



Measurement Report: Rapid decline of aerosol absorption coefficient and aerosol optical properties effects on radiative forcing in urban areas of Beijing from 2018 to 2021

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Abstract

15 Reliable observations of aerosol optical properties are crucial for quantifying the radiative forcing of climate. The simultaneous measurements of aerosol optical properties at three wavelengths for PM₁ and PM₁₀ were conducted in urban Beijing from March 2018 to February 2022. The results showed considerable reductions in aerosol absorption coefficient (σ_{ab}) at 550 nm of PM₁₀ and PM₁ by 55.0% and 53.5% from 2018 to 2021. SSA
20 increased from 0.89 ± 0.04 for PM₁₀ (0.87 ± 0.05 for PM₁) in 2018 to 0.93 ± 0.03 for PM₁₀ (0.91 ± 0.04 for PM₁) in 2021. These results indicated that the absorbing aerosols were more effectively controlled than scattering aerosols due to pollution control measure-taking.



The annual average submicron absorption ratio (R_{ab}) increased from 86.1% in 2018 to 89.2% in 2021, suggesting that fine particles are the main contributors to total PM_{10} absorption and that the fine particles to absorption became more important. Absorption Angstrom exponent (AAE) in winter decreased from 2018 to 2021, implying a decreasing contribution from brown carbon to light absorption, which may relate to the decreased emissions of biomass burning and coal combustion. During the study period, aerosol radiative forcing efficiency became more negative and were -27.0 and $-26.2 \text{ W m}^{-2} \text{ AOD}^{-1}$ for PM_{10} and PM_1 in 2021, which was mainly influenced by increasing SSA. Higher σ_{ab} and $PM_{2.5}$ mass concentrations were mainly distributed in clusters 4 and 5 transported from the south and the west of Beijing in each year. σ_{ab} and $PM_{2.5}$ corresponding to clusters 4 and 5 decreased evidently from 2018 to 2021, which may result from the control of source emissions in surrounding regions of Beijing. The 4-year data presented in this study provide critical optical parameters for radiative forcing assessment within two size ranges and are helpful for evaluating the effectiveness of clean air action.

1 Introduction

Atmospheric aerosols perturb the Earth's atmospheric radiation balance and climate forcing by directly affecting the scattering and absorption of solar radiation (Charlson et al., 1992; Jacobson, 2001) but also indirectly affecting cloud reflectivity and precipitation processes (Twomey, 2007). Light-scattering aerosols contribute to offsetting the warming effect of CO_2 , while absorbing aerosols contribute to the heating of the atmosphere (Bond and Bergstrom, 2007), and produce a positive radiative forcing (Segura et al., 2016). The largest contribution to aerosol absorption is from black carbon (BC), which absorbs strongly over the entire solar spectrum (Bond and Bergstrom, 2007). Dust and brown



carbon (BrC) are also light absorption aerosols, which strongly absorb in the ultraviolet (UV) spectrum. Globally, aerosols contributed an effective radiative forcing (ERF) of $-1.3 \pm 0.7 \text{ W m}^{-2}$, and the ERF due to emissions of BC is now estimated to be $0.11 (-0.20 \text{ to } 0.42) \text{ W/m}^2$ between 1750 to 2019 (Szopa et al., 2021). However, aerosol properties are highly spatial and temporal variation, which results in radiative forcing variation from local to global scales and creates an observational challenge (Collaud Coen et al., 2013; Ealo et al., 2018; Andrews et al., 2011). Therefore, reliable observations of aerosol optical properties are crucial for quantifying the radiative forcing of climate.

In order to assess the role of aerosols on climate forcing accurately, a set of parameters that describe aerosol's optical properties are needed, such as scattering coefficient (σ_{sp}), absorption coefficient (σ_{ab}), and single scattering albedo (SSA). SSA is a key variable that determines the magnitude and the sign of the aerosol forcing (J. Hansen et al., 1997; Lee et al., 2007; Li et al., 2022a; Zhang et al., 2020). Previous study found that typical values of SSA are between 0.8 and 1 at worldwide locations (Dubovik et al., 2002), with higher values indicating a tendency towards a cooling effect (Li et al., 2022a). Besides, aerosol optical properties are wavelength-dependent, absorption Angstrom exponent (AAE) describes the spectral dependence of light absorption by aerosols and is typically used to differentiate between different aerosol types (Helin et al., 2021). The AAE for fresh BC is ~ 1 , indicating "weak" spectral dependence of light absorption (Bond et al., 2013; Bond and Bergstrom, 2007), and the AAE > 1 indicates the presence of BrC or dust, which tend to exhibit absorption that increases sharply as wavelength decreases (Moosmüller et al., 2009; Lack and Cappa, 2010). Thus, obtaining the aerosol absorption coefficient at different wavelengths is essential and can be helpful to differentiate between different aerosol types.



