O3-precursor relationship over multiple patterns of time scale:

A case study in Zibo, Shandong Province, China

 the four patterns of time scale, which was mainly due to the embedded uncertainty in model input dataset when averaging individual daily pattern into different timescales. This implies that utilizing narrower time scale (i.e., daily pattern) is useful to derive reliable and robust O3-precursor relationship. Our results highlight the importance of quantifying the impact of different time scales to constrain the photochemical regime, which can formulate more accurate policy-relevant guidance for O³ pollution control.

1 Introduction

46 Since 2013, the ambient PM_{2.5} concentration in China has dramatically declined by implementing Clean Air Action (Lu et al., 2018; Wang et al., 2020b; Zhang et al., 2019). However, national ground surface ozone concentrations increased over the same period (Xue et al., 2020) and became a major air quality problem that needed to be addressed in China (Li et al., 2019; Wang et al., 2019). It is well-known that ground surface ozone is formed mainly by complex nonlinear photochemical oxidation of 52 volatile organic compounds (VOCs) in the presence of nitrogen oxides (NO_x = NO + NO2) and sunlight (Blanchard, 2000; Hidy, 2000; Kleinman, 2000), which adversely influences human health, vegetation and corps (Brunekreef and Holgate, 2002; Vingarzan, 2004).

 Given the complex non-linear relationship between $O₃$ formation and its 57 precursors (VOCs and NO_x), challenges in mitigating its severity lie primarily in comprehensively understanding of O3-precursor relationship (Su et al., 2018a; Tan et al., 2018a). It is commonly recognized that regional-scale air quality models and the 0- D box model are two mainstream approaches to investigate the increasingly severe ozone problem (Blanchard, 2000; Cardelino and Chameides, 1995; Hidy, 2000; Liu et al., 2019). Unlike the complicated 3-D air quality models, the 0-D box model is an observation-based model that implemented with gas-phase chemical mechanism, and 64 has been widely used to diagnose O_3 -precursor relationship in various locations (Liu et al., 2021a; Sun et al., 2016; Tan et al., 2019b; Xue et al., 2014a; Yu et al., 2020a). Some previous studies (Li et al., 2021; Lu et al., 2010a; Sicard et al., 2020; Yu et al., 2020b) 67 have reported a large variability of O_3 -precursor relationship in spatiotemporal scales in many cities of China, which indicates great challenges in current O_3 pollution control (Wang et al., 2017a; Xue et al., 2014b).

 Table 1 summarizes the published studies of O3-precursor relationship using the 0-D box model (implemented with different gas-phase chemical mechanisms) at diversified patterns of time scale in many places of China. The observational period in most previous studies was short-term (i.e., less than one month), while medium-term (i.e., from one to several months), and long-term (i.e., multiple years) periods were limited. As shown in **Table 1**, we find that model input datasets with different timescales have been employed in previous studies to identify the campaign-averaging O₃ formation regime, but there is a lack of comparison among these different timescales. We also find that more than half of the studies using the averaged diurnal patterns as box model input, which is particularly common for those medium and long-term measurements. For example, a 10 years long-term observational study by Wang et al., (2017a) adopted monthly pattern of time scale for model simulation with the reason of saving computing resources, and it also revealed a substantial temporal variability of O3-precursor relationship. In addition, it is believed that long-term (measurements of at least several months) and multiple-site continuous online measurements can provide opportunity to develop O³ control strategy more comprehensively over a wider spatiotemporal scale (Li et al., 2021; Wang et al., 2017b; Wang et al., 2017b). However, such measurements have been quite rare in China, limiting the present understanding of O3-precursor relationship (Lu et al., 2019; Wang et al., 2017b).

 In this study, a five-month field campaign was conducted in the summer of 2019 to investigate the ozone formation chemistry at 3 sites in Zibo, a major prefecture-level Chinese city in Shandong province. According to our measurements at the three sites in 92 Zibo, the averaged O_3 concentration during the whole observational period was around 50 ppbv, while the daily maximum of O³ concentrations for some extremely polluted periods were nearly 120-150 ppbv (see details in **Section 3.1**). Here we developed an approach that integrated multiple patterns of time scale for box model simulation, which 96 aimed at illustrating the non-linearity of O_3 -precursor relationship driven by its actual daily / weekly / monthly variability. Our results can be conducive to interpreting variations of O3-precursor relationship over a wider spatiotemporal scale, and they 99 provide implications for developing more precise and constrained O_3 control strategies 100 in other regions.

2 Methods

2.1 Study sites and measurements

 Field measurements were conducted in a major prefecture-level city (Zibo), which is in the middle of Shandong Province, northern China, from 1 May to 30 September, 2019. **Figure S1** shows the surrounding environment and geographical locations at the three sampling sites; a detailed description of the Tianzhen (TZ), Beijiao (BJ) and Xindian (XD) sites can be found in our previous study (Li et al., 2021). Briefly, TZ contains a mixture of crude oil processing and operation stations and farming areas, and is classified as suburban area; XD contains a mixture of residential and heavy industrial zones, and is considered as a suburban area; BJ is in the urban area of Zibo.

