Sensitivity of convectively driven tropical tropopause cirrus properties to ice habit in high-resolution simulations

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Abstract. Cirrus clouds that form in the tropical tropopause layer (TTL) can play a key role in vertical transport through the upper troposphere and lower stratosphere, which can significantly impact the radiative energy budget and stratospheric chemistry. The lack of realistic representation of natural ice clouds habit in microphysical parameterizations can lead to uncertainties in cloud-related processes and cloud-climate feedbacks. The main goal is the investigation of the role of different cloud regimes and the associated ice habits in regulating the properties of the TTL. We compare aircraft measurements from the StratoClim field campaign to a set of numerical experiments at the scales of large-eddy simulations (LES) of the same case study that employ different microphysics schemes. Aircraft measurements from the southern slopes of the Himalayas captured high ice water content (HIWC) up to 2400 ppmv and ice particle aggregates exceeding 700 μm in size with unusually long residence times. The observed ice particles were mainly of liquid origin, with a small amount formed in situ. The corresponding profile of IWC from the ERA5 reanalysis corroborates the presence of HIWC detrained from deep convective plumes in the TTL but underestimates HIWC by an order of magnitude. In the TTL, only the scheme that predicts ice habits can reproduce the observed HIWC, ice number concentration, and bimodal ice particle size distribution. The lower range of particle sizes is mostly represented by planar and columnar habits, while the upper range is dominated by aggregates. Large aggregates with sizes between 600 and 800 μm have fall speeds of less than 20 cm s^{-1}, which explains the long residence time of the aggregates in the TTL. Planar ice particles of liquid origin contribute substantially to HIWC. The columnar and aggregates habits are in the in-situ range with lower IWC and number concentration. For all habits, the ice number concentration increases with decreasing temperature. For the planar ice habit, relative humidity is inversely correlated with fall speed. This correlation is less evident for the other two ice habits. In the lower range of supersaturation with respect to ice, columnar habit has the highest fall speed. The difference in ice number concentration across habits can be up to four orders of magnitude, and aggregates

occurring in much smaller numbers. We demonstrate and quantify the linear relationship between the differential sedimentation of pristine ice crystals and the size of the aggregates that form when pristine crystals collide. The slope of this relationship depends on which pristine ice habit sediments faster. Each simulated ice habit is associated with distinct radiative and latent heating rates. This study suggests that a model configuration nested down to LES scales with a microphysical parameterization that predicts ice shape evolution is crucial to provide an accurate representation of the microphysical properties of TTL cirrus, and thus the associated (de)hydration process.

1 Introduction

Stratosphere-troposphere exchange plays an important role in the composition of the upper troposphere and lower stratosphere (Dessler and Sherwood, 2004; Stohl et al., 2003). Trace gases such as stratospheric water vapor can directly and indirectly impact climate, radiation (Forster and Shine, 1999; Solomon et al., 2010; Gettelman et al., 2011), chemistry (Dvorstov and Solomon, 2001) and dynamics (eg. Manney et al 2014). On a global scale, air enters the stratosphere via the tropical tropopause (Brewer, 1949). In the upwelling branch of the Brewer-Dobson circulation (BDC), the water vapor injected in the tropical lower stratosphere is transported meridionally to the temperate stratosphere. Due to the slow nature of BDC, this process results in a longer residence time of the injected stratospheric water vapor. The tropical tropopause layer (TTL), which reaches from the tropically averaged level of all-sky zero net radiative heating (Gettelman and Forster, 2002; Sherwood and Dessler, 2001,2002) (14 km, 150 hPa, 355 K) to the upper boundary of 18.5 km (70 hPa, 425 K), is the gateway to the stratosphere (Fu et al., 2007, Fueglistaler et al., 2009). In addition to large-scale processes, determining how convection may contribute to cross-tropopause transport is essential for understanding Earth's climate (Anderson et al., 2017). To identify the key mechanisms that control the stratospheric water vapor, considerable attention has been given to the role of cross-tropopause convective outflow (e.g., Homeyer et al., 2017; Smith et al., 2017; Ueyama et al., 2018). The Asian summer monsoon (ASM) is the dominant contributor to annual water vapor peaks in the lower stratosphere and the associated Asian summer monsoon anticyclone (AMA) (eg Fueglistaler et al., 2005; Ueyama et al.,2018; Ploeger et al., 2015). The impact of deep convection on hydration (e.g., Khaykin et al., 2009) and dehydration of the stratosphere (e.g., Sherwood and Dessler, 2001) over the southern slopes of the Himalayas, the Bay of Bengal, and the South China Sea (Fu et al., 2006; Devasthale and Fueglistaler, 2010; Wright et al., 2011; Tissier and Legras, 2016; Legras and Bucci, 2020) is topic of a debate. The amount of water vapor injected into the stratosphere by overshooting convection is regulated by the cold point temperature (CPT) in a process known as freeze-drying (Avery et al. 2017, Brewer, 1949; Evan et al. 2013; Fueglistaler et al. 2005; Holton and Gettelman, 2001; Randel et al. 2006; Read et al. 2004; Schiller et al., 2009). Ice particles injected into the lower stratosphere bypass the cold trap which is the coldest region in the TTL that modulates the water vapor entering the lower stratosphere via cirrus formation (Jensen et al., 1996) and ice crystals sedimentation (Holton and Gettelman 2001), sublimate above the level of CPT, and contribute to stratospheric hydration (Anderson et al.,2017; Herman et al., 2017; Chemel et al., 2009; Khaykin et

al., 2009). Cirrus clouds trap heat originating from the Earth's surface, scatter incoming shortwave radiation and affect the intensity of large-scale circulation in the tropics, making them an important regulator (e.g., Randall et al., 1989; Dodion et al., 2008) and one of the largest source uncertainties in the prediction (Boucher et al., 2013) of Earth's climate. Spaceborne active remote sensing (i.e., CloudSat and Calipso) reported 35% of cirrus coverage within ±15° latitude and 56% within ±30° latitude of the equator (*Sassen et al.*, 2008, 2009). According Baran et al (2012), satellite measurements in the tropics revealed that cirrus cloud cover can be 60–80% at any given time. The net radiative effect of cirrus clouds depends on their macroscopic properties and the size distribution and shape of their ice particles (e.g., Jensen et al., 2009; Lawson et al., 2010; Baran, 2012). A comparison of 19 GCMs (General Circulation Models) revealed that the most significant discrepancies were in the upper troposphere where ice clouds are dominant (Jiang, 2012). Therefore, an accurate understanding of cirrus-related microphysical processes is vital for predicting climate with confidence (e.g., Schmitt et al., 2016).

Current state-of-the-art microphysical parameterizations are primarily tested and designed for the lower and mid troposphere rather than the cold point tropopause and lower stratosphere, which have very different and complex environmental conditions. The growth of ice particles by vapor deposition in a supersaturated environment is sensitive to the temperature regime, resulting in a distinct ice habit and fall speed (e.g., Phillips et al., 2015). The treatment of ice particles and their shapes in models is very challenging (Pruppacher and Klett, 2010) and became predominantly independent of measurements (Waliser et al., 2009; Schmitt et al., 2016). Moreover, the microphysical processes of pure ice clouds at these high altitudes are poorly understood due to the difficulty of measuring the key parameters with the required accuracy (Krämer et al., 2016). Most widely used microphysics schemes (e.g., Lin et al., 1983; Thompson et al., 2008) rely on traditional mass–dimensional (m–D) and terminal velocity–dimensional (V-D) relationships which have inherent shortcomings and lead to an inaccurate evolution of the density and shape of ice particles (Sulia et al., 2013; Jensen and Harrington, 2015). Other recent attempts (e.g., Morrison and Milbrandt 2015) included aggregates but without prediction or evolution of ice density. Many recent laboratory (e.g., Järvinen et al., 2018; Harrington and Pokrifka, 2021; Pokrifka et al., 2020) and theoretical (e.g., Przybylo et al., 2019, 2022a, 2022b; Schmitt and Heymsfield, 2010, 2014) studies have examined processes of ice particle growth and provided adequate details on ice crystal complexity, ice habits and aggregates that are overlooked in most sophisticated microphysical parameterizations. According to ice habit diagrams (e.g., Bailey and Hallett, 2009), the shape of ice can change significantly in response to a temperature change of a few degrees Celsius.

In this study, we explore the ability of commonly used microphysics schemes to reproduce measured peaks of ice water content, ice number concentration, and large ice aggregates in the TTL. For ice crystals, we also explore how their fall speed depends on their habit, how relative humidity over ice impacts on the shape ice crystals attain. Thus, we aim to identify key parameters of these schemes and model configuration that can lead to an adequate representation of TTL cirrus. For model setup, we stretch our computational resources to its limits and configure a real case study at very high spatial (LES-scale) and temporal resolutions. We compare a set of simulations using different schemes to aircraft measurements from the StratoClim campaign 2017.

125 The remainder of this paper is organized as follows. Section 2 provides an overview of the StratoClim field campaign and aircraft measurements used in this study. Section 3 presents the numerical experiments, including details on the microphysical parameterizations involved. Section 4 discusses the results of the model and the comparison with the measurements. Section 5 summarizes the results and presents recommendations for modelers to adequately simulate TTL cirrus.

2 Data and methods

2.1 Field campaign and observational data sets

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The in-situ measurements used in this study are aircraft data from the StratoClim campaign, which took place in the Indian subcontinent from July 27 to August 10, 2017. The M55-Geophysica high-flying research aircraft deployed in Kathmandu, Nepal, was equipped with a comprehensive set of instruments to capture high quality datasets which include but are not limited to, water vapor, number densities, size distributions of ice particles, as well as meteorological parameters in the subtropical

135 upper troposphere/lower stratosphere (UTLS) region (Khaykin et al., 2022).

In this study, we use in-situ measurements of temperature and pressure by TDC (ThermoDynamic Complex) (Shur et al., 2007) as well as measurements of the cloud particle number size distribution in the size range of 3–937 μm diameter at a time resolution of 1 Hz (Krämer at al., 2016; Luebke et al., 2016) by NIXE-CAPS (New Ice eXpEriment: Cloud and Aerosol Particle Spectrometer). The ice water content (IWC) is derived from the particle size distribution with a lower detection limit

140 of ~ 0.05 ppmv (Krämer at al., 2020; Afchine et al., 2018). Gas-phase water was measured with the sampling rate of 1 Hz by means of the Lyman-Alpha hygrometer FLASH (Khaykin et al., 2022).

2.2 Case study description

While the aircraft completed eight successful flights during the 2017 Asian monsoon season (Krämer et al., 2020, main article and supplementary material), this study focuses on one specific flight (Flight 8, hereafter F8) on August 10 (Fig.1), which was

145 characterized by a stable anticyclone and intercepted strong overshooting deep convection, high ice water contents and very large aggregates of ice particles that persisted for an unusually long duration. Moreover, this case study is characterized by overshooting of ice crystals into a supersaturated environment above the tropopause (Khaykin et al., 2022), resulting in their rapid growth via uptake of water vapor, sedimentation, and hence irreversible dehydration (Schoeberl et al., 2018).

