

1 **The density of ambient black carbon retrieved by a new method:**
2 **implications to CCN prediction**

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25 **Abstract.**

26 The effective density of black carbon (BC) is a crucial factor relevant to its aging
27 degree that would add uncertainty in evaluating its climate effect. Here, we have
28 developed a new method to retrieve the effective density of internally mixed BC in the
29 atmosphere combining field observations conducted from 15 November to 14
30 December 2016 in urban Beijing with the Köhler theory. The uncertainty of the retrieval
31 method was evaluated within $\pm 30\%$, which was primarily caused by assumptions on
32 both the hygroscopic parameter of organics and the proportional distribution of primary
33 organic aerosols in different hygroscopic modes. Using the method, we obtain that the
34 ambient internally mixed BC, accounting for $80\pm 20\%$ of total BC aerosol particles,
35 was retrieved with a campaign mean density of $1.1\pm 0.6\text{ g cm}^{-3}$ during the observed
36 periods. The retrieved result was comparable with that reported in the literature. By
37 applying a lower (0.14 g cm^{-3}) and upper (2.1 g cm^{-3}) limit of the retrieved BC density
38 in cloud condensation nuclei (CCN) number concentration (N_{CCN}) estimation, we
39 derived that neglect of such variations in BC density would lead to an uncertainty of -
40 $28\% \sim 11\%$ in predicting N_{CCN} at supersaturations of 0.23% and 0.40% . We also find
41 that the N_{CCN} was more sensitive to the variations of BC density when it was $< 1.0\text{ g cm}^{-3}$.
42 This illustrates a necessity of accounting for the effect of BC density on CCN activity
43 closer to source regions where the BC particles are mostly freshly emitted. The CCN
44 closure achieved when introducing the retrieved real-time BC density and mixing state.
45 This study provides a unique way of utilizing field measurements to infer ambient BC

46 density and highlights the importance of applying variable BC density values in models
47 when predicting CCN and assessing its relevant climate effect.

48 **1 Introduction**

49 Black carbon (BC) aerosols, as the major absorber of solar radiation, play a vital
50 role in energy budget and climate of the earth-atmosphere system by affecting the
51 radiative forcing and cloud properties (Flanner et al., 2007; Ramanathan and
52 Carmichael, 2008). The light-absorbing capability induced by BC is related to its
53 density and morphology (Zhang et al., 2008; Rissler et al., 2014), which can be
54 modified after mixing with other atmospheric aerosol particles (Khalizov et al., 2009;
55 Xue et al., 2009). Changes in its physicochemical properties or the aging process would
56 also regulate its ability to serve as cloud condensation nuclei (CCN) and further affect
57 the CCN number concentrations (Zhang et al., 2016a, 2017; Ren et al., 2023) and the
58 radiative balance by affecting the clouds process (Yuan et al., 2008; Wang et al., 2011).
59 Owing to the complex evolution of the mixing state, density and morphology of BC,
60 the contribution of BC particles to CCN budgets is still not well understood.

61 BC particles, with diesel vehicles, industrial and residential coal combustion as
62 major sources, are ubiquitous in urban environments (Bond et al., 2013; Dameto et al.,
63 2017; Li et al., 2017; Liu et al., 2019a). The mixing state of BC describes the
64 distribution of the bare BC and coating masteries among the aerosol population.
65 Typically, freshly generated BC exists in the form of chain aggregates and initially
66 uncoated, which is known as externally mixed BC (Ex-BC). When the BC particles

67 were emitted, they generally mix with other materials by condensation, coagulation,
68 and other processes (Riemer et al., 2004; Zhang et al., 2008; Liu et al., 2013; Zhang et
69 al., 2020a), forming the internally mixed BC (In-BC) particles consisting of BC core
70 and other chemical components (Cheng et al., 2006; Zhang et al., 2016b). The BC
71 structure would be more compact with regular shapes (Pagels et al., 2009; Zhang et al.,
72 2008; Wang et al., 2017), and the effective density of internally mixed BC is changed
73 accordingly with the reconstruction (Liu et al., 2019b). The density and morphology of
74 BC particles are closely related to its sources, mobility size, coating thickness, coating
75 material and its chemical composition (Pagels et al., 2009; Zhang et al., 2022). A wide
76 range of BC density has been reported in previous studies (Lide 1992; McMurry et al.,
77 2002; Park et al., 2004; Kiselev et al., 2010). Recent field measurement has indicated
78 that the average BC density is $\sim 1.2 \text{ g cm}^{-3}$ in the ambient atmosphere (Zhang et al.,
79 2016b). Field measurements have also indicated that a considerable fraction of
80 externally mixed/uncoated BC exists (Clarke et al., 2004; Cheng et al., 2012), although
81 a higher proportion of internally mixed/aged BC particles in the ambient atmosphere
82 were observed (Schwarz et al., 2008; Massoli et al., 2015; Chen et al., 2020). In climate
83 models, the BC was generally assumed completely internally mixed and treated to have
84 a void-free spherical structure and a density value of 1.8 g cm^{-3} (Bond et al., 2013). This
85 may lead to bias in estimating the climate effect driven by BC.

86 Previous study based on a case study has shown that when the aging degree of
87 ambient particles was low, the BC density ($\sim 1.8 \text{ g cm}^{-3}$) under the spherical assumption
88 would lead to the overestimation of particle hygroscopicity by 40-50 % and the

89 overestimation could be explained almost 100 % using the effective density of fresh BC
90 ($\sim 0.45 \text{ g cm}^{-3}$) (Fan et al. 2020). This indicates the importance of using reasonable BC
91 density values in the calculation of particle hygroscopicity. In addition, when estimating
92 the CCN number concentration, a significant bias of $-35 \% \sim +20 \%$ was found due to
93 the assumption of particle mixing state (Ren et al., 2018). However, these studies have
94 not yet accounted for such impact of BC density and mixing state on CCN prediction
95 due to lack of real time measurement data. Moreover, although the BC accounts for
96 very small mass fractions (5~10 %) in total fine aerosols, according to our previous
97 field observed results, the BC-containing particles could contribute 60 %-78 % toward
98 the total number concentration in urban Beijing (Chen et al., 2020). This is comparable
99 to the other results using SP2 instrument, which measured that the number fractions of
100 the coated BC-containing aerosols could be as high as about 50-80 % at the field sites
101 in north China (Liu et al., 2019b; Zhao et al., 2022). Therefore, the effect of BC density
102 on the uncertainty of CCN prediction should be concerned carefully.

103 The mixing state and the density of BC particles are usually directly measured by
104 several techniques, such as an integrated system of a volatility tandem differential
105 mobility analyzer and a single particle soot photometer (VTDMA-SP2) (Zhang et al.,
106 2016b), or a differential mobility analyzer with a SP2 (DMA-SP2) (Olfert et al., 2007;
107 Rissler et al., 2014; Wu et al., 2019), and a differential mobility analyzer–centrifugal
108 particle analyzer–single-particle soot photometer (DMA–CPMA–SP2) system (Liu et
109 al., 2019b; Yu et al., 2020), etc. However, such techniques or measurements are not
110 available in many previously conducted field campaigns. In this study, we develop a

111 novel method for retrieving the mixing state and effective density of ambient BC
112 particles by combining field measured hygroscopic growth factor and aerosol chemical
113 composition and Köhler theory (Petters and Kreidenweis, 2007). The uncertainty of the
114 new retrieval method was evaluated. The retrieved results were also compared and
115 validated with existing observations. In addition, the effect of BC density and mixing
116 state on prediction of CCN number concentrations was further evaluated through a
117 sensitivity and closure test by accounting for the retrieved real-time variations of BC
118 density and mixing state.

