**New particle formation and growth during summer in an urban environment: a dual chamber study**

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**Abstract**

Nucleation and subsequent growth are a major source of new particles in many environments, but the pollutants involved, and the details of the corresponding processes are still under debate. While sulfuric acid has a major role in new particle formation under a lot of conditions, the role of ammonia, amines and organic vapors is less clear. In most continental areas new particle formation is quite frequent especially in relatively clean, sunny days when there is some sulfur dioxide available. In parts of Eastern Mediterranean even if all the previous requirements are satisfied, new particle formation events are relatively rare during summertime.

In this work, we take advantage of this unexpected low new particle formation frequency in Greece and use a dual atmospheric simulation chamber system with starting point ambient air in an effort to gain insights about the chemical species that is limiting nucleation in this area. A potential nucleation precursor, ammonia, was added in one of the chambers while the other one was used as a reference. Three different types of outcomes were observed: new particle formation only in the perturbed chamber, new particle formation in both chambers, and no observed new particle formation. The addition of ammonia assisted in new particle formation in almost 50% of

27  the conducted experiments. The growth rate of the newly formed particles ranged from 3 – 11 nm

28  h$^{-1}$ with particles reaching a diameter of 20-25 nm after a few hours. The nucleation rate was

29  estimated using an aerosol dynamics model and was found to be in the range of 500 to 25000

30  particles cm$^{-3}$ h$^{-1}$ for the different experiments. These results support the hypothesis that ammonia

31  at levels of several ppb can accelerate or even cause new particle formation at least in the

32  environment of the Eastern Mediterranean.

## 1. Introduction

35  Atmospheric aerosol can be produced from many natural or anthropogenic sources and

36  plays a significant role in Earth's climate and in public health (Haywood and Boucher, 2000; Pope

37  et al., 2002). Aerosols can affect climate either by scattering and absorbing incoming solar

38  radiation (direct effect) or by acting as cloud condensation nuclei (CCN) thus affecting reflectivity

39  and lifetimes of clouds (indirect effect). New particle formations (NPF) through nucleation of low-

40  volatility vapors can be an important source of atmospheric aerosols and is responsible for close

41  to 50% of the global CCN (Merikanto et al., 2009). Newly formed particles either grow to larger

42  sizes through condensation or are scavenged by larger preexisting particles through coagulation.

43  Self-coagulation is another growth process for the newly formed particles. The competition

44  between these processes determines how many of those new particles will grow to become CCN

45  and how fast this will happen. NPF has been observed in many areas around the world including

46  all types of environments (urban, rural, forests, remote, marine, etc.) (Kulmala et al., 2007;

47  Kerminen et al., 2010; Wang et al., 2017; Yao et al., 2018; Zhu et al., 2019; Saha et al., 2019).

48  Previous studies have underlined the importance of sulfuric acid for NPF in most

49  environments (Jaecker-Voirol and Mirabel, 1989; Weber et al., 1996; Laaksonen et al., 2000;

50  Sipila et al., 2010). Additional studies have shown the importance of ammonia and amines as

51  vapors that can accelerate the nucleation rate of sulfuric acid with water by stabilizing the initial

52  clusters of sulfuric acid. (Weber et al., 1998; Kirkby et al., 2011; Jen et al., 2014; Glasoe et al.,

53  2015). Low and extremely low volatility organic vapors play a major role in the growth of the new

54  particles and may be also participating in the nucleation process itself  (Yli-Juuti et al., 2011; Zhao

et al., 2014; Ehn et al., 2014; Mohr et al., 2019). In marine environments iodine compounds have been identified as vapors that can form new particles (McFiggans et al., 2010; Sipilä et al., 2016; He et al., 2021) Wang et al. (2020) have recently reported fast growth rates of newly formed particles at some atmospheric conditions due to the condensation of ammonium nitrate. The preexisting aerosol (condensation sink), the availability of gaseous precursors and the meteorological conditions all affect the intensity and frequency of NPF events in the atmosphere.

Extensive monitoring of NPF events has taken place in many sites in Europe (Manninen et al., 2010; Dinoi et al., 2021)  and  the eastern Mediterranean (Pikridas et al., 2012; Berland et al., 2017; Kalkavouras et al., 2017; Kalivitis et al., 2019; Hussein et al., 2020; Brilke et al., 2020). Siakavaras et al. (2016) reported frequent nucleation events in Thessaloniki, a major urban center in northern Greece. On the other hand the nucleation frequency in southern Greece is relatively low (compared to central and northern Europe) especially during the summer (Kalivitis et al., 2008; 2019). Kopanakis et al. (2013) observed nucleation events only in 13 out of the 157 days of measurements in the Akrotiri station, in Crete. Kalkavouras et al. (2020) reported a relatively low 20% nucleation frequency during the summer in Finokalia, Crete. Particle size distribution measurements in five stations in Greece (Athens, Patras, Thessaloniki and Finokalia) during the summer of 2012, showed low NPF frequency in Patras and Finokalia (Vratolis et al., 2019). Pikridas et al. (2012) provided evidence that ammonia or amines may be the missing reactants that are responsible for the lack of nucleation in this sunny relatively clean area with available sulfur dioxide.

In this work we test the hypothesis that in an environment such as the Eastern Mediterranean during the summer, in which despite the high sunlight intensity, rapid photochemistry, moderate to low particle levels, always available sulfur dioxide and reasonable levels of both biogenic and anthropogenic VOCs, nucleation events are rare due to the relatively low ammonia levels. The experiments took place during summer in Patras, Greece in an environment with low nucleation and growth frequency (Patoulias et al., 2018; Vratolis et al., 2019) using a dual chamber system. The use of this innovative experimental set up, in a location in which nucleation is infrequent, allows us to perturb the atmosphere (at least a few cubic meters of it) in order to identify the reactant that is limiting the new particle formation. The use of two chambers to correct for various complications arising from these challenging measurements is an

85 additional novelty of this work. Both chambers were filled with ambient air, ammonia was added

86 to one of them, and the evolution of the aerosol was followed in both chambers.

87

88 **2. Methods**

89 **2.1 Dual chamber system**

90      A dual chamber system was deployed as part of the 2019 summer PANACEA

91 (PANhellenic infrastructure for Atmospheric Composition and climatE change) campaign in

92 Patras, Greece. Measurements were conducted in the outskirts of Patras (population 200.000

93 people) between July 15 until August 15, 2019, in the Institute of Chemical Engineering Sciences

94 (ICE-HT), approximately 8 km from the city center. The dual chamber system consisted of two

95 identical Teflon chambers (1.3 m$^3$ each) located inside a structure that included the chambers and

96 five panels of UV lights used for illumination purposes ($J_{NO2}$=0.25 min$^{-1}$). Ammonia was added in

97 one of the chambers (perturbation chamber) while the other one was used as the reference. The top

98 of the structure can be removed, and natural sunlight was used if the weather conditions allowed

99 it. Details about the design and testing of the dual chamber system can be found in Kaltsonoudis

100 et al. (2019).

101      The major difference between the conditions in the reference chamber and the ambient air

102 is that the chamber has a little lower concentrations due to losses in the sampling system. The UV

103 light is also a different in some experiments in which artificial light was used. Other experiments

104 used natural sunlight, so this was not an issue. Other differences include the interactions of the

105 pollutants inside the reference chamber with the walls of the chamber (for example losses of

106 particles but also some vapors to the walls during the experiment). Finally, the pollutant levels in

107 ambient air in the site may change as the wind brings new air masses to the area, while the air mass

108 inside the reference chamber remains the same as that present in the site at the time of the filling

109 of the chamber.

110

111 **2.2 Experimental procedure**

112      Before the beginning of each experiment both chambers were flushed with ambient air for

113 approximately 2 hours. The main purpose of the flushing is the conditioning of the chambers and

114 the sampling lines to the environmental conditions and composition thus minimizing losses of

115 volatile or semivolatile compounds to the walls of the system. During this preparation period both

116 chambers were swept for 20 min using an ionizer fan (Dr Schneider PC, Model SL-001) to reduce

117 the charges on the chamber walls thus reducing the particle wall losses (Jorga et al., 2020). After

118 the chambers were ready, they were filled with ambient air using a metal-bellows pump (Senior

119 Aerospace, model MB-302). The concentrations of pollutants in both the gas and particulate phases

120 were then characterized for one hour. A 0.25 in copper tube was used for the particle phase

121 measurements and a 0.25 in PTFE tube was used for the gas measurements. The instruments were

122 located inside a room next to the chambers to avoid their exposure to high temperatures. The

123 distance from the chambers to the instruments was approximately 4 m. An automated valve was

124 used to alternate sampling between the two chambers. The valve was synchronized with the

125 sampling periods of the various instruments and sufficient time was allowed between each

126 sampling change to flush any remaining air from the previous measurement.

127 After the characterization phase, ammonia was injected through a heated line into the

128 perturbation chamber using a glass syringe. The concentration of the injected ammonia was

129 estimated using the volume of the chamber and the amount of liquid ammonia injected and it varied

130 from 20 to 200 ppb. These estimates are probably upper limits because losses of ammonia in the

131 inlet system and the walls of the chamber are expected Even though the ammonia levels used in

132 this study are relatively high, they are still in the range of observed concentrations in the

133 atmosphere. For example, Dammers et al. (2017) measured ammonia concentrations in the

134 Netherlands up to 300 ppb. The fact that we have several measurements at concentrations higher

135 than those used by Kirkby et al. (2011), but still relevant to the atmosphere is a nice feature of this

136 work. After the ammonia injection, the top cover of the system was removed, and the chambers

137 were exposed to natural sunlight. If the wind speed is high, even if only the top cover is off, the

138 chambers may be destroyed by the wind. So, during these conditions the chamber system was kept

139 completely covered and the UV lights were used instead of natural sunlight for the corresponding

140 experiment.

141 At the end of the experiment, ammonium sulfate seeds were injected into both chambers

142 to measure the size dependent particle wall-loss rate constants using the method described in Wang

143 et al. (2018). After the end of the particle wall-losses period both chambers were flushed once

144 again with ambient air for approximately 2 hours, to remove the ammonium sulfate and any

145 remaining pollutants and to prepare them for the next experiment. The potential interactions of our

146 experimental system with the chamber walls are the reason for the use of the second (reference)

147  chamber. Any interactions will also be present there and will be observable and therefore we can

148  correct for them. Desorption of ammonia from the walls was tested with blank experiments the

149  following day from a perturbation experiment. The system was filled with ambient air, with no

150  addition of ammonia and the system response was tested. In all the blank experiments we did not

151  observe any nucleation in the perturbed chamber due to the ammonia that had been added in past

152  experiments.

