

# Three dominant synoptic atmospheric circulation patterns influencing severe winter haze in eastern China

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**Abstract.** Previous studies have indicated that, on a synoptic scale, severe haze in eastern China (EC) is affected by atmospheric circulation variations. However, it is still unclear what the dominant atmospheric circulation patterns influencing the severe winter haze conditions in EC and the differences between them are. To systematically determine the dominant synoptic atmospheric circulation patterns of severe haze in different regions of EC, we use the hierarchical clustering algorithm (HCA) to classify the local geopotential height anomalies at 500 hPa over stations with severe haze and obtained three dominant synoptic atmospheric circulation types based on observed particulate matter with an aerodynamic diameter less than 2.5  $\mu\text{m}$  ( $\text{PM}_{2.5}$ ) concentrations and NCEP/NCAR reanalysis. Circulation Type1 is accompanied by significant northerly wind component anomalies over northern China and causes severe haze pollution over the Yangtze River valley. Although the local meteorological conditions are not conducive to haze formation and accumulation, the severe haze in the Yangtze River valley is related to the pollution transportation caused by the northerly wind anomalies. During the haze days with circulation Type2, the joint affection of the East Atlantic/West Russia teleconnection pattern and winter East Asia subtropical jet stimulate and maintain the anticyclonic anomalies over northeast Asia, which provides meteorological conditions conducive to the occurrence of severe haze over the whole EC. Circulation Type3 mainly caused severe haze events in northeast China through the establishment of blocking high over the Sea of Okhotsk. The results provide a basis for establishing haze prediction and management policies applicable to different regions in EC.

## 1 Introduction

Severe haze could increase the risk of traffic accidents by reducing visibility and could harm human health by causing respiratory diseases (Xie et al., 2014; Hu et al., 2015; Wang et al., 2016). Haze events in China are mainly caused by particulate matter with an aerodynamic diameter less than 2.5  $\mu\text{m}$  ( $\text{PM}_{2.5}$ ; Cai et al., 2017; Shen et al., 2018; Wang et al., 2021). Research shows that the distribution of haze days in China has characteristics of uneven spatial distribution, with

more spatial distribution in the economically developed eastern region and less in the economically underdeveloped region (Wu et al., 2013; Liu et al., 2015; Xu et al., 2015). With the rapid development of industrialization and an increase in urbanization and anthropogenic emission, eastern China (EC) has experienced more severe haze events with longer durations and larger spatial scales, causing serious harm in the past few decades (Monks et al., 2009; Qian et al., 2009; Wang et al., 2009). Since the beginning of the 21st century, the uneven spatial distribution of haze events in China has

become more obvious (Sun et al., 2016), which has led to the increasing rate of mortality related to respiratory diseases in Beijing–Tianjin–Hebei, the Yangtze River valley (YRV), and the Pearl River Delta (Tsaia et al., 2014; Ding et al., 2016; Fan and Sun, 2019). Although haze pollution control in China has been improved to some extent with the strict implementation of energy conservation and emission reduction policies after 2013 (Wang et al., 2021), haze still affects various socio-economic sectors and human health.

In addition to human activities, meteorological conditions are also considered as one of the most important factors for determining regional air quality. Previous studies have indicated that, on a weather scale, the formation and maintenance of haze days in eastern China ( $HD_{EC}$ ) are closely related to favorable weather conditions (Niu et al., 2010; Cai et al., 2017), including a strong thermal inversion potential, high relative humidity, negative sea level pressure anomaly, and weak wind speed. Furthermore, the anticyclonic anomaly could lead to a sinking movement and a weaker thermal inversion potential, which inhibit the vertical diffusion of pollutants and affect the air quality of the local or larger region (Wu et al., 2013; Xu et al., 2015). Many studies have investigated the key circulation system affecting  $HD_{EC}$  on an interannual scale or intraseasonal scale and suggested that the weak East Asian winter monsoon (Li et al., 2015; Yin et al., 2015; Zhang et al., 2022), the positive phase of Arctic Oscillation (Wang et al., 2015; Yin et al., 2015), and the positive phase of the East Atlantic/West Russia (EA/WR) teleconnection pattern (Yin et al., 2017) could result in more haze days in China. On a synoptic scale, meteorological conditions could also significantly regulate  $HD_{EC}$ . The weak synoptic circulation with a high-pressure or continuous low-pressure system is beneficial for the accumulation of pollution, while the strong weather phenomena with a large pressure gradient encourage the diffusion of pollutants (Li et al., 2019; Cai et al., 2020). Furthermore, studies have shown that cold surges can dissipate and reduce local air pollutants by bringing dry and clean cold air (Wu et al., 2017; Leung et al., 2018; Zhang et al., 2021).

A recent study classified the daily winter circulation anomalies and suggested that there are two dominant climate drivers (i.e., the EA/WR teleconnection pattern and the Victoria mode of sea surface temperature anomalies) conducive to the severe haze occurrence in northern China (Li et al., 2022). Existing studies have also investigated the synoptic circulation patterns conducive to haze pollution in different regions of China (Chang and Zhan, 2017; Li et al., 2019; Liu et al., 2019; Liao et al., 2020; Sun et al., 2020; Yang et al., 2021; Gong et al., 2022). Most of these studies produced the classification based on low-level circulation anomalies, while the upper-level circulation also plays an important role in the generation and accumulation of haze (Wang et al., 2015; Yin and Wang, 2017; Zhong et al., 2019). In addition, due to the large spatial span in EC, if we assess the classification of synoptic circulation patterns in a fixed region, it may lead to

different effects of the same classification pattern in different regions. Therefore, we classify the circulation anomalies with severe haze days of each station in EC, and finally obtain the dominant synoptic atmospheric circulation pattern of each station. In general, the present study addresses the following scientific questions: (1) what are the synoptic atmospheric circulation patterns that dominate severe haze pollution in EC? (2) What are the differences in the action ranges of each circulation pattern? (3) What are their possible mechanisms? These issues are addressed using a modified classification algorithm (hierarchical clustering algorithm, HCA) that is more suitable for studying the classification of synoptic patterns in a large spatial range.

The remaining sections of this paper are structured as follows: data and methods are introduced in Sect. 2. Section 3 shows the dominant synoptic circulation patterns of severe  $HD_{EC}$ . In Sect. 4, we compare different circulation types associated with severe  $HD_{EC}$ . Finally, the discussion and main conclusions are given in Sect. 5.

