

27 the conducted experiments. The growth rate of the newly formed particles ranged from $3 - 11$ nm h⁻¹ with particles reaching a diameter of 20-25 nm after a few hours. The nucleation rate was estimated using an aerosol dynamics model and was found to be in the range of 500 to 25000 30 particles $cm^{-3} h^{-1}$ for the different experiments. These results support the hypothesis that ammonia at levels of several ppb can accelerate or even cause new particle formation at least in the environment of the Eastern Mediterranean.

1. Introduction

 Atmospheric aerosol can be produced from many natural or anthropogenic sources and plays a significant role in Earth's climate and in public health (Haywood and Boucher, 2000; Pope et al., 2002). Aerosols can affect climate either by scattering and absorbing incoming solar radiation (direct effect) or by acting as cloud condensation nuclei (CCN) thus affecting reflectivity and lifetimes of clouds (indirect effect). New particle formations (NPF) through nucleation of low- volatility vapors can be an important source of atmospheric aerosols and is responsible for close to 50% of the global CCN (Merikanto et al., 2009). Newly formed particles either grow to larger sizes through condensation or are scavenged by larger preexisting particles through coagulation. Self-coagulation is another growth process for the newly formed particles. The competition between these processes determines how many of those new particles will grow to become CCN and how fast this will happen. NPF has been observed in many areas around the world including all types of environments (urban, rural, forests, remote, marine, etc.) (Kulmala et al., 2007; Kerminen et al., 2010; Wang et al., 2017; Yao et al., 2018; Zhu et al., 2019; Saha et al., 2019).

 One of the challenges in studying the new particle formation process is the ability of the available instrumentation to measure such small particles (Kulmala et al., 2012). Previous studies have underlined the importance of sulfuric acid for NPF in most environments (Jaecker-Voirol and Mirabel, 1989; Weber et al., 1996; Laaksonen et al., 2000; Sipila et al., 2010). Additional studies have shown the importance of ammonia and amines as vapors that can accelerate the nucleation rate of sulfuric acid with water by stabilizing the initial clusters of sulfuric acid. (Weber et al., 1998; Kirkby et al., 2011; Jen et al., 2014; Glasoe et al., 2015). Low and extremely low volatility

 organic vapors play a major role in the growth of the new particles and may be also participating in the nucleation process itself (Yli-Juuti et al., 2011; Zhao et al., 2014; Ehn et al., 2014; Mohr et al., 2019). In marine environments iodine compounds have been identified as vapors that can form new particles (McFiggans et al., 2010; Sipilä et al., 2016; He et al., 2021) Wang et al. (2020) have recently reported fast growth rates of newly formed particles at some atmospheric conditions due to the condensation of ammonium nitrate. The preexisting aerosol (condensation sink), the availability of gaseous precursors and the meteorological conditions all affect the intensity and frequency of NPF events in the atmosphere.

 Extensive monitoring of NPF events has taken place in many sites in Europe (Manninen et al., 2010; Dinoi et al., 2021) and the eastern Mediterranean (Pikridas et al., 2012; Berland et al., 2017; Kalkavouras et al., 2017; Kalivitis et al., 2019; Hussein et al., 2020; Brilke et al., 2020). Siakavaras et al. (2016) reported frequent nucleation events in Thessaloniki, a major urban center in northern Greece. On the other hand the nucleation frequency in southern Greece is relatively low (compared to central and northern Europe) especially during the summer (Kalivitis et al., 2008; 2019). Kopanakis et al. (2013) observed nucleation events only in 13 out of the 157 days of measurements in the Akrotiri station, in Crete. Kalkavouras et al. (2020) reported a relatively low 20% nucleation frequency during the summer in Finokalia, Crete. Particle size distribution measurements in five stations in Greece (Athens, Patras, Thessaloniki and Finokalia) during the summer of 2012, showed low NPF frequency in Patras and Finokalia (Vratolis et al., 2019). Pikridas et al. (2012) provided evidence that ammonia or amines may be the missing reactants that are responsible for the lack of nucleation in this sunny relatively clean area with available sulfur dioxide.

 In this work we test the hypothesis that in an environment such as the Eastern Mediterranean during the summer, under conditions favorable for nucleation (intense sunlight, low to moderate particle concentrations, adequate sulfur levels) nucleation events are rare due to the relatively low ammonia levels. The experiments took place during summer in Patras, Greece in an environment with low regional nucleation frequency (Patoulias et al., 2018; Vratolis et al., 2019) using a dual chamber system. Both chambers were filled with ambient air, ammonia was added to one of them, and the evolution of the aerosol was followed in both chambers.

-
-

2. Methods

2.1 Dual chamber system

 A dual chamber system was deployed as part of the 2019 summer PANACEA (PANhellenic infrastructure for Atmospheric Composition and climatE change) campaign in Patras, Greece. Measurements were conducted in the outskirts of Patras (population 200.000 people) between July 15 until August 15 2019, in the Institute of Chemical Engineering Sciences (ICE-HT), approximately 8 km from the city center. The dual chamber system consisted of two 93 identical Teflon chambers (1.3 $m³$ each) located inside a structure that included the chambers and Five panels of UV lights used for illumination purposes $(J_{NQ2}=0.25 \text{ min}^{-1})$. Ammonia was added in one of the chambers (perturbation chamber) while the other one was used as the reference. The top of the structure can be removed and natural sunlight was used if the weather conditions allowed it. Details about the design and testing of the dual chamber system can be found in Kaltsonoudis et al. (2019).