153

154  **2.3 Instrumentation**

155      The chemical composition of the aerosol was monitored using a High-Resolution Time-of-

156  Flight Aerosol Mass Spectrometer (HR-ToF-AMS) from Aerodyne Research Inc. Two Scanning

157  Mobility Particle Sizers (SMPS) were used to measure the number size distributions from 9-160

158  nm (classifier model 3080, CPC model 3775) and from 14-730 nm (classifier model 3080, CPC

159  model 3025A) respectively. The sample flow was dried before reaching the AMS and SMPS

160  systems using a Nafion dryer. A suite of gas monitors was used to measure the concentrations of

161  $NO_x$ (Teledyne 201E/501), $SO_2$ (Thermo Scientific Model 43i), and $O_3$ (Teledyne 400E).

162

163  **2.4 Aerosol dynamics model**

164      A zero-dimensional aerosol dynamic model was used for the simulation of nucleation,

165  condensation and coagulation inside the perturbation chamber (Capaldo et al., 1999). The

166  multicomponent aerosol size distribution is described using 270 size sections covering the

167  diameter range from 1 nm to 1 μm. The aerosol components include sulfate, ammonium, organics,

168  and others with the latter assumed to be non-volatile and inert during the few hours of the

169  simulation period.

170      The condensation rate of $H_2SO_4$ to a particle of diameter $D_p$ is described using the modified

171  form of the Fuchs-Sutugin equation (Hegg et al., 1991; Kreidenweis et al., 1991) given by:

172

173  $$J = 2\pi D D_p F(Kn) A(P - P_o) \tag{1}$$

174  where $D$ is the diffusivity of the vapor air (set to 0.1 cm$^2$ s$^{-1}$ in this application), $Kn$ is the Knudsen

175  number (that is the ratio of the air mean free path to the particle radius), $F(Kn)$ is a coefficient

176  correcting for free molecular effects given by:

177  $$F(Kn) = \frac{1+Kn}{1+1.71Kn+1.33Kn^2} \tag{2}$$

178    and *A* is a coefficient correcting for the interfacial mass transport limitations described by the

179    accommodation $\alpha_e$,

180    $$A = \left[1 + 1.33 Kn \, F(Kn) \left(\frac{1}{a_e} - 1\right)\right]^{-1} \qquad (3)$$

181    Finally, *P* is the bulk $H_2SO_4$ vapor partial pressure and $P_o$ is its partial pressure at the particle

182    surface. An accommodation coefficient of 0.02 for the condensation of $H_2SO_4$ on the aerosol

183    particles is assumed (Van Dingenen and Raes, 1991). The vapor pressure of $H_2SO_4$ at the aerosol

184    surface can be estimated from the data of Bolsaitis and Elliott (1990). Values less than $10^{-3}$ ppt

185    were found for the conditions of our experiments and therefore the surface vapor pressure of $H_2SO_4$

186    in our mass transfer calculations was assumed to be zero.

187         Brownian coagulation between all particles is simulated solving the discrete coagulation

188    equation (Seinfeld and Pandis, 2016):

189    $$\frac{dN_k(t)}{dt} = \frac{1}{2}\sum_{j=1}^{k-1} K_{j,k-j} \, N_j \, N_{k-j} - N_k \sum_{j=1}^{\infty} K_{k,j} \, N_j \qquad k \geq 2 \qquad (4)$$

190    The generalized coagulation coefficient $K_{1,2}$ for the collision of two particles is calculated as:

191    $$K_{1,2} = 2\pi(D_1 + D_2)(D_{p1} + D_{p2})\beta \qquad (5)$$

192    where $D_1$, $D_2$ are the individual Brownian diffusion coefficient for the particles, $D_{p1}$, $D_{p2}$ are the

193    particle diameters and $\beta$ is the Fuchs correction factor (Fuchs, 1964). Because of the high

194    resolution of the size distribution, coagulation can be simulated accurately by directly calculating

195    the coagulation rates between each of the size sections and moving the particles to the

196    corresponding size bin.

197

198    **2.5 Data analysis**

199         The condensation sink (CS) is a metric of the ability of the pre-existing aerosol population

200    to remove vapors from the system by condensation. The CS values were calculated using the

201    aerosol distribution between 14-700 nm, as measured by the SMPS and the properties of sulfuric

202    acid as the condensing vapor. The CS is given by:

203    $$CS = 2\pi D \sum_i \beta_{mi} D_{pi} N_i \qquad (6)$$

204    where *D* is the diffusion coefficient of sulfuric acid, $\beta_m$ is the transition-regime correction factor,

205    $D_p$ the diameter of the particle and *N* the respective number concentration in each size bin of the

206    SMPS.

207        Using the initial $SO_2$ concentration in the perturbed chamber and the condensation sink we

208   can estimate the sulfuric acid concentration according to:

209 $$[H_2SO_4] = k_{OH} \frac{[SO_2][OH]}{CS} \tag{7}$$

210   where $k_{OH}$ is the reaction constant of $SO_2$ and OH which is equal to $8.5 \times 10^{-13}$ $cm^3$ molecule$^{-1}$ s$^{-1}$

211   at 298 K (Demore et al., 1997), $[SO_2]$ and $[OH]$ are the concentrations of sulphur dioxide and

212   hydroxyl radicals respectively, and CS is the condensational sink as calculated from Equation 7.

213   For the hydroxyl radical concentration, we assumed an average value of $5\times10^6$ molecules $cm^{-3}$.

214   Equation (7) is based on the assumptions that the only sulfuric acid source is the oxidation of $SO_2$

215   from OH radicals, its major sink is its condensation onto the aerosol surface, and the system is at

216   pseudo-steady state.

217

218   **3. Results and discussion**

219   **3.1 Initial conditions**

220        Thirteen perturbation experiments were conducted during the study (Table 1). Two of them

221   took place during the night using UV lights and the rest during midday. Natural sunlight was used

222   in two experiments while UV lights were used during the rest.

223        The main components of non-refractory $PM_1$ in the beginning of our experiments were

224   organics (46.6±6.5%) followed by sulfate (37.1±4.5%), ammonium (14.3±1.8%), nitrate

225   (1.5±0.5%) and chloride (0.5±0.4%). The average oxygen to carbon ratio (O:C) (Canagaratna et

226   al., 2015) in the chambers after filling them with ambient air was 0.68±0.1, indicating an already

227   oxidized OA. In order to check if any contamination was occurring during the flushing and filling

228   processes we calculated the theta angle (Kostenidou et al., 2009) between the organic mass spectra

229   of the ambient air and the OA in the two chambers. The theta angles were less than 6$^o$ in all

230   experiments, indicating negligible contamination during the filling process.

231        The initial concentration of the $SO_2$, $NH_3$, $O_3$ and $NO_x$ inside the chambers after the filling

232   process were approximately within 10% of their ambient values. The initial concentrations of these

233   gases in the two chambers differed by less than 3%. More than 70% of the ambient $PM_1$ was

234   transferred in the chambers in most experiments. The initial $PM_1$ levels were quite low ranging

235   from 0.6 to 4.2 µg m$^{-3}$. The atmosphere of Patras was quite clear during these experiments. The

236   initial conditions in all experiments are summarized in Table 1.

**3.2 New particle formation and growth**

The conducted experiments were classified in three different classes based on the observations of new particle formation in the two chambers. In class A experiments, nucleation and particle growth occurred in only the perturbed chamber, in class B nucleation and particle growth happened in both chambers and in class C when there was no detection of new particle formation in either chamber.

**3.2.1 Nucleation and growth only in the perturbed chamber**

Nucleation and growth of the new particles to sizes above 9 nm only in the perturbed and not in the reference chamber were observed in 6 out of the 13 performed experiments. Figure 1 depicts the particle number concentration $N_9$ ($D_p$>9 nm) after corrections for particle losses to the chamber walls and the sampling lines, inside the two chambers for a typical Class A experiment. During Exp. A1 the initial concentration of $SO_2$ was 0.7 ppb and of $O_3$ equal to 58 ppb. After the injection of ammonia (approximately 150 ppb) the UV lights were turned on (t=0 h) illuminating both chambers. The $N_9$ particle number concentration start increasing in the perturbed chamber approximately 1.5 h after the lights were turned on, reaching close to 4000 $cm^{-3}$, almost double its initial concentration value. The observed delay of the $N_9$ particle number concentration in the perturbed chamber is attributed to the time needed for the particles to grow to larger sizes so that they could be detected by our available instrumentation. The $N_9$ concentration in the control chamber remained within 5% of the initial levels. Figure 2 shows the measured number distributions in the two chambers after correction for particle losses. The formation and growth of the new particles in the perturbed chamber is evident. With a CS of 0.0026 $s^{-1}$ the $H_2SO_4$ concentration was calculated to be of the order of $3\times10^7$ molecules $cm^{-3}$. Assuming that nucleation started at t=0 the measured initial growth rate in the perturbed chamber was on average 4 nm $h^{-1}$. This rate is based on the time of growth of the nucleation mode to 15 nm. For this experiment 3.8 h were required for the growth of the new particles to 15 nm so the estimated growth rate is approximately 4 nm $h^{-1}$. The newly formed particles at the end of the experiment (after 5 h from the illumination) grew to approximately 20 nm. The experiment was stopped at that point because a significant fraction of the air in the two chambers had been sampled.

The estimated initial growth rates (refers to the average rate for the period between the start of the experiment and the time the particles reach a diameter of 9 nm) for the daytime experiments

268   A1, A2, A4, A5, and A6 varied from 3.5 to 8 nm $h^{-1}$ and were correlated with the estimated $H_2SO_4$

269   levels ($R^2=0.76$) (Table 2). Experiments A1, A5 and A6 had similar CS and $H_2SO_4$ levels and

270   resulted in similar initial growth rates (Table 2). The slope of the growth rate versus sulphuric acid

271   linear regression for the daytime experiments was 1.4 (nm $h^{-1}$ molecule$^{-1}$ s) and the intercept was

272   low and equal to 0.05 nm $h^{-1}$.

