

1 **Measurement report: Changes in light absorption and molecular**  
2 **composition of water-soluble humic-like substances during a winter**  
3 **haze bloom-decay process in Guangzhou, China**

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17

18 **Abstract**

19 Water-soluble humic-like substances (HULIS) absorb light in near-UV and visible  
20 wavelengths and exert significant influence on the atmospheric environment and climate.  
21 However, knowledge on HULIS evolution during haze bloom-decay process is limited.  
22 Herein, PM<sub>2.5</sub> samples were obtained during a winter haze event in Guangzhou, China,  
23 and light absorption and molecular composition of HULIS were investigated by UV-vis  
24 spectrophotometry and ultrahigh-resolution mass spectrometry. Compared with HULIS  
25 in clean days, the absorption coefficients ( $Ab_{S_{365}}$ ) of HULIS in haze days were  
26 significantly higher but the mass absorption efficiencies ( $MAE_{365}$ ) were relatively lower,  
27 suggesting diverse and dynamic absorption properties of HULIS during haze episodes.  
28 The CHO and CHON compounds were the most abundant components in HULIS,  
29 followed by CHOS, CHONS, and CHN. Haze HULIS presented comparatively higher  
30 molecular weight, lower aromaticity index ( $AI_{mod}$ ), and higher  $O/C_w$ ,  $O/N_w$ , and  $O/S_w$   
31 ratios, indicating that HULIS fractions undergo relatively higher oxidation during haze  
32 days than clean days. Moreover, CHON and CHO compounds with high  $AI_{mod}$  were the  
33 major potential chromophores in HULIS and significantly contributed to HULIS light  
34 absorption. It's worth noting that the proportions of these chromophores were decreased  
35 during haze event, mainly owing to their higher oxidation during haze episode. Besides,  
36 accumulated contribution of organic compounds emitted from vehicles and formed from  
37 reactions of bio-VOCs also diluted light-absorbing compounds in haze HULIS. These  
38 findings help to understand HULIS evolution during haze bloom-decay process in the  
39 subtropic region of China.

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## 41 **1. Introduction**

42 Water-soluble humic-like substances (HULIS), belonging to a class of highly  
43 complex organic compounds with physical/chemical properties similar to humic  
44 substances in natural aquatic/soil environments, which constitute 30%–70% of water-  
45 soluble organic compounds in ambient aerosols and are responsible for > 70% of light  
46 absorption (at 365 nm) in water-soluble brown carbon (BrC) (Graber and Rudich, 2006;  
47 Laskin et al., 2015; Huang et al., 2018). They are thought to be comprised of aromatic  
48 structures containing aliphatic side chains and oxygenated functional groups such as  
49 hydroxyl, carboxyl, nitrate, and organosulfate groups (Lin et al., 2012; Song et al., 2018;  
50 Zeng et al., 2020). HULIS are ubiquitously identified in atmospheric aerosols, fog, cloud,  
51 and rain water, and have been demonstrated to play significant effects on both  
52 atmospheric environment and climate (Bianco et al., 2018; Wu et al., 2018; Zeng et al.,  
53 2021). In addition, HULIS exert adverse health effects because they can enhance the  
54 oxidative potential of organic aerosols (Chen et al., 2019; Ma et al., 2019).

55 In recent years, severe particulate pollution (i.e., haze events) frequently occur in  
56 some developing country such as China, which has drawn extensive public and scientific  
57 concerns (Huang et al., 2014; An et al., 2019; Zhang et al., 2020). According to An et al.  
58 (2019), contributions of organic aerosols, including primary organic aerosols and  
59 secondary organic aerosols (SOA), are significant for severe haze events; in particular,  
60 the contribution of SOA in China is expected to continuously increase because of  
61 stronger chemical reactions in the atmosphere. HULIS are an important component in  
62 organic aerosols, which originate from a variety of primary emissions (e.g., biomass  
63 burning (BB), coal combustion, off-road engine emission) (Fan et al., 2016; Cui et al.,

64 2019; Tang et al., 2020) and secondary chemical oxidation of biogenic and anthropogenic  
65 volatile organic compounds (VOCs) (Yu et al., 2016; Tomaz et al., 2018) and soot (Fan  
66 et al., 2020). During the haze episode, a number of chemical processes occur in aqueous  
67 phase (Wong et al., 2017, 2019; Wu et al., 2018) and gas phase (Sumlin et al., 2017),  
68 which lead to significant changes in chemical composition and light absorption properties  
69 of HULIS. For instance, recent studies on oxidation of BB-derived BrC have indicated  
70 that although both enhancement and bleaching of BrC occur during aging, bleaching of  
71 BrC becomes dominant over a long period (Fan et al., 2020; Wong et al., 2017, 2019; Ni  
72 et al., 2021). However, multiphase reaction between carbonyl and amine has  
73 demonstrated rapid formation of light-absorbing organic compounds (Kampf et al., 2016).  
74 Nevertheless, it should be noted that these results were mainly obtained from laboratory  
75 experiments and may not reflect the complex evolution behavior of BrC in atmospheric  
76 environment.

77 High concentrations of HULIS have been determined during typical haze episodes  
78 in northern, eastern, and southern China (Fan et al., 2016; Win et al., 2020; Zhang et al.,  
79 2020; Wang et al., 2020), and have been demonstrated to significantly influence  
80 atmospheric visibility, environment, and photochemical process. Guangzhou is the  
81 biggest city in the Pearl River Delta (PRD), one of the most developed regions in China,  
82 and is located in the subtropical zone with a population of over 18 million people (Yu et  
83 al., 2017). Although a remarkable decline in atmospheric particulate matter (PM<sub>2.5</sub>)  
84 pollution has been observed in recent years owing to strict regulatory controls, O<sub>3</sub> and  
85 VOCs still remain at higher levels and severe haze pollution caused by fine particulate  
86 matter frequently occur in winter (Huang et al., 2014; An et al., 2019; Li et al., 2019;

87 Yang et al., 2022). Several studies have investigated the optical, chemical, and molecular  
88 properties of HULIS in the PRD region (Lin et al., 2010, 2012; Fan et al., 2016; Liu et al.,  
89 2018; Jiang et al., 2020, 2021a,b). For example, the studies on the temporal variations of  
90 water-soluble HULIS in Guangzhou indicated that HULIS had higher concentrations and  
91 mass absorption efficiencies ( $MAE_{365}$ ) in the winter, which were attributed to the  
92 increasing contribution of BB and secondary nitrate formation in the winter monsoon  
93 period (Fan et al., 2016; Jiang et al., 2020, 2021a). In addition, the molecular composition  
94 of HULIS (and BrC) in the PRD region were also investigated and demonstrated that the  
95 levels of unsaturated and aromatic structures are the important factor influencing their  
96 light absorption properties (Jiang et al., 2020, 2021b). However, detailed information  
97 regarding the evolution of light absorption and molecular composition of HULIS during  
98 haze events is still scarce.

99 Recently, ultrahigh-resolution Fourier transform ion cyclotron resonance mass  
100 spectrometry (FT-ICR MS) coupled with electrospray ionization (ESI) sources has been  
101 frequently employed to investigate the molecular characteristics of HULIS in ambient  
102 aerosols (Song et al., 2018, 2022; Tang et al., 2020; Zeng et al., 2021). Owing to its  
103 extremely high mass resolution and accuracy, this technique allows further exploration of  
104 the evolution of HULIS during haze event. The present study performed comprehensive  
105 characterization of HULIS in  $PM_{2.5}$  collected during a haze event in Guangzhou, China.  
106 The abundances and light absorption properties of HULIS were first measured, and  
107 carbonaceous fractions, water-soluble ions, and levoglucosan (Lev) were determined.  
108 Subsequently, four HULIS samples collected during different haze stages were analyzed  
109 using FT-ICR MS operated in both ESI<sup>-</sup> and ESI<sup>+</sup> modes. To the best of our knowledge,

110 the present study is the first to apply a combination of optical properties and molecular  
111 characterization by FT-ICR MS to investigate HULIS in a haze event in the subtropical  
112 zone of China. The results obtained provide novel insights into the evolution of HULIS  
113 during haze event, and are important for predicting the environmental and climatic effects  
114 of HULIS in South China.

## 115 **2. Material and Methods**

### 116 **2.1. Aerosol sampling**

117 The PM<sub>2.5</sub> samples were collected on the campus of Guangzhou Institute of  
118 Geochemistry, Chinese Academy of Sciences, Guangzhou, China (23.14N, 113.35E),  
119 which is an academic and residential region. Traffic emissions and residential activities  
120 are the potential pollution sources in the sampling area. The 24-h PM<sub>2.5</sub> sampling was  
121 conducted using a high-volume sampler (Tianhong Intelligent Instrument Plant, Wuhan,  
122 China, with a flow rate of 1.0 m<sup>3</sup> min<sup>-1</sup>) during 7 to 30 January of 2018, and a total of 24  
123 samples were collected on the prebaked quartz filters (20.3 × 25.4 cm<sup>2</sup>, Whatman,  
124 Maidstone, UK). Field blank samples were collected by keeping a blank filter in the  
125 sampler without pumping air. Before sampling, the filters were wrapped in aluminum foil  
126 and prebaked at 450 °C for 6 h to remove carbonaceous impurities. Before and after  
127 sampling, the filters were weighed at 25 °C and 50% RH on a microbalance (Sartorius  
128 Model BP210D), with an accuracy of 0.01 mg. The PM<sub>2.5</sub> concentrations were  
129 determined by weighing the filters before and after collection. Finally, all filter samples  
130 were stored in a refrigerator at -20 °C until analysis. Meteorological data  
131 (<http://www.wunderground.com/history/airport/ZGGG>), including wind speed,

132 temperature, relative humidity, and concentrations of SO<sub>2</sub>, O<sub>3</sub>, and NO<sub>2</sub>, for the sampling  
133 days are presented in Figure 1 and Table S1.

