Supplementary Material for

Source Apportionment and Evolution of N-containing Aerosols at a Rural Cloud Forest in Taiwan by Isotope Analysis

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## $\delta^{18}$ O variation in NO<sub>3</sub><sup>-</sup> oxidation processes

The oxidation processes for aerosol  $NO_3^-$  formation can be represented as follows and shown as R1 to R10 in Fig. 1:

$$NO + O_3 \rightarrow NO_2 + O_2$$
 (R1)

$$NO_2 + hv \rightarrow NO + O$$
 (R2)

$$O + O_2 \longrightarrow O_3$$
 (R3)

OH radical oxidation:

$$NO_2 + OH \longrightarrow HNO_3$$
 (R4)

N<sub>2</sub>O<sub>5</sub> hydrolysis at nighttime:

$NO_2 + O_3 - $	$\rightarrow NO_3 + O_2$	(R5)
$NO_2 + O_3 - $	$\rightarrow NO_3 + O_2$	(R5

$NO_2 + NO_3 \xrightarrow{M} N_2O_5$	(R6)

 $N_2O_5 + H_2O \longrightarrow 2HNO_3$  (R7)

**OH formation:** 

$$O_3 + h\nu \longrightarrow O + O_2$$
 (R8)

$$O + H_2 O \longrightarrow 2OH$$
 (R9)

peroxyl radical oxidation:

 $NO + RO_2 \longrightarrow NO_2 + RO$  (R10)

If both oxygen atoms of NO<sub>2</sub> are from O<sub>3</sub> oxidation (R1), the O of NO<sub>3</sub><sup>-</sup> would be 2/3 from O<sub>3</sub> and 1/3 from OH via OH oxidation (R4). On the other hand, if NO<sub>3</sub><sup>-</sup> is formed via N<sub>2</sub>O<sub>5</sub> hydrolysis during nighttime (R7), O atoms on NO<sub>3</sub><sup>-</sup> would be 5/6 from O<sub>3</sub> and 1/6 from H<sub>2</sub>O. Hastings et al. (2003) suggested that the  $\delta^{18}$ O of OH radical was close to that of atmospheric H<sub>2</sub>O (from -15‰ to 0‰ over the Asian continent, Global Network of Isotopes in Precipitation, IAEA); therefore,  $\delta^{18}$ O of NO<sub>3</sub><sup>-</sup> from OH oxidation (where  $\delta^{18}$ O = +55‰ ~ +81.3‰) would be lower than that from N<sub>2</sub>O<sub>5</sub> hydrolysis at night (where  $\delta^{18}$ O = +72.5‰ ~ +101.67‰), as shown as yellow blocks in Fig. S6 and S7 (Hastings et al., 2003; Fang et al., 2011). However, if atmospheric OH radicals are mainly produced via R8 & R9, the  $\delta^{18}$ O value of atmospheric OH could be the mean value of O<sub>3</sub> and H<sub>2</sub>O in the range of +37.5‰ ~ +61.0‰ (half from O<sub>3</sub> and the other half from H<sub>2</sub>O). The calculation would result in the same  $\delta^{18}$ O (+72.5‰ ~ +101.67 ‰) of NO<sub>3</sub><sup>-</sup> formed through the OH oxidation and N<sub>2</sub>O<sub>5</sub> hydrolysis pathways.

If NO<sub>2</sub> has a significant contribution via the reaction of NO with peroxyl radicals (hydroperoxyl radical HO<sub>2</sub> or organic peroxyl radical RO<sub>2</sub>) as R10, O atoms of HO<sub>2</sub> or RO<sub>2</sub> are supposed to have a similar  $\delta^{18}$ O with atmospheric O<sub>2</sub> ( $\delta^{18}$ O = +23.5‰, Luz and Barkan (2011)) as the O atom of the peroxyl radical is originated from atmospheric O<sub>2</sub> (H + O<sub>2</sub> + M  $\rightarrow$  HO<sub>2</sub> + M). Therefore, NO<sub>2</sub> from R10 reacting OH can produce NO<sub>3</sub><sup>-</sup> with a lower  $\delta^{18}$ O (could be low as +28.2‰ ~ +36.0‰) compared to that of NO<sub>2</sub> from R5 (+90‰ to +122‰). If the NO  $\leftrightarrow$  NO<sub>2</sub> cycle has rapidly exchanged through O<sub>3</sub> oxidation and NO<sub>2</sub> photolysis during the daytime,  $\delta^{18}$ O of NO<sub>2</sub> would be close to that of O<sub>3</sub> after 7-8 times of reactions (Fig. S8). Therefore,  $\delta^{18}$ O of formed NO<sub>3</sub><sup>-</sup> would be +72.50-101.67‰ after either OH radical oxidation or nighttime N<sub>2</sub>O<sub>5</sub> hydrolysis shown as

(P4) and (P6) in Fig. S6. If RO<sub>2</sub> radicals involve the oxidation processes, NO<sub>3</sub><sup>-</sup>  $\delta^{18}$ O could be low to +28.17‰ shown as (P7) in Fig. S7.

