

# Methane emissions responsible for record-breaking atmospheric methane growth rates in 2020 and 2021

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**Abstract.** The global atmospheric methane growth rates reported by NOAA for 2020 and 2021 are the largest since systematic measurements began in 1983. To explore the underlying reasons for these anomalous growth rates we use newly available methane data from the Japanese Greenhouse gases Observing SATellite (GOSAT) to estimate methane surface emissions. Relative to baseline values in 2019 we see the largest annual increases in methane emissions during 2020 over Eastern Africa ( $13\pm 3$  Tg), tropical Asia ( $4\pm 3$  Tg), tropical South America ( $3\pm 3$  Tg), and temperate Eurasia ( $3\pm 3$  Tg), and the largest reductions over China ( $-6\pm 2$  Tg) and India ( $-2\pm 2$  Tg). We find comparable emission changes in 2021, relative to 2019, except for tropical and temperate South America where emissions increased by  $9\pm 3$  Tg and  $5\pm 3$  Tg, respectively, and tropical Asian emissions increased by  $8\pm 2$  Tg. The elevated contributions we saw in 2020 over the western half of Africa ( $-5\pm 3$  Tg) and Europe ( $-3\pm 3$  Tg) are substantially reduced in 2021, compared to our 2019 baseline. We find statistically significant positive correlations between anomalies of tropical methane emissions and groundwater, consistent with recent studies that have highlighted a growing role for microbial sources over the tropics. Emission reductions over India and China are expected in 2020 due to the Covid-19 shutdown but continued in 2021, which we do not currently understand. Based on a range of sensitivity studies for which we assume different spatial and temporal distributions of OH reduction in 2020, due to reduced pollutant emissions during the Covid-19 shutdown, and on our results from an inversion that includes OH scaling factors, we find that our *a posteriori* emissions in 2020 are 22-30% lower than our control calculation. We conclude therefore that most of the observed increase in atmospheric methane during 2020 and 2021 is due to increased emissions rather than reduced levels of OH.

## 1 Introduction

The atmospheric growth rate of methane in the 21<sup>st</sup> century has defied a definitive explanation: following a period of near-zero growth during 2000-2007 (Rigby et al. 2008) growth rates have accelerated, with values reported by NOAA for 2020 ( $15.29\pm 0.38$  ppb) and 2021 ( $16.94\pm 0.38$  ppb) exceeding all prior values since their records began in 1983. The underlying

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reasons for these anomalous growth rates in 2020 and 2021 are currently subject to intense debate with some studies attributing most of the growth in 2020 to a reduction in the hydroxyl radical (OH) sink of methane due to global-scale reductions in nitrogen oxides due to pandemic-related industry shutdowns (Laughner et al. 2021). On the face of it, this appears to be a reasonable explanation, but recent studies have used satellite observations of atmospheric methane to reveal regional hotspots over the tropics that are responding to changes in climate and have global significance (Pandey et al. 2021; Lunt et al. 2019; 2021; Pandey et al. 2017; Feng, Palmer, Zhu, et al. 2022; Palmer et al. 2021; Wilson et al. 2020). Here, we use satellite observation of methane from the Japanese Greenhouse gases Observing SATellite (GOSAT) to document global and regional changes in emissions, extending a recent study (Feng, Palmer, Zhu, et al. 2022). In the next section we describe the data and methods used to infer methane emissions. In section 3 we describe our results and conclude the study in section 4.

## 2 Data and Methods

We follow closely the methodology from a recent study (Feng, Palmer, Zhu, et al. 2022), in which we simultaneously infer methane and CO<sub>2</sub> fluxes in 2020 and 2021 by directly assimilating proxy GOSAT XCH<sub>4</sub>:XCO<sub>2</sub> retrievals (X denotes atmospheric column-averaged dry-air mole fraction). These data are anchored by surface methane and CO<sub>2</sub> measurements from an *in situ* observation network. The main advantage of this approach is that it does not rely on assumed model CO<sub>2</sub> concentrations to extract XCH<sub>4</sub> from the proxy ratio. For the sake of brevity, we include only details relevant to the calculations shown here.

### 2.1 GOSAT methane proxy data

We use version 9.0 of the proxy GOSAT XCH<sub>4</sub>:XCO<sub>2</sub> retrievals, from the University of Leicester (R. J. Parker et al. 2020; R. Parker and Boesch 2020), including both nadir observations over land and glint observations over the ocean. Analyses have shown that these retrievals have a bias of 0.2%, with a single-sounding precision of ~0.72%. We globally remove a slightly larger 0.3% bias from the GOSAT proxy data to improve the comparison with independent *in situ* observations (Feng et al. 2017; Feng, Palmer, Parker, et al. 2022). We assume that each single GOSAT proxy XCH<sub>4</sub>:XCO<sub>2</sub> ratio retrieval has an uncertainty of 1.2% to account for possible model errors, including the errors in atmospheric chemistry and transport, which helps to prevent model overfitting to observations.

### 2.2 *In situ* data

To anchor the GOSAT proxy ratio observations (Fraser et al. 2014b), we also ingest simultaneously the CO<sub>2</sub> and methane mole fraction data at surface-based sites, chosen from the NOAA compilation of the multi-laboratory *in-situ* measurements (Di Sarra et al. 2021; 2022; Cox et al. 2021; 2022). We include the same subset of the surface sites used by a recent study that documented year to year variations of methane emissions during 2010-2019 (Feng, Palmer, Zhu, et al. 2022). We assume uncertainties of 0.5 ppm and 8 ppb for these *in situ* observations of CO<sub>2</sub> and methane respectively (Feng, Palmer, Zhu, et al. 2022).

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### 75 2.3 GEOS-Chem atmospheric chemistry transport model

We use the GEOS-Chem model of atmospheric chemistry and transport at a horizontal resolution of 2° (latitude) × 2.5° (longitude), driven by the MERRA2 meteorological reanalyses from the Global Modeling and Assimilation Office Global Circulation Model based at NASA Goddard Space Flight Center.

80 Our CO<sub>2</sub> and methane model calculations are described in recent study (Feng, Palmer, Zhu, et al. 2022). ~~The *a priori* CO<sub>2</sub> flux inventory includes monthly biomass burning emission (van der Werf et al. 2017); monthly fossil fuel emissions for 2019 in the absence of more recent data (Oda and Maksyutov 2021); monthly climatological ocean fluxes (Takahashi et al. 2009); and 3-hourly terrestrial biosphere fluxes (Randerson et al. 1996).~~

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85 ~~The *a priori* methane fluxes from nature include monthly wetland emissions, including rice paddies (Bloom et al. 2017); monthly fire methane emissions (van der Werf et al. 2017); and termite emissions (Fung et al. 1991). We include emissions from geological macroseeps (Kvenvolden and Rogers 2005; Etiope 2015). For *a priori* anthropogenic emissions we use the EDGAR v4.32 global emission inventory for 2012 (Janssens-Maenhout et al. 2019) that includes various sources related to human activities (e.g., oil and gas industry, coal mining, livestock, and waste).~~

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We use monthly 3-D fields of OH, consistent with observed values for the lifetime of methyl chloroform, from the GEOS-Chem full chemistry simulation (Mao et al. 2013; Turner et al. 2015) to describe the main oxidation sink of methane. Using pre-computed fields of OH greatly simplifies our calculations, but in sensitivity calculations described below we examine the sensitivity of our results to different OH distributions. We also include net microbial consumption of methane in soil (Fung et al. 1991) and reaction with chlorine atoms (Thanwerdas et al. 2019).

