

1 **Transport pathways of carbon monoxide from Indonesian fire**
2 **pollution to a subtropical high-altitude mountain site in western**
3 **North Pacific**

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12 **Abstract:** Dry conditions associated with El Niño and a positive Indian Ocean Dipole (IOD) are
13 known to have caused major fire pollution events and intense carbon emissions over a vast spatial
14 expanse of Indonesia in October 2006 and 2015. During these two events, a substantial increase in
15 carbon monoxide (CO) mixing ratio was detected by in-situ measurements at Lulin Atmospheric
16 Background Station (LABS, 23.47°N 120.87°E, 2,862 m ASL) in Taiwan, the only background
17 station in the subtropical western North Pacific region. Compared to the long-term October mean
18 (2006-2021), CO was elevated by ~47.2 ppb (37.2%) and ~36.7 ppb (28.9%) in October 2006 and
19 2015, respectively. This study delineates plausible pathways for CO transport from Indonesia to
20 LABS using MOPITT CO observations and MERRA-2 reanalysis products (winds and
21 geopotential height (GpH)). Two simultaneously occurring transport pathways were identified: (i)
22 horizontal transport in the free troposphere and (ii) vertical transport through the Hadley
23 circulation (HC). The GpH analysis of both events revealed the presence of a high-pressure
24 anticyclone over the northern part of the South China Sea (SCS), which played an important role
25 in the free tropospheric horizontal transport of CO. In this scenario, CO in the free troposphere is
26 transported on the western edge of the high-pressure system and then driven by subtropical
27 westerlies to LABS. Simultaneously, uplifted CO over Indonesia can enter the HC and transfer to
28 subtropical locations such as LABS. The vertical cross-section of MOPITT CO and MERRA-2
29 vertical pressure velocity supported the transport of CO through the HC. Further, the results
30 revealed a distinct HC strength in two events (higher in 2006 compared to 2015) due to the

31 different El Niño conditions. Overall, the present findings can provide some insights into
32 understanding the regional transport of pollution over Southeast Asia and the role of climate
33 conditions on transport pathways.

34 **Keywords:** Indonesian fire pollution; Carbon monoxide; Lulin Atmospheric Background Station;
35 Hadley circulation

36 1. Introduction

37 Fire activity over Southeast Asia (SEA), particularly over the Maritime Continent (MC,
38 including Indonesia), is a severe environmental problem that causes widespread regional pollution
39 in the lower troposphere and impacts atmospheric chemistry, air quality, and climate at regional
40 to global scales. Over the MC, fires occur predominately in the dry season (August to October)
41 and particularly during the periods of drought, often associated with the positive phase of El Niño-
42 Southern Oscillation (ENSO) events (Duncan et al., 2003a; van der Werf et al., 2008, 2017; Field
43 et al., 2009, 2016). A recent study has also highlighted the role of the Indian Ocean Dipole on MC
44 fire activity (Pan et al., 2018). For example, dry conditions associated with the positive IOD during
45 the the extreme 2015/16 El Niño and weak 2006/07 El Niño events led to increased fire activity
46 over Indonesia and the wider MC (van der Werf et al., 2008; Chandra et al., 2009; Nassar et al.,
47 2009; Huijnen et al., 2016; Field et al., 2016). Due to these intense fires, an enormous amount
48 of carbon emissions was released into the atmosphere in the form of carbon dioxide (CO₂),
49 carbon monoxide (CO), and methane (CH₄) (Huijnen et al., 2016; Field et al., 2016; Parker et
50 al., 2016; Heymann et al., 2017). The impact of these two Indonesian fire events on carbon
51 emissions, tropospheric trace gases, aerosol composition, and air quality has been extensively
52 discussed in the literature. These two Indonesian fire events and the associated impacts on carbon
53 emissions, trace gas and aerosol composition, and air quality has been extensively discussed in the
54 literature (Chandra et al., 2006; Logan et al., 2008; Chandra et al., 2009; Nassar et al., 2009;
55 Huijnen et al., 2016; Field et al., 2016; Heymann et al., 2017; Ravindra Babu et al., 2019). For
56 example, the fire carbon emissions during September-October 2015 over Maritime SEA were the
57 largest since 1997 (Huijnen et al., 2016). By using Greenhouse gases Observing SATellite
58 (GOSAT) data, Parker et al. (2016) reported the strong enhancement of CO₂ and CH₄ over the
59 Indonesian region.

60 CO is a significant emission from the combustion of fossil fuels and biomass (forest and
61 savanna fires, biofuel use, and waste burning) and is widely used as a tropospheric tracer for these
62 sources (Ou-Yang et al., 2014; Pani et al., 2019). Inter-annual variability of CO in the tropics and
63 sub-tropics is largely linked to year-by-year changes in biomass burning (BB) emissions.
64 Indonesian fires often emit large quantities of CO by incomplete combustion associated with the
65 occurrence of peat fire pollution. Although CO is not a direct greenhouse gas (GHG), it does have
66 a global warming potential due to its chemical reactions in the atmosphere. ~~CO is also an ozone~~
67 ~~(O₃) precursor in the troposphere, and indirectly increases radiative forcing (0.23 +/- 0.05 W m⁻²)~~
68 ~~through the production of O₃ and CO₂ and depletion of hydroxyl radical, the primary chemical~~
69 ~~reactant with CH₄ in the atmosphere (IPCC, 2013). For instance, CO oxidation produces CO₂ and~~
70 ~~indirectly ozone (O₃), both of which are GHGs, and depletes hydroxyl radical (OH) concentrations,~~
71 ~~thus extending the lifetime of CH₄, another GHG in the atmosphere (IPCC, 2013).~~ The lifetime of
72 CO in the free troposphere is ~ two months, thus can be a tracer from polluted upwind regions to
73 remote downwind areas (Cooper et al., 2012). Some of the studies reported the influence of
74 Indonesian fire activity and the transport of CO from Indonesia to the Indian Ocean, Southern
75 Pacific, and western Pacific Ocean (Matsueda and Inoue, 1999; Pochanart and Akimoto, 2003;
76 Nara et al., 2011; Matsueda et al., 2002, 2019). However, the underlying transport mechanisms
77 sending this fire pollution to downwind northern hemisphere subtropical locations, particularly
78 transport to high-altitude background locations in the western north Pacific are still unclear.

79 ~~As shown in Figure 1, the Lulin Atmospheric Background Station (LABS, 23.47°N~~
80 ~~120.87°E, 2862 m ASL) located in central Taiwan was constructed in 2006 and Taiwan is located~~
81 ~~downwind of East Asia and Southeast Asia, which are major air pollutant source regions. As result,~~
82 ~~Lulin Atmospheric Background Station (LABS, 23.47°N 120.87°E, 2862 m above sea level), was~~
83 ~~constructed in 2006 to study the transboundary transport of these air pollutants and their impact~~
84 ~~on Taiwan. is the only high altitude background station in the western Pacific region for~~
85 ~~monitoring the long term variability of atmospheric compositions and also studying the influence~~
86 ~~of continental outflow and long range transported pollution (Lin et al., 2013; Ou Yang et al., 2014,~~
87 ~~2022; Ravindra Babu et al., 2022). The LABS is not affected by local sources, i.e., industrial and~~
88 ~~traffic emissions often found within the free troposphere, making it an ideal site for measuring~~
89 long-range transport of air pollutants, complementing the global network of the Global
90 Atmospheric Watch (GAW) in the East Asia region where no other high-altitude background

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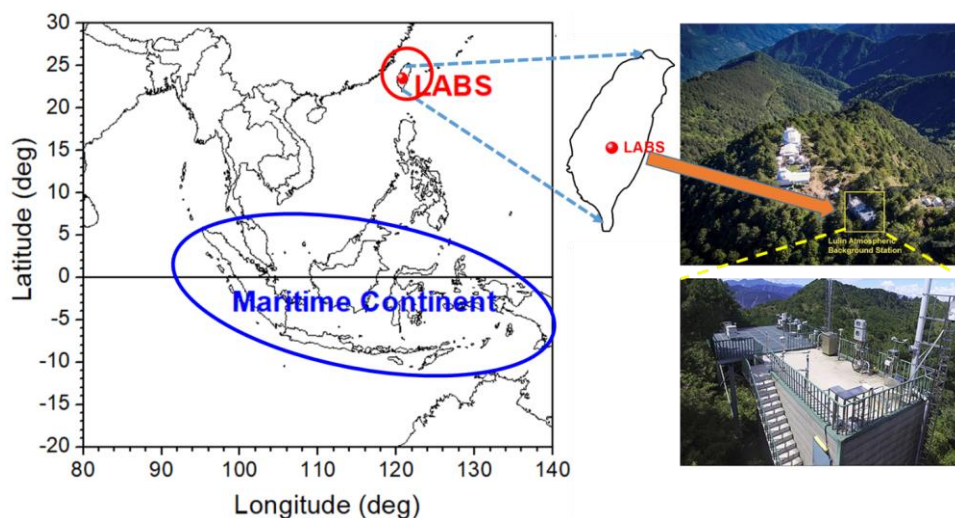
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91 station is available (Ou-Yang et al., 2014, 2022). In the framework of Seven South-East Asian
92 Studies (7-SEAS, Reid, et al., 2013; Lin et al., 2013; Wang et al., 2015), several studies at LABS
93 have reported on the long-range transport of northern peninsular Southeast Asia (PSEA) BB
94 pollutants to Taiwan through the low-level jet (LLJ) and the related impacts on air quality and
95 chemistry over Taiwan (Ou-Yang et al., 2012, 2014; Lin et al., 2013; Chuang et al., 2016; Chi et
96 al., 2016; Tsay et al., 2016; Hsiao et al., 2016; Lin et al., 2017; Park et al., 2019; Pani et al., 2016,
97 2019; Huang et al., 2019; Huang et al., 2020; Ravindra Babu et al., 2022). However, to date, no
98 studies have shown the potential influence of Indonesian fire activities on LABS measurements
99 and the BB pollution from Indonesian fires reaching LABS. Surprisingly, the extensive fire events
100 in 2006 and 2015 allowed us to track CO concentrations from the Indonesian peat fires to LABS
101 in Taiwan. By combining in-situ and satellite CO measurements and large-scale circulation
102 parameters from reanalysis products, we identified plausible transport pathways from Indonesia to
103 LABS.