As one of the world's most populous and rapidly developing megacities, Beijing
70 experienced rapid economic growth and urbanization, accompanied by severe air pollution.
Many in-situ measurements of aerosol optical properties have been conducted in Beijing
(Bergin et al., 2001; He et al., 2009; Garland et al., 2009; Jing et al., 2015; Wang et al.,
2019; Zhao et al., 2019; Xia et al., 2020). Previous studies found that high aerosol loading
leads to large σ_{ab} in Beijing (Jing et al., 2015; Garland et al., 2009; Bergin et al., 2001).
75 Moreover, the AAE showed significant seasonal variations in Beijing. Significantly higher
AAE in winter than in summer highlights the important role of absorption of non-BC
components (e.g. BrC) in winter (Xie et al., 2020; Xia et al., 2020). In order to reduce
emissions and improve air quality, the government implemented strict pollution control
measures (Xu and Zhang, 2020). Significant decreases in PM_{2.5} mass concentrations were
80 found in Beijing and the annual mean elemental carbon (EC) concentrations declined from
4.0 to 2.6 $\mu\text{g m}^{-3}$ from March 2013 to February 2018 in Beijing (Ji et al., 2019). Xia et al.
(2020) separated and quantified the effects of emission control and meteorological
transport variability on BC loading from 2015 to 2017 in north China Plain. However, the
environmental effects caused by emission controls are related to not only their mass
85 concentrations, but also their optical properties and radiative effect (Luo et al., 2020).
Therefore, it's necessary to investigate the multiple-year variations in aerosol optical
properties and radiative effect in providing a comprehensive understanding of the effects
of emission control. Wang et al. (2019) found that σ_{ap} for PM_{2.5} decreased from 2014 to
2017, with a significant decrease of σ_{ap} in autumn. Sun et al. (2022) estimated that the
90 direct radiative forcing of BC decreased by 67% from +3.36 Wm^{-2} in 2012 to +1.09 Wm^{-2}
in 2020. However, these studies were mostly conducted with conventional total suspended



particulate (TSP) cyclone, $PM_{2.5}$ size cut, or PM_{10} size cut. Few studies focused on the sub-
micron and super-micron particle optical properties and estimated aerosol radiative effect
in the post-“Action Plan on Prevention and Control of Air Pollution” era. Acquiring the
95 aerosol optical for the total ($< 10\mu\text{m}$ diameter) and submicron aerosol is also in line with
the aerosol advisory group of the Global Atmosphere Watch recommendation
(WMO/GAW, 2016).

In this study, the simultaneous measurements of aerosol optical properties at three
wavelengths for PM_1 and PM_{10} were conducted in urban Beijing from March 2018 to
100 February 2022. The annual, seasonal, and diurnal variations of aerosol optical properties
for two size cuts were investigated. The scattering properties of aerosols for two size ranges
(PM_{10} and PM_1) under dry conditions observed in Beijing have been analyzed in detail by
Hu et al. (2021). Thus, this study mainly focused on the variation of aerosol absorption
coefficient, single scattering albedo, and absorption Angstrom Exponent for PM_{10} and PM_1 .
105 Moreover, the aerosol radiative effects in two size cuts were estimated. Finally, the
transport and its impact on aerosol optical properties were analyzed. The 4-year data
presented in this study provide key optical parameters for radiative forcing assessment
within two size ranges and are helpful for evaluating the effectiveness of clean air action.

2 Instrumentation and methods

110 2.1 Site description

The sampling site in this study is located on the roof of the Chinese Academy of
Meteorological Sciences (CAMS, $116^\circ 19' \text{ E}$, $39^\circ 57' \text{ N}$) in Beijing, which is a typical urban
site in the northwest of Beijing between the 2nd and 3rd ring roads. The laboratory is



approximately 53 m above ground level. The site is mainly influenced by local emission
115 from residential living and traffic pollution (Xia et al., 2019).

2.2 Aerosol absorption measurements

The ambient air was sampled into a PM₁₀ impactor with 16.7 LPM and then to an
adsorption aerosol dryer, which controlled the relative humidity (RH) of sample air below
30% (Tuch et al., 2009). The dried aerosol sample passes through switched impactors that
120 toggle the aerosol size cut between 1.0 μm (<1 μm) and 10 μm (<10 μm) aerodynamic
particle diameters every 30 min, thus allowing to measure both fine and coarse particles
(Hu et al., 2021). The sample aerosol was then passed into the Nephelometer (TSI Inc.,
Model 3563) and Tricolor Absorption Photometer (TAP, Brechtel Manufacturing, Inc.,
Hayward, CA, USA).

125 TAP measures absorption coefficient (σ_{ab}) at 465, 520, and 640 nm with the 47 mm
diameter, glass-fiber filter and is a commercially available version of the continuous light
absorption photometer (CLAP), which is low cost and high sensitivity (Ogren et al., 2017).
The TAP comprises eight sample spots and two reference spots. The aerosol-laden air
passes through one sample spot at a time, which allows for 8 times the filter lifetime
130 compared to single-spot photometers (Davies et al., 2019). Unlike the Multi-Angle
Absorption Photometer (MAAP), TAP does require a co-located aerosol light scattering or
extinction measurement to derive aerosol light absorption (Ogren et al., 2017). Thus,
simultaneous observation of aerosol light scattering has been measured and used to correct
absorption data. When the Nephelometer and TAP were calibrated or malfunctioning, no
135 data are available. During this study, 84% of the data was effective.



2.3. Data processing

The TAP measures the light transmitted through a filter as particles are deposited onto the filter. The filter attenuation coefficient (σ_{atn}), at a specific wavelength (λ), can be determined as:

$$140 \quad \sigma_{\text{atn}}(\lambda) = \frac{A}{Q} \times \frac{\Delta \text{atn}(\lambda)}{\Delta t} \quad (1)$$

where $\Delta \text{atn}(\lambda)$ is the filter attenuation at times t_1 and t_2 , A is the area of on the filter, and Q is the sample flow rate through the filter.