Figure 1 shows the trajectory of F8 along the southern slopes of the Himalayas and the corresponding time series of water

150 vapor, ice water content (IWC), and altitude. High ice water content (HIWC) peaks up to 2400 ppmv are indicated by yellow and red arrows at 9:40 UTC and 9:50 UTC, respectively. In addition to HIWC, large aggregates of ice particles exceeding 700 μm in size and unusually long residence time are measured during F8. With the use of particle images NIXE-CIPg (Cloud Imaging Probe – grayscale) which is part of NIXE-CAPS, ice particles with diameters between 3 and 937 μm are sized and counted (Krämer et al., 2016; Costa et al., 2017). Evidence of these large aggregates is shown in Figure 2 with 6 snapshots of

ice particles from CIPG probe during F8. These crystals have a highly irregular structure due to the inherent randomness in their growth via the process of aggregation (e.g., Schmitt and Heymsfield, 2010).

3 Numerical experiments

3.1 Model setup

All three simulations of the same real case study associated with F8 were performed using Advanced Research version of the Weather Research and Forecasting Model (WRF-ARW) version 4.2 which is a nonhydrostatic, compressible model that uses a hybrid terrain-following coordinate system (Skamarock et al. 2008; Powers et al., 2017). The simulations are configured with four one-way nested domains with 9-, 3-, 1- and 0.3 km horizontal grid spacings, encompassing the trajectory of F8 over India (Fig. 3). As our focus is on the TTL, the vertical level configuration has been optimized to achieve 100-m grid spacing for the TTL and high spatial resolutions near the surface and the boundary layer to appropriately simulate fluxes that trigger and sustain deep convection. The 173 vertical levels begin with a 50-m grid spacing near the surface which gradually increases upward until the level of 6 km to reach a 250-m grid spacing which remains constant upward until the 10-km level. Then, the grid spacing decreases from 250 m to reach 100 m in the layer between the levels of 13 and 19 km. Beyond that, it increases again to gradually reach 900 m at model top of 30 km. A sponge layer is employed in the upper 5 km of all domains to damp out any reflection of waves and prevent their buildup below the model top. We apply spectral nudging for the outer domain D01 above the boundary layer to reduce the distortion of large-scale circulation (e.g., Miguez-Macho et al., 2004; Lamraoui et al., 2018). Due to the limitation of storage and computational resources, each of the simulations is a 12-h integration from 0000 UTC to 1200 UTC 10 August 2017 with a 5-min output frequency. For the analysis, we consider only the last three hours which correspond to the time when the aircraft captured peaks of IWC and large ice aggregates. The simulations were driven using the fifth-generation European Centre for Medium-Range Weather Forecasts reanalysis (ERA5) fields (Hersbach et al. 2020).

Ideally, the WRF-ARW simulations could be compared not only to the Stratoclim aircraft measurements but also to concurrent satellite observations of cloud properties, such as those offered by the CloudSat (Stephens et al., 2002) and CALIPSO (Winker et al., 2009) satellites. While CloudSat and CALIPSO did pass over the study region on the same day as Stratoclim F8 (Figure 3), they did so nine hours after F8. Because this time gap exceeds the convective and microphysical timescales relevant to our analysis, a comparison to CloudSat or CALIPSO retrievals would not be appropriate. While geostationary satellites offer measurements that are concurrent with F8 in both space and time, they cannot provide vertically resolved retrievals of cloud microphysical properties and would therefore be of little utility here. We thus focus solely on comparing the simulations to F8 and on the sensitivity of the simulations to the choice of microphysics scheme.

Except for the microphysics scheme option, the simulations are carried out using the same model configuration, which includes the Rapid Radiative Transfer Model for General circulation models (RRTMG; Iacono et al., 2008) for longwave and shortwave

190   radiation, the Eta Similarity Scheme (Janjic, 2002) for the surface layer, the Unified Noah Land Surface Model (Tewari et al., 2004) for land surface physics and the Shin-Hong Scale–aware Scheme (Shin and Hong, 2015) for the planetary boundary layer (PBL). The New Tiedtke Scheme (Zhang and Wang, 2017) for the convection parameterization is enabled only with the outer domain.

3.2 Model microphysics

195   We investigate the sensitivity of simulated TTL cirrus clouds to microphysical parameterizations with predefined and non-predefined frozen hydrometeor categories, as well as to ice morphology. For this reason, we use the following three bulk microphysics schemes (BMS): the aerosol-aware Thompson scheme (Thompson and Eidhammer, 2014), the predicted particle properties (P3) scheme (Morrison et al., 2015; Milbrandt and Morrison, 2016) and the Jensen Ice-Spheroids Habit Model with Aspect-Ratio Evolution (Ishmael) Scheme (Jensen et al., 2017). Size distributions of hydrometeors in BMS are assumed to be

200   a generalized gamma function. The Thompson and P3 schemes use simple tunable power laws to relate mass and fall speed to the size of each hydrometeor. Schemes with a free-category approach of frozen hydrometeors such as P3 and Jensen are advantageous over classical multicategory schemes (e.g., Chen and Tsai, 2016; Thompson and Eidhammer, 2014; Morrison et al., 2009; Milbrandt and Yau, 2005) that require arbitrary and artificial conversion between categories (e.g., Milbrandt et al., 2021) which relies on ad-hoc tunable parameters that cannot be directly measured. The Thompson scheme is quasi-two-

205   moment scheme that predicts the mass mixing ratios of cloud water, rain, cloud ice, snow, and graupel, as well as the number concentrations of cloud ice and rain. Artificial transitions between predefined ice categories and the corresponding ad-hoc conversion parameters can systematically bias the model output (Morrison and Grabowski, 2008; Lin and Colle, 2009). In addition to mass and number concentration, P3 predicts several other bulk physical ice properties such as rime mass, rime density and mean dimension. In P3, the changes of riming state and rime density are transitional but the growth of ice particle

210   by vapor deposition remains constrained by mass–dimensional (m–D) and fall speed–dimensional (vt–D) relationships that are chosen a priori as in traditional schemes. With the use of m–D and vt–D relationships, it is assumed that all modeled ice particles evolve and sediment along the same trajectory in mass–size and fall speed-size spaces. These relationships do not capture the proper transitions in ice particle mass and fall speed that occur during riming (Jensen and Harrington, 2015). The Ishmael scheme predicts the evolution of ice properties, including aspect ratio, mass mixing ratio, number concentration, size,

215   and ice density. Ice particles are modeled as spheroids and the nucleation temperature controls the number of nucleated species. The evolution of ice shape during vapor growth is verified against wind tunnel data (Jensen and Harrington 2015). Unlike the other two schemes, this scheme simulates measured habit-dependent riming (e.g., Fukuta and Takahashi, 1999) and its effect on fall speed. The evolution of ice shape from unrimed to lightly rimed to densely rimed is modeled without converting ice mass between predefined ice categories. The shape of the ice particles is controlled by the riming rate, resulting in a reduced

220   vapor growth rate and an increased riming rate compared with traditional BMSs (Jensen and Harrington, 2015). Light riming can expand the minor axis of ice crystals or fill in gaps in their shapes, leading to increased density and fall speed without affecting maximum dimension. The simulated ice habit is determined based on the growth rate which mainly depends on the

225 capacitance and ventilation coefficient (e.g., Liu et., 2003a). This realistic process conflicts with the widely used m–D and vt–D relationships.

4 Results and discussion

4.1 Convective activity and HIWC at the TTL

230 Observations have indicated strong convective activity and aircraft-captured peaks of HIWC detrained at the TTL level (Khaykin et al., 2022). We therefore opt for two sanity checks. First, we want to ensure that the model can reproduce the observed convective activity. For this reason, the simulated outgoing longwave radiation (OLR) corresponding to the time of the first HIWC peak measured at 0940 UTC is shown in Figure 5. Multiple cells of very low OLR are simulated reflecting strong convective activity. Second, the aircraft was at an approximately constant level of 17 km when the HIWC values were
235 detected. The corresponding IWC profile is used to provide evidence of HIWC plumes in the TTL. Therefore, we use IWC from ERA5 reanalysis within the inner domain encompassing the trajectory of F8. The 99th percentile is used to present HIWC which refers the top 1% of simulated or observed IWC values. The 99th percentile of IWC and domain-mean temperature profiles are shown in Figure 6. Although the ERA5 HIWC is about an order of magnitude lower than the measurements (Fig.6), it identifies the presence of a HIWC plume at ~125 hPa in the TTL. The profile of domain-mean temperature shown in Figure
240 6 indicates a CPT of 192 K at the pressure level of 100 hPa which is about 5 K warmer than the measured value.

4.2 Sensitivity of simulated microphysical properties to model configuration

In this section, we aim to test the ability of the model to reproduce the microphysical properties observed during F8, including HIWC peaks, ice number concentration and large ice particle aggregates. To examine the sensitivity of simulated HIWC to changes in model grid spacing and microphysical parameterizations, we performed a set of simulations configured with four
245 nested domains and three microphysics schemes. All simulations share the same high vertical spatial resolution of 100 m in the TTL (Fig.4). HIWC in the upper troposphere and TTL, shown in Figures 7, 8 and 9, indicates sensitivity to grid spacing that varies from one scheme to another. In response to overshooting deep convection, the simulated HIWC profiles shown in these figures reveal a gradual increase over 9 hours.

Despite the high spatial resolution in the TTL, the Thompson scheme (Fig.7) simulates a detrained HIWC in the TTL at ~ 9:00
250 UTC but with significant underestimation of more than an order of magnitude relative to the HIWC captured by aircraft at 9:40 UTC (Fig.1) and even that of ERA5 (Fig.6).

For the P3 scheme (Fig.8), the IWC profiles show no significant change between the 9- and 3-km grid spacings. On the other hand, with 1- and 0.3-km grid spacings, the difference becomes significant, especially the accelerated increase in the TTL which is higher in the inner domain. This scheme simulates very high IWC peaks between 8:00 and 10:00 UTC that exceed
255 3000 ppmv, which is overestimated compared to the measured IWC peak of 2400 ppmv at 9:40 UTC.

The Ishmael scheme (Fig.9) shows improvement of simulated IWC only with 1-km grid spacing like P3 but with less accelerated increase in the TTL. What is unique about the Ishmael scheme in this simulation among the other two schemes is that it can approximately reproduce the IWC peak that is observed during F8 at the same time ~09:40 UTC. No overestimation of HIWC is simulated before ~09:40 UTC or after ~10:00 UTC.

