



1     **Unambiguous identification of N-containing oxygenated organic molecules**  
2                     **using CI-Orbitrap in an eastern Chinese megacity**

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## ABSTRACT

13

14 Oxygenated organic molecules (OOMs) are dominated by the N-containing species in polluted urban  
15 environment. As N-containing OOMs, especially those with more than one nitrogen atoms, prevailed in  
16 the high  $m/z$  range ( $m/z > 350$ Th), unambiguous identification of N-containing OOMs is highly desirable  
17 for understanding of their formation processes, precursors and influencing factors. To achieve this, we  
18 applied an ultra-high-resolution chemical ionization-orbitrap (CI-Orbitrap) in a field campaign and found  
19 that OOMs contain one (1N-OOMs), two (2N-OOMs) and three (3N-OOMs) nitrogen atoms respectively  
20 comprised 50%, 26% and 4% of total OOMs. More interestingly, the fraction of 2N-OOMs increased with  
21 the increase of carbon number ( $n_C$ ) and were dominated by the ones derived from aliphatic precursors  
22 (2N-OOM<sub>Ali</sub>, 64.2%), indicating the importance of multistep oxidation. Plausible precursors of 2N-OOMs  
23 were aliphatics (2N-OOM<sub>Ali</sub>, 64.2%), aromatics (2N-OOM<sub>Aro</sub>, 16%), and monoterpenes (2N-OOM<sub>MT</sub>,  
24 15.4%). The 2N-OOM<sub>Ali</sub> was the most abundant 2N-OOMs and its fraction even increased in the polluted  
25 day with enhanced proportion of the ones with  $n_C > 10$ . While 2N-OOM<sub>Ali</sub> and 2N-OOM<sub>Aro</sub> were  
26 dominated by daytime photochemical production, nighttime NO<sub>3</sub>-initiated oxidation dominated the  
27 formation of 2N-OOM<sub>MT</sub>. 2N-OOM<sub>Aro</sub> were of highest oxygenation level, followed by 2N-OOM<sub>MT</sub> and  
28 2N-OOM<sub>Ali</sub>, which were affected by photochemistry and NO<sub>x</sub> concentrations. These results highlight the  
29 significant formation of 2N-OOMs and the influencing factors, on their formation in polluted urban  
30 environment, where various VOC precursors and atmospheric oxidants present.



## 31 **1. Introduction**

32 Secondary organic aerosol (SOA) accounts for a significant fraction of particulate matters (Donahue  
33 et al., 2009; Ehn et al., 2014; Hallquist et al., 2009; Jimenez et al., 2009). Volatile organic compounds  
34 (VOCs) and their oxidation products, *i.e.*, OVOCs, are important precursors of SOA in the atmosphere  
35 (Atkinson and Arey, 2003; Bianchi et al., 2019; Ehn et al., 2014; Nie et al., 2022). The N-containing  
36 oxygenated organic molecules (OOMs) have been identified as the important products upon VOC  
37 oxidation. Especially at high NO<sub>x</sub> levels, these products become more dominant while the others  
38 (*i.e.*, alcohols, hydroperoxides and RO<sub>2</sub> cross-reaction products) are likely suppressed (Bianchi et al., 2019;  
39 Zhao et al., 2018). The nitrogen atoms in OOM molecules are assumed to be mainly associated with nitrate  
40 group (-ONO<sub>2</sub>) formed from bi-molecular reaction between RO<sub>2</sub> radical and NO. Field measurements also  
41 observed that up to 54 % of molecules in organic aerosol (OA) contain nitrate functional groups under  
42 different atmosphere (Ditto et al., 2020; Lee et al., 2016; Lin et al., 2021; Rollins et al., 2013; Xu et al.,  
43 2015; Ye et al., 2021; Yu et al., 2019).

44 The N-containing OOM molecules can be classified into 1N-OOMs, 2N-OOMs, and 3N-OOMs,  
45 according to the number of nitrate functional group in the molecule. The chemical composition of N-  
46 containing OOMs is determined by their precursors, formation pathways and NO<sub>x</sub> level in the atmosphere  
47 (Bianchi et al., 2019; Ehn et al., 2014; Nie et al., 2022; Pye et al., 2019; Riva, 2016; Yan et al., 2016).  
48 Recent observations in megacities of China indicated that 2N-OOMs accounted for significant fractions  
49 (about 30-33%) among total N-containing OOMs besides 1N-OOMs (66-70%) due to the high NO<sub>x</sub>  
50 concentrations in polluted urban environment (Nie et al., 2022; Yan et al., 2021). Some laboratory studies  
51 also proposed that the potential formation pathways of 2N-OOMs, such as the multiple-step OH oxidation  
52 (Garmash et al., 2020) or the NO<sub>3</sub>-initiated oxidation followed by NO termination (Kiendler-Scharr et al.,  
53 2016; Liebmann et al., 2019), suggesting the increased importance of multi-step bimolecular oxidation in  
54 the formation of 2N-OOMs. On the other hand, it was also found that the formation of 2N-OOMs showed  
55 the clear preference of specific precursors compared to 1N-OOMs, *i.e.*, significantly higher branch ratio  
56 of 2N-OOMs from aliphatic hydrocarbons than those from aromatics (Nie et al., 2022), suggesting



57 considerable difference from 1N-OOMs in terms of formation pathway. Determining the formation  
58 pathway of N-containing OOM molecules, especially those containing two to three nitrogen atoms, in real  
59 atmosphere, is challenging. Identification of their chemical compositions in molecular level is the key for  
60 advancing our understanding in the precursor, formation and sources of N-containing OOMs in polluted  
61 atmosphere, where thousands of oxidation products exist and evolve constantly.

62 Traditionally, a chemical ionization atmospheric pressure interface time-of-flight mass spectrometer  
63 (CI-APi-TOF) has been used to measure the gaseous OOMs (Berndt et al., 2016; Ehn et al., 2014; Jokinen  
64 et al., 2014; Rissanen et al., 2014). Using a CI-APi-TOF, increasing number of studies have been reporting  
65 the formation of OOMs through the oxidation of various VOC precursors in chamber or flow tube (Berndt  
66 et al., 2016, 2018; Ehn et al., 2014; Garmash et al., 2020; Jokinen et al., 2014, 2015; Rissanen et al., 2014;  
67 Wang et al., 2020; Zhao et al., 2018). While 2N-OOMs in real ambient almost exclusively located in high  
68  $m/z$  (mass-to-charge) range (*i.e.*, 300 - 500 Th), a CI-APi-TOF with highest mass resolution of 12,000  
69 ( $m/\Delta m$ , in full width at half maximum), can hardly identify the molecular compositions of 2N-OOMs  
70 unambiguously. This is because low mass resolving power imposes significant uncertainties on separating  
71 overlapping peaks, which increases rapidly with increasing  $m/z$  and decreasing mass resolution. Taken the  
72 integer  $m/z$  of 342 as an example, multiple peaks overlap at this nominal mass, *i.e.*,  $C_7H_8O_{10}N_2(NO_3)^-$   
73 (342.0057 Th),  $C_8H_{12}O_9N_2(NO_3)^-$  (342.0421 Th),  $C_9H_{16}O_8N_2(NO_3)^-$  (342.0785 Th),  $C_{10}H_{20}O_7N_2(NO_3)^-$   
74 (342.1149 Th). The adjacent peaks are of mass differences ( $\Delta m$ ) of 0.0364 and a good peak separation of  
75 these peaks ( $4\sigma$ ) requires mass resolutions of at least 16,000. Therefore, development and application of  
76 mass spectrometry techniques with extremely high performance in detection limit, time resolution, and  
77 mass resolution, are highly desirable.

78 To achieve accurate identification of molecular formula from the extremely complex mass spectra, a  
79 CI (nitrate) inlet had also been coupled to an orbitrap mass spectrometer (CI-Orbitrap) to measure the  
80 OOMs in ultra-high mass resolution ( $m/\Delta m > 100,000$ ) (Riva et al., 2019a; Zhang et al., 2022). The ultra-  
81 high mass resolving power of CI-Orbitrap will undoubtedly provide significant improvements in  
82 molecular identification, separation, and quantification. Herein, we applied a CI-Orbitrap in a field



83 campaign for the measurements of OOMs, with a special focus on 2N-OOMs, in molecular level in urban  
84 Shanghai. The site represents a typical eastern Chinese megacity characterized by intense human activities,  
85 multiple anthropogenic emissions and high NO<sub>x</sub> concentrations. Based on the measurement results as well  
86 as our current knowledge on N-containing OOM formation, we classify the observed 2N-OOMs into  
87 different precursor groups and explore the potential influencing factors on their formation. Furthermore,  
88 supported by positive matrix factorization (PMF), sources and gas-phase oxidation processes for 2N-OOM  
89 formation in urban Shanghai were identified.

## 90 **2. Ambient measurement and methodology**

### 91 **2.1 Measurements**

92 The field campaign was carried out from 31<sup>th</sup> October to 18<sup>th</sup> November, 2020 on the top-floor of an  
93 8-story building in Shanghai Academy of Environmental Sciences (31° 18' N, 121° 43' E), which sits in a  
94 densely populated region surrounded by commercial properties and residential dwellings without  
95 significant industrial sources nearby. The site can represent a typical urban area of Shanghai affected by  
96 severe local emissions from vehicular traffic, commercial, and residential activities. Our campaign was  
97 carried out in autumn which represents a typical transition period from strong photochemistry in summer  
98 to intense regional transport in winter. At times, air masses transported from the neighboring provinces or  
99 even further from the northern China can also affect the air quality of the site.

100 The 2N-OOMs as well other OOMs molecules were measured in real time with a nitrate-Orbitrap.  
101 The operation of nitrate-Orbitrap has been detailed in previous studies as well as in one of our companion  
102 studies (Zhang et al., 2022), thus is only briefly described here. Ambient air was drawn into the ionization  
103 source through a 1m stainless-steel tube (3/4 inch). The reagent ion was produced by passing nitric acid  
104 in sheath flow (20L/min) into a PhotoIonizer (Model L9491, Hamamatsu, Japan) and was then introduced  
105 into a co-axial laminar flow reactor, in which the reagent ions interact with the air samples. The charged  
106 species were detected by an orbitrap mass analyzer with a mass resolution of about 140,000. Mass-  
107 dependent transmission calibrations was also performed using a depletion method (Heinritzi et al., 2016).  
108 Other ancillary measurements, including the PM<sub>2.5</sub> concentrations, trace gases (SO<sub>2</sub>, O<sub>3</sub> and NO<sub>x</sub>), volatile



109 organic compounds, as well as meteorological parameters (wind direction and speed, solar radiation, etc.)  
110 were detailed in Supporting Information (SI-1). An overview of the measurement data, illustrating the air  
111 quality as well as the meteorological conditions during the campaign, is provided in SI-2 and Figure S1.

