



- 1 Technical note: Identification of chemical composition and source of fluorescent
- 2 components in atmospheric water-soluble brown carbon by excitation-emission
- 3 matrix with parallel factor analysis: Potential limitation and application
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19 Abstract

Three-dimensional excitation-emission matrix (EEM) fluorescence spectroscopy 20 is an important method for identification of occurrences, chemical composition, and 21 sources of atmospheric chromophores. However, current knowledge on identification 22 23 and interpretation of fluorescent components is mainly based on aquatic dissolved organic matter and might not be applicable to atmospheric samples. Therefore, this 24 25 study comprehensively investigated EEM data of different types of strong 26 light-absorbing organic compounds, water-soluble organic matter (WSOM) in 27 different aerosol samples (combustion source samples and ambient aerosols), soil dust, and purified fulvic and humic acids by an EEM-parallel factor method. The 28 results demonstrated that organic compounds with high aromaticity and strong 29 30 electron-donating groups generally present strong fluorescence spectra at longer 31 emission wavelength, whereas organic compounds substituted with electron-withdrawing groups have relatively weaker fluorescence intensity. In 32 particular, aromatic compounds containing nitro groups (i.e., nitrophenols), which 33 34 show strong absorption and are the major component of atmospheric brown carbon, exhibited no significant fluorescence. Although fluorescent component 1 (235, 35 270/330 nm) in ambient WSOM is generally considered as protein-like groups, our 36 findings suggested that it is mainly composed of aromatic acids, phenolic 37 compounds, and their derivatives, with only traces of amino acids. Principal 38 component analysis and Pearson correlation coefficients between mass absorption 39 efficiency at 365 nm (MAE₃₆₅) and humification index (HIX), C1, C2, and C3 40





- indicated that the highly aromatic and oxidized fluorescent component 3 may be an
 important contributor to the light-absorption capacity of ambient WSOM. These
 findings provide new insights for the analysis of chemical properties and sources of
 atmospheric fluorophores using the EEM method.
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46 **1. Introduction**

47 Brown carbon (BrC) is a class of organic compounds with light absorption in 48 near-ultraviolet and visible wavelength, and is ubiquitous in ambient aerosols, cloud 49 or fog, and rainwater (Fu et al., 2015; Laskin et al., 2015). Owing to its strong light-absorption capacity, BrC can cause up to 45 % solar radiation absorption by 50 atmospheric aerosols, and have potential effects on regional and even global climate 51 52 (Zhang et al., 2013). In addition, BrC also participates in atmospheric photochemical reactions, affects the physicochemical properties of atmospheric aerosols (Laskin et 53 al., 2015; Tang et al., 2020b), and potentially, can be activated to form reactive 54 oxygen species that cause adverse effects on human health (Bates et al., 2019; Cao et 55 56 al., 2021).

57 Excitation-emission matrix (EEM) fluorescence spectroscopy is a highly 58 sensitive and widely used analytical technique for the identification of chemical 59 characteristics and sources of chromophores in dissolved organic matter (DOM) in 60 aquatic environments (Murphy et al., 2010, 2013; Zhang et al., 2014). Recently, 61 EEM has been further extended and frequently applied for the investigation of 62 water-soluble organic matter (WSOM), such as light-absorbing organic compounds





63	in atmospheric aerosols and fine particles from combustion process (Chen et al.,
64	2020; Fan et al., 2016; Wu et al., 2020). For instance, humic-like substances (HULIS)
65	and protein-like substances (PRLIS) have been identified as important fluorescent
66	components in combustion-derived particles and ambient aerosols (Cao et al., 2021;
67	Matos et al., 2015; Wu et al., 2020). Chen et al. (2016) used EEM coupled with
68	parallel factor analysis (PARAFAC) and high-resolution mass spectrometry to
69	identify chromophores in ambient aerosols, and proposed that fluorescent
70	components with longer excitation (Ex)/emission (Em) wavelengths comprise more
71	highly oxygenated groups (Chen et al., 2016b). In addition, further application of the
72	EEM method has also revealed that the concentration and types of fluorophores
73	obviously vary during atmospheric processes, such as photolytic aging of biomass
74	burning(BB)-derived chromophores (Aftab et al., 2018; Fan et al., 2020; Tang et al.,
75	2020b). Therefore, the EEM method has significant potential for the characterization
76	(types, sources, and evolution) of atmospheric BrC.