119 **2 Field measurements and methodology**

120 **2.1 Field measurements**

121 Measurements in this study were conducted from 15 November to 14 December
122 2016 at a typical urban site of Beijing (39.97°N, 116.37°E, 49 m above sea level). The
123 site locates at the Institute of Atmospheric Physics, Chinese Academy of Sciences,
124 which is mainly influenced by the surrounding cooking, road traffic and residential coal
125 burning emissions during the home heating periods (Sun et al., 2016). The detailed
126 information about the sampling site was presented in previous studies (Sun et al., 2016;
127 Zhang et al., 2019). The number concentration of condensation nuclei (CN) at each size
128 was measured by a scanning mobility particle sizer, which was equipped with a
129 differential mobility analyzer (DMA; model 3081, TSI) and a condensation particle
130 counter (CPC; model 3772, TSI). Subsequently, the mono-dispersed particles were
131 introduced into a Droplet Measurement Technologies CCN counter (CCNc, DMT;

132 Lance et al., 2006) to measure CCN number concentration. A hygroscopic tandem
133 differential mobility analyzer (HTDMA) system was used to measure the hygroscopic
134 growth factor (Gf) (Tan et al., 2013). Here, four diameters of 40, 80, 110, 150, and 200
135 nm were selected in the campaign. Gf is defined as the ratio of the mobility diameter at
136 the given RH to the dry diameter (Petters and Kreidenweis, 2007). The nonrefractory
137 submicron aerosol chemical composition was measured by an Aerodyne high-
138 resolution time-of-flight aerosol mass spectrometer (HR-AMS; Xu et al., 2019),
139 including sulfate, nitrate, ammonium, chloride, and organics. Two factors, including a
140 non-hygroscopic primary organic aerosol (POA) and hygroscopic secondary organic
141 aerosol (SOA) were classified by positive matrix factorization (PMF) with PMF
142 algorithm (v4.2) method (Paatero and Tapper, 1994) and followed the procedures
143 reported in Ulbrich et al. (2009). The refractory black carbon mass loading was
144 measured by an aethalometer (model AE33, Magee Scientific Corporation). Both the
145 nonrefractory materials and BC mass concentration were measured with diameters <
146 1.0 μm . The detailed description of the instrument operation and data process have been
147 described in details elsewhere (Ren et al., 2018; Xu et al., 2019; Zhang et al., 2019; Fan
148 et al., 2020).

149 **2.2 Retrieving the mixing state and density of BC**

150 2.2.1 Retrieving the mixing state of BC

151 The Gf probability distribution function (Gf-PDF) for a specified diameter can be
152 retrieved firstly based on the TDMA_{inv} algorithm (Gysel et al., 2009). The κ -PDF can

153 be further calculated based on the Gf-PDF (Fan et al., 2020). Size-resolved κ is derived
 154 using κ -Köhler theory based on hygroscopic growth factor (Gf) (Petters and
 155 Kreidenweis, 2007),

$$156 \quad \kappa_{gf} = (Gf^3 - 1) \cdot \left[\frac{1}{RH} \exp\left(\frac{4\sigma_{s/a}M_w}{RT\rho_w D_d Gf}\right) - 1 \right] \quad (1)$$

157 where Gf is hygroscopic growth factor, RH is the relative humidity in the HTDMA
 158 (90 %), D_d is the dry diameter, $\sigma_{s/a}$ is assumed to be the surface tension of pure water,
 159 R is the universal gas constant, T is the temperature, M_w and ρ_w is the molecular mass,
 160 and the density of water, respectively.

161 The κ -PDF patterns of particles in different sizes always present two modes: nearly
 162 hydrophobic (NH) mode with $\kappa_{gf} \leq 0.1$ and more hygroscopic (MH) mode with $\kappa_{gf} >$
 163 0.1 (Fig. S1). Firstly, based on the κ -PDF patterns, the number fraction (NF) of the total
 164 nearly hydrophobic group with the boundary of $[0, 0.1]$ was calculated according to the
 165 following equation:

$$166 \quad NF = \int_0^{0.1} c(\kappa, D_p) d\kappa \quad (2)$$

167 here, the κ -PDF, represented by $c(\kappa, D_p)$, was normalized as $\int c(\kappa, D_p) d\kappa = 1$, where
 168 κ can be replaced by κ_{gf} , D_p is the selected electrical mobility diameter in the campaign.

169 The nearly hydrophobic mode consists of both externally mixed POA (Ex-POA or
 170 bare POA) and externally mixed BC (Ex-BC). Since the number fraction of the nearly-
 171 hydrophobic POA would change with the emission and aging processes, in this study,
 172 we have applied different values for the number fractions of hydrophobic POA (NH-
 173 POA) under clean (91 %), moderately polluted (70 %), and heavily polluted conditions

174 (31 %) by referring to the literature (Liu et al., 2021a), as shown in Fig. S2. The number
 175 concentration of Ex-BC was then calculated using the total number fraction of NH
 176 mode minus the number of NH-POA.

$$\begin{aligned}
 177 \quad N_{POA-containing} &= N_{total} \times NF_{POA-containing} \\
 178 \quad N_{bare-POA} &= N_{POA-containing} \times NF_{bare-POA} \\
 179 \quad N_{Ex-BC} &= N_{NH} - N_{bare-POA} \quad (3)
 \end{aligned}$$

180 where $N_{POA-containing}$ and $NF_{POA-containing}$ are the number concentration and fraction of
 181 POA-containing particles, N_{total} is the total number concentration, $N_{bare-POA}$ and NF_{bare-}
 182 POA are the number concentration and fraction of bare POA particles, and N_{NH} is the
 183 number of nearly hydrophobic group.

184 The number size distribution of the externally mixed BC ($n_{Ex-BC}(\log D_p)$) can be
 185 calculated based on the particle number size distribution (PNSD) and the number
 186 fraction of the hydrophobic mode of BC (NF_{Ex-BC}) as follows:

$$187 \quad n_{Ex-BC}(\log D_p) = NF_{Ex-BC} \times n(\log D_p) \quad (4)$$

188 where $n(\log D_p)$ is the function of the aerosol number size distribution, D_p is the
 189 mobility diameter.

190 By assuming that the particles were spherical (Rader and McMurry, 1986), the
 191 mass size distribution of Ex-BC (M_{Ex-BC}) was obtained as follows:

$$192 \quad M_{Ex-BC}(\log D_p) = \frac{\pi}{6} D_p^3 \rho n_{Ex-BC}(\log D_p) \quad (5)$$

193 where D_p is the mobility diameter, ρ is the effective density of Ex-BC, and $n_{Ex-BC}(\log$
 194 $D_p)$ is the function of the number size distribution of Ex-BC, respectively. By reviewing
 195 and summarizing the existing results, we show that typical values of density for the

196 freshly emitted or externally mixed BC observed in the winter of urban Beijing or North
 197 China Plain span over 0.14-0.50 g cm⁻³, with a mean of ~0.40±0.10 g cm⁻³ (Fig. S3), in
 198 the size range of 100 to 300 nm, where the mass concentration of externally mixed BC
 199 mostly concentrated (Geller et al., 2006; Wu et al., 2019; Liu et al., 2020; Zhao et al.,
 200 2022). Therefore, an average $\rho_{\text{Ex-BC}}$ of 0.4 g cm⁻³ was used for calculating the mass
 201 concentration of externally mixed BC in our study. Uncertainty analyses due to the
 202 variations of $\rho_{\text{Ex-BC}}$ were given in section 2.3.

203 The mass size distribution of Ex-BC was fit using the log-normal distribution as
 204 shown in Fig. S4 (Wu et al., 2017; Liu et al., 2019a; Zhao et al., 2022). Thus, the bulk
 205 mass concentration of Ex-BC can be calculated from the integration of the mass size
 206 distribution:

$$207 \quad m_{\text{Ex-BC}} = \int_{D_{\text{start}}}^{D_{\text{end}}} M_{\text{Ex-BC}}(\log D_p) d \log(D_p) \quad (6)$$

$$208 \quad m_{\text{In-BC}} = m_{\text{BC}} - m_{\text{Ex-BC}} \quad (7)$$

209 where D_{start} and D_{end} are the lower and upper size limit, $M_{\text{Ex-BC}}(\log D_p)$ is the function
 210 of the Ex-BC mass size distribution. We then obtained the bulk mass concentration of
 211 internally mixed BC ($m_{\text{In-BC}}$) by subtracting $m_{\text{Ex-BC}}$ from the bulk BC mass
 212 concentration measured by AE33 in equation (7). It should be noted that the mass
 213 concentration of BC obtained from AE33 based on aerosol light absorption may lead
 214 some uncertainties, as has been further addressed in section 2.3.