273        Experiment A3 was conducted during the night (the chambers were filled with ambient air

274   at 21:00 so it has relatively different behaviour than the rest. If this experiment is included in the

275   growth rate versus $H_2SO_4$ correlation, the $R^2$ drops to 0.27. This probably suggests that the

276   estimated OH levels are not accurate in this case and therefore the $H_2SO_4$ is probably also a lot

277   more uncertain than in the other experiments. Also, the growth process may be different with

278   organic vapors playing a more significant role.

279

280   **3.2.2 Nucleation and growth in both chambers**

281        In 4 (B1 to B4) out of the 13 perturbation experiments, new particle formation and growth

282   was observed in both chambers (Table 2). This suggests that the ambient air had already the

283   potential to form new particles without the addition of ammonia.

284        Figure 1b shows the $N_9$ concentration in the two chambers during Exp. B1. The initial

285   levels of $SO_2$ in both chambers were 0.8 ppb and approximately 150 ppb of ammonia were added

286   to the perturbation chamber. Half hour after the exposure of the chambers to UV light the $N_9$ in

287   both chambers started increasing with higher concentrations in the perturbed chamber. The number

288   concentration of the particles in the perturbed chamber reached close to 6000 cm$^{-3}$ almost three

289   times the initial levels. The concentration in the baseline chamber increased by approximately 50%

290   to 3000 cm$^{-3}$. The newly formed particles in the perturbed chamber at the end of the experiment

291   reached a mode diameter of 26 nm (Fig. 3) with an initial GR of 5.5 nm $h^{-1}$. The growth rate of the

292   particles in the reference chamber was only 10% lower than in the perturbed one (Fig. 4)

293   suggesting that the addition of ammonia probably significantly influenced the nucleation rate but

294   had a small effect on the growth rate.

295        We tested the hypothesis that the appearance of the new particles in the reference chamber

296   was due to a sampling error, caused by some cross contamination of the two samples as the same

297   sampling line was used. We compared the shape of aerosol size distributions in the two chambers.

298 The nucleation mode distribution in the reference chamber was wider (Figure S1), a strong

299 indication that these were different particle populations sampled by our system.

300     The condensation sink in the class B experiments was on average 0.0024 $s^{-1}$ quite similar

301 to the 0.0023 $s^{-1}$ in the A experiments so the pre-existing particle mass was quite similar in the two

302 classes of experiments. Also, the average $SO_2$ was practically the same (0.83 ppb for the B

303 experiments and 0.82 ppb for the A experiments). The average ammonium concentration for the

304 class A experiments was only 20% higher than that of Class B-experiments. The ammonium levels

305 in this area are determined to a large extent by the sulfate levels. Adding the similarity of the UV

306 intensity, it is clear that the major factors (sunlight, condensational sink, $SO_2$ availability) usually

307 determining nucleation rates were not the reason for the weak nucleation and growth in the

308 reference chamber in these experiments. The presence of sufficient ammonia levels is one of the

309 possible explanations for this behaviour.

310     The observed growth rates in these B-class experiments varied from 3.5 to 11.3 nm $h^{-1}$ and

311 were a little higher on average than those in the A group of experiments. Adding these four

312 experiments to the linear regression of the growth rate versus sulphuric acid reduced the $R^2$ to 0.43,

313 but the slope remained the same, while the intercept increased to 0.7 nm $h^{-1}$. These provide some

314 weak evidence of the involvement of more compounds, probably organics, in the growth of these

315 newly formed particles in this second group of experiments.

316

**3.2.2 Nucleation and growth not observed**

318     In three of the experiments C1-C3, we did not observe growth of new particles to the size

319 of 9 nm. It is still possible that there was nucleation, but the growth may have been too slow (less

320 than 2 nm $h^{-1}$). We were able to measure only particles larger than 9 nm. If the growth rate in those

321 experiments was less than 2 nm $h^{-1}$ then the particles would not be reaching the 9 nm threshold

322 during the experiment. Figure 1c depicts the $N_9$ number concentration during Exp. C1. The initial

323 concentration of $SO_2$ was 1.3 ppb in this experiment. The number concentration after correction

324 for particle losses remained constant at close to 2200 $cm^{-3}$ and the corresponding number

325 distributions changed little during the experiment (Fig. 4). Exp. C1 had the lowest initial levels of

326 ozone of all the experiments (Table 1), around 12 ppb, a factor of four lower than the average

327 concentration. The low $O_3$ levels were probably due to the highest $NO_x$ levels (27 ppb) in this air

328 mass.

329     Exp. C2 was conducted in the early evening (the chamber was filled with ambient air
330     around 19:00 LT) and the lowest detectable particle size for this experiment was 14 nm because
331     of technical difficulties. Finally, Exp. C3 had relatively low levels of sulfuric acid $(2.5\times10^7$
332     molecules cm$^{-3}$) compared to the rest of the experiments, a low estimated level of injected ammonia
333     (20 ppb) and natural sunlight.

334

335     **3.3 Particle composition**

336     The mass concentration of the major components of PM$_1$ (sulfate, organics, nitrate,
337     ammonium) in the two chambers, after correcting for chamber particle wall losses, remained
338     practically constant during all experiments. The corrected for wall losses mass concentration inside
339     both chambers during Exp. B1 are shown in Figure 5. Considering the uncertainty of the wall loss
340     correction, the maximum increase of the concentration of the corresponding secondary PM
341     components during the few hours of the experiments should have been a few percent or less. This
342     will be an important constraint for the analysis of these experimental results with the aerosol
343     dynamics model in the next section.

344

345     **3.4 Estimation of nucleation rate using an aerosol dynamics model**

346     We used our aerosol dynamics model to simulate the growth and coagulation of the
347     particles in the perturbed chamber assuming a nucleation rate. Our goal is to use the observations
348     to constrain the nucleation rate that could not be measured directly. The model uses as inputs the
349     temperature and relative humidity during the experiments and is initialized with the measured
350     particle number distribution at time zero. There are three adjustable parameters in the model:  the
351     duration of nucleation, the nucleation rate, and the condensation rate. Nucleation is assumed to
352     start at time zero and a constant nucleation rate is assumed for the duration of the event. This is a
353     necessary assumption given the available measurements. This constant rate is in practice an
354     average rate for the estimated duration of the event. The condensing components are assumed to
355     have practically zero vapor pressure. The three parameters were chosen so that the model
356     predictions were in good agreement with the observations of particle number concentration and
357     size distribution and also the mass concentration. The change in the condensation sink during the
358     experiments was modest (reduction 10-30%) however other important parameters like the

359  concentration of the species participating in the nucleation process were probably changing at the
360  same time.

361      Figure 6 shows the measured and the predicted particle number, surface and volume
362  concentrations in the perturbed chamber for Exp. A1. A nucleation event with rate equal to
363  $J_1$=9500 cm$^{-3}$ h$^{-1}$ and duration of 3 h together with a condensation rate of 3.2 ppt h$^{-1}$ was needed
364  to reproduce the observations. For much lower condensation rates the particles did not grow to
365  detectable sizes and for higher condensation rates the predicted PM mass increase was not
366  consistent with the small observed mass concentration change. We performed sensitivity analysis
367  around these central values and values of $J_1$=9500±600 cm$^{-3}$ h$^{-1}$ remained consistent with the
368  observations. The average errors during the simulation were 6% for the number concentration,
369  16% for the surface concentration and 17% for the mass concentration. Other effective nucleation
370  rates (e.g., $J_3$ or $J_9$) can be estimated with our approach, but this would require reconfiguration of
371  the model so that the size distribution would start at the corresponding diameter threshold.

372      The predicted and observed evolution of the aerosol number distributions are shown in
373  Figure S2. The differences for the smaller particle sizes are partially due to the losses of these
374  particles in the sampling and measurement systems.

375      Table 3 summarizes the estimated nucleation rates together with the corresponding
376  durations of the nucleation events and the required condensation rates for all experiments in groups
377  A and B in which nucleation and growth were observed. The estimated $J_1$ rates varied from 500 to
378  25000 cm$^{-3}$ h$^{-1}$. These values are between those in ambient measurements and those of the CLOUD
379  laboratody experiments (Fig. S3).

380      These results can be roughly compared to the CLOUD measurements for sulfuric acid-
381  ammonia nucleation (Kirkby et al., 2011) in the $2\times10^7$-$10^8$ molecules cm$^{-3}$ H$_2$SO$_4$ concentration
382  range that was estimated for our experiments. The CLOUD measurements for the highest ammonia
383  levels used suggested a $J_{1.7}$ rate of approximately 500 cm$^{-3}$ h$^{-1}$ for H$_2$SO$_4$ concentration equal to
384  $5\times10^7$ molecules cm$^{-3}$. For experiment A4 we estimated the same H$_2$SO$_4$ concentration and a
385  nucleation rate of 400±200 cm$^{-3}$ h$^{-1}$ (Fig. S3). While this agreement is probably fortuitous, overall,
386  our estimated nucleation rates are in general consistent (considering their uncertainties) with the
387  CLOUD measurements for the ammonia-sulfuric acid system assuming that the rate does not
388  increase further as ammonia increases above 1 ppb.

389

**4. Conclusions**

A dual chamber system was used to investigate the hypothesis that ammonia is often the limiting reactant for new particle formation in the Eastern Mediterranean, using a new experimental approach in one of the areas with the lowest new particle formation frequency in Europe during the summer Ambient air characterized by relatively aged air masses in southern Greece was used as the starting point of our experiments. Ammonia was added in one chamber while the second was used as a reference. Using two chambers adds to the novelty of this work, allowing for corrections due to the interactions between the chamber walls and the reacting gases and particles. To the best of our knowledge this is the first study that uses such an experimental set up for determining the role of a reactant, like ammonia, in new particle formation under realistic environmental conditions.

In 6 out of the 13 experiments (46%) the addition of ammonia led to the formation and then growth to detectable size (approximately 10 nm) of new particles, while no formation of particles was observed in the reference chamber. In another 4 experiments (31%) the addition of ammonia significantly enhanced the formation of new particles, but new particles were formed also in the reference chamber. Finally, in the remaining 3 experiments (23%) we could not observe new particle formation. New particles may have been formed and may have not grown to detectable sizes in these experiments. The formed particles grew to sizes around 20-25 nm after 5 hours, with an estimated initial growth rate ranging from 3 to 11 nm h$^{-1}$. These results suggest that the presence of ammonia, at least at the high levels used in our study, allowed almost half of the time the formation and growth of particles that would not be formed otherwise. In one quarter of the cases ammonia significantly increased the nucleation rate compared to the ambient conditions. Finally, in the last quarter of the cases the high ammonia levels did not cause nucleation and growth to detectable sizes.

