1	Climate-driven deterioration of future ozone pollution in
2	Asia predicted by machine learning with multisource data
3	
4	Huimin Li <sup>1</sup> , Yang Yang <sup>1*</sup> , Jianbing Jin <sup>1</sup> , Hailong Wang <sup>2</sup> , Ke Li <sup>1</sup> , Pinya Wang <sup>1</sup> ,
5	Hong Liao <sup>1</sup>
6	
7	
8	
9	<sup>1</sup> Jiangsu Key Laboratory of Atmospheric Environment Monitoring and
10	Pollution Control, Jiangsu Collaborative Innovation Center of Atmospheric
11	Environment and Equipment Technology, School of Environmental Science
12	and Engineering, Nanjing University of Information Science and Technology,
13	Nanjing, Jiangsu, China
14	<sup>2</sup> Atmospheric Sciences and Global Change Division, Pacific Northwest
15	National Laboratory, Richland, Washington, USA
16	
17	
18	
19	*Correspondence to yang.yang@nuist.edu.cn

20 Abstract

Ozone (O<sub>3</sub>) is a secondary pollutant in the atmosphere formed by 21 22 photochemical reactions that endangers human health and ecosystems. O<sub>3</sub> has aggravated in Asia in recent decades and will vary in the future. In this study, to 23 24 quantify the impacts of future climate change on  $O_3$  pollution, near-surface  $O_3$ concentrations over Asia in 2020–2100 are projected using a machine learning 25 (ML) method along with multisource data. The ML model is trained with 26 27 combined O<sub>3</sub> data from a global atmospheric chemical transport model and real-28 time observations. The ML model is then used to estimate future O<sub>3</sub> with meteorological fields from multi-model simulations under various climate 29 scenarios. The near-surface O<sub>3</sub> concentrations are projected to increase by 5-30 31 20% over South China, Southeast Asia, and South India and less than 10% over North China and Gangetic Plains under the high forcing scenarios in the 32 last decade of 21<sup>st</sup> century, compared to the first decade of 2020–2100. The O<sub>3</sub> 33 34 increases are primarily owing to the favorable meteorological conditions for O<sub>3</sub> 35 photochemical formation in most Asian regions. We also find that the summertime  $O_3$  pollution over eastern China will expand from North China to 36 South China and extend into the cold season in a warmer future. Our results 37 demonstrate the important role of climate change penalty on Asian O<sub>3</sub> in the 38 future, which provides implications for environmental and climate strategies of 39 40 adaptation and mitigation.

41 **1. Introduction** 

Tropospheric ozone (O<sub>3</sub>) is a primary secondary air pollutant, formed by 42 43 photochemical oxidation of nonmethane volatile organic compounds (NMVOCs) and carbon monoxide (CO) in the presence of nitrogen oxides (NO<sub>x</sub> = NO + 44 NO<sub>2</sub>) and sunlight. It has adverse effects on human health (Malley et al., 2017; 45 Cakmak et al., 2018), vegetation growth (Yue et al. 2017; Mills et al., 2018) and 46 climate change (Checa-Garcia et al., 2018; Gaudel et al., 2018). A better 47 understanding of the causes of changes in O<sub>3</sub> concentrations is useful for 48 49 developing effective environment and climate strategies.

Since mid-1990s, Asian regions, including South Asia, East Asia and 50 Southeast Asia, have experienced the fastest O<sub>3</sub> increase rate of 2-8 51 52 ppb/decade at remote surface sites and in the lower free troposphere across the world (IPCC, 2021). A number of air quality monitoring stations-administered 53 by China National Environmental Monitoring Center (CNEMC) have been 54 55 established in China since 2013 to measure real-time near-surface particulate 56 matter, O<sub>3</sub>, and other air pollutants. The measurements showed an increasing trend of urban warm-season daily maximum 8-hour average (MDA8) O<sub>3</sub> 57 concentrations of 2.4 ppb (5%) yr<sup>-1</sup> that is faster than any other regions 58 59 worldwide during 2013–2019 (Lu et al., 2020). However, many regions in Asia lack O<sub>3</sub> observations with sufficient spatial and temporal coverage. Also, most 60 61 of the present regional observations are collected only near population clusters, which are not representative of the entire region (Zhou et al., 2022). 62

To supplement the limited near-surface O<sub>3</sub> measurements, many studies 63 utilized global and regional models with comprehensive physical and chemical 64 processes to simulate O<sub>3</sub> concentrations (Zhu et al., 2017; Gao et al., 2020; 65 Yang et al., 2022). Moreover, statistical models have also been used to estimate 66 67 O<sub>3</sub> concentrations (Chen et al., 2020; Zhang et al., 2020). In recent years, machine learning (ML) approaches, such as random forest (Xue et al., 2020; 68 Wei et al., 2022), neural network (Di et al., 2017), support vector machine (Su 69 70 et al., 2020), extreme gradient boosting (LiLiu et al., 2020), and ensemble 71 learning (Liu et al., 2022), were widely applied to estimate O<sub>3</sub> levels based on potential influential factors (e.g., precursor emissions, meteorological 72 73 conditions, land use, surface elevation, gross domestic product, population 74 density, and geographical variables). The abovementioned previous studies utilizing the ML methods showed high computational efficiency and accuracy, 75 with an overall R<sup>2</sup> between the observed and predicted O<sub>3</sub> concentrations in the 76 77 range of 0.7–0.9.

Meteorological factors and synoptic conditions play important roles in 78 affecting O<sub>3</sub> pollution (Fu and Tai, 2015; Gong and Liao, 2019; Yin et al., 2019; 79 Liu et al., 2020; Dang et al., 2021). Gong et aland Liao. (2019) illustrated that 80 81 hot, dry, and stagnant weather conditions are favorable for the formation and persistence of severe O<sub>3</sub> pollution over northern China. High air temperature 82 83 along with intense incoming shortwave radiation accelerates both photochemical reaction rates and natural precursor emissions for O<sub>3</sub> production 84

(Jacob and Winner, 2009). Under high relative humidity conditions, O<sub>3</sub> 85 concentrations decrease due to many complex physical and chemical 86 87 mechanisms (Jeong and Park, 2013; Kavassalis and Murphy., 2017; Lu et al., 2019; Li M. et al., 20212021a). Cloud and precipitation impact O3 levels through 88 89 reducing the downwelling solar radiation and washout of pollutants (Toh et al., 2013). Anomalous sea level pressure patterns can affect the long-range 90 transport of O<sub>3</sub> by influencing atmospheric circulation (Santurtún et al., 2015). 91 92 By changing the air stagnant condition and transport of pollutants, wind fields 93 can also affect O<sub>3</sub> concentrations in local and downwind areas of emission sources (Doherty et al., 2013). 94

Future climate change corresponding to the different climate scenarios can 95 96 impact O<sub>3</sub> through altering meteorological conditions (Wang et al., 2013, Fu and Tian et al., 2019). Using regional climate fields downscaled from general 97 circulation models to investigate potential O<sub>3</sub> variations in the U.S. due to 98 99 changing climate, Fann et al. (2015) projected the MDA8 O<sub>3</sub> to increase by 1-5 ppb as daily maximum average temperature increases by 1-4°C in 2030 100 101 relative to 2000. Colette et al. (2015) estimated that the climate penalty for 102 future summertime near-surface O<sub>3</sub> reaches 0.99–1.5 ppb by the end of the 21<sup>st</sup> century (2071-2100) in Europe compared to present-day levels using an 103 ensemble of eight global coupled climate-chemistry models under the RCP 104 (Representative Concentration Pathway) 8.5 scenario. Through fixing sea 105 surface temperature at present-day and future conditions in five atmospheric-106

only models as part of the AerChemMIP (Aerosol Chemistry Model 107 Intercomparison Project), Zanis et al. (2022) projected the climate change 108 penalties and benefits on global near-surface O<sub>3</sub> concentrations from 2015 to 109 2100 under the SSP3 scenarios of Shared Socioeconomic Pathways (SSPs) 3-110 111 7.0 scenario. They found O<sub>3</sub> reductions in most regions of the globe, except a 112 robust O<sub>3</sub> climate penalty of 1–2 ppb °C<sup>-1</sup> in South and East Asia under global warming following the SSP3-7.0 pathway. However, SSP3-7.0 is not a good 113 representative scenario for both air quality and climate in Asia. The emissions 114 115 of greenhouse gases (GHGs) and air pollutants over East Asia in SSP3-7.0 are assumed to significantly increase in the near future and keep at high levels in 116 the middle of the 21<sup>st</sup> century among all SSPs (Li et al., 2022), while the 117 118 emissions of air pollutants have been cut by a lot since 2010s in the real world (Wang et al., 2021). The GHGs and pollutant emissions are very likely to 119 continually decline in the future related to the carbon neutrality commitment 120 121 (Cheng et al., 2021).

