

1 **Projected increases in wildfires may challenge regulatory curtailment of PM<sub>2.5</sub> over the**  
2 **eastern US by 2050**

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4 Chandan Sarangi<sup>1,2\*</sup>, Yun Qian<sup>2\*</sup>, L. Ruby Leung<sup>2</sup>, Yang Zhang<sup>3</sup>, Yufei Zou<sup>4,2</sup>, Yuhang  
5 Wang<sup>4</sup>

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7 <sup>1</sup> Indian Institute of Technology Madras, Chennai, India

8 <sup>2</sup> Pacific Northwest National Laboratory, Richland, WA, USA

9 <sup>3</sup> Department of Civil and Environmental Engineering, Northeastern University, Boston, MA

10 <sup>4</sup> School of Earth and Atmospheric Sciences, Georgia Institute of Technology, Atlanta, GA,  
11 USA.

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14 \*Corresponding Authors:

15 [chandansarangi@iitm.ac.in](mailto:chandansarangi@iitm.ac.in)

16 [yun.qian@pnnl.gov](mailto:yun.qian@pnnl.gov)

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21 **Abstract**

22 Anthropogenic contribution to the overall fine particulate matter (PM<sub>2.5</sub>) concentrations has  
23 been declining sharply in North America. In contrast, a steep rise in wildfire-induced air  
24 pollution events with recent warming is evident in the region. Here, based on coupled fire-  
25 climate-ecosystem model simulations, summertime wildfire-induced PM<sub>2.5</sub> concentrations are  
26 projected to nearly double in North America by the mid-21st century compared to the  
27 present. More strikingly, the projected enhancement in fire-induced PM<sub>2.5</sub> (~ 1-2 µg/m<sup>3</sup>) and  
28 its contribution (~15-20%) to the total PM<sub>2.5</sub> are distinctively significant in the eastern US.  
29 This can be attributed to downwind transport of smoke from future enhancement of wildfires  
30 in North America to the eastern US and associated positive climatic feedback on PM<sub>2.5</sub> i.e.  
31 perturbations in circulation, atmospheric stability and precipitation. Therefore, the anticipated  
32 reductions in PM<sub>2.5</sub> from regulatory controls on anthropogenic emissions could be  
33 significantly compromised in the future in the densely populated eastern US.

34 **Key points:**

- 35 1) Wildfire-PM<sub>2.5</sub> associations studied based on unprecedented two-way coupled fire-  
36 climate-ecosystem model simulations
- 37 2) A steep rise in wildfire-induced air pollution events with recent warming is evident in  
38 the region
- 39 3) The transported smoke from enhanced wildfires in North America can severely affect  
40 air quality over Eastern US

41

42 **Keywords:** wildfire emissions, climate change, air quality, smoke transport, wildfire-climate-  
43 ecosystem interactions

## 44 **1. Introduction**

45 Wildfires are widespread burning events in forests, shrub lands, and grazing lands. In  
46 North America (mainly Canada and the US), particulate matter emissions from wildfires are a  
47 significant source of regional air pollution (Shi et al., 2019; McClure and Jaffe, 2018; Van  
48 Der Werf et al., 2010; Jaffe et al., 2008). Since the 1980s, the number of large wildfires and  
49 the length of wildfire season have been increasing, and the trends are projected to continue in  
50 the future over the western US, Alaska and Canada (Kitzberger et al., 2017; Kirchmeier-  
51 Young et al., 2017; Abatzoglou and Williams, 2016; Partain et al., 2016; Jolly et al., 2015;  
52 Westerling et al., 2006; Gillett et al., 2004). Accordingly, particulate emissions from wildfires  
53 are also anticipated to increase in North America in the 21st century (Knorr et al., 2017; Liu  
54 et al., 2016; Val Martin et al., 2015). Human exposure to high concentrations of wildfire-  
55 emitted airborne particulate matter of diameter  $\leq 2.5 \mu\text{m}$  (PM<sub>2.5</sub>) is known to have substantial  
56 adverse effects on pulmonary and cardiovascular functioning (Anjali et al., 2019; Black et al.,  
57 2017), which contribute significantly to global and regional all-cause mortality (Zhang et al.,  
58 2020; Hong et al., 2019; Yang et al., 2019; Ford et al., 2018; Johnston. et al., 2012).  
59 Therefore, a better understanding of the future changes in wildfire-induced PM<sub>2.5</sub> and its  
60 contribution to the total surface PM<sub>2.5</sub> is essential.

61 In the last two decades, ambient air quality in the US has substantially improved due  
62 to a decline in PM<sub>2.5</sub> by ~ 40 % (US EPA, 2018). The decrease in PM<sub>2.5</sub> is primarily due to  
63 curtailment of anthropogenic emissions resulting from US-based efforts to meet regulations  
64 such as the Clean Air Act (US EPA, 2009), Cross-State Air Pollution Rule, Regional Haze  
65 Rule, and the motor vehicles emissions standards. Consequently, air quality over the  
66 contiguous US (CONUS) and Canada has improved steadily such that it is predicted to  
67 achieve the targeted National Ambient Air Quality Standards in the future (Nolte et al.,

68 2018). Under this promising scenario, the influence of wildfire-emissions on the total PM<sub>2.5</sub>  
69 becomes even more crucial. Depending on the competition between climate-induced increase  
70 in wildfires and the regulatory control on anthropogenic emissions, future enhancement in  
71 wildfire-induced PM<sub>2.5</sub> may compromise the reduction in anthropogenic PM<sub>2.5</sub> concentrations  
72 in certain regions. In agreement, recent studies have highlighted the potential for future  
73 enhancement in wildfire-induced pollution to diminish the reducing trend in PM<sub>2.5</sub>, primarily  
74 over the western US (O'Dell et al., 2019; Ford et al., 2018; Val Martin et al., 2015; Yue et al.,  
75 2013).

76         While the fractional wildfire burnt area and fire intensities are the greatest over the  
77 western US and Canadian regions within North America, anthropogenic emissions dominate  
78 the ambient PM<sub>2.5</sub> concentration over the eastern US. The inherent geographical separation  
79 between the regions with large wildfire emissions and anthropogenic emissions leads to a  
80 pertinent question: will future enhancement in wildfires over the western US and Canada  
81 have significant effects on PM<sub>2.5</sub> over the eastern US? Addressing this question is crucial  
82 because the declining trend in PM<sub>2.5</sub> over the eastern US is the major contributor to the  
83 observed 40% decrease in PM<sub>2.5</sub> over the US in the last two decades (US EPA, 2018).  
84 Eastward advection of wildfire smoke from Canada and the western US has been found to  
85 severely hamper the surface air quality of the central and eastern US under the influence of  
86 the prevailing westerlies during the summer months (Brey et al., 2018; Wu et al., 2018;  
87 Gunsch et al., 2018; Kaulfus et al., 2017; Dempsey, 2013). The transported wildfire smoke  
88 can influence the meteorology and climate via the radiative impact of carbonaceous  
89 emissions, changes in land albedo and cloud system perturbations (Ward et al., 2012; Liu et  
90 al., 2014). These fire-weather interactions can have positive feedback on the locally-emitted  
91 PM<sub>2.5</sub> in the eastern US by surface cooling and boundary layer suppression(Guan et al.,  
92 2020). At the same time, fire-triggered ecosystem changes can induce negative feedback on

93 PM<sub>2.5</sub> by reducing the future wildfires over North America (Zou et al., 2020). Thus, two-  
94 way interactions between fires and climate that are important for predicting future changes in  
95 wildfire locations, intensities, and durations (Harris et al., 2016) as well as associated  
96 particulate emissions is essential. However, past studies have mostly employed simple  
97 statistical models based on statistical regressions of present-day fire burnt area on the  
98 meteorological fields (Liu et al., 2016; Spracklen et al., 2009; Yue et al., 2013; Val Martin et  
99 al., 2015), and more recently, one-way coupled modelling (Ford et al., 2018; O’Dell et al.,  
100 2019).

101 Here, based on new two-way coupled fire-climate-ecosystem simulations, we  
102 demonstrate the significance of wildfire-induced contributions to ambient PM<sub>2.5</sub> over the  
103 eastern US due to enhanced wildfire smoke transportation and smoke-induced changes in  
104 weather in eastern US. This enhancement in wildfire-induced PM<sub>2.5</sub> may potentially challenge  
105 the targeted policy-driven reduction of PM<sub>2.5</sub> in the eastern US. Next, our model setup,  
106 experiments and methodology are explained in Section 2, followed by results and discussion  
107 in Section 3. The study is summarized in Section 4.