2.2 Experimental procedure

 Before the beginning of each experiment both chambers where flushed with ambient air for approximately 2 hours. The main purpose of the flushing is the conditioning of the chambers and the sampling lines to the environmental conditions and composition thus minimizing losses of volatile or semivolatile compounds to the walls of the system. During this preparation period both chambers were swept for 20 min using an ionizer fan (Dr Schneider PC, Model SL-001) to reduce the charges on the chamber walls thus reducing the particle wall losses (Jorga et al., 2020). After the chambers were ready, they were filled with ambient air using a metal-bellows pump (Senior Aerospace, model MB-302). The concentrations of pollutants in both the gas and particulate phases were then characterized for one hour. A 0.25 in copper tube was used for the particle phase measurements and a 0.25 in PTFE tube was used for the gas measurements. The instruments were located inside a room next to the chambers to avoid their exposure to high temperatures. The distance from the chambers to the instruments was approximately 4 m. An automated valve was used to alternate sampling between the two chambers. The valve was synchronized with the sampling periods of the various instruments and sufficient time was allowed between each sampling change to flush any remaining air from the previous measurement.

 After the characterization phase, ammonia was injected through a heated line into the perturbation chamber using a glass syringe. The concentration of the injected ammonia was estimated using the volume of the chamber and the amount of liquid ammonia injected and it varied from 20 to 200 ppb. After the ammonia injection, the top cover of the system was removed and the chambers were exposed to natural sunlight. If the wind speed was high, the UV lights were turned on, illuminating both chambers which remained covered.

 At the end of the experiment, ammonium sulfate seeds were injected into both chambers to measure the size dependent particle wall-loss rate constants using the method described in Wang et al. (2018). After the end of the particle wall-losses period both chambers were flushed once again with ambient air, to remove the ammonium sulfate and any remaining pollutants and to prepare them for the next experiment.

2.3 Instrumentation

 The chemical composition of the aerosol was monitored using a High Resolution Time-of- Flight Aerosol Mass Spectrometer (HR-ToF-AMS) from Aerodyne Research Inc. Two Scanning Mobility Particle Sizers (SMPS) were used to measure the number size distributions from 9-160 nm (classifier model 3080, CPC model 3775) and from 14-730 nm (classifier model 3080, CPC model 3025A) respectively. The sample flow was dried before reaching the AMS and SMPS systems using a nafion dryer. A suite of gas monitors were used to measure the concentrations of 135 NO_x (Teledyne 201E/501), SO₂ (Thermo Scientific Model 43i), and O₃ (Teledyne 400E).

2.4 Aerosol dynamics model

 A zero-dimensional aerosol dynamic model was used for the simulation of nucleation, condensation and coagulation inside the perturbation chamber (Capaldo et al., 1999). The multicomponent aerosol size distribution is described using 270 size sections covering the diameter range from 1 nm to 1 μm. The aerosol components include sulfate, ammonium, organics and others with the latter assumed to be non-volatile and inert during the few hours of the simulation period.

144 The condensation rate of H₂SO₄ to a particle of diameter D_p is described using the modified form of the Fuchs-Sutugin equation (Hegg et al., 1991; Kreidenweis et al., 1991) given by:

$$
J = 2\pi D D_p F(Kn) A(P - P_o) \tag{1}
$$

148 where *D* is the diffusivity of the vapor air (set to 0.1 cm² s⁻¹ in this application), *Kn* is the Knudsen 149 number (that is the ratio of the air mean free path to the particle radius), *F(Kn)* is a coefficient 150 correcting for free molecular effects given by:

151
$$
F(Kn) = \frac{1+Kn}{1+1.71Kn+1.33Kn^2}
$$
 (2)

152 and *A* is a coefficient correcting for the interfacial mass transport limitations described by the 153 accommodation α_e ,

154
$$
A = \left[1 + 1.33Kn F(Kn) \left(\frac{1}{a_e} - 1\right)\right]^{-1}
$$
 (3)

155 Finally, *P* is the bulk H2SO⁴ vapor partial pressure and *P^o* is its partial pressure at the particle 156 surface. An accommodation coefficient of 0.02 for the condensation of $H₂SO₄$ on the aerosol 157 particles is assumed (Van Dingenen and Raes, 1991). The vapor pressure of H_2SO_4 at the aerosol 158 surface can be estimated from the data of Bolsaitis and Elliott (1990). Values less than 10^{-3} ppt 159 were found for the conditions of our experiments and therefore the surface vapor pressure of H₂SO₄ 160 in our mass transfer calculations was assumed to be zero.

161 Brownian coagulation between all particles is simulated solving the discrete coagulation 162 equation (Seinfeld and Pandis, 2016) :

163
$$
\frac{dN_k(t)}{dt} = \frac{1}{2} \sum_{j=1}^{k-1} K_{j,k-j} N_j N_{k-j} - N_k \sum_{j=1}^{\infty} K_{k,j} N_j \quad k \ge 2
$$
 (4)

164 The generalized coagulation coefficient *K1,2* for the collision of two particles is calculated as:

165
$$
K_{1,2} = 2\pi (D_1 + D_2)(D_{p1} + D_{p2})\beta
$$
 (5)

166 where D_1 , D_2 are the individual Brownian diffusion coefficient for the particles, D_p , D_p are the 167 particle diameters and β is the Fuchs correction factor (Fuchs, 1964). Because of the high 168 resolution of the size distribution, coagulation can be simulated accurately by calculating directly 169 the coagulation rates between each of the size sections and moving the particles to the 170 corresponding size bin.

