxal with Ammonium Sulfate, *Environ Sci Technol*, 49, 14257-14266, 10.1021/acs.est.5b03608, 2015.

634 Ling, Z., Xie, Q., Shao, M., Wang, Z., Wang, T., Guo, H., and Wang, X.: Formation and sink of glyoxal and methylglyoxal
635 in a polluted subtropical environment: observation-based photochemical analysis and impact evaluation, *Atmospheric
636 Chemistry and Physics*, 20, 11451-11467, 10.5194/acp-20-11451-2020, 2020.

637 Loeffler, K. W., Koehler, C. A., Paul, N. M., and Haan, D. D.: Oligomer formation in evaporating aqueous glyoxal and
638 methyl glyoxal solutions, *Environmental Science & Technology*, 40, 6318, 2006.

639 Lv, S., Gong, D., Ding, Y., Lin, Y., Wang, H., Ding, H., Wu, G., He, C., Zhou, L., Liu, S., Ristovski, Z., Chen, D., Shao,
640 M., Zhang, Y., and Wang, B.: Elevated levels of glyoxal and methylglyoxal at a remote mountain site in southern China:
641 Prompt in-situ formation combined with strong regional transport, *Science of the Total Environment*, 672, 869-882,
642 10.1016/j.scitotenv.2019.04.020, 2019.

643 Mader, B. T. and Pankow, J. F.: Gas/Solid Partitioning of Semivolatile Organic Compounds (SOCs) to Air Filters. 3. An
644 Analysis of Gas Adsorption Artifacts in Measurements of Atmospheric SOC and Organic Carbon (OC) When Using Teflon
645 Membrane Filters and Quartz Fiber Filters, *Environmental Science & Technology*, 35, 3422-3432, 2001.

646 McNeill, V. F.: Aqueous organic chemistry in the atmosphere: sources and chemical processing of organic aerosols,
647 *Environ Sci Technol*, 49, 1237-1244, 10.1021/es5043707, 2015.

648 Michailoudi, G., Lin, J. J., Yuzawa, H., Nagasaka, M., Huttula, M., Kosugi, N., Kurtén, T., Patanen, M., and Prisle, N. L.:
649 Aqueous-phase behavior of glyoxal and methylglyoxal observed with carbon and oxygen K-edge X-ray absorption
650 spectroscopy, *Atmospheric Chemistry and Physics*, 21, 2881-2894, 10.5194/acp-21-2881-2021, 2021.

651 Mitsuishi, K., Iwasaki, M., Takeuchi, M., Okochi, H., Kato, S., Ohira, S.-I., and Toda, K.: Diurnal Variations in
652 Partitioning of Atmospheric Glyoxal and Methylglyoxal between Gas and Particles at the Ground Level and in the Free
653 Troposphere, *ACS Earth and Space Chemistry*, 2, 915-924, 10.1021/acsearthspacechem.8b00037, 2018.

654 Odabasi, M. and Seyfioglu, R.: Phase partitioning of atmospheric formaldehyde in a suburban atmosphere, *Atmospheric
655 Environment*, 39, 5149-5156, 10.1016/j.atmosenv.2005.05.006, 2005.

656 Odum, J. R., Hoffmann, T., Bowman, F., Collins, D., Flagan, R. C., and Seinfeld, J. H.: Gas/Particle Partitioning and
657 Secondary Organic Aerosol Yields, *Environmental Science & Technology*, 30, 2580-2585, 1996.

658 Ortiz-Montalvo, D. L., Hakkinen, S. A., Schwier, A. N., Lim, Y. B., McNeill, V. F., and Turpin, B. J.: Ammonium addition
659 (and aerosol pH) has a dramatic impact on the volatility and yield of glyoxal secondary organic aerosol, *Environmental Science
660 & Technology*, 48, 255-262, 10.1021/es4035667, 2014.

661 Pankow and James, F.: An absorption model of the gas/aerosol partitioning involved in the formation of secondary organic
662 aerosol, *Atmospheric Environment*, 28, 189-193, 1994.

663 Pankow, J. F.: An absorption model of GAS/Particle partitioning of organic compounds in the atmosphere, 1994.

664 Paul, Davidovits, Charles, E., Kolb, Leah, R., Williams, John, and T.: Mass Accommodation and Chemical Reactions at
665 Gas/Liquid Interfaces, *Chemical Reviews*, 111, 2011.