In this study, we aim to better characterize the impact from future climate change on Asian O<sub>3</sub> pollution using multiple state-of-the-art modeling tools and data. It is important for policy-makers that mitigating global climate change potentially has positive benefits to surface air quality through meteorological factors, not only the reduction in fossil fuel co-emissions. The near-surface O<sub>3</sub> concentrations covering 2020–2100 in Asia are projected using a ML method integrated with multisource data, including assimilated O<sub>3</sub> data that combine

129 ground observations across China and simulations from a global 3-D chemical transport model (GEOS-Chem), meteorological fields under various climate 130 131 scenarios from the latest Coupled Model Intercomparison Project Phase 6 (CMIP6) multi-model simulations, and other auxiliary data (e.g., emissions, land 132 133 use, topography, population density, and spatiotemporal information). ML approach gives the capacity to explore many scenarios more rapidly and for 134 longer time periods than the chemical transport model process-based modeling. 135 Details of the data and methodology used in this study are described in section 136 137 2. Section 3 analyzes the results of climate-driven O<sub>3</sub> variations over different key regions of Asia. Section 4 summarizes the main conclusions and discusses 138 potential uncertainties in this study. 139

#### 140 **2. Materials and Methods**

#### 141 **2.1 GEOS-Chem model description**

142 Figure 1 illustrates the procedures for predicting future near-surface O<sub>3</sub> 143 over Asia under four scenarios. To assimilate O<sub>3</sub> data for the ML model training, the near-surface O3 concentrations over Asia from 2014 to 2019 are firstly 144 simulated using the nested-grid version of the 3-D GEOS-Chem model (version 145 12.9.3), driven by the Modern-Era Retrospective analysis for Research and 146 Applications, Version 2 (MERRA-2) reanalysis meteorological data (Gelaro et 147 al., 2017). The nested GEOS-Chem has 47 vertical layers from the surface up 148 to 0.01 hPa, with a horizontal resolution of 0.5° latitude × 0.625° longitude over 149 the Asia domain (11°S–55°N, 60–150°E). The lateral boundaries of chemical 150

tracer concentrations are provided by global simulations at 2° latitude x 2.5° 151 longitude horizontal resolution. The model includes fully coupled aerosol-O<sub>3</sub>-152 153 NO<sub>x</sub>-hydrocarbon chemical mechanisms (Park et al., 2004; Pye et al., 2009; Mao et al., 2013), with about 300 species participated in over 400 kinetic and 154 photochemical reactions (Bey et al., 2001). The stratospheric O<sub>3</sub> chemistry is 155 simulated through linearized O<sub>3</sub> parameterization scheme (LINOZ; Mclinden et 156 al., 2000), and the planetary boundary layer mixing is calculated by a nonlocal 157 scheme (Lin and McElroy, 2010). GEOS-Chem has shown a good performance 158 159 in reproducing spatiotemporal distributions of O<sub>3</sub> concentrations (e.g., Ni et al., 2018; Li et al., 2019). 160

The historical (2014–2019) anthropogenic emissions of O<sub>3</sub> precursor 161 162 gases, including NO<sub>x</sub>, NMVOCs, and CO, utilized in the nested domain are obtained from the Community Emissions Data System (CEDS; Hoesly et al., 163 2018) version 2021\_04\_21, which fully considered the recent emission 164 reductions in China related to clean air measures. The biomass burning 165 emissions are acquired from the Global Fire Emissions Database version 4 166 (GFED4; van der Werf et al., 2017). Biogenic emissions of NMVOCs from the 167 Model of Emissions of Gases and Aerosols from Nature (MEGAN) version 2.1 168 are employed, with updates from Guenther et al. (2012). Soil NO<sub>x</sub> sources are 169 calculated with an updated version of the Berkeley-Dalhousie Soil NOx 170 Parameterization scheme (Hudman et al., 2012). NOx emissions from lightning 171 are as described by Murray et al. (2012), and the vertical distribution of 172

173 emissions follows Ott et al. (2010).

### 174 **2.2 Ground O3 observations**

175 To improve the performance of the ML model in predicting O<sub>3</sub> concentrations, the nationwide hourly near-surface O<sub>3</sub> concentrations in China 176 177 during 2014–2019 are obtained from the CNEMCChina Ministry of Ecology and Environment (MEE) and used for O<sub>3</sub> data assimilation, which has been widely 178 used to examine pollution over China in previous studies (Li K. et al., 2020, 179 180 2021; Qian et al., 2022). The observational network had about 500 monitoring 181 sites in 2013, and expanded to more than 1500 sites after 2019, covering 360 182 cities in mainland China. Theln this study, the quality controlled hourly O<sub>3</sub> observations in 360 cities are averaged within each 0.5° latitude x 0.625° 183 184 longitude grid of the GEOS-Chem model.

#### 185 **2.3 Data assimilation**

The assimilation system, which is used to combine the  $O_3$  observations across China with results from GEOS-Chem simulations, is based on a threedimensional variational (3DVar) data assimilation<del>.</del> (Kalnay, 2003; Evensen et al., 2022). The goal of the 3DVar is to find the maximum likelihood estimation of a state vector x, which is the  $O_3$  concentrations here in this study, given the available observations y through minimizing the cost function:

192 
$$J(x) = \frac{1}{2} (x - x^b)^{\mathrm{T}} \mathbf{B}^{-1} (x - x^b) + \frac{1}{2} (y - \mathrm{H}(x))^{\mathrm{T}} \mathbf{0}^{-1} (y - \mathrm{H}(x))$$

Here  $x^{b}$  represents the priori simulation<sub>7</sub>. **B** is the empirical background covariance matrix representing formulated as a product of the uncertainty. in

195 the simulated value and a distance-based correlation matrix C, and the individual element is calculated as: 196  $\mathbf{B}_{i,i} = 0.2 * \mathbf{x}_i^b * 0.2 * \mathbf{x}_i^b * \mathbf{C}_{i,i}$ 197 Here we have used 20% choice to characterize uncertainty of the O<sub>3</sub> 198 199 simulation, the correlation matrix is empirically set as:  $\mathbf{C}_{i,i} = e^{-(\frac{d_{i,j}}{200 km})^2/2}$ 200 <u>Here</u>  $d_{i,j}$  represents the spatial distance between the grid cell i and j. 201 202 H denotes the linear observation operator that converts the simulation 203 results into the observation observational space, and. Here all observations are assumed to be independent, and therefore **O** is a diagonal covariance 204 matrix storing the square of the observation uncertainty of the measurements 205 used, which is also set as 20% similarly. 206 Comparisons between observed and assimilated O<sub>3</sub> concentrations over 207 208 2014–2019 are shown in Figure 42. The overall correlation coefficient (R) is

209 0.94, and the normalized mean bias (NMB) is -0.1%, suggesting that the 210 assimilated data have an excellent representation of O<sub>3</sub> observations and 211 minimize the uncertainties of GEOS-Chem simulations in China.

## 212 **2.4 Predicting O<sub>3</sub> using a machine learning method**

In this study, a random forest (RF) model is used to predict  $O_3$ concentrations, similar to our previous studies (Li <u>H.</u> et al., 2021, 2022), with input data of assimilated  $O_3$  concentrations <u>in China</u> that combine observations and results from GEOS-Chem model simulations, <u>GEOS-Chem simulated O\_3</u> 217 <u>concentrations outside of China</u>, MERRA-2 meteorological variables, O<sub>3</sub> 218 precursor emissions, land cover (LC), normalized difference vegetation index 219 (NDVI), topography (TOPO), population density (POP), and the month of the 220 year (MOY) and geographic location of each model grid as spatiotemporal 221 information. Details of the datasets are summarized in Table 1.

For predicting future climate-driven near-surface  $O_3$  concentrations, the ML 222 model is trained with samples over 2014-2018 and the remaining 2019 data 223 are used for model validation. To obtain an optimal ML model, hyperparameters 224 225 are firstly tuned using the 10-fold cross-validation (Rodroguez et al., 2010). Rodriguez et al., 2010). The best hyperparameters (n estimators=200, 226 min samples split=2, max features= "sqrt", bootstrap= "True") of the ML 227 228 model are utilized. Several statistical metrics, including coefficient of determination (R<sup>2</sup>), mean absolute error (MAE), root mean square error (RMSE) 229 and mean relative error (MRE) are used to evaluate the performance of ML 230 231 model. Then the climate-driven near-surface O<sub>3</sub> concentrations during 2020-2100 under four SSPs (SSP1-2.6, SSP2-4.5, SSP3-7.0 and SSP5-8.5) in Asia 232 can be estimated using the trained ML model with varying meteorological 233 factors under the climate change scenarios. Both anthropogenic and natural 234 235 emissions of O<sub>3</sub> precursors are fixed at the present-day levels for the prediction.

#### 236 **2.5 Meteorological fields from CMIP6 multi-model simulations**

237 Monthly meteorological parameters under four different future climate 238 scenarios, including SSP1-2.6, SSP2-4.5, SSP3-7.0 and SSP5-8.5, (a)

239 representation of low, intermediate, medium to high, and high forcing levels, respectively), are fed to a ML model to predict ground-level near-surface O<sub>3</sub> 240 241 concentrations. The Scenario Model Intercomparison Project (ScenarioMIP) as part of CMIP6 provides multi-model projections of climate variables driven by 242 243 future emission and land use changes under different SSPs (O'Neill et al., 2016). In our study, meteorological fields, such as air temperature (at 2m, 850 hPa, 244 and 500 hPa), wind fields (at 850 and 500 hPa), surface relative humidity, 245 incoming shortwave radiation at the surface, total cloud cover, precipitation rate, 246 247 and sea level pressure, are chosen as the key meteorological predictors for ground-levelnear-surface O<sub>3</sub> concentrations, which are obtained from 18 global 248 climate models, i.e., ACCESS-CM2, ACCESS-ESM1-5, CanESM5, CESM2-249 250 WACCM, CMCC-CM2-SR5, EC-Earth3-Veg, EC-Earth3, FGOALS-f3-L, FGOALS-g3, GFDL-ESM4, INM-CM5-0, IPSL-CM6A-LR, MIROC6, MPI-251 ESM1-2-HR, MPI-ESM1-2-LR, MRI-ESM2-0, NorESM2-LM, and NorESM2-252 253 MM. Before being applied to the ML model, future meteorological fields from ScenarioMIP are adjusted to minimize by their potential bias, characterized as 254 the difference in their historical climatological mean (2014-2019) and MERRA-255 2 following Li et al. (2022). It minimizes the inconsistencies in the initial 256 257 conditions in models and reanalysis data following Li et al. (2022).

