

1 Secondary Aerosol Formation in Incense Burning Particles by 2 Ozonolysis and Photochemical Oxidation via Single Particle Mixing 3 State Analysis

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18 **Abstract.** Incense burning is a common religious activity that emits abundant gaseous and particulate pollutants into the
19 atmosphere. During their atmospheric lifetime, these gases and particles are subjected to oxidation, leading to the formation
20 of secondary pollutants. We examined the oxidation of incense burning plumes under O₃ exposure and dark condition using
21 an oxidation flow reactor connected to a single particle aerosol mass spectrometer (SPAMS). Nitrate formation was observed
22 in incense burning particles, mainly attributable to the ozonolysis of nitrogen-containing organic compounds. With UV on,
23 nitrate formation was significantly enhanced, likely due to HNO₃/HNO₂/NO_x uptake triggered by OH chemistry, which is
24 more effective than ozone oxidation. The extent of nitrate formation is insensitive to O₃ and OH exposure, which can be
25 explained by the diffusion limitation on interfacial uptake. The O₃-UV-aged particles are more oxygenated and functionalized
26 than O₃-Dark-aged particles. Oxalate and malonate, two typical secondary organic aerosols (SOA), were found in O₃-UV-aged
27 particles. Our work reveals that nitrate, accompanied by SOA, can rapidly form in incense-burning particles upon
28 photochemical oxidation in the atmosphere, which could deepen our understanding of air pollution caused by religious
29 activities.

30 1 Introduction

31 Incense burning is a common religious ritual, especially in Asian and African communities (Ye et al., 2016; Khezri et al.,
32 2015; Sidibe et al., 2022), with a massive amount of particles emitted (Lyu et al., 2021; See et al., 2011). The PM_{2.5}
33 concentration at a shrine area reached 6-8 times higher than usual during the Chinese New Year in Chiang Mai, Thailand,
34 mainly due to incense burning (Bootdee et al., 2016). The particle emission factor (i.e., the mass ratio of the emitted particles

35 to the total material burnt) from incense burning could be up to 10 times higher than those from burning of various types of
36 biomasses such as rice straw (Akagi et al., 2011; See et al., 2011; Goel et al., 2020).

37 Previous research mainly focused on the chemical compositions and potential health impacts of fresh incense particles (Li et
38 al., 2012; Wang et al., 2006; Chuang et al., 2011; Lee et al., 2004). However, it was rarely considered that fresh particles would
39 also be exposed to other atmospheric pollutants and light, which could initiate chemical reactions. The formation of secondary
40 particulate pollutants could take place during the atmospheric aging of particles (Hodshire et al., 2019; Kumar et al., 2018;
41 Rudich et al., 2007). For example, our recent work reveals rapid sulfate formation in fresh incense burning particles upon SO₂
42 exposure in the dark, and it can be accelerated under light (Liang et al., 2022). Sulfate formation in incense burning particles
43 under dark and light was mainly attributed to gaseous oxidants and particulate photosensitizers, respectively.

44 Ozone and OH radicals are two of the most common oxidants in the atmosphere, contributing to secondary inorganic and
45 organic aerosol formation (Volkamer et al., 2006; Kroll et al., 2006; Chen et al., 2011; Liu et al., 2019; Liu et al., 2018).
46 Incense burning plume has been reported to contain various volatile organic compounds (VOCs) and NO_x (Lee et al., 2004;
47 Ho et al., 2002), in addition to particulate pollutants. Their interactions with ozone and OH radicals may lead to secondary
48 aerosol formation. [For instance, ozone and OH oxidations of NO_x were considered primary sources of particulate nitrate](#)
49 [\(Seinfeld et al., 2008; Liang et al., 2021; Gen et al., 2022\)](#). Nevertheless, there is still a lack of understanding of the secondary
50 aerosol formation in incense burning particles upon atmospheric aging of the plume, which could potentially worsen air quality,
51 especially near areas of intense religious activities.

52 This study examines the secondary aerosol formation in fresh incense burning particles under ozone and OH exposure using a
53 Gothenburg Potential Aerosol Mass (Go: PAM) flow reactor. We first characterized single fresh incense burning particles,
54 followed by aged particles, with a single-particle aerosol mass spectrometer (SPAMS). Control experiments were performed
55 to get insight into the possible secondary aerosol formation pathways and their significance. Then, we discussed nitrate
56 formation as a function of particle size, ozone, and OH exposure. The fragmentation severely hinders the characterization of
57 secondary organic aerosol (SOA). Nevertheless, -89[C₂HO₄] (oxalate) and -103[C₃H₃O₄] (malonate), two commonly
58 considered SOA(Cheng et al., 2017; Sullivan et al., 2007), were found.

59 **2 Experimental**

60 **2.1 Aging of incense-burning particles**

61 The schematic of the experimental set-up can be found in Figure S1. [In brief, we burnt an incense stick \(Figure S2, Kwok Tin](#)
62 [Heung, Hong Kong\) in a 20 L glass burning bottle for each experiment. The air exchange rate per hour \(ACH\) is 0.3,](#)
63 [comparable to the typical natural ventilation conditions \(Lee et al., 2004\). The relative humidity \(RH\) and the temperature](#)
64 [inside the burning bottle were 56 ± 9 % RH and 22 ± 2.7 °C.](#) The burning was rapidly converted from flaming to smoldering

65 after ignition. A two-stage system diluted the emissions with an overall dilution of around 1600. Compressed air ($\sim 0.1 \text{ L min}^{-1}$)
66 $)$ was used to introduce the diluted incense burning particles to the PAM reactor equipped with two UVC light tubes (30W,
67 Philips TUV, $\lambda_{\text{max}} = 254\text{nm}$). The spectrum of the lamp was shown in Supplementary information (Figure S3). In the control
68 experiments, a charcoal absorber or HEPA filter was used to remove the gaseous pollutants or particles prior to the introduction
69 to the Go:PAM. The data of these control experiments can be found in the supplementary information (Figure S4, S5). All
70 these control experiments were aged experiments. The removal efficiency of NO_x, VOCs, and particles were $\sim 85\%$, $\sim 90\%$,
71 and $\sim 100\%$, respectively, determined by the concentration reduction after applying a HEPA filter or charcoal absorber at the
72 exhaust of the burning bottle, using a NO_x analyzer (T200, Teledyne) or a Total VOC analyzer (Yuante) (Figure S1). A
73 controlled dry-wet mixed carrier flow of compressed air ($\sim 4 \text{ L min}^{-1}$) and a flow of O₃ ($\sim 0.1 \text{ L min}^{-1}$) generated by passing O₂
74 (99.995%, Linde) to an O₃ generator (Model 610, Jelight Company Inc, USA) were introduced into the Go:PAM. The
75 compressed air was treated by a HEPA filter and a charcoal absorber prior to the experiment system. [O₃] ranged from 300 to
76 1500 ppb, equivalent to an atmospheric ozone exposure of 10-50 min, assuming ambient concentration of 60 ppb (Xia et al.,
77 2021). The RH at the exit of the PAM was monitored by an RH sensor (M170, Vaisala, Finland). All the experiments were
78 conducted at 80% RH and $22 \pm 1.7 \text{ }^\circ\text{C}$. The residence time in the Go:PAM was $\sim 100 \text{ s}$. The exhaust of the PAM was
79 characterized by an O₃ analyzer (106L, 2B technology, USA), a water-based condensation particle counter (WCPC, Aerosol
80 dynamics Inc, USA), and a SPAMS (Hexin Analytical Instrument Co., Ltd, China). Watne et al. suggested that the penetration
81 of the particles is close to 100% for particles larger than 100 nm. Hence the wall loss is negligible for the 0.2-2 μm particles
82 that SPAMS measures. Besides, a control experiment measuring the total VOCs at the entrance and the exhaust of the Go:PAM
83 suggested that the gas wall loss was also minor ($6 \pm 4\%$). The particles passed through a diffusion dryer before entering the
84 Go:PAM to reduce the matrix effects from water (Neubauer et al., 1998). The RH at the exhaust of the diffusion dryer was
85 $\sim 15\%$. The residence time of the particles in the dryer was estimated to be 5 s and the particle loss was $\sim 4\%$ according to the
86 CPC measurements. We also collected particles on 47 mm quartz filters (PALL, USA) at the exhaust of the Go:PAM reactor
87 for offline mass and chemical analysis. The number of particles collected on the filters was estimated by the total WCPC
88 counts during the sampling period. The particle number concentration from the WCPC was $6100 \pm 2400 \text{ \# cm}^{-3}$ and the
89 estimated number of collected particles was around 10^8 \# . The filter sample was extracted by deionized (DI) water for analyzing
90 water-soluble ions (e.g., nitrate, formate, potassium) by Ion chromatography (IC) using the same protocol reported in our
91 previous work (Liang et al., 2022). The total organic content (TOC) of the water-extract was analyzed by a TOC analyzer
92 (Shimadzu TOC-L).