666 Qian, X., Shen, H., and Chen, Z.: Characterizing summer and winter carbonyl compounds in Beijing atmosphere,
667 *Atmospheric Environment*, 214, 10.1016/j.atmosenv.2019.116845, 2019.

668 Qiu, X., Wang, S., Ying, Q., Duan, L., Xing, J., Cao, J., Wu, D., Li, X., Chengzhi, X., Yan, X., Liu, C., and Hao, J.:
669 Importance of Wintertime Anthropogenic Glyoxal and Methylglyoxal Emissions in Beijing and Implications for Secondary

670 Organic Aerosol Formation in Megacities, *Environmental Science & Technology*, 54, 11809-11817, 10.1021/acs.est.0c02822,
671 2020.

672 Rao, Z., Chen, Z., Liang, H., Huang, L., and Huang, D.: Carbonyl compounds over urban Beijing: Concentrations on
673 haze and non-haze days and effects on radical chemistry, *Atmospheric Environment*, 124, 207-216,
674 10.1016/j.atmosenv.2015.06.050, 2016.

675 Sander, R.: Compilation of Henry's law constants (version 4.0) for water as solvent, *Atmospheric Chemistry and Physics*,
676 15, 4399-4981, 10.5194/acp-15-4399-2015, 2015.

677 Sareen, N., Schwier, A. N., Shapiro, E. L., Mitroo, D., and McNeill, V. F.: Secondary organic material formed by
678 methylglyoxal in aqueous aerosol mimics, *Atmospheric Chemistry and Physics*, 10, 997-1016, 2010.

679 Saxena, P. and Hildemann, L. M.: Water-soluble organics in atmospheric particles: A critical review of the literature and
680 application of thermodynamics to identify candidate compounds, *Journal of Atmospheric Chemistry*, 24, 57-109, 1996.

681 Schweitzer, F., Magi, L., Mirabel, P., and George, C.: Uptake Rate Measurements of Methanesulfonic Acid and Glyoxal
682 by Aqueous Droplets, *Journal of Physical Chemistry A*, 102, 593-600, 1998.

683 Schwier, A. N., Sareen, N., Mitroo, D., Shapiro, E. L., and McNeill, V. F.: Glyoxal-methylglyoxal cross-reactions in
684 secondary organic aerosol formation, *Environmental Science & Technology*, 44, 6174-6182, 2010.

685 Shen, H., Chen, Z., Li, H., Qian, X., Qin, X., and Shi, W.: Gas-Particle Partitioning of Carbonyl Compounds in the
686 Ambient Atmosphere, *Environmental Science & Technology*, 52, 10997-11006, 10.1021/acs.est.8b01882, 2018.

687 Shen, X., Wu, H., Zhao, Y., Huang, D., Huang, L., and Chen, Z.: Heterogeneous reactions of glyoxal on mineral particles:
688 A new avenue for oligomers and organosulfate formation, *Atmospheric Environment*, 131, 133-140,
689 10.1016/j.atmosenv.2016.01.048, 2016.

690 Shi, Q., Zhang, W., Ji, Y., Wang, J., Qin, D., Chen, J., Gao, Y., Li, G., and An, T.: Enhanced uptake of glyoxal at the acidic
691 nanoparticle interface: implications for secondary organic aerosol formation, *Environmental Science: Nano*, 7, 1126-1135,
692 10.1039/d0en00016g, 2020.

693 Tan, Y., Lim, Y. B., Altieri, K. E., Seitzinger, S. P., and Turpin, B. J.: Mechanisms leading to oligomers and SOA through
694 aqueous photooxidation: insights from OH radical oxidation of acetic acid and methylglyoxal, *Atmospheric Chemistry &
695 Physics*, 12, 801-813, 2012.

696 Teich, M., Pinxteren, D. V., Kecorius, S., Wang, Z., and Herrmann, H.: First Quantification of Imidazoles in Ambient
697 Aerosol Particles: Potential Photosensitizers, Brown Carbon Constituents, and Hazardous Components, *Environmental
698 Science & Technology*, 50, 1166-1173, 2016.

699 Volkamer, R., San Martini, F., Molina, L. T., Salcedo, D., Jimenez, J. L., and Molina, M. J.: A missing sink for gas-phase
700 glyoxal in Mexico City: Formation of secondary organic aerosol, *Geophysical Research Letters*, 34, 10.1029/2007gl030752,
701 2007.

702 Wang, H., Zhang, X., and Chen, Z.: Development of DNPH/HPLC method for the measurement of carbonyl compounds
703 in the aqueous phase: applications to laboratory simulation and field measurement, *Environmental Chemistry*, 6,
704 10.1071/en09057, 2009.