171

173 **2.5 Data analysis**

 The condensation sink (CS) is a metric of the ability of the pre-existing aerosol population to remove vapors from the system by condensation. The CS values were calculated using the aerosol distribution between 14-700 nm, as measured by the SMPS and the properties of sulfuric acid as the condensing vapor. The CS is given by:

$$
CS = 2\pi D \sum_{i} \beta_{mi} D_{pi} N_{i}
$$
 (6)

179 where D is the diffusion coefficient of sulphuric acid, β_m is the transition-regime correction factor, 180 D_p the diameter of the particle and N the respective number concentration in each size bin of the 181 SMPS.

182 Using the initial SO² concentration in the perturbed chamber and the condensation sink we 183 can estimate the sulfuric acid concentration according to:

184
$$
[H_2SO_4] = k_{OH} \frac{[SO_2][OH]}{CS}
$$
 (7)

185 where k_{OH} is the reaction constant of SO₂ and OH which is equal to 8.5 \times 10⁻¹³ cm³ molecule⁻¹ s⁻¹ 186 at 298 K (Demore et al., 1997), [SO₂] and [OH] are the concentrations of sulphur dioxide and 187 hydroxyl radicals respectively, and CS is the condensational sink as calculated from Equation 7. 188 For the hydroxyl radical concentration, we assumed an average value of 5×10^6 molecules cm⁻³. 189 Equation (7) is based on the assumptions that the only sulfuric acid source is the oxidation of $SO₂$ 190 from OH radicals, its major sink is its condensation onto the aerosol surface, and the system is at 191 pseudo-steady state.

192

193 **3. Results and discussion**

194 **3.1 Initial conditions**

195 Thirteen perturbation experiments were conducted during the study (Table 1). Two of them 196 took place during the night using UV lights and the rest during midday. Natural sunlight was used 197 in two experiments while UV lights were used during the rest.

198 The main components of non-refractory $PM₁$ in the beginning of our experiments were organics (46.6±6.5%) followed by sulfate (37.1±4.5%), ammonium (14.3±1.8%), nitrate $(1.5\pm0.5\%)$ and chloride $(0.5\pm0.4\%)$. The average oxygen to carbon ratio (O:C) (Canagaratna et al., 2015) in the chambers after filling them with ambient air was 0.68±0.1, indicating an already

 oxidized OA. In order to check if any contamination was occurring during the flushing and filling processes we calculated the theta angle (Kostenidou et al., 2009) between the organic mass spectra 204 of the ambient air and the OA in the two chambers. The theta angles were less than 6° in all experiments, indicating negligible contamination during the filling process.

206 The initial concentration of the SO_2 , NH₃, O_3 and NO_x inside the chambers after the filling 207 process were approximately within 10% of their ambient values. The initial concentrations of these 208 gases in the two chambers differed by less than 3%. More than 70% of the ambient PM_1 was 209 transferred in the chambers in most experiments. The initial $PM₁$ levels were quite low ranging 210 from 0.6 to 4.2 μ g m⁻³. The atmosphere of Patras was quite clear during these experiments. The 211 initial conditions in all experiments are summarized in Table 1.

212

213 **3.2 New particle formation and growth**

 The conducted experiments were classified in three different classes based on the observations of new particle formation in the two chamber. In class A experiments, nucleation and particle growth occurred in only the perturbed chamber, in class B nucleation and particle growth happened in both chambers and in class C when there was no detection of new particle formation in either chamber.

219

220 **3.2.1 Nucleation and growth only in the perturbed chamber**

221 Nucleation and growth of the new particles to sizes above 9 nm only in the observations of 222 new particle formation. Figure 1 depicts the particle number concentration N_9 ($D_p > 9$ nm) after 223 corrections for losses to the chamber walls and the sampling lines, inside the two chambers for a 224 typical Class A experiment. During Exp. A1 the initial concentration of $SO₂$ was 0.7 ppb and of 225 O_3 equal to 58 ppb. After the injection of ammonia (approximately 150 ppb) the UV lights were 226 turned on (t=0 h) illuminating both chambers. The *N*⁹ particle number concentration start 227 increasing in the perturbed chamber approximately 1.5 h after the lights were turned on, reaching 228 close to 4000 cm⁻³, almost double its initial concentration value. The *N*₉ concentration in the 229 control chamber remained within 5% of the initial levels. Figure 2 shows the measured number 230 distributions in the two chambers after correction for particle losses. The formation and growth of 231 the new particles in the perturbed chamber is evident. With a CS of 0.0026 s⁻¹ the H₂SO₄ 232 concentration was calculated to be of the order of 3×10^7 molecules cm⁻³. Assuming that nucleation

233 started at t=0 the measured initial growth rate in the perturbed chamber was on average 4 nm h^{-1} . This rate is based on the time of growth of the nucleation mode to 15 nm. For this experiment 3.8 h were required for the growth of the new particles to 15 nm so the estimated growth rate is 236 approximately 4 nm h^{-1} . The newly formed particles at the end of the experiment (after 5 h from the illumination) grew to approximately 20 nm. The experiment was stopped at that point because a significant fraction of the air in the two chambers had been sampled.

239 The estimated initial growth rates for the daytime experiments A1, A2, A4, A5, and A6 240 varied from 3.5 to 8 nm h⁻¹ and were correlated with the estimated H₂SO₄ levels (R^2 =0.76) (Table 241 2). Experiments A1, A5 and A6 had similar CS and H₂SO₄ levels and resulted in similar initial 242 growth rates (Table 2). The slope of the growth rate versus sulphuric acid linear regression for the 243 daytime experiments was 1.4 (nm h⁻¹ molecule⁻¹ s) and the intercept was low and equal to 0.05 nm 244 h^{-1} .

