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**Abstract.** Refractory black carbon (*r*BC) aerosols play an important role in air quality and climate change, yet high time-resolved and detailed investigation on the physicochemical properties of *r*BC and its associated coating is still scarce. In this work, we used a laser-only Aerodyne soot particle aerosol mass spectrometer (SP-AMS) to exclusively measure the *r*BC-containing (*r*BCc) particles, and compared their properties with the total non-refractory submicron particles (NR-PM<sub>1</sub>) measured in parallel by a high-resolution AMS (HR-AMS) in Shanghai. The observation shows that *r*BC was overall thickly coated with an average mass ratio of coating to *r*BC core ( $R_{BC}$ ) of  $\sim 5.0$ . However, mass of *r*BC coating species only occupied 19.1% of those in NR-PM<sub>1</sub>; sulfate tended to condense preferentially on non-*r*BC particles therefore its portion on *r*BC was only 7.4%, while the majority of primary organic aerosols (POA) were associated with *r*BC (72.7%). Positive matrix factorization reveals that cooking emitted organics did not coat on *r*BC, and a portion of organics coated on *r*BC was from biomass burning which was unidentifiable in NR-PM<sub>1</sub> organics. Small *r*BCc particles were predominantly from traffic, while large-sized ones were often mixed with secondary components and typically had thick coating. During this campaign, sulfate and secondary organic aerosol (SOA) species were generated mainly through daytime photochemical oxidation (SOA formation likely involved with in-situ chemical conversion of traffic-related POA to SOA), while nocturnal heterogeneous formation was dominant for nitrate; we also estimated the average times of 5–19 hours for those secondary species to coat on *r*BC. Particles during a short period that was affected by ship emissions, were characterized with a high vanadium concentration (on average 5.8 ng m<sup>-3</sup>) and a vanadium/nickel mass ratio of 2.0. Furthermore, the size-resolved hygroscopicity parameter ( $\kappa_{rBCc}$ ) of *r*BCc particles was obtained based on its fully chemical characterization, and was parameterized as  $\kappa_{rBCc}(x) = 0.29 - 0.14 \times \exp(-0.006 \times x)$  ( $x$  is from 150 to 1000 nm). Under critical supersaturations ( $SS_c$ ) of 0.1% and 0.2%, the  $D_{50}$  values were  $166 \pm 16$  and  $110 \pm 5$  nm, respectively, and with  $16 \pm 3\%$  and  $59 \pm 4\%$  of *r*BCc in number could be activated into cloud condensation nuclei (CCN). Our findings are valuable to advance the understanding of BC chemistry as well as the effective control of atmospheric BC pollution.



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## 1 Introduction

Refractory black carbon (*r*BC) aerosols can directly absorb solar radiation, indirectly change the nature of cloud and alter the albedo of snow or glaciers (Jacobi et al., 2015), resulting in a positive radiative forcing that is second only to carbon dioxide on both regional and global scales (Ramanathan and Carmichael, 2008; Bond et al., 2013). The fresh *r*BC particles produced by incomplete combustion of biomass and fossil fuel tend to be fractal in morphology and can mix with many other components (Peng et al., 2016; Li et al., 2021). After entering into the atmosphere, fresh *r*BC can further externally or internally mix with organic/inorganic species which are primarily emitted or secondarily formed, and such aged *r*BC-containing (*r*BCc) particles (Chen et al., 2017; Lee et al., 2017) might have contrasting chemical properties and morphologies (or mixing states) (Liu et al., 2017a; Lee et al., 2019; Xie et al., 2019). In addition, when *r*BC mixes with hydrophilic materials, its hygroscopicity, cloud condensation nuclei (CCN) activity and size distribution, etc., can be significantly changed, which subsequently affect its atmospheric behavior, impact and lifecycle (Liu et al., 2013; Lambe et al., 2015). Therefore, it is necessary to elucidate the physicochemical characteristics and sources of *r*BC cores and associated coating materials, so as to better understand their influences on climate and air quality.

Chemical composition of ambient *r*BCc particles is largely dependent upon atmospheric conditions and emission sources. In general, the thickness of coating, mass contribution of secondary components (such as sulfate, nitrate and secondary organic aerosol (SOA) species) and oxidation degree of the coated organics of *r*BCc particles, increase with the aging time or oxidation capacity of ambient environment (Cappa et al., 2012; Liu et al., 2015; Wang et al., 2017; Collier et al., 2018; Wang et al., 2019), except in some specific cases that thickly coated *r*BCc might be dominated by primarily emitted particles (such as from biomass burning (Wang et al., 2017)). Recent field observations report that SOA species coated on *r*BC cores could account for 35% and 41% of the total SOA mass near traffic emission sources and in a polluted offshore environment, respectively (Massoli et al., 2012; Massoli et al., 2015). A study of *r*BCc particles in Singapore finds that over 90% of *r*BC derived from local combustion



sources (mainly traffic), while 30% of *r*BC was associated with fresh SOA generated under the influences of daytime shipping and industrial emissions (Rivellini et al., 2020). The SOA material concentrated on the surface of *r*BC was found to be chemically different from the SOA that was externally mixed with *r*BC (Lee et al., 2017) in Fontana, California, and another study in Shenzhen, China, reveals that more oxidized SOA preferred to mix with *r*BC due to that abundant transition metals detected on *r*BC cores might act as catalysts to convert less oxidized SOA to more oxidized SOA in aerosol aqueous phase (Cao et al., 2022); the *r*BC could catalyze SO<sub>2</sub> to form sulfate as well, as observed in Beijing (Zhang et al., 2020) and Guangzhou (Zhang et al., 2021), China. Besides SOA, cooking-related OA is found to be externally mixed with *r*BC (Lee et al., 2017; Wang et al., 2019), and a unique biomass burning related OA factor was identified and was only present in *r*BCc rather than non-*r*BC particles during summertime in Beijing (Wang et al., 2020a).

Size distribution of *r*BCc particles is also modulated greatly by their original sources and ageing processes. For example, a study in Shanghai shows a bimodal size distribution of *r*BCc, with a condensation mode dominated by traffic emissions (small core size, thin coating) and a droplet mode including highly aged biomass burning particles (large core size and thick coating) and highly aged traffic particles (small core size and very thick coating) (Gong et al., 2016). Another study in Beijing (Liu et al., 2019) further resolves four size modes of *r*BCc, relevant with traffic (small core, thin coating), coal or biomass burning (moderate coating, both small and large cores), coal combustion (large core, thick coating) and secondary process (thick coating, both small and large cores).

Moreover, water uptake and CCN activity of *r*BCc particles can increase with the encapsulation of water-soluble substances such as sulfate, nitrate, and SOA (Liu et al., 2013; Wu et al., 2019). Based on the measured chemical composition of *r*BCc, our previous work has established a method for calculating size-resolved hygroscopicity parameters of *r*BCc ( $K_{rBCc}$ ), and determined the CCN activation diameters of *r*BCc particles for given critical supersaturation ( $SS_c$ ) values (Wu et al., 2019).

Highly time-resolved chemical characterization of *r*BCc particles were seldom



110 reported in China and is still lacking in Shanghai. In this study, we utilized an Aerodyne  
111 soot particle aerosol mass spectrometer (SP-AMS) to determine the concentration,  
112 composition and size distribution of  $rBC_c$  particles exclusively (technical details in  
113 Section 2.1) in urban Shanghai for the first time. We also compared the SP-AMS  
114 measurement results with those from a co-located Aerodyne high-resolution time-of-  
115 flight aerosol mass spectrometer (HR-AMS), to comprehensively investigate the  
116 characteristics of  $rBC_c$  particles. We analyzed  $K_{rBC_c}$  and estimated the proportions of  
117 activated  $rBC_c$  numbers at given  $SSC$  as well.

118

## 119 **2 Experimental methods**

### 120 **2.1 Sampling site and instrumentation**

121 The field measurement was conducted from October 31 to December 2, 2018,  
122 during which the instruments were deployed on 8<sup>th</sup> floor of the building of Shanghai  
123 Academy of Environmental Sciences (SAES) (31°10'33.348" N, 121°26'10.978" E).  
124 East of the sampling site is a large commercial shopping center, and the site is  
125 surrounded by residential areas with two busy arterial roads directly to the east (~450  
126 m) and south (~150 m), respectively (Figure S1 in the supporting information). In  
127 addition, the adjacent areas are densely populated with roadside residents, office  
128 workers, and market traders, as well as crowds in and out of the Caobao Road Metro  
129 station (~100 m). The measurement period was dominated by northeasterly winds,  
130 while many international freight companies located on northeastern side of the site, and  
131 many freighters were reposing on the Huangpu River. Overall, the sampling site was  
132 probably influenced by vehicular emissions, residential activities and the northeast  
133 Cargo ship emission plumes, etc.