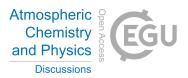
77 However, application of the EEM method for the identification of atmospheric BrC has some limitations. It is well known that the present identification, 78 classification, and interpretation of fluorescent components in atmospheric WSOM 79 80 are mainly based on the fluorescence peak position of DOM in aquatic environments (Coble, 1996; Wünsch et al., 2019). Nonetheless, the chemical and molecular 81 composition and source of WSOM in atmospheric aerosols significantly vary from 82 those of WSOM in aquatic environments (Graber and Rudich, 2006; Laskin et al., 83 2015); hence, the current fluorescence criterion derived from aquatic environments 84





could lead to some inaccurate description of the fluorescent components in 85 atmospheric WSOM. For instance, the EEM region at Ex/Em = 235(270)/330 nm is 86 assigned to PRLIS and/or tryptophan-like substances in aquatic environment (Coble, 87 1996), but is also associated with non-nitrogen species such as polyphenols in 88 89 atmospheric WSOM (Chen et al., 2016b). The EEM region at peak M (290-315/370-420 nm) is considered as a typical signal of marine-derived HULIS (Coble, 90 91 2007; Zhao et al., 2019), but the source of this peak should be cautiously 92 investigated when interpreting BrC in continental aerosols. In addition, the 93 intensities of fluorescent species are not always linearly correlated with their concentrations, which can be affected by the aromatic ring system, number and types 94 of functional groups, and inner-filter effects (IFEs), thereby leading to a greater 95 uncertainty in intensity measurements (Andrade-Eiroa et al., 2013; Chen et al., 2020; 96 97 Wang et al., 2020). Atmospheric BrC is composed of complex organic molecules with various properties (Chen et al., 2020), and only a subset of the BrC molecules 98 that contain functional groups are capable of fluorescence emission upon relaxation 99 from an excited state (Andrade-Eiroa et al., 2013). Hence, interpretation of 100 fluorescence data may only correspond to fluorescent chromophores and may not be 101 representative of BrC as a whole (Chen et al., 2020; Wang et al., 2020). All these 102 factors limit further application of the EEM method for the analysis of atmospheric 103 104 BrC. Therefore, it is essential to investigate the light-absorbing species that can be detected by EEM and obtain important information for identifying the chemical 105 compositions and possible sources of these species. 106





107	Accordingly, in the present study, the EEM profiles of a series of BrC model
108	compounds and WSOM isolated from primary combustion samples, soils, and
109	atmospheric aerosols were investigated. The chemical characteristics and sources of
110	the main fluorophores were interpreted according to the fluorescence location and
111	intensity, and the chemical structure of the model compounds and source samples
112	analyzed. Then, atmospheric aerosols in Guangzhou (GZ) and Chuzhou (CZ) cities
113	were collected and fluorescent chromophores within the water-soluble fraction were
114	identified to estimate the application of the EEM-PARAFAC method in
115	characterizing atmospheric BrC. The results obtained help to broaden application of
116	the EEM-PARAFAC method to study atmospheric BrC.

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118 2. Materials and methods

119 2.1. Materials

For accurate identification of the chemical composition and structures of 120 fluorophores in atmospheric BrC and for assessment of application of the EEM 121 method to examine atmospheric BrC, a total of 136 samples were investigated in this 122 study. The samples comprised: (1) 35 BrC model compounds, including phenolic 123 compounds, aromatic acids, nitroaromatic compounds (NACs), PRLIS, 124 N-heterocyclic compounds, and polycyclic aromatic hydrocarbons (PAHs) and their 125 derivatives (Table S1). These compounds are usually detected in ambient samples 126 and have been considered as the typical BrC model compounds; (2) 13 primary 127 combustion source samples collected from biomass burning, coal combustion(CC), 128





129	and vehicle emission (VE); (3) five soil samples obtained from the rural area of
130	Guangdong Province, China, with different vegetation; (4) six purified fulvic and
131	humic acids (FAs and HAs, respectively) kindly provided by Professor Weilin Huang
132	(Rutgers, The State University of New Jersey, NJ, USA); and (5) 34 diurnal fine
133	particulate matter ($PM_{2.5}$) samples collected from 6 to 22 April, 2021 at GZ and CZ,
134	respectively. In addition, 43 annual $PM_{2.5}$ samples were collected from February
135	2018 to January 2019 at the GZ site and classified as wet and dry season
136	atmospheric $PM_{2.5}$ samples (for detailed information, see Test S1 of supporting
137	information (SI)).

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139 2.2. Standard solution and aqueous extraction of ambient samples

Solutions of model organic compounds were prepared by dissolving a certain amount of dried solid or liquid samples in Milli-Q water or methanol. The ambient aerosol and soil samples were ultrasonically extracted with ultrapure water for three times, and the supernatants were filtered using a 0.22-µm PTFE syringe filter to isolate the WSOM. The specific separation and purification methods have been published in previous studies (Chen et al., 2020; Fu et al., 2015; Wang et al., 2020; Yan and Kim, 2017) and presented in SI (Test S2).

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148 **2.3. EEM-PARAFAC analysis**

The EEM fluorescence spectra of the aqueous extraction of the samples in 1-cm
quartz cuvettes were recorded using a three-dimensional fluorescence
spectrophotometer (Aqualog; HORIBA Scientific, USA) at room temperature. The





152	scanning ranges for Ex and Em was 200-500 and 250-550 nm, respectively. The
153	wavelength increment of the Ex and Em scans was 5 nm, integration time 0.5 s, and
154	Milli-Q water (18.2 Ω) used as blank reference. The absorbance measurements were
155	used to correct the EEM for IFEs as described previously (Fu et al., 2015) if the
156	absorbance was > 0.05 at 250 nm (Murphy et al., 2013; Tang et al., 2020a).
157	Background samples were also analyzed and the background values were subtracted
158	from the values obtained for all the samples. To avoid concentration effects, the
159	fluorescence spectra were normalized by the water Raman area to produce Raman
160	Unit (R.U.) and further by organic carbon concentration of the samples to the
161	normalized fluorescence intensities (R.U./(mg C/L)) (Yang et al., 2022) are shown in
162	Table S2.

