Review of "Modelling the atmospheric ³⁴S-sulfur budget in a column model under volcanically quiescent conditions" by Juhi Nagori et al.

The study attempts to include ${}^{34}S/{}^{32}S$ isotope fractionation of sulfur species and of the chemical, physical and biological processes affecting them in an atmospheric sulfur model. It does so by a simple 1D column model approach and concentrates on key sulfur species of atmospheric relevance.

In general, this is an interesting concept and a useful approach that should be developed further by all means. However, at this point and with the model and information used, the study falls short of addressing the questions given in the introduction and, overall, to provide new information that goes beyond the current state of the art. In my opinion, results and conclusions are to a large extent imported as *a priori* information that goes into the model or by nudging parameters in the model to make key numbers agree with previous literature. I try to explain what I mean for each of the three research questions given at the end of the introduction:

1. What is the COS contribution to SSA?

First of all, I don't think there is still much debate about the contribution of carbonyl sulfide to SSA in volcanically quiescent times (see my comment regarding page 2, lines 16 - 23 further below). But even if there was, I don't see that this study provides any new knowledge or arguments in that respect.

In Section 2.1 (paragraph starting on page 5, line 17) it is stated that the eddy diffusion coefficient is doubled in order to match the stratospheric COS turnover of 40 Gg S yr⁻¹ reported by Brühl et al. (2012). If I understand this correctly, you nudge your model to simulate exactly the stratospheric COS turnover given by Brühl et al. (2012), right? By doing so, you essentially prescribe the COS flux to the stratosphere. And because the stratospheric chemistry included in your model will inevitably turn any COS reaching the stratosphere into SO₂ and subsequently SSA, the absolute COS contribution to SSA is fixed at 40 Gg S yr⁻¹. The relative contribution of 77 % to SSA depends on how much tropospheric SO₂ your model transports to the stratosphere, and for that, you can hardly expect to obtain better or more realistic numbers than the two studies by Brühl et al., 2012, and Sheng et al., 2016, that use full atmospheric models.

Therefore, other than your statement "In the next section we explore the full stratospheric sulfur budget, to investigate the importance of COS for SSA formation." (page 16, line 25) suggests, there is little "exploration", and the result with respect to the COS contribution to SSA given in Section 3.3 (Page 17, lines 1 - 4) is, to a large extent, prescribed. Therefore, when you write "We clearly demonstrate that COS is an important precursor for SSA during non-volcanic periods" (page 23, line 31, in your conclusions), I consider this hardly a conclusion but rather a feature of your model that is directly imported from previous knowledge.

2. How can isotopic information help constrain the sulfur budget?

Here, I'm not sure what exactly you mean by "*the sulfur budget*". By focusing mainly on the sulfur compounds contributing to stratospheric sulfur (i.e. SO_2 , H_2SO_4 and COS), and in the troposphere only really discuss a *budget* when it comes to COS, you are obviously not looking at the global sulfur cycle at large.

As explained above, the stratospheric sulfur budget appears to be fixed in your model. You do, for the first time, simulate a corresponding sulfur isotope budget, but with available knowledge and observations, that does not tell us much, except that a large positive fractionation for COS photolysis proposed by some studies is probably unrealistic (cf. some of my specific comments below).

With respect to the COS budget, a number of studies have recently addressed this question (Angert et al., 2019; Hattori et al., 2020; Davidson et al., 2021), so the idea is hardly new. And while your approach bears some potential to constrain the COS budget and processing in the stratosphere, the 1D approach is not suitable to go beyond the mass balance approach employed in these studies that also consider regional and temporal variabilities to isolate the fractionation for individual COS sources and sinks (also see my specific comment below regarding page 19, line 3 to page 20, line 12).

I'm not saying that the use of isotopic information isn't valuable, and at the end of your conclusions, you provide a useful discussion of this question. But this is more an outlook rather than anything conclusive: large uncertainties in the fractionation parameters of the many processes involved and too few observations, particularly in the stratosphere, prevent your 1D model approach from really adding new knowledge or reducing uncertainties. With more information on the fractionation of sources, sinks and processes, and more observations of sulfur isotope ratios in the atmosphere, there is certainly a lot of potential (although I'm not convinced that the 1D approach is the most efficient one, and that it is not worth the – probably substantial – extra effort implementing isotope fractionation in a full 3D atmospheric model).

3. What are the largest uncertainties in the COS isotopic budget?

Uncertainties in the COS isotopic budget are discussed in your introduction, methods, and in the context of your sensitivity studies. From what I can see, the discussion adequately reflects the state of the art (i.e. the information available from the literature) and in your sensitivity runs, you encompass most of these uncertainties.

