- 1 Reconciling the total carbon budget for boreal forest wildfire emissions using airborne
- 2 observations
- 3

4	Katherine L. Hayden <sup>1*</sup> , Shao-Meng Li <sup>2</sup> , John Liggio <sup>1</sup> , Michael J. Wheeler <sup>1</sup> , Jeremy J.B. Wentzell <sup>1</sup> , Amy
5	Leithead <sup>1</sup> , Peter Brickell <sup>1</sup> , Richard L. Mittermeier <sup>1</sup> , Zachary Oldham <sup>1,6</sup> , Cris Mihele <sup>1</sup> , Ralf M. Staebler <sup>1</sup> ,
6	Samar G. Moussa <sup>1</sup> , Andrea Darlington <sup>1</sup> , Mengistu Wolde <sup>3</sup> , Daniel Thompson <sup>4</sup> , Jack Chen <sup>1</sup> , Debora
7	Griffin <sup>1</sup> , Ellen Eckert <sup>1</sup> , Jenna C. Ditto <sup>5</sup> , Megan He <sup>5</sup> and Drew R. Gentner <sup>5</sup>
8 9 10 11 12 13 14	<ul> <li>[1]{Air Quality Research Division, Environment Canada, Toronto, ON, Canada}</li> <li>[2]{College of Environmental Sciences and Engineering, Peking University, Beijing, China}</li> <li>[3]{National Research Council of Canada, Ottawa, ON, Canada}</li> <li>[4]{Canadian Forest Service, Natural Resources Canada, Edmonton, AB, Canada}</li> <li>[5]{Yale University, New Haven, CT, USA}</li> <li>[6]{University of Waterloo, Waterloo, ON, Canada}</li> </ul>
15	*Correspondence to: Katherine Hayden (katherine.hayden@ec.gc.ca)

# 17 Abstract

18 Wildfire impacts on air quality and climate are expected to be exacerbated by climate change with the 19 most pronounced impacts in the boreal biome. Despite the large geographic coverage, there is limited 20 information on boreal forest wildfire emissions, particularly for organic compounds, which are critical 21 inputs for air quality model predictions of downwind impacts. In this study, airborne measurements of 22 193 compounds from 15 instruments, including 173 non-methane organics compounds (NMOG), were 23 used to provide the most detailed characterization, to date, of boreal forest wildfire emissions. Highly 24 speciated measurements showed a large diversity of chemical classes highlighting the complexity of 25 emissions. Using measurements of the total NMOG carbon (NMOG<sub>T</sub>), the  $\Sigma$ NMOG was found to be 26  $50\pm3$  to  $53\pm3$  % of NMOG<sub>T</sub>, of which, the intermediate- and semi-volatile organic compounds (I/SVOCs) 27 were estimated to account for 7 to 10 %. These estimates of I/SVOC emission factors expand the 28 volatility range of NMOG typically reported. Despite extensive speciation, a substantial portion of 29 NMOG<sub>T</sub> remained unidentified (47±15 to 50±15 %), with expected contributions from more highly-

30	functionalized VOCs and I/SVOCs. The emission factors derived in this study improve wildfire chemical
31	speciation profiles and are especially relevant for air quality modelling of boreal forest wildfires. These
32	aircraft-derived emission estimates were further linked with those derived from satellite observations
33	demonstrating their combined value in assessing variability in modelled emissions. These results
34	contribute to the verification and improvement of models that are essential for reliable predictions of
35	near-source and downwind pollution resulting from boreal forest wildfires.
36	
37	
38	
39	
40	

## 42 **1 Introduction**

43 Wildfires play a natural role in maintaining forest health and diversity through the release of 44 nutrients, seed germination, removal of aging vegetation, and reducing the spread of forest diseases. 45 Wildfires are, however, one of the largest global sources of trace gases and aerosols to the atmosphere 46 (Andreae, 2019; Yu et al., 2019) and can have deleterious impacts on human health (Cascio, 2018; Cherry 47 and Haynes, 2017; Reid et al., 2016; Finlay et al., 2012), air quality (Landis et al., 2018; Miller et al., 48 2011; Rogers et al., 2020), ecosystems (Kou-Giesbrecht et al., 2019; Campos et al., 2019; Kallenborn et 49 al., 2012; Johnstone et al., 2010) and climate (Randerson et al., 2006). Not only can wildfire pollutants 50 fumigate local source areas, they can be transported over long distances resulting in degraded air quality 51 in locations far from fire sources (Miller et al., 2011; Rogers et al., 2020), and pose threats to downwind 52 ecosystems through wet and dry deposition processes (Kou-Giesbrecht et al., 2019; Kallenborn et al., 53 2012; Campos et al., 2019).

54 The severity and frequency of wildfires is expected to increase in response to climate change 55 (Bush and Lemmen, 2019; Seidl et al., 2017; Whitman et al., 2019) with evidence to suggest that such 56 impacts are expected to be most pronounced in the boreal biome (Seidl et al., 2017; Whitman et al., 57 2019). The boreal forest zone is the most northerly of all forest biomes accounting for 1.2 billion ha of mostly coniferous forest and comprising about 33 % of the global forest area, or 14 % of the earth's land 58 59 surface (https://www.nrcan.gc.ca/our-natural-resources/forests/sustainable-forest-management/boreal-60 forest/8-facts-about-canadas-boreal-forest/17394). On a global basis, boreal forest wildfires are 61 responsible for an estimated 20 % of yearly global biomass burning emissions (van der Werf et al., 2006). 62 Canada's boreal forests account for  $\sim 28$  % of the global boreal zone area and encompasses 75 % of 63 Canada's 347 million ha of forested land (Fig. S1) (https://www.nrcan.gc.ca/our-natural-64 resources/forests/sustainable-forest-management/boreal-forest/8-facts-about-canadas-boreal-forest/17394. 65 In the past decade, Canada has experienced unprecedented fire seasons, with large numbers of 66 evacuations, major property damage, poor air quality and significant economic impacts (NRCan, 2018;

67	Landis et al., 2018; McGee et al., 2015). Model predictions have suggested that Canadian fire
68	occurrences will increase by 25 % by 2030 from a 1975 to 1990 baseline scenario (Wotton et al., 2010).
69	To adequately assess and mitigate the risks of wildfire emissions to human and ecosystem health,
70	reliable pollutant predictions are required which depend on accurate and detailed fire emissions data.
71	Such emissions data are developed by multiplying emission factors with the mass of biomass burned
72	(Chen et al., 2019). In Canada, Environment and Climate Change Canada (ECCC) provides predictions
73	of particulate matter (PM) (<2.5 $\mu$ m in diameter) from wildfire smoke to the public using the FireWork
74	modelling system that combines forecast meteorology, emissions inputs (e.g. emission factors), forest fire
75	and fuel data (e.g. fuel maps, plume height parameterization), and a regional air quality model, GEM-
76	MACH (details in Chen et al., 2019). FireWork is also used for air quality research studies with
77	significantly more complex chemical mechanisms for emissions characterization and detailed physical
78	processes. Wildfire field studies, as well as prescribed burns and laboratory work, have resulted in
79	valuable global databases of fire emission factors covering a broad range of ecosystems and geographic
80	areas (e.g. Andreae, 2019; Akagi et al., 2011), however, they are primarily concentrated on the temperate
81	forests of the American mid-west and savannah/grasslands of Africa (e.g. Andreae 2019; Permar et al.,
82	2021; Palm et al., 2020; Lindaas et al., 2020; Roberts et al., 2020; Juncosa-Calaharrano et al., 2021;
83	Coggon et al., 2019; Koss et al., 2018; Hatch et al., 2017). Until now, the most complete characterization
84	of boreal forest wildfire emissions in Canada was provided by Simpson et al. (2011) which relied on
85	whole air canisters with offline analysis for organic compounds. Due to limited comprehensive emission
86	data specific for boreal wildfires, air quality models for northern regions face significant challenges
87	resulting in uncertain predictions of emissions, exposure and associated impacts.
88	In the summer 2018, a research aircraft was deployed to measure emissions and subsequent
89	transformation processes from a boreal forest wildfire in western Canada (Fig. 1; Fig. S1). In this paper,
90	measurements of a comprehensive suite of gas- and particle-phase compounds are used to provide a
91	detailed characterization of smoldering wildfire emissions. The highly speciated non-methane organic
92	gas (NMOG) measurements are described by broad chemical classes and across a range of volatilities

93 extending from VOCs to SVOCs. The wide range of measured NMOGs, along with concurrent total 94 NMOG carbon (NMOG<sub>T</sub>) measurements, provides a unique opportunity to reconcile the total carbon 95 budget. Emission factors are derived for 193 compounds which represents the most extensive chemical 96 speciation of wildfire emissions to date, almost tripling the number of reported values for the boreal forest 97 ecosystem in the Andreae (2019) compilation paper. Emission estimates are also combined with those 98 from satellite observations to evaluate modelled diurnal variability. The purpose of this work is to 99 provide relevant emissions information for boreal forest wildfires to ultimately contribute towards 100 improved emissions quantification and chemical speciation representations in air quality models.

101

#### 102 **2 Methods**

## 103 **2.1 Aircraft measurements**

104 The NRC's Convair-580 research aircraft was deployed on June 25, 2018 to sample a wildfire 105 detected to the east of the Alberta/Saskatchewan border (56.4°N, 109.7°W) (Fig. 1). Measurements of a 106 comprehensive suite of trace gases, particles and meteorology were made with high time resolution. 107 Meteorological measurements including relative humidity, temperature, wind direction and speed, as well 108 as aircraft state parameters such as altitude (masl) and geographic coordinates were conducted at 1 sec 109 time resolution. A detailed description of the various measurement methods with references is provided 110 in the supporting information (SI Sect. 1.1, Table S1, S2), with only a brief description provided here. 111 **2.1.1 Trace gas measurements** In-situ measurements of NO, NO<sub>2</sub>, NO<sub>y</sub>, O<sub>3</sub> and SO<sub>2</sub> were conducted 112 using commercial instruments (Thermo Scientific Inc.) modified to measure at 1 sec time resolution. 113 Ammonia (NH<sub>3</sub>) measurements were made at 1 sec time resolution using a Los Gatos Research (LGR) 114 NH<sub>3</sub>/H<sub>2</sub>S Analyzer, model 911-0039. Gas phase elemental Hg (GEM) was measured with a Tekran 115 237X instrument (Tekran Instruments Corporation) modified to allow a reduced sampling time of 2 min 116 (McLagan et al., 2021; Cole et al., 2014). CO, CO<sub>2</sub> and CH<sub>4</sub> were measured with a Cavity Ring Down 117 spectroscopy instrument (Picarro G2401-m). A second Picarro G2401-m instrument was used to measure 118 Total Carbon (TC, in units of ppm C) by passing the sample air through a catalyst to convert all carbon

119 species to  $CO_2$ . Total non-methane organic gases (NMOG<sub>T</sub>), in mixing ratios units of ppm C, was

- 120 quantified by subtracting the ambient  $CH_4$ , CO and  $CO_2$  measurements (instrument without the catalyst)
- 121

from the TC measurements (see SI Methods for more details).

122 Individually speciated NMOGs (as well as some inorganic species) were measured with a 123 Chemical Ionization Mass Spectrometer (CIMS), a Proton Transfer Time-of-Flight Mass Spectrometer 124 (PTRMS), and through whole air sampling using canisters (Advanced Whole Air Sampler; AWAS). In 125 addition, integrated cartridge-based samples were taken. The CIMS (a modified Tofwerk/Aerodyne Api-126 ToF) was operated using iodide as the reagent ion providing 1 sec time resolved measurements for 30 127 compounds (Table S2). The PTRMS (Ionicon Analytik GmbH, Austria) used chemical ionization with 128  $H_3O^+$  as the primary reagent ion providing 1 sec measurements for a suite of organic compounds. For 129 those compounds with no available gas standard, a relative response factor was calculated with reaction 130 rate constants using the method described in Sekimoto et al. (2017) and guided by the work of Koss et al. 131 (2018) ('calculated' compounds). Integrated 'grab' samples (20-30 sec) were collected from the aircraft 132 using the Advanced Whole Air Sampler (AWAS) with offline analysis. The AWAS provided speciated 133 measurements of hydrocarbons ( $\leq$ C10), but no oxygenates. Overlapping compounds/isomers that were 134 measured by both the PTRMS and AWAS, as well as between the PTRMS and CIMS, were handled as 135 described in SI Sect. 1.1.4. Integrated gas phase samples were collected using an automated adsorbent 136 tube (i.e. cartridge) sampling assembly with offline analysis (Ditto et al., 2021; Sheu et al., 2018; Khare et 137 al., 2019). These samples provided targeted measurements of gas-phase compounds ranging in volatility 138 from  $C_{10}$  volatile organic compounds (VOCs) to  $C_{25}$  semivolatile organic compounds (SVOCs) including 139 hydrocarbons (CH), and functionalized compounds containing 1 oxygen atom (CHO<sub>1</sub>), and 1 sulfur atom 140 (CHS<sub>1</sub>).

141 **2.1.2 Particle measurements** 

142 Particle chemistry was obtained with a high resolution aerosol mass spectrometer (AMS) 143 (Aerodyne) providing mass concentrations of particle species including total organics (OA), NO<sub>3</sub>, SO<sub>4</sub> 144 and NH<sub>4</sub> for particles less than  $\sim 1 \,\mu m$ . Particle size distributions were measured between 60 and 1000 145 nm at 1 sec time resolution using the Ultra High Sensitivity Aerosol Spectrometer (UHSAS; Droplet

146 Measurement Technologies). Refractory black carbon (rBC) was measured using a single particle soot

147 photometer (SP2; Droplet Measurement Technologies).

148

149 **2.2 Flight and fire description** 

150 A wildfire located near Lac La Loche in Saskatchewan (56.40°N 109.90°W) was detected by 151 satellite on June 23 (Fig. 1; Fig. S1). The fire was ignited by lightning on June 23, 2018 at 19:45 UTC 152 and lasted 50 hrs to June 25 21:41 UTC burning an estimated 10,000 ha before being extinguished by 153 rain. The area burned was mostly mature Jack pine and boreal spruce forest with a smaller fraction of 154 boreal mixed-wood forest. Satellite images from the VIIRS spectroradiometer on the Suomi NPP and 155 NOAA-20 satellites taken on June 25 showed merged fire hot spots with a visible smoke plume moving 156 in a north-westerly direction (Fig. 1; see SI Sect. 2.0 for more details). Lagrangian flight tracks were 157 flown downwind of the wildfire to follow the fire plumes. Multiple horizontal transects, vertically 158 stacked and perpendicular to the plume direction were made at different altitudes from 640 to 1460 m asl 159 (~220 – 1040 m agl, based on 420 m asl at Lac La Loche) forming virtual screens. Five screens were 160 completed over two flights with the closest screen ~10 km and the farthest screen 164 km downwind of 161 the fire, with the screens spaced such that the instruments sampled the same air parcels as they were 162 transported downwind. A vertical profile which typically reached ~2500 m asl was conducted in the 163 plume at each screen to gather information on its vertical structure and the height of the plume. As 164 demonstrated by the elevated CO mixing ratios in Fig. 2, two distinct plumes were identified - a south 165 plume (SP) and north plume (NP), that were transported in parallel in a northwesterly direction. The SP 166 is estimated to be ~42 min old based on the measured wind speed at Screen 1 and the distance from the 167 closest edge of the VIIRS fire hot spots (~10 km). The NP is estimated to be an additional 30 min older 168 than the SP (further details in SI Sect. 2.0). For the purposes of this investigation, only data from Screen 169 1 are used to characterize the direct emissions from this fire. There were no significant anthropogenic 170 sources like upwind urban or industrial areas, impacting the Screen 1 measurements. Plume evolution

during transport from Screen 1 to downwind Screens 2 to 5 is discussed in other papers (Ditto et al., 2021;
McLagan et al., 2021).

173

## 174 **2.3 Emission ratios**

175 Emission ratios (ERs) were calculated using an integration method (e.g. Yokelson et al., 2009; 176 Garofalo et al., 2019) using the in-plume measurements for the SP and NP. The integration method was 177 carried out for the real-time measurements by first subtracting a background from the in-plume 178 measurements. Background measurements were defined as the average over short time segments (~30 179 sec) outside and at the same altitude as inside the plume, and typically selected at the ends of the 180 horizontal transects. The background-subtracted plume measurements yielded enhanced plume values 181 (e.g.  $\Delta X(t)$ ) which were then integrated using the plume start and end times guided by when CO mixing 182 ratios were above the CO background. Nominal plume time periods are indicated by the vertical grey 183 bars in Fig. 3 which shows time series for CO, NMOG, OA and acetonitrile for the first 4 of 5 transects 184 on Screen 1. Integrated pollutant values were subsequently normalized by the integrated values of CO 185 (Eq. 1) to account for changes due to dilution producing emission ratios (ER) for the SP and NP for each 186 transect on Screen 1.

187

188 
$$ER = \frac{\int_{start}^{end} \Delta X(t) (dt)}{\int_{start}^{end} \Delta CO(t) (dt)}$$
(1)

189

190 CO is known to be a suitable dilution tracer as it has a long atmospheric lifetime of 1-4 months (Seinfeld 191 and Pandis, 1998), is unreactive on the time scale of the measurements. , In this study, ERs were 192 calculated using CO as it was enhanced above a background of ~0.119±0.005 ppmv for the plumes 193 measured, there were no other significant CO sources in the study area, and CO is a particularly good 194 tracer for smoldering fires (e.g. Simpson et al., 2011).

195 ERs for the AWAS compounds were determined using the average mixing ratio of 3 samples 196 taken in the SP and two in the NP, and the average mixing ratio of two background samples. CO mixing 197 ratios were averaged across the AWAS sample time period. For the integrated cartridges, samples were 198 collected over the lower set of aircraft transects ('LOW') and higher set of transects ('HIGH'), resulting 199 in two integrated cartridge samples for each screen. The HIGH sample was used as the background. The 200 HIGH sample was collected largely outside the wildfire plume, but may have been influenced to some 201 extent from emissions. However, this impact is expected to be minimal as average CO mixing ratios 202 during the HIGH sample were at background levels (~0.14 ppmv). Nevertheless, to address the potential 203 for influence of the plume in the HIGH sample, the ERs are presented as ranges with the lower estimates 204 derived by subtracting the HIGH background sample, and the upper estimates without subtracting the 205 HIGH sample. This calculation is described in Eq. 2 where  $Cartridge_{LOW}$  represents the LOW cartridge 206 sample measurements, *Cartridge<sub>BKGD</sub>* is the background derived from the HIGH cartridge sample 207 measurements, and COLOW, COBKGD are the average CO concentrations during the respective LOW and 208 HIGH cartridge integration time periods. The uncertainty with this bounding analysis is acknowledged, 209 but the I/SVOCs ERs within a plume are likely to vary similar to other work (Hatch et al., 2018).

210 
$$ER = \frac{Cartridge_{LOW} - Cartridge_{BKGD}}{CO_{LOW} - CO_{BKGD}} to \frac{Cartridge_{LOW} - 0}{CO_{LOW} - CO_{BKGD}}$$
(2)

211

# 212 **2.4 Emission factors**

Emission factors (EFs) were determined as the mass of species X emitted per unit mass of dry fuel burned in g kg<sup>-1</sup> assuming that all of the carbon in the fuel was released into the atmosphere and measured (Ward and Radke, 1993; Yokelson et al., 2007), and that the mass fraction of carbon in the fuel is constant. EFs were determined using Eq. 3 where  $F_c$  is the mass fraction of carbon in the fuel and estimated to be 0.5 (de Groot et al., 2009 and references therein), mm<sub>x</sub> is the molar mass of the compound of interest, and mm<sub>c</sub> is the molar mass of carbon, 12 g mol<sup>-1</sup>,  $\Delta X$  is the integrated background-subtracted TC. 220 TC (see Sect. 2.1) was directly measured and includes all the carbon mass in CO<sub>2</sub>, CO, CH<sub>4</sub>, and

221 NMOG<sub>T</sub>, as well as that from particulate black carbon (rBC) and particulate organic carbon (OC) (which

were added to the TC), for a complete accounting of all the emitted carbon. For species measured in mass concentration units, Eq. 3 was modified by converting TC to mass concentrations using the measured temperature and pressure, and removing the molar mass ratio term. The EFs for the AWAS and the cartridge samples were derived using the average measurements as discussed for the ER, but with TC as the denominator.

227

228 
$$EF\left(\frac{g}{kg}\right) = F_c \ x \ 1000 \ \left(\frac{g}{kg}\right) x \ \frac{mm_X}{mm_c} \ x \ \frac{\int_{start}^{end} \Delta X(t) \ (dt)}{\int_{start}^{end} \Delta TC(t) \ (dt)}$$
(3)

229

EFs were determined for the SP and NP for each transect, and then averaged to obtain screen-averagedEFs for the SP and the NP, as well as for both plumes together.

## 232 **2.5 Emissions Uncertainties**

233 There is the potential for inherent uncertainties using a plume integration method for calculating 234 EFs and ERs as the ratios derived this way represent the average plume composition and ignore the 235 spatial heterogeneity in wildfire plumes (Palm et al., 2021; Decker et al., 2021; Garofalo et al., 2019), 236 chemical transformation processes, and can also be affected by changing background levels. Pollutants 237 released by wildfires can be influenced by photochemical and physical changes that may take place 238 between the time of emission and the time of measurement, particularly for more reactive compounds 239 (e.g. Palm et al., 2021; Lindaas et al., 2020; Peng et al., 2020; Akagi et al., 2011). Although controlled 240 laboratory studies are well suited to examine direct emissions with minimal aging, they cannot reproduce 241 realistic burning conditions. Field measurements are critical to understand emissions that are impacted by 242 factors such as complex burning dynamics, fuel moisture, temperature and winds (Andreae 2019). 243 Recognizing the challenges of measuring primary emissions by aircraft, at 10 km (<1 hr) away from the

fire source, Screen 1 measurements represent some of the freshest emissions measured under wildfireconditions, thus providing best estimates of initial conditions.

246 Uncertainties in the EFs and ERs are estimated by summing in quadrature the standard error of 247 the average EF (or ER) and the propagated measurement uncertainties. The standard error is used as 248 description of the uncertainty on the average EF (and ER) characterizing repeated transects across the SP 249 and NP for a total of 20 min of in-plume sampling. The standard error is expected to at least partially 250 capture uncertainties associated with plume aging and vertical plume heterogeneity. As many compounds 251 exhibited significant in-plume enhancements above background levels, uncertainties in the integrated  $\Delta X$ , 252  $\Delta CO$  and  $\Delta TC$  values were assumed to be dominated by instrumental (measurement) uncertainties (Table 253 S1, S2). Emissions are not reported for compounds where the average mixing ratios were within  $1\sigma$  of 254 the background average. The low and high I/SVOCs EFs (and ERs) are provided as estimates of their 255 uncertainties (as described in Sect. 2.3). The derivation of AWAS and cartridge EFs (and ERs) may have 256 potential limitations as they rely on a limited number of samples, with the potential of the AWAS discrete 257 samples capturing only part of a plume.

## 258 **2.6 Combustion efficiency**

259 Combustion efficiency (CE) is a useful indicator of the relative proportion of flaming vs 260 smoldering stages of combustion which has a significant influence on the chemical composition of the 261 smoke (see SI Sect. 3.0 for further details). Flaming fires have CE > 0.90 (Yokelson et al., 1996) and 262 smoldering fires are typically ~0.8 with a range of 0.65 to 0.85 reported in the literature (Akagi et al., 263 2011; Yokelson et al., 2003). A modified combustion efficiency (MCE) is commonly calculated 264 assuming that  $CO_2+CO$  adequately represents all of the fuel carbon that has been volatilized and detected 265 in ambient air. Here, as the TC in the plume was directly measured,  $\Delta TC$  was used in Eq. 4 to improve 266 on the estimation of the CE by accounting for all the sources of carbon.  $\Delta CO_2$  and  $\Delta TC$  in Eq. 4 are the 267 integrated, background-subtracted mixing ratios.

268

$$269 \quad CE = \frac{\Delta CO_2}{\Delta TC}$$

#### **3 Results and Discussion**

# 272 **3.1 Fire combustion state**

273 The plume-averaged CE for the SP (transects 1 to 4) was  $0.84\pm0.04$  and for the NP (transects 1 to 274 3)  $0.82\pm0.01$ . Transect 4 was excluded from the calculations for the NP because only a portion of the 275 plume was detectable at this altitude (Fig. 3). The derived CE indicates that the fire was predominantly in 276 a low intensity smoldering phase which is consistent with the satellite-derived fire intensities during the 277 flight (see Fig. 10) and ground-based meteorological observations, and may reflect some residual 278 smoldering combustion (RSC). It is estimated that emissions from this fire were sampled 14 hrs post 279 flaming. Other chemical measurements from this flight also support that the fire was largely smoldering 280 including the detection of elevated  $C_2H_4O_2^+$  (levoglucosan fragment from the AMS), and no detectable 281  $K^+$  (from the AMS) (Lee et al., 2010). Significant spatial variability in the concentrations of many of the 282 measured species were observed closest to the fire source, while the plumes became more well-mixed as 283 they were transported downwind (Fig. S6). This highlights the complexities of assessing wildfire 284 combustion processes (Ward and Radke, 1993), and in particular, boreal forests have been observed to 285 exhibit greater variability in combustion efficiencies than for other vegetation types (Urbanski et al., 286 2009).