111 Typical inorganic gases of O₃, NO, NO₂, CO and SO₂ were measured using online commercial gas analysers (Thermo Scientific 49i, 42i, 48i and 43i, USA) at the three sites. Following the Chinese meteorological monitoring regulation (GB/T 35221-2017), we continuously monitored the meteorological parameters (i.e., temperature, relative humidity, UV-A solar radiation, precipitation, wind speed, and wind direction) at the three sites (Li et al., 2021). Two online GC systems (gas chromatography–flame ionisation detector, GC-FID, Thermo Scientific GC5900) were deployed at TZ and BJ respectively to measure VOC species. For C2-C⁵ VOCs, desorption and separation were performed using a GC with pre-concentration on a combination of two columns, 120 followed by a FID detector. For C_6-C_{12} VOCs, air sample was pre-concentrated on Tenax GR cartridges and subsequently separated by chromatographic column, then detected by another FID detector. Similarly, one online system (gas chromatography– flame ionisation detector/photoionisation detector, GC-FID/PID, Syntech Spectras GC 124 955-615/815) was deployed at XD site. For C_2-C_6 VOCs, the hydrocarbons were concentrated on a Tenax GR carrier, then thermally desorbed and separated on a DB-1 126 column, and finally detected by FID and PID detectors. For C_6-C_{12} VOCs, the air sample was concentrated on a Carbosieves SIII carrier at 5℃, then thermally desorbed and separated on a combination of two columns, and FID and PID detectors were employed for subsequent detection. These systems measured 55 VOC species at a 1-h resolution, and more detailed descriptions can be found elsewhere (Chien, 2007; Jiang et al., 2018; Xie et al., 2008).

 Table S1 summarized the limit of detection, accuracy, precision of the instruments at the three sites, and all the measurement instruments were regularly subjected to the service of checking and maintenance during the whole campaign. Unfortunately, we did not conduct the inter-comparison between the GC-FID and GC-FID/PID instruments at the same site due to practical reasons, as these VOC instruments were separately deployed at the three different sites for continuous routine operation. To ensure the quality assurance / quantity control (QA/QC) of online VOC measurement, two five- point calibrations (i.e., 2, 4, 6, 8, 10 ppbv, dilution from one cylinder) for standard gases with 55 VOC species (Linde Co., Ltd, USA) were carried out in May and August of 141 2019 at the three sites. **Table S2** showed that the calibration linearity (R^2) of all measured VOCs were nearly 0.9990. Additionally, a single-point calibration (i.e., 6 ppbv) was regularly performed every month during the whole campaign. As shown in **Figure S2** (a case from TZ), the retention time, peak fitting and baseline of the chromatogram were manually checked and adjusted on a daily basis.

2.2 0-D box model and design of four patterns of time scale

 The 0-D box model integrated with the latest Master Chemical Mechanism of MCMv3.3.1 [\(http://mcm.york.ac.uk;](http://mcm.york.ac.uk/) last access: 27 January 2023) has been widely utilized in many regions (He et al., 2019; Jenkin et al., 2015; Liu et al., 2019; Whalley et al., 2021). Unlike the lumped chemical mechanisms such as CB05 (Wang et al., 2017a; Yarwood et al., 2005), CB6 (Yarwood et al., 2010), RACM/RACM2 (Goliff et

 al., 2013; Stockwell et al., 1997, 2020) and SAPRC-07 (Carter, 2010), the MCMv3.3.1 is a near-explicit chemical mechanism consisting of over 5,800 species and 17,000 reactions (Jenkin et al., 2015; Saunders et al., 2003), which can be used to describe the gas-phase chemistry (i.e., in-situ photochemistry). In this study, the box model (based on the Framework for 0-D Atmospheric Modeling, F0AM) (Wolfe et al., 2016) was applied and constrained by the mean diurnal profiles of meteorological data (i.e., 158 temperature, relative humidity, and photolysis rates), 4 inorganic gases (i.e., SO₂, CO, NO, and NO2), and 45 speciated VOCs (in MCMv3.3.1 species list; see **Table S3**). Since measured photolysis rates (*J* values) were not available, the measured UV-A solar radiation was used to scale the photolysis rates calculated from the Tropospheric Ultraviolet and Visible Radiation model (TUVv5.2; [https://www.acom.ucar.edu/Models/TUV/Interactive_TUV;](https://www.acom.ucar.edu/Models/TUV/Interactive_TUV/) last access: 27 January 2023) following the approach of recent studies (Lyu et al., 2019; Lyu et al., 2016). Specifically, the geographical coordinates, date and time were initialized into the TUV model to derive photolysis rates and solar radiation. We obtained the scaling factor by comparing the observed with modeled solar radiation, and used this scaling factor to 168 scale the TUV model derived photolysis rates. A dilution rate of 3/86400 s⁻¹ was applied for all non-constraint species and simulation days through a stepwise sensitivity test by 170 adjusting it from $1/86400 \text{ s}^{-1}$ to $5/86400 \text{ s}^{-1}$ (see details in **Text S1**) for the best 171 reproduction of O_3 . For each model run (i.e., each daily model simulation), it was performed on a daily basis with intervals of 24 hours spanning from 0:00 to 23:00, and each individual model simulation was run to reach one-day diurnal steady state. The detailed descriptions of box model operation were provided in our previous study (Li et al., 2021).

