1	The density of ambient black carbon retrieved by a new method:
2	implications to CCN prediction
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25 Abstract.

26 The effective density of black carbon (BC) is a crucial factor relevant to its aging degree that would add uncertainty in evaluating its climate effect. Here, we have 27 developed a new method to retrieve the effective density of internally mixed BC in the 28 atmosphere combining field observations conducted duringfrom 15 November -to 14 29 December 2016 in urban Beijing with the Köhler theory. The uncertainty of the retrieval 30 31 method was evaluated within ±30 %, which iswas primarily caused by assumptions ofon both the hygroscopic parameter of organics and the fractionproportional 32 distribution of primary organic aerosols in non-different hygroscopic or hygroscopic 33 modemodes. Using the method, we obtain that the ambient internally mixed BC, 34 accounting for 80±20 % of total BC aerosol particles, iswas retrieved with a campaign 35 mean density of 1.1±0.6 g cm⁻³ during the observed periods. The retrieved result iswas 36 37 comparable with that reported in the Literatureliterature. By applying a lower (0.14 g 38 cm⁻³) and upper (2.1 g cm⁻³) limit of the retrieved BC density in cloud condensation 39 nuclei (CCN) number concentrations concentration (N_{CCN}) estimation, we derived that neglect of such variations in BC density would lead to an uncertainty of -28 %~11 % 40 in predicting N_{CCN} at supersaturations of 0.23 % and 0.40 %. We also find that the N_{CCN} 41 42 iswas more sensitive to the variations of BC density when it iswas <1.0 g cm⁻³. This illustrates a necessity of accounting for the effect of BC density on CCN activity closer 43 to source regions where the BC particles are mostly freshly emitted. The CCN closure 44 achieves achieved when introducing the retrieved real-time BC density and mixing state. 45

This study provides a unique way of utilizing field measurements to infer ambient BC
density and highlights the importance of applying variable BC density values in models
when predicting CCN and assessing its relevant climate effect.

49 1 Introduction

Black carbon (BC) aerosols, as the major absorber of solar radiation, play a vital 50 role in energy budget and climate of the earth-atmosphere system by affecting the 51 52 radiative forcing and cloud properties (Flanner et al., 2007; Ramanathan and Carmichael, 2008). The light-absorbing capability induced by BC is related to its 53 density and morphology (Zhang et al., 2008; Rissler et al., 2014), which can be 54 55 modified after mixing with other atmospheric aerosol particles (Khalizov et al., 2009; Xue et al., 2009). Changes in its physicochemical properties or the aging process would 56 also regulate its ability to serve as cloud condensation nuclei (CCN) and further 57 indirectly affect the CCN number concentrations (Zhang et al., 2016a, 2017; Ren et al., 58 59 2023) and the radiative balance by affecting the clouds process (Yuan et al., 2008; Wang et al., 2011). Owing to the complex evolution of the mixing state, density and 60 morphology of BC, the contribution of BC particles to CCN budgets is still not well 61 62 understood.

BC particles, with diesel vehicles, industrial and residential coal combustion as major sources, are ubiquitous in urban environments (Bond et al., 2013; Dameto et al., 2017; Li et al., 2017; Liu et al., 2019a). The mixing state of BC describes the distribution of the bare BC and coating masteries among the aerosol population.

67	Typically, freshly generated BC exists in the form of chain aggregates and initially
68	uncoated, which is known as externally mixed BC (Ex-BC). When the BC particles
69	were emitted, they generally mix with other materials by condensation, coagulation,
70	and other processes (Riemer et al., 2004; Zhang et al., 2008; Liu et al., 2013; Zhang et
71	al., 2020a), forming the internally mixed BC (In-BC) particles consisting of BC core
72	and other chemical components (Cheng et al., 2006; Zhang et al., 20162016b). The BC
73	structure would be more compact with regular shapes (Pagels et al., 2009; Zhang et al.,
74	2008; Wang et al., 2017), and the effective density of internally mixed BC areis changed
75	accordingly with the reconstruction (Liu et al., 2019b). The density and morphology of
76	BC particles are closely related to its sources, mobility size, coating thickness, coating
77	material and its chemical composition (Zhang et al., 2008; Pagels et al., 2009; Peng et
78	al., 2016; Zhang et al., 2022). A wide range of BC density has been reported in previous
79	studies (Lide 1992; McMurry et al., 2002; Park et al., 2004; Kiselev et al., 2010). Recent
80	field measurements have measurement has indicated that the average BC density is ~1.2
81	g cm ⁻³ in the ambient atmosphere (Zhang et al., $\frac{20162016b}{2016b}$). Field measurements have
82	also indicated that a considerable fraction of externally mixed/uncoated BC exists
83	(Clarke et al., 2004; Cheng et al., 2012), although a higher proportion of internally
84	mixed/aged BC particles in the ambient atmosphere were observed (Schwarz et al.,
85	2008; Massoli et al., 2015; Chen et al., 2020). In climate models, the BC was generally
86	assumed completely internally mixed and treated to have a void-free spherical structure
87	and a density value of 1.8 g cm ⁻³ (Bond et al., 2013). This may lead to bias in estimating
88	the climate effect driven by BC.

89	Previous study based on a case study showhas shown that when the aging degree	
90	of ambient particles iswas low, the BC density (~1.8 g cm ⁻³) under the spherical	
91	assumption willwould lead to the overestimation of particle hygroscopicity by 40-50 %	
92	and the overestimation eancould be explained almost 100 % using the effective density	
93	of fresh BC (~0.45 g cm ⁻³) (Fan et al. 2020). This indicates the importance of using	
94	reasonable BC density values in the calculation of particle hygroscopicity. In addition,	
95	when estimating the CCN number concentration, a significant bias of –35 % \sim +20 %	
96	was found due to the assumption of particle mixing state (Ren et al., 2018). However,	
97	these studies have not yet accounted for such impact of BC density and mixing state on	
98	CCN prediction due to lack of real time measurement data. Moreover, although the BC	
99	accounts for very small mass fractions (5~10 %) in total fine aerosols, according to our	
100	previous field observed results, the BC-containing particles could contribute 60 %-78 %	
101	toward the total number concentration in urban Beijing (Chen et al., 2020). This is	
102	comparable to the other results using SP2 instrument, which measured that the number	
103	fractions of the coated BC-containing aerosols could be as high as about 50-80 % at the	
104	field sites in north China (Liu et al., 2019b; Zhao et al., 2022). Therefore, the effect of	
105	BC density on the uncertainty of CCN prediction should be concerned carefully,	带格式的:字体颜色:自动设置
106	The mixing state and the density of BC particles are usually directly measured by	
107	several techniques, such as an integrated system of a volatility tandem differential	
108	mobility analyzer and a single particle soot photometer (VTDMA-SP2) (Zhang et al.,	
109	20162016b), or a differential mobility analyzer with a SP2 (DMA-SP2) (Olfert et al.,	
110	2007; Rissler et al., 2014; Wu et al., 2019), and a differential mobility analyzer-	

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111 centrifugal particle analyzer-single-particle soot photometer (DMA-CPMA-SP2) system (Liu et al., 2019b; Yu et al., 2020), etc. However, such techniques or 112 measurements are not available in many previously conducted filed campaigns. In this 113 study, we develop a novel method for retrieving the mixing state and effective density 114 115 of ambient BC particles by combining field measured hygroscopic growth factor and 116 aerosol chemical composition and Köhler theory (Petters and Kreidenweis, 2007). The 117 uncertainty of the new retrieval method was evaluated. The retrieved results were also compared and validated with existing observations. In addition, the effect of BC density 118 119 and mixing state on prediction of CCN number concentrations iswas further evaluated through a sensitivity and closure test by accounting for the retrieved real-time variations 120 of BC density and mixing state. 121

122 2 Field measurements and methodology

123 2.1 Field measurements

Measurements in this study were conducted from 15 November to 14 December 124 125 2016 at a typical urban site of Beijing (39.97°N, 116.37°E, 49 m above sea level). The site locates at the Institute of Atmospheric Physics, Chinese Academy of Sciences, 126 which is mainly influenced by the surrounding cooking, road traffic and residential coal 127 128 burning emissions during the home heating periods (Sun et al., 2016). The detailed information about the sampling site was presented in previous studies (Sun et al., 129 130 20152016; Zhang et al., 2019). The number concentration of condensation nuclei (CN) 131 at each size was measured by a scanning mobility particle sizer, which iswas equipped

132	with a differential mobility analyzer (DMA; model 3081, TSI) and a condensation
133	particle counter (CPC; model 3772, TSI). Subsequently, the mono-dispersed particles
134	were introduced into a Droplet Measurement Technologies CCN counter (CCNc, DMT;
135	Lance et al., 2006) to measure CCN number concentration. A hygroscopic tandem
136	differential mobility analyzer (HTDMA) system was used to measure the hygroscopic
137	growth factor (Gf) (Tan eret al., 2013). Here, four diameters of 40, 80, 110, 150, and
138	200 nm arewere selected in the campaign. Gf is defined as the ratio of the mobility
139	diameter at the given RH to the dry diameter (Petters and Kreidenweis, 2007). The
140	nonrefractory submicron aerosol chemical composition was measured by an Aerodyne
141	high-resolution time-of-flight aerosol mass spectrometer (HR-AMS; Xu et al., 2019),
142	including sulfate, nitrate, ammonium, chloride, and organics. Two factors, including a
143	non-hygroscopic primary organic aerosol (POA) and hygroscopic secondary organic
144	aerosol (SOA) were classified by positive matrix factorization (PMF) with PMF
145	algorithm (v4.2) method (Paatero and Tapper, 1994) and followed the procedures
146	reported in Ulbrich et al. (2009). The refractory black carbon mass loading was
147	measured by an aethalometer (model AE33, Magee Scientific Corporation). Both the
148	nonrefractory materials and BC mass concentration were measured with diameters <
149	$1.0\mu\text{m}.$ The detailed description of the instrument operation and data process have been
150	described in details elsewhere (Ren et al., 2018; Xu et al., 2019; Zhang et al., 2019; Fan
151	et al., 2020).

152 **2.2 Retrieving the mixing state and density of BC**

153 2.2.1 Retrieving the mixing state of BC

154 The Gf probability distribution function (Gf-PDF) for a specified diameter can be 155 retrieved firstly based on the TDMAinv algorithm (Gysel et al., 2009). The κ -PDF can 156 be further calculated based on the Gf-PDF (Fan et al., 2020). Size-resolved κ is derived 157 using κ -Köhler theory based on hygroscopic growth factor (Gf) (Petters and 158 Kreidenweis, 2007),

159
$$\kappa_{gf} = (Gf^3 - 1) \cdot \left[\frac{1}{RH} \exp\left(\frac{4\sigma_{s/a}M_w}{RT\rho_w D_d Gf} \right) - 1 \right]$$
(1)

160 where Gf is hygroscopic growth factor, RH is the relative humidity in the HTDMA 161 (90 %), D_d is the dry diameter, $\sigma_{s/a}$ is assumed to be the surface tension of pure water, 162 *R* is the universal gas constant, *T* is the temperature, M_w and ρ_w is the molecular mass, 163 and the density of water, respectively.