## 108 **2. Materials and Methods**

### 109 **2.1. RESFire-CESM Model description**

110 We employ the open-source REgion-Specific ecosystem feedback fire (RESFire)  
111 model coupled with the Community Land Model version 4.5 and the Community  
112 Atmosphere Model version 5 (CAM5) of the Community Earth System Model (CESM)  
113 version 1 (Zou et al., 2019; Neale et al., 2013) to perform two-way coupled simulations.  
114 RESFire provides state-of-the-art capabilities to simulate the complex fire-climate-ecosystem  
115 interactions globally for fires occurring over wildland, cropland, and peatland. Although  
116 wildfires dominate in the North American region, RESFire simulates both wildfires and

117 prescribed fires. Moreover, this integrated setup includes climatic feedback from fire-induced  
118 aerosol direct and indirect radiative effects and associated weather changes. It also includes  
119 feedback from fire-induced vegetation distribution changes and associated biophysical  
120 processes such as evapotranspiration and surface albedo. Sofiev et al. (2012) described the  
121 fire plume rise parameterization. Other features in CLM4.5 and CAM5, such as the  
122 photosynthesis scheme (Sun et al., 2012), the MAM3 aerosol module (Liu et al., 2012), and  
123 the cloud macrophysics scheme (Park et al., 2014), allow for more comprehensive  
124 assessments of the climate effects of fires through their interactions with vegetation and  
125 clouds. Fire-ecosystem interactions are modelled by simulating fire-induced vegetation  
126 mortality and regrowth (and associated land cover change) in RESFire. This approach has  
127 been introduced in Zou et al. (2019) and the simulated ecological and climatic effects of  
128 wildfires have been evaluated in two sets of sensitivity experiments in Zou et al. (2020).  
129 Although fire-climate-ecosystem interactions are considered in this study, our focus is on the  
130 fire-induced changes in PM<sub>2.5</sub> over Canada and the US, so the two vegetation-focused  
131 sensitivity experiments reported in Zou et al. (2020) are not included in this paper. Please  
132 refer to Zou et al. (2019) and Zou et al. (2020) for more details about the simulation of fire-  
133 ecosystem interactions.

## 134 **2.2 Numerical Experiment and Methodology**

135 We designed two sets of simulations for the present day and future scenarios to  
136 quantify the impacts of fire-climate-ecosystem interactions (Table 1). The spatial resolution is  
137 0.9° (lat) × 1.25° (lon) with a time step of 30 min. In each set of simulations, we conducted a  
138 default all emission included control run (X<sub>ALL</sub>, where x=2000 or 2050 indicates the present  
139 day or future, respectively) and a sensitivity run with no wildfire emissions to the atmosphere  
140 (X<sub>WEF</sub>, where X is the same as for the control runs). The ALL runs are designed to simulate  
141 fully interactive fire disturbances such as fire emissions with plume rise and fire induced land

142 cover changes of the present day (representative of the 2000s, 2000<sub>ALL</sub>) and a moderate future  
 143 emission scenario (representative of the 2050s, 2050<sub>ALL</sub>) via the RCP4.5. The only difference  
 144 between the ALL and WEF scenario is that wildfire emissions are absent in the WEF  
 145 scenario. Specifically, in the WEF runs, the online simulated fire emissions are not passed to  
 146 the CAM5 atmosphere model so that the difference between the ALL and WEF runs can be  
 147 used to isolate the atmospheric impacts of fire-climate interactions.

148 Table 1: Summary of the sensitivity simulations performed

Scenario	Present-day		Future	
Experiment Name	2000 <sub>ALL</sub>	2000 <sub>WEF</sub>	2050 <sub>ALL</sub>	2050 <sub>WEF</sub>
Simulation years	2001-2010	2001-2010	2051-2060	2051-2060
Atmosphere	CAM5	CAM5	CAM5	CAM5
Land	CLM4.5	CLM4.5	CLM4.5	CLM4.5
Ocean	Climatology	Climatology	RCP4.5	RCP4.5
Sea ice	Climatology	Climatology	RCP4.5	RCP4.5
Non-fire emissions	ACCMIP	ACCMIP	RCP4.5	RCP4.5
Fire emissions	Online fire aerosols with plume rise	—	Online fire aerosols with plume rise	—
Land cover	Fire disturbances on present-day condition	Fire disturbances on present-day condition	Fire disturbances on RCP4.5 condition	Fire disturbances on RCP4.5 condition

149  
 150 For the present-day experiments, we used the spun-up states from Zou et al. (2019) as  
 151 initial conditions for both meteorological and chemical variables. Sea surface temperature  
 152 (SST) for the present day was obtained from the Met Office Hadley Centre (HadISST).  
 153 Present-day non-fire emissions from anthropogenic and other sources were based on  
 154 ACCMIP (Lamarque et al., 2010) for the year 2000. We replaced the prescribed GFED2 fire  
 155 emissions (van der Werf et al., 2006) in the default setting of CESM with the online-coupled  
 156 fire emissions generated by the RESFire model. Zou et al. (2019) provided more details of  
 157 the physics parameterizations and modeling experiment settings used in these simulations.

158 Land use and land cover data for 2000 and 2050 from the Land-Use History A product (Hurtt  
159 et al., 2006) are used to initialize the 2000<sub>ALL</sub>/2000<sub>WEF</sub> and 2050<sub>ALL</sub>/2050<sub>WEF</sub> simulations,  
160 respectively. Following the above setup, the future scenario 2050<sub>ALL</sub> experiment accounts for  
161 both fuel load changes associated with the projected land use and land cover change  
162 (LULCC) in the 2050s and fire weather changes driven by the SST and sea ice forcing from a  
163 coupled CESM simulation following the greenhouse gas (GHG) forcing of the RCP4.5  
164 scenario. The global mean GHG mixing ratios in the CAM5 atmosphere model were fixed at  
165 the year 2000 levels in all the present-day experiments and they were replaced by those of the  
166 RCP4.5 scenario with the well-mixed assumption and monthly variations. However, the  
167 future population and socioeconomic conditions were identical to those of the present day so  
168 there was no explicit impact of human-induced mitigation/enhancement effects on wildfires  
169 in the future projection in all the future experiments. Future human impacts were considered  
170 implicitly in LULCC-induced fuel load changes in the RCP4.5 scenario.

171 The net projected changes by 2050s in emissions, meteorology and air quality during  
172 summer (JJA: June, July, August) months are estimated by comparing decadal-mean values  
173 simulated by 2000<sub>ALL</sub> with 2050<sub>ALL</sub>. Wildfire-induced enhancement in PM<sub>2.5</sub> concentration in  
174 the present day and mid-21<sup>st</sup> century is estimated by comparing 2000<sub>ALL</sub> with 2000<sub>WEF</sub> and  
175 2050<sub>ALL</sub> with 2050<sub>WEF</sub>, respectively. Further, the projected increase in wildfire-induced  
176 PM<sub>2.5</sub> in the future is calculated by comparing the simulated wildfire effect of the 2050s  
177 (2050<sub>ALL</sub>-2050<sub>WEF</sub>) with that of the 2000s (2000<sub>ALL</sub>-2000<sub>WEF</sub>). With large spatiotemporal  
178 variability, the projected changes in transported fire-emissions from the western US and  
179 Canada to the eastern US by the 2050s and the corresponding impacts are summarized using  
180 probability distribution functions. The latter provide information not only for the mean but  
181 also variability and extreme values to quantify the simulated changes for the three subregions.

182



### 183 **3. Results and Discussion**

#### 184 **3.1 Model Evaluation**

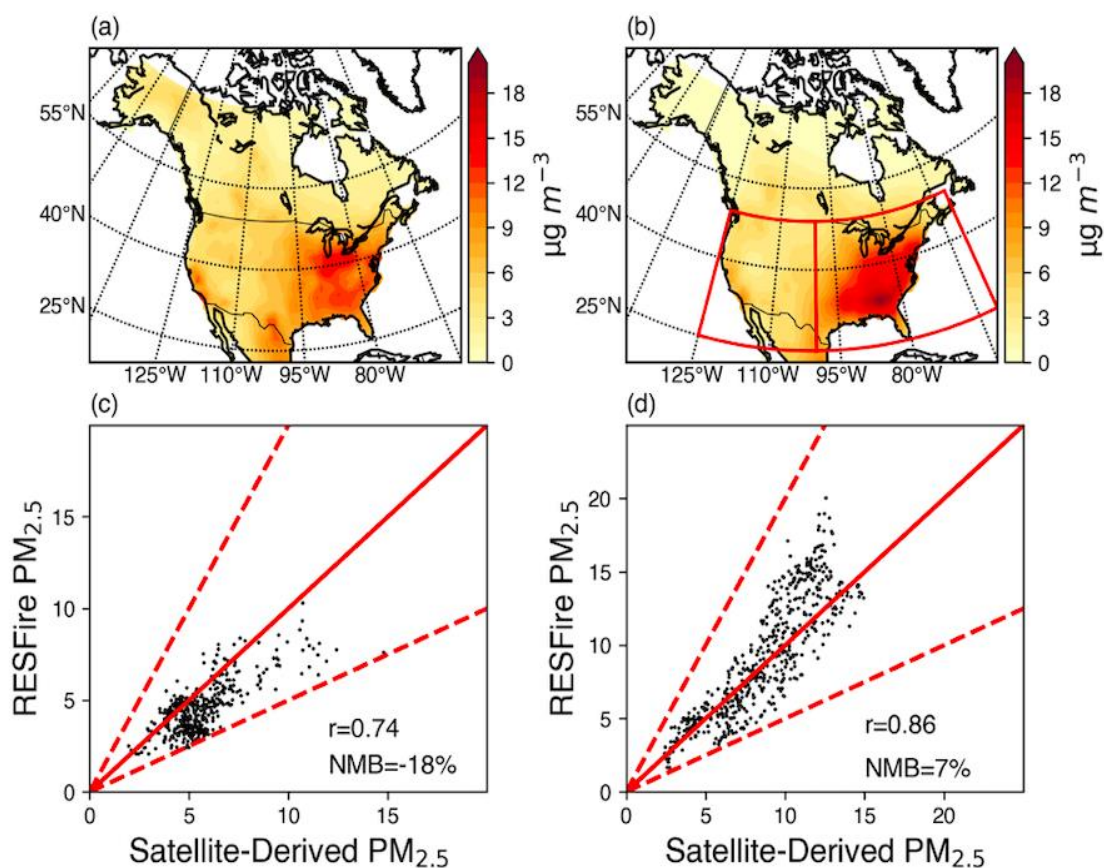
185 Zou et al. (2019) performed comprehensive evaluation of the RESFire simulated  
186 wildfire burnt area distribution, associated carbon emissions and terrestrial carbon balance to  
187 demonstrate reasonable model skill. Zou et al. (2020) compares global fire simulations by  
188 CESM-RESFire with modeling results reported in the literature to show better agreement  
189 with the GFED4.1s benchmark data and predicts more prominent changes in the future than  
190 those predicted by Kloster et al. (2010, 2012). These differences might come from differences  
191 in the climate sensitivities of the fire models and scenarios and other input data used to make  
192 future projections.