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To explore the sensitivity of our methane emission estimates for 2020 due to inferred reductions in OH due to large-scale industrial shutdown due to Covid-19 (Cooper et al. 2022), we scale down our baseline monthly 3-D OH fields by 5% where combustion emissions of CO<sub>2</sub> (Oda and Maksyutov 2021) were larger than the mean emissions over Africa, resulting in reductions mainly between 15°N and 65°N. Our choice of 5% represents a ~~reduction based on a recent study (Laughner et al. 2021). A recent study that accounted for reductions in nitrogen oxide emissions estimated a global OH reduction of ~4% due to Covid-19 lockdowns in 2020 (Miyazaki et al. 2021), which showed strong spatial and temporal variations, with localized reductions peaking at 20-30%. In the absence of direct measurements of OH and without considering co-reductions in non-methane hydrocarbons, these (and similar) results have large uncertainties. To further examine the impact of reduced OH levels on the atmospheric growth of methane during 2020, we have reported in Appendix D the results of two additional experiments that assumed different distributions of the OH reduction, guided by observed changes in tropospheric ozone (Ziemke et al.~~

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2022). [In Appendix E we report results from simultaneously inferring large-scale OH changes and methane emissions from atmospheric methane observations. These additional results are summarized in section 3.](#)

#### 115 2.4 Ensemble Kalman filter inverse method

We use an ensemble Kalman Filter (EnKF) framework to estimate simultaneously CO<sub>2</sub> and methane fluxes from satellite measurements of the atmospheric CO<sub>2</sub> and methane (Feng, Palmer, Zhu, et al. 2022). Our state vector includes monthly scaling factors for 487 regional pulse-like basis functions (Figure A1) that describe CO<sub>2</sub> and methane fluxes, including 476 land regions and 11 oceanic regions. We define these land sub-regions by dividing the 11 TransCom-3 land regions into 42 nearly  
120 equal sub-regions, with the exception for temperate Eurasia that has been divided into 56 sub-regions due to its large landmass. We use the 11 oceanic regions defined by the TransCom-3 experiment. We use a 4-month moving lag window to reduce the computational costs for projecting the flux perturbation ensemble into observation space long after (>4 months) their emissions, beyond which time it is difficult to distinguish between the emitted signal from variations in the ambient background atmosphere (Fraser et al. 2014a; Feng et al. 2017). Our *a priori* fluxes are described above. For simplicity we assume a fixed  
125 uncertainty of 40% for coefficients corresponding to the *a priori* CO<sub>2</sub> fluxes over each sub-region, and a larger uncertainty (60%) for the corresponding methane emissions. We also assume that *a priori* errors for the same gas are correlated with a spatial correlation length of 300 km and with a temporal correlation of one month.

As a sensitivity test, we also report methane fluxes inferred using the same EnKF approach but using the proxy GOSAT XCH<sub>4</sub>  
130 data [and \*in situ\* methane data](#). These GOSAT XCH<sub>4</sub> retrievals are calculated from the XCH<sub>4</sub>:XCO<sub>2</sub> ratio by applying an ensemble mean of model XCO<sub>2</sub>, and then bias-corrected according to comparison with TCCON XCH<sub>4</sub> retrievals (R. J. Parker et al. 2020).

#### 2.5 Correlative data

To help interpret the changes in methane emission estimates we use additional datasets that are relevant to microbial or  
135 pyrogenic production of methane. We use monthly surface temperature fields at a spatial resolution of 2°x2.5° from the Modern-Era Retrospective Analysis for Research and Applications, version 2 (MERRA2) developed by the Global Modeling and Assimilation Office, NASA Goddard Space Flight Center (Bosilovich *et al.*, 2015). Precipitation data are taken from the NOAA CMAP (CPC Merged Analysis of Precipitation) long-term global rainfall dataset (Xie and Arkin 1997) that provides near-global monthly coverage at a spatial resolution of 2.5° × 2.5°, from 1979 to near-present. In addition, we use monthly total  
140 water storage (liquid water equivalent depth, LWE) on a 1° × 1° global grid from the NASA/DLR Gravity Recovery and

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Climate Experiment Follow-on (GRACE-FO) (Landerer et al. 2020). Finally, we explore monthly biomass burning emissions from the Global Fire Emissions Database (GFED v4) (van der Werf et al. 2017).

## 2 Results

145 Table 1 summarizes our global emission estimates inferred from GOSAT for 2020 and 2021; and for 2019, which we use as our baseline year throughout this study (Figure A2). The largest change in our global *a posteriori* emissions occurs during 2019—2020 (27 Tg) from 583.7 to 610.7 Tg/yr. Our *a posteriori* emission estimates for 2019 and 2020 are within 2% of values reported by an independent study (Qu et al. 2022), consistent with our reported uncertainties. These elevated emissions are sustained, but not further increased, during 2021 (609.5 Tg/yr).

150 The 27 Tg emission increase in 2020 and the lack of further emissions growth in 2021 may appear inconsistent with the NOAA global annual mean growth rates of 15.3 and 16.9 ppb in 2020 and 2021, respectively (Table 1). Based on these reported atmospheric growth rates, and after considering the effects of methane sinks, we find that a one-box model calculation predicts an increase in emissions of 12.6 Tg between 2019 and 2020 and a further 15.1 Tg increase in 2021 (see Appendix B and Figure B1). These calculations use annual mean values that effectively represent the emissions increase between the middle of each  
155 successive year rather than the beginning and end. After considering the increases in monthly mean NOAA data, we find that the simple box model predicts a similar increase in emissions between December 2019 (583 Tg yr<sup>-1</sup>) and December 2020 (610 Tg yr<sup>-1</sup>) of 27 Tg yr<sup>-1</sup>, with emissions thereafter stabilizing, with mean emissions of 610 Tg yr<sup>-1</sup> in 2021. As such, we conclude that the global mean emission results inferred from GOSAT are consistent with those inferred from NOAA surface data, assuming a fixed methane atmospheric lifetime.

160 Figure 1 shows the broad geographical breakdown for our reported global changes in methane emissions. Relative to 2019, we find widespread increased emissions during 2020 except for China and India. Relative to baseline values in 2019 we see the largest annual increases in methane emissions during 2020 over Eastern Africa (13±3 Tg), tropical Asia (4±3 Tg), tropical South America (3±3 Tg), and temperate Eurasia (3±3 Tg), and the largest reductions over China (-6±2 Tg) and India (-2±2 Tg). Emission changes relative to 2019 are comparable in 2021, except for tropical and temperate South America increases to 9±3 Tg and 5±3 Tg, respectively, and tropical Asia increases to 8±2 Tg. The elevated contributions we saw in 2020 over the western half of Africa (-5±3 Tg) and Europe (-3±3 Tg) are substantially reduced in 2021, compared to our 2019 baseline. This is broadly consistent with a recent study that showed that methane emissions inferred from TROPOMI were significantly higher in the first half of 2020 than during 2019 (McNorton et al. 2022). This study focused mainly on major countries, while we find the largest changes are over tropical latitudes where emissions in the second half of 2020 make significant contributions (Figure A3).

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185 [Figure 2](#), shows the distribution of methane emissions from 2020 and 2021 and the relative changes from 2019 (Figure A1).

During 2020, there are significant decreases (20-30%) over the manufacturing regions such as eastern China, India, central America, and eastern Europe. There are also significant increases across eastern Africa (30-40%), eastern North America (30%), and maritime Southeast Asia (30%). During 2021, we see similar changes in emissions, but they are typically exaggerated. There is more of a pronounced increase over East Africa (>50%), southern Brazil (50%), and eastern North America (up to 40%), and large decreases over eastern China (-50%) and western Russia (-50%). During 2021 there is also a large decrease over equatorial West Africa and eastern Europe ([Figure 1](#)). There are substantial seasonal changes in methane emissions (Figure A3) that are broadly consistent with seasonal changes in temperature and rainfall (not shown).