134 An Aerodyne SP-AMS and an HR-AMS were operated in parallel during the  
135 campaign. The two AMSs shared a same sampling line with a  $PM_{2.5}$  cyclone (Model  
136 URG-2000-30EN) in front to remove coarse particles. Ambient air pulled through the  
137 sampling line was dried using a diffusion dryer filled with silicon gel and was  
138 subsequently drawn into both instruments. Due to the transmission efficiency of the  
139 inlet lens, both AMSs measured mainly particles of 30-1200 nm (denoted as  $PM_1$ ).



140 The working principle of SP-AMS has been described in detail previously (Onasch  
141 et al., 2012). However, in this work, we used only the intracavity infrared laser  
142 vaporizer to selectively measure *r*BCc particles (*r*BC cores and associated coating  
143 materials), as *r*BC can absorb 1064 nm laser light. The thermal tungsten vaporizer had  
144 to be physically detached otherwise non-*r*BC particles can still be detected as the  
145 filament can heat the vaporizer to ~200 °C even if it was turned off. Before sampling,  
146 the SP-AMS was tuned and calibrated following the steps described previously (Lee et  
147 al., 2015; Willis et al., 2016; Wang et al., 2017). During sampling, due to relatively low  
148 *r*BCc mass loadings, the SP-AMS was operated with two mass sensitive V modes (2.5  
149 minutes per cycle), one with a particle time-of-flight (PToF) mode (30 s) and another  
150 one (120 s) with a mass spectral mode with mass-to-charge ( $m/z$ ) ratio up to 500.  
151 Filtered air was also measured in the middle of campaign (for 60 min) to determine the  
152 limits of detection (LOD, three times the standard deviation) of various aerosol species  
153 and to adjust the air-influenced mass spectral signals (Zhang et al., 2005).

154 Before removal of the tungsten heater, the calibrations of ionization efficiency (IE)  
155 for nitrate and relative ionization efficiency (RIE) of sulfate were performed by using  
156 pure ammonium nitrate and ammonium sulfate particles (Jayne et al., 2000), and the  
157 values were assumed to be unchanged throughout the whole campaign (Willis et al.,  
158 2016). RIE of *r*BC to nitrate was calibrated by using size-selected (300 nm) BC  
159 particles (REGAL 400R pigment black, Cabot Corp.) (Onasch et al., 2012), and the  
160 average ratio of  $C_1^+$  to  $C_3^+$  was calculated to be 0.584 to correct the interference on  $C_1^+$   
161 from other organics. RIEs of ammonium, nitrate, sulfate and *r*BC were determined to  
162 be 4.53, 1.10, 1.01, and 0.17, respectively, and RIE of organics used the default value  
163 of 1.4 (Canagaratna et al., 2007). The size was calibrated by Polystyrene latex (PSL)  
164 spheres (100–700 nm) (Duke Scientific Corp., Palo Alto, CA) before the measurement.  
165 This study applied a collection efficiency (CE) of 0.5 for SP-AMS.

166 The co-located HR-AMS (DeCarlo et al., 2006) was used to measure all  $PM_{10}$   
167 including both *r*BCc and non-*r*BC particles, but it detected only non-refractory species  
168 (NR- $PM_{10}$  species) as its 600 °C thermal heater is unable to vaporize *r*BC and other  
169 refractory components. In addition, mass concentrations of gaseous pollutants, carbon



monoxide (CO), ozone (O<sub>3</sub>), nitrogen dioxide (NO<sub>2</sub>), and sulfur dioxide (SO<sub>2</sub>) were measured by the Thermo Scientific analyzers provided by SAES. Meteorological parameters including air temperature (T), relative humidity (RH), wind speed (WS), wind direction (WD) and precipitation, were obtained from Xujiahui Environmental Monitoring Station of Shanghai (31°11'49.1424"N, 121°26'34.44" E)(~2400 m away from the site). The concentrations of particle-phase vanadium (V) and nickel (Ni) that were used to investigate ship influence were measured independently by an Atmospheric heavy metal analyzer (XHAM-2000A, SAIL HERO., China).

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## 2.2 Data analysis

The AMS data (both SP-AMS and HR-AMS) were analyzed using standard ToF-AMS data analysis tool (Squirrel version 1.59D and Pika version 1.19D), based on Igor Pro 6.37 (Wavemetrics, Lake Oswego, OR, USA). The mass concentrations and high resolution mass spectra (HRMS) of *r*BC and coating species (*r*BC<sub>CT</sub>) were calculated from high-resolution (HR) fitting of V-mode data. Size distributions of *r*BC<sub>C</sub> components were determined by the PToF data with unit mass resolution and were scaled to their mass concentrations obtained above. In particular, size distribution of *r*BC was scaled to that of *m/z* 24 (C<sub>2</sub><sup>+</sup>) (the scaling factor is *r*BC mass concentration to that of calculated based on its size distribution), because *m/z* 24 as a *r*BC fragment, has least interference from other organic or inorganic species; such treatment was adopted in earlier studies too (Collier et al., 2018; Wang et al., 2019; Wang et al., 2016).

The HR ion fitting of AMS data is able to distinguish various ions and isotopic ions and calculate elemental ratios of organics such as oxygen-to-carbon (O/C), hydrogen-to-carbon (H/C), nitrogen-to-carbon (N/C), and organic mass to organic carbon (OM/OC) ratios, via the original Aiken-ambient (A-A) method (Aiken et al., 2008) and the improved method (I-A) (Canagaratna et al., 2015b). Outcomes of both methods correlated well. Average O/C, H/C, and OM/OC ratios from the I-A method used in this work were 24.9%, 7.3%, and 5.6%, respectively, higher than those from the A-A method.

Furthermore, we performed Positive matrix factorization (PMF) (Paatero and



200 Tapper, 1994) analysis on the HRMS of organics measured by the SP-AMS via the PMF  
 201 Evaluation Tool (Ulbrich et al., 2009). The PMF solutions were thoroughly evaluated  
 202 following the protocols documented in Zhang et al. (2011). Finally, a 6-factor solution  
 203 was chosen as the optimal one. The final result included four primary OA (POA) factors,  
 204 namely hydrocarbon-like species enriched OA (HOA-rich), *r*BC-enriched OA (*r*BC-  
 205 rich), biomass burning OA (BBOA), water-soluble hydrocarbon-like species enriched  
 206 OA (WS-HOA), and two secondary OA (SOA) factors including a less oxidized  
 207 oxygenated OA (LO-OOA<sub>*r*BC</sub>), and a more oxidized oxygenated OA (MO-OOA<sub>*r*BC</sub>)  
 208 (Key diagnostic plots in Fig. S2). The HR-AMS data were processed in a similar way,  
 209 and four factors of NR-PM<sub>1</sub> organics were resolved, including hydrocarbon-related OA  
 210 (HOANR-PM<sub>1</sub>), cooking OA (COANR-PM<sub>1</sub>), less oxidized oxygenated OA (LO-OOANR-  
 211 PM<sub>1</sub>), and more oxidized oxygenated OA (MO-OOANR-PM<sub>1</sub>) (Mass spectra and time  
 212 series shown in Fig. S3).

213

### 214 2.3 Calculation of size-resolved hygroscopicity of *r*BCc

215 The hygroscopicity parameter  $\kappa$  is a single parameter representing the  
 216 hygroscopicity of particles, calculated based on essentially the chemical composition  
 217 (Petters and Kreidenweis, 2007). SP-AMS measured size-resolved chemical  
 218 compositions of *r*BCc can thus lead to size-resolved hygroscopicity of *r*BCc ( $\kappa_{rBCc}$ )  
 219 (Wu et al., 2019; Hu et al., 2021). This study applied a similar method proposed by Wu  
 220 et al. (2019) and the critical parameters involved in calculation are detailed in Table S1.  
 221 The procedures are briefly described here: (1) obtain the matrix of size-resolved mass  
 222 concentrations of each *r*BCc component from SP-AMS analysis; (2) convert the size-  
 223 resolved concentration matrix of inorganic ions ( $\text{SO}_4^{2-}$ ,  $\text{NO}_3^-$  and  $\text{NH}_4^+$ ) to  
 224 corresponding matrix of inorganic salts ( $\text{NH}_4\text{NO}_3$ ,  $\text{NH}_4\text{HSO}_4$ ,  $(\text{NH}_4)_2\text{SO}_4$ ) using a  
 225 simplified solution of ion pairs (Gysel et al., 2007); (3) convert the size-resolved mass  
 226 matrix of all components (inorganic salts, OA and *r*BC) to those of size-resolved  
 227 volume fractions (Gysel et al., 2007; Chang et al., 2010; Wu et al., 2016); (4) combine  
 228 with previously reported hygroscopic parameters (Gysel et al., 2007; Chang et al.,  
 229 2010; Wu et al., 2016) to obtain the volumetric contributions of each component to the





hygroscopicity of  $rBC_c$ ; (5) use the Zdanovskii-Stokes-Robinson (ZSR) rule to derive the size-resolved hygroscopicity of  $rBC_c$  ( $\kappa_{rBC_c}$ ) (Topping et al., 2005a, b).