176 Since the box model simulations are conducted with intervals of 24 hours spanning from 0:00 to 23:00 local standard time (Wang et al., 2018), the entire campaign observations were taken into four patterns of time scale (i.e., five-month, monthly, weekly, and daily) as diurnal average format for model input (**Figure 1**). Note that some days or weeks were not modeled due to some missing data in the measurements. Nevertheless, the total simulation number at the daily (i.e., 100, 81, and 114 days for TZ, BJ and XD respectively) or weekly (i.e., 21, 20, and 19 weeks for TZ, BJ, and XD respectively) scale was representative of the five-month campaign. Specifically, the entire campaign data classified as four patterns of time scale were modeled as base runs. Then we performed the sensitivity modeling to calculate the relative incremental reactivity (RIR) of precursors by adjusting the input concentrations in the base runs (see next section) (Lu et al., 2010a).

188 **2.3** Calculation of net O_x production rate $P(O_x)$ and Relative incremental 189 **reactivity (RIR)**

190 Considering the rapid chemical titration of NO to $NO₂$ in the presence of $O₃$, the 191 concept of 'total oxidant' $(O_x = O_3 + NO_2)$ has been widely used to represent the actual 192 photochemical production of O_3 (Lu et al., 2010). Similar to those described in previous 193 studies using the 0-D box model (He et al., 2019; Lyu et al., 2016), the net or in-situ O_x 194 production rate $(P(O_x))$ is defined as the difference between the O_x gross production 195 rate $(G(O_x))$ and the O_x destruction rate $(D(O_x))$, which is formulated in accordance 196 with Eq. (1):

197
$$
P(0_x) = G(0_x) - D(0_x)
$$
 (1)

198 The O_x gross production rate $(G(O_x))$, or the total chemical production of O_x , is 199 calculated by summing the rates of oxidation of NO by $HO₂$ and RO₂ radicals in 200 accordance with Eq. (2):

201
$$
G(O_x) = k_{HO_2 + NO} [HO_2][NO] + \sum k_{RO_{2,i} + NO} [RO_{2,i}][NO]
$$
 (2)

202 The O_x destruction rate $(D(O_x))$, or total chemical loss of O_x , is calculated by 203 summing O_3 photolysis, the reaction of O_3 with OH, HO_2 and alkenes, as well as the 204 reaction between $NO₂$ and OH, as described by Eq. (3):

205
$$
D(O_x) = k_{O^1D + H_2O}[O^1D][H_2O] + k_{OH + O_3}[OH][O_3] + k_{HO_2 + O_3}[HO_2][O_3] + k_{\text{alkenes} + O_3}[\text{alkenes}][O_3] + k_{OH + NO_2}[OH][NO_2]
$$
\n(3)

 Concentrations of radicals and intermediates are obtained from the outputs of the 0-D box model. The *k* values in Eq. (2) and (3) represent the rate constants of the corresponding reactions, respectively. The subscript '*i*' in Eq. (2) represents the 210 individual RO₂ species.

211 Additionally, relative incremental reactivity (RIR) has been widely used as a 212 metric to quantify the O₃-precursor relationship, and it can be derived from the 0-D box 213 model (MCMv3.3.1) by changing the input mixing ratios of its precursors (Sillman, 214 2010; Xue et al., 2014a). The RIR is defined as the ratio of percentage change in net O_x 215 $(O_x = O_3 + NO_2)$ production rate $P(O_x)$ (Li et al., 2021) to percentage change of 216 concentration of precursor X. The RIR of a specific precursor X is described in Eq. (4) :

$$
RIR(X) = \frac{[P\text{O}_X(CX) - P\text{O}_X(CX - \Delta CX)]/P\text{O}_X(CX)}{\Delta CX/CX}
$$
(4)

218 Here, X is a specific precursor (i.e., NO_x , CO or grouped / individual VOC species), 219 CX is the measured concentration of precursor X, and Δ CX is the hypothetical 220 concentration change $(\Delta CX/CX = 10\%$ in this study in accordance with the previous 221 studies (Lyu et al., 2016; Wang et al., 2018)). $PO_x(CX)$ represents the simulated O_x 222 production rate in a base run, whereas $PO_X(CX-\Delta CX)$ is the simulated O_X production 223 in a second run with a hypothetical concentration change of species X. Obviously, a 224 higher positive value of $RIR(X)$ suggests a more effective way of reducing the ambient

225 O_3 production rate by reducing X (Ling et al., 2011; Zhang et al., 2008a).

226 In this study, the O_3 precursors were divided into four major categories, including 227 anthropogenic VOC (AVOC), biogenic VOC (BVOC, only isoprene in this study), CO 228 and NO_x (Tan et al., 2019b). AVOC was further divided into three subcategories: alkanes, 229 aromatics and alkenes* (the asterisk denotes anthropogenic alkenes, excluding isoprene 230 in this study) (Yu et al., 2020a). As mentioned, RIR method was applied mainly to 231 evaluate the O_3 -NO_x-VOC sensitivity and determine the photochemical regimes among 232 four patterns of time scale. Thus, we calculated the RIR values of major precursor 233 groups (i.e., AVOC, BVOC, CO, NO_x , alkanes, alkenes^{*} and aromatics) to further 234 quantify the O_3 -precursor relationship.

235 In general, O₃ formation chemistry is usually classified into three regimes (i.e., 236 VOC-limited, transitional and NO_x -limited) (He et al., 2019; Wang et al., 2018). In this 237 study, RIR_{NOX}/RIR_{AVOC} (the ratio of two RIR values) was used as a metric to classify 238 the photochemical regimes (Li et al., 2021). Specifically, RIR_{NOx}/RIR_{AVOC} value of less 239 than 0.5 was defined as VOC-limited regime, greater than 2 as NO_x -limited regime, and 240 from 0.5 to 2 as transitional regime (see **Text S2** and **Table S4**) (Li et al., 2021).