105 **Figure 1.** Geographic location of the Maritime Continent and Lulin Atmospheric Background
106 Station (LABS, 23.47°N 120.87°E, 2862 m ASL), Taiwan.

107 **2. Site description, Data data and methodology**

108 **2.1 Site description**

LABS is located on the summit of Mount Lulin and is shown in Figure 1, along with the location of the Maritime Continent. Hiking is the only way to access LABS, taking about 30 minutes from the nearest parking lot. There are no known point emission sources at the summit or in the surrounding area with the exception of the occasional maintenance activity at the Lulin Observatory. Because of the high altitude of LABS, measurements there are not affected by local pollution from factories, traffic, and other domestic sources; rather, it is strategically located to monitor long-range transported air pollutants from the Asian continent. More details about the instruments and their specifications can be found in Sheu et al., 2009.

2.2 In-situ measurements

Details of the CO and various meteorological measurements at LABS employed in the current study have been previously described in detail (Sheu et al., 2009; Ou-Yang et al., 2014; Ravindra Babu et al., 2022) and are thus only briefly described here. The long-term monthly mean of various meteorological parameters such as temperature, relative humidity, wind speed, and wind direction along with CO at LABS can be found in **Sup. Figure 1**. The overall mean temperature (relative humidity) was about 10.5°C (~80%), with monthly mean temperatures ranging between ~5 and 14°C. Local wind direction is mostly from the southwest and to a lesser extent from the northeast. Long-term monthly mean in CO shows distinct seasonal patterns with a springtime maximum and a summertime minimum at LABS. CO ~~mixing ratios~~ measurements were measured by a nondispersive infrared (NDIR) analyzer (APMA-360, Horiba, Japan) at LABS. Hourly averages of the 6-s data were analyzed in this study. The detection limit of the NDIR is ~20 ppb (1σ) (Zellweger et al., 2009); more details about CO measured at LABS can be found in Ou-Yang et. (2014). The magnitude of the CO concentration enhancement in 2006 and 2015 above the long-term background was determined by comparing a 16-year average (2006-2021) of October CO data at LABS. We obtained the percentage change in CO relative to the respective background using Equation 1:

$$\text{Relative change in percentage} = \left(\frac{x_i - \bar{x}}{\bar{x}} \right) \times 100 \quad (\text{Eq. 1})$$

where x_i represents the monthly mean of October in 2006 and 2015, and \bar{x} is the corresponding monthly long-term mean calculated using the data from 2006 to 2021 (Ou-Yang et al., 2014).

2.2.3 Satellite measurements

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138 CO observations from the Measurement of Pollution in the Troposphere (MOPITT, version
139 8) instrument were also utilized in this study (Worden et al., 2010; Deeter et al., 2019). MOPITT
140 is a multi-channel Thermal InfraRed (TIR) and Near InfraRed (NIR) instrument operating onboard
141 ~~the~~ sun-synchronous polar-orbiting NASA Terra satellite. V8 CO products, consisting of a CO
142 profile at ten pressure levels, have been validated; more details about the retrieval algorithm,
143 validation, and ~~the~~ uncertainties of MOPITT CO can be found in Deeter et al. (2019). In addition
144 to the MOPITT measurements, we utilized CO from the Atmospheric Infrared Sounder (AIRS) on
145 the NASA Aqua satellite, which provides CO at different vertical levels twice daily and near-
146 global coverage. AIRS uses wavenumbers 2,183-2,200 cm^{-1} (4.58-4.5 μm) for retrieving CO
147 (McMillan et al., 2005). Version 8, level 3 CO product, available at $1^\circ \times 1^\circ$ resolution at various
148 pressure levels, was utilized in the present study. AIRS data were downloaded from the following
149 website https://disc.gsfc.nasa.gov/datasets/AIRS3STM_7.0 (AIRS project., 2019). AIRS
150 sensitivity to CO is broad and optimal in the mid-troposphere between approximately 300 and 600
151 hPa (Warner et al., 2007; Warner et al., 2013; AIRS project., 2019). CO retrievals have a bias of
152 6-10% between 900 hPa and 300 hPa with a root mean square error of 8-12 % (McMillan et al.,
153 2011).

154 Apart from MOPITT and AIRS CO data, we used Moderate Resolution Imaging
155 Spectroradiometer (MODIS) collection 6.1 daily active fire hot spot data from 2006-~~2020-2021~~
156 over Indonesia (Giglio et al., 2016).

157 **2.3 MERRA-2 Reanalysis products**

158 ~~We also utilized monthly mean geopotential height (GPH), wind vectors (zonal and meridional~~
159 ~~wind speed), and pressure vertical velocity from the Modern-Era Retrospective Analysis for~~
160 ~~Research and Applications, version 2 (MERRA-2). Modern-Era Retrospective Analysis for~~
161 ~~Research and Applications, version 2 (MERRA-2) monthly mean geopotential height (GpH) wind~~
162 ~~vectors (zonal and meridional) and vertical pressure velocity (omega) during the study period were~~
163 ~~utilized.~~ MERRA-2 is the latest atmospheric reanalysis data produced by the NASA Global
164 Modeling and Assimilation Office (GMAO) (Gelaro et al., 2017). The horizontal resolution of
165 MERRA-2 reanalysis is $0.5^\circ \times 0.625^\circ$. MERRA-2 data are available online through the NASA
166 Goddard Earth Sciences Data Information Services Center (GES DISC; <https://disc.gsfc.nasa.gov/>,
167 last access: 11 September 2022).

168 **3. Results and Discussion**

169 **3.1 Higher CO mixing ratios in October 2006 and 2015 over Maritime Continent and at**
170 **LABS**

171 **Figure 2** shows the height-time cross-section of monthly mean CO over the Maritime Continent
172 (MC) obtained from MOPITT and AIRS from 2003 to 2021. There is a significant inter-annual
173 variability in the CO time series in **Figure 2** as observed by both instruments. The maximum CO
174 mixing ratio for this time period was observed in the fall of 2006 and 2015; both were tied to El
175 Niño events (Field et al., 2016; Ravindra Babu et al., 2019). Several studies have reported on the
176 impact of the intense fire activity in 2006 and 2015 and on the release of significant carbon
177 emissions and poor air quality over the wider Equatorial Asia region (Logan et al., 2008; Chandra
178 et al., 2009; Field et al., 2016; Huijnen et al., 2016; Ravindra Babu et al., 2019). Even though 2009
179 and 2014 were El Niño years, the CO over MC was not as high as observed in 2006 and 2015. The
180 weaker and shorter duration of fire activities could largely explain the lower CO over the MC in
181 2009 and 2014 in contrast to those in 2006 and 2015. Furthermore, **Figure 3** shows the temporal
182 variability of monthly mean CO from MOPITT and AIRS from January through December in both
183 years 2006 and 2015, respectively. Both instruments show maximum CO enhancement in October
184 compared to the remaining months in 2006 and 2015. Overall, it is clear from **Figures 2 and 3**
185 that in October 2006 and 2015, CO over the Maritime Continent in the entire troposphere increased
186 dramatically due to increased CO emissions near the surface from extreme fire activity (**Fig. 4c**).