705 Waxman, E. M., Elm, J., Kurten, T., Mikkelsen, K. V., Ziemann, P. J., and Volkamer, R.: Glyoxal and Methylglyoxal
706 Setschenow Salting Constants in Sulfate, Nitrate, and Chloride Solutions: Measurements and Gibbs Energies, *Environmental
707 Science & Technology*, 49, 11500-11508, 10.1021/acs.est.5b02782, 2015.

708 Waxman, E. M., Dzepina, K., Ervens, B., Lee-Taylor, J., Aumont, B., Jimenez, J. L., Madronich, S., and Volkamer, R.:
709 Secondary organic aerosol formation from semi- and intermediate-volatility organic compounds and glyoxal: Relevance of
710 O/C as a tracer for aqueous multiphase chemistry, *Geophysical Research Letters*, 40, 978-982, 10.1002/grl.50203, 2013.

711 Williams, B. J., Goldstein, A. H., Kreisberg, N. M., and Hering, S. V.: In situ measurements of gas/particle-phase
712 transitions for atmospheric semivolatile organic compounds, *Proceedings of the National Academy of Sciences*, 107, 6676-
713 6681, 10.1073/pnas.0911858107, 2010.

714 Xie, M., Hannigan, M. P., and Barsanti, K. C.: Gas/particle partitioning of 2-methyltetrols and levoglucosan at an urban
715 site in Denver, *Environmental Science & Technology*, 48, 2835-2842, 10.1021/es405356n, 2014.

716 Xu, R., Li, X., Dong, H., Wu, Z., Chen, S., Fang, X., Gao, J., Guo, S., Hu, M., Li, D., Liu, Y., Liu, Y., Lou, S., Lu, K.,
717 Meng, X., Wang, H., Zeng, L., Zong, T., Hu, J., Chen, M., Shao, M., and Zhang, Y.: Measurement of gaseous and particulate
718 formaldehyde in the Yangtze River Delta, China, *Atmospheric Environment*, 224, 10.1016/j.atmosenv.2019.117114, 2020.

719 Ying, Q., Li, J., and Kota, S. H.: Significant Contributions of Isoprene to Summertime Secondary Organic Aerosol in
720 Eastern United States, *Environmental Science & Technology*, 49, 7834-7842, 10.1021/acs.est.5b02514, 2015.

