Dear Dr. Eimear Dunne,

Thank you for your comments and suggestions. Below I provide a point-by-point response to individual comments (referee comments and suggestions are in italics, responses and revisions are in plain font; revised sections in the manuscript text in response to the comments are marked with red color).

Comments and suggestions from Dr. Eimear Dunne

Comment 1:

The value $1 e^{-3}$ on line 190 is definitely wrong somehow, and probably meant to read 1E-3 or equivalent

Responses and revisions 1:

Thanks for pointing out this error. Now we have changed $1 e^{-3}$ to " 1×10^{-3} " (line 191).

Comment 2:

Figure A4 (a) would benefit from having a log scale on the x-axis, maybe as an extra panel

Responses and revisions 2:

Thanks for the suggestion. We have added an extra panel using a log scale to Figure A4 (a), so that the vertical patterns in the middle and upper troposphere are clearer now:

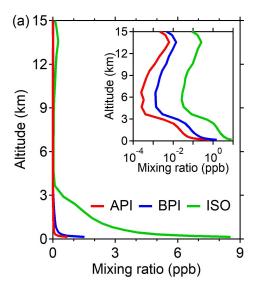


Figure A4. Simulated (a) vertical profiles and (b) time series of α -pinene (API), β -pinene (BPI), and isoprene (ISO) mixing ratios (STP) at the location of ATTO. The embedded figure in (a) is the same as the outer figure but on a log scale.

Comment 3:

It is claimed that Figure A5 shows a fair agreement with observations, but A5 only shows simulated values - please support your claim by also plotting the relevant observations

Responses and revisions 3:

Thanks for the comment. Here we did not make a point-to-point comparison between simulated and observed SO₂ due to the unavailability of the aircraft observation of SO₂ during the ACRIDICON-CHUVA campaign. The observed SO₂ concentrations used for the comparison are from published papers. Andreae & Andreae (1988) observed an SO₂ concentration of 18 ppt in the free troposphere (FT) based on aircraft measurements over the Amazon Basin during July and August, which represents a background condition with little impact from anthropogenic plumes. The simulated SO₂ concentration in this study is generally around 21 ppt in the free troposphere (except the high value at ~6 km) and decreases near the tropopause. The simulated result is 3 ppt higher than the observation but basically captures the magnitude of the SO₂ in the free troposphere. Therefore we consider this simulation result is in fair agreement with observations. We accept the referee's suggestion and have now explicitly added the observed results from Andreae & Andreae (1988) in the manuscript so that the comparison is clearer:

Line 691:

"The simulated SO₂ concentration of around 21 ppt throughout most of the FT at the location of the ATTO site (Fig. A5) is in fair agreement with an observed background SO₂ concentration of 18 ppt in the FT over the Central Amazon (Andreae & Andreae, 1988)."

As for the SO₂ in the planetary boundary layer (PBL), Andreae & Andreae (1988) observed an SO₂ concentration of 27 ppt in the PBL from aircraft measurements; Ramsay et al. (2020) observed an SO₂ concentration of 80 ppt at ATTO at 60 m. The simulated SO₂ concentration is around 68 ppt at the higher part of the PBL (above 300 m) but shows a high value of 130 ppt near the ground surface. Therefore the simulated SO₂ in the PBL is relatively high compared to observations, especially for near surface. Then we conducted a sensitivity study to constrain the whole PBL SO₂ concentration to 30 ppt to test the influence of the SO₂ overestimation in PBL on the aerosol simulation. The sensitivity study (Fig. A6) shows that the SO₂ overestimation in PBL only causes a minor difference in the aerosol concentration and does not substantially affect the aerosol simulation results in this study.

For the SO₂ comparison in the PBL, we have also added the observation results to make the comparison clearer:

Line 694:

"Compared to the observed SO₂ concentration of 27 ppt in the PBL (Andreae & Andreae, 1988) and 80 ppt near the ground surface (Ramsay et al., 2020), the modeled SO₂ concentration in the PBL, especially near the ground surface, was relatively higher (Fig. A5)."

Comment 4:

With those issues out of the way, it's time to address my main concern: the Wexler et al. (1994) nucleation parameterisation. I understand why the paper did not update the default nucleation scheme in WRF-Chem; after all, the authors were already implementing a new nucleation scheme, and it makes sense to compare it to the existing set-up. This is especially true for a model like WRF, where there are so many different configurations available. So I don't think it would be even slightly reasonable to suggest rejecting the paper on these grounds, but I do think that there needs to be more acknowledgement of how a nucleation parameterisation affects CCN in the simulated upper troposphere. The Wexler et al. (1994) publication explicitly states:

"The number of particles produced by this nucleation operator is somewhat arbitrary [...] Any error produced by this treatment is mitigated in the SoCAB because the vast majority of the aerosol loading is due to primary emission and condensation of secondary organic compounds. In locations where nucleation is more significant, this treatment may not be sufficiently accurate."

In the upper troposphere, nucleation is the only local source of aerosols. If the real conditions being simulated are actually saturated with respect to freshly nucleated particles, and the default nucleation parameterisation under-predicts the true nucleation rate significantly, then changing to any parameterisation which predicts a value close to the saturation limit will improve predictions; but it cannot then be concluded that the nucleation pathway is the dominant one in that region, even if that is the case in reality.