We should note that we did not observe ammonium nitrate formation in any of our experiments despite the high ammonia levels. This is probably due to the combination of relatively low nitric acid levels and high temperatures during our study. This suggests that ammonium nitrate was not formed in the perturbed chamber after the ammonia injection and did not contribute to the particle growth in our experiments.

An aerosol dynamics model was used to estimate the $J_1$ nucleation rate constrained by the measured aerosol number distribution and mass concentrations. The used box model does not

421 directly include the ammonia concentration or a nucleation parameterization but is used instead to
422 provide estimates of the nucleation and growth rates that are consistent with the measurements.
423 New particle formation occurred even at the lowest ammonia levels (20 ppb) used in these
424 experiments. The nucleation rate in the perturbed chamber ranged from 500 $cm^{-3} h^{-1}$ up to 25000
425 $cm^{-3} h^{-1}$. Coupled with the estimated sulfuric acid concentrations these rates are in general
426 consistent (within one order of magnitude) with both ambient measurements and those of the
427 CLOUD lab experiments for the nucleation rates in the sulfuric acid-ammonia-water system.
428 Nucleation was observed even at the lower ammonia levels used in this work (20 ppb), and the
429 estimated nucleation rate was quite high. This result is applicable to environments with high
430 ammonia levels like the Netherlands or the Po Valley. The ammonia levels in these areas are often
431 similar to those in our experiments.

432 The two major new advances of this work are first the use of a new experimental technique
433 that allowed us to test the hypothesis of Pikridas et al. (2012) which was based on circumstantial
434 evidence (the ratio of ammonium to sulfate in the particles and the wind trajectories) and second
435 the results of the experiments that strongly support the hypothesis. A technique to estimate the
436 nucleation and growth rates from these data even without measurements of sub-10 nm particles is
437 an additional contribution.

438 Experiments in which new particles formation was observed in both chambers show one
439 of the advantages of using a dual chamber system in such experiments. The use of the reference
440 chamber can help verify if the conducted perturbation was responsible for the observed change.
441 Future experiments with this system should include measurements of the sub-10 nm part of the
442 aerosol size distribution and accurate measurements of the $NH_3$ concentration.

443

444 **Data and code availability.** The laboratory results and the aerosol dynamics code are available
445 from the authors (spyros@chemeng.upatras.gr).

446

447 **Supplement.**

448

449 **Author contributions.** SDJ performed the experiments, analyzed the results, and wrote the paper.
450 KF helped in the performance of the experiments. DP wrote the aerosol dynamics code. SNP was

451 responsible for the design of the study and the synthesis of the results and contributed to the writing
452 of the paper.

453

454 **Competing interests.** The authors declare that they have no conflict of interest.

455

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## 5. References

466 Berland, K., Rose, C., Pey, J., Culot, A., Freney, E., Kalivitis, N., Kouvarakis, G., Cerro, J. C.,
467        Mallet, M., Sartelet, K., Beckmann, M., Bourriane, T., Roberts, G., Marchand, N.,
468        Mihalopoulos, N., and Sellegri, K.: Spatial extent of new particle formation events over
469        the Mediterranean Basin from multiple ground-based and airborne measurements, Atmos.
470        Chem. Phys., 17, 9567–9583, https://doi.org/10.5194/acp-17-9567-2017, 2017.

471 Bolsaitis, P. and Elliott, J. F.: Thermodynamic activities and equilibrium partial pressures for
472        aqueous sulfuric acid solutions, J. Chem. Eng. Data, 35, 69–85,
473        https://doi.org/10.1021/je00059a022, 1990.

474 Brilke, S., Fölker, N., Müller, T., Kandler, K., Gong, X., Peischl, J., Weinzierl, B., and Winkler,
475        P. M.: New particle formation and sub-10 nm size distribution measurements during the
476        A-LIFE field experiment in Paphos, Cyprus, Atmos. Chem. Phys., 20, 5645–5656,
477        https://doi.org/10.5194/acp-20-5645-2020, 2020.

478 Canagaratna, M. R., Jimenez, J. L., Kroll, J. H., Chen, Q., Kessler, S. H., Massoli, P., Hildebrandt
479        Ruiz, L., Fortner, E., Williams, L. R., Wilson, K. R., Surratt, J. D., Donahue, N. M., Jayne,
480        J. T., and Worsnop, D. R.: Elemental ratio measurements of organic compounds using
481        aerosol mass spectrometry: characterization, improved calibration, and implications,

482    Atmos. Chem. Phys., 15, 253–272, https://doi.org/10.5194/acp-15-253-2015, 2015.

483  Capaldo, K. P., Kasibhatla, P., and Pandis, S. N.: Is aerosol production within the remote marine
484    boundary layer sufficient to maintain observed concentrations?, J. Geophys. Res. Atmos.,
485    104, 3483–3500, https://doi.org/10.1029/1998JD100080, 1999.

486  Dammers, E., Schaap, M., Haaima, M., Palm, M., Wichink Kruit, R. J., Volten, H., Hensen, A.,
487    Swart, D., and Erisman, J. W.: Measuring atmospheric ammonia with remote sensing
488    campaign: Part 1 – Characterisation of vertical ammonia concentration profile in the centre
489    of The Netherlands, Atmos. Environ., 169, 97–112, https://doi.org/10.1016/j.atmosenv.
490    2017.08.067, 2017.

491  Demore, W. B., Sander, S. P., Golden, D. M., Hampson, R. F., Kurylo, M. J., Howard, C. J.,
492    Ravishankara, A. R., Kolb, C. E., and Molina, M. J.: Chemical Kinetics and Photochemical
493    Data for Use in Stratospheric Modeling Evaluation Number 12 NASA Panel for Data
494    Evaluation, 1997.

495  Van Dingenen, R. and Raes, F.: Determination of the condensation accommodation coefficient of
496    sulfuric acid on water-sulfuric acid aerosol, Aerosol Sci. Technol., 15, 93–106,
497    https://doi.org/10.1080/02786829108959516, 1991.

498  Dinoi, A., Weinhold, K., Wiedensohler, A., and Contini, D.: Study of new particle formation
499    events in southern Italy, Atmos. Environ., 244, 117920,
500    https://doi.org/10.1016/j.atmosenv.2020.117920, 2021.

501  Ehn, M., Thornton, J. A., Kleist, E., Sipilä, M., Junninen, H., Pullinen, I., Springer, M., Rubach,
502    F., Tillmann, R., Lee, B., Lopez-Hilfiker, F., Andres, S., Acir, I. H., Rissanen, M., Jokinen,
503    T., Schobesberger, S., Kangasluoma, J., Kontkanen, J., Nieminen, T., Kurtén, T., Nielsen,
504    L. B., Jørgensen, S., Kjaergaard, H. G., Canagaratna, M., Maso, M. D., Berndt, T., Petäjä,
505    T., Wahner, A., Kerminen, V. M., Kulmala, M., Worsnop, D. R., Wildt, J., and Mentel, T.
506    F.: A large source of low-volatility secondary organic aerosol, Nature, 506, 476–479,
507    https://doi.org/10.1038/nature13032, 2014.

508  Glasoe, W. A., Volz, K., Panta, B., Freshour, N., Bachman, R., Hanson, D. R., McMurry, P. H.,
509    and Jen, C.: Sulfuric acid nucleation: An experimental study of the effect of seven bases,
510    J. Geophys. Res. Atmos., 120, 1933–1950, https://doi.org/10.1002/2014JD022730, 2015.

511  Haywood, J. and Boucher, O.: Estimates of the direct and indirect radiative forcing due to
512    tropospheric aerosols: A review, Rev. Geophys., 38, 513–543,

https://doi.org/10.1029/1999RG000078, 2000.

He, X.-C., Tham, Y. J., Dada, L., Wang, M., Finkenzeller, H., Stolzenburg, D., Iyer, S., Simon, M., Kürten, A., Shen, J., Rörup, B., Rissanen, M., Schobesberger, S., Baalbaki, R., Wang, D. S., Koenig, T. K., Jokinen, T., Sarnela, N., Beck, L. J., Almeida, J., Amanatidis, S., Amorim, A., Ataei, F., Baccarini, A., Bertozzi, B., Bianchi, F., Brilke, S., Caudillo, L., Chen, D., Chiu, R., Chu, B., Dias, A., Ding, A., Dommen, J., Duplissy, J., El Haddad, I., Gonzalez Carracedo, L., Granzin, M., Hansel, A., Heinritzi, M., Hofbauer, V., Junninen, H., Kangasluoma, J., Kemppainen, D., Kim, C., Kong, W., Krechmer, J. E., Kvashin, A., Laitinen, T., Lamkaddam, H., Lee, C. P., Lehtipalo, K., Leiminger, M., Li, Z., Makhmutov, V., Manninen, H. E., Marie, G., Marten, R., Mathot, S., Mauldin, R. L., Mentler, B., Möhler, O., Müller, T., Nie, W., Onnela, A., Petäjä, T., Pfeifer, J., Philippov, M., Ranjithkumar, A., Saiz-Lopez, A., Salma, I., Scholz, W., Schuchmann, S., Schulze, B., Steiner, G., Stozhkov, Y., Tauber, C., Tomé, A., Thakur, R. C., Väisänen, O., Vazquez-Pufleau, M., Wagner, A. C., Wang, Y., Weber, S. K., Winkler, P. M., Wu, Y., Xiao, M., Yan, C., Ye, Q., Ylisirniö, A., Zauner-Wieczorek, M., Zha, Q., Zhou, P., Flagan, R. C., Curtius, J., Baltensperger, U., Kulmala, M., Kerminen, V.-M., Kurtén, T., et al.: Role of iodine oxoacids in atmospheric aerosol nucleation, Science, 371, 589–595, 2021.