The temporal evolution of the storm is well captured by the model and the simulation reproduces strong deep convection cells (Fig.5) in the domain D4 comparable to those observed during the field campaign at the same time. Therefore, for the comparison with F8, we use the simulation data of the entire inner domain D4 which encompasses most of the F8 trajectory. Following Krämer et al. (2016, 2020), we examine the measured and simulated microphysical properties of cirrus clouds in the TTL using the ice water content–temperature (IWC–T) parameter space (Figs 8, 9) and ice particle sizes larger than 3 μm. The IWC-T figures enable distinguishing the in-situ origin from the liquid origin cirrus clouds. These two types of cirrus clouds are roughly separated by the median IWC (green line, Schiller et al., 2008) in Figures 8 and 9. Above the green line, the ice particles are of liquid origin and strongly ventilated upwards by deep convection. Below this line, ice particles are associated with weak updrafts and represent mostly in-situ cirrus clouds that likely originate from heterogeneous nucleation, and less likely form homogeneous nucleation (Krämer et al., 2016; Cziczo et al., 2013; Jensen et al., 2013). Comparison between the simulation and the observation in Figures 8 and 9 reveals that most of measured ice particles are of liquid origin with only a small amount in the in-situ range. In-situ simulated cirrus clouds in all schemes (Figs. 8,9) are associated with low values of IWC and number concentration, which makes them less influential in regulating HIWC. These two figures support the results shown in Figures 7, 8 and 9 that Ishmael, P3 and Thompson schemes yield reasonable, overestimated, and underestimated values of HIWC, respectively.

For all schemes, higher ice number concentrations are associated with ice particles of liquid origin and strong updrafts. Higher values of IWC are mostly in the lower temperature range. In the temperature range between 187 and 192 K shown in Figures 8 and 9, the ice number concentration is correlated with IWC, as discussed by Kramer et al. (2020, their Figure 6). The simulated ice number concentration has a reasonable upper limit of about 10 cm^{-1} which is comparable to observations (Fig. 10). The measurements indicate lower temperatures reaching 187 K, which is not the case for the simulations, in particular with the Ishmael and P3 schemes. Only The Thompson and P3 schemes indicate events at T<188 K which can be explained by the simulated CPT levels and RHi. The Thompson and P3 schemes overestimate the CPT level (Fig.12), allowing more simulated ice particles to reach higher and colder levels (T<188). Both schemes have lower supersaturation with respect to ice compared to the Ishmael scheme. This lower supersaturation slows the growth of ice particles and prevents them from reaching the sedimentation threshold, thus allowing ice particles to reach higher and colder levels. The Ishmael scheme simulates a lower CPT level and higher supersaturation with respect to ice. This leads to accelerated growth rates, more sedimentation and cold trap at lower levels. The aircraft cruising level corresponds approximately to the observed CPT level.

The results with the Ishmael scheme shown in Figures 9 reveal that the ice particles of liquid origin and which contribute substantially to HIWC are of planar shape. Columnar and aggregate habits are mostly in the in-situ range with very low IWC and number concentration values.

Also, planar ice particles have the highest number concentration among the three habits, especially at lower temperatures (Fig. 9a). For all habits involved in this study, the number concentration increases with decreasing temperature (Figs. 9 and 10), which corroborates previous studies (e.g., Gryspeerdt et al., 2018; Sourdeval et al., 2018).

In addition, the distribution of ice number concentration shown in Figure 11 exhibits a large spread across ice habits that can reach four orders of magnitude. The planar habit takes the lead in terms of high ice number concentration with a 99th percentile of 8 cm^{-3}, followed by the columnar and then aggregates with the lowest number concentration. As expected, aggregates are much less frequent than other habits (e.g., Lawson et al., 2006).

In the comparison illustrated in Figure 12, we examine the relative humidity with respect to ice (RHi) which is a requisite environmental condition necessary for the nucleation and growth of ice crystals (Pruppacher and Klett, 2010; Heymsfield and Miloshevich, 2013). In the TTL, the supersaturation with respect to ice can be reached through different processes such as moisture lifting by strong updraft (Evans and Marchand, 2013) or temperature fluctuations due to gravity waves (Haag et al., 2003). In the inner domain, RHi saturation is mainly caused by moisture uplifting induced by deep convection. This comparison demonstrates that the CPT from the P3 and Thompson schemes are at potential temperature levels of 380 K and above. On the other hand, the Ishmael scheme produces a CPT at the level of 375 K which is closer to the observed levels of ~371 K and 379 K at 09:40 UTC and 09:50 UTC respectively.

Figure 12 shows that around the CPT, the Ishmael scheme produces larger differences between the RHi inside and outside the clouds than the other two schemes. Furthermore, Ishmael indicates very high values of RHi around CPT, which favors in-situ and homogeneous nucleation. Unfortunately, at and around the CPT level, the observation does not provide enough data outside clouds for an adequate comparison with the simulations. According to previous laboratory studies (e.g., Bailey and Hallett, 2009), RHi can reach high values of ~ 160 % at very low temperatures. Also, high RHi values of ~200 % at the tropopause level were reported in previous field campaign studies (Krämer et al., 2020; Dekoutsidis et al. 2022) but with very low frequency of occurrence of ~2%. Simulations and especially Ishmael scheme overestimate RHi. The reason that may explain the existence of such high supersaturations is the low ice number concentration, which is mainly present at low temperature (Fig. 10). Previous studies (eg., Jensen et al. 2013) have also indicated that increasing the number of ice crystals leads to decrease in RHi until reaching the saturation. The Thompson and P3 schemes produce smaller ice particles compared to Ishmael (Fig 13) and high ice number concentration (Fig 8). For this reason, the Ishmael scheme produces a higher RHi (Fig. 12). With high RHi and lower temperatures more in-situ nucleation, faster growth rate of crystals and larger ice particles are expected. The comparison between the three habits shown in Figure 11 indicate that large aggregates are associated mainly with low number concentration which demonstrate the correlation between aggregates and high supersaturation (Jensen et al. 2013). RHi values of 120% and 140% can be considered approximate thresholds for heterogeneous and homogeneous ice

350 nucleation respectively (eg., Koop et al., 2000; Kärcher, 2012). Therefore, due to the high supersaturation with respect to ice, more homogeneous and in-situ nucleation is associated with Ishmael scheme than the other two schemes.

4.3 Ice particles size distribution

In terms of ice particle sizes, Figure 13 shows that all simulated effective radius distributions are within smaller size ranges of
355 ice particles compared to the observations, in which the 99th percentile is 131 μm. The effective radii in every grid box are considered, respectively, and then the histogram is constructed from these values. The 99th percentiles of effective radius simulated using the Ishmael, P3 and Thompson schemes are 70 μm, 57 μm and 25 μm, respectively. The Ishmael scheme produces a bimodal distribution with two peaks at about 5 μm and 25 μm which have the highest frequencies. For the smaller peak, this can be explained by freshly nucleated ice particles due to high RHi (Fig. 12) that are of in-situ origin. On the other
360 hand, for the larger peak it can be due to the rapid growth of existing particles that are of liquid origin before sedimentation (Fig. 9). The distributions of the Thompson and P3 schemes are shifted toward smaller sizes, especially for Thompson. It can be explained by the slightly warmer CPT (Fig. 12) compared to the other two schemes which can limit ice crystal growth (Fig. 13) and LWC (Figs. 7 and 8). The distribution of P3 dominates the range of small sizes. In combination with a higher CPT level and lower RHi (Fig. 12), ice particles are more likely to grow slowly, resulting in lower fall speeds, longer residence time
365 and higher IWC. This explains the overestimation of the simulated IWC using the P3 scheme shown in Figures 8.
In this section, we focus on the simulation with the Ishmael scheme to examine the sensitivity of TTL cirrus clouds to different ice habits. We use the maximum dimension D_{max} (Jensen et al., 2017), which is produced only by the Ishmael scheme among all the schemes involved in this study. The observed ice particles shown in Figure 2 illustrate very large sizes captured during F8. For a fair comparison with the images captured during the field campaign, Figure 14 shows the distribution of D_{max} from
370 measurements and the Ishmael simulation. It demonstrates the bimodality (e.g., Ivanova et al., 2001; Mitchell et al., 2008; Zhao et al., 2011) of the observed and simulated cirrus particle size distributions (PSD), with an approximate 300-μm shift of the simulated size range towards larger values. The small particle mode is in the sub-100-μm and the sub-400-μm size ranges for the observations and simulation, respectively. The bimodality of the PSD over a very wide range of particle sizes indicates a difference in the physics controlling its evolution (Field and Heymsfield, 2003; Sourdeval et al., 2018). Bimodal cirrus PSDs
375 can be explained by aggregation processes and depositional growth. Depending on RHi, bimodality evolves naturally due to condensational growth and sedimentation (e.g., Mitchell et al., 1996). It can be difficult to distinguish aggregation from complex diffusional growth patterns (e.g., Lawson et al., 2006), using only in-situ observations of microphysical properties. The lower range of particle size mainly represents nucleation and diffusional growth, while the higher range is dominated by aggregation (Field, 2000). The simulated PSD shown in Figure 14 provides additional details that distinguish the different ice
380 habits. The first simulated mode of the PSD is dominated by the planar habit followed by the columnar one with a narrower range for sizes <300 μm. The second mode, however, is only associated with the aggregates that dominate the large end of the PSD (Kajikawa and Heymsfield, 1989; Heymsfield et al., 2002). The large end of the observed PSD > 400 μm shares the same

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range of simulated aggregates. Observed ice particles with sizes larger than 700 μm only account for 6.4 %, which corroborates the low ice number concentration of the simulated aggregates shown in Figures 9, 10 and 11. The maximum size of the ice particles captured by F8 is approximately neighboring the simulated median and the most frequent size range of the simulated aggregates. More specifically, in the anvil region, the median value of aggregate sizes is 828 μm, which corresponds approximately to the maximum value observed during F8. Moreover, the simulated PSDs for all ice habits shown in Figure 14 demonstrate a difference between ice particle sizes in the convective and anvil regions. To distinguish between the three life cycle stages of deep convection, three ranges of ice water path are used: deep convection (IWP>1000 g m^{-2}), thick anvil (1000>IWP>10 g m^{-2}), and thin cirrus (10>IWP>0.1 g m^{-2}) (Sokol and Hartmann, 2020). In reference to PSDs in the convective region, the larger end of the PSD in the anvil is smaller for the columnar and aggregate habits. For the planar habit, Figure 14 shows a broadening of the distribution in the anvil. Despite the overestimation of the simulated particle sizes, the Ishmael scheme is the only scheme used in this study that can reproduce the large particle sizes captured by F8 and identify their shape. While the range of particle sizes for each habit is similar in anvil and deep convective clouds, the largest particles are more frequently found within deep convection.