## 134 **2.2. Isolation of HULIS**

135 HULIS were isolated using a water extraction and solid-phase extraction (SPE)  
136 procedure as described previously (Zou et al., 2020). This method has been used in most  
137 previous studies because of its easy operation and high reliability and reproducibility and  
138 low limit of detection (Fan et al., 2002), therefore, it was also used in this study. Briefly,  
139 portions of the PM<sub>2.5</sub> samples (100 cm<sup>2</sup>) were ultrasonically extracted with 50 mL of  
140 ultrapure water for 30 min. The extracts were filtered through a 0.22- $\mu$ m PTFE syringe  
141 filter to remove the suspended insoluble particles. About 50 mL of water extracts were  
142 obtained from each sample, of which 20 mL was used for the isolation and analysis of  
143 HULIS, 20 mL for analysis of water-soluble organic carbon (WSOC), and the remained  
144 extracts for the analysis of inorganic ions, respectively. Then, the 20 mL water extracts  
145 were adjusted to pH of 2 with HCl, and loaded on a preconditioned SPE cartridge (Oasis  
146 HLB, 200 mg/6 mL, Waters, USA). The hydrophilic fraction (i.e., inorganic ions, high-  
147 polar organic acids, etc) was removed with ultrapure water, whereas the relatively  
148 hydrophobic HULIS fraction was retained and eluted with 2% (v/v) ammonia/methanol.  
149 Finally, HULIS solution was evaporated to dryness with a gentle N<sub>2</sub> stream and  
150 redissolved with ultrapure water for the analysis.

151 It is noted that the HULIS here is the hydrophobic portion of water-soluble organic  
152 matter, which can be isolated with different types of SPE columns (e.g., HLB, C-18,  
153 DEAE, XAD-8, and PPL) (Fan et al., 2012, 2013; Lin et al., 2012; Zou et al., 2020; Jiang  
154 et al., 2020; Qin et al., 2022). Although each resin type has its special chemical properties,

155 the hydrophobic HULIS isolated with different sorbents were similar in chemical,  
156 molecular properties based on previous studies (Fan et al., 2012, 2013; Zou et al., 2020).  
157 Therefore, for better comparison with other studies, the hydrophobic fractions isolated by  
158 SPE methods were all termed as HULIS in the present paper.

### 159 **2.3. Light absorption analysis**

160 The absorption spectra of the WSOC and HULIS fractions were measured by a UV-  
161 vis spectrophotometer (UV-2600, Shimadzu) between 200 to 700 nm. Each spectrum was  
162 corrected for the filter blanks. The light absorption coefficients, absorption Ångström  
163 exponent (AAE) and mass absorption efficiency ( $MAE_{\lambda}$ ) were calculated and the detailed  
164 methods are presented in the Supporting Information (SI).

### 165 **2.4. Chemical analysis**

166 For FT-ICR MS analysis, the HULIS samples were isolated from  $PM_{2.5}$  collected  
167 during four periods: before haze days (clean-I days, 7–12 January), haze bloom days  
168 (haze-I days, 13–18 January), haze decay days (haze-II days, 19–24 January), and after  
169 haze days (clean-II days, 25–30 January). A filter punch (18 cm in diameter) was taken  
170 from every sample, and all the six samples in each period was combined for the isolation  
171 of HULIS fractions. The obtained HULIS samples were measured with an ESI FT-ICR  
172 MS (Bruker Daltonik GmbH, Bremen, Germany) equipped with a 9.4 T refrigerated  
173 actively shielded superconducting magnet. The system was operated in both ESI<sup>-</sup> and  
174 ESI<sup>+</sup> modes. The scan range was set to  $m/z$  from 100 to 1000, with a typical mass-  
175 resolving power >450,000 at  $m/z$  319 with <0.2 ppm absolute mass error. The mass  
176 spectra were calibrated externally with arginine clusters and internally recalibrated with



177 typical O<sub>5</sub>-class species peaks in DataAnalysis 4.4 (Bruker Daltonics). Due to the  
178 inherent differences in the ionization mechanisms between ESI- and ESI+ modes, the  
179 data detected by the two ionization modes can provide complementary information on the  
180 molecular composition of atmospheric HULIS (Lin et al., 2012; Lin et al., 2018). The  
181 details of data analysis are provided in the SI.

182 The amounts of organic carbon (OC) and elemental carbon (EC) were determined by  
183 a OC/EC analyzer (Sunset Laboratory Inc., USA) (Mo et al., 2018). The concentrations  
184 of WSOC and HULIS were determined by a TOC analyzer (Shimadzu TOC\_VCPH,  
185 Kyoto, Japan). The water-soluble inorganic species (NO<sub>3</sub><sup>-</sup>, SO<sub>4</sub><sup>2-</sup>, Cl<sup>-</sup>, NH<sub>4</sub><sup>+</sup>, K<sup>+</sup>, Na<sup>+</sup>,  
186 Ca<sup>2+</sup>, Mg<sup>2+</sup>,) were measured with a Dionex ICS-900 ion chromatography system (Thermo  
187 Fisher Scientific, USA) as described previously (Huang et al., 2018). The concentrations  
188 of Lev were analyzed with a gas chromatography–MS after derivatization with BSTFA  
189 and pyridine at 70 °C for 3 h (Huang et al., 2018). Detailed information regarding these  
190 measurements is provided in the SI.

### 191 **3. Results and Discussion**

#### 192 **3.1. Abundance and chemical composition of PM<sub>2.5</sub>**

193 Figure 1 shows the meteorological conditions, PM<sub>2.5</sub> concentration, and  
194 concentrations of major chemical constituents, including carbon fractions and water-  
195 soluble inorganic ions in PM<sub>2.5</sub> samples obtained during a haze bloom-decay process.  
196 Based on the variation in PM<sub>2.5</sub> concentration, these samples were categorized into four  
197 groups: clean-I days (before haze, 14–24 µg m<sup>-3</sup>), haze-I days (haze bloom, 45–114 µg  
198 m<sup>-3</sup>), haze-II days (haze decay, 58–115 µg m<sup>-3</sup>), and clean-II days (after haze, 9–35 µg

199  $\text{m}^{-3}$ ). As indicated in Table S1 and Figure 1, the  $\text{PM}_{2.5}$  concentrations increased from 18  
200  $\pm 3.3 \mu\text{g m}^{-3}$  in clean-I days to  $82 \pm 26$  and  $84 \pm 22 \mu\text{g m}^{-3}$  in haze-I and haze-II days,  
201 respectively, and then decreased to  $21 \pm 10 \mu\text{g m}^{-3}$  in clean-II days. This finding  
202 obviously indicated that the average  $\text{PM}_{2.5}$  concentrations during the examined haze  
203 episode are higher than the second-grade national ambient air quality standard in China  
204 ( $75 \mu\text{g m}^{-3}$ , 24 h), whereas those during clean days are lower than the first-grade national  
205 ambient air quality standard in China ( $35 \mu\text{g m}^{-3}$ , 24 h). However, the average  $\text{PM}_{2.5}$   
206 concentrations during the haze event are lower than those in the cities in winter haze,  
207 including Shenyang ( $108 \mu\text{g m}^{-3}$ ) (Zhang et al., 2020), and Nanjing ( $123 \pm 28.5 \mu\text{g m}^{-3}$ )  
208 (Li et al., 2020), Beijing ( $158 \mu\text{g m}^{-3}$ ) and Xi'an ( $345 \mu\text{g m}^{-3}$ ) (Zhang et al., 2018).

209 As shown in Table S1, the average concentrations of OC and EC were 2.2–15 and  
210  $0.36\text{--}2.7 \mu\text{gC m}^{-3}$  in the four stages, respectively, implying that the distinct changes in  
211 OC and EC were higher during haze episodes than those in clear days. During the entire  
212 study period, WSOC concentration ranged from 0.5 to  $12.5 \mu\text{gC m}^{-3}$  ( $4.3 \pm 1.2 \mu\text{gC m}^{-3}$ ),  
213 which contributed to 53%–57% of OC in  $\text{PM}_{2.5}$ . The HULIS concentration noted in the  
214 present study ranged from 0.15 to  $6.1 \mu\text{gC m}^{-3}$  ( $2.2 \pm 1.9 \mu\text{gC m}^{-3}$ ), which was  
215 comparable to those observed in the PRD region, such as Hong Kong ( $2.38 \pm 1.62 \mu\text{gC}$   
216  $\text{m}^{-3}$ ) (Ma et al., 2019), Guangzhou ( $2.4 \pm 1.6 \mu\text{gC m}^{-3}$ ) (Fan et al., 2016), and Heshan  
217 ( $2.08 \pm 1.16 \mu\text{gC m}^{-3}$ ) (Jiang et al., 2020), but lower than those in northern cities of  
218 China, such as Xi'an ( $12.4 \pm 6.5 \mu\text{gC m}^{-3}$ ) (Huang et al., 2020), Beijing ( $3.79 \pm 3.03$   
219  $\mu\text{gC m}^{-3}$ ) (Mo et al., 2018), and Lanzhou ( $4.7 \mu\text{gC m}^{-3}$ ) (Tan et al., 2016). As shown in  
220 Figure 1, HULIS also exhibited obvious variations during the entire sampling period. The  
221 average HULIS concentration was  $0.46 \pm 0.22 \mu\text{gC m}^{-3}$  in clean-I days, which sharply

222 increased to  $4.5 \pm 1.2 \mu\text{gC m}^{-3}$  in haze-I days, then decreased to  $3.1 \pm 1.2 \mu\text{gC m}^{-3}$  in  
223 haze-II days, and rapidly declined to  $0.75 \pm 0.52 \mu\text{gC m}^{-3}$  in clean-II days. This result  
224 was consistent with the changing trend of WSOC, OC, and EC. In addition, the  
225 HULIS/WSOC ratios were about  $0.50 \pm 0.13$  in the  $\text{PM}_{2.5}$  samples, which are in broad  
226 agreement with other studies showing that HULIS is the major fraction of WSOC (Fan et  
227 al., 2016; Ma et al., 2019; Jiang et al., 2020).

