# Elucidate the Formation Mechanism of Particulate Nitrate Based on Direct Radical Observations in the Yangtze River Delta summer 2019

4 Tianyu Zhai<sup>a</sup>, Keding Lu<sup>a, b\*</sup>, Haichao Wang<sup>c</sup>, Shengrong Lou<sup>d</sup>, Xiaorui Chen<sup>a, f</sup>, Renzhi

5 Hu<sup>e</sup>, Yuanhang Zhang<sup>a, b\*</sup>

<sup>a</sup> State Key Joint Laboratory of Environmental Simulation and Pollution Control,
College of Environmental Sciences and Engineering, Peking University, Beijing
100871, China.

<sup>b</sup> Collaborative Innovation Center of Atmospheric Environment and Equipment
Technology, Nanjing University of Information Science & Technology, Nanjing
210044, China.

<sup>c</sup> School of Atmospheric Sciences, Sun Yat-sen University, Guangzhou 510275, China.

<sup>13</sup> <sup>d</sup> State Environmental Protection Key Laboratory of Formation and Prevention of the

14 Urban Air Complex, Shanghai Academy of Environmental Sciences, Shanghai, 200223,

15 China.

16 <sup>e</sup> Key Laboratory of Environmental Optics and Technology, Anhui Institute of Optics

17 and Fine Mechanics, Chinese Academy of Sciences, Hefei, 230031, China.

<sup>f</sup> Now at: Department of Civil and Environmental Engineering, The Hong Kong
 Polytechnic University, Hong Kong, China.

20

21 *\* Correspondence to:* 

22 Keding Lu (k.lu@pku.edu.cn), Yuanhang Zhang (yhzhang@pku.edu.cn)

23

24 Abstract. Particulate nitrate  $(NO_3)$  is one of the dominant components of fine particles 25 in China, especially during pollution episodes, and has a significant impact on human 26 health, air quality, and climate. Here a comprehensive field campaign that focuses on 27 the atmospheric oxidation capacity and aerosol formation and their effects in the 28 Yangtze River Delta (YRD) was conducted from May to June 2019 at a regional site in 29 Changzhou, Jiangsu province in China. The concentration of NO<sub>3</sub><sup>-</sup>, OH radical, N<sub>2</sub>O<sub>5</sub>, 30 NO<sub>2</sub>, O<sub>3</sub>, and relevant parameters were measured simultaneously. We showed a high NO<sub>3</sub><sup>-</sup> mass concentration with 10.6  $\pm$  8.9 µg m<sup>-3</sup> on average, which accounted for 38.3 % 31 32 of total water-soluble particulate components and 32.0 % of total PM<sub>2.5</sub>, followed by 33 the proportion of sulfate, ammonium, and chloride by 26.0 %, 18.0 %, and 2.0 %, 34 respectively. This result confirmed that the heavy nitrate pollution in eastern China not 35 only happened in winter but also in the summertime. This study's high nitrate oxidation 36 ratio (NOR) emphasizes the solid atmospheric oxidation and fast nitrate formation 37 capacity in YRD. It is found that OH + NO<sub>2</sub> during daytime dominates nitrate formation 38 on clean days while N<sub>2</sub>O<sub>5</sub> hydrolysis vastly enhanced and became comparable with that 39 of OH + NO<sub>2</sub> during polluted days (67.2 % and 30.2 %, respectively). An updated 40 observed-constrain Empirical Kinetic Modeling Approach (EKMA) was used to assess 41 the kinetic controlling factors of both local O<sub>3</sub> and NO<sub>3</sub><sup>-</sup> productions, which indicated that the O<sub>3</sub>-targeted scheme (VOCs:  $NO_x = 2$ : 1) is adequate to mitigate the O<sub>3</sub> and 42 43 nitrate pollution coordinately during summertime in this region. Our results promote 44 the understanding of nitrate pollution mechanisms and mitigation based on field 45 observation and model simulation and call for more attention to nitrate pollution in the 46 summertime.

#### 47 Keywords:

48 Nitrate pollution; Dinitrogen pentoxide; Nitrate formation; Pollution mitigation

# 49 **1 Introduction**

50 Chemical compositions of fine particles have been measured in China during the past 51 twenty years, and secondary inorganic aerosol is regarded as one of the dominant 52 species in aerosol (Cao et al., 2012; Hagler et al., 2006; Zhao et al., 2013; Andreae et 53 al., 2008). Since the Air Pollution Prevention and Control Action Plan, there has been 54 a significant decrease in SO<sub>2</sub>, NO<sub>2</sub>, and PM<sub>2.5</sub> concentration in China, while the 55 inorganic nitrate ratio in PM<sub>2.5</sub> increased and became the considerable component in 56 PM<sub>2.5</sub> (Shang et al., 2021; Zhang et al., 2022). Therefore, a comprehensive 57 understanding of the particlate nitrate formation mechanism is essential and critical to 58 mitigating haze pollution in China.

59 Massive research has been done in China to investigate nitrate formation 60 mechanisms, and a basic framework has been established (Sun et al., 2006; Chang et 61 al., 2018; Wu et al., 2019). In the daytime,  $NO_2 + OH$  radical oxidation (Reaction 1) is the major particulate nitrate formation pathway. The product (HNO<sub>3</sub>) reacts with 62 63 alkaline substances in aerosol, generating particulate nitrate. This pathway is mainly 64 controlled by precursors concentration as well as the gas-particle partition of gaseous 65 nitric acid, and particulate nitrate depends on temperature, relative humidity (RH), NH<sub>3</sub> 66 concentration, and aerosol acidity (Wang et al., 2009; Song and Carmichael, 2001; 67 Meng et al., 2020; Zhang et al., 2021). At night, N<sub>2</sub>O<sub>5</sub> uptake is a vital nitrate formation pathway (Reaction 4)(Chen et al., 2020; Wang et al., 2022). N<sub>2</sub>O<sub>5</sub> is formed through 68  $NO_2 + NO_3$  (Reaction 3) and there exists a quick thermal equilibrium balance (K<sub>eq</sub> = 5.5 69  $\times$  10<sup>-17</sup> cm<sup>-3</sup> molecule<sup>-1</sup> s<sup>-1</sup>, 298 K). However, two problems remain ambiguous in 70 71 quantifying the contribution of N<sub>2</sub>O<sub>5</sub> uptake to nitrate formation. The first is the N<sub>2</sub>O<sub>5</sub> 72 heterogeneous uptake coefficient ( $\gamma$ ) on ambient aerosol is highly varied with the range from 10<sup>-4</sup> to 10<sup>-1</sup> based on previous lab and field measurements (Bertram and Thornton, 73 74 2009; Brown et al., 2009; Wang et al., 2017c; Wang and Lu, 2016). The other one is 75 CINO<sub>2</sub> production yield which influences nitrate contribution due to the extensive 76 variation range (Phillips et al., 2016; Staudt et al., 2019; Tham et al., 2018). Both two parameters are complex to well-predicted by current schemes. NO<sub>2</sub> heterogeneous
uptake has been found nonnegligible for nitrate formation, which can be a vital pathway
during heavy haze events, according to recent study (Qiu et al., 2019; Chan et al., 2021).
The uptake coefficient and nitrate yield remain uncertain, as same as the N<sub>2</sub>O<sub>5</sub>
heterogeneous reaction. Besides, N<sub>2</sub>O<sub>5</sub> homogeneous hydrolysis and NO<sub>3</sub> radical
oxidation have a minor contribution to particulate nitrate under ambient conditions
(Brown et al., 2009; Seinfeld and Pandis, 2016).

$$NO_2 + OH \rightarrow HNO_3$$
 R1

$$NO_2 + O_3 \rightarrow NO_3 + O_2$$
 R2

$$NO_2 + NO_3 + M \rightarrow N_2O_5 + M$$
 R3.1

$$N_2O_5 + M \rightarrow NO_2 + NO_3 + M$$
 R3.2

$$N_2O_5 + (H_2O \text{ or } Cl^-) \rightarrow (2 - \varphi) NO_3^- + \varphi ClNO_2$$
 R4

84 As a critical area of China's economy and industry, Yangtze River Delta (YRD) has suffered severe air pollution during past decades, and fine particle pollution in YRD 85 86 has raised a widespread concern (Guo et al., 2014; Zhang et al., 2015; Zhang et al., 2017; Ming et al., 2017; Xue et al., 2019). However, most research focuses on 87 88 wintertime PM<sub>2.5</sub> pollution and lacks measurements of critical intermediate species and 89 radicals to assess the importance of each nitrate formation pathway. In this study, with 90 the direct measurements of hydroxyl radical and the reactive nitrogen compounds and 91 chemical box model analysis, we explore the characteristics of nitrate and precursors in 92 YRD in the summer of 2019, the importance of particulate nitrate formation pathways 93 is quantified, and the controlling factors are explored. A further suggestion for summer 94 pollution prevention and control in the local area is proposed.

# 95 2 Site description and methods

# 96 2.1 The campaign site

This campaign took place at a suburban sanatorium from May 30<sup>th</sup> to June 18<sup>th</sup>, 2019, 97 98 in Changzhou, China. Changzhou (119.95 °E, 31.79 °N) is located in Jiangsu province 99 and about 150 km northwest of Shanghai. The sanatorium, located 420 m east of Lake 100 Ge (one of the largest lakes in Jiangsu province, 164 square kilometers), is surrounded 101 by farmland and fishponds. With the closest arterial traffic 1 km away, several industry 102 zones are 4 km to the east. The prevailing wind was from the south and southeast sectors 103 (about 30 % of the time) compared to 20 % from the west sector, of which only 15 % came from the east. The wind speed was usually lower than 5 m s<sup>-1</sup> with faster speed 104 from the west. This site was influenced by anthropogenic and biological sources with 105 106 occasional biomass burning.