## 112 **2.2 Data analysis of nitrate CI-Orbitrap**

113 The raw mass spectra were first extracted by Orbitool (Cai et al., 2020) and the molecular information  
114 was then achieved by applying a homemade toolkit based on the MATLAB software. The toolkit drew on  
115 the idea from “tofTools” package which is used for analyzing the mass spectral data obtained from the  
116 TOF analyzer, such as nitrate CI-API-TOF (Junninen et al., 2010). The concentrations of the detected  
117 species are then determined as follows:

$$118 \quad [X] = \frac{i[X^-]}{NO_3^- \cdot (HNO_3)_{0-2}} \cdot C \quad (1)$$

119 where  $i[X^-]$  is the transmission-corrected signal intensity of ion  $X$  in unit of counts per second (cps),  $C$   
120 represents the calibration factor.  $C$  is determined from the collision frequency of target species with the  
121 nitrate ions (cluster) during its residence in the charger, taking into account of the losses onto the walls of  
122 the reactor and the tube (Eq. 2):

$$123 \quad C = C_{H_2SO_4} = \frac{1}{k_{ion} \times RT \times f_{inlet}} \quad (2)$$

124 where  $k_{ions}$  is the ion collision frequency in the range of  $(1.7 - 2.3) \times 10^{-9} \text{ cm}^3 \text{ s}^{-1}$  (Ehn et al., 2014);  $RT$  is  
125 the residence time in the charger and  $f_{inlet}$  represents the fractions of target species that passed through the  
126 inlet.

127 Herein, we apply the  $C$  determined for sulfuric acid ( $H_2SO_4$ ) of  $3.4 \times 10^9 \text{ molecules cm}^{-3} \text{ ncps}^{-1}$  to  
128 semi-quantify the concentrations of OOMs, which is widely used in previous studies. Among the low  
129 volatility vapors, it had been demonstrated that nitrate ions exhibit highest charging efficiency toward  
130  $H_2SO_4$  (Ehn et al., 2014; Hyttinen et al., 2015, 2018; Riva et al., 2019b). The estimated concentrations of  
131 OOMs thus can be considered as the lower limits with an uncertainty of  $\pm 50\%$  according to error  
132 propagation. Positive matrix factorization (PMF) was also performed for the measured species using  
133 Source Finder (SoFi, v6.3) based on Igor and run by the multilinear engine (ME-2) as detailed in SI-S3  
134 and Figure S3-S6 (Canonaco et al., 2013).



### 135 **3. Results and discussion**

#### 136 **3.1 Chemical characteristic of OOMs**

137 In total, we have identified 562 OOMs, which concentrated in the nC range of 5 to 10, taking up 84.6%  
138 of total OOMs. Possible precursors of C<sub>5-10</sub> OOMs include isoprene (C<sub>5</sub>), benzene/alkyl benzenes (C<sub>6-10</sub>),  
139 aliphatic VOCs (C<sub>5-10</sub>) and monoterpene (C<sub>10</sub>) according to previous studies (Bianchi et al., 2019; Nie et  
140 al., 2022). C<sub>≤4</sub> OOMs only took up a small fraction of 6.7% among total OOMs and were likely a result  
141 of the decomposition from OOMs with large carbon numbers as suggested by one of our companion  
142 studies. The rest 8.7% were C<sub>>10</sub> OOMs which accounted for a dominating fraction (70%, Figure S2)  
143 among the extremely low-volatility organic compounds (ELVOC, C\* < 3 × 10<sup>-5</sup> μg m<sup>-3</sup>) (Donahue et al.,  
144 2011, 2012; Schervish and Donahue, 2020), and potentially poses larger impacts on SOA formation owing  
145 to their lower volatility.

146 We further classified the detected OOMs into four groups based on the number of nitrate groups they  
147 possessed, including non-nitrogen (0N-) OOMs, 1N-OOMs, 2N-OOMs, and 3N-OOMs. Their average  
148 fractional contributions to total OOM concentrations as well as the carbon number (nC) distributions are  
149 shown in Figure 1. We found 1N-OOMs dominated the total OOM concentration with an average fraction  
150 of 50%, followed by 2N-OOMs (26%), demonstrating the dominance of N-containing OOMs among total  
151 OOMs. The 3N-OOMs only took up a small fraction (4%) of total OOMs and the rest 20% was 0N-OOMs.

152 More interestingly, we found 1N-OOMs prevailed among the OOMs with nC ≤ 10, yet 2N-OOMs  
153 dominated the C<sub>>10</sub> OOMs (41.8-84.2%) and its fraction increased stepwise with the increase of nC (Figure  
154 1b). This is likely because more carbon atoms in the molecular skeleton can potentially provide more  
155 active sites for multi-step oxidation, followed by NO termination. Our observation suggests the increased  
156 importance of multi-step bimolecular oxidation in the formation of 2N-OOMs with nC > 10, which is also  
157 consistent with those reported in recent laboratory studies during the oxidation of aromatic and long chain  
158 alkanes with the presence of NO. Furthermore, these 2N-OOMs with nC > 10 had an average molecular  
159 composition of C<sub>12.5</sub>H<sub>22.7</sub>O<sub>2.1</sub>(NO<sub>3</sub>)<sub>2</sub>. Assuming the nitrogen atoms are only associated with nitrate group  
160 (-ONO<sub>2</sub>), the mean double bond equivalent value (DBE) (Nie et al., 2022; Xu et al., 2021) was 1.15 on



161 the carbon skeleton, suggesting its origination from aliphatic compounds, such as alkanes, alkenes, etc.  
162 (Gong et al., 2005; Mentel et al., 2015; Wang and Hildebrandt Ruiz, 2018).

163 We thus further classified the 2N-OOMs to their possible VOC precursors following a recently  
164 developed workflow based on the up-to-date understanding of VOC oxidation as well as PMF results (Nie  
165 et al., 2022), *i.e.*, aromatics (2N-OOM<sub>Aro</sub>), aliphatics (2N-OOM<sub>Ali</sub>), and monoterpene (2N-OOM<sub>MT</sub>). Note  
166 that we group isoprene 2N-OOMs (2N-OOM<sub>Iso</sub>) into 2N-OOM<sub>Ali</sub> as well because of the low concentration  
167 of isoprene in cold season. As a result, the average fractions of 2N-OOM<sub>Aro</sub>, 2N-OOM<sub>Ali</sub> and 2N-OOM<sub>MT</sub>  
168 among total 2N-OOMs were 16.0%, 64.2% and 15.3%, respectively (Figure 2), suggesting significant  
169 contribution of aliphatic compounds to 2N-OOMs formation. Taken together, the increased fraction of  
170 2N-OOMs with the increase of nC and the dominant fraction of 2N-OOM<sub>Ali</sub> highlight the significant  
171 contribution of high-molecular-weight aliphatic precursors (*i.e.*, intermediate volatility or semi-volatile  
172 organic compounds, I/SVOCs) to high-molecular-weight 2N-OOM formation, which were potentially  
173 important SOA material. We thus focus our attention on the formation of 2N-OOMs in the following  
174 sections.

### 175 3.2 2N-OOM formation in PM episodes

176 To investigate the formation mechanisms and factors that may affect the 2N-OOM formation, one  
177 clean day (4<sup>th</sup>-5<sup>th</sup> November) and one polluted day (7<sup>th</sup>-8<sup>th</sup> November) based on the pollution levels, *i.e.*,  
178 PM<sub>2.5</sub> concentrations, were selected for further analysis. Since OOM formation is directly mediated by  
179 photochemistry or nighttime chemistry, the clean and polluted cases were thus split into one clean daytime  
180 case (CL<sub>day</sub>), one clean nighttime case (CL<sub>night</sub>), one polluted daytime case (PL<sub>day</sub>) and one polluted  
181 nighttime case (PL<sub>night</sub>). Detailed information on durations, pollution levels, meteorological conditions  
182 and 2N-OOM concentrations during these four cases were summarized in Table 1.

183 During the whole campaign, the concentrations of 2N-OOMs ranged from  $1.1 \times 10^6$  to  $42.0 \times 10^6$   
184 molecule cm<sup>-3</sup> as shown in Figure 2. We found the concentrations of 2N-OOMs in the polluted cases were  
185 1.7-2.7 times higher than those in clean cases (Table 1). Among the 2N-OOMs, 2N-OOM<sub>Ali</sub> occupied the  
186 largest fractions, which were even higher in polluted cases (66-66%) than those in clean cases (56-61%),





187 Figure 2). Especially for the 2N-OOM<sub>Ali</sub> with  $nC > 10$ , its concentration in polluted cases increased by a  
188 factor of 2.3-4.8 compared to the clean cases (Figure 3). From PMF analysis, we also identified a factor  
189 characterized by a series of 2N-OOM<sub>Ali</sub> (*i.e.*,  $C_nH_{2n-2}O_8N_2$ ,  $n=5-11$ ) as the fingerprint molecules (Table  
190 S1). This factor tracks the PM<sub>2.5</sub> concentration well especially during PM episodes (Figure S7), likely due  
191 to the availability of adequate aliphatic precursors during pollution episode. Furthermore, 2N-OOM<sub>Ali</sub>  
192 with  $nC > 10$  presented both higher concentrations and fractions during daytime than nighttime cases  
193 (Figure 3), suggesting that the photochemical formation of 2N-OOM<sub>Ali</sub> prevailed compared to nighttime  
194 formation. To compare CL<sub>night</sub> and PL<sub>night</sub>, it was also found that the pollution case would lead enhanced  
195 importance of nighttime formation pathways of 2N-OOM<sub>Ali</sub> with  $nC > 10$ .