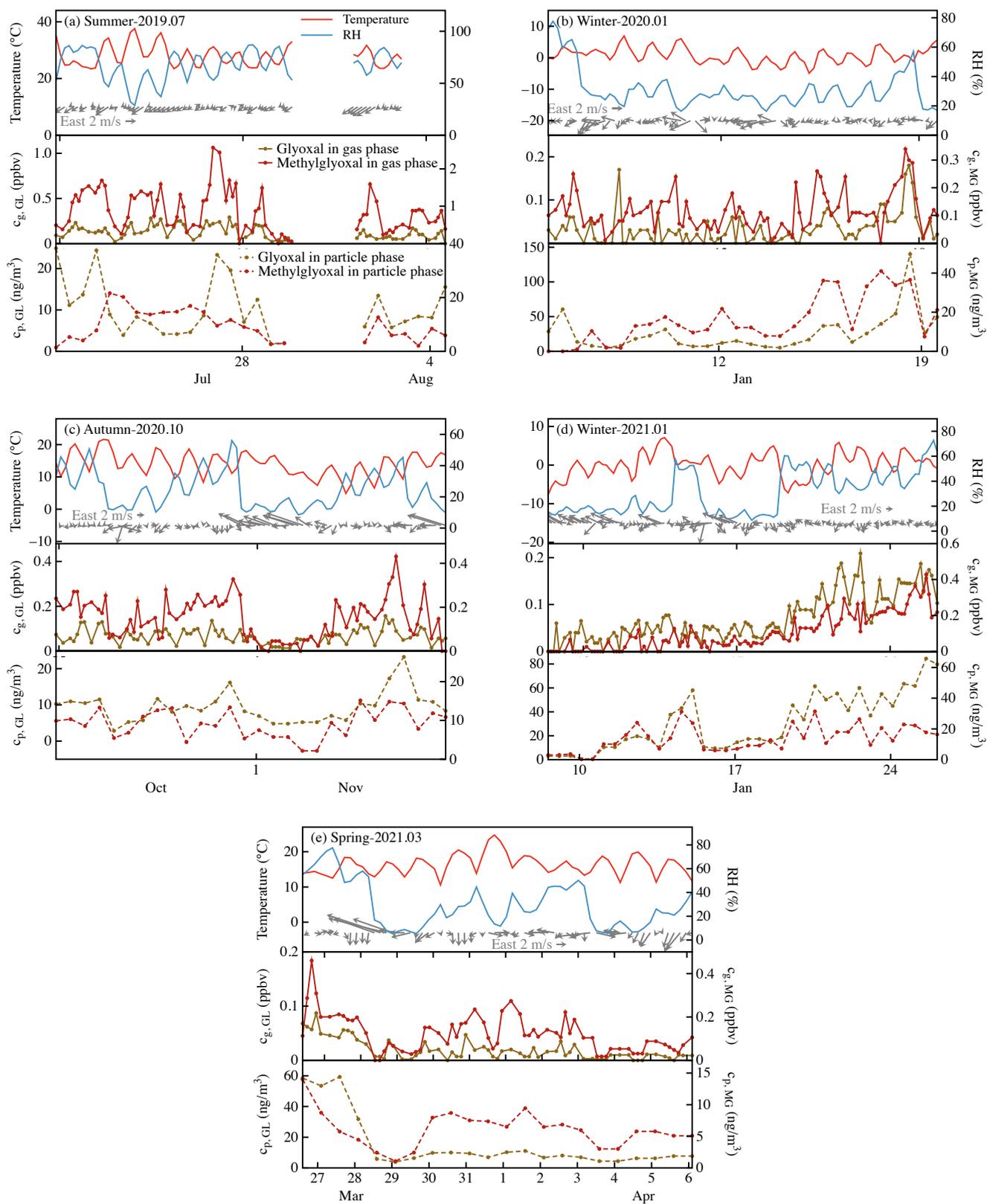
721 Zarzana, K. J., Selimovic, V., Koss, A. R., Sekimoto, K., Coggon, M. M., Yuan, B., Dubé, W. P., Yokelson, R. J., Warneke,
722 C., de Gouw, J. A., Roberts, J. M., and Brown, S. S.: Primary emissions of glyoxal and methylglyoxal from laboratory
723 measurements of open biomass burning, *Atmospheric Chemistry & Physics*, 18, 15451-15470, 10.5194/acp-18-15451-2018,
724 2018.

725 Zarzana, K. J., Min, K. E., Washenfelder, R. A., Kaiser, J., Krawiec-Thayer, M., Peischl, J., Neuman, J. A., Nowak, J. B.,
726 Wagner, N. L., Dube, W. P., St Clair, J. M., Wolfe, G. M., Hanisco, T. F., Keutsch, F. N., Ryerson, T. B., and Brown, S. S.:
727 Emissions of Glyoxal and Other Carbonyl Compounds from Agricultural Biomass Burning Plumes Sampled by Aircraft,
728 *Environmental Science & Technology*, 51, 11761-11770, 10.1021/acs.est.7b03517, 2017.

729 Zhao, J., Levitt, N. P., Zhang, R., and Chen, J.: Heterogeneous reactions of methylglyoxal in acidic media: implications
730 for secondary organic aerosol formation, *Environmental Science & Technology*, 40, 7682-7687, 2006.

731 Zhu, Y., Yang, L., Chen, J., Kawamura, K., Sato, M., Tilgner, A., van Pinxteren, D., Chen, Y., Xue, L., Wang, X., Simpson,
732 I. J., Herrmann, H., Blake, D. R., and Wang, W.: Molecular distributions of dicarboxylic acids, oxocarboxylic acids and α -
733 dicarbonyls in PM_{2.5} collected at the top of Mt. Tai, North China, during the wheat burning season of 2014, *Atmospheric
734 Chemistry and Physics*, 18, 10741-10758, 10.5194/acp-18-10741-2018, 2018.