107

Figure 1 The location of the campaign site (red star), Changzhou, is 150 km on thenorthwest side of Shanghai.

# 110 **2.2 The instrumentation**

111 Multiple gaseous and particlate parameters were measured simultaneously during the

112 campaign to comprehensively interpret the nocturnal atmospheric capacity and aerosol 113 formation. The related instruments are listed in Table 1.  $N_2O_5$  and Particle Number and 114 Size Distribution (PNSD) were measured on the fourth floor of the sanatorium, which 115 is the top of the building. Other instruments were placed in containers on the ground 116 170 m northeast of the building and sampling inlets at circa 5 m above the ground 117 through the containers' roof.

118 N<sub>2</sub>O<sub>5</sub> was measured by Cavity Enhanced Absorption Spectrometer (CEAS) based 119 on Lambert-Beer's law which was developed by (Wang et al., 2017b). Briefly, air 120 samples were drawn through the window and reached out of the wall 30 cm to prevent 121 influence from surface deposition. The aerosol membrane filter was deployed before 122 the PFA sampling tube and changed every 2 hours at night to avoid a decrease in N<sub>2</sub>O<sub>5</sub> 123 transmission efficiency due to the increased loss of N<sub>2</sub>O<sub>5</sub> from the accumulated aerosols 124 on the filter. N<sub>2</sub>O<sub>5</sub> was decomposed to NO<sub>3</sub> and NO<sub>2</sub> through preheating tube heat at 125 130 °C and detected within a PFA-coated resonator cavity heated at 110 °C to prevent the formation of N<sub>2</sub>O<sub>5</sub> by reversible reaction subsequently. At the end of each sampling 126 cycle (5 min), a 30 s injection of high concentration NO (10 ppm, 20 ml min<sup>-1</sup>) mixed 127 128 with sample air was set to eliminate NO<sub>3</sub>-N<sub>2</sub>O<sub>5</sub> in the system. The NO titration 129 spectrums were adopted as the dynamic background spectrum by assuming no H<sub>2</sub>O 130 concentration variation in a single sampling cycle. The loss of N<sub>2</sub>O<sub>5</sub> in the sampling 131 system and filter was also considered during data correction. The limit of detection 132 (LOD) was estimated to be 2.7 pptv  $(1 \sigma)$  with an uncertainty of 19 %.

OH radical measurement was conducted by Fluorescence Assay by Gas Expansion Laser-Induced Fluorescence techniques (FAGE-LIF). Ambient air was expanded through a 0.4 mm nozzle to low pressure in a detection chamber, where the 308 nm laser pulse irradiated OH radical at a repetition rate of 8.5 kHz (Chen et al., 2018). NO<sub>x</sub> and O<sub>3</sub> were monitored by commercial monitors (Thermo-Fisher 42i and 49i). Volatile organic compounds (VOCs) were measured by an automated Gas Chromatograph equipped with a Mass Spectrometer and flame ionization detector (GC-MS) with a time resolution of 60 min. The photolysis frequencies were determined from the spectralactinic photon flux density measured by a spectroradiometer (Bohn et al., 2008).

142 PM<sub>2.5</sub> concentration was obtained by Tapered Element Oscillating Microbalance 143 (TEOM 1405, Thermo Scientific Inc). Aerosol surface concentration (Sa) was 144 converted from particle number and size distribution, which was measured by Scanning 145 Mobility Particle Sizer (SMPS, TSI 3936) and Aerosol Particle Sizer (APS, TSI 3321) 146 and modified to the wet particle-state S<sub>a</sub> with a hygroscopic growth factor (Liu et al., 147 2013). The uncertainty of the wet  $S_a$  was ~ 30 %. Meanwhile, water-soluble particulate 148 components and their gaseous precursors were analyzed through the Monitor for 149 AeRosols and GAses in ambient air (MARGA, Chen et al. (2017)). Meteorological data 150 were also available, including the temperature, relative humidity (RH), pressure, wind 151 speed, and wind direction.

Parameters	Detection of limit	Method	Accuracy
N <sub>2</sub> O <sub>5</sub>	2.7 pptv (1 σ, 1 min)	CEAS	$\pm 19\%$
ОН	$1.6 \times 10^5 \mathrm{cm}^{-3} (1 \sigma, 60 \mathrm{s})$	LIF <sup>a</sup>	$\pm 21 \%$
NO	60 pptv (2 σ, 1 min)	$PC^{c}$	$\pm \ 10 \ \%$
NO <sub>2</sub>	0.3 ppbv (2 σ, 1 min)	$PC^{c}$	$\pm10$ %
O <sub>3</sub>	0.5 ppbv (2 σ, 1 min)	UV photometry	± 5 %
VOCs	20-300 pptv (60 min)	GC-MS	$\pm 15 \%$
PM <sub>2.5</sub>	0.1 μg m <sup>-3</sup> (1 min)	TEOM <sup>d</sup>	$\pm$ 5 %
Photolysis frequencies	5×10 <sup>-5</sup> s <sup>-1</sup> (1 min)	SR <sup>e</sup>	$\pm10$ %
PNSD	14 nm -700 nm (4 min)	SMPS, APS	$\pm \ 10 \ \%$
HNO3, NO3, HCl	0.06 ppbv (30 min)	MARGA <sup>f</sup>	$\pm 20 \%$
NH4 <sup>+</sup> , NO3 <sup>-</sup> , Cl <sup>-</sup> , SO4 <sup>2-</sup>	0.05 µg m <sup>-3</sup> (30 min)	MARGA <sup>f</sup>	$\pm 20 \%$

152 **Table 1** The observed gas and particle parameters during the campaign.

<sup>153</sup> <sup>a</sup> Laser-induced fluorescence; <sup>b</sup> Chemiluminescence; <sup>c</sup> Photolytic converter; <sup>d</sup> Tapered

154 Element Oscillating Microbalance; <sup>e</sup> Spectroradiometer; <sup>f</sup> the Monitor for AeRosols and

155 GAses in ambient air.

### 156 **2.3 The empirical kinetic modeling approach**

157 A box model coupled with the Regional Atmospheric Chemical Mechanism version 2

158 (RACM2, Goliff, Stockwell & Lawson, 2013) is used to conduct the mitigation

159 strategies studies. The model is operated in one-hour time resolution with measurement

160 results of temperature, relative humidity, pressure, CO, NO<sub>2</sub>, H<sub>2</sub>O, photolysis frequencies, and aggregated VOCs input to constrain the model. It should be noted that 161 162 HONO concentration is calculated by NO<sub>2</sub> times 0.02, as suggested by Elshorbany et al. 163 (2012), and has been used in the box model before (Lou et al., 2022). Long-lived species 164 such as H<sub>2</sub> and CH<sub>4</sub> are assumed as constants (550 ppbv and 1900 ppbv, respectively). Moreover, a 13-hour constant loss rate of unconstrained intermediate and secondary 165 166 products, the result of synthetic evaluating secondary simulation of secondary species, 167 is set for representing the multi-effects of deposition, transformation, and transportation. 168 The approaches to the chemical production of  $O_3$  (P( $O_3$ )) and inorganic nitrate 169 (P(NO<sub>3</sub><sup>-</sup>)) are described in previous articles (Tan et al., 2021; Tan et al., 2018) and 170 expressed as Equation 1 and 4:

$$P(O_3) = F(O_3) - D(O_3)$$
 Eq1

$$F(O_3) = k_{HO_2+NO}[NO][HO_2] + k_{(RO_2+NO)eff}[NO][RO_2]$$
 Eq2

$$D(O_3) = k_{OH+NO_2}[OH][NO_2] + (k_{OH+O_3}[OH] + k_{HO_2+O_3}[HO_2] + k_{alkenes+O_3}[alkenes])[O_3]$$
Eq3

$$P(NO_3) = P(HNO_3) + P(pNO_3)$$
 Eq4

$$P(HNO_3) = k_{OH+NO_2} [OH][NO_2]$$
Eq5

$$P(pNO_3) = 0.25(2 - \varphi) C \gamma S_a [N_2O_5]$$
 Eq6

171 briefly,  $P(O_3)$  is net ozone production, which is calculated by peroxyl radial + NO 172 oxidation (Eq. 2) minus the chemical loss of  $O_3$  and  $NO_2$  (Eq. 3). P( $NO_3^-$ ) is constituted by reaction  $OH + NO_2$  (Eq. 5) and  $N_2O_5$  heterogeneous uptake (Eq. 6). Here, rate 173 174 constants of reactions are obtained from NASA JPL Publication (Burkholder et al., 175 2015) or RACM2 (Goliff et al., 2013).  $\gamma$  is the N<sub>2</sub>O<sub>5</sub> uptake coefficient calculated from 176 parameterization ( $\gamma_{P}$ , more details in chapter 3.3).  $\varphi$  represents CINO<sub>2</sub> production yield 177 through N<sub>2</sub>O<sub>5</sub> hydrolysis, and the mean value reported by Xia et al. (2020) is used in this work. 178

179 The empirical Kinetic Modeling Approach (EKMA) was innovated to study the 180 effects of precursors (VOCs and NO<sub>x</sub>) reactivity on the region's ozone pollution by 181 Kanaya et al., which helps recognize the region's susceptibility to precursors by weight 182 and become a prevalent tool to study the process of ozone formation (Tan et al., 2018; 183 Yu et al., 2020b; Kanaya et al., 2008). The prevention and control problem of pollutant 184 generation can be transformed through the EKMA curve to reduce its precursors' emissions. Furthermore, the precursor reduction scheme needed for total pollutant 185 control is given qualitatively. P(NO<sub>3</sub><sup>-</sup>) can also be analyzed through EKMA for the 186

- 187 nonlinear secondary formation relationship with precursor reactivity. Here, an isopleth
- 188 diagram of the net ozone production rate as functions of the reactivities of  $NO_x$  and
- 189 VOCs can be derived from EKMA. In detail, 0.01 to 1.2 emission reduction strategy
- 190 assumptions are exponential interpolation into 20 kinds of emission situations of  $NO_x$
- and VOCs, respectively, which in total counts 400 scenarios.