196 We note that the fraction of 2N-OOM<sub>Ali</sub> increased during CL<sub>night</sub> primarily due to the more evident  
197 decrease of 2N-OOM<sub>Aro</sub> (Table 1), whose formation is dominated by photochemistry. On the other hand,  
198 the decrease of 2N-OOM<sub>Aro</sub> concentrations at PL<sub>night</sub> was not as obvious as those on CL<sub>night</sub>. Due to the  
199 significant increase of 2N-OOM<sub>Ali</sub> concentration, the fraction of 2N-OOM<sub>Aro</sub> decreased in pollution cases,  
200 but their absolute concentrations only had few changes in the daytime. 2N-OOM<sub>MT</sub> showed overall higher  
201 concentrations during nighttime than those in daytime, suggesting the importance of nighttime chemistry  
202 in 2N-OOM<sub>MT</sub> formation, which will be discussed in later sections.

203 To summarize, the formation of 2N-OOM<sub>Ali</sub> were significantly promoted by pollution levels, whereas  
204 the 2N-OOM<sub>Aro</sub> were predominantly affected by photochemical production, whose formation was less  
205 sensitive to pollution levels compared to 2N-OOM<sub>Ali</sub>. In contrast, the concentrations of 2N-OOM<sub>MT</sub>  
206 significantly influenced by pollution levels and seem not solely/almost depend on daytime/nighttime  
207 formation pathway. In addition, we note that both daytime photochemistry and nighttime chemistry had  
208 profound effects on 2N-OOMs formation under different pollution levels, presumably depending on  
209 availability of the precursors as well as the oxidants. We thus focus our attention on the formation of 2N-  
210 OOMs during daytime versus nighttime in the following sections.

### 211 3.3 Daytime vs. nighttime formation of 2N-OOMs

212 We thus investigate the effects of photochemistry and nighttime chemistry on the formation of



213 individual 2N-OOMs. While the former is dominated by OH radical oxidation, the latter involves NO<sub>3</sub>  
214 radical oxidation as well as reactions with ozone or other oxidants, *e.g.*, halogen. Herein, we use solar  
215 radiation as a proxy of photochemical reactivity, and the concentrations of NO<sub>3</sub> radical were estimated  
216 assuming that NO<sub>3</sub>, NO<sub>2</sub> and N<sub>2</sub>O<sub>5</sub> were under fast equilibration in troposphere (Brown and Stutz, 2012).  
217 The correlation coefficients (Spearman type) between individual 2N-OOM molecules and solar radiation  
218 ( $R_{2N-OOMs-solar}$ ) or NO<sub>3</sub> radical ( $R_{2N-OOMs-NO3}$ ) derived from different precursors during the whole campaign  
219 were shown in Figure 4a.

220 Both 2N-OOM<sub>Aro</sub> and 2N-OOM<sub>Ali</sub> showed stronger correlations with solar radiation over NO<sub>3</sub>  
221 radicals, indicating their association with daytime photochemistry since benzene/alkyl benzenes and  
222 aliphatic VOCs rapidly react with OH radicals compared to other oxidants, such as NO<sub>3</sub> radicals. This is  
223 also supported by the observation that both 2N-OOM<sub>Aro</sub> and 2N-OOM<sub>Ali</sub> peaked during noontime (12:00-  
224 14:00) as shown in Figure 4b. Similarly, the PMF analysis also distinguished two daytime factors. The  
225 daytime factor-1 peaked at around 12:00-14:00 (Table S1) and highly correlated with solar radiation  
226 ( $R=0.57$ ). The fingerprint-molecules of daytime factor-1 are C<sub>n</sub>H<sub>2n-4</sub>O<sub>10</sub>N<sub>2</sub> ( $n=8-10$ ) with average DBE  
227 values of 2 on the carbon skeleton, suggesting the dominance of 2N-OOMs likely formed from aromatic  
228 precursors. Since each step of OH oxidation of aromatics followed by RO<sub>2</sub>+NO<sub>x</sub> termination would  
229 increase the nH by one, this factor is likely dominated by 2N-OOMs formed from two steps of OH-  
230 initiated oxidation from alkylbenzenes given the carbon numbers ranged from 8 to 10.

231 The key fingerprint molecule of daytime factor-2 is C<sub>n</sub>H<sub>2n</sub>O<sub>8</sub>N<sub>2</sub> ( $n=4-5$ ) (accounting for 30.8% in the  
232 factor profile), followed by C<sub>n</sub>H<sub>2n-2</sub>O<sub>8</sub>N<sub>2</sub> ( $n=5-6$ ) (accounting for 9.7% in the factor profile), which is  
233 likely a result of the decomposition from 2N-OOM<sub>Ali</sub> with large carbon numbers, according to their DBE  
234 values of 0-1 on the carbon skeleton. This aliphatic factor presented even higher correlation with solar  
235 radiation ( $R=0.65$ ), peaking at around 12:00-14:00. Strong daytime peaks together with the good  
236 correlations with irradiation suggest the dominance of photochemical oxidation in the formation of 2N-  
237 OOM<sub>Ali</sub>. For 2N-OOM<sub>Ali</sub>, although it showed strong daytime peak, weak nighttime peak was still observed.  
238 This indicates that although daytime formation of 2N-OOM<sub>Ali</sub> prevails, their nighttime formation still



239 existed. For example, we have obtained a nighttime factor from PMF analysis (nighttime factor-2), whose  
240 fingerprint molecules are  $C_5H_8O_9N_2$  and  $C_nH_{2n}O_7N_2$  ( $n=5-8$ ).  $C_5H_8O_9N_2$  was likely originated from  
241 isoprene and  $C_nH_{2n}O_7N_2$  were likely from anthropogenic aliphatic precursors.

242 Besides the nighttime formation of  $2N-OOM_{Ali}$ , nighttime chemistry also dominated the formation  
243 of  $2N-OOM_{MT}$ , which showed higher concentrations during the nighttime compared to daytime. This is  
244 further supported by the stronger correlation between  $2N-OOM_{MT}$  and  $NO_3$  radicals than solar radiation.  
245 Formation of  $2N-OOM_{MT}$  is likely a result of  $NO_3$  radical initiated oxidation. As shown in Figure 5, we  
246 have identified a series of  $2N-OOM_{MT}$  molecules with molecular composition of  $C_{10}H_{16}O_{7,9,11}N_2$ , which  
247 showed strong positive correlations with  $NO_3$  radical. The occurrence of propagation reaction from  $RO_2$   
248 to  $RO$  was critical to the formation of odd oxygen as proposed in previous chamber studies (Boyd et al.,  
249 2015; Clafin and Ziemann, 2018). Therefore, formation of  $C_{10}H_{16}O_{7,9,11}N_2$  likely started with the reaction  
250 of monoterpene with  $NO_3$  radicals forming a  $NO_3-C_{10}H_{16}$  alkyl radical, followed by the formation of  
251 organic peroxy radical ( $RO_2$ ) upon addition of  $O_2$ . The  $RO_2$  is then converted to an alkoxy radical ( $RO$ )  
252 upon reaction with  $NO$ . The autoxidation process would then start and stepwise introduce  $O_2$  into the  
253 molecule, forming a series of more oxygenated  $RO_2$  radicals, *i.e.*,  $NO_3-C_{10}H_{16}(O)(OO)_n$ . The  $NO$   
254 termination reaction of these  $RO_2$  radicals would finally result in  $ON$ s with chemical composition of  $NO_3-$   
255  $C_{10}H_{16}(O)(OO)_nO(NO)O$  ( $n=0, 1, 2$ ).

256 On the other hand, the reaction rate between monoterpenes (*i.e.*, alpha-pinene, beta-pinene and  
257 limonene) and  $NO_3$  are about 60,000-140,000 times faster than that between monoterpenes and  $O_3$  at 293K  
258 (MCMv3.1), but the averaged nighttime concentrations of  $O_3$  (22.8 ppb) was only about 18,000 times  
259 higher than that of  $NO_3$  (1.3 ppt). Therefore,  $NO_3$ -initiated oxidation process posed significant impacts on  
260  $2N-OOM_{MT}$  formation during nighttime. The  $2N-OOM_{MT}$  resulted from  $NO_3$  oxidation is also resolved as  
261 a nighttime factor (nighttime factor-1) from PMF analysis, which tracked the  $NO_3$  concentrations very  
262 well (Figure S8) and peaked at around 19:00-23:00. Fingerprint molecule of nighttime factor-1 was mainly  
263 including  $C_{10}H_{16}O_9N_2$  and  $C_{10}H_{16}O_8N_2$ , which is generated from  $NO_3$ -initiated oxidation followed by  $NO$   
264 termination and this process will not change the  $nH$  of the parent monoterpene molecule.



### 265 3.4 Oxygenation level of 2N-OOMs

266 We then calculated the average effective oxygen number ( $nO_{\text{eff}} = nO - 2nN$ ) of 2N-OOMs, which is  
267 used to indicate the oxidation state of carbon by excluding the oxygen atoms bonded with nitrogen atoms.  
268 Note that calculation of  $nO_{\text{eff}}$  assumes that the nitrogen atoms are only associated with nitrate group (-  
269  $ONO_2$ ), which is reasonable after excluding nitrophenol peaks. The average  $nO_{\text{eff}}$  of 2N-OOMs from  
270 different precursors in  $CL_{\text{day}}$ ,  $CL_{\text{night}}$ ,  $PL_{\text{day}}$  and  $PL_{\text{night}}$  were shown in Figure 6 and summarized in Table  
271 S2.  $2N\text{-OOM}_{\text{Aro}}$  had the highest  $nO_{\text{eff}}$  (4.8-5.6), followed by  $2N\text{-OOM}_{\text{MT}}$  (4.5-4.9) and  $2N\text{-OOM}_{\text{Ali}}$  were  
272 of lowest  $nO_{\text{eff}}$  (3.9-4.0). Difference in the oxygenation level of different types of OOMs can be attributed  
273 to the difference in oxidation mechanisms of the initiation reactions. For example, the OH-initiated  
274 oxidation of alkanes, aromatics and monoterpene/alkenes would form a  $C_xH_yO_2$  radical,  $C_xH_yO_3$  radical  
275 and  $C_xH_yO_3$  radical, respectively, incorporating different number of oxygen atoms into the original  
276 precursor molecules at the first step of oxidation (MCMv3.1). On the other hand, during the multiple-step  
277 oxidation in daytime, aromatics could still provide more C=C bonds than other precursors after the initial  
278 step which is plausibly capable to further react with OH,  $O_3$  and others oxidants.