190 [Figure 3](#), shows different annual surface temperature warming patterns in 2020 and 2021. During 2020 the high northern

latitudes are dominated by summer warming over Siberia (2-3 K on an annual scale) that has been linked to greenhouse gas emissions (Ciavarella et al. 2021) and surface temperatures over Alaska were 2-3 K cooler than baseline values in 2019, where there were comparatively small changes in groundwater (< 5 cm). North America, western Europe and Scandinavia also experienced anomalously warm annual mean temperatures (typically within  $\pm 2$  K of 2019 values). There were smaller changes in temperatures at low latitudes (typically  $\pm 1$  K of 2019 values), but larger increases in groundwater ( $\pm 10$ -20 cm) associated with higher changes in rainfall (Figure A4), particularly over East Africa and eastern Brazil. During 2021, high northern latitudes were cooler than 2019 (<2-3 K) except for the contiguous US and Canada (higher than 2019 values by 2-3 K). Midlatitudes and low latitudes generally did not experience the warm temperatures of 2020. Elevated groundwater was sustained in 2021 over East and Southern Africa, eastern tropical South America (principally Brazil but stretching up to Venezuela), central America, India, maritime Southeast Asia, and north and southeast Australia. Groundwater decreased over the contiguous US, part of tropical South America and parts of Eurasia. We find generally stronger annual and seasonal relationships between methane emission anomalies and hydrological anomalies (rainfall and groundwater) for 2020 and 2021 ([Figure A5](#)) than for temperature anomalies. Particularly, we find statistically significant large-scale positive correlations (typically 0.5-0.6;  $p < 0.001$ ) for all seasons between methane and groundwater anomalies over Eastern Africa, tropical South America, and tropical Asia, but no significant correlation between methane and surface temperature anomalies. This is consistent with recent studies that have highlighted an increasing role for microbial sources in the tropical methane budget (Lunt et al. 2019; Feng, Palmer, Zhu, et al. 2022; Wilson et al. 2020). Over North America, we find a significant negative correlation (from -0.3 to -0.5;  $p < 0.001$ ) with rainfall during MAM and JJA and a significant positive correlation with temperature during JJA (0.3;  $p < 0.001$ ), which we do not currently understand. Fire emissions did not increase much where we report the largest increases in methane emissions in 2020 or 2021, except over central Canadian provinces. (Figure A4).

215 To explore the robustness of our results, we explore [a range](#) sensitivity runs. First, we ran our inversion for 2020 with OH concentrations decreased by 5% over regions with large anthropogenic emissions (as described in section 2.3), [an estimate](#) for the value we expect based on widespread reductions in nitrogen oxide emissions due to shutting manufacturing and other

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industries during Covid-19 lockdowns (Laughner et al. 2021). A reduction in OH, assuming constant emissions, would lead to an increase in atmospheric methane. Ignoring that possibility would result in a positive bias for reported methane emissions estimates. We find that our imposed 5% decrease in OH concentrations results in an emission reduction of 6.0 Tg/yr, representing 22% of the emission increase, relative to 2019, in the control run. [Figure 4](#) show that the largest impacts of reducing OH on *a posteriori* methane emissions are of the order of  $\pm 10\%$  over regions where there are the largest decreases in nitrogen oxide emissions (Cooper et al. 2022), including China, India, Europe, North America. The regions where values are higher than the control are a consequence of mass balance. Our 5% OH decrease cannot fully explain the increase in atmospheric methane, and consequently regions where OH were not decreased have had to compensate. The results for a smaller 1% decrease in OH have similar distributions but smaller changes (not shown). Second, we ran an experiment in which we used the methane columns determined by the proxy data, assuming model values for CO<sub>2</sub> (R. J. Parker et al. 2020). Using these data, we find our results for 2020 and 2021 are usually within a few percent of the values we report using the proxy data directly (Figure A6 and Figure A7). We found that assuming a different distribution for the OH reduction, based on observed changes in tropospheric ozone (Ziemke et al. 2022), did not significantly affect the results reported here (Appendix D). Including large, localized but short-term OH reductions also did not significantly impact our results. By including OH scaling factors into our state vector, we were able to infer simultaneously OH distributions and methane emissions (Appendix E). We find that the increased atmospheric growth rate of methane in 2020 that can be attributed to OH reductions is consistent with the values we reported from our experiments that used simpler *ad hoc* methods to describe reduced OH. Our *a posteriori* emission estimates from the control and the OH scaling inversions are consistent with independent observations from the TCCON network (Appendix C and E, respectively). Our sensitivity experiments provide confidence in our *a posteriori* emission estimates and in our statement that most of the atmospheric methane growth in 2020 was due to an increase in emissions, with the OH reduction associated with the Covid-19 lockdowns representing up to 30% of the global atmospheric methane growth.

### 3 Concluding Remarks

We reported regional emission estimates of methane during 2020 and 2021, two years with record-breaking atmospheric growth rates, inferred from satellite observations of methane from the Japanese Greenhouse gases Observing SATellite.

We find in both years that emissions from Eastern Africa, tropical Asia, and tropical South America dominate the global atmospheric growth rate, increasing by 11-13 Tg/yr, 4-8 Tg/yr, and 3-9 Tg relative to the 2019 baseline year, respectively. During 2020, we also find substantial increased emissions, relative to the 2019 baseline values, from Australia ( $3\pm 2$  Tg) and temperate Eurasia ( $3\pm 5$  Tg). Emission changes relative to 2019 are comparable in 2021, except for temperate South America that increases to  $5\pm 3$  Tg and temperate North America that increases to  $3\pm 3$  Tg. The elevated contributions we saw in 2020 over the western half of Africa ( $-5\pm 3$  Tg) and Europe ( $-3\pm 3$  Tg) are substantially reduced in 2021, compared to our 2019 baseline. We find statistically significant positive correlations between tropical methane emission and hydrological anomalies,

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consistent with recent studies that have highlighted a growing role for microbial sources over the tropics (Lunt et al. 2019; Feng, Palmer, Zhu, et al. 2022; Wilson et al. 2020). Our results are broadly consistent with a recent study of the 2020 period (Qu et al. 2022), including the magnitude of change associated with a change in OH, albeit concluded using an independent method.

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Substantial, widespread reductions in nitrogen oxides during 2020 associated with the shutdown of manufacturing and other industries will have perturbed atmospheric concentrations of the OH loss of methane. A reduction in OH could also help explain, in principle, the record-breaking atmospheric increase in methane. Here we explore that issue by scaling back our prescribed OH fields by 5% over fossil fuel combustion regions to assess the impact of our reported increases in methane emissions. We find a 5% decrease in OH results in an *a posteriori* emission reduction of 6.0 Tg compared to our control run that does not consider a decrease in OH, equivalent to 22% of the emission increase in our control run for 2020. These changes are mainly focused over fossil fuel emitting regions, as expected. Other sensitivity experiments that explored alternative distributions of OH reductions and an inversion that included OH scaling factors found similar results, with OH reductions explaining up to 30% of the global atmospheric methane growth rate in 2020. Our results suggest that an increase in emissions was predominately responsible for the observed growth in atmospheric methane during 2020.

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This study highlights the tremendous value of using satellite observations to understand rapid changes in atmospheric methane. They provide crucial information not only to identify regional column hotspots associated with emissions but also provide correlative information to help attribute those hotspots to specific anthropogenic or natural emissions.