In order to correct the error caused by multiple scattering and filter loading, the aerosol light absorption coefficient ($\sigma_{\text{ab}}(\lambda)$) was corrected based on the methods of Bond et al. (1999) and Ogren et al. (2017). First, the effect of filter loading was calibrated based on Eq. (2):

$$145 \quad \sigma_{\text{ab}}(\lambda)_{\text{raw}} = \frac{0.85 \times \sigma_{\text{atn}}(\lambda)}{K_2 \times (1.0796 \times \text{Tr}(\lambda) + 0.71)} \quad (2)$$

Then, $\sigma_{\text{ab}}(\lambda)_{\text{raw}}$ at 465, 520, and 640 nm were adjusted to the wavelength of the light scattering coefficient based on the calculated AAE. Finally, the multiple scattering effect was corrected based on Eq. (3):

$$150 \quad \sigma_{\text{ab}}(\lambda) = \sigma_{\text{ab}}(\lambda)_{\text{raw}} - \frac{K_1 \times \sigma_{\text{sp}}(\lambda)}{K_2} \quad (3)$$

where $\text{Tr}(\lambda)$ is the normalized filter transmittance at time t relative to transmittance at the start of sampling ($t=0$) and σ_{sp} is the aerosol light-scattering coefficient at 450, 550, and 700 nm measured by the nephelometer. K_1 and K_2 were derived by Bond et al. (1999) as $K_1 = 0.02 \pm 0.02$ and $K_2 = 1.22 \pm 0.20$, where the uncertainties are given for the 95% confidence level.



Using the corrected absorption coefficient data, the following parameters were calculated:

Absorption Angstrom exponent (AAE) describes the spectral dependence of light
160 absorption.

$$AAE = -\frac{\ln(\sigma_{ab}^{\lambda_1}/\sigma_{ab}^{\lambda_2})}{\ln(\lambda_1/\lambda_2)} \quad (4)$$

The submicron absorption ratio (Rab) is determined as the ratio of the absorption coefficients for PM₁ and PM₁₀.

$$Rab = \frac{\sigma_{ab}(D<1\mu m)}{\sigma_{ab}(D<10\mu m)} \quad (5)$$

165 where $\sigma_{ab}(D<1\mu m)$ and $\sigma_{ab}(D<10\mu m)$ are σ_{ab} for particle diameters $<1\mu m$ and $10\mu m$, respectively.

Aerosol radiative forcing efficiency (RFE) at top-of-the-atmosphere (TOA) is a simplified formula that describes how large of an impact the aerosols would make to the aerosol radiative forcing (ΔF) per unit of aerosol optical depth (AOD) (Sheridan and Ogren,
170 1999) and we estimated the RFE at TOA as the Eq.6 (Haywood and Shine, 1995; Sheridan and Ogren, 1999):

$$RFE = \frac{\Delta F}{AOD} = -DS_0T_{at}^2(1-A_C) \times SSA \times \beta \times ((1-R_s)^2 - (\frac{2R_s}{\beta}) \times (\frac{1}{SSA} - 1)) \quad (6)$$

175 where D is the fractional day length, S_0 is the solar constant, T_{at} is the atmospheric transmission, A_c is the fractional cloud amount, and R_s is the surface reflectance. The constants used were $D = 0.5$, $S_0 = 1370 \text{ Wm}^{-2}$, $T_{at} = 0.76$, $A_c = 0.6$, and $R_s = 0.15$ as suggested by Haywood and Shine (1995), and upper scatter fraction β was calculated from $\beta = 0.0817 + 1.8495 \times b - 2.9682 \times b^2$. backscatter ratio (b) was calculated based on scattering



coefficient (σ_{sp}) and backscattering coefficient (σ_{bsp}) measured by Nephelometer as $b =$
180 $\sigma_{bsp} / \sigma_{sp}$. Equation (6) has been widely used to assess the intrinsic radiative forcing
efficiency of aerosols at the top of the atmosphere (Sheridan and Ogren, 1999; Virkkula et
al., 2011; Shen et al., 2018). In this study, the values of ΔF at TOA were also calculated by
multiplying the RFE for PM_{10} with the AOD of ambient atmospheric aerosols observed at
the CAMS site during the study periods. AOD can be downloaded from Aerosol Robotic
185 Network (AERONET). Note that RFE was at a dry state, thus the ΔF at TOA here may be
slightly underestimated.

2.4. Other data used

The hourly $PM_{2.5}$ and PM_{10} mass concentrations were measured at Guan yuan station,
which is about 3km from the CAMS site. The data can be derived from the national air
190 quality real-time publishing platform (<http://106.37.208.233:20035/>).

2.5. Back trajectories analysis

To investigate the influence of air mass origins on aerosol optical properties, 48-h
backward trajectories arriving at Beijing at a height of 500 m above ground level were
calculated from 0:00 to 23:00 local time each day from March 2018 to February 2022,
195 using the Trajstat Software, combined with HYSPLIT 4 model (Hybrid Single-Particle
Lagrangian Integrated Trajectory), and the NCEP Global Data Assimilation System
(GDAS) data with a $1^\circ \times 1^\circ$ resolution (Draxler and Hess, 1998; Wang et al., 2009).

In this study, four seasons are defined as follows: spring from March to May, summer
from June to August, autumn from September to November, and winter from December to
200 the following February, and all data are reported in Beijing time (UTC+8).