93

94 We studied the aging of the particles under 'UV', 'O₃ and dark', and 'O₃ and UV' in the PAM. We named these aged particles
95 UV-aged, O₃-Dark-aged, and O₃-UV-aged, respectively. Although 254 nm is not atmospherically relevant, UV-aged particles
96 are used as a reference in the discussions of the properties of O₃-UV-aged particles. The OH exposure, equivalent to the product
97 of gas-phase OH concentration and residence time, was determined by introducing a stream of SO₂ to the PAM for consuming
98 OH radicals and monitoring the [SO₂] decay, following a well-established approach in the literature (Kang et al., 2007). [SO₂]

99 was almost constant under UV on but without O₃, suggesting that the photochemistry of incense plume does not affect our
100 estimation of OH exposure. The upper limit of OH exposure used in this study varied from 1×10^{10} to 5×10^{10} molecules cm⁻³
101 s, equivalent to 2~10 hours of photochemical aging, assuming an ambient OH concentration of 1.5×10^6 molecules cm⁻³
102 (Mao et al., 2009).

103 2.2 SPAMS analysis

104 A detailed description of the SPAMS can be found in Li et al (Li et al., 2011). After the particle flow exits the PAM reactor,
105 it first passes a PM_{2.5} cyclone to avoid clogging before entering the SPAMS through a 0.1 mm critical orifice at 80 mL min⁻¹
106 flow. Particles achieved a terminal velocity in the supersonic expansion airflow and were detected and aerodynamically sized
107 by two continuous diode Nd: YAG laser beams (532 nm). They were then ionized by a pulsed Nd: YAG laser (266 nm)
108 triggered based on the velocity of a specific particle. The positive and negative ions produced were detected according to the
109 different mass-to-charge ratios (m/z). The energy of the ionization laser was kept at ~0.6 mJ (Cheng et al., 2017). Spectra of
110 more than 3000 individual particles collected for ~15 min were used for further analysis for each experiment. [The size accuracy](#)
111 [of the instrument was routinely calibrated with polystyrene latex spheres of 0.2-2.5 μm diameter \(Nanosphere Size Standards,](#)
112 [Duke Scientific Corp., Palo Alto\)](#). An adaptive resonance theory method (ART-2a) based on MATLAB was used to categorize
113 the incense particles of similar SPAMS spectral characteristics into different particle groups (Phares et al., 2001). [In the ART-](#)
114 [2a analysis, we used a vigilance factor of 0.85, and more than 98% of the particles were classified. Note that there is no general](#)
115 [rule for the vigilance factor in ART-2a. Zhao et al. \(2008\) reported that both a small vigilance factor \(e.g., 0.5 or 0.6\) and a](#)
116 [relatively high vigilance factor \(e.g., 0.8\) show very similar clustering accuracies \(± ~5%\) \(Zhao et al., 2008\).](#)

117 3 Results and discussions

118 3.1 Single-particle characteristics of incense burning particles

119 [The relative peak area \(RPA\), defined as the peak area of a specific peak divided by the total positive or negative mass spectral](#)
120 [peak area, can reflect the relative abundance of particulate components \(Liang et al., 2022\). Note that RPA is not equivalent](#)
121 [to concentration due to the matrix effect. Furthermore, control experiments using different laser fluences indicate that partial](#)
122 [ablation was insignificant \(Table S1, will be discussed later\). The average spectra of the incense burning particles \(Figure 1a\)](#)
123 [are similar to our previous work on incense burning at 50% RH \(Liang et al., 2022\). +39\[K\] dominates the positive spectra,](#)
124 [and organic nitrogen \(ON\) peaks \(i.e., -26\[CN\] and -42\[CN\]\) from nitrogen-containing organics \(NOC\) dominate the](#)
125 [negative spectra \(Zhang et al., 2020; Zhai et al., 2015; Zhang et al., 2021\). These features are also found in biomass burning](#)
126 [particles \(Bi et al., 2011; Peng et al., 2019; Luo et al., 2020\). Despite that particulate polyaromatic hydrocarbons \(PAHs\) were](#)
127 [found in a previous incense combustion study \(Ji et al., 2010\) and a recent study of ambient particles based on a single-particle](#)
128 [mass spectrometer with ART-2a \(Passig et al., 2022\), none of the fresh incense burning particles in our experiments contained](#)

129 the PAHs peaks ($m/z = 178, 189$ (fragment of alkylated phenanthrenes), 202, 220, 228, and 252). Regardless of the presence
130 of PAHs or not, our conclusion on nitrate formation does not depend on the detection of specific chemicals such as PAHs.

131
132 ART-2a categorizes fresh incense burning particles into K-ON, K-ONEC, K-Cl, and OC-ON. EC, Cl and OC are abbreviations
133 of elemental carbon, chloride and organic carbon, respectively. Briefly, the "K" and "OC" before the hyphen indicate the
134 characteristics of the positive spectra, while "ON", "ONEC" and "Cl" after the hyphen indicate the characteristics of the
135 negative spectra. "K-" particles contain a dominant +39 peak and a small +41 peak attributed to isotopic potassium (Silva et
136 al., 1999). On the other hand, the "OC-" particles are rich in typical organic fragments such as +27[C₂H₃] (Silva et al., 2000).
137 According to the negative spectra, "-ON" particles have dominant ON signals. "-ONEC" particles have elemental carbon (EC)
138 peaks of -12n[C_n⁻], with intensities comparable to typical ON peaks (Whiteaker et al., 2002). "-Cl" particles have prominent
139 Cl⁻ ($m/z=-35, -37$ (isotopic)) and KCl₂⁻ ($m/z=-109, -111$ (isotopic)) peaks (Guazzotti et al., 2001; Dall'osto et al., 2004). The
140 average spectra of each category can be found in Figure S6. There are slightly fewer K-ON particles and more K-ONEC
141 particles observed at 80% RH (this work) than at 50% RH in Liang et al (Liang et al., 2022), probably due to the lower organic
142 concentrations at higher RH to limit particle-phase partitioning of volatile organic compounds (Donaldson et al., 2006; Mcfall
143 et al., 2020; Chan et al., 2011; Chan et al., 2010). Overall, the number fraction (NF) of each category is similar to our previous
144 work, with a descending order of K-ON (47.3±5.2%) >> OC-ON (25.7±4.7%) ≈ K-ONEC (20.2±2.8%) > K-Cl (5.1±1.1%)
145 (Figure 1c), reflecting the fresh incense burning particles are organic-rich (Li et al., 2012; Zhang et al., 2022).

146 3.2 Ozonolysis of the incense burning particles

147 Figure 1a also shows the average spectra of aged incense burning particles under 800 ppb O₃. Qualitatively, the major peaks
148 are similar to those in fresh incense burning particles, except for the rise of -62[NO₃⁻] and -46[NO₂⁻], which indicates the
149 formation of nitrate and probably nitrite. The formation of organo-nitrate is not considered significant due to the decreased -
150 26[CN] and -42[CNO].

151
152 To compare the changes in the organic signals, we first excluded all inorganics and EC peaks (Table S2). Control experiments
153 atomizing KNO₃ solution (as K⁺ is the main inorganic cation found in incense burning particles) showed the RPA ratio of -
154 16[O] to nitrate peaks is (6 ± 1.7) % due to fragmentation. Sulfate shows negligible fragmentation under our experimental
155 conditions (Liang et al., 2022). Thus, we subtracted the RPA of -16[O] by 6% RPA of nitrate. Then, we recalculated the RPA
156 of the organic peaks only, defined as "organic spectra". Figure 1b shows the differences in the organic spectra of the aged and
157 fresh particles. The positive difference spectra show an RPA increase in the hydrocarbon +37[C₃H] but an RPA decrease in
158 +51[C₄H₃](Dall'osto et al., 2013). Besides, the increase of +28[CO], +42[C₂H₂O], and +43[C₂H₃O] indicates the formation of
159 oxidized organics in the particles during ozonolysis. The negative difference spectra show a decrease in ON peaks, possibly
160 due to the destruction of C-N bonds under ozonolysis, and an increase of -45[CHO₂] formate peak.