241 **3 Results and discussion**

242 **3.1 Overview of the field campaign**

 Figure 2 shows the time series of measured meteorological parameters and O³ as well as its precursors at the three sites during the whole campaign. In general, the temperature (*T*) and relative humidity (RH) were basically consistent at the three sites, while the wind speeds were different, which suggests that the three sites had an overall consistent meteorological condition. In addition, the time series of UV-A radiation was shown in **Figure 2d**, which was only available from one urban site of Zibo but expected to represent the whole Zibo city in this study. Following the protocol of the previous studies (Lyu et al., 2019; Wang et al., 2017b; Xue et al., 2014), the time series of 251 photolysis rates (e.g., J_{NO2} (**Figure 2e**) and $J_{O¹D}$ (**Figure 2f**)) were calculated from TUVv5.2 model and further scaled from UV-A radiation measurement.

253 As shown in **Figure 2g**, we found that severe O³ pollution was observed at the 254 three sites throughout the whole campaign. According to our measurements at the three 255 sites in Zibo, the averaged O_3 concentration during the whole observational period was 256 around 50 ppby, while the daily maximum of $O₃$ concentrations for some extremely 257 polluted periods were nearly 120-150 ppby (**Figure 2g**). Interestingly, the O_3 258 concentrations at the three sites were generally consistent, while the levels of its 259 precursors (e.g., VOC, NOx) were obviously different (**Figure 2h-k**), which implies the 260 site-to-site variation of O_3 formation chemistry for the whole Zibo city.

261 Generally, OH reactivity (or OH loss rate, k_{OH}) is widely applied to quantify the 262 capacity of OH consumption by VOCs (Tan et al., 2019a). According to **Table S3**, the

263 BVOC reactivity $(k_{\text{BVOC}}, 3.5 \pm 4.1 \text{ s}^{-1})$ in TZ were highest among the three sites. As BJ was mainly influenced by the emission from urban region, it showed the highest AVOC 265 reactivity (k_{AVOC} , 6.8 ± 6.3 s⁻¹) and NO_x level (31.1 \pm 28.6 ppbv). In addition, XD 266 showed the highest level of alkenes* reactivity of 4.0 ± 3.2 s⁻¹ within the three sites, and the local petrochemical industry nearby XD area may explain such characteristic (Li et al., 2021).

3.2 Evaluation of box model performance

270 The measured O_3 concentrations were not constrained in our MCMv3.3.1 box model calculation, thus the model performance could be quantitatively assessed by 272 comparing the modeled O_3 (from base runs) with the measured O_3 . **Figure S3-S8** show 273 the time series of simulated and observed O_3 concentrations at four patterns of time scale. In most cases, the box model simulation could accurately capture the level and 275 variation trend of the observed O_3 . However, on some days the modeling results 276 underestimated or overestimated the O_3 concentrations, particularly the 277 underestimation of nocturnal O_3 concentrations. Such discrepancies between the 278 simulated and observed O_3 were likely due to limitations in explicit representations of atmospheric and transport processes (i.e., the horizontal and vertical transport process of ground ozone) by 0-D modeling approach (Lyu et al., 2019; Yu et al., 2020b). Specifically, ozone simulated by the 0-D box model is considered as in-situ 282 photochemical processes from its precursors. Unlike the 3-D air quality model, 0-D box model usually simplifies the representation of the physical processes (i.e., deposition and advection) (Lu et al., 2010a; Sillman, 2010). Note that some adjustable parameters (e.g., radiation scheme, dilution rate) were remained consistent in all of our model calculations, which ensured the comparability of model results to the greatest extent.

 The index of agreement (*IOA*) (Li et al., 2021; Lyu et al., 2016), Pearson's correlation coefficient (*r*) and root mean square error (*RMSE*) were jointly used as statistical metrics to quantify the goodness-of-fit between the simulated and observed O³ concentrations. **Table S5** summarizes these statistical metrics for each site at various patterns of time scale. Because any single statistical metric has its own limitations, using these three indicators conjointly provided a more comprehensive evaluation of the model performance (Su et al., 2018b). Generally, higher *IOA* and *r* as well as lower *RMSE* indicate better agreement between the simulated and observed values (Wang et al., 2018; Willmott, 1982). As shown in **Table S5**, slightly reduced correlation was observed as the time scale changed from the wider (i.e., five-month scale) to the narrower (i.e., daily scale) pattern, which is understandable because of the enlarged statistical samples in the narrower pattern of time scale.

 In summary, TZ showed the best performance of the box model simulation, followed by XD and BJ, regardless of any statistical metrics or different patterns of time scale, which may be associated with the optimized dilution rate for non-constraint species in model configuration. The overall model performance in this study (i.e., a day- to-day *IOA* of approximately 0.90 for TZ) was close to or slightly better than those reported in previous studies, such as *IOA* = 0.74 in Hong Kong (Liu et al., 2019), *IOA* = 0.74 in Wuhan (Lyu et al., 2016) and *IOA* = 0.90 in Jiangmen (He et al., 2019). According to the above evaluation of base runs, our modeled results were acceptable for the subsequent O3-precursor relationship analysis described in the following sections.