#### 192 **2.4 The calculation of aerosol liquid water content**

193 Aerosol liquid water content (ALWC) is calculated through ISORROPIA II

194 (Fountoukis and Nenes, 2007). Forward mode is applied in this study. Furthermore,

195 water-soluble particulate components in  $PM_{2.5}$  and gaseous species ( $NH_3 + HNO_3 + HCl$ )

196 obtained from MARGA, along with RH and T, are input as initial input. In addition,

197 metastable aerosol state is chosen due to high RH during this campaign.

### 198 **3 Result and discussion**

# 199 **3.1 Overview of measurements**

200 The time used in this study is China Standard Time (UTC + 8) and the local sunrise and 201 sunset time during the campaign were around 5 am and 7 pm, respectively. The whole 202 campaign period is divided into four PM<sub>2.5</sub> clean periods and four PM<sub>2.5</sub> polluted 203 periods (9 out of 14 days, the latter polluted periods days day refer to PM<sub>2.5</sub> pollution 204 except specified description) according to the Chinese National Air Quality Standard (CNAAQS) Grade I of daily PM<sub>2.5</sub> concentrations ( $< 35.0 \ \mu g \ m^{-3}$ ). Figure 2 shows the 205 206 meteorological parameters, and gas-phase and particulate species timeseries during the 207 observation. During the campaign, the temperature was high; the maximum reached 208 34.5 °C, with an average of  $25.1 \pm 3.7$  °C. RH changed drastically from 21 % to 88 %, with a mean value of  $58.9 \pm 14.0$  %. The mean NO<sub>2</sub> concentration was  $14.8 \pm 9.5$  ppbv. 209 210 Meanwhile, the  $O_3$  average was 54.6  $\pm$  28.8 ppbv, exceeding CNAAQS Grade II for a maximum daily average of 8 h ozone (160  $\mu$ g m<sup>-3</sup>) on 14 out of 19 days and exceeding 211  $200 \ \mu g \ m^{-3}$  on six days. 212



Figure 2 Timeseries of NO<sub>2</sub>, O<sub>3</sub>, N<sub>2</sub>O<sub>5</sub>, OH radical, PM<sub>2.5</sub>, and water-soluble particulate components, temperature, and RH. The vertical dotted line represents the zero clock. The black horizontal solid line in O<sub>3</sub> and PM<sub>2.5</sub> panels represents Chinese national air quality standards for O<sub>3</sub> and PM<sub>2.5</sub>, respectively. The top panel color blocks represent the PM<sub>2.5</sub> clean day (light green) and PM<sub>2.5</sub> polluted day(salmon).

219

213

Daytime OH radical ranged from  $2 \times 10^6$  to  $8 \times 10^6$  molecular cm<sup>-3</sup> with a daily 220 peak over  $3 \times 10^6$  molecular cm<sup>-3</sup>. Maximum OH radical reached  $8.18 \times 10^6$  molecular 221 222 cm<sup>-3</sup> in this campaign. Compared with other summertime OH radical observed in China, 223 OH radical concentration in this site is relatively low but still on the same order of 224 magnitude (Lu et al., 2012; Lu et al., 2013; Ma et al., 2022; Tan et al., 2017; Woodward-Massey et al., 2020; Yang et al., 2021). N<sub>2</sub>O<sub>5</sub> mean concentration was  $21.9 \pm 39.8$  pptv 225 226 with a nocturnal average of  $61.0 \pm 63.1$  pptv and a daily maximum of over 200 pptv at 227 eight nights. The maximum concentration of N<sub>2</sub>O<sub>5</sub> (477.2 pptv, 5 min resolution) appeared at 20:47 on June 8<sup>th</sup>. The average NO<sub>3</sub> radical production rate P(NO<sub>3</sub>) is 2.1 228  $\pm$  1.4 ppbv h<sup>-1</sup> with nocturnal average P(NO<sub>3</sub>) 2.8  $\pm$  1.6 ppbv h<sup>-1</sup> and daytime P(NO<sub>3</sub>) 229 230  $2.2 \pm 1.4$  ppbv h<sup>-1</sup>. P(NO<sub>3</sub>) is about twice of documented value in Taizhou and North 231 China Plain (Wang et al., 2017a; Wang et al., 2018b; Wang et al., 2020a), but close to

232 another result in YRD before (Chen et al., 2019). The average  $PM_{2.5}$  was  $34.6 \pm 17.8$  $\mu g m^{-3}$  with a maximum reach of 163.0  $\mu g m^{-3}$ . The water-soluble particulate 233 components of PM<sub>2.5</sub> are displayed as well. The average NO<sub>3</sub><sup>-</sup> concentration was 10.6 234 235  $\mu$ g m<sup>-3</sup>, which accounts for 38.3 % mass concentration of water-soluble particulate components and 32.0 % total PM<sub>2.5</sub>, while the proportion of sulfate, ammonium, and 236 237 chloride is 26.0 %, 18 %, and 2.0 % respectively. To sum up, during the campaign 238 period, the pollution of PM<sub>2.5</sub> would be generally exacerbated on high O<sub>3</sub> and NO<sub>2</sub> days. 239 Precipitation occurred during four clean processes receded pollutant concentration; 240 otherwise, the pollution condition remained severe.

241 The mean diurnal variations (MDC) of temperature, RH, NO<sub>2</sub>, O<sub>3</sub>, P(NO<sub>3</sub>), N<sub>2</sub>O<sub>5</sub>, 242 OH radical, and PM<sub>2.5</sub> in different air quality are shown in Figure 3. The temperature, 243 RH, and OH radical MDC show indistinctive differences between clean days (CD) and 244 polluted days (PD). The MDC of NO<sub>2</sub> has two concentration peaks that appear at 06:00 245 and 21:00 on CD, while at PD, its peak appears at 20:00 and maintains a high level 246 during the whole night. O<sub>3</sub> diurnal pattern reflects a typical urban-influenced character 247 with a maximum O<sub>3</sub> peak that lasts four hours from 14:00 to 17:00, while polluted-day 248  $O_3$  peak concentration is 1.2 higher than clean-day. P(NO<sub>3</sub>) grows after the  $O_3$  peak and maximum P(NO<sub>3</sub>) shows at 19:00 with an average value of 1.7 ppbv h<sup>-1</sup> on CD. By 249 250 contrast, the mean polluted-day P(NO3) is 2.6 ppbv h-1, and the maximum value reaches 4.7 ppbv  $h^{-1}$ . In contrast, the clean-day N<sub>2</sub>O<sub>5</sub> has a higher average and maximum 251 252 concentration than PD, which suggests a faster removal process during PD. PM<sub>2.5</sub> has 253 a similar trend with  $P(NO_3)$  and has a higher concentration during nighttime.

254



Figure 3 The mean diurnal variations of temperature, RH, NO<sub>2</sub> (Salmon), O<sub>3</sub>, P(NO<sub>3</sub>),
N<sub>2</sub>O<sub>5</sub>, OH radical(orange), and PM<sub>2.5</sub> of clean day and polluted day.

# **3.2 The evolution of nitrate pollution**

259 Figure 4 (a) shows the relationship between nitrate and sulfate with water-soluble 260 particulate components. Nitrate positively correlates with total water-soluble particulate components, while the sulfate ratio has an inverse correlation. With PM<sub>2.5</sub> 261 262 concentration increasing, nitrate proportion increases rapidly and keeps high weight at 263 heavy PM<sub>2.5</sub> period while sulfate ratio pears opposite phenomenon. Once the mass concentration of total water-soluble particulate component is over 30  $\mu$ g m<sup>-3</sup>, the mass 264 265 fraction of nitrate in total water-soluble particulate components is up to 50 % on average. 266 This result illustrates that particulate nitrate is one of the vital sources of explosive 267 growth particulate matter.



Figure 4 (a) Particulate ion mass concentration ratio of nitrate and sulfate to water
soluble ion. (b) NOR against RH, colored with temperature. (c) SOR against RH,
colored with temperature.