215 2.2.2 Retrieving the density of BC

216 For retrieval of the density of BC, the principal idea is to use the measured κ_{gf} to

217 calculate the density of BC based on the Zdanovskii–Stokes–Robinson (ZSR) mixing
218 rule (Stokes and Robinson, 1966; Zdanovskii, 1948) with the chemical composition
219 measured by AMS (Petters & Kreidenweis, 2007). In the retrieval, several aspects are
220 concerned. First, since the ZSR rule assumes the aerosol particles are internally mixed,
221 the κ_{gf} value of the more MH mode ($\kappa_{\text{gf-MH}}$) is thus applied for retrieving the density of
222 internally mixed BC. Second, since the size distribution of BC number concentration is
223 usually with peaks between 100 and 200 nm (Liu et al., 2019a; Yu et al., 2020; Zhao et
224 al., 2022), the $\kappa_{\text{gf-MH}}$ value of particles in accumulation mode was averaged and applied
225 for the retrieval. Previous studies showed an independence of $\kappa_{\text{gf-MH}}$ on particle size
226 when the $D_p > 100$ nm during the campaign period (Fan et al., 2020). Therefore, the
227 average of $\kappa_{\text{gf-MH}}$ in accumulation mode is reasonable for the determination of the In-
228 BC density. Third, since only one hydrophobic and/or one hygroscopic mode was
229 observed by the HTDMA in most cases during the campaign (Fig. S1, S5), the chemical
230 components of the more hygroscopic (MH) mode at a given diameter should contain
231 both these hygroscopic non-BC and the coatings on BC-containing particles, which
232 would be measured by the HR-AMS instrument together. Therefore, by subtracting the
233 externally mixed POA in non-hygroscopic mode (see section 2.3), the concentration
234 and mass fraction of each component measured by HR-AMS can represent the overall
235 chemical composition of MH modes, and thus was applied in the ZSR mixing rule for
236 the retrieval of the density of internally mixed BC in this study. In addition, because the
237 inversion including measurements from HTDMA and HR-AMS, a total mass closure
238 of the measured aerosol particles was conducted between the two techniques by

239 comparing the mass concentration of PM₁ and the results are well consistent (Fig. S6).

240 The density of internally mixed BC (In-BC), ρ_{In-BC} is then derived from the following

241 equations:

$$242 \quad \kappa_{gf-MH} = \kappa_{chem} = \sum_i \varepsilon_i \kappa_i = \frac{v_{inorg}}{v_{total}} \kappa_{inorg} + \frac{v_{SOA}}{v_{total}} \kappa_{SOA} + \frac{v_{In-POA}}{v_{total}} \kappa_{POA} + \frac{v_{In-BC}}{v_{total}} \kappa_{BC} \quad (8)$$

243 where κ_{gf-MH} is the hygroscopic parameter of the more hygroscopic (MH) mode, κ_{chem}

244 is the hygroscopic parameter of aerosol particles in the mixed composition and can be

245 calculated based on chemical volume fractions using a simple rule (Stokes and

246 Robinson, 1966; Petters & Kreidenweis, 2007), κ_i is the hygroscopic parameter of each

247 pure composition and ε_i is the volume fraction of the individual components in the

248 internally mixed particle. v_{inorg} , v_{SOA} and v_{In-POA} are the volume of the inorganic, SOA

249 and internally mixed POA species, and can be calculated as follows: $v_{inorg} = \frac{m_{inorg}}{\rho_{inorg}}$,

250 $v_{SOA} = \frac{m_{SOA}}{\rho_{SOA}}$, and $v_{In-POA} = \frac{m_{In-POA}}{\rho_{POA}}$. v_{total} is the total volume of all the species and can be

251 written as $v_{total} = \frac{m_{inorg}}{\rho_{inorg}} + \frac{m_{SOA}}{\rho_{SOA}} + \frac{m_{In-POA}}{\rho_{POA}} + \frac{m_{In-BC}}{\rho_{In-BC}}$. In equation (8), κ_{BC} and κ_{POA} are

252 assumed to be 0. So, the total volume v_{total} can be further written as $v_{total} =$

253 $\frac{v_{inorg} \kappa_{inorg} + v_{SOA} \kappa_{SOA}}{\kappa_{gf-MH}}$. The volume of internally mixed v_{In-BC} can be calculated as follows,

$$254 \quad v_{In-BC} = \frac{v_{inorg} \kappa_{inorg} + v_{SOA} \kappa_{SOA}}{\kappa_{gf-MH}} - v_{inorg} - v_{SOA} - v_{In-POA}$$

$$255 \quad = \frac{\frac{m_{inorg}}{\rho_{inorg}} \kappa_{inorg} + \frac{m_{SOA}}{\rho_{SOA}} \kappa_{SOA}}{\kappa_{gf-MH}} - \frac{m_{inorg}}{\rho_{inorg}} - \frac{m_{SOA}}{\rho_{SOA}} - \frac{m_{In-POA}}{\rho_{POA}} \quad (9)$$

256 then, the ρ_{In-BC} can be calculated based on its mass concentration and volume as follows:

$$257 \quad \rho_{In-BC} = \frac{m_{In-BC}}{\left(\frac{m_{inorg} \kappa_{inorg} + m_{SOA} \kappa_{SOA}}{\kappa_{gf-MH}} - \frac{m_{inorg}}{\rho_{inorg}} - \frac{m_{SOA}}{\rho_{SOA}} - \frac{m_{In-POA}}{\rho_{POA}} \right)} \quad (10)$$

258 where, m_{In-BC} is the mass concentration of internally mixed BC, m_{inorg} and m_{SOA} are the

259 mass concentrations of the inorganic species and SOA, which are measured by the AMS.

260 m_{In-POA} is the mass concentrations of internally mixed POA and can be calculated

261 through subtracting the mass fraction of NH-POA from the total mass concentrations
 262 of POA. ρ_{inorg} , ρ_{SOA} and ρ_{POA} are the density of the inorganic species, SOA and POA.
 263 Since the AMS measures the concentrations of the organic and inorganic ions, including
 264 SO_4^{2-} , NO_3^- , NH_4^+ , Cl^- . Here inorganic species were derived by applying a simplified
 265 ion pairing scheme (Gysel et al., 2007) to convert mass concentrations of ions to the
 266 inorganic salts as follows:

$$\begin{aligned}
 268 \quad n_{\text{NH}_4\text{NO}_3} &= n_{\text{NO}_3^-} \\
 269 \quad n_{\text{NH}_4\text{HSO}_4} &= \min(2n_{\text{SO}_4^{2-}} - n_{\text{NH}_4^+} + n_{\text{NO}_3^-}, n_{\text{NH}_4^+} - n_{\text{NO}_3^-}) \\
 270 \quad n_{(\text{NH}_4)_2\text{SO}_4} &= \max(n_{\text{NH}_4^+} - n_{\text{NO}_3^-} - n_{\text{SO}_4^{2-}}, 0) \\
 267 \quad n_{\text{H}_2\text{SO}_4} &= \max(0, n_{\text{SO}_4^{2-}} - n_{\text{NH}_4^+} + n_{\text{NO}_3^-}) \quad (11)
 \end{aligned}$$

271 where n represents the number of moles, and the mass concentrations were obtained by
 272 the number of moles times the molar mass of each inorganic salts. Because the value of
 273 the $n_{\text{H}_2\text{SO}_4}$ was zero in this campaign. Three inorganic salts including NH_4HSO_4 ,
 274 $(\text{NH}_4)_2\text{SO}_4$, and NH_4NO_3 were applied in our study. The densities for inorganic salts
 275 were taken from previous studies (Gysel et al., 2007; Wu et al., 2016). Here the densities
 276 for three inorganics are 1.78, 1.77 and 1.72 g cm^{-3} , respectively. By summarizing the
 277 previous studies (Gysel et al., 2007; Dinar et al., 2006), 1.4 g cm^{-3} was selected as the
 278 density of SOA (ρ_{SOA}). The density of POA (ρ_{POA}) is assumed to be 1.0 g cm^{-3} for urban
 279 environments, which is similar to the that of the lubricating oil (Wu et al., 2016). Since
 280 the cooking organic aerosols represent a high contribution to POA in urban
 281 environments, we choose the mean density of the rapeseed oil and oleic acid ($\sim 0.85 \text{ g}$
 282 cm^{-3}) (Reyes-Villegas et al., 2018) to evaluated the result as shown in section 2.3. The

283 values of κ for inorganic components are 0.56 for NH_4HSO_4 , 0.48 for $(\text{NH}_4)_2\text{SO}_4$ and
284 0.58 for NH_4NO_3 , along with the best-fit values for the three inorganic salts (Petters &
285 Kreidenweis, 2007 and Gunthe et al., 2009). The κ_{SOA} is assumed to be 0.15 according
286 to the field studies in urban areas (Chang et al., 2010; Kawana et al., 2016).

287 Note that the method fails to retrieve the BC density when organic accounts for a
288 large fraction (>60 %). This is because that a higher fraction of OA usually corresponds
289 to lower total volume of all the species (Fig. S7), yielding negative values for $v_{\text{In-BC}}$
290 introduced in equation (9). As a result, 61 % of the data observed during the campaign
291 were valid for calculating the BC density.