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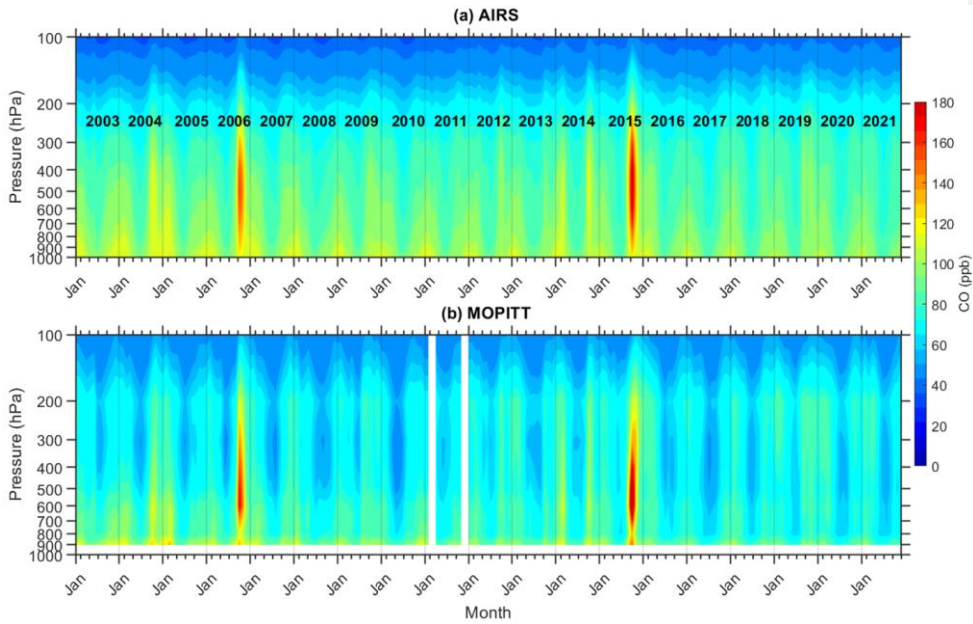
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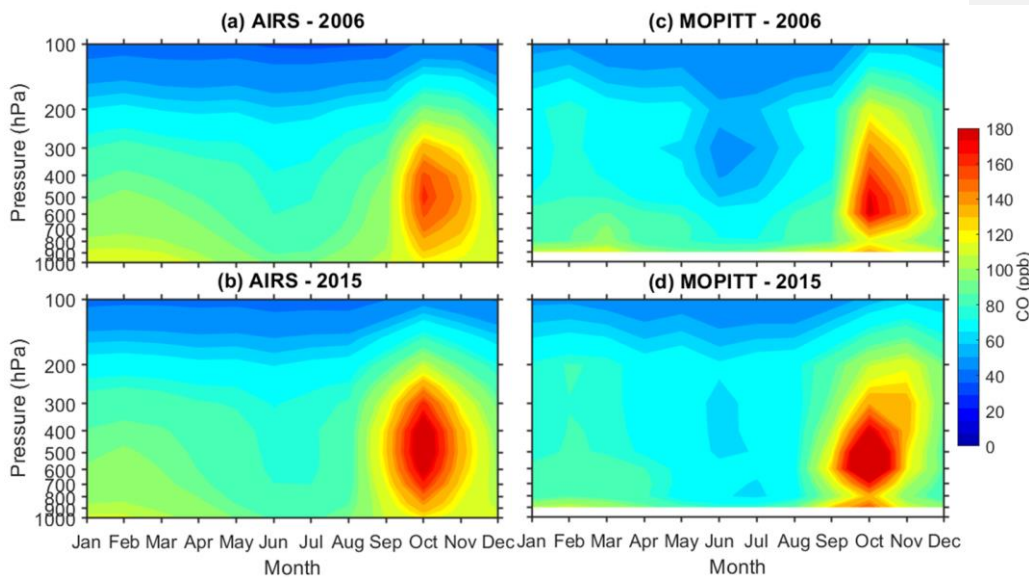
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187
 188 **Figure 2.** Pressure-time cross-section of monthly mean carbon monoxide observed over the
 189 Maritime continent (average over 90E-140E,10S-10N) during 2003-2021 obtained from (a) AIRS,
 190 and (b) MOPITT satellite measurements.



191

192 **Figure 3.** Pressure-time cross-section of monthly mean carbon monoxide observed over the
 193 Maritime continent (average over 90E-140E,10S-10N) in (a) 2006, (b) 2015 obtain from AIRS
 194 satellite measurements. Subplots (c) and (d) are the same as subplots (a) and (b) but for the
 195 MOPITT satellite measurements.

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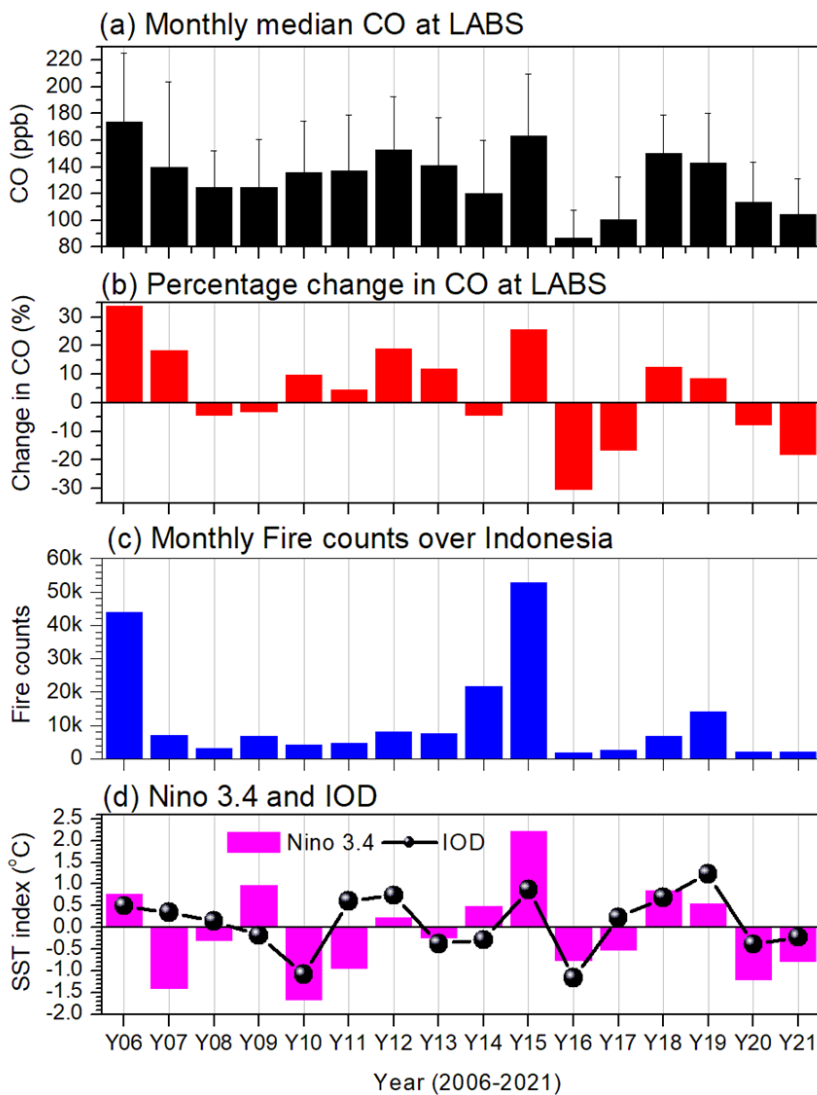
196 **Higher CO mixing ratios in October 2006 and 2015 at LABS**

197 **Figure 2-4** summarizes the inter-annual variations of CO in October observed at LABS
 198 along with MODIS active fire counts over Indonesia and the observed Niño 3.4 and the IOD index
 199 values, which helped to motivate this study. The highest CO mixing ratios for this period were
 200 observed in 2006 and 2015, well over the long-term means of 132.1 ± 23.3 ppb when including all
 201 points and 126.8 ± 19.6 ppb when excluding 2006 and 2015. A significant enhancement of CO, over
 202 the latter mean calculation, of more than 47.2 ppb (37.2%) in 2006 and 36.7 ppb (28.9%) in 2015
 203 was observed, with the value in 2006 (2015) more significant than the $\pm 2\sigma$ ($\pm 1\sigma$) standard
 204 deviation of the long-term mean (**Table 1**). Higher CO mixing ratios in 2006 and 2015 at LABS
 205 were also evident from the MOPITT and AIRS satellite measurements obtained over a 1-degree
 206 radius around the LABS location (**Sup. Fig. 31**).

207 Unprecedented CO values in 2006 and 2015 at LABS could be due to the transport of CO
208 from large-scale forest fires that were intense during the same period in the Indonesian ~~region~~
209 ~~region~~. It is clear from **Figure 24**, that the higher values of CO at LABS in 2006 and 2015
210 coincided with more intense fire activity over Indonesia along with warm phases of ENSO and
211 IOD (**Fig. 2e-4c and 2d4d**), which have been extensively studied due to the induced drought
212 conditions in those years (Field et al., 2016; Huijnen et al., 2016; Pan et al., 2018). Previous studies
213 (e.g., Logan et al., 2008; Zhang et al., 2011; Field et al., 2016; Pan et al., 2018) have demonstrated
214 the direct relationship between strong Indonesian fires and El Niño events. ~~Several studies have~~
215 ~~reported on the impact of the intense BB in 2006 and 2015 on the release of significant carbon~~
216 ~~emissions and the air quality over the wider Equatorial Asian region (Logan et al., 2008; Chandra~~
217 ~~et al., 2009; Field et al., 2016; Huijnen et al., 2016; Ravindra Babu et al., 2019).~~ The enhanced
218 CO values from the 2006 and 2015 events at LABS in the present study complement the findings
219 of Matsueda and Inoue (1999) in the case the of 1997 El Niño event and Nara et al. (2011) in the
220 case of 2006 El Niño event. However, the impact on CO at LABS occurred significantly further
221 north of the source region than in either of the aforementioned studies. Based on aircraft
222 measurements, Matsueda and Inoue (1999) reported the enhancement of CO₂, CO, and CH₄ in the
223 upper troposphere (at 9-12 km) over the South China Sea (SCS) during October 1997 Indonesian
224 fire event. However, this large CO increase appeared only over the SCS west of Kalimantan and
225 not in the subtropics between 10°N and 26°N. Nara et al. (2011) reported a substantial increase in
226 CO mixing ratios over the Western Tropical Pacific Ocean (between 15°N and the Equator) by
227 shipboard observations routinely operated between Japan and Australia and New Zealand during
228 October and November of 2006. Similarly, Pochanart and Akimoto (2003) also reported the
229 influence of the 1997 Indonesian fire event on CO enhancement at the rural station Srinakarin
230 (14°22'N, 99°07'E, 296 m above sea level) in Thailand.