279 Furthermore, we also found that regardless of the pollution level, the  $nO_{\text{eff}}$  was considerably higher  
280 in daytime cases than that in nighttime cases particularly for  $2N\text{-OOM}_{\text{Aro}}$  and  $2N\text{-OOM}_{\text{MT}}$ , suggesting a  
281 profound effects of photochemistry on the formation of highly oxygenated 2N-OOMs. This is likely  
282 because of the high  $NO_x$  concentrations during the nighttime (Table 1), which could efficiently suppress  
283 the  $RO_2$  radicals from autoxidation reactions forming overall less oxygenated OOM molecules. The effect  
284 of  $NO_x$  on oxygenation levels would be discussed in a subsequent paragraph. The average  $nO_{\text{eff}}$  of 2N-  
285  $OOM_{\text{Ali}}$  in four sub-periods were similar without significant daytime and nighttime difference, ranging  
286 from 3.9-4.0. This could be partly explained by the fact that reactions with oxidants such as OH and  
287 halogen radicals will similarly result in the addition of oxygen atoms by two for alkanes during the first  
288 step of oxidation. Thus, the oxygenation levels of  $2N\text{-OOM}_{\text{Ali}}$  were supposed to be insensitive to the  
289 oxidants in daytime or nighttime.

290 It is known that  $NO_x$  is also critical in determining the fate of  $RO_2$  radical during the oxidation,



291 forming RO radicals or organonitrates. Formation of RO radicals and organonitrates will have opposite  
292 effects on the oxidation state of the termination products since the former will significantly increase the  
293 oxygenation state of carbon through initiating propagation reactions before termination. We thus explore  
294 the effects of  $\text{NO}_x$  on the average oxygenation levels of 2N-OOMs from different precursors during the  
295 whole campaign (Figure 7). Consistent with previous studies in polluted urban environment (Qiao et al.,  
296 2021; Yan et al., 2021), the detected 2N-OOMs were also of low oxygenation with  $n\text{O}_{\text{eff}}$  of 3.9-5.4 (25-  
297 75% percentile) compared to those measured in forest or in laboratory studies (Berndt et al., 2016; Ehn et  
298 al., 2014; Jokinen et al., 2014; Rissanen et al., 2014; Yan et al., 2016).  $n\text{O}_{\text{eff}}$  of  $2\text{N-OOM}_{\text{Aro}}$  and  $2\text{N-}$   
299  $\text{OOM}_{\text{MT}}$  increased with the decrease of  $\text{NO}_x$  concentrations. This is likely due to the prevailing of  $\text{NO}_x$   
300 termination reactions because the maximum autoxidation rate constant of alkylbenzenes with long-chain  
301 substituents (e.g., isopropyl-benzene, ethyl-benzene) and monoterpene are comparable to the bimolecular  
302 reaction rate between  $\text{RO}_2$  and  $\text{NO}$  (Bianchi et al., 2019). The oxygenation levels of  $2\text{N-OOM}_{\text{Ali}}$  appears  
303 to be insensitive to the pollution levels and  $\text{NO}_x$  concentrations, which should be further investigated in  
304 future studies.

#### 305 **4 Conclusion**

306 We report the unambiguous identification of 2N-OOMs as well as other OOMs using an ultra-high-  
307 resolution orbitrap coupled with a nitrate inlet. We found that OOMs distributed in a wide range of carbon  
308 numbers ( $n\text{C} = 4 - 16$ ), among which the 2N-OOMs occupied a considerable fraction (26%) of the total  
309 observed OOMs. During the whole campaign, the 2N-OOM concentrations ranged from  $1.1 \times 10^6$  to  
310  $42.0 \times 10^6$  molecule  $\text{cm}^{-3}$  and concentrated in the  $n\text{C}$  range of 5 to 10 with high molecular weight ( $m/z >$   
311  $350$  Th), suggesting their low volatilities and thus potentially high contribution to local SOA formation.

312 Aliphatic, aromatics, and monoterpenes were plausible precursors of 2N-OOMs with a fraction of  
313 64.2%, 16% and 15.4%, respectively. The  $2\text{N-OOM}_{\text{Ali}}$  was found to be the most abundant 2N-OOMs and  
314 its fraction even increased in the polluted day with enhanced proportion of ones with  $n\text{C} > 10$ , probably  
315 due to the high concentrations of aliphatic precursors accompanied with PM episodes. Significant  
316 contribution of long-chain aliphatic compounds ( $n\text{C} > 10$ ) to 2N-OOM formation is also supported by the



317 observation that 2N-OOM fraction increased with the increase of nC and they are of low DBE values,  
318 likely through multistep bimolecular oxidation. 2N-OOM<sub>Ali</sub> and 2N-OOM<sub>Aro</sub> mainly peaked in daytime  
319 and showed stronger correlations with solar radiation over NO<sub>3</sub> radicals, indicating their association with  
320 daytime photochemistry since benzene/alkyl benzenes and aliphatic hydrocarbons rapidly react with OH  
321 radicals compared with other oxidants, such as NO<sub>3</sub> radicals. In contrast, 2N-OOM<sub>MT</sub> prevailed during  
322 nighttime and also showed higher correlations with estimated NO<sub>3</sub> radical concentrations than solar  
323 radiation, suggesting that NO<sub>3</sub>-initiated oxidation of monoterpene dominated the formation of 2N-OOM<sub>MT</sub>  
324 during nighttime even under the polluted urban environment, *i.e.*, high NO<sub>x</sub>. In terms of oxygenation levels,  
325 we found that 2N-OOM<sub>Aro</sub> had highest averaged nO<sub>eff</sub> followed by 2N-OOM<sub>MT</sub>. Daytime photochemistry  
326 and low NO<sub>x</sub> concentrations had profound effects on the formation of more oxygenated 2N-OOMs. 2N-  
327 OOM<sub>Ali</sub> had the lowest nO<sub>eff</sub> and had negligible changes under different pollution levels. These results  
328 demonstrate the preference of 2N-OOM formation and the influencing factors in a Chinese megacity  
329 involving various VOC precursors (biogenic VOCs such as monoterpene and anthropogenic VOCs such  
330 as aromatics, aliphatic hydrocarbons) and various atmospheric oxidants (such as OH radical and NO<sub>3</sub>  
331 radicals), and highlight the influence of PM episode.

### 332 **Code/Data availability**

333 Data presented in this paper are available upon request to the corresponding author.

### 334 **Author contributions**

335 CH designed this study. YL, YM, DDH, SL, SJ, and YG conducted the field campaign. YL analyzed  
336 data with contributions from CH and all the other co-authors. YL wrote the manuscript with contributions  
337 from all the other co-authors.

### 338 **Competing interests**

339 The authors declare that they have no conflict of interest.

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## 343 Reference

- 344 Atkinson, R. and Arey, J.: Atmospheric Degradation of Volatile Organic Compounds, *Chem. Rev.*, 103(3),  
345 4605–4638, doi:10.1021/cr0206420, 2003.
- 346 Berndt, T., Richters, S., Jokinen, T., Hyttinen, N., Kurtén, T., Otkjær, R. V., Kjaergaard, H. G., Stratmann,  
347 F., Herrmann, H., Sipilä, M., Kulmala, M. and Ehn, M.: Hydroxyl radical-induced formation of  
348 highly oxidized organic compounds, *Nat. Commun.*, 7(13677), doi:10.1038/ncomms13677, 2016.
- 349 Berndt, T., Scholz, W., Mentler, B., Fischer, L., Herrmann, H., Kulmala, M. and Hansel, A.: Accretion  
350 Product Formation from Self- and Cross-Reactions of RO<sub>2</sub> Radicals in the Atmosphere, *Angew.*  
351 *Chemie - Int. Ed.*, 57(14), 3820–3824, doi:10.1002/anie.201710989, 2018.
- 352 Bianchi, F., Kurtén, T., Riva, M., Mohr, C., Rissanen, M. P., Roldin, P., Berndt, T., Crouse, J. D.,  
353 Wennberg, P. O., Mentel, T. F., Wildt, J., Junninen, H., Jokinen, T., Kulmala, M., Worsnop, D. R.,  
354 Thornton, J. A., Donahue, N., Kjaergaard, H. G. and Ehn, M.: Highly Oxygenated Molecules (HOM)  
355 from Gas-Phase Autoxidation Involving Organic Peroxy Radicals: A Key Contributor to  
356 Atmospheric Aerosol, *Chem. Rev.*, 119, 3472–3509, doi:10.1021/acs.chemrev.8b00395, 2019.
- 357 Boyd, C. M., Sanchez, J., Xu, L., Eugene, A. J., Nah, T., Tuet, W. Y., Guzman, M. I. and Ng, N. L.:  
358 Secondary organic aerosol formation from the  $\beta$ -pinene+NO<sub>3</sub> system: Effect of humidity and peroxy  
359 radical fate, *Atmos. Chem. Phys.*, 15(13), 7497–7522, doi:10.5194/acp-15-7497-2015, 2015.
- 360 Brown, S. S. and Stutz, J.: Nighttime radical observations and chemistry, *Chem. Soc. Rev.*, 41, 6405–  
361 6447, doi:10.1039/c2cs35181a, 2012.
- 362 Cai, R., Li, Y., Clément, Y., Li, D., Dubois, C., Fabre, M., Besson, L., Perrier, S., George, C., Ehn, M.,  
363 Huang, C., Yi, P., Ma, Y. and Riva, M.: Orbitool: A software tool for analyzing online Orbitrap mass  
364 spectrometry data, *Atmos. Meas. Tech.*, 14, 2377–2387, doi:10.5194/amt-2020-267, 2020.
- 365 Canonaco, F., Crippa, M., Slowik, J. G., Baltensperger, U. and Prévôt, A. S. H. H.: SoFi, an IGOR-based  
366 interface for the efficient use of the generalized multilinear engine (ME-2) for the source  
367 apportionment: ME-2 application to aerosol mass spectrometer data, *Atmos. Meas. Tech.*, 6(12),  
368 3649–3661, doi:10.5194/amt-6-3649-2013, 2013.