3 Results and discussion

3.1 Temporal variation of aerosol optical properties

Figure 1 shows the annual variation of σ_{ab} , SSA, Rab, and PM_{2.5} mass concentration from 2018 to 2021. During the study period, the annual mean PM_{2.5} in 2018 was 54.7 $\mu\text{g m}^{-3}$, and it decreased by 34.4% (35.9 $\mu\text{g m}^{-3}$) in 2021, which suggested that the strict pollution control measures are effective in reducing the PM loadings in Beijing (Lei et al., 2021). σ_{ab} at 550 nm of PM₁₀ and PM₁ showed similar annual variations. The annual mean σ_{ab} at 550 nm of PM₁₀ and PM₁ decreased by 55.0% and 53.5%, respectively. A significant decrease in σ_{ab} indicated the effective control of absorbing aerosols, which was mainly related to decreasing biomass burning and coal consumption, increasing usage of natural gases, and the implementation of a stricter vehicular emission standard in recent years (Sun et al., 2022). The annual mean σ_{ab} for PM₁₀ and PM₁ in 2021 was 9.8 Mm⁻¹ and 8.7 Mm⁻¹, which were both lower than the result observed in Nainital, in the GH region, India (Dumka et al., 2015), but both higher than SMEAR II station located in Hyytiälä, southern Finland (Luoma et al., 2019). SSA is a key variable in assessing the aerosol radiative forcing. The variation of SSA also reflects the the ratio of aerosol scattering to total extinction with aerosol composition changes. The annual variations of SSA for PM₁₀ and PM₁ were similar. During 2018-2021, annual mean SSA at 550 nm increased from 0.89 ± 0.04 for PM₁₀ (0.87 ± 0.05 for PM₁) in 2018 to 0.93 ± 0.03 for PM₁₀ (0.91 ± 0.04 for PM₁) in 2021, which demonstrated that the absorbing aerosols were effectively controlled compared to scattering aerosols during the past four years. On the other hand, increasing SSA also implied that scattering aerosols play a more important role in radiative forcing in urban Beijing. The mean submicron absorption ratio (Rab) increased yearly during the same



period. It was from 86.1% in 2018 to 89.2% in 2021, suggesting that fine particles are the
225 main contributors to total PM_{10} absorption, and the contributions from fine particles to
absorption became more important.

The σ_{ab} , SSA, and AAE for PM_1 and PM_{10} showed similar annual variations in all
seasons (Fig. 2 and Fig. S1). Thus, if not stated otherwise, the following discussion takes
the aerosol optical properties of PM_{10} as an example. As shown in Fig. 2 seasonal average
230 of σ_{ab} presented a decreasing trend during all seasons from 2018 to 2021, reflecting the
reduction of absorbing aerosols which were related to effective control of absorbing
aerosols emissions in Beijing. σ_{ab} decreased by half in autumn and winter during the study
period, which was probably due to reducing coal consumption as a heating source and the
reduction of biomass burning. In particular, σ_{ab} for PM_1 and PM_{10} decreased even up to 63%
235 and 67% in the summer from 2018 to 2021. Traffic is a relatively stable source of
absorption aerosols in summer (Li et al., 2022b). The largest deduction of σ_{ab} was in
summer and could be related to more strict vehicle emission standards (Zhang et al., 2019).

In general, AAE was lowest in summer and highest in winter. The mean values of
AAE for PM_{10} were 1.13 and 1.41 in summer and winter, respectively, similar to result at
240 an urban site in Beijing in 2018 (Xie et al., 2020). During summer, the average AAE was
generally close to 1, which suggested that BC from traffic emissions was the major
component of light-absorbing aerosols. Li et al. (2022b) found that the percentage of liquid
fuel (traffic) contributing to the total BC was 86.8% in summer in Beijing. The highest
AAE suggested that BrC contributed to light absorption strongest in winter, which is due
245 to enhanced emissions from biomass burning and coal combustion in winter (Sun et al.,
2018). Notably, AAE decreased in winter from 1.48 for PM_{10} (1.48 for PM_1) in 2018 to



1.37 for PM_{10} (1.34 for PM_1) in 2021 (Fig. 2 and Fig. S1), indicating a decreasing contribution from BrC to light absorption, which may relate to the effect control of biomass burning and coal combustion caused by changes in heating energy structure (Ji et al., 2022).

250 To improve air quality, the Beijing-Tianjin-Hebei region adjusted the energy structure during the heating period and developed clean heating projects, such as the “coal to gas” project (Zhao et al., 2020; Liu et al., 2019). During the whole period, AAE was similar in spring and autumn indicating that light-absorbing aerosols were from similar emission sources in spring and autumn (Ran et al., 2016). AAE slightly increased in spring and

255 autumn from 2018 to 2021. Part of the reason was the occurrence of multiple fugitive dust in spring and autumn (Yi et al., 2021; Gui et al., 2022). On the other hand, BrC could also be formed from secondary reactions (Bond et al., 2013; Wang et al., 2022). A slight increase in AAE in spring and autumn may also have been caused by a greater amount of secondary organic aerosol formation as a result of an increased atmospheric oxidation

260 capacity (Ji et al., 2019; Lei et al., 2021).

The seasonal mean SSA showed an increasing trend in all seasons from 2018 to 2021, indicating that the contribution of scattering aerosols to extinction increased. This suggested that more effective control of scattering aerosols should be attached more importance in order to improve visibility in the future. In particular, SSA in winter

265 increased significantly from 0.88 in 2018 to 0.93 in 2021, which revealed that the proportion of absorbing aerosols decreases considerably in winter. This is consistent with recent research which suggests that air pollution control measures has been more effective in reducing the primary pollution emissions than secondary species (Vu et al., 2019; Sun et al., 2020). On the other hand, seasonal mean SSA for PM_{10} was 0.94 ± 0.04 , 0.94 ± 0.04 ,



270 0.92 ± 0.04 , 0.93 ± 0.03 in spring, summer, autumn, and winter 2021. Similar SSA suggests that the proportions of light absorbing and scattering components became relatively stable in four seasons.