193 Here, we evaluate the simulated surface PM<sub>2.5</sub> against satellite-estimates (Figure 1)  
194 over North America. The PM<sub>2.5</sub> concentration is calculated as the sum of sulfate, nitrate, fine  
195 sea salt (first 2 size bins), fine dust (first size bin), black carbon (BC), and organic aerosol  
196 (OC) at the surface-level of model. OC is the sum of primary organic matter (POM) and  
197 secondary organic aerosol (SOA), and SOA is the sum of secondary species formed from  
198 toluene, monoterpenes, isoprene, benzene, and xylene. Figure 1 compares the observed and  
199 simulated mean annual PM<sub>2.5</sub> averaged over 2001-2010. The 10-year average satellite AOD-  
200 derived annual mean surface PM<sub>2.5</sub> concentrations (Van Donkelaar et al., 2018) are regridded  
201 to the model grid (Figure 1A) and then compared with the RESFire simulations in the  
202 2000<sub>ALL</sub> present-day run (Figure 1B). The spatial distribution of annual surface PM<sub>2.5</sub> is  
203 reasonably well simulated but also have some biases. To quantify the biases, we also  
204 estimated the correlation coefficient as well as normalized mean biases (NMB) of the  
205 simulated values compared against the satellite retrieved values over two subregions.  
206 Quantitatively, the NMB values over the western US (WUS) and eastern US (EUS) are 18%  
207 and 7%, respectively (Figure 1C-D). In addition, the spatial variability of the 2001-2010

208 averaged annual AOD distribution (Supplementary Figure 1) is also well represented in our  
209 simulation, although the model underestimates high AOD values. Similar spatial variability  
210 and biases in AOD and PM<sub>2.5</sub> were also found when a comparison was performed for only  
211 summer months (June through August; JJA). The simulated PM<sub>2.5</sub> has also been evaluated  
212 against the ground-based Interagency Monitoring of Protected Visual Environments  
213 (IMPROVE) data, showing similar spatial pattern and biases (10-25%) (Supplementary  
214 Figure 2). The biases are smaller over Eastern US and Southwestern US region. The  
215 simulated PM<sub>2.5</sub> values over California matches quite well with the observed annual mean  
216 values. However, the biases over Northwestern US region are ~30-40%, a portion of which  
217 could be attributed to possible biases in model's meteorology in northwestern US region.  
218 Nonetheless, both satellite and in situ evaluation indicate that our simulation biases are  
219 largely within the uncertainty range among the various satellite and ground-based datasets,  
220 which have normalized mean biases ranging from -3.3% to 33.3% when benchmarked against  
221 the ground-based IMPROVE data over the contiguous US (Diao et al., 2019; Val Martin et al.  
222 (2015)).

223         Discrepancies between the simulated and observed PM<sub>2.5</sub> values may be attributed to  
224 several potential reasons. First, the satellite-derived data has a non-zero lower bound of PM<sub>2.5</sub>  
225 concentrations, so the ambient background concentrations for relatively cleaner regions such  
226 as the western US may be overestimated (Figure 1C), also the sampling frequency between  
227 these datasets are different. Second, year 2000-based constant non-fire emissions were used  
228 in the RESFire simulation, which may result in overestimation of the PM<sub>2.5</sub> concentrations  
229 from non-fire sources during 2001-2010 when anthropogenic emissions and PM<sub>2.5</sub>  
230 concentrations continue to decrease (US EPA, 2018). This overestimation is prominent in  
231 regions dominated by non-fire sources such as the southeastern US. Third, large uncertainties  
232 in fuel consumption and emission factors preclude an accurate estimation of the primary fire

233 emissions in the model, especially for the eastern US where large fractions of low-intensity  
 234 prescribed fires consume only under-canopy fuels such as litter and duff layers. The fire  
 235 model may fail to capture the subtle distinctions between low-intensity prescribed fires and  
 236 forest fires, so more fuels are consumed and result in higher emissions. Lastly, comparison of  
 237 a coarsely resolved simulation against in-situ observations also contributes to uncertainty.  
 238 Differences in the degree to which fire-climate interactions and other physical processes and  
 239 feedbacks are represented by the models can explain the slight differences in estimating the  
 240 present day mean wildfire-induced change in PM<sub>2.5</sub> over local and downwind regions  
 241 between our simulations and previous studies. Nonetheless, reasonable simulation of the  
 242 spatial distribution of wildfire burnt area, AOD, and near surface particulate concentration  
 243 (mean bias of ~10-20 %) instills confidence about the fidelity of our model setup in  
 244 particulate pollution simulation, which is the focus of this study.



245

246

247 Figure 1: Comparison of the 10-year (2001-2010) averaged annual mean surface PM<sub>2.5</sub>  
248 concentration between observations and RESFire simulations. (a) Satellite-derived surface  
249 PM<sub>2.5</sub> concentrations (with dust and sea-salt removed) estimated by Donkelaar et al., 2018  
250 (available at [https://sedac.ciesin.columbia.edu/data/set/sdei-global-annual-gwr-pm2-5-modis-](https://sedac.ciesin.columbia.edu/data/set/sdei-global-annual-gwr-pm2-5-modis-misr-seawifs-aod)  
251 [misr-seawifs-aod](https://sedac.ciesin.columbia.edu/data/set/sdei-global-annual-gwr-pm2-5-modis-misr-seawifs-aod); last access: 5 November, 2021); (b) 2000<sub>ALL</sub> Simulated surface PM<sub>2.5</sub>  
252 concentrations (with dust and sea-salt removed) averaged over 2001-2010; The red boxes  
253 denote the two subregions (EUS and WUS) shown in Fig. 2 in the main text. (c) comparison  
254 of simulated and satellite based gridded surface PM<sub>2.5</sub> concentrations in the WUS subregion;  
255 Number of samples is equal to the number of land grids ~450 (d) same as (c) but in the EUS  
256 subregion. Number of samples is equal to the number of land grids ~375 The red solid and  
257 dashed lines denote the 1:1 ratio line and  $\pm 100\%$  biases, respectively. The correlation  
258 coefficients and NMB values are shown at the lower-right corner of each subplot.