 Experiment A3 was conducted during the night (the chambers were filled with ambient air at 21:00 so it has relatively different behaviour than the rest. If this experiment is included in the 247 growth rate versus H₂SO₄ correlation, the R^2 drops to 0.27. This probably suggests that the 248 estimated OH levels are not accurate in this case and therefore the $H₂SO₄$ is probably also a lot more uncertain than in the other experiments. Also, the growth process may be different with organic vapors playing a more significant role.

251

252 **3.2.2 Nucleation and growth in both chambers**

253 In 4 (B1 to B4) out of the 13 perturbation experiments, new particle formation and growth 254 was observed in both chambers (Table 2). This suggests that the ambient air had already the 255 potential to form new particles without the addition of ammonia.

256 Figure 1b shows the *N*⁹ concentration in the two chambers during Exp. B1. The initial 257 levels of SO² in both chambers were 0.8 ppb and approximately 150 ppb of ammonia were added 258 to the perturbation chamber. Half hour after the exposure of the chambers to UV light the N_9 in 259 both chambers started increasing with higher concentrations in the perturbed chamber. The number 260 concentration of the particles in the perturbed chamber reached close to 6000 cm^{-3} almost three 261 times the initial levels. The concentration in the baseline chamber increased by approximately 50% 262 to 3000 cm^{-3} . The newly formed particles in the perturbed chamber at the end of the experiment 263 reached a mode diameter of 26 nm (Fig. 3) with an initial GR of 5.5 nm h^{-1} . The growth rate of the

 particles in the reference chamber was only 10% lower than in the perturbed one (Fig. 4) suggesting that the addition of ammonia probably influenced significantly the nucleation rate but had a small effect on the growth rate.

 We tested the hypothesis that the appearance of the new particles in the reference chamber was due to a sampling error, caused by some cross contamination of the two samples as the same sampling line was used. We compared the shape of aerosol size distributions in the two chambers. The nucleation mode distribution in the reference chamber was wider (Figure S1), a strong indication that these were different particle populations sampled by our system.

272 The condensation sink in the class B experiments was on average 0.0024 s⁻¹ quite similar 273 to the 0.0023 s⁻¹ in the A experiments so the pre-existing particle mass was quite similar in the two 274 classes of experiments. Also the average $SO₂$ was practically the same (0.83 ppb for the B experiments and 0.82 ppb for the A experiments). Adding the similarity of the UV intensity, it is 276 clear that the major factors (sunlight, condensational sink, SO_2 availability) usually determining nucleation rates were not the reason for the weak nucleation and growth in the reference chamber in these experiments. Unfortunately, we did not have available accurate ammonia measurements, because the presence of sufficient ammonia levels is one of the possible explanations for this behaviour.

281 The observed growth rates in these B-class experiments varied from 3.5 to 11.3 nm h^{-1} and were a little higher on average than those in the A group of experiments. Adding these four 283 experiments to the linear regression of the growth rate versus sulphuric acid reduced the R^2 to 0.43, 284 but the slope remained the same, while the intercept increased to 0.7 nm h⁻¹. These provide some weak evidence of the involvement of more compounds, probably organics, in the growth of these newly formed particles in this second group of experiments.

3.2.2 Nucleation and growth not observed

 In three of the experiments C1-C3, we did not observe growth of new particles to the size of 9 nm. It is still possible that there was nucleation, but the growth may have been too slow (less 291 than 2 nm h^{-1}). Figure 1c depicts the N_9 number concentration during Exp. C1. The initial 292 concentration of SO_2 was 1.3 ppb in this experiment. The number concentration after correction 293 for particle losses remained constant at close to 2200 cm^{-3} and the corresponding number distributions changed little during the experiment (Fig. 4). Exp. C1 had the lowest initial levels of

 ozone of all the experiments (Table 1), around 12 ppb, a factor of four lower than the average 296 concentration. The low O_3 levels were probably due to the highest NO_x levels (27 ppb) in this air mass.

 Exp. C2 was conducted in the early evening (the chamber was filled with ambient air around 19:00 LT) and the lowest detectable particle size for this experiment was 14 nm because 300 of technical difficulties. Finally, Exp. C3 had relatively low levels of sulfuric acid $(2.5\times10^{7}$ molecules cm⁻³ compared to the rest of the experiments, a low estimated level of injected ammonia (20 ppb) and natural sunlight.

3.3 Particle composition

 The mass concentration of the major components of PM¹ (sulfate, organics, nitrate, ammonium) in the two chambers, after correcting for chamber particle wall losses, remained practically constant during all experiments. The corrected for wall losses mass concentration inside both chambers during Exp. B1 are shown in Figure 5. Taking into account the uncertainty of the wall loss correction, the maximum production of the corresponding secondary PM components during the few hours of the experiments should have been a few percent or less. This will be an important constraint for the analysis of these experimental results with the aerosol dynamics model in the next section.

3.4 Estimation of nucleation rate using an aerosol dynamics model

 We used our aerosol dynamics model to simulate the growth and coagulation of the particles in the perturbed chamber assuming a nucleation rate. Our goal is to use the observations to constrain the nucleation rate that could not be measured directly. The model uses as inputs the temperature and relative humidity during the experiments and is initialized with the measured particle number distribution at time zero. There are three adjustable parameters in the model: the duration of nucleation, the nucleation rate and the condensation rate. Nucleation is assumed to start at time zero and a constant nucleation rate is assumed for the duration of the event. The condensing components are assumed to have practically zero vapor pressure. The three parameters were chosen so that the model predictions were in good agreement with the observations of particle number concentration and size distribution and also the mass concentration.