*Data availability.* The University of Leicester GOSAT Proxy v9.0 XCH<sub>4</sub> data are available from the Centre for Environmental Data Analysis data repository at <https://doi.org/10.5285/18ef8247f52a4cb6a14013f8235cc1eb>. Precipitation, temperature, and the GRACE datasets are available at <http://grace.jpl.nasa.gov>. The community-led GEOS-Chem model of atmospheric chemistry and model is maintained centrally by Harvard University (<http://geos-chem.seas.harvard.edu>), and is available on request. The ensemble Kalman filter code is publicly available as PyOSSE (<https://www.nceo.ac.uk/data-tools/atmospheric-tools/>). The TCCON data were obtained from the TCCON Data Archive hosted by CaltechDATA at <https://tccodata.org>.

Field Code Changed

*Author contributions.* LF and PIP designed the research; LF and MFL prepared the calculations; RJP and HB provided the GOSAT data and expert advice on its usage; PIP wrote the paper, with comments from LF, RJP, MFL, and HB.

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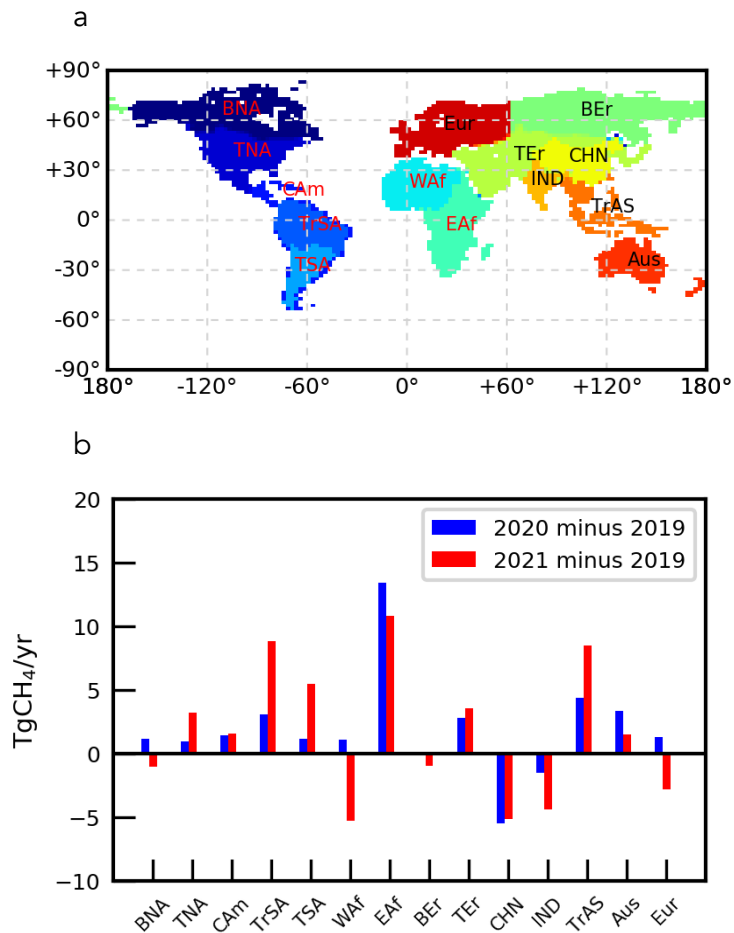
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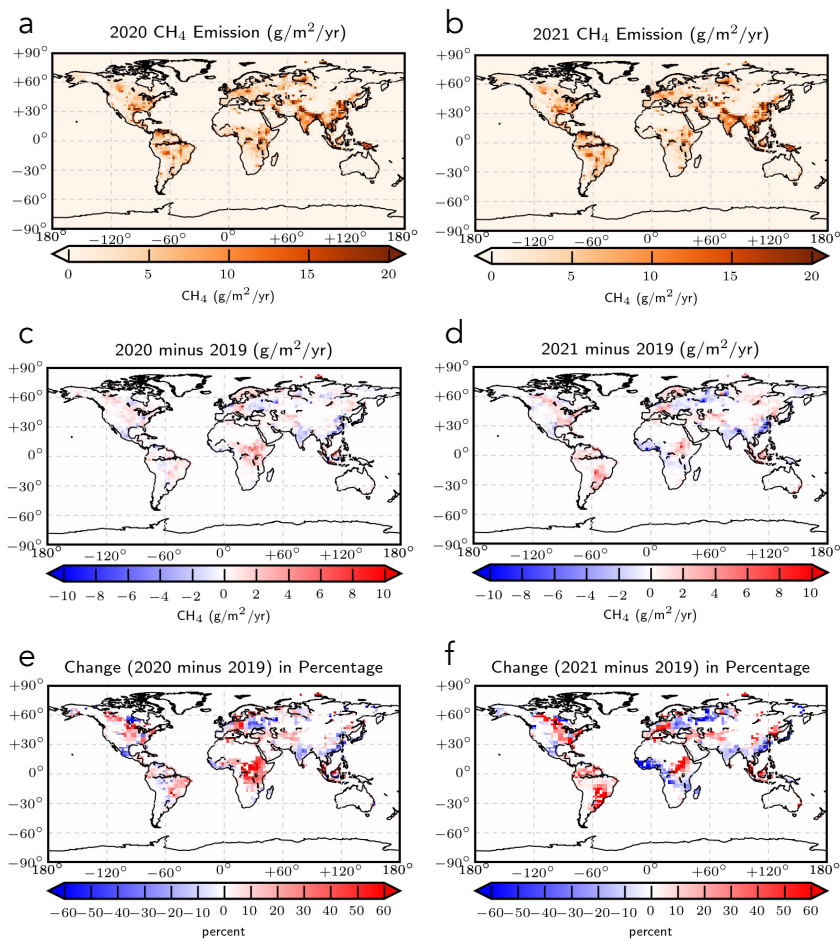
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Figures



480 **Figure 1** Large-scale geographical regions a) for which we report methane changes (TgCH<sub>4</sub>/yr) in 2020 and 2021 b) relative to our 2019 baseline year. Geographical regions, informed by TransCom-3 experiments (Gurney et al. 2004), include boreal North America (BNA), temperate North America (TNA), central America (Cam), tropical South America (TrSA), temperate South America (TSA), Europe (Eur), western Africa (WAF), eastern Africa (EAF), boreal Eurasia (BEr), temperate Eurasia (TEr), India (IND), China (CHN), tropical Asia (TrAs), and Australia (Aus).