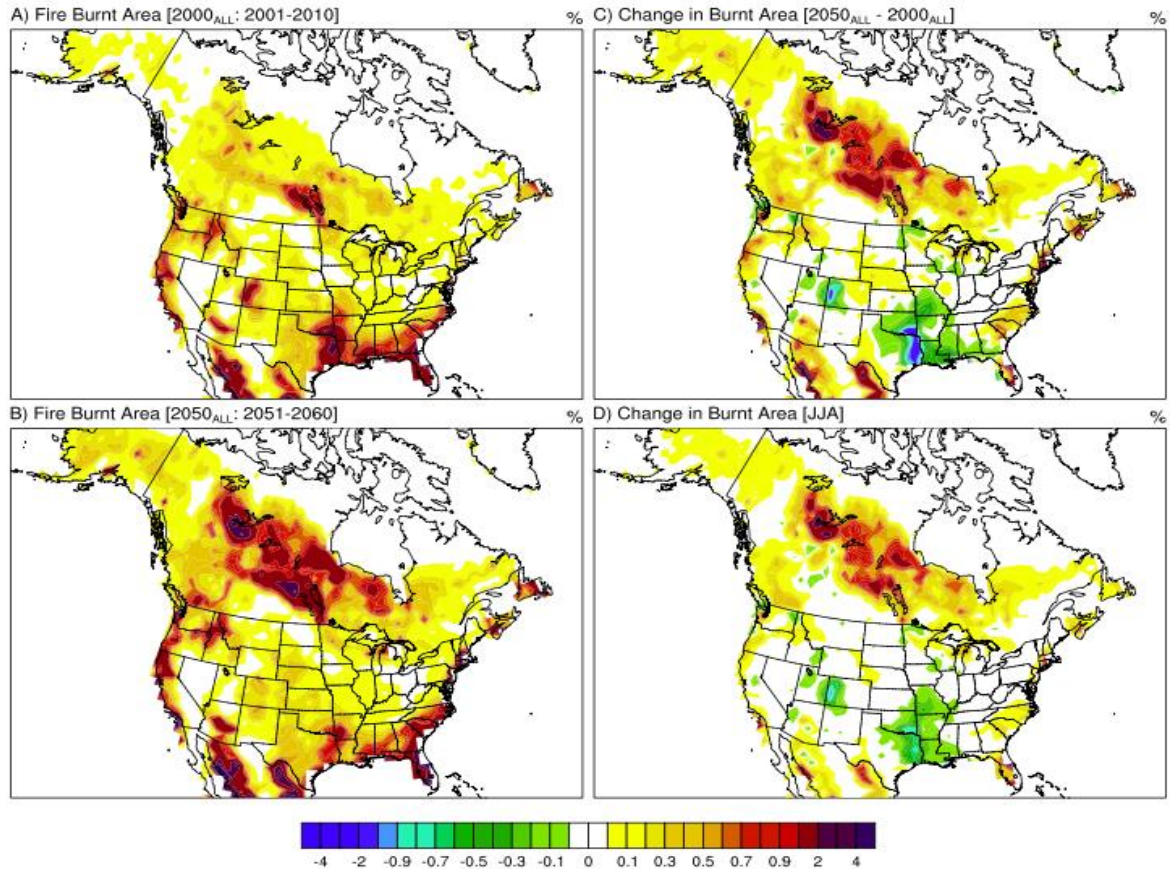
259

### 260 **3.2 Fire-induced changes in burnt area and PM<sub>2.5</sub>**

261 The decadal-mean annual fire burnt area simulated for the present day shows  
262 widespread wildfires over the entire North America (Figure 2A). Specifically, Canada and the  
263 forested areas of the northwestern ( $> 36$  N latitude) and southeastern ( $< 36$  N latitude) US are  
264 most intensely affected by wildfires in the present day. By the mid-21<sup>st</sup> century, a striking  
265 increase of 2-5 times in fire burnt area is projected over Canada, Alaska, the Pacific  
266 Northwest and portions of the western US by the 2050s (Figure 2B). A distinct positive shift  
267 in the probability density function (PDF) of annual fire burnt area is evident in the future,  
268 with the decadal-mean difference statistically significant at the 99% confidence level (Zou et  
269 al., 2020). A small and statistically insignificant change in interannual variability ( $\sim 0.4$  Mha  
270  $\text{yr}^{-1}$ ) of fire burnt areas is also simulated between the present and future. Specifically, our  
271 model predicts more than a doubling of burnt area in boreal regions of Canada in the future,  
272 in line with a previous projection for Canada (Wotton et al., 2017). Future enhancement in  
273 fire burnt area is  $\sim 20$ -50% in most fire grids over the western coast of US, which is higher  
274 than that over the eastern US (Figures 2A and 2C). The increase over the western US is closer  
275 to the lower bound of that derived from statistical model ensemble projections for the western  
276 US in the mid-21<sup>st</sup> century (Yue et al., 2013). The statistics-based projections of future burnt

277 area over North America were likely too high because fire-induced land cover change, fuel  
278 load reduction and factors could induce a negative fire feedback, which was not considered in  
279 previous fire projection studies (Zou et al. 2020).

280 Annual fire burnt area in the southeastern US shows a decline in the future (Figure  
281 2C), as precipitation is projected to increase in that region (discussed later). Note that all  
282 future fire changes between 2050<sub>ALL</sub> and 2000<sub>ALL</sub> are primarily associated with climate  
283 warming in response to the increase in greenhouse gas (GHG) concentrations in the RCP4.5  
284 scenario. No direct impacts of population and socioeconomic changes on wildfires are  
285 included in our simulations, although these factors contribute to changes in GHG emissions  
286 (via the RCP scenario) that influence the climate simulated in 2000<sub>ALL</sub> and 2050<sub>ALL</sub>. As about  
287 80% of the projected fire changes in the future is restricted to the summer season (June  
288 through August; JJA) (Figure 2D), we focus on analysis of the summer-mean wildfire-  
289 induced PM<sub>2.5</sub> and its projected future changes over North America.



290

291 **Figure 2: Spatial distribution of fire burnt area.** **A-D**, Spatial distribution of simulated  
 292 decadal-mean annual burnt area (as percentage) over North America for present day (A),  
 293 mid-21<sup>st</sup> century (B) and the net change between the 2050s and the 2000s (C). **D**, same as  
 294 (C), but for wildfire burnt area during summer only (June through August; JJA). The colorbar  
 295 illustrate grid fraction of area burnt.

296

297

The simulated 10-year averaged summer-mean wildfire-induced PM<sub>2.5</sub> values in

298 2000<sub>ALL</sub> are more than 0.5 μg/m<sup>3</sup> over a large part of North America in the present day, with

299 noticeably larger values (> 1 μg/m<sup>3</sup>) in Canada and the northwestern, central, and

300 southeastern US (Figure 3A). Interestingly, the spatial distribution of wildfire-induced

301 PM<sub>2.5</sub> > 1 μg/m<sup>3</sup> resembles an inverted horse-shoe shape. The inverted horse-shoe shaped

302 spatial distribution is also consistent with the wildfire-smoke climatology derived from the

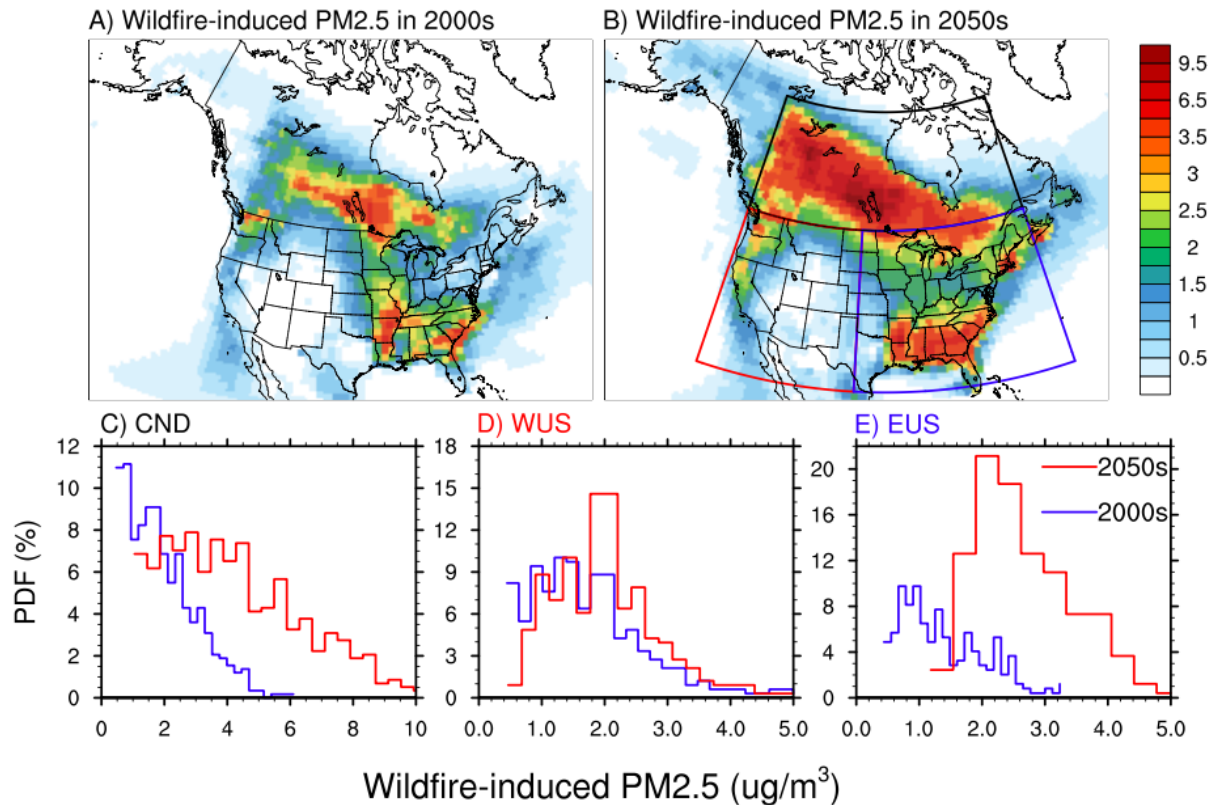
303 satellite-guided operational smoke product of the Hazard Mapping System (HMS) during

304 2005-2015 (Brey et al., 2018; Kaulfus et al., 2017). By the mid-21<sup>st</sup> century, the spatial extent

305 of the horse-shoe shape for areas with wildfire-induced PM<sub>2.5</sub> enhancement > 1 μg/m<sup>3</sup>

306 expands significantly to span most regions of North America, with the most pronounced  
307 enhancement occurring over Canada (Figure 3B). The PDFs of the spatial distribution for the  
308 three regions can be seen in Figure 3C-E. Specifically, wildfire induced PM<sub>2.5</sub> in the 2000s  
309 over Canada, WUS and EUS during summer is ~ 1-3 µg/m<sup>3</sup>, 1-3 µg/m<sup>3</sup> and 0.6-1.2 µg/m<sup>3</sup>,  
310 respectively. Maximum values within the WUS region are found over the Pacific Northwest,  
311 with most areas having wildfire induced PM<sub>2.5</sub> values of ~ 2-3 µg/m<sup>3</sup>. Similarly, the southern  
312 states have relatively high wildfire induced PM<sub>2.5</sub> concentrations of ~ 2-4 µg/m<sup>3</sup> within the  
313 EUS in the present-day simulation.