163 The PARAFAC modeling procedure was conducted in MATLAB 2014b (Mathwork.Inc, USA) using the drEEM toolkit (Murphy et al., 2018; Wünsch et al., 164 2019). The PARAFAC was computed using two to nine component models, with 165 non-negativity constraints and residual analysis, and split-half analysis was 166 employed to validate the number of fluorescent components. Based on the results of 167 the split-half and core consistency analyses, three-component models were chosen 168 for further investigation. The relative contribution of individual chromophores was 169 estimated by calculating the maximum fluorescence intensities (Fmax: maximum 170 fluorescence intensity of the identified fluorescent components; relative content (%) 171 $= F_{max}/\Sigma F_{max}$) (Chen et al., 2020; Fan et al., 2020). 172

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175 **3. Results and discussion**

176 **3.1. Fluorescence properties of BrC model compounds**

To identify whether the light-absorbing species possess fluorescence, a series of BrC model compounds were tested by the EEM method, and the fluorescence profiles are shown in Fig. S1. The results revealed that the location and intensity of the fluorescence peaks of different compounds were different, which varied with the distinct functional groups and aromatic conjugate system.

182 Although phenolic compounds are important light-absorbing species in atmospheric BrC (Smith et al., 2016; Yu et al., 2014, 2016), not all of them exhibit 183 strong fluorescence. As shown in Fig. S1a, a strong fluorescence peak in the EEM 184 185 spectrum of phenol was observed at Ex/Em = 270/295 nm. When the phenol compounds were substituted with electron-donating groups (e.g., hydroxyl), all of 186 the stronger fluorescence peaks were obviously red-shifted to 310-320 nm (e.g., 187 catechol, hydroquinone, and 2-methoxyphenol). However, phenolic compounds 188 189 substituted with electron-withdrawing groups (e.g., carboxyl and aldehyde) displayed weaker or even no fluorescence (Fig. S1a). These differences could be 190 owing to the ability of the electron-donating groups to form a larger conjugate 191 system coupled with the benzene ring and decrease the $\pi \rightarrow \pi^*$ transition energy, thus 192 193 leading to an increase in the Em wavelength (i.e., red shift) and variation in fluorescence intensity. In contrast, the electron-withdrawing group can reduce the 194 conjugate structure formed by the benzene ring and hydroxyl group, reducing the 195

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197 Aromatic acid and its derivatives are also important light-absorbing organic compounds in atmospheric BrC. Owing to the negative effects of carboxyl group, a 198 weak fluorescence peak (275/315 nm) was identified for benzoic acid and no 199 200 fluorescence was detected for benzene polycarboxylic acids, such as phthalic acid, terephthalic acid, and trimesic acid (Fig. S1b). However, when benzoic acid was 201 202 substituted with electron-donating groups (e.g., hydroxyl, methoxy), higher intensity 203 fluorescence peaks were observed. Two strong fluorescence peaks at 230/405 and 204 290/405 nm were identified for 2-hydroxybenzoic acid substituted with only one hydroxyl group. These peaks could have been the result of the ortho structure of the 205 hydroxy and carboxyl groups, which is favorable for the formation of intramolecular 206 207 hydrogen bond and generates a double-ring conjugate structure, reducing the transition energy and thereby presenting strong UV absorption and fluorescence 208 (Andrade-Eiroa et al., 2013). 209

fluorescence intensity (Andrade-Eiroa et al., 2010; Andrade-Eiroa et al., 2013).

The N-containing compounds, especially NACs, have strong light absorption, 210 211 and have been reported to be the major components of atmospheric BrC, accounting for more than 60 % of the total light absorption intensity at 300-500 nm (Huang et 212 al., 2021; Lin et al., 2016; Lin et al., 2017). However, most of the NACs did not 213 exhibit any fluorescence (Fig. S1c), similar to that reported in a previous study by 214 215 Chen et al. (2020), which could be owing to the significant reduction in the electron density of benzene ring by the nitro (-NO₂) group—strong electron-withdrawing 216 group-thereby weakening the fluorescence. 217





218	Tryptophan and tyrosine are the two most studied PRLIS species, and their
219	EEM spectra are generally used as standards for comparison with fluorophores in
220	atmospheric WSOM and aquatic DOM (Coble, 1996, 2007). As shown in Fig. S1d,
221	the Ex/Em peaks at 275/300 and 275/350 nm corresponded to tyrosine and
222	tryptophan, respectively. The maximum Em wavelength of phenylalanine was more
223	inclined to short wavelength (280 nm) and with much weaker fluorescence intensity.
224	Moreover, the fluorescence peaks of PRLIS were obviously overlapped with phenols
225	and aromatic acids (Fig. S1a, b). It must be noted that the concentrations of phenols
226	and aromatic acids were significantly higher than those of tryptophan and tyrosine in
227	the atmospheric samples (Table S2); therefore, the aerosol BrC fluorophores in these
228	regions are more likely to have originated from phenols and aromatic acids rather
229	than PRLIS.