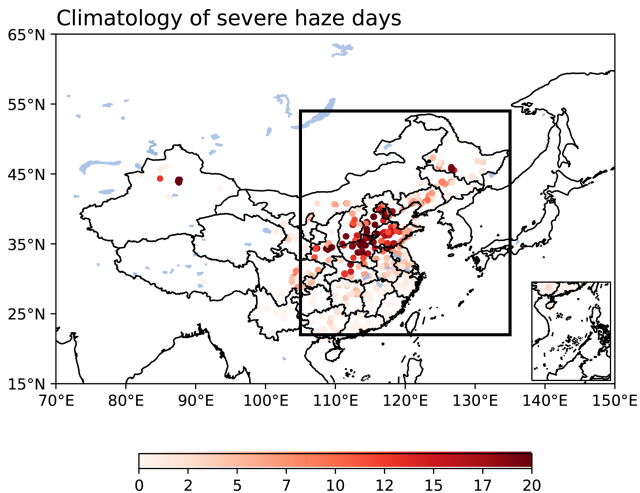
## 2 Data and methods

### 2.1 Data

In this study, the daily meteorological data and the observed  $PM_{2.5}$  concentrations from 2014 to 2021 are used to analyze the dominant circulation patterns and their main causes of severe haze in winter in EC. The daily NCEP/NCAR reanalysis was obtained from <https://psl.noaa.gov/> (last access: 16 May 2022), which includes sea level pressure (SLP), surface air temperature (SAT), the temperature in multiple pressure levels, geopotential height (GPH), three-dimensional wind, relative humidity (RH) at 1000 hPa, and vertical velocity ( $\omega$ ) at 850 hPa (Kalnay et al., 1996). The dataset has a horizontal resolution of  $2.5^\circ \times 2.5^\circ$ . In this study, we define the thermal inversion potential (TIP) as the air temperature at 850 Pa  $T_{850}$  minus SAT, referring to Yin and Wang (2019). The daily  $PM_{2.5}$  concentrations for 935 meteorological stations in China (following Yin and Wang, 2016, and Yin et al., 2021; the stations with missing data of more than 5% are dropped and the stations with data continuously lost for 3 d or more are also discarded) were obtained from the China National Environmental Monitoring Centre (<https://quotsoft.net/air/>, last access: 16 May 2022). The sporadic missing data (less than 3 d) were filled by cubic spline interpolation.

### 2.2 Definition of severe $HD_{EC}$

In this study, severe  $HD_{EC}$  is defined when  $PM_{2.5}$  concentration  $\geq 150 \mu\text{g m}^{-3}$  (Cai et al., 2017; Zhong et al., 2019). We focus on the haze days in the cool season (November to February of the following year, abbreviated as NDJF), which accounts for more than 40% of the total haze days in China in a year (Sun et al., 2013; Wang et al., 2015). Figure 1 shows the climatology of haze days in China from 2014 to 2021 in



**Figure 1.** Spatial distribution of the annual averaged severe haze days (unit: day) in China from 2014 to 2021 in NDJF. The black box represents EC.

NDJF. The severe haze days are mainly concentrated in EC (east of 105° E and south of 54° N), which is selected as the target area in the present study. Thus, a subset of 853 stations is selected.

### 2.3 Definition of blocking index

In winter, the anticyclone anomaly over the Sea of Okhotsk, usually related to atmospheric blocking, may lead to haze accumulation (Yun and Yoo, 2019; Hwang et al., 2022). Thus, based on previous studies (Tibaldi and Molteni, 1990; Fang and Lu, 2020), here we identify the blocking by northward gradients (GHGN) and southward gradients (GHGS) of  $Z_{500}$  at each grid point:

$$\text{GHGN} = \frac{z_{500}(\lambda, \phi + \Delta\phi) - z_{500}(\lambda, \phi)}{\Delta\phi}, \quad (1)$$

$$\text{GHGS} = \frac{z_{500}(\lambda, \phi) - z_{500}(\lambda, \phi - \Delta\phi)}{\Delta\phi}, \quad (2)$$

where  $\phi = 35, 35.5, \dots, 75^\circ \text{N}$  (TS2);  $\lambda = 70, 70.5, \dots, 160^\circ \text{E}$  (TS3); and  $\Delta\phi = 15^\circ$ . A given longitude is defined as “blocked” at a particular time when it satisfies the following conditions:

$$\text{GHGS} > 0, \text{GHGN} < -10 \text{ m}(\text{deg lat})^{-1}.$$

Based on these conditions, we can identify whether any grid in the range of 35–70° N is blocked at any time.

### 2.4 Plumb’s wave activity flux

Here we use the wave flux of Rossby to show the propagation of wave energy (Plumb, 1985). The two-dimensional

Plumb’s wave activity flux (WAF) can be expressed by

$$F_s = \frac{P}{P_0} \cos \varphi \times \begin{pmatrix} v'^2 - \frac{1}{2\Omega a \sin 2\varphi} \frac{\partial(v'\phi')}{\partial\lambda} \\ -u'v' + \frac{1}{2\Omega a \sin 2\varphi} \frac{\partial(u'\phi')}{\partial\lambda} \end{pmatrix}. \quad (3)$$

In Eq. (3),  $F_s$  (unit:  $\text{m}^{-2} \text{s}^{-2}$ ) denotes the horizontal stationary wave activity flux;  $P$  means the pressure;  $P_0 = 1000 \text{ hPa}$ ;  $u'$  and  $v'$  are the zonal and meridional wind deviation, respectively;  $\phi'$  is the geopotential height;  $\varphi$  ( $\lambda$ ) represents the latitude (longitude);  $a$  is the radius of Earth; and  $\Omega$  means Earth’s rotation rate.

### 2.5 Classification algorithm of synoptic atmospheric circulation

This paper uses the hierarchical clustering algorithm (HCA) to classify the severe  $\text{HD}_{\text{EC}}$  based on the associated circulation anomalies. Based on the HCA (Rokach and Maimon, 2005), we could create a clustering tree of data samples by calculating the Euclidean distance between different categories. The original data samples of different types are at the lowest level of the tree, and the root point of a cluster is at the top level of the tree.

Unlike Li et al. (2022), we only cluster the circulation anomalies of days with severe  $\text{HD}_{\text{EC}}$ . It could ensure that the  $\text{PM}_{2.5}$  at least one station in EC of each sample exceeds the standard of severe haze pollution. Secondly, the circulation samples selected are not in a fixed region, but the rectangular regions of the same size are centered on each station with severe haze. Since the upper-level circulation represented by 500 hPa GPH anomalies play an important role in the generation and accumulation of haze (Wang et al., 2015; Yin and Wang, 2017; Zhong et al., 2019), the GPH anomalies at 500 hPa in a rectangular region of 30° from east, west, north, and south with each station as the center on the day of severe  $\text{HD}_{\text{EC}}$  were taken as the samples to perform the HCA. It means that our classification results focus on the local circulation anomalies accompanied by haze, which can help us understand the impact of different local circulation patterns on different stations more accurately. Specifically, this clustering scheme can ensure that each station is located in the center of the circulation pattern when severe haze occurs and avoids the impact of circulation pattern movement. The final composite results of the same pattern can reflect the average statement of the current type of circulation anomaly, which is helpful to investigate its possible physical mechanism.

We use the silhouette coefficient to determine the optimal classification result (Rousseeuw, 1987). For any sample  $i$ , the silhouette coefficient  $s(i)$  is defined as

$$s(i) = \frac{b(i) - a(i)}{\max\{a(i), b(i)\}}, \quad (4)$$

where  $a(i)$  means the average distance from sample  $i$  to all other samples in the cluster it belongs to, and  $b(i)$  means the

lowest average distance from sample  $i$  to all samples in any other cluster. The silhouette coefficient of the clustering result is the average of the silhouette coefficients of all samples. The closer to 1, the better the classification results. Figure S1 shows the clustering tree and its associated silhouette coefficient of this study.