161

162 NO₂ emitted by incense burning may hydrolyze in the incense burning particles to form nitrite and nitrate (2NO₂ + H₂O =>
163 HNO₂ + HNO₃) (Finlayson-Pitts et al., 2003; Ramazan et al., 2006). The uptake of NO₂ is slow in deionized water ($\gamma \approx 10^{-7}$),
164 but it could be significantly promoted by chloride ($\gamma \approx 10^{-3}$ in 1mM NaCl) (Enami et al., 2009; Yabushita et al., 2009), which
165 is found as major inorganic anion in the incense burning particles. In addition, the reaction between NO₂ and O₃ produces NO₃
166 radicals, which could react with organics to form organo-nitrate (Ng et al., 2017). However, control experiments using a
167 charcoal absorber to remove NO_x only show ~20% decrease in RPA of total nitrate in O₃ aged particles (Figure S4), indicating
168 that NO₂ hydrolysis and nitration may not be the main contributor to the nitrate formation. The charcoal absorber removes
169 both NO_x and VOC. However, it was not expected to be important in the observed nitrate RPA reduction, as the content of
170 nitrogen-containing VOCs was minor in the incense burning plume, according to the literature (Manoukian et al., 2013). We
171 categorized the O₃-Dark-aged particles into 7 groups of K-ON, K-ONEC, K-ONN, K-N, K-Cl, OC-N, and OC-ON particles
172 by ART-2a. The definitions of K-, OC-, -ON, -ONEC, -Cl are the same as before. -N particles show prominent nitrate peaks
173 (-46[NO₂] and -62[NO₃]) in the negative spectra, while -ONN particles show comparable ON peaks and nitrate peaks (Figure
174 S6). The NF of different categories descends in the order of K-ON (29.0±0.7%) ≈ K-ONEC (22.8±1.8%) ≈ OC-N (21.9±1.0%)
175 > K-ONN (11.2±0.9%) > OC-ON (7.6±1.7%) > K-Cl (3.9±0.4%) > K-N (2.5±0.5%) (Figure 1c). Interestingly, the K-ONEC
176 and K-Cl NFs are similar before and after aging, whereas the K-ON NF decreased, and the decrease is comparable to the sum
177 of the K-ONN and K-N NF increases. OC-ON was the only fresh OC- particle type, but OC-N was dominant after aging. A
178 control experiment with a HEPA filter before the Go:PAM showed no detectable particles by SPAMS. Thus, we assume the
179 total SPAMS-detectable particle number was constant before and after aging, and O₃-aging may have preferentially converted
180 some -ON type particles to nitrate-containing particles (i.e., -ONN, -N). It has been reported that the HEPA filter would cause
181 the loss of semi-volatile VOC (SVOC) or less-volatile VOC (LVOC) (Schilling, 1997). However, the addition of the charcoal
182 to remove VOCs at the exhaust only caused ~6 % reduction of the NF of -ON type particles and nitrate-containing particles
183 (i.e., -ONN, -N), suggesting the roles of SVOC and LVOC were minor to our conclusion. Besides, O₃-Dark-aged particles
184 have lower ON (i.e., the sum of -26[CN] and -42[CNO]) absolute peak area (APA) and higher total nitrate APA (i.e., -46[NO₂]
185 and -62[NO₃]), than fresh particles (Figure S7). We considered -46[NO₂] mainly as a nitrate fragment but not nitrite since the
186 IC-measured [NO₂⁻]/[NO₃⁻] in the water-extract of collected particles was very low (~0.01) (Figure S8). We used APA here
187 because it reflects the total abundance of ions and would not affect the analysis of other ions (Spencer et al., 2006). The
188 formation of nitrate would increase the total peak area and decrease the RPA of other peaks, even the APA of others kept
189 constant. Figure S9 also shows a positive correlation between the total nitrate RPA and formate RPA in the aged particles.
190 Based on offline IC analysis, the water-extract of O₃-Dark-aged particles has higher [Formate]/[K⁺] and [NO₃⁻]/[K⁺] in than
191 fresh particles (Figure S10), assuming that K⁺ is not reactive and used as an internal standard (Figure S11). Taking these
192 altogether, nitrate and formate likely formed together, preferentially on K-ON and OC-ON particles. Ozonolysis of NOC has
193 been reported to generate nitrate and formate (Sharma et al., 2010; Yao et al., 2020).

194 3.3 Photochemical oxidation of incense burning particles

195 With UV (254nm) on, the 800 ppb O₃ was partly photolyzed to generate OH radicals in the presence of water vapor, resulting
196 in an OH exposure of $\sim 3 \times 10^{10}$ molecules cm⁻³ s, equivalent to a photochemical age of ~ 5 h. We will use xx ppb O₃ (initial
197 concentration) +UV, instead of OH exposure, to describe OH aging. The average spectra of O₃-UV-aged particles are generally
198 similar to that of O₃-Dark-aged particles, with potassium and nitrate peaks dominating the positive and negative spectra,
199 respectively (Figure 1a). However, the RPA of -46[NO₂] and -62[NO₃] were 0.2 and 0.4, around 2 times higher than O₃-Dark-
200 aged particles, likely indicating more nitrate formation. As will be discussed later, photochemistry triggered by light-absorbing
201 compounds such as photosensitizers and Fe salts is a possible source of nitrate formation in O₃-UV-aged particles.⁵¹⁻⁵³
202 However, its contribution is considered minor compared with OH chemistry since UV-aged particles only show a total nitrate
203 RPA of 0.05, much lower than that of O₃-UV-aged particles (~ 0.7 , will be discussed later). Control experiments using a
204 charcoal absorber to remove the NO_x significantly reduced the RPA of total nitrate by $\sim 75\%$ (Figure S4). These suggest that
205 OH chemistry involving NO_x dominated the particulate nitrate formation under OH exposure. Under 800 ppb O₃ and UV, the
206 $\sim 90\%$ reduction of [NO_x] with a simultaneous increase in total nitrate peaks under UV suggests the oxidation of NO_x by OH
207 radicals to form HNO₂ and HNO₃, which can be uptake by the particles afterward (Finlayson-Pitts et al., 1999). Reactive
208 uptake of NO_x initiated by OH chemistry cannot be excluded.

209

210 Similar to the O₃-Dark-aged particles, O₃-UV-aged particles show decreases in ON and other organic peaks (+38[C₃H₂],
211 +50[C₄H₂], and +51[C₄H₃]) in the difference organic averaged spectra (Silva et al., 2000), likely due to oxidative consumption
212 by OH radicals (Figure 1b). The ON peaks decrease in O₃-UV-aged particles was more significant than in O₃-Dark-aged
213 particles, whereas the increase in formate peak is less obvious. These indicate that NOCs can also be effectively degraded via
214 OH oxidation. Using the commonly used general markers of oxidized/aged organics in single-particle mass spectrometric
215 studies of -16[O], -17[OH], +42[C₂H₂O], and +43[C₂H₃O] as examples (Taiwo et al., 2014; Denkenberger et al., 2007; Qin et
216 al., 2006), the RPA increase in O₃-UV-aged particles are 18, 10, 3, and 17 times higher than in O₃-Dark-aged particles. This
217 suggests that OH aging produced more oxidized products than O₃ aging. The difference average organic spectra of UV-aged
218 particles almost showed no noticeable peaks, indicating that the chemistry initiated by particulate photoactive compounds may
219 not be essential to the transformation of the organics (Figure S12).

220

221 The O₃-UV-aged particles can be categorized into K-ONN, K-N, and OC-N, and they generally have more intense nitrate
222 peaks than O₃-Dark-aged particles. Still, "-ONN" particles have comparable ON and nitrate peaks, and "-N" particles have
223 dominant nitrate peaks in the negative spectra (Figure S6). The NF descends in the order of OC-N ($35.7 \pm 7.2\%$) \approx K-N
224 ($35.5 \pm 4.2\%$) $>$ K-ONN ($25.7 \pm 2.1\%$) (Figure 1c). Notably, the NF of OC- particles of O₃-UV-aged particles is 50% larger than
225 the fresh particles, likely due to the formation of additional particulate organics. We could not identify any preferential nitrate
226 formation in specific particle types since most of the particles have high RPA of nitrate. [The collection efficiency of SPAMS](#)

227 increased from ~1% at 200 nm to ~40% at 960 nm (Figure S13). Therefore, secondary aerosol formation in small particles
228 may have been underestimated. However, this would not affect the conclusion of the results that nitrate formed in incense
229 burning particles upon O₃+Dark and O₃+UV aging. A control experiment doubling the laser fluence from 0.6 to 1.2 mJ showed
230 minor differences in the classification of the aged particles and the RPA of total nitrate peaks (Table S1). This suggested that
231 the partial ablation due to the formation of nitrate coating is not significant. It is possible that there were some pure secondary
232 organic particles still not ionized at 1.2 mJ, which resulted in an underestimation of secondary organic aerosol formation.