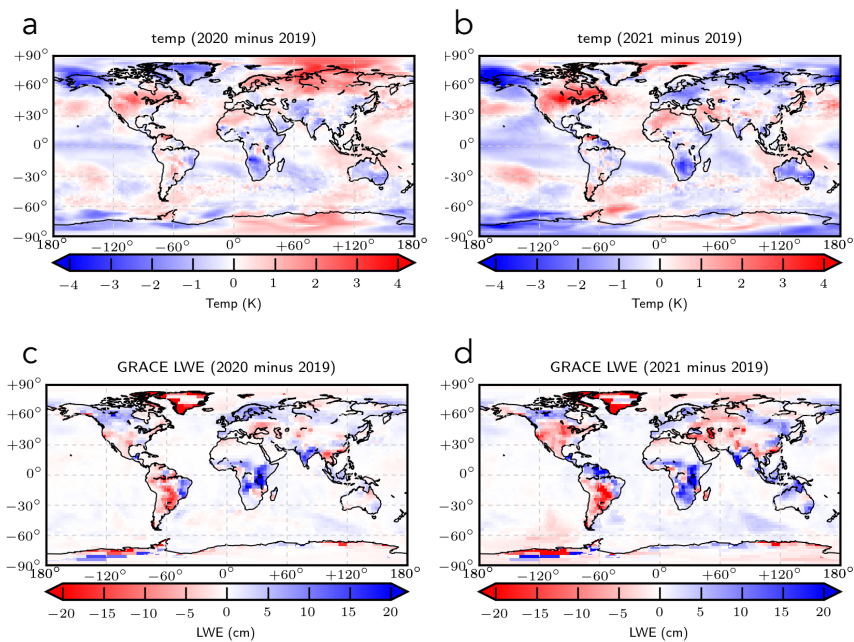


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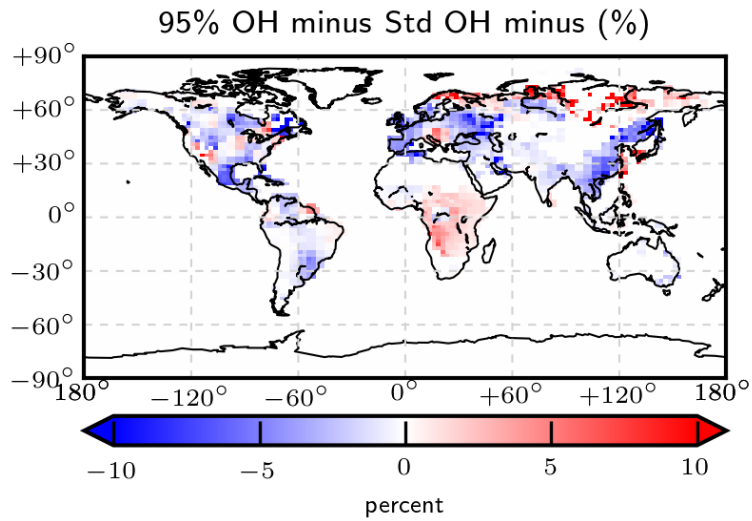
**Figure 2** Global *a posteriori* emissions of methane ( $\text{g}/\text{m}^2/\text{yr}$ ) inferred from GOSAT methane: $\text{CO}_2$  column ratio data for 2020 (panel a) and 2021 (panel b) and how they differ from the baseline year of 2019, described in terms of absolute (panels c and d, respectively) and percentage values (panels e and f).

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**Figure 3** Global annual mean surface temperature and GRACE liquid water equivalent (LWE) anomalies in 2020 (panel a and c) and 2021 (panels b and d) relative to values in 2019.



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**Figure 4** Percentage difference in global *a posteriori* methane emissions inferred from GOSAT for 2020 between a sensitivity run in which the OH field was decreased by 5% and the control run ([Figure 2](#)).

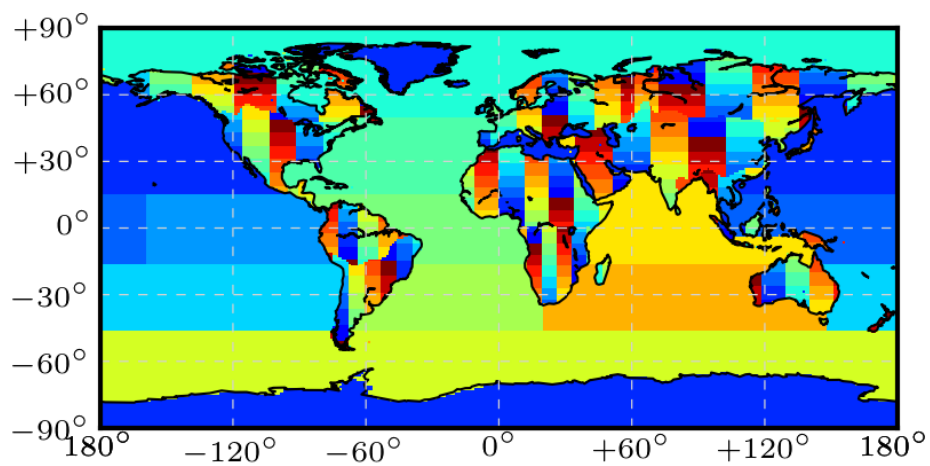
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	Global annual methane emissions (Tg/yr)		
	2019	2020	2021
GOSAT	583.7±11.2	610.7±11.3	609.5±12.0
GOSAT with OH decreased by 5%		604.7±11.3	
<i>In situ</i>	588.9±18.1	601.4±18.6	--
NOAA atmospheric growth rate (ppb/yr)	9.89±0.64	15.29±0.38	16.94±0.38

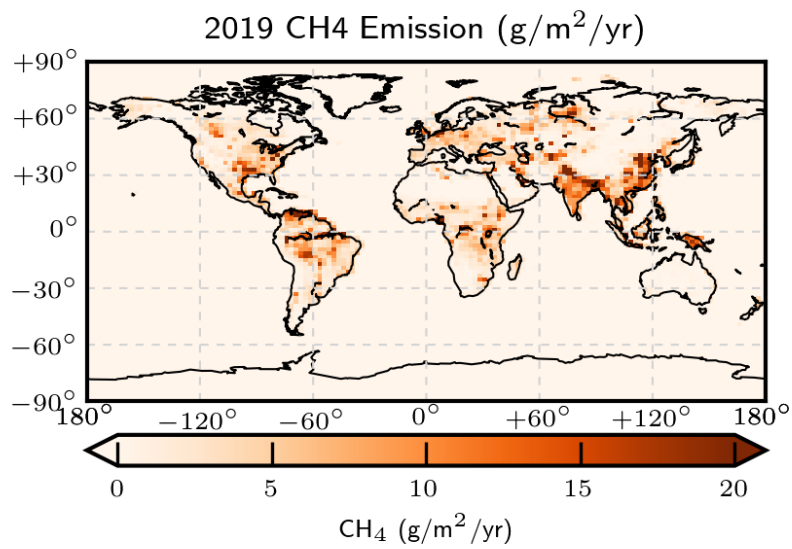
**Table 1** Global annual emission estimates of methane (Tg/yr) inferred from GOSAT (2019-2021) and *in situ* (2019-2020) atmospheric measurements of methane. The annual atmospheric methane growth rate (ppb/yr) for 2019 to 2022 reported by NOAA is also shown.