272

To further assess the conversion capacity of nitrate and sulfate in this site, the sulfur oxidation ratio (SOR) and the nitrogen oxidation ratio (NOR) are used to indicate the secondary transformation ratio of SO<sub>2</sub> and NO<sub>2</sub>, respectively (Sun et al., 2006). SOR and NOR are estimated using the formulae below:

277

$$SOR = \frac{nSO_4^{2-}}{nSO_4^{2-} + nSO_2}$$
Eq7

$$NOR = \frac{1}{nNO_3 + nNO_2} Eq8$$

278 Where n refers to the molar concentration, the higher SOR and NOR represent more 279 oxidation of gaseous species into a secondary aerosol. As depicted in Figure 4 (b-c), 280 NOR rapidly increases at RH < 45 %, remains constant at 45 % < RH < 75 %, and ends 281 with a sharp increase at RH > 75 %. During the study period, not only is the average 282 concentration of NO2 higher among PD but there is also a significant difference 283 between PD and CD NOR. The average values of NOR are 0.32 in PD and 0.25 in CD, 284 respectively, which manifests the more secondary transformation and pollution 285 potential in PD. In contrast, the SOR stays constant at a high value ( $\sim 0.5$ ) during the 286 whole RH scale, which shows a different pattern from previous research (Li et al., 2017; 287 Zheng et al., 2015). One possible explanation is that SO<sub>2</sub> concentration stays low during 288 the whole campaign ( $4.4 \pm 2.4$  ppbv on average), and SO<sub>2</sub> oxidation depends on the 289 limit of SO<sub>2</sub> instead of oxidation capability. Meanwhile, the mean SOR in both 290 situations is over 0.5 (0.52 in CD and 0.56 in PD), further supporting the SO<sub>2</sub> limited 291 hypothesis. Besides, Table 2 summarizes NOR and SOR values in YRD. NOR and 292 SOR in this study are similar to values reported in other YRD research (Shu et al., 2019; 293 Zhang et al., 2020b; Qin et al., 2021; Zhao et al., 2022), except values in 2013 (Wang 294 et al., 2016), but higher than north China study (Cao et al., 2017) which emphasize the 295 solid atmospheric oxidation capacity in YRD region.

Table 2 Statistical result of NOR and SOR in YRD

1 1	SOR			NOR				D.C		
Location and Year	Max	Min	Mean	SD	Max	Min	Mean	SD	- References	
Nanjing 2013 Winter	0.42	0.10	0.28	0.11	0.29	0.15	0.21	0.05		
Suzhou 2013 Winter	0.41	0.15	0.27	0.11	0.30	0.06	0.16	0.08		
Lin'an 2013 Winter	0.50	0.19	0.35	0.11	0.24	0.12	0.18	0.05	(Wang et al., 2016)	
Hangzhou 2013 Winter	0.30	0.14	0.21	0.06	0.11	0.06	0.09	0.02		
Ningbo 2013 Winter	0.35	0.09	0.21	0.11	0.23	0.03	0.11	0.07		
YRD 2016 Summer	-	-	0.347	-	-	-	0.11	-	(Shu et al., 2019)	
YRD 2016 Winter	-	-	0.247	-	-	-	0.15	-		
Nanjing 2019 spring	0.48	0.38	-	-	0.31	0.29	-	-	(Qin et al., 2021)	
Changzhou 2019 spring	0.35	0.3	-	-	0.27	0.23	-	-		
Changzhou 2019 Winter	0.68	0.24	0.35	0.12	0.44	0.13	0.2	0.1	(Zhang et al., 2020b)	
Changzhou 2019 Summer	0.16	0.76	0.54	0.1	0.08	0.63	0.28	0.14	This work	

# 297 **3.3** The derivation of N<sub>2</sub>O<sub>5</sub> uptake coefficient

298 Statistical analysis of the observation above highlights the rapid formation of 299 particulate nitrate. To assess the contribution of N2O5 hydrolysis to particular nitrate 300 formation, two methods are applied to calculate the  $N_2O_5$  uptake coefficient. The first 301 method is a stationary-state approximation (Brown et al., 2003). By assuming that the 302 rates of production and loss of N2O5 are approximately in balance, the total loss rate of 303  $N_2O_5$  ( $k_{N_2O_5}$ ) can be calculated through equation 9. The  $k_{N_2O_5}$  is main dominated by 304 N<sub>2</sub>O<sub>5</sub> heterogeneous uptake, since homogeneous hydrolysis of N<sub>2</sub>O<sub>5</sub> contribute tiny (Brown and Stutz, 2012). N<sub>2</sub>O<sub>5</sub> uptake coefficient through steady-state (note as  $\gamma_{-S}$ ) is 305 306 derived as equation 10. Here C is the mean molecule speed of N<sub>2</sub>O<sub>5</sub>, and S<sub>a</sub> is the aerosol 307 surface concentration.

$$\pi_{ss}(N_2O_5) = \frac{[N_2O_5]}{k_{R3.1}[NO_2][O_3]} = (k_{N_2O_5} + \frac{k_{NO_3}}{K_{eq}[NO_2]})^{-1}$$
Eq9

$$k_{N_2O_5} = 0.25 \text{ C } \gamma_{S} \text{ S}_a$$
 Eq10

308 Due to the fast variety of  $NO_3$  loss rates from VOCs, the steady-state method has been 309 unattainable in conditions affected by emission interferences. During the whole 310 campaign, we only retrieve three valid fitting results. As shown in Figure 5, the fitted

- 311  $\gamma_{s}$  ranged from 0.057 to 0.123, which is comparable with Taizhou (0.041, Wang et al.
- 312 (2020a)) and much higher than other results in China (Yu et al., 2020a; Wang et al.,
- 313 2018a; Wang et al., 2020b; Wang et al., 2017a). The calculated  $k_{NO_3}$  ranged from 0.002
- 114 to 0.16 s<sup>-1</sup>, represents drastic VOCs change during this campaign.



**Figure 5** Derived N<sub>2</sub>O<sub>5</sub> uptake coefficients from N<sub>2</sub>O<sub>5</sub> steady lifetime ( $\gamma_{-S}$ ) with NO<sub>2</sub> and S<sub>a</sub>, plots (a-c) represent the linear fitting results on the nights of 05/30, 06/10, and 06/11, respectively.

The other approach is the parameterization by (Yu et al., 2020a) which is depicted as follows:

$$\gamma_{-P} = \frac{4}{c} \frac{V_a}{S_a} K_H \times 3.0 \times 10^4 \times [H_2O] \left( 1 - \frac{1}{\left( 0.033 \times \frac{[H_2O]}{[NO_3]} \right) + 1 + \left( 3.4 \times \frac{[CI^-]}{[NO_3]} \right)} \right)$$
Eq11

321

315

Where  $V_a/S_a$  is the measured aerosol volume to surface area ratio by SMPS; K<sub>H</sub> is 322 323 Henry's law coefficient which is set as 51 as recommended; [NO<sub>3</sub><sup>-</sup>] and [Cl<sup>-</sup>] are aerosol 324 inorganic concentration measured by Marga; [H<sub>2</sub>O] is aerosol water content calculated 325 through ISORROPIA II. The valid parameterization calculated N<sub>2</sub>O<sub>5</sub> uptake coefficient (note as  $\gamma_{-P})$  from May 30<sup>th</sup> to June 08<sup>th</sup>, 2019, shows in Figure 6 a good consistency 326 327 between the trends of  $\gamma_{P}$  and aerosol water content. Nighttime  $\gamma_{P}$  varies from 0.001 to 328 0.024 with an average of  $0.069 \pm 0.0050$  in polluted condition and  $0.0036 \pm 0.0026$  in 329 clean condition. The N<sub>2</sub>O<sub>5</sub> uptake coefficient shows a good correlation between RH and 330 aerosol water content. For the N<sub>2</sub>O<sub>5</sub> uptake coefficient, although particulate nitrate mass 331 concentration increased during the pollution event, an antagonistic effect on the N<sub>2</sub>O<sub>5</sub> 332 uptake coefficient was not obvious for the nitrate molarity decreasing.

Furthermore, we compare the difference between  $\gamma_{S}$  and  $\gamma_{P}^{h}$ . Taking the night of 333 May 30<sup>th</sup> as an example, the  $\gamma_{s}$  is 0.089 while  $\gamma_{P}$  ranges from 0.024 to 0.057 with an 334 average value of 0.013  $\pm$  0.0051. The difference between steady-state and 335 336 parameterization is significant; one possible explanation is uncertainty for stationary-337 state approximation caused by local NO or VOCs emission (Brown et al., 2009; Chen et al., 2022). Another reason is that parameterization by Yu et al. ignores the impact of 338 organic matter on the fine particle. The difference in aerosol composition between this 339 work and Yu et al may also bring uncertainty. Overall consideration,  $\gamma_{P}$  will be chosen 340 341 for the  $N_2O_5$  heterogeneous uptake coefficient in later analysis and discussion.



**Figure 6** Results of N<sub>2</sub>O<sub>5</sub> uptake coefficients through parameterization ( $\gamma_{-p}$ ). (a) shows timeseries of  $\gamma_{-p}$  and ISORROPIA II results of aerosol water content (AWC). (b) is the box-plot of  $\gamma_{-p}$  on the polluted day and clean day, the hollow square represents the mean value, and the solid line across the box shows the median score for the data set, while the top and bottom whiskers represent 90 % and 10 % value of  $\gamma_{-p}$ , respectively.