233 **3.4 The formation of secondary nitrate**

234 Figure 2a shows the RPA of nitrate peaks under UV and different exposure of O₃ and OH. Since fresh particles also have high
235 NF of total nitrate, NF cannot accurately depict the effectiveness of nitrate formation. Fresh incense burning particles exhibit
236 very low RPA of total nitrate, whereas exposure to O₃ increases the RPA from almost 0 to around 0.2, irrespective of the [O₃].
237 Only a slight increase (~0.02) in total nitrate RPA was observed for UV-aged particles. However, with both O₃ and UV on,
238 the RPAs of total nitrate further increased to above 0.7, which is also independent of the initial [O₃]. Consistent with the
239 average spectra shown before, nitrate formation due to OH oxidation is likely more efficient than that by ozonolysis. Under
240 both O₃ and OH exposure, the summed APA of nitrate peaks increased as particle size increased, suggesting possibly a larger
241 total amount of nitrate formed in larger particles (Figure 2b, d). However, the RPA shows an opposite trend, which can be
242 interpreted as lower nitrate concentration in larger particles. Larger particles have larger surfaces but smaller surface-to-volume
243 ratios, which lead to a larger absolute amount of nitrate formed but a lower relative concentration of particulate nitrate (Figure
244 2c, e). It has been reported that the larger organic and inorganic mixed particles could undergo liquid-liquid phase separation
245 (LLPS) more easily than smaller ones (Kucinski et al., 2019). It is possible that the organics are mainly located at the outer
246 layers of the particles, whereas the inorganic components reside in the core. The hydrophobic organic shell may have retarded
247 the uptake of HNO₃/HNO₂/NO_x to form nitrate. Under O₃+UV, it is also possible that comparable HNO₃ was generated under
248 excess [OH] and contributed to the similar total nitrate RPA since the [NO_x] reductions under different OH exposure are
249 similarly high (Figure S14). The insensitivity of nitrate formation to O₃ and OH exposure can be potentially explained by the
250 diffusion limitation of interfacial uptake due to the poor hygroscopicity of fresh incense burning particles (Li et al., 1993;
251 Zaveri et al., 2018; Slade et al., 2014; Liang et al., 2022).

252 **3.5 The Potential formation of SOA**

253 Oxalate and malonate are two major dicarboxylates in atmospheric particles and are considered SOA (Yao et al., 2002). They
254 have been widely studied using single-particle mass spectrometry with well-validated detection efficiency, without severe
255 complications in mass spectra due to fragmentations (Cheng et al., 2017; Sullivan et al., 2007). Figure 3a shows the NF ratio
256 (aged particles to fresh particles) of oxalate and malonate. We used the NF ratio rather than the APA or RPA to avoid large
257 uncertainties in organic abundance due to the much weaker peaks of organics in the spectra, which was potentially due to the

258 matrix effects. The APA threshold for oxalate or malonate was 15 arbitrary unit (a.u.) with a baseline of zero unit, the same as
259 that adopted in an ambient study (Zhu et al., 2020).

260

261 Compared to 300 ppb O₃ and UV, the NF ratios of malonate and oxalate were 30 and 9 folds higher, at 1500 ppb O₃ and UV,
262 respectively. This trend is different from the independence of nitrate formation on OH exposure, probably because the
263 formation of SOA was slower than nitrate via multiphase uptake. These NFs are lower estimates due to the possible degradation
264 by photolysis of Fe-decarboxylate complexes to CO₂ (Gen et al., 2021). While the TOC ratio at different O₃+UV which
265 indicates the formation of SOA, shows an overall trend similar to the NF ratios of oxalate and malonate (Figure S15), the TOC
266 ratio was 1.2-7.1 folds higher than the NF ratios of oxalate and malonate. However, this could be due to the formation of many
267 other species as well as the matrix effects. SOA formation during aging of incense burning plume should be further assessed
268 by other quantitative online instruments in future work. In contrast, no oxalate and malonate were observed during ozonolysis,
269 irrespective of [O₃]. It is likely the SOA characteristics and precursors are different under O₃ and OH. Hence, we are
270 conservative and use the term “potential”, even though oxalate and malonate are commonly found highly oxidized SOA. The
271 TOC ratio of aged to fresh particles extract was higher upon OH oxidation than O₃ oxidation (Figure S16). Furthermore, UV-
272 aged particles did not show an NF increase of both, indicating that the oxalate and malonate formation were mainly due to OH
273 radicals, rather than oxidants from particulate photoactive compounds or ozone. The control experiment with a charcoal
274 absorber shows around 60% and 70% NF reduction of oxalate and malonate, suggesting that the precursors are mainly in the
275 gas phase (Figure S5). The size distribution of oxalate and malonate containing particles skewed towards the larger sizes,
276 supporting their nature of secondary formation (i.e., oxidized gaseous precursors were added to the particles that cause size
277 increase, Figure S17). Figure 3b shows the NF of oxalate and malonate in different categories of the particles. The particles in
278 the replicated experiments under the same OH exposure were combined to compensate for the low particle concentrations.
279 The error bars show one standard deviation among different OH exposures. In descending order, the NF of oxalate and
280 malonate was K-N > K-ONN > OC-N. The mass hygroscopic grow factor (i.e., the mass ratio of wet particles to dry particles)
281 of inorganic potassium salts KNO₃ and KCl at 80% RH are around 1.6 and 2.2 based on AIOMFAC model predictions (Text
282 S1, <https://aiomfac.lab.mcgill.ca/about.html> (Zuend et al., 2008)), much higher than that in the water extract of biomass
283 burning particles (1.1-1.4, including both lab-generated and ambient collected) (Rissler et al., 2006; Carrico et al., 2008; Chan
284 et al., 2005), as well as fresh incense burning particles (around 1) (Liang et al., 2022), which are organic-rich. The likely higher
285 fraction of hygroscopic inorganic of inorganic fraction allows K-N and K-ONN to retain more liquid water to dissolve gaseous
286 SOA precursors for oxalate and malonate formation. The difference of oxalate and malonate NF is statistically significantly
287 different between K-N and OC-N (i.e., P < 0.05 in ANOVA test). There are many other changes in the NF of organic fragments,
288 which suggest the oxidation of primary organics and the formation of SOA. However, further analysis was limited by the lack
289 of molecular information after severe fragmentation. The major spectral evolution and possible peak attribution can be found
290 in Text S2.

291 4 Conclusions

292 In this work, we report the single-particle mixing state characteristics of incense burning particles upon ozonolysis and
293 photochemical oxidation. Formation of secondary aerosol including nitrate and organics was found. This indicates that besides
294 the significant primary emission of particles, additional particulate pollutants could be formed upon atmospheric aging, further
295 worsening the air quality in both outdoor and indoor environments. Nitrate formation initiated by O₃ is generally considered
296 to involve the so-called N₂O₅ pathway, in which oxidation of NO_x forms NO₃ radical and then N₂O₅, which hydrolyzes to
297 form particulate nitrate (Zhao et al., 2021; Xiao et al., 2020). In our study, nitrate formation was found, as indicated by the
298 increase of total nitrate RPA from near 0 to around 0.2, upon O₃ exposure. We propose that ozonolysis of NOCs may be a
299 potential pathway for nitrate formation, in addition to the N₂O₅ pathway. With UV on, ozone was photolyzed to form OH
300 radicals, and we observed a significant increase in total nitrate RPA to above 0.7 at 300 ppb O₃ or above. Nitrate formation in
301 O₃-UV-aged particles is more prominent than in O₃-Dark-aged particles and is attributed to multiphase OH oxidation involving
302 NO_x, such as HNO₃/HNO₂/NO_x uptake (Chen et al., 2020; Lu et al., 2019). At 300 ppb O₃ and UV in this study, the equivalent
303 OH and O₃ exposure time of the incense particles is estimated to be ~2 h and ~10 min, respectively, assuming daytime OH
304 and O₃ concentration of 1.5×10^6 molecules cm⁻³ and 60 ppb (Xia et al., 2021; Mao et al., 2009). Despite the differences in
305 the estimated exposure time for OH and O₃, nitrate formation in incense particles under sunlight can be efficient.

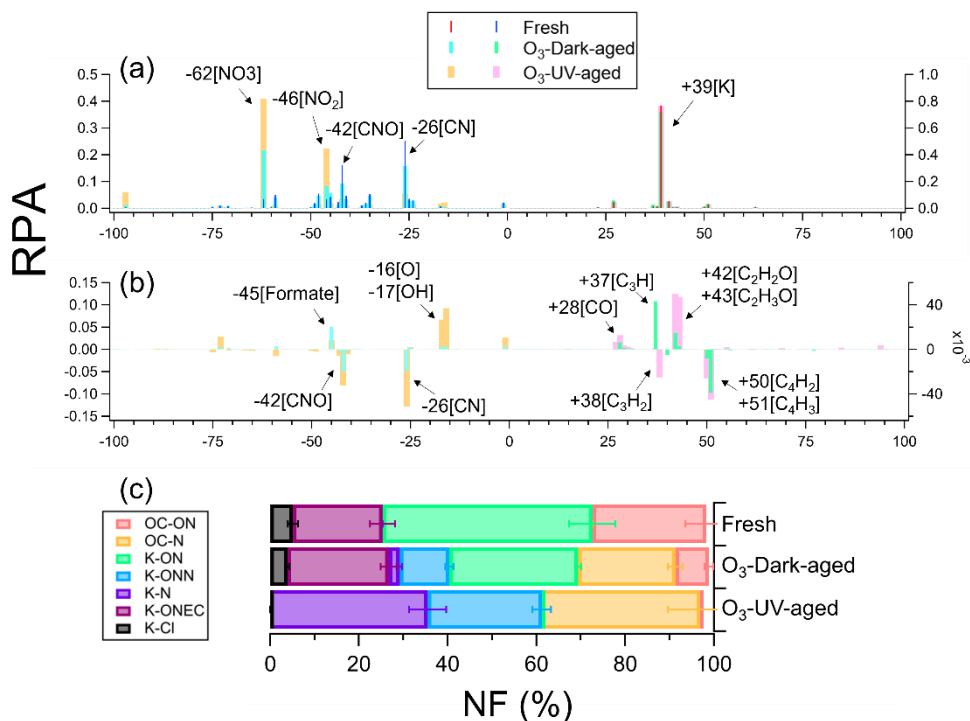
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307 We also observed various changes in organics peaks, though less apparent than nitrate in the average spectra. Overall,
308 oxygenated fragments like +42[C₂H₂O] increase, which indicates functionalization of the organics upon oxidation. The
309 increase of such oxygenated fragments is more significant in O₃-UV-aged than O₃-Dark-aged particles. -26[CN] and -42[CNO]
310 attributed to NOC decreased under O₃ and OH exposure. Apparent formate formation was observed in O₃-Dark-aged particles,
311 likely from the degradation of NOC. Production of formate in O₃-UV-aged particles was less significant than that in O₃-Dark-
312 aged particles, attributed to the photolysis of O₃. Oxalate and malonate were observed in O₃-UV-aged particles but not in O₃-
313 Dark-aged particles, and the NFs increased with OH exposure. Furthermore, oxalate and malonate preferentially formed on K-
314 N particles, followed by K-ONN and then OC-N, indicating a potentially crucial role of aerosol liquid water in SOA formation.