### 3 Dominant synoptic atmospheric circulation patterns of severe HD<sub>EC</sub>

Figure 2a shows the composite anomalies of 500 hPa GPH during all severe HD<sub>EC</sub> in 853 stations. Generally, the stations with severe haze are located in the southwestern parts of the anticyclonic anomaly center, which is consistent with previous studies (Zhong et al., 2019; Wang and Zhang, 2020). We then performed the HCA as described in Sect. 2.5 and obtained three types of dominant local circulation anomalies associated with severe HD<sub>EC</sub> (Fig. 2b, c, d). Circulation Type1 shows a wave-train structure of “+ − +”, and the stations are located in the west of the anticyclonic anomaly and the south of the cyclonic anomaly. Circulation Type2 shows the circulation anomalies similar to Fig. 2a. Finally, circulation Type3 denotes that the stations are located south of the anticyclonic anomaly, and the intensity and range of the anticyclonic anomaly are significantly stronger than those of the other two patterns. The differences between the types imply that severe HD<sub>EC</sub> may be related to different causes.

For each station, when the probability of a certain circulation type is greater than the sum of the other two types, we define this as the dominant type of the station. Figure 3 shows the leading circulation types of severe HD<sub>EC</sub> for 853 stations and the weighted probability density distribution of three circulation types (the weight of each station is the probability of the corresponding dominant type occurring at the station). Stations dominated by circulation Type1 are mainly distributed in the Yangtze River valley (YRV). The stations dominated by circulation Type2 cover almost the whole EC, with two centers in south China (SC) and the Beijing–Tianjin–Hebei region. The stations dominated by circulation Type3 are mainly located in northeast China (NEC). In general, the stations in the north of EC are accompanied by higher PM<sub>2.5</sub> concentrations and more haze days (Fig. S2). These results suggest significant differences in the circulation patterns of severe haze in different regions of EC.

### 4 Comparison of different circulation types associated with severe HD<sub>EC</sub>

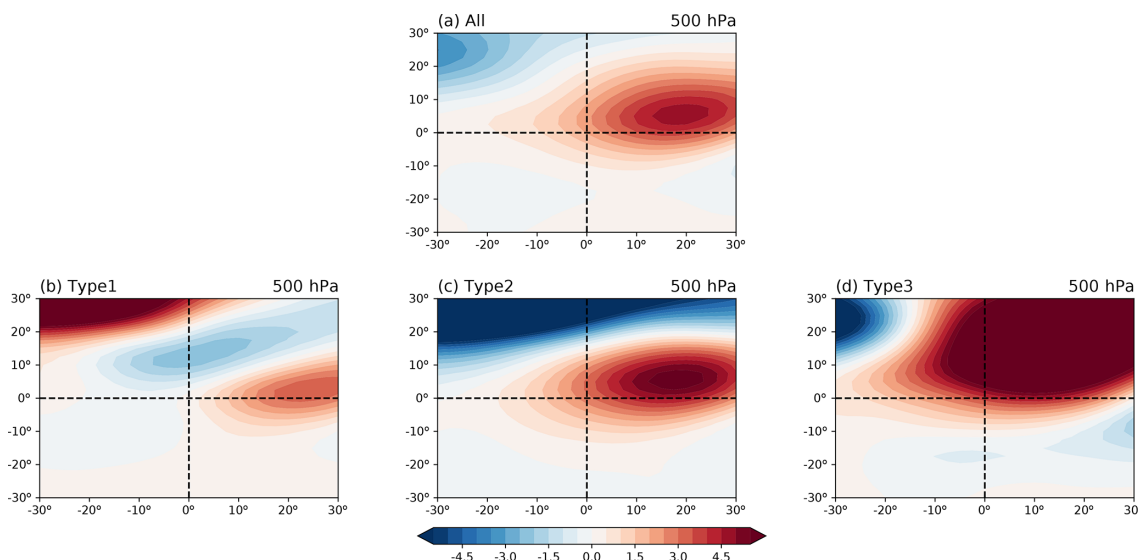
Figure 4a, b, and c show the composite anomalies of circulation Type1 at 500 and 850 hPa. Circulation Type1 is associated with the upper troposphere’s wave-train structure of “− + −”. Unlike previous studies (Zhong et al., 2019; Wang and Zhang, 2020), there are no significant anticyclonic

anomalies in the mid-troposphere over the YRV, but there is a substantial northerly wind component in the lower troposphere over northern China. The TIP, sinking movement, and RH anomalies over the YRV are weak (Fig. 4d, e, f). Therefore, it can be inferred that it is not the local circulation anomalies that promote the formation and accumulation of haze pollution but the regional haze transportation caused by the northerly wind component anomalies that leads to the severe haze in the YRV.

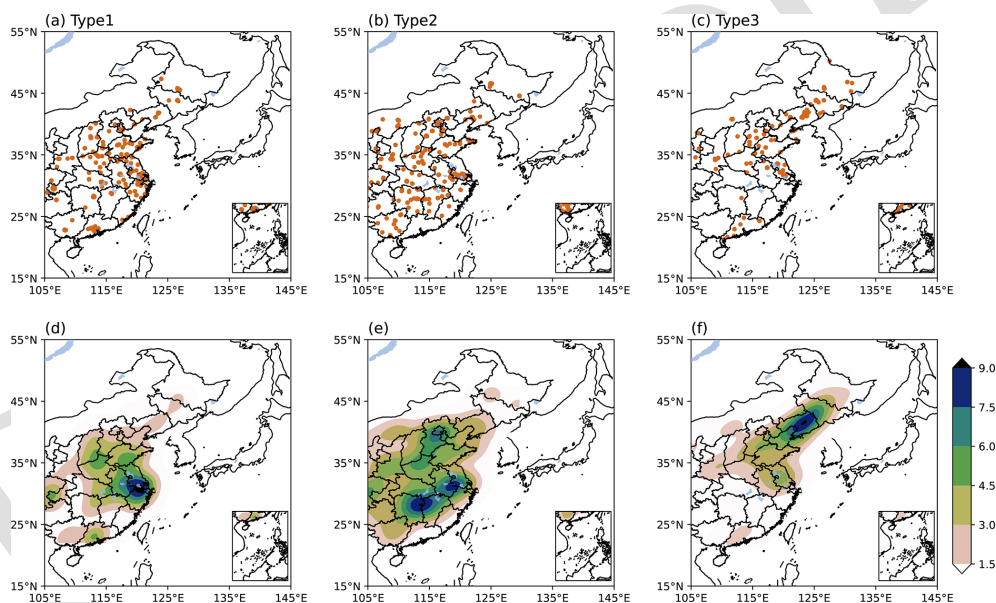
To further explore the relationship between Type1 severe HD<sub>EC</sub> and northerly wind component anomalies, we present the evolution of PM<sub>2.5</sub> concentration variations (PM<sub>2.5</sub> concentration on Day <sub>$i$</sub>  minus that on Day <sub>$i-1$</sub> ) from −3 to 2 d of Type1 severe HD<sub>EC</sub> occurrences (Fig. 5a, b, c, d, e) and the corresponding horizontal wind variations at 500 hPa (Fig. 5f, g, h, i, j). PM<sub>2.5</sub> concentrations tend to increase at first and then dissipate, showing an obvious transportation process from north to south. Accordingly, the horizontal wind changes from anticyclonic anomalies to cyclonic anomalies, with the south wind turning to the north wind. Here we average the PM<sub>2.5</sub> concentration variations in Fig. 5a, b, c, d, e and meridional wind variations in Fig. 5f, g, h, i, j along latitudes (Fig. 5k, l, m, n, o). The result shows that PM<sub>2.5</sub> concentrations gradually increased from north EC to south EC and began to decrease after severe HD<sub>EC</sub> occurred. With the variation in PM<sub>2.5</sub> concentrations, the south wind in north EC gradually weakens and turns to a north wind when severe HD<sub>EC</sub> occurs. With the dry and cold air from the north invading southward, the haze dissipates rapidly, and EC can maintain high air quality weather. Therefore, although circulation Type1 will lead to severe haze in the YRV, its circulation anomalies do not match the conditions to maintain haze pollution.