During F8, very large aggregates of ice particles were observed. For this reason, we want to check whether this tendency is also reproduced by the simulation. Figure 15 shows the time series of the change in the median value of the maximum dimension for the three ice habits, in reference to the initial value of August 10, 2017, at 9:00 UTC. The temporal evolution of the aggregates size indicates no decrease in the median value during the whole three hours (Fig.15), which backs up the observed persistence of the large aggregates. The size of the columnar habit is similarly constant, fluctuating by less than 2%. On the other hand, the size of planar ice particles experiences an approximately constant decrease of 4% per hour. From the PSDs of each ice habit shown in Figure 14, the planar shape occupies a wide range that envelops that of the columnar habit. The large planar ice particles either fall off if they become sufficiently heavy or contribute to the formation of aggregates. The lower end of the PSD for the planar habit represents mostly small, freshly nucleated ice particles. Moreover, due to the difference in RHi (Fig.12), the PSD indicates more in-situ nucleation of planar ice particles in the anvil than in the convective region (Fig.14). Therefore, the steady decrease in particle size for the planar habit shown in Figure 15 is due to the loss of large ice particles in favor of sedimentation or aggregate formation and the increasing nucleation of small ice particles. For the columnar habit, ice particles contribute to aggregate formation and experience growth or sublimation depending on the RHIi. For columnar habit, the low in-situ nucleation rate prevents the ice crystals from reaching the sedimentation threshold and from undergoing a strong decrease in the number of large ice crystals encountered by the planar habit (Fig. 15). Due to the coexistence of columnar and planar habits, ice aggregates form continuously. The aggregate habit has a hollow structure with a very low density of ice particles more than an order of magnitude smaller than the density of the planar and columnar habits. For this reason, ice aggregates can remain longer in the TTL even with large sizes until reaching the critical sedimentation size.

4.4 Fall speed and growth of ice particles across ice habits

445 The fall speed of ice particles is one of the factors that determines the lifetimes of cirrus clouds (e.g., Jiang et al., 2019) and thus indirectly affects climate sensitivity (e.g., Sanderson et al., 2008). The corresponding RHi which can vary significantly (Fig. 12) in the TTL regulates the rates of nucleation, sublimation, and growth of ice particles. Due to the difference in shapes and densities, ice particles can evolve differently in the outflow of deep convection, anvil, and in-situ thin cirrus (e.g., Lawson et al., 2019). To further explore the behaviour of each ice habit in these three regions in the TTL, Figure 16 shows scatter plots

450 of fall speed as function of RHi, color coded by D_{max}. The StratoClim field campaign and more specifically the F8 trajectory aimed to sample deep convective cases. The results of the inner domain D4 (Fig. 3) using the Ishmael scheme shown in Figure 16 are consistent with the observation that indicates the dominance of deep convective and thick anvil clouds, with a very low occurrence of thin cirrus (Fig 16). The bottom row of Figure 16 shows the small amount of thin cirrus consisting exclusively of planar ice habit with small sizes mostly less than 100 μm and low fall speeds below 20 cm s^{-1}. For higher values of RHi

455 ~130% and above, these planar ice particles are largely formed in situ, and are mostly 50 μm or smaller with fall speeds of ~10 cm s^{-1}. These small ice particles explain the extended life cycle of thin cirrus clouds.

For deep convection and thick anvil regions, Fig 16 (upper and middle rows) shows the coexistence of the three ice habits but with different characteristics. The dependence of fall speed on particle sizes is not similar from one ice habit to another, which is mainly due to the difference in particle shape and density. In terms of fall speed, the deep convection region within the TTL

460 occupies the largest range starting from a few centimetres per second for planar ice to over 150 cm s^{-1} for columnar ice. For planar ice, when RHi<130%, Figure 16 shows the smallest ice particles (<30 μm) with very slow fall speeds (<5m s^{-1}). These particles are mainly of liquid origin and freshly nucleated. On the other hand, planar ice particles associated with higher supersaturation (RHi>150%) are more likely to grow and fall in the deep convection region (upper left). In the anvil (middle left), high supersaturation (RHi>150%), very low fall speed and small particle size indicate prevalent in-situ nucleation. Also,

465 for the planar habit, RHi is inversely correlated with fall speed. This correlation is less evident for the other two ice habits. For the columnar ice habit, Figure 16 (middle column) shows a shift in fall speeds to higher values, larger particle sizes and lower supersaturation with respect to ice. These characteristics indicate slower growth and high sedimentation rates when compared with planar habit. Larger particles are more likely to sediment with fall velocities > ~1 m/s (Heymsfield and Wright, 2014) from the outflow depending on ice density, as soon as they are no longer supported by an updraft. As ice particles sediment,

470 they may contribute to aggregate formation, grow or sublimate depending on the RHi. High fall speeds and RHi > ~130% such as experienced with columnar habit (Fig 16), may correspond to regrowth downwind of aged ice particles via vapor diffusion (e.g., Gallagher et al., 2012). In contrast, the aggregates demonstrate different behavior than pristine columnar and planar crystals. Very large particle sizes between 600 and 800 μm are associated with very low fall speeds below 20 cm s^{-1}. This explains the longer residence time of aggregates observed during F8. For these ice particles to fall out with speed of about 30

475 cm s^{-1} or faster, they need to reach a size of ~900 μm or larger.

As ice crystals sediment, they may aggregate (Lawson et al., 2019) by differential sedimentation throughout the cloud layer (e.g., Schmitt and Heymsfield, 2010). Large ice crystals with higher fall speeds can collect smaller ice crystals as they fall through clouds and form aggregates. The ratio of the sizes of aggregating crystals and their terminal velocities are important in aggregation (e.g., Kajikawa and Heymsfield, 1989). To explore this relationship and quantify the correlation between maximum dimensions and corresponding fall speeds, Figure 17 shows density plots that relate these two parameters. The figure reveals the linear relationship across all ice habits and demonstrates comparable shallow slopes of about 3 μm/cm s^{-1} for columnar and planar habits, but with a larger spread for the latter. For columnar ice particles, Figure 19 (middle column) shows a sharp correlation. A small change in size can lead to a large change in fall speed, which is supported in Figure 17 (upper row, middle column) by higher fall speeds. The large spread associated with the planar habit is due to the larger aspect ratio range. The minor axis can evolve broadly and regulates ice crystal density and fall speed without affecting the maximum dimension. On the other hand, the aggregates indicate a sharp correlation and a steep slope of about 27 μm cm s^{-1}. This reveals a lower sensitivity of fall speed to change in maximum dimension which is due to very low ice densities for this ice habit.

To explore the relationship between differential sedimentation and maximum dimension (D_{max}) of aggregates, Figure 18 reveals an evident linear relationship between the D_{max} of aggregates and the difference in fall speeds between two different ice habits (columnar-planar) throughout the cloudy TTL. This relationship is consistent in the deep convection (Fig 18, right) and thick anvil (Fig 18, left) regions. Aggregates can also occur involving similar ice habits, but these density plots (Fig 18) reveal a predominance of aggregation formation through different habits in controlling the maximum dimension of aggregates. Also, Figure 18 provides additional details on the differential sedimentation associated with the case of the current study. In the formation of aggregates, which ice habit falls faster is important in determining the slope and the spread of the linear relationship. When the planar habit sediments faster than the columnar, the slope of the linear relationship is steep of about 28 μm/cm s^{-1} and the linear relationship is sharp. On the other hand, when the columnar habit sediments faster than the planar, the slope of the linear relationship is even steeper at ~87 μm/cm s^{-1}, and the maximum dimension of the corresponding aggregate is more sensitive to the difference in fall speeds. The spread of the linear relationship in this case is larger.

4.5 Sensitivity of diabatic heating in TTL cirrus to different ice habits

Although this study focuses on the sensitivity of simulated HIWC in the TTL to model configuration and different ice habits, we extend our analysis in this section to explore the sensitivity of diabatic heating to ice habits. We examine the sensitivity of radiative and latent heating rates to different ice habits, spanning hours 9–12 UTC. The distributions of the two heating rates are taken over the entire domain at each grid cell every 5 minutes which is the output frequency of the simulations. The distributions are divided according to the cloud categories that are based on ice water path thresholds. The heating rates in Figure 19 show distinct distributions between different cloud morphologies (thick anvil and thin cirrus), as well as between ice habits. Figure 19 (left column) shows the planar habit with an approximately balanced distribution with an average radiative heating rate of 1.5 K day^{-1}. Moreover, moving away from deep convection, the radiative heating rate associated with planar habit gradually shifts towards more positive radiative heating rate in thick anvils with an average of ~ 2.5 K day^{-1} and even

more in thin cirrus with average of ~ 3.5 K day^{-1}. The radiative heating rate associated with thin cirrus clouds is predominantly positive. If only planar habit is considered, this tendency is consistent with the findings of Hartmann et al. (2018) that rising and thinning anvils lead to a net radiative warming effect. Adding more realistic and complex ice habits in the microphysical parameterizations can alter the radiative heating of clouds. Radiative properties of cirrus depend on the optical properties of its ice crystals, which in turn depends on the microphysical properties such as particle shape, particle size distribution, ice water content (eg., Key et al., 2002) and number concentration (eg, Liu et al., 2003b). Previous studies (eg. *Takano and Liou* 1989) found that single-scattering albedo is related to aspect ratio and absorption coefficient of ice crystals. Also, Fu (1996) calculated the single-scattering properties of columnar and planar habits, using the geometric-optics-integration-equation technique (Yang and Liou, 1996b). Key et al (2002) quantified significant sensitivity of ice cloud radiative properties to ice crystal optical properties of seven ice particle shapes.

The current study reveals a significant change in these microphysical properties across ice habit (Fig 9, Fig 11, and Fig 16). Additionally, ice crystal capacitance is parameterized differently from one ice habit to another depending on the maximum diameter of the crystals. This affects the growth rate of ice crystals differently, which in turn affects microphysical processes to alter the size distribution of ice crystals (Fig. 14), further modifying radiative heating. For columns and aggregates, the distributions of radiative heating rate are slightly different, with averages of 4 and 4.6 K day^{-1} in deep convection and 4.6 and 4.3 K day^{-1} in thick anvil, respectively.

Comparison between the three ice habits indicates that the columnar habit takes the lead in net radiative warming, followed by the aggregates, then the planar. The contribution of the columnar habit versus the planar habit to the net radiative rate is more than double in the deep convection region and almost double in the thick anvil. Therefore, neglecting the shapes of ice particles can lead to an underestimation of the net cloud radiative rate by more than double.