164 The κ -PDF patterns of particles in different sizes always present two modes: nearly 165 hydrophobic (NH) mode with $\kappa_{gf} \leq 0.1$ and more hygroscopic (MH) mode with $\kappa_{gf} >$ 166 0.1 (Fig. S1). Firstly, based on the κ -PDF patterns, the number fraction (NF) of the total 167 nearly hydrophobic group with the boundary of [0, 0.1] was calculated according to the 168 following equation:

169

$$NF = \int_0^{0.1} c(\kappa, D_p) \, d\kappa \tag{2}$$

170 here, the κ -PDF, represented by c (κ , D_p), was normalized as $\int c(\kappa, D_p) d\kappa = 1$, where 171 κ can be replaced by κ_{gf} , D_p is the selected electrical mobility diameter in the campaign. 172 The nearly hydrophobic mode consists of both externally mixed POA (Ex-POA or 8

173	bare POA) and externally mixed BC (Ex-BC). Since the number fraction of the nearly-	
174	hydrophobic POA would change with the emission and aging processes, in this study,	
175	we have applied different values for the number fractions of hydrophobic POA (NH-	
176	POA) under clean (91 %), moderately polluted (70 %), and heavily polluted conditions	
177	(31 %) by referring to the literature (Liu et al., 2021a), as shown in Fig. S2. The number	
178	concentration of Ex-BC was then calculated using the total number fraction of NH	
179	mode minus the number of NH-POA.	
180	$N_{POA-containing} = N_{total} \times NF_{POA-containing}$	
181	$N_{bare-POA} = N_{POA-containing} \times NF_{bare-POA}$	
182	$N_{Ex-BC} = N_{NH} - N_{bare-POA} \tag{3}$	
183	where $N_{\text{POA-containing}}$ and $NF_{\text{POA-containing}}$ are the number concentration and fraction of	
184	POA-containing particles, N_{total} is the total number concentration, $N_{\text{bare-POA}}$ and $NF_{\text{bare-POA}}$	
185	$_{\rm POA}$ are the number concentration and fraction of bare POA particles, and $N_{\rm NH}$ is the	
186	number of nearly hydrophobic group.	
187	The number size distribution of the externally mixed BC (n_{Ex-BC} (log D_p)) can be	
188	calculated based on the particle number size distribution (PNSD) and the number	
189	fraction of the hydrophobic mode of BC (NF_{Ex-BC}) as follows:	
190	$n_{\rm Ex-BC}(\log Dp) = NF_{\rm Ex-BC} \times n \ (\log Dp) \tag{4}$	
191	where $n (\log D_p)$ is the function of the aerosol number size distribution, D_p is the	
192	mobility diameter.	
193	By assuming that the particles arewere spherical (Rader and McMurry, 1986), the	
10/	mass size distribution of F_{x-BC} ($M_{F_{x-BC}}$) was obtained as follows:	

194 mass size distribution of Ex-BC (M_{Ex-BC}) was obtained as follows:

$$M_{\rm Ex-BC}(\log Dp) = \frac{\pi}{6} D_p^{-3} \rho n_{\rm Ex-BC}(\log Dp)$$
(5)

where D_p is the mobility diameter, ρ is the effective density of Ex-BC, and $n_{\text{Ex-BC}}$ (log 196 $D_{\rm p}$) is the function of the number size distribution of Ex-BC, respectively. By reviewing 197 and summarizing the existing results, we show that typical values of density for the 198 199 freshly emitted or externally mixed BC observed in the winter of urban Beijing or North 200 China Plain span over 0.14-0.50 g cm⁻³, with <u>a</u> mean of ~0.40±0.10 g cm⁻³ (Fig. S3), in 201 the size range of 100 to 300 nm, where the mass concentration of externally mixed BC 202 mostly concentrated (Geller et al., 2006; Peng et al., 2016, 2017; Wu et al., 2019; Liu et al., 2020; Zhao et al., 2022). Therefore, an average $\rho_{\text{Ex-BC}}$ of 0.4 g cm⁻³ was used for 203 calculating the mass concentration of externally mixed BC in our study. Uncertainty 204 analyses due to the variations of $\rho_{\text{Ex-BC}}$ were given in section 2.3. 205

The mass size distribution of Ex-BC was fit using the log-normal distribution as shown in Fig. S4 (Wu et al., 2017; Liu et al., 2019a; Zhao et al., 2022). Thus, the bulk mass concentration of Ex-BC can be calculated from the integration of the mass size distribution:

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$$m_{\text{Ex-BC}} = \int_{D_{\text{start}}}^{D_{\text{end}}} M_{\text{Ex-BC}}(\log D_p) \ d \log(D_p) \tag{6}$$

$$m_{\rm In-BC} = m_{\rm BC} - m_{\rm Ex-BC}$$

(7)

where D_{start} and D_{end} are the lower and upper size limit, $M_{\text{Ex-BC}}$ (log D_{p}) is the function of the Ex-BC mass size distribution. We then obtained the bulk mass concentration of internally mixed BC ($m_{\text{In-BC}}$) by subtracting $m_{\text{Ex-BC}}$ from the bulk BC mass concentration measured by AE33 in equation (7-). It should be noted that the mass concentration of BC obtained from AE33 based on aerosol light absorption may lead some uncertainty uncertainties, as has been further addressed in Section 2.3.

218 2.2.2 Retrieving the density of BC

For retrieval of the density of BC, the principal idea is to use the measured κ_{gf} to 219 calculate the density of BC based on the Zdanovskii-Stokes-Robinson (ZSR) mixing 220 221 rule (Stokes and Robinson, 1966; Zdanovskii, 1948) with the chemical composition measured by AMS (Petters & Kreidenweis, 2007). In the retrieval, several aspects are 222 223 concerned. First, since the ZSR rule assumes the aerosol particles are internally mixed, the κ_{gf} value of the more MH mode (κ_{gf-MH}) is thus applied for retrieving the density of 224 internally mixed BC. Second, since the size distribution of BC number concentration is 225 226 usually with peaks between 100 and 200 nm (Liu et al., 2019a; Yu et al., 2020; Zhao et 227 al., 2022), the κ_{gf-MH} value of particles in accumulation mode was averaged and applied 228 for the retrieval. Previous studies showed an independence of κ_{gf-MH} on particle size 229 when the $D_p > 100$ nm during the campaign period (Fan et al., 2020). Therefore, the 230 average of κ_{gf-MH} in accumulation mode is reasonable for the determination of the In-231 BC density, Third, since only one hydrophobic and/or one hygroscopic mode was 带格式的: 字体颜色: 自动设置 232 observed by the HTDMA in most cases during the campaign (Fig. S1, S5), the chemical components of the more hygroscopic (MH) mode at a given diameter should contain 233 234 both these hygroscopic non-BC and the coatings on BC-containing particles, which 235 would be measured by the HR-AMS instrument together. Therefore, by subtracting the externally mixed POA in non-hygroscopic mode (see section 2.3), the concentration 236 and mass fraction of each component measured by HR-AMS can represent the overall 237

238 chemical composition of MH modes, and thus was applied in the ZSR mixing rule for the retrieval of the density of internally mixed BC in this study. In addition, because the 239 inversion including measurements from HTDMA and HR-AMS, a total mass closure 240 of the measured aerosol particles was conducted between the two techniques by 241 242 comparing the mass concentration of PM1 and the results are well consistent (Fig. S6). The density of internally mixed BC (In-BC), $\rho_{\text{In-BC}}$ is then derived from the following 243 equations: 244

245
$$\kappa_{gf-MH} = \kappa_{chem} = \sum_{i} \varepsilon_{i} \kappa_{i} = \frac{v_{inorg}}{v_{total}} \kappa_{inorg} + \frac{v_{SOA}}{v_{total}} \kappa_{SOA} + \frac{v_{In-POA}}{v_{total}} \kappa_{POA} + \frac{v_{In-BC}}{v_{total}} \kappa_{BC}$$
(8)

where κ_{gf-MH} is the hygroscopic parameter of the more hygroscopic (MH) mode, κ_{chem} 246 247 is the hygroscopic parameter of aerosol particles in the mixed composition and can be calculated based on chemical volume fractions using a simple rule (Stokes and 248 Robinson, 1966; Petters & Kreidenweis, 2007), κ_i is the hygroscopic parameter of each 249 pure composition and ε_i is the volume faction of the individual components in the 250 internally mixed particle. Vinorg, VSOA and VIn-POA are the volume of the inorganic, SOA 251 and internally mixed POA species, and can be calculated as follows: $v_{inorg} = \frac{m_{inorg}}{\rho_{inorg}}$ 252 $v_{SOA} = \frac{m_{SOA}}{\rho_{SOA}}$, and $v_{in-POA} = \frac{m_{in-POA}}{\rho_{POA}}$. v_{total} is the total volume of all the species and can be 253 written as $v_{total} = \frac{m_{inorg}}{\rho_{inorg}} + \frac{m_{SOA}}{\rho_{SOA}} + \frac{m_{In-POA}}{\rho_{POA}} + \frac{m_{In-BC}}{\rho_{In-BC}}$. In equation (8), κ_{BC} and κ_{POA} are 254 255 assumed to be 0. So, the total volume v_{total} can be further written as $v_{total} =$ $v_{inorg}\kappa_{inorg}+v_{SOA}\kappa_{SOA}$. The volume of internally mixed v_{In-BC} can be calculated as follows, 256 κ_{gf-MH}

257
$$v_{in-BC} = \frac{v_{inorg}\kappa_{inorg} + v_{SOA}\kappa_{SOA}}{\kappa_{gf-MH}} - v_{inorg} - v_{SOA} - v_{In-POA}$$
258
$$= \frac{\frac{m_{inorg}}{\rho_{inorg}}\kappa_{inorg} + \frac{m_{SOA}}{\rho_{SOA}}\kappa_{SOA}}{\rho_{SOA}} - \frac{m_{inorg}}{\rho_{morg}} - \frac{m_{SOA}}{\rho_{morg}} - \frac{m_{In-POA}}{\rho_{MA}}$$

58
$$= \frac{\frac{m_{inorg}}{\rho_{inorg}}\kappa_{inorg} + \frac{m_{SOA}}{\rho_{SOA}}\kappa_{SOA}}{\kappa_{gf-MH}} - \frac{m_{inorg}}{\rho_{inorg}} - \frac{m_{SOA}}{\rho_{SOA}} - \frac{m_{In-POA}}{\rho_{POA}}$$
(9)

259 Then<u>then</u>, the $\rho_{\text{In-BC}}$ can be calculated based on its mass concentration and volume as 260 follows:

261
$$\rho_{In-BC} = \frac{m_{In-BC}}{(\frac{m_{inorg} \kappa_{inorg} + \frac{m_{SOA}}{\rho_{SOA}} \kappa_{SOA}}{\kappa_{gf-MH}} - \frac{m_{inorg} - m_{SOA} - m_{In-POA}}{\rho_{inorg} - \rho_{SOA} - \rho_{POA}})}$$
(10)

262 where, $m_{\text{In-BC}}$ is the mass concentration of internally mixed BC, m_{inorg} and m_{SOA} are the mass concentrations of the inorganic species and SOA, which are measured by the AMS. 263 mIn-POA is the mass concentrations of internally mixed POA and can be calculated 264 through subtracting the mass fraction of NH-POA from the total mass concentrations 265 of POA. ρ_{inorg} , ρ_{SOA} and ρ_{POA} are the density of the inorganic species, SOA and POA. 266 267 Since the AMS measures the concentrations of the organic and inorganic ions, including SO4²⁻, NO3⁻, NH4⁺, Cl⁻. Here inorganic species were derived by applying a simplified 268 269 ion pairing scheme (Gysel et al., 2007) to convert mass concentrations of ions to the 270 inorganic salts as follows:

$$272 n_{\mathrm{NH}_4\mathrm{NO}_3} = n_{\mathrm{NO}_3^-}$$

273
$$n_{\rm NH_4HSO_4} = \min(2n_{\rm SO_4^{2-}} - n_{\rm NH_4^+} + n_{\rm NO_3^-}, n_{\rm NH_4^+} - n_{\rm NO_3^-})$$

$$n_{(\mathrm{NH}_4)_2\mathrm{SO}_4} = \max(n_{\mathrm{NH}_4^+} - n_{\mathrm{NO}_3^-} - n_{\mathrm{SO}_4^{2-}}, 0)$$

271
$$n_{\text{H}_2\text{SO}_4} = \max(0, n_{\text{SO}_4^{--}} - n_{\text{NH}_4^+} + n_{\text{NO}_3^-})$$
 (11)

where *n* represents the number of moles, and the mass concentrations were obtained by 275 276 the number of moles times the molar mass of each inorganic salts. Because the value of the $n_{\rm H_2SO_4}$ was zero in this campaign. Three inorganic salts including NH₄HSO₄, 277 278 (NH₄)₂SO₄, and NH₄NO₃ were applied in our study. The densities for inorganic salts 279 were taken from previous studies (Gysel et al., 2007; Wu et al., 2016). Here the densities for three inorganics are 1.78, 1.77 and 1.72 g cm⁻³, respectively. By summarizing the 280 281 previous studies (Gysel et al., 2007; Dinar et al., 2006), 1.4 g cm⁻³ was selected as the 13