## 287 **3.2 General plume features**

Table A1 shows mixing ratios (or concentrations) and background levels of 193 pollutants that were enhanced in the fire plumes. The quantification of this suite of compounds provides new and additional emission estimates to those reported in Simpson et al. (2011) and compiled in Andreae (2019) for the boreal forest ecosystem. Several sulphur-containing compounds and a few other VOCs were not detected (Table S6), and although not part of the measurement suite in the present study, Simpson et al. (2011) did not observe emissions of anthropogenic halocarbons from wildfires in the same boreal forest

294 ecosystem. In Fig. 3, the in-plume portions are highlighted by the grev vertical bars and the SP and NP 295 are indicated as the aircraft flew at increasing altitudes to complete five horizontal transects. The lowest 4 296 transects showed enhanced pollutant levels while the 5<sup>th</sup> transect (not shown) was predominantly above 297 the height of the plumes. Higher concentrations were generally observed in the SP compared to the NP, 298 possibly because of some plume dilution in the NP resulting from a change in wind direction prior to 299 sampling. The SP and NP were distinctly separated from each other, with pollutants typically dropping to 300 background levels between the plumes. NMOG<sub>T</sub> mixing ratios varied between background levels of ~375 301 ppbv to near 10 ppmv in-plume. CO and acetonitrile, often used as tracers of biomass burning (e.g. 302 Wiggins et al., 2021; Landis et al., 2018; Simpson et al., 2011; de Gouw et al., 2006), reached 6.6 ppmv 303 and 20 ppby, respectively in the SP, while maximum OA concentrations reached 276 µg m<sup>-3</sup>, above a 304 background level of  $\sim 12.5 \pm 0.83 \ \mu g \ m^{-3}$ . OA was the largest contributor to particulate mass (PM) 305 comprising over 90 % of the measured submicron mass with remaining portion comprised of BC, NO<sub>3</sub>, 306 NH<sub>4</sub>, and SO<sub>4</sub> (Fig. S6). Integrated filter samples taken from the aircraft across Screen 1 also showed the 307 presence of a diverse set of functionalized particle-phase organic compounds (Ditto et al., 2021). 308 The most abundant reactive nitrogen compounds ( $N_r$ ) were in the forms of reduced nitrogen (79%) 309 with NH<sub>3</sub> comprising 42 % of  $\Sigma N_r$  (Fig. 4) and substantially lower nitrogen oxides i.e. NO<sub>x</sub> < 1 ppbv. 310 Dominant proportions of reduced nitrogen in biomass burning emissions were also reported previously 311 (Lindaas et al., 2020; Burling et al., 2011; Yokelson et al., 1996). Nitrogen-containing organics were 312 detected in the present study totalling 3.9 ppbv and 18 % of  $\Sigma N_r$  (Fig. 4), however, other such compounds 313 that were not included with the instrument suite used in this study were also likely emitted. Such 314 compounds could include organic nitrates, amines, amides, heterocyclic compounds, nitriles and nitro 315 compounds that have been found in biomass burning emissions (Roberts et al., 2020; Lindaas et al., 2020; 316 Andreae 2019; Koss et al., 2018; Tomaz et al., 2018; Stockwell et al., 2015). Alkyl nitrates have been 317 identified in biomass burning emissions, but their contributions to total Nr appeared to be small (Juncosa-318 Calahorrano et al., 2021; Roberts et al., 2020; Lindaas et al., 2020; Simpson et al., 2011; Alvarado et al., 319 2010; Singh et al., 2010).

#### **320 3.3 Total carbon budget**

#### 321 3.3.1 NMOG chemical classes – PTRMS, CIMS, AWAS

322 In-plume mixing ratios and the relative contribution of individually measured NMOG species to 323 the sum of those species ( $\Sigma$ NMOG) are shown for 13 chemical classes in Fig. 5. (See Fig. S7 for separate 324 SP and NP chemical classes). The largest chemical classes include carbonyls (acids, aldehydes and 325 ketones), alcohols, hydrocarbons (alkanes, alkenes, alkynes), aromatics (including furans, phenol, 326 benzene and toluene), and nitriles. Hydrocarbons (i.e. C<sub>x</sub>H<sub>y</sub>, including some aromatics) were responsible 327 for just over half of the  $\Sigma$ NMOG (53 %) (Fig. S8), with 29 % identified as alkenes such as ethene, 328 propadiene, and propene, 19 % alkanes, predominantly ethane, and 3 % alkynes, almost entirely 329 acetylene. Non-aromatic oxygenates accounted for an additional 36 % of the  $\Sigma$ NMOG with roughly 330 equal contributions (9 to 12%) from acids, aldehydes and alcohols, and a smaller fraction from ketones (5 331 %). Including other oxygenated compounds such as furanoids and phenol/phenol derivatives, all 332 oxygenates ( $C_xH_yO_z$ ) comprised 42 % (Fig. S8), of the  $\Sigma$ NMOG.

333 A similar range of compound classes has been observed in previous field and laboratory studies, 334 noting that the measured compound suite between studies varies to some extent. For example, Simpson 335 et al. (2011) found a similar distribution of compound classes with 57 measured NMOG species, based on 336 discrete canister samples, in boreal forest wildfires. In that study, oxygenates (non-aromatic) comprised a 337 smaller portion of NMOG (29%) as major emitted species like acetaldehyde and acetic acid (Fig. 8) were 338 not included. Other studies have also found oxygenates to be a large portion of NMOG emissions across 339 multiple fuel types, including those similar to the current study, ranging from 51 - 68 % (Permar et al., 340 2021; Koss et al., 2018; Gilman et al., 2015; Akagi et al., 2011) with a range of 25 – 55 % reported in 341 Hatch et al. (2017). The fraction of NMOG oxygenates in the present study (42 %) was closer to those 342 reported in Hatch et al. (2017) when only the most relevant fuel types of pine and spruce were considered 343 (55 % and 43 %, respectively). Similar to previous work (Koss et al., 2018, Stockwell et al., 2015; Hatch 344 et al., 2015), emissions of substituted oxygenates like furanoids (furans+derivatives) and phenolic 345 compounds were observed. Furanoids contributed 4 % of the  $\Sigma$ NMOG mostly due to furfural, furan and

methyl furan while phenolic compounds eg. guaiacol, methyl guaiacol, contributed 0.5 % of the ΣNMOG
(Fig. S9). Although these emissions were less abundant in the present study, they represent important OH
reactants (Coggon et al., 2019; Koss et al., 2018; Gilman et al., 2015) with phenols being implicated as
precursors to brown carbon formation in secondary organic aerosol (SOA) (Palm et al., 2020).

Biogenic emissions of terpenoids including isoprene, monoterpenes, carvone, sesquiterpenes,
camphor/isomers and terpine-4-ol/cineole/isomers were elevated in the plumes collectively reaching ~2.4
ppbv, and contributing ~2 % to the ΣNMOGs (Fig. S9). Isoprene was ~66 % of these compounds with an
additional 32 % from monoterpenes. Emissions of isoprene from biomass burning has been observed
from a wide range of fuel types (Hatch et al., 2019). As isoprene is not stored by plants and the
measurements were taken ~14 hrs post flaming, it was likely emitted as a combustion product.

356 In this study, furfural was the most abundant oxygenated aromatic compound and a factor of 5 357 times higher than that of phenol. Although Koss et al. (2018) found that phenol and furfural emissions 358 were similar for most fuels tested in the laboratory, furfural emissions derived from multiple wildfires 359 sampled in Permar et al. (2021) were similar to those in the present study, and a factor of 1.6 higher for 360 phenol. As phenol emissions are associated with lignin pyrolysis (Stockwell et al., 2015; Simoneit et al., 361 1999), the lower emissions in the current study could be because the lignin content in the fuel mixture 362 was lower than fuels used in previous laboratory studies or that most of the phenolic compounds were 363 emitted during the earlier phases of the fire. Several modelling studies have indicated that aromatics and 364 terpenes are insufficient to explain SOA formation in biomass burning plumes (e.g. Hodshire et al., 2019) 365 suggesting the importance of inclusion of other aromatic species such as phenolics and furanoid 366 compounds. However, models typically do not include reactions involving phenolic and furanoids 367 species, especially substituted compounds like furfural, guaiacol, and methyl guaiacol. Box model 368 simulations have also shown that incorporation of OH oxidation of furan, 2-methyfuran, 2,5-369 dimethylfuran, furfural, 5-methylfurfural, and guaiacol, leads to 10 % more O<sub>3</sub> formed (Coggon et al., 370 2019).

## 371 **3.3.2** Intermediate-volatility and semivolatile organic compounds (I/SVOCs)

372 Offline analysis of cartridge samples showed a wider range of hydrocarbons and functionalized 373 gas-phase organic compounds not observed in the PTRMS, CIMS, and AWAS measurements, including 374 I/SVOC compounds in the wildfire plume. ERs (Table S7) for species containing carbon and hydrogen, 375 and with either sulfur or oxygen (i.e. CH (hydrocarbons),  $CHS_1$  and  $CHO_1$  type molecules)) accounted for 376 a sizeable fraction of carbon in the  $C_{10}$  to  $C_{25}$  range. Additional contributions are expected from more 377 highly functionalized organics in the gas (and particle) phase not reflected in the CH, CHO<sub>1</sub>, and CHS<sub>1</sub> 378 compound classes (e.g., gas-phase species with multiple oxygen atoms like vanillic acid or 379 acetovanillone, and gas-phase species containing combinations of oxygen and nitrogen atoms (CHON) 380 (Ditto et al., 2021; 2022)). ERs in the plume varied across the carbon number range; in general, the 381 highest ratios were observed for the complex mixture of hydrocarbons (i.e. CH compounds) broadly 382 peaking at  $C_{20}$ - $C_{23}$  in the SVOC range, with a larger contribution from  $C_{10}$  compounds including 383 monoterpenes. By comparison, the complex mixture of  $CHO_1$  compounds was slightly lower in 384 abundance than CH with contributions from  $C_{10}$  monoterpenoid emissions or monoterpene oxidation 385 products. CHS<sub>1</sub> IVOC-SVOCs were the lowest abundance species quantified. CHN<sub>1</sub> compounds 386 represent another observed contributor of IVOCs-SVOCs; the sum of all CHN<sub>1</sub> compound ion 387 abundances was two orders of magnitude smaller than the sum of all  $CHO_1$  species. We note that for 388  $CHN_1$ , this qualitative comparison is in terms of ion abundances only, given a lack of appropriate 389 standards to calibrate for the mass spectrometer's response to the complex mixture of reduced nitrogen-390 containing I/SVOCs.

EFs were estimated to be  $1.4\pm0.037 - 2.4\pm0.063$  g kg<sup>-1</sup> for CH,  $0.81\pm0.078 - 0.81\pm0.079$  g kg<sup>-1</sup> for CHO<sub>1</sub>, and  $0.21\pm0.0033 - 0.22\pm0.0060$  g kg<sup>-1</sup> for CHS<sub>1</sub> species, for a total EF of  $2.4\pm0.12 - 3.5\pm0.15$ g kg<sup>-1</sup> (Table A1). Here, the uncertainty represents measurement uncertainty associated with the conversion from signal to mass, and the reported ranges show lower and upper limit EF values that account for a contaminated background and that assume no background concentrations, respectively (as described above). These estimates accounted for C<sub>11</sub>-C<sub>25</sub> species and focused on I/SVOCs to avoid double counting the monoterpenes and C<sub>10</sub> monoterpenoid species, as they were already accounted for in

398 the PTRMS data. It is noted that the concentrations estimated for the cartridge samples may be sensitive 399 to variations in sampling efficiency within the under-wing sampling pod across  $C_{10}$ - $C_{25}$  though these 400 effects are expected to be minimal for the adsorbent tubes used in this study (Ditto et al., 2021; Sheu et al. 401 2018). These emission estimates expanded the characterized spectrum of organic species to include 402 IVOC/SVOCs in boreal forest fire emissions, which until now, had only been available from laboratory 403 measurements (Hatch et al., 2018). However, the observed emissions of the complex mixture of 404 hydrocarbons and functionalized species may include contributions from the re-volatilization of 405 compounds previously emitted from upwind oil sands operations and deposited in the forest ecosystem, as 406 noted in Ditto et al. (2021).

407

# 408 **3.3.3 Accounting for the observed carbon**

409 Measurements of TC, along with the speciated measurements from the PTRMS, CIMS, AWAS 410 and cartridges, provided a unique opportunity to reconcile the TC budget in a wildfire. Fig. 6 shows the 411 TC partitioning based on derived EFs (Sect. 3.5); overlapping compounds from the individual 412 measurement methods were handled as described in SI Sect. 1.1.4. The total EF for all carbon-containing 413 compounds was 1652 g kg<sup>-1</sup> and, as expected, CO<sub>2</sub> was the dominant contributor comprising >90 % of 414 TC. CO contributed 7.0 % followed by a contribution from NMOG<sub>T</sub> of 1.9 % with even smaller 415 contributions observed from CH<sub>4</sub> (0.5 %) followed by OC and BC (not shown) at <0.5 %. The two 416 magnified pie charts (right side), representing the low and high I/SVOC EF estimates, show the percent 417 breakdown of the measured NMOGs, and the remaining unidentified portion of NMOG<sub>T</sub>. The EF values 418 (g C kg<sup>-1</sup>) are identified in the box below. The  $\Sigma$ NMOG EFs (for PTRMS+CIMS+AWAS measurements), totalling 13.6±0.9 g C kg<sup>-1</sup>, accounted for 43±3 % of the NMOG<sub>T</sub> EF of 31.2±4.7 g C kg<sup>-1</sup> (refer to Fig. 419 420 S10 for the individual SP and NP breakdowns). The  $\Sigma$ NMOG uncertainties were estimated by summing 421 in quadrature the individual compound EF uncertainties for the SP and NP separately, with these 422 uncertainties subsequently summed in quadrature to derive the average  $\Sigma$ NMOG uncertainty (Fig. 6). 423 The cartridge data showed the presence of a range of larger molecular weight I/SVOC compounds

424 between C<sub>10</sub> and C<sub>25</sub> with an additional 2.1 to 3.0 g C kg<sup>-1</sup> representing 7 to 10 % of NMOG<sub>T</sub>. Together, 425 all of the speciated NMOG measurements in this study accounted for  $50\pm3$  % to  $53\pm3$  % of NMOG<sub>T</sub>. The 426 remaining carbon mass was unidentified comprising  $47\pm15$  % to  $50\pm15$  % of NMOG<sub>T</sub>. Despite using 427 four state-of-the-art measurement techniques resulting in an extensive measurement suite, almost half of 428 NMOG<sub>T</sub> remained unidentified. This is consistent with previous work estimating  $\sim$ 50 % of NMOG<sub>T</sub> by 429 mass as unidentified (Akagi et al., 2011). It is noted, however, that the magnitude of the unidentified 430 portion is partly affected by uncertainties in the speciated measurements. For example, many of the 431 'calculated' PTRMS compounds are uncertain by an estimated factor of  $\sim 2$  (SI Sect. 1.1.1, Table S1). 432 Nevertheless, a portion of the unidentified species likely consisted of challenging-to-measure-VOCs and 433 larger I/SVOCs that were highly functionalized or contained molecular features like reduced nitrogen 434 groups (e.g. amines) that have been observed in the gas and particle phase at various sites (Ditto et al., 435 2020; Ditto et al., 2022). The presence of I/SVOCs in biomass burning emissions has been previously 436 observed in laboratory experiments (e.g. Koss et al., 2018; Hatch et al., 2018; Hatch et al., 2017; Bruns et 437 al., 2016) with smoldering more likely to emit a higher fraction of compounds with low volatility than 438 higher temperature processes (Koss et al., 2018). The unidentified portion may also have been comprised 439 of nitrogen-containing organics (Sect. 3.1). Studies that included measurements of a larger range of 440 nitrogen-containing organics in biomass burning emissions estimated that they comprised < 5-6 % of the 441 total nitrogen budget (Lindaas et al. 2020; Gilman et al., 2015), and thus, an even smaller fraction of 442  $NMOG_{T}$ . Advancing analytical techniques to expand the suite of NMOG speciation will enable further 443 reconciliation of the TC budget which is important for assessing secondary formation processes in the 444 atmosphere.

445 **3.3.4 Volatility distribution of NMOG** 

446 Volatility distributions can help track the full range of organic species to assess their partitioning 447 between the condensed and gas phases (Donahue et al., 2011). Fig. 7 shows the fractional sum of all 448 NMOG EFs within each volatility bin in terms of saturation concentration ranges ( $log_{10}C_o$ ,  $\mu g m^{-3}$ ) for the 449 low I/SVOC EF estimate.  $C_o$  values were estimated using the parameterization developed by Li et al.

450 (2016). NMOG emissions from this fire spanned a large range of volatilities from  $\log_{10}C_0$  of -2 to 10 µg 451 m<sup>-3</sup> across SVOC to VOC categories. The bin-averaged O/C ratio based on the measurements increased 452 with reduced volatility reflecting the presence of compounds with additional oxygen-containing functional groups. The highest fraction of emissions was present as VOCs with 81 % having  $log_{10}C_0 > 6$ 453 454  $\mu$ g m<sup>-3</sup>, 9 % as IVOCs having 4 < log<sub>10</sub>C<sub>0</sub>  $\mu$ g m<sup>-3</sup> < 6  $\mu$ g m<sup>-3</sup> and 10 % as SVOCs having log<sub>10</sub>C<sub>0</sub> < 3  $\mu$ g m<sup>-3</sup> 455 <sup>3</sup>. These results align with laboratory studies showing that oxygenates comprised more than > 75 % of 456 IVOCs across a range of biomass types with IVOCs accounting for  $\sim 11$  % of the  $\Sigma$ NMOG (Hatch et al.; 457 2018). Fig. 7 encompasses the range of volatilities based on all the identified NMOGs in this study that is 458 expected to represent initial emission conditions for modelling downwind chemistry. However, improved 459 speciation, particularly of lower volatility compounds, is needed to further expand the range of volatilities 460 and advance knowledge in gas to particle partitioning processes.

461

## 462 **3.4 Emission factors and comparisons with other studies**

463 EFs (and ERs) in this study are derived for 193 compounds from 15 instruments of which 173 are 464 NMOG species (Table A1). This dataset represents the most extensive range of field-based EFs ever 465 determined for a wildfire in the boreal forest ecosystem. In Fig. 8 average EFs are shown for compounds 466 grouped by a) particles, b) gas-phase inorganics, and c) gas-phase organics. Separate EFs and ERs for the 467 SP and NP are shown in the SI (Figs. S11 to S13). In Fig. 9a-c, EFs are compared with those from other 468 relevant studies. Fig. 9a shows a comparison with boreal forest field measurements largely taken from a 469 compilation by Andreae (2019) referred to as BFF19, as well as values from Akagi et al. (2011) and Liu 470 et al. (2017). This results in a comparison for 50 compounds (35 organics and 15 inorganics/particulate 471 species) with the largest suite of EFs from one study conducted in a similar boreal region as the present 472 study (Simpson et al., 2011). EFs are also compared with laboratory-derived EFs for lodgepole pine Koss 473 et al. (2018; referred to as LAB18) (Fig. 9b), a similar fuel type in the current study, with a total of 99 474 NMOGs and 3 inorganics in common. In Fig. 9c, EFs are compared with those recently reported in 475 Permar et al. (2021) (referred to as TFF21) based on aircraft measurements of temperate forest wildfires

in areas mostly dominated by pine, fir and spruce trees, which provides the closest suitable comparison
with similar speciated NMOGs under wildfire conditions. Comparisons include 111 NMOGs, and 4
inorganics/black carbon. While the Permar et al. (2021) study was conducted in a temperate forest
region, it was at high elevation locations with similar vegetation types as the current study.

480

481 **3.4.1 Particle species** The PM<sub>1</sub> EF (6.8 $\pm$ 0.8 g kg<sup>-1</sup>) represents the total of all particle component species as measured by the AMS. The PM<sub>1</sub> EF of  $6.8\pm1.1$  g kg<sup>-1</sup> (Fig. 8a) (accounting for estimated mass 482 483 differences due to particle diameters (SI Sect. 1.1.2)) falls in the lower end of the large range previously 484 observed for boreal forest wildfires (18.7±15.9 g kg<sup>-1</sup>; Fig. 9b). The few PM EFs for BFF19 (n=5) over a 485 limited range of MCEs (i.e. 0.89 to 0.93) shows significant variability consistent with previous work 486 (Jolleys et al., 2015; Akagi et al., 2011; Cubison et al., 2011; Hosseini et al., 2013). OA, accounting for 487 90 % of PM<sub>1</sub>, has the largest EF, with comparatively lower EFs for NO<sub>3</sub>, BC, NH<sub>4</sub>, and SO<sub>4</sub> (Fig. 8a, Fig. 488 S6). This reflects the dominant particle-phase organic carbon content of the burned fuel and 489 correspondingly lower fractions of nitrogen and sulphur-containing compounds. Similar high organic 490 fractions have been previously observed in biomass burning emissions (Liu et al., 2017; May et al., 2014; 491 Hecobian et al., 2011). ERs similarly highlight the dominant OA emissions. Although the magnitude of 492 EFs between the SP and NP are within their derived uncertainties, the ERs showed differences by up to 493 70 % for NH<sub>4</sub> (Fig. S12) suggesting some differences in photochemistry between the two plumes. EFs 494 and ERs for chemically-speciated particle species derived in this study represent the first such 495 measurements under boreal forest wildfire conditions. EFs for chemically speciated compounds are not 496 found in BFF19 (except BC), but when compared with available values for U.S. temperate forest wildfires 497 (Liu et al., 2017) are found to be lower for OA (Fig. 9a), SO<sub>4</sub>, NO<sub>3</sub> and NH<sub>4</sub> by factors of 3.7, 5.0, 5.3, 498 and 3.0, respectively. Although differences in fuel type burned between the present study (mature Jack 499 pine, boreal spruce, boreal mixed-wood) and Liu et al. (2017) (mixed conifer, grass, brush and chaparral) 500 may influence the chemical composition of emissions, these large differences suggest the importance of 501 other factors in controlling OA emissions. The lower OA emissions under smoldering conditions in the

502 current study compared to Liu et al. (2017) with higher combustion efficiencies (0.877 to 0.935) conflicts 503 with some findings showing increased OA emissions with lower fire intensities (Liu et al., 2017, Burling 504 et al., 2011). However, the relationship between  $EF_{OA}$  and combustion efficiency can be impacted by 505 multiple factors such as OA loading, gas-particle partitioning related to dilution, and fuel moisture content 506 (May et al., 2014). The EF<sub>OA</sub> in the current study ( $6.6\pm 2.6$  g kg<sup>-1</sup>) lies in the range of EF<sub>OA</sub> reported for 507 prescribed burns across three temperate ecosystems  $(2.8\pm1.6 \text{ to } 11.2\pm2.7 \text{ g kg}^{-1})$  (May et al., 2014). This 508 may imply that the low intensity, surface, smoldering wildfire conditions in the present study (Sect. 3.1) 509 may be similar to prescribed burn conditions which are typically low intensity fires that are restricted to 510 the forest floor and understory, and conducted under controlled and consistent meteorological and fuel 511 moisture conditions (Yokelson et al., 2013; Carter and Foster, 2004). Inorganic PM emissions, however, 512 are likely more dependent on fuel elemental composition than combustion efficiency (Liu et al., 2017). 513 Differences in fuel composition between boreal and temperate forest ecosystems are inferred through 514 comparisons of NO<sub>x</sub> and SO<sub>2</sub> emissions. For example, the average NO<sub>x</sub> and SO<sub>2</sub> EFs for boreal forests, 515 are lower than the average EFs for temperate forests by factors of 2.5 and 3.0, respectively. The lower 516  $NO_x$  and  $SO_2$  emissions from boreal vs temperate forest wildfires are likely reflective of the reduced S 517 and N content in boreal biomass (Bond-Lamberty et al., 2006) relative to conifer (Misel, 2012) fuels in 518 the western U.S., as well as the possible influence of lower anthropogenic sources of nitrogen and sulphur 519 atmospheric deposition in boreal forests (Jia et al., 2016).

520

**3.4.2 Gas-phase inorganic species** The largest average EFs for inorganic gases (Fig. 8b; separate NP and SP Fig. S11) were from reduced nitrogen compounds dominated by NH<sub>3</sub> ( $0.63\pm0.14$  g kg<sup>-1</sup>) and followed by HCN ( $0.31\pm0.07$  g kg<sup>-1</sup>), with lower EFs for oxidized nitrogen compounds such as NO<sub>2</sub> ( $0.15\pm0.04$  g kg<sup>-1</sup>) and HONO ( $0.02\pm0.012$  g kg<sup>-1</sup>). This is consistent with previous work identifying elevated emissions of NH<sub>3</sub> and HCN during smoldering conditions, whereas emissions of HONO and NO<sub>x</sub> are primarily associated with flaming combustion (e.g. Roberts et al., 2020; Akagi et al., 2013; Yokelson et al., 1997; Griffith et al., 1991). The EFs for CO<sub>2</sub> and CO from the present study are

528 comparable within uncertainties of that previously reported for BFF19 (Table A1). However, EFs for 529 most other gaseous inorganic species were lower than the BFF19 EF average including NH<sub>3</sub>HONO, and 530  $NO_x$  by factors of 4.0, 20, and 7.1, respectively (Fig. 9a). There are only a limited number of studies 531 reporting EFs for these compounds in the BFF19 category. For example, there are only 4 previously 532 reported BFF19 EFs for NH<sub>3</sub> (2.5±1.8 g kg<sup>-1</sup>) showing a large range of values. Although these 533 comparisons are limited by the few reported values in the literature, the differences indicate a strong 534 sensitivity towards factors like fire intensity, chemical reactivity, fuel type and moisture, and 535 meteorology. In contrast, EFs for HCN derived in the current study (0.31±0.07 g kg<sup>-1</sup>) lie within the 536 range of BFF19, LAB18 and TFF21 values  $(0.28\pm0.06 \text{ to } 0.53\pm0.30 \text{ g kg}^{-1})$ , (Figs 9a, b, c, respectively) 537 and does not vary widely suggesting that HCN may be less sensitive to burning characteristics. HCN is 538 of concern due to its impacts on human health particularly since biomass burning emissions are 539 responsible for the majority of the global HCN (Moussa et al., 2016 and references therein).

540

541 3.4.3 Gas-phase organic species In Fig. 8c, the top 25 average EFs for gas-phase organic species are 542 shown in decreasing order of magnitude. The most abundant emissions were from the lower molecular 543 weight compounds; such trends are generally in agreement with previous field-based measurements for a 544 range of fuel types (e.g. Permar et al., 2021; Andreae, 2019; Liu et al., 2017; Simpson et al., 2011; 545 Urbanski et al., 2009). Excluding CH<sub>4</sub>, the largest EFs were associated with methanol, followed by 546 ethene, ethane, acetic acid, acetaldehyde, formaldehyde, and acetone ranging from  $1.9\pm0.4$  g kg<sup>-1</sup> to 547  $0.82\pm0.22$  g kg<sup>-1</sup> for these compounds. Noting some variations related to differences in measurement 548 methods, other studies have identified many of these same species as dominating biomass burning 549 emissions (e.g. Permar et al., 2021; Simpson et al., 2011; Akagi et al., 2011). For example, Simpson et al. 550 (2011) found that 5 of the same compounds in the present study including formaldehyde, methanol, 551 ethene, ethane and acetone were in the top 10 NMOG EFs from aircraft-based measurements made of 552 boreal forest wildfires in northern Saskatchewan, Canada, and within ~300 km of the current study. In 553 the present study, the top 24 NMOG compounds accounted for 81 % of the  $\Sigma$ NMOG by total molecular

554 mass with lower emissions from the remaining 149 measured compounds. In western U.S. wildfires, 555 small emissions from 151 species were found to account for almost half of  $\Sigma$ NMOG (Permar et al., 556 2021). EFs for the NP and SP generally agreed within their uncertainties with larger differences for some 557 of the more reactive species like isoprene, monoterpenes, and furan. For example, the SP EF for isoprene 558 was a factor of 3.4 lower than that for the NP ( $0.64\pm0.34$  g kg<sup>-1</sup>) (Fig. S13). Although the reasons for 559 these differences are not yet known, observations of higher  $O_3$  in the SP (52.4 $\pm$ 3.0 ppbv) compared to the 560 NP (44.7 $\pm$ 3.6 ppbv) suggest the influence of higher oxidant chemistry in the SP emissions compared to 561 the NP.