Appendix A: supplementary figures

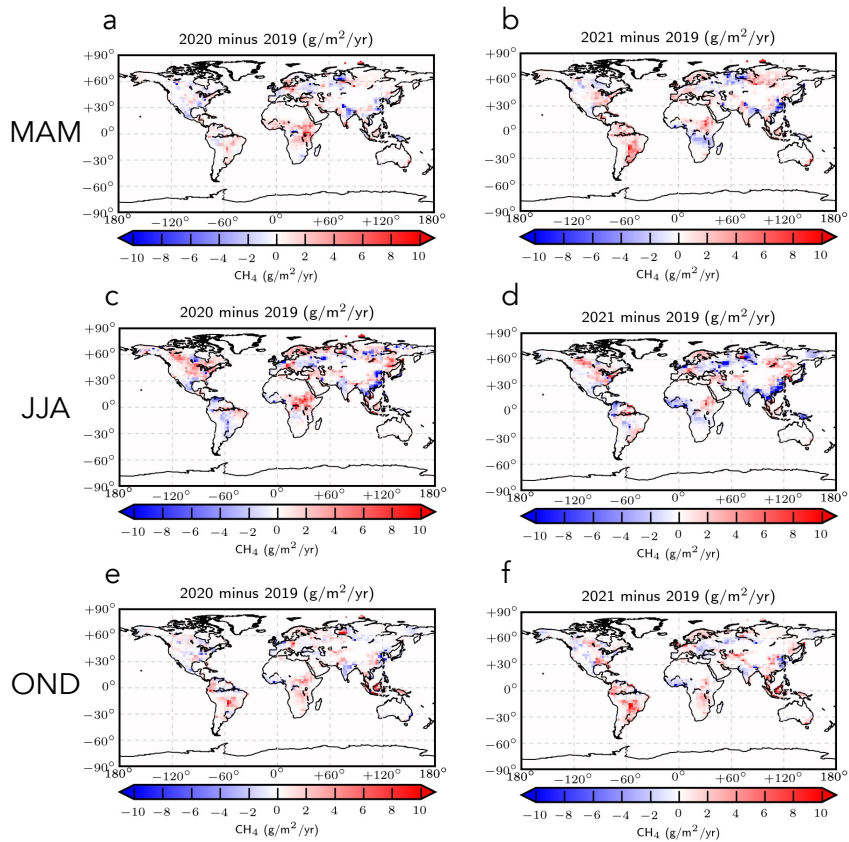


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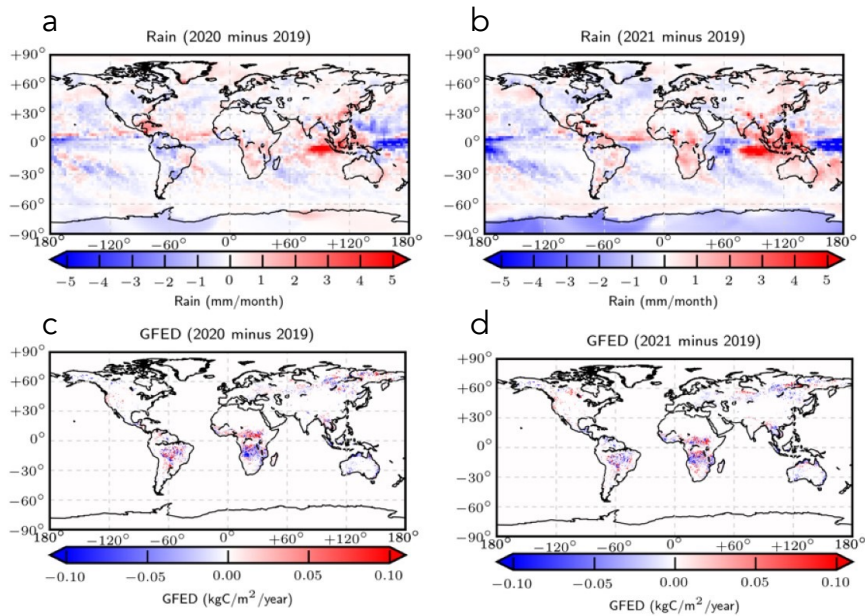
**Figure A1** Basis functions that describe the 487 regions where we estimate methane emissions, including 476 land regions and 11 oceanic regions.



510 **Figure A2** Global *a posteriori* emissions of methane (g/m<sup>2</sup>/yr) inferred from GOSAT methane:CO<sub>2</sub> column ratio data for the baseline year of 2019 (Feng, Palmer, Zhu, et al. 2022).

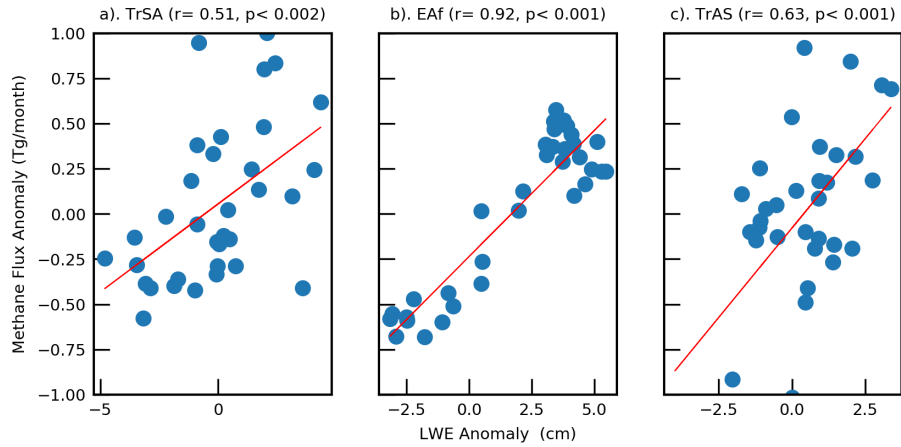


515 **Figure A3** Global seasonal *a posteriori* emissions of methane ( $\text{g}/\text{m}^2/\text{yr}$ ) inferred from GOSAT methane: $\text{CO}_2$  column ratio data for 2020 (l.h.s. panels) and 2021 (r.h.s. panels) relative to the baseline year of 2019, described in terms of absolute values. Seasons are based on rainfall changes over the tropics.



**Figure A4** Global annual mean NOAA CMAP precipitation (mm/month/yr) and GFED fire emission (kgC/m<sup>2</sup>/yr) anomalies in 2020 (panel a and c) and 2021 (panels b and d) relative to values in 2019.

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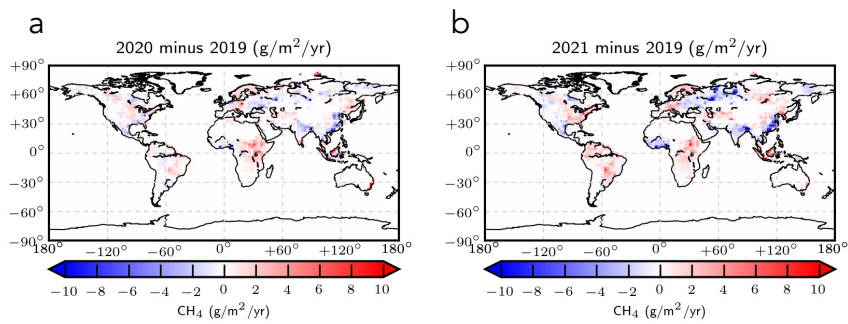


**Figure A5** Scatter plot of monthly GRACE-FP LWE anomalies (cm) and methane flux anomalies, 2018-2021, over a) Tropical South America, b) Eastern Africa, and c) Tropical Asia. Red lines denote the linear regression. Numbers atop of each panel denote the Pearson correlation coefficient  $r$  and the  $p$ -value.

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530 **Figure A6** Global *a posteriori* emissions of methane (g/m<sup>2</sup>/yr) inferred from GOSAT methane column data for 2020 (panel a) and 2021 (panel b).

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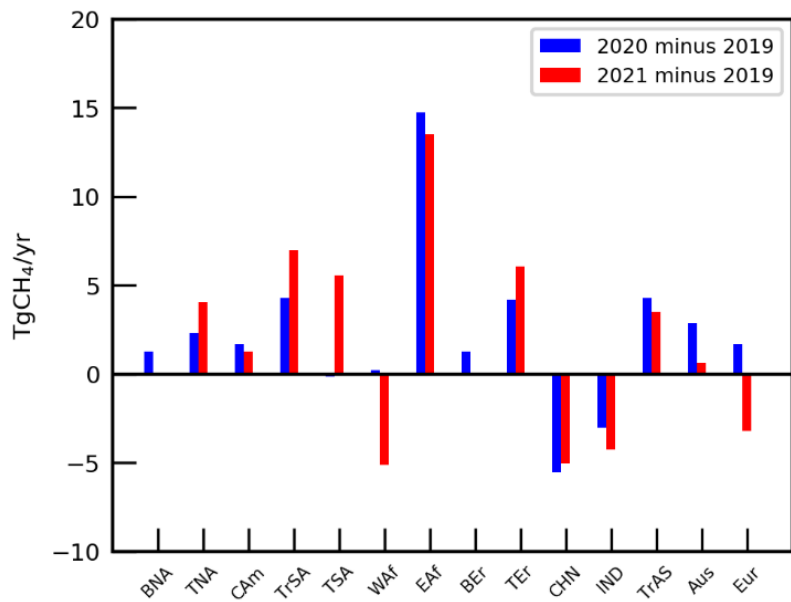


Figure A7 As Figure 1b but for an inversion that uses GOSAT proxy XCH<sub>4</sub> and *in situ* methane data.