### 3 Results and discussion

#### 3.1 Overview of chemical characteristics of $rBC_c$ and NR-PM<sub>1</sub> particles

Figure 1 presents the time series of meteorological parameters, concentrations of gaseous pollutants (CO, NO<sub>2</sub>, O<sub>3</sub>, and SO<sub>2</sub>), ship emission tracers (vanadium and nickel),  $rBC$  and  $rBC_{CT}$  species and their mass percentages to the total  $rBC_c$  mass, PMF-resolved OA factors and their corresponding fractional contributions. The sampling period was featured by relatively moderate temperatures and very stagnant conditions with average ( $\pm 1\sigma$ ) temperature of  $15.3 \pm (2.9)^\circ\text{C}$  and wind speed (WS) of  $0.16 \pm (0.29) \text{ m s}^{-1}$ . Calm wind ( $<0.5 \text{ m s}^{-1}$ ) dominated most of the sampling days and 42% of sampling time was with near-zero wind, therefore overall influence of WS and WD on surface mass loadings of  $rBC_c$  was insignificant. Yet one should keep in mind that WD can affect  $rBC_c$  sources, and WS in higher altitude might be stronger therefore long-range transport of air pollutants was still possible. The average concentrations of CO, NO<sub>2</sub>, O<sub>3</sub>, SO<sub>2</sub>, V, and Ni were determined to be 0.60 ppm, 29.20 ppt, 27.10 ppt, 1.27 ppt, 4.05 ng m<sup>-3</sup>, and 3.06 ng m<sup>-3</sup>, respectively.

The mass loadings of  $rBC$  and  $rBC_{CT}$  ranged from 0.04 to 11.00  $\mu\text{g m}^{-3}$  and 0.37 to 30.47  $\mu\text{g m}^{-3}$  with campaign-mean values ( $\pm 1\sigma$ ) of  $0.92 \pm (0.81) \mu\text{g m}^{-3}$  and  $4.55 \pm (4.40) \mu\text{g m}^{-3}$ . The coating materials accounted for 83.0% of the total  $rBC_c$  mass, of which organics was the most abundant species ( $2.54 \pm 2.52 \mu\text{g m}^{-3}$ , 46.0%), followed by nitrate ( $1.20 \pm 1.30 \mu\text{g m}^{-3}$ , 20.0%), ammonium ( $0.44 \pm 0.40 \mu\text{g m}^{-3}$ , 9.0%), sulfate ( $0.30 \pm 0.19 \mu\text{g m}^{-3}$ , 6.0%), and chloride ( $0.07 \pm 0.05 \mu\text{g m}^{-3}$ , 1.0%). The mass ratio of  $rBC_{CT}$  to  $rBC$  ( $R_{BC}$ ) ranged from 2.2 to 9.0, with an average of  $\sim 5.0 (\pm 1.7)$ . The average  $R_{BC}$  was higher than that in California ( $R_{BC} = 2.3$ ) (Collier et al., 2018) and in Shenzhen ( $R_{BC} = 2.5$ ) (Cao et al., 2022), lower than that in Tibetan Plateau ( $R_{BC} = 7.7$ ) (Wang et al., 2017) and similar to that in Beijing ( $R_{BC} = 5.0$ ) (Wang et al., 2019), suggesting  $rBC$  was relatively thickly coated throughout the campaign. Correlation between  $rBC$  and  $rBC_{CT}$  was moderate (Pearson's  $r^2 = 0.58$ ). Correlation coefficients ( $r^2$ ) of chloride, nitrate,



sulfate and organics with  $r_{BC}$  were 0.52, 0.75, 0.51 and 0.53, respectively, suggesting variability of sources among different coating components.

Figure 2 compares the campaign-averaged diurnal patterns of  $r_{BCc}$  and  $NR-PM_{10}$  species, chemical compositions of  $r_{BCc}$  and  $NR-PM_{10}$ , and mass ratios of the species coated on  $r_{BC}$  to those of  $NR-PM_{10}$ . We found that the diurnal variations of nitrate, sulfate, ammonium, chloride were very similar ( $r^2 > 0.86$ ) between the two particle groups, while apparent difference was found for  $r_{BCc}$  organics with a much obvious morning rush hour peak. The results indicate that the formation processes of inorganic salts coated on  $r_{BC}$  were similar to those uncoated on  $r_{BC}$ , but there were large difference regarding sources/processes existed for organics. For  $r_{BC}$  itself, the diurnal cycle presented clearly a morning peak and an evening peak, likely relevant with rush hour traffic emissions (CO showed similar pattern). On the contrary,  $R_{BC}$  had a minimum in the morning and dropped to a low level in later afternoon, probably due to influence from traffic-emitted fresh and barely coated  $r_{BC}$  particles (details in Section 3.2.1).

Distributions of species between  $r_{BCc}$  and non- $r_{BC}$  particles were also different, leading to different chemical compositions (Figs. 2g and 2h). Sulfate tended to preferentially condense on non- $r_{BC}$  particles, as its mass contribution to total  $r_{BCc}$  mass was only 6.5%, while its contribution to total  $NR-PM_{10}$  was 17.6%. Apportionment of nitrate between  $r_{BCc}$  and non- $r_{BC}$  particles was about even as it both occupied ~26% of the total  $r_{BCc}$  and  $NR-PM_{10}$  masses. Organics occupied 55.9% of  $r_{BCc}$  mass, larger than it in  $NR-PM_{10}$  (43.7%), due to that primary OA species preferentially associated with  $r_{BC}$ . Such result is similar to that observed in winter in Beijing but contrary to the result that SOA was more abundant in  $r_{BCc}$  in Shenzhen (Cao et al., 2022).

On average,  $r_{BCc}$  accounted for 19.1% of the total  $NR-PM_{10}$  mass loading ( $21.61 \pm 15.80 \mu\text{g m}^{-3}$ ) (Fig. 2i), comparable to that in Fontana, California (Lee et al., 2017). The finding reveals that significant fractions of aerosol species were externally mixed with  $r_{BC}$ . Individually, sulfate captured by  $r_{BC}$  only represented 7.4% of  $NR-PM_{10}$  sulfate, similar to the earlier results (Lee et al., 2017; Wang et al., 2020a; Cao et al.,



2022); mass fractions of *rBCc* nitrate (20.1%) and chloride (20.4%) in NR-PM<sub>1</sub> were similar to the portion of total *rBCc* (19.1%), while the fraction of organics was higher (26.1%). The relatively high ratio of organics was attributed to the fact that majority of POA species were coated on *rBC* (average ratio of 72.7%), while *rBC*-related SOA was 21.8 % of the total. Note the *rBCc* POA here included all four POA factors, and CO<sub>ANR-PM1</sub> did not coat on *rBC* thus was not included in the calculation.

### 3.2 Distinctive sources of OA in *rBCc* and in NR-PM<sub>1</sub>

As shown previously, source apportionment results of OA in *rBCc* and NR-PM<sub>1</sub> were different. This section discusses in details the characteristics of OA sources in *rBCc* and in bulk NR-PM<sub>1</sub>.