# 348 **3.4 Quantifying the contribution of nitrate formation pathways**

After the N<sub>2</sub>O<sub>5</sub> uptake coefficient is counted, nitrate production potential (P(NO<sub>3</sub><sup>-</sup>)) can be calculated. Here N<sub>2</sub>O<sub>5</sub> uptake coefficient is set as 0.036 on clean day and 0.069 on polluted day, respectively, which are the average value derived from parameterization. The production ratio of NO<sub>3</sub><sup>-</sup> (by considering ClNO<sub>2</sub> yield of 0.54) is set as 1.46 in the former study (Xia et al., 2020). Gas particle distribution is considered by the result of particular nitrate and gas-phase nitrate by MARGA (input HNO<sub>3</sub>/NO<sub>3</sub><sup>-</sup> ratio to the model as OH + NO<sub>2</sub> nitrate production rate). NO<sub>2</sub> heterogeneous uptake coefficient is set as  $5.8 \times 10^{-6}$  depending on the report by Yu et al. (2021) which is the result of 70% RH on urban grime.

The mean diurnal variations of the nitrate production potential of clean and polluted day are depicted in Figure 7. The OH + NO<sub>2</sub> pathway shows no significate difference between clean and polluted day and dominates clean day nitrate formation potential. Since the level of OH and NO<sub>2</sub> is less affected by the fine particle level. However, the rapid increase of the N<sub>2</sub>O<sub>5</sub> heterogeneous uptake pathway on polluted day is fatal, and its peak formation rate at night over the OH + NO<sub>2</sub> pathway can be used to explain nighttime nitrate explosive growth.

365 As shown in Figure 7c,  $OH + NO_2$  dominates nitrate production on clean day, while the N<sub>2</sub>O<sub>5</sub> uptake pathway only contributes 13.6  $\mu$ g m<sup>-3</sup>. On polluted days, the ability of 366 N<sub>2</sub>O<sub>5</sub> uptake grows fast, reaching 50.1 µg m-3, while the OH pathway doesn't change 367 368 much. There is no distinct difference in the daytime pathway  $(OH + NO_2)$  between clean 369 day and polluted day, while the nighttime pathway ratio rises from 38.1 % on clean day to 67.2 % on polluted day.NO<sub>2</sub> heterogeneous uptake increases from 0.93  $\mu$ g m<sup>-3</sup> on 370 clean day to 2.0 µg m<sup>-3</sup> on polluted day, but the contribution proportion does not change 371 372 obviously. Both the higher N<sub>2</sub>O<sub>5</sub> uptake coefficient and higher S<sub>a</sub> on polluted day 373 increase the contribution of N<sub>2</sub>O<sub>5</sub> hydrolysis on particular nitrate at pollution condition. 374



376 **Figure 7** The mean diurnal variations of the nitrate production potential of clean day(a) 377 and polluted day (b) and the  $P(NO_3^-)$  distribution of clean and polluted day (c).

# 378 **3.5** Mitigation strategies of particulate nitrate and ozone productions

We selected two pollution episodes (Episode I (2019.05.30 00:00 - 2019.06.02 00:00) 379 380 and IV (2019.06.14 17:30 - 2019.06.17 12:00)) to explore the mitigation way of ozone 381 and nitrate pollution. Figure 8 shows the EKMA of  $P(O_3)$  and  $P(NO_3)$  of these two 382 periods, O<sub>3</sub> located at VOCs controlling area in the two pollution episodes, which 383 consist with previous YRD urban ozone sensitivity study (Jiang et al., 2018; Zhang et 384 al., 2020a; Xu et al., 2021). The best precursor reduction for  $O_3$  is VOCs:  $NO_x = 2:1$ while nitrate is located at the transition area, which means either of the precursors 385 386 reduction will mitigate nitrate pollution. For the regional and complex air pollution 387 characteristics in this region, a fine particle-targeting reduction scheme will aggravate 388  $O_3$  pollution. In contrast, the  $O_3$ -targeting scheme can mitigate  $O_3$  and fine particle 389 simultaneously.



390

391 Figure 8 Isogram of  $P(O_3)$  and  $P(NO_3^-)$  of polluted episode I (2019.05.30 00:00 -

392 2019.06.02 00:00) and IV (2019.06.14 17:30 - 2019.06.17 12:00) with different NOx



# 394 4 Conclusion

395 A comprehensive campaign was conducted to interpret the atmospheric oxidation capacity and aerosol formation from May 30<sup>th</sup> to June 18<sup>th</sup>, 2019, in Changzhou, China. 396 397 The high O<sub>3</sub> and PM<sub>2.5</sub> concentrations confirm complex air pollution characteristics in 398 Changzhou, and nitrate accounts for 38.3 % mass concentration of total water-soluble 399 particulate components and 32.0 % of total PM<sub>2.5</sub>. In addition, the average values of 400 NOR are 0.32 in PD and 0.25 in CD. The positive correlation between NOR and RH 401 and inverse correlation refer to the contribution of N2O5 heterogeneous uptake to nitrate 402 formation.

Based on field observations of OH and related parameters, we show OH oxidation of the NO<sub>2</sub> pathway steadily contributes to nitrate formation no matter the clean or polluted period and domination clean day nitrate production (about 22  $\mu$ g m<sup>-3</sup>). N<sub>2</sub>O<sub>5</sub> heterogeneous uptake contribution proliferated on polluted day, from 13.6  $\mu$ g m<sup>-3</sup> (38.1 %) on clean days to 50.1  $\mu$ g m<sup>-3</sup> (67.2 %) on polluted days. NO<sub>2</sub> heterogeneous uptake contributes minor to nitrate formation (2.6 %).

The precursor reduction simulation suggests the reduction ratio of VOCs:  $NO_x$ equals 2:1 can simultaneously and effectively mitigate  $O_3$  and fine particle pollution during the summertime complex pollution period in Changzhou. To more precisely and delicately establish a cooperative control scheme for regional  $O_3$  and nitrate, the regional and long-time field campaigns are needed in the future to analyze the seasonal and interannual variation of  $O_3$  and nitrate and relevant parameters.

415

416 **Code/Data availability.** The datasets used in this study are available from the 417 corresponding author upon request (k.lu@pku.edu.cn).

418

419 Author contributions. K.D.L. and Y.H.Z. designed the study. T.Y.Z analyzed the data420 and wrote the paper with input from all authors.

422 **Competing interests**. The authors declare that they have no conflicts of interest.

423

421

424 Acknowledgments. This project is supported by the National Natural Science 425 Foundation of China (21976006); the Beijing Municipal Natural Science Foundation 426 for Distinguished Young Scholars (JQ19031); the National Research Program for Key 427 Issue in Air Pollution Control (DQGG0103-01, 2019YFC0214800). Thanks for the data 428 contributed by field campaign team.

#### 429 References

- 430 Andreae, M. O., Schmid, O., Yang, H., Chand, D., Yu, J. Z., Zeng, L.-M., and Zhang, 431 Y.-H.: Optical properties and chemical composition of the atmospheric aerosol in urban Guangzhou, China, Atmospheric Environment, 42, 6335-6350, 432 10.1016/j.atmosenv.2008.01.030, 2008. 433
- 434 Bertram, T. H. and Thornton, J. A.: Toward a general parameterization of N2O5 435 reactivity on aqueous particles: the competing effects of particle liquid water, 436 nitrate and chloride, Atmospheric Chemistry and Physics, 9, 8351-8363, 437

10.5194/acp-9-8351-2009, 2009.

- 438 Bohn, B., Corlett, G. K., Gillmann, M., Sanghavi, S., Stange, G., Tensing, E., 439 Vrekoussis, M., Bloss, W. J., Clapp, L. J., Kortner, M., Dorn, H. P., Monks, P.
- 440 S., Platt, U., Plass-Dülmer, C., Mihalopoulos, N., Heard, D. E., Clemitshaw, K. 441 C., Meixner, F. X., Prevot, A. S. H., and Schmitt, R.: Photolysis frequency
- 442 measurement techniques: results of a comparison within the ACCENT project, 443 ACP, 8, 5373-5391, 10.5194/acp-8-5373-2008, 2008.
- Brown, S. S. and Stutz, J.: Nighttime radical observations and chemistry, Chemical 444 445 Society Reviews, 41, 6405-6447, 10.1039/c2cs35181a, 2012.
- 446 Brown, S. S., Stark, H., and Ravishankara, A. R.: Applicability of the steady state 447 approximation to the interpretation of atmospheric observations of NO3 and N2O5, Journal of Geophysical Research-Atmospheres, 108, 10, 448
- 449 10.1029/2003jd003407, 2003.
- 450 Brown, S. S., Dube, W. P., Fuchs, H., Ryerson, T. B., Wollny, A. G., Brock, C. A., 451 Bahreini, R., Middlebrook, A. M., Neuman, J. A., Atlas, E., Roberts, J. M.,
- 452 Osthoff, H. D., Trainer, M., Fehsenfeld, F. C., and Ravishankara, A. R.: Reactive
- uptake coefficients for N2O5 determined from aircraft measurements during the 453 454 Second Texas Air Quality Study: Comparison to current model
- 455 parameterizations, Journal of Geophysical Research-Atmospheres, 114, 456 10.1029/2008jd011679, 2009.
- 457 Burkholder, J. B., Sander, S. P., Abbatt, J., Barker, J., Huie, R., Kolb, C. E., Kurylo, M., Orkin, V., Wilmouth, D. M., and Wine, P.: Chemical Kinetics and 458
- 459 Photochemical Data for Use in Atmospheric Studies, Evaluation Number 18, 460 10.13140/RG.2.1.2504.2806, 2015.
- 461 Cao, J.-J., Shen, Z.-X., Chow, J. C., Watson, J. G., Lee, S.-C., Tie, X.-X., Ho, K.-F., Wang, G.-H., and Han, Y.-M.: Winter and Summer PM2.5 Chemical 462