- 369 Claflin, M. S. and Ziemann, P. J.: Identification and Quantitation of Aerosol Products of the Reaction of  
370  $\beta$ -Pinene with NO<sub>3</sub> Radicals and Implications for Gas- and Particle-Phase Reaction Mechanisms, *J.*  
371 *Phys. Chem. A*, 122(14), 3640–3652, doi:10.1021/acs.jpca.8b00692, 2018.
- 372 Ditto, J. C., Joo, T., Slade, J. H., Shepson, P. B., Ng, N. L. and Gentner, D. R.: Nontargeted Tandem Mass  
373 Spectrometry Analysis Reveals Diversity and Variability in Aerosol Functional Groups across  
374 Multiple Sites, Seasons, and Times of Day, *Environ. Sci. Technol. Lett.*, 7(2), 60–69,  
375 doi:10.1021/acs.estlett.9b00702, 2020.
- 376 Donahue, N. M., Robinson, A. L. and Pandis, S. N.: Atmospheric organic particulate matter: From smoke  
377 to secondary organic aerosol, *Atmos. Environ.*, 43(1), 94–106, doi:10.1016/j.atmosenv.2008.09.055,  
378 2009.
- 379 Donahue, N. M., Epstein, S. A., Pandis, S. N. and Robinson, A. L.: A two-dimensional volatility basis set:  
380 1. organic-aerosol mixing thermodynamics, *Atmos. Chem. Phys.*, 11(7), 3303–3318,  
381 doi:10.5194/acp-11-3303-2011, 2011.
- 382 Donahue, N. M., Kroll, J. H., Pandis, S. N. and Robinson, A. L.: A two-dimensional volatility basis set-  
383 Part 2: Diagnostics of organic-aerosol evolution, *Atmos. Chem. Phys.*, 12(2), 615–634,  
384 doi:10.5194/acp-12-615-2012, 2012.
- 385 Ehn, M., Thornton, J. A., Kleist, E., Sipilä, M., Junninen, H., Pullinen, I., Springer, M., Rubach, F.,  
386 Tillmann, R., Lee, B., Lopez-Hilfiker, F., Andres, S., Acir, I.-H. H., Rissanen, M., Jokinen, T.,  
387 Schobesberger, S., Kangasluoma, J., Kontkanen, J., Nieminen, T., Kurtén, T., Nielsen, L. B.,  
388 Jørgensen, S., Kjaergaard, H. G., Canagaratna, M., Maso, M. D., Berndt, T., Petäjä, T., Wahner, A.,  
389 Kerminen, V.-M. M., Kulmala, M., Worsnop, D. R., Wildt, J. and Mentel, T. F.: A large source of  
390 low-volatility secondary organic aerosol, *Nature*, 506(7489), 476–479, doi:10.1038/nature13032,  
391 2014.
- 392 Garmash, O., Rissanen, M. P., Pullinen, I., Schmitt, S., Kausiala, O., Tillmann, R., Zhao, D., Percival, C.,  
393 Bannan, T. J., Priestley, M., Hallquist, A. M., Kleist, E., Kiendler-Scharr, A., Hallquist, M., Berndt,  
394 T., McFiggans, G., Wildt, J., Mentel, T. F. and Ehn, M.: Multi-generation OH oxidation as a source





- 395 for highly oxygenated organic molecules from aromatics, *Atmos. Chem. Phys.*, 20(1), 515–537,  
396 doi:10.5194/acp-20-515-2020, 2020.
- 397 Gong, H., Matsunaga, A. and Ziemann, P. J.: Products and mechanism of secondary organic aerosol  
398 formation from reactions of linear alkenes with NO<sub>3</sub> Radicals, *J. Phys. Chem. A*, 109(19), 4312–  
399 4324, doi:10.1021/jp058024l, 2005.
- 400 Hallquist, M., Wenger, J. C., Baltensperger, U., Rudich, Y., Simpson, D., Claeys, M., Dommen, J.,  
401 Donahue, N. M., George, C., Goldstein, A. H., Hamilton, J. F., Herrmann, H., Hoffmann, T., Iinuma,  
402 Y., Jang, M., Jenkin, M. E., Jimenez, J. L., Kiendler-Scharr, A., Maenhaut, W., McFiggans, G.,  
403 Mentel, T. F., Monod, A., Prévôt, A. S. H., Seinfeld, J. H., Surratt, J. D., Szmigielski, R. and Wildt,  
404 J.: The formation, properties and impact of secondary organic aerosol: Current and emerging issues,  
405 *Atmos. Chem. Phys.*, 9(14), 5155–5236, doi:10.5194/acp-9-5155-2009, 2009.
- 406 Heinritzi, M., Simon, M., Steiner, G., Wagner, A. C., K<sup>??</sup>rten, A., Hansel, A. and Curtius, J.:  
407 Characterization of the mass-dependent transmission efficiency of a CIMS, *Atmos. Meas. Tech.*, 9(4),  
408 1449–1460, doi:10.5194/amt-9-1449-2016, 2016.
- 409 Hyttinen, N., Kupiainen-Määttä, O., Rissanen, M. P., Muuronen, M., Ehn, M. and Kurtén, T.: Modeling  
410 the Charging of Highly Oxidized Cyclohexene Ozonolysis Products Using Nitrate-Based Chemical  
411 Ionization, *J. Phys. Chem. A*, 119(24), 6339–6345, doi:10.1021/acs.jpca.5b01818, 2015.
- 412 Hyttinen, N., Otkjær, R. V., Iyer, S., Kjaergaard, H. G., Rissanen, M. P., Wennberg, P. O. and Kurtén, T.:  
413 Computational Comparison of Different Reagent Ions in the Chemical Ionization of Oxidized  
414 Multifunctional Compounds, *J. Phys. Chem. A*, 122(1), 269–279, doi:10.1021/acs.jpca.7b10015,  
415 2018.
- 416 Jimenez, J. L., Canagaratna, M. R., Donahue, N. M., Prevot, A. S. H. H., Zhang, Q., Kroll, J. H., DeCarlo,  
417 P. F., Allan, J. D., Coe, H., Ng, N. L., Aiken, A. C., Docherty, K. S., Ulbrich, I. M., Grieshop, A. P.,  
418 Robinson, A. L., Duplissy, J., Smith, J. D., Wilson, K. R., Lanz, V. A., Hueglin, C., Sun, Y. L., Tian,  
419 J., Laaksonen, A., Raatikainen, T., Rautiainen, J., Vaattovaara, P., Ehn, M., Kulmala, M., Tomlinson,  
420 J. M., Collins, D. R., Cubison, M. J., Dunlea, J., Huffman, J. A., Onasch, T. B., Alfarra, M. R.,



- 421 Williams, P. I., Bower, K., Kondo, Y., Schneider, J., Drewnick, F., Borrmann, S., Weimer, S.,  
422 Demerjian, K., Salcedo, D., Cottrell, L., Griffin, R., Takami, A., Miyoshi, T., Hatakeyama, S.,  
423 Shimono, A., Sun, J. Y., Zhang, Y. M., Dzepina, K., Kimmel, J. R., Sueper, D., Jayne, J. T., Herndon,  
424 S. C., Trimborn, A. M., Williams, L. R., Wood, E. C., Middlebrook, A. M., Kolb, C. E., Baltensperger,  
425 U., Worsnop, D. R., Dunlea, E. J., Huffman, J. A., Onasch, T. B., Alfarra, M. R., Williams, P. I.,  
426 Bower, K., Kondo, Y., Schneider, J., Drewnick, F., Borrmann, S., Weimer, S., Demerjian, K.,  
427 Salcedo, D., Cottrell, L., Griffin, R., Takami, A., Miyoshi, T., Hatakeyama, S., Shimono, A., Sun, J.  
428 Y., Zhang, Y. M., Dzepina, K., Kimmel, J. R., Sueper, D., Jayne, J. T., Herndon, S. C., Trimborn, A.  
429 M., Williams, L. R., Wood, E. C., Middlebrook, A. M., Kolb, C. E., Baltensperger, U. and Worsnop,  
430 D. R.: Evolution of Organic Aerosols in the Atmosphere, *Science* (80-. ), 326(5959), 1525–1529,  
431 doi:10.1126/science.1180353, 2009.
- 432 Jokinen, T., Sipilä, M., Richters, S., Kerminen, V. M., Paasonen, P., Stratmann, F., Worsnop, D., Kulmala,  
433 M., Ehn, M., Herrmann, H. and Berndt, T.: Rapid autoxidation forms highly oxidized RO<sub>2</sub> radicals  
434 in the atmosphere, *Angew. Chemie Int. Ed.*, 53, 14596–14600, doi:10.1002/anie.201408566, 2014.
- 435 Jokinen, T., Berndt, T., Makkonen, R., Kerminen, V.-M., Junninen, H., Paasonen, P., Stratmann, F.,  
436 Herrmann, H., Guenther, A. B., Worsnop, D. R., Kulmala, M., Ehn, M. and Sipilä, M.: Production  
437 of extremely low volatile organic compounds from biogenic emissions: Measured yields and  
438 atmospheric implications, *Proc. Natl. Acad. Sci.*, 112(23), 7123–7128,  
439 doi:10.1073/pnas.1423977112, 2015.
- 440 Junninen, H., Ehn, M., Petäjä, Luosujärvi, L., Kotiaho, T., Kostianen, R., Rohner, U., Gonin, M., Fuhrer,  
441 K., Kulmala, M. and Worsnop, D. R.: A high-resolution mass spectrometer to measure atmospheric  
442 ion composition, *Atmos. Meas. Tech.*, 3(4), 1039–1053, doi:10.5194/amt-3-1039-2010, 2010.
- 443 Kiendler-Scharr, A., Mensah, A. A., Friese, E., Topping, D., Nemitz, E., Prevot, A. S. H., Äijälä, M.,  
444 Allan, J., Canonaco, F., Canagaratna, M., Carbone, S., Crippa, M., Dall'Osto, M., Day, D. A., De  
445 Carlo, P., Di Marco, C. F., Elbern, H., Eriksson, A., Freney, E., Hao, L., Herrmann, H., Hildebrandt,  
446 L., Hillamo, R., Jimenez, J. L., Laaksonen, A., McFiggans, G., Mohr, C., O'Dowd, C., Otjes, R.,