	02D	02N	13N	14D	14N	15D	20N	21D	22D	22N
$\delta^{15}$ N, NH <sub>4</sub> <sup>+</sup>	14.38	9.93	13.33	16.30	12.30	14.92	11.32	7.61	10.43	8.94
δ <sup>15</sup> N, PM <sub>1-10</sub> -NO <sub>3</sub> -	2.72			3.70		4.24	1.85	5.20		
δ <sup>15</sup> N, PM <sub>1</sub> -NO <sub>3</sub> <sup>-</sup>			1.46	1.50		1.93				
δ <sup>18</sup> O, PM <sub>1-10</sub> -NO <sub>3</sub> -	70.05			70.75		72.45	74.82	72.52		
δ <sup>18</sup> O, PM <sub>1</sub> -NO <sub>3</sub> <sup>-</sup>			79.81	69.07		72.01				
δ <sup>15</sup> N, NO <sub>3</sub> <sup>-</sup>	2.72		1.46	2.29		2.83	1.85	5.20		
δ <sup>18</sup> O, NO <sub>3</sub> <sup>-</sup>	70.05		79.81	69.67		72.18	74.82	72.52		

Table S1. Mass weighted isotope value (‰) of each group of samples.



Figure S1. IR spectra of 3 collected samples on 2 December 2018. The absorption peaks for curve fitting are 2-curve fit for  $NH_4^+$ ,  $NO_3^-$ ; 3-curve fit for polytetrafluoroethylene (PTFE) and  $SO_4^{2-}$ ; the absorbance of BC is determined by the average absorption of  $3950 \pm 5$  cm<sup>-1</sup> where have less interference by the absorption of other chemical species.



Figure S2. The Lorentzian fitting curve of the absorption curve of  $NH_4^+$  and  $NO_3^-$ . The yellow thick curve is the background-corrected IR spectrum. Red and green curves are the fitted  $NH_4^+$  and  $NO_3^-$  absorption curves, respectively. The black curve is the add-up of  $NH_4^+$  and  $NO_3^-$  absorption, and the orange one represents the absolute value of the fitting curve minus the original curve.



Figure S3. The Lorentzian fitting curve of the absorption curve of  $SO_4^{2-}$  and polytetrafluoroethylene (PTFE). The red curve is the baseline-corrected absorption curve. Two gray curves are the fitting curve of PTFE filter absorption, and the green line is that of  $SO_4^{2-}$ . The black thick curve is the add-up of absorption curves, and the orange one represents the absolute value of the fitting curve minus the original curve.

Concentration comparison between different methods



Figure S4. Comparison between the measured water-soluble (TN minus NN) and FTIR measured  $NH_4^+$  concentration.



Figure S5. Statistic box plot of concentration of (A)  $SO_4^{2-}$ , (B) black carbon (BC) in non-foggy daytime and nighttime, and (C)  $SO_4^{2-}$ , (D) BC in the foggy and non-foggy daytime condition in different cut-off sizes. (diamond: mean value; outliers: < 1<sup>st</sup> quartile Q1-1.5 interquartile range (IQR) or > 3<sup>rd</sup> quartile Q3+1.5 IQR).



Figure S6. The predicted  $\delta^{18}$ O (gray shaded bottom-right boxes) of products during HNO<sub>3</sub> formation pathways assuming the  $\delta^{18}$ O of initial NO is 90~122‰ (i.e., fully reacted with O<sub>3</sub>). Products are in darkblue font color, reactants in light-blue, and by-products in green. The yellow-shaded boxes are based on the assumption of water-originated O atom of OH radical, whereas the O atoms of P1 and P4 are assuming that the OH is from O<sub>3</sub> and H<sub>2</sub>O. The reactions in the dashed boxes are related to N<sub>2</sub>O<sub>5</sub> hydrolysis occurring during nighttime.



Figure S7. The predicted  $\delta^{18}$ O (gray shaded bottom-right blocks) of products during HNO<sub>3</sub> formation pathways assuming the  $\delta^{18}$ O of initial NO is +23.5‰ (i.e., freshly emitted and close to that of atmospheric O<sub>2</sub>). Products are in dark-blue font color, reactants in light-blue, and by-products in green. The yellowshaded blocks are based on the assumption of water-originated O atom of OH radical, whereas the O atoms of P7 and P10 are assuming that the OH is from O<sub>3</sub> and H<sub>2</sub>O. The reactions in the dashed squares are related to N<sub>2</sub>O<sub>5</sub> hydrolysis, where the reactions only occur at night.



Figure S8. The predicted  $\delta^{18}O$  of freshly formed NO (assume the  $\delta^{18}O$  close of atmospheric  $O_2 = 23.5\%$ ) after exchanging with  $O_3$  (where  $\delta^{18}O = 90 \sim 122\%$ ).

## References

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