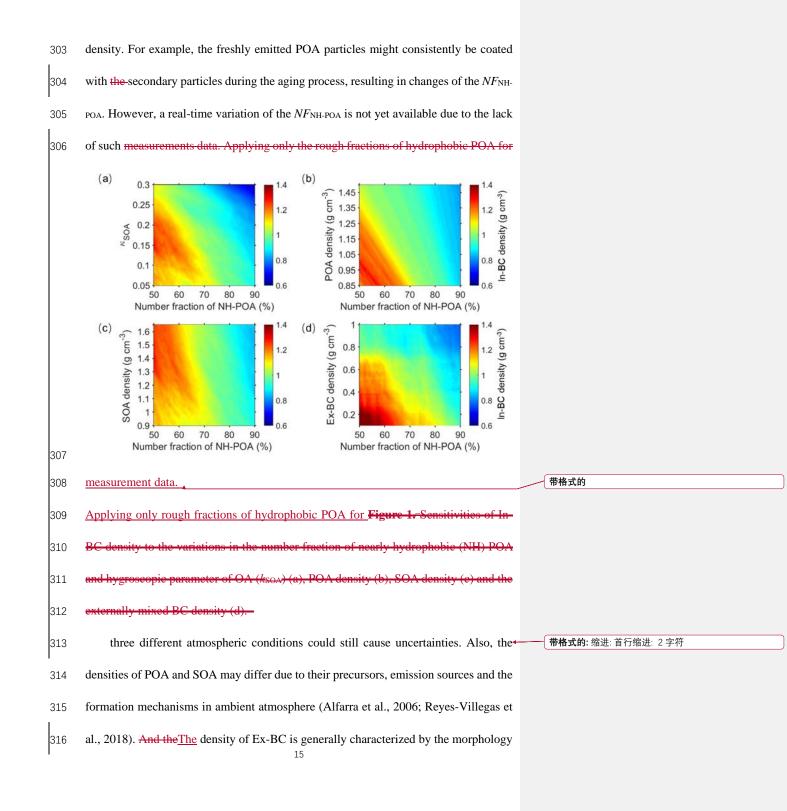
282	density of SOA (ρ_{SOA}). The density of POA (ρ_{POA}) is assumed to be 1.0 g cm ⁻³ for urban
283	environments, which is similar to the that of the lubricating oil (Wu et al., 2016). Since
284	the cooking organic aerosols represent a high contribution to POA in urban
285	environments, we choose the mean density of the rapeseed oil and oleic acid (~0.85 g $$
286	cm ⁻³) (Reyes-Villegas et al., 2018) to evaluated the result as shown in section 2.3. The
287	values of κ for inorganic components are 0.56 for NH ₄ HSO ₄ , 0.48 for (NH ₄) ₂ SO ₄ and
288	$0.58 \mbox{ for NH}_4 NO_3,$ along with the best-fit values for the three inorganic salts (Petters &
289	Kreidenweis, 2007 and Gunthe et al., 2009). The κ_{SOA} is assumed to be 0.15 according
290	to the field studies in urban areas (Chang et al., 2010; Kawana et al., 2016).

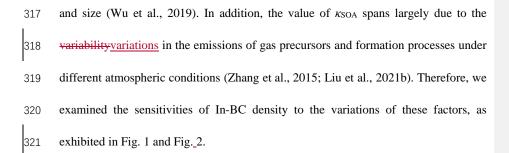
Note that the method fails to retrieve the BC density when organic accounts for a large fraction (>60 %). This is because that a higher fraction of OA usually corresponds to lower total volume of all the species (Fig. S7), yielding negative values for v_{In-BC} introduced in equation (9-). As a result, 61 % of the data observed during the campaign were valid for calculating the BC density.

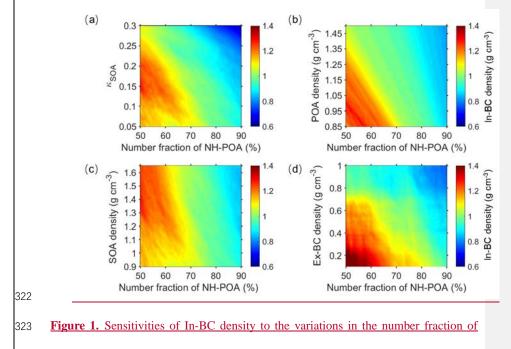
296 Similarly, the bulk density of BC ($\rho_{bulk-BC}$) is calculated with the same method as 297 that for calculating the ρ_{In-BC} . When calculating the $\rho_{bulk-BC}$, the bulk κ_{gf} value measured 298 by HTDMA is applied assuming that all the aerosol particles are internally mixed.

299 2.3 Uncertainties and limitations

300 For the retrieval, the assumptions on the values of κ_{SOA} , ρ_{POA} , ρ_{SOA} and ρ_{Ex-BC} as 301 well as the fraction of primary organic aerosols in non-hygroscopic or hygroscopic 302 mode would add uncertainty in the inferred values of ambient internally mixed BC

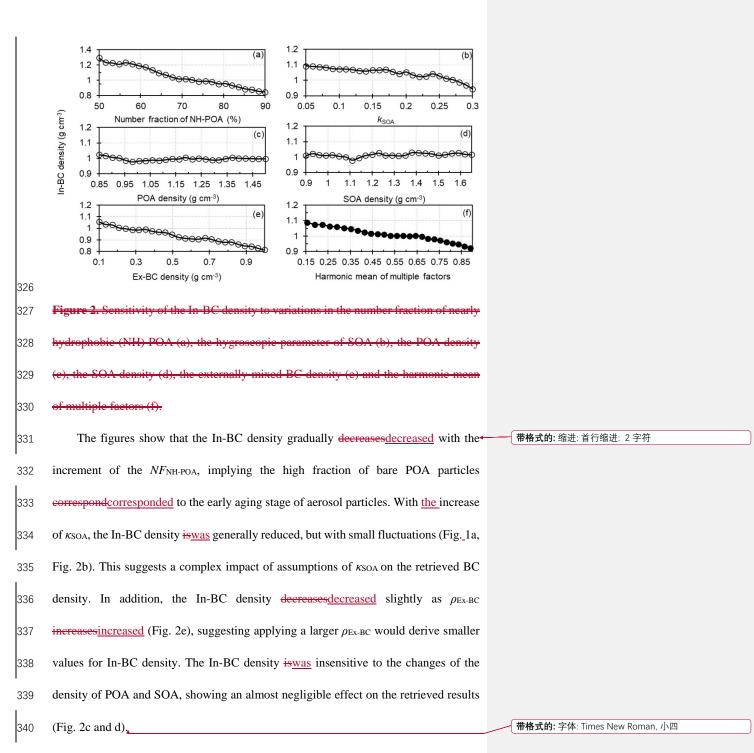






324 nearly hydrophobic (NH) POA and hygroscopic parameter of OA (ksoa) (a), POA

325 <u>density (b), SOA density (c) and the externally mixed BC density (d).</u>



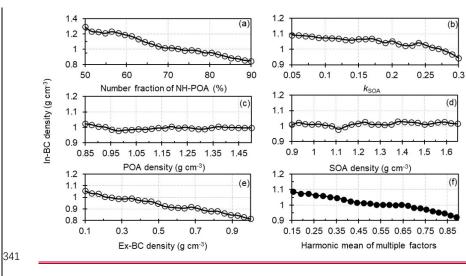
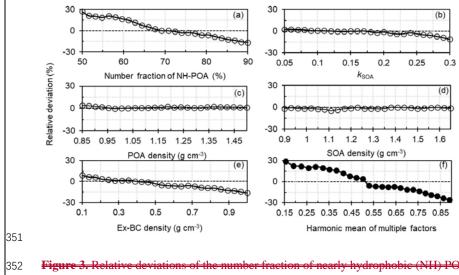


Figure 2. Sensitivity of the In-BC density to variations in the number fraction of nearly
hydrophobic (NH) POA (a), the hygroscopic parameter of SOA (b), the POA density
(c), the SOA density (d), the externally mixed BC density (e) and the harmonic mean
of multiple factors (f).
The uncertainty analysis shows that, by comparing the results based on the mean

fractions<u>fraction</u> of the *NF*_{NH-POA} with a typical atmospheric observed range of 50-90 % for the *NF*_{NH-POA} (Liu et al., 2021a), we show that the assumption on *NF*_{NH-POA} can lead to relative deviations (uncertainty) of -17 %~+27 % for the retrieved BC density (Fig. 3a).

18



to the In BC density (a), the hygroscopic parameter of OA to the In BC density (b), the
 POA density to the In BC density (c), the SOA density to the In BC density (d), the
 externally mixed BC density to In BC density (c) and the combined deviations based
 on multiple factors mentioned above (f).

In addition, unlike inorganics (eg., NH4HSO4, (NH4)2SO4 and NH4NO3), for which 357 the hygroscopicity has been already well-understood (Petters and Kreidenweis, 2007), 358 the hygroscopicity of organic species varies largely due to the complexity in organic 359 aerosol constituents. Therefore, the assumption of the values of κ_{SOA} will add the 360 361 uncertainty in the calculation of BC density. Previous studies have suggested that the organics hashave a wide range of k values ranging from 0.05 to 0.3 (Jimenez et al., 2009; 362 363 Mei et al., 2013). Thus, the sensitivity test has also been done to examine the effect due to changes in KSOA on calculating the density of BC (Fig. 1a). The result shows that the 364 365 assumption of κ_{SOA} values can cause an average relative deviation of -10 % ~ +3 %

366 in calculating the density of In-BC (Fig. 3b).

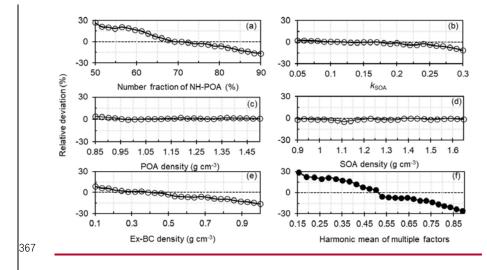


Figure 3. Relative deviations of the number fraction of nearly hydrophobic (NH) POA
to the In-BC density (a), the hygroscopic parameter of OA to the In-BC density (b), the
POA density to the In-BC density (c), the SOA density to the In-BC density (d), the
externally mixed BC density to In-BC density (e) and the combined deviations based
on multiple factors mentioned above (f).
However, the sensitivity test shows that the impact of both the *ρ*POA and *ρ*SOA

variations on the BC density estimation iswas very small or even negligible (Fig. 1b, c). By varying the ρ_{POA} from 0.85 to 1.5 g cm⁻³ and the ρ_{SOA} from 0.9 to 1.65 g cm⁻³ according to the literatures literature (Noureddini et al., 1992; Alfarra et al., 2006; Reyes-Villegas et al., 2018; Cai et al., 2020; Kostenidou et al., 2007), the retrieval uncertainties in the BC density arewere both within ±5 % (Fig. 3c, d). For ρ_{Ex-BC} , it exhibits exhibited that the evolution of the ρ_{Ex-BC} could lead to an average deviation of -16 %~+9 % in calculating In-BC density (Fig. 3e) when increasing the values of ρ_{Ex-BC} BC from 0.1 to 1.0 g cm⁻³, which represents a typical range in ambient atmosphere (Wu et al., 2019; Liu et al., 2020). A combined uncertainty (δ) caused by the multiple factors (δ_i), which iswas calculated by equation (12, is), was -26 %~+29 % as shown in Fig. 384 3f.

$$\delta = \sqrt{\sum_{i=1}^{n} {\delta_i}^2} \tag{12}$$

386 In addition, it should be noted that the mass concentration of BC obtained from 387 AE33 based on aerosol light absorption may lead some uncertainty.uncertainties. 388 However, the comparison of the simultaneously measured data by SP2 with thatthose 389 by AE33 during the campaign shows that the temporal variations of BC mass 390 concentrations measured by the two techniques arewere well consistent (Fig-S5. S8). Note that the BC mass measured by SP2 is occasionally low probably because of the 391 low detection efficiency in small size (McMeeking et al., 2010; Schwarz et al., 2006). 392 In addition, the SP2 is unable to quantify the BC mass beyond a certain limit because 393 394 of the saturation of electronic devices recording signals (Pileci et al., 2021). We show 395 that, compared the results that were retrieved if applying the BC mass measured by SP2, 396 the BC density retrieved based on AE33 can be 18_% higher. Given that the measurement bias from SP2, this overestimation indicates an upper limit of the 397 398 uncertainty.

