

May 26, 2022
Ref.: acp-2022-202

Author Response (AR): We thank all three reviewers for providing thoughtful guidance and insight, improving the current manuscript. We are excited to share these results with the greater community.

AR: All of the below comments have been addressed in the final manuscript, unless otherwise noted.

Interactive comment on “Measurement Report: Tropospheric and Stratospheric Ozone Profiles during the 2019 TROPomi vaLidation eXperiment (TROLIX-19)” by John Sullivan et al.,

Anonymous Referee #1

The paper submitted to ACP by J.F. Sullivan and co-workers is a measurement report of data collected during a joint NASA-KNMI field experiment designed to validate ozone retrieval by the TROPOMI spaceborne mission. On 11 days, lidar O₃ observations in the troposphere and the stratosphere are carried out for several hours at the Cabauw site. The two NASA lidars are indeed very valuable tools for the assessment of TROPOMI. Observations of total ozone column (TOC), stratospheric limb sounders (MLS), and nearby ECC launched at De Bilt are also analyzed in this work, in addition to the daily ozone simulations of the GEOS-CF model. The time evolution of TOC and tropospheric O₃ partial columns (0-2km, 0km-Tropopause level) is used to assess TROPOMI observations and GEOS-CF model ozone mapping. Considering the value of satellite validation exercises and comparisons between model and observations, the paper is appropriate for publication in ACP (or AMT considering the focus on experimental data analysis). It is well written with some remaining errors in figure legends. I fully agree that there is a general good agreement between the different observations and that the GEOS-CF model performs quite well.

AR: Thank you for understanding the value of utilizing many different platforms to better evaluate the satellite products and their physical interpretation. We hope to continue using the lidars, models, and other ground-based observations to evaluate future satellites as well. We also agree this is well within the scope of ACP (especially the special issue) and think has broad enough scope to be quite impactful.

My main criticism would be that detailed discussions of the differences observed in the troposphere are sometimes missing and that these differences need to be acknowledged in the conclusions: TROPOMI overestimate in the lowermost stratosphere (10-15 km), GEOS-CF low performance during the 9/20-21 episode (not only at 4 km but also below the tropopause leading to small differences of tropospheric columns without a good agreement in the vertical profile), ECC/lidar positive differences in the free troposphere on 9/12,17,21 (De Bilt/Cabauw spatial differences ?).

AR: We appreciate the thoughtful review and have answered the below comments that should be more than adequate to alleviate any concerns.

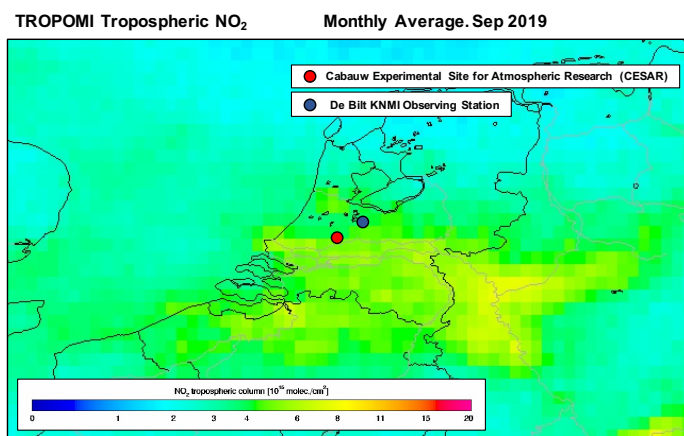
Specific comments

Line 45 : Lidar is also sensitive to cloud cover. Are missing lidar data during the campaign related to cloud cover ?

AR: Yes and this has now been mentioned in the paragraph beginning on Line 116.

Fig. 1 : Fig. 1 is not referenced in the text. What is the purpose of this figure ? Please also show De Bilt and Cabauw on the NO₂ map.

AR: Figure 1 is now reference in line 50. The purpose of this figure is to provide the read with geographic context for the site. Rather than focus on a specific day in the original submission (Sep 14, 2019), the image now serves as a monthly mean for September 2019 (the TROLIX-19 period).



Line 108 : Addition of lidar O₃ measurement accuracy and vertical resolution would be useful in Table 1. Since partial tropospheric columns and TOCs are discussed in section 3.2 and 5, the expected accuracy of the lidar retrieval on these columns would help.

AR: The below text has been added to the paper near line 125 to better characterize the uncertainty estimates.

A brief description and community standardized definitions of the uncertainty budget of the lidar measurements presented in this paper can be found in Sullivan *et al.*, 2014, Leblanc *et al.*, 2016 and Leblanc *et al.*, 2018. The maximum statistical uncertainties for the two GSFC lidars vary from night to night depending on atmospheric conditions and laser power fluctuations. They are mostly within 10-20% for 5 min and 5-8% for 30 min integrations throughout the atmosphere within their measurable ranges although they are different at the same altitude due to laser performance and telescope/detector efficiency differences.

Line 151 : Give a more recent reference i.e. Smit and Thompson 2021 GAW report 268 describing ASOPOS 2.0 error calculations.

AR: This reference has been added and we appreciate the attention to detail relating to contemporary references on this subject matter.

Line 171 : What is the expected difference when using the 2.5 PVu instead the 3 PVu tropopause definition taken for the lidar tropospheric ozone column calculation ?

AR: The tropopause height was calculated from the GEOS-CF and was then used consistently for all measurements in the trop. ozone column calculations. We have confidence that this is appropriate by visually comparing the tropopause heights compared to the model and lidar in Figure 2 and later on with ozonesonde values. Within the community, there are several definitions of the tropopause height, and we have done our best effort to stay consistent with the 2.5 PVU value and have described the method used to calculate this in the literature.

Line 206 : What is the typical GEOS-CF model vertical resolution in the UTLS ?

AR: There are 72 total vertical layers in the GEOS-CF model, scaled by atmospheric pressure. From 10 mb to 375mb there are 20 vertical layers. In general this is roughly ~1.1-1.2 km and a visual reference is shown in Fig. 3 here :<https://acp.copernicus.org/articles/17/1417/2017/acp-17-1417-2017.pdf>

You can find much of the details (also cited in paper as Keller et al., 2021) pertaining to the GEOS-CF system here:

<https://agupubs.onlinelibrary.wiley.com/doi/10.1029/2020MS002413>

Line 225 : Add 2f in the list of panels with large lidar/model differences.

AR: Done.

Line 229 : It is also true on 9/13-14.

AR: Added.

Line 233 : What is the thickness of the 4-km layer used for plotting ozone in Fig. 3 ?

AR: The layer was calculated to match the closest vertical layer of the GEOS-CF for consistent intercomparison. The top of model layer 53 is at 3.94 km (~605 mb, for the specific files used in this work) and the top of model layer 54 is 4.44 km (~570 mb, for the specific files used in this work). This is roughly a 500m or 40mb layer thickness.

Figure 2 : Ozone unit is not specified in the caption or color scale. Please add the pink tropopause altitude on the ozone plot and the vertical limits of the 4-km layer used in Fig. 3.

What is the time end of panel e ? This plot is very nice. It's a shame not including the two days with ECC at De Bilt on 9/12 and 9/17.

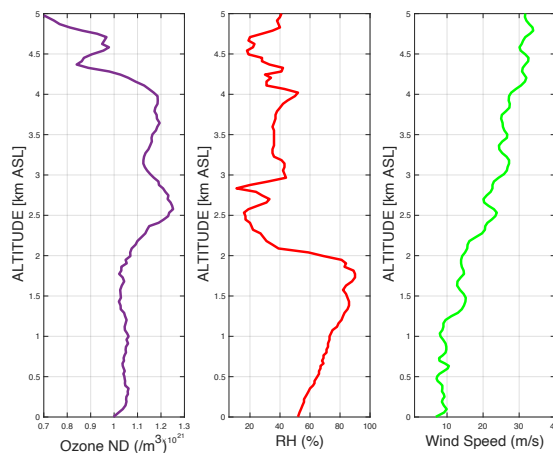
AR: The ozone unit in Figure 2 has been added – Ozone mixing ratio [ppbv]. Also the tropopause heights have been added to the ozone lidar measurements. The details of the layer thickness have also been added to the caption.

Line 243 : The O3 vertical structure in the UTLS is also missed by the model (no low O3 at 11 km). It is then likely that the mesoscale ozone 3D transport in the frontal system is not very well resolved by the model for this specific event.

AR: A comment reflecting this has been added to the revised manuscript starting near Line 286.

Line 246 : Why are 9/17 ECC ozone concentrations significantly larger than model, TROPOMI and lidar O3 values ? Did you see any horizontal O3 gradients near Cabauw in the model or TROPOMI mapping ?

AR: Below is a quick plot of the 9/17 ECC (however, not added to current manuscript). You can see the sampled ozone, RH, and wind speed in their respective units. There is no sign of any pump current overheating or other QC/QA issues that may cause invalid data within the sonde chamber. In particular, this feature at ~4km near 1.2×10^{21} mol/m³ is associated with a much drier air mass and is notable traveling upwards of 30 m/s during the sampling period. The wind direction was consistently coming from the WNW during this period. At this speed and direction, the sonde (already launched at De Bilt) may have just been sampling a similar but not identical air mass. The differences begin to decrease in this profile compared to the lidar near about 11km (as shown in the current manuscript Figures).



Line 271-274 : This discussion is unclear. The difference between the tropospheric columns are in fact not too bad but are fortuitous because the effect of the missing layer at 4 km is cancelled by the high model ozone at 11 km. Please reformulate this part of the discussion.

AR: A clarification regarding this has been added to the text near Line 408.

Line 279 Figure 4 low panel (not top panel)

AR: Done.

Line 282 I do not understand this sentence.

AR: A clarification regarding this has been added to the text and discussion in the same paragraph has been added from the previous comment (starting near line 408).

Line 285 What are the reasons for low 0-2 km column in some lidar data on 9/12 and 14 and in TROPOMI data on 9/17 ?

AR: These values were driven by cloud contamination in the lower levels of the troposphere for all observations.

For the 09/12 and 09/14 lidar data, this is due to some data points being cloud contaminated. A stricter (and more manual) cloud screen using lidar backscatter SNR has been implemented now.

For the low TROPOMI data on 09/17 (and by extension a few other days in the second half of the domain) are also likely due to cloud interference. There were several TROPOMI pixels that were not retrieved because they did not match the quality requirements (noise too large). By using the OMPS reflectivity data, I have screened additionally for the tropospheric and 0-2km column data in Figure 4. This was changed from a 0.8 reflectivity to 0.6 reflectivity and a note has been made in the caption.

Line 288 A discussion on the TROPOMI retrieval of the partial column is needed in this section. While the 0-2 km columns remain within the range of the diurnal variability, several full tropospheric columns (9 /12,13, 18,19,20) are well above this diurnal variability. It is likely related to the limited vertical resolution in the UTLS. The TROPOMI overestimates are also quite clear in the ozone profiles shown in Fig. 5.

AR: This is an important point, thank you for this. A short discussion has been added, and the reader is directed to the Mettig et al., work that systematically goes through the retrieval from first principles within the spectral range of the ozone profile retrieval (<https://amt.copernicus.org/articles/14/6057/2021/amt-14-6057-2021.pdf>).

Line 299 Reading this sentence, I am not sure which altitude range is critically needed for monitoring of the column

AR: This has been specified in the context of the previous comment.

Line 308 What is OMPS-LP ? The stratospheric MERRA-2 profile used for the OMPS tropospheric ozone column calculation ? Do hybrid lidar profiles combine daytime TROPOZ data with nighttime STROZ-LITE data or are nighttime profiles only considered in the hybrid version ? Please clarify this point.

AR: From Section 2.3.1, this is the OMPS Limb Profiler ozone profile product and a clarifying statement has been added to start Sec 2.3.1.

The vertical distribution of ozone in the stratosphere and lower mesosphere is obtained from the OMPS Limb-Profiler (LP) sensor on the Suomi-NPP satellite merging the UV (29.5-52.5 km) and VIS (12.5-35.5 km) bands to provide a full profile from 12.5km to 52.5km (Kramarova *et al.*, 2018).

AR: I have specified that this is daytime or nighttime TROPOZ data, and only nighttime measurements for the STROZ. This was ultimately manually selected each day since this was such a focused intensive campaign and we were able to do so. Laser performance and cloud considerations also drove the decisions.

Fig. 5 I believe that the titles of the bottom panels are wrong otherwise Fig.3 and Fig.5 are not consistent.

AR: Great catch – the ozonesonde read file was pointing to opposite directories for those two days (so the titles were correct, but the ECC profile is swapped). This has been corrected and actually shows a much better comparison to the lidar and model, while simultaneously highlighting the feature described in Fig 3.

Line 320 Comparisons of the stratospheric profiles are extensively discussed while the general agreement is quite good in the stratosphere. The discussion of the tropospheric differences is however limited while there are some interesting differences (TROPOMI in the UTLS, lidar ECC differences).