562 To compare the total NMOG derived in the present study with those from previous studies that 563 typically sum up their speciated measurements i.e.  $\Sigma NMOG$ , estimates were made using two methods: 1. 564 increasing the  $\Sigma$ NMOG to account for the unidentified portion of NMOG<sub>T</sub>; and 2. adjusting the NMOG<sub>T</sub> 565 to reflect the total molecular mass (not just the carbon portion). For method 1, the  $\Sigma NMOG EF$ (including the I/SVOCs) in this study (24.5±1.6 to 25.6±1.6 g kg<sup>-1</sup>) was increased by 50 and 47 % (Fig. 566 567 6), respectively, equalling  $36.8\pm11.3$  to  $37.6\pm12.2$  g kg<sup>-1</sup>. This estimate assumes that the carbon 568 distribution is the same as the identified, speciated measurements. For method 2, based on the speciated 569 measurements, the average molecular mass was 100 g mol<sup>-1</sup> and the average carbon number was 6 570 resulting in  $\sim 28\pm 24$  % of the molecular fraction represented by atoms other than carbon. Adjusting the 571 NMOG<sub>T</sub> of 31.2±3.8 g C kg<sup>-1</sup> upwards by 28±24 % to reflect the additional molecular mass results in a 572 NMOG of 39.9±5.8 g kg<sup>-1</sup>. The resulting estimated NMOG in this study of 36.8±11.3 to 39.9±5.8 g kg<sup>-1</sup> 573 lies between the estimated average of 58.7 g kg<sup>-1</sup> for the BFF19 (Fig. 9a) and those estimated from the 574 ΣNMOG EFs of 25.0 g kg<sup>-1</sup> (LAB18) (Fig. 9b), and 26.1±6.9 g kg<sup>-1</sup> (TFF21) (Fig. 9c) derived from 575 laboratory- and field-based studies (Table A1). In contrast to the current work, previous estimates of 576  $NMOG_T$  are likely to underestimate total NMOG emissions as they typically represent the sum of 577 measured species only. Some studies have attempted to account for NMOG<sub>T</sub> by including the sum of 578 measured plus estimates of 'unknown' portions of NMOGs ( $\Sigma$ NMOGs) (Permar et al., 2021; Koss et al., 579 2018; Stockwell et al., 2015; Gilman et al., 2015). The BFF19 EF was recently doubled from  $29.3 \pm 10.1$  g  $kg^{-1}$  to 58.7 g kg<sup>-1</sup> to account for unidentified NMOGs where the  $\Sigma$ NMOGs were measured by FTIR, GC and PTRMS (Andreae, 2019; Akagi et al., 2011). These results support that doubling the  $\Sigma$ NMOG provides a reasonable estimate the NMOG<sub>T</sub>. It is noted, however, that the average BFF19 NMOG EF is ~1.5 times higher than that derived in the present study, however, this may reflect variability in NMOG emissions even within the same boreal biome.

585

586 Although it is known that acidic compounds are emitted from biomass burning, few studies have 587 quantified their emissions, particularly under field conditions (Andreae, 2019; Veres et al., 2010; 588 Yokelson et al., 2009; Goode et al.; 2000). In this study, EFs for 22 organic acidic compounds were 589 derived (Table A1) representing the most detailed set of organic acid EFs from biomass burning for any 590 ecosystem (Andreae, 2019). The largest EFs for these compounds include acetic acid, C4 oxo-carboxylic 591 acids, and pyruvic acid, all of which are found among the top 24 NMOGs (Fig. 8c). For those 592 measurements that are available for comparison, EFs in the present study were lower for formic acid and 593 acetic acid compared toBFF19, and were also lower than in LAB18, and TFF21, ranging from factors of 594 1.7 to 8.8 (Figs. 9a, b, c). Organic acids that were in common with TFF21 and LAB18 (Table A1) had 595 lower EFs, with the exception of pyruvic acid, which was substantially higher (> factor of 37) in the 596 present study. Differences in fuel type may be an important factor in the variability of these comparisons. 597 Based on laboratory experiments, Veres et al. (2010) found a large range (factor of 5 to 13) of organic 598 acid emissions with different fuel types suggesting that the lignin content of the fuel could be a source of 599 biomass burning organic acid emissions. Emissions for 10 organic acids that have not previously been 600 reported, as well as several inorganic acids including nitrous acid, isocyanic acid, and peroxynitric acid, 601 are included in Table A1. These acids, representing 9 % of the  $\Sigma$ NMOGs (Fig. 5), are an important class 602 of oxygenates as they can form additional PM (Reid et al., 2005) and influence the hygroscopicity of 603 smoke particles (Rogers et al., 1991; Kotchenruther and Hobbs, 1998). 604 Isoprene and monoterpenes, with similar EFs  $\sim 0.41\pm0.19$  g kg<sup>-1</sup>, represented 16<sup>th</sup> and 18<sup>th</sup>,

respectively, of the top 24 NMOG EFs in this study. Terpenes are known to be emitted from a range of

606 biomass burning fuels (Andreae, 2019 and references therein), but there have been few measurements in 607 boreal forest wildfire plumes (Simpson et al., 2011; Andreae, 2019). It is noted that PTRMS 608 measurements of IVOCs like sesquiterpenes likely represent lower limits as they tend to be easily lost to 609 sample inlet lines due to their low volatility. The isoprene average EF of  $0.42\pm0.26$  g kg<sup>-1</sup> was more than 610 a factor of 5 higher, while the monoterpenes EF, 0.41±0.19 g kg<sup>-1</sup>, was substantially lower than the only 611 reported EF for boreal forest wildfires (Simpson et al., 2011). The difference in EFs for isoprene would 612 be even greater if only the NP EF  $(0.64\pm0.34 \text{ g kg}^{-1})$  is compared (if it is assumed that isoprene emissions 613 were influenced by photochemical losses in the SP). As the present study and the Simpson et al. (2011) 614 study were conducted in similar locations (i.e. boreal forest region within ~300 km of each other), and 615 comparable background levels, these differences are likely driven by fire combustion state, despite having 616 similar study-averaged MCEs. The majority of monoterpenes are stored in plant tissues (resin stores) for 617 long periods of time, but isoprene is synthesized and immediately released by plants, and can also be 618 emitted as a combustion product (Ciccioli et al., 2014; Akagi et al., 2013). Hatch et al. (2019) found that 619 a wide range of terpenoids are released across a variety of biomass types with variable emissions that 620 were dependent on plant species, and specifically related to their fuel resin stores. In the present study, 621 monoterpenes may have 'boiled-off' through distillation processes in the early stages of the fire resulting 622 in lower monoterpenes emissions at the aircraft sampling time,  $\sim 14$  hrs post-flaming. In contrast, the 623 Simpson et al. (2011) study sampled comparatively earlier and more intense fire stages where higher 624 monoterpene emissions were likely released from live or recently fallen trees that still contained 625 significant resin stores. The monoterpenes EF reported by Simpson et al. (2011) was likely even higher 626 given only two monoterpenes were speciated and emissions of other terpenes were likely (Hatch et al., 627 2019). Higher isoprene emissions in the present study compared to Simpson et al. (2011) could be related 628 to the comparatively larger smoldering component. Although limited data exist on the release of isoprene 629 as a function of fire intensity, negative relationships between isoprene and MCE were observed in 630 Australian temperate forest fires (Guérette et al., 2018) and wheat fields (Kumar et al., 2018).

631 Several furanoid compounds also exhibited significant emissions (Fig. 8c) including furfural, 632 furan, and methyl furan ranking 12<sup>th</sup>, 19<sup>th</sup>, and 22<sup>nd</sup> of the top 24 organic compounds, respectively. 633 Emissions of furanoids have been observed for a wide range of fuel types (Hatch et al., 2017; Simpson et 634 al., 2011). Agreement within uncertainties was found with BFF19 for furfural, and furan (Fig 9a). The EFs for furan  $(0.39\pm0.19 \text{ g kg}^{-1})$  and furfural  $(0.65\pm0.31 \text{ g kg}^{-1})$  were also similar to that in LAB18 (Fig. 635 636 9b), and TFF21 (Fig. 9c), as well as other ecosystems (Andreae, 2019) suggesting their emissions were 637 relatively insensitive to fire intensity and fuel mixture. The comparisons in Fig. 9 indicate that for the 638 higher emitting species, the current results are fairly similar, but for the lower emitting species, these 639 results are lower than previous reported values. These comparisons provide context for the emissions 640 reported in the present study and moves towards improved statistics to better constrain wildfire emissions. 641 Additional factors are considered to explain variability in emissions between this study and other reported 642 values, as well as within this study (NP vs SP). Differences and variability in burn conditions (e.g. fire 643 intensity, winds, fuel density, flame dynamics, fuel moisture) likely influence these comparisons; the 644 Screen 1 measurements in the present study were taken from 9-10 am LT when the fire was in a low 645 intensity, smoldering state, while those in Permar et al. (2021) and Simpson et al. (2011) took place 646 during mid-day under active wildfire conditions. Aircraft measurements in general have a higher 647 probability of sampling variable burn conditions compared to laboratory studies (Hodshire et al., 2019), 648 and as such, aircraft-derived EFs are likely to reflect variability for reactive species as speculated earlier 649 with isoprene. Particularly for reactive species that can exhibit complex variation across plumes, EFs 650 (and ERs) derived by integrating across plumes can be biased low, (Sect. 2.5; Peng et al., 2021; Decker et 651 al., 2021). Also, EFs derived using TC in this study may result in lower, albeit small, EFs compared to 652 reported values that do not account for all the carbon (estimated to be 1-2 % (Akagi et al., 2011)). 653 3.5 Evaluation of emissions models

654 **3.5.1** Comparison of EFs with the model emissions speciation profile

655

EFs derived in the present study are compared with those that are currently incorporated into the emissions component of the FireWork modelling system using the Forest Fire Emissions Prediction

658 System (CFFEPS). CFFEPS uses EFs allocated for 3 combustion states (flaming, smoldering and 659 residual) and for 8 species including lumped non-methane hydrocarbons (NMHC) based on United States 660 vegetation data compiled in Urbanski et al. (2014) (Table 3 in Chen et al., 2019). Fig. 9d (bolded 661 compounds) shows that the smoldering EFs in the present study were comparable for CO, CH<sub>4</sub> and 662 NMOG, but lower for PM<sub>1</sub> (PM<sub>2.5</sub>), NH<sub>3</sub>, SO<sub>2</sub> and NO<sub>x</sub> by factors of 3.4, 2.4, 6.6 and 17, respectively. 663 The lower EFs for these species implies that the CFFEPS EFs would not adequately capture their total 664 emissions under smoldering conditions for the boreal fuel in the current measurement study. 665 For incorporation into numerical air quality models, total organic gas (TOG=NMOG+CH<sub>4</sub>) 666 emissions are typically split into detailed chemical components using chemical mass speciation profiles, 667 and converted to lumped chemical mechanism species. In the FireWork modelling system, the 668 smoldering combustion TOG is split into components based on EPA's SPECIATEv4.5 profile (#95428) 669 (US EPA 2016, Urbanski et al.; 2014 - supplement Table A.2, Boreal Forest Duff/Organic soil). This 670 profile is ultimately compiled using laboratory data from Yokelson et al. (2013), Bertschi et al. (2003), 671 and Yokelson et al. (1997) based entirely on U.S. fuel types. EFs in the present study were found to be 672 generally lower than the laboratory-based EFs for 74 species in common ranging from factors of 1.7 to 673 8.5 including for monoterpenes, formic acid, phenol, and furan (Fig. 9d). A few species including 674 furfural, propane nitrile and ethyl styrene are comparable, while isoprene and pyruvic acid are notably

higher by factors 2 to 5.3.

676 For a research version of the FireWork system, the component speciation is mapped to the 677 SAPRC-11 chemical mechanism species (Carter and Heo, 2013) with detailed oxygenated compounds 678 and aromatic species, largely to better represent SOA formation processes. For comparison with the 679 measurement derived speciation profile in this study, EFs were first mapped to SAPRC-11 species and 680 normalized by the total identified mass species fraction without unknowns to obtain mass fractions of 681 relevant model mechanism species (Table S9). Comparing the normalized mass fractions for similar 682 mechanism species (Fig. S14) showed a substantially lower fraction of reactive alkanes (ALK5) with an 683 estimated 5 % in this study compared to 28 % in the SPECIATEv4.5 wildfire smoldering profile. Mass

684 fractions in this study are notably higher for the ACYL, ETHE, and ISOP lumped model species by 685 factors of 13, 7 and 51. The mass fraction of CH<sub>4</sub> is also different with 24 % of TOG in this study 686 compared to 4 % from the SPECIATE4.5 profile. The measurement derived chemical speciation profile 687 is expected to be different from the average speciation profile from EPA's SPECIATEv4.5 due to 688 differences in chemical species identification, fuel type, fire and measurement conditions, and 689 uncertainties on how measured compounds are mapped to lumped mechanism species. The emissions 690 profile developed in the present study can be used to improve predictions of wildfire smoldering 691 emissions specific to the Canadian boreal forest.

692

#### 693 **3.5.2** Linking aircraft and satellite observations to evaluate modelled emissions diurnal variability

694 Wildfires generally exhibit a diurnal cycle with fire intensities maximizing late afternoon and 695 diminishing at night having important implications for fire emissions (Chen et al., 2019). Evaluating 696 modelled emissions throughout the diurnal cycle with observations is a critical step in verifying smoke 697 predictions. Emissions models mostly parameterize diurnal fire emissions with prescribed profiles that 698 distribute daily total emissions to hourly. In CFFEPS, a diurnal profile is applied to allocate daily burn 699 area to hourly intervals, with highest activity in the late afternoon. The actual fuel consumed, and thus, 700 hourly emissions, is then calculated with depth of burn estimates driven by hourly meteorology (Chen et 701 al., 2019). In Fig. 10, for the wildfire in the present study, the hourly CFFEPS-predicted emissions 702 (orange dots) for selected compounds are shown between 2018-06-24 17:00 UTC and 2018-06-25 21:00 703 UTC, spanning the aircraft sample time (red arrow at 15:00 UTC). After 21:00 UTC, the discrepancy 704 between the CFFEPS-predicted emissions and FRP increased as a result of rain that passed through the 705 area that is not considered in the model bottom-up emission estimates (not shown in the figure). The 706 burning phases are outlined in the figure where flaming (light pink background) is assumed to occur when 707 the atmospheric conditions alongside fire behaviour and emissions model outputs infer a fireline intensity 708 >4,000 kW m<sup>-1</sup> (Forestry Canada Fire Danger Rating Group, 1992), and a smoldering fire (blue 709 background) for intensity <4000 kW m<sup>-1</sup>. The fire intensity distinction between flaming and smoldering

710 roughly aligns with the observed minimum for this particular fire with the fire radiative power (FRP, grev 711 dots) retrieval from the GOES-16 satellite sensor of 500 MW where smoldering occurs <500 MW and 712 flaming for >500 MW. The 500 MW threshold over the approximately 1,700 ha of actively smoldering 713 area observed by overnight VIIRS thermal detections gives an estimated energy density of 0.29 MW ha<sup>-1</sup>. 714 This energy density threshold for smoldering < 0.29 MW ha<sup>-1</sup> found in this study is in agreement with 715 O'Brien et al. (2015) who found flaming combustion at >0.4 MW ha<sup>-1</sup> for lower intensity flaming fires 716 and smoldering combustion at lower energy densities. The FRP represents the sum over all hotspots of 717 this fire for each 15-min observation period. Emission rates in metric tonnes per hour (t h<sup>-1</sup>) were derived 718 from selected aircraft measurements using a mass balance method that was designed to estimate pollutant 719 transfer rates through virtual screens using aircraft flight data (Gordon et al., 2015) (see SI Methods). 720 Emission rates were estimated to be  $29\pm2.1$  t h<sup>-1</sup> for PM<sub>1</sub>,  $433\pm26.7$  t h<sup>-1</sup> for CO,  $0.65\pm0.03$  t h<sup>-1</sup> for NO<sub>x</sub> 721 (as NO), and  $2.7\pm0.16$  t h<sup>-1</sup> for NH<sub>3</sub> (red arrows). Emission rates were also derived from satellite 722 observations (black arrows) for CO, NO<sub>x</sub>, and NH<sub>3</sub>. Emissions of CO were estimated using a flux method 723 as described in Stockwell et al. (2021) using TROPOMI satellite observations yielding  $1670\pm670$  t h<sup>-1</sup> at 724 19:06 UTC and 4050 $\pm$ 1620 t h<sup>-1</sup> at 20:48 UTC. NO<sub>x</sub> emissions (9.1 $\pm$ 3.4; scaled to t NO h<sup>-1</sup> at 19:06 UTC 725 (not enough high-quality observations for the 20:48 UTC overpass) were derived from the TROPOMI 726 NO<sub>2</sub> dataset using an Exponentially Modified Gaussian approach (Griffin et al., 2021). NH<sub>3</sub> emission 727 rates  $(5.6\pm3.9 \text{ t} \text{ h}^{-1})$  were derived from CRIS satellite observations at the satellite overpass time of 19:00 728 UTC by applying a flux method (Adams et al., 2019).

The aircraft measurements were taken when the FRP was low reflecting a smoldering surface fire. However, the satellite overpass occurred ~4 hrs later than the aircraft measurements close to the FRP daily maximum, after which rain passed through the area. The CFFEPS model, exhibiting a prescribed diurnal pattern, captures the increase in NO<sub>x</sub> and NH<sub>3</sub> emissions between that derived from the aircraft and satellites transitioning from a smoldering to predominantly flaming fire; NO<sub>x</sub> emissions increased by a factor >10, whereas the NH<sub>3</sub> emissions increased by a factor of approximately 2. This is in agreement with recent laboratory measurements that found that the release of NO<sub>x</sub> is favoured during the

flaming stage and the release of reduced forms of nitrogen, such as NH<sub>3</sub>, is favoured during the

smoldering phase (Roberts et al., 2020) (also see Fig. 4). However, the CFFEPS CO emission rates do
not track the increase in CO emissions between the aircraft-derived value and the two TROPOMI values,
indicating that the CO EF for flaming is low in the model. This highlights the need to validate model
emission rates with measurements to adjust and update the EFs accordingly.

741 The aircraft- and satellite-derived emission rates for CO, NO<sub>x</sub> and NH<sub>3</sub> were each ratioed to FRP 742 (in units of t h<sup>-1</sup> MW<sup>-1</sup>, referred to as R<sub>species/FRP</sub>) to represent the two end burning states ie. smoldering and 743 flaming conditions. R<sub>species/FRP</sub> values were estimated for the flaming and smoldering phases of the fire and 744 it was assumed that flaming occurred for FRP >500 MW and smoldering for FRP < 500 MW. . The CO 745  $R_{\text{species/FRP}}$  values were roughly twice as large during smoldering compared to flaming. For the two 746 satellite overpasses during the flaming phase of the fire, the R<sub>CO/FRP</sub> values were within the uncertainties 747  $(19:06 \text{ UTC } R_{CO/FRP} = 0.47 \pm 0.25 \text{ t} \text{ h}^{-1} \text{ MW}^{-1}; 20:48 \text{ UTC } R_{CO/FRP} = 0.43 \pm 0.23 \text{ t} \text{ h}^{-1} \text{ MW}^{-1}).$  The R<sub>NOx/FRP</sub> 748 value for NO<sub>x</sub> is also twice as large for flaming compared to smoldering, and for NH<sub>3</sub>, the R<sub>NH3/FRP</sub> value 749 is ~5 times larger for smoldering than flaming. Total emissions were then estimated by integrating the 750 GOES FRP over the period 2018-06-24 17:00 UTC to 2018-06-25 23:00 UTC (after which no more hot 751 spots were detected by GOES and the fire presumably extinguished), and applying the derived smoldering 752 and flaming R<sub>species/FRP</sub> values. Assuming that the fire went out when GOES did not observe any hot spots, 753 total emissions for this fire of CO, NO<sub>x</sub> and NH<sub>3</sub> are estimated at  $22,000\pm8700, 104\pm42$ , and  $84\pm33$ 754 tonnes, respectively. If the fire is assumed to have continued burning when GOES did not detect any fire 755 hot spots (between 22:00 - 04:00 UTC and 07:00 - 15:00 UTC, with an FRP of 150 MW (~GOES 756 detection limit; Roberts et al., 2015), the emissions increase to 24,000±9600, 106±43 and 98±39 tonnes, 757 respectively, providing an upper limit of emissions. The combination of aircraft and satellite-derived 758 emission estimates for multiple species helps to obtain the diurnal variability of emissions and to obtain 759 more complete details on the emission information across different burning stages.

760 **4. Summary and Implications** 

761 This study provides detailed emissions information for boreal forest wildfires under smoldering 762 combustion conditions. Consistent with previous results, highly speciated airborne measurements showed 763 a large diversity of chemical classes highlighting the complexity of emissions. Despite extensive 764 speciation across a range of NMOG volatilities, a substantial portion of NMOG<sub>T</sub> remained unidentified 765  $(47\pm15 \text{ to } 50\pm15 \text{ \%})$  and is expected to be comprised of more highly functionalized VOCs and I/SVOCs. 766 Although these compounds are challenging to measure, their characterization is necessary to more fully 767 understand particle-gas partitioning processes related to the formation of SOA. Methodological 768 advancements to achieve higher time resolution speciated measurements of I/SVOCs would move 769 towards further NMOG<sub>T</sub> closure and span a more complete range of volatilities. A detailed suite of EFs 770 that were derived in this study builds on previous work (e.g. Simpson et al., 2011; Andreae 2019) and can 771 be used to improve chemical speciation profiles that are relevant for air quality modelling of boreal forest 772 wildfires. Aircraft-derived emission estimates were paired with those from satellite observations 773 demonstrating their combined usefulness in assessing modelled emissions diurnal variability. As satellite 774 instrumentation and methodologies advance, linking emissions derived from aircraft (and ground) 775 observations for additional compounds will improve the ability to simulate and predict the diurnal 776 variation in wildfire emissions. 777 The results presented here represent only one smoldering boreal forest wildfire with limited in-plume 778 sampling times. Additional measurements are needed under a variety of fire conditions (combustion 779 state, fire stage, biomass mixtures, time of day, etc) in order to elucidate the major controlling factors and 780 improve statistical representation for constraining and modelling these sources. For example, 781 measurements are needed to assess dark chemistry reactions in biomass burning emissions which have

783 2020). In addition, reduced actinic flux associated with high particle loadings in biomass burning

782

emissions can influence plume chemistry (e.g. Juncosa-Calahorrano et al., 2021; Parrington et al., 2013).

been shown to be important in the formation of OA (Kodros et al., 2020) and brown carbon (Palm et al.;

785 The emissions information in this work can be used for evaluation and improvements of models that are

essential for reliable predictions of boreal forest wildfire pollutants and their downwind chemistry.