 Figure 6 shows the measured and the predicted particle number, surface and volume concentrations in the perturbed chamber for Exp. A1. A nucleation event with rate equal to $J_1=9500 \text{ cm}^{-3} \text{ h}^{-1}$ and duration of 3 h together with a condensation rate of 3.2 ppt h⁻¹ was needed to reproduce the observations. For much lower condensation rates the particles did not grow to detectable sizes and for higher condensation rates the predicted PM mass increase was not consistent with the small observed mass concentration change. We performed sensitivity analysis 331 around these central values and values of $J_1=9500\pm600$ cm⁻³ h⁻¹ remained consistent with the observations. The average errors during the simulation were 6% for the number concentration, 16% for the surface concentration and 17% for the mass concentration.

 The predicted and observed evolution of the aerosol number distributions are shown in Figure S2. The differences for the smaller particle sizes are partially due to the losses of these particles in the sampling and measurement systems.

 Table 3 summarizes the estimated nucleation rates together with the corresponding durations of the nucleation events and the required condensation rates for all experiments in groups 339 A and B in which nucleation and growth were observed. The estimate J_1 rates varied from 500 to $25000 \text{ cm}^{-3} \text{ h}^{-1}$.

 These results can be roughly compared to the CLOUD measurements for sulfuric acid-342 ammonia nucleation (Kirkby et al., 2011) in the 2×10^7 -10⁸ molecules cm⁻³ H₂SO₄ concentration range that was estimated for our experiments. The CLOUD measurements for the highest ammonia 344 levels used suggested a $J_{1.7}$ rate of approximately 500 cm⁻³ h⁻¹ for H₂SO₄ concentration equal to $5x10^7$ molecules cm⁻³. For experiment A4 we estimated the same H₂SO₄ concentration and a 346 nucleation rate of 400 ± 200 cm⁻³ h⁻¹ (Fig. S3). While this agreement is probably fortuitous, overall our estimated nucleation rates are in general consistent (considering their uncertainties) with the CLOUD measurements for the ammonia-sulfuric acid system assuming that the rate does not increase further as ammonia increases above 1 ppb.

4. Conclusions

 A dual chamber system was used to investigate the hypothesis that ammonia is often the limiting reactant for new particle formation in the Eastern Mediterranean. Ambient air characterized by relatively aged air masses in southern Greece was used as the starting point of our experiments. Ammonia was added in one chamber while the second was used as a reference.

 In 6 out of the 13 experiments (46%) the addition of ammonia led to the formation and then growth to detectable size (approximately 10 nm) of new particles, while no formation of particles was observed in the reference chamber. In another 4 experiments (31%) the addition of ammonia enhanced significantly the formation of new particles, but new particles were formed also in the reference chamber. Finally, in the remaining 3 experiments (23%) we could not observe new particle formation. New particles may have been formed and may have not grown to detectable sizes in these experiments. The formed particles grew to sizes around 20-25 nm after 5 hours, with an estimated initial growth rate ranging from 3 to 11 nm h^{-1} . These results suggest that the presence of ammonia, at least at the high levels used in our study, allowed almost half of the time the formation and growth of particles that would not be formed otherwise. In one quarter of the cases ammonia increased significantly the nucleation rate compared to the ambient conditions. Finally, in the last quarter of the cases the high ammonia levels did not cause nucleation and growth to detectable sizes.

 We should note that we did not observe ammonium nitrate formation in any of our experiments despite the high ammonia levels. This is probably due to the combination of relatively low nitric acid levels and high temperatures during our study. This means that ammonium nitrate was not involved in the formation and growth processes in this environment.

 An aerosol dynamics model was used to estimate the *J*¹ nucleation rate constrained by the measured aerosol number distribution and mass concentrations. The nucleation rate in the 375 perturbed chamber ranged from 500 cm⁻³ h⁻¹ up to 25000 cm⁻³ h⁻¹. Coupled with the estimated sulfuric acid concentrations these rates are in general consistent (within one order of magnitude) with the CLOUD measurements for the nucleation rates in the sulfuric acid-ammonia-water system.

 Experiments in which new particles formation was observed in both chambers show one of the advantages of using a dual chamber system in such experiments. The use of the reference chamber can help verify if the conducted perturbation was responsible for the observed change. Future experiments with this system should include measurements of the sub-10 nm part of the aerosol size distribution and accurate measurements of the NH³ concentration.

 Data and code availability. The laboratory results and the aerosol dynamics code are available from the authors (spyros@chemeng.upatras.gr).