AR: A clarification regarding this has been added to the text throughout Lines 352-373 and is now reiterated in the concluding remarks.

Fig. 7 Color scale is missing

AR: This has been added and this was a major oversight to miss on initial submission.

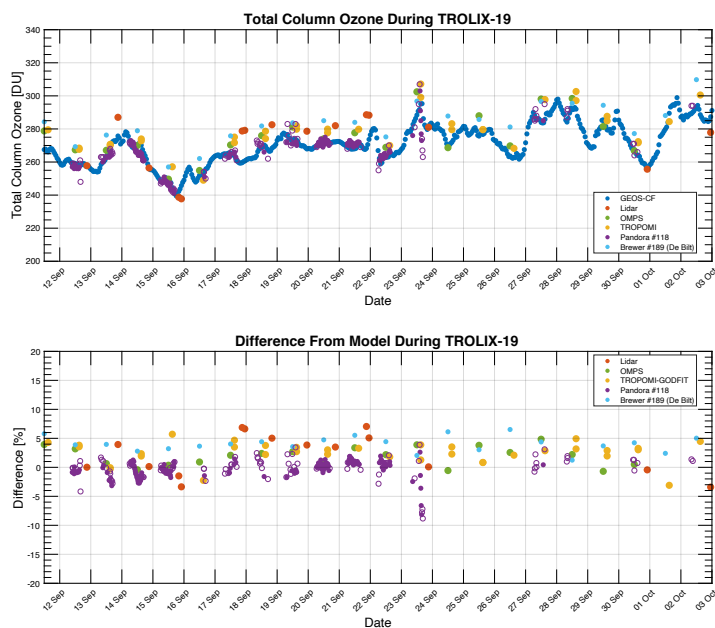
Line 339: Why do you say that OMPS-LP underestimates O₃ concentrations at altitudes below 20 km ? I see both positive and negative differences. Regarding TROPOMI, it looks like TROPOMI versus MLS is better between 10 and 20 km than the large lidar/TROPOMI differences in the UTLS. Is it related to similar UTLS vertical resolutions for MLS and TROPOMI ?

AR: A clarification regarding this has been added to the text throughout Lines 352-373 and is now reiterated in the concluding remarks.

Line 358: A 7% difference on TOC is already quite significant. What are the reasons for the low TOCs given by the hybrid lidar retrieval on 9/14-15 or by Pandora on 9/23 ?

In looking further at the mismatch in lidar total column ozone values, we discovered the calculation was performed and plotted with the corresponding TROPOZ integration time of day which was predominantly during the daytime observations. For this specific section describing total column ozone, the majority of that information is coming from the STROZ LITE nighttime data. Therefore, more optimal merges (particularly in the UTLS) were performed to include as much STROZ data as possible and shift the red data points to a more representative time of the lidar observations. This results in a much closer comparison to the observations as compared to the previous figure for the time in question.

Pandora flags equal to 10 (purple filled dots) and 11 (purple open dots) are shown. On 09/23 there were many observations mixed with both flags throughout the day, likely due to cloud contamination. Only data with the highest QC/QA level (i.e. 10) is plotted in the updated final figure. This results in less data overall.



Line 371: GEOS-CF indeed performs well reproducing the ozone downward transport in the UTLS, but a sentence could be added about GEOS-CF failure to resolve some high resolution laminae related to specific mesoscale ozone 3D transport from the UTLS, e.g. 09/19-20.

AR: A clarification regarding this has been added to the text starting near Line 288.

Line 374: The overestimate of the TROPOMI retrieval between 10 and 15 km needs to be mentioned in the conclusions.

AR: A clarification regarding this has been added throughout the previous recommendations and line numbers above as well as additional context in the conclusions.

We thank you for such a thoughtful review (including a few typos/errors) and know the manuscript has improved greatly from your suggestions.

Anonymous Referee #2

Review of “Measurement Report: 1 Tropospheric and Stratospheric Ozone Profiles during the 2019 TROPomi vaLidation eXperiment (TROLIX-19)” by J.T. Sullivan et al.

In this manuscript the authors describe the many ozone measurements made from multiple platforms during the TROPomi vaLidation eXperiment (TROLIX-19) campaign in fall 2019. The campaign was designed to support satellite validation efforts with the emphasis on understanding the vertical profile retrievals of ozone. Instrumentation included ozone lidars, pandora spectrometer, Brewer Spectrophotometer and ozonesondes. Satellite data used include OMPS, MERA-2, MLS and TROPOMI. In addition, measurement data is compared with GEOS-CF model output.

The analysis focuses on two main goals; one to evaluate ozone retrievals in relation to current and future (TEMPO) satellites. And the combination of the tropospheric ozone lidar and the stratospheric lidar providing hybrid ozone profiles from ~0.2 km to ~50 km.

The article is clearly written and provides a comprehensive presentation of data from a number of measurement platforms. With figures and discussion of the comparisons across measurement platforms and model within the full column ozone and 0-2 km tropospheric column ozone. The study makes several important observations as to the structure of the ozone within the atmospheric column and the differences in instrument/model performances.

AR: Thank you for understanding the value of utilizing many different platforms to better evaluate the satellite products and their physical interpretation. We hope to continue using the lidars, models, and other ground-based observations to evaluate future satellites as well. We also agree this is well within the scope of ACP (especially the special issue) and think has broad enough scope to be quite impactful.

Specific comments:

I would suggest that the conclusions section needs to be expanded upon. This section emphasizes the importance of observations and for the site itself (which I agree on) but I would like to see more quantitative evaluation of the data here to back up the statements, adding some percentage differences etc would make this an easy reference source for future readers. For example, there is a statement (line 74-75): “ TROPOMI ozone profile products are able to accurately reproduce ozone quantities in the lower troposphere...” with reference to Figure 3 but this only shows TROPOMI compared with observations vertical layer at 4km. In addition, Figure 7 indicates that TROPOMI generally overestimates, especially within the troposphere. Expanding the conclusions to include some of the quantitative results would help to firm up the concluding statements.

AR: The conclusion section has been broadened beyond the importance of observational platforms. We have added several discussion clarification on TROPOMI overestimations per the previous reviewer that are intended to improve the communication surrounding this issue.

Figure 7 was intended to be the quantitative resultant figure that could be referenced or looked up easily and this has been elaborated on in the conclusions. We feel this figure better illustrates both the temporal (over the course of the campaign) and vertical differences observed in the retrievals and reducing that down to a simple number or XX % may underserve the push for the vertical profiling needed for these types of efforts. (And hopefully the colorbar now added helps the reader too!)

Figure 1. Add the CESAR site also be added to the image on the right or add lat/lon information to maps.

AR: CESAR has been added.

Figure 7. Needs a color bar with values for the differences in ozone number densities

AR: Done, thanks!

Reviewer 3:

Review of “Measurement Report: Tropospheric and Stratospheric Ozone Profiles during the 2019 TROPomi vaLIdation eXperiment (TROLIX-19)” by J.T. Sullivan et al, submitted to ACP

In this manuscript, a measurement report is given on the various ground-based measurements deployed during the TROLIX-19 field campaign in Cabauw, Netherlands during September 2019. The focus of the campaign and this paper was to provide comprehensive validation of the TROPOMI ozone profile retrievals using ozone lidar measurements, the measurement focused on in this paper, as well as ozonesondes, Pandora, Brewer, etc. Ozone lidar measurements were also compared against other satellite and model datasets, such as OMPS (satellite), MLS (satellite), and GEOS-CF (model). The authors also analyzed the temporal variability of full tropospheric and 0-2 km ozone columns using GEOS-CF, TROPOMI, ozonesondes, and ozone lidar.

The article provides a comprehensive analysis of the measurements and ancillary datasets used to demonstrate the variability of tropospheric and stratospheric ozone profiles during TROLIX-19. The authors clearly and concisely state the impact of the ground-based measurements have on understanding the capabilities and limitations of current polar orbiting and future geostationary ozone retrievals.

Specific comments:

Line 95: It would be worth mentioning how many stations are in the network

AR: Done.

Line 145: Correct spelling of campaign

AR: Done.

Line 147: Specify which radiosonde manufacturer was used and the specific model

AR: Done.

Line 159: Put space between “are” and “used”

AR: Done.

Figure 2: Add label to the color bar, include units

AR: Done.

Line 233: It can be assumed that 4 km is chosen to investigate the model and observation difference observed between 3-5 km on 20-21 September. However, it should be specifically mentioned at the beginning of this paragraph why 4 km is chosen.

AR: A clarification has been added.

Line 279: Correct to say Figure 4, bottom panel

AR: Done.

~~Measurement Report~~: Tropospheric and Stratospheric Ozone Profiles during the 2019 TROPomi vaLIdation eXperiment (TROLIX-19)

John T. Sullivan¹, Arnoud Apituley², Nora Mettig³, Karin Kreher⁴, K. Emma Knowland^{1,5}, Marc Allaart², Ankie Piters², Michel Van Roozendaal⁶, Pepijn Veeffkind², Jerry R. Ziemke^{1,5}, Natalya Kramarova¹, Mark Weber³, Alexei Rozanov³, Laurence Twigg^{1,7}, Grant Sumnicht^{1,7}, and Thomas J. McGee^{1*}

¹ NASA Goddard Space Flight Center, Greenbelt, MD 20771

² Royal Netherlands Meteorological Institute (KNMI), De Bilt, Netherlands

³ Institute of Environmental Physics, University of Bremen, Bremen, Germany

⁴ BK Scientific GmbH, Mainz, Germany

⁵ Morgan State University/GESTAR-II, Baltimore, MD 21251

⁶ Belgian Institute for Space Aeronomie (BIRA), Ukkel, Belgium

⁷ Science Systems and Applications Inc., Lanham, MD, 20706

*Now Emeritus

Correspondence to: John Sullivan (john.t.sullivan@nasa.gov)

Abstract A TROPospheric Monitoring Instrument (TROPOMI) validation campaign was held in the Netherlands based at the CESAR (Cabauw Experimental Site for Atmospheric Research) Observatory during September 2019. The TROPomi vaLIdation eXperiment (TROLIX-19) consisted of active and passive remote sensing platforms in conjunction with several balloon-borne and surface chemical (e.g. ozone and nitrogen dioxide) measurements. The goal of this joint NASA-KNMI geophysical validation campaign was to make intensive observations in the TROPOMI domain in order to be able to establish the quality of the L2 satellite data products under realistic conditions, such as non-idealized conditions with varying cloud cover and a range of atmospheric conditions at a rural site. The research presented here focuses on using ozone lidars from NASA's Goddard Space Flight Center to better evaluate the characterization of ozone throughout TROLIX-19. Results of comparisons to the lidar systems with balloon, space-borne, and ground-based passive measurements are shown. In addition, results are compared to a global coupled chemistry meteorology model to illustrate the vertical variability and columnar amounts of both tropospheric and stratospheric ozone during the campaign period.

31 **1 Introduction**

32 In September 2019, a joint Royal Netherlands Meteorological Institute (KNMI) and the U.S. National
33 Aeronautics and Space Administration (NASA) field campaign was performed in the Netherlands, based
34 at the Cabauw Experimental Site for Atmospheric Research (CESAR, 51.97° N, 4.93° E), to provide the
35 scientific community with additional information to further understand and evaluate the Copernicus
36 Sentinel-5 Precursor mission (S-5P) TROPOspheric Monitoring Instrument (TROPOMI) instrument
37 (<https://sentinels.copernicus.eu/web/sentinel/missions/sentinel-5p>). The main objective of the
38 Copernicus Sentinel-5P mission is to perform atmospheric measurements with high spatio-temporal
39 resolution, to be used for scientific studies and monitoring of air quality and chemical transport
40 (https://www.esa.int/Applications/Observing_the_Earth/Copernicus/Sentinel-5P).

41 To properly support satellite evaluation, the 2019 TROPomi vaLIdation eXperiment (TROLIX-19)
42 campaign was designed to bring together many active and passive remote sensing platforms in
43 conjunction with several balloon-borne, airborne and surface measurements. Specifically, the
44 observations were established to provide geophysical verification in order to establish the quality of
45 TROPOMI Level 2 (L2) main data products under realistic non-idealized conditions with varying cloud
46 cover and a wide range of atmospheric conditions. Cabauw, using its comprehensive in-situ and remote
47 sensing observation program in and around the 213 m meteorological tower ([https://ruisdael-
48 observatory.nl/trolix19-tropomi-validation-experiment-2019/](https://ruisdael-observatory.nl/trolix19-tropomi-validation-experiment-2019/)) was the main site of the campaign with
49 focus on vertical profiling using lidar instruments for aerosols, clouds, water vapor, tropospheric and
50 stratospheric ozone, as well as balloon-borne sensors for nitrogen dioxide (NO₂) and ozone (Figure 1).
51 Although this work focuses primarily on the ozone lidar profiling during the study, the larger campaign
52 overview, background, and motivation can be found in Apituley *et al.* (2019; 2020) or Kreher *et al.*
53 (2020a).