Latent heating rates (Figure 19, right column) also indicate sensitivity to ice habit and cloud morphology. All distributions are skewed to positive values, except for the planar habit which indicates an approximate balanced distribution in the deep convection region. Across different stages of cloud morphology, the positive skewness of the latent heating distributions reflects a predominance of ice particle growth and nucleation. In the deep convection region and within the TTL, columnar habit growth and sublimation is dominant over the other two habits, with average latent heating rate of 2.3 K day^{-1}. The average values for the planar and aggregates habits are 1.2 and 2.1 K day^{-1}, respectively. The transition from deep convection to thick anvil indicates a significant narrowing of the latent heating rate distribution associated with columnar habit and decreasing its average value to 1.52 K day^{-1}. The decrease in the rate of sublimation and the growth of columnar habit is consistent with the higher sedimentation shown in Figure 16 (middle column). Also, Figure 19 (right column, middle row) shows dominant positive heating rates associated with aggregates which reflects continuous aggregates growth in thick anvil. The average latent heating rate associated with aggregates in the anvil is 1.54 K day^{-1}. For thin cirrus clouds with an average latent heat of 0.1 K day^{-1}, Figure 19 mainly shows lower positive latent heating rates with negligible negative values. Insignificant sublimation and low growth rate will allow ice particles to persist without reaching the sedimentation threshold, resulting in a prolonged residence time in the TTL. This supports the continued nucleation of small ice particles and the slower growth rate

demonstrated in Figures 15 and 16. Comparison of latent warming associated with the three ice habits indicates that aggregates
and columnar habits contribute more than twice as much as planar habit. For diabatic heating in thick anvils, columnar habit
dominates the net radiative warming and aggregates dominate net latent warming. The results shown in Figure 19 quantitatively
confirm the major impact that different ice particle habits can bring to diabatic heating and consequently to TTL cirrus life
cycle.

5 Summary and conclusions

Many previous studies have addressed the impact of ice habit in mixed-phase clouds (e.g., Avramov and Harrington, 2010;
Hashino et al., 2020; Harrington et al., 2013; Sulia et al., 2021), but few have investigated its impact on TTL cirrus clouds,
which occur in frequently different environments. The properties of cirrus clouds may differ depending on geographical
location (e.g., Heymsfield et al., 2017) and proximity to convection (e.g., Lawson et al., 2019). This study focuses on a specific
case study associated with deep convection during the Asian summer monsoon.

The main findings of this study reveal that:

- Measurements in the TTL during F8 of the StratoClim campaign captured peaks of high ice water content up to 2400 ppmv and ice particle aggregates exceeding 700 µm.
- The ERA5 reanalysis indicates the presence of HIWC in the TTL but with an order of magnitude less than what was measured.
- In this study, we used high-resolution (LES-scale) simulations of a real case study to examine the sensitivity of TTL cirrus clouds to ice habit and cloud regimes.
- Comparison of simulated microphysical properties to Stratoclim measurements shows that simulated HIWC in the TTL is significantly sensitive to horizontal grid spacing only with the Ishmael scheme, with HIWC increasing substantially as grid spacing decreases from 3 to 0.3 km.
- The microphysics scheme that predicts ice habit (Ishmael) is the only scheme that reproduces the observed HIWC peak and ice number concentrations.
- Particular attention is given to Ishmael scheme, the microphysical parameterization predicting the ice habit. The corresponding fall speed and diabatic heating rates for each ice particle shape are compared between convective, thick anvil, and thin cirrus clouds.
- Stratoclim measurements are dominated by ice particles of liquid origin with a small amount in the in-situ range.
- Simulated in-situ cirrus clouds are associated with low IWC and number concentration. Higher values of IWC are mainly of liquid origin and occur at lower temperatures.
- The simulated ice habits reveal that the ice particles of liquid origin, which substantially contribute to the HIWC, are of planar shape.
- The columnar and aggregates habits are mostly in the in-situ range with small IWC and number concentration.

- For all habits, ice number concentration increases with decreasing temperature, which corroborates previous studies (e.g., Gryspeerdt et al., 2018; Sourdeval et al., 2018). The disparity in ice number concentration across ice habits can be up to four orders of magnitude, and aggregates are much less frequent than other habits.

- CPT and RHi, which regulate cirrus clouds, vary across schemes, and the Ishmael scheme simulates them adequately.

- The Ishmael scheme can produce the observed bimodal distribution of ice particle sizes. The lower range of the distribution is mainly dominated by nucleation and diffusional growth, while the higher range is dominated by aggregation processes and depositional growth. The simulated bimodality is shifted towards larger values compared to the measurement.

- The microphysical evolution of TTL cirrus clouds is strongly influenced by the type of ice habit in the cloud. The simulated ice habits evolve differently. The frequency of small planar ice particles increases faster due to enhanced in-situ nucleation.

- High RHi values above 130% with fall speeds below 10 cm s^{-1} are only associated with the planar ice habit, leading to an extended life cycle of cirrus clouds.

- For the planar ice habit, RHi is found to be inversely correlated with fall speed. This correlation is less evident for the other two ice habits.

- Columnar ice particles sediment rapidly without significant regrowth of aged ice particles, especially in the lower range of supersaturation with respect to ice. As these ice particles sediment, they may contribute to aggregate formation.

- Simulated aggregates exhibit different characteristics compared to pristine crystals. The large aggregates with sizes between 600 and 800 μm have fall speeds of less than 20 cm s^{-1}, which leads to a long residence time of the aggregates.

- We demonstrated and quantified the relationship between differential sedimentation of pristine ice crystals and the maximum dimension of aggregates. We found that the slope of this relationship depends on which pristine ice habit sediments faster. Therefore, ice habit could have a significant impact on the cirrus life cycle and the (de)hydration of the TTL by regulating the sedimentation rate and the freeze-drying process.

- Ice habit also influences in-cloud radiative and latent heating rates, which are known to influence cloud dynamics at a variety of scales (Gasparini et al, 2022; Ramaswamy and Ramanathan, 1989; Sherwood et al., 1994) scales. These results demonstrate the importance of ice habit in shaping the properties and evolution of tropical cirrus clouds.

6 Code availability

WRF is available from the University Corporation for Atmospheric Research following the instructions at https://www2.mmm.ucar.edu/wrf/users/download/get_sources_new.php (University Corporation for Atmospheric Research, 2022).

7 Data availability

All the data used in this study are openly available at Harvard Dataverse repository at https://doi.org/10.7910/DVN/PKDDEI.

8 Author contributions

FL, MK and ZK conceptualized the study. FL ran the model and postprocessed the data. MK, AA and SK provided measurements from the campaign, details about instruments onboard of aircraft, and contributed to the interpretation. FL performed the analysis and writing of the original draft manuscript. FL, MK, AA, ABS, SK, AP and ZK edited the manuscript and provided feedback on the interpretation of the results.

9 Competing interests

Martina Krämer is a member of the editorial board of *Atmospheric Chemistry and Physics*. The remaining authors declare that they have no conflicts of interest.

10 Special issue statement

This article is part of the special issue "StratoClim stratospheric and upper tropospheric processes for better climate predictions (ACP/AMT inter-journal SI)".

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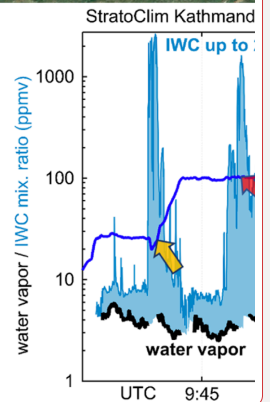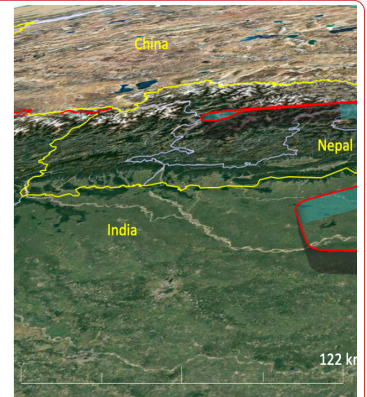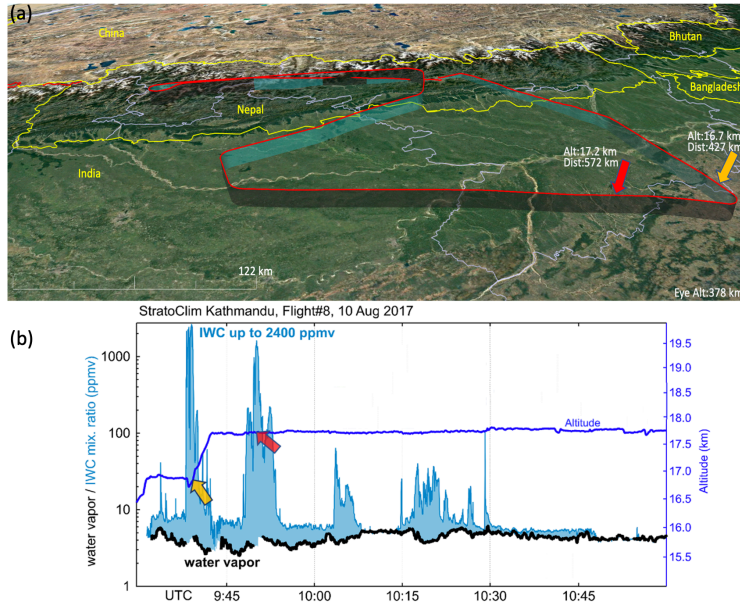
Figure 1: (a) Trajectory of Flight 8 and (b) the corresponding time series of water vapor (black), ice water content (light blue), altitude (dark blue). Yellow and red arrows in both subfigures indicate high ice water content peak's location (upper) and timing (lower). The corresponding altitude and distance from origin are indicted near each arrow (upper). The map was made using ©Google Earth 2022.