258 **3. Results** 

## 259 **3.1 Predictive capability of the machine learning model**

260 The ML predicted monthly O<sub>3</sub> concentrations over Asia in 2019 by the ML

261 model are in good agreement with the assimilated O<sub>3</sub> data constructed with observations and GEOS-Chem model results (Fig. 23). The overall R<sup>2</sup> between 262 263 the predicted and assimilated O<sub>3</sub> concentrations is as high as 0.92 and the ML model has a low MRE of 9% in predicting O<sub>3</sub> concentrations over the Asia 264 265 domain. Overall, these statistical indices indicate that the RF model is promising for predicting the spatial distributions and temporal variations of near-surface 266 O<sub>3</sub> concentrations over Asia, which can provide a practical means for studying 267 long-term variations in  $O_3$  under the future climate change. 268

269 Meanwhile, the ML model predictive capability for each grid cell over the entire domain during 2014-2019 is further evaluated and demonstrated in 270 271 Figure 34. Regarding the spatial performance, the estimated O<sub>3</sub> concentrations 272 are highly correlated to the assimilated data in most regions of Asia with small biases, indicating a strong spatial predictive ability of the RF model. More than 273 80% of land areas have a R<sup>2</sup> greater than 0.9. In terms of model uncertainties, 274 275 about 95% of land areas have a RMSE (MAE) less than 3 (2) parts per billion (ppb). Furthermore, approximately 86% of land areas show small modeling bias 276 with MRE below 5%. Note that several grid cells show MRE over 5% but still 277 below 15%, which is related to the data assimilation using monitored and 278 279 simulated O<sub>3</sub> concentrations in China and the coarse resolution for coastal areas and islands over Southeast Asia. 280

Figure 4<u>5</u> shows the importance score of independent variables that contribute to the prediction of trained ML model, which called Gini importance

283 and implies the influence of input features on the target variable in the ML model. The results suggest that among all the input predictors, relative humidity, 284 285 incoming solar radiation at the surface, and topography are the top-three most influential variables for the model construction of near-surface O<sub>3</sub> in Asia, with 286 287 importance scores of 15%, 12% and 10%, respectively. The primary importance of relative humidity has also been reported in previous studies (e.g., Han et al., 288 2020; Qian et al., 2022). Other meteorological parameters, such as cloud cover, 289 sea level pressure, air temperature, precipitation, also have a substantial 290 291 impact on the O<sub>3</sub> estimates, with importance scores ranging from 4% to 8%. In the ML model, the emissions of three primary O<sub>3</sub> precursors, including NMVOCs, 292 293 NO<sub>x</sub>, and CO, have a relatively low importance score of 4–5% individually due 294 to the spatiotemporal diversity of O<sub>3</sub> production regimes. However, it is noted that the O<sub>3</sub> variations in different regions are dominated by different 295 meteorological factors (Weng et al., 2022). The importance score of each 296 297 independent feature quantified in this study can only reflect the overall 298 importance across Asia, which is less representative of any specific regions.

**3.2 Predicted future climate-driven O<sub>3</sub> variations** 

Figure 56 shows the predicted absolute and percentage changes in annual mean <u>near</u>-surface O<sub>3</sub> concentrations in response to climate change between the first and last decades of 2020–2100 based on the future meteorological fields from the 18 CMIP6 models. Fig. 67 shows the time series of the regional averaged values over six sub-regions of Asia during 2020–2100. Under the

global warming trends of all future scenarios, the climate-driven near-surface 305 O<sub>3</sub> concentrations increase constantly from 2020 to 2100 over many key 306 307 regions in Asia, such as North China (NC), South China (SC), Southeast Asia (SEA), South India (SI) and Gangetic Plains (GP), except the Tibetan Plateau 308 309 (TP). The O<sub>3</sub> concentrations over SC, SEA, and SI are projected to increase considerably with the maximum increase up to 5 ppb (20%) in 2095 (2091-310 2100 mean) compared to 2025 (2020-2029 mean) under the SSP5-8.5 311 scenario, revealing a strong O<sub>3</sub>-climate penalty in most Asian regions. The 312 313 climate-driven changes in O<sub>3</sub> concentrations are smaller under the less warming scenarios, especially in SSP1-2.6 that has O<sub>3</sub> changes less than 5% across 314 Asia. These suggest that future climate following low emissions and sustainable 315 316 pathways is more favorable for the mitigation of O<sub>3</sub> pollution in Asia than high forcing scenarios. 317

The strong O<sub>3</sub>-climate penalty over eastern China can be attributed to the 318 319 particularly high O<sub>3</sub> precursor emissions (Fig. S1), relative to western China, 320 which lead to a positive local net O<sub>3</sub> production close to sources in a warming climate (Fig. S2) (Zanis et al., 2022). The absolute and percentage changes in 321 322 regional averaged <u>near-surface</u> O<sub>3</sub> concentrations between 2025 and 2095 323 under the four scenarios are shown in Figure  $\frac{78}{2}$ . The climate-driven changes in  $O_3$  concentrations are gradually stronger from north (2–3%) to south (3–8%) 324 325 of China, which demonstrates that the changes in meteorology exert a greater impact on ground-levelnear-surface O<sub>3</sub> concentrations over SC than NC under 326

future climate change. By the end of the 21<sup>st</sup> century, the relative humidity will decrease (Fig. S3) and downward solar radiation will increase (Fig. S4) over SC compared to those in 2025, which are conducive to the O<sub>3</sub> productions, while NC has the opposite changes. Moreover, cloud cover will decrease more over SC than NC (Fig. S5), contributing to the larger increase in O<sub>3</sub> productions and concentrations over SC than NC in a warming climate.

In South Asia, climate change also enhances  $O_3$  concentrations by <5%333 over GP and SI (Fig. 78), due to the massive precursor emissions (Fig. S1) and 334 335 O<sub>3</sub> productions. Over SI, the decreases in relative humidity (Fig. S3) and cloud amount (Fig. S5), and increases in downward solar radiation at the surface (Fig. 336 337 S4) favor photochemical production of  $O_3$  and induce the large increases in  $O_3$ 338 concentrations in this region. Averaged over SEA, O<sub>3</sub> concentrations driven by higher temperature (Fig. S2), more downward solar radiation (Fig. S4), and 339 lower relative humidity (Fig. S3) and cloud cover (Fig. S5) in 2095 are projected 340 341 to increase O<sub>3</sub> concentrations by 5–7% in SSP3-7.0 and SSP5-8.5 and 0–3% 342 in SSP1-2.6 and SSP2-4.5 scenarios, relative to 2025 (Fig. 78).

The Tibetan Plateau (TP), known as the highest topography in China with more solar radiation at the surface, has strong stratosphere-troposphere exchanges of O<sub>3</sub> compared with other regions leading to high O<sub>3</sub> concentrations over this region (Fig. S6). Climate-driven O<sub>3</sub> concentrations are projected to decline by less than 2% over TP from 2025 to 2095 (Fig. <u>78</u>). It is likely because less solar radiation (Fig. S4) and more frequent occurrence of rainy weather

349 (Fig. S7) in the future would reduce the local chemical production of O<sub>3</sub>.

#### 350 **3.3 The seasonality of future climate-driven O<sub>3</sub> variations**

351 Climate over Asia has obvious seasonal variation related to the Asian 352 monsoon system. Figure 89 shows the spatial distributions of percentage 353 changes in projected climate-driven O<sub>3</sub> concentrations in spring (March-April-May, MAM), summer (June-July-August, JJA), autumn (September-October-354 November, SON), and winter (December-January-February, DJF) between 355 2025 and 2095 under the four scenarios. In general, air quality in many regions 356 357 of Asia will deteriorate in all seasons associated with intensified O<sub>3</sub> pollution under climate change. 358

In eastern China, O<sub>3</sub> pollution occurs most frequently in summer and is 359 360 more severe in NC than SC currently (Li et al., 2019). Under future climate warming, JJA O<sub>3</sub> concentrations will increase by 5–20% in SC under the high 361 forcing scenarios, while the changes in NC are less than 5%. It suggests that 362 363 future climate change will expand the summertime O<sub>3</sub> pollution from NC to SC over eastern China. Another feature is the strong increases in O<sub>3</sub> 364 concentrations by 10-20% throughout eastern China and exceeding 20% over 365 Sichuan Basin in SON, which relate to the significant increases in temperature 366 367 (Fig. S8) and solar radiation (Fig. S9) in this season over central-eastern China under the high forcing scenarios. It further indicates that future climate change 368 369 will extend the O<sub>3</sub> pollution from summer into autumn.