 Canagaratna, M. R., Jimenez, J. L., Kroll, J. H., Chen, Q., Kessler, S. H., Massoli, P., Hildebrandt Ruiz, L., Fortner, E., Williams, L. R., Wilson, K. R., Surratt, J. D., Donahue, N. M., Jayne, J. T., and Worsnop, D. R.: Elemental ratio measurements of organic compounds using aerosol mass spectrometry: characterization, improved calibration, and implications, Atmos. Chem. Phys., 15, 253–272, https://doi.org/10.5194/acp-15-253-2015, 2015. Capaldo, K. P., Kasibhatla, P., and Pandis, S. N.: Is aerosol production within the remote marine boundary layer sufficient to maintain observed concentrations?, J. Geophys. Res., 104, 3483–3500, https://doi.org/10.1029/1998JD100080, 1999. Demore, W. B., Sander, S. P., Golden, D. M., Hampson, R. F., Kurylo, M. J., Howard, C. J., Ravishankara, A. R., Kolb, C. E., and Molina, M. J.: Chemical Kinetics and Photochemical Data for Use in Stratospheric Modeling Evaluation Number 12 NASA Panel for Data Evaluation, 1997. Van Dingenen, R. and Raes, F.: Determination of the condensation accommodation coefficient of sulfuric acid on water-sulfuric acid aerosol, Aerosol Sci. Technol., 15, 93–106, https://doi.org/10.1080/02786829108959516, 1991. Dinoi, A., Weinhold, K., Wiedensohler, A., and Contini, D.: Study of new particle formation events in southern Italy, Atmos. Environ., 244, 117920, https://doi.org/10.1016/ j.atmosenv.2020.117920, 2021. Ehn, M., Thornton, J. A., Kleist, E., Sipilä, M., Junninen, H., Pullinen, I., Springer, M., Rubach, F., Tillmann, R., Lee, B., Lopez-Hilfiker, F., Andres, S., Acir, I. H., Rissanen, M., Jokinen, T., Schobesberger, S., Kangasluoma, J., Kontkanen, J., Nieminen, T., Kurtén, T., Nielsen, L. B., Jørgensen, S., Kjaergaard, H. G., Canagaratna, M., Maso, M. D., Berndt, T., Petäjä, T., Wahner, A., Kerminen, V. M., Kulmala, M., Worsnop, D. R., Wildt, J., and Mentel, T. F.: A large source of low-volatility secondary organic aerosol, Nature, 506, 476–479, https://doi.org/10.1038/nature13032, 2014. Glasoe, W. A., Volz, K., Panta, B., Freshour, N., Bachman, R., Hanson, D. R., McMurry, P. H., and Jen, C.: Sulfuric acid nucleation: An experimental study of the effect of seven bases, J. Geophys. Res. Atmos., 120, 1933–1950, https://doi.org/10.1002/2014JD022730, 2015. Haywood, J. and Boucher, O.: Estimates of the direct and indirect radiative forcing due to tropospheric aerosols: A review, Rev. Geophys., 38, 513–543, https://doi.org/10.1029/ 1999RG000078, 2000.

- E., and Mihalopoulos, N.: Particle size distributions in the Eastern Mediterranean troposphere, Atmos. Chem. Phys., 8, 6729–6738, https://doi.org/10.5194/acp-8-6729- 2008, 2008. Kalivitis, N., Kerminen, V. M., Kouvarakis, G., Stavroulas, I., Tzitzikalaki, E., Kalkavouras, P., Daskalakis, N., Myriokefalitakis, S., Bougiatioti, A., Manninen, H. E., Roldin, P., Petäjä, T., Boy, M., Kulmala, M., Kanakidou, M., and Mihalopoulos, N.: Formation and growth of atmospheric nanoparticles in the eastern Mediterranean: Results from long-term
- measurements and process simulations, Atmos. Chem. Phys., 19, 2671–2686, https://doi.org/10.5194/acp-19-2671-2019, 2019.
- Kalkavouras, P., Bossioli, E., Bezantakos, S., Bougiatioti, A., Kalivitis, N., Stavroulas, I., Kouvarakis, G., Protonotariou, A. P., Dandou, A., Biskos, G., Mihalopoulos, N., Nenes, A., and Tombrou, M.: New particle formation in the southern Aegean Sea during the Etesians: Importance for CCN production and cloud droplet number, Atmos. Chem. Phys., 17, 175–192, https://doi.org/10.5194/acp-17-175-2017, 2017.
- Kalkavouras, P., Bougiatioti, A., Grivas, G., Stavroulas, I., Kalivitis, N., Liakakou, E., Gerasopoulos, E., Pilinis, C., and Mihalopoulos, N.: On the regional aspects of new particle formation in the Eastern Mediterranean: A comparative study between a background and an urban site based on long term observations, Atmos. Res., 239, 104911, https://doi.org/10.1016/j.atmosres.2020.104911, 2020.
- Kaltsonoudis, C., Jorga, S. D., Louvaris, E., Florou, K., and Pandis, S. N.: A portable dual-smog- chamber system for atmospheric aerosol field studies, Atmos. Meas. Tech., 12, 2733–2743, https://doi.org/10.5194/amt-12-2733-2019, 2019.
- Kerminen, V. M., Petäjä, T., Manninen, H. E., Paasonen, P., Nieminen, T., Sipilä, M., Junninen, H., Ehn, M., Gagné, S., Laakso, L., Riipinen, I., Vehkam̈ aki, H., Kurten, T., Ortega, I. K., Dal Maso, M., Brus, D., Hyv̈arinen, A., Lihavainen, H., Leppä, J., Lehtinen, K. E. J.,
- Mirme, A., Mirme, S., Hõrrak, U., Berndt, T., Stratmann, F., Birmili, W., Wiedensohler,
- A., Metzger, A., Dommen, J., Baltensperger, U., Kiendler-Scharr, A., Mentel, T. F., Wildt,
- J., Winkler, P. M., Wagner, P. E., Petzold, A., Minikin, A., Plass-Dülmer, C., Pöschl, U., Laaksonen, A., and Kulmala, M.: Atmospheric nucleation: Highlights of the EUCAARI project and future directions, Atmos. Chem. Phys., 10, 10829–10848, https://doi.org/10.5194/acp-10-10829-2010, 2010.