228 As illustrated in Figure 1, obvious variations in chemical compositions were also  
229 observed in these  $\text{PM}_{2.5}$  samples. Secondary inorganic aerosols (SIA) (i.e.,  $\text{SO}_4^{2-}$ ,  $\text{NO}_3^-$ ,  
230 and  $\text{NH}_4^+$ ), OC, and EC exhibited a similar variation during the entire study period, and  
231 their contents sharply increased from 10 January in clean-I days to 13–18 January in  
232 haze-I days, then slowly decreased in haze-II days, and finally reached lower levels in  
233 clean-II days. It must be noted that the increasing rate of EC was similar to that of SIA in  
234 haze-I days, indicating that direct emissions and atmospheric reactions may play similar  
235 roles in  $\text{PM}_{2.5}$  increase during this haze bloom period. As indicated in Figure 1f, the  
236 highest values of  $\text{NO}_3^-/\text{SO}_4^{2-}$  were observed in haze-I days, implying the important  
237 influence of traffic exhausts in the haze bloom period (Mo et al., 2018). In addition, the  
238 high  $\text{NO}_2$  and  $\text{O}_3$  concentrations and the stable meteorological condition with high  
239 temperature also led to the outburst of fine particulate pollution in this period. During  
240 haze-II days, the SIA and OM contents in  $\text{PM}_{2.5}$  slowly decreased, whereas the  
241 concentrations of  $\text{Na}^+$ ,  $\text{Cl}^-$ , and unidentified materials in  $\text{PM}_{2.5}$  increased (Figure 1e,h),  
242 suggesting that local contribution weakened and regional contribution via sea salt became  
243 more important (Jiang et al., 2021a). This phenomenon was also observed to be  
244 consistent with the changes in the pollutant sources transported by air masses. As

245 indicated in Figure S1, the PM<sub>2.5</sub> samples in haze-II days included some contributors  
246 transported from coastal area of eastern Guangdong Province and Fujian Province, and  
247 the PM<sub>2.5</sub> are likely to be enriched with sea salt materials and mineral dusts.

### 248 **3.2. Light absorption**

249 The light absorption properties of WSOC and HULIS (Figure 1d, i, j and Table S2)  
250 exhibited obvious temporal variations during the sampling period. The AAE values of  
251 WSOC and HULIS ranged from 4.1 to 6.4 and 5.6 to 6.6, respectively. The AAE values  
252 for HULIS were obviously higher than those for WSOC in the same sample (Figure 1i),  
253 indicating that light absorption of HULIS is more wavelength-dependent than that of  
254 WSOC. This difference may be related with the light-absorbing organic species in the  
255 isolated HULIS fractions have strong wavelength dependence than those in the original  
256 WSOC. Moreover, the AAE values of HULIS did not present significant variation during  
257 the entire haze process.

258 Light absorption at 365 nm ( $Abs_{365}$ ) for WSOC and HULIS were  $2.5 \pm 2.0$  and  $1.8 \pm$   
259  $1.6 \text{ M m}^{-1}$ , respectively (Table S2). HULIS contributed to about 72% of light absorption  
260 coefficients by WSOC, implying that they enriched the major light-absorbing  
261 components in WSOC. As shown in Figure 1d, the  $Abs_{365}$  values for HULIS presented  
262 obvious temporal variations. The  $Abs_{365,HULIS}$  value was  $0.55 \pm 0.06 \text{ M m}^{-1}$  in clean-I  
263 days, which first increased to  $3.4 \pm 1.5 \text{ M m}^{-1}$  in haze-I days and then slowly decreased to  
264  $2.6 \pm 0.85 \text{ M m}^{-1}$  in haze-II days, and finally rapidly declined to  $0.64 \pm 0.32 \text{ M m}^{-1}$  in  
265 clean-II days. This result was similar to the variations in the mass concentration of  
266 HULIS. Furthermore, the  $Abs_{365}$  values for HULIS in Guangzhou were found to be  
267 higher than those observed in southeastern Tibetan Plateau ( $0.38\text{--}1.0 \text{ M m}^{-1}$ ) (Zhu et al.,

268 2018), but obviously lower than those in Xi'an ( $7.6\text{--}36 \text{ M m}^{-1}$ ) (Shen et al., 2017) and  
269 Beijing, ( $3.7\text{--}10.1 \text{ M m}^{-1}$ ) (Du et al., 2014).

270 In general,  $\text{MAE}_{365}$  value can be used to assess the light absorption capacity of target  
271 organic compounds (Li et al., 2019). As shown in Figure 1j and Table S2, the average  
272  $\text{MAE}_{365}$  value for WSOC was  $1.0 \pm 0.21 \text{ m}^2 \text{ gC}^{-1}$  ( $0.68\text{--}1.3 \text{ m}^2 \text{ gC}^{-1}$ ), nearly same to  
273  $1.1 \pm 0.27 \text{ m}^2 \text{ gC}^{-1}$  ( $0.77\text{--}1.8 \text{ m}^2 \text{ gC}^{-1}$ ) for HULIS, during the entire sampling period.  
274 Moreover, the  $\text{MAE}_{365}$  values for HULIS measured in the present study were noted to be  
275 dropped in the ranges of those determined in Beijing ( $1.43 \pm 0.33 \text{ m}^2 \text{ g C}^{-1}$ ) (Mo et al.,  
276 2018), Xi'an ( $0.91\text{--}1.85 \text{ m}^2 \text{ g C}^{-1}$ ) (Yuan et al., 2021), and Hong Kong ( $1.84 \pm 0.77 \text{ m}^2$   
277  $\text{gC}^{-1}$ ) (Ma et al., 2019). The average  $\text{MAE}_{365}$  values for HULIS exhibited some temporal  
278 variations. The  $\text{MAE}_{365}$  values for HULIS were  $0.91 \pm 0.03$  and  $0.95 \pm 0.11 \text{ m}^2 \text{ gC}^{-1}$  in  
279 haze-I and haze-II days, respectively, which were lower than those ( $1.3 \pm 0.22$  and  $1.3 \pm$   
280  $0.27 \text{ m}^2 \text{ gC}^{-1}$ , respectively) observed in clean-I and clean-II days, suggesting that HULIS  
281 have a relatively weaker light absorption capability in haze days. This finding is  
282 consistent with the results reported by Zhang et al. (2017), who found that the  $\text{MAE}_{365}$   
283 values in the heating or non-heating seasons during hazy days were lower than those in  
284 clean days. These differences in  $\text{MAE}_{365}$  values may potentially contribute to the  
285 enhanced oxidation reaction that was derived by the increased  $\text{O}_3$  levels and high  
286 temperature and relative humidity (RH) during haze days (Figure 1). This oxidation  
287 process would lead the chromophores containing C=C unsaturated bond to be severely  
288 degraded (Wang et al., 2017a; Zhang et al., 2017). Besides, an increase in additional  
289 sources for HULIS in the study area, such as weaker or non-light-absorbing compounds

290 formed by atmospheric oxidation, could also result in weaker light absorption of HULIS  
291 during the haze episode (Liu et al., 2018).

### 292 **3.3. Molecular evolution of HULIS during the haze process**

293 For an in-depth understanding of the variation in HULIS at molecular level during  
294 the haze process, the four HULIS samples collected in different stages of the haze  
295 process were analyzed by ESI FT-ICR MS in both negative and positive modes. As  
296 shown in Figure 2, thousands of peaks were detected in the mass range between  $m/z$  100  
297 and  $m/z$  700, with the high intensity ions noted within  $m/z$  150–400. It is obvious that  
298 some organic compounds with stronger arbitrary abundance were labeled, and their  
299 formulas, double bond equivalent (DBE), modified aromaticity index ( $AI_{\text{mod}}$ ), and  
300 potential sources were listed in Table S3. Compounds a ( $C_7H_7NO_3$ ) and b ( $C_8H_6O_4$ ), both  
301 have high DBE values, which might be assigned to aromatics such as methylnitrophenol  
302 and phthalic acid, whereas compound d ( $C_8H_{18}O_4S$ ) with low DBE value and high O/S  
303 ratio was probably aliphatic organosulfate. According to previous studies, these organic  
304 molecules might be derived from BB and diesel fuel and thereby these results suggested  
305 that both BB and vehicular emissions are important sources of BrC in ambient aerosols  
306 (Mohr et al., 2013; Riva et al., 2015; Blair et al., 2017). Furthermore, compound e  
307 ( $C_{10}H_{17}NO_7S$ ) and compound f ( $C_{10}H_{18}N_2O_{11}S$ ) in Table S3 were found to be identical to  
308 the oxidation products of monoterpenes, suggest that biogenic sources could contribute to  
309 the formation of HULIS (Surratt et al., 2008; Wang et al., 2019). Thus, HULIS could be  
310 affected by multiple sources during the haze process, possibly including BB, biogenic  
311 sources, and anthropogenic emissions.