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08102017   Probe: CIPG   Resolution: 15 microns   Buffer width: 960 microns

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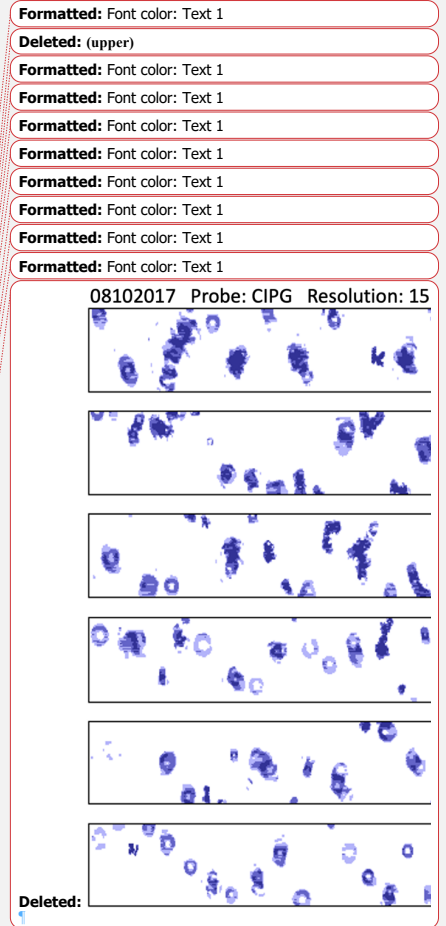
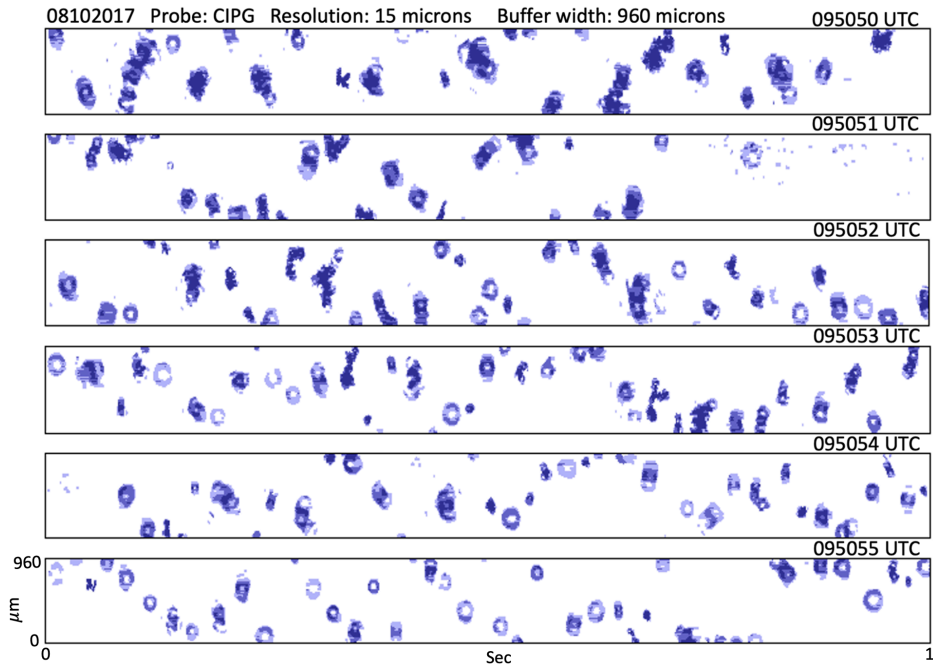


Figure 2: Snapshots of ice particles images from CIPG probe during Flight 8 (F8). Each slide has a height of 960 microns and length of 1 sec.

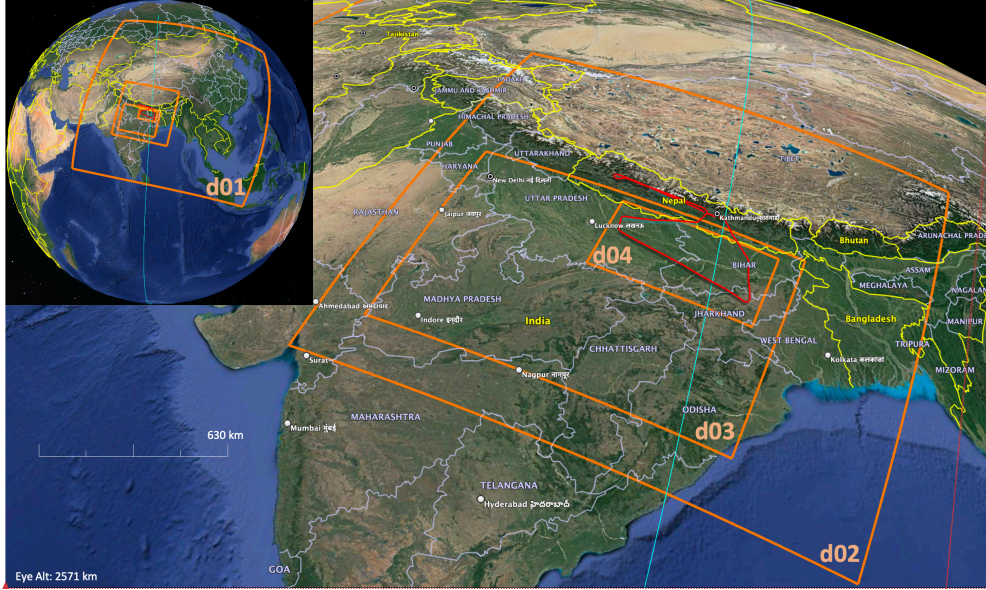



Figure 3: WRF 4 nested domains (orange), Flight 8 trajectory (red) and closest CALIPSO passage (cyan). The map was made using ©Google Earth 2022.

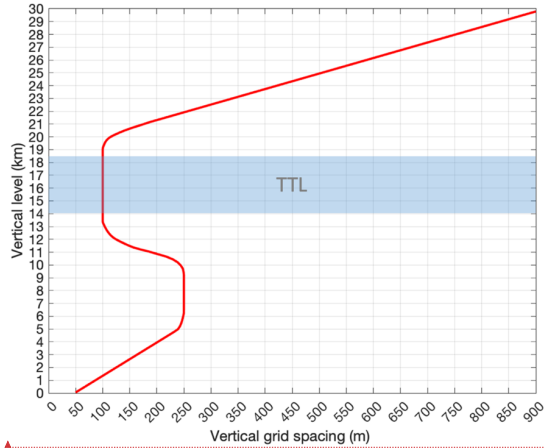



Figure 4: Vertical resolution as a function of model vertical level. The blue shaded area represents the TTL.

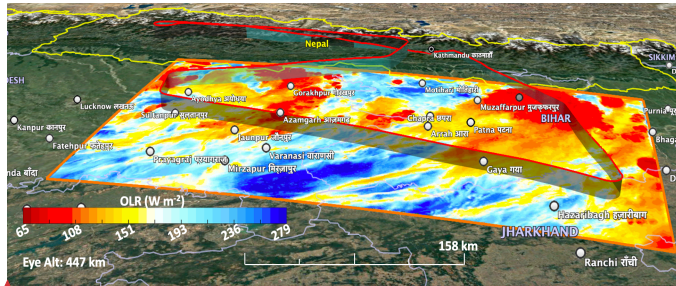



Figure 5: High-resolution (0.3 km) simulated OLR on 1000 UTC 10 Aug 2017. The map was made using ©Google Earth 2022.

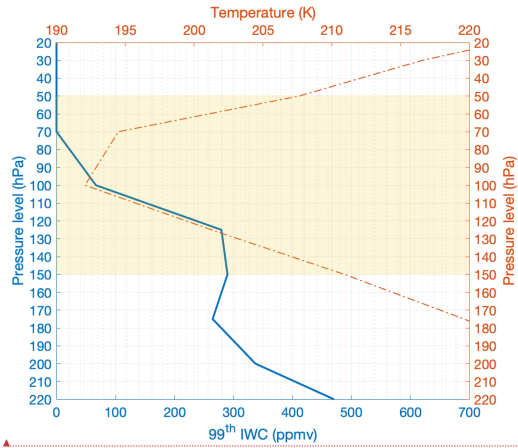



Figure 6: Profiles below and within the TTL (yellow shade) in the inner-domain D4 from ERA5 99th of ice water content (blue solid line) and domain-mean temperature (red dash-dotted line), on 1000 UTC 10 Aug 2017.

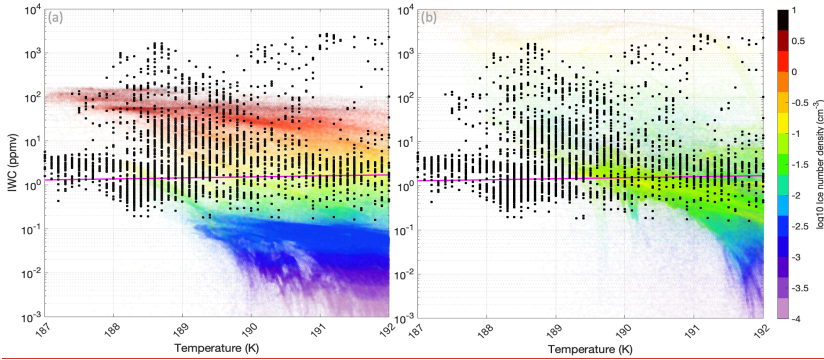
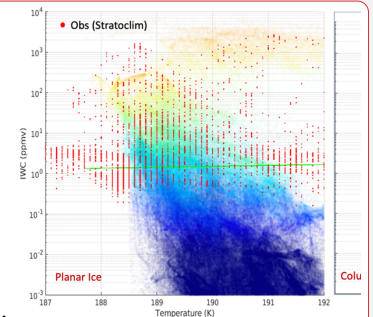
Figure 7: Profile of 99th percentile of simulated IWC below and within the TTL (cyan shared area) using grid spacing of 0.3 km and three microphysics schemes (a) Ishmael (b) P3 and (c) Thompson (TOM), color coded by time (UTC) from 2 to 11 AM. Black dash-dotted line represents the simulated profile at 09:40 UTC (timing of the first captured peak of IWC by aircraft). The red double arrow represents the observed value at the corresponding observation level.

070 **Figure 8: Simulated (shaded) and observed (black dots) IWC versus Temperature color-coded by log$_{10}$ of ice number concentration at 17 km and for D_ice>3µm using (a) Thompson scheme and (b) P3 scheme.**

Figure 9: Simulated (shaded) and observed (black dots) IWC versus Temperature color-coded by log$_{10}$ of ice number concentration
075 **at 17 km and D_ice>3µm using the Ishmael scheme from three ice habits (a) Planar, (b) columnar and (c) aggregates.**

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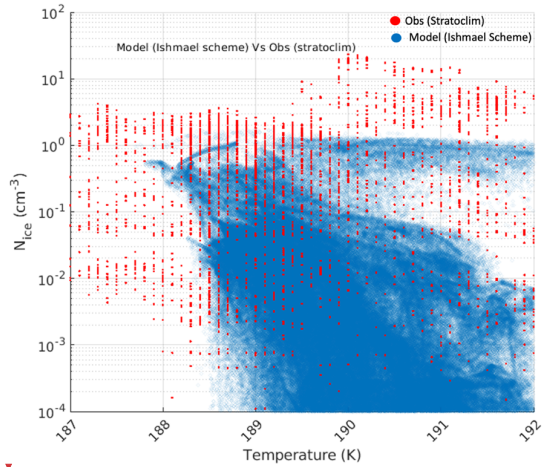
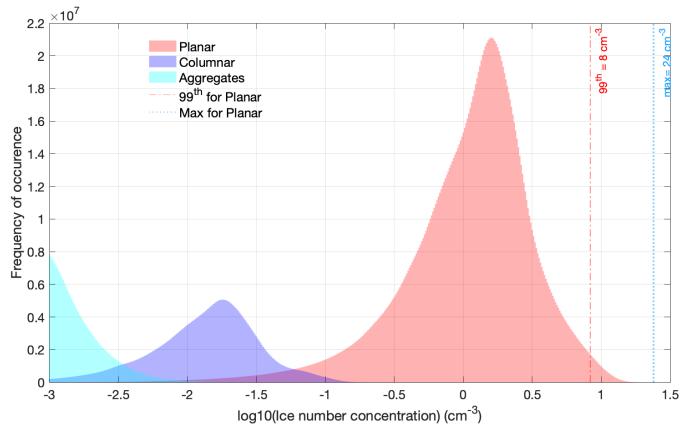


Figure 10: Simulated (blue) and observed (red dots) N_{ice} versus Temperature at 17 km using the Ishmael scheme.