#### 3.2.1 Source apportionment of OA in *rBCc*

Figure 3 shows the HRMS and temporal variations of the six OA factors resolved from PMF analysis of *rBCc* organics. Note the PMF analysis included *rBC* signals (i.e., C<sub>x</sub><sup>+</sup> ions) to aid identification of different factors, yet calculations of elemental ratios of these OA factors did not include C<sub>x</sub><sup>+</sup> ions in order to explicitly explore the properties of organic coating. The HRMS of HOA-rich and *rBC*-rich were similar to the OA previously reported in urban environments near traffic emissions and/or in gasoline/diesel vehicle exhaust (Massoli et al., 2012; Lee et al., 2015; Enroth et al., 2016; Saarikoski et al., 2016; Willis et al., 2016; Lee et al., 2017), therefore both factors were traffic-related. The HOA-rich mass spectrum was featured by intense hydrocarbon ion series (i.e., C<sub>n</sub>H<sub>2n+1</sub><sup>+</sup> and C<sub>n</sub>H<sub>2n-1</sub><sup>+</sup> ions in Fig. 3c), and a lowest O/C ratio of 0.07. Mass fraction of *rBC* signals (i.e., C<sub>n</sub><sup>+</sup> ions, such as *m/z* 12, 24, 36, 48, and 60, etc.) in HOA-rich was 8.1%. Mass spectrum of *rBC*-rich factor had remarkable contribution from *rBC* (24.2%). Beside C<sub>n</sub><sup>+</sup> ions, the *rBC*-rich factor contained more oxygenated organic fragments and presented a higher O/C ratio of 0.21 than that of HOA-rich, similar to previous studies (Willis et al., 2016; Lee et al., 2017). This result is reasonable as previous studies (Corbin et al., 2014; Malmberg et al., 2017) found that refractory organics could generate oxygenated ion fragments (such as CO<sup>+</sup> and CO<sub>2</sub><sup>+</sup> derived from oxygenated species on soot surface or inside soot nanostructure). HOA-rich factor



correlated very well with the common AMS tracer of vehicular OA,  $\text{C}_4\text{H}_9^+$  ( $r=0.95$ , Fig. 3i), while  $r\text{BC}$ -rich factor did match the variation of  $r\text{BC}$  well ( $r=0.90$ , Fig. 3g). Since diesel combustion often releases more  $r\text{BC}$  than that of gasoline, it is likely that the  $r\text{BC}$ -rich factor is representative of diesel vehicle exhaust while HOA-rich factor represents gasoline combustion emissions. This result demonstrates that laser-only SP-AMS is capable of distinguishing diesel and gasoline burning particles which typically cannot be separated by other AMS measurements. Further verification should be subject of future work.

In this work, a multiple linear regression for the three-dimension size-resolved mass spectra according to the method provided in Ulbrich et al. (2012) was used to resolve the average size distributions of six OA factors. The diagnostic plots are shown in Fig. 4. Overall, the lumped size distribution of six OA factor could reproduce well that of total OA (except for a few size bins, most deviations are within 10%). Correlation between measured and reconstructed OA of all size bins was very tight with  $r$  of 0.99 and a slope of 0.97, indicating the robustness of the regressed size distributions of all OA factors. The results together with size distributions of  $r\text{BCc}$  components, and corresponding mass fractional contributions of different components in all size bins are illustrated in Fig. 5. The average HOA-rich size distribution peaked around 150 nm ( $D_{\text{va}}$ , vacuum aerodynamic diameter), generally matching with previously reported size distribution of HOA (Sun et al., 2012; Ulbrich et al., 2012; Zhou et al., 2016). Interestingly, size distribution of  $r\text{BC}$ -rich factor presented two modes, with one peaking  $\sim 260$  nm, and a more pronounced one peaking  $\sim 580$  nm (Fig. 5a).

The BBOA was identified owing to its obviously higher signals of  $\text{C}_2\text{H}_4\text{O}_2^+$  (2.03%) and  $\text{C}_3\text{H}_5\text{O}_2^+$  (1.62%) than those in other factors, as these two ions are well-known AMS fragments of the biomass burning tracer, levoglucosan (Mohr et al., 2009; Cubison et al., 2011). In addition, the time series of BBOA correlated particularly tightly with both marker ions ( $r$  of 0.86 and 0.80, respectively); it in fact also correlated well with  $\text{K}^+$  ( $r=0.79$ ), another tracer of biomass burning emission. The O/C and H/C ratios of BBOA were 0.12 and 1.78, and  $\text{C}_n^+$  ions accounted for 9.1% of BBOA, all suggesting that the factor was fresh and might be co-emitted with  $r\text{BC}$ . A relatively high N/C ratio (0.033)



was found for BBOA, which could be attributed to the large amounts of nitrogen-containing organic species enriched in biomass burning OA (Laskin et al., 2009). In addition, the size distribution of BBOA (Fig. 5a) (peak size ~500 nm) was similar to that of biomass burning BC-containing particles obtained using single particle mass spectrometry in Shanghai (Gong et al., 2016).

The PMF analysis deconvoluted a unique OA factor coated on *r*BC, namely WS-HOA. Firstly, the WS-HOA mass spectrum had a series of hydrocarbon ion fragments and its time series correlated well with them (e.g., *r* of 0.90 and 0.92 for WS-HOA vs.  $C_4H_7^+$  and  $C_4H_9^+$ , respectively). Secondly, this factor correlated the best (*r*=0.57) with aerosol liquid water content (ALWC) (Fig. 3j) compared with the other five OA factors (all *r*<0.2). Thirdly, a previous study (Ye et al., 2017) that investigated specially the water-soluble fraction of OA via PMF analysis, separated also a HOA factor that contained significant nitrogen-containing organic fragments, with a highest N/C ratio among all other factors, and correlated well those nitrogenated fragments. The WS-HOA defined here showed similar characteristics with the highest N/C (0.037) among all factors and tight correlations with nitrogen-containing fragments (*r*>0.80). At last, although WS-HOA had a relatively high O/C (0.31) with remarkable contributions from  $C_2H_3O^+$  and  $CO_2^+$ , yet its correlations with these two ions were in fact not strong (*r* of 0.46 and 0.44, respectively); and WS-HOA had the least fraction of *r*BC fragments (0.9%) (note *r*BC is hydrophobic), even less than the two SOAs (Fig. 3d). Both results suggest that this factor is a collection of water-soluble primary OA species. The peak of WS-HOA size distribution was ~150 nm, close to that of HOA-rich (Figs. 5a). Aqueous-phase processed SOA (aqSOA) were typically with very high O/C ratio (Xu et al., 2019), and size distribution of aqSOA often presented a droplet mode, peaking in relatively large sizes (such as 500–600 nm) (Gilardoni et al., 2016; Wang et al., 2021; Ge et al., 2012). Therefore, the moderate O/C (0.31) and small mode size of WS-HOA again manifest it was not from aqueous-phase reactions but more likely the water-soluble fraction of POA.

The PMF analyses separated two SOA factors, LO-OOA<sub>*r*BC</sub> and MO-OOA<sub>*r*BC</sub>. Mass spectral features of the two SOAs were consistent with previous studies: The LO-



OOA<sub>rBC</sub> was rich in C<sub>x</sub>H<sub>y</sub>O<sub>1</sub><sup>+</sup> ions (38.7%) (such as C<sub>2</sub>H<sub>3</sub>O<sup>+</sup> at *m/z* 43) but with less contribution from C<sub>x</sub>H<sub>y</sub>O<sub>2</sub><sup>+</sup> ions (7.6%) and an overall moderate O/C (0.25), while MO-OOA<sub>rBC</sub> had much more contribution from C<sub>x</sub>H<sub>y</sub>O<sub>2</sub><sup>+</sup> ion family (22.7%) (such as CO<sub>2</sub><sup>+</sup> at *m/z* 44) and C<sub>x</sub>H<sub>y</sub>O<sub>1</sub><sup>+</sup> ions (44.7%), with a high O/C ratio (0.56). In addition, LO-OOA<sub>rBC</sub> correlated better with nitrate (*r*=0.83) than it with sulfate (*r*=0.69), while the correlation between MO-OOA<sub>rBC</sub> and sulfate (*r*=0.84) is better than it with nitrate (*r*=0.76). Size distributions of the two SOAs were also in accord with their secondary behaviors, both accumulating at larger sizes (~450 nm for LO-OOA<sub>rBC</sub>, and a bit larger mode size of ~500 nm for MO-OOA<sub>rBC</sub>), in agreement with previous observations (Sun et al., 2012; Ulbrich et al., 2012; Zhou et al., 2016).

In total, traffic-related POA (sum of HOA-rich, *r*BC-rich and WS-HOA) was the most abundant source of *r*BCc organics (39.1%); BBOA occupied ~18.4%; the contributions of two SOAs were on par with each other (20.2% for LO-OOA<sub>rBC</sub>, and 22.3% for MO-OOA<sub>rBC</sub>) (Fig. 2g). Among traffic POA, gasoline derived HOA-rich factor outweighed the diesel derived *r*BC-rich factor (11.4% vs. 6.3% of the total *r*BCc). Contributions of different factors varied greatly for different sizes of *r*BCc particles (Fig. 5b). Small-sized particles were overwhelmingly dominated by traffic POA; SOA contributions increased with increase of size, and dominated over POA for 300-800 nm ones; contribution of BBOA was also relative larger for 300-800 nm than for other-sized ones; the very large particles (800-1000 nm) were found to be affected mainly by traffic POA in this work. Correspondingly, for the total *r*BCc particles, *r*BC cores peaked at ~170 nm, while other secondary inorganic components, behaving like SOA factors, all peaked at a big size (~550 nm) (Fig. 5c) and their mass percentages were also large for large-sized particles (Fig. 5d).