Because they are all taken from the literature, the uncertainties themselves are *a priori* rather than *a posteriori* knowledge in your study. And because the combined uncertainties related to different terms in your model offer more degrees of freedom than you have constraints (e.g. from observations), you can hardly draw any conclusions with respect to the significance or importance of individual uncertainties (again, see my specific comment below regarding page 19, line 3 to page 20, line 12).

Compared to the previous studies by Brühl et al. (2012) and Sheng et al. (2016), I don't see that your 1D column model approach offers any new information in terms of science, and I would hardly call it "*pio-neering*" (page 4, line 2) in spite of the fact that the isotope information is added as a novel feature. Therefore, I do not consider the manuscript publishable in ACP in its present form.

This does not mean that the work that you put into implementing and parameterizing the isotope information in a model isn't useful and important, and I strongly encourage you to rewrite and publish it as a proof-of-concept-study, for example as an ACP Technical Note or as a GMD Article. The results on the COS photolysis fractionation (where you can reduce uncertainties to some extent, see some of my specific comments below) could still be included to show the potential when the necessary information (in this case the SSA fractionation observed by Castleman et al., 1974) is available.

Below, I include a number of specific comment that either illustrate my criticism above further, and/or may help you in rewriting the manuscript.

Specific comments:

- <u>Page 1, line 17</u>: It should be "stratospheric sulfate aerosol" (you should consistently use the same term throughout the paper)
- <u>Page 2, line 2</u>: This statement is somewhat inaccurate. COS is the most abundant reduced sulfur compound in the troposphere (at all times) and it is the largest contributor to SSA in volcanic quiescent times.
- Page 2, line 5: The expression "above the ozone layer" is not accurate. COS photolysis can occur only when there is sufficient UV radiation, and if you look at the stratospheric COS distribution (e.g. in Barkley et al., 2008) and at the modelling study by Brühl et al. (2012), then it is most relevant in the tropics above ~ 20 km altitude.
- <u>Page 2, lines 7 8</u>: I suggest to elaborate more on the open questions regarding COS sources and sinks; "still poorly constrained" is not a precise term, and personally, I would not use it in the context of the COS budget. Issues related to the COS budget were also discussed in some detail by Kremser et al. (2016), and I think we have come one or two steps further (towards better constraining the budget) since that review was published.
- <u>Page 2, lines 13 14</u>: Again, this statement is inaccurate. In current models, a 0.7 % yield is assumed everywhere. The older laboratory studies that you cite already hypothesized that this yield could be zero in polluted conditions (based on efficient quenching by NOx), and some rather recent evidence (e.g. Jernigan et al., 2022) shows that COS production from DMS could potentially (!) be even higher than 0.7 % under "extremely pristine conditions".
- Page 2, lines 16 23: For the first sentence, Notholt et al. (2003) is not a good reference, because the statement was merely an introductory sentence in that paper. The first to suggest this was, in fact, Crutzen (1976). And I think we can now say that COS is considered not only a major but the most important source of SSA under volcanically quiescent conditions (Kremser et al., 2016, and references therein, which are the ones you also cite later in the paragraph). I really don't think there is still much debate about the contribution of COS to SSA in volcanically quiescent times (that concerns also the same statement in your abstract on page 1, line 4). While the Kjellström (1997) reference was important at the time (25 years ago) by showing that the net COS flux to the stratosphere was probably lower than some people thought and that some SO₂ could at least make it to the lower stratosphere, there now exist comprehensive studies (cited in your manuscript) that give us a pretty good overall picture, even though there are still some uncertainties on the exact numbers.
- <u>Page 3, line 6</u>: Given the evidence in the work cited (in particular Hattori et al., 2020), "*it is expected that anthropogenic COS is likely to be more depleted than oceanic COS*" is probably an understatement. I suggest to make this statement more affirmative. And it would be logical to discuss the role of COS sink processes already here: what do we know about the fractionation related to the vegetation and soil sinks on land? As a side note in this context, I'm not convinced that I we can expect a clear ocean source signature with one single number for δ^{34} S. For direct COS exchange between the ocean and the atmosphere, while the ocean is a net source, regional, diurnal and seasonal fluxes vary and are often enough negative because of a complex interplay between COS production and removal processes in the water. The magnitude of the net flux varies with region and season, and I would expect the net isotope fractionation effect to so as well.
- Page 3, lines 11 12: Better write "...within the range of -5 to 0 ‰ calculated by Schmidt et al. (2012) for this reaction."