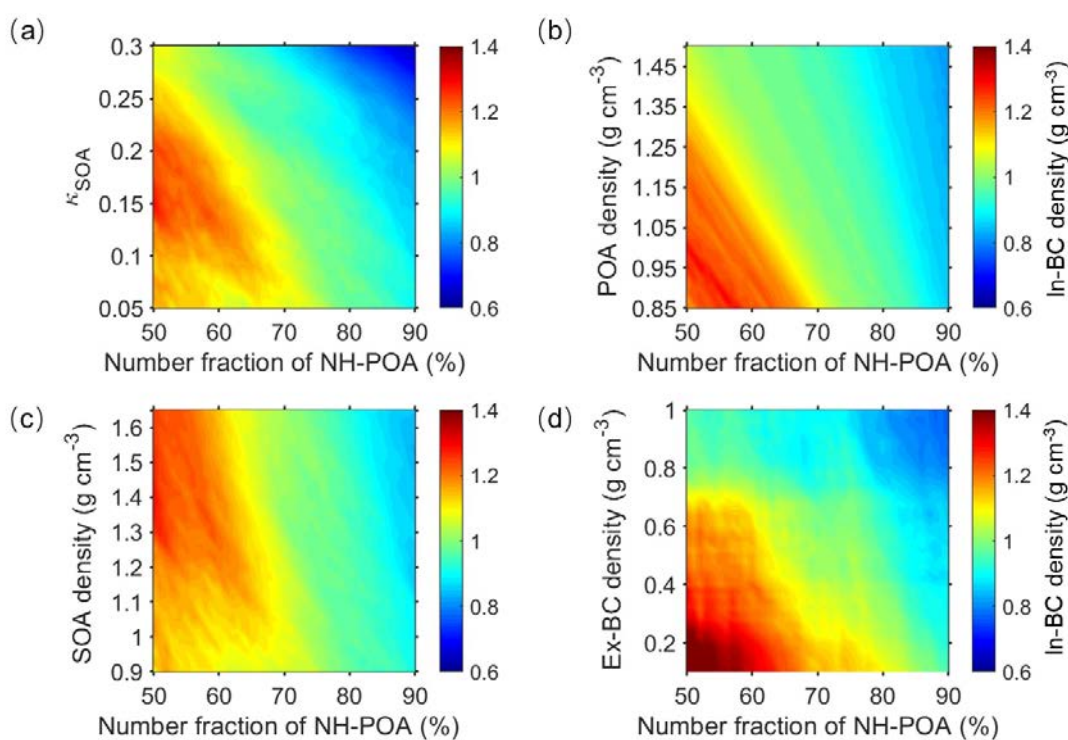
292 Similarly, the bulk density of BC ($\rho_{\text{bulk-BC}}$) is calculated with the same method as
293 that for calculating the $\rho_{\text{In-BC}}$. When calculating the $\rho_{\text{bulk-BC}}$, the bulk κ_{gf} value measured
294 by HTDMA is applied assuming that all the aerosol particles are internally mixed.

295 **2.3 Uncertainties and limitations**

296 For the retrieval, the assumptions on the values of κ_{SOA} , ρ_{POA} , ρ_{SOA} and $\rho_{\text{Ex-BC}}$ as
297 well as the fraction of primary organic aerosols in non-hygroscopic mode would add
298 uncertainty in the inferred values of ambient internally mixed BC density. For example,
299 the freshly emitted POA particles might consistently be coated with secondary particles
300 during the aging process, resulting in changes of the $NF_{\text{NH-POA}}$. However, a real-time
301 variation of the $NF_{\text{NH-POA}}$ is not yet available due to the lack of such measurement data.

302 Applying only rough fractions of hydrophobic POA for three different
303 atmospheric conditions could still cause uncertainties. Also, the densities of POA and

304 SOA may differ due to their precursors, emission sources and the formation
 305 mechanisms in ambient atmosphere (Alfarra et al., 2006; Reyes-Villegas et al., 2018).
 306 The density of Ex-BC is generally characterized by the morphology and size (Wu et al.,
 307 2019). In addition, the value of κ_{SOA} spans largely due to the variations in the emissions
 308 of gas precursors and formation processes under different atmospheric conditions
 309 (Zhang et al., 2015; Liu et al., 2021b). Therefore, we examined the sensitivities of In-
 310 BC density to the variations of these factors, as exhibited in Fig. 1 and Fig. 2.

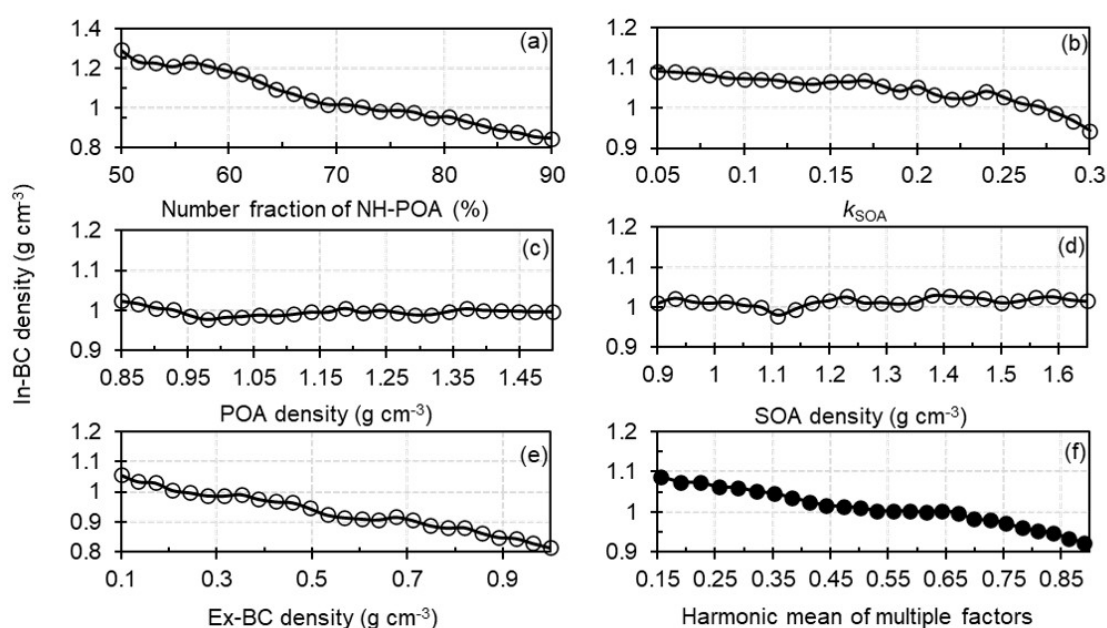


311

312 **Figure 1.** Sensitivities of In-BC density to the variations in the number fraction of
 313 nearly hydrophobic (NH) POA and hygroscopic parameter of OA (κ_{SOA}) (a), POA
 314 density (b), SOA density (c) and the externally mixed BC density (d).

315 The figures show that the In-BC density gradually decreased with the increment
 316 of the $NF_{\text{NH-POA}}$, implying the high fraction of bare POA particles corresponded to the
 317 early aging stage of aerosol particles. With the increase of κ_{SOA} , the In-BC density was