3.3 Month-to-month

 Figure 3a-b presents the monthly RIR values of the major precursor groups at 311 each site, and the large variability of O_3 -precursor relationship at spatiotemporal scale (i.e., site-to-site and month-to-month) was observed. Specifically, in most months, XD generally showed the highest RIRAVOC among the three sites, followed by BJ and TZ. 314 In addition, RIR_{BVOC} showed similar level to RIR_{AVOC} in TZ, but much less than RIRAVOC in BJ and XD, which can be explained by the observed higher BVOC reactivity in TZ than the other two sites (see **Figure S9** and **Table S3**). Also, almost all 317 the precursor groups showed positive RIR values, except negative RIR_{NOx} appeared in 318 BJ and XD in September. In addition, the RIR_{CO} values at the three sites suggested its limited role in O³ formation at the three sites, compared with other major categories of O³ precursors. Among the three subcategories of AVOC, alkenes* always had the highest RIR values, followed by aromatics, while the contribution of alkanes to $O₃$ formation can be ignored due to their near-zero RIR values. That sequence of O₃-AVOC sensitivity (alkenes* > aromatics > alkanes) indicated by the RIR analysis was consistent with previous studies in some other Chinese cities (Su et al., 2018b; Tan et al., 2019b). Significant monthly variations of O3, NOx, CO, VOC reactivity and TVOC/NO_x ratios (in ppbC/ppbv, as a widely used simple metric to determine the photochemical regime) (National Research Council, 1991) were also observed from May to September (see **Figure S9** and **Table S3**) at the three sites. For example, the BVOC reactivity in TZ showed highest level among the three sites during the whole campaign, and the AVOC reactivity in BJ showed more considerable variations in different months, which indicated spatial and temporal variations of local primary emission for O³ precursors in Zibo city.

Figure 3c shows monthly RIR_{NOX}/RIR_{AVOC} at each site, which clearly reveals the spatial and temporal variations in photochemical regimes. For instance, the photochemical regime at the TZ site was considered to be transitional regime in May, NOx-limited regime in June and July, and VOC-limited regime in August and September; whereas for a specific month like June, NOx-limited, VOC-limited, and transitional regimes were generally identified for TZ, BJ, and XD respectively. **Figure**

339 **5b** shows good consistency between monthly TVOC/NO_x and $\text{RIR}_{\text{NOx}}/\text{RIR}_{\text{AVOC}}$, suggesting that the changes of local emissions for O_3 precursors may partially explain the considerable variation of O_3 formation chemistry in different months.

3.4 Week-to-week

 Figure 4 shows the time series of week-to-week RIR values of major precursor 344 groups and RIR_{NOX}/RIR_{AVOC} at three sites in Zibo. Compared with month-to-month results, **Figure 4** further reveals the O3-precursor relationship with more information in temporal trends. The temporal variations in weekly RIR_{AVOC} at the three sites generally decreased and then increased, whereas weekly RIR_{NQx} represented an opposite temporal 348 variation during the entire campaign. Additionally, weekly RIR_{BVOC} showed a trend of first decrease and then increase at TZ, while it did not show clear temporal variation at BJ and XD due to low values (**Figure 4a-c**). In general, RIRalkanes, RIRalkenes* and RIRaromatics showed a tendency consistent with that of the RIRAVOC at three sites (**Figure 4d-f**). Overall, these phenomena were consistent among the three sites, though the 353 magnitude of RIR values varied site-to-site. In parallel, the temporal changing of $O₃$ 354 precursor (e.g., AVOC, NO_x) was also observed at the three sites during the entire campaign (see **Figure S10**). For example, the weekly NO^x concentration showed an overall trend of first decrease and then increase, while the AVOC reactivity showed a different temporal variation. Given the moderate correlation between weekly TVOC/NO_x and RIR_{AVOC}/RIR_{NOx} (**Figure 5c**), the temporal variations of RIR values and O_3 formation chemistry at the three sites may be partially elucidated by the emission changes of O³ precursors.

 As shown in **Figure 4g-i**, all the three sites showed similar temporal trends of RIR_{NOx}/RIR_{AVOC}, as it increased first and then decreased, though the magnitude of RIRNOx/RIRAVOC varied largely at each site. Such site-to-site variability of RIR_{NOx}/RIR_{AVOC} suggests that the photochemical regime in a local scale was mainly influenced by local emissions. By contrast, the site-to-site synchronization in temporal 366 trend of $RIR_{\text{NOx}}/RIR_{\text{AVOC}}$ suggests that the photochemical regime in a local scale may also be influenced by the emissions in a regional area. Therefore, the long-term, week- to-week $\text{RIR}_{\text{NOX}}/\text{RIR}_{\text{AVOC}}$ of multiple sites can further reflect the variability of ozone formation regime at a large geographic scale.

3.5 Day-to-day

 In this section, O₃-precursor relationship at the narrowest pattern of time scale was identified in detail. **Figure S11-S12** shows the time series of daily RIR values at three sites in Zibo, where the temporal trend of RIR values was consistent with that at weekly 374 scale (**Figure** 4). Additionally, the time series of daily RIR_{NOx}/RIR_{AVOC} (**Figure S13**) showed more irregular variations in temporal trends during the entire campaign, though

 such temporal trends were overall consistent with that of weekly scale in **Figure 4 g-i**. In summary, the time series of RIR values from the daily scale can provide more informative variations and characteristics of O3-precursor relationship in temporal trends.