312 The identified formulas could be divided into seven compound categories, namely,  
313 CHO<sup>-</sup>, CHON<sup>-</sup>, CHOS<sup>-</sup>, and CHONS<sup>-</sup> detected in ESI<sup>-</sup> mode and CHO<sup>+</sup>, CHN<sup>+</sup>, and  
314 CHON<sup>+</sup> detected in ESI<sup>+</sup> mode. As illustrated in Figure 2, the CHO compounds were the  
315 most abundant group in all the HULIS, accounting for 43%–50% and 51%–57% of the  
316 overall compounds detected in the ESI<sup>-</sup> and ESI<sup>+</sup> modes, respectively. It must be noted  
317 that relatively lower contents of CHO<sup>-</sup> were detected during the haze episode (haze-I and  
318 haze-II days) and CHO<sup>+</sup> molecules in haze-I HULIS. The CHON compounds were the  
319 second most abundant group in all the HULIS. As shown in Figure 2, the relative content  
320 of CHON<sup>-</sup> was 23% in clean-I days, which slightly increased to 24%–25% in haze  
321 episode, and then decreased to 23% in clean-II days. In contrast, the relative content of  
322 CHON<sup>+</sup> compounds was 41% in clean-I days, which increased to 45% in haze-I days,  
323 then fell to 42% in haze-II days and 41% in clean-II days. Both CHOS<sup>-</sup> and CHONS<sup>-</sup>  
324 compounds were identified in all the four HULIS, accounting for 19%–22% and 8%–11%  
325 of the total identified compounds, respectively. The CHN<sup>+</sup> compounds were the least  
326 abundant (1.3%–3.6%) in the four HULIS samples, and were relatively higher during the  
327 haze episode, especially in haze-I days.

328 Tables S4 and S5 show the relative abundance weighted elemental ratios, molecular  
329 weight (MW), DBE, AI<sub>mod</sub>, and carbon oxidation state (OS<sub>C</sub>) for the identified  
330 compounds in HULIS. The MW<sub>w</sub> values for HULIS determined in the ESI<sup>-</sup> mode in  
331 haze-I and haze-II days were 302 and 283, respectively, which were higher than those in  
332 clean-I and clean-II days (266 and 264, respectively). Similar variation was also observed  
333 for MW<sub>w</sub> for HULIS detected in ESI<sup>+</sup> mode (Table S5). These results clearly indicated  
334 that more higher MW compounds constituted HULIS obtained during the haze episode. It

335 has been reported that the low MW compounds (provided by size exclusion  
336 chromatography) are more susceptible to atmospheric oxidation processes, while the high  
337 MW compounds have relatively higher chemical resistance for BB aerosols (Di Lorenzo,  
338 et al., 2017; Wong et al., 2017; Dasari et al., 2019). Although the HULIS samples in this  
339 study were more complex than those in BB aerosols, it is expected that the high MW  
340 molecules mostly were the recalcitrant fraction in HULIS. Therefore, the HULIS  
341 compounds undergo higher oxidation during haze episode, and are thereby characterized  
342 by relatively high MW values.

343 Furthermore, the molecular properties of HULIS in different stages of haze process  
344 also exhibited some observable differences. As shown in Table S4, the HULIS samples in  
345 haze episode detected by ESI<sup>-</sup> mode presented relatively lower  $AI_{\text{mod,w}}$  values and  
346 relatively higher  $O/C_w$ ,  $O/N_w$ , and  $O/S_w$  ratios than those in clean days, indicating that  
347 haze HULIS exhibited relatively lower aromaticity and higher oxidation degree than  
348 clean HULIS. These differences can be attributed to the enhanced oxidation degradation  
349 of aromatic compounds (e.g., phenols, nitroaromatic compounds and polycyclic aromatic  
350 hydrocarbons (PAHs)) during the haze process. In addition, increased contribution from  
351 traffic emission and secondary reactions of bio-VOCs also decreased the aromaticity and  
352 increased the oxidation degree of HULIS (Liu et al., 2016; Tang et al., 2020). These  
353 changes in HULIS compounds led to the decrease in their  $MAE_{365}$  values during the haze  
354 episode, as described above ( Zhong and Jang, 2014; Song et al., 2019).

### 355 **3.3.1. CHO Compounds**

356 The CHO compounds bear O-containing functional groups, and have been  
357 frequently detected in ambient aerosols. As shown in Figure 2, the CHO compounds were



358 the predominant component in the four HULIS samples, and the  $MW_w$  values for CHO<sup>-</sup>  
359 and CHO<sup>+</sup> compounds were 247–288 and 236–272, respectively, with relatively higher  
360  $MW_w$  values observed for the CHO group (CHO<sup>-</sup> and CHO<sup>+</sup>) in haze HULIS, especially  
361 in haze-I samples. This finding may be related to the stronger oxidation of HULIS during  
362 haze days, because the aqueous oxidation of biomass burning aerosols was found to yield  
363 high MW of organic products (Tomaz et al., 2018; Yu et al., 2016).

364 The  $OS_C$  is often used to describe the degree of oxidation of organic species in the  
365 atmosphere (Kroll et al., 2011; Tong et al., 2019). Figure 3 shows plots of  $OS_C$  versus  
366 carbon number for the CHO compounds. As indicated in the figure, CHO compounds  
367 exhibited  $OS_C$  from  $-2$  to  $+1$  with up to 40 carbon atoms. Kroll et al. (2011) proposed  
368 that compounds with  $OS_C$  between  $-0.5$  and  $+1$  and  $< 18$  carbon atoms can be attributed  
369 to semi-volatile and low-volatile oxidized organic aerosols (SV-OOA and LV-OOA),  
370 which are mainly formed by complex oxidation reactions in atmosphere. Compounds  
371 with  $OS_C$  between  $-0.5$  and  $-1.5$  and 6–23 carbon atoms are related to primary biomass  
372 burning organic aerosol (BBOA). In addition, compounds with  $OS_C$  between  $-1$  and  $-2$   
373 and  $\geq 18$  carbon atoms have been suggested to be hydrocarbon-like organic aerosols  
374 (HOA), which are regarded as primary combustion surrogate (Zhang et al., 2005; Kroll et  
375 al., 2011; Wang et al., 2017b).

376 As illustrated in Figure 3 and Table S6, most of the CHO<sup>-</sup> compounds clustered in  
377 the BBOA region, accounting for 40%–46% of the total CHO<sup>-</sup> compounds, thus  
378 suggesting that BB may be a major contributor to CHO compounds in HULIS. Figure 3  
379 clearly indicates that the majority of aromatic and condensed aromatic compounds  
380 produced signals in the  $OS_C$  region between  $-0.5$  and  $1.0$  and carbon number of 3–18

381 (Figure 3), which corresponded to SV-OOA and LV-OOA. The proportions of SV-OOA  
382 and LV-OOA accounted for 23%–28% and 1.9%–2.4% of the total CHO<sup>-</sup> compounds,  
383 respectively, and presented no significant variation. In contrast, the HOA components in  
384 haze-I days showed the highest abundance (18%), which were much higher than those  
385 (3.5%–4.5%) in haze-II, clean-I, and clean-II days. This finding indicated that the  
386 increase in the primary source is associated with fossil fuel combustion such as vehicle  
387 emissions during the haze bloom period (Zhang et al., 2005).

388 As shown in Figure 3, CHO<sup>+</sup> compounds presented lower OS<sub>C</sub> (from -2.0 to 1.0)  
389 than CHO<sup>-</sup> compounds. Most of the CHO<sup>+</sup> compounds occurred in the BBOA region in  
390 all four HULIS samples, making up to 60%–72% of the total CHO<sup>+</sup> compounds, which  
391 again suggesting that BB is an important contributor to CHO compounds in HULIS. The  
392 HOA among CHO<sup>+</sup> compounds showed the same changing trends as those among CHO<sup>-</sup>  
393 compounds, and higher HOA abundance was observed during haze-I days. In addition,  
394 some high AI<sub>mod</sub> values of aromatics were found in the regions A1+ and A2+ (Figure 3),  
395 which implied that the highest AI<sub>mod</sub> values (AI ≥ 0.67) with DBE ≥ 22 were only  
396 detected during the haze days possibly owing to soot-derived materials or oxidized PAHs  
397 (Decesari et al., 2002; Kuang and Shang, 2020). It must be noted that the sampling site in  
398 the present study is influenced by traffic sources, the enhanced oxidation of vehicle-  
399 exhausted soot also results in the accumulation of water-soluble high aromatic organic  
400 species (Decesari et al., 2002).

### 401 **3.3.2. CHON Compounds**

402 In the present study, 1379–2217 and 2008–2943 formulas were assigned to CHON  
403 compounds identified in the ESI<sup>-</sup> and ESI<sup>+</sup> spectra, respectively, which accounted for

404 23%–25% (ESI–) and 41%–45% (ESI+) of total identified compounds, respectively.  
405 Relatively higher contents of CHON– compounds were obviously detected in HULIS  
406 samples obtained during haze-I days, suggesting the occurrence of more N-containing  
407 components in HULIS during haze bloom days. As shown in Tables S4 and S5, the  
408 average  $MW_w$  values for CHON– and CHON+ compounds were 328 and 317 in haze-I  
409 days, respectively, which were slightly higher than those determined in haze-II days and  
410 all higher than those observed in clean-I and clean-II days. Meanwhile, the  $AI_{mod,w}$  values  
411 for CHON– in haze days were 0.31–0.34, which were slightly lower than those in clean  
412 days (0.37 and 0.40). These findings indicated that more high MW CHON compounds  
413 with lower aromatic structures were formed during the haze episode.