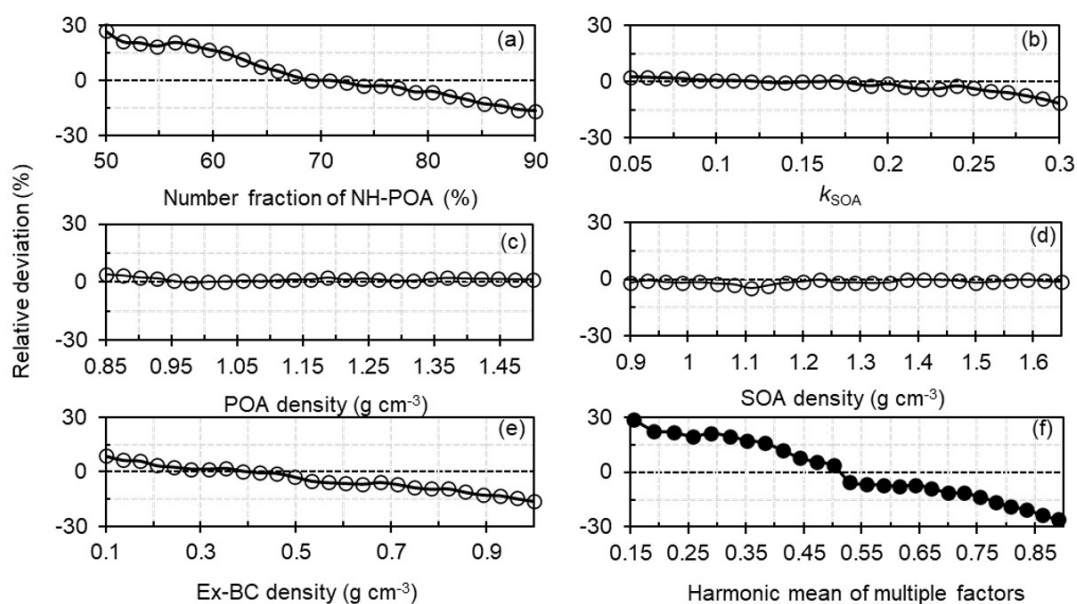
318 generally reduced, but with small fluctuations (Fig. 1a, Fig. 2b). This suggests a
 319 complex impact of assumptions of κ_{SOA} on the retrieved BC density. In addition, the In-
 320 BC density decreased slightly as $\rho_{\text{Ex-BC}}$ increased (Fig. 2e), suggesting applying a larger
 321 $\rho_{\text{Ex-BC}}$ would derive smaller values for In-BC density. The In-BC density was
 322 insensitive to the changes of the density of POA and SOA, showing an almost negligible
 323 effect on the retrieved results (Fig. 2c and d).



324
 325 **Figure 2.** Sensitivity of the In-BC density to variations in the number fraction of nearly
 326 hydrophobic (NH) POA (a), the hygroscopic parameter of SOA (b), the POA density
 327 (c), the SOA density (d), the externally mixed BC density (e) and the harmonic mean
 328 of multiple factors (f).

329 The uncertainty analysis shows that, by comparing the results based on the mean
 330 fraction of the $NF_{\text{NH-POA}}$ with a typical atmospheric observed range of 50-90 % for the
 331 $NF_{\text{NH-POA}}$ (Liu et al., 2021a), the assumption on $NF_{\text{NH-POA}}$ can lead to relative deviations
 332 (uncertainty) of -17 % ~ +27 % for the retrieved BC density (Fig. 3a).

333 In addition, unlike inorganics (eg., NH_4HSO_4 , $(\text{NH}_4)_2\text{SO}_4$ and NH_4NO_3), for which
 334 the hygroscopicity has been already well-understood (Petters and Kreidenweis, 2007),
 335 the hygroscopicity of organic species varies largely due to the complexity in organic
 336 aerosol constituents. Therefore, the assumption of the values of κ_{SOA} will add the
 337 uncertainty in the calculation of BC density. Previous studies have suggested that the
 338 organics have a wide range of κ values ranging from 0.05 to 0.3 (Jimenez et al., 2009;
 339 Mei et al., 2013). Thus, the sensitivity test has also been done to examine the effect due
 340 to changes in κ_{SOA} on calculating the density of BC (Fig. 1a). The result shows that the
 341 assumption of κ_{SOA} values can cause an average relative deviation of -10 % ~ +3 % in
 342 calculating the density of In-BC (Fig. 3b).



343
 344 **Figure 3.** Relative deviations of the number fraction of nearly hydrophobic (NH) POA
 345 to the In-BC density (a), the hygroscopic parameter of OA to the In-BC density (b), the
 346 POA density to the In-BC density (c), the SOA density to the In-BC density (d), the
 347 externally mixed BC density to In-BC density (e) and the combined deviations based

348 on multiple factors mentioned above (f).

349 However, the sensitivity test shows that the impact of both the ρ_{POA} and ρ_{SOA}
350 variations on the BC density estimation was very small or even negligible (Fig. 1b, c).
351 By varying the ρ_{POA} from 0.85 to 1.5 g cm⁻³ and the ρ_{SOA} from 0.9 to 1.65 g cm⁻³
352 according to the literature (Noureddini et al., 1992; Alfarra et al., 2006; Reyes-Villegas
353 et al., 2018; Kostenidou et al., 2007), the retrieval uncertainties in the BC density were
354 both within $\pm 5\%$ (Fig. 3c, d). For $\rho_{\text{Ex-BC}}$, it exhibited that the evolution of the $\rho_{\text{Ex-BC}}$
355 could lead to an average deviation of $-16\% \sim +9\%$ in calculating In-BC density (Fig.
356 3e) when increasing the values of $\rho_{\text{Ex-BC}}$ from 0.1 to 1.0 g cm⁻³, which represents a
357 typical range in ambient atmosphere (Wu et al., 2019; Liu et al., 2020). A combined
358 uncertainty (δ) caused by the multiple factors (δ_i), which was calculated by equation
359 (12), was $-26\% \sim +29\%$ as shown in Fig. 3f.

$$360 \quad \delta = \sqrt{\sum_{i=1}^n \delta_i^2} \quad (12)$$

361 In addition, it should be noted that the mass concentration of BC obtained from
362 AE33 based on aerosol light absorption may lead some uncertainties. However, the
363 comparison of the simultaneously measured data by SP2 with those by AE33 during the
364 campaign shows that the temporal variations of BC mass concentrations measured by
365 the two techniques were well consistent (Fig. S8). Note that the BC mass measured by
366 SP2 is occasionally low probably because of the low detection efficiency in small size
367 (McMeeking et al., 2010; Schwarz et al., 2006). In addition, the SP2 is unable to
368 quantify the BC mass beyond a certain limit because of the saturation of electronic
369 devices recording signals (Pileci et al., 2021). We show that, compared the results that

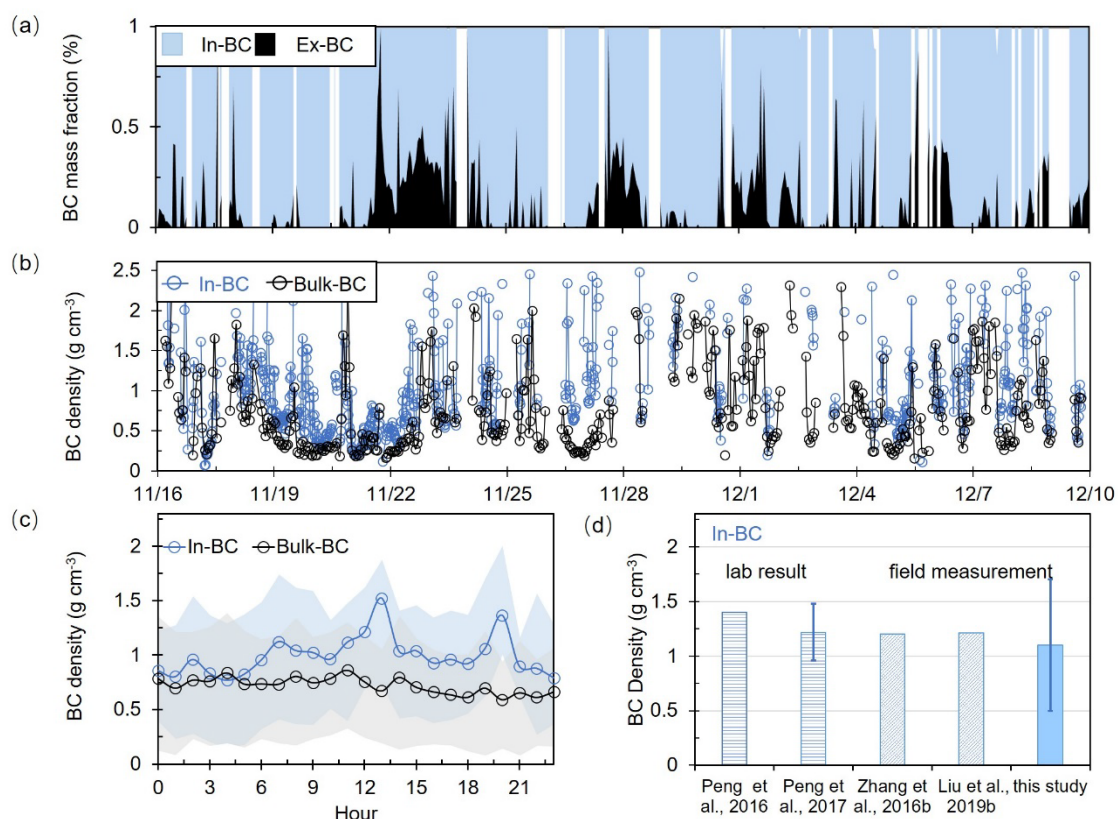
370 were retrieved if applying the BC mass measured by SP2, the BC density retrieved
371 based on AE33 can be 18 % higher. Given the measurement bias from SP2, this
372 overestimation indicates an upper limit of the uncertainty.