787 Acknowledgements

789	The authors acknowledge the significant technical and scientific contributions towards the success of this
790	study from the AQRD technical and data teams, the NRC team, and excellent program management by
791	Stewart Cober. The authors would like to thank Mark Shephard for his work on the CRIS NH <sub>3</sub> retrievals
792	and making those available. JCD, MH, and DRG acknowledge support from the National Science
793	Foundation (AGS1764126) and GERSTEL for their collaboration with the thermal desorption unit used
794	as part of this study, and MH also acknowledges the Goldwater Scholarship Foundation. SM.L.
795	acknowledges the support of the Ministry of Science and Technology of China (Grant
796	2019YFC0214700).
797	
798	Author contribution
799	KH, SML, JL, MJW, JJBW, AL, PB, RLM, CM, AS, RMS, SM, AD, and MW all contributed to the
800	collection and analyses of the aircraft observations in the field. JCD, MH, and DRG analysed the
801	cartridge samples. ZO contributed to the analyses and created many of the figures. DT contributed to the
802	analyses of the physical and combustion state of the wildfire fire. DG and EE provided the satellite
803	observations and DG wrote the satellite comparison section. JC contributed to the comparisons with the
804	model emission speciation profile. KH wrote the paper with input from all co-authors.
805	
806	Competing interests
807	The authors declare that they have no substantive conflicts of interest, but acknowledge that
808	DRG and JL are associate editors with Atmospheric Chemistry and Physics.
809	
810	Data availability
811	All data used in this publication are available upon request.
812	

# 813 **References**

- Adams, C., McLinden, C. A., Shephard, M. W., Dickson, N., Dammers, E., Chen, J., Makar, P., Cady-
- 815 Pereira, K. E., Tam, N., Kharol, S. K., Lamsal, L. N., and Krotkov, N. A.: Satellite-derived emissions of
- 816 carbon monoxide, ammonia, and nitrogen dioxide from the 2016 Horse River wildfire in the Fort
- 817 McMurray area, Atmos. Chem. Phys., 19, 2577-2599, https://doi.org/10.5194/acp-19-2577-2019, 2019.
- 818
- Ahern, A. T., Robinson, E. S., Tkacik, D. S., Saleh, R., Hatch, L. E., Barsanti, K. C., Stockwell, C. E.,
- 820 Yokelson, R. J., Presto, A. A., Robinson, A. L., Sullivan, R. C., and Donahue, N. M.: Production of
- 821 secondary organic aerosol during aging of biomass burning smoke from fresh fuels and its relationship to
- 822 VOC Precursors, J. Geophys. Res.-Atmos, 124, 3583-3606, https://doi.org/10.1029/2018JD029068, 2019.
- 823
- Akagi, S. K., Yokelson, R. J., Wiedinmyer, C., Alvarado, M. J., Reid, J. S., Karl, T., Crounse, J. D., and Wennberg, P. O.: Emission factors for open and domestic biomass burning for use in atmospheric models,
- 826 Atmos. Chem. Phys., 11, 4039-4072, https://doi.org/10.5194/acp-11-4039-2011, 2011.
- 827
- 828 Akagi, S. K., Craven, J. S., Taylor, J. W., McMeeking, G. R., Yokelson, R. J., Burling, I. R., Urbanski, S.
- 829 P., Wold, C. E., Seinfeld, J. H., Coe, H., Alvarado, M. J., and Weise, D. R.: Evolution of trace gases and
- particles emitted by a chaparral fire in California, Atmos. Chem. Phys., 12, 1397-1421,
- 831 https://doi.org/10.5194/acp-12-1397-2012, 2012.
- 832

Akagi, S. K., Yokelson, R. J., Burling, I. R., Meinardi, S., Simpson, I., Blake, D. R., McMeeking, G. R.,

- 834 Sullivan, A., Lee, T., Kreidenweis, S., Urbanski, S., Reardon, J., Griffith, D. W. T., Johnson, T. J., and
- 835 Weise, D. R.: Measurements of reactive trace gases and variable O3 formation rates in some South
- 836 Carolina biomass burning plumes, Atmos. Chem. Phys., 13, 1141-1165, https://doi.org/10.5194/acp-13837 1141-2013, 2013.
- 837 838
- Alvarado, M. J., Logan, J. A., Mao, J., Apel, E., Riemer, D., Blake, D., Cohen, R. C., Min, K. E., Perring,
- A. E., Browne, E. C., Wooldridge, P. J., Diskin, G. S., Sachse, G. W., Fuelberg, H., Sessions, W. R.,
- 841 Harrigan, D. L., Huey, G., Liao, J., Case-Hanks, A., Jimenez, J. L., Cubison, M. J., Vay, S. A.,
- 842 Weinheimer, A. J., Knapp, D. J., Montzka, D. D., Flocke, F. M., Pollack, I. B., Wennberg, P. O., Kurten,
- A., Crounse, J., Clair, J. M. S., Wisthaler, A., Mikoviny, T., Yantosca, R. M., Carouge, C. C., and Le
- Sager, P.: Nitrogen oxides and PAN in plumes from boreal fires during ARCTAS-B and their impact on
  ozone: an integrated analysis of aircraft and satellite observations, Atmos. Chem. Phys., 10, 9739-9760,
- 846 https://doi.org/10.5194/acp-10-9739-2010, 2010.
- 847
- Andreae, M. O.: Emission of trace gases and aerosols from biomass burning an updated assessment,
  Atmos. Chem. Phys., 19, 8523-8546, https://doi.org/10.5194/acp-19-8523-2019, 2019. Biomass burning
  emission factors https://edmond.mpgl.ee/imeji/collection/op2vVE8m0us\_gcGC, ver 14 Apr 2021.
- 851
- 852 Andreae, M. O. and Merlet, P.: Emission of trace gases and aerosols from biomass burning, Global
- 853 Biogeochem. Cy., 15, 955-966, https://doi.org/10.1029/2000GB001382, 2001.
- 854
- 855 Baumgardner, D., Kok, G., and Raga, G.: Warming of the Arctic lower stratosphere by light absorbing
- 856 particles, Geophys. Res. Lett., 31, L06117, https://doi.org/10.1029/2003GL018883, 2004.
- 857
- 858 Bertschi, I., Yokelson, R. J., Ward, D. E., Babbitt, R. E., Susott, R. A., Goode, J. G., and Hao, W. M.:
- Trace gas and particle emissions from fires in large diameter and belowground biomass fuels, J. Geophys.
  Res.-Atmos, 108, 8472, https://doi.org/10.1029/2002JD002100, 2003.
- 861

- 862 Bond-Lamberty, B., Gower, S. T., Wang, C., Cyr, P., and Veldhuis, H.: Nitrogen dynamics of a boreal
- black spruce wildfire chronosequence, Biogeochemistry, 81, 1-16, https://doi.org/10.1007/s10533-0069025-7, 2006.
- 865
- 866 Bruns, E. A., El Haddad, I., Slowik, J. G., Kilic, D., Klein, F., Baltensperger, U., and Prevot, A. S. H.:
- 867 Identification of significant precursor gases of secondary organic aerosols from residential wood
   868 combustion, Sci. Rep., 6, https://doi.org/10.1038/srep27881, 2016.
- 869
- 870 Burling, I. R., Yokelson, R. J., Akagi, S. K., Urbanski, S. P., Wold, C. E., Griffith, D. W. T., Johnson, T.
- J., Reardon, J., and Weise, D. R.: Airborne and ground-based measurements of the trace gases and
  particles emitted by prescribed fires in the United States, Atmos. Chem. Phys., 11, 12197-12216,
  https://doi.org/10.5194/acp-11-12197-2011, 2011.
- 874
- Bush, E. and Lemmen, D. S.: Canada's changing climate report, Government of Canada, Ottawa, ON.,
  444 pp., www.ChangingClimate.ca/CCCR2019, 2019.
- 878 Campos, I., Abrantes, N., Pereira, P., Micaelo, A. C., Vale, C., and Keizer, J. J.: Forest fires as potential
  879 triggers for production and mobilization of polycyclic aromatic hydrocarbons to the terrestrial ecosystem,
  880 Land Degrad, Dev., 30, 2360-2370, https://doi.org/10.1002/ldr.3427, 2019.
- Land Degrad. Dev., 30, 2360-2370, https://doi.org/10.1002/ldr.3427, 2019.
  881
- Carter, W. P. L. and Heo, G.: Development of revised SAPRC aromatics mechanisms, Atmos. Environ.,
  77, 404-414, https://doi.org/10.1016/j.atmosenv.2013.05.021, 2013.
- Carter, M.C., and Foster, C.D., Prescribed burning and productivity in southern pine forests: a review:
  Forest Ecol. Mgmt., 191, 93-109, 2004.
- Cascio, W. E.: Wildland fire smoke and human health, Sci. Total Environ., 624, 586-595,
  https://doi.org/10.1016/j.scitotenv.2017.12.086, 2018.
- 890

- 891 Chen, J., Anderson, K., Pavlovic, R., Moran, M. D., Englefield, P., Thompson, D. K., Munoz-Alpizar, R.,
- and Landry, H.: The FireWork v2.0 air quality forecast system with biomass burning emissions from the
- Canadian Forest Fire Emissions Prediction System v2.03, Geosci. Model Dev., 12, 3283-3310,
  https://doi.org/10.5194/gmd-12-3283-2019, 2019.
- 895
- Cherry, N. and Haynes, W.: Effects of the Fort McMurray wildfires on the health of evacuated workers:
  follow-up of 2 cohorts, Can. Med. Assoc. J., 5, E638-E645, https://doi.org/10.9778/cmajo.20170047,
  2017.
- 899
- 900 Ciccioli, P., Centritto, M., and Loreto, F.: Biogenic volatile organic compound emissions from vegetation
  901 fires, Plant, Cell Environ., 37, 1810-1825, https://doi.org/10.1111/pce.12336, 2014.
- 902
- Coggon, M. M., Lim, C. Y., Koss, A. R., Sekimoto, K., Yuan, B., Gilman, J. B., Hagan, D. H., Selimovic,
  V., Zarzana, K. J., Brown, S. S., Roberts, J. M., Müller, M., Yokelson, R., Wisthaler, A., Krechmer, J. E.,
- Jimenez, J. L., Cappa, C., Kroll, J. H., de Gouw, J., and Warneke, C.: OH chemistry of non-methane
   (1) 1000 (1000)
- 906 organic gases (NMOGs) emitted from laboratory and ambient biomass burning smoke: evaluating the
- 907 influence of furans and oxygenated aromatics on ozone and secondary NMOG formation, Atmos. Chem.
  908 Phys., 19, 14875-14899, https://doi.org/10.5194/acp-19-14875-2019, 2019.
- 909
- 910 Cole, A. S., Steffen, A., Eckley, C. S., Narayan, J., Pilote, M., Tordon, R., Graydon, J. A., St. Louis, V.
- L., Xu, X., and Branfireun, B. A.: A survey of mercury in air and precipitation across Canada: Patterns
- 912 and trends, Atmosphere, 5, 635-668, https://doi.org/10.3390/atmos5030635, 2014.

- 913
- 914 Cubison, M.J., Ortega, A.M., Haves, P.L., Farmer, D.K., Day, D., Lechner, M.J., Brune, W.H., Apel, E.,
- 915 Diskin, G.S., Fisher, J.A., Fuelberg, H.E., Hecobian, A., Knapp, D.J., Mikoviny, T., Riemer, D., Sachse,
- 916 G.W., Sessions, W., Weber, R.J., Weinheimer, A.J., Wisthaler, A., and J.L. Jimenez: Effects of aging on
- 917 organic aerosol from open biomass burning smoke in aircraft and laboratory studies, Atmos. Chem. Phys.,
- 918 11, 12049-12064, doi:10.5194/acp-11-12049-2011, 2011.
- 919
- 920 de Gouw, J. A., Warneke, C., Stohl, A., Wollny, A. G., Brock, C. A., Cooper, O. R., Holloway, J. S.,
- 921 Trainer, M., Fehsenfeld, F. C., Atlas, E. L., Donnelly, S. G., Stroud, V., and Lueb, A.: The VOC
- 922 composition of merged and aged forest fire plumes from Alaska and Western Canada, J. Geophys. Res. 923 Atmos, 111, D10303, https://doi.org/10.1029/2005JD006175, 2006.
- 924
- de Groot, W. J., Pritchard, J. M., and Lynham, T. J.: Forest floor fuel consumption and carbon emissions
  in Canadian boreal forest fires, Can. J. Forest Res., 39, 367-382, https://doi.org/10.1139/x08-192, 2009.
- 927
- 928 Decker, Z. C. J., Zarzana, K. J., Coggon, M., Min, K.-E., Pollack, I., Ryerson, T. B., Peischl, J., Edwards,
- P., Dubé, W. P., Markovic, M. Z., Roberts, J. M., Veres, P. R., Graus, M., Warneke, C., de Gouw, J.,
- Hatch, L. E., Barsanti, K. C., and Brown, S. S.: Nighttime chemical transformation in biomass burning
- plumes: A box model analysis Initialized with aircraft observations, Environ. Sci. Technol., 53, 2529-
- 932 2538, https://doi.org/10.1021/acs.est.8b05359, 2019.
- 933
- Decker, Z.C., Wang. S., Novel analysis to quantify plume crosswind heterogeneity applied to biomass
   burning smoke, EST, 2021.
- 936
- 937 Ditto, J. C., Joo, T., Slade, J. H., Shepson, P. B., Ng, N. L., and Gentner, D. R.: Nontargeted Tandem
- 938 Mass Spectrometry Analysis Reveals Diversity and Variability in Aerosol Functional Groups across
- Multiple Sites, Seasons, and Times of Day, Environmental Science and Technology Letters, 7, 60 69, https://doi.org/10.1021/acs.estlett.9b00702, 2020.
- 941
- Ditto, J. C., He, M., Hass-Mitchell, T. N., Moussa, S. G., Hayden, K., Li, S. M., Liggio, J., Leithead, A.,
- Lee, P., Wheeler, M. J., Wentzell, J. J. B., and Gentner, D. R.: Atmospheric evolution of emissions from a
  boreal forest fire: the formation of highly functionalized oxygen-, nitrogen-, and sulfur-containing organic
- 945 compounds, Atmos. Chem. Phys., 21, 255-267, https://doi.org/10.5194/acp-21-255-2021, 2021.
- 946
- Ditto, J. C., Machesky, J., and Gentner, D. R.: Analysis of reduced and oxidized nitrogen-containing
- 948 organic compounds at a coastal site in summer and winter, Atmos. Chem. Phys., 22, 3045949 3065, https://doi.org/10.5194/acp-22-3045-2022, 2022.
- 950
- Donahue, N. M., Epstein, S. A., Pandis, S. N., and Robinson, A. L.: A two-dimensional volatility basis
- 952 set: 1. organic-aerosol mixing thermodynamics, Atmos. Chem. Phys., 11, 3303-3318,
- 953 https://doi.org/10.5194/acp-11-3303-2011, 2011.
- 954
- EPA: Air method, toxic organics-15 (TO-15): Compendium of methods for the determination of toxic
  organic compounds in ambient air, second edition: Determination of volatile organic compounds (VOCs)
  in air collected in specially-prepared canisters and analyzed by gas chromatography/mass spectrometry
  (GC/MS)." EPA 625/R-96/010b, 1999.
- 959
- 960 Finlay, S. E., Moffat, A., Gazzard, R., Baker, D., and Murray, V.: Health impacts of wildfires, PLoS
- 961 Curr., 4, e4f959951cce959952c, https://doi.org/10.1371/4f959951cce2c, 2012.
- 962

- 963 Forestry Canada Fire Danger Rating Group: Development and structure of the Canadian Forest Fire
- 964 Behaviour Prediction System, Forestry Canada, Headquarters, Fire Danger Group and Science and
- 965 Sustainable Development Directorate, Ottawa, Information Report ST-X-3, 64 p., 992.966
- Garofalo, L. A., Pothier, M. A., Levin, E. J. T., Campos, T., Kreidenweis, S. M., and Farmer, D. K.:
  Emission and evolution of submicron organic aerosol in smoke from wildfires in the Western United
  States, ACS Earth Space Chem., 3, 1237-1247, https://doi.org/10.1021/acsearthspacechem.9b00125,
  2019.
- 970 971
- 972 Gilman, J. B., Lerner, B. M., Kuster, W. C., Goldan, P. D., Warneke, C., Veres, P. R., Roberts, J. M., de
- Gouw, J. A., Burling, I. R., and Yokelson, R. J.: Biomass burning emissions and potential air quality
- 974 impacts of volatile organic compounds and other trace gases from fuels common in the US, Atmos.
- 975 Chem. Phys., 15, 13915-13938, https://doi.org/10.5194/acp-15-13915-2015, 2015.
- 976
- Goode, J. G., Yokelson, R. J., Ward, D. E., Susott, R. A., Babbitt, R. E., Davies, M. A., and Hao, W. M.:
  Measurements of excess O<sub>3</sub>, CO<sub>2</sub>, CO, CH<sub>4</sub>, C<sub>2</sub>H<sub>4</sub>, C<sub>2</sub>H<sub>2</sub>, HCN, NO, NH<sub>3</sub>, HCOOH, CH<sub>3</sub>COOH, HCHO,
- and CH<sub>3</sub>OH in 1997 Alaskan biomass burning plumes by airborne fourier transform infrared spectroscopy
- 980 (AFTIR), J. Geophys. Res., 105, 22147-22166, https://doi.org/10.1029/2000JD900287, 2000.
- 981
- Gordon, M., Li, S. M., Staebler, R., Darlington, A., Hayden, K., O'Brien, J., and Wolde, M.: Determining
  air pollutant emission rates based on mass balance using airborne measurement data over the Alberta oil
  sands operations, Atmos. Meas. Tech., 8, 3745-3765, https://doi.org/10.5194/amt-8-3745-2015, 2015.
- 985
- 986 Griffin, D., McLinden, C. A., Dammers, E., Adams, C., Stockwell, C., Warneke, C., Bourgeois, I.,
- 987 Peischl, J., Ryerson, T. B., Zarzana, K. J., Rowe, J. P., Volkamer, R., Knote, C., Kille, N., Koenig, T. K.,
- Lee, C. F., Rollins, D., Rickly, P. S., Chen, J., Fehr, L., Bourassa, A., Degenstein, D., Hayden, K.,
- 989 Mihele, C., Wren, S. N., Liggio, J., Akingunola, A., and Makar, P.: Biomass burning nitrogen dioxide
- 990 emissions derived from space with TROPOMI: methodology and validation, Atmos. Meas. Tech.
- 991 Discuss. [preprint], https://doi.org/10.5194/amt-2021-223, in review, 2021.
- 992
- Griffith, D. W. T., Mankin, W. G., Coffey, M. T., Ward, D. E., and Riebau, A.: FTIR remote sensing of
  biomass burning emissions of CO<sub>2</sub>, CO, CH<sub>4</sub>, CH<sub>2</sub>O, NO, NO<sub>2</sub>, NH<sub>3</sub>, and N<sub>2</sub>O, in: Global Biomass
  Burning: Atmospheric, Climatic, and Biospheric Implications, edited by: Levine, J. S., MIT Press,
  Cambride, MA, United States, 230-241, 1991.
- 997
- 998 Guérette, E. A., Paton-Walsh, C., Desservettaz, M., Smith, T. E. L., Volkova, L., Weston, C. J., and
- 999 Meyer, C. P.: Emissions of trace gases from Australian temperate forest fires: Emission factors and 1000 dependence on modified combustion efficiency, Atmos. Chem. Phys., 18, 3717-3735,
- 1001 https://doi.org/10.5194/acp-18-3717-2018, 2018.
- 1002
- Hatch, L. E., Luo, W., Pankow, J. F., Yokelson, R. J., Stockwell, C. E., and Barsanti, K. C.: Identification and quantification of gaseous organic compounds emitted from biomass burning using two-dimensional
- 1005 gas chromatography–time-of-flight mass spectrometry, Atmos. Chem. Phys., 15, 1865-1899,
  1006 https://doi.org/10.5194/acp-15-1865-2015, 2015.
- 1006 1007
- 1008 Hatch, L. E., Rivas-Ubach, A., Jen, C. N., Lipton, M., Goldstein, A. H., and Barsanti, K. C.:
- 1009 Measurements of I/SVOCs in biomass-burning smoke using solid-phase extraction disks and two-
- 1010 dimensional gas chromatography, Atmos. Chem. Phys., 18, 17801-17817, https://doi.org/10.5194/acp-18-1011 17801-2018, 2018.
- 1012

- 1013 Hatch, L. E., Yokelson, R. J., Stockwell, C. E., Veres, P. R., Simpson, I. J., Blake, D. R., Orlando, J. J.,
- 1014 and Barsanti, K. C.: Multi-instrument comparison and compilation of non-methane organic gas emissions 1015
- from biomass burning and implications for smoke-derived secondary organic aerosol precursors, Atmos.
- 1016 Chem. Phys., 17, 1471-1489, https://doi.org/10.5194/acp-17-1471-2017, 2017. 1017
- 1018 Hatch, L. E., Jen, C. N., Kreisberg, N. M., Selimovic, V., Yokelson, R. J., Stamatis, C., York, R. A.,
- 1019 Foster, D., Stephens, S. L., Goldstein, A. H., and Barsanti, K. C.: Highly speciated measurements of
- 1020 terpenoids emitted from laboratory and mixed-conifer forest prescribed fires, Environ. Sci. Technol., 53, 1021 9418-9428, https://doi.org/10.1021/acs.est.9b02612, 2019.
- 1022
- 1023 Hecobian, A., Liu, Z., Hennigan, C. J., Huey, L. G., Jimenez, J. L., Cubison, M. J., Vay, S., Diskin, G. S.,
- 1024 Sachse, G. W., Wisthaler, A., Mikoviny, T., Weinheimer, A. J., Liao, J., Knapp, D. J., Wennberg, P. O., 1025 Kürten, A., Crounse, J. D., Clair, J. S., Wang, Y., and Weber, R. J.: Comparison of chemical
- 1026 characteristics of 495 biomass burning plumes intercepted by the NASA DC-8 aircraft during the
- 1027 ARCTAS/CARB-2008 field campaign, Atmos. Chem. Phys., 11, 13325-13337,
- 1028 https://doi.org/10.5194/acp-11-13325-2011, 2011.
- 1029
- 1030 Hodshire, A. L., Akherati, A., Alvarado, M. J., Brown-Steiner, B., Jathar, S. H., Jimenez, J. L.,
- 1031 Kreidenweis, S. M., Lonsdale, C. R., Onasch, T. B., Ortega, A. M., and Pierce, J. R.: Aging effects on
- 1032 biomass burning aerosol mass and composition: A critical review of field and laboratory studies, Environ.
- 1033 Sci. Technol., 53, 10007-10022, https://doi.org/10.1021/acs.est.9b02588, 2019.
- 1034
- 1035 Hosseini, S., Urbanski, S., Dixit, P., Li, Q., Burling, I., Yokelson, R., Johnson, T.E., Sharivastava, M., 1036 Jung, H., Weise, D.R., Miller, W., and Cocker, D.: Laboratory characterization of PM emissions from
- 1037 combustion of wildland biomass fuels, J. Geophys. Res., 118, 9914-9929,
- 1038 https://doi.org/10.1002/jgrd.50481, 2013. 1039
- 1040 Jia, Y., Yu, G., Gao, Y., He, N., Wang, Q., Jiao, C., and Zuo, Y.: Global inorganic nitrogen dry 1041 deposition inferred from ground- and space-based measurements, Sci. Rep., 6, 19810, 1042 https://doi.org/10.1038/srep19810, 2016.
- 1043
- 1044 Johnstone, J. F., Hollingsworth, T. N., Chapin III, F. S., and Mack, M. C.: Changes in fire regime break 1045 the legacy lock on successional trajectories in Alaskan boreal forest, Global Change Biol., 16, 1281-1295, 1046 https://doi.org/10.1111/j.1365-2486.2009.02051.x, 2010. 1047
- 1048 Jolley, M.D., Coe, H., McFiggans, G., Taylor, J.W., O'Shea, S.J., Le Breton, M., Bauguitte, S.J.-B.,
- 1049 Moller, S., Di Carlo, P., Aruffo, E., Palmer, P.I., Lee, J.D., Percival, C.J., and Gallagher, M.W.:
- 1050 Properties and evolution of biomass burning organic aerosol from Canadian boreal forest fires, Atmos. 1051 Chem. Phys., 15, 3077-3095, doi:10.5194/acp-15-3077-2015, 2015.
- 1052
- 1053 Juncosa Calahorrano, J. F., Lindaas, J., O'Dell, K., Palm, B. B., Peng, Q., Flocke, F., Pollack, I. B.,
- 1054 Garofalo, L. A., Farmer, D. K., Pierce, J. R., Collett Jr., J. L., Weinheimer, A., Campos, T., Hornbrook,
- 1055 R. S., Hall, S. R., Ullmann, K., Pothier, M. A., Apel, E. C., Permar, W., Hu, L., Hills, A. J., Montzka, D.,
- 1056 Tyndall, G., Thornton, J. A., and Fischer, E. V.: Daytime oxidized reactive nitrogen partitioning in
- 1057 western U.S. wildfire smoke plumes, J. Geophys. Res.-Atmos, 126, e2020JD033484,
- 1058 https://doi.org/10.1029/2020JD033484, 2021.
- 1059
- 1060 Kaiser, J. W., Heil, A., Andreae, M. O., Benedetti, A., Chubarova, N., Jones, L., Morcrette, J.-J.,
- 1061 Razinger, M., Schultz, M. G., Suttie, M., and van der Werf, G. R.: Biomass burning emissions estimated
- 1062 with a global fire assimilation system based on observed fire radiative power, Biogeosciences, 9, 527-
- 1063 554, https://doi.org/10.5194/bg-9-527-2012, 2012.