- <u>Page 3, lines 14 19</u>: This is a good discussion of the large uncertainty range of COS photolysis. This uncertainty (here, the term "*not well constrained*" would be more than appropriate) of the key parameter for the fractionation of SSA produced from COS is absolutely worthy to be addressed, and I see it as one of the few questions where your modelling study in its present form can make a significant contribution to reduce uncertainties (see comments further below and my general comments above).
- <u>Page 4, lines 10 11</u>: This sentence is odd and "*explore*" does not seem to be the right word here. After all, you designed the model and parameterized the sulfur chemistry.
- <u>Page 4, lines 15 16</u>: these uncertainties are indeed rather large. In my opinion, they give the model far too many degrees of freedom for it to produce meaningful outcome, and you do not have enough information to resolve uncertainties except for the fractionation of COS photolysis.
- <u>Page 4, lines 21 23</u>: How do you treat the vast spatial and temporal differences? Do you just average everything into one global number?
- <u>Page 5, lines 4 6</u>: When I read this, my first feeling is that this approach oversimplifies things and may not fully represent the real processes in the atmosphere, e.g. in terms of contribution from different oxidation pathways. It is clear that with the 1D approach, you do not account for regional differences, but does your model account for seasonal and diurnal variability?
- <u>Page 5, line 19 page 6, line 3</u>: I suspect that the tuning of the eddy diffusivity to nudge the model to the stratospheric COS turnover given in the previous literature (cf. my general comment above) is needed to account for tropical upwelling that obviously is important for the upward COS flux in the full 3D models. If I understand this correctly, then this means that, in terms of vertical transport, you are tuning your 1D model to represent a tropical regime, while in terms of photochemistry, it simulates more or less mid-latitude conditions, i.e. with the standard US oxidant profiles (page 4, line 5) and using a mean zenith angle of 57.3 ° (page 6, line 25).
- Page 6, lines 25 27: Continuing from my last point above, I don't find the overall setup convincing, because the overall rate balance between photolytic processing and ascend/mixing in the 1D model does not, as far as I can see, represent anything I would expect in the real atmosphere, which in turn poses questions with respect to the results, at least when you interpret them quantitatively.
- <u>Page 9, lines 6 7</u>: Please give the references that your assumption for these numbers is based upon!
- <u>Page 9, lines 12 13</u>: Is the overhead burden of sulfur gases really a key factor here? I would expect overhead ozone to be the more important factor determining photolysis rates.
- <u>Section 2.2</u>: If I understand this correctly, you combine all COS sources and sinks into one integrated global source and one sink (that you term "*dry deposition*"), each with one respective average δ^{34} S that you derive from literature information on isotope signatures and source split. With the purpose of trying to keep things simple, this approach is fine, but it makes me wonder how you can then possibly derive and information or constraints on the COS budget with your model? My impression is that any information on the COS budget and isotope effects related to the budget is taken from the recent literature.
- <u>Section 3.1, first paragraph, top panel of Fig. 1</u>: While you investigate in some detail how well your model fits observed isotope ratios, there is no comparison for the mixing ratio profiles. This is surprising, because representative profiles can be derived from a wealth of satellite data. With such a comparison, you could test if your model really does represent global averages, and how it compares to real profiles when different latitude bands and seasons are considered. That could set the scene for looking at observed isotope ratios, which I consider meaningful only if the concentrations are also reasonably well represented.

In other words: why would you trust the simulated isotope fractionations and compare them to observations if you are not confident that the mixing ratios are represented well?