- 447 Ovadnevaite, J., Pandis, S. N., Poulain, L., Schlag, P., Sellegri, K., Swietlicki, E., Tiitta, P.,  
448 Vermeulen, A., Wahner, A., Worsnop, D. and Wu, H. C.: Ubiquity of organic nitrates from nighttime  
449 chemistry in the European submicron aerosol, *Geophys. Res. Lett.*, 43(14), 7735–7744,  
450 doi:10.1002/2016GL069239, 2016.
- 451 Lee, B. H., Mohr, C., Lopez-Hilfiker, F. D., Lutz, A., Hallquist, M., Lee, L., Romer, P., Cohen, R. C.,  
452 Iyer, S., Kurtén, T., Hu, W., Day, D. A., Campuzano-Jost, P., Jimenez, J. L., Xu, L., Ng, N. L., Guo,  
453 H., Weber, R. J., Wild, R. J., Brown, S. S., Koss, A., De Gouw, J., Olson, K., Goldstein, A. H., Seco,  
454 R., Kim, S., McAvey, K., Shepson, P. B., Starn, T., Baumann, K., Edgerton, E. S., Liu, J., Shilling,  
455 J. E., Miller, D. O., Brune, W., Schobesberger, S., D'Ambro, E. L. and Thornton, J. A.: Highly  
456 functionalized organic nitrates in the southeast United States: Contribution to secondary organic  
457 aerosol and reactive nitrogen budgets, *Proc. Natl. Acad. Sci. U. S. A.*, 113(6), 1516–1521,  
458 doi:10.1073/pnas.1508108113, 2016.
- 459 Liebmann, J., Sobanski, N., Schuladen, J., Karu, E., Hellén, H., Hakola, H., Zha, Q., Ehn, M., Riva, M.,  
460 Williams, J., Fischer, H., Lelieveld, J. and Crowley, J. N.: Alkyl nitrates in the boreal forest:  
461 Formation via the NO<sub>3</sub>, OH and O<sub>3</sub> induced oxidation of BVOCs and ambient lifetimes, *Atmos.*  
462 *Chem. Phys. Discuss.*, (3), 1–23, doi:10.5194/acp-2019-463, 2019.
- 463 Lin, C., Huang, R. J., Duan, J., Zhong, H. and Xu, W.: Primary and Secondary Organic Nitrate in  
464 Northwest China: A Case Study, *Environ. Sci. Technol. Lett.*, 8(11), 947–953,  
465 doi:10.1021/acs.estlett.1c00692, 2021.
- 466 Mentel, T. F., Springer, M., Ehn, M., Kleist, E., Pullinen, I., Kurtén, T., Rissanen, M., Wahner, A. and  
467 Wildt, J.: Formation of highly oxidized multifunctional compounds: autoxidation of peroxy radicals  
468 formed in the ozonolysis of alkenes—deduced from structure–product relationships, *Atmos. Chem.*  
469 *Phys. Discuss.*, 15(2), 2791–2851, doi:10.5194/acpd-15-2791-2015, 2015.
- 470 Nie, W., Yan, C., Huang, D. D., Wang, Z., Liu, Y., Qiao, X., Guo, Y., Tian, L., Zheng, P., Xu, Z., Li, Y.,  
471 Xu, Z., Qi, X., Sun, P., Wang, J., Zheng, F., Li, X., Yin, R., Dallenbach, K. R., Bianchi, F., Petäjä,  
472 T., Zhang, Y., Wang, M., Schervish, M., Wang, S., Qiao, L., Wang, Q., Zhou, M., Wang, H., Yu, C.,



- 473 Yao, D., Guo, H., Ye, P., Lee, S., Li, Y. J., Liu, Y., Chi, X., Kerminen, V.-M., Ehn, M., Donahue, N.  
474 M., Wang, T., Huang, C., Kulmala, M., Worsnop, D., Jiang, J. and Ding, A.: Secondary organic  
475 aerosol formed by condensing anthropogenic vapours over China's megacities, *Nat. Geosci.*, 15,  
476 255–261, doi:10.1038/s41561-022-00922-5, 2022.
- 477 Pye, H. O. T., D'Ambro, E. L., Lee, B. H., Schobesberger, S., Takeuchi, M., Zhao, Y., Lopez-Hilfiker, F.,  
478 Liu, J., Shilling, J. E., Xing, J., Mathur, R., Middlebrook, A. M., Liao, J., Welti, A., Graus, M.,  
479 Warneke, C., de Gouw, J. A., Holloway, J. S., Ryerson, T. B., Pollack, I. B. and Thornton, J. A.:  
480 Anthropogenic enhancements to production of highly oxygenated molecules from autoxidation, *Proc.*  
481 *Natl. Acad. Sci. U. S. A.*, 116(14), 6641–6646, doi:10.1073/pnas.1810774116, 2019.
- 482 Qiao, X., Yan, C., Li, X., Guo, Y., Yin, R., Deng, C., Li, C., Nie, W., Wang, M., Cai, R., Huang, D., Wang,  
483 Z., Yao, L., Worsnop, D. R., Bianchi, F., Liu, Y., Donahue, N. M., Kulmala, M. and Jiang, J.:  
484 Contribution of Atmospheric Oxygenated Organic Compounds to Particle Growth in an Urban  
485 Environment, *Environ. Sci. Technol.*, doi:10.1021/acs.est.1c02095, 2021.
- 486 Rissanen, M. P., Kurtén, T., Sipilä, M., Thornton, J. A., Kangasluoma, J., Sarnela, N., Junninen, H.,  
487 Jørgensen, S., Schallhart, S., Kajos, M. K., Taipale, R., Springer, M., Mentel, T. F., Ruuskanen, T.,  
488 Petäjä, T., Worsnop, D. R., Kjaergaard, H. G. and Ehn, M.: The formation of highly oxidized  
489 multifunctional products in the ozonolysis of cyclohexene, *J. Am. Chem. Soc.*, 136(44), 15596–  
490 15606, doi:10.1021/ja507146s, 2014.
- 491 Riva, M.: Multiphase Chemistry of Highly Oxidized Molecules: The Case of Organic Hydroperoxides,  
492 *Chem*, 1(4), 526–528, doi:10.1016/j.chempr.2016.09.015, 2016.
- 493 Riva, M., Ehn, M., Li, D., Tomaz, S., Bourgain, F., Perrier, S. and George, C.: CI-Orbitrap: An Analytical  
494 Instrument to Study Atmospheric Reactive Organic Species, *Anal. Chem.*, 91, 9419–9423,  
495 doi:10.1021/acs.analchem.9b02093, 2019a.
- 496 Riva, M., Rantala, P., Krechmer, J. E., Peräkylä, O., Zhang, Y., Heikkinen, L., Garmash, O., Yan, C.,  
497 Kulmala, M., Worsnop, D. and Ehn, M.: Evaluating the performance of five different chemical  
498 ionization techniques for detecting gaseous oxygenated organic species, *Atmos. Meas. Tech.*, 12,



- 499 2403–2421, doi:10.5194/amt-2018-407, 2019b.
- 500 Rollins, A. W., Pusede, S., Wooldridge, P., Min, K. E., Gentner, D. R., Goldstein, A. H., Liu, S., Day, D.  
501 A., Russell, L. M., Rubitschun, C. L., Surratt, J. D. and Cohen, R. C.: Gas/particle partitioning of  
502 total alkyl nitrates observed with TD-LIF in Bakersfield, J. Geophys. Res. Atmos., 118(12), 6651–  
503 6662, doi:10.1002/jgrd.50522, 2013.
- 504 Schervish, M. and Donahue, N. M.: Peroxy radical chemistry and the volatility basis set, Atmos. Chem.  
505 Phys., 20(2), 1183–1199, doi:10.5194/acp-20-1183-2020, 2020.
- 506 Wang, D. S. and Hildebrandt Ruiz, L.: Chlorine-initiated oxidation of n-alkanes under high NO<sub>x</sub>  
507 conditions: Insights into secondary organic aerosol composition and volatility using a FIGAERO-  
508 CIMS, Atmos. Chem. Phys. Discuss., (x), 1–26, doi:10.5194/acp-2018-443, 2018.
- 509 Wang, Y., Mehra, A., Krechmer, J., Yang, G., Hu, X., Lu, Y., Lambe, A., Canagaratna, M., Chen, J.,  
510 Worsnop, D., Coe, H. and Wang, L.: Oxygenated products formed from OH-initiated reactions of  
511 trimethylbenzene: Autoxidation and accretion, Atmos. Chem. Phys., 20, 9563–9579,  
512 doi:10.5194/acp-2020-165, 2020.
- 513 Xu, L., Suresh, S., Guo, H., Weber, R. J. and Ng, N. L.: Aerosol characterization over the southeastern  
514 United States using high-resolution aerosol mass spectrometry: Spatial and seasonal variation of  
515 aerosol composition and sources with a focus on organic nitrates, Atmos. Chem. Phys., 15(13), 7307–  
516 7336, doi:10.5194/acp-15-7307-2015, 2015.
- 517 Xu, Z. N., Nie, W., Liu, Y. L., Sun, P., Huang, D. D., Yan, C., Krechmer, J., Ye, P. L., Xu, Z., Qi, X. M.,  
518 Zhu, C. J., Li, Y. Y., Wang, T. Y., Wang, L., Huang, X., Tang, R. Z., Guo, S., Xiu, G. L., Fu, Q. Y.,  
519 Worsnop, D., Chi, X. G. and Ding, A. J.: Multifunctional Products of Isoprene Oxidation in Polluted  
520 Atmosphere and Their Contribution to SOA, Geophys. Res. Lett., 48(1), 1–10,  
521 doi:10.1029/2020GL089276, 2021.
- 522 Yan, C., Nie, W., Äijälä, M., Rissanen, M. P., Canagaratna, M. R., Massoli, P., Junninen, H., Jokinen, T.,  
523 Sarnela, N., Häme, S. A. K., Schobesberger, S., Canonaco, F., Yao, L., Prévôt, A. S. H., Petäjä, T.,  
524 Kulmala, M., Sipilä, M., Worsnop, D. R. and Ehn, M.: Source characterization of highly oxidized