231 In addition, due to La Niña and the negative phase IOD, the fire activity in Indonesia during
232 2016 was much less intense than in 2006 and 2015 (**Fig. 2e-4c and 2d4d**). Interestingly, CO at
233 LABS during 2016 exhibited the lowest October values in the entire data period, ~39.8 ppb
234 (31.4%) lower than the long-term October mean (2006-2021). It is well known that the major
235 sources of CO at LABS are BB from peninsular SEA in spring and industrial emissions from
236 continental Asia in winter (Ou-Yang et al., 2014; Pani et al., 2019; Ravindra Babu et al., 2022;
237 Ou-Yang et al., 2022). However, October is a transition month from the summer to winter at

238 LABS, when air masses can still arrive from the Pacific Ocean. Our analysis (**Fig. 24**) suggests
239 that the extensive fires that occurred during the 2006 and 2015 El Niño events over Indonesia may
240 have yielded the unprecedented CO mixing ratios at LABS in October of those years. Combined
241 El Niño and IOD-related changes in the large-scale dynamics and circulations may have promoted
242 CO emissions from Indonesian fires to transport to LABS.



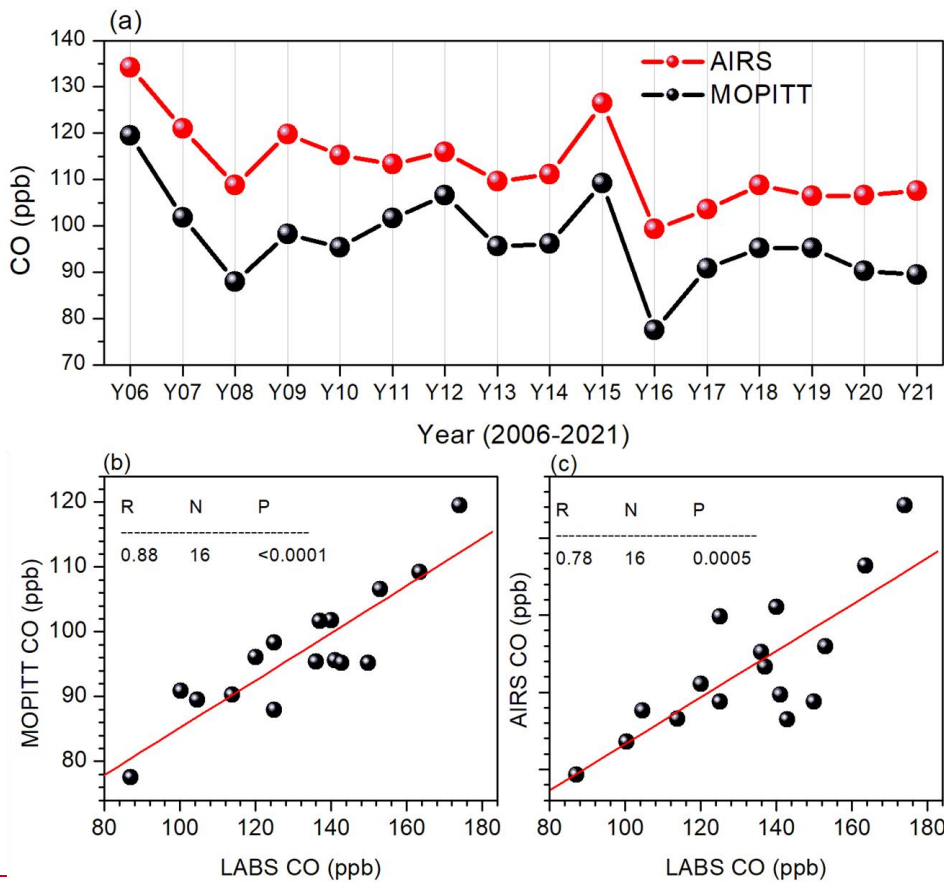
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244 **Figure 24.** Inter-annual variations in October of the (a) monthly median of CO, (b) percentage
 245 change in CO from the long-term mean at LABS, (c) MODIS (Moderate Resolution Imaging
 246 Spectroradiometer) total active fire counts (only fires tagged with >30 % confidence) over

247 Indonesia, (d) sea surface temperature index for Niño 3.4 (magenta) and Indian ocean dipole
 248 (black) during 2006 to 2021.

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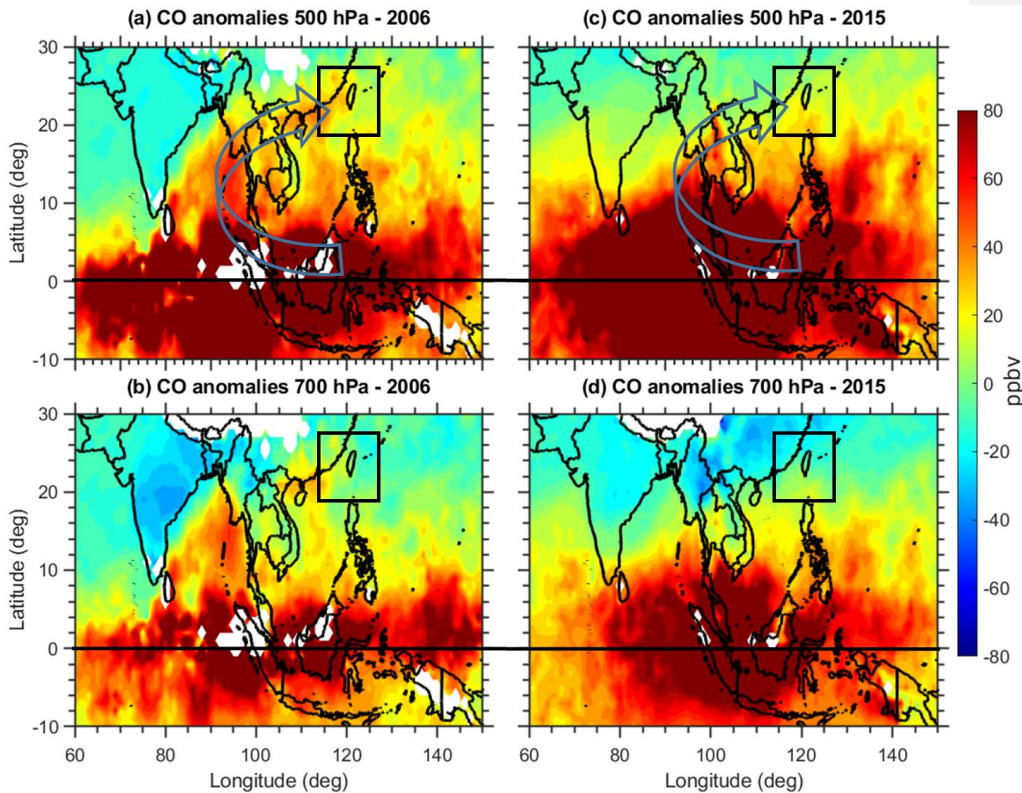
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251 -
 252 **Figure 3.** (a) MOPITT (black) and AIRS (red) satellite-observed CO mixing ratios within the 1-
 253 degree radius around the LABS location, (b) correlation plot between in-situ CO at LABS and
 254 MOPITT CO, and (c) correlation plot between in-situ CO at LABS and AIRS CO in October
 255 month during 2006 to 2021. (R is the correlation coefficient; N is the sample size; P is the
 256 significance value)

257

258 To confirm the impact of Indonesian fire pollution on LABS CO, we further checked the
259 spatial distribution of CO in 2006 and 2015 from the MOPITT satellite CO observations. An inter-
260 comparison between October monthly mean CO at LABS (2006-2021) and MOPITT and AIRS
261 CO data at 700 hPa within the 1-degree radius around the LABS location yielded correlation
262 coefficients of 0.88 and 0.78 ($p < 0.01$), respectively (Sup. Fig. 32). We then used the MOPITT
263 satellite data to track the spatial and vertical CO changes in October 2006 and 2015; first, we
264 examined the distribution of the CO anomalies at free tropospheric heights in those years. **Figure**
265 **4-5** shows these anomalies compared to the long-term mean (2001-2021) at 700 hPa and 500 hPa,
266 revealing extensive enhancements of CO mixing ratios over most of equatorial Asia in 2006 and
267 2015. **Figure 4-5** indicates that CO from the Indonesian fires affected both the Indian Ocean to the
268 west and South Pacific and the northern Pacific to the east. Furthermore, these outflows of CO
269 split northwestward into the Bay of Bengal and northeastward into the western North Pacific. It is
270 also worth noting that the anomalies were significantly higher at 500 hPa than 700 hPa. Elevated
271 CO is visible in the Taiwan region at 700 hPa and 500 hPa in both years. This further provides a
272 clear signature of the impact of Indonesian fire activity on enhanced CO in 2006 and 2015 at
273 LABS. Overall, from **Figure 4-5**, MOPITT CO data shows the Indonesia fires transported CO
274 vertically and horizontally in all directions. We further investigated the associated dynamics and
275 large-scale circulations supporting the transport of Indonesian pollution to LABS.