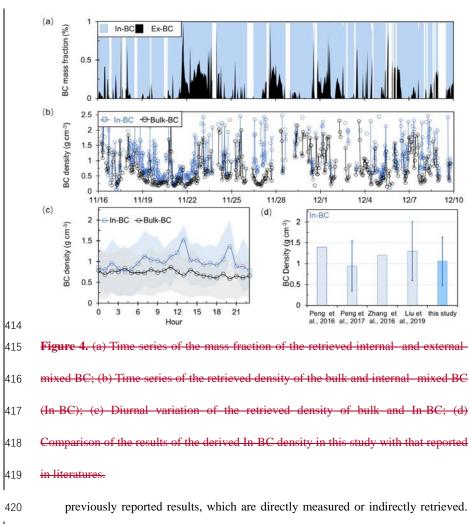
399 3 Results and Discussion

400 **3.1 Retrieved A comparison and validation of retrieved mixing state and density of**

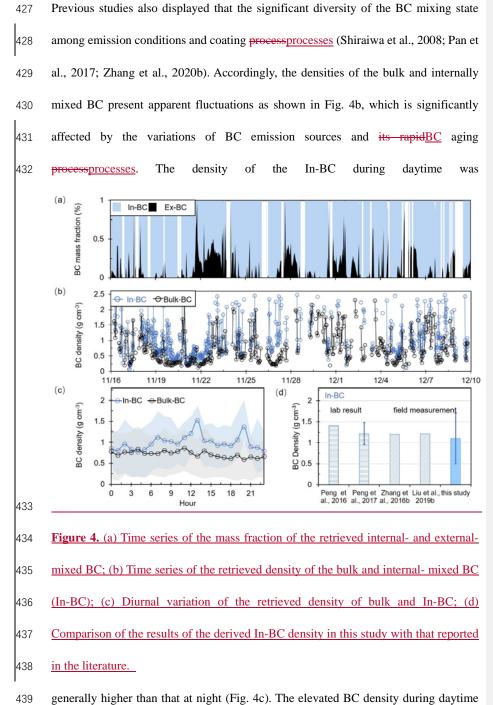
401 BC: comparison and validation

402 Figure 4a shows retrieved time series of the mixing state of ambient BC during the 403 campaign. Large temporal variations of the mass fraction of internally and externally 404 mixed BC are presented present during the observed period at the sites. The temporal changes should be related to the atmospheric aging process or diurnal variations of 405 406 emissions (Liu et al., 2019a; Fan et al., 2020). Statistically, the average mass 407 fraction fractions of externally and internally mixed BC is were 20±18 % and 80±20 % 408 respectively, showing that most of the BC particles were aged and internally mixed with 409 other components. Previous studies at urban sites have shown that the co-existence of the externally mixed BC in the ambient atmosphere (Schwarz et al., 2008; Cheng et al., 410 411 2012; Chen et al., 2020) due to continuous combustion processes (e.g., vehicle exhaust 412 and residential sector) (Wang et al., 2017; Liu et al., 2019a). Our results are basically 413 comparable with those

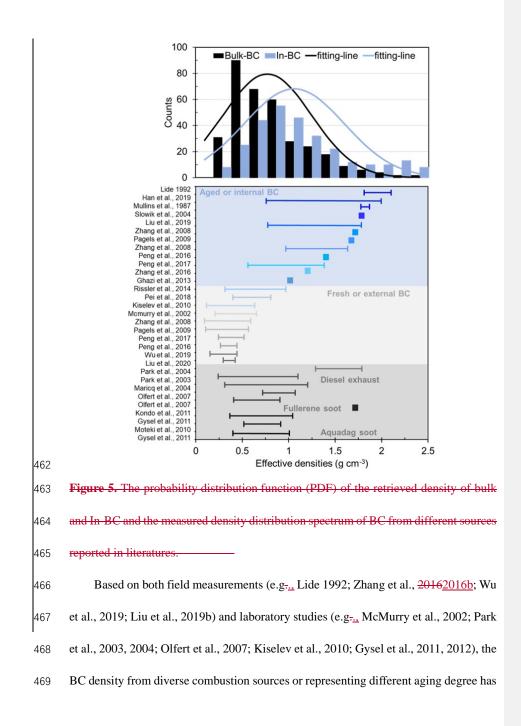
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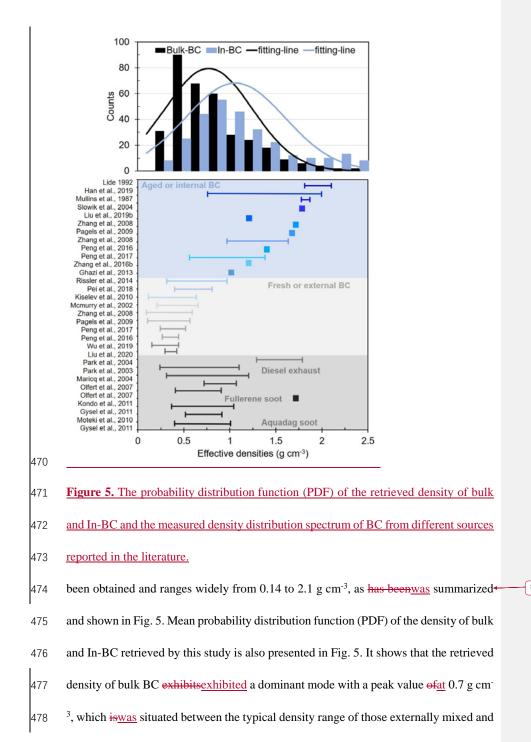


For example, Chen et al., (2020) found that the mass fraction of internally mixed BC particles was nearly ~80–90 % in summer of Beijing based on VTDMA measurements. Liu et al. (2020), using a tandem system with an aerodynamic aerosol classifier and SP2, reported that the mass fraction of internally BC-containing particles would increase with increasing size and reach ~70 % in Beijing. Overall, the mass fraction obtained in our study iswas comparable with thosethat reported in urban Beijing.



1	
440	iswas likely due to that the strong photochemical processes promote the aging of BC
441	particles, which resulted in a conversion from uncompacted structure to compact and
442	regular spherical shapes of BC (Qiao et al., 2018; Liu et al., 2019b; Zhou et al., 2022).).
443	The liftrising in BC density around 20:00 LT might indicate that the BC particles would
444	be rapidly coated with the secondary inorganic aerosol (SIA) particles and continuously
445	aged in the polluted period due to the heterogeneous reactions of SIA in urban regions
446	(Zhang et al., 20162016b; Peng et al., 2017). Actually, following the haze evolution, the
447	fraction of nearly hydrophobic group reduced rapidly (Fig. <u>\$859</u>). Consequently, the
448	average density of In-BC increased obviously from the clean conditions to the polluted
449	periods (Fig. <u>\$9\$5</u>). A slight decrease was observed in the bulk BC density during
450	traffic hours. This is likely associated with the continuescontinuous emissions (e.g.,
451	vehicle exhaust) that lead to uncoated or uncompacted BC particles in this period. The
452	diurnal cycle in In-BC density iswas consistent with the coating thickness measured by
453	a tandem CPMA-SP2-DMA-SP2 (Liu et al., 2020), demonstrating that the new method
454	can derive the density of ambient BC particles reasonably. Averagely, the campaign
455	average values of the bulk and internally mixed BC densities are with campaign
456	averaged values of were 0.7 \pm 0.5 and 1.1 \pm 0.6 g cm ⁻³ respectively, which are were much
457	less than 1.8 g cm ⁻³ , implying that the BC particles is are not a-void-free spheres in the
458	urban atmosphere. The results of In-BC density arewere comparable with thatthose
459	observed at the other sites in North China Plain (NCP) as shown in Fig. 4d, illustrating
460	that the BC effective density retrieved by this method $\frac{1}{100}$ within the range of $\frac{1}{100}$
461	from field measurements.





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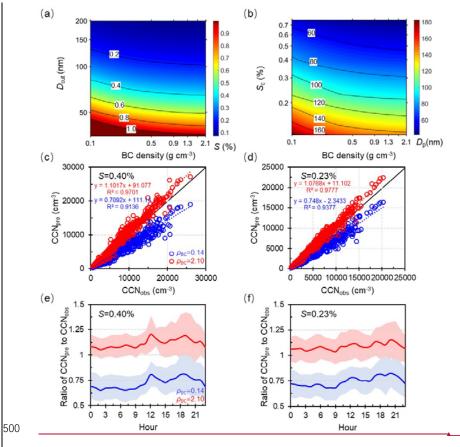
479	internally mixed BC measured previously. For the In-BC, the PDF iswas with a peak
480	value at 1.1 g cm ⁻³ , but ranges ranged widely from ~0.5 to 2.5 g cm ⁻³ , which
481	indicates indicated various morphologies, different aging degree and compositions of
482	ambient BC particles due to the complex impact of multiple local sources and aging
483	processes during the observed period in urban Beijing. Overall, the retrieved values for
484	In-BC density fall within the range of typical internalinternally mixed BC reported in
485	the literatures literature, verifying the reliability of our inversion results.

486 **3.2 Sensitivity of predicted** *N*_{CCN} **to changes of BC density**

487 A previous study showed that the use of an inaccurate density value of BC particles 488 would result in large <u>biasbiases</u> in estimating κ of ambient aerosol particles with the 489 ZSR mixing rule (Fan et al., 2020), as would further lead to uncertainties in prediction 490 of *N*_{CCN} and relevant climate effects. Considering the large variation range of BC 491 density

492	during the campaign, which is closely associated with <u>itsBC</u> morphology or degree
493	of \underline{itsBC} aging, we further examine the sensitivity of critical supersaturation (S _c),
494	critical diameter (D_{cut}) and predicted N_{CCN} to variations of BC density (Fig. 6). Here,
495	we use the critical diameter and particle number size distribution to calculate $N_{\rm CCN}$. The
496	method The method to derive the critical diameter is based on Köhler theory and ZSR
497	rule. Three CCN closure studies were assumed to evaluate the effect of BC density and
498	mixing state on prediction of CCN number concentrations. Closure studies provide a
499	useful way to

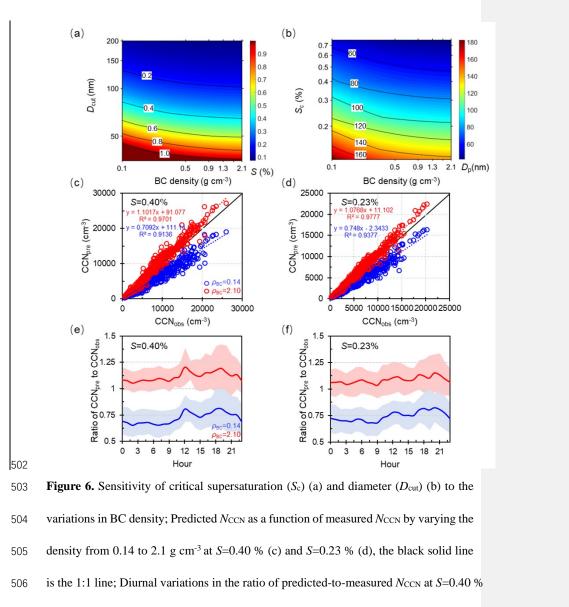
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501 to derive the critical diameter is based on Köhler theory and ZSR rule.



507 (e) and *S*=0.23 % (f).

I

508	investigate the importance of aerosol properties to CCN concentration prediction. If the 带格式的 : 缩进: 首行	缩进: 0 字符
509	2 closure study is achieved, it means the bias between the predicted and measured CCN	
510	concentrations is within ±15 % (Chang et al., 2007). The results show that detailed 带格式的	

511	calculation methods are presented in the supporting information (SI: Methods) or the
512	reference in Ren et al. (2018). The results show that, by varying the value of density
513	from 0.14 to 2.1 g cm ⁻³ that represents the lower and upper limitrange of BC density in
514	the atmosphere, the $D_{\text{cut}} \frac{\text{reduces}}{\text{reduced}}$ apparently at a given supersaturation (S) (Fig.
515	6a), or similarly, the S_c decreases decreased rapidly for a given particle size (Fig. 6b).
516	The results show that the changes of the D_{cut} and S_c are were more sensitive when the
517	BC density is was below 1.0 g cm ⁻³ . And the The effects on the D_{cut} and S_c both gradually
518	weakened with the increase of BC density. This shows that it is critical to apply more
519	accurate BC density for the aerosol particles with low aging degree in predicting CCN
520	and its climate effect. Accordingly, the ratios of predicted-to-measured $N_{\rm CCN}$ ranged
521	from 0.72 to 1.11 by varying the BC density from 0.14 to 2.1 g cm ⁻³ at the typical S of
522	0.23 % and 0.40 % (Fig. 6c, 6d), showing an estimation uncertainty of -28 % \sim 11 %
523	in <i>N</i> _{CCN} prediction.

The diurnal variations in the ratio of predicted-to-measured N_{CCN} at S=0.40 % and 524 0.23 % are shown to examine the response of the BC density on N_{CCN} prediction at 525 526 different time periods (Fig. 6e, 6f). By applying the lower limit of density value of 0.14 g cm⁻³, the prediction iswas much worse compared tothan the use of the density of 2.1 527 528 g cm⁻³ at nighttimenight (00:00-06:00 LT), when the latter iswas much closer to the real 529 density of ambient BC (Fig. 4c). The prediction iswas improved substantially by 530 applying the value of 0.14 g cm⁻³ during evening rush hours (18:00-20:00 LT), during 531 which the ambient BC particles iswere disturbed by the traffic emissions (Fig. 4c). And now, the <u>The</u> prediction <u>becomes</u> became worse by applying the value of 2.1 g cm⁻³, and 532

an obvious overestimation by up to ~40 % iswas shown. The results further illustrate that it is critical to account for the real-time mixing state and density of BC particles in $N_{\rm CCN}$ prediction, particularly in those-regions with heavy traffic and residential coal emissions.