- 525 multifunctional compounds in a boreal forest environment using positive matrix factorization, *Atmos.*  
526 *Chem. Phys.*, 16, 12715–12731, doi:10.5194/acp-16-12715-2016, 2016.
- 527 Yan, C., Yin, R., Lu, Y., Dada, L., Yang, D., Fu, Y., Kontkanen, J., Deng, C., Garmash, O., Ruan, J.,  
528 Baalbaki, R., Schervish, M., Cai, R., Bloss, M., Chan, T., Chen, T., Chen, Q., Chen, X., Chen, Y.,  
529 Chu, B., Dällenbach, K., Foreback, B., He, X., Heikkinen, L., Jokinen, T., Junninen, H.,  
530 Kangasluoma, J., Kokkonen, T., Kurppa, M., Lehtipalo, K., Li, H., Li, H., Li, X., Liu, Y., Ma, Q.,  
531 Paasonen, P., Rantala, P., Pileci, R. E., Rusanen, A., Sarnela, N., Simonen, P., Wang, S., Wang, W.,  
532 Wang, Y., Xue, M., Yang, G., Yao, L., Zhou, Y., Kujansuu, J., Petäjä, T., Nie, W., Ma, Y., Ge, M.,  
533 He, H., Donahue, N. M., Worsnop, D. R., Veli-Matti Kerminen, Wang, L., Liu, Y., Zheng, J.,  
534 Kulmala, M., Jiang, J. and Bianchi, F.: The Synergistic Role of Sulfuric Acid, Bases, and Oxidized  
535 Organics Governing New-Particle Formation in Beijing, *Geophys. Res. Lett.*, 48, 2020GL091944,  
536 doi:10.1029/2020gl091944, 2021.
- 537 Ye, C., Yuan, B., Lin, Y., Wang, Z., Hu, W., Li, T., Chen, W., Wu, C., Wang, C., Huang, S., Qi, J., Wang,  
538 B., Wang, C., Song, W., Wang, X., Zheng, E., Krechmer, J. E., Ye, P., Zhang, Z., Wang, X., Worsnop,  
539 D. R. and Shao, M.: Chemical characterization of oxygenated organic compounds in the gas phase  
540 and particle phase using iodide CIMS with FIGAERO in urban air, *Atmos. Chem. Phys.*, 21(11),  
541 8455–8478, doi:10.5194/acp-21-8455-2021, 2021.
- 542 Yu, K., Zhu, Q., Du, K. and Huan, X. F.: Characterization of nighttime formation of particulate organic  
543 nitrates based on high-resolution aerosol mass spectrometry in an urban atmosphere in China, *Atmos.*  
544 *Chem. Phys.*, 19(7), 5235–5249, doi:10.5194/acp-19-5235-2019, 2019.
- 545 Zhang, Y., Li, D., Ma, Y., Dubois, C., Wang, X., Perrier, S., Chen, H., Wang, H., Jing, S., Lu, Y., Lou,  
546 S., Yan, C., Nie, W., Chen, J., Huang, C., George, C. and Riva, M.: Field Detection of Highly  
547 Oxygenated Organic Molecules in Shanghai by Chemical Ionization–Orbitrap, *Environ. Sci.*  
548 *Technol.*, doi:10.1021/acs.est.1c08346, 2022.
- 549 Zhao, Y., Thornton, J. A. and Pye, H. O. T.: Quantitative constraints on autoxidation and dimer formation  
550 from direct probing of monoterpene-derived peroxy radical chemistry, *Proc. Natl. Acad. Sci. U. S.*



551 A., 115(48), 12142–12147, doi:10.1073/pnas.1812147115, 2018.

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Table 1 Summary of the four cases including the meteorological conditions and concentrations of trace gases and 2N-OOMs

Case	Time	$\overline{PM_{2.5}}$ ( $\mu\text{g m}^{-3}$ )	$\overline{Solar}$ ( $\text{W m}^{-2}$ )	$\overline{T}$ ( $^{\circ}\text{C}$ )	$\overline{RH}$ (%)	$[\text{O}_3]$ (ppb)	$[\text{NO}]$ (ppb)	$[\text{NO}_2]$ (ppb)	$[\text{NO}_3]$ (ppt)	$[\text{2N-OOM}_{Ar6}]$ ( $\times 10^6 \text{ cm}^{-3}$ )	$[\text{2N-OOM}_{Ar1}]$ ( $\times 10^6 \text{ cm}^{-3}$ )	$[\text{2N-OOM}_{Ar7}]$ ( $\times 10^6 \text{ cm}^{-3}$ )	$[\text{2N-OOM}_{total}]$ ( $\times 10^6 \text{ cm}^{-3}$ )
CL <sub>day</sub>	Nov. 4 <sup>th</sup> 12:00 - 14:00	7.5	635.6	18.9	35.2	41.9	3.2	8.2	0.1	6.3	11.7	2.3	20.8
CL <sub>night</sub>	Nov. 4 <sup>th</sup> 23:00 - Nov. 5 <sup>th</sup> 01:00	9.5	2.4	13.0	64.1	8.0	2.9	40.5	0.2	1.3	6.0	2.3	9.8
PL <sub>day</sub>	Nov. 7 <sup>th</sup> 12:00 - 14:00	44.0	384.5	23.9	30.5	73.9	2.2	20.6	0.3	6.4	23.8	4.3	36.2
PL <sub>night</sub>	Nov. 7 <sup>th</sup> 23:00 - Nov. 8 <sup>th</sup> 01:00	60.5	2.5	17.8	44.9	27.2	2.1	38.7	6.2	3.1	17.4	5.3	26.5





## Figure Captions

**Figure 1.** (a) Average mass spectrum of the detected OOMs during the whole campaign. The pie chart shows the fractions of OOMs with different number of nitrogen and carbon atoms; (b) The fractions of 0N-OOMs, 1N-OOMs, 2N-OOMs and 3N-OOMs among total OOMs as a function of carbon number (nC).

**Figure 2.** The timeseries of 2N-OOMs originated from different precursors. Four sub-periods were selected to further investigate the fractional distribution of different type of OOM molecules as shown in the pie chart, including a clean daytime case (12:00 to 14:00 on November 4<sup>th</sup>,  $PM_{2.5}=7.5 \mu\text{g m}^{-3}$ ,  $CL_{\text{day}}$ ), a clean nighttime case (23:00 on November 4<sup>th</sup> to 01:00 on November 5<sup>th</sup>,  $PM_{2.5}=9.5 \mu\text{g m}^{-3}$ ,  $CL_{\text{night}}$ ), a daytime case in a PM episode (12:00 to 14:00 on November 7<sup>th</sup>,  $PM_{2.5}=44.0 \mu\text{g m}^{-3}$ ,  $PL_{\text{day}}$ ) and a nighttime case in a PM episode (23:00 on November 7<sup>th</sup> to 01:00 on November 8<sup>th</sup>,  $PM_{2.5}=60.5 \mu\text{g m}^{-3}$ ,  $PL_{\text{night}}$ ). The sizes of pie charts are scaled to the total concentrations of 2N-OOMs.

**Figure 3.** The fractions of  $2N\text{-OOM}_{\text{Ali}}$  with different carbon numbers in the four cases.

**Figure 4.** (a) Statistical distribution of the correlation coefficients (Spearman type) between 2N-OOMs and solar radiation ( $R_{2N\text{-OOMs} - \text{solar}}$ ) in red and the correlation coefficients between 2N-OOMs and  $[\text{NO}_3]$  ( $R_{2N\text{-OOMs} - \text{NO}_3}$ ) in blue for 2N-OOMs from different precursors. The horizontal lines are the median values, boxes denote the 25th- and 75th- percentile values, and whiskers represent the 10th- and 90th- percentile values. (b) The diel patterns of 2N-OOMs from different precursors.

**Figure 5.** Scatter plot of  $R_{2N\text{-OOMs} - \text{NO}_3}$  against  $R_{2N\text{-OOMs} - \text{solar}}$  for specific 2N-OOM species.

**Figure 6.** The  $nO_{\text{eff}}$  of 2N-OOMs derived from different precursors in the four cases, the error bars represent the standard deviations.

**Figure 7.** Effective oxygen number ( $nO_{\text{eff}}$ ) of 2N-OOMs against  $\text{NO}_x$  concentration, the lines represent the median value while the shadows denote the 25th- and 75th- percentile values.