Figure 6a further demonstrates the changes of mass fractions of each *r*BCc component as a function of *R*<sub>BC</sub>. *R*<sub>BC</sub> is a proxy of coating thickness. It was found that the thinly coated *r*BCc (*R*<sub>BC</sub><3) were dominated (up to ~80%) by traffic-related POA. With the increase of *R*<sub>BC</sub>, contribution of secondary components increased gradually, especially the two SOAs and nitrate (sulfate showed little changes across the whole *R*<sub>BC</sub> range); SOA and nitrate contributions reached 40.2% and 31.3% at *R*<sub>BC</sub> = 8, respectively.



Accordingly, the oxidation degree ( $Osc = 2 \cdot O/C-H/C$ ) (Kroll et al., 2011) of coated organics increased with  $R_{BC}$ . In Fig. 6b, we presented the mass contributions of OA factors to the  $r_{BC}$  mass at different  $R_{BC}$  values. The  $r_{BC}$  was as expected, predominantly associated with POA (from 94% at  $R_{BC} < 2$  to 66% at  $R_{BC} > 8$ ), similar to those from Fontana (Lee et al., 2017). Contribution of  $r_{BC}$ -rich factor decreased obviously, and those of SOA factors increased with  $R_{BC}$ . The contributions of HOA-rich, WS-HOA and BBOA factors changed little.

417

### 3.2.2 Comparisons with NR-PM<sub>1</sub> organics

As shown in Fig. S3, PMF analysis separated four OA factor for NR-PM<sub>1</sub> organics. Two SOA factors (LO-OOA and MO-OOA) were resolved for both  $r_{BCc}$  and NR-PM<sub>1</sub>, and their contributions to them were also close (Figs. 2g and 2h). Correlations of time series between the two LO-OOA factors and between the two MO-OOA factors were also tight ( $r$  of 0.94 and 0.90, respectively), indicating similar formation processes for each SOA. But, of course, the SOAs from  $r_{BCc}$  and from NR-PM<sub>1</sub> were not entirely the same, as later ones had higher O/C ratios (0.52 of LO-OOA<sub>NR-PM1</sub> and 0.62 of MO-OOA<sub>NR-PM1</sub>). On average, the portion of LO-OOA coated on  $r_{BC}$  took up 21.6% mass of LO-OOA in total NR-PM<sub>1</sub>, and the portion was 26.0% for MO-OOA. This result suggests that there were some but not big differences regarding the partitioning of LO-OOA and MO-OOA onto  $r_{BCc}$  and non- $r_{BC}$  particles.

Compared with SOAs, source apportionment results of POA were quite distinct. Firstly, there was only one HOA factor resolved for NR-PM<sub>1</sub>, while three HOA factors were separated for  $r_{BCc}$ . Note the  $r_{BC}$ -rich and WS-HOA factors occupied merely 3.1% and 2.1% of NR-PM<sub>1</sub> OA mass, respectively, probably the cause that they were not identified in NR-PM<sub>1</sub> OA. Nevertheless, mass loadings of the sum of HOA-rich,  $r_{BC}$ -rich and WS-HOA factors (termed as HOA <sub>$r_{BC}$</sub> ) agreed quite well with that of HOA<sub>NR-PM1</sub> ( $r=0.95$ ) (Fig. 7a), verifying both source apportionment results. And, HOA <sub>$r_{BC}$</sub>  took up 63.7% of HOA<sub>NR-PM1</sub>, while previous studies reported that 81% (Massoli et al., 2012) and 87 % (Massoli et al., 2015) of HOA were associated with  $r_{BC}$ . These results imply that HOA species in NR-PM<sub>1</sub> were largely internally mixed



with  $rBC$  affected by vehicular emissions. Secondly,  $CO_{NR-PM_1}$  was only identified in  $NR-PM_1$  OA. AMS-resolved  $CO_{NR-PM_1}$  is mainly from cooking oil and food itself, therefore it negligibly internally mixes with  $rBC$ . This result is consistent with previous observations (Lee et al., 2015; Willis et al., 2016; Lee et al., 2017; Collier et al., 2018). At last, BBOA was identified in  $rBCc$  OA but not in  $NR-PM_1$  OA. One plausible reason was that the BBOA mass contribution was minor (equivalent to <5% of  $NR-PM_1$  OA mass); another speculation is that laser only SP-AMS can detect refractory species that HR-AMS cannot, some of them might be originated from biomass burning (Wang et al., 2020a).

Diurnal cycles of the POA and SOA factors from both PMF analyses are compared in Figs. 7b and 7c. The diurnal pattern of stacked  $HOA_{rBC}$  indeed agreed with that of  $HOA_{NR-PM_1}$ , both with two peaks in the morning and evening rush hours, and, the patterns of  $rBC$ -rich, HOA-rich, and WS-HOA factors showed similar behaviors individually (Fig. 7b). The diurnal variation of  $CO_{NR-PM_1}$  had pronounced peaks during lunch and dinner times, and its percentage in  $NR-PM_1$  OA reached 54% at night. Diurnal patterns of two LO-OOA factors were somewhat different ( $r=0.35$ ). LO- $OOA_{rBC}$  has a major peak in the afternoon, while though LO- $OOA_{NR-PM_1}$  concentration rose in the afternoon too but peaked in early evening (~8 pm). The daily variations of two MO-OOA factors were similar ( $r=0.83$ ), both peaking at 16:00. The afternoon increases of both SOAs indicate an important role of photochemical reactions, yet differences in formation mechanisms, volatilities and partitioning behaviors of SOA products could lead to diversities of their diurnal patterns and HRMS.

### 3.3 Evolution and formation $rBCc$ components

#### 3.3.1 Behaviors of $rBC$ cores

Size distribution of  $rBC$  cores shown in Fig. 5c was relatively wide. Beside the main peak at ~170 nm, it extended significantly into large sizes and had a small peak at ~550 nm, which was close to the peak of secondary components. With results shown in Fig. 5a, we found that small  $rBC$  cores were often thinly coated, while thickly coated  $rBCc$  particle were often highly aged and a portion of them also had large-sized  $rBC$





470 cores. This result is likely owing to that oxidation of *r*BC-bounded organics and/or  
471 condensation of secondary species onto *r*BC surface can induce restructuring of soot  
472 aggregates to form compact and large cores (Chen et al., 2018; Chen et al., 2016). Such  
473 phenomenon is in line with earlier studies (Liu et al., 2019; Gong et al., 2016). We  
474 further show the image plot of size distributions of *r*BC at different  $R_{BC}$  in Fig. S5a. It  
475 can be found that the *r*BC mass in a large part concentrated in particles with  $R_{BC}$  of 5-  
476 8, and there was indeed a significant portion of *r*BC appearing in large size range (400-  
477 800 nm) with very thick coating ( $R_{BC}$  of 8-9).

478

### 479 3.3.2 Formation of inorganic salts

480 Sulfate and nitrate both peaked at a big size ~550 nm (Fig. 5c) and were mainly  
481 associated with thickly coated *r*BCc ( $R_{BC}>6$ , Figs. S5d and S5e). To investigate the  
482 impacts of photochemistry and aqueous/heterogeneous chemistry on the formation of  
483 *r*BC<sub>CT</sub> species, we plotted the image plots of size distributions of nitrate, sulfate and  
484 organics versus  $O_x$  ( $O_3 + NO_2$ ) and relative humidity (RH) in Fig. 8. Here  $O_x$  is used  
485 as a proxy of photochemical activity (Xu et al., 2017), and RH is an indicator of aqueous  
486 reactions (Wu et al., 2018). Nitrate significantly concentrated in 65-100 ppb  $O_x$  range  
487 but there was a weak accumulation in low  $O_x$  as well (Fig. 8a), while in Fig. 8d, nitrate  
488 had a prominent hotspot in  $RH>85\%$ . Generally, both strong photochemical activity  
489 and high RH could promote nitrate formation. For sulfate, although the distribution was  
490 scattered due to low level of mass loadings, high  $O_x$  level seemed to favor sulfate  
491 formation (Fig. 8b); sulfate was scattered in the whole RH range and there were some  
492 enhancements at high RH ( $>80\%$ ) but was much less clear-cut (Fig. 8e). Therefore  
493 aqueous-phase production of sulfate was not important in this campaign.