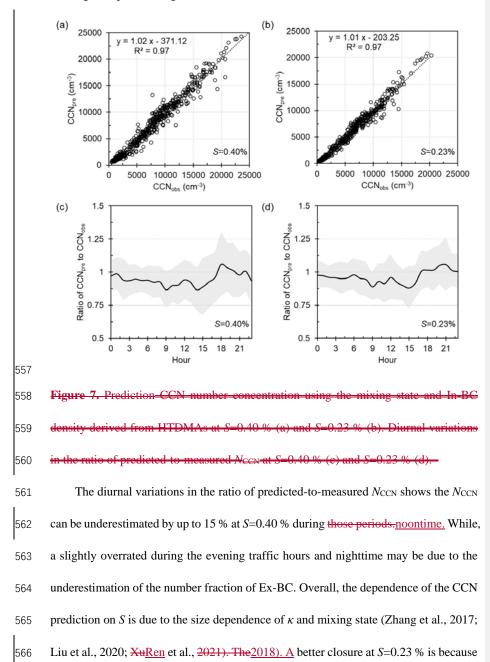
537 It should be noted that the assumption of the surface tension of water would overestimate the critical diameter and underpredict CCN number concentration. While 538 539 the surface tension depression might be more obvious for the small size particles (<60 540 nm), as the fraction of organics areis higher at small particles size (Meng et al., 2014; Cai et al., 2018). Here, in this study, we calculated the critical diameters at 541 supersaturations of 0.40 % and 0.23 %, typical values in cloud, corresponding to larger 542 sizes (> 70 nm and 90 nm) of aerosols. Therefore, the uncertainties from the application 543 544 of the surface tension of pure water should be negligible (< 10 %). Here, three schemes 545 were assumed to evaluate the effect of BC density and mixing state on prediction of 546 CCN number concentrations. The detailed calculation methods are presented in the supporting information (SI: Methods) or referenced from Ren et al. (2018). 547

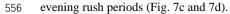
3.3 UsingN_{CCN} prediction based on the real-time variations of BC density and mixing state to predict N_{CCN}

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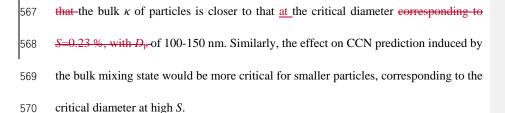
Figure 7 exhibits the comparisons between predicted and measured $N_{\rm CCN}$ at *S* of 0.23 % and 0.40 % by accounting for the retrieved real-time variations of BC density and mixing state. It shows that the $N_{\rm CCN}$ can be well predicted with a slope of 1.01 and 1.02 at *S* of 0.23 % and 0.40 % respectively (Fig. 7a, 7b), only presenting a slight deviation. The slight deviation is primarily due to the fixed value of the density for the 32

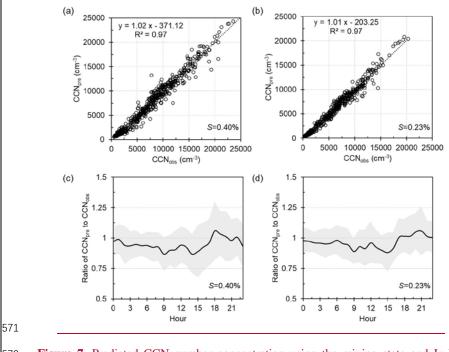
555 externally mixed BC caused by the retrieved method, especially during noontime and





³³





572 **Figure 7.** Predicted CCN number concentration using the mixing state and In-BC

573 <u>density derived from HTDMAs at S=0.40 % (a) and S=0.23 % (b). Diurnal variations</u>

574 <u>in the ratio of predicted-to-measured N_{CCN} at S=0.40 % (c) and S=0.23 % (d).</u>

575 Overall, when considering the effective density of BC relevant to its mixing state, 576 the CCN closure achieves. Previous studies have shown that the fresh emitted BC 577 particles may convert from fractal-like aggregates to a compact structure and its density 578 would increase with the aging process (Pagels et al., 2009; Rissler et al., 2014; Peng et 579 al., 2016; Liu et al., 2019b; Zhang et al., 2020a, 2022), but the actual density of In-BC 34

580	may be lower than 1.8 g $\rm cm^{\text{-}3}$ in the ambient atmosphere according to this study.
581	Therefore, the currently applied value represents a density of the void-free structure of
582	BC particles may cause an overestimation in CCN prediction. In addition, although the
583	BC accounts for small mass fractions in ambient fine aerosols, according to the
584	measurements simultaneously conducted at the site, the BC-containing particles could
585	contribute 60 %-78 % toward the total number concentration in urban Beijing (Chen et
586	al., 2020). Our results further highlight the effect of BC density on the uncertainty of
587	CCN prediction should be concerned carefully,

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588 4 Conclusions

The mixing state and effective density of BC changed through heterogenous 589 chemistry process and thus would cause uncertainty in evaluating its CCN activity. In 590 this study, we develop a new method to retrieve the mixing state and effective density 591 of ambient BC using field measurements and the Köhler theory. The uncertainty of the 592 new retrieval method was evaluated within ±30 %, which iswas primarily caused by 593 assuming the value of κ_{SOA} and the fraction of primary organic aerosols in non-594 595 hygroscopic mode. The retrieved results show that most of the BC particles were aged and internally mixed with other components, with mean mass fraction of 80±20 %. 596 597 Averagely, the retrieved densities of the bulk and internally mixed BC arewere 0.7±0.5 and 1.1±0.6 g cm⁻³ respectively, but rangesranged widely from ~0.1 to 2.5 g cm⁻³, 598 indicating various morphologies, different aging degree and compositions of ambient 599 BC particles due to the complex impact of multiple local sources and aging processes 600

601 during the observed period. The retrieved results are result was basically comparable with-the previous observations in North China Plain. 602 Further examination shows the uncertainties of the NCCN prediction is with 603 uncertainties of were, -28 % ~11 % at the typical S of 0.23 % and 0.40 % by varying the 604 BC density from 0.14 to 2.1 g cm⁻³ that represents represented the lower and upper 605 limitrange of ambient BC particles. Moreover, the prediction iswas found more 606 607 sensitive to the variability variations of BC density when it is was <1.0 g cm⁻³, suggesting a great significance to account of accounting for the effect of BC density for the aerosol 608 particles with low aging degree when evaluating the climate effect. The CCN closure 609 610 achievesachieved when introducing the retrieved real-time BC density relevant to its mixing state. This work provides a unique way of utilizing field observations to infer 611 612 ambient BC density and highlights the current assumption of a void-free structure of BC particles in models would cause large uncertainties in CCN prediction and in the 613 relevant climate effect evaluation. 614

The method used to derive the ambient BC density has limitations. Since the 615 616 assumptions on the values of κ_{SOA} , ρ_{POA} , ρ_{SOA} and ρ_{Ex-BC} as well as the fraction of primary organic aerosols in non-hygroscopic or hygroscopic mode would add 617 618 uncertainty in the inferred values of ambient internally mixed BC density. It is thus 619 necessary to examine observational data to verify this methodology in furtherfuture 620 studies. However, the method and results of this study could provide the way for a more 621 comprehensive understanding of the variability variations in BC density in Beijing. Additionally, it has the potential to reveal the uncertainties of usage of void-free 622

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623 structure of BC density in accessingassessing the climate effects.

624 Data availability.

- All data needed to evaluate the conclusions in the paper are present in the paper and/or
- 626 the Supplement. All data used in the study are also available from the corresponding
- 627 author upon request (zhangfang2021@hit.edu.cn).

628 Author contributions.

- 629 FZ and JR conceived the conceptual development of the manuscript. JR directed and
- 630 performed of the experiments with JL, LC, and FZ. JR conducted the data analysis and
- 631 wrote the draft of the manuscript. All authors edited and commented on the various
- 632 sections of the manuscript.

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639 **Competing interests.**

- 640 The contact author has declared that neither they nor their co-authors have any
- 641 competing interests.

642 **References**

- Alfarra, M. R., Paulsen, D., Gysel, M., Garforth, A. A., Dommen, J., Prévôt, A. S. H.,
 Worsnop, D. R., Baltensperger, U., and Coe, H.: A mass spectrometric study of
 secondary organic aerosols formed from the photooxidation of anthropogenic and
 biogenic precursors in a reaction chamber, Atmos. Chem. Phys., 6, 5279– 5293,
 https://doi:10.5194/acp-6-5279-2006, 2006.
- Bond, T. C., Doherty, S. J., Fahey, D., Forster, P., Berntsen, T., DeAngelo, B., Flanner,
 M., Ghan, S., Kärcher, B., and Koch, D.: Bounding the role of black carbon in the
 climate system: A scientific assessment, J. Geophys. Res.-Atmos., 118(11), 5380–
 5552, https://doi.org/10.1002/jgrd.50171, 2013
- Cheng, Y. F., Eichler, H., Wiedensohler, A., Heintzenberg, J., Zhang, Y. H., Hu, M.,
 Herrmann, H., Zeng, L. M., Liu, S., Gnauk, T., Brüggemann, E., and He, L. Y.:
 Mixing state of elemental carbon and non-light-absorbing aerosol components
 derived from in situ particle optical properties at Xinken in Pearl River Delta of China,
 J. Geophys. Res., 111, D20204, doi:10.1029/2005JD006929, 2006.
- Clarke, A.D., Shinozuka, Y., Kapustin, V.N., Howell, S., Huebert, B., Doherty, S.,
 Anderson, T., Covert, D., Anderson, J., Hua, X., Moore II, K.G., McNaughton, C.,
 Carmichael, G., Weber, R.: Size distributions and mixtures of dust and black carbon
 aerosol in Asian outflow: physiochemistry and optical properties, J. Geophys. Res.Atmos., 109, D15S09, https://doi.org/10.1029/2003JD004378, 2004.
- Cheng, Y. F., Su, H., Rose, D., Gunthe, S. S., Berghof, M., Wehner, B., Achtert, P.,
 Nowak, A., Takegawa, N., Kondo, Y., Shiraiwa, M., Gong, Y. G., Shao, M., Hu, M.,
 Zhu, T., Zhang, Y. H., Carmichael, G. R., Wiedensohler, A., Andreae, M. O., and
 Pöschl, U.: Size-resolved measurement of the mixing state of soot in the megacity
 Beijing, China: diurnal cycle, aging and parameterization, Atmos. Chem. Phys., 12,
 4477–4491, https://doi.org/10.5194/acp-12-4477-2012, 2012.
- Cheng, Y. F., Eichler, H., Wiedensohler, A., Heintzenberg, J., Zhang, Y. H., Hu, M.,
 Herrmann, H., Zeng, L. M., Liu, S., Gnauk, T., Brüggemann, E., and He, L. Y.:
 Mixing state of elemental earbon and non light absorbing aerosol components
 derived from in situ particle optical properties at Xinken in Pearl River Delta of China.
 J. Geophys. Res., 111, D20204, doi:10.1029/2005JD006929, 2006.
- Chen, L., F. Zhang, P. Yan, X. Wang, L. Sun, Y. Li, X. Zhang, Y. Sun, and Z. Li.: The
 large proportion of black carbon (BC)-containing aerosols in the urban atmosphere,