In South Asia, the climate-driven increases in O<sub>3</sub> concentrations vary from

JJA over SI to DJF over GP. Relative to 2025, in summer of 2095, anomalous 371 high pressure (Fig. S10) along with anticyclone (Figs. S11 and S12) dominates 372 373 South Asia, which is not conducive to O<sub>3</sub> diffusion, leading to increases in JJA O<sub>3</sub> concentrations over SI. The intensified O<sub>3</sub> pollution across GP in DJF under 374 375 climate change is related to the strong surface warming (Fig. S8), decreases in relative humidity (Fig. S13), cloud cover (Fig. S14) and rainfall (Fig. S15), as 376 well as increases in solar radiation at the surface (Fig. S9), favoring the 377 photochemical production of O<sub>3</sub>. In north part of Southeast Asia, JJA has the 378 379 largest O<sub>3</sub> rise via the same mechanism as for SI, while O<sub>3</sub> increases by the same magnitude in all seasons in south part of Southeast Asia driven by future 380 climate change. 381

#### 382 4. Conclusions and discussion

383 Ground-level The O<sub>3</sub> pollution has been increasing over Asia in recent decades, which harms human health and vegetations. In the future warmer 384 385 climate, O<sub>3</sub> pollution over Asia can be modulated by changes in meteorological fields. In this study, to examine the variations in O<sub>3</sub> concentrations over Asia 386 387 due to the future climate change, monthly near-surface O<sub>3</sub> concentrations from 2020 to 2100 under four climate scenarios (SSP1-2.6, SSP2-4.5, SSP3-7.0, 388 and SSP5-8.5) are predicted using a ML model with input data from assimilated 389 O3 combining GEOS-Chem simulations and real-time observations, future 390 meteorological parameters from CMIP6 multi-model simulations, emissions of 391 O<sub>3</sub> precursors, land use, topography, population density and spatiotemporal 392

information. Our results suggest that the future  $O_3$  pollution over Asia will be significantly exacerbated in a warming climate, especially under high forcing scenarios.

Trained by the assimilated O<sub>3</sub> concentrations and reanalysis data, the ML model can well predict O<sub>3</sub> over Asia with the coefficient determination of 0.92 between assimilated and predicted O<sub>3</sub> concentrations and relative error of 9%. Then the future Asian O<sub>3</sub> concentrations from 2020 to 2100 driven by climate change are projected in the ML model with varying meteorological fields from 18 CMIP6 models under four future climate scenarios.

The climate penalty on O<sub>3</sub> is robust over most regions of Asia. The annual 402 mean O<sub>3</sub> levels in 2095 are projected to increase by 5–20% relative to 2025 403 404 under the high forcing scenarios over South China, Southeast Asia, and South India and less than 10% over North China and Gangetic Plains, due to more 405 favorable meteorological conditions for O<sub>3</sub> photochemical production, while 406 407 there is a decrease of <5% over the Tibetan Plateau. The climate-driven changes in O<sub>3</sub> concentrations are smaller under the less warming scenarios, 408 suggesting that future climate following low emissions and sustainable 409 pathways would be more effective in the mitigation of O<sub>3</sub> pollution in Asia than 410 411 the high forcing scenarios. Seasonal variation analysis reveals that the summertime O<sub>3</sub> pollution over eastern China will expand from North China to 412 413 South China and extend into the cold season under the future climate change. In addition, South Asian O<sub>3</sub> pollution will increase over South India in summer 414

415 and over Gangetic Plains in winter.

Zanis et al. (2022) analyzed the global climate change benefit and penalty 416 417 on O<sub>3</sub> based on sensitivity simulations from five CMIP6 models under the SSP3-7.0 scenario. They showed positive changes in JJA O<sub>3</sub> concentrations by less 418 419 than 1 ppb from 2010 to 2095 over East Asia and South Asia driven by climate change, but with large uncertainties due to the model diversity. The ML method 420 in this study gives similar positive changes in O<sub>3</sub> as Zanis et al. (2022). Pommier 421 et al. (2018) applied the EMEP chemical transport model driven by the 422 423 downscaled meteorological data from the NorESM1-M to investigate the 424 impacts of regional climate change on near-surface O<sub>3</sub> over India. They showed that near-surface O<sub>3</sub> would increase by up to 4% over Northern India and 425 426 decrease by 3% over Southern India from 2050 to 2100 under the RCP8.5 scenario. We show that the climate-driven O<sub>3</sub> in this study would increase over 427 both Gangetic Plains (0.2%) and South India (3%) under the SSP5-8.5 scenario 428 429 in 2050 relative to 2016 (2014–2019 mean). The discrepancies may rise from that the results of Pommier et al. (2018) were based on NorESM1-M simulated 430 climate alone, while the climate change predicted by 18 CMIP6 models were 431 applied in this study and the ensemble mean O<sub>3</sub> concentrations were shown 432 433 here.

There are a few uncertainties and limitations in the projected near-surface
O<sub>3</sub> concentrations over Asia in terms of input data for the ML model, GEOSChem simulations, and CMIP6 multi-model simulations, and the ML model. First,

only observational data over 2014–2019 across China were used for the O<sub>3</sub> 437 assimilation. Longer-term measurements with broader spatial coverage are 438 more desirable to improve the model performance. Land use data and 439 population density are fixed at present-day conditions when predicting the 440 future O<sub>3</sub> since we focus on the variations in meteorological parameters under 441 climate change, which will vary in the future. In addition, natural O<sub>3</sub> precursor 442 emissions such as biogenic emissions of NMVOCs, and NO<sub>x</sub> from soil and 443 lightning sources are fixed at year-2016 levels in the future estimates, which 444 445 can induce biases in the O<sub>3</sub> projections since climate change can strongly influence natural emissions of O<sub>3</sub> precursors (Liu et al., 2019). Although the 446 447 climate influence of methane is considered concentrations in the future 448 predictions, GEOS-Chem model are prescribed and its role in the O<sub>3</sub> production is not considered in the ML model, the climate influence of methane is included 449 in the ML model. The CMIP6 multi-model simulations. Consequently, the impact 450 451 of future changes in methane on O<sub>3</sub> concentrations via climate change are 452 considered in the future projections.

453 <u>Second, the</u> GEOS-Chem model has been demonstrated to well capture
454 the magnitude of and spatiotemporal variations in O<sub>3</sub>, with an average bias of
455 about 10% over China (Lou et al., 2014) and Southeast Asia (Marvin et al.,
456 2021), and less than 20% over India (David et al., 2019). The future decrease
457 in relative humidity will cause stomatal closure and also increase <u>near</u>-surface
458 O<sub>3</sub>. The O<sub>3</sub>-vegetation interactions are not represented in the default GEOS-

Chem model. A newly coupled global atmospheric chemistry-vegetation model
(Lei et., 2020) could be applied in the future study. Additionally

461 Third, the meteorological parameters characterizing future climate change from the CMIP6 multi-model simulations can also give rise to uncertainties in 462 463 this study (Xu et al., 2021). Moreover, the spatial autocorrelation in random split of training data for cross-validation would lead to the overly optimistic statistics 464 of ML model predictive power (Ploton et al., 2020). Additionally, the overall 465 466 importance scores of the features in this study can only reflect that from the 467 whole study domain. Further investigations are required to identify and quantify the importance score of each local variable contributed to the near-surface O<sub>3</sub> 468 predictions in different specific regions. Also, the good ability of the ML model 469 470 for the present-day condition may not imply a satisfactorily extrapolation under the future warming condition, which can bias our results and deserves further 471 investigation in future studies. 472

473 <u>Last but not least, the near-surface O<sub>3</sub> have increased rapidly in China</u>
474 <u>since 2013 owing to both precursor emission changes and atmospheric</u>
475 <u>warming (Li M. et al., 2021b), which significantly affect human health (Lu et al., 2020) and also requires further studies.</u>

477 Overall, our study provides a framework of combining real-time
478 observations, chemical transport model simulations and multi-climate model
479 predictions with data assimilation and machine learning methods to estimate
480 future climate driven near-surface O<sub>3</sub> concentrations. The emphasis of this work

- 481 is to quantify the impacts of future climate change on  $O_3$  pollution in Asia, which
- 482 is of great significance for the future O<sub>3</sub> pollution mitigation strategies.

#### 483 Author contributions

YY designed the research. HL performed the model simulations, analyzed data
and wrote the initial draft. JJ designed the data assimilation. YY, JJ, HW, and
KL helped edit and review the manuscript. All the authors discussed the results
and contributed to the final manuscript.

#### 488 Code and data availability

The GEOS-Chem model available 489 is at https://zenodo.org/record/3974569#.YTD81NMzagR (last access: 1 August 490 491 2022). MERRA-2 reanalysis data can be downloaded at https://gmao.gsfc.nasa.gov/reanalysis/MERRA-2/ (last access: 1 August 2022). 492 Multi-model projections of climate variables are from Scenario Model 493 494 Intercomparison Project in Phase 6 of the Coupled Model Intercomparison Project https://esgf-node.llnl.gov/search/cmip6/ (last access: 1 August 2022). 495 Land cover is derived from http://maps.elie.ucl.ac.be/CCI/viewer/download.php 496 497 (last access: 1 August 2022). Hourly O<sub>3</sub> concentrations are obtained from the public website of the China National Environmental Monitoring Centre 498 499 httpMinistry of Ecology and Environment https://www.cnemcmee.gov.cn/en/ (last access: 1 August 2022). Normalized difference vegetation index is 500 501 obtained from https://www.ncei.noaa.gov/data/avhrr-land-normalizeddifference-vegetation-index/access/ (last access: 1 August 2022). Topography 502 503 is collected from https://cgiarcsi.community/data/srtm-90m-digital-elevation-

504	database-v4-1/ (last access: 1 August 2022). Population density is acquired
505	from https://landscan.ornl.gov/landscan-datasets (last access: 1 August 2022).