Hegg, D. A., Radke, L. F., and Hobbs, P. V.: Measurements of Aitken nuclei and cloud condensation nuclei in the marine atmosphere and their relation to the DMS-Cloud-climate hypothesis, J. Geophys. Res., 96, 18727, https://doi.org/10.1029/91JD01870, 1991.

Hussein, T., Atashi, N., Sogacheva, L., Hakala, S., Dada, L., Petäjä, T., and Kulmala, M.: Characterization of urban new particle formation in Amman-Jordan, Atmosphere (Basel)., 11, 79, https://doi.org/10.3390/atmos11010079, 2020.

Jaecker-Voirol, A. and Mirabel, P.: Heteromolecular nucleation in the sulfuric acid-water system, Atmos. Environ., 23, 2053–2057, https://doi.org/10.1016/0004-6981(89)90530-1, 1989.

Jen, C. N., McMurry, P. H., and Hanson, D. R.: Stabilization of sulfuric acid dimers by ammonia, methylamine, dimethylamine, and trimethylamine, J. Geophys. Res. Atmos., 119, 7502–7514, https://doi.org/10.1002/2014JD021592, 2014.

Jorga, S. D., Kaltsonoudis, C., Liangou, A., and Pandis, S. N.: Measurement of Formation Rates of Secondary Aerosol in the Ambient Urban Atmosphere Using a Dual Smog Chamber System, Environ. Sci. Technol., 54, 1336–1343, https://doi.org/10.1021/acs.est.9b03479,

544     2020.

545    Kalivitis, N., Birmili, W., Stock, M., Wehner, B., Massling, A., Wiedensohler, A., Gerasopoulos,
546        E., and Mihalopoulos, N.: Particle size distributions in the Eastern Mediterranean
547        troposphere, Atmos. Chem. Phys., 8, 6729–6738, https://doi.org/10.5194/acp-8-6729-
548        2008, 2008.

549    Kalivitis, N., Kerminen, V. M., Kouvarakis, G., Stavroulas, I., Tzitzikalaki, E., Kalkavouras, P.,
550        Daskalakis, N., Myriokefalitakis, S., Bougiatioti, A., Manninen, H. E., Roldin, P., Petäjä,
551        T., Boy, M., Kulmala, M., Kanakidou, M., and Mihalopoulos, N.: Formation and growth
552        of atmospheric nanoparticles in the eastern Mediterranean: Results from long-term
553        measurements and process simulations, Atmos. Chem. Phys., 19, 2671–2686,
554        https://doi.org/10.5194/acp-19-2671-2019, 2019.

555    Kalkavouras, P., Bossioli, E., Bezantakos, S., Bougiatioti, A., Kalivitis, N., Stavroulas, I.,
556        Kouvarakis, G., Protonotariou, A. P., Dandou, A., Biskos, G., Mihalopoulos, N., Nenes,
557        A., and Tombrou, M.: New particle formation in the southern Aegean Sea during the
558        Etesians: Importance for CCN production and cloud droplet number, Atmos. Chem. Phys.,
559        17, 175–192, https://doi.org/10.5194/acp-17-175-2017, 2017.

560    Kalkavouras, P., Bougiatioti, A., Grivas, G., Stavroulas, I., Kalivitis, N., Liakakou, E.,
561        Gerasopoulos, E., Pilinis, C., and Mihalopoulos, N.: On the regional aspects of new particle
562        formation in the Eastern Mediterranean: A comparative study between a background and
563        an urban site based on long term observations, Atmos. Res., 239, 104911,
564        https://doi.org/10.1016/j.atmosres.2020.104911, 2020.

565    Kaltsonoudis, C., Jorga, S. D., Louvaris, E., Florou, K., and Pandis, S. N.: A portable dual-smog-
566        chamber system for atmospheric aerosol field studies, Atmos. Meas. Tech., 12, 2733–2743,
567        https://doi.org/10.5194/amt-12-2733-2019, 2019.

568    Kerminen, V. M., Petäjä, T., Manninen, H. E., Paasonen, P., Nieminen, T., Sipilä, M., Junninen,
569        H., Ehn, M., Gagné, S., Laakso, L., Riipinen, I., Vehkaṁaki, H., Kurten, T., Ortega, I. K.,
570        Dal Maso, M., Brus, D., Hyṿarinen, A., Lihavainen, H., Leppä, J., Lehtinen, K. E. J.,
571        Mirme, A., Mirme, S., Hõrrak, U., Berndt, T., Stratmann, F., Birmili, W., Wiedensohler,
572        A., Metzger, A., Dommen, J., Baltensperger, U., Kiendler-Scharr, A., Mentel, T. F., Wildt,
573        J., Winkler, P. M., Wagner, P. E., Petzold, A., Minikin, A., Plass-Dülmer, C., Pöschl, U.,
574        Laaksonen, A., and Kulmala, M.: Atmospheric nucleation: Highlights of the EUCAARI

project and future directions, Atmos. Chem. Phys., 10, 10829–10848, https://doi.org/10.5194/acp-10-10829-2010, 2010.

Kirkby, J., Curtius, J., Almeida, J., Dunne, E., Duplissy, J., Ehrhart, S., Franchin, A., Gagné, S., Ickes, L., Kürten, A., Kupc, A., Metzger, A., Riccobono, F., Rondo, L., Schobesberger, S., Tsagkogeorgas, G., Wimmer, D., Amorim, A., Bianchi, F., Breitenlechner, M., David, A., Dommen, J., Downard, A., Ehn, M., Flagan, R. C., Haider, S., Hansel, A., Hauser, D., Jud, W., Junninen, H., Kreissl, F., Kvashin, A., Laaksonen, A., Lehtipalo, K., Lima, J., Lovejoy, E. R., Makhmutov, V., Mathot, S., Mikkilä, J., Minginette, P., Mogo, S., Nieminen, T., Onnela, A., Pereira, P., Petäjä, T., Schnitzhofer, R., Seinfeld, J. H., Sipilä, M., Stozhkov, Y., Stratmann, F., Tomé, A., Vanhanen, J., Viisanen, Y., Vrtala, A., Wagner, P. E., Walther, H., Weingartner, E., Wex, H., Winkler, P. M., Carslaw, K. S., Worsnop, D. R., Baltensperger, U., and Kulmala, M.: Role of sulphuric acid, ammonia and galactic cosmic rays in atmospheric aerosol nucleation, Nature, 476, 429–433, https://doi.org/10.1038/nature10343, 2011.

Kopanakis, I., Chatoutsidou, S. E., Torseth, K., Glytsos, T., and Lazaridis, M.: Particle number size distribution in the eastern Mediterranean: Formation and growth rates of ultrafine airborne atmospheric particles, Atmos. Environ., 77, 790–802, https://doi.org/10.1016/j.atmosenv.2013.05.066, 2013.

Kostenidou, E., Lee, B. H., Engelhart, G. J., Pierce, J. R., and Pandis, S. N.: Mass spectra deconvolution of low, medium, and high volatility biogenic secondary organic aerosol, Environ. Sci. Technol., 43, 4884–4889, https://doi.org/10.1021/es803676g, 2009.

Kreidenweis, S. M., Yin, F., Wang, S.-C., Grosjean, D., Flagan, R. C., and Seinfeld, J. H.: Aerosol formation during photooxidation of organosulfur species, Atmos. Environ. Part A. Gen. Top., 25, 2491–2500, https://doi.org/10.1016/0960-1686(91)90165-4, 1991.

Kulmala, M., Riipinen, I., Sipila, M., Manninen, H. E., Petaja, T., Junninen, H., Maso, M. D., Mordas, G., Mirme, A., Vana, M., Hirsikko, A., Laakso, L., Harrison, R. M., Hanson, I., Leung, C., Lehtinen, K. E. J., and Kerminen, V.-M.: Toward Direct Measurement of Atmospheric Nucleation, Science, 318, 89–92, https://doi.org/10.1126/science.1144124, 2007.

Laaksonen, A., Pirjola, L., Kulmala, M., Wohlfrom, K. H., Arnold, F., and Raes, F.: Upper tropospheric so2 conversion into sulfuric acid aerosols and cloud condensation nuclei, J.

Geophys. Res. Atmos., 105, 1459–1469, https://doi.org/10.1029/1999JD900933, 2000.

Manninen, H. E., Nieminen, T., Asmi, E., Gagné, S., Häkkinen, S., Lehtipalo, K., Aalto, P., Vana, M., Mirme, A., Mirme, S., Hõrrak, U., Plass-Dülmer, C., Stange, G., Kiss, G., Hoffer, A., Töro, N., Moerman, M., Henzing, B., De Leeuw, G., Brinkenberg, M., Kouvarakis, G. N., Bougiatioti, A., Mihalopoulos, N., O'Dowd, C., Ceburnis, D., Arneth, A., Svenningsson, B., Swietlicki, E., Tarozzi, L., Decesari, S., Facchini, M. C., Birmili, W., Sonntag, A., Wiedensohler, A., Boulon, J., Sellegri, K., Laj, P., Gysel, M., Bukowiecki, N., Weingartner, E., Wehrle, G., Laaksonen, A., Hamed, A., Joutsensaari, J., Petäjä, T., Kerminen, V. M., and Kulmala, M.: EUCAARI ion spectrometer measurements at 12 European sites-analysis of new particle formation events, Atmos. Chem. Phys., 10, 7907–7927, https://doi.org/10.5194/acp-10-7907-2010, 2010.

McFiggans, G., Bale, C. S. E., Ball, S. M., Beames, J. M., Bloss, W. J., Carpenter, L. J., Dorsey, J., Dunk, R., Flynn, M. J., Furneaux, K. L., Gallagher, M. W., Heard, D. E., Hollingsworth, A. M., Hornsby, K., Ingham, T., Jones, C. E., Jones, R. L., Kramer, L. J., Langridge, J. M., Leblanc, C., LeCrane, J. P., Lee, J. D., Leigh, R. J., Longley, I., Mahajan, A. S., Monks, P. S., Oetjen, H., Orr-Ewing, A. J., Plane, J. M. C., Potin, P., Shillings, A. J. L., Thomas, F., Von Glasow, R., Wada, R., Whalley, L. K., and Whitehead, J. D.: Iodine-mediated coastal particle formation: An overview of the Reactive Halogens in the Marine boundary layer (RHaMBLe) Roscoff coastal study, Atmos. Chem. Phys., 10, 2975–2999, https://doi.org/10.5194/acp-10-2975-2010, 2010.