 Table 2 summarizes the number of days and proportions that were classified into the three photochemical regimes across each site and each pattern of time scale. Near-382 consistent proportions of O_3 formation regimes (using RIR_{NOx}/RIR_{AVOC} as a metric) were shown among multiple patterns of time scale, whereas a variability of proportion occurred among the three sites. The proportions of photochemical regimes changed accordingly along with the time scale varied from wider to narrower pattern. Taking TZ as an example, 20% (monthly) and 26% (daily) of the time was considered as VOC- limited regime. The number of days and proportions for photochemical regimes summarized at four patterns of time scales can reveal a more plausible and comprehensive variation in ozone formation chemistry. Compared with patterns of monthly and weekly scales, the results derived at a daily scale can reveal the temporal variability of photochemical regimes more comprehensively. Note that the photochemical regime proportion obtained from the day-to-day scale has an advantage due to the large number of statistical samples.

3.6 Comparison among different patterns of time scale

 This section gives a more comprehensive understanding of the campaign- averaging O3-precursor relationship by comparing the similarities and differences of the results from various patterns of time scale. The overall O₃-precursor relationship for the entire campaign can be quantified by averaging the RIR values from the individual simulation runs depending on the chosen time scale (e.g., five simulation runs for monthly scale in this study). Therefore, four sets of logical and comparable results can be derived to represent the campaign-averaging O3-precursor relationship, as four patterns of time scale (i.e., five-month, monthly, weekly, and daily) were treated in this study.

 Figure 6 shows the averaged RIR values of the major precursor groups at different patterns of time scale. As the time scale changed from wider (i.e., five-month scale) to narrower (i.e., daily scale) pattern, all three sites showed increases in the means of RIR_{AVOC} and RIR_{alkenes}* as well as decreases in averaged RIR_{NOx}, whereas the averaged RIR of other precursors (i.e., BVOC, CO, alkanes and aromatics) did not vary obviously 409 (see **Table S6**). Comparing with the O_3 -VOC-NO_x sensitivity at the daily scale, the results obtained at the five-month scale underestimated O3-AVOC sensitivity (indicated by averaged RIR values) by 48% (TZ), 66% (BJ), and 49% (XD), and overestimated O_3-NO_x sensitivity by 37% (TZ), 142% (BJ), and 144% (XD). We performed comprehensive uncertainty analysis for model input and output results, which was

 assessed through statistical methods (see details in **Section 3.7**). We found that the model-derived RIR values may become more uncertain when the input dataset was averaged into a wider diurnal pattern (i.e., five-month scale), which may explain the discrepancy of RIR values between five-month scale and daily scale. We expect that such discrepancies derived from different patterns of time scale could widely exist in many other world areas. Note that the mean RIR values were generally consistent 420 among the four patterns of time scale within a reasonable range (within $25\text{-}75^{\text{th}}$ quantile and standard deviation, see **Figure 6** and **Table S4**), suggesting that any selected pattern of time scale could reasonably derive the campaign-averaging O₃-precursor relationship.

 Figure 7 further shows the variations in photochemical regimes (defined by RIRNOx/RIRAVOC; see **Text S2** and **Table S4** for details) for each pattern of time scale. Specifically, TZ was mainly considered as transitional regime for the entire campaign period, whereas its variations covered three photochemical regimes, which was consistent with the results from **Table S6**. BJ was generally identified as VOC-limited regime, whereas some days were also grouped into transitional regime. XD was considered as primarily between VOC-limited and transitional regime, and its variations also spanned three photochemical regimes. Compared with the five-month 431 pattern, it was further found that the averaged $RIR_{\text{NOx}}/RIR_{\text{AVOC}}$ from other time scale patterns (i.e., monthly, weekly, and daily) were higher (12% to 20% for TZ; 38% to 153% for XD) or lower (21% to 65% for BJ) than that from five-month scale. Note that the above discrepancies in photochemical regime derived from multiple patterns of time 435 scale may influence the development of targeted O_3 control strategies. In summary, the 436 photochemical regime derived by averaging RIR_{NOX}/RIR_{AVOC} from the daily scale (see **Table S6**) suggests that the three sites mainly followed the sequence of TZ (1.34 \pm 1.39 > XD (0.67 \pm 1.49) > BJ (0.16 \pm 0.65).

439 In addition, the temporal variations of $TVOC/NO_x$ in different timescales were identified during the whole campaign, and good correlations between observed 441 TVOC/NO_x and model derived RIR_{NOX}/RIR_{AVOC} at four patterns of time scale were also found (see **Figure 5**). Such consistency suggests that both metrics can reasonably reflect the variation of photochemical regimes, which can also improve the reliability of our box model simulation.

 The consistency and difference of model output (summarized in **Table S7**) are quantified by the statistical methods of Pearson's correlation coefficient (Hu et al., 2018) and paired-samples *t*-test analysis (Wang et al., 2016). In particular, we assess and compare the degree of significance of differences among multiple patterns of time scale 449 by the *p* values (a statistical significance assuming at $p < 0.05$) through paired-samples *t*-test and Wilcoxon matched-paired signed-rank test (non-parametric statistics) (Chiclana et al., 2013). **Figure 8a** shows high Pearson's correlation coefficients (with values all above 0.85, *p* < 0.01) were found among four patterns of time scale, and the

 higher correlation coefficient was identified between the two closer patterns. **Figure 8b-c** shows that the differences among multiple patterns of time scale were non- significant using Paired-samples t-test analysis and Wilcoxon matched-pair signed-rank test respectively. Furthermore, their results indicate that more significant difference was recognized between the two distant patterns (e.g., daily and five-month), which is consistent with the results of Pearson's correlation analysis. Noted that the discrepancy 459 between the two distant patterns was not significant but non-negligible (e.g., $p = 0.092$) of Wilcoxon matched-paired signed-rank test between five-month and daily patterns).