The N-heterocyclic compounds such as pyrrole, pyridine, and imidazole are 230 commonly identified in atmospheric samples (Dou et al., 2015; Jiang et al., 2019; 231 Kosyakov et al., 2020). However, no fluorescence was observed for these three 232 species in the present study, indicating that the absorbed energy may have been 233 consumed relaxation vibration (Fig. S1e). 234 by or Nevertheless, imidazole-2-formaldehyde produced two strong fluorescence peaks at 290/440 and 235 350/440 nm, formed from the oxidation of imidazole, suggesting that some 236 N-heterocyclic compounds from secondary reaction may exhibit strong fluorescence 237 at higher wavelength in atmospheric BrC (Ackendorf et al., 2017). 238

239 PAHs and its derivatives are mainly formed from incomplete combustion





processes, and are important components of BrC (Chen et al., 2020; Lin et al., 2017; 240 Mahamuni et al., 2020). As shown in Fig. S1f, all PAHs exhibited strong 241 fluorescence emission, with its peak location associated with the conjugated 242 aromatic system. Naphthalene presented a fluorescence band located at the 243 244 maximum Em wavelength of about 325 nm. As expected, with the increasing size of the π -bond system and degree of conjugation, the fluorescence band moved toward 245 246 the longer wavelength range, and a new Em band was observed at 360-390 nm for 247 3-4-ring phenanthrene and pyrene, and at 400–500 nm for \geq 5-ring PAHs 248 (Mahamuni et al., 2020). The fluorescence spectra of high-ring PAHs were more complex because of more types of double bonds. As shown in Fig. S1f, the intensity 249 and location of the fluorescence peaks were also significantly changed when 250 251 different types of groups were substituted with PAHs. For example, 1-naphthol exhibited stronger EEM peak at a relatively longer wavelength (230, 290/460 nm) 252 owing to its high conjugated structure, when compared with naphthalene. This EEM 253 spectrum was located in the EEM region of FAs, implying that FAs are composed of 254 255 aromatic units and O-containing groups. In contrast, relatively weaker fluorescence was observed for 9-fluorenone, anthraquinone, and 2-naphthalenecarboxylic acid, 256 and no EEM signals were observed for 2-nitronaphthol (Fig. S1c), which was 257 substituted with a strong electron-withdrawing group (-NO₂). 258

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3.2. Fluorescence properties of BrC from different sources

As shown in Fig. S4a and S4b, BB and CC WSOM exhibited similar





262	fluorescence spectra, with two types of fluorescence peaks at Ex/Em \approx (230–
263	240)/(340–400) nm (peak A) and Ex/Em \approx (260–280)/(330–360) nm (peak B),
264	respectively. The two fluorescence peaks were similar to those previously reported
265	for BB WSOM and HULIS (Fan et al., 2020; Tang et al., 2020a; Yang et al., 2022).
266	In general, peak A mainly corresponds to the protein-like UV region, with a minor
267	contribution from fulvic-like substances, whereas peak B could be attributed to
268	tryptophan-like fluorophores. However, based on the results of the present study,
269	these two peaks could be mainly attributed to aromatic species such as aromatic
270	acids, phenolic compounds, and minor quantity of PAHs (e.g., naphthalene) (Fig. 1).
271	The fluorescence spectra of WSOM from two types of vehicles (diesel and gasoline)
272	also presented two fluorophores. A relatively strong fluorescence peak was observed
273	at the low Ex wavelength (Ex/Em $\approx 230/350$ nm) and a relatively weaker peak was
274	detected at the high Ex wavelength (Ex/Em $\approx 270/350$ nm) (Fig. S4c). These results
275	are consistent with those reported in previous studies on VE (Chen et al., 2020; Tang
276	et al., 2020a; Yang et al., 2022), and similar to the EEM fluorescence spectra of BB
277	and CC WSOM. However, the fluorescence ranges of vehicles WSOM were
278	obviously narrower, suggesting that BB and CC WSOM fluorescent components are
279	more complex.

Soil-derived DOM is also a primary source of atmospheric WSOM. As shown in Fig. S5a, two main fluorescence peaks located at Ex/Em = 230/430 and 320/430nm, respectively, were detected in the fluorescence spectra of soil DOM, which are similar to those reported in previous studies (Ge et al., 2021; Liu et al., 2009) and





284 particularly close to the position of FAs (Fig. S5b).