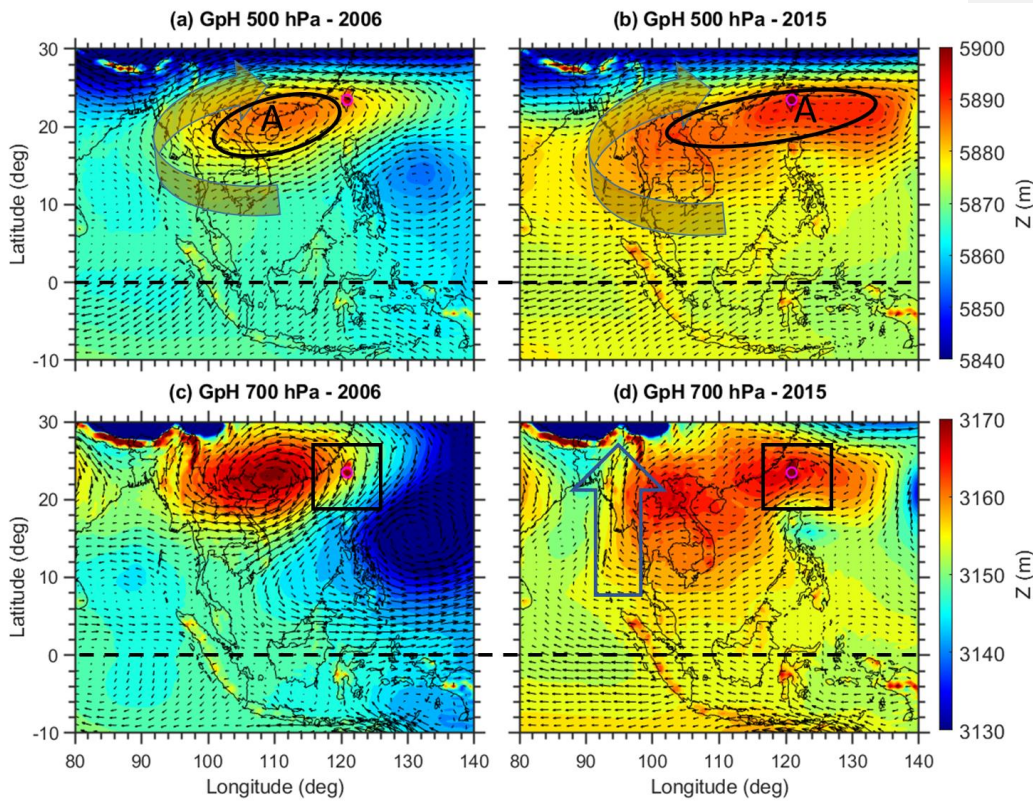


276
 277 **Figure 45.** Monthly mean CO anomalies obtained from MOPITT satellite observations (a) at 500
 278 hPa and (b) at 700 hPa during October 2006. Subplots (c) and (d) are the same as subplots (a) and
 279 (b) but for October 2015, respectively. The anomalies are obtained by subtracting the 2006 and
 280 2015 data from the long-term mean of MOPITT CO data from ~~2001 to~~ 2001 to 2021.

281 3.2 Role of large-scale dynamics and atmospheric circulations

282 Large-scale dynamics and circulations can play a crucial role in transporting Indonesian
 283 pollution to long-distance downwind regions (Bowman, 2006; Nara et al., 2011; Matsueda et al.,
 284 2019). To understand the plausible mechanisms behind the transport of Indonesian fire pollution
 285 to LABS, we further examined the MERRA-2 reanalysis of geopotential height (GpH) and wind
 286 distribution in 2006 and 2015. The spatial distribution of GpH at two pressure levels (700 and 500
 287 hPa) in both events is shown in **Figure 56**. The GpH and wind vectors in the two event years

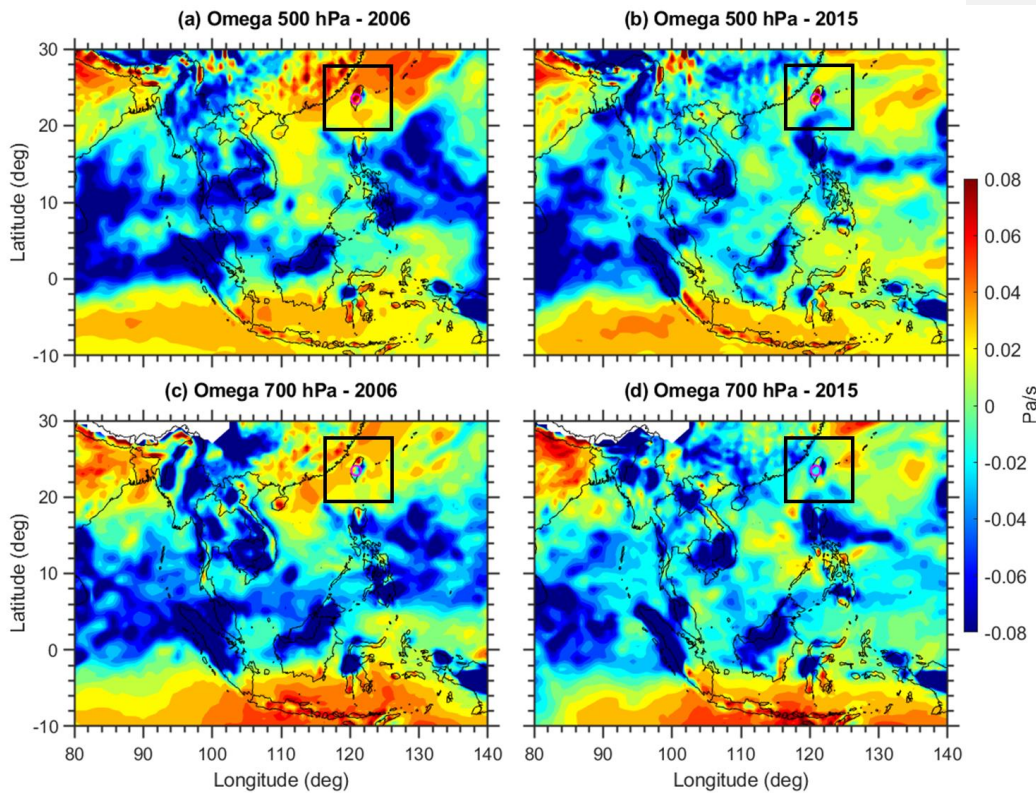
288 exhibited quite different patterns in relation to a high-pressure system over the northern parts of
 289 the SCS. A high-pressure anti-cyclonic circulation center extended from the Indo-China Peninsula
 290 to the SCS in October 2006 with LABS located precisely on the eastern edge of the anticyclone.
 291 In 2015, the anticyclone extended from the Indo-China Peninsula to the western North Pacific
 292 region and over Taiwan.



293
 294 **Figure 56.** Monthly mean Geopotential height (GpH) obtained from MERRA-2 reanalysis (a) at
 295 500 hPa and (b) at 700 hPa during October 2006. Subplots (c) and (d) are the same as subplots (a)
 296 and (b) but for October 2015.

297 During both event years, strong southerlies at 500 hPa were evident due to the high-
 298 pressure anticyclone system in the northern SCS. It is assumed that the northern edge of the
 299 Indonesian fire pollution plume can be carried out by the southerlies and around the western edge

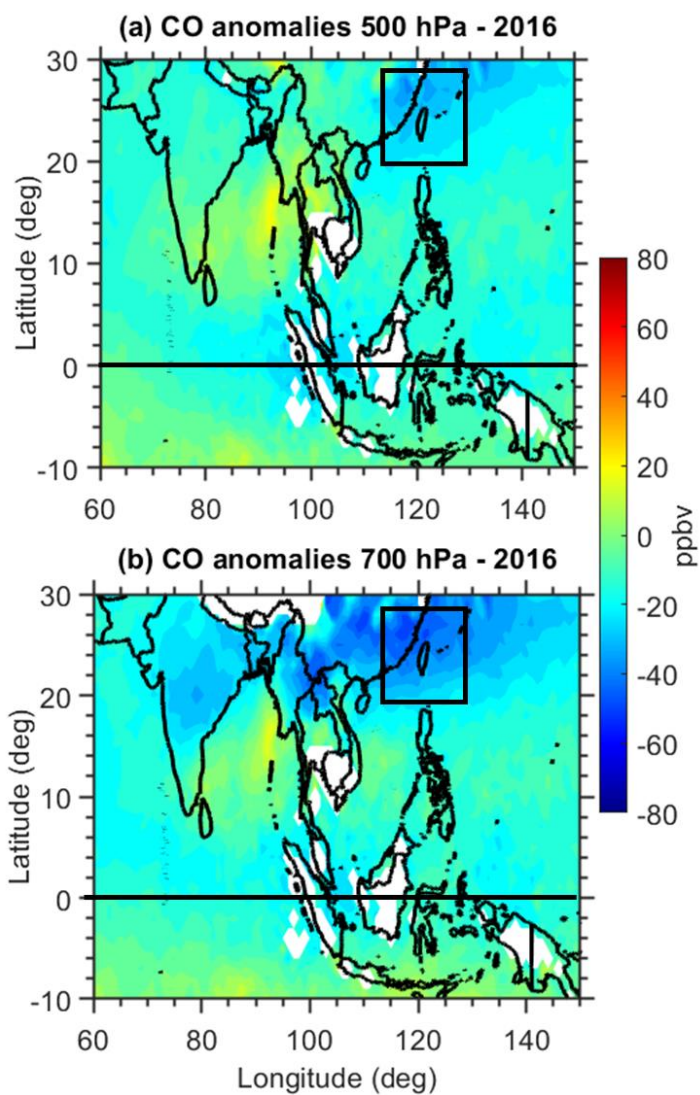
300 of the high-pressure anti-cyclone over SCS. An apparent merging of the southerlies from the
301 equator with the subtropical westerlies in the northern PSEA region subsequently led to the
302 transport of CO to downwind LABS. Overall, in both events, there was a significant anticyclone
303 over the SCS. El Niño and the positive IOD-induced high-pressure anticyclone over SCS
304 strengthen the southerlies from the equator, consequently bringing higher amounts of CO to LABS.
305 We further investigated the vertical pressure velocity (ω) behavior in both events ([Sup. Fig](#)
306 [3Fig-6](#)), where negative (positive) values represent upward (downward) winds. Significant
307 upward wind in both events was evident over equatorial MC, while vertical pressure velocity over
308 Taiwan and surrounding regions at both pressure levels were mostly downwards in 2006 and 2015.
309 The presence of a downwind will provide downward transport of any pollutant presence in the
310 upper troposphere over that region. Also, the downward wind was relatively higher in 2006
311 compared to 2015. The center of the downward wind was shifted eastwards in the western North
312 Pacific in 2015. The distinct behavior of vertical pressure velocity [at-around the LABS region](#)
313 during [these](#) two events might be due to the associated climate conditions in the two periods; more
314 discussion will be provided in section 3.4.