314           Compared to the 2000s, the wildfire induced JJA-averaged PM<sub>2.5</sub> values are almost  
315 doubled to ~ 3-6 µg/m<sup>3</sup> over Canada in the 2050s (Figure 3B and Figure 3C). Consistently,  
316 the values of wildfire induced PM<sub>2.5</sub> over WUS (mainly coastal) also doubled in 2050s  
317 compared to 2000s, with modal values of ~ 2-2.5 µg/m<sup>3</sup> (Figure 3D). Most interestingly, the  
318 enhancement in wildfire-induced summer-mean PM<sub>2.5</sub> over the northern EUS is also  
319 significant by the 2050s (Figures 3B). Largely, the summer-mean wildfire-induced PM<sub>2.5</sub>  
320 concentration over EUS increases from ~0.8 to ~2 µg/m<sup>3</sup> in the mid-century to values of 1.2-  
321 3.0 µg/m<sup>3</sup> (Figure 3E). The summer-mean wildfire-induced PM<sub>2.5</sub> is thus projected to double  
322 in North America by the 2050s compared to the 2000s, with a substantial coverage over the  
323 EUS. An important finding from these PDFs appears to be that there are fewer grids with < 1  
324 µg/m<sup>3</sup> wildfire induced PM<sub>2.5</sub>, or alternatively, that more regions are being influenced by  
325 PM<sub>2.5</sub>, and many areas that were already seeing wildfire impacts are seeing enhanced  
326 impacts. Such enhancement is found not only at the surface but also in an elevated  
327 atmospheric layer over EUS between 900 and 700 hPa. This is nonintuitive given the fact that  
328 the increase in fire-burnt area by mid-century over the EUS is not substantial.



329

330 Figure 3: **Spatial distribution of PM<sub>2.5</sub> concentrations.** Spatial distribution of decadal-mean  
 331 wildfire-induced enhancement in summer (June through August; JJA) PM<sub>2.5</sub> concentration  
 332 over North America for present day (A, 2000<sub>ALL</sub>-2000<sub>WEF</sub>) and future (B, 2050<sub>ALL</sub>-2050<sub>WEF</sub>).  
 333 C-E, Probability density functions (PDFs) of wildfire contribution within the three regions  
 334 shown in Figure 2B for Canada (CND: black box) (C), WUS (red box) (D), and EUS (blue  
 335 box) (E), respectively, for 2000s (blue) and 2050s (red). Only grids over land in North  
 336 America are used to generate the PDFs. The y-axis indicates the probability of occurrence of  
 337 different PM<sub>2.5</sub> values shown in the x-axis. The colorbar illustrates PM<sub>2.5</sub> in ug/m<sup>3</sup>.

338

339 As anthropogenic- and wildfire-induced PM<sub>2.5</sub> concentrations may change differently

340 with time across North America, next, we investigate the relative contribution of wildfire-

341 induced PM<sub>2.5</sub> to the total PM<sub>2.5</sub> in the future. Prominent enhancement of the wildfire

342 contribution is apparent in the entire domain by the 2050s (Figures 4A-B). Largely, during

343 the 2000s, the simulated fractional contribution of wildfires to PM<sub>2.5</sub> is ~15-50 % in Canada

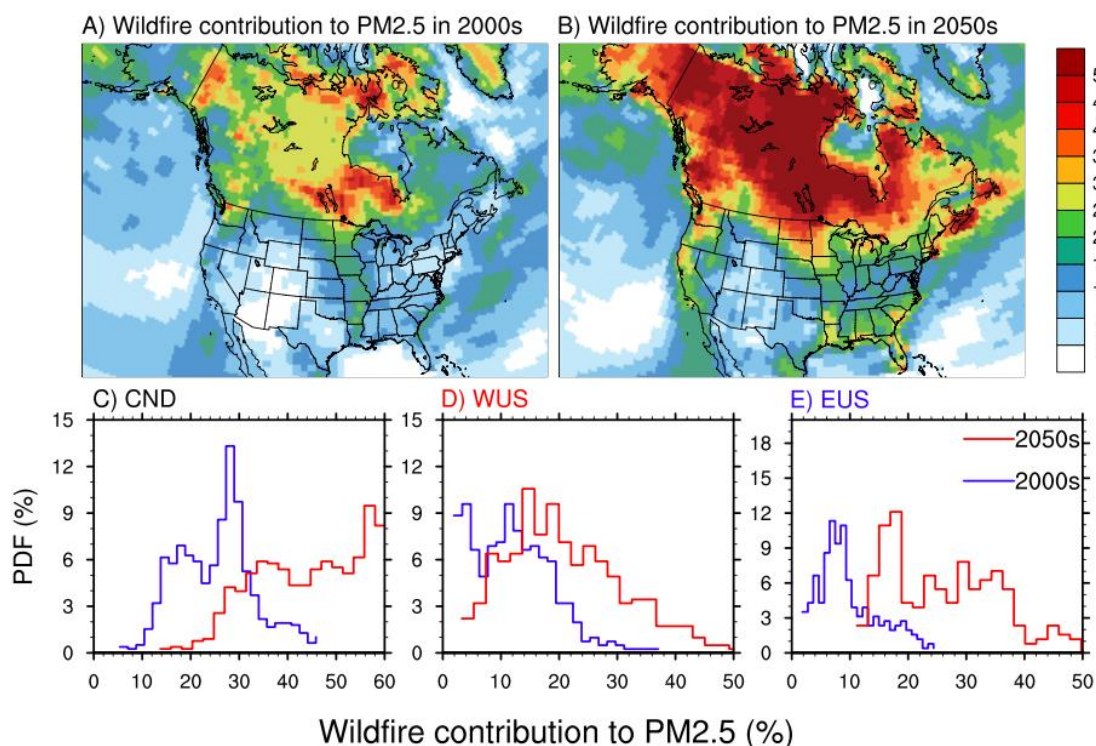
344 (Figure 4A). Specifically, a bi-modal distribution is simulated over Canada with modal values

345 around 18% and 30% (Figure 4C). Over WUS, the present day simulated percentage

346 contributions of wildfire-induced values are 5-25% (Figure 4A), with modal values between



347 10-20% (Figure 4D). Note that many areas located in the Pacific Northwest have higher  
 348 values of ~30-40% (Figure 4A). At the same time, the fractional contribution by wildfire-  
 349 induced PM<sub>2.5</sub> is ~5-10% in most areas of EUS in present day (Figure 4F). Nevertheless,  
 350 some areas in the central US also have higher values of ~10-25% (Figure 4A).



351

352 **Figure 4: Spatial distribution and probability density function of the percentage**  
 353 **contribution of wildfire emissions. A-B,** Spatial distribution of the percentage contribution  
 354 of wildfire emissions to decadal-averaged summer (June through August; JJA) mean PM<sub>2.5</sub>  
 355 concentrations over North America during present day (A) and future (B). The percentage  
 356 contribution of wildfire-induced PM<sub>2.5</sub> to the total PM<sub>2.5</sub> concentrations is calculated as  
 357  $([2000_{ALL}-2000_{WEF}]/2000_{ALL})$  and  $([2050_{ALL}-2050_{WEF}]/2050_{ALL})$  for the present and future,  
 358 respectively. **C-E,** Probability density functions (PDFs) of the percentage wildfire  
 359 contribution within the three regions shown in Figure 2D for Canada (CND: black box) (C),  
 360 WUS (red box) (D), and EUS (blue box) (E), respectively, for the 2000s (blue) and the 2050s  
 361 (red). Only grids over land in North America are used to generate the PDFs. The y-axis  
 362 indicates the probability of occurrence of different PM<sub>2.5</sub> values shown in the x-axis.