Figure 3 shows the diurnal variations of σ_{ab} and SSA at 550 nm for PM_{10} , which are similar to those for PM_1 (Figure. S2). In the past four years, σ_{ab} was lower during the day
275 and higher at night in four seasons. This was consistent with that observed at an urban site in Beijing during 2014–2016 (Wang et al., 2019). The evolution of the planetary boundary layer had an important influence on the diurnal variation of the σ_{ab} . With stronger solar radiation, the boundary layer was more fully developed during the daytime, and after sunset, the convective boundary layer underwent a transition to the nocturnal stable boundary layer
280 (Guo et al., 2016). Furthermore, emissions also affected the diurnal variation of the σ_{ab} . For example, heavy-duty diesel trucks and heavy-duty vehicles were only allowed to enter urban areas from 23:00 to the following day 06:00 (Hu et al., 2021). As a response, the minimum σ_{ab} occurred during 12:00–18:00, when the planetary boundary layer was well-developed, and truck emission was lower. With shallow boundary layer height and
285 enhanced emissions from heavy-duty trucks, σ_{ab} reached the maximum at night. During the study period, SSA showed a significant peak in the afternoon in four seasons, which was similar to previous studies in urban Beijing (Zhao et al., 2019; Wang et al., 2019). Higher SSA was shown in the afternoon, which was mainly related to the reduction of absorbing aerosols emission, and more secondary scattering aerosol produced by strong chemical
290 reactions under intensive solar radiation and high temperature in the afternoon (Han et al., 2017).



3.2 Aerosol radiative effect

To study the climate impact of the aerosol particles, we investigated the variation of aerosol radiative forcing efficiency (RFE) at the top-of-the-atmosphere (TOA) variations.

295 As seen in Fig. 4, RFE for PM_{10} and PM_1 were always negative during the whole observation period, suggesting that the aerosols measured in urban Beijing have a stable cooling effect on the climate. RFE for PM_{10} and PM_1 were -27.0 and $-26.2 \text{ W m}^{-2} \text{ AOD}^{-1}$ in 2021 in urban Beijing, which was slightly negative than that of $-24.9 \text{ W m}^{-2} \text{ AOD}^{-1}$ in Nanjing (Shen et al., 2018) and highly negative than that of $-19.9 \text{ W m}^{-2} \text{ AOD}^{-1}$ in Finland

300 (Virkkula et al., 2011). This suggested that the aerosols in urban Beijing have a higher cooling efficiency. RFE was affected by SSA and backscattering ratio (b) and we investigated the RFE variations with SSA and b in Beijing. As shown in Fig. 5, When SSA increases from 0.7 to 0.92, the mean RFE increases by 1.59 times, suggesting that SSA plays an important role in strengthening cooling efficiency. When $SSA > 0.92$, the mean

305 RFE relatively keeps constant. The approximate constant RFE does not mean that the absolute aerosol radiative forcing is constant; it just suggests that the intrinsic nature of the aerosol will not significantly affect the calculation of RFE (Andrews et al., 2011). Also, the backscattering ratio has a negative relationship with RFE. A lower values of backscattering ratio corresponds to larger particles (Luoma et al., 2019). RFE became more

310 negative with increasing b , suggesting that smaller particles would cool the atmosphere more efficiently. During the study period, SSA increased from 0.89 to 0.93, while the yearly mean value of b was 0.13 every year during the study period. RFE became more negative from 2018 to 2021, suggesting that the efficiency of the aerosol cooling atmosphere was higher, which was mainly influenced by increasing SSA.



315 The ratio of $\Delta F/AOD$ is known as the aerosol radiative forcing efficiency (RFE) and
 ΔF at TOA was calculated by multiplying the RFE for PM_{10} with the AOD of ambient
atmospheric aerosols observed at the CAMS site during the study periods. The mean value
of ΔF from 2018–2021 was -15.0 W m^{-2} , -12.5 W m^{-2} , -12.1 W m^{-2} , and -11.8 W m^{-2} ,
respectively. Although RFE became more negative, the annual mean ΔF in 2021
320 corresponding to lower columnar aerosol loading became less negative than that of 2018
corresponding to higher columnar aerosol loading (Fig. S3) which was consistent with the
analysis that aerosol loading was an essential factor for the estimation of ΔF (Andrews et
al., 2011; Delene and Ogren, 2002).

3.3 Transport and its impact on aerosol optical properties in Beijing

325 In order to explore the influence of regional transports and the potential source regions
of pollutants, we analyzed the air mass back-trajectories. The air mass back trajectories
during 2018–2021 were calculated and clustered (Fig. 7); then, we statistic the aerosol
optical properties of each cluster from 2018–2021 (Fig. 8). Based on the Euclidean distance,
the back trajectories were classified into five clusters, in which clusters 1, 2 and 3, which
330 originated from the clean areas in Mongolia and eastern Inner Mongolia, and transported
to Beijing along the pathway with low emissions, were corresponded to low σ_{ab} and low
 $PM_{2.5}$ (Fig. 8a, d). Cluster 4 from the south of Beijing and cluster 5 from the west of Beijing
were referred to as the polluted air masses, and the average $PM_{2.5}$ concentrations and σ_{ab}
of clusters 4 and 5 were higher than those of clusters 1, 2, and 3 in each year (Fig. 8a, d).
335 Cluster 4 passed through Shandong and Hebei Province, which was heavily polluted before
arriving in Beijing. Cluster 5 passed through polluted Shanxi and Hebei during transport.
SSA of cluster 4 from the south was higher (Fig. 8b), which may relate to low $BC/PM_{2.5}$



ratios in south air masses (Xia et al., 2020). Zhang et al. (2013) found that high levels of secondary inorganic aerosols related to high humidity were transported by southern air masses, which enhanced heterogeneous reaction and led to relatively low BC/PM_{2.5} ratios. Fig. 7b showed percentage of each cluster accounting for the total back trajectories in each year. The results indicated that variation in each cluster fraction from 2018 to 2021 was slight. In general, cluster 1-5 accounted for 19%-21%, 13%-17%, 16%-20%, 29%-36%, 12%-20% of total back trajectories, respectively. Notably, the percentage of polluted-relevant air masses (cluster 4 and cluster 5) was ~50% each year, indicating that the transport from the south and the west of has a considerable impact on the aerosol optical properties. Higher σ_{ab} and PM_{2.5} mass concentrations were mainly distributed in clusters 4 and 5 in each year. σ_{ab} and PM_{2.5} corresponding to clusters 4 and 5 decreased evidently from 2018 to 2021 (Fig. 8a, d), which may result from the control of source emissions in surrounding regions of Beijing. Therefore, the comprehensive control of atmospheric pollution in Beijing and surrounding regions would be highly effective in reducing air pollution in Beijing.