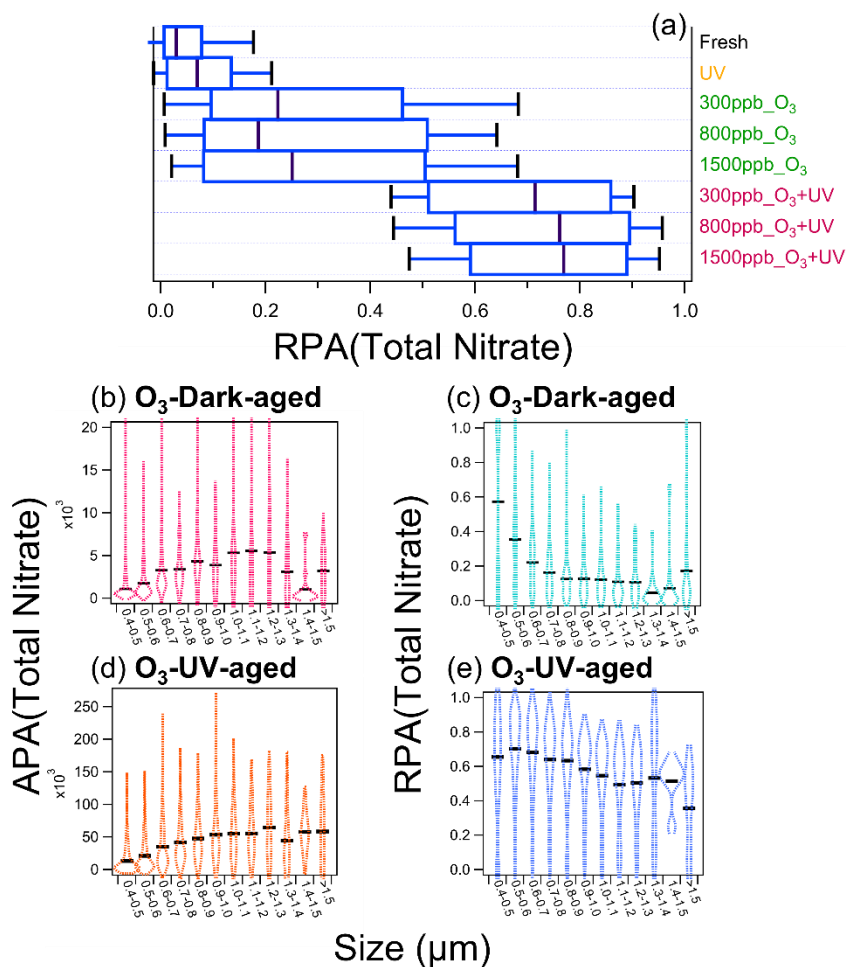
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316 Though the molecular characterization of SOA is beyond the focus of this work, the formation of oxalate and malonate shed
317 light on the SOA formation upon photochemical oxidation of the incense burning plumes. Formate and dicarboxylates are
318 important hygroscopic organics in atmospheric particles, which can potentially act as cloud condensation nuclei (Yao et al.,
319 2002; Peng et al., 2001). Incense burning particles were often used as biomass burning particle surrogates (Li et al., 2012;
320 Schurman et al., 2017; Zhang et al., 2014; Kuwata et al., 2017), due to their similar physicochemical properties and overall
321 composition (Li et al., 2012; Zhang et al., 2022). Our work sheds light on the secondary aerosol formation in biomass burning
322 particles upon exposure of atmospheric oxidants, despite that the detailed composition of incense burning plume may be
323 different from biomass burning plume because of the manufacturing process of incense sticks. For instance, incense burning

324 sticks may contain additives such as adhesives beyond biomass constituents (Lin et al., 2007). Future works are encouraged to
 325 explore the formation mechanism and kinetics of secondary pollutants in the incense burning and biomass burning particles.
 326 Due to the short residence time in a PAM reactor, assumption in the interchangeability of oxidant concentration and reaction
 327 time in estimating total exposure was made. However, Chu et al. (2019) challenged such interchangeability in ozonolysis
 328 reaction of linoleic acid (Chu et al., 2019). Aging at ambient concentrations of oxidants should also be investigated.
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 332 **Figure 1.** (a) The average spectra of fresh and aged incense burning particles at 800 ppb O₃(+UV); (b) The difference (aged
 333 minus fresh) of the average organic spectra of incense burning particles at 800 ppb O₃(+UV). The left axis and right axis are
 334 for negative spectra and positive spectra, respectively. (c) Number fraction of different categories in fresh, O₃-Dark-aged, and
 335 O₃-UV-aged particles.
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337

338 **Figure 2.** (a) The whisker-box plot of total nitrate RPA of fresh and aged particles. The violin plots of (b, d) APA and (c, e)
 339 RPA of total nitrate in O₃- and O₃-UV-aged particles as a function of size (unit: μm) of particles aged at 1500 ppb O₃ (+ UV).
 340 The medians are shown as the lines, and the kernel densities represent the probability density of the data at different values.

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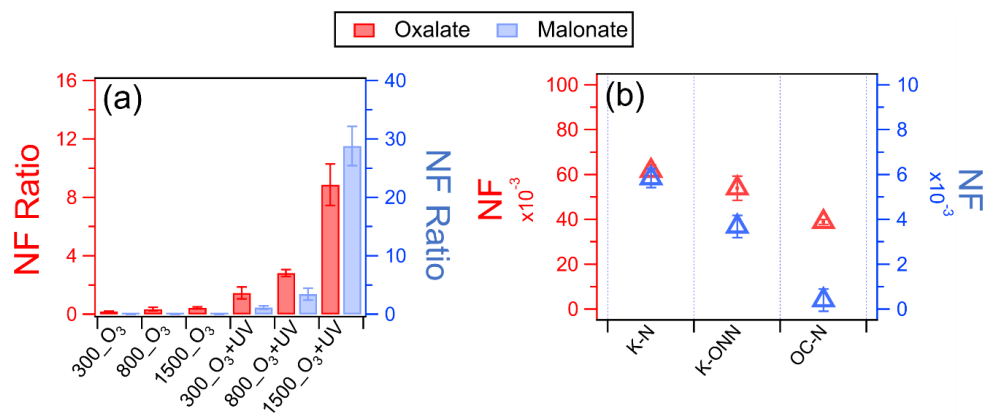
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352 **Figure 3.** (a) The NF ratios of oxalate and malonate under different conditions. 300_O₃ denotes aging at 300 ppb O₃; (b) The
 353 NF of oxalate (left axis) and malonate (right axis) in different categories of aged particles.

354

355

356 *Data availability.* The supplement provides additional figures and tables.

357

358 *Competing interests.* The contact author has declared that neither they nor their co-authors have any competing interests.

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365 References

366 Akagi, S. K., Yokelson, R. J., Wiedinmyer, C., Alvarado, M. J., Reid, J. S., Karl, T., Crouse, J. D., and Wennberg, P. O.:
 367 Emission factors for open and domestic biomass burning for use in atmospheric models, *Atmos. Chem. Phys.*, 11, 4039-4072,
 368 10.5194/acp-11-4039-2011, 2011.

369 Bi, X., Zhang, G., Li, L., Wang, X., Li, M., Sheng, G., Fu, J., and Zhou, Z.: Mixing state of biomass burning particles by single
 370 particle aerosol mass spectrometer in the urban area of PRD, China, *Atmospheric Environment*, 45, 3447-3453,
 371 <https://doi.org/10.1016/j.atmosenv.2011.03.034>, 2011.