During the occurrence of circulation Type2, there was an anticyclonic anomaly with a quasi-barotropic structure over northeast Asia, and EC was located in the southwest of the anticyclone (Fig. 6a, b, c). The significant positive TIP, sinking movement, and positive RH anomalies control the region over EC (Fig. 6d, e, f). With the increase in TIP and the warm, humid air from the sea that is transported to EC, the horizontal and vertical dispersion of pollutants was restrained, while higher-surface RH exacerbated the formation of particulates. Such circulation anomalies are beneficial for the formation and maintenance of haze pollution.

Here we investigate the dynamic mechanism of circulation Type2 by compositing the GPH and WAF anomalies in the upper troposphere. The circulation anomalies show two quasi-zonal wave trains over the mid–high latitudes. One is characterized by a “− + − +” pattern of GPH anomalies from the south of Greenland across Siberia to northeast China, with positive GPH anomalies in the second and fourth centers. Such anomalies are similar to the positive phase of EA/WR teleconnection, which can strengthen stable weather conditions over EC (Wu et al., 2016; Yin and Wang, 2016) by causing weak wind speed, higher RH, and strong TIP



**Figure 2.** (a) Composite anomalies of GPH at 500 hPa (units: gpm) during all severe HD<sub>EC</sub> in 853 stations. (0°, 0°) represents the location of stations. Panels (b), (c), and (d) are the same as (a) but for three sub-types.



**Figure 3.** Distribution of stations dominated by (a) Type1, (b) Type2, and (c) Type3 synoptic circulation patterns. Weighted probability density distribution of stations dominated by (d) Type1, (e) Type2, and (f) Type3 synoptic circulation patterns.

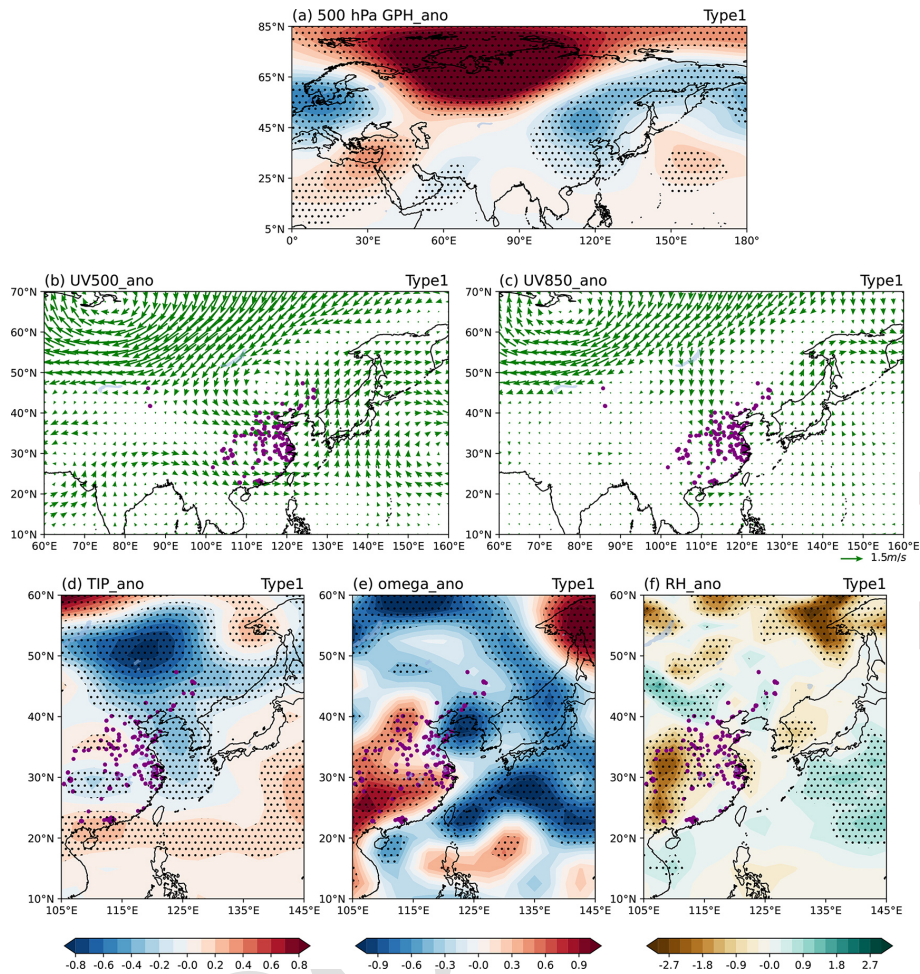
(Niu et al., 2010; Ding and Liu, 2014; Cai et al., 2017). Figure 7c shows the correlation coefficients between PM<sub>2.5</sub> concentrations during the occurrence of circulation Type2 and the EA/WR index. (The EA/WR index was computed by the NOAA Climate Prediction Center according to the rotated principal component analysis used by Barnston and Livezey, 1987.) The results show significant positive correlations between the two in north EC and weak negative correlations in south EC. However, circulation Type2 caused severe HD<sub>EC</sub> for almost the whole EC, which is not completely consistent

with the results of Fig. 7c. Therefore, we speculate that the other wave train may lead to haze pollution in south EC.

It was found that the second wave train reaches EC from Europe along with southern Asia, forming a “+ - + - +” pattern of GPH anomalies. The formation of such a wave train is closely related to the winter East Asia subtropical jet (EASJ) (Xiao et al., 2016; An et al., 2020; Zhang et al., 2022). Here we use an empirical orthogonal function (EOF) analysis of zonal wind from 1980 to 2021 to determine the leading modes of winter EASJ (Xiao et al., 2016). The vari-

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**Figure 4.** Composite anomalies of (a) GPH at 500 hPa (unit: gpm), horizontal wind (unit:  $\text{m s}^{-1}$ ) at (b) 500 hPa and (c) 850 hPa, (d) TIP (unit: K), (e) omega (unit:  $10^{-2} \text{ Pa s}^{-1}$ ), and (f) RH (unit: %) for circulation Type1. Dotted areas are statistically significant at the 95 % confidence level. The purple dots represent the stations dominated by circulation Type1.