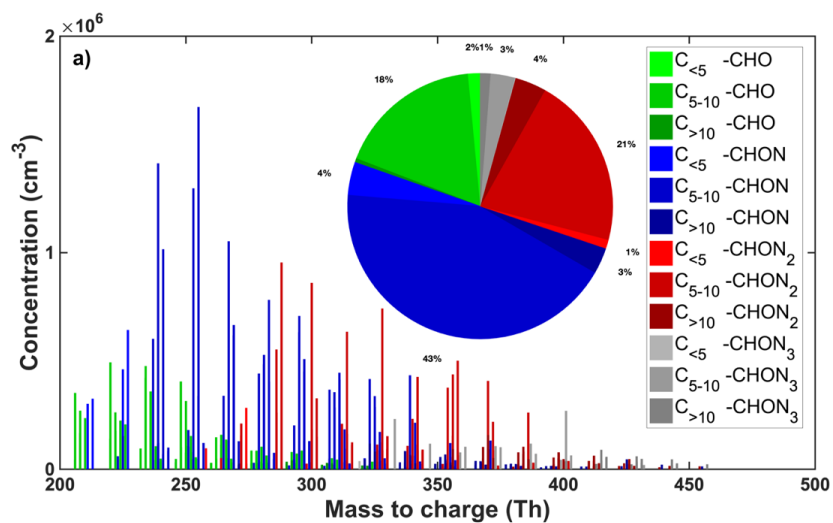


Figure 1a

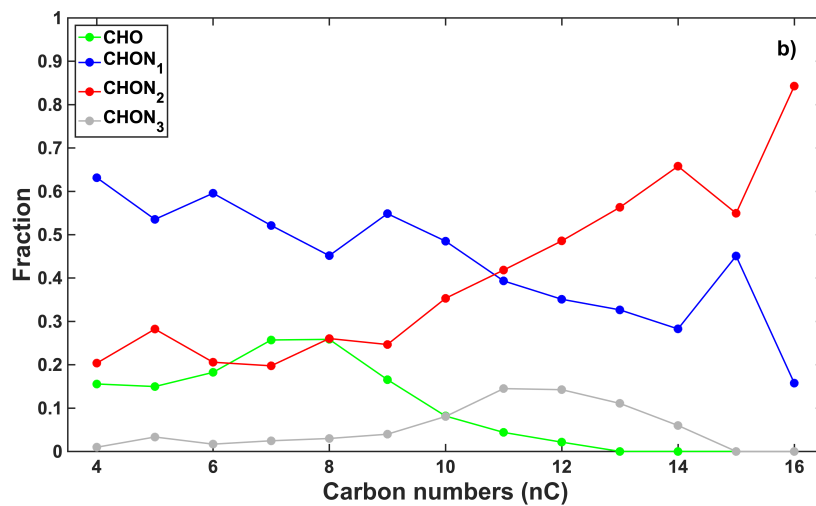


Figure 1b

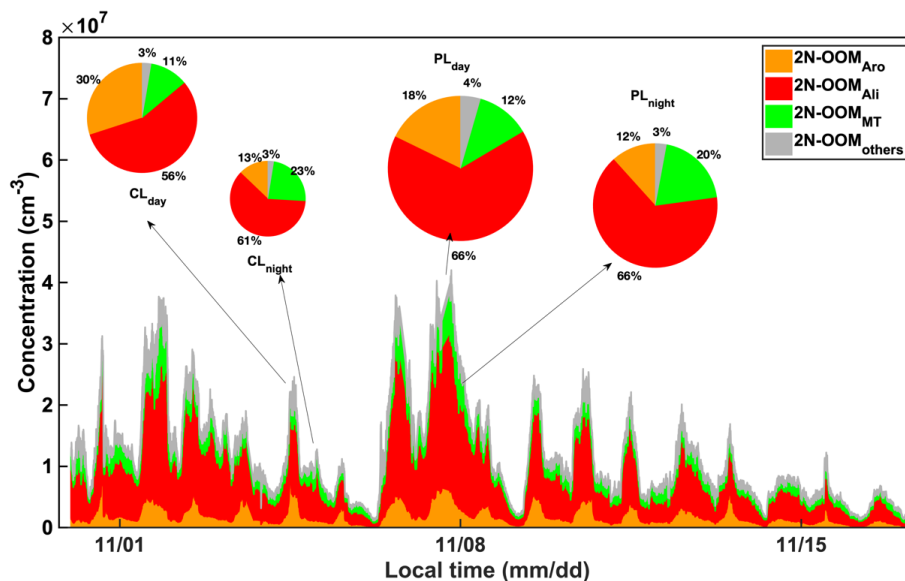


Figure 2

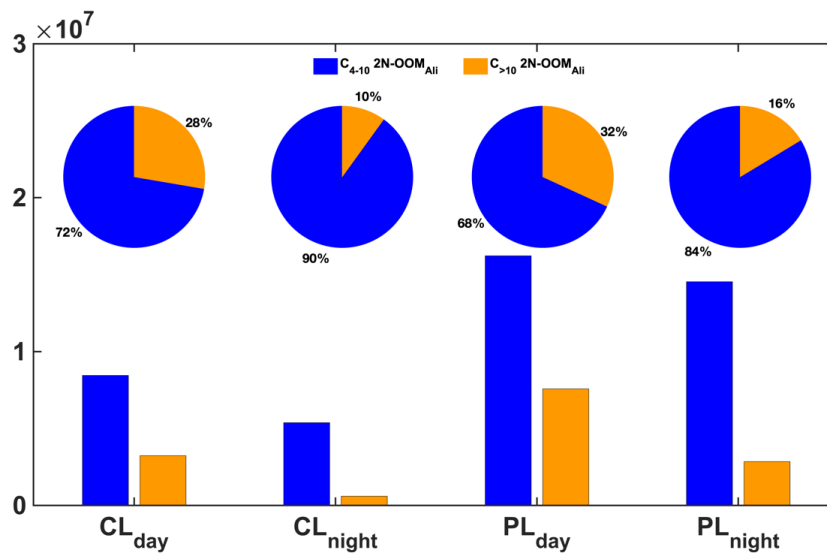


Figure 3

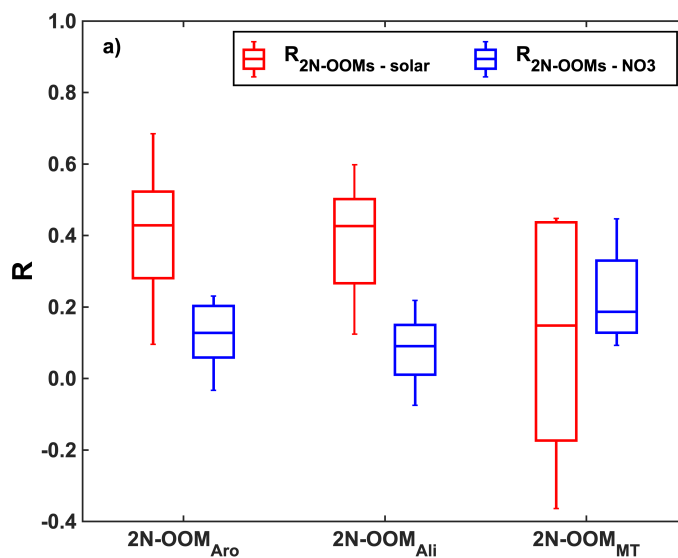


Figure 4a

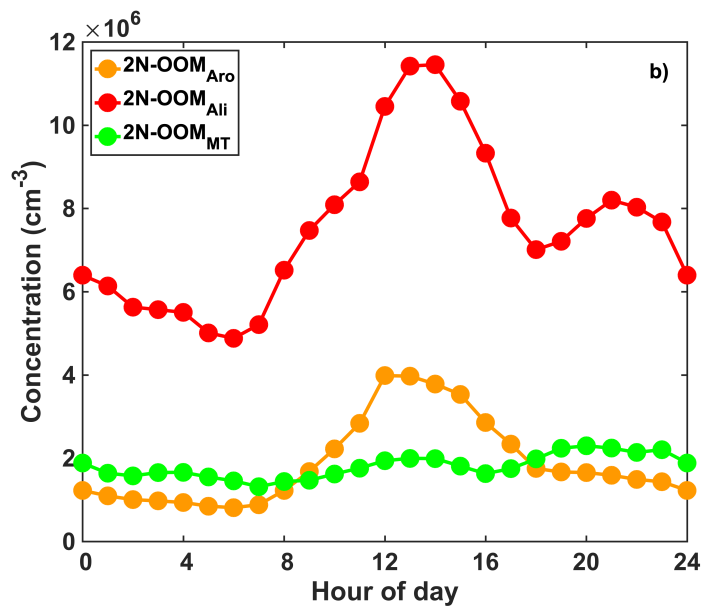


Figure 4b

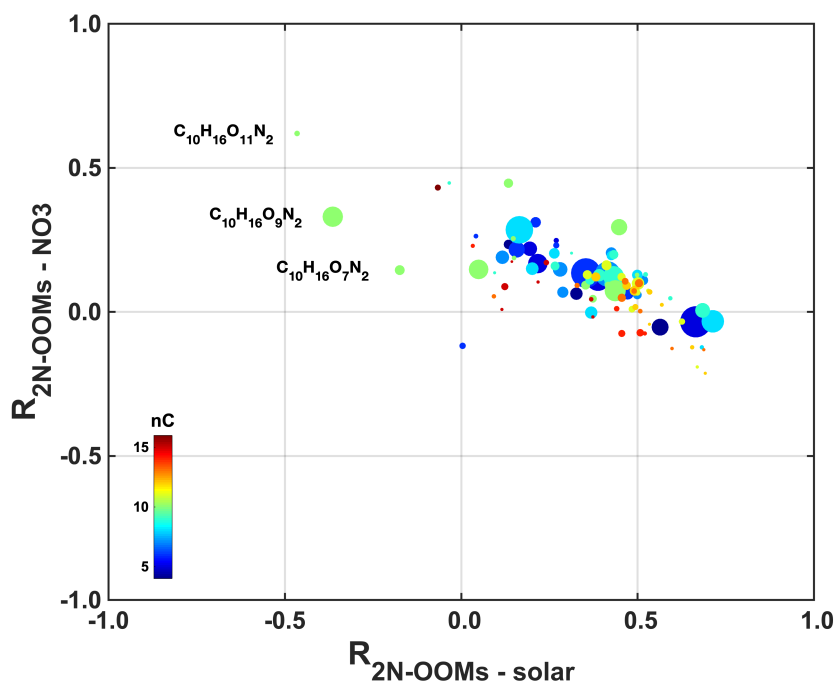


Figure 5

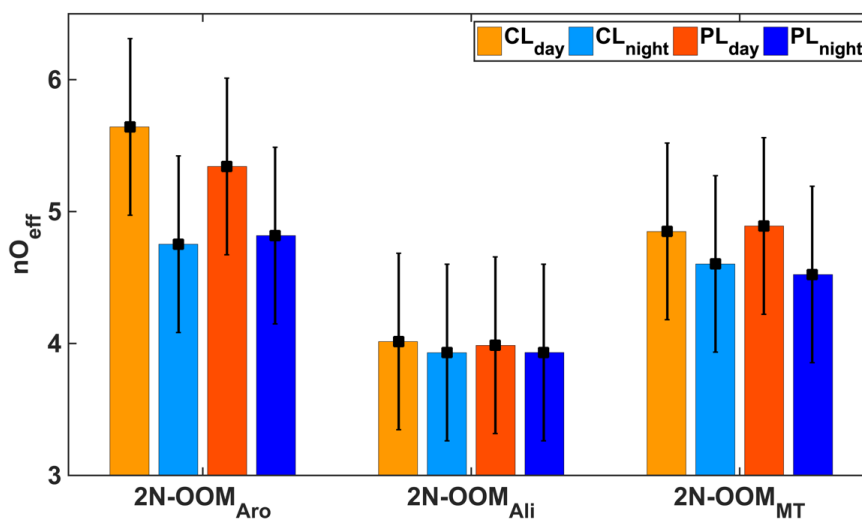


Figure 6

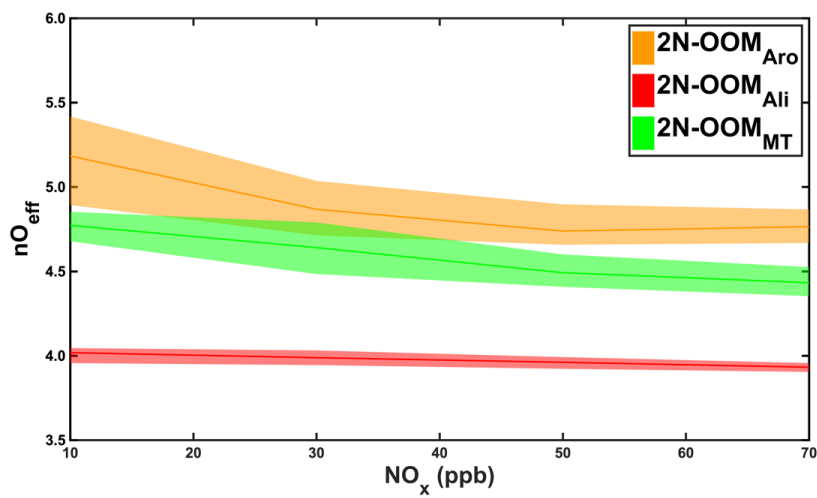


Figure 7