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## Appendix B: Description of box model calculation

540 To calculate global emissions of methane from the NOAA global mean data we use a simple one-box model. In this model, the change in global mean methane concentration over time is given by:

$$\frac{dB}{dt} = Q - kB$$

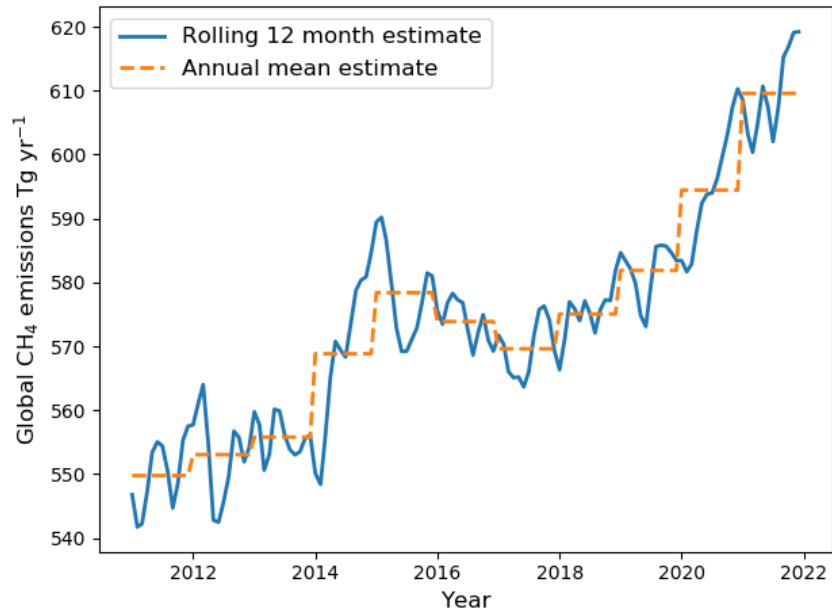
where B is the atmospheric mass of methane in Tg, k is the loss rate given as 1/lifetime and Q is the emissions rate. From this, after integration, the annual emissions rate can be calculated as:

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$$Q_t = \frac{k(B_t - B_{t-1} \cdot e^{-k})}{(1 - e^{-k})}$$

The loss rate was tuned to match a steady state concentration of 1775 ppb during 2000-2006 based on constant emissions of 530 Tg yr<sup>-1</sup> during this period. We calculated the rolling 12-month annual emissions to track the progression of global emissions between 2019 and 2021. We used the difference between the atmospheric concentration in January 2019 and January 2020, February 2019 to February 2020, etc. to calculate the change in emissions in the intervening 12 months. Figure B1 shows the increase in emissions throughout 2020 followed by more variable month-to-month changes in 2021. The large increase in emissions primarily occurs in 2020, with emissions at the 12-month period ending in Dec 2020 being 27 Tg yr<sup>-1</sup> larger than the emissions one year earlier. In contrast if emissions are calculated using annual mean concentrations, it appears as if there is a larger emission increase in 2021. The box model results show that the highly simplified calculation based on global average data is consistent with the more complex inverse modelling approach applied to the GOSAT data.

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**Figure B1** Global box model methane emission estimates between 2011 and 2021, respectively. Emission estimates are based on NOAA global mean surface data. The blue line denotes the rolling 12-month annual emissions, and the orange line denotes the emissions based on annual mean concentrations.

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### Appendix C: Evaluation of *a posteriori* flux estimates

We indirectly evaluate our *a posteriori* methane fluxes by comparing the GEOS-Chem methane distribution, driven by the *a posteriori* fluxes, with independent  $\text{XCH}_4$  retrievals from Total Carbon Column Observing Network (TCCON) of Fourier transform spectrometers (Wunch et al., 2011). We use bias-corrected TCCON  $\text{XCH}_4$  data from the latest GGG2020 public release of the TCCON data set from 2019 to 2021, including updates until October 2022. For a comprehensive description of the network and the available data from each TCCON site, we refer the reader to the TCCON project page. Here we use a subset of available TCCON data, dependent on their availability between 2018 and 2021 (Buschmann et al. 2022; De Mazière et al. 2022; García et al. 2022; Hase et al. 2022; Kivi, Heikkinen, and Kyrö 2022; Liu et al. 2022; Morino, Ohyama, et al. 2022b; Morino, Velazco, et al. 2022; Morino, Ohyama, et al. 2022a; Notholt et al. 2022; Petri et al. 2022; Pollard, Robinson, and Shiona 2022; Warneke et al. 2022; Shiomi et al. 2022; Té, Jeseck, and Janssen 2022; Wennberg, Wunch, et al. 2022; Wennberg, Roehl, Wunch, Blavier, et al. 2022; Wennberg, Roehl, Wunch, Toon, et al. 2022; Wunch et al. 2022; Zhou et al. 2022). For further details about the data we direct the reader to the TCCON project page: <http://tccodata.org/>.

Figure C1 shows the mean and standard deviation of the differences between our *a posteriori* model simulation and TCCON GGG2020 data in 2019 and 2020. The *a posteriori* model simulation is driven by our *a posteriori* methane emission estimates. We sample the associated model 3-D atmospheric methane distributions at the time and location of each TCCON site used. We then convolve the sampled vertical profile with site- and time-dependent TCCON instrument averaging kernels, which describes the altitude-dependent instrument sensitivity to changes in atmospheric methane concentration.

We do not report results for 2021 due to data availability; for data that are available, we find the mean statistics (not shown) are similar to those we report here for 2019 and 2020. For most sites, we find the mean bias is typically smaller than  $\pm 10$  ppb and the standard deviation has a range 5-15 ppb with typically values smaller than 10 ppb. We find the largest differences at northern high latitudes where the model has a large overestimate ( $\sim 10$  ppb), consistent with previous studies (Feng et al., 2017, 2022), due to poor coverage of GOSAT data during the boreal winter and to model error.

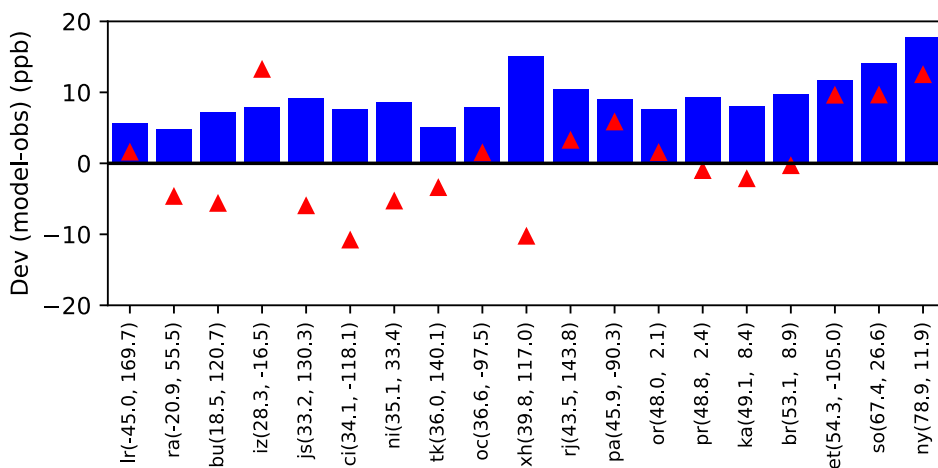


Figure C1 Statistical comparison of the GEOS-Chem *a posteriori* methane distribution and TCCON  $\text{XCH}_4$  data (v GGG2020) in 2019 and 2020. Red upward triangles denote the mean bias and the blue bars denote the corresponding 1s values.