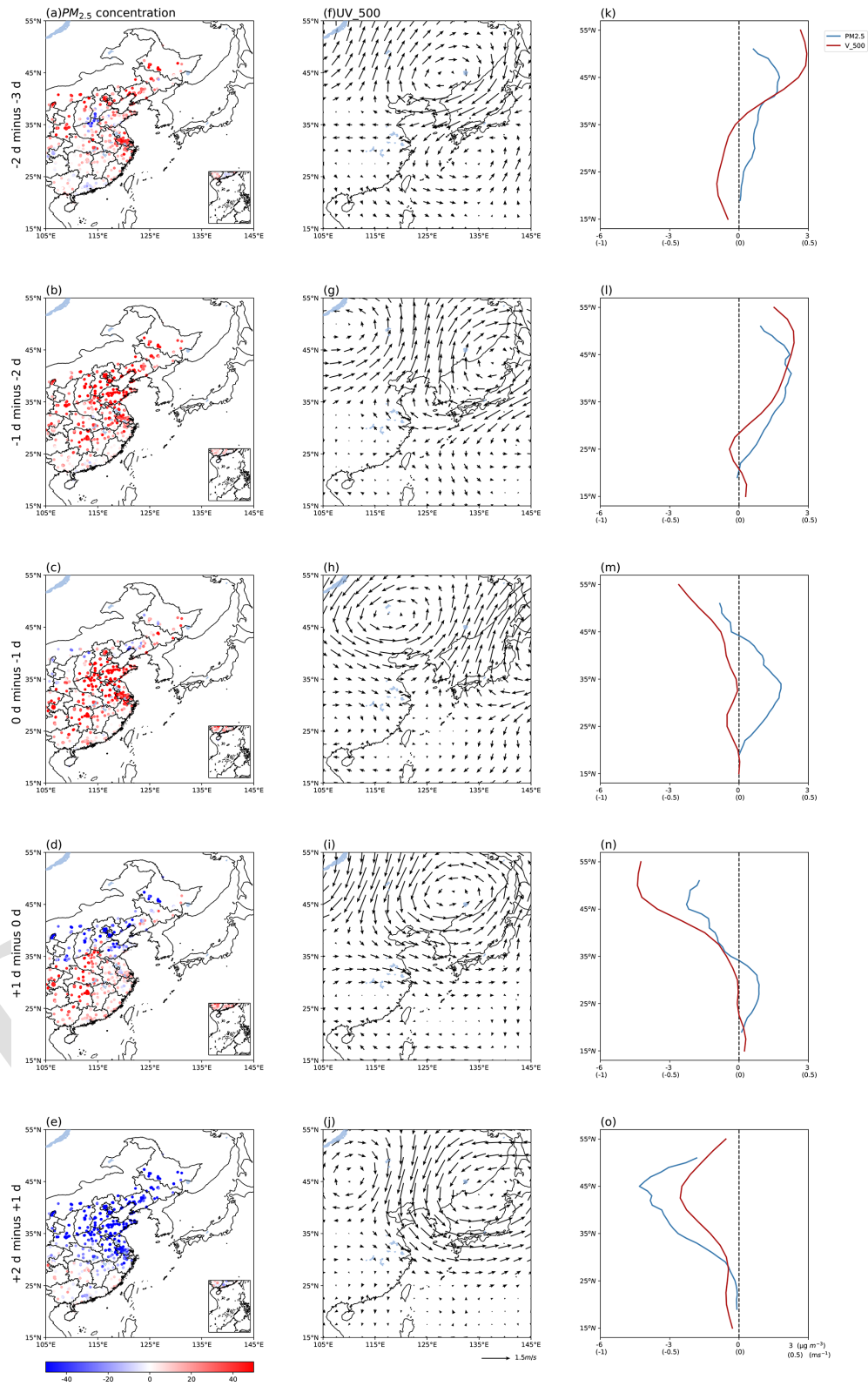
ance of the first mode (EOF1) accounts for 57.4 % of the total variance and indicates the intensity of EASJ (Fig. 8a), which could significantly affect the haze pollution in EC (An et al., 2020; Zhang et al., 2022).

5 The correlation coefficients between daily  $\text{PM}_{2.5}$  concentrations and the first principle component ( $\text{PC1}_{\text{jet}}$ ) during the occurrence of circulation Type2 are shown in Fig. 8b, which have significant positive correlations in south EC and negative correlations in north EC. It indicates that circulation Type2 may cause severe haze pollution in most areas of EC under the joint affection of EA/WR teleconnection and winter EASJ. The results suggest that when discussing the impact of an anticyclonic anomaly in northeast Asia on haze pollution in EC, we should comprehensively consider  
10 the joint affection of signals from high and middle latitudes.

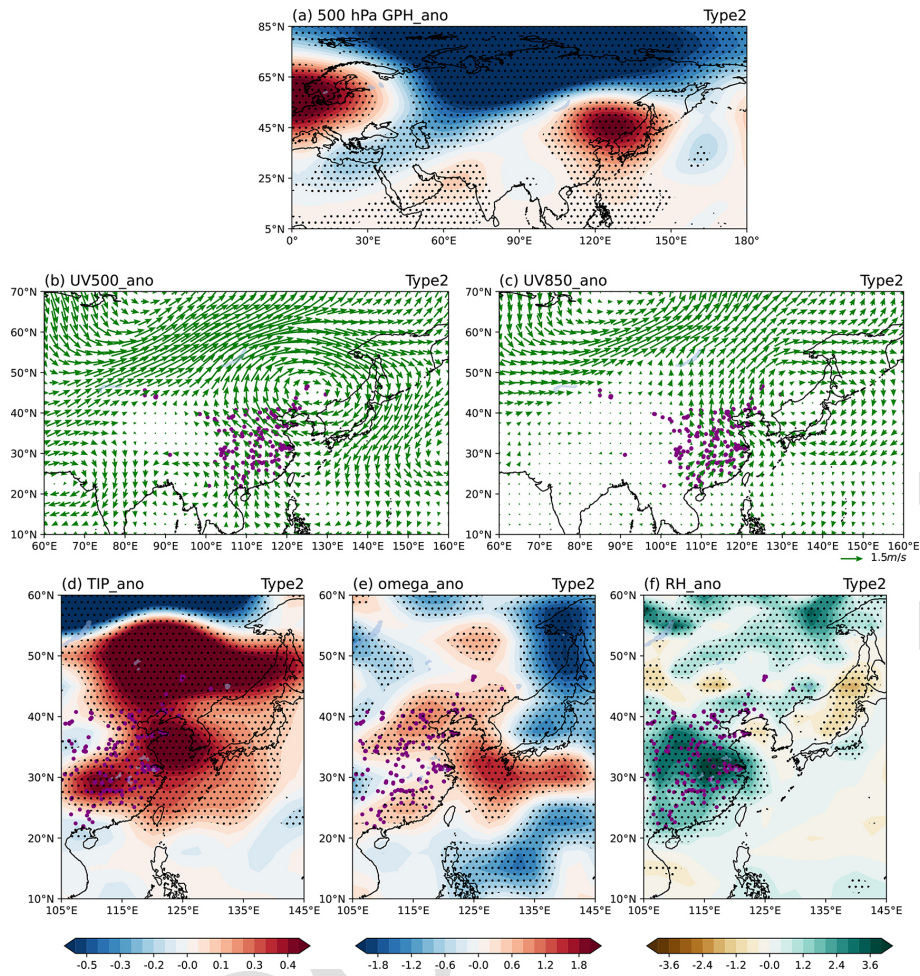
Compared with circulation Type2, the range and intensity of anticyclonic anomalies in northeast Asia in circulation Type3 are more robust, and the location is more northerly (Fig. 9a). Such circulation anomalies lead to southeasterly