 The influence of different patterns of time scale on deriving RIR values from individual AVOC species was further investigated. Briefly, quantifying the relative 463 contribution of individual AVOC on O_3 formation based on RIR calculation is beneficial to the development of cost-effective AVOC control strategies (Zhang et al., 2021). **Figure 9** shows the averaged RIR values of individual AVOC species (i.e., top 10) at different patterns of time scale (i.e., five-month, month-to-month, week-to-week) at three sites in Zibo. As shown in **Figure 9**, the 10 individual AVOC species at the three sites were selected according to the top 10 highest RIR from five-month pattern. All three sites showed that the RIR of individual AVOC species increased gradually as the 470 time scale changed from the wider (i.e., five-month) to narrower (i.e., weekly) pattern, which was consistent with the earlier discussion (see **Figure 6** and **Table S6**) of O3- AVOC sensitivity derived from four patterns of time scale. The results also indicate that the choice of time scale pattern has a limited effect on deriving high-ranking AVOC species (i.e., top 10) based on RIR calculations.

3.7 Uncertainty analysis

 The uncertainty of model input was quantified in this section, which is embedded in pre-processed dataset with multiple patterns of time scale. As showed in **Figure 1**, the daily simulation used the individual daily pattern to constrain model, while the input dataset of averaged diurnal patterns (i.e., weekly, monthly, and five-month) is treated by averaging individual daily pattern into different timescales. This averaging approach will conceal the temporal variations of O³ precursors and meteorological factors, particularly for a long-term observational campaign. **Figure S14** shows the 483 distributions of the standard deviations for OH reactivity (k_{OH}) or concentration of O₃ precursor groups at three averaged patterns of time scale at the three sites. As the time scale changed from wider (i.e., five-month scale) to narrower (i.e., weekly scale) pattern, the uncertainty (indicated by the average, median and 25%-75% quantile) decreased accordingly. In addition, meteorological factors such as temperature and irradiation also play an important role on $O₃$ formation, especially these meteorological parameters can vary greatly over a long observational period (Boleti et al., 2020; Liu et al., 2019b; Weng et al., 2022). Therefore, the masked temporal variation of these meteorological

factors behind the averaged input dataset would also result in model uncertainty.

 Moreover, it has been widely recognized that the uncertainty for 0-D box model simulation mainly arises from the constraint of observation dataset and the configuration of model scheme. Note that constraints with more species from measurements (or including as many species as possible) would lower its uncertainty from the chemical box model simulation (Wolfe et al., 2011, 2016). Nevertheless, due to the measurement limitation in our field campaign, we are unable to measure some important atmospheric species (i.e., HONO and oxygenated VOC (OVOC)), and these may arise uncertainty in box model simulation. For instance, Xue et al., (2021) performed a sensitivity test for HONO constraint in their box model simulation, and they showed that without HONO constraint would lead to $O₃$ photochemical production rate decreasing by 42%. More recently, Wang et al., (2022) obtained a comprehensive VOC dataset at Guangzhou, and their results showed that box model simulation without OVOCs constraints would underestimate the productions of ROx and O3. Besides, both 505 gaseous HNO₃ and organic nitrates can result in interferences on NO_x measurement by chemiluminescence technique, which may arise uncertainty in our box modelling (Ge 507 et al., 2022; Uno et al., 2017; Xu et al., 2013). Since the accurate NO_x measurement is 508 essential in determining the photochemical regime, more in-depth studies on NO_x measurement uncertainty in box model simulation are required in the future. In addition, the parameter configuration of model scheme is essential to derive a reliable and valid model output, such as dilution rate as an important model technical parameter. We performed a stepwise sensitivity test for this parameter to obtain an optimized dilution rate, and assigned it to all non-constraint species, which can reduce uncertainty in box model simulation (see details in **Text S1**). Also, the dry and/or wet deposition of pollutants is an important atmospheric physical process, which has been mostly parameterized in emission-based chemical transport modeling but very limited in box model, as most of the primarily emitted species are already constrained from 518 measurements. Xue et al., (2014) considered O_3 deposition into box model simulation, 519 and their result showed negligible contribution of O_3 deposition to total O_3 destruction rates. As for this work, we are unable to consider the deposition due to the difficulty in representing and parameterizing this term in the 0-D box model. Nevertheless, deposition of O³ and other species may be one of the uncertainties during box model simulation, which is worth further study in the future.

4 Summary and implications

 Our present results suggest that comprehensively understanding of multiple 526 patterns of time scale is conductive to formulating a more accurate and robust O_3 control strategy. Specifically, as identified from the narrower patterns of time scale (i.e., weekly 528 and daily), the site-to-site photochemical regime indicated by RIR_{NOx}/RIR_{AVOC} showed 529 various magnitudes but a synchronous temporal trend. This indicates that the $O₃$ formation regime in a city area can be influenced by local and regional emissions jointly. The reason behind this phenomenon is not clear at present, and we believe that further investigation on the synergetic effect of local and regional emission reduction for O³ control would help elucidating this observation. It was also found that the campaign- averaging photochemical regimes showed overall consistency but non-negligible variability among the four patterns of time scale, which was mainly due to the embedded uncertainty in model input dataset with averaged diurnal patterns. This implies that comparison among multiple patterns of time scale based on RIR analysis 538 is useful to derive the O_3 -precursor relationship more accurately and reliably.