54 One main goal of this work is also to understand ozone profile retrievals as they relate to upcoming
55 satellite endeavors. As NASA prepares to launch its first geostationary air quality satellite “Tropospheric

56 Emissions: Monitoring of Pollution” (TEMPO₂) this work also specifically establishes a paradigm of
57 evaluation for TEMPO-derived products such as tropospheric ozone columns and a 0-2km tropospheric
58 ozone product. An analogous geo-stationary air quality satellite, the Copernicus Sentinel-4 mission (S-
59 4), <https://sentinels.copernicus.eu/web/sentinel/missions/sentinel-4>, will provide hourly data on
60 tropospheric constituents over Europe and the CESAR site is directly within the satellite’s field of
61 regard. Due to the finer spatial footprint, increased temporal frequency and vertical extent of
62 TEMPO’s tropospheric ozone retrievals, ozone lidars are an ideal platform to perform future
63 evaluations of the products, which builds from recent work done in Johnson *et al.*, 2018. Specifically,
64 this work will investigate the results from the combination of having a co-located NASA tropospheric
65 (Sullivan *et al.*, 2014) and stratospheric ozone lidars (McGee *et al.*, 1991) in order to obtain an entire
66 vertical profile of ozone from ~0.2km to 50km.

67 For the first time, this transportable combination of lidars is able to explicitly derive diurnally varying
68 tropospheric and total ozone columns, which are compared directly to measurements obtained by
69 ground-based passive sensors, current satellite instrumentation and chemical transport models. In
70 Section 2 we present all available data and methods used in this work, across the various platforms
71 during the TROLIX-19 study. Section 3 focuses on comparisons of the tropospheric ozone retrievals of
72 the vertical profiles of ozone within the troposphere and columnar reductions of 0-10 km and 0-2 km.
73 Comparisons of lidar data with available complete ozone profiles (Sec 4) and columnar amounts (Sec 5)
74 from several platforms and chemical transport models are also presented to further understand the
75 quality of satellite derived ozone profiles during the TROLIX-19 period.

76

77 2 Data and Methods

78 Descriptions of the various observational and model data sets and used in this study are below, including
79 a summary table (Table 1).

80 2.1 NASA Ozone Differential Absorption Lidars (DIAL)

81 NASA deployed and operated two ozone lidars during TROLIX-19 at the Cabauw site near the CESAR
82 tower to observe temporal and vertical gradients in tropospheric and stratospheric ozone. This was the
83 first dual-deployment of these lidars, in which the tropospheric ozone lidar measured between the near
84 surface (about 0.2 km) to a height of about 18 km and the stratospheric lidar during night-time from 15
85 km upwards to nearly 50 km, providing complete hybrid ozone profiles for the campaign period.

86 Measurements were made during periods of mostly clear skies, although occasional cloud cover did
87 enter the measurement period.

88 The NASA GSFC Mobile Stratospheric Ozone Lidar Trailer Experiment (STROZ-LITE) has been a
89 participant in the Network for the Detection of Atmospheric Composition Change (NDACC) since its
90 inception and is housed in a 12.5m container allowing for transport around the world (McGee et al.,
91 1991). The lidar instrument transmits two wavelengths, 308 nm from a XeCl excimer laser, and 355 nm
92 from a ND:YAG laser to derive ozone number density profiles, which have historically served as an
93 intercomparison data set for other NDACC ozone lidars (recent intercomparison can be found at Wing *et*
94 *al.*, 2020; 2021).

95 The NASA GSFC TROPOZ has been developed in a transportable 13.5m trailer to take routine
96 measurements of tropospheric ozone near the Baltimore–Washington, D.C. area as well as various
97 campaign locations (Sullivan *et al.*, 2014; 2015,2019, Leblanc *et al.*, 2018). This instrument, which
98 utilizes a ND:YAG laser and Raman cell, has been developed as part of the ground-based Tropospheric
99 Ozone Lidar NETwork (TOLNet, <https://www-air.larc.nasa.gov/missions/TOLNet/>), which currently
100 consists stations across the North America (<http://www-air.larc.nasa.gov/missions/TOLNet/>). The
101 primary purposes of the instruments within TOLNet are to provide regular, high-fidelity profile
102 measurements of ozone within the troposphere for satellite and model evaluation. This lidar also
103 operates routinely for the Network for the Detection of Atmospheric Composition Change (NDACC).

104 [More than thirty NDACC ground-based Lidar instruments](#)
105 [\(https://lidar.jpl.nasa.gov/ndacc/index_ndacc.php\)](https://lidar.jpl.nasa.gov/ndacc/index_ndacc.php) deployed worldwide from Pole to Pole are
106 [monitoring atmospheric ozone, temperature, aerosols, water vapor, and polar stratospheric clouds.](#)

107 Both lidars collect backscattered radiation with a large primary telescope and a 10cm telescope for near
108 field channels. Spectral separation is accomplished using dichroic beam-splitters and interference filters.
109 For the stratospheric system, five return wavelengths are recorded: the two transmitted wavelengths, and
110 the nitrogen Raman scattered radiation from each of the transmitted beams 332 nm and 382 nm, and the
111 408 nm water vapor channel. In this arrangement for TROLIX-19, the tropospheric system pumped the
112 Raman cell with the fourth harmonic (266 nm), which resulted in conversion to 289 nm and 299 nm
113 using a single hydrogen/deuterium Raman cell. All of the signals are further split to improve the
114 dynamic range of the respective lidar optical detection chains and are then amplified, discriminated and
115 recorded using photon counting techniques.

116 During TROLIX-19, the STROZ-LITE was operated on cloud free nights, with measurements lasting
117 between 2-4 hours to obtain enough signal to properly retrieve the entire stratospheric ozone profile. The
118 TROPOZ was operated during daytime and night time to provide tropospheric ozone profiles. For
119 instances of TROPOMI overpasses, campaign ozonesondes, or coincident stratospheric ozone lidar
120 measurements, the TROPOZ reported data is averaged for 30 minutes, centered around the satellite
121 overpass or launch time. This temporal period of averaging has been optimized in several cases to avoid
122 cloud contamination. For all other times during the TROPOZ operation, the data has been averaged to
123 10 minutes, which is suitable under most clear sky conditions to retrieve ozone information within the
124 entire troposphere. [A brief description and community standardized definitions of the uncertainty budget](#)
125 [of the lidar measurements presented in this paper can be found in Sullivan *et al.*, 2014, Leblanc *et al.*,](#)
126 [2016 and Leblanc *et al.*, 2018. The maximum statistical uncertainties for the two GSFC lidars vary from](#)
127 [night to night depending on atmospheric conditions and laser power fluctuations. They are mostly within](#)
128 [10-20% for 5 min and 5-8% for 30 min integrations throughout the atmosphere. Within overlapping](#)

129 measurement regions in the upper troposphere/lower stratosphere, they are different at the same altitude
130 due to laser performance and telescope/detector efficiency differences, and are therefore joined
131 manually for this work based on appropriate signal to noise and uncertainty estimates.

132 **2.2 Ground Based Passive Sensors and Ozonesondes**

133 **2.2.1 Pandora Spectrometer Instrument**

134 A Pandora spectrometer instrument (#118) has been used to measure columnar amounts of trace gases in
135 the atmosphere at 3–5-minute resolution at the Cabauw site since 2016 and previously used for the
136 second Cabauw Intercomparison of Nitrogen Dioxide (CINDI-2) campaign (Kreher et al., ~~2020b~~2020).
137 Using the theoretical solar spectrum as a reference, Pandora determines trace gas amounts using
138 differential optical absorption spectroscopy (DOAS). This attributes in principal these differences in
139 spectra measured by Pandora to the presence of trace gases within the atmosphere (*i.e.* the difference
140 between the theoretical solar spectrum and measured spectrum is caused by absorption of trace gas
141 species). For this study, L2 direct sun columnar values of ozone are used, although retrievals of nitrogen
142 dioxide are also operationally acquired. Data used passed the strictest QC/QA estimate (Flags = 10) and
143 was obtained from the Pandonia Global Network (<http://data.pandonia-global-network.org/>).

144 **2.2.2 Brewer MKIII Spectrophotometer**

145 A Brewer MKIII spectrometer instrument (#189) has been used to measure daily columnar amounts of
146 ozone in the atmosphere at the KNMI/De Bilt (30km NE of Cabauw) ~~site since 2007.~~, 52.10° N, 5.18°
147 E). Brewer #189 has been operated continuously since 1 October 2006. It replaced Brewer #100 which
148 provided observations since 1 January 1994. De Bilt has the longest continuous record of ozone
149 measured with an MKIII instrument in the World Ozone and Ultraviolet Radiation Data Centre
150 (WOUDC) database.

151 The Brewer is specifically designed to provide high accuracy measurement of spectrally resolved UV
152 for satellite evaluation, climatology monitoring and public health to international standards. Similar to
153 Pandora spectrometers, these measurements of total column of trace gases are compared to the measured
154 UV spectrum with the known solar output, and modeling the scattering properties of the atmosphere and
155 have been historically used to evaluate columnar satellite products (McPeters *et al.*, 2007; Wenig *et al.*,
156 2008; Garane *et al.*, 2019). The Brewer is the standard instrument used in the [World Meteorological
157 Organization](#) ozone monitoring network and for NDACC. This data was obtained at the NDACC
158 website (<https://www-air.larc.nasa.gov/missions/ndacc/data.html>).

159 **2.2.3 Ozonesondes**

160 Ozonesondes have been used to measure vertical profiles of ozone in the atmosphere at the KNMI/De
161 Bilt (30km NE of Cabauw) site since November 1992, and measurements are made weekly, historically
162 at 12 UTC on Thursdays. Description of the Electro Chemical Cell (ECC) details and metadata are
163 summarized in Malderen *et al.*, 2016, which also describes the importance of understanding and
164 reporting changes in ozonesonde operation procedures. During the ~~campagin~~[campaign](#), in situ
165 measurements of ozone were made using a balloon-borne payload consisting of an ECC ozonesonde
166 (Science Pump Corporation, [Serial Numbers: 6A35438, 6A35447, 6A35448, 6A35441](#)) coupled with a
167 radiosonde ([Vaisala RS41](#)) and have been used to evaluate TROPOMI tropospheric ozone products in
168 the tropics (Hubert *et al.*, 2021). The ECC technique is widely used for the high vertical resolution
169 measurements of O₃. The ECC consists of two chambers with platinum electrodes immersed in
170 potassium iodide (KI) solutions at different concentrations. The accuracy in the O₃ concentration
171 measured by an ECC ozonesonde is $\pm 5\%$ – 10% up to an altitude of 30 km (Smit *et al.*, 2007); [Smit and
172 Thompson *et al.*, 2021](#)). This data was obtained at the NDACC website ([https://www-
air.larc.nasa.gov/missions/ndacc/data.html](https://www-
173 air.larc.nasa.gov/missions/ndacc/data.html)).

174 2.3 Satellite Observations and Products

175 Satellite data used in this work was selected based on the closest retrieval (*i.e.* column, profile) to the
176 CESAR station within +/-2.5 degrees latitude and +/-10 degrees in longitude.

177 2.3.1 Ozone Mapping and Profiling Suite (OMPS) and MERRA-2 products

178 ~~Daily total column ozone overpasses over Cabauw station from the~~ The Ozone Mapping and
179 ~~Profiling Profiler~~ Suite (OMPS) ~~Nadir Mapper (NM) instrument~~ on the Suomi National Polar-orbiting
180 Partnership (S-NPP) platform ~~are used~~ consists of three sensors to measure the total column and the
181 vertical distribution of ozone with high spatial and vertical resolutions (Flynn et al., 2006). Daily total
182 column ozone overpasses over Cabauw station from the OMPS Nadir-Mapper (NM) instrument are used
183 in this study. The vertical distribution of ozone in the stratosphere and lower mesosphere is obtained
184 from the OMPS Limb-Profiler (LP) sensor on the Suomi-NPP satellite merging the UV (29.5-52.5 km)
185 and VIS (12.5-35.5 km) bands to provide a full profile from 12.5km to 52.5km (Kramarova *et al.*, 2018).
186 Variations of this merged OMPS-LP retrieval were considered, however the work shown in Arosio *et*
187 *al.*, 2018, indicates the same overall conclusions would be reached. Further work beyond this manuscript
188 may involve comparing this TROLIX-19 measurement data set to specific experimentally performed
189 satellite retrievals.