- Kallenborn, R., Halsall, C., Dellong, M., and Carlsson, P.: The influence of climate change on the global
  distribution and fate processes of anthropogenic persistent organic pollutants, J. Environ. Monitor., 14,
  2854-2869, https://doi.org/10.1039/c2em30519d, 2012.
- 1068 Khare, P., Marcotte, A., Sheu, R., Walsh, A. N., Ditto, J. C., and Gentner, D. R.: Advances in offline 1069 approaches for trace measurements of complex organic compound mixtures via soft ionization and high-1070 resolution tandem mass spectrometry, J. Chromatogr. A, 1598, 163-
- 1071 174, https://doi.org/10.1016/j.chroma.2019.03.037, 2019.
- 1072

- Kodros, J. K., Papanastasiou, D. K., Paglione, M., Masiol, M., Squizzato, S., Florou, K., Skyllakou, K.,
  Kaltsonoudis, C., Nenes, A., and Pandis, S. N.: Rapid dark aging of biomass burning as an overlooked
  source of oxidized organic aerosol, P. Natl. Acad. Sci. USA, 117, 33028-33033,
- 1076 https://doi.org/10.1073/pnas.2010365117, 2020. 1077
- 1078 Kondo, Y., Sahu, L., Moteki, N., Khan, F., Takegawa, N., Liu, X., Koike, M., and Miyakawa, T.:
- 1079 Consistency and traceability of black carbon measurements made by laser-induced incandescence,
- thermal-optical transmittance, and filter-based photo-absorption techniques, Aerosol Sci. Tech., 45, 295312, https://doi.org/10.1080/02786826.2010.533215, 2011.
- 1082
- 1083 Koss, A. R., Sekimoto, K., Gilman, J. B., Selimovic, V., Coggon, M. M., Zarzana, K. J., Yuan, B.,
- 1084 Lerner, B. M., Brown, S. S., Jimenez, J. L., Krechmer, J., Roberts, J. M., Warneke, C., Yokelson, R. J.,
- and de Gouw, J.: Non-methane organic gas emissions from biomass burning: identification,
- quantification, and emission factors from PTR-ToF during the FIREX 2016 laboratory experiment,
   Atmos. Chem. Phys., 18, 3299-3319, https://doi.org/10.5194/acp-18-3299-2018, 2018.
- 1088 1089 Kotchenruther, R. A. and Hobbs, P. V.: Humidification factors of aerosols from biomass burning in 1000 Brazil, J. Courthur Bog. Atmos. 102, 22081 22080, https://doi.org/10.1020/08JD00240, 1008
- 1090 Brazil, J. Geophys. Res.-Atmos, 103, 32081-32089, https://doi.org/10.1029/98JD00340, 1998. 1091
- Kou-Giesbrecht, S. and Menge, D.: Nitrogen-fixing trees could exacerbate climate change under elevated
  nitrogen deposition, Nat. Commun., 10, 1493, https://doi.org/10.1038/s41467-019-09424-2, 2019.
- Kumar, V., Chandra, B. P., and Sinha, V.: Large unexplained suite of chemically reactive compounds
  present in ambient air due to biomass fires, Sci. Rep., 8, 626, https://doi.org/10.1038/s41598-017-191393, 2018.
- 1098
  1099 Laborde, M., Mertes, P., Zieger, P., Dommen, J., Baltensperger, U., and Gysel, M.: Sensitivity of the
  1100 Single Particle Soot Photometer to different black carbon types, Atmos. Meas. Tech., 5, 1031-1043,
  - 1101 https://doi.org/10.5194/amt-5-1031-2012, 2012.
  - 1102
  - Landis, M. S., Edgerton, E. S., White, E. M., Wentworth, G. R., Sullivan, A. P., and Dillner, A. M.: The impact of the 2016 Fort McMurray Horse River Wildfire on ambient air pollution levels in the Athabasca
- 1105 Oil Sands Region, Alberta, Canada, Sci. Total Environ., 618, 1665-1676,
- 1106 https://doi.org/10.1016/j.scitotenv.2017.10.008, 2018.
- 1107
- 1108 Lee, T., Sullivan, A. P., Mack, L., Jimenez, J. L., Kreidenweis, S. M., Onasch, T. B., Worsnop, D. R.,
- Malm, W., Wold, C. E., Hao, W. M., and Collett Jr, J. L.: Chemical smoke marker emissions during
   flaming and smoldering phases of laboratory open burning of wildland fuels, Aerosol Sci. Tech., 44, i-v,
- 1111 https://doi.org/10.1080/02786826.2010.499884, 2010.
- 1112
- 1113 Leifer, I., Melton, C. Tratt, D.M., Buckland, K.N., Clarisse, L., Coheur, P., Frash, J., Gupta, M., Johnson,
- 1114 P.D., Leen, J.B., Van Damme, M., Whitburn, S., and Yurganov, L.: Remote sensing and in situ

- 1115 measurements of methane and ammonia emissions from a megacity dairy complex: Chino, CA, Environ.
- 1116 Poll., 221, 37-51, https://doi.org/10.1016/j.envpol.2016.09.083, 2017.
- 1117
- 1118 Lerner, B. M., Gilman, J. B., Aikin, K. C., Atlas, E. L., Goldan, P. D., Graus, M., Hendershot, R.,
- 1119 Isaacman-VanWertz, G. A., Koss, A., Kuster, W. C., Lueb, R. A., McLaughlin, R. J., Peischl, J., Sueper,
- 1120 D., Ryerson, T. B., Tokarek, T. W., Warneke, C., Yuan, B., and de Gouw, J. A.: An improved, automated
- 1121 whole air sampler and gas chromatography mass spectrometry analysis system for volatile organic
- 1122 compounds in the atmosphere, Atmos. Meas. Tech., 10, 291-313, https://doi.org/10.5194/amt-10-291-2017, 2017.
- 1123
- 1124
- 1125 Li, Y., Poschl, U., and Shiraiwa, M.: Molecular corridors and parameterizations of volatility in the
- 1126 chemical evolution of organic aerosols, Atmos. Chem. Phys., 16, 3327-3344, https://doi.org/10.5194/acp-1127 16-3327-2016, 2016.
- 1128
- 1129 Lindaas, J., Pollack, I. B., Garofalo, L. A., Pothier, M. A., Farmer, D. K., Kreidenweis, S. M., Campos, T.
- 1130 L., Flocke, F., Weinheimer, A. J., Montzka, D. D., Tyndall, G. S., Palm, B. B., Peng, Q., Thornton, J. A.,
- 1131 Permar, W., Wielgasz, C., Hu, L., Ottmar, R. D., Restaino, J. C., Hudak, A. T., Ku, I.-T., Zhou, Y., Sive,
- 1132 B. C., Sullivan, A., Collett Jr, J. L., and Fischer, E. V.: Emissions of reactive nitrogen from western U.S.
- 1133 wildfires during summer 2018, J. Geophys. Res.-Atmos, 125, e2020JD032657,
- 1134 https://doi.org/10.1029/2020JD032657, 2020.
- 1135
- 1136 Liu, X., Huey, L. G., Yokelson, R. J., Selimovic, V., Simpson, I. J., Müller, M., Jimenez, J. L.,
- 1137 Campuzano-Jost, P., Beyersdorf, A. J., Blake, D. R., Butterfield, Z., Choi, Y., Crounse, J. D., Day, D. A.,
- 1138 Diskin, G. S., Dubey, M. K., Fortner, E., Hanisco, T. F., Hu, W., King, L. E., Kleinman, L., Meinardi, S.,
- 1139 Mikoviny, T., Onasch, T. B., Palm, B. B., Peischl, J., Pollack, I. B., Ryerson, T. B., Sachse, G. W.,
- 1140 Sedlacek, A. J., Shilling, J. E., Springston, S., St. Clair, J. M., Tanner, D. J., Teng, A. P., Wennberg, P.
- 1141 O., Wisthaler, A., and Wolfe, G. M.: Airborne measurements of western U.S. wildfire emissions:
- 1142 Comparison with prescribed burning and air quality implications, J. Geophys. Res.-Atmos., 122, 6108-
- 1143 6129, https://doi.org/10.1002/2016JD026315, 2017.
- 1144
- 1145 Loehman, R. A., Reinhardt, E., and Riley, K. L.: Wildland fire emissions, carbon, and climate: Seeing the
- 1146 forest and the trees – A cross-scale assessment of wildfire and carbon dynamics in fire-prone, forested 1147 ecosystems, For. Ecol. Manag., 317, 9-19, https://doi.org/10.1016/j.foreco.2013.04.014, 2014.
- 1148
- 1149 Matz, C.J., Egyed, M., Xi, G., Racine, J., Pavlovic, R., Rittmaster, R., Henderson, S.B., and Stieb, D.M.:
- 1150 Health impact analysis of PM<sub>2.5</sub> from wildfire smoke in Canada (2013-2015, 2017-2018), Sci. Total
- 1151 Environ., 725(10), https://doi.org/10.1016/j.scitotenv.2020.138506, 2020.
- 1152

1153 May, A. A., McMeeking, G. R., Lee, T., Taylor, J. W., Craven, J. S., Burling, I., Sullivan, A. P., Akagi,

- 1154 S., Collett Jr., J. L., Flynn, M., Coe, H., Urbanski, S. P., Seinfeld, J. H., Yokelson, R. J., and Kreidenweis,
- 1155 S. M.: Aerosol emissions from prescribed fires in the United States: A synthesis of laboratory and aircraft 1156 measurements, J. Geophys. Res.-Atmos, 119, 11826-11849, https://doi.org/10.1002/2014JD021848,
- 1157 2014.
- 1158
- 1159 McGee, T., McFarlane, B., and Tymstra, C.: Chapter 3 - Wildfire: A Canadian Perspective, in: Wildfire
- 1160 Hazards, Risks and Disasters, edited by: Shroder, J. F., and Paton, D., Elsevier, Amsterdam, The
- 1161 Netherlands, 35-58, https://doi.org/10.1016/B978-0-12-410434-1.00003-8, 2015.
- 1162
- 1163 McLagan, D. S., Stupple, G. W., Darlington, A., Hayden, K., and Steffen, A.: Where there is smoke there
- 1164 is mercury: Assessing boreal forest fire mercury emissions using aircraft and highlighting uncertainties

- associated with upscaling emissions estimates, Atmos. Chem. Phys., 21, 5635-5653,
- 1166 https://doi.org/10.5194/acp-21-5635-2021, 2021.
- 1167
- 1168 McMeeking, G. R., Kreidenweis, S. M., Baker, S., Carrico, C. M., Chow, J. C., Collett Jr., J. L., Hao, W.
- 1169 M., Holden, A. S., Kirchstetter, T. W., Malm, W. C., Moosmüller, H., Sullivan, A. P., and Wold, C. E.:
- 1170 Emissions of trace gases and aerosols during the open combustion of biomass in the laboratory, J.
- 1171 Geophys. Res.-Atmos, 114, D19210, https://doi.org/10.1029/2009JD011836, 2009. 1172
- Miller, D. J., Sun, K., Zondlo, M. A., Kanter, D., Dubovik, O., Welton, E. J., Winker, D. M., and Ginoux,
  P.: Assessing boreal forest fire smoke aerosol impacts on U.S. air quality: A case study using multiple
- 1175 data sets, J. Geophys. Res.-Atmos, 116, D22209, https://doi.org/10.1029/2011JD016170, 2011.
- 1176
- Moteki, N. and Kondo, Y.: Dependence of laser-induced incandescence on physical properties of black
  carbon aerosols: Measurements and theoretical interpretation, Aerosol Sci. Tech., 44, 663-675,
  https://doi.org/10.1080/02786826.2010.484450, 2010.
- 1180
- 1181 Moussa, S. G., Leithead, A., Li, S. M., Chan, T. W., Wentzell, J. J. B., Stroud, C., Zhang, J. H., Lee, P.,
- Lu, G., Brook, J. R., Hayden, K., Narayan, J., and Liggio, J.: Emissions of hydrogen cyanide from on-
- 1183 road gasoline and diesel vehicles, Atmos. Environ., 131, 185-195,
- 1184 https://doi.org/10.1016/j.atmosenv.2016.01.050, 2016.
- 1185
- NRCan, Blueprint for wildland fire science in Canada (2019-2029), Sankey, S., Technical
  Coordinator. Canadian Forest Service, Northern Forestry Centre, Edmonton, AB, 45p,
- 1188 https://cfs.nrcan.gc.ca/publications?id=39429, 2018.
- 1189
- 1190 O'Brien, J. J., Loudermilk, E. L., Hornsby, B. S., Hudak, A. T., Bright, B. C., Dickinson, M. B., Hiers, J.
- 1191 K., Teske, C., and Ottmar, R. D.: High-resolution infrared thermography for capturing wildland fire
- behaviour: RxCADRE 2012, Int. J. Wildland Fire, 25, 62-75, https://doi.org/10.1071/WF14165, 2015.
- Palm, B. B., Peng, Q., Fredrickson, C. D., Lee, B. H., Garofalo, L. A., Pothier, M. A., Kreidenweis, S.
  M., Farmer, D. K., Pokhrel, R. P., Shen, Y., Murphy, S. M., Permar, W., Hu, L., Campos, T. L., Hall, S.
- R., Ullmann, K., Zhang, X., Flocke, F., Fischer, E. V., and Thornton, J. A.: Quantification of organic
  aerosol and brown carbon evolution in fresh wildfire plumes, P. Natl. Acad. Sci. USA, 117, 29469-29477,
- 1198 https://doi.org/10.1073/pnas.2012218117, 2020.
- 1199
- 1200 Peng, Q., Palm, B. B., Melander, K. E., Lee, B. H., Hall, S. R., Ullmann, K., Campos, T., Weinheimer, A.
- 1201 J., Apel, E. C., Hornbrook, R. S., Hills, A. J., Montzka, D. D., Flocke, F., Hu, L., Permar, W., Wielgasz,
- 1202 C., Lindaas, J., Pollack, I. B., Fischer, E. V., Bertram, T. H., and Thornton, J. A.: HONO Emissions from
- 1203 Western U.S. Wildfires Provide Dominant Radical Source in Fresh Wildfire Smoke, Environ. Sci.
- 1204 Technol., 54, 5954-5963, https://doi.org/10.1021/acs.est.0c00126, 2020.
- 1205
- 1206 Permar, W., Wang, Q., Selimovic, V., Wielgasz, C., Yokelson, R. J., Hornbrook, R. S., Hills, A. J., Apel,
- 1207 E. C., Ku, I.-T., Zhou, Y., Sive, B. C., Sullivan, A. P., Collett Jr, J. L., Campos, T. L., Palm, B. B., Peng,
- 1208 Q., Thornton, J. A., Garofalo, L. A., Farmer, D. K., Kreidenweis, S. M., Levin, E. J. T., DeMott, P. J., 1200 Electron E. V. and Hu, L. Emissions of trace arguing structure IVS with the second structure in the second
- 1209 Flocke, F., Fischer, E. V., and Hu, L.: Emissions of trace organic gases from western U.S. wildfires based
- 1210 on WE-CAN aircraft measurements, J. Geophys. Res.-Atmos, 126, e2020JD033838,
- 1211 https://doi.org/10.1029/2020JD033838, 2021.
- 1212
- 1213 Randerson, J. T., Liu, H., Flanner, M. G., Chambers, S. D., Jin, Y., Hess, P. G., Pfister, G., Mack, M. C.,
- 1214 Treseder, K. K., Welp, L. R., Chapin, F. S., Harden, J. W., Goulden, M. L., Lyons, E., Neff, J. C., Schuur,

- 1215 E. A., and Zender, C. S.: The impact of boreal forest fire on climate warming, Science, 314, 1130-1132, 1216 https://doi.org/10.1126/science.1132075, 2006.
- 1217
- 1218 Reid, C. E., Brauer, M., Johnston, F. H., Jerrett, M., Balmes, J. R., and Elliott, C. T.: Critical review of
- 1219 health impacts of wildfire smoke exposure, Environ. Health Persp., 124, 1334-1343,
- 1220 https://doi.org/10.1289/ehp.1409277, 2016.
- 1221 1222 Reid, J. S., Koppmann, R., Eck, T. F., and Eleuterio, D. P.: A review of biomass burning emissions part 1223 II: intensive physical properties of biomass burning particles, Atmos. Chem. Phys., 5, 799-825, 1224 https://doi.org/10.5194/acp-5-799-2005, 2005.
- 1225
- 1226 Roberts, J. M., Stockwell, C. E., Yokelson, R. J., de Gouw, J., Liu, Y., Selimovic, V., Koss, A. R.,
- 1227 Sekimoto, K., Coggon, M. M., Yuan, B., Zarzana, K. J., Brown, S. S., Santin, C., Doerr, S. H., and
- 1228 Warneke, C.: The nitrogen budget of laboratory-simulated western US wildfires during the FIREX 2016 1229 Fire Lab study, Atmos. Chem. Phys., 20, 8807-8826, https://doi.org/10.5194/acp-20-8807-2020, 2020.
- 1230
- 1231 Roberts, G., Wooster, M. J., Xu, W., Freeborn, P. H., Morcrette, J. J., Jones, L., Benedetti, A., Jiangping,
- 1232 H., Fisher, D., and Kaiser, J. W.: LSA SAF Meteosat FRP products – Part 2: Evaluation and 1233
- demonstration for use in the Copernicus Atmosphere Monitoring Service (CAMS), Atmos. Chem. Phys., 1234 15, 13241-13267, https://doi.org/10.5194/acp-15-13241-2015, 2015.
- 1235
- 1236 Rogers, C. F., Hudson, J. G., Hallett, J., and Penner, J. E.: Cloud Droplet Nucleation by Crude-Oil Smoke 1237 and Coagulated Crude-Oil Wood Smoke Particles, Atmos. Environ. a-Gen, 25, 2571-2580, 1238 https://doi.org/10.1016/0960-1686(91)90174-6, 1991.
- 1239
- 1240 Rogers, H. M., Ditto, J. C., and Gentner, D. R.: Evidence for impacts on surface-level air quality in the 1241 northeastern US from long-distance transport of smoke from North American fires during the Long Island
- 1242 Sound Tropospheric Ozone Study (LISTOS) 2018, Atmos. Chem. Phys., 20, 671-
- 1243 682, https://doi.org/10.5194/acp-20-671-2020, 2020.
- 1244
- 1245 Schwarz, J. P., Gao, R. S., Fahey, D. W., Thomson, D. S., Watts, L. A., Wilson, J. C., Reeves, J. M.,
- 1246 Darbeheshti, M., Baumgardner, D. G., Kok, G. L., Chung, S. H., Schulz, M., Hendricks, J., Lauer, A., 1247 Kärcher, B., Slowik, J. G., Rosenlof, K. H., Thompson, T. L., Langford, A. O., Loewenstein, M., and 1248 Aikin, K. C.: Single-particle measurements of midlatitude black carbon and light-scattering aerosols from
- 1249 the boundary layer to the lower stratosphere, J. Geophys. Res.-Atmos, 111, D16207,
- 1250 https://doi.org/10.1029/2006JD007076, 2006. 1251
- 1252 Seidl, R., Thom, D., Kautz, M., Martin-Benito, D., Peltoniemi, M., Vacchiano, G., Wild, J., Ascoli, D., 1253 Petr, M., Honkaniemi, J., Lexer, M. J., Trotsiuk, V., Mairota, P., Svoboda, M., Fabrika, M., Nagel, T. A., and Reyer, C. P. O.: Forest disturbances under climate change, Nat. Clim. Change, 7, 395-402, 1254
- 1255 https://doi.org/10.1038/nclimate3303, 2017. 1256
- 1257 Seinfeld, J. H. and Pandis, S. N.: Atmospheric chemistry and physics: from air pollution to climate 1258 change, John Wiley & Sons, New York, 1998.
- 1259
- 1260 Sekimoto, K., Li, S.-M., Yuan, B., Koss, A., Coggon, M., Warneke, C., and de Gouw, J.: Calculation of
- 1261 the sensitivity of proton-transfer-reaction mass spectrometry (PTR-MS) for organic trace gases using
- 1262 molecular properties, Int. J. Mass Spectrom., 421, 71-94, https://doi.org/10.1016/j.ijms.2017.04.006, 2017.
- 1263

- 1265 Sheu, R., Marcotte, A., Khare, P., Charan, S., Ditto, J. C., and Gentner, D. R.: Advances in offline
- 1266 approaches for chemically speciated measurements of trace gas-phase organic compounds via adsorbent
- 1267 tubes in an integrated sampling-to-analysis system, J. Chromatogr. A, 1575, 80-
- 1268 90, https://doi.org/10.1016/j.chroma.2018.09.014, 2018. 1269
- Simoneit, B. R. T., Schauer, J. J., Nolte, C. G., Oros, D. R., Elias, V. O., Fraser, M. P., Rogge, W. F., and
  Cass, G. R.: Levoglucosan, a tracer for cellulose in biomass burning and atmospheric particles, Atmos.
- 1272 Environ., 33, 173-182, https://doi.org/10.1016/S1352-2310(98)00145-9, 1999.
- 1273
- 1274 Simpson, I. J., Akagi, S. K., Barletta, B., Blake, N. J., Choi, Y., Diskin, G. S., Fried, A., Fuelberg, H. E.,
- 1275 Meinardi, S., Rowland, F. S., Vay, S. A., Weinheimer, A. J., Wennberg, P. O., Wiebring, P., Wisthaler,
- 1276 A., Yang, M., Yokelson, R. J., and Blake, D. R.: Boreal forest fire emissions in fresh Canadian smoke
- 1277 plumes:  $C_1$ - $C_{10}$  volatile organic compounds (VOCs),  $CO_2$ , CO,  $NO_2$ , NO, HCN and  $CH_3CN$ , Atmos. 1278 Chem. Phys., 11, 6445-6463, https://doi.org/10.5194/acp-11-6445-2011, 2011.
- 1278 Chem. Phys., 11, 6445-6463, https://doi.org/10.5194/acp-11-6445-20. 1279
- 1280 Singh, H. B., Anderson, B. E., Brune, W. H., Cai, C., Cohen, R. C., Crawford, J. H., Cubison, M. J.,
- 1281 Czech, E. P., Emmons, L., Fuelberg, H. E., Huey, G., Jacob, D. J., Jimenez, J. L., Kaduwela, A., Kondo,
- 1282 Y., Mao, J., Olson, J. R., Sachse, G. W., Vay, S. A., Weinheimer, A., Wennberg, P. O., and Wisthaler, A.:
- 1283 Pollution influences on atmospheric composition and chemistry at high northern latitudes: Boreal and
- 1284 California forest fire emissions, Atmos. Environ., 44, 4553-4564,
- 1285 https://doi.org/10.1016/j.atmosenv.2010.08.026, 2010. 1286
- Stephens, M., Turner, N., and Sandberg, J.: Particle identification by laser-induced incandescence in a
  solid-state laser cavity, Appl. Optics, 42, 3726-3736, https://doi.org/10.1364/ao.42.003726, 2003.
- Stockwell, C. E., Veres, P. R., Williams, J., and Yokelson, R. J.: Characterization of biomass burning
  emissions from cooking fires, peat, crop residue, and other fuels with high-resolution proton-transferreaction time-of-flight mass spectrometry, Atmos. Chem. Phys., 15, 845-865, https://doi.org/10.5194/acp15-845-2015, 2015.
- 1294
- 1295 Stockwell, C. E., Kupc, A., Witkowski, B., Talukdar, R. K., Liu, Y., Selimovic, V., Zarzana, K. J.,
- Sekimoto, K., Warneke, C., Washenfelder, R. A., Yokelson, R. J., Middlebrook, A. M., and Roberts, J.
  M.: Characterization of a catalyst-based conversion technique to measure total particulate nitrogen and
  organic carbon and comparison to a particle mass measurement instrument, Atmos. Meas. Tech., 11,
  2749-2768, https://doi.org/10.5194/amt-11-2749-2018, 2018.
- 1300
- Stockwell, C. E., Bela, M., Coggon, M. M., Gkatzelis, G. I., Wiggins, E. B., Gargulinski, E. M., Shingler,
  T., Fenn, M., Griffin, D., Holmes, C. D., Ye, X., Saide, P. E., Bourgeois, I., Peischl, J., Womack, C. C.,
- 1303 Washenfelder, R. A., Veres, P. R., Neuman, J. A., Gilman, J. B., Lamplugh, A., Schwantes, R. H.,
- 1304 McKeen, S. A., Wisthaler, A., Piel, F., Guo, H., Campuzano-Jost, P., Jimenez, J. L., Fried, A., Hanisco,
- 1305 T. F., Huey, L. G., Kondragunta, S., Zhang, X., Perring, A., Katich, J. M., Diskin, G. S., Nowak, J. B.,
- 1306 Bui, T. P., Halliday, H. S., Pereira, G., James, E. P., Ahmadov, R., McLinden, C. A., Soja, A. J., Moore,
- 1307 R. H., Hair, J. W., and Warneke, C.: Airborne emission rate measurements validate remote sensing
- observations and emission inventories of western U.S. wildfires, ES&T, 56, 7564-7577,
   https://doi.org/10.1021/acs.est.1c07121, 2021.
- 1310
- 1311 Urbanski, S.: Wildland fire emissions, carbon, and climate: Emission factors, For. Ecol. Manag., 317, 51-
- 1312 60, https://doi.org/10.1016/j.foreco.2013.05.045, 2014.
- 1313

- 1314 Urbanski, S. P.: Combustion efficiency and emission factors for wildfire-season fires in mixed conifer
- 1315 forests of the Northern Rocky Mountains, US, Atmos. Chem. Phys., 13, 7241-7262,
- 1316 https://doi.org/10.5194/acp-13-7241-2013, 2013.
- 1317
- 1318 Urbanski, S. P., Hao, W. M., and Baker, S.: Chapter 4 Chemical Composition of Wildland Fire
- 1319 Emissions, in: Developments in Environmental Science, edited by: Bytnerowicz, A., Arbaugh, M. J.,
- 1320 Riebau, A. R., and Andersen, C., Elsevier, Amsterdam, The Netherlands, 79-107,
- 1321 https://doi.org/10.1016/S1474-8177(08)00004-1, 2009. 1322
- 1323 Uresk, D. W., Cline, J. F., and Rickard, W. H.: Growth rates of a cheatgrass community and some 1324 associated factors, J. Range Manage., 32, 168-170, https://doi.org/10.2307/3897114, 1979.
- 1325
- US EPA, SPECIATE Version 4.5 Database Development Documentation, Final Report EPA/600/R 16/294, September 2016. Available https://www.epa.gov/air-emissions-modeling/speciate.
- van der Werf, G. R., Randerson, J. T., Giglio, L., Collatz, G. J., Kasibhatla, P. S., and Arellano Jr, A. F.:
  Interannual variability in global biomass burning emissions from 1997 to 2004, Atmos. Chem. Phys., 6,
  3423-3441, https://doi.org/10.5194/acp-6-3423-2006, 2006.
- van der Werf, G. R., Randerson, J. T., Giglio, L., van Leeuwen, T. T., Chen, Y., Rogers, B. M., Mu, M.,
  van Marle, M. J. E., Morton, D. C., Collatz, G. J., Yokelson, R. J., and Kasibhatla, P. S.: Global fire
  emissions estimates during 1997–2016, Earth Syst. Sci. Data, 9, 697–720, https://doi.org/10.5194/essd-9697-2017, 2017.
- Veres, P., Roberts, J. M., Burling, I. R., Warneke, C., de Gouw, J., and Yokelson, R. J.: Measurements of
  gas-phase inorganic and organic acids from biomass fires by negative-ion proton-transfer chemicalionization mass spectrometry, J. Geophys. Res.-Atmos, 115, D23302,
- 1341 https://doi.org/10.1029/2010JD014033, 2010.
- 1342

Ward, D. E. and Radke, L. F.: Emissions measurements from vegetation fires: A comparative evaluation
of methods and results, in: Fire in the Environment: The Ecological, Atmospheric, and Climatic
Importance of Vegetation Fires. Dahlem Workshop Reports: Environmental Sciences Research Report
added by: Crutzen, P. J., and Goldammer, J. G., John Wiley & Sons, Chischester, England, 53-76,
1993.