414 The  $O/N_w$  ratios for CHON– and CHON+ during haze-I and haze-II days were 5.3–  
415 5.7 and 3.8, respectively, which were higher than those determined during the two clean  
416 periods, confirming that these compounds were highly oxidized during the haze episode  
417 (Tables S4 and S5). In general, compounds with  $O/N \geq 3$  may indicate oxidized N groups  
418 such as nitro ( $-NO_2$ ) or nitrooxy ( $-ONO_2$ ), whereas compounds with  $O/N < 3$  may denote  
419 the reduced N compounds (i.e., amines) (Lin et al., 2012; Song et al., 2018). In the  
420 present study, most of the CHON compounds (79%–91% of CHON– compounds and  
421 61%–64% of CHON+ compounds) exhibited  $O/N \geq 3$ , suggesting that high  
422 concentrations of nitro compounds or organonitrates were contained in the CHON  
423 compounds. Moreover, these compounds were more abundant in the CHON– group  
424 during the haze episode (87%–91%), when compared with those during clean-I and  
425 clean-II days (79%–82%), again implying that CHON– compounds undergo relatively  
426 higher oxidization during the haze episode. As indicated in Figure 1, the increase in  $NO_2$

427 was consistent with increased production of highly oxidized N-containing organic  
428 compounds (NOCs) during the haze episode, which suggested the significant contribution  
429 of NO<sub>3</sub>-related multigenerational chemistry to organonitrate aerosol formation  
430 (Berkemeier et al., 2016).

431 The majority of aromatics and condensed aromatics produced clear signals in  
432 regions associated with SV-OOA and LV-OOA (Figure 4). BBOA also constituted a  
433 significant proportion (33%–39%) in the CHON<sup>-</sup> group, and a relatively lower BBOA  
434 content was observed in haze-I days. The abundance of HOA was relatively lower,  
435 accounting for 2.3%–7.8% of the total CHON compounds, and the relative abundance of  
436 HOA in haze-I days was much higher than that in haze-II, clean-I, and clean-II days,  
437 suggesting the accumulation of primary fossil fuel combustion during haze-I days.

438 The CHON<sup>+</sup> compounds mainly occurred at the range of  $-2.0 < OS_C < 1.5$ , with  
439 average OS<sub>C</sub> values of around  $-1.0$  for each sample, clearly indicating that CHON<sup>+</sup>  
440 compounds were relatively lower than CHON<sup>-</sup> compounds. Most of the CHON<sup>+</sup>  
441 compounds were detected in the BBOA region, accounting for 60%–76% of the total  
442 CHON<sup>+</sup> compounds. The relative contribution of BBOA in haze-I days was lower than  
443 that in haze-II and clean days. Moreover, a large number of aromatic species were  
444 observed at the region B1<sup>+</sup> (Figure 4), demonstrating that higher aromatic compounds  
445 were only detected in haze-I days, which may be related to soot or BC. Similar trend was  
446 also exhibited by CHO<sup>+</sup> compounds, indicating the contribution of local combustion  
447 sources (e.g., traffic emission) during haze-I days.

### 448 3.3.3. CHOS and CHONS Compounds

449 In this study, 478–696 CHOS compounds and 306–589 CHONS compounds were  
450 identified in ESI– mode (Table S4). Among these S-containing compounds, >86% of the  
451 CHOS compounds had O/S ratios >4, whereas > 89% of the CHONS compounds  
452 presented O/S ratios >7, suggesting that these S-containing compounds were possibly  
453 organosulfates and nitrooxyorganosulfates. As listed in Table S4, the  $AI_{mod,w}$  values for  
454 CHOS and CHONS were about 0.02 and 0.01 in the HULIS fraction, which were much  
455 lower than those for CHO and CHON. Almost 99% of the CHOS and CHONS  
456 compounds in the HULIS fraction had  $AI_{mod}$  values <0.5, while >93% of the CHONS  
457 compounds had  $AI_{mod} = 0$ , indicating that they were mainly comprised of aliphatic and  
458 olefinic organosulfates. These results are consistent with the previous findings that the  
459 major S-containing compounds among organic aerosols in Guangzhou are organosulfates  
460 formed by secondary oxidation reaction of long-chain alkenes/fatty acids with  $SO_2$  (Jiang  
461 et al., 2020), which generally possessed long aliphatic carbon chains and a higher degree  
462 of oxidation. However, these compounds are different from the S-containing compounds  
463 detected during the hazy days in Beijing (Jiang et al., 2016; Mo et al., 2016), which were  
464 determined to be aliphatic organosulfates with low degree of oxidation and higher  
465 amounts of aromatics and PAH-derived organosulfates, having a strong correlation with  
466 anthropogenic emissions.

467 As described earlier, CHOS– and CHONS– compounds might be related to  
468 organosulfates or nitrooxyorganosulfates, which have been observed to be derived from  
469 atmospheric reactions of bio-VOCs such as  $\alpha$ -pinene, limonene, and isoprene (Huang et  
470 al., 2018; Surratt *et al.*, 2008) and fossil fuel combustion including coal combustion, off-  
471 road engine emissions (Song et al., 2018, 2019; Cui et al., 2019). In the present study, the

472 relative contents of S-containing compounds (CHOS+CHONS) in the HULIS fraction in  
473 haze days were all higher than those in clean days (Figure 2). Moreover, the CHOS and  
474 CHONS compounds in haze HULIS always have relatively high relatively O/S ratios  
475 than those in clean HULIS. These findings suggested the relatively higher contribution of  
476 SO<sub>2</sub>-related chemical oxidation during the haze event.

#### 477 **3.3.4. CHN Compounds**

478 The N-bases (CHN) are usually identified in ambient aerosols and smokes from BB.  
479 In the present study, 110–165 CHN<sup>+</sup> compounds were identified in ESI<sup>+</sup> mode, with  
480 most of them (>86%) presenting DBE ≥ 2, suggesting that they might be nitrile and  
481 amine species (Lin et al., 2012). As shown in Figure 2, the abundances of CHN<sup>+</sup>  
482 compounds were 2.0%–3.6% in the haze days, which were much higher than those noted  
483 in clean days (1.3%–1.4%), indicating higher contribution of CHN<sup>+</sup> compounds to the  
484 HULIS fraction during the haze episode. The MW<sub>w</sub> values for CHN<sup>+</sup> compounds were  
485 204–223, which were lower than those for the other groups (i.e., CHO<sup>+</sup>, CHON<sup>+</sup>) (Table  
486 S5). However, the average AI<sub>mod</sub> values for N-bases (0.37–0.48) detected in the ESI<sup>+</sup>  
487 mode were much higher than those for CHO<sup>+</sup> (0.11–0.12) and CHON<sup>+</sup> (0.20–0.22)  
488 compounds, implying that these reduced CHN<sup>+</sup> compounds exhibited more unsaturated  
489 or aromatic structures.

490 To further understand the molecular distribution of CHN<sup>+</sup> compounds during the  
491 haze process, van Krevelen (VK) diagrams were constructed by plotting the H/C ratio  
492 versus N/C ratio (Figure S2). It was obvious that this plot could separate the compound  
493 classes with different degree of AI. As shown in Figure S2, compounds (denoted in black  
494 color) in the upper region of the VK diagram had one N atom with DBE = 0, indicating

495 that they are aliphatic amines. It can be noted from Table S7 that the aliphatic group  
496 presented the lowest abundance in all the samples, suggesting that the CHN<sup>+</sup> compounds  
497 possessed comparatively lower aliphatic structures. Olefinic compounds showed the  
498 highest abundance in the four samples, which accounted for 37%–51% of the total CHN<sup>+</sup>  
499 compounds. Importantly, a large proportion of the compounds (>39%) exhibited high  
500 degree of AI (AI > 0.5) (Figure S2 and Table S7), suggesting a large amounts of aromatic  
501 structure and N-heterocyclic ring in HULIS. Moreover, the CHN<sup>+</sup> compounds in haze-I  
502 days presented obviously lower content of aromatic structures than those in haze-II,  
503 clean-I, and clean-II days, signifying the relatively high contribution of fossil fuel  
504 combustion (which generally emits more low-aromatic CHN compounds) during the haze  
505 bloom episode(Song et al., 2022). In addition, the CHN<sup>+</sup> group also constituted a large  
506 proportion of BBOA (Table S6), which indicated the significant contribution of BB.  
507 However, it must be noted that a relatively lower content of BBOA was detected during  
508 haze-I days, which was consistent with the changing trends of CHON<sup>-</sup> or CHON<sup>+</sup>  
509 compounds during the haze episode. These results suggested the relatively lower  
510 contribution of BB during haze-I days, because quiet and stable weather conditions can  
511 prevent regional transport of BB sources during this stage (Wu et al., 2018).