190 The Modern-Era Retrospective analysis for Research and Applications, Version 2 (MERRA-2) provides
191 data beginning in 1980 and since August 2004 assimilates NASA's satellite ozone profile observations
192 from Aura Microwave Limb Sounder (MLS) (Livesey et al, 2008) to more comprehensively characterize
193 stratospheric ozone abundance. A residual tropospheric ozone product (Ziemke *et al.*, 2019) is derived
194 using the OMPS NM total column ozone minus the co-located MERRA-2 stratospheric column ozone.
195 Tropopause pressure is derived from MERRA-2 potential vorticity (2.5 PVU) and potential temperature
196 (380 K).

197 **2.3.3 MLS**

198 NASA's Aura Microwave Limb Sounder (MLS) uses microwave emission to measure stratospheric
199 and upper tropospheric constituents, such as ozone. Ozone data (v5) used in this study is binned on
200 various vertical grids and are converted from volume mixing ratio to number density using the
201 coincident MERRA-2 atmosphere state parameters. Both daytime and nighttime data are used in this
202 study and the corresponding closest profile is utilized for comparison.

203 **2.3.4 TROPOMI**

204 In October 2017, the Sentinel-5 Precursor (S5P) mission was launched, carrying the TROPospheric
205 Monitoring Instrument (TROPOMI), which is a nadir-viewing 108° Field-of-View push-broom grating
206 hyperspectral spectrometer. Starting in August 2019, Sentinel-5P TROPOMI along-track high spatial
207 resolution (approximately 5.5 km at nadir) has been implemented and total ozone columns values used
208 in this work are subsetted from the NASA GES DISC
209 (https://tropomi.gesdisc.eosdis.nasa.gov/data/S5P_TROPOMI_Level2/S5P_L2__O3_TOT_HiR.1/) to
210 provide the Offline 1-Orbit L2 (S5P_L2__O3_TOT_HiR), which is based on the Direct-fitting algorithm
211 (S5P_TO3_GODFIT), comprising a non-linear least squares inversion by comparing the simulated and
212 measured backscattered radiances.

213 Tropospheric Ozone vertical profiles were retrieved using the TOPAS (Tikhonov regularized Ozone
214 Profile retrieval with SCIATRAN) algorithm and were applied to the TROPOMI L1B spectral data
215 version 2, using spectral data between 270 and 329 nm for the retrieval (Mettig *et al.*, 2021). This data
216 set will cover the TROLIX-19 period from 09 September until 28 September; however, it is available
217 outside of this work for specific weeks between June 2018 and October 2019. Since the ozone profiles
218 are very sensitive to absolute calibration at short wavelengths, a re-calibration of the measured radiances
219 is required using comparisons with simulated radiances with ozone limb profiles from collocated
220 satellites used as input. The a priori profiles for ozone are taken from the ozone climatology of Lamsal

221 et al. (2004) and the calibration correction spectrum is determined using the radiances modelled with
222 ozone information from collocated MLS/Aura measurements as described in depth throughout Mettig *et*
223 *al.*, 2021.

224

225 **2.4 Coupled Chemistry and Meteorology Model**

226 The GEOS Composition Forecasting (GEOS-CF, https://gmao.gsfc.nasa.gov/weather_prediction/GEOS-
227 [CF/](https://gmao.gsfc.nasa.gov/weather_prediction/GEOS-CF/), Keller *et al.*, 2021, Knowland et al., 2021) system was chosen to serve as a comparison simulation
228 for this effort, based on its altitude coverage (up to 80 km) and implications for future geostationary
229 satellite use. The system produces global, three-dimensional distributions of atmospheric composition
230 with a spatial resolution of 25km. Using meteorological analyses from other GEOS systems, the GEOS-
231 CF products include a running atmospheric replay to provide near-time estimates of surface pollutant
232 distributions and the composition of the troposphere and stratosphere. Individual case study evaluations
233 using ozone lidar of the GEOS-CF meteorological replay have recently been performed in Dacic et al.
234 (2020), Gronoff et al. (2021) and Johnson et al. (2021). These results will also be used to better evaluate
235 the GEOS-CF as the source of a priori ozone profiles for use in the TEMPO tropospheric ozone
236 retrievals. Model output for this work is used from the closest GEOS-CF model grid cell to the CESAR
237 observatory.

238 **3 Tropospheric Ozone Comparisons**

239 **3.1 Vertical Profiles**

240 Example tropospheric ozone profile observations are presented in **Figure 2** for 7 individual observation
241 periods during the TROLIX-19 campaign. Each of the panels show the cloud screened TROPOZ lidar
242 retrievals (top panels) and the corresponding GEOS-CF model output (bottom panels). Pink dots are
243 overlaid to indicate the simulated tropopause altitude based on a blended estimate (TROPPB) which

244 meets criteria of the lowest altitude bin corresponding with either a pressure level above the thermal
245 tropopause (380K) or dynamical (3 PVU) tropopause.

246 In general, the observations and simulations agree quite well in characterizing the broad features that
247 impacted the CESAR site during the TROLIX-19 campaign. However, in each panel there are ozone
248 laminae within the lower troposphere that are not replicated in the model simulation, most notably the
249 underestimation of ozone during the September 20-21 period from 3-5km (Figure 2d-e). ~~The~~
250 ~~observations indicate increased ozone levels as compared to the model during this period, centered~~
251 ~~around the 3-5km and 8-10km region of the atmosphere (this is explored in more detail below, black~~
252 ~~dashed box~~). However, the model does simulate well the lowered tropopause height and abundance of
253 lower stratospheric ozone observed in the 2 October observations (see Fig. 2g), which is an indication of
254 the model ~~well representing the dynamical variability that affects the lowering of the tropopause~~
255 ~~height~~ representing the dynamical variability that affects the lowering of the tropopause height. This
256 suggests the model is appropriately capturing the complex dynamics during this period near the upper
257 troposphere, but may not have been initialized with the correct boundary conditions or is too spatially
258 coarse to allow for simulation of the layer emphasized with the black dashed box. However, this is an
259 important altitude region for identifying long range transport of aged stratospheric air and inter-
260 continental transport that may be downward mixing towards the surface layer and will be explored in
261 more detail below.

262 To bring in additional platforms and to better understand these differences throughout the campaign at
263 discrete altitudes, **Figure 3** shows the ozone number density values for the TROPOZ lidar, GEOS-CF
264 model, TROPOMI and ECC ozonesondes at the average 4 km vertical level for the entire TROLIX-19
265 campaign period. Within ~~the 4km~~ this layer, the platforms are all characterizing the general ozone
266 features throughout the campaign at an altitude that frequently is associated with aged transported
267 layering laminae. There is a noticeable difference between the observations and model during the
268 previously described 20-21 September period. On 21 September at 12 UT, the lidar and ECC sonde

269 quantify an elevated layer ($1.2-1.3 \times 10^{21}$ molecules m^{-3}) -into the region that is not simulated by model
270 ($0.75-0.9 \times 10^{21}$ molecules m^{-3}), resulting in an approximately 30% ~~difference in ozone abundance within~~
271 ~~layer. Since the model correctly simulated many other ozone features during this time period within the~~
272 ~~upper tropospheric region, this is likely aged transport into the domain that was not available during~~
273 ~~model initialization. Back trajectories were performed to better identify the source of this air mass,~~
274 ~~however nothing conclusive can be reported. The layer is not associated with any increase in lidar~~
275 ~~attenuated backscatter within the associated altitude, suggesting it was not urban in origin and therefore~~
276 ~~more likely aged stratospheric air mixing down to the lower free troposphere. Outside of this Sep 21~~
277 ~~period, there is generally good agreement between the observations and model, indicating the~~
278 ~~combination of observations and modeling are able to represent the rural conditions and ozone~~
279 ~~perturbations at the CESAR site.~~underestimation in ozone abundance within the layer.

280 **3.2 Columnar Data Reduction**

281 There continues to be a need within the atmospheric and satellite community to understand the
282 variability of ozone as it pertains to both the tropospheric column (*i.e.* the Earth's surface to the
283 tropopause height) and the 0-2km tropospheric column (*i.e.* the Earth's surface to the 2 km height). The
284 0-2 km region is of particular interest as it is projected to be delivered hourly from the North American
285 geo-stationary satellite: Tropospheric Emissions: Monitoring of Pollution (TEMPO). Due to the
286 increased temporal frequency and vertical extent of TEMPO's tropospheric ozone retrievals, ozone
287 lidars, such as those from TOLNet (<https://www-air.larc.nasa.gov/missions/TOLNet/>) used in this work,
288 are an ideal platform to perform future evaluations of the products, ~~which build from Johnson *et al.*,~~
289 ~~2018.~~

290 Full tropospheric columns (**Figure 4, top panel**) are consistently calculated from each platform using
291 the blended tropopause height (TROPPB) produced by the GEOS-CF and described above (*c.f.* pink dots
292 in **Figure 2**) and are then converted to Dobson Units (DU). The tropospheric and 0-2km columns are

293 calculated explicitly by integrating the ozone number density from the lowest data bin of usable data to
294 the TROPB or 2km layer height produced in the nearest model temporal output. The exception to this
295 is the OMPS/MERRA-2 tropospheric column using the residual method described above (subtracting
296 the MERRA-2 stratospheric column from the OMPS-NM total ozone column). ~~For the 0-2km~~
297 ~~tropospheric column (Figure 4, bottom panel), there were no major surface layer pollution events at the~~
298 ~~CESAR observatory during TROLIX-19.~~

299 For the full tropospheric column (Figure 4, top panel), the campaign variability ranges from
300 approximately 20-55 DU; based on the lidar observations. The model, lidar, and ECC sonde
301 observations agree quite well throughout the 12 Sep to 23 Sep time frame when looking at day-to-day
302 variability. However, when assessing the variability on a single day for 21 Sep, full tropospheric
303 columns reduced from the lidar observations are some of the largest observed during this TROLIX-19
304 period (reaching nearly 5046 DU), ~~while the~~ but mostly staying between 34-40 DU. During this time,
305 the model mainly ranges between 35-37 DU ~~or a difference, resulting in differences within 10% for~~
306 most of ~~upwardsthe~~ observations (albeit closer to 30% for the peak on this day).

307 When looking at Figure 2d-f, the lower ozone values just below the tropopause during this period are not
308 simulated in the model, which may be an indication that the mesoscale ozone transport in the frontal
309 system is not very well resolved by the model for this specific event. Since the model correctly
310 simulated many other ozone features during this time period within the upper tropospheric region, this
311 may also be attributed to aged transport into the domain that was not available during model
312 initialization. Back-trajectories were performed to better identify the sources of 40%. Therefore, this
313 these air masses, however nothing conclusive can be reported. The layer is not associated with any
314 increase ~~identified and discussed in Fig 3, not only results an~~ in lidar attenuated backscatter within the
315 associated altitude-specific difference, but ultimately results in a large overall impact to the full,
316 suggesting it was not urban in origin and therefore more likely aged stratospheric air mixing down to the
317 lower free troposphere. Outside of this Sep 21 period, there is generally good agreement between the

318 observations (including the OMPS-MERRA2 product) and model, indicating the combination of
319 observations and modeling is able to represent the rural conditions and ozone perturbations at the
320 CESAR site.

321 When assessing these tropospheric column values from the TROPOMI ozone profile
322 observations, it is important to mention the vastly different vertical resolution or averaging kernel
323 schemes as compared to the independent observations near the tropopause. The ECC samples an
324 instantaneous observation with a vertical resolution generally less than 100m, while the lidar is
325 averaging over 500-750m of atmosphere for each data point near the tropopause. However, the vertical
326 resolution near the tropopause for TROPOMI using the TOPAS algorithm (Mettig *et al.*, 2021) is nearly
327 6 km, indicating it is not able to completely represent sharp gradients that may occur near the tropopause
328 layer and the lower stratosphere (where ozone content sharply increases). This lack of degrees of
329 independent information is evident in the relatively higher TROPOMI tropospheric column ozone values
330 as compared to the other independent measurements presented in this work. This suggests ground-based
331 profiling observations are still critically needed to confirm large deviations from a priori and
332 climatology in order to evaluate the atmospheric chemistry models, especially in the upper tropospheric
333 region and within the boundary layer.