372 Bootdee, S., Chantara, S., and Prapamontol, T.: Determination of PM_{2.5} and polycyclic aromatic hydrocarbons from incense
373 burning emission at shrine for health risk assessment, *Atmospheric Pollution Research*, 7, 680-689,
374 <https://doi.org/10.1016/j.apr.2016.03.002>, 2016.

375 Carrico, C. M., Petters, M. D., Kreidenweis, S. M., Collett Jr., J. L., Engling, G., and Malm, W. C.: Aerosol hygroscopicity
376 and cloud droplet activation of extracts of filters from biomass burning experiments, *Journal of Geophysical Research:*
377 *Atmospheres*, 113, <https://doi.org/10.1029/2007JD009274>, 2008.

378 Chan, L. P. and Chan, C. K.: Enhanced Reactive Uptake of Nonanal by Acidic Aerosols in the Presence of Particle-Phase
379 Organics, *Aerosol Science and Technology*, 45, 872-883, 10.1080/02786826.2011.567314, 2011.

380 Chan, L. P., Lee, A. K. Y., and Chan, C. K.: Gas-Particle Partitioning of Alcohol Vapors on Organic Aerosols, *Environmental*
381 *Science & Technology*, 44, 257-262, 10.1021/es9018018, 2010.

382 Chan, M. N., Choi, M. Y., Ng, N. L., and Chan, C. K.: Hygroscopicity of Water-Soluble Organic Compounds in Atmospheric
383 Aerosols: Amino Acids and Biomass Burning Derived Organic Species, *Environmental Science & Technology*, 39, 1555-
384 1562, 10.1021/es049584l, 2005.

385 Chen, X., Hopke, P. K., and Carter, W. P.: Secondary organic aerosol from ozonolysis of biogenic volatile organic compounds:
386 chamber studies of particle and reactive oxygen species formation, *Environmental science & technology*, 45, 276-282, 2011.

387 Chen, X., Wang, H., Lu, K., Li, C., Zhai, T., Tan, Z., Ma, X., Yang, X., Liu, Y., Chen, S., Dong, H., Li, X., Wu, Z., Hu, M.,
388 Zeng, L., and Zhang, Y.: Field Determination of Nitrate Formation Pathway in Winter Beijing, *Environmental Science &*
389 *Technology*, 54, 9243-9253, 10.1021/acs.est.0c00972, 2020.

390 Cheng, C., Li, M., Chan, C. K., Tong, H., Chen, C., Chen, D., Wu, D., Li, L., Wu, C., and Cheng, P.: Mixing state of oxalic
391 acid containing particles in the rural area of Pearl River Delta, China: implications for the formation mechanism of oxalic acid,
392 *Atmospheric Chemistry and Physics*, 17, 9519-9533, 2017.

393 Chu, Y., Cheng, T. F., Gen, M., Chan, C. K., Lee, A. K., and Chan, M. N.: Effect of ozone concentration and relative humidity
394 on the heterogeneous oxidation of linoleic acid particles by ozone: an insight into the interchangeability of ozone concentration
395 and time, *ACS Earth and Space Chemistry*, 3, 779-788, 2019.

396 Chuang, H.-C., Jones, T., Chen, Y., Bell, J., Wenger, J., and Bérubé, K.: Characterisation of airborne particles and associated
397 organic components produced from incense burning, *Analytical and bioanalytical chemistry*, 401, 3095-3102, 2011.

398 Dall'Osto, M., Beddows, D. C. S., Kinnersley, R. P., Harrison, R. M., Donovan, R. J., and Heal, M. R.: Characterization of
399 individual airborne particles by using aerosol time-of-flight mass spectrometry at Mace Head, Ireland, *Journal of Geophysical*
400 *Research: Atmospheres*, 109, <https://doi.org/10.1029/2004JD004747>, 2004.

401 Dall'Osto, M., Ovadnevaite, J., Ceburnis, D., Martin, D., Healy, R. M., O'Connor, I. P., Kourtchev, I., Sodeau, J. R., Wenger,
402 J. C., and O'Dowd, C.: Characterization of urban aerosol in Cork city (Ireland) using aerosol mass spectrometry, *Atmospheric*
403 *Chemistry and Physics*, 13, 4997-5015, 2013.

404 Denkenberger, K. A., Moffet, R. C., Holecek, J. C., Rebotier, T. P., and Prather, K. A.: Real-time, single-particle measurements
405 of oligomers in aged ambient aerosol particles, *Environmental Science & Technology*, 41, 5439-5446, 2007.

406 Donaldson, D. J. and Vaida, V.: The Influence of Organic Films at the Air–Aqueous Boundary on Atmospheric Processes,
407 *Chemical Reviews*, 106, 1445-1461, 10.1021/cr040367c, 2006.

408 Enami, S., Hoffmann, M. R., and Colussi, A.: Absorption of inhaled NO₂, *The Journal of Physical Chemistry B*, 113, 7977-
409 7981, 2009.

410 Finlayson-Pitts, B., Wingen, L., Sumner, A., Syomin, D., and Ramazan, K.: The heterogeneous hydrolysis of NO₂ in
411 laboratory systems and in outdoor and indoor atmospheres: An integrated mechanism, *Physical Chemistry Chemical Physics*,
412 5, 223-242, 2003.

413 Finlayson-Pitts, B. J. and Pitts Jr, J. N.: *Chemistry of the upper and lower atmosphere: theory, experiments, and applications*,
414 Elsevier1999.

415 Gen, M., Zhang, R., and Chan, C. K.: Nitrite/Nitrous Acid Generation from the Reaction of Nitrate and Fe(II) Promoted by
416 Photolysis of Iron–Organic Complexes, *Environmental Science & Technology*, 55, 15715-15723, 10.1021/acs.est.1c05641,
417 2021.

418 Gen, M., Liang, Z., Zhang, R., Go Mabato, B. R., and Chan, C. K.: Particulate nitrate photolysis in the atmosphere,
419 *Environmental Science: Atmospheres*, 10.1039/D1EA00087J, 2022.

420 Goel, A., Mundra, R., and Ola, D.: Examination of Particle Characteristics and Quantification of Emission Factors for Smoke
421 Generated from a Popular Indian Incense Burnt in an Experimental Chamber, in: *Indoor Environmental Quality*, Springer, 77-
422 84, 2020.

423 Guazzotti, S. A., Coffee, K. R., and Prather, K. A.: Continuous measurements of size-resolved particle chemistry during
424 INDOEX-Intensive Field Phase 99, *Journal of Geophysical Research: Atmospheres*, 106, 28607-28627,
425 <https://doi.org/10.1029/2001JD900099>, 2001.

426 Ho, S. S. H. and Yu, J. Z.: Concentrations of formaldehyde and other carbonyls in environments affected by incense burning,
427 *Journal of Environmental Monitoring*, 4, 728-733, 10.1039/B200998F, 2002.

428 Hodshire, A. L., Akherati, A., Alvarado, M. J., Brown-Steiner, B., Jathar, S. H., Jimenez, J. L., Kreidenweis, S. M., Lonsdale,
429 C. R., Onasch, T. B., and Ortega, A. M.: Aging effects on biomass burning aerosol mass and composition: A critical review
430 of field and laboratory studies, *Environmental science & technology*, 53, 10007-10022, 2019.

431 Ji, X., Le Bihan, O., Ramalho, O., Mandin, C., D'Anna, B., Martinon, L., Nicolas, M., Bard, D., and Pairon, J.-C.:
432 Characterization of particles emitted by incense burning in an experimental house, *Indoor Air*, 20, 147-158,
433 <https://doi.org/10.1111/j.1600-0668.2009.00634.x>, 2010.

434 Kang, E., Root, M., Toohey, D., and Brune, W.: Introducing the concept of potential aerosol mass (PAM), *Atmospheric
435 Chemistry and Physics*, 7, 5727-5744, 2007.

436 Khezri, B., Chan, Y. Y., Tiong, L. Y. D., and Webster, R. D.: Annual air pollution caused by the Hungry Ghost Festival,
437 *Environmental Science: Processes & Impacts*, 17, 1578-1586, 10.1039/C5EM00312A, 2015.

438 Kroll, J. H., Ng, N. L., Murphy, S. M., Flagan, R. C., and Seinfeld, J. H.: Secondary organic aerosol formation from isoprene
439 photooxidation, *Environmental science & technology*, 40, 1869-1877, 2006.

440 Kucinski, T. M., Dawson, J. N., and Freedman, M. A.: Size-Dependent Liquid-Liquid Phase Separation in Atmospherically
441 Relevant Complex Systems, *The Journal of Physical Chemistry Letters*, 10, 6915-6920, 2019.

442 Kumar, N. K., Corbin, J. C., Bruns, E. A., Massabó, D., Slowik, J. G., Drinovec, L., Močnik, G., Prati, P., Vlachou, A., and
443 Baltensperger, U.: Production of particulate brown carbon during atmospheric aging of residential wood-burning emissions,
444 *Atmospheric Chemistry and Physics*, 18, 17843-17861, 2018.