#### 512 **3.4. Factors influencing light absorption and molecular characteristics of HULIS** 513 **during the haze bloom-decay process**

514 As described earlier, the light absorption properties of HULIS exhibited obvious  
515 variation during the haze bloom-decay process. The average Abs<sub>365</sub> value for HULIS was  
516  $0.55 \pm 0.06 \text{ M m}^{-1}$  in clean-I days, which first increased to  $3.4 \pm 1.5 \text{ M m}^{-1}$  in haze-I days,  
517 then slowly decreased to  $2.6 \pm 0.85 \text{ M m}^{-1}$  in haze-II days, and finally rapidly declined to

518  $0.64 \pm 0.32 \text{ M m}^{-1}$  in clean-II days. In general, the light absorption of HULIS can be  
519 related to their chemical and molecular properties that are influenced by factors such as  
520 sources, secondary formation, and aging process. The results of principal component  
521 analysis (PCA) obviously showed a positive loading for principal component 1 (PC1),  
522 and the  $\text{Abs}_{365}$  values for HULIS were clustered with EC,  $\text{K}_{\text{bb}}^+$ , Lev,  $\text{NH}_4^+$ , and  $\text{NO}_3^-$   
523 (Figure 5). These results suggested that BB and other sources such as new particle  
524 formation could contribute to light absorption of HULIS (Huang et al., 2014; An et al.,  
525 2019; Song et al., 2019). Similarly, the findings of Pearson correlation coefficient  
526 analysis revealed that the  $\text{Abs}_{365}$  values for HULIS exhibited significant positive  
527 correlations with  $\text{K}_{\text{bb}}^+$  ( $r = 0.728$ ,  $p < 0.01$ ) and Lev ( $r = 0.800$ ,  $p < 0.01$ ) (Table S8). As  
528 Lev and  $\text{K}_{\text{bb}}^+$  are generally considered as tracers derived from BB, these results suggested  
529 the significant contribution of BB to light absorption of HULIS. This observation was  
530 also supported by the abundance of BBOA compounds detected in all the four HULIS  
531 samples (Table S6). The significant positive relationships between the  $\text{Abs}_{365}$  values for  
532 HULIS and secondary ions (i.e.,  $\text{NO}_3^-$  ( $r = 0.702$ ,  $p < 0.01$ ),  $\text{SO}_4^{2-}$  ( $r = 0.554$ ,  $p < 0.05$ ),  
533 and  $\text{NH}_4^+$  ( $r = 0.899$ ,  $p < 0.01$ )) indicated the important impact of secondary formation on  
534 the light absorption of HULIS. Besides, the  $\text{Abs}_{365}$  values for HULIS were also strongly  
535 correlated with  $\text{NO}_2$ ,  $\text{O}_3$ , and  $\text{NO}_2$ , which confirmed the important impact of atmospheric  
536 oxidation reactions on the light absorption of HULIS.

537 It must be noted that  $\text{MAE}_{365}$  is a key parameter signifying the light absorption  
538 ability of HULIS. As listed in Table S2, the  $\text{MAE}_{365}$  values for HULIS varied in different  
539 stages, and were lower in haze days owing to the variation in the chemical and molecular  
540 composition of HULIS during the haze bloom-decay process. Furthermore, the  $\text{AI}_{\text{mod}}$



541 values for HULIS varied in different stages (Tables S4), and were relatively lower in  
542 haze days, indicating that haze HULIS have comparatively lower degree of conjugation  
543 or aromaticity. This finding suggested that the HULIS compounds may undergo higher  
544 oxidation during the haze episode, causing a decline in chromophores and reduction in  
545 the light absorption capacity of HULIS (Lin et al., 2017). Besides, the accumulated  
546 contribution of organic compounds from vehicle emission and secondary chemical  
547 reactions of bio-VOCs may also dilute light-absorbing compounds in haze HULIS (Tang  
548 et al., 2020; Liu et al., 2016).

549 Lin et al. (2018) reported that potential light-absorbing chromophores can be  
550 determined in the region between  $DBE = 0.5 \times C$  (linear conjugated polyenes  $C_xH_yC_2$ )  
551 and  $DBE = 0.9 \times C$  (fullerene-like hydrocarbons). In the present study, most of the high-  
552 intensity CHON, CHO, and CHN compounds with high AI values were clustered in  
553 potential BrC chromophore region (Figures S3 and S4), which mainly comprised CHON  
554 (46%–50% in ESI- mode and 56%–62% in ESI+ mode, respectively) and CHO (44%–48%  
555 in ESI- mode and 29%–38% in ESI+ mode, respectively) compounds (Table 1).  
556 Although the contribution of CHN+ compounds to BrC was relatively lower, the content  
557 of potential chromophores among the total CHN+ compounds was higher than those in  
558 CHON+ and CHO+ compounds. Therefore, these three groups of light-absorbing  
559 compounds (i.e., CHON+, CHN+, and CHO+ compounds) were further examined. As  
560 shown in Table 1, the  $Int_C/Int_{BrC}$  values of CHO- (i.e., content of CHO- chromophores  
561 in the total chromophores) decreased from 48% to 44% whereas the  $Int_C/Int_{BrC}$  values of  
562 CHON- increased from 46% to 50% during the haze bloom process. These findings  
563 indicated that more NOCs chromophores were formed during this stage in which higher

564 NO<sub>2</sub> concentration may be preferred for the formation of N-containing chromophores  
565 such as nitrophenols. However, it must be noted that the proportions of both CHO<sup>-</sup> and  
566 CHON<sup>-</sup> chromophores among the total identified compounds decreased from clean-I to  
567 haze-I days, suggesting the occurrence of stronger photo-bleaching process during the  
568 haze bloom stage (Zeng et al., 2020). Likewise, both CHO<sup>+</sup> and CHON<sup>+</sup> compounds  
569 presented similar variation during the entire study period. In addition, the CHN<sup>+</sup>  
570 compounds also exhibited higher Int<sub>C</sub>/Int<sub>BrC</sub> values during the haze bloom process and  
571 suggesting the accumulated contribution from local combustion process. Furthermore, the  
572 proportion of CHON<sup>+</sup> chromophores in the total CHON<sup>+</sup> compounds increased with the  
573 decreasing content of CHN<sup>+</sup> chromophores, may implying that some aromatic CHN<sup>+</sup>  
574 compounds were transformed to CHON<sup>+</sup> compounds during the aging process.

575

#### 576 **4. Conclusions**

577 This study investigated the evolution of light absorption and molecular properties of  
578 HULIS during a winter haze bloom-decay process, and examined the key factors  
579 affecting the light absorption of HULIS in Guangzhou, China. The results showed that  
580 HULIS exhibited significant variation in light absorption during the haze bloom-decay  
581 process. First, higher Abs<sub>365</sub> values were observed in haze days, indicating the presence  
582 of significant amounts of light-absorbing organic compounds during the haze episode.  
583 However, the MAE<sub>365</sub> values for HULIS in haze days were relatively lower than those in  
584 clean days, suggesting the light absorption capabilities of HULIS were weakened during  
585 the haze event. Furthermore, CHON and CHO compounds, exhibiting relatively higher  
586 degree of conjugated structure, were the most abundant groups in all the HULIS samples,

587 and were also the major contributors to light absorption capacity of HULIS. Importantly,  
588 the molecular properties of HULIS dynamically varied during the entire haze episode.  
589 When compared with HULIS in clean days, those in haze days presented relatively lower  
590  $AI_{\text{mod}}$  values and higher  $O/C_w$ ,  $O/N_w$ , and  $O/S_w$  ratios, suggesting the predominance of  
591 compounds with low aromaticity and higher oxidation in HULIS during haze episode.  
592 These results indicated that HULIS compounds undergo relatively stronger oxidation  
593 during the haze days. Moreover, PCA and Pearson correlation analysis revealed that BB  
594 and secondary chemical formation both contributed to the variation in the light absorption  
595 properties of HULIS. Both primary sources (such as accumulated contribution of organic  
596 compounds formed from local traffic emission) and secondary sources (such as stronger  
597 chemical reactions) led to the rapid increase in HULIS during the haze bloom days.  
598 However, stronger oxidation of HULIS compounds were observed during the haze  
599 episode, and some potential BrC chromophores were degraded. In addition, the chemical  
600 reactions of bio-VOCs such as isoprene also diluted the light-absorbing compounds in  
601 HULIS.

602 Thus, the present study provides novel insights into the light and molecular  
603 evolution of HULIS during haze event, which are important for predicting the  
604 environmental and climatic effects of HULIS. However, as this study examined only one  
605 haze bloom-decay process in winter in Guangzhou, the results obtained may be not  
606 adequate for understanding all the haze episodes in South China. Therefore, there is a  
607 need for a comprehensive investigation of haze episode in different seasons and regions  
608 in future.

609

610 **Data availability**

611 The research data are available in the Harvard Dataverse  
612 (<https://doi.org/10.7910/DVN/DYGYQT>, Song, 2022).

613

614 **Author contributions.** J. Song and P. Peng designed the research together. C, Zou, T.  
615 Cao, and M. Li carried out the PM<sub>2.5</sub> sampling experiments. C, Zou and T. Cao extracted  
616 and analyzed the WSOC and HULIS samples. B. Jiang analyzed the HULIS samples by  
617 FT-ICR MS. C. Zou and J. Song wrote the paper. J. Li, X. Ding, Z Yu, and G. Zhang  
618 commented and revised the paper.

619

620 **Competing interests.** The authors declare that they have no conflict of interest

621

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627 **References**

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926 **Table 1.** Formular number of potential BrC chromophores and the intensity ratios of each group of  
 927 potential BrC in total potential BrC and each group of total identified formulas, respectively.