334 There exists both diurnal and day-to-day variability of the 0-2 km ozone, ranging from 4-10 DU (Figure
335 4, ~~top~~bottom panel). In the 0-2 km ozone reduction, the lidar and model are critically needed to
336 understand ozone variability on a continuous scale. For instance, on 15 Sep the 0-2 km ozone column
337 was near 9 DU at 03 UT and finished near 5.5 DU at 16 UT, resulting in a -60% change in DU within a
338 ~~single day. Furthermore, the gradient of the 21 Sep ozone column change was similar in scale to the~~
339 ~~entire campaign variability, indicating that there is a significant amount of information gained in the~~
340 ~~understanding of the variability in ozone from continuous measurements. Although a daily snapshot of~~
341 ~~OMPS-MERRA-2 residuals and TROPOMI ozone profile observations are critical for their vast spatial~~
342 ~~coverage, ground-based observations such as ozone lidar and ECC sondes are critically needed to~~

343 ~~quantify measurement gaps.~~ 13 hours. The need for continuous measurements during highly variable
344 days are further emphasized by the fact that this gradient in 0-2km ozone for this single day (15 Sep, 5.5
345 DU – 9 DU) was comparable to the variance of 0-2 km ozone values throughout the entire campaign.
346 In summary, we find that the ozone columns evaluated in this study generally reproduced the structure
347 of the TROLIX-19 ozone lidar observations for N=835 coincidences. For the full tropospheric column,
348 the lidar calculated median was 30.9 ± 4.7 DU, compared to 33.4 ± 3.9 DU for the GEOS-CF. This
349 indicates a difference of 2.5 DU or 7.9 %, which is well within the lidar uncertainty of around 10 %
350 throughout the tropospheric column, and as we described above is likely driven by select days rather
351 than an overall bias between the measurements. For the 0-2 km tropospheric column, the lidar
352 calculated median was $5.8 \text{ DU} \pm 0.9 \text{ DU}$, compared to ~~6.9~~7.8 $\text{DU} \pm 0.7 \text{ DU}$ for the corresponding
353 GEOS-CF measurements. ~~This indicates a difference between observations and model of 1.1 DU or 18.9~~
354 ~~%, which is higher than the lidar uncertainty of around 5–10 % throughout the column.~~ For the
355 TROLIX-19 campaign, a 0-2 km tropospheric column accounts for approximately 20% of the
356 tropospheric column as detailed in Figure 4 (top panel), indicating measurements above the surface are
357 critically needed at understanding ozone variability at rural sites such as Cabauw, NL, where free
358 tropospheric ozone features dominate the column.

359 **4 Full Profile Ozone Comparisons**

360 **4.1 Hybrid Tropospheric/Stratospheric Ozone Comparisons**

361 To better understand differences in ozone retrievals from multiple platforms, it is important to assess the
362 entire vertical distribution of ozone. To characterize the vertical distribution throughout the entire
363 troposphere and stratosphere, hybrid ozone profiles were created from longer (integrations of 60-120
364 ~~minute~~minutes vs 10 minutes in Sec 3) temporal retrievals from the closed co-located daytime/nighttime
365 TROPOZ and nighttime STROZ ~~lidar~~lidar data, which were then interpolated to the GEOS-CF model
366 vertical grid levels. **Figure 5** compares these results to the GEOS-CF, OMPS-LP, TROPOMI, MLS and

367 the ECC ozonesonde profiles for 12 Sep, 17 Sep, 19 Sep, and 21 Sep 2019. These days were selected as
368 days within the campaign that had an ECC launch from De Bilt, NL (30 km from Cabauw).
369 For each observation period in **Figure 5**, all platforms manage to characterize a similar shape and extent
370 of the ozone maxima between 2.5-4.5 molecules m⁻³ throughout the vertical layer between 20-25 km. In
371 each case, there are differences between the platforms in characterizing the vertical variability and
372 extent of the ozone maxima, which will be quantified in the following section. One notable feature that
373 emphasizes the cross-platform ability to illustrate ozone variability in the stratosphere is from the 19 and
374 21 Sep profiles. A dual ozone maximum is observed quite remarkably by the merged lidar, ECC, MLS,
375 OMPS-LP and simulated by the GEOS-CF centered around 20 km and then again at 25km. The wind
376 observations from the ozonesonde payload (not shown) indicate a wind shear within the two ozone
377 layers, suggesting this feature was dynamically driven. The TROPOMI retrieval is not able to retrieve
378 this vertical features due to its coarser vertical resolution and appears to average through the layers.

379 **4.2 Difference Profiles**

380 To quantitatively compare the ozone retrievals and simulations, **Figure 6** displays the ozone
381 values for the TROLIX-19 time period from the hybrid lidar dataset (**Figure 6a**), GEOS-CF (**Figure**
382 **6b**), OMPS-LP (**Figure 6c**), MLS (**Figure 6d**) and TROPOMI (**Figure 6e**). This double ozone maxima,
383 starting after 20 September serves as a geophysical marker to visually compare the ozone products. The
384 lidar, model, and OMPS-LP all capture this feature, but with varying ozone abundances and altitudes.
385 From **Figure 6d**, it appears as if TROPOMI retrievals are not able to resolve this feature. The percent
386 differences, as compared to the lidar observations, are displayed in **Figure 7a-d**. These percent
387 differences are calculated using (1)

$$(1) \text{ Percent Difference} = \frac{(E_1 - E_2)}{\frac{1}{2}(E_1 + E_2)} \times 100$$

390
391

392 where E_2 are the lidar observations and E_1 are the respective ozone values from the various platforms in
393 **Figure 6**.

394 The percent differences in **Figure 7a** indicate the GEOS-CF model from 20-45 km generally represents
395 the lidar observations, but are generally 0-10 % lower in abundance. The percent differences in **Figure**
396 **7b** indicate OMPS-LP is also representing the ozone maxima and altitude above 25 km. There are larger
397 differences below 20 km, which indicates the OMPS-LP retrieval ~~is generally underestimating~~
398 ~~how~~ worsens (in both directions) as compared to the ozone abundance below 20 km as shown in the
399 profiles in **Figure 5**. The percent differences in **Figure 7c** indicate the MLS data, especially that within
400 the 20-40km region, perform quite well as compared to the lidar observations. The percent differences in
401 **Figure 7d** indicate the TROPOMI retrieval is generally over representing the ozone concentrations
402 throughout the atmosphere, which worsens within the troposphere- and has been discussed earlier for
403 the tropospheric ozone column as a result of a much larger vertical resolution in this region. In all cases,
404 the most variability in the differences occur within the active region from 10-20 km that is driven by the
405 dynamical tropopause height and lower stratospheric ozone abundance. Within each satellite dataset, we
406 find larger biases in the lower stratosphere and upper troposphere below 18km, which has been
407 previously described in the literature for the OMPS-LP dataset in Kramarova *et al.*, 2018 and were
408 improved in the updated version 2.5 algorithm used in this work.

409 **5 Total Column Ozone**

410 Similar to the troposphere, to better understand to what extent the vertical distribution of ozone impacts
411 the atmospheric column, **Figure 8 (top panel)** shows the various platforms and their retrieved total
412 column ozone. For this analysis, the GEOS-CF, lidar, OMPS-NM, TROPOMI (GODFIT) are shown, in
413 addition to local ground-based measurements from a Pandora instrument and Brewer. The total column
414 values range from 230-300 DU throughout the campaign period, with the median total column ozone of
415 271 DU. With the previous analyses from Sec 3.2, this indicates the median total tropospheric column of

416 33 DU and 0-2km boundary layer column of 6 DU result in percentages of the entire ozone column of
417 12% and 2.3%, respectively. Similar to the full tropospheric ozone columns, larger total ozone columns
418 were observed towards the end of the TROLIX-19 period, suggesting this variability was partly due to a
419 larger abundance of ozone in the lower stratosphere.

420 **Figure 8 (bottom panel)** shows the various platforms as a percent differences from the model. In
421 general, the various platforms are all mostly within ±05 % of each other, with most differences being
422 within ±53%. This analysis emphasizes the stability and maturity of the Pandora and Brewer systems for
423 monitoring the total column ozone amounts. Interestingly, the double maxima feature in vertical ozone
424 distribution in the stratosphere (with local minima between) described in Sec 4.1 on 21 Sep does not
425 severely impact the total column ozone.

426 **6 Conclusions**

427 This work has highlighted the various differences in retrieved ozone quantities during the TROLIX-19
428 campaign. This has emphasized the importance of ground-based ozone lidars and other measurements in
429 understanding the vertical variability of ozone and how it relates to the column reduction. This work
430 also shows the first effort to directly resolve both tropospheric columns and 0-2km ozone columns from
431 the NASA TROPOZ lidar. Other TOLNet lidars are able to perform this data reduction and future work
432 will be to expand this effort to the other TOLNet locations. This work indicates the level of performance
433 of the GEOS-CF modeling system as compared to the other platforms, which ultimately performs
434 extremely well both in the stratosphere (Figure 6 and Figure 7) and within the troposphere, ~~as~~
435 emphasized (Figure 2 and Figure 4).

436 One takeaway message or point of caution for future efforts is that although there are situations
437 identified where the vertical profile and the model disagree in Figure 6 and Figure 7, a certain altitude
438 range (Figure 3), when the data is reduced to a columnar product, compensating over/under-estimations
439 may cancel out and produce a more accurate value when only looking at the resultant as compared to

440 observations. For this reason, it is essential when doing data columnar reduction for the troposphere, and
441 even more so in the 0-2km column or planetary boundary layer, that observations of the vertical profile
442 be used to evaluate the representativeness of the model and auxillary data sets.

443 In looking towards the NASA TEMPO ~~mission~~, this work indicates that the GEOS-CF, with its global
444 coverage, hourly resolution, and adequate vertical information to resolve most atmospheric features, is
445 an appropriate choice for the a priori profiles for the TEMPO ozone retrievals. Continued investigations
446 are needed with high resolution observations, as presented in this work, to better evaluate the GEOS-CF,
447 especially in these common transport regions of the atmosphere. Although the GEOS-CF performed
448 well in reproducing the ozone downward transport throughout the upper troposphere and lower
449 stratosphere, the model did fail to resolve some high-resolution laminae deeper into the lower
450 troposphere related to specific mesoscale ozone transport in this region as evidenced in Figure 2 and
451 Figure 3.

452 This work shows the TROPOMI TOPAS ozone profile algorithm products are able to accurately
453 reproduce ozone quantities in the lower troposphere at various atmospheric levels. In particular, **Figure**
454 **3** shows and Figure 4 show promising results that indicate the TROPOMI satellite observations compare
455 well with the observations from ground-based measurements (lidar, sonde) of specific elevated ozone
456 features. However, there is an observed overestimate of the TROPOMI retrieval in the upper
457 troposphere and lower stratosphere (between 10 and 15 km) associated with a larger vertical resolution
458 that needs to also be further evaluated to better understand the representativeness of the retrieval in this
459 region.

460
461 Figure 7 was presented as a quantitative resultant figure to illustrate both the temporal (i.e., throughout
462 the course of the TROLIX-19 campaign) and vertical differences observed in the retrievals from each
463 observational platform. This serves as a rare opportunity to cross-evaluate multiple satellited based
464 observations, a global chemical transport model, ozonesondes and a high-resolution ozone lidar suite.

465 The authors feel that this figure has served to point out the strengths of each platform and present careful
466 considerations for areas of under/over estimation, Furthermore, we feel reducing these comparisons
467 down to a specific percentage may underserve the community push for supporting the vertical profiling
468 needed for these types of efforts.

469 The CESAR Observatory continues to be a critical landmark for campaigns that revolve around
470 atmospheric composition measurements for satellite validation and evaluation beyond this effort, such as
471 CINDI and CINDI-2 (Kreher *et al.*, [2020](#); Wang *et al.*, 2020; Tirpitz *et al.*, 2021). As the European
472 Commission (EC) in partnership with the European Space Agency (ESA) continues to launch
473 tropospheric composition satellites, including the upcoming geo-stationary Sentinel-4 satellite, we
474 expect this observatory will continue to host and maintain critical atmospheric sampling for future
475 validation efforts.

476

477

478

479 **Data Availability.**

480

481 1. MLS ozone profiles can be downloaded from the NASA Goddard Space Flight Center Earth
482 Sciences Data and Information Services Center (GES DISC; Schwartz *et al.*, 2020,
483 <https://doi.org/10.5067/Aura/MLS/DATA2516>, last access: 29 March 2022).

484 2. The Pandora data is available at the Pandonia Global Network Archive ([http://data.pandonia-](http://data.pandonia-global-network.org/Cabauw/)
485 [global-network.org/Cabauw/](http://data.pandonia-global-network.org/Cabauw/), last access 29 March 29, 2022).

486 3. The OMPS LP version 2.5 ozone profiles can be downloaded from the NASA Goddard Space
487 Flight Center Earth Sciences Data and Information Services Center (GES DISC;
488 at <https://doi.org/10.5067/X1Q9VA07QDS7> (Deland, 2017, last access: 29 March 2022).