#### 506 Acknowledgments

507 H.W. acknowledges the support by the U.S. Department of Energy (DOE),

508 Office of Science, Office of Biological and Environmental Research (BER), as 509 part of the Earth and Environmental System Modeling program. The Pacific 510 Northwest National Laboratory (PNNL) is operated for DOE by the Battelle 511 Memorial Institute under contract DE-AC05-76RLO1830. The projected O<sub>3</sub> 512 concentrations in this study are available upon request.

#### 513 Competing Interest

514 The contact author has declared that neither they nor their co-authors have any 515 competing interests.

#### 516 *Financial support.*

517 This study was supported by the National Key Research and Development 518 Program of China (grant 2019YFA0606800 and 2020YFA0607803) and the 519 National Natural Science Foundation of China (grant 41975159) and Jiangsu 520 Science Fund for Distinguished Young Scholars (grant BK20211541).

# 521 **Reference**

527

533

538

- Bey, I., Jacob, D. J., Yantosca, R. M., Logan, J. A., Field, B. D., Fiore, A. M., Li, 522 Q., Liu, H., Mickley, L. J., and Schultz, M. G.: Global modeling of 523 tropospheric chemistry with assimilated meteorology: Model description 524 525 and evaluation, J. Geophys. Res. Atmos., 106, 23073-23095, https://doi.org/10.1029/2001JD000807, 2001. 526
- Cakmak, S., Hebbern, C., Pinault, L., Lavigne, E., Vanos, J., Crouse, D. L., and
  Tjepkema, M.: Associations between long-term PM<sub>2.5</sub> and ozone exposure
  and mortality in the Canadian Census Health and Environment Cohort
  (CANCHEC), by spatial synoptic classification zone, Environ. Int., 111, 200–
  211, https://doi.org/10.1016/j.envint.2017.11.030, 2018.
- Checa-Garcia, R., Hegglin, M. I., Kinnison, D., Plummer, D. A., and Shine, K.
  P.: Historical tropospheric and stratospheric ozone radiative forcing using
  the CMIP6 database, Geophys. Res. Lett., 45, 3264–3273,
  https://doi.org/10.1002/2017GL076770, 2018.
- Chen, L., Liang, S., Li, X., Mao, J., Gao, S., Zhang, H., Sun, Y., Vedal, S., Bai,
  Z., Ma, Z., Haiyu., and Azzi, M.: A hybrid approach to estimating long-term
  and short-term exposure levels of ozone at the national scale in China using
  land use regression and Bayesian maximum entropy, Sci. Total Environ.,
  752, 141780, https://doi.org/10.1016/j.scitotenv.2020.141780, 2020.
- Cheng, J., Tong, D., Zhang, Q., Liu, Y., Lei, Y., Yan, G., Yan, L., Yu, S., Cui, R.
  Y., Clarke, L., Geng, G., Zheng, B., Zhang, X., Davis, S. J., and He, K.:
  Pathways of China's PM<sub>2.5</sub> air quality 2015–2060 in the context of carbon
  neutrality, Natl. Sci. Rev., 8, nwab078, https://doi.org/10.1093/nsr/nwab078,
  2021.
- 550
- Colette, A., Andersson, C., Baklanov, A., Bessagnet, B., Brandt, J., Christensen,
  J. H., Doherty, R., Engardt, M., Geels, C., Giannakopoulos, C., Hedegaard,
  G. B., Katragkou, E., Langner, J., Lei, H., Manders, A., Melas, D., Meleux,
  F., Rouïl, L., Sofiev, M., Soares, J., Stevenson, D. S., Tombrou-Tzella, M.,
  Varotsos, K. V., and Young, P.: Is the ozone climate penalty robust in
  Europe? Environ. Res. Lett., 10, 084015, http://dx.doi.org/10.1088/17489326/10/8/084, 2015.
- 558
- Dang, R., Liao, H., and Fu, Y.: Quantifying the anthropogenic and
  meteorological influences on summertime surface ozone in China over
  2012–2017, Sci. Total Environ., 754, 142394,
  https://doi.org/10.1016/j.scitot, 2021.

- David, L. M., Ravishankara, A., Brewer, J. F., Sauvage, B., Thouret, V.,
  Venkataramani, S., and Sinha, V.: Tropospheric ozone over the Indian
  subcontinent from 2000 to 2015: Data set and simulation using GEOSChem chemical transport model, Atmos. Environ., 219, 117039,
  https://doi.org/10.1016/j.atmosenv.2019.117039, 2019.
- Di, Q., Rowland, S., Koutrakis, P., and Schwartz, J.: A hybrid model for spatially
  and temporally resolved ozone exposures in the continental United States,
  J. Air Waste Manage. Assoc., 67, 39–52,
  https://doi.org/10.1080/10962247.2016.1200159, 2017.
- Doherty, R. M., Wild, O., Shindell, D. T., Zeng, G., MacKenzie, I. A., Collins, W.
  J., Fiore, A. M., Stevenson, D. S., Dentener, F. J., Schultz, M. G., Hess, P.,
  Derwent, R. G., and Keating, T. J.: Impacts of climate change on surface
  ozone and intercontinental ozone pollution: A multi-model study, J.
  Geophys. Res., 118, 3744–3763, https://doi.org/10.1002/jgrd.50266, 2013.
- Evensen, G., Vossepoel, F. C., and van Leeuwen, P. J.: Data Assimilation
   Fundamentals: A Unified Formulation of the State and Parameter
   Estimation Problem, Springer Nature, https://doi.org/10.1007/978-3-030 96709-3, 2022.
- Fann, N., Nolte, C. G., Dolwick, P., Spero, T. L., Brown, A. C., Phillips, S., and
  Anenberg, S.: The geographic distribution and economic value of climate
  change-related ozone health impacts in the United States in 2030, J. Air
  Waste Manag. Assoc., 65, 570–580,
  https://doi.org/10.1080/10962247.2014.996270, 2015.
- Fu, T.-M., and Tian, H.: Climate Change Penalty to Ozone Air Quality: Review
  of Current Understandings and Knowledge Gaps, Curr. Pollut. Rep., 5,
  159–171, https://doi.org/10.1007/s40726-019-00115-6, 2019.
- Fu, Y., and Tai, A. P. K.: Impact of climate and land cover changes on
  tropospheric ozone air quality and public health in East Asia between 1980
  and 2010, Atmos. Chem. Phys., 15, 10093–10106,
  https://doi.org/10.5194/acp-15-10093-2015, 2015.
- 600

569

574

580

585

591

- Gao, M., Gao, J., Zhu, B., Kumar, R., Lu, X., Song, S., Zhang, Y., Jia, B., Wang,
  P., Beig, G., Hu, J., Ying, Q., Zhang, H., Sherman, P., and McElroy, M. B.:
  Ozone pollution over China and India: seasonality and sources, Atmos.
  Chem. Phys., 20, 4399–4414, https://doi.org/10.5194/acp-20-4399-2020,
  2020.
- 607 Gaudel, A., Cooper, O. R., Ancellet, G., Barret, B., Boynard, A., Burrows, J. P.,

Clerbaux, C., Coheur, P. F., Cuesta, J., Cuevas, E., Doniki, S., Dufour, G., 608 Ebojie, F., Foret, G., Garcia, O., Granados-Muñoz, M. J., Hannigan, J. W., 609 Hase, F., Hassler, B., Huang, G., Hurtmans, D., Jaffe, D., Jones, N., 610 Kalabokas, P., Kerridge, B., Kulawik, S., Latter, B., Leblanc, T., Le 611 Flochmoën, E., Lin, W., Liu, J., Liu, X., Mahieu, E., McClure-Begley, A., 612 613 Neu, J. L., Osman, M., Palm, M., Petetin, H., Petropavlovskikh, I., Querel, R., Rahpoe, N., Rozanov, A., Schultz, M. G., Schwab, J., Siddans, R., 614 Smale, D., Steinbacher, M., Tanimoto, H., Tarasick, D. W., Thouret, V., 615 Thompson, A. M., Trickl, T., Weatherhead, E., Wespes, C., Worden, H. M., 616 Vigouroux, C., Xu, X., Zeng, G., and Ziemke, J.: Tropospheric Ozone 617 Assessment Report: Present-day distribution and trends of tropospheric 618 ozone relevant to climate and global atmospheric chemistry model 619 620 evaluation, Elem. Sci. Anth., 6, 39, https://doi.org/10.1525/elementa.291, 621 2018.