 Moreover, the high-ranking AVOC species (i.e., top 10) based on RIR calculations were overall consistent from the narrow to wide patterns of time scale. **Table S8** summarizes the total run number of box model for different patterns of time scale. It is known that large-scale computing capacity and computational efficiency were required in the narrower pattern of time scale (e.g., 2760 simulation runs in weekly scale in this study). Considering the difficulties of performing long-term and continuous online measurements in some environments, it is also advisable to identify the high-ranking VOC species from the campaign-averaging diurnal pattern in box model simulation.

547 In this study, we explored the non-linearity of O_3 -precursor relationship in a way driven by the actual daily / weekly / monthly variability around the distribution. Our results highlight the importance to quantitatively test the impact of different timescales on photochemical regime determination, as there is uncertainty embedded in model input dataset when averaging individual daily pattern into different timescales. Such understanding would be complementary in developing more accurate $O₃$ pollution 553 control strategy, particularly as the long-term O_3 -precursor observations (e.g., from several months to years) are becoming more available than before in many places of China. In addition, site-to-site differences of model-derived photochemical regimes 556 also underlines the importance of developing target O_3 control strategy for different 557 areas in a city scale. Specifically, according to the averaged RIR_{NOx}/RIR_{AVOC} at daily pattern, the derived photochemical regime was transitional for TZ (suburban) and XD (suburban), while VOC-limited for BJ (urban). This implies that for mitigating ozone pollution in Zibo city, more endeavors should be devoted to the anthropogenic VOC reduction in urban areas, while strengthening the synergetic mitigation of VOC and NO_x emissions at the same time in other suburban areas. Although the above implications for O³ control were derived from a case study in a major prefecture-level city (Zibo) of northern China, the developed approach by integrating multiple patterns of time scale in the present work can be used to other regions, particularly the on-going "One City One Policy" campaign (2021-2023) for O³ control in many cities in China.

Code availability

- The code for the Master Chemical Mechanism (MCMv3.3.1) can be achieved from
- http://mcm.york.ac.uk (last access: 27 January 2023).

Data availability

- The datasets generated and analysed during the current study are available on
- reasonable request from the corresponding author (Kangwei Li).

Supplement

The supplement related to this article is available online.

Author contribution

- KL conceived and led the study. ZZ performed the modeling. ZZ, KL, and ZB analyzed
- the data. BX, JD, LL, GZ, SL, CG, and WY conducted the field measurement. ZZ and
- KL wrote the paper. MA and ZB commented on the manuscript.

Competing interests

The authors declare that they have no conflicts of interest.

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606 published 0-D box model studies in China. 909 Table 1. Summary of relevant published 0-D box model studies in China. Summary of relevant

- 911 **Table 2.** Summary of the number of days (for model calculation) and proportions that
- 912 were classified into the three photochemical regimes across each site and multiple 913 patterns of time scale.

916 **Figure 1.** Schematic diagram of the dataset treatment to derive four patterns of time scale for 0-D box

917 model input. Note that the four patterns (i.e., five-month, monthly, weekly, and daily) were the diurnal

918 average of the initial dataset. This diagram takes one site and several input measurements (temperature,

919 toluene, and $NO₂$) as examples.

920
921 Figure 2. Time series of meteorological parameters, O₃ and its precursors (i.e., CO, NO_x, VOCs)

922 throughout the whole campaign at the three sites in Zibo.

925 **Figure 3.** Time series of month-to-month RIR values of major precursor groups and RIRNOX/RIRAVOC at 926 three sites (TZ, BJ and XD) in Zibo. The green dash line indicates to $RIR_{NOx}/RIR_{AVOC} = 0.5$ and 2.

928 Figure 4. Time series of week-to-week RIR values of major precursor groups and RIR_{NOx}/RIR_{AVOC} at 930 three sites (TZ, BJ, and XD) in Zibo. The blue lines in (g)-(i) are the three points moving average of 931 RIR_{NOx}/RIR_{AVOC} value.

932
933 Figure 5. The correlations of TVOC/NO_x with RIR_{NOX}/RIR_{AVOC} at multiple patterns of time scale at the

934 three sites in Zibo.

936
937 Figure 6. Distribution of RIR values of major precursor groups in multiple patterns of time scale at three

938 sites (TZ, BJ, and XD) in Zibo.

940 **Figure 7.** Distribution of RIR_{NOx}/RIR_{AVOC} (indicator of photochemical regime) in multiple patterns of

941 time scale at three sites (TZ, BJ, and XD) in Zibo.

942
943 Figure 8. The statistical analysis results of RIR values (from Table S6) at multiple patterns of time scale:

944 (a) Pearson's r correlation analysis (all the results have passed statistical significance assumed at $p <$ 945 0.01), (b) Paired-samples t-test analysis (p values refer to differences with a statistical significance

946 assumes at $p \le 0.05$), (c) Wilcoxon matched-pair signed-rank test (\dot{p} values refer to differences with a

947 statistical significance assumes at $p < 0.05$).

948
949 Figure 9. Averaged RIR values of individual AVOC species (top 10) at different patterns of time scale at

950 three sites (TZ, BJ, and XD) in Zibo. The error bars represent the standard deviations of the mean.