- <u>Page 15, lines 8 9</u>: Here, it would be interesting to state how much SSA at 18.5 km originates from COS and the tropospheric SO₂ source in your model (when looking at the top panels of Fig. 1, both seem to be significant, and based on your Figure S10, the tropospheric SO₂ contribution dominates below 20 km). If there is still a significant contribution from SO₂, then this would add yet another uncertain tuning parameter (that is independent on COS) when it comes to bring your model close to the observed δ^{34} S.
- <u>Page 16, lines 3-5</u>: This statement is even more inaccurate, unclear and confusing than the earlier one (page 2). If you use the 0.7 % as one global number, then simply say so, cite the references and leave it at that!
- <u>Page 16, line 8</u>: What do you mean by "*adjusted*"? First you include an "*additional source*" that was introduced to compensate for the larger flux to the biosphere, and then you adjust the flux to the biosphere in your model. To me, that does not make sense. And why did you adjust to 527 ppt and not 500 ppt?
- <u>Page 16, lines 9 11</u>: Could this be an effect of the OH profile chosen in your 1D model, and that in other models and the real atmosphere, OH and other compounds as well as physical conditions vary in a complex manner?
- <u>Page 16, line 19</u>: Note that this is a little less than half the burden of 283.1 Gg S given in Sheng et al. (2015), which is in excellent agreement with a burden of 280 Gg S derived from satellite observations (Kloss, 2017).
- <u>Page 18, lines 6-8</u>: This is true, and the sensitivity test for the photolysis reaction is the one point where I see this study making a contribution towards reducing uncertainties.
- <u>Page 18, lines 10 12</u>: The described transport enrichment is a response to δ^{34} S gradients in the model at steady state and the model does not include any *mechanistic* fractionation during transport, right?
- <u>Page 18, line 19</u>: This should be "...the subsequent sulfate $\delta^{34}S$ signature is affected..."
- <u>Page 18, lines 20 28</u>: As stated above, the sensitivity study on the photolysis fractionation appears to yield a viable result, although it is to some extent compromised by the likely sensitivity of the Castleman et al. (1974) observations towards tropospheric SO₂ (at least in your 1D model, see my comment above related to page 15, lines 8 9).
- Page 19, line 3 to page 20, line 12: These sensitivity tests yield expected, almost trivial, results. I wonder, what we can really learn from them. What you show is that the fractionation of stratospheric SO₂ and sulfate respond to the fractionation of tropospheric COS, which is expected and rather trivial. And when you state that the "set of sensitivity simulations very clearly shows that COS is a source of stratospheric SO₂ and sulfate" then, again, this is a direct result of the design of the model with a prescribed COS flux to the stratosphere. Clearly, this does not in any way go beyond what was already known. And when it comes to looking at uncertainties in the various fluxes, reaction rates and fractionation effects, you have at least three rather uncertain parameters, and within the margins explored in your sensitivity runs, they all have very similar effects (e.g., the top and bottom panels in Figure 4 look almost exactly the same, except for the colors and legends), so the comparison to stratospheric observations does not help to reduce any of these uncertainties. This would be true even if you had perfectly accurate profiles of observed δ^{34} S in stratospheric SO₂ and sulfate. Clearly, to explore and hopefully reduce uncertainties of

the various parameters affecting δ^{34} S of COS in the atmosphere, you need to look at regional and temporal variabilities in order to separate the effects of individual sources and processes, and a 1D model is not the right tool to do so.

- <u>Page 20, line 15</u>: I'd say the current knowledge of the system was imported rather than integrated (see my general comments)!
- <u>Page 22, lines 18 19</u>: In my opinion, that is the one viable conclusion that can currently be drawn from your study. And even that depends on the accuracy of an average δ^{34} S in SSA from a few measurements in the 1970s between 18 and 19 km, where, at least in your model, the contribution from tropospheric SO₂ is large if not dominant (cf. my other comments above).
- <u>Page 22, lines 21 34</u>: This is an interesting discussion on its own, but it does not help the overall purpose of the paper. And besides the fractionation of the $CS_2 \rightarrow COS$ reaction, the $\delta^{34}S$ of the CS_2 source may be equally or even more important.
- Page 23, line 7: At least for Lennartz et al. (2017), it is not true that DMS oxidation was "*postulated*" as the missing COS source. It was rather not ruled out as one possibility to explain part of the missing source, but more arguments were given against than for this possibility.

References not already listed in the Nagori et al manuscript:

Barkley, M. P., Palmer, P. I., Boone, C. D., Bernath, P. F., and Suntharalingam, P.: Global distributions of carbonyl sulfide in the upper troposphere and stratosphere, Geophys. Res. Lett., 35, L14810, https://doi.org/10.1029/2008GL034270, 2008.

Jernigan, C. M., Fite, C. H., Vereecken, L., Berkelhammer, M. B., Rollins, A. W., Rickly, P. S., et al.: Efficient production of carbonyl sulfide in the low-NOx oxidation of dimethyl sulfide. Geophys. Res. Lett., 49, e2021GL096838, <u>https://doi.org/10.1029/2021GL096838</u>, 2022.

Kloss, C.: Carbonyl Sulfide in the stratosphere: airborne instrument development and satellite based data analysis, Ph.D., Chemistry Department, Bergische Unviersität Wuppertal, Wuppertal, <u>http://elpub.bib.uni-wuppertal.de/servlets/DocumentServlet?id=7570&lang=en</u>, 2017.