489 4. The tropospheric ozone lidar data used in this publication were obtained from the Cabauw
490 Experimental Site for Atmospheric Research (CESAR) as part of a campaign involving the

- 491 Network for the Detection of Atmospheric Composition Change (NDACC) and NASA's
492 Tropospheric Ozone Lidar Network (TOLNet) and are publicly available ([https://www-
493 air.larc.nasa.gov/cgi-bin/ArcView.1/TOLNet?NASA-GSFC=1](https://www-air.larc.nasa.gov/cgi-bin/ArcView.1/TOLNet?NASA-GSFC=1), last access: 29 March 2022).
- 494 5. The ozonesonde and Brewer data used in this publication were obtained from the De Bilt, NL
495 site as part of a campaign involving the Network for the Detection of Atmospheric Composition
496 Change (NDACC) and are publicly available (<ftp://ftp.cpc.ncep.noaa.gov/ndacc/station/debilt/>,
497 last access: 29 March 2022).
- 498 6. The stratospheric ozone lidar data used in this publication were obtained from the Cabauw
499 Experimental Site for Atmospheric Research (CESAR) as part of a campaign involving the
500 Network for the Detection of Atmospheric Composition Change (NDACC) and are publicly
501 available (<ftp://ftp.cpc.ncep.noaa.gov/ndacc/station/cabauw/>, last access: 29 March 2022).
- 502 7. The TROPOMI TOPAS Ozone Profile data and source codes are available upon request from
503 Nora Mettig (mettig@iup.physik.uni-bremen.de) or Mark Weber (weber@uni-bremen.de). The
504 L1B version of the S5P data is available upon request to the S5P Validation Team.
- 505 8. The Tropospheric Ozone Column from OMPS-NM/MERRA-2 Daily measurements data are
506 available upon request from Jerry Ziemke (Jerald.r.ziemke@nasa.gov).
- 507 9. The NASA GEOS-CF simulations are available at the data sharing portal
508 (<https://portal.nccs.nasa.gov/datashare/gmao/geos-cf/v1/forecast/>, last access 29 March 2022).

509

510 Author contributions. JS drafted the original manuscript. JS, LT, GS, and TM deployed and operated the
511 NASA ozone lidars and provided expertise on use of measurements. NM, AR, and MW provided
512 TOPAS ozone profile data and guidance on how best to use the measurements. AA and KK provided
513 overall context as principal investigators of the TROLIX-19 campaign and coordinated science team
514 meetings to foster this collaboration. KEK provided GEOS-CF data and insight on its use in this work.
515 MA, AP, MvR, and PV provided expertise and data for the ground-observations for ozonesondes,

516 Brewer, and historical data for the Cabauw site. JZ provided data for the OMPS-MERRA-2 tropospheric
517 column data. NK provided Aura MLS, OMPS-LP merged data and further insight into the use of the
518 data.

519

520

521 Competing interests. The authors declare that they have no conflict of interest.

522

523 Disclaimer. Publisher's note: Copernicus Publications remains neutral with regard to jurisdictional
524 claims in published maps and institutional affiliations.

525

526 Acknowledgements. NASA data has been provided through the Tropospheric Composition and Upper
527 Atmosphere Research Programs. We acknowledge all additional data providers and their funding
528 agencies for performing regular measurements and retrievals.

529

530

531

532

533

534

535

536

537

538

539

540

541

542

543

544

545

546

547

548

549

550

552 **References**

553

554 Apituley, Arnoud, Karin Kreher, Michael Van Roozendaal, John Sullivan, Thomas J. McGee, Marc
 555 Allaart, Ankie PETERS et al. "Overview of activities during the 2019 TROPOMI validation experiment
 556 (TROLIX'19)." In *AGU Fall Meeting Abstracts*, vol. 2019, pp. A43J-2958. 2019

557

558 Apituley, Arnoud, Karin Kreher, Ankie PETERS, John Sullivan, Michel van Roozendaal, Tim Vlemmix,
 559 Mirjam den Hoed et al. "Overview of the 2019 Sentinel-5P TROPOMI validation experiment
 560 (TROLIX)." In *EGU General Assembly Conference Abstracts*, p. 10539. 2020.

561

562 Arosio, Carlo, Alexei Rozanov, Elizaveta Malinina, Kai-Uwe Eichmann, Thomas von Clarmann, and
 563 John P. Burrows. "Retrieval of ozone profiles from OMPS limb scattering observations." *Atmospheric
 564 Measurement Techniques* 11, no. 4 (2018): 2135-2149.

565

566 Copernicus Sentinel data processed by ESA, German Aerospace Center (DLR) (2019), Sentinel-5P
 567 TROPOMI Total Ozone Column 1-Orbit L2 5.5km x 3.5km, Greenbelt, MD, USA, Goddard Earth
 568 Sciences Data and Information Services Center (GES DISC), Accessed: [10 December 2021],
 569 [10.5270/S5P-fqouvyz](https://doi.org/10.5270/S5P-fqouvyz)

570

571 Dacic, Natasha, John T. Sullivan, K. Emma Knowland, Glenn M. Wolfe, Luke D. Oman, Timothy A.
 572 Berkoff, and Guillaume P. Gronoff. "Evaluation of NASA's high-resolution global composition
 573 simulations: Understanding a pollution event in the Chesapeake Bay during the summer 2017 OWLETS
 574 campaign." *Atmospheric Environment* 222 (2020): 117133.

575

576 [Flynn, L. E., Seftor, C. J., Larsen, J. C., and Xu, P.: The Ozone Mapping and Profiler Suite, in: Earth
 577 Science Satellite Remote Sensing, edited by: Qu, J. J., Gao, W., Kafatos, M., Murphy, R. E., and
 578 Salomonson, V. V., Springer, Berlin, 279–296, doi:10.1007/978-3-540-37293-6, 2006.](#)

579

580 Garane, Katerina, Maria-Elissavet Koukouli, Tjil Verhoelst, Christophe Lerot, Klaus-Peter Heue, Vitali
 581 Fioletov, Dimitrios Balis et al. "TROPOMI/S5P total ozone column data: global ground-based
 582 validation and consistency with other satellite missions." *Atmospheric Measurement Techniques* 12, no.
 583 10 (2019): 5263-5287.

584

585 Gronoff, G., T. Berkoff, K. E. Knowland, L. Lei, M. Shook, B. Fabbri, W. Carrion, and A. O. Langford.
 586 "Case study of stratospheric intrusion above Hampton, Virginia: lidar-observation and modeling
 587 analysis." *Atmospheric Environment* (2021): 118498.

588

589 Hubert, Daan, Klaus-Peter Heue, Jean-Christopher Lambert, Tjil Verhoelst, Marc Allaart, Steven
 590 Compennolle, Patrick D. Cullis et al. "TROPOMI tropospheric ozone column data: geophysical
 591 assessment and comparison to ozonesondes, GOME-2B and OMI." *Atmospheric Measurement
 592 Techniques* 14, no. 12 (2021): 7405-7433.

593

594 Johnson, M. S., Liu, X., Zoogman, P., Sullivan, J., Newchurch, M. J., Kuang, S., Leblanc, T., and
 595 McGee, T.: Evaluation of potential sources of a priori ozone profiles for TEMPO tropospheric ozone
 596 retrievals, *Atmos. Meas. Tech.*, 11, 3457–3477, <https://doi.org/10.5194/amt-11-3457-2018>, 2018.

597

598 Keller, Christoph A., K. Emma Knowland, Bryan N. Duncan, Junhua Liu, Daniel C. Anderson, Sampa
599 Das, Robert A. Lucchesi et al. "Description of the NASA GEOS Composition Forecast Modeling
600 System GEOS-CF v1. 0." *Journal of Advances in Modeling Earth Systems* 13, no. 4 (2021):
601 e2020MS002413.
602

603 Kramarova, Natalya A., Pawan K. Bhartia, Glen Jaross, Leslie Moy, Philippe Xu, Zhong Chen, Matthew
604 DeLand et al. "Validation of ozone profile retrievals derived from the OMPS LP version 2.5 algorithm
605 against correlative satellite measurements." *Atmospheric Measurement Techniques* 11, no. 5 (2018):
606 2837-2861.
607

608 Kreher, Karin, Michel Van Roozendaal, Francois Hendrick, Arnoud Apituley, Ermioni Dimitropoulou,
609 Udo Frieß, Andreas Richter et al. "Intercomparison of NO₂, O₄, O₃ and HCHO slant column
610 measurements by MAX-DOAS and zenith-sky UV–visible spectrometers during CINDI-2."
611 *Atmospheric Measurement Techniques* 13, no. 5 (2020): 2169-2208.
612

613 Lamsal, L. N., M. Weber, S. Tellmann, and J. P. Burrows. "Ozone column classified climatology of
614 ozone and temperature profiles based on ozonesonde and satellite data." *Journal of Geophysical*
615 *Research: Atmospheres* 109, no. D20 (2004).
616

617 Leblanc, Thierry, Mark A. Brewer, Patrick S. Wang, Maria Jose Granados-Muñoz, Kevin B.
618 Strawbridge, Michael Travis, Bernard Firanski et al. "Validation of the TOLNet lidars: the Southern
619 California Ozone Observation Project (SCOOP)." *Atmospheric measurement techniques* 11, no. 11
620 (2018): 6137-6162.
621

622 [Leblanc, Thierry, Robert J. Sica, Joanna AE Van Gijssel, Sophie Godin-Beekmann, Alexander Haefele,](#)
623 [Thomas Trickl, Guillaume Payen, and Gianluigi Liberti. "Proposed standardized definitions for vertical](#)
624 [resolution and uncertainty in the NDACC lidar ozone and temperature algorithms–Part 2: Ozone DIAL](#)
625 [uncertainty budget." *Atmospheric Measurement Techniques* 9, no. 8 \(2016\): 4051-4078.](#)
626

627 Livesey, N. J., M. J. Filipiak, L. Froidevaux, W. G. Read, A. Lambert, M. L. Santee, J. H. Jiang et al.
628 "Validation of Aura Microwave Limb Sounder O₃ and CO observations in the upper troposphere and
629 lower stratosphere." *Journal of Geophysical Research: Atmospheres* 113, no. D15 (2008).
630

631 Malderen, Roeland Van, Marc AF Allaart, Hugo De Backer, Herman GJ Smit, and Dirk De Muer. "On
632 instrumental errors and related correction strategies of ozonesondes: possible effect on calculated ozone
633 trends for the nearby sites Uccle and De Bilt." *Atmospheric Measurement Techniques* 9, no. 8 (2016):
634 3793-3816.
635

636 McGee, Thomas J., David N. Whiteman, Richard A. Ferrare, James J. Butler, and John F. Burris.
637 "STROZ LITE: stratospheric ozone lidar trailer experiment." *Optical Engineering* 30, no. 1 (1991): 31-
638 39.
639

640 McPeters, Richard D., Gordon J. Labow, and Jennifer A. Logan. "Ozone climatological profiles for
641 satellite retrieval algorithms." *Journal of Geophysical Research: Atmospheres* 112, no. D5 (2007).
642

643 Mettig, N., Weber, M., Rozanov, A., Arosio, C., Burrows, J. P., Veefkind, P., Thompson, A. M., Querel,
644 R., Leblanc, T., Godin-Beekmann, S., Kivi, R., and Tully, M. B.: Ozone profile retrieval from nadir
645 TROPOMI measurements in the UV range, *Atmos. Meas. Tech.*, 14, 6057–6082,
646 <https://doi.org/10.5194/amt-14-6057-2021>, 2021.
647

648 Mettig, Nora, Mark Weber, Alexei Rozanov, John P. Burrows, Pepijn Veefkind, Nadia Smith, Anne M.
649 Thompson et al. "Combined UV and IR ozone profile retrieval from TROPOMI and CrIS
650 measurements." *Atmospheric Measurement Techniques Discussions* (2021): 1-33.
651

652 Peters, A. J. M., Boersma, K. F., Kroon, M., Hains, J. C., Van Roozendaal, M., Wittrock, F., Abuhassan,
653 N., Adams, C., Akrami, M., Allaart, M. A. F., Apituley, A., Beirle, S., Bergwerff, J. B., Berkhout, A. J.
654 C., Brunner, D., Cede, A., Chong, J., Clémer, K., Fayt, C., Frieß, U., Gast, L. F. L., Gil-Ojeda, M.,
655 Goutail, F., Graves, R., Griesfeller, A., Großmann, K., Hemerijckx, G., Hendrick, F., Henzing, B.,
656 Herman, J., Hermans, C., Hoexum, M., van der Hoff, G. R., Irie, H., Johnston, P. V., Kanaya, Y., Kim,
657 Y. J., Klein Baltink, H., Kreher, K., de Leeuw, G., Leigh, R., Merlaud, A., Moerman, M. M., Monks, P.
658 S., Mount, G. H., Navarro-Comas, M., Oetjen, H., Pazmino, A., Perez-Camacho, M., Peters, E., du
659 Piesanie, A., Pinardi, G., Puentedura, O., Richter, A., Roscoe, H. K., Schönhardt, A., Schwarzenbach,
660 B., Shaiganfar, R., Sluis, W., Spinei, E., Stolk, A. P., Strong, K., Swart, D. P. J., Takashima, H.,
661 Vlemmix, T., Vrekoussis, M., Wagner, T., Whyte, C., Wilson, K. M., Yela, M., Yilmaz, S., Zieger, P.,
662 and Zhou, Y.: The Cabauw Intercomparison campaign for Nitrogen Dioxide measuring Instruments
663 (CINDI): design, execution, and early results, *Atmos. Meas. Tech.*, 5, 457–485, doi:10.5194/amt-5-457-
664 2012, 2012.
665