445 Kuwata, M. and Lee, W.-C.: 1-octanol-water partitioning as a classifier of water soluble organic matters: Implication for
446 solubility distribution, *Aerosol Science and Technology*, 51, 602-613, 10.1080/02786826.2017.1283004, 2017.

447 Lee, S.-C. and Wang, B.: Characteristics of emissions of air pollutants from burning of incense in a large environmental
448 chamber, *Atmospheric Environment*, 38, 941-951, <https://doi.org/10.1016/j.atmosenv.2003.11.002>, 2004.

449 Li, L., Huang, Z., Dong, J., Li, M., Gao, W., Nian, H., Fu, Z., Zhang, G., Bi, X., Cheng, P., and Zhou, Z.: Real time bipolar
450 time-of-flight mass spectrometer for analyzing single aerosol particles, *International Journal of Mass Spectrometry*, 303, 118-
451 124, <https://doi.org/10.1016/j.ijms.2011.01.017>, 2011.

452 Li, W. and Hopke, P. K.: Initial Size Distributions and Hygroscopicity of Indoor Combustion Aerosol Particles, *Aerosol*
453 *Science and Technology*, 19, 305-316, 10.1080/02786829308959638, 1993.

454 Li, Y. J., Yeung, J. W. T., Leung, T. P. I., Lau, A. P. S., and Chan, C. K.: Characterization of Organic Particles from Incense
455 Burning Using an Aerodyne High-Resolution Time-of-Flight Aerosol Mass Spectrometer, *Aerosol Science and Technology*,
456 46, 654-665, 10.1080/02786826.2011.653017, 2012.

457 Liang, Z., Zhang, R., Gen, M., Chu, Y., and Chan, C. K.: Nitrate Photolysis in Mixed Sucrose–Nitrate–Sulfate Particles at
458 Different Relative Humidities, *The Journal of Physical Chemistry A*, 125, 3739-3747, 10.1021/acs.jpca.1c00669, 2021.

459 Liang, Z., Zhou, L., Infante Cuevas, R. A., Li, X., Cheng, C., Li, M., Tang, R., Zhang, R., Lee, P. K. H., Lai, A. C. K., and
460 Chan, C. K.: Sulfate Formation in Incense Burning Particles: A Single-Particle Mass Spectrometric Study, *Environmental*
461 *Science & Technology Letters*, 9, 718-725, 10.1021/acs.estlett.2c00492, 2022.

462 Lin, T.-C., Yang, C.-R., and Chang, F.-H.: Burning characteristics and emission products related to metallic content in incense,
463 *Journal of hazardous materials*, 140, 165-172, 2007.

464 Liu, T., Wang, Z., Huang, D. D., Wang, X., and Chan, C. K.: Significant Production of Secondary Organic Aerosol from
465 Emissions of Heated Cooking Oils, *Environmental Science & Technology Letters*, 5, 32-37, 10.1021/acs.estlett.7b00530,
466 2018.

467 Liu, T., Zhou, L., Liu, Q., Lee, B. P., Yao, D., Lu, H., Lyu, X., Guo, H., and Chan, C. K.: Secondary Organic Aerosol Formation
468 from Urban Roadside Air in Hong Kong, *Environmental Science & Technology*, 53, 3001-3009, 10.1021/acs.est.8b06587,
469 2019.

470 Lu, K., Fuchs, H., Hofzumahaus, A., Tan, Z., Wang, H., Zhang, L., Schmitt, S. H., Rohrer, F., Bohn, B., Broch, S., Dong, H.,
471 Gkatzelis, G. I., Hohaus, T., Holland, F., Li, X., Liu, Y., Liu, Y., Ma, X., Novelli, A., Schlag, P., Shao, M., Wu, Y., Wu, Z.,
472 Zeng, L., Hu, M., Kiendler-Scharr, A., Wahner, A., and Zhang, Y.: Fast Photochemistry in Wintertime Haze: Consequences
473 for Pollution Mitigation Strategies, *Environ Sci Technol*, 53, 10676-10684, 10.1021/acs.est.9b02422, 2019.

474 Luo, J., Zhang, J., Huang, X., Liu, Q., Luo, B., Zhang, W., Rao, Z., and Yu, Y.: Characteristics, evolution, and regional
475 differences of biomass burning particles in the Sichuan Basin, China, *Journal of Environmental Sciences*, 89, 35-46,
476 <https://doi.org/10.1016/j.jes.2019.09.015>, 2020.

477 Lyu, X., Huo, Y., Yang, J., Yao, D., Li, K., Lu, H., Zeren, Y., and Guo, H.: Real-time molecular characterization of air
478 pollutants in a Hong Kong residence: Implication of indoor source emissions and heterogeneous chemistry, *Indoor Air*, 31,
479 1340-1352, <https://doi.org/10.1111/ina.12826>, 2021.

480 Manoukian, A., Quivet, E., Temime-Roussel, B., Nicolas, M., Maupetit, F., and Wortham, H.: Emission characteristics of air
481 pollutants from incense and candle burning in indoor atmospheres, *Environmental Science and Pollution Research*, 20, 4659-
482 4670, 10.1007/s11356-012-1394-y, 2013.

483 Mao, J., Ren, X., Brune, W., Olson, J., Crawford, J., Fried, A., Huey, L., Cohen, R., Heikes, B., and Singh, H.: Airborne
484 measurement of OH reactivity during INTEX-B, *Atmospheric Chemistry and Physics*, 9, 163-173, 2009.

485 McFall, A. S., Johnson, A. W., and Anastasio, C.: Air–water partitioning of biomass-burning phenols and the effects of
486 temperature and salinity, *Environmental Science & Technology*, 54, 3823-3830, 2020.

487 Neubauer, K. R., Johnston, M. V., and Wexler, A. S.: Humidity effects on the mass spectra of single aerosol particles,
488 *Atmospheric Environment*, 32, 2521-2529, 1998.

489 Ng, N. L., Brown, S. S., Archibald, A. T., Atlas, E., Cohen, R. C., Crowley, J. N., Day, D. A., Donahue, N. M., Fry, J. L., and
490 Fuchs, H.: Nitrate radicals and biogenic volatile organic compounds: oxidation, mechanisms, and organic aerosol, *Atmospheric*
491 *chemistry and physics*, 17, 2103-2162, 2017.

492 Passig, J., Schade, J., Irsig, R., Kröger-Badge, T., Czech, H., Adam, T., Fallgren, H., Moldanova, J., Sklorz, M., and Streibel,
493 T.: Single-particle characterization of polycyclic aromatic hydrocarbons in background air in northern Europe, *Atmospheric*
494 *Chemistry and Physics*, 22, 1495-1514, 2022.

495 Peng, C. and Chan, C. K.: The water cycles of water-soluble organic salts of atmospheric importance, *Atmospheric*
496 *Environment*, 35, 1183-1192, 2001.

497 Peng, X., Liu, X., Shi, X., Shi, G., Li, M., Liu, J., Huangfu, Y., Xu, H., Ma, R., Wang, W., and Feng, Y.: Source apportionment
498 using receptor model based on aerosol mass spectra and 1 h resolution chemical dataset in Tianjin, China, *Atmospheric*
499 *Environment*, 198, 387-397, <https://doi.org/10.1016/j.atmosenv.2018.11.018>, 2019.

500 Phares, D. J., Rhoads, K. P., Wexler, A. S., Kane, D. B., and Johnston, M. V.: Application of the ART-2a algorithm to laser
501 ablation aerosol mass spectrometry of particle standards, *Analytical Chemistry*, 73, 2338-2344, 2001.

502 Qin, X. and Prather, K. A.: Impact of biomass emissions on particle chemistry during the California Regional Particulate Air
503 Quality Study, *International Journal of Mass Spectrometry*, 258, 142-150, <https://doi.org/10.1016/j.ijms.2006.09.004>, 2006.

504 Ramazan, K., Wingen, L. M., Miller, Y., Chaban, G. M., Gerber, R. B., Xantheas, S. S., and Finlayson-Pitts, B. J.: New
505 experimental and theoretical approach to the heterogeneous hydrolysis of NO₂: Key role of molecular nitric acid and its
506 complexes, *The Journal of Physical Chemistry A*, 110, 6886-6897, 2006.

507 Rissler, J., Vestin, A., Swietlicki, E., Fisch, G., Zhou, J., Artaxo, P., and Andreae, M. O.: Size distribution and hygroscopic
508 properties of aerosol particles from dry-season biomass burning in Amazonia, *Atmos. Chem. Phys.*, 6, 471-491, 10.5194/acp-
509 6-471-2006, 2006.

510 Rudich, Y., Donahue, N. M., and Mentel, T. F.: Aging of organic aerosol: Bridging the gap between laboratory and field
511 studies, *Annu. Rev. Phys. Chem.*, 58, 321-352, 2007.

512 Schilling, J. B.: Extraction of semivolatile organic compounds from high-efficiency particulate air (HEPA) filters by
513 supercritical carbon dioxide, Argonne National Lab., Analytical Chemistry Lab., IL (United States), 1997.