315
 316 **Figure 6.** Monthly mean vertical pressure velocity obtained from MERRA-2 reanalysis (a) at 500
 317 hPa and (b) at 700 hPa during October 2006. Subplots (c) and (d) are the same as subplots (a) and
 318 (b) but for October 2015.

319 We further showed CO deviations at both pressure levels in October 2016 when there was
 320 very low fire activity in Indonesia (Fig. 78). Interestingly, there was a significant lowering of CO
 321 over the Taiwan region in 2016, which agrees with the observed low CO values from the in-situ
 322 measurements at LABS (Fig. 2b4b). Also in agreement, 2016 was a La Niña and negative IOD
 323 year and fire activity was much weaker (Fig. 2e-4c and 2d4d). During the La Niña years, large-
 324 scale dynamical processes are greatly reversed with respect to El Niño years. We further analyzed
 325 the GpH and wind circulation patterns in 2016 (Fig. 89). A significant high-pressure system
 326 (western North Pacific subtropical High) was present over the western North Pacific region in

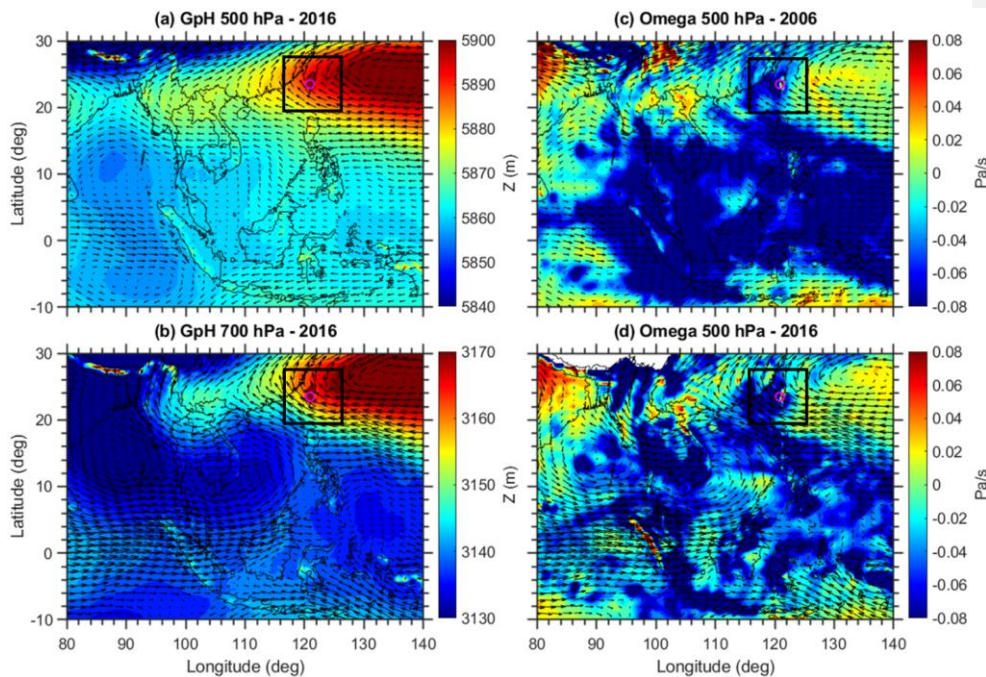
327 2016, which was shifted considerably further eastward compared to ~~over~~ the SCS in 2006 and
328 2015. The wind vectors also highlighted the transport of a clean marine air mass from the Pacific
329 Ocean to LABS in 2016. Interestingly, the vertical pressure velocity exhibited a pronounced
330 upward wind over Taiwan in 2016, in contrast to the downward wind in 2006 and 2015. This
331 indicates that dominant clean marine air reached LABS in 2016 resulting in the lowest CO mixing
332 ratio in the entire dataset at LABS.



333

334 **Figure 78.** Monthly mean CO deviations from the long-term mean (2001-2021) were obtained

335 from MOPITT satellite observations (a) at 500 hPa and (b) at 700 hPa during October 2016.



336

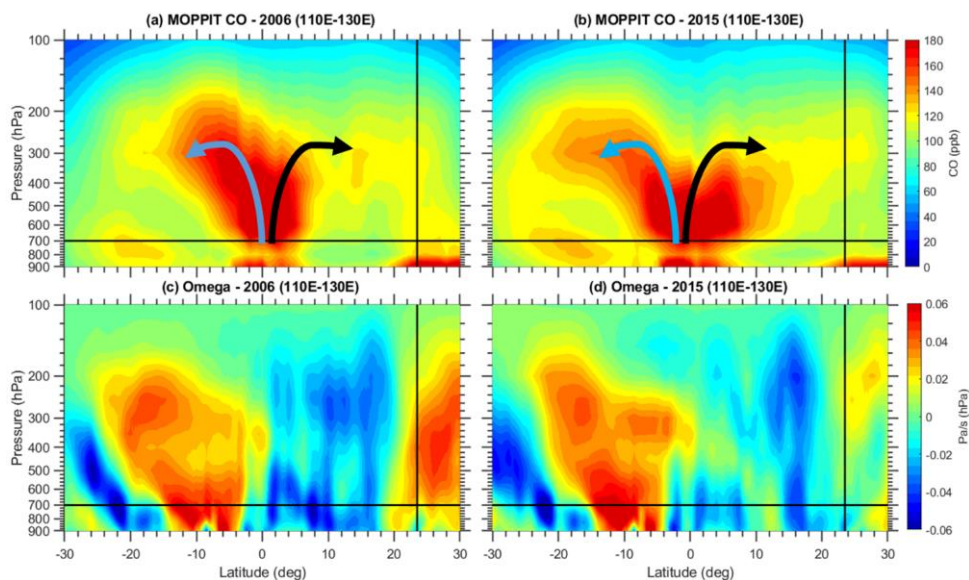
337 **Figure 89.** Monthly mean Geopotential height (GpH) obtained from MERRA-2 reanalysis (a) at
 338 500 hPa, and (b) at 700 hPa during October 2016. The subplots (c) and (d) are the same as subplots
 339 (a) and (b) but for the observed vertical pressure velocity (Omega).

340 3.4 Role of Hadley circulation

341 The Hadley circulation (HC) is a crucial component of the climate system, which is
 342 characterized by a thermally driven large-scale meridional circulation (Hadley, 1735). This
 343 circulation links the troposphere and stratosphere and the tropics and extra-tropics, through
 344 horizontal and vertical motions, transporting moisture, heat, and momentum to regulate Earth's
 345 energy budget. As the CO sources (Indonesia) in this study were close to the equator, it is expected
 346 that air tends to rise more or less directly over the CO sources. **Figure 9-10** shows the vertical-
 347 meridional cross-section of CO and vertical pressure velocity in separate panels averaged along
 348 110°–130°E in October 2006 and 2015. The black-colored vertical line in all the panels in **Figure**
 349 **9-10** shows the location of LABS and the horizontal line represents the 700 hPa. The vertical cross-
 350 section of CO highlights the uplifting of CO into the upper troposphere over the equator, followed

351 by southward and northward movement in both 2006 and 2015 (**Fig. 9a-10a** and **9b10b**). A clear
352 transport of CO from the source region to the sub-tropics via meridional transport was evident in
353 both events. It is noted that the higher CO observed between 20–30°N latitude below ~700 hPa is
354 related to anthropogenic emissions and not due to the Indonesian fires. To confirm the lofted CO
355 from Indonesia ~~is~~ really descended in the subtropics due to the Hadley circulation, we looked into
356 the vertical cross-section of vertical pressure velocity in both events. From **Figure 810**, it is
357 suggested that large amounts of CO from Indonesia were transferred into the free troposphere by
358 the strong upward air motion in this region. Similarly, there was a pronounced descending motion
359 (positive values of vertical pressure velocity) during October 2006 (**Fig. 9e10c**) in the northern
360 hemisphere subtropics around 20–30°N latitude, which corresponds well with the location of
361 LABS. However, in October 2015, the descending motion was not significant compared to 2006.
362 This may be due to the different El Niño conditions in 2006 and 2015. While IOD conditions were
363 indeed similar between 2006 and 2015 (**Fig. 2d4d**), the higher descending motions in 2006 can be
364 explained in part by the moderate El Niño conditions during that year. A well-developed El Niño
365 condition was already established in 2015 compared to 2006. In October 2006, the observed Niño
366 3.4 value was around 0.7 whereas in 2015 it was around 2.21. These values indicate that the El
367 Niño conditions were already well established in October 2015 whereas, in 2006, the conditions
368 were not developed as El Niño. It is reported that in El Niño conditions, the western Pacific HC is
369 observed to be weakened whereas the eastern Pacific HC is strengthened (Wang, 2004). This is
370 supported by the observed lesser descending motions in 2015 from the present study. These
371 differences in the descending motions likely influenced the greater CO enhancement in 2006
372 compared to 2015 at LABS (**Fig. 2b-4b** and **Table 1**). Overall, it is clearly illustrated from the
373 MOPITT CO vertical cross-section and the MERRA-2 vertical pressure velocity that the CO
374 emitted from the Indonesian fire was transported vertically through the Hadley circulation to the
375 LABS location.