wind anomalies at 850 hPa, strong TIP, and abundant moisture that induces severe haze over NEC (Fig. 9d, f). In addition, the ascending motion over south EC and the descending motion over the Beijing–Tianjin–Hebei region and NEC formed meridional circulation cell anomalies (Fig. 9e), which are conducive to the accumulation of severe  $\text{HDE}_{\text{EC}}$   
25 over the NEC.

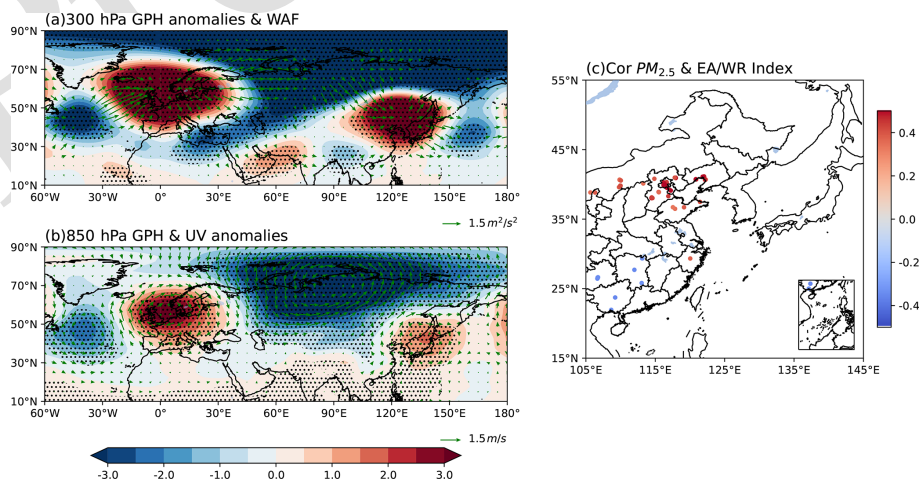
In winter, the anticyclonic anomalies over the Sea of Okhotsk are usually related to atmospheric blocking (Yun and Yoo, 2019; Fang and Lu, 2020; Hwang et al., 2022). Therefore, we calculated the daily atmospheric blocking introduced in Sect. 2.3 to investigate its relationship with Type3 severe  $\text{HDE}_{\text{EC}}$ . Figure 10 shows that when Type3 severe  $\text{HDE}_{\text{EC}}$  occurs, the  $\text{PM}_{2.5}$  concentration increases with the blocking anomalies in the high-latitudes build-up, dissipating with the blocking anomalies' crash. The blocking anomalies strengthen the TIP and provide sufficient RH in the lower atmosphere (Fig. 11), causing severe  $\text{HDE}_{\text{EC}}$  in NEC.  
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**Figure 5.** Composite anomalies of (a–e) the spatial distribution of  $PM_{2.5}$  concentrations (unit:  $\mu g m^{-3}$ ) from  $-3$  to  $2$  d related to Type1 severe  $HD_{EC}$  occurrences and (f–j) the corresponding horizontal wind (unit:  $m s^{-1}$ ) at 500 hPa. Panels (k)–(o) show the zonal averaged  $PM_{2.5}$  concentration variations (unit:  $\mu g m^{-3}$ ) and meridional wind variations (unit:  $m s^{-1}$ ) in the range of  $15$ – $55^\circ N$ ,  $105$ – $135^\circ E$ .

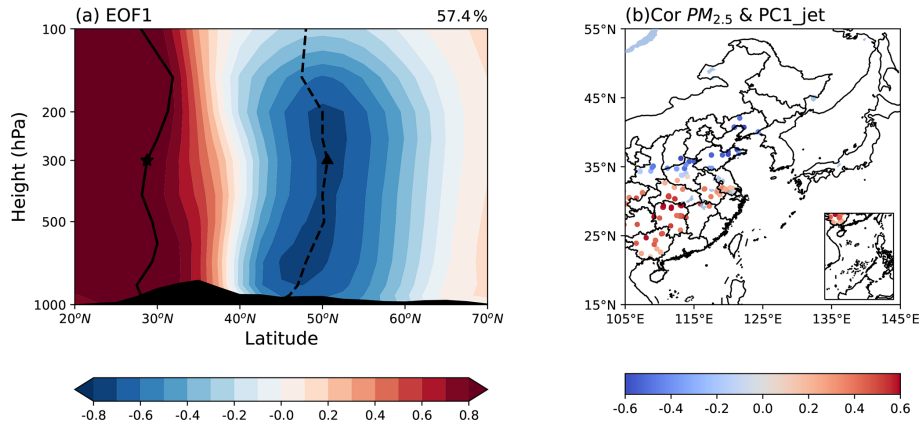


**Figure 6.** Composite anomalies of (a) GPH at 500 hPa (unit: gpm), horizontal wind (unit:  $\text{m s}^{-1}$ ) at (b) 500 hPa and (c) 850 hPa, (d) TIP (unit: K), (e) omega (unit:  $10^{-2} \text{ Pa s}^{-1}$ ), and (f) RH (unit: %) for circulation Type2. Dotted areas are statistically significant at the 95 % confidence level.