 Kirkby, J., Curtius, J., Almeida, J., Dunne, E., Duplissy, J., Ehrhart, S., Franchin, A., Gagné, S., Ickes, L., Kürten, A., Kupc, A., Metzger, A., Riccobono, F., Rondo, L., Schobesberger, S., Tsagkogeorgas, G., Wimmer, D., Amorim, A., Bianchi, F., Breitenlechner, M., David, A., Dommen, J., Downard, A., Ehn, M., Flagan, R. C., Haider, S., Hansel, A., Hauser, D., Jud, W., Junninen, H., Kreissl, F., Kvashin, A., Laaksonen, A., Lehtipalo, K., Lima, J., Lovejoy, E. R., Makhmutov, V., Mathot, S., Mikkilä, J., Minginette, P., Mogo, S., Nieminen, T., Onnela, A., Pereira, P., Petäjä, T., Schnitzhofer, R., Seinfeld, J. H., Sipilä, M., Stozhkov, Y., Stratmann, F., Tomé, A., Vanhanen, J., Viisanen, Y., Vrtala, A., Wagner, P. E., Walther, H., Weingartner, E., Wex, H., Winkler, P. M., Carslaw, K. S., Worsnop, D. R., Baltensperger, U., and Kulmala, M.: Role of sulphuric acid, ammonia and galactic cosmic rays in atmospheric aerosol nucleation, Nature, 476, 429–433, https://doi.org/ 10.1038/nature10343, 2011. Kopanakis, I., Chatoutsidou, S. E., Torseth, K., Glytsos, T., and Lazaridis, M.: Particle number size distribution in the eastern Mediterranean: Formation and growth rates of ultrafine airborne atmospheric particles, Atmos. Environ., 77, 790–802, https://doi.org/10.1016/ j.atmosenv.2013.05.066, 2013. Kostenidou, E., Lee, B. H., Engelhart, G. J., Pierce, J. R., and Pandis, S. N.: Mass spectra deconvolution of low, medium, and high volatility biogenic secondary organic aerosol, Environ. Sci. Technol., 43, 4884–4889, https://doi.org/10.1021/es803676g, 2009. Kreidenweis, S. M., Yin, F., Wang, S.-C., Grosjean, D., Flagan, R. C., and Seinfeld, J. H.: Aerosol formation during photooxidation of organosulfur species, Atmos. Environ. Part A. Gen. Top., 25, 2491–2500, https://doi.org/10.1016/0960-1686(91)90165-4, 1991. Kulmala, M., Riipinen, I., Sipila, M., Manninen, H. E., Petaja, T., Junninen, H., Maso, M. D., Mordas, G., Mirme, A., Vana, M., Hirsikko, A., Laakso, L., Harrison, R. M., Hanson, I., Leung, C., Lehtinen, K. E. J., and Kerminen, V.-M.: Toward Direct Measurement of Atmospheric Nucleation, Science, 318, 89–92, https://doi.org/10.1126/science.1144124, 2007. Kulmala, M., Petäjä, T., Nieminen, T., Sipilä, M., Manninen, H. E., Lehtipalo, K., Dal Maso, M., Aalto, P. P., Junninen, H., Paasonen, P., Riipinen, I., Lehtinen, K. E. J., Laaksonen, A., and Kerminen, V. M.: Measurement of the nucleation of atmospheric aerosol particles, Nat. Protoc., 7, 1651–1667, https://doi.org/10.1038/nprot.2012.091, 2012.

 Laaksonen, A., Pirjola, L., Kulmala, M., Wohlfrom, K. H., Arnold, F., and Raes, F.: Upper tropospheric SO² conversion into sulfuric acid aerosols and cloud condensation nuclei, J. Geophys. Res. Atmos., 105, 1459–1469, https://doi.org/10.1029/1999JD900933, 2000. Manninen, H. E., Nieminen, T., Asmi, E., Gagné, S., Häkkinen, S., Lehtipalo, K., Aalto, P., Vana, M., Mirme, A., Mirme, S., Hõrrak, U., Plass-Dülmer, C., Stange, G., Kiss, G., Hoffer, A., Töro, N., Moerman, M., Henzing, B., De Leeuw, G., Brinkenberg, M., Kouvarakis, G. N., Bougiatioti, A., Mihalopoulos, N., O'Dowd, C., Ceburnis, D., Arneth, A., Svenningsson, B., Swietlicki, E., Tarozzi, L., Decesari, S., Facchini, M. C., Birmili, W., Sonntag, A., Wiedensohler, A., Boulon, J., Sellegri, K., Laj, P., Gysel, M., Bukowiecki, N., Weingartner, E., Wehrle, G., Laaksonen, A., Hamed, A., Joutsensaari, J., Petäjä, T., Kerminen, V. M., and Kulmala, M.: EUCAARI ion spectrometer measurements at 12 European sites-analysis of new particle formation events, Atmos. Chem. Phys., 10, 7907– 7927, https://doi.org/10.5194/acp-10-7907-2010, 2010. McFiggans, G., Bale, C. S. E., Ball, S. M., Beames, J. M., Bloss, W. J., Carpenter, L. J., Dorsey, J., Dunk, R., Flynn, M. J., Furneaux, K. L., Gallagher, M. W., Heard, D. E., Hollingsworth, A. M., Hornsby, K., Ingham, T., Jones, C. E., Jones, R. L., Kramer, L. J., Langridge, J. M., Leblanc, C., LeCrane, J. P., Lee, J. D., Leigh, R. J., Longley, I., Mahajan, A. S., Monks, P. S., Oetjen, H., Orr-Ewing, A. J., Plane, J. M. C., Potin, P., Shillings, A. J. L., Thomas, F., Von Glasow, R., Wada, R., Whalley, L. K., and Whitehead, J. D.: Iodine-mediated coastal particle formation: An overview of the Reactive Halogens in the Marine boundary layer (RHaMBLe) Roscoff coastal study, Atmos. Chem. Phys., 10, 2975–2999, https://doi.org/10.5194/acp-10-2975-2010, 2010. Merikanto, J., Spracklen, D. V., Mann, G. W., Pickering, S. J., and Carslaw, K. S.: Impact of nucleation on global CCN, Atmos. Chem. Phys., 9, 8601–8616, https://doi.org/ 10.5194/acp-9-8601-2009, 2009. Mohr, C., Thornton, J. A., Heitto, A., Lopez-Hilfiker, F. D., Lutz, A., Riipinen, I., Hong, J., Donahue, N. M., Hallquist, M., Petäjä, T., Kulmala, M., and Yli-Juuti, T.: Molecular identification of organic vapors driving atmospheric nanoparticle growth, Nat. Commun., 10, 1–7, https://doi.org/10.1038/s41467-019-12473-2, 2019. Patoulias, D., Fountoukis, C., Riipinen, I., Asmi, A., Kulmala, M., and Pandis, S. N.: Simulation of the size-composition distribution of atmospheric nanoparticles over Europe, Atmos.