- Whitman, E., Parisien, M. A., Thompson, D. K., and Flannigan, M. D.: Short-interval wildfire and
  drought overwhelm boreal forest resilience, Sci. Rep., 9, 18796, https://doi.org/10.1038/s41598-01955036-7, 2019.
- 1352
- 1353 Wiggins, E. B., Andrews, A., Sweeney, C., Miller, J. B., Miller, C. E., Veraverbeke, S., Commane, R.,
- Wofsy, S., Henderson, J. M., and Randerson, J. T.: Boreal forest fire CO and CH<sub>4</sub> emission factors
  derived from tower observations in Alaska during the extreme fire season of 2015, Atmos. Chem. Phys.,
- 1356 21, 8557-8574, https://doi.org/10.5194/acp-21-8557-2021, 2021.
- 1357
- 1358 Wiggins, E. B., Soja, A. J., Gargulinski, E., Halliday, H. S., Pierce, R. B., Schmidt, C. C., Nowak, J. B.,
- 1359 DiGangi, J. P., Diskin, G. S., Katich, J. M., Perring, A. E., Schwarz, J. P., Anderson, B. E., Chen, G.,
- 1360 Crosbie, E. C., Jordan, C., Robinson, C. E., Sanchez, K. J., Shingler, T. J., Shook, M., Thornhill, K. L.,
- 1361 Winstead, E. L., Ziemba, L. D., and Moore, R. H.: High Temporal Resolution Satellite Observations of
- 1362 Fire Radiative Power Reveal Link Between Fire Behavior and Aerosol and Gas Emissions, Geophys. Res.
- 1363 Lett., 47, https://doi.org/10.1029/2020GL090707, 2020.
- 1364

- Wotton, B. M., Nock, C. A., and Flannigan, M. D.: Forest fire occurrence and climate change in Canada,
  Int. J. Wildland Fire, 19, 253-271, https://doi.org/10.1071/WF09002, 2010.
- 1367
- 1368 Yokelson, R. J., Bertschi, I. T., Christian, T. J., Hobbs, P. V., Ward, D. E., and Hao, W. M.: Trace gas
- 1369 measurements in nascent, aged, and cloud-processed smoke from African savanna fires by airborne
- 1370 Fourier transform infrared spectroscopy (AFTIR), J. Geophys. Res.-Atmos, 108, 8478,
- 1371 https://doi.org/10.1029/2002JD002322, 2003.
- 1372
- 1373 Yokelson, R. J., Burling, I. R., Gilman, J. B., Warneke, C., Stockwell, C. E., de Gouw, J., Akagi, S. K.,
- 1374 Urbanski, S. P., Veres, P., Roberts, J. M., Kuster, W. C., Reardon, J., Griffith, D. W. T., Johnson, T. J.,
- Hosseini, S., Miller, J. W., Cocker, D. R., Jung, H., and Weise, D. R.: Coupling field and laboratory
  measurements to estimate the emission factors of identified and unidentified trace gases for prescribed
- 1377 fires, Atmos. Chem. Phys., 13, 89-116, https://doi.org/10.5194/acp-13-89-2013, 2013.
- 1378
- 1379 Yokelson, R. J., Crounse, J. D., DeCarlo, P. F., Karl, T., Urbanski, S., Atlas, E., Campos, T., Shinozuka,
- 1380 Y., Kapustin, V., Clarke, A. D., Weinheimer, A., Knapp, D. J., Montzka, D. D., Holloway, J., Weibring,
- 1381 P., Flocke, F., Zheng, W., Toohey, D., Wennberg, P. O., Wiedinmyer, C., Mauldin, L., Fried, A., Richter,
- 1382 D., Walega, J., Jimenez, J. L., Adachi, K., Buseck, P. R., Hall, S. R., and Shetter, R.: Emissions from
- biomass burning in the Yucatan, Atmos. Chem. Phys., 9, 5785-5812, https://doi.org/10.5194/acp-9-57852009, 2009.
- 1385
- Yokelson, R. J., Goode, J. G., Ward, D. E., Susott, R. A., Babbitt, R. E., Wade, D. D., Bertschi, I.,
  Griffith, D. W. T., and Hao, W. M.: Emissions of formaldehyde, acetic acid, methanol, and other trace
- 1388 gases from biomass fires in North Carolina measured by airborne Fourier transform infrared
- spectroscopy, J. Geophys. Res.-Atmos, 104, 30109-30125, https://doi.org/10.1029/1999JD900817, 1999.
- 1391Yokelson, R. J., Griffith, D. W. T., and Ward, D. E.: Open-path Fourier transform infrared studies of1392large-scale laboratory biomass fires, J. Geophys. Res.-Atmos, 101, 21067-21080,
- 1393 https://doi.org/10.1029/96JD01800, 1996.
- 1394
- Yokelson, R. J., Karl, T., Artaxo, P., Blake, D. R., Christian, T. J., Griffith, D. W. T., Guenther, A., and
  Hao, W. M.: The tropical forest and fire emissions experiment: Overview and airborne fire emission
  factor measurements, Atmos. Chem. Phys., 7, 5175-5196, https://doi.org/10.5194/acp-7-5175-2007, 2007.
- 1398
- Yokelson, R. J., Susott, R., Ward, D. E., Reardon, J., and Griffith, D. W. T.: Emissions from smoldering
  combustion of biomass measured by open-path Fourier transform infrared spectroscopy, J. Geophys.
  Res.-Atmos, 102, 18865-18877, https://doi.org/10.1029/97JD00852, 1997.
- 1402
- Yu, S.: Role of organic acids (formic, acetic, pyruvic and oxalic) in the formation of cloud condensation
  nuclei (CCN): a review, Atmos. Res., 53, 185-217, https://doi.org/10.1016/S0169-8095(00)00037-5,
  2000.
- 1406
- Yu, P., Toon, O. B., Bardeen, C. G., Zhu, Y., Rosenlof, K. H., Portmann, R. W., Thornberry, T. D., Gao,
  R.-S., Davis, S. M., Wolf, E. T., Gouw, J. d., Peterson, D. A., Fromm, M. D., and Robock, A.: Black
  carbon lofts wildfire smoke high into the stratosphere to form a persistent plume, Science, 365, 587-590,
  https://doi.org/10.1126/science.aax1748, 2019.
- 1412 Zhang, R., Suh, I., Zhao, J., Zhang, D., Fortner, E. C., Tie, X., Molina, L. T., and Molina, M. J.:
- 1413 Atmospheric new particle formation enhanced by organic acids, Science, 304, 1487-1490,
- 1414 https://doi.org/10.1126/science.1095139, 2004.
- 1415



Figure 1. Corrected reflectance satellite image from the VIIRS spectroradiometer on the SuomiNPP and NOAA-20 satellites taken on June 25, 2018. The fire hot spots for the wildfire of

1419 interest are indicated by the red dots. Flight tracks were flown at Lagrangian distances

1420 downwind of the wildfire. Multiple transects at varying altitudes perpendicular to the plume

1421 direction formed 5 virtual screens. Plume direction of travel is indicated by the red arrow. The

1422 location of the Alberta oil sands mining facilities are shown in white.



Figure 2. Flight tracks coloured by CO mixing ratio (ppmv) for Screens 1 to 4. The two plumes are identified as south plume (SP) and north plume (NP). The fire perimeter surrounding the detected MODIS-derived 'hot spots' on June 25, 2018 is shown in the green hatched area. The

source of the NP is expected to be the same hot spots as the SP but ~ 30 min older; see Sect. 2.2.

1430 The small blue arrows along the flight tracks indicate the aircraft measured wind direction with

the average wind direction depicted with the large grey arrow. Distances between screens are

- shown in the grey arrows.
- 1433



1435 Figure 3. Time series of NMOGs (ppmv), acetonitrile (C<sub>2</sub>H<sub>3</sub>N; ppbv) and CO (ppmv), as well as OA concentrations (µg m<sup>-3</sup>) and altitude for Screen 1. The in-plume portions are indicated by the vertical grey bars. The aircraft flew back and forth across the plumes at increasing altitudes to complete five transects; a transect represents one pass across the SP and NP at the same altitude. 





1442 Figure 4. Background-subtracted average Screen 1 in-plume mixing ratios of measured gas- and

1443particle-phase N-containing species  $(N_r)$  and their fractional contribution to the total summed  $N_r$ 1444species. The  $N_r$  species are grouped into categories of reduced inorganics, reduced organics,1445oxidized inorganics and oxidized organics with reduced species in shades of red and oxidized1446species in shades of blue.



1456 Figure 5. Background-subtracted average mixing ratios of individually measured NMOGs

shown for thirteen chemical classes. In some cases, compounds are double- (or triple-) counted if

they can be identified in more than one category. For example, phenol is an alcohol + an

aromatic; guaiacol is an alcohol + an ether + an aromatic. In the pie chart, the *Other* category

1460 includes amides, amines, ethers, thiols and sulfides. The unidentified category contains

1461 molecular formulas detected, but the compound(s) could not be identified.



Figure 6. Total carbon (TC) partitioning based on EFs (carbon fraction). The bar chart shows
the stacked EFs for carbon-containing compounds with the middle pie chart showing their
percent contributions to the TC. The two magnified pie charts (right side), representing the low and
high I/SVOC EF estimates, show the percent breakdown of the measured NMOGs and the
remaining unidentified portion. The EF values (g C kg<sup>-1</sup>) are provided in the box. Note that all the

1468 EFs shown in Table A1 were converted to  $g C kg^{-1}$  for this breakdown



Figure 7. Fraction of total ΣNMOG emissions in each volatility bin, as well as the bin-averaged
O/C ratio spanning VOCs, IVOCs and SVOCs. Data is included from PTRMS, CIMS, AWAS
and cartridge measurements. The O/C ratio is derived for only the PTRMS, CIMS and AWAS
measurements and the errors bars indicate the standard deviation of the average O/C ratio.







Figure 8. Average emission factors (g kg<sup>-1</sup>) of a) particle species; b) inorganic gas-phase species, and c)

1481 the top 25 measured gas-phase organic species. C4 acids = C4 oxo-carboxylic acids; propadiene =

1482 fragments/propadiene; hydroxy acetone = hydroxy acetone/ ethyl formate. Organic species

1483 measurements are from the PTRMS, CIMS and AWAS.



Figure 9. Comparison of averaged emission factors with a) boreal forest field-based
measurements (Andreae, 2019; Akagi et al., 2011; Liu et al., 2017), b) laboratory-based
measurements of lodgepole pine (Koss et al., 2018), c) temperate forest field-based
measurements (Permar et al., 2021), and d) those used in CFFEPS (Urbanski et al., 2014). See

- 1489 Table S8 for compound comparisons that don't have exact matches.



Figure 10. Fire radiative power (FRP; in MW) from GOES-R (grey dots) and emissions from 1497 1498 the CFFEPS model (orange dots) from 2018-06-24 17:00 UTC to 2018-06-25 21:00 UTC. Local 1499 time = UTC - 6 hrs. Aircraft-derived emission rates are shown for a) PM<sub>1</sub>, b) CO, c) NO<sub>x</sub> (as NO) and d) NH<sub>3</sub> (in t h<sup>-1</sup>; red arrow) at 15:00 UTC when the aircraft flew closest to the fire. The 1500 1501 corresponding TROPOMI satellite-derived emission rates are also shown (in t h<sup>-1</sup>; black arrows). Note, the aircraft flight time occurred when the fire intensity reflected a surface, smoldering fire 1502 1503 and the satellite overpass time took place when the fire had transitioned to a crown (flaming) 1504 fire. The smoldering and flaming time periods are coloured in blue and pink, respectively. 1505

# 1506 **Table A1**

# 1507

1508 Summary of in-plume and background average mixing ratios (or concentrations), emission

- 1509 factors  $(g kg^{-1})$  (EF) and emission ratios (ppbv ppmv<sup>-1</sup> except CO<sub>2</sub> which is in units of ppmv 1510 ppmv<sup>-1</sup>; particulates in µg m<sup>-3</sup> ppm<sup>-1</sup> and GEM in ng m<sup>-3</sup> ppmv<sup>-1</sup>) (ER) for the SP, NP, and the EF
- 1510 ppinv ; particulates in µg in <sup>2</sup> ppin<sup>2</sup> and GEW in ig in <sup>2</sup> ppinv<sup>2</sup>) (EK) for the SP, NP, and the Er 1511 average of the two plumes. In-plume and background averages are in units of ug m<sup>-3</sup> for
- 1512 particulates, ppbv for gas-phase compounds, except GEM which is ng  $m^{-3}$ , and CO<sub>2</sub> is ppmv.
- 1513 Compounds are grouped by particulate species, and inorganic and organic gas-phase species and
- 1514 sorted by increasing molecular weight. PM<sub>1</sub> is the sum of the AMS-derived particulate species.
- 1515 The CE was 0.84±0.04 and 0.82±0.01 for the SP and NP, respectively. For comparison, EFs are
- 1516 also included from previously published literature including Andreae  $(2019)^1$ , Koss et al.
- 1517  $(2018)^2$ , Permar et al.  $(2021)^3$ , and Liu et al.  $(2017)^4$ . The Andreae (2019) PM EF represents
- 1518 PM<sub>2.5</sub>. See Table S8 for compounds that did not have exact matches for comparison to literature 1519 values. To derive the EF for species measured in mass concentration units, Eq. 3 was modified
- 1517 values. To derive the EF for species measured in mass concentration units, Eq. 5 was modified 1520 by converting TC to mass concentrations using the measured temperature and pressure, and
- removing the molar mass ratio term. \* Indicates that the compound was 'calculated' (SI Sect
- 1522 1.1.1) while the remaining compounds were calibrated. Uncertainties were estimated by
- 1522 1.1.1) while the remaining compounds were canorated. Uncertainties were estimated by 1523 summing in quadrature the standard error of the average EF (or ER) and the measurement
- 1524 uncertainties (see Sect. 2.5).
- 1525 1526

Molecu-	Comp-	Compou	Instru-	SP	NP	Backgr-	Average	NP EF (g	SP EF (g	Literature	NP ER	SP ER
lar	ound	nd Name	ment	Averag	Average	ound	EF (g kg-1)	kg⁻¹)	kg⁻¹)	EF		
Weight				е						(g kg⁻¹)		
Particulate	S										I	
		particulat										
		e matter			75.5±29.					18.7±15.9 <sup>1</sup>	58.8±1.	
	PM <sub>1</sub>	(<1µm)	AMS	112±35	3	13.2±0.9	6.8±0.8	7.1±0.3	6.6±1.1	26.0±6.24	0	65.1±7.3
		black			0.74±0.3	0.11±0.0	0.13±0.0	0.11±0.0	0.14±0.0	$0.43 \pm 0.21^{1}$	0.55±0.	
	BC	carbon	SP2	1.3±0.4	0	6	3	2	4	0.39±0.17 <sup>3</sup>	08	0.58±0.19
		p-										
		ammoniu				0.21±0.0	0.11±0.0	0.11±0.0	0.12±0.0			
	NH <sub>4</sub>	m	AMS	2.3±1.0	1.2±0.4	3	3	3	4	0.34±0.15 <sup>4</sup>	1.1±0.3	1.9±0.6
						0.078±0.	0.17±0.0	0.14±0.0	0.19±0.0		0.90±0.	
	NO₃	p-nitrate	AMS	3.2±1.5	1.4±0.5	017	4	3	5	0.87±0.13 <sup>4</sup>	16	1.2±0.3
		p-			0.98±0.3	0.39±0.0	0.060±0.	0.066±0.	0.055±0.		0.035±	
	SO <sub>4</sub>	sulphate	AMS	1.7±0.7	1	3	022	020	023	0.30±0.16 <sup>4</sup>	0.011	0.054±0.020
		p-total			72.6±27.	12.5±0.8					57.5±1	
	OA	organics	AMS	101±34	0	3	6.6±2.6	6.9±2.4	6.3±2.8	24.3±0.21 <sup>4</sup>	9.4	61.7±27.1
Gas												
Inorganic	1	r	1	1	1	T	T	r	Т	Т	1	1
				45 410		-	0.6210.4	0.45.0.0	0.02+0.4	2 5 14 01		
17 021	NUL			15.4±9.	F 2 1 2 1	0.039±2.	0.63±0.1	0.45±0.0	0.82±0.1	$2.5\pm1.8^{-1}$	5 0 0 0	12 612 5
17.031	NH <sub>3</sub>	ammonia	LGK	6	5.2±2.1	2	4	4	9	0.68±0.19 <sup>2</sup>	5.8±0.6	12.0±2.5
										$0.53\pm0.30^{-1}$		
		hudrogon				0 19+0 0	0.21+0.0	0.24+0.0	0.20+0.0	0.28±0.060		
27 026		nyurogen	CIMS	2 2+1 6	2 2+1 0	0.18±0.0	0.31±0.0	0.34±0.0	0.29±0.0	0 42+0 173	2 8+0 5	2 9+0 7
27.020	TICN	cyanide	CIIVIS	5.211.0	2.3±1.0	/	/	0	· /	0.45±0.17	2.810.5	2.9±0.7
		monovid		991+44						121+47 <sup>1</sup>		
28.01	co	e	Picarro	3	819+327	119+5	116+6	127+4	104+7	99.3+19.73	108+39	126+52
20.01		nitric	. 100110	0 14+0	515_527	0.0014+0	0.016+0	0.016+0	0 14+0 0	0.291	100100	120252
30.006	NO	oxide	TECOs	05		.086	006	006	5	0.25		0.14±0.046
111000		isocvanic		0.52±0.	0.44±0.1	0.068±0.	0.083±0.	0.091±0.	- 0.076±0.	0.57±0.24 <sup>2</sup>	0.46±0.	
43.025	HNCO	acid	CIMS	17	3	024	029	027	031	$0.16\pm0.04^{3}$	13	0.47±0.92

Molecu- lar Weight	Comp- ound	Compou nd Name	Instru- ment	SP Averag e	NP Average	Backgr- ound	Average EF (g kg <sup>-1</sup> )	NP EF (g kg <sup>-1</sup> )	SP EF (g kg <sup>-1</sup> )	Literature EF (g kg <sup>-1</sup> )	NP ER	SP ER
		carbon		414±0.				1481±10		1529±135 <sup>1</sup>		
44.009	CO <sub>2</sub>	dioxide	Picarro	4	411±0.2	405±0.4	1496±92	3	1511±80	1413±61 <sup>3</sup>	7.4±0.5	9.4±0.45
46.005	NO <sub>2</sub>	nitrogen dioxide	TECOs	0.88±0. 17		0.39±0.1 9	0.15±0.0 4	0.15±0.0 4	0.88±0.1 7	1.0 <sup>1</sup>		0.83±0.21
46.005	NOx	sum (NO+NO <sub>2</sub> )	TECOs	1.0±0.2		0.39±0.2 0	0.17±0.0 4	0.17±0.0 4	1.0±0.2	1.2±0.9 <sup>1</sup>		0.97±0.58
47.013	ноло	nitrous acid	CIMS	0.22±0. 04		0.098±0. 038	0.020±0. 012	0.020±0. 012	0.22±0.0 4	0.60+0.20 <sup>2</sup>		0.11+0.061
64.064	50.	sulphur	TECOS	1 2+0 2		0.19±0.4	0.26±0.0	0.26±0.0	1 2+0 2	0.22+0.211		1 1+0 16
04.004	302	uloxide	TLCOS	1.5±0.5		0	5	5	1.5±0.5	0.2210.31	0.0008	1.110.10
79.011	HNO₄	pernitric acid	CIMS	0.036±	0.032±0. 0043	0.020±0. 007	0.0010±0	0.00047±	0.00085±		9±0.00	0.0028±0.003
75.011	11104	gaseous	CINIS	0.0015	0013	007	0.000087	0.000082	0.000092		0.0006	3
		elementa					±0.00001	±0.00001	±0.00001	0.00023±0.	8±0.00	0.00091±0.00
200.59	GEM	I mercury	Tekran	1.6±0.2	1.4±0.1	1.2±0.03	7	7	6	00030 <sup>1</sup>	014	014
Gas Organic												
			PTRMS									
			+CIMS									
		non	+AWAS				24 5 14 6					
	ΣΝΜΟ	organic	+ cartrid				24.5±1.6					
	G	gases	ges				25.6±1.6	26.2±2.1	25.4±5.8			
			PTRMS									
	Estim.		+CIMS									
	NMOG	non	+AWAS				26 8+11			59 71		
	T(See Sect	organic	cartrid				3 to			$25.0^2$		
	3.4.3)	gases	ges				39.9±5.8			26.1 <sup>3</sup>		
		carbon										
	NMOG	fraction		936±34							680±11	
	т	of NMOG	Picarro	1	649±225	375±85	31.2±3.8	36.8±5.1	25.5±5.6	E E+2 E1	1	580±92
16.043	CH₄	methane	Picarro	4	1982±35	1911±8	8.3±0.9	7.8±0.4	8.7±1.1	5.9±1.8 <sup>3</sup>	107±5	146±16
				-		0.34±0.0	0.27±0.0	0.20±0.0	0.34±0.1			
26.038	$C_2H_2$	acetylene	AWAS	6.9±2.5	1.8±0.7	041	8	5	1	0.31±0.17 <sup>3</sup>	2.2±0.9	4.0±1.1
				32.4±1		0.64±0.0					12.9±3.	
28.054	C <sub>2</sub> H <sub>4</sub>	ethene	AWAS	4.0	9.3±3.8	1	1.5±0.4	1.3±0.3	1.7±0.5	$1.5\pm1.0^{3}$	5	18.3±5.0
		formalde		13.9+4.					0.93+0.3	$1.8\pm0.4^{-1}$ $1.9\pm0.7^{2}$		
30.026	CH <sub>2</sub> O	hyde	PTR	9	10.1±2.8	4.4±2.0	1.0±0.3	1.1±0.3	6	$1.9\pm0.4^3$	8.1±2.2	8.9±3.2
				27.1±1		1.9±0.01					12.5±2.	
30.07	C <sub>2</sub> H <sub>6</sub>	ethane	AWAS	2.4	10.1±3.4	6	1.3±0.4	1.3±0.3	1.4±0.5	1.1±0.84 <sup>3</sup>	7	13.8±3.8
				21 0+7						$2.3\pm1.0^{1}$	14 0+2	
32.042	CH₄O	methanol	PTR	21.9±7. 9	15.9±4.9	6.8±0.9	1.9±0.4	2.2±0.4	1.6±0.4	$1.5\pm0.4^3$	14.9±2. 9	13.4±3.6
										0.060±0.03		
		fragment								01		
		s/propadi				0.39±0.3	0.53±0.1	0.64±0.1	0.42±0.1	0.088±0.04		
40.065	C <sub>3</sub> H <sub>4</sub>	ene*	PTR	3.7±1.6	3.0±1.3	5	1	1	1	$1^{2}$	3.5±0.6	2.8±0.7
						0.4010.5	0.000	0.4010.0	0.4010.5	0.31±0.10 <sup>2</sup> 0.086±0.02		
41.050		acetonitri	DTP	20+14	2 0+0 9	0.10±0.0	0.44±0.0	0.48±0.0	0.40±0.0	74	2 6+0 2	2 6+0 5
41.053	C2H3N	ie cvanamid	FIK	2.8±1.4	2.0±0.8 0.40+0.1	ס 0 10+0 2	/ 0.064+0	0 0.067+0	o 0.061+0	0.31±0.15	∠.0±0.3	2.0±0.5
42.041	$CH_2N_2$	e*	PTR	22	2	0	042	042	042		1.4±0.9	1.3±0.9
						0.12±0.0	0.68±0.2	0.62±0.1	0.73±0.2			
42.081	C <sub>3</sub> H <sub>6</sub>	propene	AWAS	9.7±4.8	2.9±1.3	1	1	5	5	0.74±0.62 <sup>3</sup>	4.2±1.0	5.2±1.5

Molecu- lar	Comp- ound	Compou nd Name	Instru- ment	SP Averag	NP Average	Backgr- ound	Average EF (g kg <sup>-1</sup> )	NP EF (g kg <sup>-1</sup> )	SP EF (g kg <sup>-1</sup> )	Literature EF	NP ER	SP ER
Weight				e						(g kg <sup>-1</sup> )		
		acetalde				0.96±0.2				$0.81\pm0.23^{1}$ $0.92\pm0.32^{2}$		
44.053	$C_2H_4O$	hyde	PTR	7.3±3.2	5.2±2.0	0	1.1±0.2	1.2±0.2	1.0±0.2	1.7±0.4 <sup>3</sup>	6.3±0.9	6.3±1.2
44.097	C₃H <sub>8</sub>	propane	AWAS	6.7±3.2	2.5±1.0	0.28±0.0 3	0.52±0.1 4	0.53±0.1 2	0.50±0.1 6	0.46±0.18 <sup>3</sup>	3.4±0.7	3.4±1.0
		formic					0.17±0.0	0.17±0.0	0.17±0.0	1.0±0.9 <sup>1</sup> 0.28±0.14 <sup>2</sup>		
46.025	CH <sub>2</sub> O <sub>2</sub>	acid	CIMS	3.0±1.2	2.8±0.9	2.4±0.1	5	4	6	1.5±0.6 <sup>3</sup>	1.2±2.6	0.56±1.3
48.103	CH₄S	thiol*	PTR	0.074± 0.039	0.049±0. 022	0.0024±0 .028	0.014±0. 0086	0.015±0. 009	0.013±0. 008	0.011±0.00 6 <sup>2</sup>	0.068± 0.043	0.073±0.043
50.057	CH <sub>6</sub> O <sub>2</sub>	methanol hydrate*	PTR	0.25±0. 10	0.16±0.0 6	0.062±0. 057	0.028±0. 020	0.034±0. 023	0.022±0. 017		0.15±0. 10	0.12±0.09
52.070	<u></u>	buten- yne/frag	DTD	0.11±0.	0.080±0.	0.011±0.	0.018±0.	0.020±0.	0.016±0.	0.052±0.01	0.086±	0.004+0.050
52.076	C4H4	ments*	PIK	05	034	044	010	011	010	8 <sup>3</sup> 0.025±0.01	0.046	0.081±0.050
E2 064	C-H-N	acrylonitr	DTD	0.17±0.	0.12±0.0	0.0024±0	0.036±0.	0.040±0.	0.032±0.	2 <sup>2</sup> 0.044±0.01	0.17±0.	0 16+0 00
55.004	C3H3N	iie	FIN	08	0	-	018	018	018	0.034±0.01 4 <sup>2</sup>	07	0.1010.09
54.048	C <sub>3</sub> H <sub>2</sub> O	propynal *	PTR	0.053± 0.033	0.023±0. 0054	0.013±0. 019	0.0087±0 .0053	0.0045±0 .0031	0.013±0. 007	0.037±0.01 5 <sup>3</sup>	0.018± 0.013	0.062±0.032
		butadien				-				0.089±0.03 0 <sup>1</sup>		
54.092	C <sub>4</sub> H <sub>6</sub>	e/fragme nts*	PTR	0.74±0. 38	0.47±0.2 4	0.070±0. 17	0.15±0.0 8	0.15±0.0 8	0.15±0.0 8	$\begin{array}{c} 0.34 {\pm} 0.18^2 \\ 0.27 {\pm} 0.10^3 \end{array}$	0.62±0. 30	0.73±0.37
		1,3-		0.7410	0.2010.0	0.004110	0.005+0	0.055+0	0.075+0	0.089±0.03 0 <sup>1</sup>	0.2010	
54.092	$C_4H_6$	e	AWAS	0.74±0. 38	0.20±0.0 9	.00041±0	0.065±0. 022	0.055±0. 016	0.075±0. 026	$0.34\pm0.18^2$ $0.27\pm0.10^3$	0.29±0. 09	0.41±0.12
										0.012±0.00 5 <sup>2</sup>		
55.08	C₃H₅N	propane nitrile*	PTR	0.11±0. 05	0.080±0. 032	0.0097±0 .019	0.022±0. 012	0.025±0. 012	0.019±0. 012	0.037±0.01 8 <sup>3</sup>	0.10±0. 05	0.094±0.057
						0.17±0.0	0.28±0.0	0.29±0.0	0.26±0.0	0.34 <sup>1</sup> 0.97±0.50 <sup>2</sup>	0.82±0.	
56.064	C <sub>3</sub> H <sub>4</sub> O	cis-2-	PTR	1.5±0.6 0.16±0.	1.0±0.4	9 0.016±0.	5 0.015±0.	4	6 0.015±0.	0.40±0.18 <sup>3</sup>	12	0.83±0.15
56.108	C <sub>4</sub> H <sub>8</sub>	butene isobuten	AWAS	08	0.34+0.1	008	006	0.082+0.	006		0.41+0.	0.078±0.023
56.108	$C_4H_8$	e	AWAS	49	2	0022	023	008	032		03	0.45±0.13
56.108	C <sub>4</sub> H <sub>8</sub>	t-2- butene	AWAS	0.13±0. 07		0.010±0. 003	0.012±0. 005		0.012±0. 005			0.063±0.018
56.108	$C_4H_8$	1-butene	AWAS	1.4±0.7	0.41±0.1 7	0.014±0. 005	0.13±0.0 3	0.12±0.0 1	0.14±0.0 4		0.60±0. 05	0.74±0.12
	C₂H₃N	hydroxy acetonitri		0.021±	0.0078±0	0.00035±	0.0035±0	0.0025±0	0.0044±0	0.033±0.00	0.0095	
57.052	0	le methyl	CIMS	0.029	.014	0.00014	.0031	.0028	.0034	9 <sup>3</sup>	±0.011	0.021±0.016
57.052	C₂H₃N O	isocyanat e*	PTR	0.074± 0.029		0.0067±0 .06	0.0052±0 .0032		0.0052±0 .0032	0.033±0.00 9 <sup>3</sup>		0.024±0.015
							0.82±0.2	0.99±0.2	0.65±0.1	1.6±1.6 <sup>1</sup> 0.34±0.12 <sup>2</sup>	0.065±	
58.08	C₃H <sub>6</sub> O	acetone	PTR	6.0±1.8	4.7±1.2	2.5±0.3	2	5	9	0.84±0.22 <sup>3</sup> 0.11±0.06 <sup>1</sup>	0.018	0.072±0.028
58.124	C4H10	n-butane	AWAS	1.5±0.7	0.62±0.2 2	0.098±0. 013	0.15±0.0 4	0.16±0.0 4	0.14±0.0 5	0.12±0.06 <sup>3</sup>	0.79±0. 17	0.73±0.20
60.052	$C_2H_4O_2$	acetic acid	CIMS	8.8±7.5	6.0±3.9	2.1±0.8	1.3±0.8	1.1±0.5	1.6±0.9	3.8±2.0 <sup>1</sup> 2.4±0.6 <sup>3</sup>	7.4±3.5	8.9±5.2