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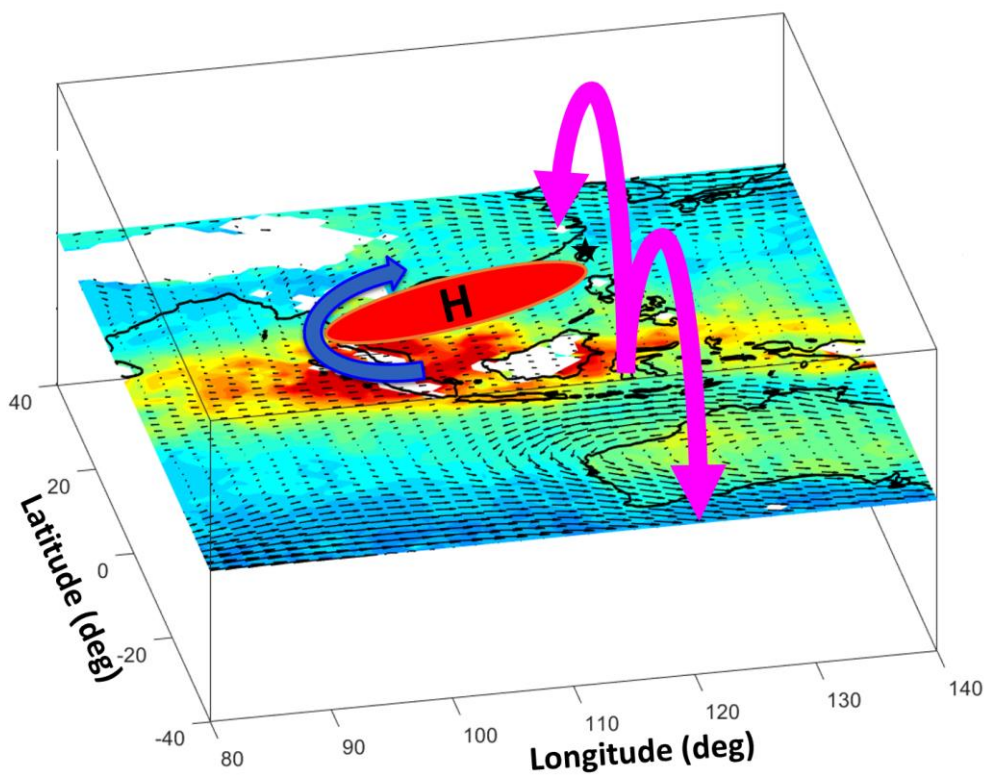


376
 377 **Figure 910.** Vertical pressure-latitude meridional cross-section of MOPITT CO averaged along
 378 110°–130°E (a) for October 2006 and (b) for October 2015. Subplots (c) and (d) are the same as
 379 subplots (a) and (b) but for the MERRA-2 reanalysis vertical pressure velocity. Positive (negative)
 380 values represent the downward (upward) wind.

381 One of the worst fire events in Indonesia’s history occurred in October 1997 and was
 382 associated with an El Niño event and a positive IOD (Duncan et al., 2003a). In order to see any
 383 similarities between 1997 and 2006 and 2015, particularly in large-scale circulations, we further
 384 checked the MERRA-2 GpH and wind circulation pattern in October 1997. To note, none of the
 385 satellite measurements of CO (either AIRS or MOPITT) are available during the 1997 event, and
 386 measurements at LABS didn’t start until 2006. Hence, we only cross-checked the large-scale
 387 circulations that are observed in October 1997, respectively. Sup. Figure 4a shows the spatial
 388 distribution of GpH observed at 500 hPa and Sup. Figure 4b shows the vertical-meridional cross-
 389 section of vertical pressure velocity averaged along 110°–130°E in October 1997. Interestingly,
 390 we noticed a significant high-pressure anti-cyclonic circulation over northern parts of the SCS in
 391 October 1997 as observed in 2006 and 2015. Also evident is the merging of the southerlies from
 392 the MC with the subtropical westerlies in the northern PSEA region. Similarly, the vertical

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393 pressure velocity also shows the upward wind over the MC and the downward wind over the
394 northern hemisphere subtropics around 20–30°N latitude. From Sup. Figure 4b, it is very clear that
395 a significant descending wind was evident around the LABS location in 1997 similar to 2006 and
396 2015. Overall, it is very clear from the present results that it is possible to transport pollution from
397 Indonesia to sub-tropical East Asia during extreme and higher-duration fire events like 1997, 2006,
398 and 2015.



399
400 **Figure 11.** Schematic diagram of CO transport from Indonesian fires to subtropical East Asian
401 region. The horizontal transport of CO due to the high-pressure anticyclone is denoted by the blue-
402 colored arrow. H denotes a high-pressure anticyclone over northern parts of the South China Sea.
403 Magenta-colored arrows indicate the transport of CO through the local Hadley circulation (over
404 110°–130°E). Black-colored star symbol represents the LABS location.

405 The major transport pathways of CO from Indonesia to subtropical East Asia are illustrated in a
406 schematic diagram (Figure 11). Illustrated mechanisms include horizontal transport in the free
407 troposphere due to El Niño and positive IOD-induced high-pressure anticyclone circulation, and
408 vertical transport through the Hadley circulation. The southerlies on the southwest flank of the
409 anticyclone merged with the subtropical westerlies over PSEA and then transported polluted air to
410 LABS. Apart from this horizontal transport, CO was transported through the Hadley circulation to
411 LABS in both events. However, there was a distinctly different Hadley Circulation strength in
412 2006 compared to 2015 due to the different El Niño conditions. These two events were strongly
413 associated with positive IOD, but in 2006, the El Niño conditions were not well-developed,
414 whereas in 2015 well-developed El Niño conditions were evident. These El Niño conditions
415 further suppressed the HC over the western Pacific in 2015 compared to 2006. This suggested the
416 importance of the background climate conditions (ENSO and IOD) on the pollutant transport
417 process.

419 **4. Summary and Conclusions**

420 Due to the combined impact of positive phase IOD and El Niño-induced drought
421 conditions in 2006 and 2015, Indonesia experienced extreme fire activity. MODIS active fire
422 counts showed the largest fires in October 2006 and 2015 compared to the other years in the 16-
423 year period in Indonesia. These record fires reflected two of the largest carbon emissions in the
424 Indonesian region since 1997.– Changes in the background climate will inevitably impact
425 meteorological transport processes and the concentrations of pollutants arriving at downwind
426 regions.–Lulin Atmospheric Background Station (LABS, 23.47°N 120.87°E, 2862 m ASL) is the
427 only high-altitude background station located in the western North Pacific region, and is optimally
428 located to study some of these transport processes, including long-range transport of pollution in
429 the free troposphere and stratospheric intrusions. Interestingly, during these two events (October
430 2006 and 2015), we noticed an abnormal enhancement of CO compared to other years at LABS
431 from the in-situ measurements. In the present study, for the first time, the impact of Indonesia fire
432 pollution on CO measurements at LABS and the plausible transport pathways for the transport of
433 CO from Indonesia to sub-tropical East Asia were investigated. The main findings are summarized
434 below:

435 1. Compared to 16-year (2006-2021) means, a substantial increase in CO mixing ratios of
436 about ~47.2 ppb (37.2%) in October 2006 and ~36.7 ppb (28.9%) in October 2015 was
437 observed at LABS.

438 2. By comparing the CO and atmospheric large-scale circulation data, we found two plausible
439 transport pathways of CO from Indonesia to LABS. i.e. horizontal transport in the free
440 troposphere and vertically through the Hadley Circulation.

441 3. El Niño and positive IOD-induced high-pressure anticyclone circulation over northern
442 parts of the South China Sea play an important role in the horizontal transport of CO.

443 4. Distinct strength of the Hadley circulation over the western Pacific was observed in
444 October 2006 (stronger) and 2015 (weaker). Well-developed El Niño conditions in
445 October 2015 suppressed the strength of the Hadley Circulation over the western Pacific.