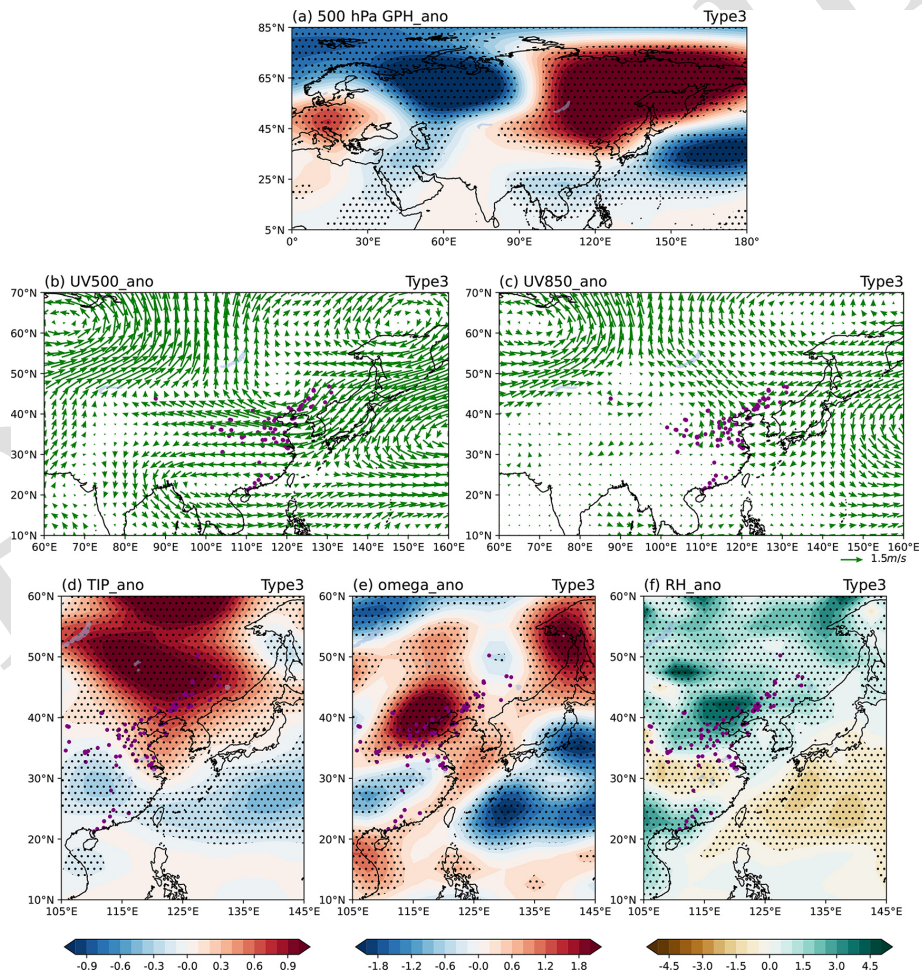


**Figure 7.** Composite anomalies of (a) GPH (shading, gpm) and WAF (vectors,  $\text{m}^2 \text{ s}^{-2}$ ) at 300 hPa, and (b) GPH (shading, gpm) and horizontal wind (vectors,  $\text{m s}^{-1}$ ) at 850 hPa for Type2. Dotted areas are statistically significant at the 95 % confidence level. (c) Correlation coefficients between Type2  $\text{PM}_{2.5}$  concentrations and the EA/WR index.

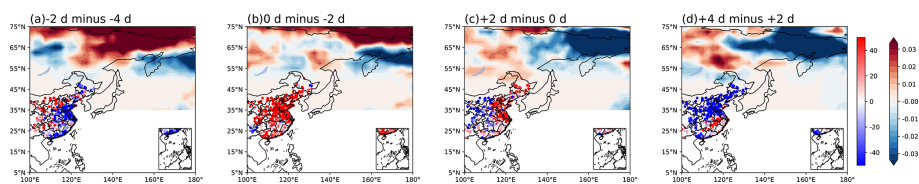




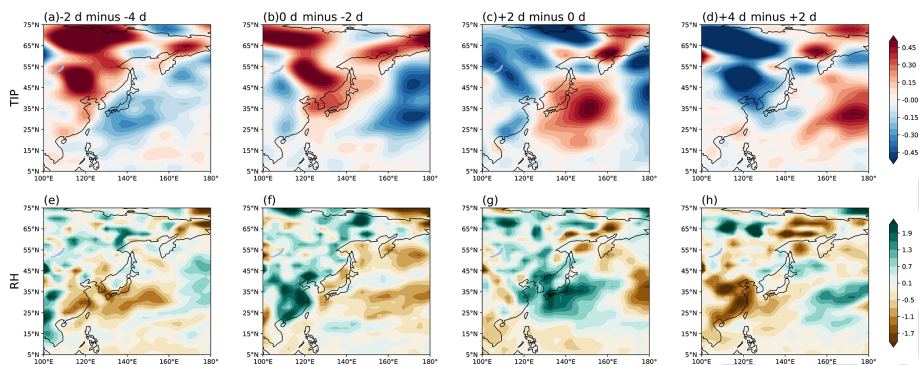
**Figure 8.** (a) The first EOF mode of zonal wind (EOF1,  $\text{m s}^{-1}$ ) averaged from 60 to 160° E in NDJF. The star and triangle at 300 hPa denote the subtropical jet and polar-front jet cores, respectively. The zonal mean orography is dark-shaded. (b) Correlation coefficients between Type2  $\text{PM}_{2.5}$  concentrations and PC1\_jet.



**Figure 9.** Composite anomalies of (a) GPH at 500 hPa (unit: gpm), horizontal wind (unit:  $\text{m s}^{-1}$ ) at (b) 500 hPa and (c) 850 hPa, (d) TIP (unit: K), (e) omega (unit:  $10^{-2} \text{ Pa s}^{-1}$ ), and (f) RH (unit: %) for circulation Type3. Dotted areas are statistically significant at the 95 % confidence level.



**Figure 10.** Composite anomalies of (a–d) the spatial distribution of  $\text{PM}_{2.5}$  concentration variations (unit:  $\mu\text{g m}^{-3}$ ) and blocking from  $-4$  to  $4$  d related to Type3 severe  $\text{HDEC}$  occurrences.



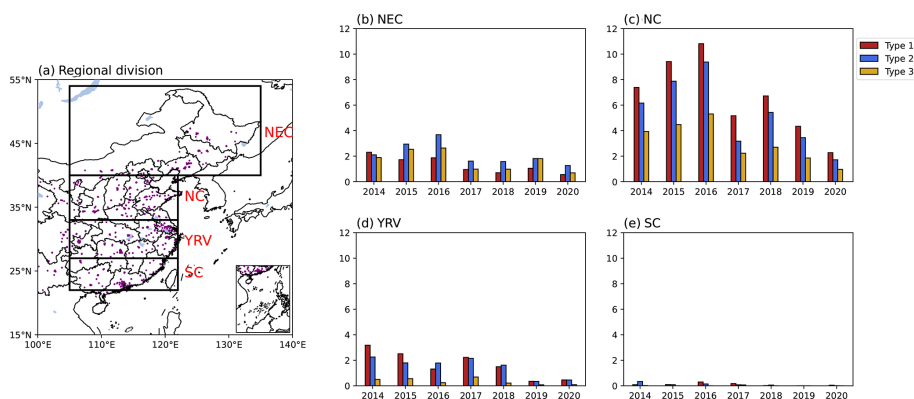
**Figure 11.** Composite anomalies of (a–d) TIP variations (unit: K) and (e–h) RH variations (unit: %) from  $-4$  to  $4$  d related to Type3 severe  $\text{HDEC}$  occurrences.

Based on previous studies and the differences in the influence ranges of the three circulations types in this study, we divided EC into NEC ( $40\text{--}54^\circ\text{N}$ ,  $105\text{--}135^\circ\text{E}$ ), north China (NC;  $33\text{--}40^\circ\text{N}$ ,  $105\text{--}122^\circ\text{E}$ ), the YRV ( $27\text{--}33^\circ\text{N}$ ,  $105\text{--}122^\circ\text{E}$ ), and SC ( $22\text{--}27^\circ\text{N}$ ,  $105\text{--}122^\circ\text{E}$ ) to analyze the temporal characteristics of three  $\text{HDEC}$  types in different subregions of EC (Fig. 12a). Figure 12b, c, d, and e display the annual regional-averaged frequency of the three  $\text{HDEC}$  types in the four subregions. The results show that severe haze pollution mainly occurs in NC and less in SC. The frequency of severe haze generally shows a downward trend in the four subregions.