- Gelaro, R., McCarty, W., Suárez, M. J., Todling, R., Molod, A., Takacs, L., 623 Randles, C. A., Darmenov, A., Bosilovich, M. G., Reichle, R., Wargan, K., 624 Coy, L., Cullather, R., Draper, C., Akella, S., Buchard, V., Conaty, A., da 625 Silva, A. M., Gu, W., Kim, G.-K., Koster, R., Lucchesi, R., Merkova, D., 626 Nielsen, J. E., Partyka, G., Pawson, S., Putman, W., Rienecker, M., 627 Schubert, S. D., Sienkiewicz, M., and Zhao, B.: The Modern-Era 628 Retrospective Analysis for Research and Applications, Version 2 (MERRA-629 2), J. Clim., 30, 5419-5454, https://doi.org/10.1175/JCLI-D-16-0758.1, 630 2017. 631
- Gong, C., and Liao, H.: A typical weather pattern for ozone pollution events in
  North China, Atmos. Chem. Phys., 19, 13725–13740,
  https://doi.org/10.5194/acp-19-13725-2019, 2019.
- Guenther, A. B., Jiang, X., Heald, C. L., Sakulyanontvittaya, T., Duhl, T.,
  Emmons, L. K., and Wang, X.: The Model of Emissions of Gases and
  Aerosols from Nature version 2.1 (MEGAN2.1): an extended and updated
  framework for modeling biogenic emissions, Geosci. Model Dev., 5, 1471–
  1492, https://doi.org/10.5194/gmd-5-1471-2012, 2012.
- 642

622

632

- Han, H., Liu, J., Shu, L., Wang, T., and Yuan, H.: Local and synoptic
  meteorological influences on daily variability in summertime surface ozone
  in eastern China, Atmos. Chem. Phys., 20, 203–222,
  https://doi.org/10.5194/acp-20-203-2020, 2020.
- 647
- Hoesly, R. M., Smith, S. J., Feng, L., Klimont, Z., Janssens-Maenhout, G.,
  Pitkanen, T., Seibert, J. J., Vu, L., Andres, R. J., Bolt, R. M., Bond, T. C.,
  Dawidowski, L., Kholod, N., Kurokawa, J.-I., Li, M., Liu, L., Lu, Z., Moura,
  M. C. P., O'Rourke, P. R., and Zhang, Q.: Historical (1750–2014)

anthropogenic emissions of reactive gases and aerosols from the
Community Emissions Data System (CEDS), Geosci. Model Dev., 11,
369–408, https://doi.org/10.5194/gmd-11-369-2018, 2018.

- Hudman, R. C., Moore, N. E., Mebust, A. K., Martin, R. V., Russell, A. R., Valin,
  L. C., and Cohen, R. C.: Steps towards a mechanistic model of global soil
  nitric oxide emissions: implementation and space based-constraints, Atmos.
  Chem. Phys., 12, 7779–7795, https://doi.org/10.5194/acp-12-7779-2012,
  2012.
- IPCC: Climate change 2021: The physical science basis. Contribution of
   working group I to the sixth assessment report of the intergovernmental
   panel on climate change. Cambridge, UK: Cambridge University Press,
   2021.
- Jacob, D. J., and Winner, D. A.: Effect of climate change on air quality, Atmos.
  Environ., 43, 51–63, https://doi.org/10.1016/j.atmosenv.2008.09.051, 2009.
- Jeong, J. I., and Park, R. J.: Effects of the meteorological variability on regional
  air quality in East Asia, Atmos. Environ., 69, 46–55,
  https://doi.org/10.1016/J.Atmosenv.2012.11.061, 2013.
- Kalnay, E.: Atmospheric Modeling, Data Assimilation and Predictability,
   Cambridge University Press, Cambridge, United Kingdom, 2003.
- Kavassalis, S. C., and Murphy, J. G.: Understanding ozone-meteorology
  correlations: A role for dry deposition, Geophys. Res. Lett., 44, 2922–2931,
  https://doi.org/10.1002/2016gl071791, 2017.
- Lei, Y., Yue, X., Liao, H., Gong, C., and Zhang, L.: Implementation of Yale
  Interactive terrestrial Biosphere model v1.0 into GEOS-Chem v12.0.0: a
  tool for biosphere– chemistry interactions, Geosci. Model Dev., 13, 1137–
  1153, https://doi.org/10.5194/gmd-13-1137-2020, 2020.
- Li, H., Yang, Y., Wang, H., Li, B., Wang, P., Li, J., and Liao, H.: Constructing a
  spatiotemporally coherent long-term PM<sub>2.5</sub> concentration dataset over
  China during 1980–2019 using a machine learning approach, Sci. Total
  Environ., 765, 144263, https://doi.org/10.1016/j.scitotenv.2020.144263,
  2021.
- 691

655

661

666

669

673

676

680

Li, H., Yang, Y., Wang, H., Wang, P., Yue, X., and Liao, H.: Projected Aerosol
Changes Driven by Emissions and Climate Change Using a Machine
Learning Method, Environ. Sci. Technol., 56, 7, 3884–3893,
https://doi.org/10.1021/acs.est.1c04380, 2022.

701

706

712

718

726

- Li, K., Jacob, D. J., Liao, H., Shen, L., Zhang, Q., and Bates, K. H.:
  Anthropogenic Drivers of 2013–2017 Trends in Summer Surface Ozone in
  China, P. Natl. Acad. Sci. USA., 116, 422–427,
  https://doi.org/10.1073/pnas.1812168116, 2019.
- Li, K., Jacob, D. J., Shen, L., Lu, X., De Smedt, I., and Liao, H.: Increases in surface ozone pollution in China from 2013 to 2019: anthropogenic and meteorological influences, Atmos. <u>Chem. Phys.</u>, 20, 11423–11433, https://doi.org/10.5194/acp-20-11423-2020, 2020.
- Li, K., Jacob, D. J., Liao, H., Qiu, Y., Shen, L., Zhai, S., Bates, K. H., Sulprizio,
   M. P., Song, S., Lu, X., Zhang, Q., Zheng, B., Zhang, Y., Zhang, J., Lee, H.
   C., and Kuk, S. K.: Ozone pollution in the North China Plain spreading into
   the late-winter haze season, P. Natl. Acad. Sci. USA, 118, 1–7,
   https://doi.org/10.1073/pnas.2015797118, 2021.
- Li, M., Yu, S., Chen, X., Li, Z., Zhang, Y., Wang, L., Liu, W., Li, P., Lichtfouse,
  E., Rosenfeld, D., and Seinfeld, J. H.: Large scale control of surface ozone
  by relative humidity observed during warm seasons in China, Environ.
  Chem. Lett., 19, 3981–3989, https://doi.org/10.1007/s10311-021-01265-0,
  20212021a.
- Li, R., ZhaoM., Wang, T., Shu, L., Qu, Y., Zhou, W., Meng, Y., Zhang, ZXie, M.,
  Liu, J., Wu, H., and Fu, H.: Developing a novel hybrid model for the
  estimation of Kalsoom, U.: Rising surface 8h ozone (O<sub>3</sub>) across the remote
  Tibetan Plateau during 2005–2018, in China from 2013 to 2017: A response
  to the recent atmospheric warming or pollutant controls?, Atmos. Environ.,
  246, 118130 Chem. Phys., 20, 6159–6175, https://doi.org/10.5194/acp-206159-20, 1016/j.atmosenv.2020.118130, 2021b.
- Lin, J.-T., and McElroy, M. B.: Impacts of boundary layer mixing on pollutant
  vertical profiles in the lower troposphere: Implications to satellite remote
  sensing, Atmos. Environ., 44, 1726–1739,
  https://doi.org/10.1016/j.atmosenv.2010.02.009, 2010.
- Tai
  Liu, R., Ma, Z., Liu, Y., Shao, Y., Zhao, W., and Bi, J.: Spatiotemporal distributions of surface ozone levels in China from 2005 to 2017: a machine learning approach, Environ. Int., 142, 105823, https://doi.org/10.1016/j.envint.2020.105823, 2020.
- Liu, S., Xing, J., Zhang, H., Ding, D., Zhang, F., Zhao, B., Sahu, S. K., and
  Wang, S.: Climate-driven trends of biogenic volatile organic compound
  emissions and their impacts on summertime ozone and secondary organic

- aerosol in China in the 2050s, Atmos. Environ., 218, 117020,
  https://doi.org/10.1016/j.atmosenv.2019.117020, 2019.
- 742
- Liu, X., Zhu, Y., Xue, L., Desai, A. R., and Wang, H.: Cluster-enhanced
  ensemble learning for mapping global monthly surface ozone from 2003 to
  2019, Geophys. Res. Lett., 49, e2022GL097947,
  https://doi.org/10.1029/2022GL097947, 2022.
- Liu, Y., and Wang, T.: Worsening urban ozone pollution in China from 2013 to
  2017–Part 1: The complex and varying roles of meteorology, Atmos. Chem.
  Phys., 20, 6305–6321, https://doi.org/10.5194/acp-20-6305-2020, 2020.
- 751

- Lou, S., Liao, H., and Zhu, B.: Impacts of aerosols on surface-layer ozone
  concentrations in China through heterogeneous reactions and changes in
  photolysis rates, Atmos. Environ., 85, 123–138,
  http://dx.doi.org/10.1016/j.atmosenv.2013.12.004, 2014.
- Lu, X., Zhang, L., Chen, Y., Zhou, M., Zheng, B., Li, K., Liu, Y., Lin, J., Fu, T.M., and Zhang, Q.: Exploring 2016–2017 surface ozone pollution over
  China: source contributions and meteorological influences, Atmos. Chem.
  Phys., 19, 8339–8361, https://doi.org/10.5194/acp-19-8339-2019, 2019.
- 761
- Lu, X., Zhang, L., Wang, X., Gao, M., Li, K., Zhang, Y., Yue, X., and Zhang, Y.:
  Rapid Increases in Warm-Season Surface Ozone and Resulting Health
  Impact in China since 2013, Environ. Sci. Technol. Lett., 7, 240–247,
  https://doi.org/10.1021/acs.estlett.0c00171, 2020.
- Malley, C. S., Henze, D. K., Kuylenstierna, J. C. I., Vallack, H., Davila, Y.,
  Anenberg, S. C., Turner, M. C., and Ashmore, M.: Updated Global
  Estimates of Respiratory Mortality in Adults ≥ 30 Years of Age Attributable
  to Long-Term Ozone Exposure, Environ. Health Perspect., 125, 087021,
  https://doi.org/10.1289/EHP1390, 2017.
- 772
- Mao, J., Paulot, F., Jacob, D. J., Cohen, R. C., Crounse, J. D., Wennberg, P. 773 774 O., Keller, C. A., Hudman, R. C., Barkley, M. P., and Horowitz, L. W.: Ozone and organic nitrates over the eastern United States: sensitivity to isoprene 775 776 chemistry. J. Geophys. Res. Atmos. 118, 11256-68. https://doi.org/10.1002/jgrd.50817, 2013. 777
- 778
- Marvin, M. R., Palmer, P. I., Latter, B. G., Siddans, R., Kerridge, B. J., Latif, M.
  T., and Khan, M. F.: Photochemical environment over Southeast Asia
  primed for hazardous ozone levels with influx of nitrogen oxides from
  seasonal biomass burning, Atmos. Chem. Phys., 21, 1917–1935,
  https://doi.org/10.5194/acp-21-1917-2021, 2021.