Compositions in Fourteen Chinese Cities, J Air Waste Manage, 62, 1214-1226, 463 10.1080/10962247.2012.701193, 2012. 464 465 Cao, Z., Zhou, X., Ma, Y., Wang, L., Wu, R., Chen, B., and Wang, W.: The Concentrations, Formations, Relationships and Modeling of Sulfate, Nitrate and 466 Ammonium (SNA) Aerosols over China, Aerosol and Air Quality Research, 17, 467 468 84-97, 10.4209/aagr.2016.01.0020, 2017. 469 Chan, Y. C., Evans, M. J., He, P. Z., Holmes, C. D., Jaegle, L., Kasibhatla, P., Liu, X. 470 Y., Sherwen, T., Thornton, J. A., Wang, X., Xie, Z. Q., Zhai, S. T., and 471 Alexander, B.: Heterogeneous Nitrate Production Mechanisms in Intense Haze 472 Events in the North China Plain, Journal of Geophysical Research-Atmospheres, 473 126, 10.1029/2021jd034688, 2021. 474 Chang, Y. H., Zhang, Y. L., Tian, C. G., Zhang, S. C., Ma, X. Y., Cao, F., Liu, X. Y., 475 Zhang, W. Q., Kuhn, T., and Lehmann, M. F.: Nitrogen isotope fractionation during gas-to-particle conversion of NOx to NO3- in the atmosphere -476 477 implications for isotope-based NOx source apportionment, Atmospheric 478 Chemistry and Physics, 18, 11647-11661, 10.5194/acp-18-11647-2018, 2018. 479 Chen, H., Hu, R., Xie, P., Xing, X., Ling, L., Li, Z., Wang, F., Wang, Y., Liu, J., and 480 Liu, W.: A hydroxyl radical detection system using gas expansion and fast gating 481 laser-induced fluorescence techniques, J. Environ. Sci., 65, 190-200, 482 10.1016/j.jes.2017.03.012, 2018. 483 Chen, X., Walker, J. T., and Geron, C.: Chromatography related performance of the 484 Monitor for AeRosols and GAses in ambient air (MARGA): laboratory and 485 field-based evaluation, Atmos. Meas. Tech., 10, 3893-3908, 10.5194/amt-10-486 3893-2017, 2017. 487 Chen, X., Wang, H., and Lu, K.: Interpretation of NO3-N2O5 observation via steady 488 state in high-aerosol air mass: the impact of equilibrium coefficient in ambient 489 conditions, Atmospheric Chemistry and Physics, 22, 3525-3533, 10.5194/acp-490 22-3525-2022, 2022. 491 Chen, X. R., Wang, H. C., Liu, Y. H., Su, R., Wang, H. L., Lou, S. R., and Lu, K. D.: Spatial characteristics of the nighttime oxidation capacity in the Yangtze River 492 493 Delta, China, Atmospheric Environment, 208, 150-157, 494 10.1016/i.atmosenv.2019.04.012, 2019. 495 Chen, X. R., Wang, H. C., Lu, K. D., Li, C. M., Zhai, T. Y., Tan, Z. F., Ma, X. F., 496 Yang, X. P., Liu, Y. H., Chen, S. Y., Dong, H. B., Li, X., Wu, Z. J., Hu, M., 497 Zeng, L. M., and Zhang, Y. H.: Field Determination of Nitrate Formation 498 Pathway in Winter Beijing, Environmental Science & Technology, 54, 9243-499 9253, 10.1021/acs.est.0c00972, 2020. Elshorbany, Y. F., Steil, B., Brühl, C., and Lelieveld, J.: Impact of HONO on global 500 501 atmospheric chemistry calculated with an empirical parameterization in the EMAC model, ACP, 12, 9977-10000, 10.5194/acp-12-9977-2012, 2012. 502 503 Fountoukis, C. and Nenes, A.: ISORROPIA II: a computationally efficient 504 thermodynamic equilibrium model for K+-Ca2+-Mg2+-Nh(4)(+)-Na+-SO42--

505	NO3ClH2O aerosols, Atmospheric Chemistry and Physics, 7, 4639-4659,
506	2007.
507	Goliff, W. S., Stockwell, W. R., and Lawson, C. V.: The regional atmospheric
508	chemistry mechanism, version 2, Atmospheric Environment, 68, 174-185,
509	10.1016/j.atmosenv.2012.11.038, 2013.
510	Guo, L., Hu, Y., Hu, Q., Lin, J., Li, C., Chen, J., Li, L., and Fu, H.: Characteristics
511	and chemical compositions of particulate matter collected at the selected metro
512	stations of Shanghai, China, Science of the Total Environment, 496, 443-452,
513	10.1016/j.scitotenv.2014.07.055, 2014.
514	Hagler, G. S. W., Bergin, M. H., Salmon, L. G., Yu, J. Z., Wan, E. C. H., Zheng, M.,
515	Zeng, L. M., Kiang, C. S., Zhang, Y. H., Lau, A. K. H., and Schauer, J. J.:
516	Source areas and chemical composition of fine particulate matter in the Pearl
517	River Delta region of China, Atmos. Environ., 40, 3802-3815,
518	10.1016/j.atmosenv.2006.02.032, 2006.
519	Jiang, M., Lu, K., Su, R., Tan, Z., Wang, H., Li, L., Fu, Q., Zhai, C., Tan, Q., Yue, D.,
520	Chen, D., Wang, Z., Xie, S., Zeng, L., and Zhang, Y.: Ozone formation and key
521	VOCs in typical Chinese city clusters, Chinese Sci Bull, 63, 1130-1141, 2018.
522	Kanaya, Y., Fukuda, M., Akimoto, H., Takegawa, N., Komazaki, Y., Yokouchi, Y.,
523	Koike, M., and Kondo, Y.: Urban photochemistry in central Tokyo: 2. Rates and
524	regimes of oxidant (O-3+NO2) production, Journal of Geophysical Research-
525	Atmospheres, 113, 10.1029/2007jd008671, 2008.
526	Li, H. Y., Zhang, Q., Zhang, Q., Chen, C. R., Wang, L. T., Wei, Z., Zhou, S.,
527	Parworth, C., Zheng, B., Canonaco, F., Prevot, A. S. H., Chen, P., Zhang, H. L.,
528	Wallington, T. J., and He, K. B.: Wintertime aerosol chemistry and haze
529	evolution in an extremely polluted city of the North China Plain: significant
530	contribution from coal and biomass combustion, Atmospheric Chemistry and
531	Physics, 17, 4751-4768, 2017.
532	Liu, X., Gu, J., Li, Y., Cheng, Y., Qu, Y., Han, T., Wang, J., Tian, H., Chen, J., and
533	Zhang, Y.: Increase of aerosol scattering by hygroscopic growth: Observation,
534	modeling, and implications on visibility, Atmos. Res., 132-133, 91-101,
535	https://doi.org/10.1016/j.atmosres.2013.04.007, 2013.
536	Lou, S., Tan, Z., Gan, G., Chen, J., Wang, H., Gao, Y., Huang, D., Huang, C., Li, X.,
537	Song, R., Wang, H., Wang, M., Wang, Q., Wu, Y., and Huang, C.: Observation
538	based study on atmospheric oxidation capacity in Shanghai during late-autumn:
539	Contribution from nitryl chloride, Atmospheric Environment, 271, 118902,
540	https://doi.org/10.1016/j.atmosenv.2021.118902, 2022.
541	Lu, K. D., Hofzumahaus, A., Holland, F., Bohn, B., Brauers, T., Fuchs, H., Hu, M.,
542	Haseler, R., Kita, K., Kondo, Y., Li, X., Lou, S. R., Oebel, A., Shao, M., Zeng,
543	L. M., Wahner, A., Zhu, T., Zhang, Y. H., and Rohrer, F.: Missing OH source in
544	a suburban environment near Beijing: observed and modelled OH and HO2
545	concentrations in summer 2006, Atmospheric Chemistry and Physics, 13, 1057-
546	1080, 10.5194/acp-13-1057-2013, 2013.