 Figure 11: Distribution of ice number concentration on a log10 scale using the Ishmael scheme for different ice habits planar (red), columnar (blue) and aggregates (gray).

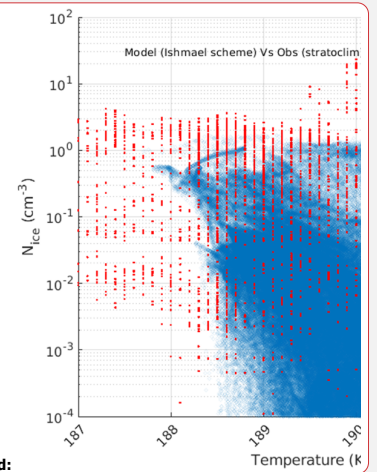
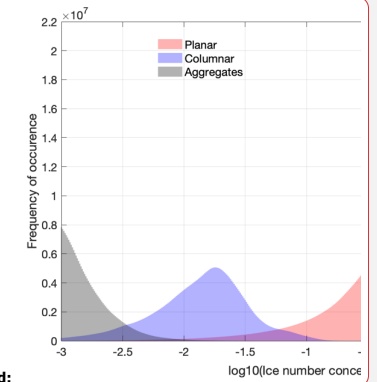
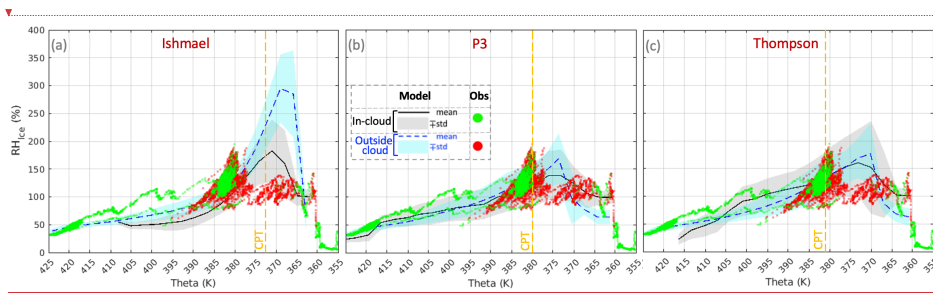



Figure 12: Relative humidity with respect to ice versus potential temperature in-cloud and outside cloud from observation (red and green) and simulation (gray and blue) using Ishmael (left), P3 (middle) and Thompson (right) schemes. The vertical orange dashed line represent the level of cold point temperature (CPT)

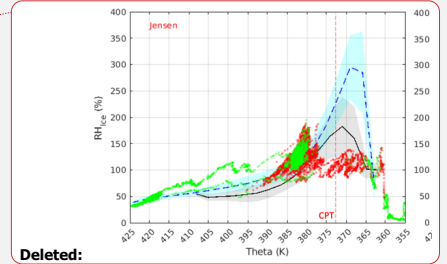
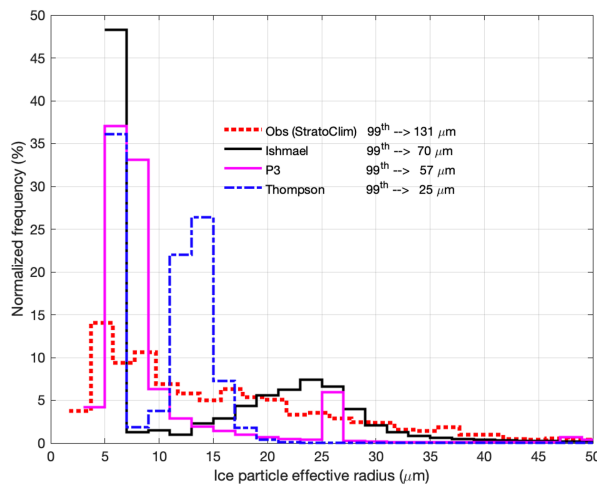



Figure 13: Effective radius distribution at 17 km from observation (red) and simulation using Ishmael (Black), P3 (purple) and Thompson (blue) schemes.

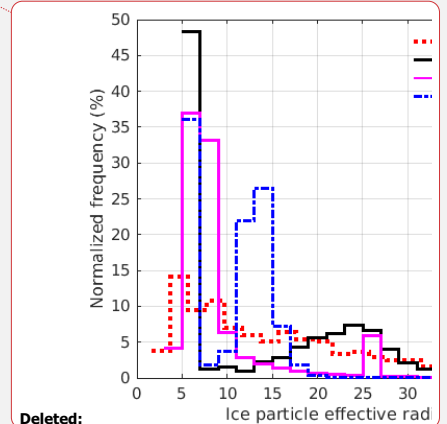
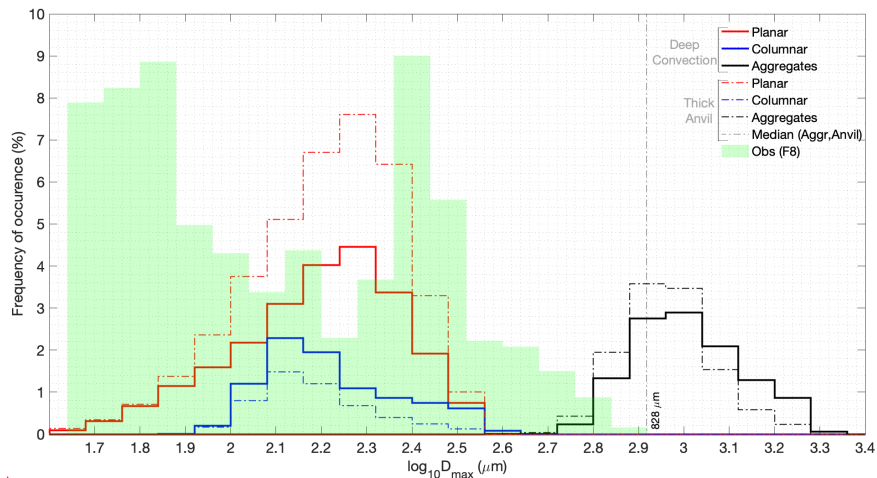



Figure 14: Distribution of ice particles maximum dimension (D_max) from observation (green shade) and the Ishmael scheme within the TTL in the deep convection (solid line) and anvil (dash-dotted) regions for ice habits; planar (red), columnar (blue) and aggregates (black).

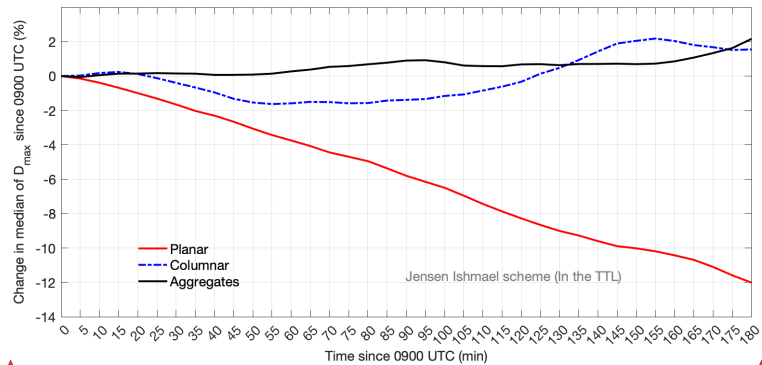



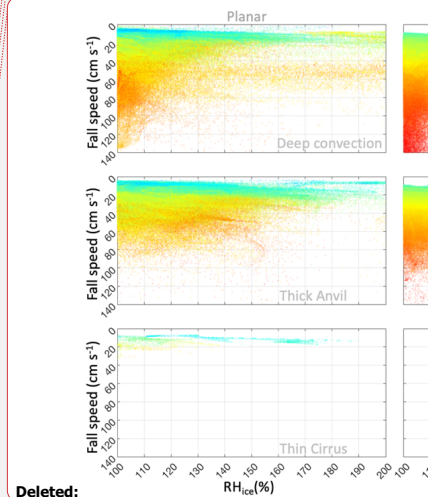
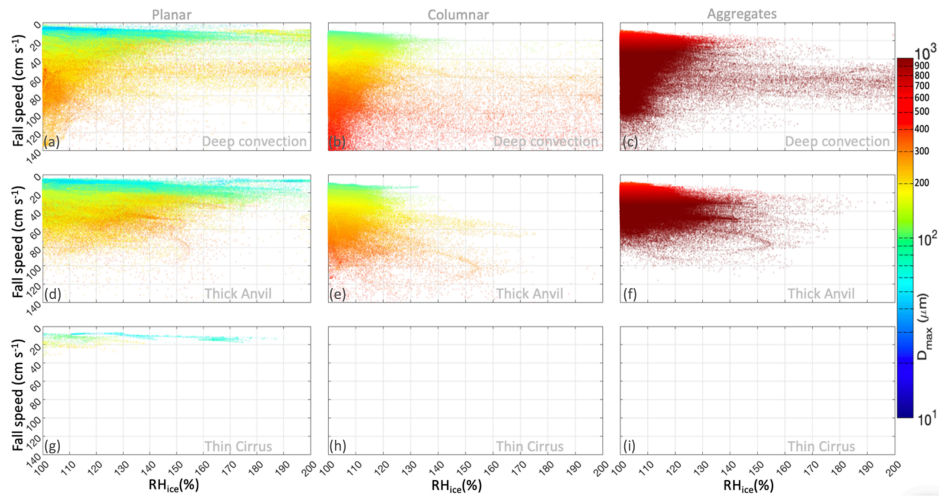
Figure 15: Time series (starting from 2017-08-10 09:00 UTC) of the change in median value of simulated maximum dimension in the TTL using the Ishmael scheme for ice habits; planar (red), columnar (dash-dotted blue) and aggregates (black).

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Figure 16: Scatter plots of ice particle fall speed as a function of relative humidity with respect to ice (RHi), color coded by maximum dimension (D_{max}) for three different ice habits, planar (a,d,g), columnar (b,e,h), aggregates (c,f,i), and three different regions, deep convection (a,b,c), thick anvil (d,e,f) and thin cirrus (g,h,i).