Molecu- lar Weight	Comp- ound	Compou nd Name	Instru- ment	SP Averag e	NP Average	Backgr- ound	Average EF (g kg <sup>-1</sup> )	NP EF (g kg <sup>-1</sup> )	SP EF (g kg <sup>-1</sup> )	Literature EF (g kg <sup>-1</sup> )	NP ER	SP ER
	CH <sub>4</sub> N <sub>2</sub>			0.44±0.	0.28±0.0	0.067±0.	0.078±0.	0.079±0.	0.076±0.		0.29±0.	
60.056	0	Urea*	PTR	18	8	13	052	054	049	0.07410.02	20	0.34±0.21
										0.074±0.03 0 <sup>2</sup>		
61 04	CH₃NO	nitromet	PTR	0.055±	0.038±0.	0.0051±0	0.011±0.	0.010±0.	0.011±0.	0.078±0.00 93	0.036±	0.048+0.030
01.01	2	ethylene		0.023±	020	0.0036±0	0.0036±0		0.0036±0	5	0.021	0.01020.000
62.068	C <sub>2</sub> H <sub>6</sub> O <sub>2</sub>	glycol*	PTR	0.0077		.018	.0023		.0023	0.0016±0.0		0.015±0.010
										008 <sup>2</sup>		
		dimethyl		0.051±		0.011±0.	0.0067±0		0.0067±0	0.080±0.08 3 <sup>3</sup>		
62.13	C <sub>2</sub> H <sub>6</sub> S	sulfide	PTR	0.022		034	.0047		.0047	0.00474		0.029±0.020
		cyclopent		0.13±0.	0.12±0.0	0.025±0.	0.032±0.	0.041±0.	0.023±0.	5 <sup>3</sup>	0.14±0.	
66.103	C₅H <sub>6</sub>	andiene*	PTR	05	4	039	019	022	016	0.054±0.02	07	0.096±0.064
				0.4010	0.06710	-	0.02610	0.00710	0.025+0	9 <sup>2</sup>	0.000.	
67.091	C₄H₅N	pyrrole*	PTR	0.10±0. 06	0.067±0. 033	0.00073± 0.013	0.026±0. 014	0.027±0. 014	0.025±0. 014	0.039±0.02 1 <sup>3</sup>	0.090± 0.046	0.098±0.055
						- 0.0083+0	0 39+0 1	0 43+0 1	0 35+0 2	0.36±0.44 <sup>1</sup> 0.36±0.11 <sup>2</sup>		
68.075	C4H4O	furan*	PTR	1.5±0.8	1.1±0.5	.035	9	7	0	0.43±0.19 <sup>3</sup>	1.4±0.6	1.4±0.8
										0.074 <sup>1</sup> 0.22±0.11 <sup>2</sup>		
C0 110	C 11		DTD	1 7 0 7	1.010.0	0.52±0.2	0.42±0.2	0.64±0.3	0.19±0.1	0.082±0.09	2 1 1 1 1	0 4710 47
68.119	C5H8	Isoprene	PIR	1./±0./	1.8±0.8	5	6	4	5	0.074 <sup>1</sup>	2.1±1.1	0.4/±0.4/
				0.82+0	0 35+1 6		0 30+0 1	0.40+0.1	0 20+0 1	$0.22\pm0.11^{2}$		
68.119	C₅H <sub>8</sub>	isoprene	AWAS	46	5	1.3±0.4	8	7	9	5 <sup>3</sup>	2.0±1.0	0.18±0.18
69.083	C₄H₅O	*	PTR	0.017± 0.0064	0.013±0. 0063	0.00079± 0.01	0.0043±0 .0025	0.0047±0 .0028	0.0038±0 .0022		0.015± 0.009	0.015±0.008
										0.011±0.00		
		butane		0.030±	0.022±0.	0.0041±0	0.0077±0	0.0088±0	0.0065±0	5 <sup>2</sup> 0.020±0.01	0.028±	
69.107	C <sub>4</sub> H <sub>7</sub> N	nitrile*	PTR	0.014	009	.011	.0051	.0059	.0042	0 <sup>3</sup>	0.019	0.025±0.016
		methacro										
		lein, crotonald				0.91±0.1	0.19±0.0	0.20±0.0	0.18±0.1	$0.11\pm0.12^{1}$ $0.34\pm0.15^{2}$	0.66±0.	
70.091	$C_4H_6O$	ehyde	PTR	1.3±0.3	1.3±0.3	7	9	8	1	0.39±0.15 <sup>3</sup>	26	0.68±0.38
										0.046±0.02 5 <sup>1</sup>		
70 135	CrHao	c-2-	۵۱۸/۵۵	0.040±	0.013±0.	0.0040±0	0.0040±0	0.0033±0	0.0048±0	0.015±0.00	0.013±	0 021+0 0057
70.135	051110	pentene	AWAS	0.017	0034	.0000	.0012	.0007	.0010	0.046±0.02	0.004	0.02110.0037
		cyclopent		0.031+	0.015+0.	0.0052+0	0.0035+0	0.0038+0	0.0031+0	5 <sup>1</sup> 0.015+0.00	0.016+	
70.135	C <sub>5</sub> H <sub>10</sub>	ane	AWAS	0.013	0022	.0003	.0009	.0009	.0009	8 <sup>3</sup>	0.004	0.014±0.005
										0.046±0.02 5 <sup>1</sup>		
70 135	CrHao	1-	Διλ/Δς	0.42±0.	0.15±0.0	0.0053±0	0.052±0.	0.053±0.	0.052±0.	0.015±0.00 8 <sup>3</sup>	0.21±0.	0 22+0 06
70.135	05110	pentelle	AVVAJ	~ 1	0	.0012	015	015	010	0.046±0.02	05	0.22±0.00
		t-2-		0.13±0	0.068±0.	0.0094±0	0.018±0.	0.0049±0	0.031±0.	5 <sup>1</sup> 0.015±0.00	0.013±	
70.135	$C_5H_{10}$	pentene	AWAS	12	020	.0058	013	.0040	018	8 <sup>3</sup>	0.010	0.063±0.035
70.135	$C_5H_{10}$	2-me-1- butene	AWAS	0.12±0. 061	0.047±0. 018	0.0088±0 .0017	0.014±0. 005	0.014±0. 004	0.015±0. 005	0.046±0.02 5 <sup>1</sup>	0.056± 0.013	0.062±0.018

Molecu- lar	Comp- ound	Compou nd Name	Instru- ment	SP Averag	NP Average	Backgr- ound	Average EF (g kg <sup>-1</sup> )	NP EF (g kg <sup>-1</sup> )	SP EF (g kg <sup>-1</sup> )	Literature EF	NP ER	SP ER
Weight				е						(g kg <sup>-1</sup> )		
										0.015±0.00 8 <sup>3</sup>		
										0.046±0.02 5 <sup>1</sup>	0.0068	
70 135	CEH10	2-me-2- butene	Δ₩Δ	0.019±	0.0077±0 0016	0.0034±0	0.0019±0 0006	0.0017±0	0.0022±0	0.015±0.00 8 <sup>3</sup>	±0.003	0.0095±0.002 7
70.155	031110	butche	///////////////////////////////////////	0.0071						0.046±0.02	0	,
70 125	C II	3-me-1-	A)A/A C	0.10±0.	0.045±0.	0.029±0.	0.016±0.	0.0078±0	0.025±0.	0.015±0.00	0.045±	0.05.8±0.028
70.135	C51110	acrylic	CINC	0.28±0.	0.21±0.1	0.060±0.	0.096±0.	0.13±0.0	0.062±0.	0 22 4 0 003	0.25±0.	0.05810.028
72.063	C3H4U2	MEK, 2-	CIIVIS	24	5	046	048	б	035	0.22±0.08°	11	0.35±0.20
		methyl acetate,										
72.107	C <sub>4</sub> H <sub>8</sub> O	ethyl formate	PTR	0.80±0. 35	0.57±0.2 3	0.097±0. 051	0.18±0.0 7	0.22±0.0 8	0.14±0.0 6		0.67±0. 23	0.54±0.20
72.151	C5H12	n- pentane	AWAS	0.59±0. 28	0.26±0.0 9	0.035±0. 005	0.078±0. 021	0.086±0. 019	0.070±0. 023	0.057±0.02 8 <sup>3</sup>	0.34±0. 07	0.29±0.08
		2- methylbu		0.21±0.	0.11±0.0	0.051±0.	0.022±0.	0.024±0.	0.021±0.	0.057±0.02	0.097±	
72.151	C <sub>5</sub> H <sub>12</sub>	tane propanoi	AWAS	08 0.81±0.	1 0.70±0.1	001 0.49±0.1	006 0.13±0.0	0049 0.12±0.0	007 0.14±0.0	8 <sup>3</sup>	0.019 0.35±0.	0.086±0.027
74.079	$C_3H_6O_2$	c acid hvdroxv	CIMS	24	5	7	8	8	9	0.57±0.20 <sup>3</sup>	24	0.51±0.30
		acetone/ ethyl										
74 079	CaHeOa	formate *	PTR	1.5±0.6	1 1+0 4	0.30±0.1	0.32±0.2	0.35±0.2	0.30±0.2		1 0+0 8	1 1+0 7
74.075	0311602			1 4+0 6	1.110.4	0.054+0	0.41±0.0	0.47±0.0	0 36+0 0	0.57±0.21 <sup>1</sup>	1.010.0	1.110.7
78.114	C <sub>6</sub> H <sub>6</sub>	benzene	PTR	9 9	1.0±0.5	045	6	6	7	0.42±0.23 0.50±0.14 <sup>3</sup>	1.3±0.2	1.2±0.2
		andienon		0.040	0.02410	-	0.011+0	0.0000.0	0.012+0	0.007+0.01	0.0261	
80.086	C₅H₄O	s*	PTR	0.049± 0.026	0.024±0. 012	0.00054± 0.0061	0.011±0. 006	0.0093±0 .0047	0.012±0. 007	0.027±0.01 7 <sup>3</sup>	0.026± 0.013	0.040±0.023
		cyclohex adiene/m										
		onoterpe ne										
80.13	C <sub>6</sub> H <sub>8</sub>	fragment *	PTR	0.45±0. 18	0.40±0.2 0	0.040±0. 040	0.14±0.0 7	0.17±0.0 7	0.10±0.0 6		0.48±0. 19	0.34±0.20
		pentene nitriles/										
81.118	C₅H7N	methyl pyrrole*	PTR	0.018± 0.0093	0.013±0. 0053	0.0015±0 .0055	0.0047±0 .0032	0.0050±0 .0036	0.0044±0 .0028	0.020±0.01 1 <sup>3</sup>	0.014± 0.010	0.015±0.009
82.102	C₅H6O	methyl furan*	PTR	1.1±0.5	0.69±0.3 2	0.042±0. 065	0.30±0.1 5	0.31±0.1 3	0.29±0.1 7	0.32±0.11 <sup>2</sup> 0.28±0.13 <sup>3</sup>	0.84±0. 34	0.96±0.55
		cyclohex		0.14±0.	0.093±0.	0.015±0.	0.054±0.	0.075±0.	0.033±0.	0.015±0.01 1 <sup>3</sup>	0.20±0.	
82.146	C <sub>6</sub> H <sub>10</sub>	ene*	PTR	06 0.0096	044	035	030	038	020		10	0.11±0.07
83.09	C₄H₅N O	methylox azole*	PTR	±0.004 4		0.00012± 0.0083	0.0020±0 .0011		0.0020±0 .0011			0.0066±0.003 7
83.134	C5H9N	pentanen itriles*	PTR	0.049± 0.025	0.037±0. 017	0.0024±0 .0088	0.016±0. 008	0.019±0. 009	0.013±0. 008	0.021±0.01 1 <sup>3</sup>	0.049± 0.024	0.042±0.025
84.074	C4H4O2	*	CIMS	1.7±0.7	0.79±0.2 4	0.29±0.1 3	0.38±0.1 4	0.23±0.0 4	0.52±0.2 0	0.32±0.11 <sup>3</sup>	0.61±0. 20	1.3±0.6
		furanone		0 54+0	0 37+0 1	- - 0.0030+0	0 16+0 0	0 16+0 0	0 15+0 0		0.42+0	
84.074	$C_4H_4O_2$	*	ptr	25	6	.041	9	9	8		23	0.48±0.26

Molecu- lar Weight	Comp- ound	Compou nd Name	Instru- ment	SP Averag e	NP Average	Backgr- ound	Average EF (g kg <sup>-1</sup> )	NP EF (g kg <sup>-1</sup> )	SP EF (g kg <sup>-1</sup> )	Literature EF (g kg <sup>-1</sup> )	NP ER	SP ER
		cyclopent anone/is		0.23±0.	0.16±0.0	0.017±0.	0.069±0.	0.073±0.	0.065±0.	0.087±0.03	0.19±0.	0.01.0.10
84.118	C₅H <sub>8</sub> O	omers*	PTR	11	7	026	036	032	039	8 <sup>3</sup> 0.008+0.01	09	0.21±0.12
84.162	C <sub>6</sub> H <sub>12</sub>	hexene*	PTR	0.0231	0.02110.	.015	0.01510.	0.02010.	.0065	4 <sup>3</sup>	0.032	0.031±0.019
											0.0069	
94 162	Cellin	c-2-	A\A/AS	0.019±	0.0079±0	0.0031±0	0.0020±0	0.0021±0	0.0020±0	0.008±0.01	±0.004	0.0064±0.005
04.102	C6H12	nexene	AWAS	0.012	.0024	.0002	.0014	.0010	.0011	4	4 0.0064	5
84,162	CeH12	cyclohex ane	AWAS	0.022±	0.0089±0	0.0051±0	0.0022±0	0.0019±0	0.0026±0	0.008±0.01 4 <sup>3</sup>	±0.005	0.0097±0.005 8
0.1102	00.112	butanedi			10010	10000	10010	10010	10010			0
		one/isom		0.57±0.	0.39±0.1	0.089±0.	0.13±0.0	0.13±0.0	0.13±0.0		0.33±0.	
86.09	C <sub>4</sub> H <sub>6</sub> O <sub>2</sub>	ers	ptr	23	4	077	9	9	9		23	0.41±0.26
		pentanon		0.14±0.	0.11±0.0	0.013±0.	0.046±0.	0.053±0.	0.038±0.	0.062±0.02	±0.0093	0.0080±0.004
86.134	C5H10O	e	PTR	07	4	020	025	026	024	3 <sup>3</sup>	6	9
06 170	C II		A)A/AC	0.31±0.	0.13±0.0	0.013±0.	0.049±0.	0.053±0.	0.044±0.	0.050±0.03	0.17±0.	0.1610.07
86.178	C <sub>6</sub> H <sub>14</sub>	n-nexane	AWAS	14	5 0.013+0	0.0012	0.0031+0	019	0.0022+0	6 <sup>3</sup> 0.050+0.03	06	0.16±0.07
86.178	C <sub>6</sub> H <sub>14</sub>	2,3-DMB	AWAS	0.0091	001	.0001	.0012	0012	.0013	6 <sup>3</sup>	0.004	9
		2,3-										
96 179	C.H.	methylpe	A\A/AS	0.090±	0.026±0.	0.011±0.	0.010±0.	0.0089±0	0.011±0.	0.050±0.03	0.032±	0.020+0.021
80.178	C61114	Intalle	AWAS	0.047	004	0003	005	.0041	000	0.012±0.00	0.011	0.03910.021
										5 <sup>2</sup>		
88.063	C 11 O	pyruvic	CIME	1 1+2 1	2 4+1 0	2 2+0 2	0.72±0.7	0.56±0.5	0.89±0.8	0.019±0.00	0.022±	-
88.002	C3H4O3	methyl	CIIVIS	4.4±2.4	3.4±1.0	2.310.3	1	4	5	83	0.022	0.0025±0.010
		propanoa		0.24±0.	0.17±0.0	0.021±0.	0.070±0.	0.075±0.	0.065±0.		0.19±0.	
88.106	C <sub>4</sub> H <sub>8</sub> O <sub>2</sub>	te *	ptr	11	7	040	043	047	040		12	0.20±0.12
		oxathian		0.012+	0.0090+0	- 0.00061+	0.0031+0	0.0023+0	0.0040+0		0.0058	
88.168	C <sub>4</sub> H <sub>8</sub> OS	e*	PTR	0.0049	.0030	0.0073	.0024	.0021	.0026		3	0.012±0.008
				0.026±		0.0012±0	0.0074±0		0.0074±0			
90.125	C <sub>7</sub> H <sub>6</sub>	* diethyl	PTR	0.014		.018	.0011		.0011			0.022±0.013
		sulfide,										
		butaneth		0.31±0.	0.22±0.1	0.036±0.	0.077±0.	0.083±0.	0.071±0.		0.20±0.	
90.184	C <sub>4</sub> H <sub>10</sub> S	iol	PTR	13	0	12	015	014	015	0 25+0 111	04	0.21±0.05
				0.62±0.	0.48±0.2	0.034±0.	0.26±0.0	0.26±0.0	0.26±0.0	$0.35\pm0.11$ $0.25\pm0.13^{2}$	0.63±0.	
92.141	C <sub>7</sub> H <sub>8</sub>	toluene	PTR	30	1	037	7	4	9	0.42±0.16 <sup>3</sup>	08	0.71±0.23
	CILN			0.0071		0.0012+0	0.0025+0		0.0025+0			0.0070+0.004
93.082	C <sub>2</sub> H <sub>7</sub> N O <sub>3</sub>	*	PTR	±0.003 4		0.0012±0 .0046	0.0025±0 .0012		0.0025±0 .0012			0.0070±0.004 8
										0.0026±0.0		
				0.055	0.000.0	0.00000.	0.000.0	0.000.0	0.01010	010 <sup>2</sup>	0.050.	
93 085	C₅H₃N O	furancarb	PTR	0.056±	0.038±0. 017	0.00022±	0.020±0. 011	0.022±0. 010	0.018±0. 011	0.0088±0.0 037 <sup>3</sup>	0.053±	0.053+0.030
55.005	-			0.001		0.0010				3.0 <sup>1</sup>	0.025	0.000_0.000
				0.42±0.	0.27±0.1	0.0026±0	0.12±0.0	0.12±0.0	0.12±0.0	0.57±0.36 <sup>2</sup>	0.28±0.	
94.113	C <sub>6</sub> H <sub>6</sub> O	phenol*	PTR	0.080+	3	.030	6 0.021+0	5	7 0.020+0	0.33±0.13 <sup>3</sup>	11	0.35±0.20
94.157	C <sub>7</sub> H <sub>10</sub>	adiene*	PTR	0.035	028	020	011	011	012		0.025	0.056±0.033
				0.010		-	0.0000		0.0000	0.0024±0.0		
94 10	CaHeSa	dimethyl disulfide*	PTR	0.012±		0.0012±0	0.0039±0 .0022		0.0039±0 .0022	0092		0.011+0.006
54.15	0211652	alsaillae		0.0071		.012	.0022		.0022		0.0099	5.01120.000
				0.014±	0.0094±0	0.0012±0	0.0030±0	0.0043±0	0.0038±0		±0.006	
95.077	$C_5H_3O_2$	*	PTR	0.0085	.0043	.0070	.0021	.0028	.0012		3	0.011±0.007

Molecu- lar	Comp- ound	Compou nd Name	Instru- ment	SP Averag	NP Average	Backgr- ound	Average EF (g kg <sup>-1</sup> )	NP EF (g kg <sup>-1</sup> )	SP EF (g kg <sup>-1</sup> )	Literature EF	NP ER	SP ER
Weight				e					0,	(g kg <sup>-1</sup> )		
-	C-H-N	pyridinal		0.0066	0.0045±0	-	0.0022+0	0.0021+0	0.0023+0	0.0099±0.0	0.0048	0 0063+0 003
95.101	0	*	PTR	±0.002 6	.0026	0.000871	.0015	.0017	.0012	034	1 1	2
		C2		0.0068 ±0.003		- 0.000091	0.0021±0		0.0021±0			0.0060±0.003
95.145	C <sub>6</sub> H <sub>9</sub> N	pyrrole*	PTR	4		±0.0054	.0012		.0012	0.611		3
0.0005		с с 1 <b>ж</b>		2.0.1.0	1 2 2 2 5	0.0040±0	0.65±0.3	0.67±0.2	0.64±0.3	$0.54\pm0.17^{2}$	15.00	10110
96.085	C5H4O2	furfural*	PIR	2.0±1.0	1.3±0.6	029	1	6	6	0.53±0.213	1.5±0.6	1.8±1.0
96.129	C6H8O	C2- furan*	PTR	0.32±0. 16	0.20±0.0 9	0.00016± 0.024	0.087±0. 044	0.086±0. 037	0.087±0. 050	0.20±0.10 <sup>3</sup>	0.20±0. 09	0.24±0.14
06 172	C-H	cyclohept	DTD	0.042±	0.035±0.	0.0049±0	0.022±0.	0.033±0.	0.011±0.		0.076±	0.021+0.020
50.175	C/112	ene	FIN	0.02	013	.017	015	020	007		0.0098	0.03110.020
97.073	C <sub>4</sub> H <sub>3</sub> N O <sub>2</sub>	*	PTR	0.010± 0.005	0.0093±0 .0026	0.0012±0 .0075	0.0030±0 .0021	0.0044±0 .0027	0.0036±0 .0012		±0.007 3	0.0096±0.007 1
		hovanoni		0.011+	0.0077+0	0.00021+	0.0040+0	0.0041+0	0.0040+0	0.0088±0.0	0.0093	
97.161	$C_6H_{11}N$	trile*	PTR	0.0053	.0041	0.0040	.0026	.0028	.0023	047	3	0.011±0.006
		maleic anhydrid		0.21±0.	0.16±0.0	0.011±0.	0.070±0.	0.072±0.	0.068±0.	0.14±0.07 <sup>3</sup>	0.16±0.	
98.057	$C_4H_2O_3$	e* furan	PTR	08	5	032	036	031	040		07	0.18±0.11
		methanol		0.28+0	0.20+0.0	0.021+0	0.05.8±0	0.061+0	0.05.4+0	$0.38\pm0.15^{2}$	0.14+0	
98.101	$C_5H_6O_2$	*	PTR	0.28±0. 13	0.20±0.0 8	0.021±0. 047	0.058±0. 030	0.081±0. 025	0.054±0. 034	0.090±0.04 3 <sup>3</sup>	0.14±0. 06	0.15±0.09
		methyl cyclopent								0.022±0.00 9 <sup>2</sup>		
98.145	C <sub>6</sub> H <sub>10</sub> O	anone/is omers*	PTR	0.052± 0.023	0.036±0. 015	0.0027±0 .015	0.015±0. 008	0.017±0. 009	0.013±0. 008	0.034±0.01 5 <sup>2</sup>	0.038± 0.019	0.035±0.020
		unsaturat										
		carboxyli		0.20±0.	0.25±0.0	0.078±0.	0.072±0.	0.10±0.0	0.045±0.	0.11±0.04 <sup>3</sup>	0.22±0.	
100.117	C <sub>5</sub> H <sub>8</sub> O <sub>2</sub>	c acids methyl	CIMS	03	7	043	027	3	021		07	0.13±0.06
		methacry late/isom		0.15±0.	0.12±0.0	0.017±0.	0.036±0.	0.035±0.	0.037±0.		0.078±	
100.117	$C_5H_8O_2$	ers*	ptr	06	4	041	023	022	024	0.0046+0.0	0.049	0.098±0.062
		hexanal/								0.0046±0.0 029 <sup>2</sup>		
100.161	C <sub>6</sub> H <sub>12</sub> O	hexanon e*	PTR	0.022± 0.0077	0.018±0. 008	0.003±0. 010	0.0065±0 .0043	0.0074±0 .0049	0.0057±0 .0035	0.013±0.00 6 <sup>3</sup>	0.016± 0.011	0.015±0.009
		C4 oxo- carboxyli					0.74+0.3	0.57+0.3	0.92+0.4	0.044±0.02 0 <sup>3</sup>		
102.089	$C_4H_6O_3$	c acids	CIMS	4.7±1.2	3.8±0.7	2.3±0.2	7	0	3		1.2±0.7	2.4±1.1
		acetic anhydrid		0.033±	0.022±0.	0.0020±0	0.0075±0	0.0078±0	0.0072±0		0.017±	
102.089	C <sub>4</sub> H <sub>6</sub> O <sub>3</sub> C <sub>5</sub> H <sub>10</sub> O	e * valeric	ptr	0.016 0.083±	008 0.057±0.	.018 0.0037±0	.0046 0.024±0.	.0046 0.027±0.	.0045 0.020±0.		0.010 0.059±	0.019±0.012
102.133	2	acid*	ptr	0.042	025	.016	014	016	012		0.035	0.052±0.031
103.121	C4H9N O2	*	PTR	0.017± 0.0063	0.013±0. 005	.0032	.0020	.0026	0.0064±0 .0011		0.016± 0.009	0.016±0.010
										0.021±0.00 4 <sup>2</sup>		
103 124	C <sub>7</sub> H <sub>5</sub> N	benzonitr ile*	PTR	0.15±0. 08	0.11±0.0	0.00028±	0.060±0. 029	0.065±0. 026	0.054±0. 031	0.055±0.02	0.14±0. 06	0.14+0.08
103.124	0,150			0.0073	0.0050.0	-	0.0000.0	0.0000.00	0.000	-	0.0069	0.0000.0000
104.149	C5H12O	pentaned iol*	PTR	±0.003 4	0.0052±0 .0016	0.00087± 0.0057	0.0029±0 .0025	0.0033±0 .0033	0.0024±0 .0013		±0.006 9	0.0060±0.003 2