514 Schurman, M. I., Kim, J. Y., Cheung, H. H., and Chan, C. K.: Atmospheric particle composition-hygroscopic growth
515 measurements using an in-series hybrid tandem differential mobility analyzer and aerosol mass spectrometer, *Aerosol Science
516 and Technology*, 51, 694-703, 2017.

517 See, S. W. and Balasubramanian, R.: Characterization of fine particle emissions from incense burning, *Building and
518 Environment*, 46, 1074-1080, <https://doi.org/10.1016/j.buildenv.2010.11.006>, 2011.

519 Seinfeld, J. and Pandis, S.: *Atmospheric chemistry and physics*. 1997, New York, 2008.

520 Sharma, V. K. and Graham, N. J. D.: Oxidation of Amino Acids, Peptides and Proteins by Ozone: A Review, *Ozone: Science
521 & Engineering*, 32, 81-90, 10.1080/01919510903510507, 2010.

522 Sidibe, A., Sakamoto, Y., Murano, K., Sato, K., Yuba, A., Futami, M., Koita, O. A., Traore, I., and Kajii, Y.: Chemical
523 Characterization and Health Risk Assessment of Particulate Matter from Household Activities in Bamako, Mali, Western Sub-
524 Saharan Africa, *Atmosphere*, 13, 1290, 2022.

525 Silva, P. J. and Prather, K. A.: Interpretation of Mass Spectra from Organic Compounds in Aerosol Time-of-Flight Mass
526 Spectrometry, *Analytical Chemistry*, 72, 3553-3562, 10.1021/ac9910132, 2000.

527 Silva, P. J., Liu, D.-Y., Noble, C. A., and Prather, K. A.: Size and Chemical Characterization of Individual Particles Resulting
528 from Biomass Burning of Local Southern California Species, *Environmental Science & Technology*, 33, 3068-3076,
529 10.1021/es980544p, 1999.

530 Slade, J. H. and Knopf, D. A.: Multiphase OH oxidation kinetics of organic aerosol: The role of particle phase state and relative
531 humidity, *Geophysical Research Letters*, 41, 5297-5306, 2014.

532 Spencer, M. T. and Prather, K. A.: Using ATOFMS to determine OC/EC mass fractions in particles, *Aerosol Science and*
533 *Technology*, 40, 585-594, 2006.

534 Sullivan, R. C. and Prather, K. A.: Investigations of the diurnal cycle and mixing state of oxalic acid in individual particles in
535 Asian aerosol outflow, *Environmental Science Technology*, 41, 8062-8069, 2007.

536 Taiwo, A. M., Harrison, R. M., Beddows, D. C., and Shi, Z.: Source apportionment of single particles sampled at the
537 industrially polluted town of Port Talbot, United Kingdom by ATOFMS, *Atmospheric Environment*, 97, 155-165, 2014.

538 Volkamer, R., Jimenez, J. L., San Martini, F., Dzepina, K., Zhang, Q., Salcedo, D., Molina, L. T., Worsnop, D. R., and Molina,
539 M. J.: Secondary organic aerosol formation from anthropogenic air pollution: Rapid and higher than expected, *Geophysical*
540 *Research Letters*, 33, 2006.

541 Wang, B., Lee, S. C., and Ho, K. F.: Chemical composition of fine particles from incense burning in a large environmental
542 chamber, *Atmospheric Environment*, 40, 7858-7868, <https://doi.org/10.1016/j.atmosenv.2006.07.041>, 2006.

543 Whiteaker, J. R., Suess, D. T., and Prather, K. A.: Effects of Meteorological Conditions on Aerosol Composition and Mixing
544 State in Bakersfield, CA, *Environmental Science & Technology*, 36, 2345-2353, 10.1021/es011381z, 2002.

545 Xia, N., Du, E., Guo, Z., and de Vries, W.: The diurnal cycle of summer tropospheric ozone concentrations across Chinese
546 cities: Spatial patterns and main drivers, *Environmental Pollution*, 286, 117547, <https://doi.org/10.1016/j.envpol.2021.117547>,
547 2021.

548 Xiao, H. W., Zhu, R. G., Pan, Y. Y., Guo, W., Zheng, N. J., Liu, Y. H., Liu, C., Zhang, Z. Y., Wu, J. F., and Kang, C. A.:
549 Differentiation between nitrate aerosol formation pathways in a southeast Chinese city by dual isotope and modeling studies,
550 *Journal of Geophysical Research: Atmospheres*, 125, e2020JD032604, 2020.

551 Yabushita, A., Enami, S., Sakamoto, Y., Kawasaki, M., Hoffmann, M., and Colussi, A.: Anion-catalyzed dissolution of NO₂
552 on aqueous microdroplets, *The Journal of Physical Chemistry A*, 113, 4844-4848, 2009.

553 Yao, X., Fang, M., and Chan, C. K.: Size distributions and formation of dicarboxylic acids in atmospheric particles,
554 *Atmospheric Environment*, 36, 2099-2107, 2002.

555 Yao, Y., Xie, Y., Zhao, B., Zhou, L., Shi, Y., Wang, Y., Sheng, Y., Zhao, H., Sun, J., and Cao, H.: N-dependent ozonation
556 efficiency over nitrogen-containing heterocyclic contaminants: A combined density functional theory study on reaction
557 kinetics and degradation pathways, *Chemical Engineering Journal*, 382, 122708, <https://doi.org/10.1016/j.cej.2019.122708>,
558 2020.

559 Ye, C., Chen, R., and Chen, M.: The impacts of Chinese Nian culture on air pollution, *Journal of Cleaner Production*, 112,
560 1740-1745, 2016.

561 Zaveri, R. A., Shilling, J. E., Zelenyuk, A., Liu, J., Bell, D. M., D'Ambro, E. L., Gaston, C. J., Thornton, J. A., Laskin, A., and
562 Lin, P.: Growth kinetics and size distribution dynamics of viscous secondary organic aerosol, *Environmental science &*
563 *technology*, 52, 1191-1199, 2018.

564 Zhai, J., Wang, X., Li, J., Xu, T., Chen, H., Yang, X., and Chen, J.: Thermal desorption single particle mass spectrometry of
565 ambient aerosol in Shanghai, *Atmospheric Environment*, 123, 407-414, 2015.

566 Zhang, G., Lian, X., Fu, Y., Lin, Q., Li, L., Song, W., Wang, Z., Tang, M., Chen, D., Bi, X., Wang, X., and Sheng, G.: High
567 secondary formation of nitrogen-containing organics (NOCs) and its possible link to oxidized organics and ammonium, *Atmos.*
568 *Chem. Phys.*, 20, 1469-1481, 10.5194/acp-20-1469-2020, 2020.

569 Zhang, S., Wang, Z., Zhang, J., Guo, D., and Chen, Y.: Inhalable cigarette-burning particles: Size-resolved chemical
570 composition and mixing state, *Environmental Research*, 202, 111790, 2021.

571 Zhang, X., Xu, J., Zhai, L., and Zhao, W.: Characterization of Aerosol Properties from the Burning Emissions of Typical
572 Residential Fuels on the Tibetan Plateau, *Environmental Science & Technology*, 56, 14296-14305, 10.1021/acs.est.2c04211,
573 2022.

574 Zhang, Y., Zhang, X., Sun, J., Hu, G., Shen, X., Wang, Y., Wang, T., Wang, D., and Zhao, Y.: Chemical composition and
575 mass size distribution of PM 1 at an elevated site in central east China, *Atmospheric Chemistry and Physics*, 14, 12237-12249,
576 2014.

577 Zhao, W., Hopke, P. K., and Prather, K. A.: Comparison of two cluster analysis methods using single particle mass spectra,
578 *Atmospheric Environment*, 42, 881-892, <https://doi.org/10.1016/j.atmosenv.2007.10.024>, 2008.

579 Zhao, Z.-Y., Cao, F., Fan, M.-Y., Zhai, X.-Y., Yu, H.-R., Hong, Y., Ma, Y.-J., and Zhang, Y.-L.: Nitrate aerosol formation
580 and source assessment in winter at different regions in Northeast China, *Atmospheric Environment*, 267, 118767, 2021.

581 Zhu, S., Li, L., Wang, S., Li, M., Liu, Y., Lu, X., Chen, H., Wang, L., Chen, J., Zhou, Z., Yang, X., and Wang, X.: Development
582 of an automatic linear calibration method for high-resolution single-particle mass spectrometry: improved chemical species
583 identification for atmospheric aerosols, *Atmos. Meas. Tech.*, 13, 4111-4121, 10.5194/amt-13-4111-2020, 2020.

584 Zuend, A., Marcolli, C., Luo, B. P., and Peter, T.: A thermodynamic model of mixed organic-inorganic aerosols to predict
585 activity coefficients, *Atmos. Chem. Phys.*, 8, 4559-4593, 10.5194/acp-8-4559-2008, 2008.

586