Secondary chemical formation is another important source of atmospheric 285 WSOM. For example, the aqueous-phase reactions of aldehydes with ammonium 286 sulfate (AS) can produce highly fluorescent species (Hawkins et al., 2016). 287 288 Glyoxal-AS and glyoxal/glycine reaction products fluoresce at 340/450 nm, whereas formaldehyde-AS reaction product fluoresce at 250/430 nm. Secondary organic 289 290 aerosols (SOAs) produced in the limonene/O₃ system have been reported to strongly 291 fluoresce in the presence of NH₃ (Bones et al., 2010). In addition, aging of primary 292 organic compounds has also been found to change the fluorescence spectra (Lee et al., 2013; Li et al., 2021; Powelson et al., 2014). For instance, aging of syringic acid 293 with OH radicals caused the initial fluorescence band to move toward the long 294 295 wavelength range, producing a new band at a broad Em band at 400-600 nm. Similarly, the fluorescence peaks red-shifted (e.g., from 260-270/360 nm to 280-296 290/390-400 nm) during the O₃ aging process (Fan et al., 2020), suggesting the 297 degradation of the initial compound and formation of new secondary organic 298 299 compounds generally located at longer wavelengths, possibly with a high degree of aromaticity or highly oxidized functional groups (Chen et al., 2016a). 300

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302 3.3. Identification of chemical species and potential sources of fluorescent 303 components in ambient aerosols

The typical EEM spectra of atmospheric water-soluble light-absorbing compounds are shown in Fig. 2. Three fluorescence peaks were identified in the





306	aerosol WSOM samples: a stronger fluorescence peak at $Ex/Em = 230-250/360-420$
307	nm, and two relative weaker fluorescence peaks at $Ex/Em = 270-290/340-370$ nm
308	and 300-320/360-420 nm, respectively. Similar fluorescence bands have been
309	previously identified in the EEM fluorescence spectra of WSOM from $\text{PM}_{2.5}$ in
310	Xi'an of Northwest China and in Godavari of Nepal (Qin et al., 2018; Wu et al.,
311	2019). Although the fluorescence intensities varied with different sites and seasons,
312	the EEM spectra of WSOM were very similar, making it difficult to directly
313	distinguish the different samples solely based on the characteristics of the EEM
314	profiles. Therefore, a more powerful protocol named the PARAFAC method was
315	employed to identify the individual fluorophores in ambient WSOM.

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317 3.3.1. Identification and quantification of fluorescent components by the

PARAFAC method 318 To better explain the various fluorophores in different atmospheric WSOM 319 samples, the EEM spectra were resolved with the EEM-PARAFAC tool 320 (Andrade-Eiroa et al., 2013; Murphy et al., 2013). As shown in Fig. 3, three 321 fluorescent components (C1, C2, and C3) were identified in the atmospheric samples; 322 C1 occurred at relatively lower Em wavelength, exhibiting two fluorescence peaks at 323 Ex/Em = 235(270)/330 nm, C2 presented fluorescence peaks at around Ex/Em = 324 235(320)/390 nm, and C3 had a longer Em wavelength than C1 and C2, which was 325 located at Ex/Em = about 250(355)/455 nm. In general, these fluorescent 326 components have been interpreted based on the knowledge of fluorescence 327

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characteristics of aquatic DOM. Accordingly, C1 is considered to belong to the 328 329 typical PRLIS (Chen et al., 2016a; Chen et al., 2020), C2 is associated with fulvic-like substances or less-oxygenated HULIS (Chen et al., 2016a; Wang et al., 330 2020), and C3 is usually considered to correspond to terrestrial HULIS that are 331 332 highly oxygenated organic matter (Table S3) (Chen et al., 2016a; Zhou et al., 2017). However, it must be noted that the sources and transformation process are 333 334 significantly different for WSOM in aerosols and DOM in aquatic environment; 335 therefore, the fluorescence classifications of DOM might not be applicable to 336 atmospheric WSOM.

In general, the Ex and Em wavelengths of fluorescent components are mainly 337 associated with their chemical characteristics and structures (Table S1 and Fig. 1). In 338 339 the present study, C1 was similar to tryptophan-like fluorophore associated with PRLIS in rainwater (Zhang et al., 2014; Zhou et al., 2017) and fog water (Bianco et 340 al., 2014; Bianco et al., 2016). However, this fluorophore might also be possibly 341 related to the small molecular aromatic compounds, such as aromatic acids (e.g., 342 343 3,5-dihydroxybenzoic acid and 2-naphthalenecarboxylic acid) and PAHs (e.g., naphthalene, phenanthrene, and anthraquinone) (Fig. 1a) (Miyakawa et al., 2015; Wu 344 et al., 2019). In addition, this fluorophore could also contain traces of some phenolic 345 compounds, including catechol, hydroquinone, and 2-methoxyphenol. These organic 346 347 species might have probably been generated by various types of combustion process and atmospheric oxidation reaction. It must be noted that investigations of the 348 fluorescent components in atmospheric WSOM should not only consider their 349





position in the fluorescence spectrum, but also their concentration and possibility of 350 351 trapping. Many previous studies have reported that the concentration of amino acids in the atmospheric aerosols is almost negligible, when compared with that of lower 352 molecular weight aromatic compounds such as aromatic acids and phenolic 353 354 compounds (Table S2). Therefore, fluorescent components in this Ex/Em region could be attributed to non-nitrogen aromatic species, rather than PRLIS. Moreover, 355 356 this fluorophore overlapped with that of WSOM from combustion process such as 357 BB, CC, and VE (Fig. 1b), suggesting significant contribution of combustion 358 process.