McLinden, C. A., Olsen, S. C., Hannegan, B., Wild, O., Prather, M. J., and
Sundet, J.: Stratospheric ozone in 3-D models: A simple chemistry and the
cross-tropopause flux, J. Geophys. Res. Atmos., 105, 14653–14665,
https://doi.org/10.1029/2000jd900124, 2000.

789

784

- Mills, G., Pleijel, H., Malley, C. S., Sinha, B., Cooper, O. R., Schultz, M. G.,
  Neufeld, H. S., Simpson, D., Sharps, K., Feng, Z., Gerosa, G., Harmens,
  H., Kobayashi, K., Saxena, P., Paoletti, E., Sinha, V., and Xu, X.:
  Tropospheric ozone assessment report: Present-day tropospheric ozone
  distribution and trends relevant to vegetation, Elem. Sci. Anth., 6,
  47, https://doi.org/10.1525/elementa.302, 2018.
- 796

802

806

- Murray, L. T., Jacob, D. J., Logan, J. A., Hudman, R. C., and Koshak, W. J.:
  Optimized regional and interannual variability of lightning in a global
  chemical transport model constrained by LIS/OTD satellite data, J.
  Geophys. Res. Atmos., D20307, https://doi.org/10.1029/2012jd017934,
  2012.
- Ni, R., Lin, J., Yan, Y., and Lin, W.: Foreign and domestic contributions to
  springtime ozone over China, Atmos. Chem. Phys., 18, 11447–11469,
  https://doi.org/10.5194/acp-18-11447-2018, 2018.
- O'Neill, B. C., Tebaldi, C., van Vuuren, D. P., Eyring, V., Friedlingstein, P., Hurtt,
  G., Knutti, R., Kriegler, E., Lamarque, J.-F., Lowe, J., Meehl, G. A., Moss,
  R., Riahi, K., and Sanderson, B. M.: The Scenario Model Intercomparison
  Project (ScenarioMIP) for CMIP6, Geosci. Model Dev., 9, 3461–3482,
  https://doi.org/10.5194/gmd-9-3461-2016, 2016.
- 812
- Ott, L. E., Pickering, K. E., Stenchikov, G. L., Allen, D. J., DeCaria, A. J., Ridley,
  B., Lin, R.-F., Lang, S., and Tao, W.-K.: Production of lightning NO<sub>x</sub> and its
  vertical distribution calculated from three-dimensional cloud-scale chemical
  transport model simulations, J. Geophys. Res., 115, D04301,
  https://doi.org/10.1029/2009JD011880, 2010.
- 818

- Park, R. J., Jacob, D. J., Field, B. D., Yantosca, R. M., and Chin, M.: Natural
  and transboundary pollution influences on sulfate-nitrate-ammonium
  aerosols in the United States: Implications for policy, J. Geophys. Res.
  Atmos., 109, 20, https://doi.org/10.1029/2003jd004473, 2004.
- Ploton, P., Mortier, F., Réjou-Méchain, M., Barbier, N., Picard, N., Rossi, V.,
   Dormann, C., Cornu, G., Viennois, G., Bayol, N., Lyapustin, A., Gourlet Fleury, S., and Pélissier, R.: Spatial validation reveals poor predictive
   performance of large-scale ecological mapping models, Nat. Commun., 11,

# 828 <u>1–11, https://doi.org/10.1038/s41467-020-18321-y, 2020.</u>

- Pommier, M., Fagerli, H., Gauss, M., Simpson, D., Sharma, S., Sinha, V.,
  Ghude, D. S., Landgren, O., Nyiri, A., and Wind, P.: Impact of regional
  climate change and future emission scenarios on surface O<sub>3</sub> and PM<sub>2.5</sub> over
  India, Atmos. Chem. Phys., 18, 103–27, https://doi.org/10.5194/acp-18103-2018, 2018.
- Pye, H. O. T., Liao, H., Wu, S., Mickley, L. J., Jacob, D. J., Henze, D. K., and
  Seinfeld, J. H.: Effect of changes in climate and emissions on future sulfatenitrate-ammonium aerosol levels in the United States, J. Geophys. Res.
  Atmos., 114, D01205, https://doi.org/10.1029/2008jd010701, 2009.
- Qian, J., Liao, H., Yang, Y., Li, K., Chen, L., and Zhu, J.: Meteorological
  influences on daily variation and trend of summertime surface ozone over
  years of 2015–2020: Quantification for cities in the Yangtze River Delta, Sci.
  Total Environ., 834, 155107,
  https://doi.org/10.1016/j.scitotenv.2022.155107, 2022.
- Rodriguez, J. D., Perez, A., and Lozano, J. A.: Sensitivity analysis of k-fold
  cross validation in prediction error estimation, IEEE T. Pattern Anal., 32,
  569–575, https://doi.org/10.1109/TPAMI.2009.187, 2010.
- Santurtún, A., González-Hidalgo, J. C., Sanchez-Lorenzo, A., and Zarrabeitia,
  M. T.: Surface ozone concentration trends and its relationship with weather
  types in Spain (2001–2010), Atmos. Environ., 101, 10–22,
  https://doi.org/10.1016/j.atmosenv.2014.11.005, 2015.
- Su, X., An, J., Zhang, Y., Zhu, P., and Zhu, B.: Prediction of ozone hourly
  concentrations by support vector machine and kernel extreme learning
  machine using wavelet transformation and partial least squares methods,
  Atmos. Pollut. Res., 6, 51–60, https://doi.org/10.1016/j.apr.2020.02.024,
  2020.
- Toh, Y. Y., Lim, S. F., and von Glasow, R.: The influence of meteorological
  factors and biomass burning on surface ozone concentrations at Tanah
  Rata, Malaysia, Atmos. Environ., 70, 435–446,
  https://doi.org/10.1016/j.atmosenv.2013.01.018, 2013.
- 866

861

829

835

840

846

850

855

van der Werf, G. R., Randerson, J. T., Giglio, L., van Leeuwen, T. T., Chen, Y., 867 Rogers, B. M., Mu, M., van Marle, M. J. E., Morton, D. C., Collatz, G. J., 868 Yokelson, R. J., and Kasibhatla, P. S.: Global fire emissions estimates 869 1997-2016, Earth Syst. Sci. Data, 9, 697-720, 870 during https://doi.org/10.5194/essd-9-697-2017, 2017. 871

- Wang, Y., Shen, L., Wu, S., Mickley, L. J., He, J., and Hao, J.: Sensitivity of
  surface ozone over China to 2000–2050 global changes of climate and
  emissions, Atmos. Environ., 75, 374–382,
  https://doi.org/10.1016/j.atmosenv.2013.04.045, 2013.
- 877
- Wang, Z., Lin, L., Xu, Y., Che, H., Zhang, X., Dong, W., Wang, C., Gui, K., and
  Xie, B.: Incorrect Asian aerosols affecting the attribution and projection of
  regional climate change in CMIP6 models, npj Clim. Atmos. Sci., 4, 2,
  https://doi.org/10.1038/s41612-020-00159-2, 2021.
- 882

887

892

- Wei, J., Li, Z., Li, K., Dickerson, R., Pinker, R., Wang, J., Liu, X., Sun, L., Xue,
  W., and Cribb, M.: Full-coverage mapping and spatiotemporal variations of
  ground-level ozone (O<sub>3</sub>) pollution from 2013 to 2020 across China, Remote
  Sens. Environ., 270, 112775, https://doi.org/10.1016/j.rse.2, 2022.
- Weng, X., Forster, G. L., and Nowack, P.: A machine learning approach to
   guantify meteorological drivers of ozone pollution in China from 2015 to
   2019, Atmos. Chem. Phys., 22, 8385–8402, https://doi.org/10.5194/acp-22 8385-2022, 2022.
- Xu, Z., Han, Y., Tam, C. Y., Yang, Z., and Fu, C.: Bias-corrected CMIP6 global
  dataset for dynamical downscaling of the historical and future climate
  (1979–2100), Sci. Data, 8, 293, https://doi.org/10.1038/s41597-02101079-3, 2021.
- 897

903

- Xue, T., Zheng, Y., Geng, G., Xiao, Q., Meng, X., Wang, M., Li, X., Wu, N.,
  Zhang, Q., and Zhu, T.: Estimating Spatiotemporal Variation in Ambient
  Ozone Exposure during 2013–2017 Using a Data-Fusion Model, Environ.
  Sci. Technol., 54, 14877–14888,
  https://dx.doi.org/10.1021/acs.est.0c03098, 2020.
- Yang, Y., Li, M., Wang, H., Li, H., Wang, P., Li, K., Gao, M., and Liao, H.: ENSO
  modulation of summertime tropospheric ozone over China, Environ. Res.
  Lett., 17, 034020, https://doi.org/10.1088/1748-9326/ac54cd, 2022.
- Yin, Z., Cao, B., and Wang, H.: Dominant patterns of summer ozone pollution
  in eastern China and associated atmospheric circulations, Atmos. Chem.
  Phys., 19, 13933–13943, https://doi.org/10.5194/acp-19-13933-2019, 2019.
- Yue, X., Unger, N., Harper, K., Xia, X., Liao, H., Zhu, T., Xiao, J., Feng, Z., and
  Li, J.: Ozone and haze pollution weakens net primary productivity in China,
  Atmos. Chem. Phys., 17, 6073–6089, https://doi.org/10.5194/acp-17-60732017, 2017.