547	Lu, K. D., Rohrer, F., Holland, F., Fuchs, H., Bohn, B., Brauers, T., Chang, C. C.,
548	Haseler, R., Hu, M., Kita, K., Kondo, Y., Li, X., Lou, S. R., Nehr, S., Shao, M.,
549	Zeng, L. M., Wahner, A., Zhang, Y. H., and Hofzumahaus, A.: Observation and
550	modelling of OH and HO2 concentrations in the Pearl River Delta 2006: a
551	missing OH source in a VOC rich atmosphere, Atmospheric Chemistry and
552	Physics, 12, 1541-1569, 10.5194/acp-12-1541-2012, 2012.
553	Ma, X. F., Tan, Z. F., Lu, K. D., Yang, X. P., Chen, X. R., Wang, H. C., Chen, S. Y.,
554	Fang, X., Li, S. L., Li, X., Liu, J. W., Liu, Y., Lou, S. R., Qiu, W. Y., Wang, H.
555	L., Zeng, L. M., and Zhang, Y. H.: OH and HO2 radical chemistry at a suburban
556	site during the EXPLORE-YRD campaign in 2018, Atmospheric Chemistry and
557	Physics, 22, 7005-7028, 2022.
558	Meng, Z. Y., Wu, L. Y., Xu, X. D., Xu, W. Y., Zhang, R. J., Jia, X. F., Liang, L. L.,
559	Miao, Y. C., Cheng, H. B., Xie, Y. L., He, J. J., and Zhong, J. T.: Changes in
560	ammonia and its effects on PM2.5 chemical property in three winter seasons in
561	Beijing, China, Science of the Total Environment, 749, 2020.
562	Ming, L., Jin, L., Li, J., Fu, P., Yang, W., Liu, D., Zhang, G., Wang, Z., and Li, X.:
563	PM2.5 in the Yangtze River Delta, China: Chemical compositions, seasonal
564	variations, and regional pollution events, Environmental Pollution, 223, 200-212,
565	10.1016/j.envpol.2017.01.013, 2017.
566	Phillips, G. J., Thieser, J., Tang, M. J., Sobanski, N., Schuster, G., Fachinger, J.,
567	Drewnick, F., Borrmann, S., Bingemer, H., Lelieveld, J., and Crowley, J. N.:
568	Estimating N2O5 uptake coefficients using ambient measurements of NO3,
569	N2O5, ClNO2 and particle-phase nitrate, Atmospheric Chemistry and Physics,
570	16, 13231-13249, 10.5194/acp-16-13231-2016, 2016.
571	Qin, Y., Li, J. Y., Gong, K. J., Wu, Z. J., Chen, M. D., Qin, M. M., Huang, L., and
572	Hu, J. L.: Double high pollution events in the Yangtze River Delta from 2015 to
573	2019: Characteristics, trends, and meteorological situations, Science of the Total
574	Environment, 792, 10.1016/j.scitotenv.2021.148349, 2021.
575	Qiu, X. H., Ying, Q., Wang, S. X., Duan, L., Zhao, J., Xing, J., Ding, D., Sun, Y. L.,
576	Liu, B. X., Shi, A. J., Yan, X., Xu, Q. C., and Hao, J. M.: Modeling the impact of
577	heterogeneous reactions of chlorine on summertime nitrate formation in Beijing,
578	China, Atmospheric Chemistry and Physics, 19, 6737-6747, 10.5194/acp-19-
579	6737-2019, 2019.
580	Seinfeld, J. H. and Pandis, S. N.: Atmospheric chemistry and physics: from air
581	pollution to climate change, Third;3rd;, Book, Whole, Wiley, Hoboken, New
582	Jersey2016.
583	Shang, D. J., Peng, J. F., Guo, S., Wu, Z. J., and Hu, M.: Secondary aerosol formation
584	in winter haze over the Beijing-Tianjin-Hebei Region, China, Front. Env. Sci.
585	Eng., 15, 13, 10.1007/s11783-020-1326-x, 2021.
586	Shu, L., Wang, T. J., Xie, M., Li, M. M., Zhao, M., Zhang, M., and Zhao, X. Y.:
587	Episode study of fine particle and ozone during the CAPUM-YRD over Yangtze
588	River Delta of China: Characteristics and source attribution, Atmospheric

589	Environment, 203, 87-101, 10.1016/j.atmosenv.2019.01.044, 2019.
590	Song, C. H. and Carmichael, G. R.: Gas-particle partitioning of nitric acid modulated
591	by alkaline aerosol, Journal of Atmospheric Chemistry, 40, 1-22, 2001.
592	Staudt, S., Gord, J. R., Karimova, N. V., McDuffie, E. E., Brown, S. S., Gerber, R. B.,
593	Nathanson, G. M., and Bertram, T. H.: Sulfate and Carboxylate Suppress the
594	Formation of ClNO2 at Atmospheric Interfaces, Acs Earth and Space Chemistry,
595	3, 1987-1997, 2019.
596	Sun, Y. L., Zhuang, G. S., Tang, A. H., Wang, Y., and An, Z. S.: Chemical
597	characteristics of PM2.5 and PM10 in haze-fog episodes in Beijing,
598	Environmental Science & Technology, 40, 3148-3155, 10.1021/es051533g,
599	2006.
600	Tan, Z., Wang, H., Lu, K., Dong, H., Liu, Y., Zeng, L., Hu, M., and Zhang, Y.: An
601	Observational Based Modeling of the Surface Layer Particulate Nitrate in the
602	North China Plain During Summertime, Journal of Geophysical Research:
603	Atmospheres, 126, e2021JD035623, https://doi.org/10.1029/2021JD035623,
604	2021.
605	Tan, Z. F., Lu, K. D., Dong, H. B., Hu, M., Li, X., Liu, Y. H., Lu, S. H., Shao, M., Su,
606	R., Wang, H. C., Wu, Y. S., Wahner, A., and Zhang, Y. H.: Explicit diagnosis of
607	the local ozone production rate and the ozone-NOx-VOC sensitivities, Science
608	Bulletin, 63, 1067-1076, 10.1016/j.scib.2018.07.001, 2018.
609	Tan, Z. F., Fuchs, H., Lu, K. D., Hofzumahaus, A., Bohn, B., Broch, S., Dong, H. B.,
610	Gomm, S., Haseler, R., He, L. Y., Holland, F., Li, X., Liu, Y., Lu, S. H., Rohrer,
611	F., Shao, M., Wang, B. L., Wang, M., Wu, Y. S., Zeng, L. M., Zhang, Y. S.,
612	Wahner, A., and Zhang, Y. H.: Radical chemistry at a rural site (Wangdu) in the
613	North China Plain: observation and model calculations of OH, HO2 and RO2
614	radicals, Atmospheric Chemistry and Physics, 17, 663-690, 10.5194/acp-17-663-
615	2017, 2017.
616	Tham, Y. J., Wang, Z., Li, Q., Wang, W., Wang, X., Lu, K., Ma, N., Yan, C.,
617	Kecorius, S., Wiedensohler, A., Zhang, Y., and Wang, T.: Heterogeneous N2O5
618	uptake coefficient and production yield of ClNO2 in polluted northern China:
619	roles of aerosol water content and chemical composition, Atmospheric
620	Chemistry and Physics, 18, 13155-13171, 10.5194/acp-18-13155-2018, 2018.
621	Wang, H., Zhu, B., Shen, L., Xu, H., An, J., Pan, C., Li, Y. e., and Liu, D.: Regional
622	Characteristics of Air Pollutants during Heavy Haze Events in the Yangtze River
623	Delta, China, Aerosol and Air Quality Research, 16, 2159-2171,
624	10.4209/aaqr.2015.09.0551, 2016.
625	Wang, H., Lu, K., Chen, X., Zhu, Q., Chen, Q., Guo, S., Jiang, M., Li, X., Shang, D.,
626	Tan, Z., Wu, Y., Wu, Z., Zou, Q., Zheng, Y., Zeng, L., Zhu, T., Hu, M., and
627	Zhang, Y.: High N2O5 Concentrations Observed in Urban Beijing: Implications
628	of a Large Nitrate Formation Pathway, Environmental Science and Technology
629	Letters, 4, 416-420, 10.1021/acs.estlett.7b00341, 2017a.
630	Wang, H. C. and Lu, K. D.: Determination and Parameterization of the Heterogeneous

631	Uptake Coefficient of Dinitrogen Pentoxide (N2O5), Prog. Chem., 28, 917-933,
632	10.7536/pc151225, 2016.
633	Wang, H. C., Chen, J., and Lu, K. D.: Development of a portable cavity-enhanced
634	absorption spectrometer for the measurement of ambient NO3 and N2O5:
635	experimental setup, lab characterizations, and field applications in a polluted
636	urban environment, Atmos. Meas. Tech., 10, 1465-1479, 10.5194/amt-10-1465-
637	2017, 2017b.
638	Wang, H. C., Lu, K. D., Chen, X. R., Zhu, Q. D., Wu, Z. J., Wu, Y. S., and Sun, K.:
639	Fast particulate nitrate formation via N2O5 uptake aloft in winter in Beijing,
640	Atmospheric Chemistry and Physics, 18, 10483-10495, 10.5194/acp-18-10483-
641	2018, 2018a.
642	Wang, H. C., Chen, X. R., Lu, K. D., Hu, R. Z., Li, Z. Y., Wang, H. L., Ma, X. F.,
643	Yang, X. P., Chen, S. Y., Dong, H. B., Liu, Y., Fang, X., Zeng, L. M., Hu, M.,
644	and Zhang, Y. H.: NO3 and N2O5 chemistry at a suburban site during the
645	EXPLORE-YRD campaign in 2018, Atmospheric Environment, 224, 9,
646	10.1016/j.atmosenv.2019.117180, 2020a.
647	Wang, H. C., Chen, X. R., Lu, K. D., Tan, Z. F., Ma, X. F., Wu, Z. J., Li, X., Liu, Y.
648	H., Shang, D. J., Wu, Y. S., Zeng, L. M., Hu, M., Schmitt, S., Kiendler-Scharr,
649	A., Wahner, A., and Zhang, Y. H.: Wintertime N2O5 uptake coefficients over
650	the North China Plain, Science Bulletin, 65, 765-774,
651	10.1016/j.scib.2020.02.006, 2020b.
652	Wang, H. C., Lu, K. D., Guo, S., Wu, Z. J., Shang, D. J., Tan, Z. F., Wang, Y. J., Le
653	Breton, M., Lou, S. R., Tang, M. J., Wu, Y. S., Zhu, W. F., Zheng, J., Zeng, L.
654	M., Hallquist, M., Hu, M., and Zhang, Y. H.: Efficient N2O5 uptake and NO3
655	oxidation in the outflow of urban Beijing, Atmospheric Chemistry and Physics,
656	18, 9705-9721, 10.5194/acp-18-9705-2018, 2018b.
657	Wang, S. B., Wang, L. L., Fan, X. G., Wang, N., Ma, S. L., and Zhang, R. Q.:
658	Formation pathway of secondary inorganic aerosol and its influencing factors in
659	Northern China: Comparison between urban and rural sites, Science of the Total
660	Environment, 840, 2022.
661	Wang, X. F., Zhang, Y. P., Chen, H., Yang, X., Chen, J. M., and Geng, F. H.:
662	Particulate Nitrate Formation in a Highly Polluted Urban Area: A Case Study by
663	Single-Particle Mass Spectrometry in Shanghai, Environmental Science &
664	Technology, 43, 3061-3066, 2009.
665	Wang, Z., Wang, W. H., Tham, Y. J., Li, Q. Y., Wang, H., Wen, L., Wang, X. F., and
666	Wang, T.: Fast heterogeneous N2O5 uptake and ClNO2 production in power
667	plant and industrial plumes observed in the nocturnal residual layer over the
668	North China Plain, Atmospheric Chemistry and Physics, 17, 12361-12378,
669	10.5194/acp-17-12361-2017, 2017c.
670	Woodward-Massey, R., Slater, E. J., Alen, J., Ingham, T., Cryer, D. R., Stimpson, L.
671	M., Ye, C. X., Seakins, P. W., Whalley, L. K., and Heard, D. E.: Implementation
672	of a chemical background method for atmospheric OH measurements by laser-