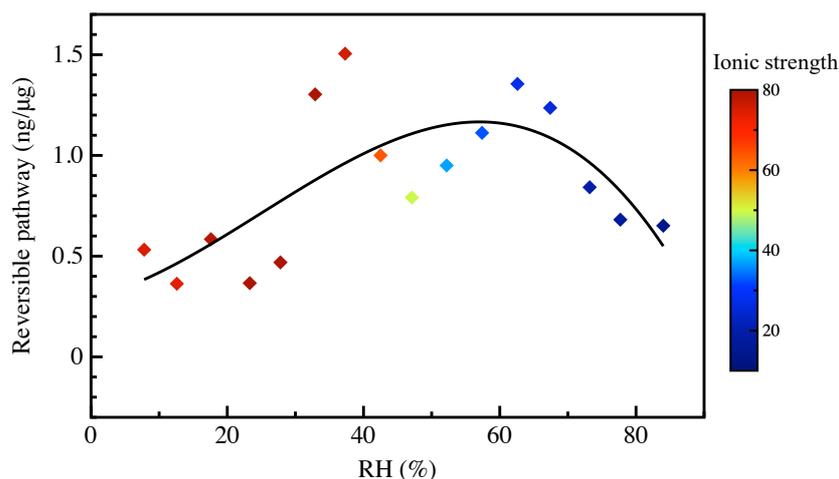
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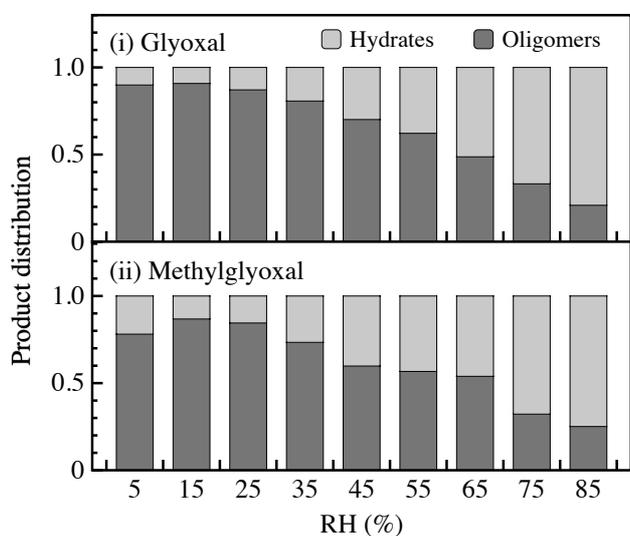
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737 Figure 1: Time series of meteorological parameters and gas- and particle-phase glyoxal and methylglyoxal observed in different
 738 seasons: (a) summer, 2019.07.20-2019.08.04; (b) winter, 2020.01.05-2020.01.19; (c) autumn, 2020.10.24-2020.11.07; (d)
 739 winter, 2021.01.08-2021.01.26; (e) spring, 2021.03.26-2021.04.06.

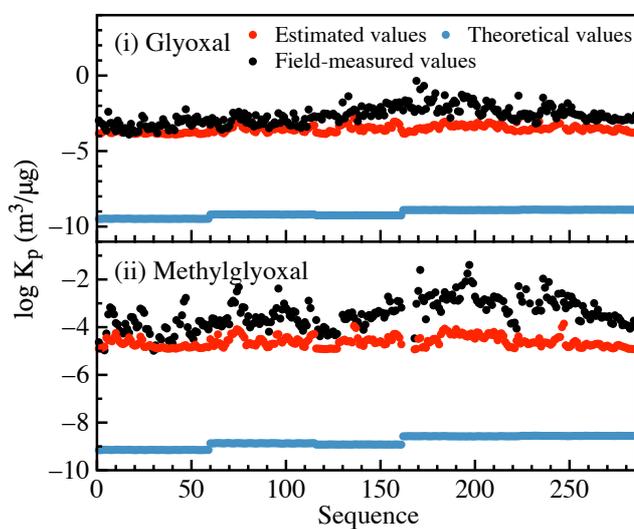
(a) RH dependence of dicarbonyl concentrations