When compared with C1, C2 exhibited a strong fluorescence peak at longer 359 Ex/Em wavelength of 235(320)/390 nm, implying that this fluorescent component 360 presented relative larger molecular size and higher aromaticity than C1 (Pöhlker et 361 362 al., 2012). As shown in Fig. 1a, the fluorescence of C2 is similar to that of aromatic 2-naphthalenecarboxylic acid, 2-hydroxybenzoic compounds acid, 363 (e.g., anthraquinone) and high-ring PAHs (e.g., pyrene, anthraquinone, anthracene, 364 chrysene) (Mahamuni et al., 2020), and overlaps with the fluorescence spectra of 365 FAs. In addition, this fluorophore has also been reported to be related to the 366 generation of SOAs from organic precursors emitted from biological/anthropogenic 367 emission and combustion process (Wang et al., 2020). For example, the 368 369 aqueous-phase reactions of aldehydes with AS has been proposed as an important source of atmospheric BrC, which present similar fluorescence spectral profiles 370 (Hawkins et al., 2016; Lee et al., 2013) (Fig. 1b). Besides, oxidative oligomerization 371





of phenols and their derivatives can also shift the Ex/Em wavelength of these substances to longer wavelength, falling into similar fluorescence region (Li et al., 2021; Tang et al., 2020a; Vione et al., 2019). As suggested by Chen et al. (2016a), this fluorescent component may be a less-oxygenated fluorescent group contributed by biomass combustion. Therefore, fluorophore C2 might be related to the derivatives of biomass burning and/or biogenic molecules, with relatively lower degree of oxidation (Chen et al., 2016a; Jiang et al., 2022).

379 C3 presented longer Em wavelength than C1 and C2, with two peaks at around 380 Ex/Em=250/455 nm and 355/455 nm (Fig. 3). This fluorescent component overlaps with the fluorescence of high-ring PAHs and their derivatives, such as fluoranthene, 381 benzo-b-fluoranthene, benzo-a-pyrene, indeno-123cd-pyrene, 1-naphthol, and 382 383 N-heterocyclic compounds, including imidazole-2-formaldehyde (Chen et al., 2020; 384 Mahamuni et al., 2020). Furthermore, this fluorescent component exhibited similar Ex/Em wavelength to that of FAs and HAs (Fig. S5b), suggesting the possible 385 contribution of soil dust, and thus could be assigned as HULIS (Lin and Guo, 2020). 386 387 Similar fluorescent substances have also been identified in the study of atmospheric aerosol fluorescent chromophores, such as "HULIS-1" at about 470 nm in Nagoya, 388 Japan (Chen et al., 2016a) and "component 2" at about 480 nm in Godavari, Nepal 389 (Wu et al., 2019). Based on the PARAFAC results with aerosol mass spectrometry 390 391 data, C3 was considered to be a fluorescent group with high oxygen content and high O/C ratio, close to that of aged organic aerosols (Chen et al., 2016a; Jiang et al., 392 2022) (Fig. 1b). It must be noted that low molecular weight organic compounds can 393





394	further undergo oligomerization to high molecular weight species with long Em
395	wavelength during the aging process (Hawkins et al., 2016; Li et al., 2021; Tang et
396	al., 2020b; Yu et al., 2016). The resulting compounds may present a more complex
397	structure than its precursor, probably owing to the presence of condensed aromatic
398	ring and other π -electron systems with a high level of conjugation; thus, atmospheric
399	aging is assumed to be a potential contributor to C3 (Barsotti et al., 2016; De
400	Laurentiis et al., 2013; Hawkins et al., 2016).

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402 3.3.2. Spatial and seasonal variations of fluorescent components in WSOM

The relative contributions of C1, C2, and C3 components to the total 403 fluorescence intensities $(F_{max}/\sum F_{max})$ were calculated (Fig. 4), and were found to be 404 405 similar for WSOM from CZ and GZ, exhibiting maximum C2 content and relatively lower C1 and C3 contents. Furthermore, these WSOM samples showed obvious 406 spatial and seasonal variations. First, CZ WSOM presented relatively higher C3 407 content, whereas GZ WSOM had relatively higher C2 content at the same sampling 408 409 period. Such differences in the composition of fluorescent components may be ascribed to the variation in the primary emission sources and atmospheric aging 410 process in the two sites. The relatively higher C3 content in CZ could be attributed to 411 the comparatively high contribution of soil dust in the suburban region. In contrast, 412 the relatively higher C2 content in GZ WSOM may be attributed to the 413 comparatively stronger atmospheric chemical reaction associated with bio-volatile 414 organic compounds (bio-VOCs) in the hot and humid region of GZ. This result was 415





416 consistent with the relatively higher humification index (HIX) and normalized
417 fluorescence volume (NFV) values (log(NFV)) of CZ WSOM (Fig. S6) (Chen et al.,
418 2020; Yang et al., 2022).