4 Conclusions

In this study, 4-year measurements of aerosol absorption properties and single scattering albedo for PM₁₀ and PM₁ in Beijing were analyzed. The annual mean σ_{ab} of PM₁₀ and PM₁ decreased by 55.0% and 53.5%, respectively, and it showed a similar decreasing trend in all seasons. SSA at 550 nm increased from 0.89 ± 0.04 for PM₁₀ (0.87 ± 0.05 for PM₁) in 2018 to 0.93 ± 0.03 for PM₁₀ (0.91 ± 0.04 for PM₁) in 2021 and the seasonal averages of SSA for two sizes also increased in four seasons. These results imply that absorbing aerosols are effectively controlled due to pollution control measure-taking. On



the other hand, increasing SSA also implied that scattering aerosols play a more important role in radiative forcing in urban Beijing. During the study period, the annual average of Rab increased year by year and was up to 89.2% in 2021, indicating that fine particles are the main contributors to the total PM₁₀ particle absorption, and the contributions from fine
365 particles to absorption became more important in Beijing.

During the study period, AAE was lowest in summer and highest in winter. Seasonal mean AAE in summer was generally close to 1 indicating that freshly emitted BC from traffic sources was a major component of light-absorbing aerosols. Highest AAE highlights the importance of BrC light absorption in winter. Notably, AAE in winter decreased from
370 2018 to 2021, implying a decreasing contribution from BrC to absorption, which may relate to the effect control of biomass burning and coal combustion caused by changes in heating energy structure. AAE in spring and autumn was similar, indicating light-absorbing aerosols were from similar emission sources in these two seasons.

Using a simple analytical equation, we investigated the aerosol radiative effect.
375 Aerosol radiative forcing efficiency (RFE) for PM₁₀ and PM₁ was always negative, suggesting that the aerosols measured in urban Beijing have a stable cooling effect on the climate. RFE for PM₁₀ and PM₁ were -27.0 and $-26.2 \text{ W m}^{-2} \text{ AOD}^{-1}$ in 2021 in urban Beijing. RFE was influenced by SSA and b. Higher b corresponds to more negative RFE suggesting that smaller particles larger would cool the atmosphere more efficiently. When
380 $\text{SSA} < 0.92$, the absolute value of mean RFE increased by 1.59 times, suggesting that SSA plays an important role in strengthening cooling efficiency. When $\text{SSA} > 0.92$, the mean RFE keeps relatively constant, suggesting that the intrinsic nature of the aerosol will not significantly affect the calculation of RFE. SSA increased from 0.89 to 0.93, while the



yearly mean value of b was 0.13 every year during the study period. RFE became more
385 negative from 2018 to 2021, suggesting that the efficiency of the aerosol cooling
atmosphere was higher, which was mainly influenced by increasing SSA.

Regional transport and its impact on aerosol optical properties were also analyzed.
The air mass back trajectories arriving at Beijing were divided into five clusters. Clusters
1, 2, and 3, which originated from the clean area in Mongolia and eastern Inner Mongolia,
390 were transported to Beijing along the pathway with low emissions, corresponding to low
 σ_{ab} and low $PM_{2.5}$. Air masses from south and west (Cluster 4 and Cluster 5), which both
crossed the polluted region, always brought high $PM_{2.5}$ concentrations and σ_{ab} . σ_{ab} and
 $PM_{2.5}$ corresponding to clusters 4 and 5 decreased evidently from 2018 to 2021, which may
result from the control of source emissions in surrounding regions of Beijing. Therefore,
395 comprehensive control of atmospheric pollution in surrounding regions of Beijing is
conductive to reducing pollution in Beijing.

Data availability.

The data in this study are available at: <https://doi.org/10.5281/zenodo.7466069> (Hu et al.,
2022)

400 **Competing interests.**

The authors declare that they have no conflict of interest.



Author contributions.

XYH analyzed the data, prepared the figures and wrote the manuscript. JYS designed the research and outlined the manuscript. XYH, CX, and JYS conducted the measurements.

405 All co-authors discussed the results and commented on the manuscript.

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410 Observation and Forecasts of MOST.