带格式的:超链接,字体:+西文正文 (等线),五号

675 676 677 678 679 680 681 682	 Environ. Pollut., 263, 114507, https://doi.org/10.1016/j.envpol.2020.114507, 2020. Chang, R. YW., Slowik, J. G., Shantz, N. C., Vlasenko, A., Liggio, J., Sjostedt, S. J., Leaitch, W. R., and Abbatt, J. P. D.: The hygroscopicity parameter (k) of ambient organic aerosol at a field site subject to biogenic and anthropogenic influences: relationship to degree of aerosol oxidation, Atmos. Chem. Phys., 10, 5047–5064, https://doi:10.5194/acp-10-5047-2010, 2010. Chang, R. Y. W., Liu, P. S. K., Leaitch, W. R., and Abbatt, J. P. D.: Comparison between measured and predicted CCN concentrations at Egbert, Ontario: Focus on 	<	带格式的
683	the organic aerosol fraction at a semirural site, Atmos. Environ., 41, 8172-8182,		
684	https://doi:10.1016/j.atmosenv.2007.06.039, 2007.		
685	Cai, M., Tan, H., Chan, C. K., Qin, Y., Xu, H., Li, F., Schurman, M. I., Liu, L., and Zhao,		
686	J.: The size-resolved cloud condensation nuclei (CCN) activity and its prediction		
687	based on aerosol hygroscopicity and composition in the Pearl Delta River (PRD) region during wintertime 2014, Atmos. Chem. Phys., 18, 16419–16437,		
688 689	https://doi.org/10.5194/acp-18-16419-2018, 2018.		
690	Cai, J., Chu, B., Yao, L., Yan, C., Heikkinen, L. M., Zheng, F., Li, C., Fan, X., Zhang,		
691	S., Yang, D., Wang, Y., Kokkonen, T. V., Chan, T., Zhou, Y., Dada, L., Liu, Y., He,		
692	H., Paasonen, P., Kujansuu, J. T., Petäjä, T., Mohr, C., Kangasluoma, J., Bianchi, F.,		
693	Sun, Y., Croteau, PL., Worsnop, D. R., Kerminen, V. M., Du, W., Kulmala, M., and		带格式的: 超链接, 字体: +西文正文 (等线), 五号
694	Daellenbach, K. R.: Size segregated particle number and mass concentrations from	\leftarrow	带格式的
695	different emission sources in urban Beijing, Atmos. Chem. Phys., 20, 12721-12740,		带格式的
696	https://doi.org/10.5194/acp-20-12721-2020, 2020.		带格式的: 超链接, 字体: +西文正文 (等线), 五号
697	Dinar, E., Mentel, T. F., and Rudich, Y.: The density of humic acids and humic like		
698	substances (HULIS) from fresh and aged wood burning and pollution acrosol		
699	particles, Atmos. Chem. Phys., 6, 5213-5224, doi:10.5194/acp-6-5213-2006, 2006,		带格式的: 超链接, 字体: +西文正文 (等线), 五号
700	Dameto de España, C., Wonaschütz, A., Steiner, G., Rosati, B., Demattio, A., Schuh,		
701	H., and Hitzenberger, R.: Long-term quantitative field study of New Particle		
702	Formation (NPF) events as a source of Cloud Condensation Nuclei (CCN) in the		
703	urban background of Vienna, Atmos. Environ., 164, 289–298,		
704	https://doi.org/10.1016/j.atmosenv.2017.06.001, 2017.		(带格式的: 下划线
705	Dinar, E., Mentel, T. F., and Rudich, Y.: The density of humic acids and humic like		
706	substances (HULIS) from fresh and aged wood burning and pollution aerosol		
707	particles, Atmos. Chem. Phys., 6, 5213–5224, doi:10.5194/acp-6-5213-2006, 2006,		带格式的:超链接,字体:+西文正文(等线),五号
708	Flanner, M. G., Zender, C. S., Randerson, J. T., and Rasch, P. J.: Present-day climate		
709	forcing and response from black carbon in snow, J. Geophys. ResAtmos., 112,		
710	D11202, https://doi.org/10.1029/2006JD008003, 2007.		
711	Fan, X., Liu, J., Zhang, F., Chen, L., Conllins, D., Xu, W., Jin, X., Ren, J., Wang, Y., Wu, H., Li, S., Sun, Y., Li, Z.: Contrasting size-resolved hygroscopicity of fine particles		
712 713	derived by HTDMA and HR-ToF-AMS measurements between summer and winter		
713 714	in Beijing: the impacts of aerosol aging and local emissions, Atmos. Chem. Phys. 20,		
714	915-929, https://doi.org/10.5194/acp-20-915-2020, 2020,	_	带格式的: 下划线
716	Gysel, M., McFiggans, G. B., and Coe, H.: Inversion of tandem differential mobility		带格式的
710	analyser (TDMA) measurements, J. Aerosol Sci., 40, 134–151,		
718	https://doi.org/10.1016/j.jaerosci.2008.07.013, 2009.	_	带格式的: 超链接, 字体: +西文正文 (等线), 五号
1. 70		_)

719	Geller, M., Biswas, S., and Sioutas, C.: Determination of particle effective density in	
720	urban environments with a differential mobility analyzer and aerosol particle mass	
721	analyzer, Aerosol Sci. Technol., 40, 709–723,	
722	https://doi.org/10.1080/02786820600803925, 2006.	
723	Gysel, M., McFiggans, G. B., and Coe, H.: Inversion of tandem differential mobility	带格式的
724	analyser (TDMA) measurements, J. Aerosol Sci., 40, 134-151,	
725	https://doi.org/10.1016/j.jaerosei.2008.07.013, 2009.	带格式的:
726	Gysel, M., Crosier, J., Topping, D. O., Whitehead, J. D., Bower, K. N., Cubison, M. J.,	带格式的
727	Williams, P. I., Flynn, M. J., McFiggans, G. B., and Coe, H.: Closure study between	
728	chemical composition and hygroscopic growth of aerosol particles during TORCH2,	
729	Atmos. Chem. Phys., 7, 6131–6144, https://doi.org/10.5194/acp-7-6131-2007, 2007.	
730	Gunthe, S. S., King, S. M., Rose, D., Chen, Q., Roldin, P., Farmer, D. K., Jimenez, J.	
731	L., Artaxo, P., Andreae, M. O., Martin, S. T., and Pöschl, U.: Cloud condensation	
732	nuclei in pristine tropical rainforest air of Amazonia: size resolved measurements and	
733	modeling of atmospheric aerosol composition and CCN activity, Atmos. Chem. Phys.,	
734	9, 7551–7575, https://doi.org/10.5194/acp-9-7551-2009, 2009.	
735	Gysel, M., Laborde, M., Olfert, J. S., Subramanian, R., & Gröhn, A. J.: Effective density	
736	of aquadag and fullerene soot black carbon reference materials used for SP2	
737	calibration, Atmos. Meas. Tech., 4(12), 4937-4955, https://doi.org/10.5194/amt-4-	
738	2851-2011, 2011.	
739	Gysel, M., Laborde, M., Mensah, A. A., Corbin, J. C., Keller, A., Kim, J., et al.:	
740	Technical note: The single particle soot photometer fails to reliably detect PALAS	
741	soot nanoparticles, Atmos. Meas. Tech., 5(12), 3099-3107,	
742	https://doi.org/10.5194/amt-5-3099-2012, 2012.	
743	Jimenez, J. L., Canagaratna, M. R., Donahue, N. M., Prevot, A. S. H., Zhang, Q., Kroll,	
744	J. H., DeCarlo, P. F., Allan, J. D., Coe, H., Ng, N. L., Aiken, A. C., Docherty, K. S.,	
745	Ulbrich, I. M., Grieshop, A. P., Robinson, A. L., Duplissy, J., Smith, J. D., Wilson,	
746	K. R., Lanz, V. A., Hueglin, C., Sun, Y. L., Tian, J., Laaksonen, A., Raatikainen, T.,	
747	Rautiainen, J., Vaattovaara, P., Ehn, M., Kulmala, M., Tomlinson, J. M., Collins, D.	
748	R., Cubison, M. J., Dunlea, E. J., Huffman, J. A., Onasch, T. B., Alfarra, M. R.,	
749	Williams, P. I., Bower, K., Kondo, Y., Schneider, J., Drewnick, F., Borrmann, S.,	
750	Weimer, S., Demerjian, K., Salcedo, D., Cottrell, L., Griffin, R., Takami, A., Miyoshi,	
751	T., Hatakeyama, S., Shimono, A., Sun, J. Y., Zhang, Y. M., Dzepina, K., Kimmel, J.	
752	R., Sueper, D., Jayne, J. T., Herndon, S. C., Trimborn, A. M., Williams, L. R., Wood,	
753	E. C., Middlebrook, A. M., Kolb, C. E., Baltensperger, U., and Worsnop, D. R.:	
754	Evolution of Organic Aerosols in the Atmosphere, Science., 326, 1525-1529,	
755	https://doi.org/10.1126/science.1180353, 2009.	
756	-	
757	Khalizov, A. F., Zhang, R., Zhang, D., Xue, H., Pagels, J., and McMurry, P. H.:	
758	Formation of highly hygroscopic soot aerosols upon internal mixing with sulfuric	
759	acid vapor, J. Geophys. ResAtmos., 114, D05208,	
760	https://doi.org/10.1029/2008jd010595, 2009.	
761	Kiselev, A., Wennrich, C., Stratmann, F., Wex, H., Henning, S., Mentel, T.F., Kiendler-	

Scharr, A., Schneider, J., Walter, S., Lieberwirth, I.: Morphological characterization 762

的: 超链接, 字体: +西文正文 (等线), 五号

of soot aerosol particles during LACIS Experiment in November (LExNo), J.

- Geophys. Res. Atmos., 115, D11204. https://doi.org/10.1029/2009jd012635, 2010.
 Khalizov, A. F., Zhang, R., Zhang, D., Xue, H., Pagels, J., and McMurry, P. H.:
 Formation of highly hygroscopic soot aerosols upon internal mixing with sulfuric
- 760Formation of highly hygroscopic soor acrossis upon internal mixing with surface767acidvapor,J.Geophys.Res. Atmos.,114,D05208,768https://doi.org/10.1029/2008jd010595, 2009.
- Kawana, K., Nakayama, T., and Mochida, M.: Hygroscopicity and CCN activity of
 atmospheric aerosol particles and their relation to organics: Characteristics of urban
 aerosols in Nagoya, Japan, J. Geophys. Res.-Atmos., 121, 4100–4121,
 https://doi.org/10.1002/2015JD023213, 2016.
- Kostenidou, E., Pathak, R. K., & Pandis, S. N.: An Algorithm for the Calculation of
 Secondary Organic Aerosol Density Combining AMS and SMPS Data, Aerosol
 Science and Technology, 41:11, 1002-1010, https://doi:
 10.1080/02786820701666270, 2007.
- Li, M., Zhang, Q., Kurokawa, J.-I., Woo, J.-H., He, K., Lu, Z., Ohara, T., Song, Y.,
 Streets, D. G., Carmichael, G. R., Cheng, Y., Hong, C., Huo, H., Jiang, X., Kang, S.,
 Liu, F., Su, H., and Zheng, B.: MIX: a mosaic Asian anthropogenic emission
 inventory under the international collaboration framework of the MICS-Asia and
 HTAP, Atmos. Chem. Phys., 17, 935–963, https://doi.org/10.5194/acp-17-935-2017,
 2017.
- Liu, D., Joshi, R., Wang, J., Yu, C., Allan, J. D., Coe, H., Flynn, M. J., Xie, C., Lee, J.,
 Squires, F., Kotthaus, S., Grimmond, S., Ge, X., Sun, Y., and Fu, P.: Contrasting
 physical properties of black carbon in urban Beijing between winter and summer,
 Atmos. Chem. Phys., 19, 6749–6769, https://doi.org/10.5194/acp-19-6749-2019,
 2019a.
- Liu, D., Allan, J., Whitehead, J., Young, D., Flynn, M., Coe, H., McFiggans, G.,
 Fleming, Z. L., and Bandy, B.: Ambient black carbon particle hygroscopic properties
 controlled by mixing state and composition, Atmos. Chem. Phys., *13*, 2015–2029,
 https://doi.org/10.5194/acp-13-2015-2013, 2013.
- Liu, H., Pan, X.L., Wu, Y., Wang, D.W., Tian, Y., Liu, X.Y., et al.: Effective densities of
 soot particles and their relationships with the mixing state at an urban site in the
 Beijing megacity in the winter of 2018, Atmos. Chem. Phys. 19, 14791–14804,
 https://doi.org/10.5194/acp-19-14791-2019, 2019b.
- Lide, D. R. (ed.). CRC Handbook of Chemistry and Physics. CRC Press: Ann Arbor,
 MI. (1992).
- Lance, S., Medina, J., Smith, J., and Nenes, A.: Mapping the operation of the DMT
 continuous flow CCN counter, Aerosol Sci. Tech., 40, 242–254,
 https://doi.org/10.1080/02786820500543290, 2006.
- Liu, H., Pan, X., Liu, D., Liu, X., Chen, X., Tian, Y., Sun, Y., Fu, P., and Wang, Z.:
 Mixing characteristics of refractory black carbon acrosols at an urban site in Beijing,
 Atmos. Chem. Phys., 20, 5771–5785, https://doi.org/10.5194/acp-20-5771-2020,
 2020.
- Liu, L, Zhang, J, Zhang, Y, Wang, Y, Xu, L, Yuan, Q, et al.: Persistent residential
 burning-related primary organic particles during wintertime hazes in North China:

带格式的: 下划线

带格式的: 超链接, 字体: +西文正文 (等线), 五号

insights into their aging and optical changes, Atmos. Chem. Phys. 21, 2251–2265,
https://doi.org/10.5194/acp-21-2251-2021, 2021a.