Merikanto, J., Spracklen, D. V., Mann, G. W., Pickering, S. J., and Carslaw, K. S.: Impact of nucleation on global CCN, Atmos. Chem. Phys., 9, 8601–8616, https://doi.org/10.5194/acp-9-8601-2009, 2009.

Mohr, C., Thornton, J. A., Heitto, A., Lopez-Hilfiker, F. D., Lutz, A., Riipinen, I., Hong, J., Donahue, N. M., Hallquist, M., Petäjä, T., Kulmala, M., and Yli-Juuti, T.: Molecular identification of organic vapors driving atmospheric nanoparticle growth, Nat. Commun., 10, 1–7, https://doi.org/10.1038/s41467-019-12473-2, 2019.

Patoulias, D., Fountoukis, C., Riipinen, I., Asmi, A., Kulmala, M., and Pandis, S. N.: Simulation of the size-composition distribution of atmospheric nanoparticles over Europe, Atmos. Chem. Phys., 18, 13639–13654, https://doi.org/10.5194/acp-18-13639-2018, 2018.

Pikridas, M., Riipinen, I., Hildebrandt, L., Kostenidou, E., Manninen, H., Mihalopoulos, N.,

Kalivitis, N., Burkhart, J. F., Stohl, A., Kulmala, M., and Pandis, S. N.: New particle formation at a remote site in the eastern Mediterranean, J. Geophys. Res. Atmos., 117, D12205, https://doi.org/10.1029/2012JD017570, 2012.

Pope, C. A., Burnett, R. T., Thun, M. J., Calle, E. E., Krewski, D., Ito, K., and Thurston, G. D.: Lung cancer, cardiopulmonary mortality, and long-term exposure to fine particulate air pollution, J. Am. Med. Assoc., 287, 1132–1141, https://doi.org/10.1001/jama.287.9.1132, 2002.

Saha, P. K., Zimmerman, N., Malings, C., Hauryliuk, A., Li, Z., Snell, L., Subramanian, R., Lipsky, E., Apte, J. S., Robinson, A. L., and Presto, A. A.: Quantifying high-resolution spatial variations and local source impacts of urban ultrafine particle concentrations, Sci. Total Environ., 655, 473–481, https://doi.org/10.1016/j.scitotenv.2018.11.197, 2019.

Siakavaras, D., Samara, C., Petrakakis, M., and Biskos, G.: Nucleation events at a coastal city during the warm period: Kerbside versus urban background measurements, Atmos. Environ., 140, 60–68, https://doi.org/10.1016/j.atmosenv.2016.05.054, 2016.

Sipila, M., Berndt, T., Petaja, T., Brus, D., Vanhanen, J., Stratmann, F., Patokoski, J., Mauldin, R. L., Hyvärinen, A. P., Lihavainen, H., and Kulmala, M.: The role of sulfuric acid in atmospheric nucleation, Science, 327, 1243–1246, https://doi.org/10.1126/science.1180315, 2010.

Sipilä, M., Sarnela, N., Jokinen, T., Henschel, H., Junninen, H., Kontkanen, J., Richters, S., Kangasluoma, J., Franchin, A., Peräkylä, O., Rissanen, M. P., Ehn, M., Vehkamäki, H., Kurten, T., Berndt, T., Petäjä, T., Worsnop, D., Ceburnis, D., Kerminen, V. M., Kulmala, M., and O'Dowd, C.: Molecular-scale evidence of aerosol particle formation via sequential addition of HIO3, Nature, 537, 532–534, https://doi.org/10.1038/nature19314, 2016.

Vratolis, S., Gini, M. I., Bezantakos, S., Stavroulas, I., Kalivitis, N., Kostenidou, E., Louvaris, E., Siakavaras, D., Biskos, G., Mihalopoulos, N., Pandis, S. N., Pilinis, C., Papayannis, A., and Eleftheriadis, K.: Particle number size distribution statistics at City-Centre Urban Background, urban background, and remote stations in Greece during summer, Atmos. Environ., 213, 711–726, https://doi.org/10.1016/j.atmosenv.2019.05.064, 2019.

Wang, M., Kong, W., Marten, R., He, X. C., Chen, D., Pfeifer, J., Heitto, A., Kontkanen, J., Dada, L., Kürten, A., Yli-Juuti, T., Manninen, H. E., Amanatidis, S., Amorim, A., Baalbaki, R., Baccarini, A., Bell, D. M., Bertozzi, B., Bräkling, S., Brilke, S., Murillo, L. C., Chiu, R.,

Chu, B., De Menezes, L. P., Duplissy, J., Finkenzeller, H., Carracedo, L. G., Granzin, M., Guida, R., Hansel, A., Hofbauer, V., Krechmer, J., Lehtipalo, K., Lamkaddam, H., Lampimäki, M., Lee, C. P., Makhmutov, V., Marie, G., Mathot, S., Mauldin, R. L., Mentler, B., Müller, T., Onnela, A., Partoll, E., Petäjä, T., Philippov, M., Pospisilova, V., Ranjithkumar, A., Rissanen, M., Rörup, B., Scholz, W., Shen, J., Simon, M., Sipilä, M., Steiner, G., Stolzenburg, D., Tham, Y. J., Tomé, A., Wagner, A. C., Wang, D. S., Wang, Y., Weber, S. K., Winkler, P. M., Wlasits, P. J., Wu, Y., Xiao, M., Ye, Q., Zauner-Wieczorek, M., Zhou, X., Volkamer, R., Riipinen, I., Dommen, J., Curtius, J., Baltensperger, U., Kulmala, M., Worsnop, D. R., Kirkby, J., Seinfeld, J. H., El-Haddad, I., Flagan, R. C., and Donahue, N. M.: Rapid growth of new atmospheric particles by nitric acid and ammonia condensation, Nature, 581, 184–189, https://doi.org/10.1038/s41586-020-2270-4, 2020.

Wang, N., Jorga, S. D., Pierce, J. R., Donahue, N. M., and Pandis, S. N.: Particle wall-loss correction methods in smog chamber experiments, Atmos. Meas. Tech., 11, 6577–6588, https://doi.org/10.5194/amt-11-6577-2018, 2018.

Wang, Z., Wu, Z., Yue, D., Shang, D., Guo, S., Sun, J., Ding, A., Wang, L., Jiang, J., Guo, H., Gao, J., Cheung, H. C., Morawska, L., Keywood, M., and Hu, M.: New particle formation in China: Current knowledge and further directions, Sci. Total Environ., 577, 258–266, https://doi.org/10.1016/j.scitotenv.2016.10.177, 2017.

Weber, R. J., McMurry, P. H., Mauldin, L., Tanner, D. J., Eisele, F. L., Brechtel, F. J., Kreidenweis, S. M., Kok, G. L., Schillawski, R. D., and Baumgardner, B.: A study of new particle formation and growth involving biogenic and trace gas species measured during ACE 1, J. Geophys. Res. Atmos., 103, 16385–16396, https://doi.org/10.1029/97JD02465, 1998.

Weber, R. J., Marti, J. J., Mcmurry, P. H., Eisele, F. L., Tanner, D. J., And Jefferson, A.: Measured atmospheric new particle formation rates: implications for nucleation mechanisms, Chem. Eng. Commun., 151, 53–64, https://doi.org/10.1080/00986449608936541, 1996.

Yao, L., Garmash, O., Bianchi, F., Zheng, J., Yan, C., Kontkanen, J., Junninen, H., Mazon, S. B., Ehn, M., Paasonen, P., Sipilä, M., Wang, M., Wang, X., Xiao, S., Chen, H., Lu, Y., Zhang, B., Wang, D., Fu, Q., Geng, F., Li, L., Wang, H., Qiao, L., Yang, X., Chen, J., Kerminen, V.-M., Petäjä, T., Worsnop, D. R., Kulmala, M., and Wang, L.: Atmospheric new particle

699        formation from sulfuric acid and amines in a Chinese megacity, Science, 361, 278–281,

700        https://doi.org/10.1126/science.aao4839, 2018.

701 Yli-Juuti, T., Nieminen, T., Hirsikko, A., Aalto, P. P., Asmi, E., Hõrrak, U., Manninen, H. E.,

702        Patokoski, J., Dal Maso, M., Petäjä, T., Rinne, J., Kulmala, M., and Riipinen, I.: Growth

703        rates of nucleation mode particles in Hyytiälä during 2003-2009: Variation with particle

704        size, season, data analysis method and ambient conditions, Atmos. Chem. Phys., 11,

705        12865–12886, https://doi.org/10.5194/acp-11-12865-2011, 2011.

706 Zhao, Y., Hennigan, C. J., May, A. A., Tkacik, D. S., De Gouw, J. A., Gilman, J. B., Kuster, W.

707        C., Borbon, A., and Robinson, A. L.: Intermediate-volatility organic compounds: A large

708        source of secondary organic aerosol, Environ. Sci. Technol., 48, 13743–13750,

709        https://doi.org/10.1021/es5035188, 2014.

710 Zhu, Y., Li, K., Shen, Y., Gao, Y., Liu, X., Yu, Y., Gao, H., and Yao, X.: New particle formation

711        in the marine atmosphere during seven cruise campaigns, Atmos. Chem. Phys., 19, 89–

712        113, https://doi.org/10.5194/acp-19-89-2019, 2019.

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714    **Table 1**: Initial conditions in the conducted experiments.

| Exp. | Category | PM$_1$ ($\mu$g m$^3$) | O$_3$ (ppb) | NH$_3$[c] (ppb) | SO$_2$ (ppb) | NO$_x$ (ppb) | RH (%) |
|---|---|---|---|---|---|---|---|
| Exp. A1 | | 1.6 | 58 | 150 | 0.7 | 3.6 | 40 |
| Exp. A2[a] | | 4.2 | 47 | 25 | 1.6 | 12.1 | 50 |
| Exp. A3[b] | | 0.9 | 49 | 200 | 0.6 | 5.6 | 42 |
| Exp. A4 | Class A | 0.6 | 52 | 120 | 0.5 | 4.4 | 45 |
| Exp. A5 | | 3.6 | 54 | 120 | 0.6 | 3.7 | 40 |
| Exp. A6 | | 3.7 | 52 | 150 | 1 | 7.2 | 40 |
| Exp. B1 | | 2.2 | 45 | 150 | 0.8 | 6.7 | 50 |
| Exp. B2 | Class B | 1.6 | 49 | 25 | 1.1 | 8.4 | 56 |
| Exp. B3 | | 1 | 41 | 200 | 0.6 | 8.3 | 58 |
| Exp. B4 | | 2.2 | 56 | 120 | 0.8 | 4 | 40 |
| Exp. C1 | | 3 | 12 | 150 | 1.3 | 27 | 48 |
| Exp. C2[b] | Class C | 2.5 | 56 | 75 | 0.6 | 7.5 | 38 |
| Exp. C3[a] | | 2.2 | 50 | 20 | 0.6 | 11 | 52 |

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716    [a] Experiments illuminated by natural sunlight

717    [b] Experiment conducted at night

718    [c] Estimated concentration in the perturbation chamber.