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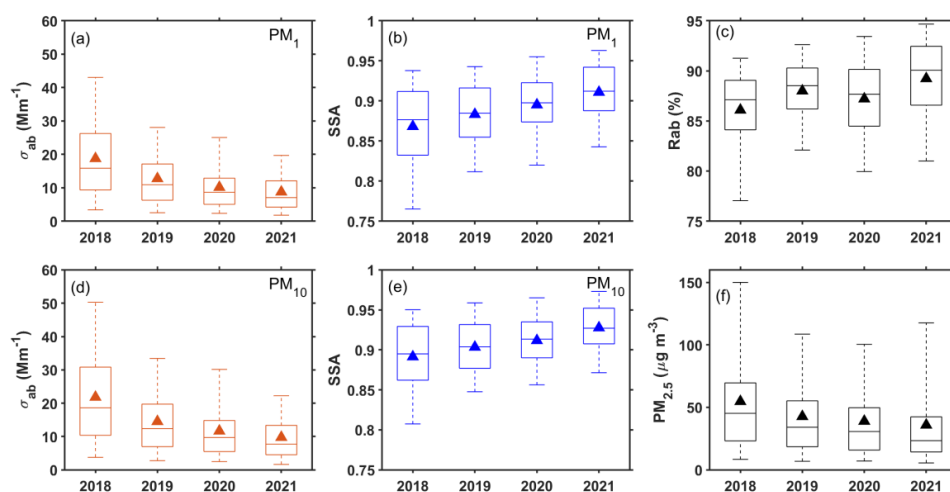
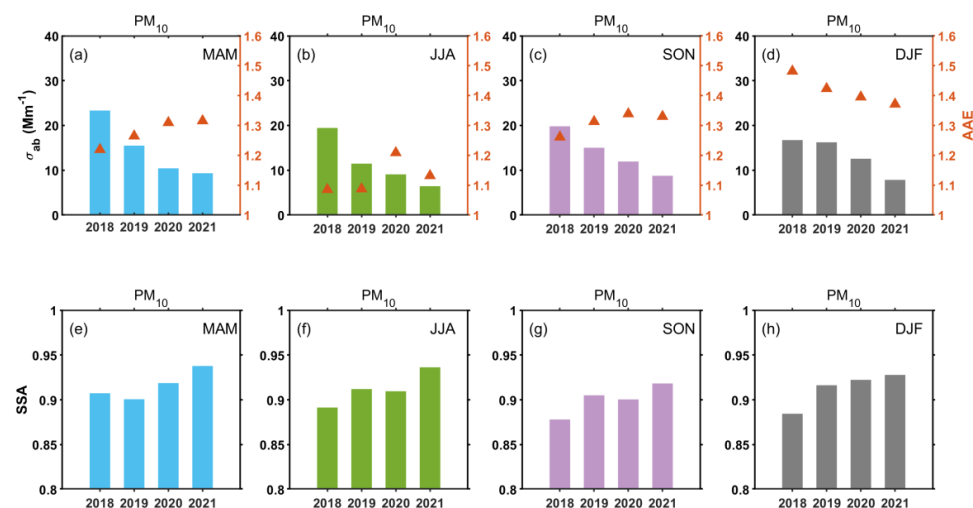


Figure 1. Annual variation of aerosol optical properties and PM_{2.5} mass concentration, absorption coefficient σ_{ab} at 550 nm for (a) PM₁ and (d) PM₁₀, SSA at 550 nm for (b) PM₁
660 and (e) PM₁₀, (c) Rab and (f) PM_{2.5} mass concentration. The solid line inside the box represents the median and the triangle indicates the mean. The box contains the range of values from 25% (bottom) to 75% (top), and the upper and lower whiskers are the 95th and 5th percentiles, respectively.



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Figure 2. Seasonal variation of aerosol optical properties of PM_{10} from 2018-2021, (a-d) σ_{ab} (bar) at 550 nm, $AAE_{450/700}$ (triangle), and (e-h) SSA (bar) at 550 nm.

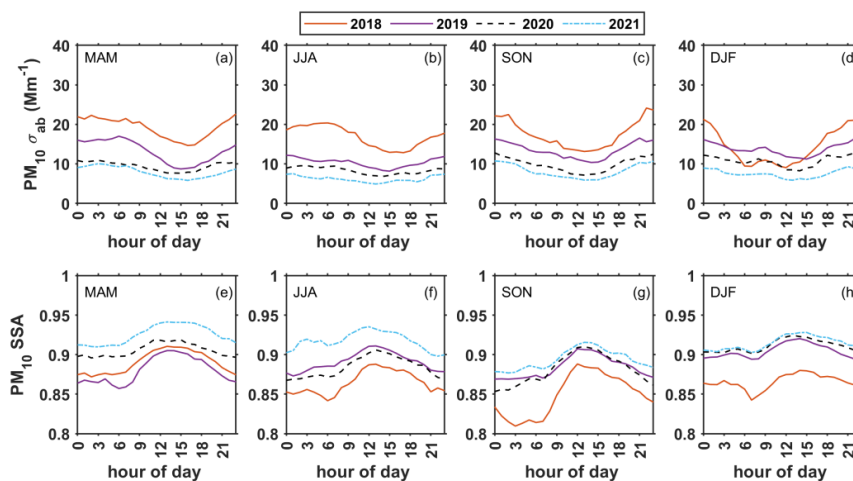


Figure 3. Diurnal variations of σ_{ab} (a-d) and SSA (e-h) at 550 nm for PM_{10} in four seasons from 2018 to 2021.

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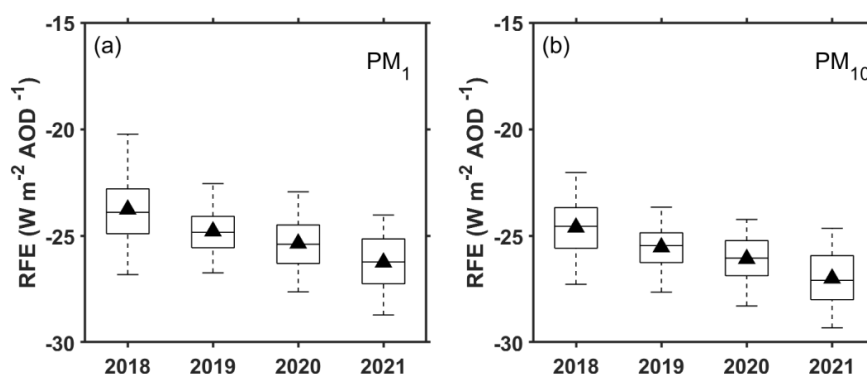


Figure 4. The annual variation of aerosol radiative forcing efficiency for PM₁ (a) and PM₁₀ (b). The solid line inside the box represents the median, and the triangle indicates the mean. The box contains the range of values from 25% (bottom) to 75% (top), and the 95th and 5th percentiles, respectively.