494 We further calculated the sulfur oxidation ratio (SOR) and nitrogen oxidation ratio  
495 (NOR) (Xu et al., 2014), and plotted their variations against  $O_x$  and RH in Figs. 9a and  
496 9e, respectively. The NOR rose substantially at  $O_x>60$  ppb but showed no increase at  
497  $O_x<60$  ppb, while it increased continuously with RH. Mass ratio of nitrate to *r*BC  
498 stayed at a high level during nighttime when RH was high as well (overall diurnal trend  
499 of  $NO_3^-/rBC$  was similar to that of RH, see Figs. S6a and S6d). This result indicates a



likely more important role of nocturnal nitrate formation ( $\text{N}_2\text{O}_5$  hydrolysis) (Pathak et al., 2011) (Sun et al., 2011) than photochemical production of nitrate during this campaign; moreover, low temperature and high RH favor nitrate partitioning into the particle phase during nighttime too (Gao et al., 2011). For sulfate, The SOR increased with  $\text{O}_x$  while it increased with RH under relatively dry conditions ( $<60\%$ ) but decreased with RH when  $\text{RH}>60\%$ . This result, on the other hand, highlights that photochemical production is more important than aqueous/heterogenous formation for sulfate. In addition, mass ratios of sulfate to  $r\text{BC}$  were enhanced remarkably during daytime and peaked in the afternoon (Fig. S6e), supporting the key role of photochemical formation of sulfate. Sulfate precursor  $\text{SO}_2$  was at a high level during daytime too. The main formation pathway of sulfate is strikingly different from that observed in winter Nanjing (Wu et al., 2018), suggesting significant seasonal variability of sulfate formation even in the same region.

513

### 514 3.3.3 Evolution of organics

Organics had a broad average size distribution (Fig. 5c), but unlike  $r\text{BC}$ , its main peak appeared at 500–600 nm. Figure S5b shows that the majority of organics partitioned in  $r\text{BCc}$  with  $R_{\text{BC}}$  of 5.0–9.0 and wide size coverage (300–800 nm). Regarding its dependences on  $\text{O}_x$  and RH, it mainly accumulated at  $\text{O}_x>70$  ppb (Fig. 8c) and very high RH ( $\sim 90\%$ ) (Fig. 8f). Consistently, O/C ratio and OSc both peaked in the afternoon (Fig. S6b), all suggesting a critical role of photochemistry in affecting the behavior of organics.

Figure 9 illustrates the mass ratios of each OA factor to  $r\text{BC}$  varying with  $\text{O}_x$  and RH. Mass ratios of all four POA factors generally presented decreasing trends (despite some fluctuations) against  $\text{O}_x$  (Fig. 9b), and the total  $\text{POA}_{r\text{BC}}$  showed evident decrease with increase of  $\text{O}_x$  (Fig. 9d). Instead, both  $\text{LO-OOA}_{r\text{BC}}$  and  $\text{MO-OOA}_{r\text{BC}}$ , as well as their sum ( $\text{SOA}_{r\text{BC}}$ ) increased continuously with  $\text{O}_x$  (Figs. 9c and 9d). This result proves that photochemical oxidation contributed significantly to both  $\text{LO-OOA}_{r\text{BC}}$  and  $\text{MO-OOA}_{r\text{BC}}$  formations. Comparatively, decreases of  $\text{POA}_{r\text{BC}}$  perhaps point to its reaction loss upon photochemical oxidation. With regard to RH, besides WS-HOA, the other



three POA<sub>rBC</sub> factors showed almost no dependences on RH (Fig. 9f); note the increase of WS-HOA with RH did not indicate the aqueous production of WS-HOA (see discussion in Section 3.2.1), but a result of enhanced dissolution with increase of moisture. Overall small increase of POA<sub>rBC</sub> (Fig. 9h) with RH then mainly attributed to WS-HOA. Contrary to the trends with O<sub>X</sub>, mass ratios of two SOA factors as well as the total SOA to *rBC* went down with increase of RH (Figs. 9g and 9h), indicating a trivial role of aqueous/heterogenous oxidation for the SOA coated on *rBC* observed during this campaign.

The aging of OA is generally characterized by the increase of O/C and decrease of H/C (Ng et al., 2011; Zhao et al., 2019). The different aging pathways of OA follow different slopes in the Van Krevelen (VK) diagram (Heald et al., 2010). For example, addition of only one oxygen atom to the carbon skeleton results in a slope equal to 0, while replacement of the hydrogen atom with a carboxylic acid group (–COOH) results in a slope of –1 without fragmentation (C–C bond breaking), and –0.5 with fragmentation (Heald et al., 2010; Ng et al., 2011; Zhao et al., 2019). As presented in Fig. 10a, fitting of all OA data yielded a slope of –0.96, very close to –1, suggesting that OA ageing process resembled the hydrogen substitution with a –COOH group (carboxylation). Interestingly, the four OA factors (HOA-rich, WS-HOA, LO-OOA<sub>rBC</sub> and MO-OOA<sub>rBC</sub>) aligned almost in a straight line with a slope of –0.77 (Fig. 10a), also close to –1. Figure S7 further reveals that there was a strong anti-correlation between mass fractions of sum of HOA-rich and WS-HOA and sum of LO-OOA<sub>rBC</sub> and MO-OOA<sub>rBC</sub> ( $r = -0.97$ ); the slope of fitted line was –0.86. All these results suggest that OA evolution may contain a channel of photochemical transformations from HOA-rich and WS-HOA to LO-OOA<sub>rBC</sub> and then to MO-OOA<sub>rBC</sub>. This result is also in line with the observed decrease of POA<sub>rBC</sub> and increase of SOA<sub>rBC</sub> against O<sub>X</sub>.

CHO<sup>+</sup>, CHO<sub>2</sub><sup>+</sup> and C<sub>2</sub>H<sub>2</sub>O<sub>2</sub><sup>+</sup> are the AMS tracer ions for carbonyl, carboxylic acid and glyoxal, respectively (Wang et al., 2020b; Canagaratna et al., 2015a; Yu et al., 2014). Mass fractions of these three fragment ions presented decreasing trends (or no clear trends) against RH (Fig. S8), suggesting again that aqueous processing is not an important pathway in OA evolution during this campaign. Conversely, fractional



contributions of these three ions presented increasing trends versus  $O_x$ , supporting the dominance of photochemical oxidation pathway (Figs. 10b-d). Figures 10e-g show the scatter plots of H/C versus O/C at different  $O_x$  concentrations. The regressed slope was  $-1.03$  for low  $O_x$  ( $<60$  ppb) conditions (Fig. 10e), indicating that the OA aging at low  $O_x$  level is mainly analogue to the carboxylation process. This result corresponds precisely to the fact that mass fractions of  $CHO_2^+$  and  $C_2H_2O_2^+$  increased gradually with  $O_x$  at low  $O_x$ , whereas that of  $CHO^+$  remained essentially unchanged (Figs. 10e-g). The fitted slope was  $-1.14$  for moderate  $O_x$  conditions (60-80 ppb), and it changed to  $-0.43$  for high  $O_x$  level ( $>80$  ppb) but correlation became weaker ( $r=-0.57$ ). This result implies that the OA evolution under high  $O_x$  conditions might include oxidations by the additions of alcohols, peroxides and carboxylation. In all, the evolution of  $rBCc$  OA in Shanghai during this campaign is governed by photochemistry rather than aqueous chemistry, but with different oxidation pathways at different  $O_x$  levels.

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### 3.3.4 Coating time of secondary species onto $rBC$

Although the  $rBCc$  organics was dominated by primary sources (Fig. 2g), the diurnal variations of  $OSc$ , O/C and H/C of the total organics, were controlled predominantly by the two SOA factors. Correlations between the diurnal cycles of  $MO-OOA_{rBC}/rBC$  and  $OSc$  were extremely well ( $r=0.97$  with  $OSc$  and  $r=0.98$  with O/C), and those of  $LO-OOA_{rBC}$  were also tight ( $r=0.91$  with  $OSc$  and  $r=0.92$  with O/C). The correlations with  $LO-OOA_{rBC}$  were a bit weaker than those of  $MO-OOA_{rBC}$ , indicating that  $MO-OOA_{rBC}$  was probably the final products and was more important in governing the overall oxidation level of organic coating. Figure 11a depicts the diurnal variations of  $SOA_{rBC}/rBC$  and  $POA_{rBC}/rBC$ . Diurnal variations of  $POA_{rBC}/rBC$  and  $rBC$  were overall similar, while the daily pattern of  $SOA_{rBC}/rBC$  was almost opposite to that of  $rBC$ . This result likely indicates that most  $POA_{rBC}$  species were co-emitted and coated on  $rBC$  cores originally, therefore the coating process during  $rBCc$  lifecycle was mainly relevant with SOA species rather than POA species. This process is mainly through photochemical reactions, including in-situ oxidation of originally coated POA species (for example, oxidation of HOA-rich plus WS-HOA to LO-



OOA<sub>rBC</sub>, then to MO-OOA<sub>rBC</sub>), and partitioning of secondary species formed in gas-phase reactions onto *r*BC surface. In addition, sulfate and nitrate were both secondarily formed, but sulfate was dominated by photochemical production while nitrate was governed by nocturnal heterogenous formation (as discussed in Section 3.3.2); different diurnal patterns of them (Fig. S6) point to different coating processes too.