In addition, the resolved Ex and Em spectra for GZ WSOM were also similar in 419 420 different seasons, implying that the types of fluorophores contributing to WSOM were predominantly the same throughout the year. However, the compositions of 421 422 fluorescent components varied in different seasons. In the dry season (October-423 March), WSOM showed relatively higher contents of C3 fluorophores, whereas in 424 the wet season (April-September), slightly higher contents of C2 fluorophores were detected (Fig. 4) (Chen et al., 2020; Wang et al., 2020). These differences might 425 possibly be associated with the variations in the source composition and aging 426 427 effects of BrC in different seasons. The higher content of C3 in WSOM in the dry season suggested the occurrence of more highly aromatic and highly oxidized 428 compounds. These results could be explained by the fact that more aged organic 429 aerosols and dust were transported from the northern region of China (Jiang et al., 430 431 2021). In contrast, the slightly higher C2 content in the wet season may be attributed to the relatively stronger secondary formation of bio-SOAs and photodegradation 432 effects in high-temperature and relative humidity season. 433

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435 **3.4.** Correlation with optical and fluorescence properties

It is well known that some organic substances that can generate fluorescenceshould absorb energy and transit from ground state to excited state and back to





ground state, which could produce corresponding fluorescence. However, it is very 438 439 difficult to clearly understand the absorption spectra and corresponding fluorescence spectrum of complex WSOM. As indicated in the present study, not all WSOM with 440 strong light absorption exhibit strong fluorescence (e.g., NACs). Hence, to elucidate 441 442 the association between light-absorbing chromophores and fluorescent components, the relationship between optical properties and fluorophores was analyzed by using 443 444 principal component analysis and Pearson correlation analysis (Fig. S7). The 445 MAE_{365} showed an obvious positive loading for principal component 1 (PC1) and was grouped with C3 and HIX. Moreover, the MAE₃₆₅ values were positively 446 correlated with C3 and HIX, indicating that light absorption by BrC is more 447 dependent on these fluorophores with long Ex and Em wavelengths and high 448 449 humification. These results also suggested that the oligomerized fluorescent products with high aromaticity and oxygenated groups may be the dominant factor affecting 450 the light absorption capacity of the fluorescent chromophores. Similar findings have 451 also been reported by Chen et al. (2020) and Tang et al. (2021). 452

453

454 4. Conclusion and future prospects

In this study, the fluorescence properties of BrC model compounds were investigated to determine the chromophoric species that can be evaluated by the EEM method. Accordingly, the aerosol WSOM in two sites (CZ and GZ) were investigated by the EEM-PARAFAC method, and the chemical characteristics and potential sources of fluorescent components were examined. The main conclusions





460 and future prospects are as follows:

461 (1) Fluorescent components have predominantly been evaluated based on the knowledge of fluorophores in aquatic DOM, which often leads to misinterpretation. 462 In the present study, the chemical characteristics of fluorophores in different Ex/Em 463 464 regions were discussed based on the fluorescence properties of BrC model compounds and their amounts in aerosols. In particular, the C1 fluorophore in 465 466 atmospheric WSOM, which has been frequently assigned to PRLIS because of the 467 similarity in fluorescence spectra, was demonstrated to mainly include aromatic 468 acids, phenolic compounds, and their derivatives, with negligible amount of amino acids. 469

(2) The fluorescence properties of target compounds are mainly influenced by 470 471 the aromatic system and characteristics of adjacent functional groups. Organic compounds with high aromaticity and strong electron-donating groups (e.g., 472 hydroxyl, methoxyl) generally exhibited strong fluorescence spectra at longer Em 473 wavelength, whereas organic compounds substituted with electron-withdrawing 474 475 groups presented relatively weaker fluorescence intensity. In particular, aromatic compounds containing nitro groups (i.e., nitrophenols) showed strong absorption and 476 were the major component of atmospheric BrC; however, they did not exhibit 477 significant fluorescence. Thus, fluorescence method could only measure a subset of 478 chromophores in aerosol BrC and should be used with caution for the investigation 479 of aerosol BrC. 480

481 (3) The EEM spectra for aerosol WSOM were very similar; however, the





relative contents of certain fluorescent components significantly varied with the sampling site and season. For example, more fluorescent components associated with dust and secondary oxidation of small molecular compounds from combustion emission were identified in GZ WSOM, whereas more fluorescent components derived from atmospheric chemical reaction of bio-VOCs were observed in CZ WSOM. In addition, GZ WSOM exhibited more highly aromatic and highly oxidized compounds in the dry season.

489 Although many studies have applied EEM-PARAFAC method to investigate 490 atmospheric WSOM and have obtained useful data, there are still challenges and gaps that must be addressed. First, caution should be taken for credible 491 interpretations of the fluorescent components in atmospheric WSOM because of the 492 493 differences in chemical characteristics of organic matter derived from different sources. In addition, the same fluorophores may exhibit different Ex/Em ranges and 494 intensities in different environmental conditions (e.g., pH, co-existing metal ions and 495 inorganic salts, etc.). Therefore, more theoretical and experimental studies are 496 497 necessary to understand the relationship between the fluorescent groups and positions of fluorescence peaks, as well as the influences of sources and chemical 498 formation process of the fluorescent groups on fluorescence peaks. 499