Zanis, P., Akritidis, D., Turnock, S., Naik, V., Szopa, S., Georgoulias, A. K.,
Bauer, S. E., Deushi, M., Horowitz, L. W., Keeble, J., Le Sager, P.,
O'Connor, F. M., Oshima, N., Tsigaridis, K., and van Noije., T.: Climate
change penalty and benefit on surface ozone: a global perspective based
on CMIP6 earth system models, Environ. Res. Lett., 17, 024014,
https://doi.org/10.1088/1748-9326/ac4a34, 2022.

- Zhang, X., Zhao, L., Cheng, M., and Chen, D.: Estimating ground-level ozone
  concentrations in eastern China using satellite-based precursors, IEEE
  Trans. Geosci. Remote Sens., 58, 4754–4763, https://doi.org/
  10.1109/TGRS.2020.2966780, 2020.
- 928

923

- Zhou, C., Gao, M., Li, J., Bai, K., Tang, X., Lu, X., Liu, C., Wang, Z., and Guo,
  Y.: Optimal Planning of Air Quality-Monitoring Sites for Better Depiction of
  PM<sub>2.5</sub> Pollution across China, Environ. Au., 2, 314–323,
  https://doi.org/10.1021/acsenvironau.1c00051, 2022.
- Zhu, J., Liao, H., Mao, Y., Yang, Y., and Jiang, H.: Interannual variation,
  decadal trend, and future change in ozone outflow from East Asia, Atmos.
  Chem. Phys., 17, 3729–3747, https://doi.org/10.5194/acp-17-3729-2017,
  2017.

938

939	able 1. S	ummary of detailed datasets	s used in this	study.		
Dataset	Variable	Description	Spatial	Temporal	Time	
type			resolution	resolution	period	Data Source
						Assimilated
0	O <sub>3</sub>	Near-surface ozone	0.5°×0.625°	Monthly	2014–2019	GEOS-Chem
03		concentrations		(historical)	(historical)	simulations and
						Observations
	T_2m	Air temperature at 2 meters	0.5°×0.625°	Monthly (historical) Monthly (future)	2014–2019 (historical) 2020–2100 (future)	
	T_850	Air temperature at 850 hPa				MERRA-2 (historical) Adjusted CMIP6 (future)
	T_500	Air temperature at 500 hPa				
	U_850	Zonal wind at 850 hPa				
	U_500	Zonal wind at 500 hPa				
	V_850	Meridional wind at 850 hPa				
Meteorology	V_500	Meridional wind at 500 hPa				
	RH	Relative humidity				
	PRECP	Precipitation rate				
	CLT	Total cloud cover				
	Pene	Incoming shortwave radiation at				
	RODO	the surface				
	SLP	Sea level pressure				
		Nitric oxide from anthropogenic	Monthly (historical) 0.5°×0.625° Monthly (future)	2014–2019		
	NOx CO	sources		Monthly (historical) Monthly (future)	(historical)	
		Nitric oxide from biomass burning			2019	
					(future)	
		Nitric oxide from soil sources			2016	
		Nitric oxide from lightning sources			2016	CEDS
		Carbon monoxide from				(Anthropogenic)
		anthropogenic sources				GFED4
Emission		Carbon monoxide from biomass			2014-2019	(Biomass burning)
		burning			(historical)	MEGAN V-2 1
		Non-methane volatile organic			2019	(Biogenic)
		compounds from anthropogenic			(future)	(9)
		sources			(	
		Non-methane volatile organic				
		compounds from biomass burning				
		Non-methane volatile organic			2016	
	۹	compounds from biogenic sources				
	LC	Land cover	300 m×300 m		2014–2019	ESA CCI
Land use	Normalized Di NDVI	Normalized Difference Vegetation	0.05°×0.05°	Monthly	(historical)	
		Index			2019	AVHRR
					(future)	
Topography	TOPO	Digital elevation model	90 m×90 m	-	2010	SRTM
Population	POP	Population density	1 km×1 km	-	2010	Land Scan

# 1 Summary of d





950

**Figure 2.** Spatial distributions of observed <u>near-surface O<sub>3</sub> concentrations</u> across China (a) and assimilated <u>O<sub>3</sub> concentrations over Asia</u> (b) <del>O<sub>3</sub></del> concentrations in 2014–2019 over Asia. Correlation coefficient (R) between observed and assimilated O<sub>3</sub> and the normalized mean bias (NMB =  $\sum$ (Observed – Assimilated) /  $\sum$ Assimilated × 100%) are given at the bottom left of panel (a).

\_\_\_\_\_



Figure 23. Density scatterplots of 10-fold cross-validation results forpredicted
vs assimilated monthly near-surface O3 concentrations (ppb) in 2019 over Asia.
The gray and red lines are the 1:1 line and linear regression line, respectively.
Statistical metrics including the number of samples (N), correlation of
determination (R<sup>2</sup>, unitless), root mean square error (RMSE, ppb), mean
absolute error (MAE, ppb), and mean relative error (MRE, %) are shown at the
top left.





Figure 34. Spatial distributions of the performance statistics of the random forest (RF)ML model with regard to (a) R<sup>2</sup> (unitless), (b) RMSE (ppb), (c) MAE
 (ppb), and (d) MRE (%) during 2014–2019 over Asia.



**Figure 45.** Importance scores of independent variables (meteorological parameters, emissions, land use, topography, and population density) used in the ML model for predicting future near-surface O<sub>3</sub> concentrations over Asia.



982 Figure 56. The spatial distributions of absolute (ppb) and percentage difference 983 (%) of surface O<sub>3</sub> level between 2025 (2020–2029 mean) and 2095 (2091–2100 984 mean) driven by climate change under four scenarios (a, e) SSP1-2.6, (b, f) 985 SSP2-4.5, (c, g) SSP3-7.0 and (d, h) SSP5-8.5. The box-outlined areas in (d) 986 are North China (NC, 35°-41°N, 105°-120°E), South China (SC, 22°-33.5°N, 987 105°-120°E), Southeast Asia (SEA, -9.5°S-19°N, 93.75°-140°E), South India 988 989 (SI, 8°-18°N, 73.125°-80.625°E), Gangetic Plains (GP, 21.5°-23.5°N, 85.625°-92.5°E, 23.5°-27°N, 76.25°-92.5°E, and 27°-30°N, 76.25°-81.25°E), 990 and Tibetan Plateau (TP, 28°-31°N, 81.875°-102.5°E and 31°-38°N, 73.125°-991 102.5°E). No overlaying hatch pattern indicates statistical significance with 95% 992 993 confidence from a two-tailed t test.



995

Figure 67. Time series (2020-2100) of annual mean near-surface O3 996 concentrations (ppb) driven by climate change under the four scenarios (SSP1-997 998 2.6, SSP2-4.5, SSP3-7.0 and SSP5-8.5) over North China (NC, 35°-41°N, 105°-120°E), South China (SC, 22°-33.5°N, 105°-120°E), Southeast Asia 999 (SEA, -9.5°S-19°N, 93.75°-140°E), South India (SI, 8°-18°N, 73.125°-1000 80.625°E), Gangetic Plains (GP, 21.5°-23.5°N, 85.625°-92.5°E, 23.5°-27°N, 1001 76.25°-92.5°E, and 27°-30°N, 76.25°-81.25°E), and Tibetan Plateau (TP, 28°-1002 31°N, 81.875°–102.5°E and 31°–38°N, 73.125°–102.5°E). The black lines are 1003 the averages of the predicted O<sub>3</sub> based on meteorological fields from 18 CMIP6 1004 1005 models.



Figure 78. Absolute (a, ppb) and percentage (b, %) changes in projected nearsurface climate-driven O3 concentrations in 2095 (2091–2100 mean) relative to

surface climate-driven  $O_3$  concentrations in 2095 (2091–2100 mean) relative to 2025 (2020–2029 mean) over the six selected key regions of Asia, including NC, SC, SEA, SI, GP, and TP under four future climate scenarios. The error bars indicate one standard deviation.



1014

1015

Figure 89. The spatial distributions of percentage differences (%) in near-1016 surface O3 concentrations between 2025 (2020-2029 mean) and 2095 (2091-1017 2100 mean) driven by climate change under four scenarios (SSP1-2.6, SSP2-1018 4.5, SSP3-7.0 and SSP5-8.5, from left to right) averaged in MAM (March-April-1019 May), JJA (June-July-August), SON (September-October-November), and 1020 DJF (December-January-February) (from top to bottom). No overlaying hatch 1021 pattern indicates statistical significance with 95% confidence from a two-tailed 1022 1023 t test.