(b) Product distribution

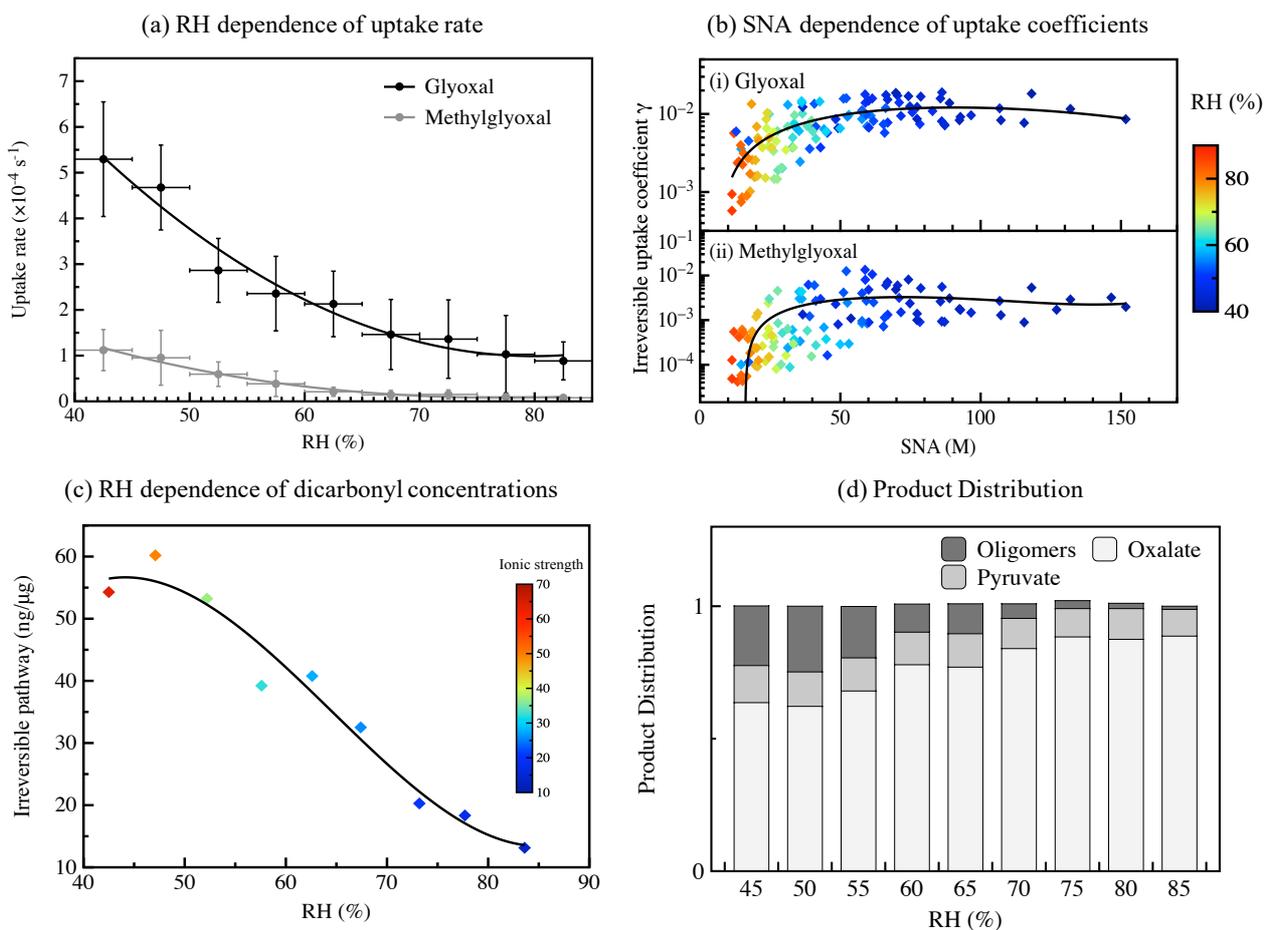


(c) Gas-particle Partitioning Coefficients



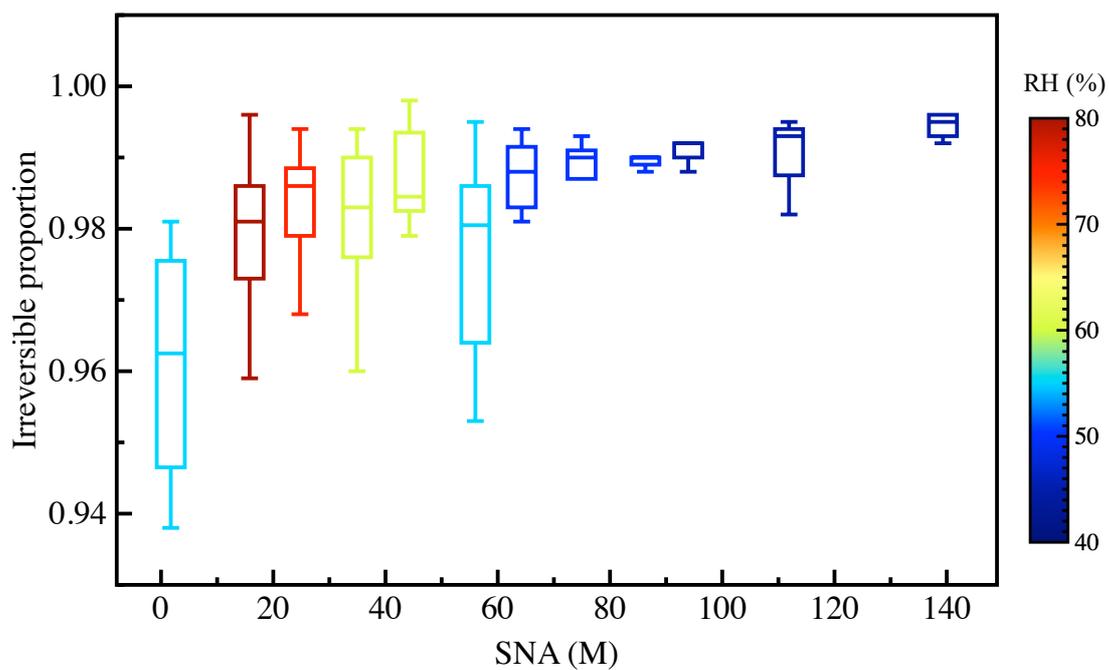
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741 Figure 2: Gas-particle partitioning of dicarbonyls via reversible pathways. (a) The RH dependence of particulate
742 concentrations of dicarbonyl via reversible pathways. (b) The product distribution for (i) glyoxal and (ii) methylglyoxal under
743 different RH conditions. (c) The gas-particle partitioning coefficients for (i) glyoxal and (ii) methylglyoxal. The black, ~~dark~~
744 grayred, and light grayblue circles refer to field-measured values, estimated values by the proposed mechanism, and theoretical
745 values calculated by Pankow's absorptive model, respectively.



746

747 Figure 3: Gas-particle partitioning of dicarbonyls via irreversible pathways. (a) The RH dependence of irreversible uptake rate
 748 for glyoxal and methylglyoxal. (b) The SNA dependence of uptake coefficients for (i) glyoxal and (ii) methylglyoxal, SNA
 749 refers to the concentration of sulfate, nitrate and ammonia in wet aerosols. (c) The RH dependence of particulate