500

501 Data availability

502 The research data can be accessed in the Harvard Dataverse (https://
503 doi.org/10.7910/DVN/ULCIU9, Song, 2022)





504

505	Author contributions. J. Song designed the research. T. Cao and C. Xu, analyzed			
506	the model compounds and WSOM samples by UV-Vis and EEM. T. Cao and X. Fan			
507	resolved the EEM by PARAFAC tool. M. Li carried out the $PM_{2.5}$ sampling			
508	experiments. T. Cao and J. Song wrote the paper. J. Li, W. Jia, and P. Peng			
509	commented and revised the paper.			
510				
511	Competing interests. The authors declare that they have no conflict of interest			
512				
513	Acknowledgments. The present work was supported by the National Natural			
514	Science Foundation of China (42192514 and 41977188), Guangdong Foundation for			
515	Program of Science and Technology Research (2020B1212060053), and Guangdong			
516	Foundation for Program of Science and Technology Research (2019B121205006).			
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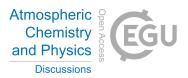
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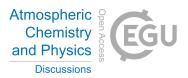
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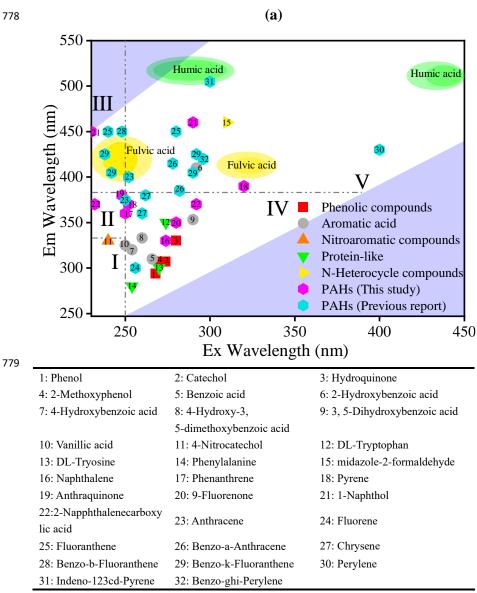




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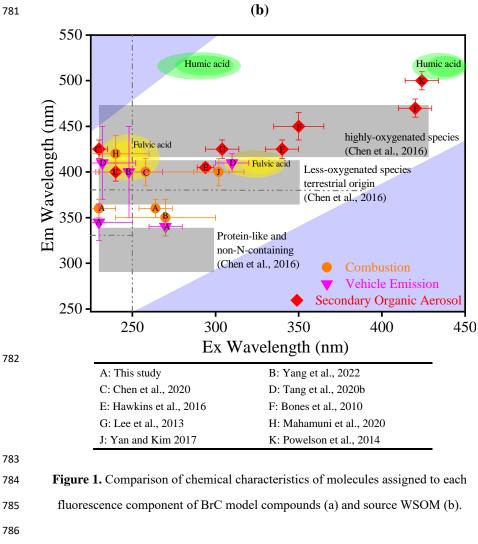




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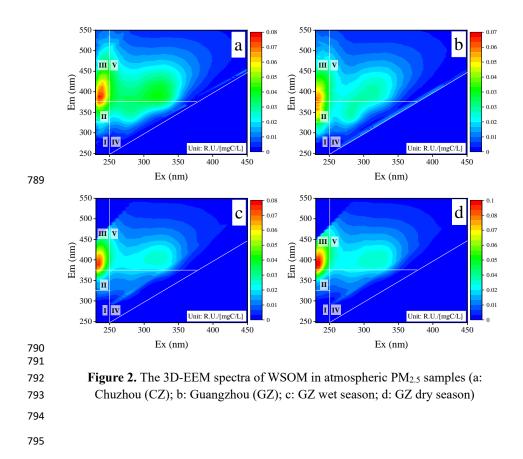


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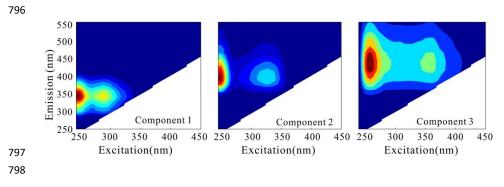












799 Figure 3. The EEM components derived from the PARAFAC model of WSOC in

atmospheric PM_{2.5} samples collected at Chuzhou (CZ) and Guangzhou (GZ) sites.

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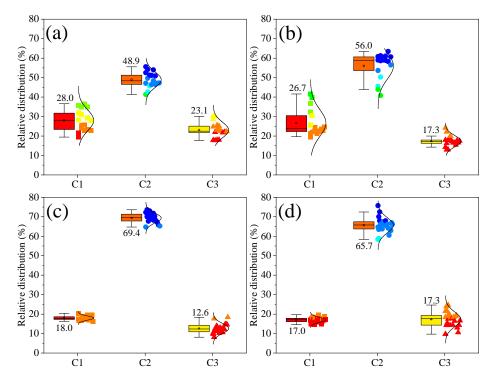


Figure 4. Relative contribution of individual fluorophores of atmospheric WSOM. (a:

805 Chuzhou (CZ); b: Guangzhou (GZ); c: wet season of GZ; d: dry season of GZ)

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