- Liu, H., Pan, X., Liu, D., Liu, X., Chen, X., Tian, Y., Sun, Y., Fu, P., and Wang, Z.:
 Mixing characteristics of refractory black carbon aerosols at an urban site in Beijing,
 Atmos. Chem. Phys., 20, 5771–5785, https://doi.org/10.5194/acp-20-5771-2020,
 2020.
- Liu, J., Zhang, F., Xu, W., Sun, Y., Chen, L., Li, S.: Hygroscopicity of organic aerosols
 linked to formation mechanisms, Geophysical Research Letters, 48, e2020GL091683,
 https://doi.org/10.1029/2020gl091683, 2021b.
- McMurry, H. Peter, Wang Xin, Park Kihong & Ehara Kensei.: The Relationship
 between Mass and Mobility for Atmospheric Particles: A New Technique for
 Measuring Particle Density, Aerosol Sci. Technol., 36:2, 227-238,
 https://doi.10.1080/027868202753504083, 2002.
- Massoli, P., Onasch, T.B., Cappa, C.D., Nuamaan, I., Hakala, J., Hayden, K., Li, S.M.,
 Sueper, D.T., Bates, T.S., Quinn, P.K., Jayne, J.T., Worsnop, D.R.: Characterization
 of black carbon-containing particles from soot particle aerosol mass spectrometer
 measurements on the R/V Atlantis during CalNex 2010, J. Geophys. Res.- Atmos.,
 120, 2575-2593, https://doi.org/10.1002/2014JD022834, 2015.
- Mei, F., Setyan, A., Zhang, Q., and Wang, J.: CCN activity of organic aerosols observed
 downwind of urban emissions during CARES, Atmos. Chem. Phys., 13, 12155–
 12169, https://doi.org/10.5194/acp-13-12155-2013, 2013.
- Meng, J. W., Yeung, M. C., Li, Y. J., Lee, B. Y. L., and Chan, C. K.: Size resolved cloud
 condensation nuclei (CCN) activity and closure analysis at the HKUST Supersite in
 Hong Kong, Atmos. Chem. Phys., 14, 10267–10282, https://doi.org/10.5194/acp-14 10267-2014, 2014.
- McMeeking, G.R., Hamburger, T., Liu, D., Flynn, M., Morgan, W.T., Northway, M.,
 Highwood, E.J., Krejci, R., Allan, J.D., Minikin, A., Coe, H.: Black carbon
 measurements in the boundary layer over western and northern Europe. Atmos.
 Chem. Phys. 10, 9393-9414, https://doi.org/10.5194/acp-10-9393-2010, 2010.
- Meng, J. W., Yeung, M. C., Li, Y. J., Lee, B. Y. L., and Chan, C. K.: Size-resolved cloud
 condensation nuclei (CCN) activity and closure analysis at the HKUST Supersite in
 Hong Kong, Atmos. Chem. Phys., 14, 10267–10282, https://doi.org/10.5194/acp-14 10267-2014, 2014.
- 840 Noureddini, H., Teoh, B. C., Davis Clements, L.: Densities of vegetable oils and fatty
- acids, J. Am. Oil Chem. Soc., 69 (12), 1184–1188, 1992.
- Olfert, J. S., Symonds, J. P. R., and Collings, N.: The effective density and fractal
 dimension of particles emitted from a light-duty diesel vehicle with a diesel oxidation
 catalyst, J. Aerosol Sci., 38, 69–82, https://doi.org/10.1016/j.jaerosci.2006.10.002,
 2007.
- Pagels, J., Khalizov, A.F., McMurry, P.H. and Zhang, R.Y.: Processing of soot by
 controlled sulphuric acid and water condensation-mass and mobility relationship,
 Aerosol Sci. Technol., 43, 629–640, https://doi.org/10.1080/02786820902810685,
 2009.

(带格式的: 字体颜色: 文字 1

带格式的:字体:+西文正文(等线),五号,下划线

850	Park, K., Ki	ittelson,	D. B.,	and McMur	ry, P. H.: S	tructur	al properties	of die	sel exhaust
851	particles	measure	ed by t	ransmission	electron n	nicrosc	opy (TEM):	Relat	ionships to
852	particle	mass	and	mobility,	Aerosol	Sci.	Technol.,	38,	881–889,
853	https://do	i.org/10	.1080/0	0278682905	05189, 200	4.			

- Pagels, J., Khalizov, A.F., McMurry, P.H. and Zhang, R.Y.: Processing of soot by
 controlled sulphuric acid and water condensation mass and mobility relationship,
 Aerosol Sei. Technol., 43, 629–640, https://doi.org/10.1080/02786820902810685,
 2009.
- Peng, J. F., Hu, M., Guo, S., Du, Z. F., Zheng, J., Shang, D. J., Zamora, M., Zeng, L.
 M., Shao, M., Wu, Y. S., Zheng, J., Wang, Y., Glen, C., Collins, D., Molina, M., and
 Zhang, R. Y.: Markedly enhanced absorption, and direct radiative foreing of black
 carbon under polluted urban environments, P. Natl. Acad. Sci. USA, 113(16), 4266–
 4271, https://doi.org/10.1073/pnas.1602310113, 2016.
- Petters, M. D. and Kreidenweis, S. M.: A single parameter representation of
 hygroscopic growth and cloud condensation nucleus activity, Atmos. Chem. Phys.,
 7, 1961–1971, https://doi.org/10.5194/acp-7-1961-2007, 2007.
- Paatero, P. and Tapper, U.: Positive matrix factorization: A nonnegative factormodel
 with optimal utilization of error estimates of data values, Environmetrics, 5, 111–126,
 1994.
- Pileci, R. E., Modini, R. L., Bertò, M., Yuan, J., Corbin, J. C., Marinoni, A., Henzing,
 B., Moerman, M. M., Putaud, J. P., Spindler, G., Wehner, B., Müller, T., Tuch, T.,
 Trentini, A., Zanatta, M., Baltensperger, U., and Gysel-Beer, M.: Comparison of colocated refractory black carbon (rBC) and elemental carbon (EC) mass concentration
 measurements during field campaigns at several European sites, Atmos. Meas. Tech.,
 14, 1379–1403, https://doi.org/10.5194/amt-14-1379-2021, 2021.
- Peng, J. F., Hu, M., Guo, S., Du, Z. F., Zheng, J., Shang, D. J., Zamora, M., Zeng, L.
 M., Shao, M., Wu, Y. S., Zheng, J., Wang, Y., Glen, C., Collins, D., Molina, M., and
 Zhang, R. Y.: Markedly enhanced absorption, and direct radiative forcing of black
 carbon under polluted urban environments, P. Natl. Acad. Sci. USA, 113(16), 4266–
 4271, https://doi.org/10.1073/pnas.1602310113, 2016.
- Peng, J. F., Hu, M., Guo, S., Du, Z. F., Zheng, J., M., Zeng, L. M., Shao, M., Wu, Y. S.,
 Collins, D., Molina, M., and Zhang, R. Y.: Ageing and hygroscopicity variation of
 black carbon particles in Beijing measured by a quasi-atmospheric aerosol evolution
 study (QUALITY) chamber, Atmos. Chem. Phys., 17(17), 10333-10348,
 https://doi.org/10.5194/acp-17-10333-2017, 2017.
- Pan, X.L., Kanaya, Y., Taketani, F., Miyakawa, T., Inomata, S., Komazaki, Y., et al.:
 Emission characteristics of refractory black carbon aerosols from fresh biomass
 burning: a perspective from laboratory experiments, Atmos. Chem. Phys., 17(21),
 13001–13016, https://doi.org/10.5194/acp-17-13001-2017, 2017.
- Park, K., Cao, F., Kittelson, D. B., & McMurry, P. H.: Relationship between particle
 mass and mobility for diesel exhaust particles, Environ. Sci. Tehnol., 37, 577–583,
 https://doi.org/10.1021/es025960v, 2003.
- Pileci, R. E., Modini, R. L., Bertò, M., Yuan, J., Corbin, J. C., Marinoni, A., Henzing,
 B., Moerman, M. M., Putaud, J. P., Spindler, G., Wehner, B., Müller, T., Tuch, T.,

带格式的:字体:+西文正文(等线),五号,下划线

 Trentini, A., Zanatta, M., Baltensperger, U., and Gysel-Beer, M.: Comparison of colocated refractory black carbon (rBC) and elemental carbon (EC) mass concentration measurements during field campaigns at several European sites, Atmos. Meas. Tech., 14, 1379–1403, https://doi.org/10.5194/amt-14-1379-2021, 2021.

- Qiao, K., Wu, Z., Pei, X., Liu, Q., Shang, D., Zheng, J., Du, Z., Zhu, W., Wu, Y., Lou, S.,
 Guo, S., Chan, C.K., Pathak, R.K., Hallquist, M., Hu, M.: Size-resolved effective
 density of submicron particles during summertime in the rural atmosphere of Beijing.
 China, J. Environ. Sci. (China) 73, 69–77. https://doi.org/10.1016/j.jes.2018.01.012,
 2018.
- Rissler, J., Nordin, E. Z., Eriksson, A. C., Nilsson, P. T., Frosch, M., Sporre, M. K.,
 Wierzbicka, A., Svenningsson, B., Londahl, J., Messing, M. E., Sjogren, S.,
 Hemmingsen, J. G., Loft, S., Pagels, J. H., and Swietlicki, E.: Effective Density and
 Mixing State of Aerosol Particles in a Near-Traffic Urban Environment, Environ. Sci.
 Technol., 48, 6300–6308, https://doi.org/10.1021/es5000353, 2014.
- Riemer, N., Vogel, H., and Vogel, B.: Soot aging time scales in polluted regions during
 day and night, Atmos. Chem. Phys., 4, 1885–1893, https://doi.org/10.5194/acp-41885-2004, 2004.
- Ramanathan, V. and Carmichael, G.: Global and regional climate changes due to black
 carbon, Nat. Geosci., 36, 221-227, https://doi.org/10.1038/ngeo156, 2008.
- Ren J, Zhang F, Chen L, et al.: Identifying the hygroscopic properties of fine aerosol
 particles from diverse sources in urban atmosphere and the applicability in prediction
 of cloud nuclei, Atmospheric Environment, 298: 119615,
 https://doi.org/10.1016/j.atmosenv.2023.119615, 2023.
- Ren, J., Zhang, F., Wang, Y., Collins, D., Fan, X., Jin, X., et al.: Using different assumptions of aerosol mixing state and chemical composition to predict CCN concentrations based on field measurements in urban Beijing, Atmos. Chem. Phys., 18, 6907–6921, https://doi.org/10.5194/acp-18-6907-2018, 2018,
- Rader, D.J., McMurry, P.H.: Application of the tandem differential mobility analyzer
 to studies of droplet growth or evaporation, J. Geophys. Res.- Atmos., 17, 771-787,
 https://doi.org/10.1016/0021-8502(86)90031-5, 1986.
- Reyes-Villegas, E., Bannan, T., Le Breton, M., Mehra, A., Priestley, M., Percival, C.,
 Coe, H., and Allan, J. D.: Online Chemical Characterization of Food-Cooking
 Organic Aerosols: Implications for Source Apportionment, Environ. Sci. Technol.,
 52, 5308–5318, https://doi.org/10.1021/acs.est.7b06278, 2018.
- 928 varz, J.P., Gao, R.S., Fahey, D.W., Thomson, D.S., Reeves, J.M., Darbeheshti, M., Baumgardner, D.G., Kok, G.L 929 Chung 930 M., Hendricks, J., Lauer, A., K Carcher, B., Slowik, J.G., Roser 931 T.L., Langford, A.O., Loewenstein, M., Aikin, K.C.: Single-particle 932 of midlatitude black carbon and light-scattering acrosols from the boundary 933 the lower stratosphere. J. Geophys. Atmosphere 111 D16207 - Doo https://doi.org/10.1029/2006JD007076, 2006. 934
- Schwarz, J. P., Gao, R. S., Spackman, J. R., Watts, L. A., Thomson, D. S., Fahey, D.
 W., Ryerson, T. B., Peischl, J., Holloway, J. S., Trainer, M., Frost, G. J., Baynard,
 T., Lack, D. A., de Gouw, J. A., Warneke, C., and Del Negro, L. A.: Measurement