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725 **Table 2:** Nucleation time, nucleation rate and condensation rate in the experiments were NPF was

726 observed in the perturbed chamber.

| Experiment | Initial GR (nm h$^{-1}$) | Condensation sink × 10$^3$ (s$^{-1}$) | H$_2$SO$_4$ × 10$^{-7}$ (molecule cm$^{-3}$) |
|---|---|---|---|
| Exp. A1 | 4 | 2.6 | 3 |
| Exp. A2[a] | 8 | 3.5 | 5 |
| Exp. A3[b] | 5.5 | 0.8 | 8 |
| Exp. A4 | 6.5 | 1.1 | 5 |
| Exp. A5 | 3.5 | 3.1 | 2 |
| Exp. A6 | 3.7 | 2.9 | 4 |
| Exp. B1 | 5.5 | 3.1 | 3 |
| Exp. B2 | 11.3 | 2.1 | 6 |
| Exp. B3 | 7 | 2.4 | 3 |
| Exp. B4 | 3.5 | 2 | 4 |
| Exp. C1 | 0 | 2.8 | 5 |
| Exp. C2[b] | 0 | 2.1 | 3 |
| Exp. C3[a] | 0 | 2.5 | 2 |

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728 [a] Experiments illuminated by natural sunlight

729 [b] Experiment conducted at night

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**Table 3:** Nucleation time, nucleation rate and condensation rate in the experiments that NPF was observed in the perturbed chamber.

| Experiment | Nucleation time (h) | Nucleation rate ($cm^{-3} h^{-1}$) | Condensation rate ($ppt\ h^{-1}$) |
|---|---|---|---|
| Exp. A1 | 3 | $9500 \pm 600$ | $3.2 \pm 0.3$ |
| Exp. A2[a] | 2 | $10000 \pm 1000$ | $4.5 \pm 0.4$ |
| Exp. A3[b] | 2 | $500 \pm 100$ | $3.8 \pm 0.4$ |
| Exp. A4 | 2 | $600 \pm 200$ | $4.5 \pm 0.3$ |
| Exp. A5 | 2 | $6500 \pm 1000$ | $2 \pm 0.2$ |
| Exp. A6 | 3 | $7000 \pm 500$ | $3 \pm 0.4$ |
| Exp. B1 | 2.5 | $15000 \pm 1500$ | $4.5 \pm 0.5$ |
| Exp. B2 | 1.9 | $25000 \pm 2000$ | $10 \pm 1$ |
| Exp. B3 | 2 | $5000 \pm 700$ | $4.5 \pm 0.5$ |
| Exp. B4 | 2.5 | $14000 \pm 1000$ | $3 \pm 0.2$ |

[a] Experiments illuminated by natural sunlight

[b] Experiment conducted at night

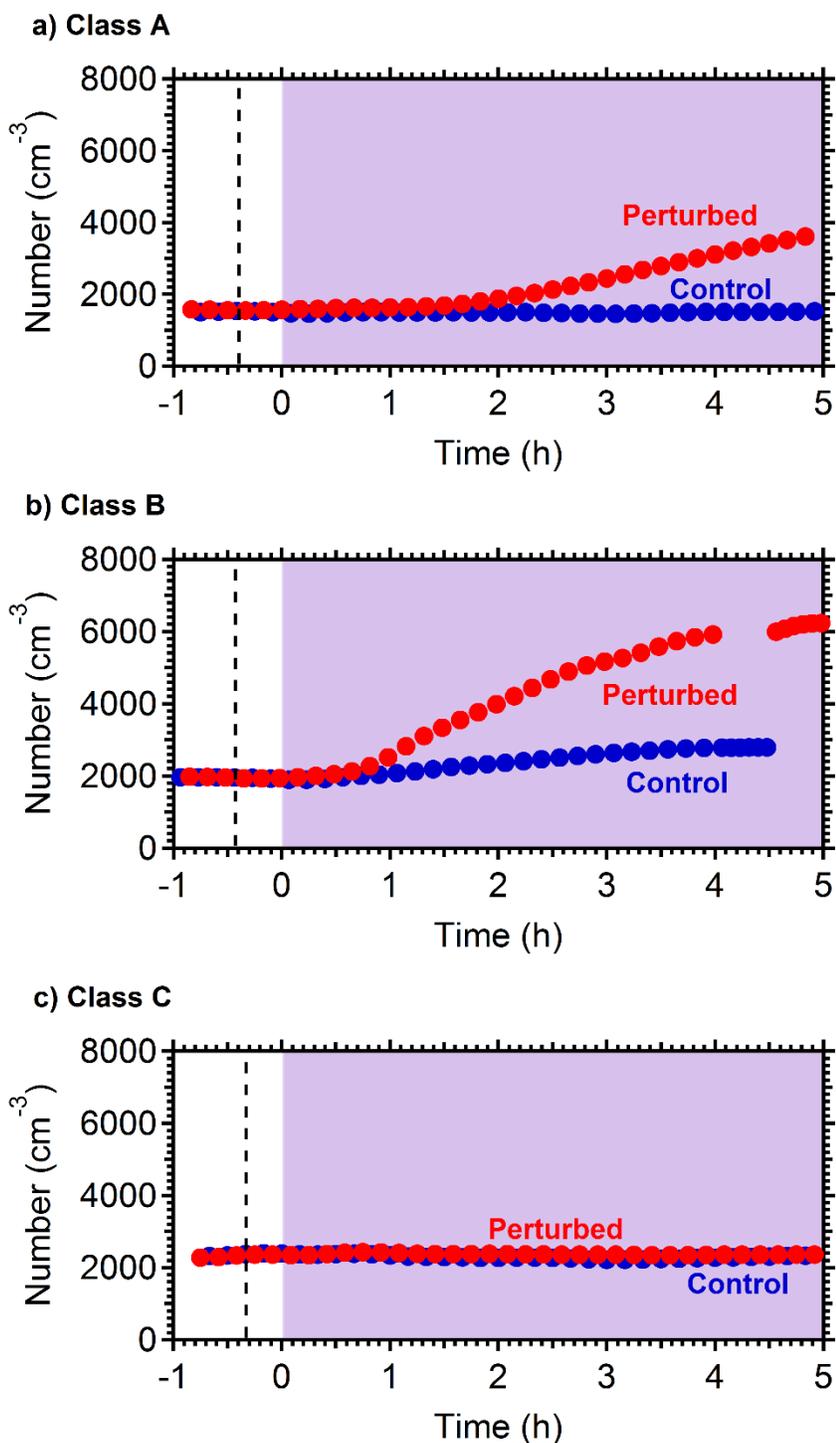
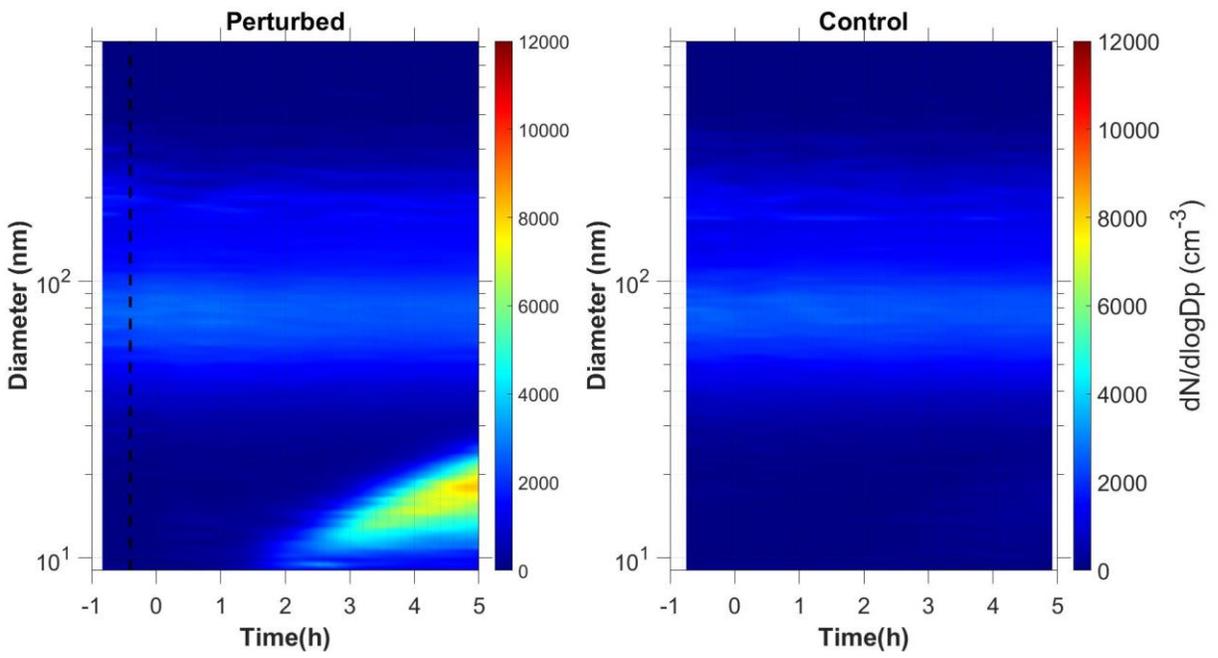



**Figure 1:** Wall loss corrected number concentration $N_9$ in the three different types of experiments, a) NPF and growth only in the perturbed chamber (Exp. A1), b) NPF and growth in both chambers (Exp. B1) and c) no NPF observed (Exp. C1). The dashed line marks the time that ammonia was injected in the perturbed chamber. At t=0 both chambers were illuminated with UV light.