 750 concentrations of dicarbonyl via irreversible pathways. (d) The corresponding modeled product distribution in wet aerosols
 751 under different RH conditions.



752

753 Figure 4: Correlation between the proportion of the irreversible pathway in gas-particle partitioning process for dicarbonyls
 754 and aqueous sulfate, nitrate, and ammonia (SNA) concentration in ambient aerosols under different relative humidity
 755 conditions.
 756

Table 1: Statistical data for the α -dicarbonyls in gas and particle phase in different seasons.

Season	<u>Dates of the measurements</u>	Gas phase (ppbv)		Particle phase (ng m ⁻³)	
		Glyoxal	Methylglyoxal	Glyoxal	Methylglyoxal
summer	<u>2019.07.20-08.04</u>	0.13 ± 0.07	0.87 ± 0.54	10.18 ± 6.63	9.50 ± 5.62
spring	<u>2021.03.26-04.06</u>	0.02 ± 0.02	0.12 ± 0.08	15.24 ± 17.50	6.07 ± 2.79
autumn	<u>2020.10.24-11.07</u>	0.07 ± 0.03	0.15 ± 0.09	9.33 ± 4.24	9.15 ± 3.62
winter	<u>2020.01.05-01.19</u> <u>2021.01.08-01.26</u>	0.06 ± 0.05	0.11 ± 0.09	28.77 ± 25.33	14.61 ± 10.15

759 Table 2: Comparison of the field-measured partitioning coefficient K^f values for the dicarbonyls and their corresponding
 760 theoretical K^t values in different seasons.