373 **3 Results and Discussion**

374 **3.1 A comparison and validation of retrieved mixing state and density of BC**

375 Figure 4a shows retrieved time series of the mixing state of ambient BC during the
376 campaign. Large temporal variations of the mass fraction of internally and externally
377 mixed BC present during the observed period at the sites. The temporal changes should
378 be related to the atmospheric aging process or diurnal variations of emissions (Liu et
379 al., 2019a; Fan et al., 2020). Statistically, the average mass fractions of externally and
380 internally mixed BC were 20 ± 18 % and 80 ± 20 % respectively, showing that most of
381 the BC particles were aged and internally mixed with other components. Previous
382 studies at urban sites have shown the co-existence of the externally mixed BC in the
383 ambient atmosphere (Schwarz et al., 2008; Cheng et al., 2012; Chen et al., 2020) due
384 to continuous combustion processes (e.g., vehicle exhaust and residential sector) (Wang
385 et al., 2017; Liu et al., 2019a). Our results are basically comparable with those
386 previously reported results, which are directly measured or indirectly retrieved. For
387 example, Chen et al. (2020) found that the mass fraction of internally mixed BC
388 particles was nearly ~80–90 % in summer of Beijing based on VTDMA measurements.
389 Liu et al. (2020) using a tandem system with an aerodynamic aerosol classifier and SP2,
390 reported that the mass fraction of internally BC-containing particles would increase

391 with increasing size and reach ~70 % in Beijing. Overall, the mass fraction obtained in
 392 our study was comparable with that reported in urban Beijing. Previous studies also
 393 displayed that the significant diversity of the BC mixing state among emission
 394 conditions and coating processes (Shiraiwa et al., 2008; Pan et al., 2017; Zhang et al.,
 395 2020b). Accordingly, the densities of the bulk and internally mixed BC present apparent
 396 fluctuations as shown in Fig. 4b, which is significantly affected by the variations of BC
 397 emission sources and BC aging processes. The density of the In-BC during daytime was

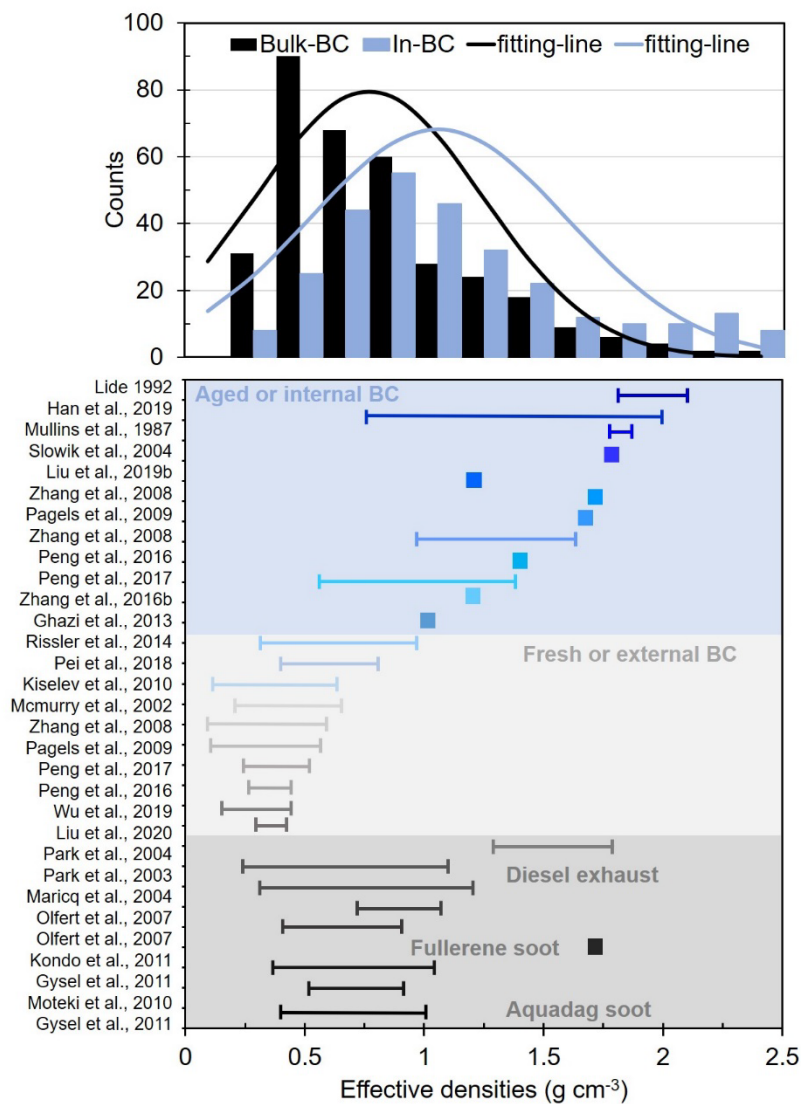


398

399 **Figure 4.** (a) Time series of the mass fraction of the retrieved internal- and external-
 400 mixed BC; (b) Time series of the retrieved density of the bulk and internal- mixed BC
 401 (In-BC); (c) Diurnal variation of the retrieved density of bulk and In-BC; (d)
 402 Comparison of the results of the derived In-BC density in this study with that reported
 403 in the literature.

404 generally higher than that at night (Fig. 4c). The elevated BC density during daytime
405 was likely due to the strong photochemical processes promote the aging of BC particles,
406 which resulted in a conversion from uncompacted structure to compact and regular
407 spherical shapes of BC (Qiao et al., 2018; Liu et al., 2019b). The rising in BC density
408 around 20:00 LT might indicate that the BC particles would be rapidly coated with
409 secondary inorganic aerosol (SIA) particles and continuously aged in the polluted
410 period due to the heterogeneous reactions of SIA in urban regions (Zhang et al., 2016b;
411 Peng et al., 2017). Actually, following the haze evolution, the fraction of nearly
412 hydrophobic group reduced rapidly (Fig. S9). Consequently, the average density of In-
413 BC increased obviously from the clean conditions to the polluted periods (Fig. S5). A
414 slight decrease was observed in the bulk BC density during traffic hours. This is likely
415 associated with the continuous emissions (e.g., vehicle exhaust) that lead to uncoated
416 or uncompacted BC particles in this period. The diurnal cycle in In-BC density was
417 consistent with the coating thickness measured by a tandem CPMA-SP2-DMA-SP2
418 (Liu et al., 2020), demonstrating that the new method can derive the density of ambient
419 BC particles reasonably. Averagely, the campaign average values of the bulk and
420 internally mixed BC densities were 0.7 ± 0.5 and 1.1 ± 0.6 g cm⁻³ respectively, which were
421 much less than 1.8 g cm⁻³, implying that the BC particles are not void-free spheres in
422 the urban atmosphere. The results of In-BC density were comparable with those
423 observed at other sites in North China Plain (NCP) as shown in Fig. 4d, illustrating that
424 the BC effective density retrieved by this method was within the range of density from
425 field measurements.

426 Based on both field measurements (e.g., Lide 1992; Zhang et al., 2016b; Wu et al.,
 427 2019; Liu et al., 2019b) and laboratory studies (e.g., McMurry et al., 2002; Park et al.,
 428 2003, 2004; Olfert et al., 2007; Kiselev et al., 2010; Gysel et al., 2011, 2012), the BC
 429 density from diverse combustion sources or representing different aging degree has