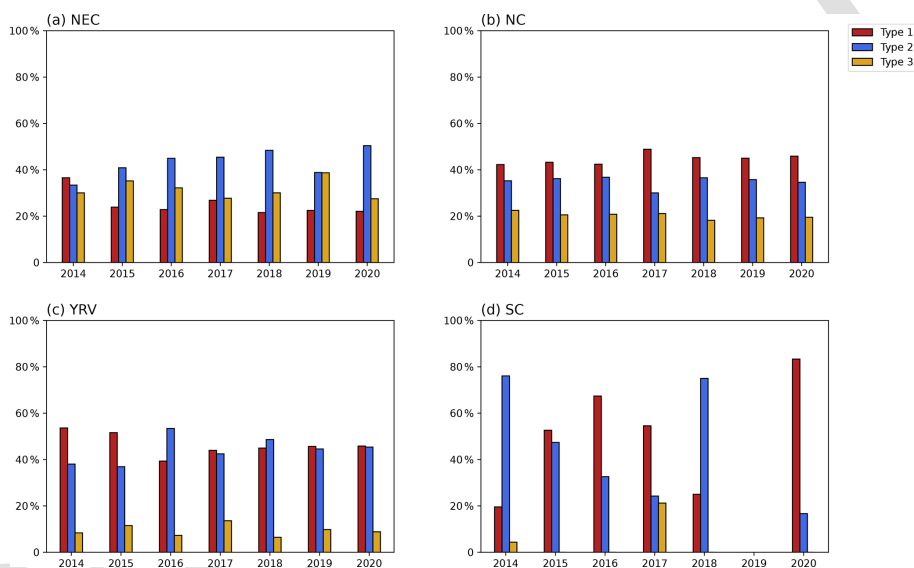
We further calculate the proportion of the frequency of each circulation type in the total annual severe haze frequency in the four subregions (Fig. 13). For NEC, the proportion of the three circulation types is almost equal. It should be noted that the proportion of circulation Type3 is much larger than in the other three subregions. In NC, the proportion of circulation Type1 is more than 40%, while the proportion of circulation Type3 is about 20%. For the YRV, circulation Type1 and Type2 lead the severe haze pollution. There is relatively little severe haze pollution in SC. Therefore, the dominant circulation type in SC has a strong interannual variation and is hardly affected by circulation Type3. Overall, on a weather scale, the  $\text{HDEC}$  is affected by a variety of synoptic circulations, and the areas affected by each synoptic circulation are also different.

## 5 Conclusions and discussion

In this study, the HCA was used to investigate three dominant circulation types that could lead to severe  $\text{HDEC}$ . We clustered the circulations over the stations in EC on severe haze days from 2014 to 2021, which eliminated the interference of the circulations of non-severe haze days on the cluster results. The results show that three dominant circulation types associated with severe  $\text{HDEC}$  are obtained, which are mainly characterized by a local anticyclonic anomaly but also present obvious spatial variation on large-scale circulations. Circulation Type1 with a wave-train structure of “ $- + -$ ” in the upper troposphere mainly causes severe haze pollution in the YRV through the low-level northerly wind anomalies over NC. Although the sinking movement, TIP, and RH anomalies over the YRV are weak or not significant, the regional haze transportation leads to severe haze in the YRV. Circulation Type2 is characterized by two quasi-barotropic Rossby wave trains at 300 hPa, which may be stimulated and sustained by the joint affection of the EA/WR teleconnection and the winter EASJ. One travels from the south of Greenland across Siberia to NEC, forming a “ $- + - +$ ” pattern of GPH anomalies, and the other travels from Europe to southern Asia, forming a “ $+ - + - +$ ” pattern of GPH anomalies, which led to an anticyclonic over northeastern Asia and is conducive to the accumulation of haze. Circulation Type3 is characterized by a blocking anomaly over the Sea of Okhotsk, which influences the severe  $\text{HDEC}$  over NEC with a southeasterly wind at 850 hPa, strong TIP, and abundant moisture. The temporal characteristics of the three cir-



**Figure 12.** (a) The four subregions of EC. The purple dots are the stations. (b–e) Frequency of three types of cool season severe  $HD_{EC}$  in NEC, NC, YRV, and SC.



**Figure 13.** Annual percentage of the three types of severe  $HD_{EC}$  in (a) NEC, (b) NC, (c) the YRV, and (d) SC.

circulation types in NEC, NC, the YRV, and SC were further analyzed. The result shows that on the synoptic scale,  $HD_{EC}$  is affected by various synoptic atmospheric circulations, and the regions affected by each synoptic atmospheric circulation are also different.

The study shows that circulation patterns and key systems that contribute to severe  $HD_{EC}$  are complex and diverse, revealing the dominant circulation patterns of severe haze in different regions of EC. These three dominant atmospheric circulation patterns could potentially be used to establish severe winter haze prediction models for different regions of EC (e.g., project the future variations of severe haze in different regions of EC by identifying similar circulation patterns through machine learning or regression fitting). Due to the limitation of data, it is difficult to carry out the work of circulation classification over a longer period. Therefore, whether there is an interannual or interdecadal connection between

the dominant circulation types of severe haze and its key circulation system needs further investigation. In addition, considering the latitude difference of  $PM_{2.5}$  concentrations in EC and the decreasing of  $PM_{2.5}$  concentrations due to the implementation of the Air Pollution Prevention and Control Action Plan since 2013, the flexible threshold to identify haze days is suggested for use in further studies. We will further carefully compare the impact of emissions and meteorological factors on haze in subsequent work.

This study shows that different circulation types may lead to severe haze in different regions of EC, and further studies are needed to investigate whether there are differences in persistence or intensity among them.

**Data availability.** The daily  $PM_{2.5}$  concentrations for 935 meteorological stations in China are collected from the China National Environmental Monitoring Centre archive at <https://soft.net/air/>

(last access: 16 May 2022). Daily mean meteorological data are obtained from the NCEP/NCAR reanalysis data archive at <http://www.esrl.noaa.gov/psd/data/gridded/data.ncep.reanalysis.html> (last access: 16 May 2022) (Kalnay et al., 1996; NCEP/NCAR, 2022).

5 The monthly EA/WR index (Barnston and Livezey, 1987; CPC, 2022) can be downloaded from NOAA's Climate Prediction Center: <http://www.cpc.ncep.noaa.gov/data/teledoc/telecontents.shtml> (last access: 16 May 2022).

**Supplement.** The supplement related to this article is available online at: <https://doi.org/10.5194/acp-22-1-2022-supplement>.

**Author contributions.** SZ and GZ put forward the conception of this paper, and TW improved the research and manuscript. SZ, XY, and VI carried out the research. SZ wrote the manuscript with contributions from all co-authors.

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