Sample	Elemental composition	Number	ESI-		Elemental composition	Number	ESI+	
			Int <sub>C</sub> /Int <sub>Br</sub>	Int <sub>BrC,i</sub> /Int <sub>bul</sub>			Int <sub>C</sub> /Int <sub>Br</sub>	Int <sub>BrC,i</sub> /Int <sub>bul</sub>
	s	r	c	k	s	r	c	k
Clean-I	CHO-	424	<b>0.48</b>	<b>0.25</b>	CHO+	263	<b>0.37</b>	0.07
	CHON-	773	<b>0.46</b>	<b>0.53</b>	CHON+	480	<b>0.56</b>	<b>0.15</b>
	CHOS-	63	0.03	0.05	CHN+	79	0.07	<b>0.56</b>
	CHONS-	43	0.03	0.08	all in ESI+	822		<b>0.11</b>
	all in ESI-	1303		<b>0.26</b>				
Haze-I	CHO-	356	<b>0.44</b>	<b>0.21</b>	CHO+	244	<b>0.29</b>	0.09
	CHON-	791	<b>0.50</b>	<b>0.45</b>	CHON+	614	<b>0.62</b>	<b>0.22</b>
	CHOS-	43	0.03	0.03	CHN+	94	0.09	<b>0.39</b>
	CHONS-	39	0.03	0.07	all in ESI+	952		<b>0.16</b>
	all in ESI-	1229		<b>0.22</b>				
Haze-II	CHO-	444	<b>0.45</b>	<b>0.26</b>	CHO+	333	<b>0.34</b>	0.06
	CHON-	941	<b>0.49</b>	<b>0.49</b>	CHON+	595	<b>0.56</b>	<b>0.13</b>
	CHOS-	67	0.03	0.03	CHN+	89	0.1	<b>0.48</b>
	CHONS-	78	0.03	0.07	all in ESI+	1017		<b>0.10</b>
	all in ESI-	1530		<b>0.25</b>				
Clean-II	CHO-	391	<b>0.46</b>	<b>0.27</b>	CHO+	234	<b>0.38</b>	0.09
	CHON-	707	<b>0.48</b>	<b>0.59</b>	CHON+	462	<b>0.56</b>	<b>0.18</b>
	CHOS-	64	0.03	0.05	CHN+	75	0.06	<b>0.57</b>
	CHONS-	49	0.03	0.10	all in ESI+	771		<b>0.13</b>
	all in ESI-	1211		<b>0.29</b>				

928 Int<sub>C</sub>: the intensity of each group of identified potential BrC;

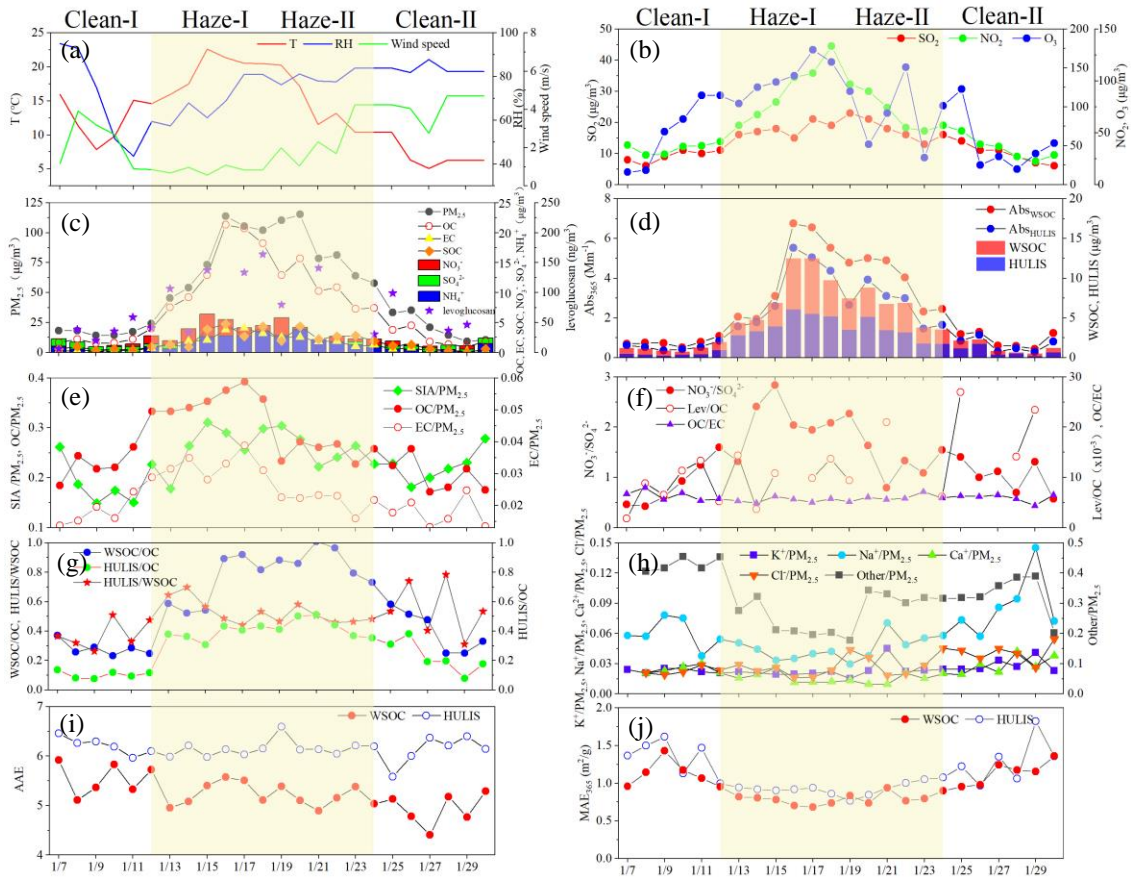
929 Int<sub>BrC</sub>: the sum intensity of identified potential BrC;

930 Int<sub>Bulk</sub>: the sum intensity of each group of total identified formulas.

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**Figure 1.** Temporal variation in meteorological parameters, concentrations of chemical composition, and optical properties ( $Abs_{365}$ ,  $MAE_{365}$ , and  $AAE$ ) of water-soluble BrC in the  $PM_{2.5}$  samples.

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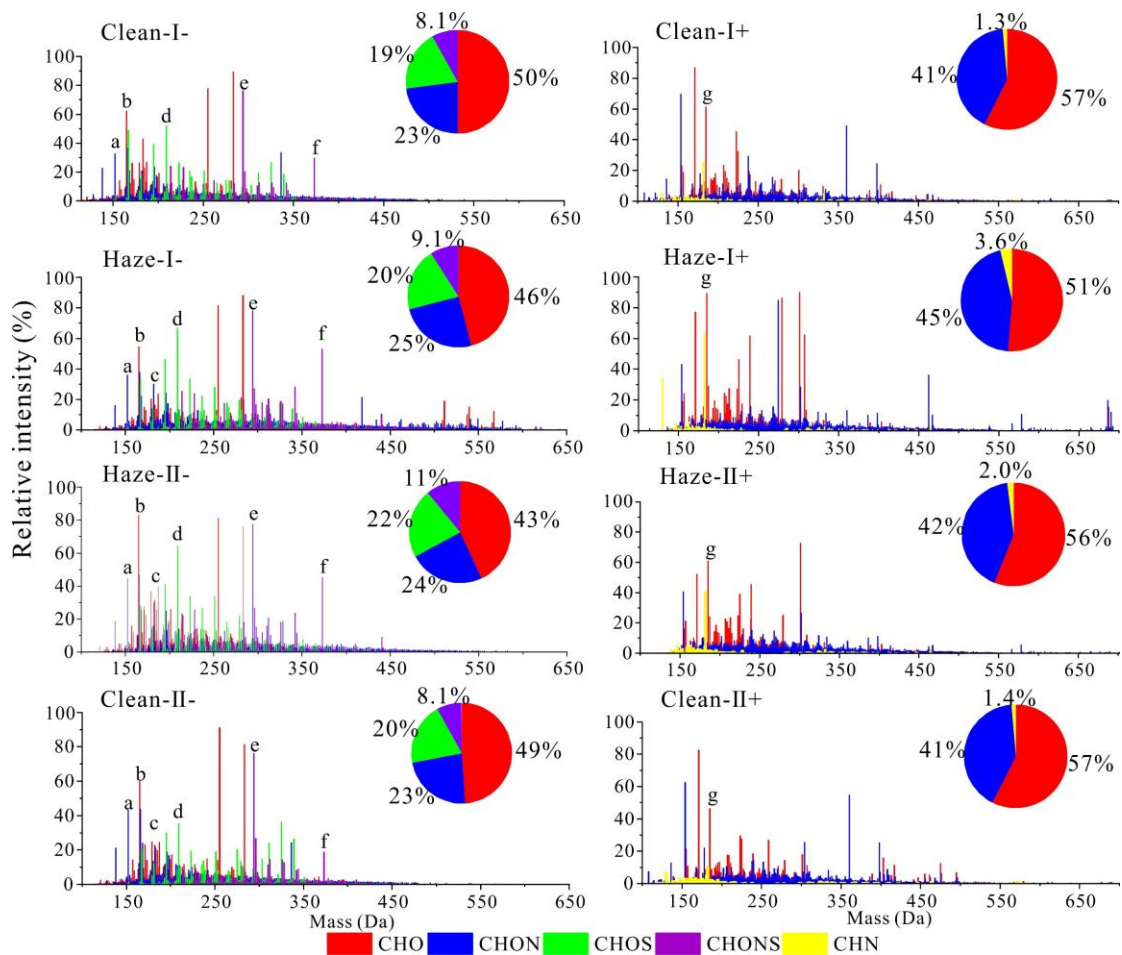
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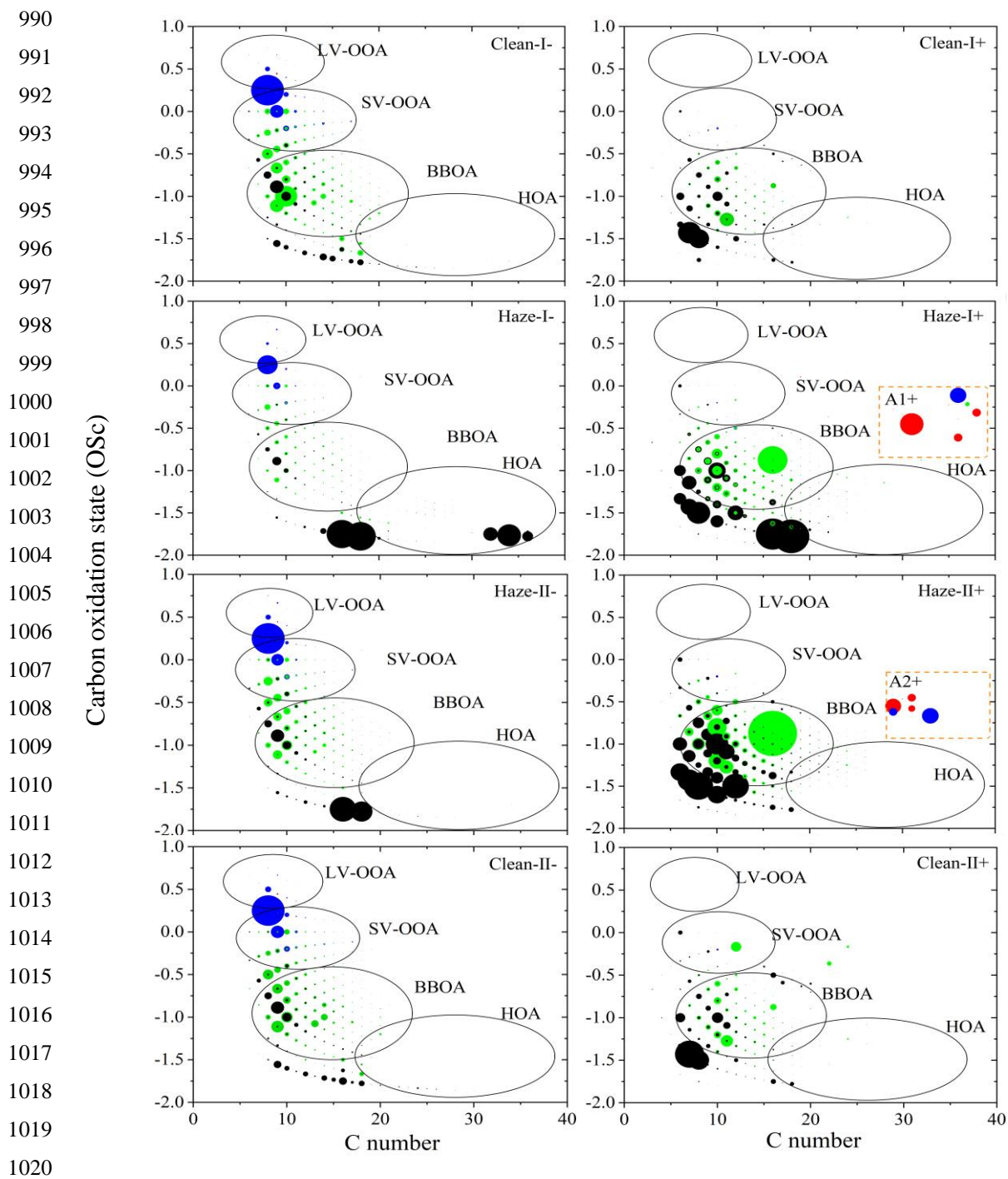
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**Figure 2.** Mass spectra of HULIS detected in ESI- and ESI+ modes during the haze process. The pie charts represent the intensity percent of different compound groups.