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776 **Figure 2:** Wall loss corrected measured number distributions in the two chambers for Exp. A1.

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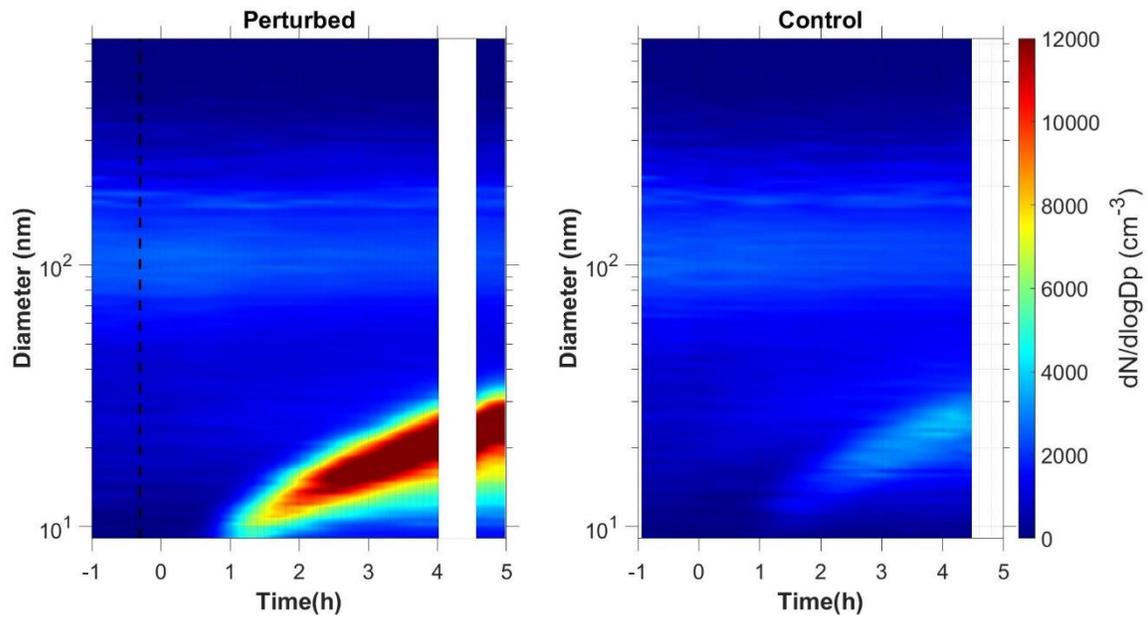
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794 **Figure 3:** Wall loss corrected measured number distributions in the two chambers for Exp. B1.

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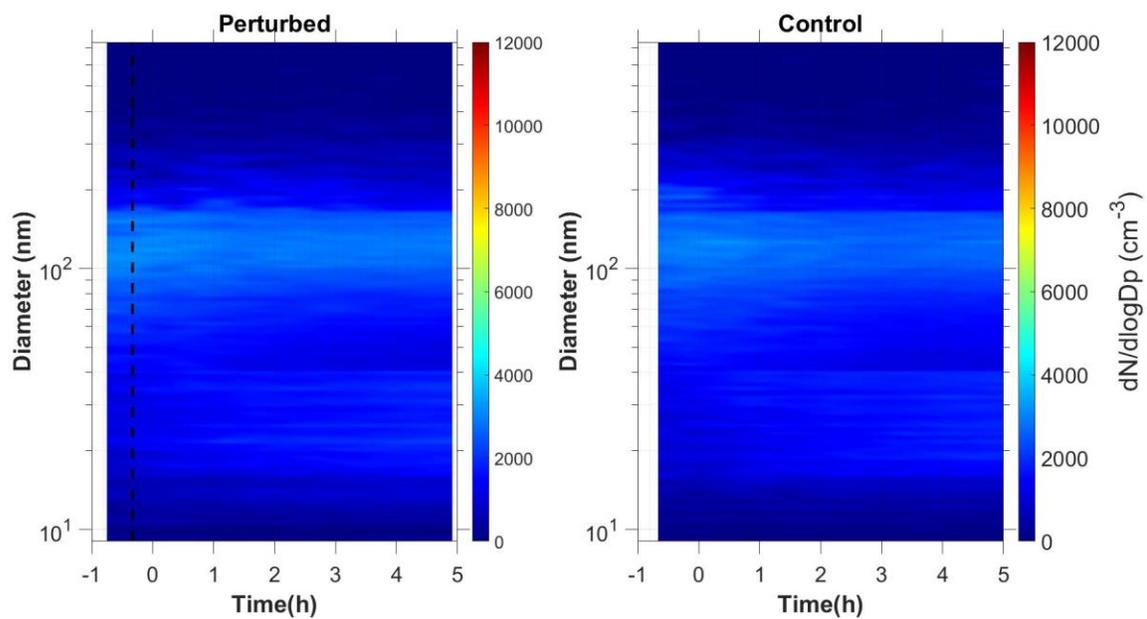
809 **Figure 4:** Wall loss corrected measured number distributions in the two chambers for Exp. C1.

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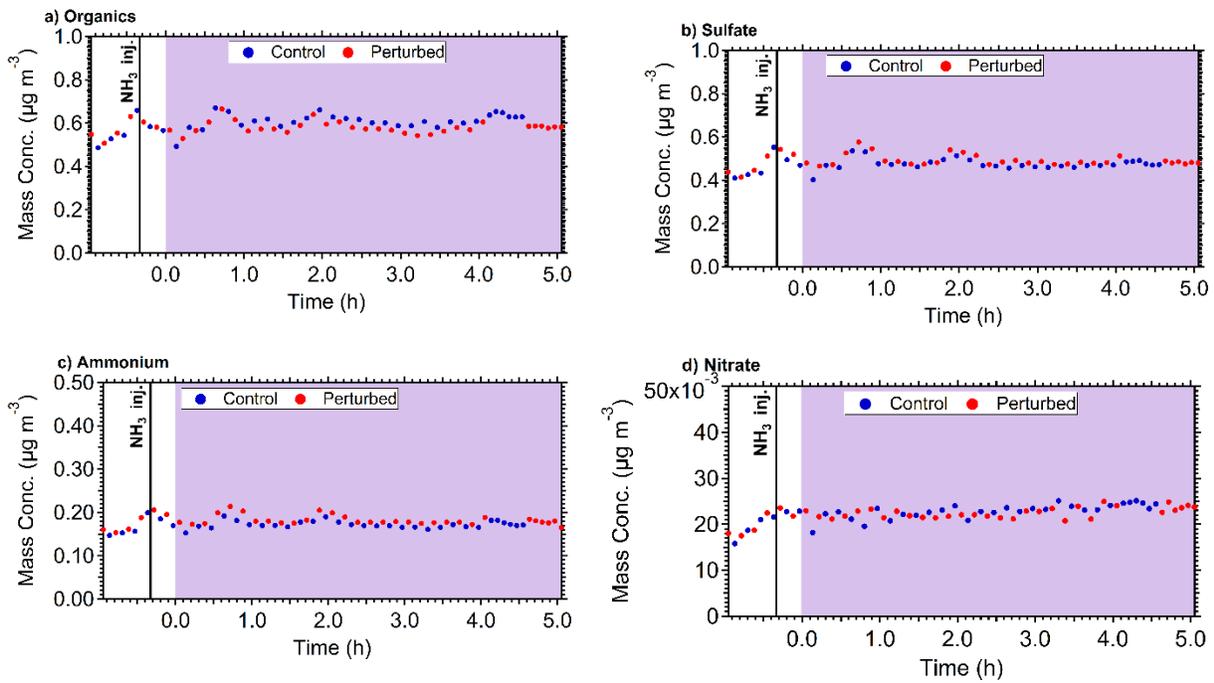
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828  **Figure 5:** The wall loss corrected mass concentration of a) organics, b) sulfate, c) ammonium and

829  d) nitrate in the control (blue dots) and perturbed (red dots) chamber during Exp. B1. The purple

830  shades region represents the time that the chambers were under UV illumination.

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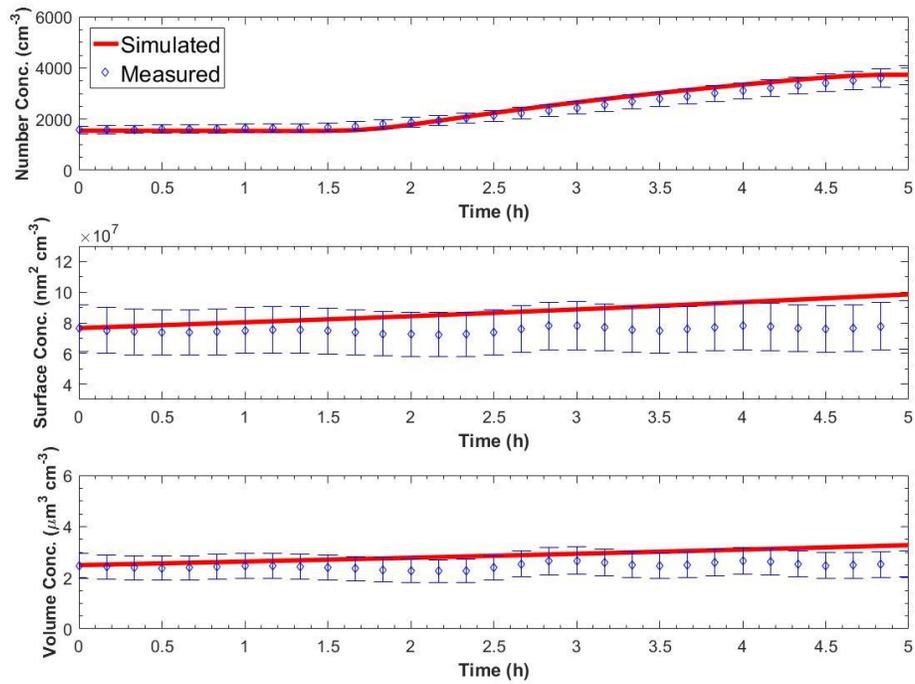
**Figure 6:** Measured and simulated number, surface, and volume concentration in the perturbed chamber after turning UV lights on for Exp. A1. The error bars in the measured values are calculated from the uncertainty in the particle wall loss correction and represent two standard deviations.