- 1998.
- Weber, R. J., Marti, J. J., Mcmurry, P. H., Eisele, F. L., Tanner, D. J., And Jefferson, A.: Measured atmospheric new particle formation rates: implications for nucleation mechanisms, Chem. Eng. Commun., 151, 53–64, https://doi.org/10.1080/00986449608936541, 1996.
- Yao, L., Garmash, O., Bianchi, F., Zheng, J., Yan, C., Kontkanen, J., Junninen, H., Mazon, S. B., Ehn, M., Paasonen, P., Sipilä, M., Wang, M., Wang, X., Xiao, S., Chen, H., Lu, Y., Zhang, B., Wang, D., Fu, Q., Geng, F., Li, L., Wang, H., Qiao, L., Yang, X., Chen, J., Kerminen,

 V.-M., Petäjä, T., Worsnop, D. R., Kulmala, M., and Wang, L.: Atmospheric new particle formation from sulfuric acid and amines in a Chinese megacity, Science, 361, 278–281, https://doi.org/10.1126/science.aao4839, 2018. Yli-Juuti, T., Nieminen, T., Hirsikko, A., Aalto, P. P., Asmi, E., Hõrrak, U., Manninen, H. E., Patokoski, J., Dal Maso, M., Petäjä, T., Rinne, J., Kulmala, M., and Riipinen, I.: Growth rates of nucleation mode particles in Hyytiälä during 2003-2009: Variation with particle size, season, data analysis method and ambient conditions, Atmos. Chem. Phys., 11, 12865–12886, https://doi.org/10.5194/acp-11-12865-2011, 2011. Zhao, Y., Hennigan, C. J., May, A. A., Tkacik, D. S., De Gouw, J. A., Gilman, J. B., Kuster, W. C., Borbon, A., and Robinson, A. L.: Intermediate-volatility organic compounds: A large source of secondary organic aerosol, Environ. Sci. Technol., 48, 13743–13750, https://doi.org/10.1021/es5035188, 2014. Zhu, Y., Li, K., Shen, Y., Gao, Y., Liu, X., Yu, Y., Gao, H., and Yao, X.: New particle formation in the marine atmosphere during seven cruise campaigns, Atmos. Chem. Phys., 19, 89– 113, https://doi.org/10.5194/acp-19-89-2019, 2019.

657 **Table 1**: Initial conditions in the conducted experiments.

658

^a 659 Experiments illuminated by natural sunlight

^b⁶⁶⁰ Experiment conducted at night

^c 661 Estimated concentration in the perturbation chamber.

662

663

664

665

666

667

669

- 670 **Table 2:** Nucleation time, nucleation rate and condensation rate in the experiments were NPF was
- 671 observed in the perturbed chamber.

672

^a 673 Experiments illuminated by natural sunlight

^b⁶⁷⁴ Experiment conducted at night

675

676

- 678 **Table 3:** Nucleation time, nucleation rate and condensation rate in the experiments that NPF was
- 679 observed in the perturbed chamber.

680

^a 681 Experiments illuminated by natural sunlight

682 b Experiment conducted at night

683

684 685

686

687

 Figure 1: Wall loss corrected number concentration *N*9 in the three different types of experiments, a) NPF and growth only in the perturbed chamber (Exp. A1), b) NPF and growth in both chambers (Exp. B1) and c) no NPF observed (Exp. C1). The dashed line marks the time that ammonia was 717 injected in the perturbed chamber. At t=0 both chambers were illuminated with UV light.

Figure 2: Wall loss corrected measured number distributions in the two chambers for Exp. A1.

-
-

Figure 3: Wall loss corrected measured number distributions in the two chambers for Exp. B1.

-
-

-
-

 Figure 5: The wall loss corrected mass concentration of a) organics, b) sulfate, c) ammonium and d) nitrate in the control (blue dots) and perturbed (red dots) chamber during Exp. B1. The purple shades region represents the time that the chambers were under UV illumination.

-
-

-
-
-

 Figure 6: Measured and simulated number, surface and volume concentration in the perturbed chamber after turning UV lights on for Exp. A1. The error bars in the measured values are calculated from the uncertainty in the particle wall loss correction and represent two standard deviations.