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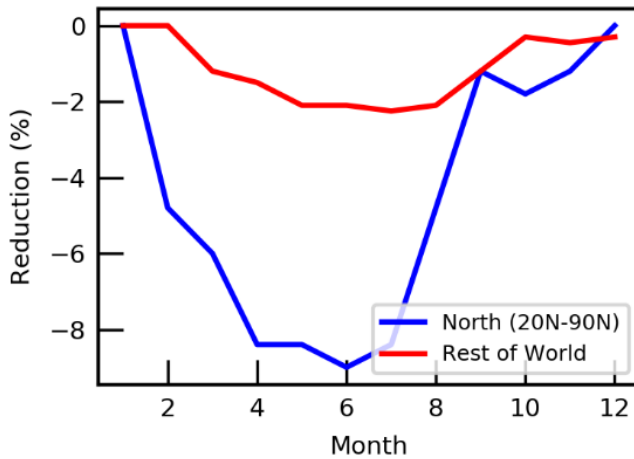
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#### Appendix D: Time-dependent OH reduction sensitivity test

We assume a temporal-spatial distribution to describe the OH reduction in 2020, following a recent study on tropospheric ozone changes in 2020 and 2021 (Ziemke et al., 2022). First, we divide the world into regions: northern hemisphere (20°N-90°N) and the rest of the world. We assume the reduction in OH in the northern hemisphere starts from the boreal Spring of 2020 and peaks during the summer with a magnitude of 9% (blue line, Figure D1), higher than the ozone reduction found by Ziemke et al. (2022). For latitudes south of 20°N the time evolution of the reduction in ozone is less clear (Ziemke et al., 2020). We assume for simplicity that the OH reduction at these latitudes (red line, Figure D1) has a smaller peak value (-2.3%) and with a time lag of one month compared to the northern hemisphere region.



**Figure D1** Assumed temporal distribution for OH reduction (%) in the northern hemisphere (20°N-90°N, blue line) and the rest of the world (red line) for our sensitivity calculation.

Our inversion that uses this assumed time-dependent OH reduction results in an annual global methane emission of  $602.2 \pm 11.2$  Tg for 2020, which is about 8.5 Tg (~30%) smaller than our control run that used a fixed OH climatology (Table 1). This is consistent with increased methane emissions being primarily responsible for the anomalously large atmospheric growth rate of methane in 2020. We find that introducing an additional localized 25% OH reduction over China in March 2020 reduces the *a posteriori* global methane emission increase (Table 1) by only 0.26 Tg.

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## Appendix E: Atmospheric OH inversion

605 Currently there is no direct observation of the global distribution of atmospheric OH. Indirect constraints on atmospheric OH from the changing lifetimes of trace gases such as CO and methane are insufficient to infer determine 3-D distributions of OH.

610 Here, as part of the state vector, we infer 18 scaling factors for monthly *a priori* OH fields from atmospheric methane observations (section 2). These scaling factors correspond to 18 geographic regions (Figure E1). We define six 25° latitude bands between 75°S and 75°N. We do not consider scaling polar OH values. We then split each of latitude band into three longitude ranges, chosen so that each region contains a different tropical land mass: -180° to -30°; -30° to 60°; and 60° to 180°.

615 The Jacobian matrix, which describes the sensitivity of model methane concentrations to regional OH fields, is calculated with GEOS-Chem forced by *a priori* methane fluxes from the control run (Table 1) but with the OH climatology reduced by 5% for each of the 18 regions. To reduce the computational cost of this calculation, we use the same four-month lag window for the OH scaling factor estimates as for the methane emission estimates. So that each monthly OH scaling factor will be constrained only by observations in the subsequent four months, but its impact will remain for the entire experimental period. We used the OH climatology as our *a priori*, and assume a uniform 5% uncertainty for each of the 18 regions so that the 2-sigma range covers possible OH changes that range  $\pm 10\%$ . We use such a simple linearization scheme to adjust surface methane emissions and the monthly tropospheric OH by optimally fitting model calculations to atmospheric methane observations.

620 We conduct the joint inversion for 2018 to 2021, including the 18 monthly OH scaling factors and methane emissions, and use the same atmospheric methane data used by the control calculation. The corresponding methane emissions estimate in 2019 is now  $606.8 \pm 13$  Tg/yr, about 23.1 Tg/yr higher than the control run. This higher emission is a result of (and largely cancelled out by) an elevated tropospheric OH loss of methane of  $24.5 \pm 8$  Tg/yr, which is simultaneously inferred from the observation datasets. As expected, the resulting *a posteriori* atmospheric methane field is consistent with the assimilated GOSAT data and with independent TCCON observations (Figure D2). Further work is required to assess whether the higher methane loss is due to errors in the OH climatology or errors in atmospheric methane corresponding to the *a priori* inventories.

625 The global *a posteriori* methane emission estimate for 2020 is also higher than the control run ( $627.3 \pm 13$  Tg/yr vs  $610.7 \pm 11.3$  Tg/yr), but the 17 Tg/yr increase is smaller than for 2019. As a result, the increase in emissions from 2019 to 2020 is 20.5 Tg/yr, which is about 75% of the control run. The corresponding additional OH loss of methane in 2020 is about 8 Tg/yr smaller than in 2019, which we believe is due primarily to the OH reduction in 2020. These results are consistent with our sensitivity calculations that used a perturbed OH distribution (Appendix D).

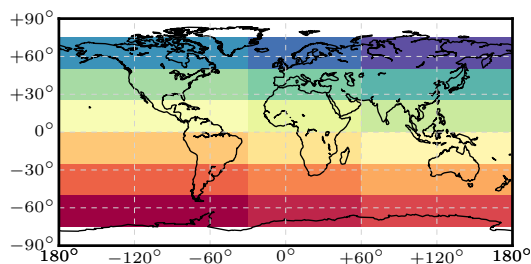


Figure E1 Definition of the 18 regions for OH scaling factors that we include in our joint OH factor-methane emission state vector.

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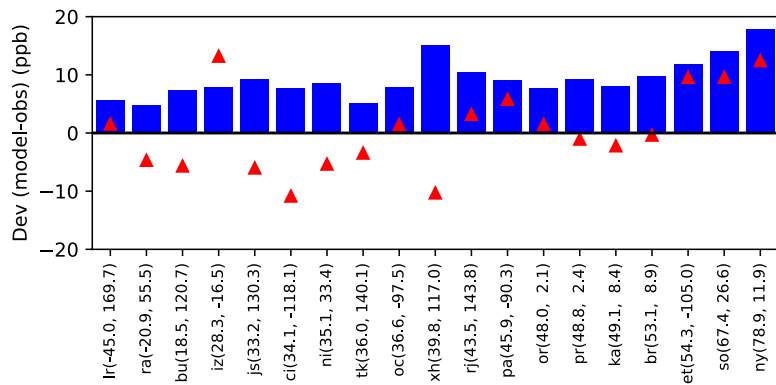


Figure E2 The same as Figure C1 but for the inversion that also includes OH scaling factors.

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