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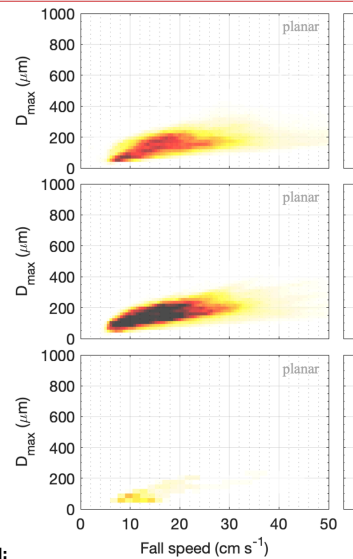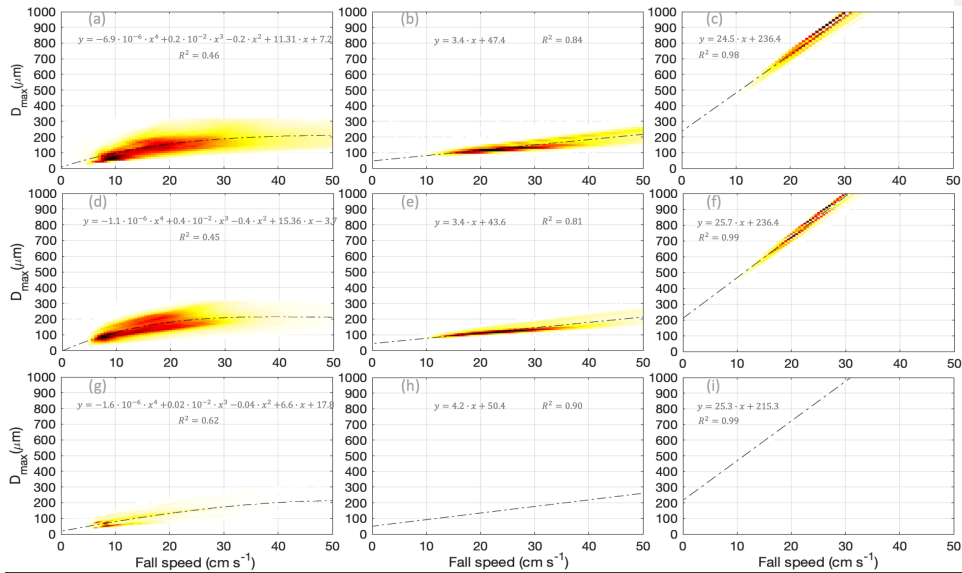
Figure 17: Density plots of maximum dimension as function of fall speed, color coded by bin counts for three different ice habits, planar (a,d,g), columnar (b,e,h), aggregates (c,f,i), and three different regions, deep convection (a,b,c), thick anvil (d,e,f) and thin cirrus (g,h,i). Dash-dotted lines represent fit lines.
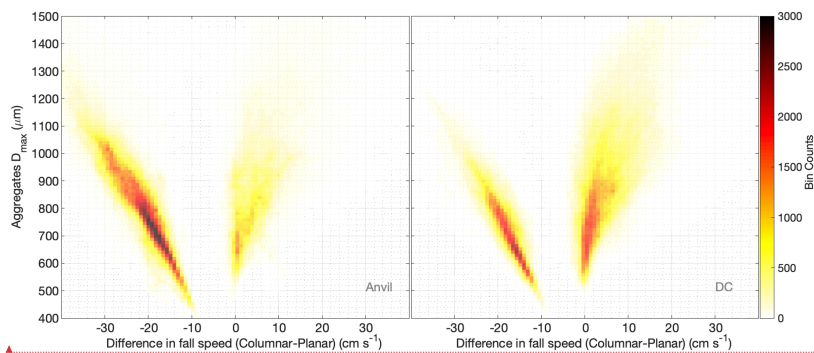
Figure 18: Density plots of maximum dimension of aggregates as function of the difference in fall speeds of columnar and planer habits, color coded by bin counts for two regions in the TTL, deep convection (right) and thick anvil (left).

Figure 19: Histograms of diabatic heating in the TTL from microphysics (right column) and radiation (left column) in the deep convection (upper row), thick anvil (middle row), and thin cirrus (bottom row) regions, and different ice habits; columnar (blue) planar (red) and aggregates (black).

APPENDIX

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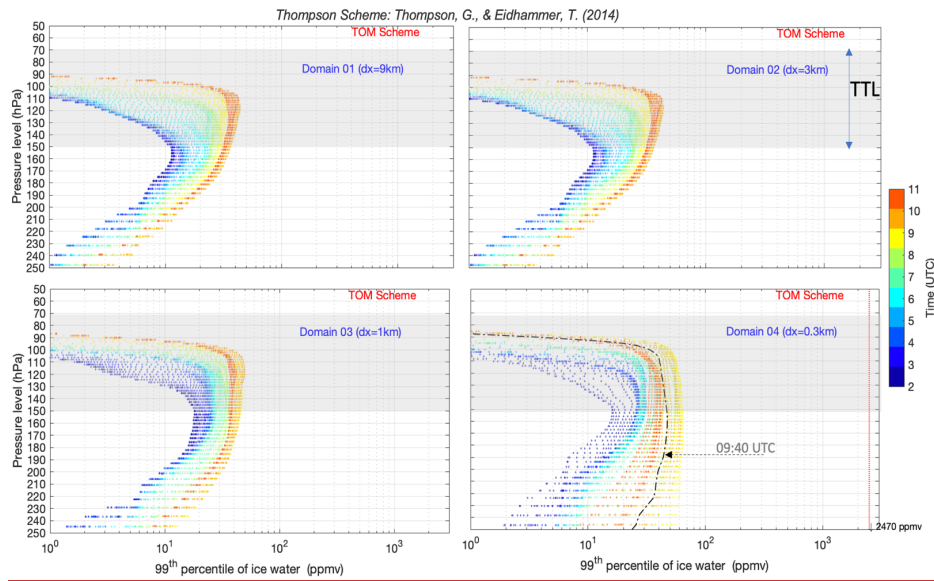
Figure A1: Profile of 99th percentile of simulated IWC below and within the TTL (gray shade) using Thompson microphysics scheme, grid spacing of 9 km (upper left), 3 km (upper right), 1km (lower left) and 0.3 km (lower right) color coded by time (UTC) from 2 to 11 AM. Black dash-dotted line (lower right) represents the simulated profile at 09:40 UTC (timing of the first captured peak of IWC by aircraft) and the red vertical dotted line represents the observed value.

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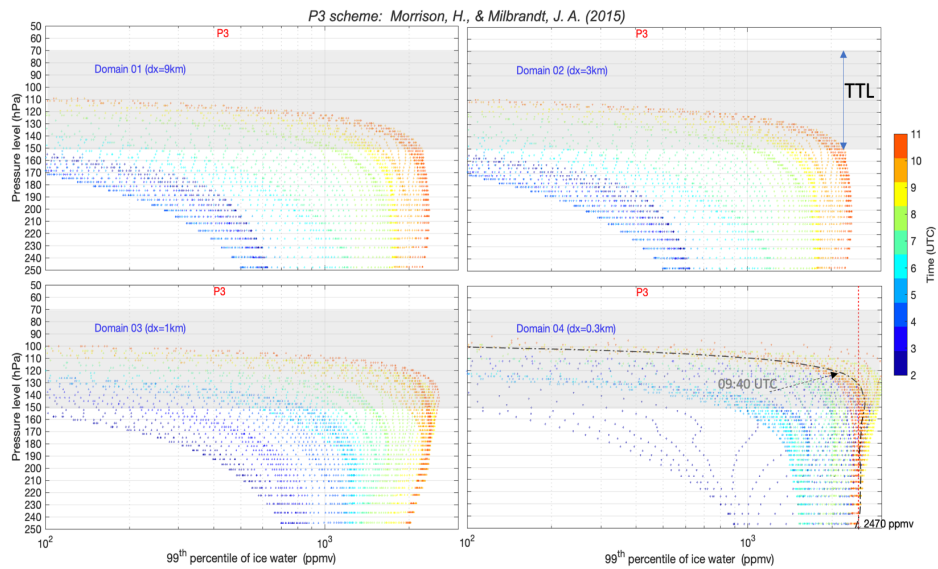
Figure A2: Profile of 99th percentile of simulated IWC below and within the TTL (gray shade) using P3 microphysics scheme, grid spacing of 9 km (upper left), 3 km (upper right), 1km (lower left) and 0.3 km (lower right) color coded by time (UTC) from 2 to 11 AM. Black dash-dotted line (lower right) represents the simulated profile at 09:40 UTC (timing of the first captured peak of IWC by aircraft) and the red vertical dotted line represents the observed value.

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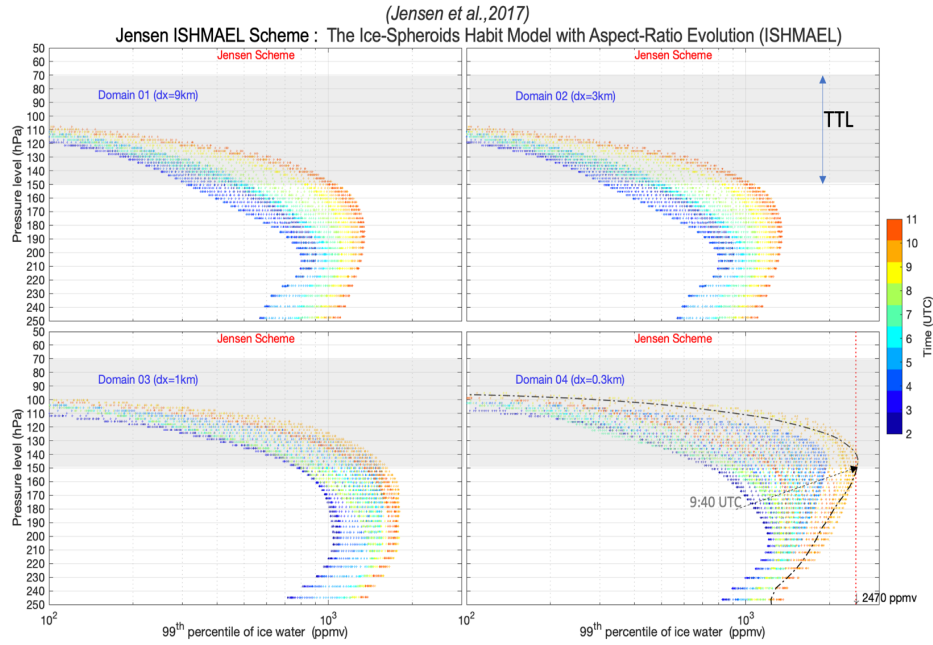
Figure A3: Profile of 99th percentile of simulated IWC below and within the TTL (gray shade) using the Ishmael microphysics scheme, grid spacing of 9 km (upper left), 3 km (upper right), 1km (lower left) and 0.3 km (lower right) color coded by time (UTC) from 2 to 11 AM. Black dash-dotted line (lower right) represents the simulated profile at 09:40 UTC (timing of the first captured peak of IWC by aircraft) and the red vertical dotted line represents the observed value.

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