666 [Smit, H. G. J., Thompson, A. M., & the Panel for the Assessment of Standard Operating Procedures for](#)
667 [Ozonesondes, v2.0 \(ASOPOS 2.0\) \(2021\). Ozonesonde Measurement Principles and Best Operational](#)
668 [Practices. World Meteorological Organization, GAW Report 268. \[Available](#)
669 [at https://library.wmo.int/doc_num.php?explnum_id=10884](https://library.wmo.int/doc_num.php?explnum_id=10884).
670

671 [Smit, Herman GJ, Wolfgang Straeter, Bryan J. Johnson, Samuel J. Oltmans, Jonathan Davies, David](#)
672 [W. Tarasick, Bruno Hoegger et al. "Assessment of the performance of ECC-ozonesondes under quasi-](#)
673 [flight conditions in the environmental simulation chamber: Insights from the Juelich Ozone Sonde](#)
674 [Intercomparison Experiment \(JOSIE\)." *Journal of Geophysical Research: Atmospheres* 112, no. D19](#)
675 [\(2007\)](#).
676

677 Sullivan, J. T., T. J. McGee, G. K. Sumnicht, L. W. Twigg, and R. M. Hoff. "A mobile differential
678 absorption lidar to measure sub-hourly fluctuation of tropospheric ozone profiles in the Baltimore–
679 Washington, DC region." *Atmospheric Measurement Techniques* 7, no. 10 (2014): 3529-3548.
680

681 Sullivan, John T., Thomas J. McGee, Anne M. Thompson, R. Bradley Pierce, Grant K. Sumnicht,
682 Laurence W. Twigg, Edwin Eloranta, and Raymond M. Hoff. "Characterizing the lifetime and
683 occurrence of stratospheric-tropospheric exchange events in the rocky mountain region using high-
684 resolution ozone measurements." *Journal of Geophysical Research: Atmospheres* 120, no. 24 (2015):
685 12410-12424.
686

687 Sullivan, John T., Timothy Berkoff, Guillaume Gronoff, Travis Knepp, Margaret Pippin, Danette Allen,
688 Laurence Twigg et al. "The ozone water–land environmental transition study: An innovative strategy for
689 understanding Chesapeake Bay pollution events." *Bulletin of the American Meteorological Society* 100,
690 no. 2 (2019): 291-306.
691

692 Tirpitz, Jan-Lukas, Udo Frieß, François Hendrick, Carlos Alberti, Marc Allaart, Arnoud Apituley, Alkis
693 Bais et al. "Intercomparison of MAX-DOAS vertical profile retrieval algorithms: studies on field data
694 from the CINDI-2 campaign." *Atmospheric Measurement Techniques* 14, no. 1 (2021): 1-35.
695

696 [van Geffen, J.H.G.M., Eskes, H.J., Boersma, K.F., Maasakkers, J.D. and Veeffkind, J.P., TROPOMI](#)
697 [ATBD of the total and tropospheric NO2 data products, Report S5P-KNMI-L2-0005-RP, KNMI, De](#)
698 [Bilt, The Netherlands; see TROPOMI ATBD list for the latest available version.](#)
699
700 [Wang, Yang, Arnoud Apituley, Alkiviadis Bais, Steffen Beirle, Nuria Benavent, Alexander Borovski,](#)
701 [Ilya Bruchkouski et al. "Inter-comparison of MAX-DOAS measurements of tropospheric HONO slant](#)
702 [column densities and vertical profiles during the CINDI-2 campaign." *Atmospheric Measurement*](#)
703 [*Techniques* 13, no. 9 \(2020\): 5087-5116.](#)
704
705 Wenig, Mark O., A. M. Cede, E. J. Bucsela, E. A. Celarier, K. F. Boersma, J. P. Veeffkind, E. J.
706 Brinksma, J. F. Gleason, and J. R. Herman. "Validation of OMI tropospheric NO2 column densities
707 using direct-Sun mode Brewer measurements at NASA Goddard Space Flight Center." *Journal of*
708 *Geophysical Research: Atmospheres* 113, no. D16 (2008).
709
710 Wing, Robin, Wolfgang Steinbrecht, Sophie Godin-Beekmann, Thomas J. McGee, John T. Sullivan,
711 Grant Sumnicht, Gérard Ancellet, Alain Hauchecorne, Sergey Khaykin, and Philippe Keckhut.
712 "Intercomparison and evaluation of ground-and satellite-based stratospheric ozone and temperature
713 profiles above Observatoire de Haute-Provence during the Lidar Validation NDACC Experiment
714 (LAVANDE)." *Atmospheric Measurement Techniques* 13, no. 10 (2020): 5621-5642.
715
716 Wing, Robin, Sophie Godin-Beekmann, Wolfgang Steinbrecht, Thomas J. Mcgee, John T. Sullivan,
717 Sergey Khaykin, Grant Sumnicht, and Laurence Twigg. "Evaluation of the new DWD ozone and
718 temperature lidar during the Hohenpeißenberg Ozone Profiling Study (HOPS) and comparison of results
719 with previous NDACC campaigns." *Atmospheric Measurement Techniques* 14, no. 5 (2021): 3773-3794.
720
721 Ziemke, Jerry R., Luke D. Oman, Sarah A. Strode, Anne R. Douglass, Mark A. Olsen, Richard D.
722 McPeters, Pawan K. Bhartia et al. "Trends in global tropospheric ozone inferred from a composite
723 record of TOMS/OMI/MLS/OMPS satellite measurements and the MERRA-2 GMI simulation."
724 *Atmospheric Chemistry and Physics* 19, no. 5 (2019): 3257-3269.
725
726
727
728
729
730
731
732
733
734
735
736
737
738
739
740
741
742
743
744
745

746
747
748
749
750
751
752
753

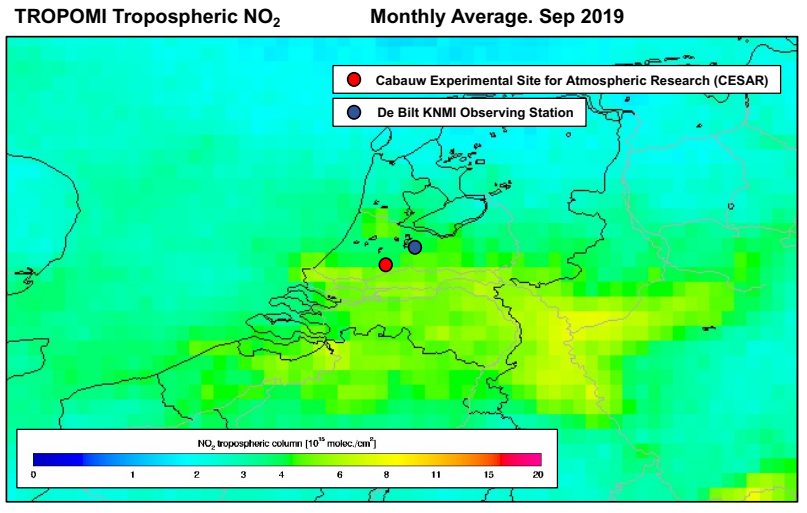
754 **Table 1: Instrument platforms, associated products, and short description used in this work during the TROLIX-19**
755 **campaign.**

756

Instrument	Products	Platform	Description
GSFC TROPOZ [NASA]	Profiles [0.2 – 18 km]	Ground-based Lidar	10 min integration; 30-90-min avg around ECC or Satellite Overpass
GSFC STROZ [NASA]	Profiles [15 - 48 km]	Ground-based Lidar	~2-4-hr avg between (20-23 UT)
ECC Ozonesondes [KNMI]	Profiles [0 – 33 km]	Balloonborne	Balloonborne, Launched at 12 UT from De Bilt (~30 km from Cabauw) on 4 days
Pandora [NASA/KNMI]	Column [TCO]	Spectrometer	L2 Pandora 118s, Data Used has QC/QA Flags = 10
Brewer [KNM]	Column [TCO]	Spectrophotometer	L2 Brewer #189m, MKIII, Located in De Bilt
S5P/TropOMI [ESA]	Column [TCL]	Satellite	L2 TOPAS Product, Overpass between 12-14 UT (5.5x3.5 km, nadir)
S5P/TropOMI [KNMI]	Column [TCO]	Satellite	L2 GODFIT v4 TO3 Product, Overpass between 12-14 UT (5.5x3.5 km, nadir)
OMPS [NASA]	Column [TCO]	Satellite	L3 NM Product, Version 2, Daily Overpass between 12-14 UT (50x50 km, nadir)
OMPS-LP [NASA]	Profiles [12-60km]	Satellite	Merged L2 v2.5 Daily Merged Product, Overpass between 12-14 UT (1km vertical bins)
OMPS/MERRA-2 [NASA]	Trop. Columns	Satellite/Assimilation	L4 Derived Product, OMPS-NM daily Overpass, MERRA-2
AURA MLS [NASA]	Profiles [12-60km]	Satellite	Merged L2 v5 Daily Daytime/Nighttime Products, Overpass between 12-14 UT (1km vertical bins) and 01-03 UT.
GEOS-CF [NASA]	Profiles [0-80km]	Global 3-D CCMM	<u>1-Hr, 72 lev, Met. Replay, (25x25km)</u>

757

758
759
760
761
762
763
764
765
766
767



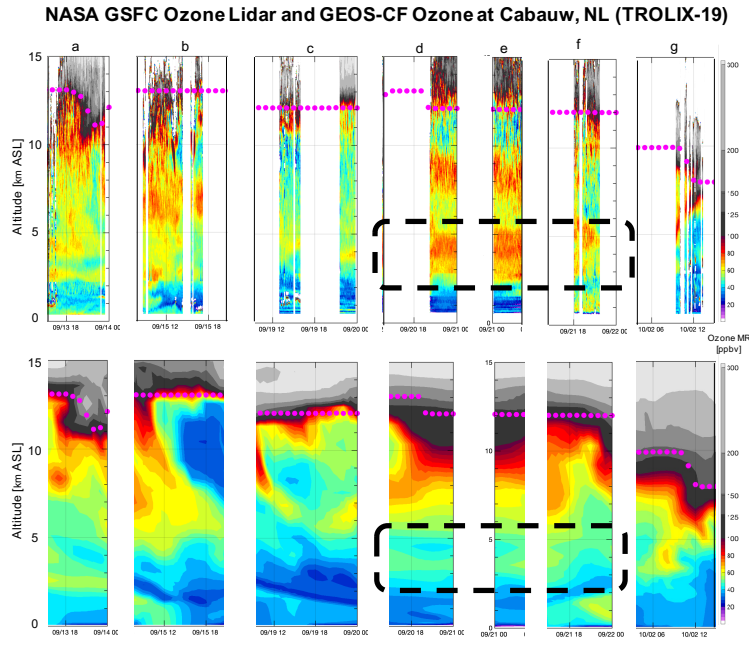
768
769
770

771 **Figure 1: ~~Aqua/MODIS True Color Corrected Reflectance (left) and TROPOMI Tropospheric~~ monthly-mean**
772 **tropospheric NO₂ (right column (version 1.0)) for 14-September 2019. The CESAR site is and De Bilt, NL sites are**
773 **indicated in the image on the left.**

774
775
776

777
778
779
780
781
782
783
784
785
786
787
788
789

790
791
792
793
794
795

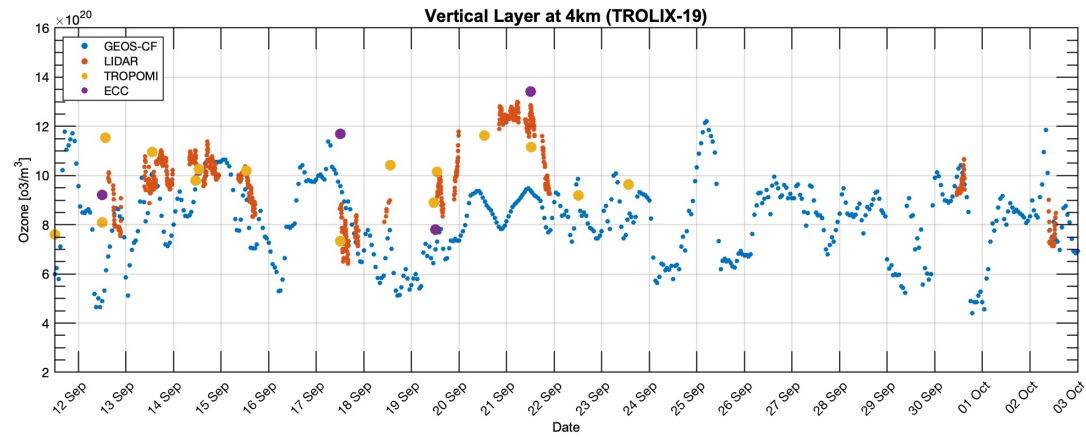


796

797 **Figure 2: Cloud screened TROPOZ lidar retrievals (top panel) and the corresponding GEOS-CF model output (bottom**
798 **panel) from the closest model grid cell to the CESAR observatory during TROLIX-19 for a) 13 Sep 14-00 UTC, b) 15**
799 **Sep 09-21 UTC, c) 19 Sep 10-00 UT, d) 20 Sep 16-00 UT, e) 21 Sep 0-3 UT, f) 21 Sep 16-00UT, and g) 02 Oct 04-14 UT.**
800 **Pink dots are overlaid to indicate the simulated tropopause altitude based on a blended estimate (TROP PB).**