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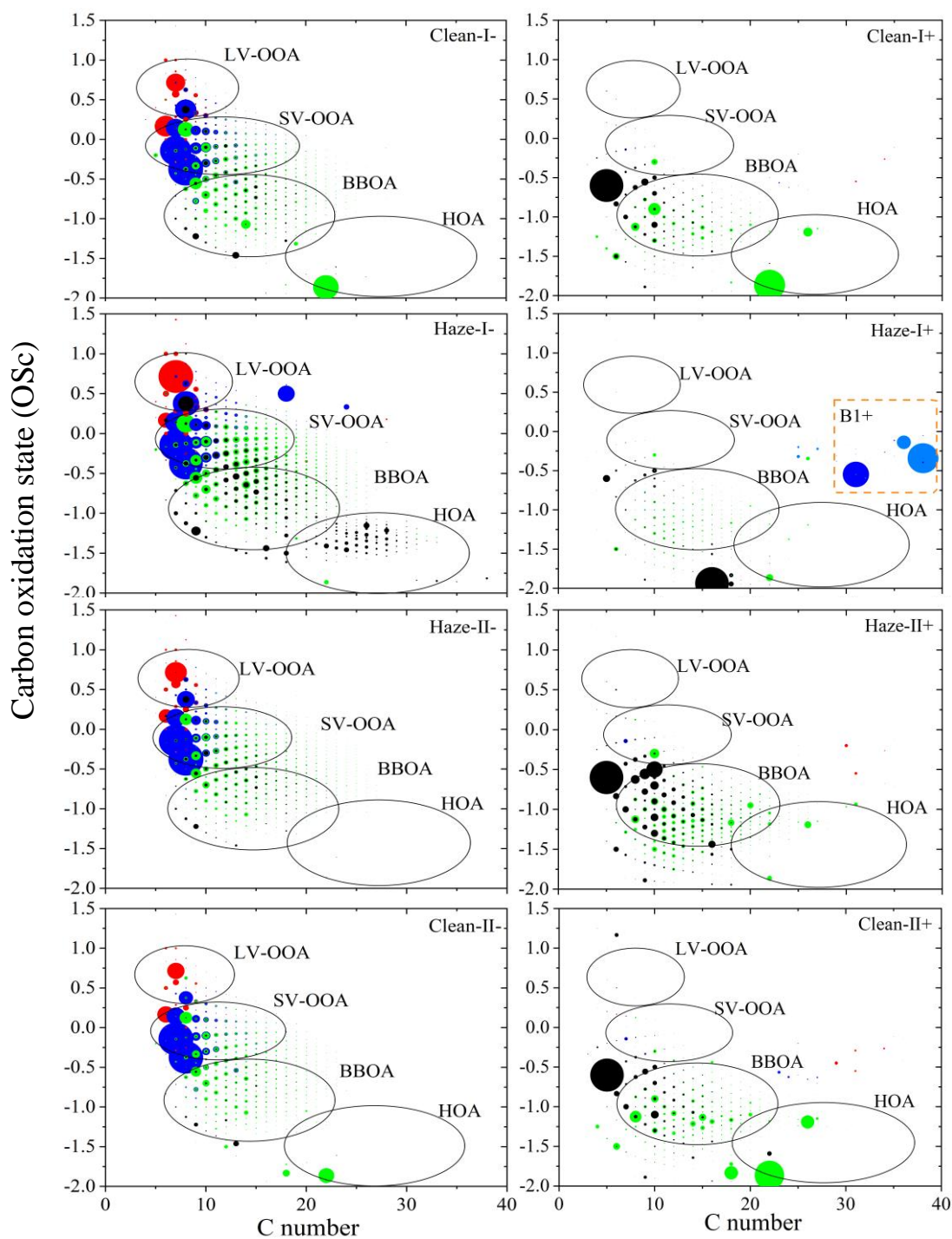
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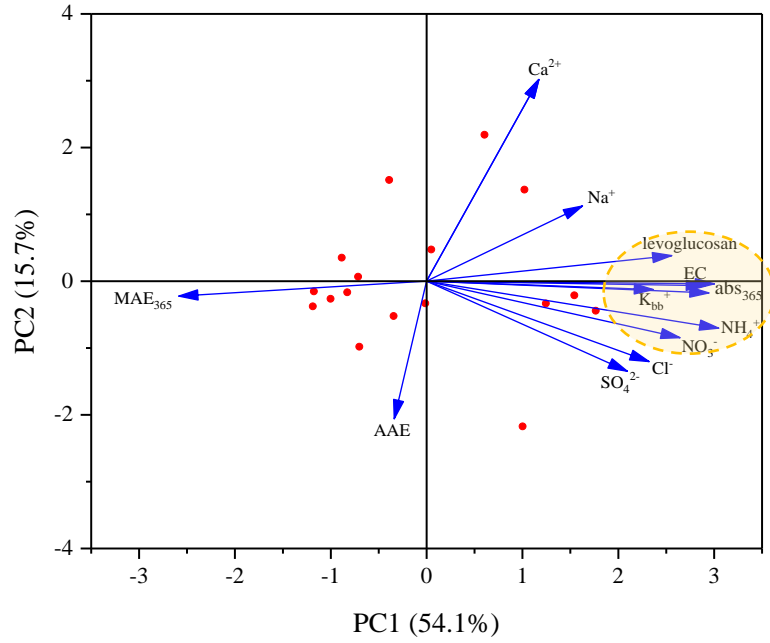
1021 **Figure 3.** Carbon oxidation state (OSc) plots for CHO<sup>-</sup> and CHO<sup>+</sup>. Formulas with black,  
 1022 green, blue, and red are assigned to aliphatic (AI = 0), olefinic (0 < AI < 0.5), aromatic  
 1023 (0.5 ≤ AI < 0.67), and condensed aromatic (AI ≥ 0.67) species (Koch and Dittmar, 2006),  
 1024 respectively.

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1057 **Figure 4.** Carbon oxidation state (OSc) plots for CHON- and CHON+. Formulas with  
 1058 black, green, blue, and red are assigned to aliphatic ( $AI = 0$ ), olefinic ( $0 < AI < 0.5$ ),  
 1059 aromatic ( $0.5 \leq AI < 0.67$ ), and condensed aromatic ( $AI \geq 0.67$ ) species (Koch and Dittmar,  
 1060 2006), respective.

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**Figure 5.** Principal component analysis results for the optical properties of HULIS and chemical compositions of PM<sub>2.5</sub>.