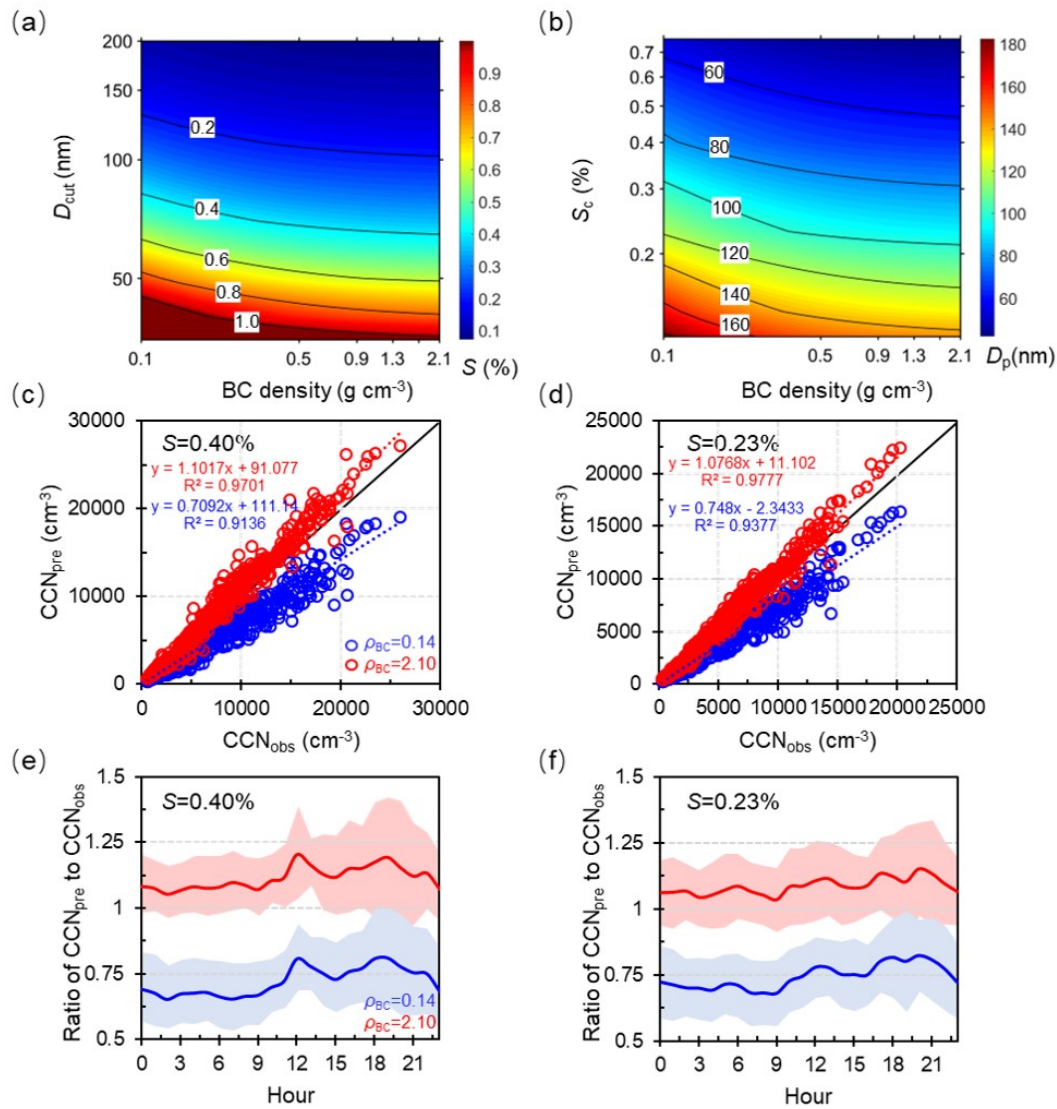
430
 431 **Figure 5.** The probability distribution function (PDF) of the retrieved density of bulk
 432 and In-BC and the measured density distribution spectrum of BC from different sources
 433 reported in the literature.
 434 been obtained and ranges widely from 0.14 to 2.1 g cm⁻³, as was summarized and shown
 435 in Fig. 5. Mean probability distribution function (PDF) of the density of bulk and In-

436 BC retrieved by this study is also presented in Fig. 5. It shows that the retrieved density
437 of bulk BC exhibited a dominant mode with a peak value at 0.7 g cm^{-3} , which was
438 situated between the typical density range of those externally mixed and internally
439 mixed BC measured previously. For the In-BC, the PDF was with a peak value at 1.1 g
440 cm^{-3} , but ranged widely from ~ 0.5 to 2.5 g cm^{-3} , which indicated various morphologies,
441 different aging degree and compositions of ambient BC particles due to the complex
442 impact of multiple local sources and aging processes during the observed period in
443 urban Beijing. Overall, the retrieved values for In-BC density fall within the range of
444 typical internally mixed BC reported in the literature, verifying the reliability of our
445 inversion results.

446 **3.2 Sensitivity of predicted N_{CCN} to changes of BC density**

447 A previous study showed that the use of an inaccurate density value of BC particles
448 would result in large biases in estimating κ of ambient aerosol particles with the ZSR
449 mixing rule (Fan et al., 2020), as would further lead to uncertainties in prediction of
450 N_{CCN} and relevant climate effects. Considering the large variation range of BC density
451 during the campaign, which is closely associated with BC morphology or degree of BC
452 aging, we further examine the sensitivity of critical supersaturation (S_c), critical
453 diameter (D_{cut}) and predicted N_{CCN} to variations of BC density (Fig. 6). Here, we use
454 the critical diameter and particle number size distribution to calculate N_{CCN} . The method
455 to derive the critical diameter is based on Köhler theory and ZSR rule. Three CCN
456 closure studies were assumed to evaluate the effect of BC density and mixing state on

457 prediction of CCN number concentrations. Closure studies provide a useful way to



458

459 **Figure 6.** Sensitivity of critical supersaturation (S_c) (a) and diameter (D_{cut}) (b) to the

460 variations in BC density; Predicted N_{CCN} as a function of measured N_{CCN} by varying the

461 density from 0.14 to 2.1 $g\ cm^{-3}$ at $S=0.40\%$ (c) and $S=0.23\%$ (d), the black solid line

462 is the 1:1 line; Diurnal variations in the ratio of predicted-to-measured N_{CCN} at $S=0.40\%$

463 (e) and $S=0.23\%$ (f).

464 investigate the importance of aerosol properties to CCN concentration prediction. If the

465 closure study is achieved, it means the bias between the predicted and measured CCN

466 concentrations is within $\pm 15\%$ (Chang et al., 2007). The detailed calculation methods

467 are presented in the supporting information (SI: Methods) or the reference in Ren et al.
468 (2018). The results show that, by varying the value of density from 0.14 to 2.1 g cm⁻³
469 that represents the range of BC density in the atmosphere, the D_{cut} reduced apparently
470 at a given supersaturation (S) (Fig. 6a), or similarly, the S_c decreased rapidly for a given
471 particle size (Fig. 6b). The results show that the changes of the D_{cut} and S_c were more
472 sensitive when the BC density was below 1.0 g cm⁻³. The effects on the D_{cut} and S_c both
473 gradually weakened with the increase of BC density. This shows that it is critical to
474 apply more accurate BC density for the aerosol particles with low aging degree in
475 predicting CCN and its climate effect. Accordingly, the ratios of predicted-to-measured
476 N_{CCN} ranged from 0.72 to 1.11 by varying the BC density from 0.14 to 2.1 g cm⁻³ at the
477 typical S of 0.23 % and 0.40 % (Fig. 6c, 6d), showing an estimation uncertainty of -28 %
478 ~ 11 % in N_{CCN} prediction.

479 The diurnal variations in the ratio of predicted-to-measured N_{CCN} at $S=0.40$ % and
480 0.23 % are shown to examine the response of the BC density on N_{CCN} prediction at
481 different time periods (Fig. 6e, 6f). By applying the lower limit of density value of 0.14
482 g cm⁻³, the prediction was much worse than the use of the density of 2.1 g cm⁻³ at night
483 (00:00-06:00 LT), when the latter was much closer to the real density of ambient BC
484 (Fig. 4c). The prediction was improved substantially by applying the value of 0.14 g
485 cm⁻³ during evening rush hours (18:00-20:00 LT), during which the ambient BC
486 particles were disturbed by the traffic emissions (Fig. 4c). The prediction became worse
487 by applying the value of 2.1 g cm⁻³ and an obvious overestimation by up to ~40 % was
488 shown. The results further illustrate that it is critical to account for the real-time mixing

489 state and density of BC particles in N_{CCN} prediction, particularly in regions with heavy
490 traffic and residential coal emissions.