Molecu-	Comp-	Compou	Instru-	SP	NP	Backgr-	Average	NP EF (g	SP EF (g	Literature	NP ER	SP ER
Weight	ound	nu name	ment	e	Average	ound	EF (g Kg )	к <u>е</u> /	к <u>е</u> )	(g kg⁻¹)		
										0.088±0.05		
										6 <sup>2</sup>		
104 152	CoHo	styrene*	PTR	0.053±	0.041±0. 020	0.000051	0.039±0. 021	0.056±0. 026	0.022±0.	0.018±0.01 2 <sup>3</sup>	0.12±0.	0.058+0.033
104.152	C <sub>8</sub> H <sub>8</sub> C <sub>4</sub> H <sub>10</sub> O	Diethylen		0.027 0.014±	020	0.0016±0	0.0036±0	020	0.0036±0	2	00	0.0088±0.006
106.121	3	e glycol*	PTR	0.0035		.011	.0026		.0026			1
										0.095±0.05		
		benzalde		0.10±0.	0.079±0.	0.011±0.	0.036±0.	0.042±0.	0.030±0.	0.084±0.02	0.087±	
106.124	C <sub>7</sub> H <sub>6</sub> O	hyde*	PTR	05	037	015	019	018	019	6 <sup>3</sup>	0.038	0.077±0.047
		C8 aromatic		0 19+0	0 13+0 0	0.0082+0	0.075+0	0.082+0	0.068+0		0 17+0	
106.168	C <sub>8</sub> H <sub>10</sub>	S	PTR	0.1510.	6	.020	012	012	013	0.21±0.08 <sup>3</sup>	02	0.17±0.03
		pyridine		0.0035		-						
107 112	C <sub>6</sub> H₅N	aldehyde *	DTR	±0.002		0.00051±	0.0015±0		0.0015±0			0.0038±0.002
107.112	0	dimethyl		1		0.0050	.0008		.0008			0
		pryidine/		0.0048								
107 156	C <sub>7</sub> H <sub>0</sub> N	heptyl nitriles*	PTR	±0.001		0.000012	0.0018±0 001		0.0018±0 0010	0.0050±0.0 033 <sup>2</sup>		0.0043±0.002
107.150	C/HgN	intrites				10.0040	.001		.0010	0.084±0.02		
		benzoqui								4 <sup>2</sup>		
108.096	CcH4O2	none/qui	PTR	0.093±	0.061±0. 023	0.0025±0 019	0.025±0. 013	0.024±0.	0.025±0. 015	0.077±0.02	0.049±	0.062+0.035
100.050	C6114O2	methyl	111	0.045	025	.015	015	011	015	0	0.022	0.00210.035
		phenol/a										
108.14	C₂H₀O	nisol/cres	PTR	0.13±0. 07	0.083±0. 043	0.00068± 0.0094	0.040±0. 020	0.040±0. 017	0.040±0. 023	0.41±0.17 <sup>2</sup> 0.23+0.11 <sup>3</sup>	0.083± 0.035	0.099+0.057
100.11	07.180	cycloocta		0.036±	0.029±0.	0.0038±0	0.015±0.	0.017±0.	0.013±0.	0.2020.22	0.034±	0.0000_0.0007
108.184	C <sub>8</sub> H <sub>12</sub>	diene*	PTR	0.016	013	.017	0089	010	008		0.020	0.032±0.020
109.104	C <sub>6</sub> H <sub>5</sub> O <sub>2</sub>	*	PTR	0.019± 0.0095	0.013±0. 005	0.0012±0 .0058	0.0030±0 .0020	0.0055±0 .0026	0.0055±0 .0011		0.011± 0.007	0.014±0.008
		benzene										
		diol/met		0.2410	0 22 10 1	-	0.11+0.0	0.1110.0	0.1110.0	0.0010.202	0.2110	
110.112	C6H6O2	furfural*	PTR	0.34±0. 17	0.22±0.1 1	0.00028± 0.016	0.11±0.0 5	0.11±0.0 4	0.11±0.0 6	$0.68\pm0.29^{2}$ $0.25\pm0.12^{3}$	0.21±0. 08	0.27±0.15
										0.079±0.03		
		norcamp		0.006+	0.062+0	0.0014+0	0 022+0	0 020+0	0.024+0	2	0.050+	
110.156	C7H10O	furan*	PTR	0.0981	0.062±0. 030	.016	0.032±0. 018	0.030±0. 015	0.034±0. 020	0.040±0.02 4 <sup>3</sup>	0.0391	0.083±0.048
		cycloocte		0.017±	0.012±0.	0.0011±0	0.0088±0	0.012±0.	0.0053±0		0.024±	
110.2	C <sub>8</sub> H <sub>14</sub>	ne* dihydrox	PTR	0.0083	004	.0085	.0071	009	.0034		0.019	0.014±0.009
		y										
		piridine/										
	CEHEN	methyl maleimid		0.0074	0.0062+0	0.00025+	0.0026+0	0.0031+0	0.0022+0	0 024+0 00	0.0061	0.0051+0.003
111.1	O <sub>2</sub>	e*	PTR	1	.0021	0.0040	.0018	.0022	.0015	8 <sup>3</sup>	3	4
		furoic										
		oxv		0.18±0.	0.14±0.0	0.071±0.	0.041±0.	0.044±0.	0.038±0.		0.087±	
112.084	C₅H₄O <sub>3</sub>	furfural*	PTR	06	4	058	036	038	035	0.12±0.03 <sup>3</sup>	0.074	0.089±0.080
		cyclohex			0.044+0	- 0.0017±0	0.014+0	0.014±0	0.014+0		0.029+	
112.128	C <sub>6</sub> H <sub>8</sub> O <sub>2</sub>	*	PTR	0.037±	0.044±0. 021	.017	0.014±0. 0072	0.014±0. 007	0.014±0. 007	0.12±0.06 <sup>3</sup>	0.028±	0.033±0.018
		ethylcycl		T				Ì	1		T	
112 172	Calling	oheptano	DTP	0.019±		0.0019±0	0.0070±0		0.0070±0	0.014±0.00 7 <sup>3</sup>		0.016+0.010
112.1/2	0/1120	sum of		0.0074		.0007				,		0.01010.010
	$C_6H_{10}O$	cyclic			0.19±0.0	0.074±0.	0.12±0.0	0.12±0.0		0.039±0.01	0.11±0.	
114.144	2	saturated	CIMS		7	086	4	44		7 <sup>3</sup>	04	

Molecu- lar Weight	Comp- ound	Compou nd Name	Instru- ment	SP Averag	NP Average	Backgr- ound	Average EF (g kg <sup>-1</sup> )	NP EF (g kg <sup>-1</sup> )	SP EF (g kg <sup>-1</sup> )	Literature EF	NP ER	SP ER
weight				е						(8 ~ 8 )		
		and n-										
		unsaturat										
		carboxyli										
		c acids										
		Caprolact										
		esters/										
		c6										
11/ 1//	C <sub>6</sub> H <sub>10</sub> O	diketone	ntr	0.033±	0.029±0.	0.0065±0	0.0068±0	0.0082±0	0.0053±0		0.016±	0.013+0.009
114.144	2	heptanon	pti	0.011	011	.015	.0040	.0037	.0037		0.011	0.01310.005
		e/heptan										
11/ 188	C-H40	al/isomer	DTR	0.010±		0.00080±	0.0039±0		0.0039±0	0.0072±0.0		0.0090±0.005
11.100	0/11140	C5 oxo-		0.000		0.0070	.0025		.0025	023		,
		carboxyli		0.18±0.	0.15±0.0	0.10±0.0	0.031±0.	0.026±0.	0.037±0.	0.034±0.01	0.050±	0.000.00.045
116.116	C5H8O3	c acids butyl	CIMS	04	3	1	017	013	020	9 <sup>3</sup>	0.025	0.083±0.045
	$C_6H_{12}O$	acetate/c		0.023±	0.019±0.	0.00086±	0.0073±0	0.0094±0	0.0052±0		0.018±	
116.16	2	6 esters *	ptr	0.010	009	0.0081	.0045	.0054	.0033		0.010	0.012±0.007
		cyclohex		0.0094	0.0073+0	- 0.00048+	0.0032+0	0.0040+0	0.0025+0		0.0075	0.0056+0.003
116.222	$C_6H_{12}S$	anethiol*	PTR	7	.0017	0.0064	.0028	.0037	.0016		9	6
				0.0026	0.0005.10	-	0.001710	0.0000010	0.00001		0.0048	0.001010.000
118.088	C₄H₅O₄	acid*	PTR	±0.001 1	0.0025±0 .0011	0.0012±0 .0031	0.0017±0 .0007	0.0026±0	0.00081± 0.00034		±0.001 6	0.0018±0.000 7
										0.037±0.02	-	
		bonzofur		0.020+	0.025±0	-	0.017+0	0.018±0	0.017+0	$0^{2}$	0.024+	
118.135	C <sub>8</sub> H <sub>6</sub> O	an*	PTR	0.039± 0.022	0.025±0. 016	0.0018±0 .0060	0.017±0. 009	0.018±0. 010	0.017±0. 009	$5^{3}$	0.034± 0.018	0.038±0.021
		methylst										
		yrene/pr										
		benzenes		0.022±	0.016±0.	0.0021±0	0.018±0.	0.024±0.	0.011±0.	0.037±0.02	0.046±	
118.179	C <sub>9</sub> H <sub>10</sub>	*	PTR	0.010	007	.0079	012	016	007	0 <sup>3</sup>	0.030	0.025±0.015
				0.0039		0.0012+0	0.0016+0		0.0016+0			0.0035+0.002
119.167	C <sub>8</sub> H <sub>9</sub> N	*	PTR	2		.0017	.0015		.0015			9
		methylbe								0.40.000		
		nzaideny de/tolual		0.064±	0.039±0.	0.0026±0	0.025±0.	0.024±0.	0.026±0.	0.13±0.08 <sup>2</sup> 0.082±0.03	0.044±	
120.151	C <sub>8</sub> H <sub>8</sub> O	dehyde*	PTR	0.031	019	.013	014	012	016	0 <sup>3</sup>	0.022	0.058±0.034
		trimethyl								0.051 <sup>1</sup>		
		C9								0.051±0.02		
		aromatic		0.070±	0.056±0.	0.0078±0	0.052±0.	0.075±0.	0.029±0.	0.069±0.03	0.14±0.	
120.195	C <sub>9</sub> H <sub>12</sub>	S*	PTR	0.031	022	.017	029	037	018	$1^3$	07	0.064±0.039
		acid/hvdr								$5^{2}$		
		oxybenza		0.068±	0.053±0.	0.0092±0	0.020±0.	0.021±0.	0.019±0.	0.065±0.02	0.037±	
122.123	C <sub>7</sub> H <sub>6</sub> O <sub>2</sub>	Idehyde*	PTR	0.025	016	.018	011	010	012	33	0.018	0.040±0.026
		2										
		phenol/		0.000	0.000 -	0.0000				0.411-0-12	0.000	
122 167	C <sub>8</sub> H <sub>10</sub> O	methylan isole*	PTR	0.033±	0.025±0. 014	0.00069±	0.015±0. 0082	0.016±0. 009	0.013±0. 008	0.11±0.04 <sup>2</sup> 0.10±0.06 <sup>3</sup>	0.029±	0.029+0.017
122.107	001100	cyclohexy		0.019±	0.014±0.	0.0027±0	0.0076±0	0.0083±0	0.0068±0	0.1010.00	0.015±	0.02320.017
122.211	$C_9H_{14}$	lallene*	PTR	0.0074	005	.0083	.0051	.0059	.0043		0.011	0.015±0.009

Molecu- lar	Comp- ound	Compou nd Name	Instru- ment	SP Averag	NP Average	Backgr- ound	Average EF (g kg <sup>-1</sup> )	NP EF (g kg <sup>-1</sup> )	SP EF (g kg <sup>-1</sup> )	Literature EF	NP ER	SP ER
Weight				е						(g kg <sup>-1</sup> )		
124.095	C <sub>6</sub> H <sub>4</sub> O <sub>3</sub>	hydroxy benzoqui none*	PTR	0.014± 0.008	0.011±0. 005	- 0.00040± 0.0051	0.0032±0 .0022	0.0029±0 .0021	0.0035±0 .0023	$\begin{array}{c} 0.073 {\pm} 0.01 \\ 8^2 \\ 0.045 {\pm} 0.02 \\ 6^3 \end{array}$	0.0051 ±0.003 7	0.0075±0.004 7
124.139	C7H8O2	guaiacol*	PTR	0.15±0. 09	0.093±0. 054	- 0.0013±0 .0073	0.052±0. 026	0.051±0. 020	0.053±0. 031	$0.37\pm0.12^2$ $0.27\pm0.17^3$	0.091± 0.037	0.12±0.07
124.183	C <sub>8</sub> H <sub>12</sub> O	acetylcycl ohexene *	PTR	0.018± 0.009	0.013±0. 005	0.00032± 0.0056	0.0078±0 .0045	0.0087±0 .0051	0.0068±0 .0039		0.015± 0.009	0.015±0.008
126.111	C <sub>6</sub> H <sub>6</sub> O <sub>3</sub>	hydroxy methylfu rfural*	PTR	0.033± 0.014	0.025±0. 010	0.00066± 0.0087	0.0096±0 .0053	0.0094±0 .0049	0.0098±0 .0056	$\begin{array}{c} 0.27 \pm 0.10^2 \\ 0.064 \pm 0.02 \\ 6^3 \end{array}$	0.016± 0.009	0.021±0.012
		unsaturat ed C6 cyclic		0.052+	0.042+0	0.0000+0	0.012+0	0.015+0	0.0097±0		0.026+	
126.155	2	c acid*	CIMS	0.032 <u>+</u> 0.005	0.043±0. 011	.0096	0.012±0. 005	0.013±0. 0057	.0044		0.020 <u>+</u> 0.010	0.019±0.009
126 155	C7H10O	cyclohex ene carboxyli c acid *	ntr	0.015±	0.013±0.	0.0026±0	0.0064±0	0.0080±0	0.0048±0		0.014±	0 010+0 007
126.199	C <sub>8</sub> H <sub>14</sub> O	octenone *	PTR	0.0064 ±0.002 9		0.00012± 0.0051	0.0027±0 .0016		0.0027±0 .0016			0.0057±0.003 3
126.217	C <sub>7</sub> H <sub>10</sub> S	trimethyl thiophen e*	PTR	0.013± 0.004	0.011±0. 005	0.00044± 0.0084	0.011±0. 011	0.016±0. 015	0.0054±0 .0030		0.028± 0.026	0.012±0.007
128.127	C <sub>6</sub> H <sub>8</sub> O <sub>3</sub>	di hydroxy methyl furan*	PTR	0.024± 0.013	0.020±0. 007	- 0.0017±0 .0087	0.0044±0 .0026	0.0059±0 .0032	0.0029±0 .0017		0.010± 0.006	0.0063±0.003 6
128.171	C7H12O	cyclohex anoic acid *	ptr	0.014± 0.004		0.0028±0 .011	0.0050±0 .0034		0.0050±0 .0034			0.010±0.007
128.174	C10H8	naphthal ene*	PTR	0.033± 0.013	0.025±0. 011	0.00075± 0.0086	0.017±0. 011	0.018±0. 012	0.015±0. 009	0.078±0.05 6 <sup>2</sup>	0.031± 0.021	0.031±0.018
128.215	C <sub>8</sub> H <sub>16</sub> O	octanone *	PTR	0.0072 ±0.003 5		0.0014±0 .0055	0.0028±0 .0020		0.0028±0 .0020			0.006±0.0042
130.187	C7H14O	amyl acetate *	ptr	0.0074 ±0.004 6	0.0056±0 .0021	0.00025± 0.0046	0.0031±0 .0019	0.0034±0 .0020	0.0028±0 .0018		0.0056 ±0.003 3	0.0058±0.003 6
132.159	C <sub>6</sub> H <sub>12</sub> O	C6 hydroxy- carboxyli c acids	CIMS	0.0090 ±0.000 6	0.010±0. 002	0.0062±0 .0014	0.0017±0 .0008	0.0027±0 .0011	0.00075± 0.00034		0.0045 ±0.001 8	0.0016±0.001 0
132 162	CallaO	methyl benzo furans*	DTR	0.023±	0.016±0.	0.00018±	0.010±0.	0.010±0.	0.011±0.	0.055±0.03 0 <sup>2</sup> 0.046±0.02	0.017±	0.021+0.012
132.102	CyrigO	ethyl styrene/ methyl		0.012	0.014±0	0.0012±0	0.0000+0	0.0000+0	0.0000+0	0.041±0.01 9 <sup>2</sup>	0.014	0.02110.012
132.206	C <sub>10</sub> H <sub>12</sub>	benzene*	PTR	0.019±	0.014±0. 007	.0063	.0050	.0050	.0050	6 <sup>3</sup>	0.014± 0.008	0.017±0.010
134.134	$C_8H_6O_2$	phthalic acid*	PTR	0.0074 ±0.002 8	0.0071±0 .0025	0.0011±0 .0051	0.0039±0 .0029	0.0044±0 .0035	0.0033±0 .0022		0.0071 ±0.005 7	0.0065±0.004 2

Molecu- lar	Comp- ound	Compou nd Name	Instru- ment	SP Averag	NP Average	Backgr- ound	Average EF (g kg <sup>-1</sup> )	NP EF (g kg <sup>-1</sup> )	SP EF (g kg <sup>-1</sup> )	Literature EF	NP ER	SP ER
Weight				e			(8-8 /			(g kg <sup>-1</sup> )		
		methylac				_				0.053±0.03		
134.178	C <sub>9</sub> H <sub>10</sub> O	etopheno ne*	PTR	0.012± 0.007	0.0085±0 .0041	0.00032± 0.004	0.0059±0 .0035	0.0062±0 .0036	0.0056±0 .0033	0.045±0.01 9 <sup>3</sup>	0.010± 0.006	0.011±0.006
124 222	Cullu	C10 Aromatic	DTD	0.030±	0.024±0.	0.0024±0	0.024±0.	0.035±0.	0.013±0.	0.040±0.02	0.058±	0.026+0.016
134.222	C101114	5	FIN	0.013	010	.0095	014	019	008	0.081±0.03	0.031	0.02010.010
136.15	C <sub>8</sub> H <sub>8</sub> O <sub>2</sub>	benzoic acid *	PTR	0.027± 0.015	0.018±0. 009	0.00058± 0.007	0.013±0. 007	0.014±0. 0069	0.012±0. 007	0- 0.066±0.02 9 <sup>3</sup>	0.022± 0.011	0.023±0.014
		monoter		0.68+0.	0.65+0.3	0.057+0.	0.41+0.1	0.49+0.2	0.33+0.1	0.41±0.06 <sup>1</sup> 0.87+0.72 <sup>2</sup>	0.80+0.	
136.238	C <sub>10</sub> H <sub>16</sub>	penes*	PTR	28	1	055	9	0	7	0.21±0.15 <sup>3</sup>	32	0.65±0.38
138.122	C <sub>7</sub> H <sub>6</sub> O <sub>3</sub>	hydroxyb enzoic acid*	PTR	0.0080 ±0.003 0	0.0076±0 .0033	- 0.000093 ±0.0050	0.0026±0 .0017	0.0039±0 .0023	0.0014±0 .0008		0.0061 ±0.003 6	0.0028±0.001 5
129.166	C <sub>8</sub> H <sub>10</sub> O	creosol/ methyl	DTD	0.016±	0.012±0.	0.000003	0.0073±0	0.0077±0	0.0069±0	0 1 4+0 113	0.012±	0.012±0.008
138.100	2	isophoro	PIK	0.0093 0.027±	0.025±0.	0.0079±0	0.0046 0.0092±0	0.0049 0.0086±0	0.0098±0	0.14±0.11	0.008 0.014±	0.013±0.008
138.21	C <sub>9</sub> H <sub>14</sub> O	ne* dimethvl	PTR	0.0075	009	.01	.0064	.0053	.0074		0.009	0.019±0.014
		benzofur		0.0008						0.043±0.01	0.0079	
	C <sub>10</sub> H <sub>10</sub>	benzofur		±0.004	0.0072±0	0.00041±	0.0048±0	0.0052±0	0.0045±0	0.051±0.02	±0.005	0.0083±0.004
146.189	0	an*	PTR	4 0.0064	.0033	0.0037	.0029	.0034	.0024	83	1	4
146.233	C <sub>11</sub> H <sub>14</sub>	*	PTR	±0.003 5		0.0012±0 .0041	0.0034±0 .0013		0.0034±0 .0013			0.0061±0.004 4
		cinnamic		0.0040 ±0.003		- 0.00094±	0.0021±0		0.0021±0			0.0037±0.002
148.161	C <sub>9</sub> H <sub>8</sub> O <sub>2</sub>	acid*	PTR	3		0.0048	.0013		.0013	0 027+0 01		2
	<b>C</b> 11	benzylac		0.0047	0.004210	0.000451	0.000010	0.0000.00	0.000410	2 <sup>2</sup>	0.0033	0.004410.000
148.205	C <sub>10</sub> H <sub>12</sub> O	etone/es tragole*	PTR	±0.002 1	0.0043±0 .0019	0.00045± 0.0030	0.0023±0 .0016	0.0022±0 .0017	0.0024±0 .0015	0.025±0.01 5 <sup>3</sup>	±0.002 6	0.0044±0.002 7
		C11 aromatic								0.014±0.00		
		s/pentam		0.0074	0.0054+0	0.00054+	0.0041+0	0.0042+0	0.0028+0	8 <sup>2</sup>	0.0064	0.0069+0.004
148.249	C <sub>11</sub> H <sub>16</sub>	zene*	PTR	2	.0027	0.000341	.0028	.0032	.0023	7 <sup>3</sup>	8	1
	C.H.O	ethyl benzoate		0.0059	0.0044+0	0.00030+	0 0028+0	0 0020+0	0.0028+0	$0.14\pm0.08^{2}$	0.0043	0.0040+0.003
150.177	2 2	guaiacol*	PTR	10.002 9	.0020	0.00039± 0.0034	.0028±0	.0023	.0017	5 <sup>3</sup>	±0.003 4	0.0049±0.003 0
150.221	C <sub>10</sub> H <sub>14</sub>	carvone*	PTR	0.0040 ±0.001 2	0.0039±0	0.00055±	0.0021±0	0.0027±0	0.0015±0		0.0039 ±0.003 0	0.0027±0.001
150.221	0			-		0.0020		.0021		0.027±0.01	Ŭ	
152.237	C <sub>10</sub> H <sub>16</sub> O	/isomers *	PTR	0.022± 0.0074	0.023±0. 010	0.0063±0 .0086	0.011±0. 007	0.013±0. 008	0.0087±0 .0066	0.025±0.01 4 <sup>3</sup>	0.020± 0.011	0.015±0.011
	C <sub>9</sub> H <sub>14</sub> O	norborna neacetic		0.0036 ±0.001		- 0.00050±	0.0022±0		0.0022±0			0.0038±0.002
154.209	2	acid*	PTR	5 0.0054		0.0041	.0012		.0012		0.0046	0
154.212	C <sub>12</sub> H <sub>10</sub>	acenapht hene*	PTR	±0.002 0	0.0040±0 .0015	0.00026± 0.0031	0.0029±0 .0022	0.0033±0 .0028	0.0025±0 .0013		±0.004 0	0.0042±0.002 3

Molecu- lar Weight	Comp- ound	Compou nd Name	Instru- ment	SP Averag e	NP Average	Backgr- ound	Average EF (g kg <sup>-1</sup> )	NP EF (g kg <sup>-1</sup> )	SP EF (g kg <sup>-1</sup> )	Literature EF (g kg <sup>-1</sup> )	NP ER	SP ER
154.253	C <sub>10</sub> H <sub>18</sub> O	terpine- 4- ol/cineol e/isomer s*	PTR	0.0024 ±0.000 9	0.0020±0 .00068	- 0.00044± 0.0022	0.0018±0 .0014	0.0019±0 .0018	0.0017±0 .0009	0.0056±0.0 021 <sup>2</sup> 0.0027±0.0 017 <sup>3</sup>	0.0029 ±0.002 8	0.0028±0.001 5
204.357	C15H24	sesquiter penes*	PTR	0.0027 ±0.001 1		0.00030± 0.0021	0.0017±0 .0011		0.0017±0 .0011	$\begin{array}{c} 0.15 \pm 0.07^2 \\ 0.029 \pm 0.02 \\ 8^3 \end{array}$		0.0022±0.001 4
239±61	C <sub>11</sub> to C <sub>25</sub>	I/SVOCs – CH	cartrid ge			5.2	1.4±0.03 7 to 2.4±0.06 3					
255±61	C <sub>11</sub> to C <sub>25</sub>	I/SVOCs – CHO1	cartrid ge			4.6	0.81±0.0 78 to 0.81± 0.079					
271±61	C <sub>11</sub> to C <sub>25</sub>	I/SVOCs – CHS1	cartrid ge			0.2	0.21±0.0 033 to 0.22±0.0 060					