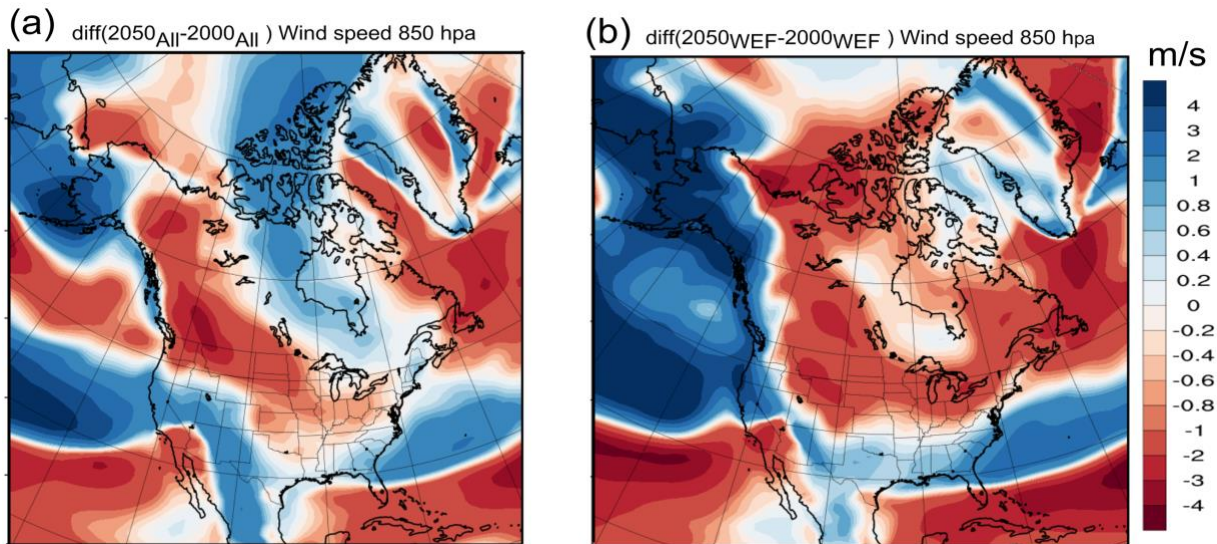
363

364 The wildfire contributions in the 2050s show a clear shift towards higher values in all  
 365 sub-regions compared to the 2000s (Figure 4B). Over Canada, the values shifted from 15-30

366 % in the 2000s to ~30-60% in the 2050s, a nearly two-fold increase in the fractional  
367 contribution of wildfire emissions to the total PM<sub>2.5</sub> concentration is simulated (Figure 4B  
368 and corresponding PDF in Figure 4C). Similarly, the contribution values increased to ~ 10-35  
369 % in the 2050s, compared to 10-20% in the 2000s over WUS (Figure 4B), thereby featuring a  
370 broadening of the bi-modal distribution of wildfire contribution (Figure 4D). The shift in the  
371 percentage contribution is most prominent for the higher values, corresponding to some areas  
372 located in the Pacific Northwest and west coast of the US (Figure 4B). Consistent with Figure  
373 3B, the shift in the contribution values over EUS is also very distinct, revealing an increase in  
374 the mode values from 6-10% in the 2000s to ~16-20 % by the 2050s (Figure 4B and Figure  
375 4E). Thus, not only in absolute values, but our results also underscore a large increase in the  
376 contribution of wildfire emissions over EUS in the future.

### 377 **3.3. Mechanistic understanding of the underlying processes**

378 The larger enhancement in the relative contribution of wildfire emissions to the total  
379 surface PM<sub>2.5</sub> in EUS in the 2050s can be explained by three mechanisms. First, due to the  
380 increase in Canadian and western US wildfires, downwind transport of wildfire smoke  
381 plumes to EUS will be enhanced by the 2050s. This long-range transport to the atmospheric  
382 column of EUS can happen within a few days of the fire occurrence (Supplementary Figures  
383 3A and 3B). Using Hazard Mapping System (HMS)-detected smoke plumes, recent studies  
384 identified a strong positive association between the transported smoke plumes in the  
385 atmospheric column and collocated surface PM<sub>2.5</sub> enhancement in EUS (Brey et al., 2018;  
386 Wu et al., 2018; Gunsch et al., 2018; Kaulfus et al., 2017; Larsen et al., 2017; Dempsey,  
387 2013). Hazard Mapping System (HMS) is an operational smoke detection product over North



388

389 Figure 5a: Spatial distribution of decadal mean summer (June through August; JJA) wildfire-  
 390 induced future changes  $[(2050_{ALL}-2050_{WEF}) - (2000_{ALL}-2000_{WEF})]$ . A) Wind speed at 850 hpa  
 391 for  $[(2050_{ALL} - 2000_{ALL})]$ , B) Wind speed at 850 hpa  $[(2050_{WEF} - 2000_{WEF})]$ .

392 America known as developed by the National Oceanic and Atmospheric Administration  
 393 (NOAA) and operated by National Environmental Satellite, Data, and Information Service  
 394 (NESDIS), available at <http://satapsanone.nesdis.noaa.gov/FIRE/fire.html>. Specifically, these  
 395 studies found that the smoke plumes transported from Canada are located at an altitude of ~  
 396 1-3 km over EUS (Colarco et al., 2004; Wu et al., 2018). Due to mixing by the daytime  
 397 boundary layer and deposition, the smoke plumes enhance the surface PM<sub>2.5</sub> concentration  
 398 over EUS (Wu et al., 2018; Colarco et al., 2004; Rogers et al., 2020; Dreessen et al., 2015).  
 399 Hence HMS smoky days may be a useful proxy for wildfire-induced surface PM<sub>2.5</sub> over  
 400 North America. In agreement, Brey et al. (2018) showed that the HMS-based smoke plumes  
 401 observed over EUS is significantly aged, suggestive of their long-range transport origin.  
 402 Consistent with the observed temporal change in HMS pattern, Xue et al. (2021) estimated  
 403 using the mid-visible Multi Angle Implementation of Atmospheric Correction (MAIAC)  
 404 satellite-derived Aerosol Optical Depth (AOD) that Canadian and western US fires have  
 405 caused an increase in the daily PM<sub>2.5</sub> over Montana, North Dakota, South Dakota and  
 406 Minnesota by 18.3, 12.8, 10.4 and 10.1  $\mu g m^{-3}$ , respectively, between August 2011 (a low

407 fire month) and August 2018 (a high fire month). In summary, the visually apparent satellite-  
408 based signatures of wildfire-smoke across Canada and EUS provide a necessary, though not  
409 sufficient, support for the influence of Canadian smoke plumes on EUS air quality. Although,  
410 the change in burnt area over northeastern EUS is negligible compared to the western US and  
411 Canadian regions, however, there are some enhancements seen over east coast of US, which  
412 can also contribute to enhanced fire emissions.

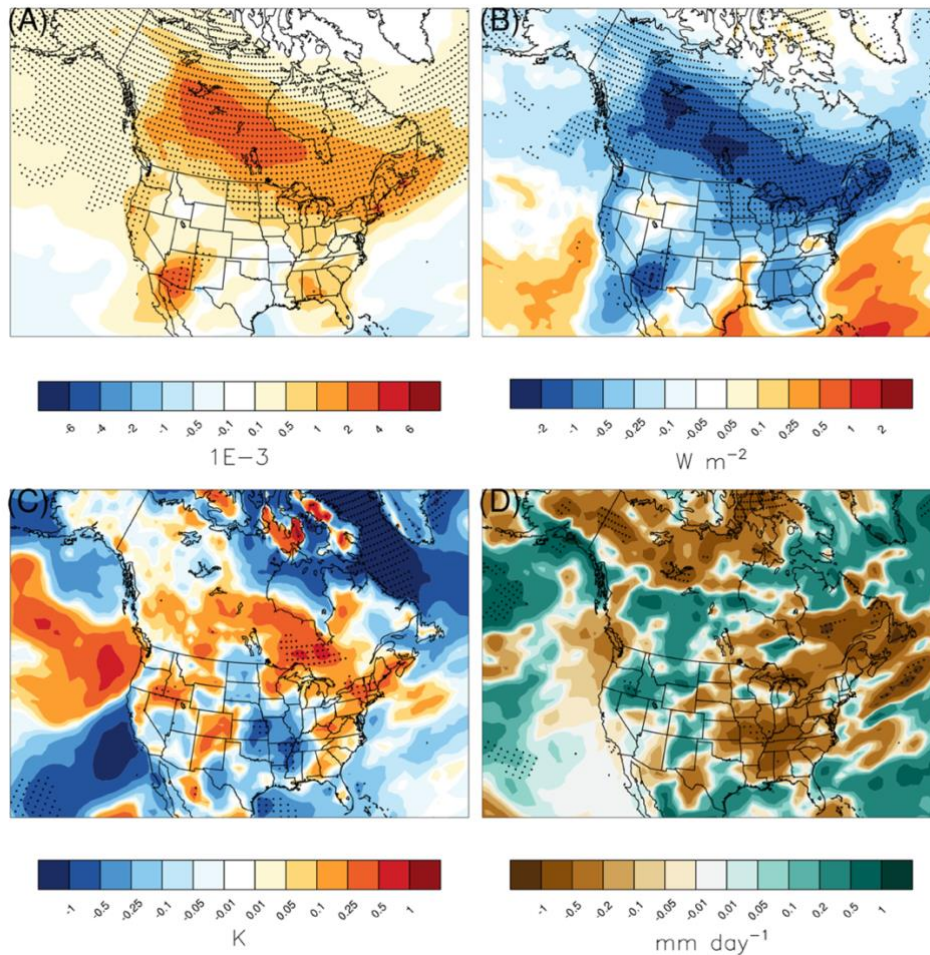
413 In future, the simulated speed of the westerly jet flows over Canada wildfire regions is  
414 reduced in both the scenarios (Figure 5 A-B). It indicates that the westerly-induced  
415 transported wildfire emissions from Canada boreal forests to the eastern half of Northern  
416 America and EUS will be slower in future compared to that in present era. On the one hand, it  
417 implies that the advection of smoke plumes will be slightly reduced in future. But on the  
418 other hand, this phenomenon can also contribute to the enhanced PM<sub>2.5</sub> values at surface as  
419 these transporting plumes will be subject to relatively more boundary layer mixing over the  
420 EUS and dry deposition/settling enhances. At the same time, the westerly winds over western  
421 US below 40 °N is strengthened in future (Figure 5 A-B) compared to present day which  
422 indicates more advection flux of wildfire emissions to EUS. Thus, the net effect is more  
423 removal of wildfire-emitted PM<sub>2.5</sub> from WUS and more influx of wildfire-emitted PM<sub>2.5</sub> in  
424 EUS.