Coefficients	Dicarbonyl	Season	K^f		^a K^t	K^f / K^t
			Average	Range		
Gas-particle partitioning coefficients ($m^3 \cdot \mu g^{-1}$)	Glyoxal	summer	8.11×10^{-4}	$(1.25-58.6) \times 10^{-4}$	3.27×10^{-10}	2.48×10^6
		autumn	2.14×10^{-3}	$(2.61-224) \times 10^{-4}$	6.27×10^{-10}	3.41×10^6
		spring	1.43×10^{-2}	$(0.08-14.6) \times 10^{-2}$	5.59×10^{-10}	3.55×10^7
		winter	1.30×10^{-2}	$(0.067-44.2) \times 10^{-2}$	1.27×10^{-9}	1.02×10^7
	Methylglyoxal	summer	1.49×10^{-4}	$(0.833-169) \times 10^{-5}$	7.10×10^{-10}	2.10×10^5
		autumn	9.55×10^{-4}	$(0.65-86.9) \times 10^{-4}$	1.35×10^{-9}	7.07×10^5
		spring	1.06×10^{-3}	$(0.42-108) \times 10^{-4}$	1.21×10^{-9}	8.77×10^5
		winter	2.60×10^{-3}	$(0.34-410) \times 10^{-4}$	2.72×10^{-9}	9.93×10^5
Henry's law coefficients ($M \cdot atm^{-1}$)	Glyoxal	summer	1.96×10^8	$(1.71-167) \times 10^7$	3.29×10^5	6.11×10^2
		autumn	5.08×10^8	$(1.88-10.2) \times 10^8$	1.14×10^6	3.63×10^3
		spring	2.53×10^9	$(1.23-139) \times 10^8$	9.03×10^5	1.92×10^4
		winter	1.04×10^9	$(1.37-55.4) \times 10^8$	4.15×10^6	2.55×10^3
	Methylglyoxal	summer	4.92×10^7	$(1.70-363) \times 10^6$	2.73×10^3	1.88×10^4
		autumn	8.52×10^7	$(3.66-15.7) \times 10^7$	9.50×10^3	2.00×10^5
		spring	1.33×10^8	$(5.22-456) \times 10^6$	7.49×10^3	1.36×10^5
		winter	2.63×10^8	$(1.03-175) \times 10^7$	3.44×10^4	9.01×10^4

761 ^a Theoretical gas-particle partitioning coefficients were calculated on the basis of Eq. (3) and theoretical Henry's law
 762 coefficients here referred to the Henry's law constant in pure water, which were calculated on the basis of Eq.(S1)-(S2) (Ip et
 763 al., 2009; Sander, 2015).

764

765 Table 3: Summary of calculated uptake coefficients γ and effective uptake rate coefficient $k_{\text{eff, uptake}}$ in different seasons for
 766 glyoxal and methylglyoxal.

Dicarbonyl	Season	T (K)	RH (%)	$\gamma (\times 10^{-3})$			$k_{\text{eff, uptake}} (\text{s}^{-1})$ ($\times 10^{-4}$)
				Average	Min	Max	
<u>Glyoxal</u> GL	Summer	301.1	67.7	4.15	0.12	7.30	1.61
	Autumn	287.2	45.4	8.62	0.29	12.9	4.83
	Spring	289.4	54.0	11.7	1.24	14.9	6.85
	Winter	273.5	54.0	10.6	0.56	14.4	3.59
	^a General	287.8	59.0	8.0	0.46	11.4	3.38
MG <u>Methylgly</u> <u>oxal</u>	Summer	301.1	67.7	1.01	0.02	2.09	0.25
	Autumn	287.2	45.4	1.83	0.04	3.94	0.92
	Spring	289.4	54.0	2.36	0.06	4.34	0.69
	Winter	273.5	54.0	3.45	0.11	5.83	0.77
	^a General	287.8	59.0	2.0	0.05	3.8	0.55

767 ^aGeneral is the average value of all the samples observed in the five field observations.

768

769 Table 4: Calculated relative importance of reversible and irreversible pathways in the gas-particle partitioning processes and
 770 their contribution to the ~~particle-particulate~~ matter.

Season	Glyoxal		Methylglyoxal		Contribution to particulate matters
	^a [X] _{P, rever}	^b [X] _{P, irrever}	^a [X] _{P, rever}	^b [X] _{P, irrever}	
Summer	0.17 (1.2%)	18.87 (98.8%)	0.25 (5.5%)	20.55 (94.5%)	3.98%
Autumn	0.14 (0.7%)	23.91(99.3%)	0.12 (0.8%)	17.02 (99.2%)	4.12%
Spring	0.26 (1.9%)	15.94 (98.1%)	0.09 (1.5%)	14.16 (98.5%)	3.05%
Winter	0.89 (7.1%)	34.70 (92.9%)	0.38 (7.2%)	12.59 (92.8%)	4.86%
^c General	0.43 (3.3%)	24.26 (96.7%)	0.25 (5.0%)	16.53 (95.0%)	4.15%

771 ^a [X]_{P, rever} is the concentration of particle-phase carbonyl via reversible pathway (ng·μg⁻¹) and its proportion (%).

772 ^b [X]_{P, irrever} is the concentration of particle-phase carbonyl via irreversible pathway (ng·μg⁻¹) and its proportion (%).

773 ^c General is the average value of all the samples observed in the five field observations.