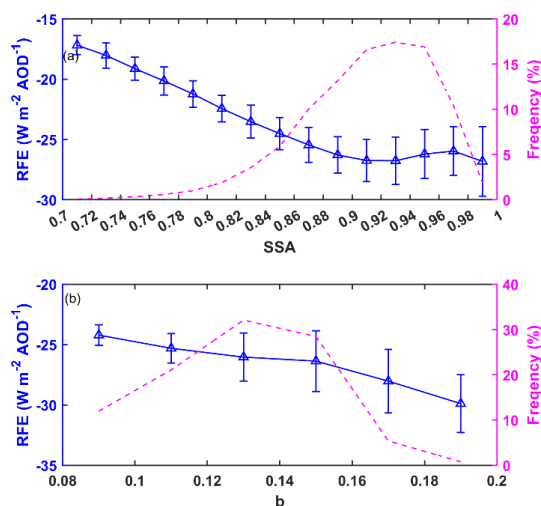
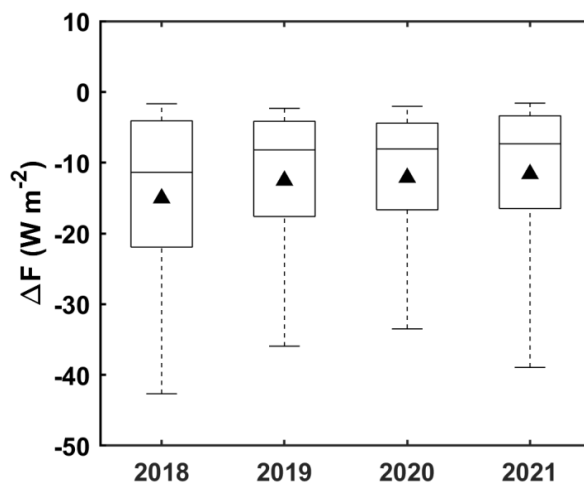
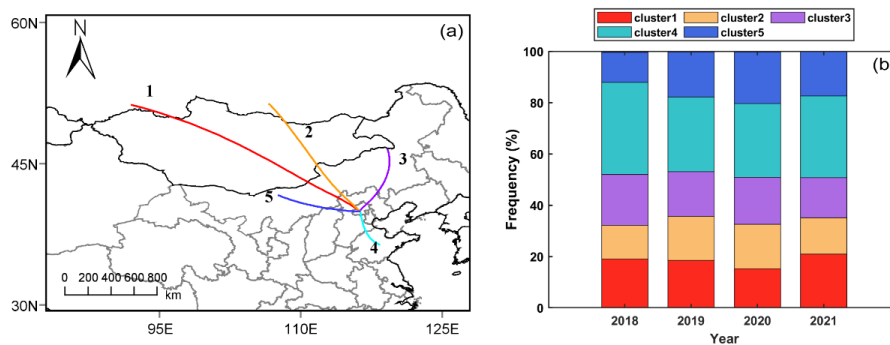


Figure 5. The relationship of RFE with (a) SSA and (b) backscattering ratio. The pink dash line represents the frequency distribution of SSA (a) and backscattering ratio (b).



680 **Figure 6.** Annual variation of aerosol radiative forcing (ΔF) at TOA from 2018 to 2021
calculated from daily mean data. The solid line inside the box represents the median, and
the triangle indicates the mean. The box contains the range of values from 25% (bottom)
to 75% (top), and the 95th and 5th percentiles, respectively.



685 **Figure 7.** (a) Air mass clusters of back trajectories arriving in Beijing during 2018–2021
and (b) the fraction of each cluster accounting for the total back trajectories in each year.

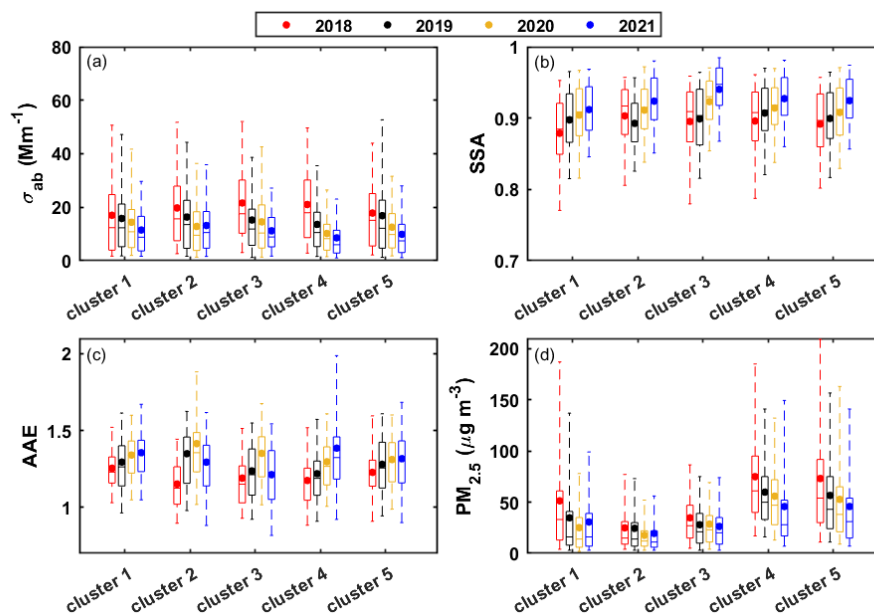


Figure 8. The variation of (a) σ_{ab} , (b) SSA, (c) AAE, and $PM_{2.5}$ mass concentration in each cluster from 2018 to 2021. The solid line inside the box represents the median and the dot indicates the mean. The box contains the range of values from 25% (bottom) to 75% (top), and the 95th and 5th percentiles, respectively.