425 Along with this dynamical changes, other climatic feedbacks simulated can also contribute to  
426 enhancement of EUS pollution. Specifically, the enhancement of wildfire-induced smoke  
427 aerosols increases solar absorption and scattering in the future (Figure 6A). This reduces the  
428 incoming solar radiation reaching at the surface (Figure 6B) and induces surface cooling.  
429 With atmospheric warming and surface cooling, lower-tropospheric stability is enhanced by  
430 wildfire aerosols in the future (Figure 6C). The smoke plumes which reaches eastern US are

431 at an elevated altitude due to self-lofting property of absorbing aerosols as they travel  
432 downwind but the smoke over western US is at near surface elevation as it is at its source  
433 region. This can explain the more significant atmospheric stability simulated over the eastern  
434 US compared to the source regions in western US and boreal forests of Canada. Relatively  
435 stronger atmospheric stability over eastern US impose a stronger thermal capping that traps  
436 more anthropogenic aerosols and particulate matter near the surface over EUS (already an  
437 emission hotspot). At the same time, future increase in wildfire emissions also leads to  
438 greater reduction of monthly rainfall (Figure 6D) over EUS, which may additionally  
439 strengthen the positive feedback to surface PM<sub>2.5</sub> over EUS by reducing wet scavenging of  
440 transported wildfire smoke to EUS. Thus, wildfire-emitted aerosols induce positive feedback  
441 on the surface PM<sub>2.5</sub> concentration over EUS through fire-climate interactions that vary on a  
442 regional scale. Moreover, the above discussed dynamical changes in future can also feedback  
443 these simulated thermodynamical and precipitation changes, exaggerating the enhancement in  
444 PM<sub>2.5</sub> values over EUS in future. However, due to computational constrains, no direct  
445 quantification of the magnitude of these feedback (with aerosol-radiation and aerosol-cloud  
446 interactions turned off) on PM<sub>2.5</sub> is performed and would be taken up in future studies.

447         Lastly, the reason of why the contribution of wildfire emissions to the total surface  
448 PM<sub>2.5</sub> in EUS is so substantial in the 2050s is the drastic reduction of anthropogenic  
449 contribution to the surface PM<sub>2.5</sub> over EUS in the future primarily due to policy-driven  
450 reduction in anthropogenic emissions under the RCP4.5 scenario. Specifically, the simulated  
451 ambient summer mean PM<sub>2.5</sub> concentration exhibits widespread declines in the future  
452 (Supplementary Figure 4), with reduction in PM<sub>2.5</sub> concentration over eastern US in the range  
453 of 4-15  $\mu\text{g}/\text{m}^3$ , which is greatest within North America. Thus, large reduction in  
454 anthropogenic contribution combined with increased downwind advection of Canadian

455 smoke to EUS and the associated positive feedbacks can explain the projected dominance of  
456 wildfire emissions over EUS in future.



457  
458 Figure 6: Spatial distribution of decadal mean summer (June through August; JJA) wildfire-  
459 induced future changes [(2050<sub>ALL</sub>-2050<sub>WEF</sub>) - (2000<sub>ALL</sub>-2000<sub>WEF</sub>)]. A) aerosol absorption  
460 optical depth at 550 nm, B) aerosol direct radiative forcing at surface, C) lower-tropospheric  
461 stability calculated as the difference between the potential temperature at 900 hPa and 1000  
462 hPa, D) summer averaged precipitation rates, over North America. Areas marked with black dots  
463 indicate grids where changes are significant at the 95% confidence level.

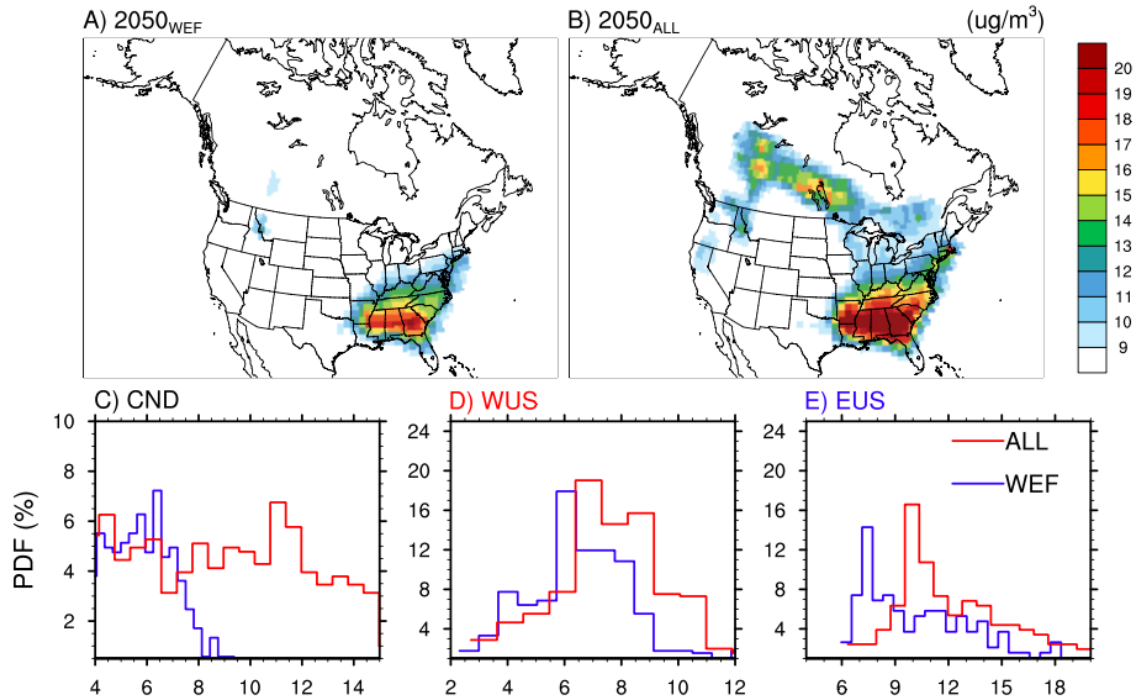
464

### 465 **3.4 Future Implications and uncertainties**

466 However, is the simulated future enhancement in wildfire contribution over EUS  
467 substantial enough to affect the surface PM<sub>2.5</sub> values over EUS in future? The World Health  
468 Organization (WHO) air quality guidelines for annual and daily PM<sub>2.5</sub> concentration are 10  
469 μg m<sup>-3</sup> and 25 μg m<sup>-3</sup>, respectively. As no specific guideline for seasonal-mean PM<sub>2.5</sub> in the

470 summer is available, we use the annual guideline value as a reference to understand the  
471 implication of wildfire emissions on ambient PM<sub>2.5</sub> concentration in the future. Interestingly,  
472 the mean summertime PM<sub>2.5</sub> concentration in the wildfire emission free (WEF) scenario is  
473 projected to remain within 10 µg m<sup>-3</sup> over most of North America, except for the  
474 southeastern US (~15% of the domain) (Figure 7A). However, the ALL-scenario projects an  
475 increase in the exposure concentration level such that values > 10 µg/m<sup>3</sup> are common in  
476 Canada and EUS in the future (Figure 7B). Quantitatively, over Canada, the entire PDF of  
477 PM<sub>2.5</sub> concentration shifts towards higher values by ~5-6 µg/m<sup>3</sup>. Specifically, the modal value  
478 shifts from ~ 6 µg/m<sup>3</sup> in 2050<sub>WEF</sub> to 11-12 µg/m<sup>3</sup> in 2050<sub>ALL</sub> (Figure 7C), so PM<sub>2.5</sub>  
479 concentration is projected to surpass the WHO guidelines over a large fraction of Canada in  
480 the future. Similarly, the entire PDF of PM<sub>2.5</sub> concentration shifts towards higher values by  
481 ~2-3 µg/m<sup>3</sup> over EUS, with the mode of the PDF increasing from ~ 7-8 µg/m<sup>3</sup> in 2050<sub>WEF</sub> to  
482 ~ 10-11 µg/m<sup>3</sup> in 2050<sub>ALL</sub> (Figure 7E). The modal value of summer mean PM<sub>2.5</sub> over WUS  
483 increases from ~ 6 µg/m<sup>3</sup> in 2050<sub>WEF</sub> to ~ 7-8 µg/m<sup>3</sup> in 2050<sub>ALL</sub> (Figure 7D), although a few  
484 grid cells show PM<sub>2.5</sub> values greater than 10 µg/m<sup>3</sup> (Figure 7B).