I would agree with Referee #1 that the conclusions ought to be softened. If any nucleation scheme that was known to be more robust in the UT had been used, I would have been happy to accept the conclusions as they stand. However, I would also be happy to discuss implementing the Dunne et al. (2016) scheme in WRF-Chem with the authors, if they would be interested in a future collaboration where their stronger conclusions might yet be validated!

Responses and revisions 4:

We thank the referee for the insightful comment and patient explanation. We agree that in this study we can not conclude a dominant role of pure organic nucleation without testing other nucleation mechanisms. In this manuscript, we describe the effect of organic nucleation on CN in the UT (i.e. an increase of 2100 cm⁻³, over one quarter of the observed total CN concentration of 7700 cm⁻³, Table A4) as "strong particle production (Line 1)", "play important roles in maintaining the particle population and size distribution in the UT (Line 128)", and "effectively increases the CN number by replenishing new nano-sized particles (Line 129)" instead of concluding it as a dominant mechanism. We realize that the sentence on Line 69 that "The organic nucleation mechanism in this study focuses on pure organic nucleation,, as it was found dominant among organic nucleation pathways in the Amazon (Zhu & Penner, 2019)" could be misleading. This is not the conclusion of this study but a result of Zhu & Penner (2019, Figure 3) who made comparisons among organics-involved nucleation pathways. Also, the short summary of the manuscript was in a relatively strong tone and has been softened from

"we show that the UT aerosol formation triggered by biogenic organics shapes the UT aerosols, and organic condensation is key for UT CCN production." to

"we show strong aerosol nucleation and condensation in the UT triggered by biogenic organics, and organic condensation is key for UT CCN production."

To address the referee's concern about the simulation performance of the 'Wexler 1994' scheme, we compared the binary nucleation simulation using the Wexler 1994 scheme in this study to the simulations in other studies. The comparison shows that the aerosol production by Wexler 1994 scheme generally agrees with the simulation by other schemes both in vertical distribution and in magnitude:

Line 716, "As shown from the nucleation rate in Fig. A12, H₂SO₄-H₂O binary nucleation mainly occurs in the free troposphere, which is consistent with the vertical distribution of binary nucleation

simulated for the Amazon region in Zhao et al. (2020). The H₂SO₄-H₂O binary nucleation causes a CN increase of over 3000 cm⁻³ in the UT under sufficient particle condensational growth as approximately estimated from the difference between OCD and BASEnoNUC (Table A4). It is of a comparable magnitude to the CN increase of 2100 cm⁻³ by organic nucleation. A higher rate of the H₂SO₄-H₂O binary nucleation over the organic nucleation was also found by Zhao et al. (2020) in the Amazon from 9 to 13 km altitude but the overall H₂SO₄-H₂O binary nucleation in the UT was insignificant in Zhao et al. (2020), which is different from the result in this study. This is expected as the result in Zhao et al. (2020) was for a low-SO₂ area and there was competition for H₂SO₄ by other H₂SO₄-involving nucleation processes in Zhao et al. (2020). In a global simulation where the inorganic nucleation was represented only by the H₂SO₄-H₂O binary nucleation, the columnintegrated H₂SO₄-H₂O binary nucleation in the Amazon is of the same magnitude as but somewhat lower than the organic nucleation (Zhu & Penner, 2019). Considering the H₂SO₄-H₂O binary nucleation occurs mainly in the upper troposphere and the organic nucleation in Zhu & Penner (2019) includes the hetero-molecular organic nucleation, the relative importance of H₂SO₄-H₂O binary nucleation to pure organic nucleation in the UT should be greater than shown in the column-integrated results. Therefore the simulated H₂SO₄-H₂O binary nucleation in this study should be generally reasonable.

For the question of how nucleation parameterisation affects CCN in the upper troposphere, we found the simulations in Westervelt et al. (2014) based on a global model GEOS-Chem-TOMAS can provide some hints. The sensitivity simulations in Westervelt et al. (2014; Figure 3) show that the CCN in the upper troposphere did not vary substantially when the nucleation scheme changed from binary nucleation to ternary nucleation or even when a tuning factor of 10⁻⁵ was applied to the nucleation rate. Based on the simulation results, the nucleation parameterisation did not seem to affect the upper tropospheric CCN in a significant way. However, these simulations were conducted by global models while sensitivity simulations by regional models where convective transport can be better resolved are still lacking. It's worthwhile to conduct such a sensitivity study, and we highly look forward to a corporation with the referee in the future. Here we accept the referee's suggestion and have added acknowledgment of the need to further investigate other nucleation schemes in order to understand their relative roles.

Line 220, "Note that although pure organic nucleation contributes importantly to the aerosol population in the UT, the relative roles of pure organic nucleation and other nucleation mechanisms, such as ternary and ion-induced inorganic nucleation (Napari et al., 2002; Yu et al., 2008), in the UT aerosol production remain to be investigated with a comprehensive consideration of nucleation parametrizations, e.g. those in Dunne et al. (2016) and Riccobono et al. (2014)."

Westervelt, D. M., Pierce, J. R., and Adams, P. J.: Analysis of feedbacks between nucleation rate, survival probability and cloud condensation nuclei formation, Atmos. Chem. Phys., 14, 5577–5597, https://doi.org/10.5194/acp-14-5577-2014, 2014.