带格式的:超链接,字体:+西文正文(等线),五号

- of the mixing state, mass, and optical size of individual black carbon particles in
 urban and biomass burning emissions, Geophys. Res. Lett., 35, L13810,
 https://doi.org/10.1029/2008GL033968, 2008.
- 941 Stokes, R. and Robinson, R.: Interactions in aqueous nonelectrolyte solutions, I. Solute 942 solvent equilibria, J. Phys. Chem. US, 70, 2126–2131, 1966.
- Sun, Y., Du, W., Fu, P., Wang, Q., Li, J., Ge, X., Zhang, Q., Zhu, C., Ren, L., Xu, W.,
 Zhao, J., Han, T., Worsnop, D. R., and Wang, Z.: Primary, and secondary aerosols
 in Beijing in winter: sources, variations, and processes, Atmos. Chem. Phys., 16,
 8309–8329, https://doi.org/10.5194/acp-16-8309-2016,2016.
- Sun, Y. L., Wang, Z. F., Du, W., Zhang, Q., Wang, Q. Q., Fu, P. Q., Pan, X. L., Li, J.,
 Jayne, J., and Worsnop, D. R.: Long term real-time measurements of aerosol particle
 composition in Beijing, China: seasonal variations, meteorological effects, and
 source analysis, Atmos. Chem. Phys., 15, 10149–10165, https://doi.org/10.5194/acp15–10149–2015, 2015.
- <u>Stokes, R. and Robinson, R.: Interactions in aqueous nonelectrolyte solutions, I. Solute-</u>
 <u>solvent equilibria, J. Phys. Chem.-US, 70, 2126–2131, 1966.</u>
- Schwarz, J.P., Gao, R.S., Fahey, D.W., Thomson, D.S., Watts, L.A., Wilson, J.C.,
 Reeves, J.M., Darbeheshti, M., Baumgardner, D.G., Kok, G.L., Chung, S.H., Schulz,
 M., Hendricks, J., Lauer, A., K€archer, B., Slowik, J.G., Rosenlof, K.H., Thompson,
 T.L., Langford, A.O., Loewenstein, M., Aikin, K.C.: Single-particle measurements
 of midlatitude black carbon and light-scattering aerosols from the boundary layer to
 the lower stratosphere. J. Geophys. Res.: Atmosphere 111, D16207,
 https://doi.org/10.1029/2006JD007076, 2006.
- Shiraiwa, M., Kondo, Y., Moteki, N., Takegawa, N., Sahu, L., Takami, A., et al.:
 Radiative impact of mixing state of black carbon aerosol in Asian outflow, J.
 Geophys. Res.- Atmos., 113, D24210, https://doi.org/10.1029/2008JD010546, 2008.
- Tan, H., Xu, H., Wan, Q., Li, F., Deng, X., Chan, P. W., Xia, D., and Yin, Y.: Design and application of an unattended multifunctional H-TDMA system, J. Atmos. Ocean.
 Tech., 30, 1136–1148, https://doi.org/10.1175/JTECH-D-12-00129.1, 2013.
- Ulbrich, I. M., Canagaratna, M. R., Zhang, Q., Worsnop, D. R., and Jimenez, J. L.:
 Interpretation of organic components from Positive Matrix Factorization of aerosol
 mass spectrometric data, Atmos. Chem. Phys., 9, 2891–2918,
 https://doi.org/10.5194/acp-9- 2891-2009, 2009.
- Wang, Y., Wan, Q., Meng, W., Liao, F., Tan, H., and Zhang, R.: Long-term impacts of
 aerosols on precipitation and lightning over the Pearl River Delta megacity area in
 China, Atmos. Chem. Phys., 11, 12421–12436, https://doi.org/10.5194/acp-1112421-2011, 2011.
- Wang, Y. Y., Liu, F. S., He, C. L., Bi, L., Cheng, T. H., Wang, Z. L., Zhang, H., Zhang,
 X. Y., Shi, Z. B., and Li, W. J.: Fractal dimensions and mixing structures of soot
 particles during atmospheric processing, Environ. Sci. Tech. Lett., 4, 487–493,
 https://doi.org/10.1021/acs.estlett.7b00418, 2017.
- Wu, Y. F., Xia, Y. J., Huang, R. J., Deng, Z. Z., Tian, P., Xia, X. G., et al.: A study of the
 morphology and effective density of externally mixed black carbon aerosols in
 ambient air using a size-resolved single-particle soot photometer (SP2), Atmos. Meas.

982 Tech., 12, 4347–4359, https://doi.org/10.5194/amt-12-4347-2019, 2019.

Wu, Y., Wang, X., Tao, J., Huang, R., Tian, P., Cao, J., Zhang, L., Ho, K.-F., Han, Z.,
and Zhang, R.: Size distribution and source of black carbon aerosol in urban Beijing
during winter haze episodes, Atmos. Chem. Phys., 17, 7965–7975,

986 https://doi.org/10.5194/acp-17-7965-2017, 2017.

- Wu, Z. J., Zheng, J., Shang, D. J., Du, Z. F., Wu, Y. S., Zeng, L. M., Wiedensohler, A.,
 and Hu, M.: Particle hygroscopicity and its link to chemical composition in the urban
 atmosphere of Beijing, China, during summertime, Atmos. Chem. Phys., 16, 1123–
 1138, https://doi.org/10.5194/acp-16-1123-2016, 2016.
- Xue, H., Khalizov, A. F., Wang, L., Zheng, J., and Zhang, R.: Effects of dicarboxylic
 acid coating on the optical properties of soot, Phys. Chem. Chem. Phys., 11, 7869–
 7875, https://doi.org/10.1039/b904129j, 2009.
- Xu, W., Sun, Y., Wang, Q., Zhao, J., Wang, J., Ge, X., et al.: Changes in aerosol
 chemistry from 2014 to 2016 in winter in beijing: Insights from high-resolution
 aerosol mass spectrometry, J. Geophys. Res.-Atmos., 124, 1132–1147.
 https://doi.org/10.1029/2018jd029245, 2019.
- Yu, W., Fossum, K. N., Ovadnevaite, J., Lin, C., Huang, R. J., O'Dowd, C., and Ceburnis, D.: The impact of aerosol size dependent hygroscopicity and mixing state on the cloud condensation nuclei potential over the north east Atlantic, Atmos. Chem. Phys., 21, 8655–8675, https://doi.org/10.5194/acp-21-8655-2021, 2021.
- Yuan, T., Li, Z., Zhang, R., and Fan, J.: Increase of cloud droplet size with aerosol optical depth: An observation and modeling study, J. Geophys. Res.-Atmos., 113, 1004
 D04201, https://doi.org/10.1029/2007JD008632, 2008.
- Yu, C., Liu, D., Broda, K., Joshi, R., Olfert, J., Sun, Y., Fu, P., Coe, H., Allan, J.D.:
 Characterising mass-resolved mixing state of black carbon in Beijing using a
 morphology-independent measurement method, Atmos. Chem. Phys., 20, 3645–
 3661. https://doi.org/10.5194/acp-20-3645-2020, 2020.
- Zhang, R. Y., Khalizov, A. F., Pagels, J., Zhang, D., Xue, H. X., and McMurry, P. H.:
 Variability in morphology, hygroscopicity, and optical properties of soot aerosols
 during atmospheric processing, P. Natl. Acad. Sci. USA, 105, 10291–10296,
 https://doi.org/10.1073/pnas.0804860105, 2008.
- Zhang, F., Li, Z., Li, Y., Sun, Y., Wang, Z., Li, P., Sun, L., Wang, P., Cribb, M., Zhao,
 C., Fan, T., Yang, X., and Wang, Q.: Impacts of organic aerosols and its oxidation
 level on CCN activity from measurement at a suburban site in China, Atmos. <u>Chem.</u>
 Phys., Y., Zhang, Q., <u>16</u>, 5413–5425, https://doi.org/10.5194/acp-16- 5413-2016,
- 1017 2016a. 1018 Zhang, Ch 1019 nd Ho K · M Wi 1020 mixed black carbon with SP2 and **VTDMA** 1021 ment of black carbon in the atmosphere 0 1922 1022 https://doi.org/10.5194/amt-9-1833-2016, 2016.
- 1023 Zdanovskii, A.: New methods for calculating solubilities of electrolytes in 1024 multicomponent systems, Zh. Fiz. Khim.C, 22, 1475–1485, 1948. –
- 1025 Zhang, F., Wang, Y., Peng, J., Ren, J., Collins, D., Zhang, R., et al.: Uncertainty in 带格式的: 字体颜色: 自动设置

带格式的: 超链接, 字体: +西文正文 (等线), 五号
 带格式的: 超链接, 字体: +西文正文 (等线), 五号
 带格式的: 超链接, 字体: +西文正文 (等线), 五号

1026	predicting CCN activity of aged and primary aerosols, J. Geophys. ResAtmos.,	
1027	122(21), 11723–11736, https://doi.org/10.1002/2017jd027058, 2017.	
1028	Zhang, F., Ren, J., Fan, T., Chen, L., Xu, W., Sun, Y., et al.: Significantly enhanced	
1029	acrosol-CCN activity and number, J. Geophys. ResAtmos., 124, 14102–14113,	
1030	https://doi.org/10.1029/2019jd031457, 2019.	带格式的:默认段落字体,字体:+西文正文(等线),五号,
1031	Zhang, F., Wang, Y., Peng, J., Chen, L., Sun, Y., Duan, L., Ge, X., Li, Y., Zhao, J., Liu,	下划线,字体颜色:自动设置
1032	C., Zhang, X., Zhang, G., Pan, Y., Wang, Y., Zhang, A. L., Ji, Y., Wang, G., Hu, M.,	
1033	Molina, M. J., Zhang, R.: An unexpected catalyst dominates formation and radiative	
1034	forcing of regional haze, P. Natl. Acad. Sci. USA, 117(8), 3960-3966,	
1035	https://doi.org/10.1073/pnas.1919343117, 2020a,	带格式的:默认段落字体,字体:+西文正文(等线),五号, 下划线,字体颜色:自动设置
1036	Zhang, Y., Zhang, Q., Cheng, Y., Su, H., Kecorius, S., Wang, Z., Wu, Z., Hu, M., Zhu,	
1037	T., Wiedensohler, A., and He, K.: Measuring the morphology and density of	
1038	internally mixed black carbon with SP2 and VTDMA: new insight into the	
1039	absorption enhancement of black carbon in the atmosphere, Atmos. Meas. Tech., 9,	
1040	<u>1833–1843, https://doi.org/10.5194/amt-9-1833-2016, 2016b.</u>	
1041	Yao, Z., Li, H.: Particle Size and Mixing State of Freshly Emitted Black Carbon from	
1042	Different Combustion Sources in China, Environ. Sci. Technol., 54(13): p. 7766-	
1043	7774, https://doi.org/10.1021/acs.est.9b07373, 2020b.	
1044	Zhang, F., Peng, J., Chen, L., Collins, D., Li, Y., Jiang, S., Liu, J., Zhang, R.: The effect	
1045	of Black carbon aging from NO2 oxidation of SO2 on its morphology, optical and	
1046	hygroscopic properties, Environ. Res., 212, 113238,	
1047	https://doi.org/10.1016/j.envres.2022.113238, 2022	带格式的:下划线
1048	Zhang, R., Wang, G., Guo, S., Zamora, M. L., Ying, Q., Lin, Y.: Formation of urban fine particulate matter, Chemical Reviews, 115(10), 3803–3855,	
1049	https://doi.org/10.1021/acs.chemrev.5b00067, 2015.	
1050		
1051	Zhou, Y., Ma, N., Wang, Q., Wang, Z., Chen, C., Tao, J., Hong, J., Peng, L., He, Y.,	
1052	Xie, L., Zhu, S., Zhang, Y., Li, G., Xu, W., et al.: Bimodal distribution of size-	
1053	resolved particle effective density: results from a short campaign in a rural environ	
1054	ment over the North China Plain, Atmos. Chem. Phys., 22, 2029-2047.	
1055	https://doi.org/10.5194/acp 22 2029 2022, 2022.	
1056	Zhao, G., Tan, T., Hu, S., Du, Z., Shang, D., Wu, Z., Guo, S., Zheng, J., Zhu, W., Li,	
1057	M., Zeng, L., and Hu, M.: Mixing state of black carbon at different atmospheres in	
1058	north and southwest China, Atmos. Chem. Phys., 22, 10861-10873,	
1059	https://doi.org/10.5194/acp-22-10861-2022, 2022.	
1060	Zhang, F., Ren, J., Fan, T., Chen, L., Xu, W., Sun, Y., et al.: Significantly enhanced	
1061	aerosol CCN activity and number, J. Geophys. ResAtmos., 124, 14102-14113,	
1062	https://doi.org/10.1029/2019jd031457, 2019,	带格式的:默认段落字体,字体:+西文正文(等线),五号, 下划线,字体颜色:自动设置
1063	Zdanovskii, A.: New methods for calculating solubilities of electrolytes in	
1064	<u>multicomponent systems, Zh. Fiz. Khim.C, 22, 1475–1485, 1948.</u>	
1065	Zhang, R., Wang, G., Guo, S., Zamora, M. L., Ying, Q., Lin, Y.: Formation of urban	带格式的:字体颜色:自动设置
1066	fine particulate matter, Chemical Reviews, 115(10), 3803-3855,	
1067	https://doi.org/10.1021/acs.chemrev.5b00067, 2015.	
1068	Zhang, Y., Zhang, Q., Yao, Z., Li, H.: Particle Size and Mixing State of Freshly Emitted	带格式的:不对齐到网格
1069 1070	Black Carbon from Different Combustion Sources in China, Environ. Sci. Technol., 54(13): p. 7766-7774, https://doi.org/10.1021/acs.est.9b07373, 2020b	
11(1)/(1)	54(13); n //66-///4 https://doi.org/10.1071/acs.est.9607373.70706	