673	induced fluorescence: characterisation and observations from the UK and China,
674	Atmos. Meas. Tech., 13, 3119-3146, 10.5194/amt-13-3119-2020, 2020.
675	Wu, S. P., Dai, L. H., Zhu, H., Zhang, N., Yan, J. P., Schwab, J. J., and Yuan, C. S.:
676	The impact of sea-salt aerosols on particulate inorganic nitrogen deposition in
677	the western Taiwan Strait region, China, Atmos. Res., 228, 68-76, 2019.
678	Xia, M., Peng, X., Wang, W., Yu, C., Sun, P., Li, Y., Liu, Y., Xu, Z., Wang, Z., Xu,
679	Z., Nie, W., Ding, A., and Wang, T.: Significant production of ClNO2 and
680	possible source of Cl-2 from N2O5 uptake at a suburban site in eastern China,
681	Atmospheric Chemistry and Physics, 20, 6147-6158, 10.5194/acp-20-6147-2020,
682	2020.
683	Xu, J. W., Huang, X., Wang, N., Li, Y. Y., and Ding, A. J.: Understanding ozone
684	pollution in the Yangtze River Delta of eastern China from the perspective of
685	diurnal cycles, Science of the Total Environment, 752,
686	10.1016/j.scitotenv.2020.141928, 2021.
687	Xue, H., Liu, G., Zhang, H., Hu, R., and Wang, X.: Similarities and differences in
688	PM10 and PM2.5 concentrations, chemical compositions and sources in Hefei
689	City, China, Chemosphere, 220, 760-765, 10.1016/j.chemosphere.2018.12.123,
690	2019.
691	Yang, X. P., Lu, K. D., Ma, X. F., Liu, Y. H., Wang, H. C., Hu, R. Z., Li, X., Lou, S.
692	R., Chen, S. Y., Dong, H. B., Wang, F. Y., Wang, Y. H., Zhang, G. X., Li, S. L.,
693	Yang, S. D., Yang, Y. M., Kuang, C. L., Tan, Z. F., Chen, X. R., Qiu, P. P.,
694	Zeng, L. M., Xie, P. H., and Zhang, Y. H.: Observations and modeling of OH
695	and HO2 radicals in Chengdu, China in summer 2019, Science of the Total
696	Environment, 772, 2021.
697	Yu, C., Wang, Z., Xia, M., Fu, X., Wang, W. H., Tham, Y. J., Chen, T. S., Zheng, P.
698	G., Li, H. Y., Shan, Y., Wang, X. F., Xue, L. K., Zhou, Y., Yue, D. L., Ou, Y.
699	B., Gao, J., Lu, K. D., Brown, S. S., Zhang, Y. H., and Wang, T.: Heterogeneous
700	N2O5 reactions on atmospheric aerosols at four Chinese sites: improving model
701	representation of uptake parameters, Atmospheric Chemistry and Physics, 20,
702	4367-4378, 10.5194/acp-20-4367-2020, 2020a.
703	Yu, C. A., Wang, Z., Ma, Q. X., Xue, L. K., George, C., and Wang, T.: Measurement
704	of heterogeneous uptake of NO2 on inorganic particles, sea water and urban
705	grime, J. Environ. Sci., 106, 124-135, 10.1016/j.jes.2021.01.018, 2021.
706	Yu, D., Tan, Z., Lu, K., Ma, X., Li, X., Chen, S., Zhu, B., Lin, L., Li, Y., Qiu, P.,
707	Yang, X., Liu, Y., Wang, H., He, L., Huang, X., and Zhang, Y.: An explicit
708	study of local ozone budget and NOx-VOCs sensitivity in Shenzhen China,
709	Atmospheric Environment, 224, 117304, 10.1016/j.atmosenv.2020.117304,
710	2020b.
711	Zhang, K., Xu, J. L., Huang, Q., Zhou, L., Fu, Q. Y., Duan, Y. S., and Xiu, G. L.:
712	Precursors and potential sources of ground-level ozone in suburban Shanghai,
713	Front. Env. Sci. Eng., 14, 10.1007/s11783-020-1271-8, 2020a.
714	Zhang, R., Han, Y. H., Shi, A. J., Sun, X. S., Yan, X., Huang, Y. H., and Wang, Y.:

715	Characteristics of ambient ammonia and its effects on particulate ammonium in
716	winter of urban Beijing, China, Environ Sci Pollut R, 28, 62828-62838, 2021.
717	Zhang, Y., Hong, Z., Chen, J., Xu, L., Hong, Y., Li, M., Hao, H., Chen, Y., Qiu, Y.,
718	Wu, X., Li, JR., Tong, L., and Xiao, H.: Impact of control measures and
719	typhoon weather on characteristics and formation of PM2.5 during the 2016 G20
720	summit in China, Atmospheric Environment, 224, 117312,
721	https://doi.org/10.1016/j.atmosenv.2020.117312, 2020b.
722	Zhang, Y., Tang, L., Yu, H., Wang, Z., Sun, Y., Qin, W., Chen, W., Chen, C., Ding,
723	A., Wu, J., Ge, S., Chen, C., and Zhou, Hc.: Chemical composition, sources and
724	evolution processes of aerosol at an urban site in Yangtze River Delta, China
725	during wintertime, Atmospheric Environment, 123, 339-349,
726	10.1016/j.atmosenv.2015.08.017, 2015.
727	Zhang, Y., Tang, L., Croteau, P. L., Favez, O., Sun, Y., Canagaratna, M. R., Wang,
728	Z., Couvidat, F., Albinet, A., Zhang, H., Sciare, J., Prevot, A. S. H., Jayne, J. T.,
729	and Worsnop, D. R.: Field characterization of the PM2.5 Aerosol Chemical
730	Speciation Monitor: insights into the composition, sources, and processes of fine
731	particles in eastern China, Atmospheric Chemistry and Physics, 17, 14501-
732	14517, 10.5194/acp-17-14501-2017, 2017.
733	Zhang, Y. Y., Tang, A. H., Wang, C., Ma, X., Li, Y. Z., Xu, W., Xia, X. P., Zheng, A.
734	H., Li, W. Q., Fang, Z. G., Zhao, X. F., Peng, X. L., Zhang, Y. P., Han, J.,
735	Zhang, L. J., Collett, J. L., and Liu, X. J.: PM (2.5) and water-soluble inorganic
736	ion concentrations decreased faster in urban than rural areas in China, J. Environ.
737	Sci., 122, 83-91, 2022.
738	Zhao, P. S., Dong, F., He, D., Zhao, X. J., Zhang, X. L., Zhang, W. Z., Yao, Q., and
739	Liu, H. Y.: Characteristics of concentrations and chemical compositions for
740	PM2.5 in the region of Beijing, Tianjin, and Hebei, China, Atmospheric
741	Chemistry and Physics, 13, 4631-4644, 10.5194/acp-13-4631-2013, 2013.
742	Zhao, Z. Z., Sun, N., Zhou, W. L., Ma, S. S., Li, X. D., Li, M. L., Zhang, X., Tang, S.
743	S., and Ye, Z. L.: Chemical Compositions in Winter PM2.5 in Changzhou of the
744	Yangtze River Delta Region, China: Characteristics and Atmospheric Responses
745	Along With the Different Pollution Levels, Front Env Sci-Switz, 10, 2022.
746	Zheng, G. J., Duan, F. K., Su, H., Ma, Y. L., Cheng, Y., Zheng, B., Zhang, Q., Huang,
747	T., Kimoto, T., Chang, D., Poschl, U., Cheng, Y. F., and He, K. B.: Exploring
748	the severe winter haze in Beijing: the impact of synoptic weather, regional
749	transport and heterogeneous reactions, Atmospheric Chemistry and Physics, 15,
750	2969-2983, 2015.