In this regard, we hereby propose a concept of average coating time (ACT), which is used to roughly estimate the timescales required for secondary components coated onto *r*BC. The specific method is listed as follows:

1. Move forward the diurnal variation of SA<sub>rBC</sub>/*r*BC (SA represents a secondary aerosol species) for *n* hours to get a new SA<sub>rBC</sub>/*r*BC diurnal pattern, labelled as “SA<sub>rBC</sub>/*r*BC(*r*-*nh*-ahead)”. Here, the *r* value is the linear correlation coefficient between the new SA<sub>rBC</sub>/*r*BC diurnal pattern with that of *r*BC.

2. Choose the best correlation coefficient (*r*-*nh*-ahead), and *nh* corresponds to the ACT for this secondary component.

Diurnal patterns of LO-OOA<sub>rBC</sub>/*r*BC and MO-OOA<sub>rBC</sub>/*r*BC are depicted in Fig. 11b. They were both opposite to the trend of *r*BC, and they were similar to each other, except that MO-OOA<sub>rBC</sub>/*r*BC peaked hours later in the afternoon, signifying that the MO-OOA<sub>rBC</sub> needs longer time to coated on *r*BC than LO-OOA<sub>rBC</sub> does, consistent with the fact that MO-OOA<sub>rBC</sub> was “more aged”. Correspondingly, we obtained an ACT of 7 hours for MO-OOA<sub>rBC</sub> (0.35-7h-ahead) and 5 hours for LO-OOA<sub>rBC</sub> (0.57-5h-ahead) (Fig. 11c) using the method described above. The ACT of sulfate (0.65-7h-ahead) and nitrate (0.30-19h-ahead) were 7 and 19 hours, respectively (Fig. 11d). The results suggest that the *r*BC emitted mainly in the morning rush hours requires a few hours to be adequately coated by LO-OOA<sub>rBC</sub>, MO-OOA<sub>rBC</sub> and sulfate, as these three species are mainly produced in the afternoon by photochemical reactions; while photochemical production of nitrate is insignificant, thus *r*BC was coated by nitrate until later night when nitrate was formed efficiently by heterogenous N<sub>2</sub>O<sub>5</sub> hydrolysis. Note the best *r* values were not high (for example, 0.35 for MO-OOA<sub>rBC</sub> and 0.30 for nitrate), yet the adjusted trends of all secondary components (Figs. 11c and 11d) matched that of *r*BC quite well during 3:00~12:00 (*r* of 0.90, 0.91, 0.84 and 0.84 for MO-OOA<sub>rBC</sub>, LO-



OOA<sub>rBC</sub>, sulfate and nitrate, respectively), likely meaning that daytime variations of two SOAs and sulfate were indeed controlled by the coating process while governing factors of their nighttime variations might be complex, and *vice versa* for nitrate.

### 3.4 Size-resolved hygroscopicity of *rBCc*

By using the method in Section 2.3, we calculated the size-resolved hygroscopicity parameters for the total *rBCc* ( $\kappa_{rBCc}$ ) and the coatings materials ( $\kappa_{CT}$ ) across the whole campaign. The image plots are illustrated in Figs. 12a and b. Generally, large  $\kappa_{rBCc}$  and  $\kappa_{CT}$  values occurred at big particle sizes, and this result was overall similar to that in Nanjing during winter (Wu et al., 2019). We further illustrate the size-resolved  $\kappa_{rBCc}$  as a function of  $R_{BC}$  in Fig. 12c. The figure shows that  $\kappa_{rBCc}$  overall became larger with increasing particle size regardless of the coating thickness. However, there were some (though not significant) relatively high  $\kappa_{rBCc}$  values in the range of 80-150 nm (bottom left and bottom right in Fig. 12c).

In Figs. 13a and b, we further determined the average size-resolved  $\kappa_{rBCc}$  and  $\kappa_{CT}$  as a function of coated diameter ( $D_{rBCc}$ ). Both  $\kappa_{rBCc}$  and  $\kappa_{CT}$  values were relatively high at  $D_{rBCc} < 100$  nm and presented slight decreases from 100 to 150 nm. This is distinctive from those observed in Nanjing, where  $\kappa_{CT}$  increased with  $D_{rBCc}$  from 50 nm (Wu et al., 2019). From Figs. 5b and 5d, it can be seen the *rBCc* with  $D_{rBCc} < 150$  nm was dominated by a portion of ammonium and sulfate (8-10%) and organics (~60%), of which organics was dominated by WS-HOA. Such composition explains the relatively high hygroscopicity at  $D_{rBCc} < 150$  nm as both ammonium sulfate and WS-HOA are hydrophilic; a slight decrease of the hygroscopicity from 100 to 150 nm  $D_{rBCc}$  was also a response of decreased mass contributions of ammonium sulfate and WS-HOA (and increase of hydrophobic HOA-rich contribution).

Figures 13a and b provide the fitted exponential functions for the mean  $\kappa_{rBCc}$  and  $\kappa_{CT}$  with  $D_{rBCc}$ . The equations are:  $\kappa_{rBCc}(x) = 0.29 - 0.14 \times \exp(-0.006 \times x)$  and  $\kappa_{CT}(x) = 0.35 - 0.09 \times \exp(-0.003 \times x)$  ( $x$  is  $D_{rBCc}$ ,  $150 < x < 1000$  nm). Here, 0.29 and 0.35 are the upper limits of  $\kappa_{rBCc}$  and  $\kappa_{CT}$ , higher than those reported in Nanjing (0.28 and 0.30 for  $\kappa_{rBCc}$  and  $\kappa_{CT}$ ); yet the increasing rates of  $\kappa_{rBCc}$  and  $\kappa_{CT}$  with  $D_{rBCc}$  are 0.14 and 0.09,



650 respectively, which are much lower than those from Nanjing (0.35 and 0.27 for  $\kappa_{rBCc}$   
 651 and  $\kappa_{CT}$ ) (Wu et al., 2019). Smaller increased contributions of hydrophilic secondary  
 652 species from 150 to 1000 nm in Shanghai than those from 100 to 1000 nm in Nanjing  
 653 are likely the cause of smaller increasing rates of hygroscopicity parameters. In addition,  
 654 it should be noted that the hygroscopicity is not only determined by the bulk  
 655 composition, but also affected by the phase state of particles. For instance, a recent  
 656 study reveals that the hygroscopic growth of phase-separated particles (with ammonium  
 657 sulfate as cores) can be reduced by the secondary organic shells and is dependent on  
 658 the thickness of organic coating (Li et al., 2021).

659 The critical supersaturation ( $SS_c$ ) for a selected dry diameter ( $D_{rBCc}$  measured by  
 660 SP-AMS) of a  $rBCc$  particle with a hygroscopicity parameter  $\kappa_{rBCc}$ , can be calculated  
 661 by the “ $\kappa$ -Kohler theory” equation (Petters and Kreidenweis, 2007). Based on the size-  
 662 resolved  $\kappa_{rBCc}$ , the CCN activation diameter ( $D_{50}$ ) of particles at a given critical  $SS_c$  can  
 663 be calculated (Wu et al., 2019). Then, by combining the measured  $rBCc$  number size  
 664 distribution and the  $D_{50}$  value, activation fraction ( $f_{AC}$ ) of  $rBCc$  number population (i.e.,  
 665 the fraction greater than the  $D_{50}$ ) can be obtained. Figure 13c shows the  $SS_c$  as a  
 666 function of  $D_{rBCc}$  for the entire sampling period to obtain the  $D_{50}$  at a specific  $SS_c$ . The  
 667  $D_{50}$  values of the  $rBCc$  particles were determined to be  $166 \pm 16$  nm and  $110 \pm 5$  nm  
 668 for  $SS_c$  of 0.1% and 0.2%, respectively. The two  $D_{50}$  values are both smaller than those  
 669 determined for  $rBCc$  particle in Nanjing (Wu et al., 2019), likely owing to the presence  
 670 of WS-HOA in Shanghai. Figure 13d shows the  $f_{AC}$  at  $SS_c$  of 0.1% ( $D_{50}$  of 166 nm) was  
 671  $16 \pm 3\%$ , and the  $f_{AC}$  at  $SS_c$  of 0.2% ( $D_{50}$  of 110 nm) was  $59 \pm 4\%$ .