801
802
803
804
805
806
807
808
809
810
811
812
813
814
815
816
817

818
819
820
821
822
823
824
825
826
827
828
829
830

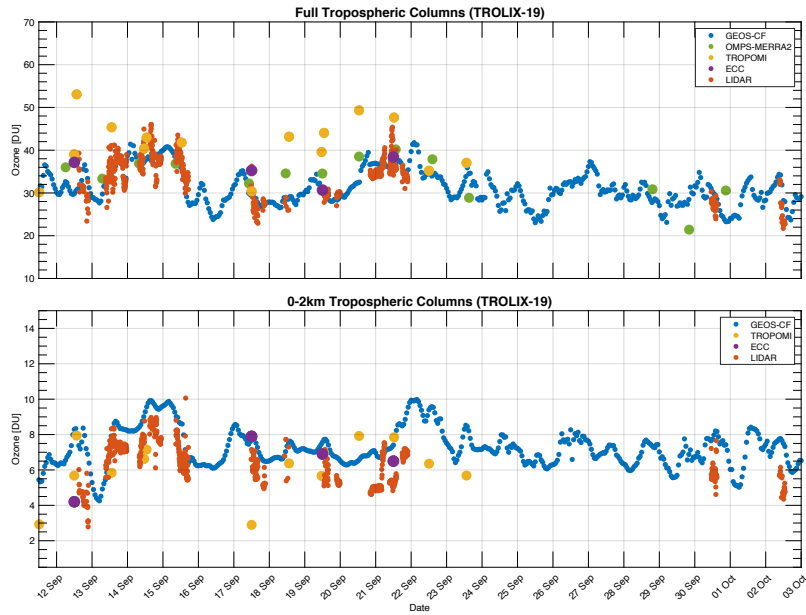


831
832

833 **Figure 3: Ozone number density values for the TROPOZ lidar, GEOS-CF mode, TROPOMI and electro-chemical cell**
834 **(ECC) ozonesondes at the 4km layers/levels. The layer was calculated to match the closest representative vertical layer**
835 **of the GEOS-CF for consistent intercomparison. Data is averaged in a 500m layer from 3.94 km to 4.44 km AGL.**

836
837
838
839
840
841
842
843

844
845
846
847
848
849
850
851
852
853
854
855
856

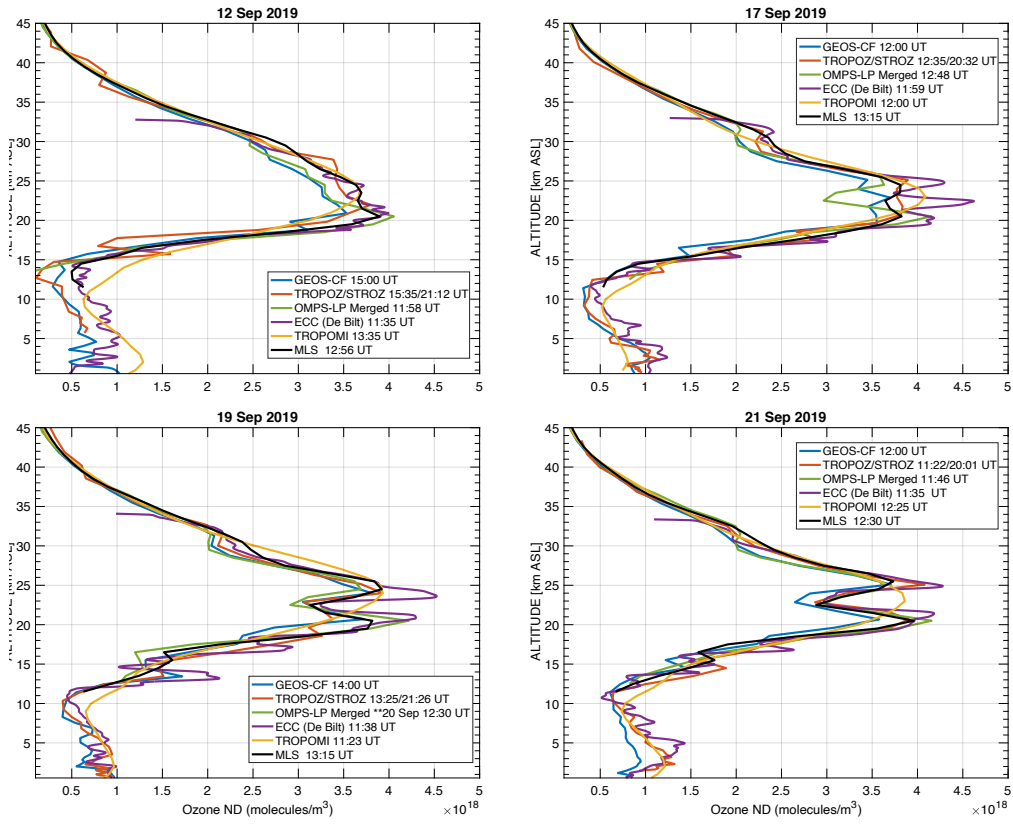


857

858 **Figure 4: Full tropospheric columns (top panel) and 0-2km tropospheric columns (bottom panel) calculated from**
859 **GEOS-CF, OMPS-MERRA2 (full column only), TROPOMI, Lidar and ECC. Data where reflectivity was greater than**
860 **0.6 was excluded to remove cloud interference.**

861
862
863
864
865
866
867
868
869
870
871
872

873
874
875
876
877
878
879
880

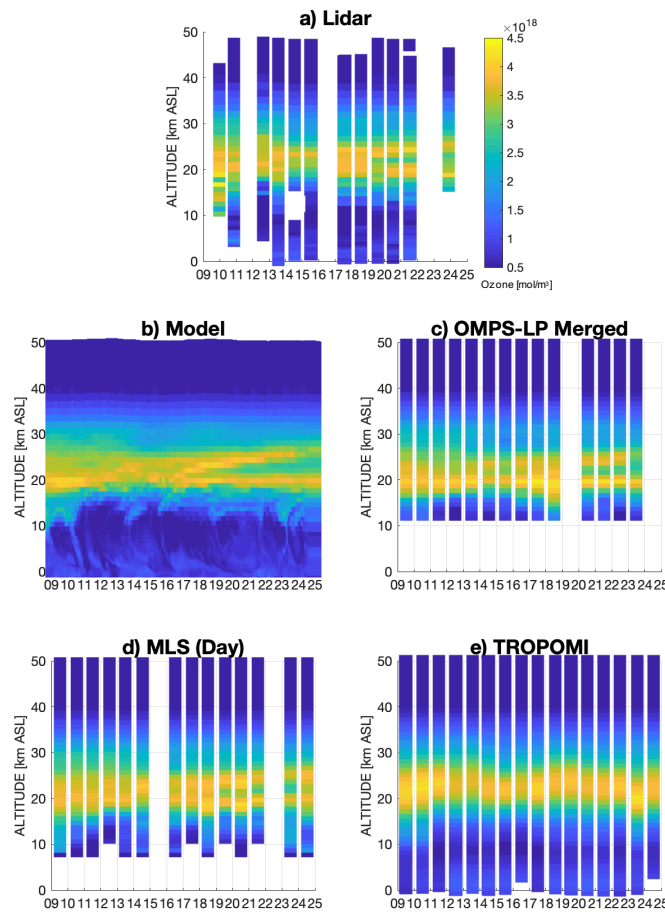


881
882

883 **Figure 5: GEOS-CF, Lidar, OMPS-LP, ECC, TROPOMI, and MLS ozone profile comparisons for 12 Sep, 17- Sep, 19**
884 **Sep, and 21 Sep 2019. These days were selected as days within the campaign that had an ECC launch from De Bilt.**

885
886
887
888
889
890

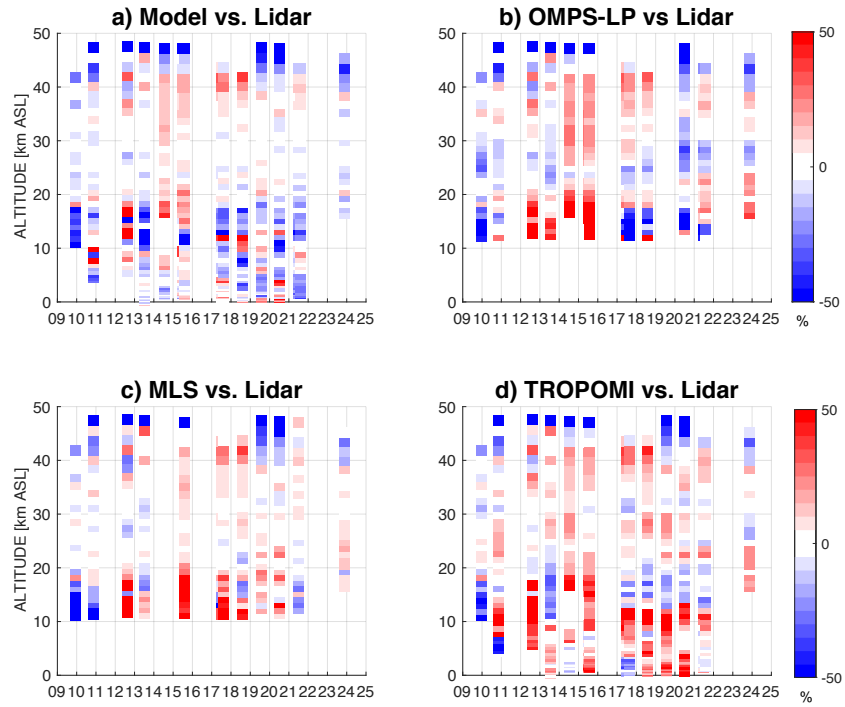
891
892
893
894
895
896
897



898
899
900
901
902

Figure 6: Ozone number densities across all platforms for the TROLIX-19 time period from the hybrid lidar dataset (Figure 6a), GEOS-CF (Figure 6b), OMPS-LP (Figure 6c), MLS (Figure 6d), TROPOMI (Figure 6e). The x-axis as day of September 2019.

903
904
905
906
907
908
909
910
911
912
913
914
915



916

917 **Figure 7: Differences in ozone number densities across all platforms for the TROLIX-19 time period for Model (Figure**
918 **7a), OMPS-LP (Figure 7b), MLS (Figure 7c), and TROPOMI (Figure 7d). The x-axis as day of September 2019.**

919

920

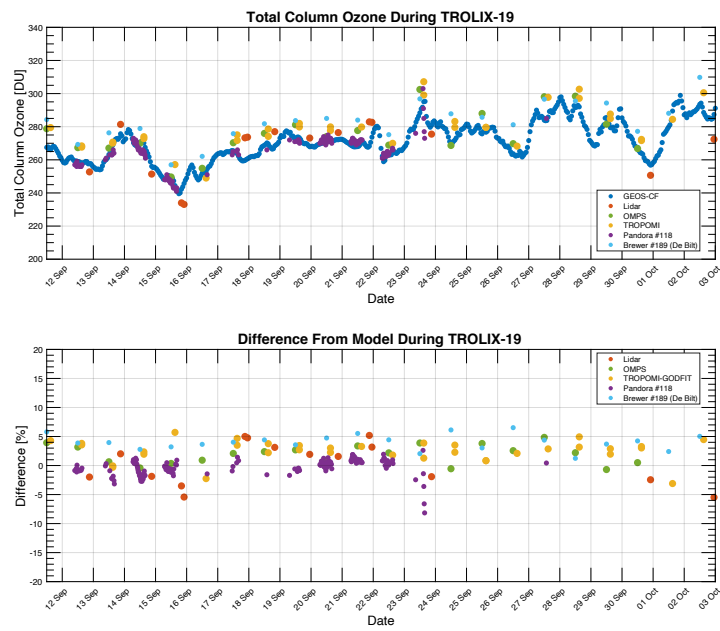
921

922

923

924

925
926
927
928
929
930
931
932



933
934
935
936

Figure 8: Total Ozone columns (top panel) and percent differences (bottom panel) as compared to the model observations for GEOS-CF, lidar, OMPS, TROPOMI, Pandora, and Brewer.