446 A changing warmer climate can influence carbon emissions and alter the transport pathways, hence
447 impacting the various scales of air pollution and climate. Changes in the background climate will
448 inevitably impact meteorological transport processes and the concentrations of pollutants arriving
449 at downwind regions. Overall, the present results further provide knowledge to the atmospheric
450 chemistry community about the different transport pathways of pollutants and the role of climate
451 conditions.

452
453 ~~During October 2006 and 2015, there were substantial increases in CO mixing ratios, ~47.2~~
454 ~~ppb (37.2%) and ~36.7 ppb (28.9%) increase compared to the 16 year (2006-2021) means at~~
455 ~~LABS. Interestingly, these two events (2006 and 2015) were strongly associated with the two~~
456 ~~major biomass burning episodes over Indonesia, which resulted from a combined impact of~~
457 ~~positive phase ENSO and IOD induced drought conditions. MODIS active fire counts showed the~~
458 ~~largest fires in October 2006 and 2015 compared to the other years in the 16 year period in~~
459 ~~Indonesia. These record fires reflected two of the largest carbon emissions in the Indonesian region~~
460 ~~since 1997. Apart from these high values in October 2006 and 2015, in October 2016, extremely~~
461 ~~low CO values were recorded at LABS (~31.4% lower compared to 2006-2021 mean). October~~
462 ~~2016 was associated with negative IOD over the Indian Ocean and La Niña in the Pacific Ocean,~~
463 ~~resulting in the lowest fire activity over the MC. Further, we found that the large scale circulations~~
464 ~~in 2016 were quite different from 2006 and 2015. In 2016, LABS was dominated by the southerlies~~

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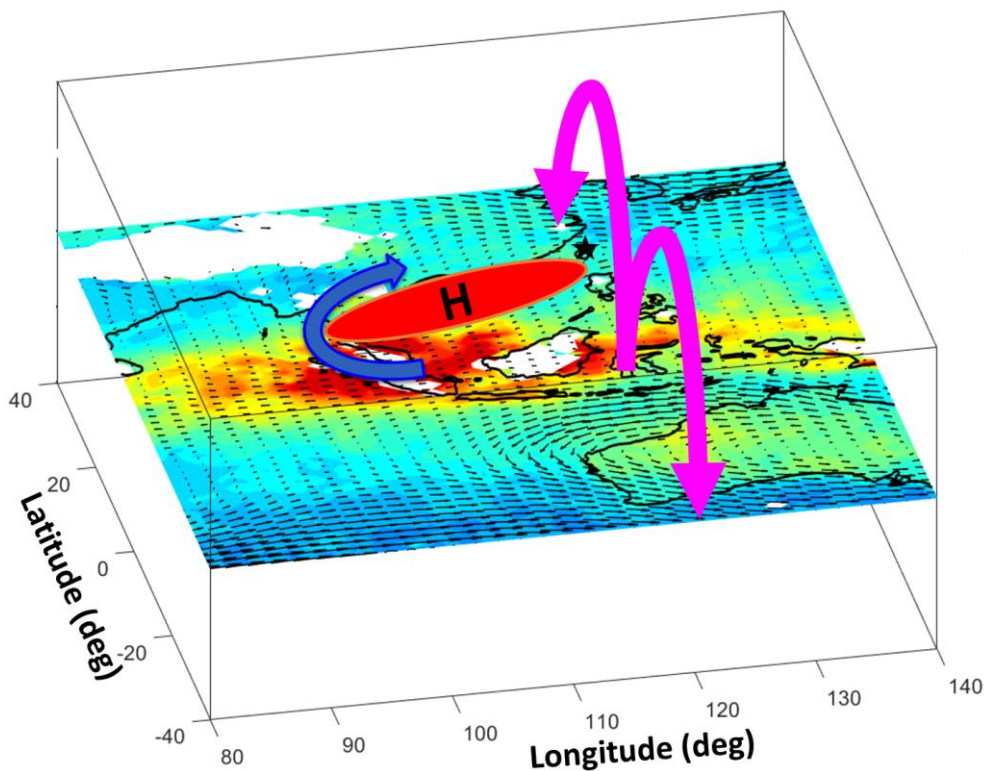
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465 due to the western Pacific subtropical High (WPSH), which transported clean marine air from the
466 Pacific Ocean and caused record low CO values at LABS. The main aim of our study was to
467 examine the transport pathways of CO from Indonesia's source region to the downwind LABS
468 region. By comparing the CO and atmospheric circulation data from the 2006 and 2015 El Niño
469 (positive IOD) years and 2016 La Niña (negative IOD), we found two plausible transport pathways
470 of CO from Indonesia to LABS.

471 **Figure 10** illustrates a schematic diagram of the major transport pathways of CO from
472 Indonesia to subtropical East Asia during the two event years. They include horizontal transport
473 in the free troposphere due to El Niño induced high pressure anticyclone circulation and vertically
474 through the Hadley circulation. For October 2006 and 2015, corresponding to El Niño and positive
475 IOD, northern SCS was influenced by the high pressure anti-cyclonic system in the free
476 troposphere. The southerlies on the southwest flank of the anticyclone further merged with the
477 subtropical westerlies over PSEA and then transported polluted air to LABS. Apart from this
478 horizontal transport, CO was transported through the Hadley circulation to LABS in both events.
479 However, there was a distinctly different HC strength in 2006 compared to 2015 due to the
480 different El Niño conditions. These two events were strongly associated with positive IOD, but in
481 2006, the El Niño conditions were not well developed, whereas in 2015 well-developed El Niño
482 conditions were evident. These El Niño conditions further suppressed the HC over the western
483 Pacific in 2015 compared to 2006. This suggested the importance of the background climate
484 conditions (ENSO and IOD) on the pollutant transport process. A changing warmer climate can
485 influence carbon emissions and alter the transport pathways, hence impacting the various scales of
486 air pollution and climate. Overall, the present results further provide knowledge to the atmospheric
487 chemistry community about the different transport pathways of pollutants and the role of climate
488 conditions.



489
 490 **Figure 10.** Schematic diagram of CO transport from Indonesian fires to subtropical East Asian
 491 region. Horizontal transport of CO due to the high pressure anticyclone is denoted by the blue-
 492 colored arrow. H denotes high pressure anticyclone over northern parts of the South China Sea.
 493 Magenta colored arrows indicate the transport of CO through the local Hadley circulation (over
 494 110°–130°E). Black colored star symbol represents the LABS location.

495 **Data availability**

496 The CO data at LABS can be assessed at http://lulin.tw/index_en.htm. The AIRS and MOPITT
 497 CO data can be downloaded from the following websites
 498 https://disc.gsfc.nasa.gov/datasets/AIRS3STM_7.0 (AIRS project., 2019) and
 499 <https://asdc.larc.nasa.gov/project/MOPITT>. MERRA-2 data are available online through the

500 NASA Goddard Earth Sciences Data Information Services Center (GES DISC;
501 <https://disc.gsfc.nasa.gov>, last access: 30 May 2022). Nino 3.4 Index and IOD data can be
502 downloaded through the following websites https://psl.noaa.gov/gcos_wgsp/Timeseries/Niño34/.
503 https://psl.noaa.gov/gcos_wgsp/Timeseries/DMI/. The MODIS fire products can be downloaded
504 from the following website https://firms.modaps.eosdis.nasa.gov/active_fire/.

505 **Author contributions**

506 **Saginela Ravindra Babu:** Conceptualization, Data curation, Formal analysis, Investigation,
507 Software, Validation, Visualization, Writing – original draft preparation, Writing – review and
508 editing; **Chang-Feng Ou-Yang:** Data curation, Software, Validation, Visualization; **Stephen M.**
509 **Griffith;** Writing – review and editing; **Shantanu Kumar Pani:** Data curation and Visualization;
510 **Steven S. Kong:** Data curation and Visualization; **Neng-Huei Lin:** Conceptualization,
511 Investigation, Funding Acquisition, Supervision, Resources, Writing – review and editing.

512 **Competing Interest**

513 The authors declare that they have no conflict of interest.

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519 AIRS satellite data. We thank NASA's Global Monitoring and Assimilation Office (GMAO) for
520 providing the Modern-Era Retrospective analysis for Research and Applications, Version 2
521 (MERRA-2) data. We also thank NOAA ESRL Physical Sciences Laboratory for providing Indian
522 Ocean Dipole and Niño 3.4 index values through the following websites
523 https://psl.noaa.gov/gcos_wgsp/Timeseries/DMI/
524 https://psl.noaa.gov/gcos_wgsp/Timeseries/Niño34/.

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744 **Table 1.** Detailed statistics of observed CO in October during 2006 to 2021 at LABS.

Year	Mean	Median	Standard Deviation	Change in CO (%)	Total data points
2006	175.8	174	51	33.9	703
2007	155.3	140	63.4	18.3	732
2008	125.5	125	26.9	-4.4	599
2009	127.1	125	35.5	-3.2	533
2010	143.9	136	38.1	9.6	739
2011	137.1	137	41.9	4.4	734
2012	155.8	153	39.4	18.7	643
2013	146.8	141	35.7	11.8	365
2014	125.6	120	39.8	-4.2	602
2015	164.8	163.5	46.2	25.6	732
2016	91.6	87	20.9	-30.2	732
2017	109.7	100.3	32.4	-16.4	744
2018	147.7	149.9	29.1	12.5	736
2019	142.4	142.8	37.7	8.5	742
2020	121.3	113.8	29.5	-7.5	742
2021	107.7	104.6	26.9	-17.9	744

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