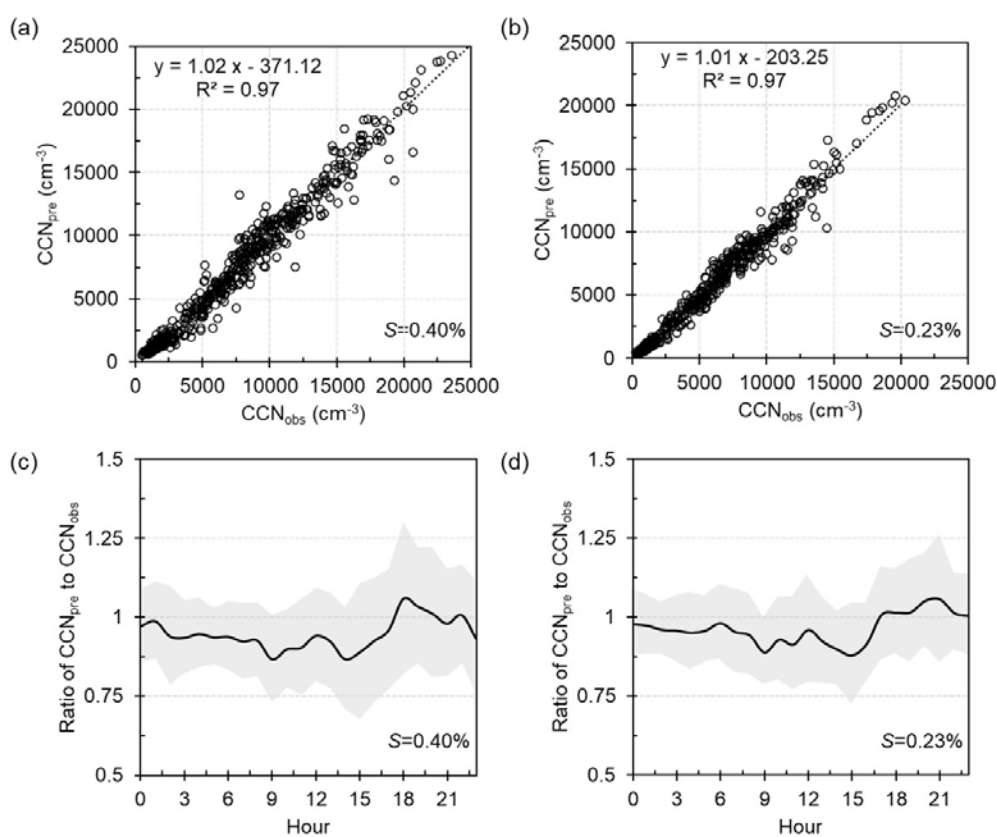
491 It should be noted that the assumption of the surface tension of water would
492 overestimate the critical diameter and underpredict CCN number concentration. While
493 the surface tension depression might be more obvious for the small size particles (<60
494 nm), as the fraction of organics is higher at small particles size (Meng et al., 2014; Cai
495 et al., 2018). Here, in this study, we calculated the critical diameters at supersaturations
496 of 0.40 % and 0.23 %, typical values in cloud, corresponding to larger sizes (> 70 nm
497 and 90 nm) of aerosols. Therefore, the uncertainties from the application of the surface
498 tension of pure water should be negligible (< 10 %).

499 **3.3 N_{CCN} prediction based on the real-time variations of BC density and mixing** 500 **state**

501 Figure 7 exhibits the comparisons between predicted and measured N_{CCN} at S of
502 0.23 % and 0.40 % by accounting for the retrieved real-time variations of BC density
503 and mixing state. It shows that the N_{CCN} can be well predicted with a slope of 1.01 and
504 1.02 at S of 0.23 % and 0.40 % respectively (Fig. 7a, 7b), only presenting a slight
505 deviation. The slight deviation is primarily due to the fixed value of the density for the
506 externally mixed BC caused by the retrieved method, especially during noontime and
507 evening rush periods (Fig. 7c and 7d).

508 The diurnal variations in the ratio of predicted-to-measured N_{CCN} shows the N_{CCN}
509 can be underestimated by up to 15 % at $S=0.40$ % during noontime. While, a slightly
510 overrated during the evening traffic hours and nighttime may be due to the

511 underestimation of the number fraction of Ex-BC. Overall, the dependence of the CCN
 512 prediction on S is due to the size dependence of κ and mixing state (Zhang et al., 2017;
 513 Liu et al., 2020; Ren et al., 2018). A better closure at $S=0.23\%$ is because the bulk κ of
 514 particles is closer to that at the critical diameter of 100-150 nm. Similarly, the effect on
 515 CCN prediction induced by the bulk mixing state would be more critical for smaller
 516 particles, corresponding to the critical diameter at high S .



517
 518 **Figure 7.** Predicted CCN number concentration using the mixing state and In-BC
 519 density derived from HTDMAs at $S=0.40\%$ (a) and $S=0.23\%$ (b). Diurnal variations
 520 in the ratio of predicted-to-measured N_{CCN} at $S=0.40\%$ (c) and $S=0.23\%$ (d).

521 Overall, when considering the effective density of BC relevant to its mixing state,
 522 the CCN closure achieves. Previous studies have shown that the fresh emitted BC
 523 particles may convert from fractal-like aggregates to a compact structure and its density

524 would increase with the aging process (Liu et al., 2019b; Zhang et al., 2020a, 2022),
525 but the actual density of In-BC may be lower than 1.8 g cm^{-3} in the ambient atmosphere
526 according to this study. Therefore, the currently applied value represents a density of
527 the void-free structure of BC particles may cause an overestimation in CCN prediction.
528 In addition, although the BC accounts for small mass fractions in ambient fine aerosols,
529 according to the measurements simultaneously conducted at the site, the BC-containing
530 particles could contribute 60 %-78 % toward the total number concentration in urban
531 Beijing (Chen et al., 2020). Our results further highlight the effect of BC density on the
532 uncertainty of CCN prediction should be concerned carefully.

533 **4 Conclusions**

534 The mixing state and effective density of BC changed through heterogenous
535 chemistry process and thus would cause uncertainty in evaluating its CCN activity. In
536 this study, we develop a new method to retrieve the mixing state and effective density
537 of ambient BC using field measurements and the Köhler theory. The uncertainty of the
538 new retrieval method was evaluated within $\pm 30 \%$, which was primarily caused by
539 assuming the value of κ_{SOA} and the fraction of primary organic aerosols in non-
540 hygroscopic mode. The retrieved results show that most of the BC particles were aged
541 and internally mixed with other components, with mean mass fraction of $80 \pm 20 \%$.
542 Averagely, the retrieved densities of the bulk and internally mixed BC were 0.7 ± 0.5 and
543 $1.1 \pm 0.6 \text{ g cm}^{-3}$ respectively, but ranged widely from ~ 0.1 to 2.5 g cm^{-3} , indicating
544 various morphologies, different aging degree and compositions of ambient BC particles

545 due to the complex impact of multiple local sources and aging processes during the
546 observed period. The retrieved result was basically comparable with previous
547 observations in North China Plain.

548 Further examination shows the uncertainties of the N_{CCN} prediction were -28 %
549 ~11 % at the typical S of 0.23 % and 0.40 % by varying the BC density from 0.14 to
550 2.1 g cm^{-3} that represented the range of ambient BC particles. Moreover, the prediction
551 was found more sensitive to the variations of BC density when it was $<1.0 \text{ g cm}^{-3}$,
552 suggesting a great significance of accounting for the effect of BC density for the aerosol
553 particles with low aging degree when evaluating the climate effect. The CCN closure
554 achieved when introducing the retrieved real-time BC density relevant to its mixing
555 state. This work provides a unique way of utilizing field observations to infer ambient
556 BC density and highlights the current assumption of a void-free structure of BC
557 particles in models would cause large uncertainties in CCN prediction and in the
558 relevant climate effect evaluation.

559 The method used to derive the ambient BC density has limitations. Since the
560 assumptions on the values of κ_{SOA} , ρ_{POA} , ρ_{SOA} and ρ_{EX-BC} as well as the fraction of
561 primary organic aerosols in non-hygroscopic mode would add uncertainty in the
562 inferred values of ambient internally mixed BC density. It is thus necessary to examine
563 observational data to verify this methodology in future studies. However, the method
564 and results of this study could provide the way for a more comprehensive understanding
565 of the variations in BC density in Beijing. Additionally, it has the potential to reveal the
566 uncertainties of usage of void-free structure of BC density in assessing the climate

567 effects.

568 **Data availability.**

569 All data needed to evaluate the conclusions in the paper are present in the paper and/or
570 the Supplement. All data used in the study are also available from the corresponding
571 author upon request (zhangfang2021@hit.edu.cn).

572 **Author contributions.**

573 FZ and JR conceived the conceptual development of the manuscript. JR directed and
574 performed of the experiments with JL, LC, and FZ. JR conducted the data analysis and
575 wrote the draft of the manuscript. All authors edited and commented on the various
576 sections of the manuscript.

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583 **Competing interests.**

584 The contact author has declared that neither they nor their co-authors have any
585 competing interests.

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