485         Clearly, the climate-induced enhancement in fires and its influence via the advected  
486 wildfire smoke to EUS can have significant implications for air quality management in the  
487 future. The PM<sub>2.5</sub> enhancement in future over the southern states within EUS is large (Figure  
488 7A-B), which is consistent with Figure 3 and 4 results. However, the future change in burnt  
489 area over the same region is negligible or mostly reducing (Figure 1C-D). Thus, it can be  
490 argued that the simulated enhancement is mostly related with the dynamic perturbations and  
491 thermodynamical feedbacks due to wildfire emissions (Figure 6). As the rate of  
492 anthropogenic emissions is also regionally highest over the Southeastern states, the impact of  
493 these wildfire-induced climatic feedbacks on local air quality is distinctly seen over the EUS.



494

495 **Figure 7: Spatial distribution and probability density function of PM<sub>2.5</sub> concentration in**  
 496 **2050s. A-B**, Spatial distribution of decadal-average summer (June through August; JJA)  
 497 mean PM<sub>2.5</sub> concentration over North America in mid-21<sup>st</sup> century from 2050<sub>WEF</sub> (wildfire  
 498 emission-free) (A) and 2050<sub>ALL</sub> (wildfire emission-inclusive) (B). **C-E**, Probability density  
 499 functions (PDFs) of the same within the three regions shown in Figure 2B for Canada; CND  
 500 (C), western US; WUS (D), and eastern US; EUS (E), respectively, for the 2050<sub>WEF</sub> (blue)  
 501 and 2050<sub>ALL</sub> (red) runs. The y-axis indicates the probability of occurrence of different PM<sub>2.5</sub>  
 502 values shown in the x-axis. Only grids over land in North America are used to generate the  
 503 PDFs. Note the different ranges of values shown in the y- and x-axis in **C-E**. The colorbar  
 504 and the x-axis for Panel C-E indicates PM<sub>2.5</sub> values.

505 Note that our simulated present-day estimates of wildfire induced PM<sub>2.5</sub> values as well  
 506 as the percentage contribution of wildfire emissions are within the range of reported values in  
 507 previous studies over the domain, which augment the fidelity our future projections.

508 Specifically, our simulated present-day estimates of wildfire induced PM<sub>2.5</sub> values are also  
 509 within the range of reported values in previous studies over the domain. Reported values of  
 510 wildfire-induced PM<sub>2.5</sub> over WUS during summertime vary from  $\sim 1 \mu\text{g}/\text{m}^3$  (Jaffe et al., 2008)  
 511 to  $\sim 2 \mu\text{g}/\text{m}^3$  (Park et al., 2007) and  $\sim 3 \mu\text{g}/\text{m}^3$  (Ford et al., 2018), with the highest values  
 512 documented over the Pacific Northwest and west coast regions ( $\sim 1-4 \mu\text{g}/\text{m}^3$ ) (O'Dell et al.,  
 513 2019). The wildfire-induced PM<sub>2.5</sub> over EUS during summertime varies from  $\sim 1 \mu\text{g}/\text{m}^3$  (Park



514 et al., 2007) to  $\sim 2.5 \mu\text{g}/\text{m}^3$  ( $\sim 3 \mu\text{g}/\text{m}^3$  in the southeastern US) (Ford et al., 2018).  
515 Consistently, our simulated present-day estimates of wildfire contribution values are also  
516 within the range of reported values in previous studies. For example, Meng et al. (2019)  
517 found that wildfires can be the largest sectoral contributor ( $\sim 18\text{-}59\%$ ) to the population-  
518 weighted  $\text{PM}_{2.5}$  in various subregions of Canada. Over WUS, the present-day percentage  
519 contribution of wildfire induced  $\text{PM}_{2.5}$  to the total  $\text{PM}_{2.5}$  is reported to be  $\sim 12\%$  (Liu et al.,  
520 2017),  $\sim 15\%$  (Park et al., 2007) and  $\sim 30\%$  (Ford et al., 2018), with higher values of  $\sim 40\%$  in  
521 the Pacific Northwest (O'Dell et al., 2019). Over EUS our simulated values are also within  
522 the range of previously reported values of  $\sim 5\%$  (Park et al., 2007) and  $\sim 15\text{-}18\%$  (Ford et al.,  
523 2018). However, our two-way coupled simulations illustrate that future enhancement in the  
524 wildfire associated  $\text{PM}_{2.5}$  over the EUS could be greater compared to the western US, which  
525 is not emphasized explicitly in any of the previous studies (although Ford et al., 2018  
526 illustrated increase in  $\text{PM}_{2.5}$  over mid and central US from Canadian fires). These could be  
527 since inclusion of the wildfire-induced climatic feedbacks in our simulation is an  
528 unprecedented exercise. Please also note that our study is focused on JJA period and the  
529 wildfires in western US mainly occurs during August-September months, so the results  
530 should be compared consciously.

531         Nonetheless, inherent limitations in our simulations may introduce uncertainties in the  
532 projected future changes. For example, our reported changes in  $\text{PM}_{2.5}$  concentrations based on  
533 relatively coarse resolution simulations and decadal averages likely represent a low-end  
534 estimate compared to changes at regional and daily/weekly scales. Moreover, our  
535 experiments do not consider the direct human influences such as population change and  
536 socioeconomic development on wildfires, which may aggravate the increase in  $\text{PM}_{2.5}$   
537 concentrations over the densely populated EUS in the future. Common sources of uncertainty  
538 in modeling burnt area and fire emission and fire aerosol and smoke are also present in our

539 model. Fire smoke, in particular, is extremely hard to measure and evaluate. Lastly, inherent  
540 uncertainties in the physics parameterizations used in the model, sensitivity of climate to  
541 GHGs emissions, and the RCP scenarios should also be noted. Thus, ensemble modeling  
542 considering different emissions scenarios, population and future time periods, and the use of  
543 a finer spatial resolution may provide a more robust and better quantification of the wildfire-  
544 induced impact on policy regulated improvements in PM<sub>2.5</sub> over EUS.

#### 545 **4. Conclusion**

546 In summary, online coupled fire-climate-ecosystem simulations project a nearly  
547 twofold increase in wildfire-induced summer-mean surface PM<sub>2.5</sub> concentration by the mid-  
548 21<sup>st</sup> century over the entire North America. In a wildfire-emission free future, a large portion  
549 of North America will have PM<sub>2.5</sub> values below the WHO guidelines. But in a future with  
550 wildfire emissions, the improvements from policy-driven reductions in anthropogenic PM<sub>2.5</sub>  
551 will be compromised by the projected doubling of PM<sub>2.5</sub> from wildfires. More strikingly,  
552 wildfire-induced enhancement in surface PM<sub>2.5</sub> values and percentage contribution of the  
553 wildfire emissions over EUS could be substantial by mid-century. This is mainly because of  
554 the large enhancement in fires over Northern America by 2050s and associated increase in  
555 amount of downwind transport of smoke to EUS. In addition, enhancement of smoke  
556 transport induces a positive climate feedback to PM<sub>2.5</sub> concentrations over EUS by increasing  
557 the lower-tropospheric stability and reducing wet scavenging rates. Despite the inherent  
558 limitations, this study highlights the natural versus anthropogenic contributions and the non-  
559 local nature of air pollution that can complicate regulatory strategies aimed at improving air  
560 quality over the eastern US in a warmer future.

#### 561 **Data availability statements**

562 The HMS data used in this paper are available free through the link  
563 <https://www.ospo.noaa.gov/Products/land/hms.html>. The model simulations data are  
564 available at <https://portal.nersc.gov/project/m1660/yang560/wildfire>

### 565 **Code availability statements**

566 The model code and scripts are available at  
567 <https://portal.nersc.gov/project/m1660/yang560/wildfire>

568

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579

### 580 **Authors Contribution**

581 YQ, CS and RL designed this study. CS did the model and satellite analysis and wrote the  
582 first draft of the manuscript. YZou performed the simulations. All authors provided inputs  
583 throughout the study and helped in the drafting and submission process.

584

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