672

### 673 3.5 A case study influenced by ship emissions

#### 674 3.5.1 Potential source areas of $rBCc$

675 To explore the potential geographic origins of  $rBCc$  at the receptor site, the hybrid  
 676 single-particle Lagrangian integrated trajectory (HYSPLIT) model (version 4.9) was  
 677 applied here. Figure 14a shows that the backward trajectories were classified into four  
 678 air mass clusters, including one long-range transport from northern sea (Cluster1), one  
 679 long-range transport from northeastern sea (Cluster2), a local one from eastern ports



(Cluster3), and one from northwestern inland region (Cluster4). The four clusters occupied 23.8%, 33.8%, 37.3%, and 5.0% of the total trajectories, respectively. It is clear that the sampling period was influenced by offshore air masses in most of the time (95%). Cluster3 had the highest mean  $rBC_c$  concentrations ( $13.2 \pm 10.9 \mu\text{g m}^{-3}$ ) while the mean concentrations of the other three clusters were apparently lower ( $4.3\sim 5.2 \mu\text{g m}^{-3}$ ). This result is plausible as Cluster3 was the shortest in length therefore was least diluted compared with other three clusters. Average chemical compositions of the  $rBC_c$  from four clusters (Fig. 14b) showed some differences as well: Cluster1 and Cluster2 had higher MO-OOA <sub>$rBC$</sub>  contributions, possibly owing to the interceptions of more aged SOA species during the long-range transports; While Cluster4 had less MO-OOA <sub>$rBC$</sub>  but a bit more POA <sub>$rBC$</sub>  contributions, likely attributing to more primary species emitted in inland regions (similarly, a higher fraction of nitrate was likely because of enhanced NO<sub>2</sub> emissions).

693

### 694 3.5.2 A typical case of $rBC_c$ affected by ship emissions

Ship emissions are found to have an important impact on the air quality of Shanghai and the Yangtze River Delta (Zhao et al., 2013; Fan et al., 2016; Liu et al., 2017b; Chen et al., 2019). The ship engines usually burn heavy fuel oil (HFO), and vanadium (V) and nickel (Ni) can be adopted as reliable tracers for the ship-emitted particles (Ault et al., 2009; Moldanová et al., 2009; Ault et al., 2010). The long-term variation of Ni/V ratio in ship-emitted particles in Shanghai has been recently reported (Yu et al., 2021). The main conclusion is that Ni/V ratio was close to 0.4 in 2018, while it became to be greater than 2.0 in 2020. Our measurement was carried out in 2018, and we therefore chose a period from November 3 to 5 as a typical case affected by ship emissions (SEP period), since the average Ni/V ratio was  $\sim 0.50$  and high concentrations of both Ni and V were found. Figure S9 shows the concentration-weighted trajectories (CWT) of ship emission tracers (V, Ni),  $rBC$  and  $rBC$ -rich factor during SEP, displaying that these components were mainly from sea. During SEP, the correlation coefficient ( $r$ ) between V(Ni) and  $rBC$  was 0.69 (0.74), indicating the SEP was indeed impacted by ship emissions.





Figure 15 displays the chemical characteristics of *r*BCc and NR-PM<sub>1</sub> components (especially the OA factors), V and Ni, gaseous pollutants and the meteorological parameters during SEP. As a comparison, we also selected a period with no ship emissions with the same time span as SEP (from 0:00 on November 10 to 0:00 on November 12, termed as non-SEP), and a similar plot during non-SEP is presented in Fig. S10. During SEP, the wind was mainly from east, and the average wind speed was  $\sim 0.5 \text{ m s}^{-1}$  (Fig. 15a); Wusong, Luojing and Waigaoqiao ports located northeast of the sampling site (Fig. S1). Instead, the wind was mainly from north during non-SEP (Fig. S10a). During SEP, the average mass concentrations of V and Ni were  $5.8 \text{ ng m}^{-3}$  and  $2.9 \text{ ng m}^{-3}$ , respectively, while those during non-SEP were only  $2.9 \text{ ng m}^{-3}$  and  $2.6 \text{ ng m}^{-3}$ . The average mass ratio of V/Ni during SEP was 2.0 in agreement with those reported in ship-influenced PM<sub>1</sub> (Mazzei et al., 2008; Mar et al., 2009), and within the range of 1.9 to 3.5 for domestic HFO (Zhao et al., 2013), while the average ratio of V/Ni (1.1) during non-SEP was outside the range. Moreover, the major air pollutants emitted from ships include nitrogen oxides (NO<sub>x</sub>), sulfur oxides (SO<sub>2</sub>), carbon monoxide (CO), hydrocarbons and primary/secondary particles (Becagli et al., 2017; Wu et al., 2021). As demonstrated in Fig. 15d and Fig. S10d, SO<sub>2</sub> concentration was overall higher during SEP (10:00–16:00 on November 11 not included); mass loading of NO<sub>2</sub> was higher during the rush hours of non-SEP, but was higher at night (no traffic) during SEP. Mass proportion of sulfate in NR-PM<sub>1</sub> during SEP was also higher than that during non-SEP (21.0% vs. 18.7%), but *vice versa* for nitrate (18.8% during SEP vs. 25.4% during non-SEP).

We further investigated the dependences of *r*BCc and NR-PM<sub>1</sub> species on V during SEP and non-SEP, as shown in Fig. S11. Here we only used V since Ni level might be influenced by other emission sources, such as refining industry (Jang et al., 2007; Kim et al., 2014) in urban Shanghai, and during non-SEP, Ni still presented a good correlation with *r*BC ( $r=0.80$ ). During SEP, V concentrations (most of them  $>4 \text{ ng m}^{-3}$ ) overall positively correlated with both sulfate and nitrate (except for *r*BCc sulfate) (Fig. S11a). Considering that V concentration was independently measured for all fine particles, a better correlation with total NR-PM<sub>1</sub> sulfate than it with *r*BCc sulfate is



reasonable. Similarly, V-rich particles had positive correlations with traffic-related OA and LO-OOA no matter in  $rBC_c$  or in total NR- $PM_{10}$  (Fig. S11b). Conversely, during non-SEP, particles with low-V content (most of them  $<4 \text{ ng m}^{-3}$ ) had no clear links with sulfate, nitrate, POA and SOA species (even anti-correlations for  $V > 2.5 \text{ ng m}^{-3}$ ) (Figs. S11c and S11d; a detailed comparison of the correlation coefficients of V with OA factors of  $rBC_c$  and NR- $PM_{10}$  organics are provided in Table S2). The results above demonstrate that during SEP, chemical properties of the particles (both fresh and aged ones) were clearly linked with ship influences.

Previous studies (Ault et al., 2009; Ault et al., 2010; Liu et al., 2017b) have shown that the fresh ship-emitted V-rich particles are typically accompanied by high sulfate contribution, while those aged V-containing ship particles are on the other hand with relatively high nitrate contribution. In order to further probe chemical characteristics and evolution processes of particles influenced by ship emissions, we divided SEP into three episodes (i.e., EP1-EP3) (marked in Fig. 15). During EP1,  $rBC$  content was highest (Fig. 15k) and coating was thinnest (Fig. S12i) indicating the particles were relatively fresh, nevertheless the SOA contents were not low (Figs. 15j and 15l), OSC was moderate (Fig. S12j), sulfate portions in NR- $PM_{10}$  and  $rBC_c$  were both the highest (26.5%), and nitrate portion was the lowest (9.8%) among three episodes (Figs. 15i and 15k). Such composition is not common in other cases, demonstrating it was a specific period impacted by fresh ship emissions. EP2 had the highest mass loadings of V, gas pollutants as well as the lowest planetary boundary layer (PBL) height ( $\sim 200$  meters) (Fig. S12). Mass contributions of nitrate increased and sulfate decreased, and  $rBC$  content decreased from those during EP1, signifying that the particle population likely contained some aged ship-emitted particles. Of course, particle composition during EP2 was also influenced by the formation mechanisms of secondary species: EP2 was mostly during nighttime, therefore photochemical formation of sulfate and SOA were weak (contributions were low as shown in Figs. 15i-l), while nitrate formation was enhanced due to the nocturnal process. During EP3,  $rBC$  was lowest, sulfate and V were moderate, nitrate and SOA were highest in both  $rBC_c$  and NR- $PM_{10}$ , OSC and  $R_{BC}$  were highest in  $rBC_c$  among the three episodes (Figs. 15i-l and Figs. S12i-k), therefore it



was also a period with influence from aged ship-emitted particles; the difference from EP2 is that photochemically formed sulfate and SOA were still significant as EP3 was in the later afternoon and early evening, though heterogeneously formed nitrate played a non-negligible role too (see the increase of RH, increase of nitrate and decrease of O<sub>3</sub> concentrations from the beginning of EP3 in Fig. 15).

#### 4. Conclusions and implications

We conducted a field measurement during November of 2018 in urban Shanghai, China, focusing on the elucidation of physical and chemical properties of the ambient particles containing *r*BC cores (*r*BCc) by using a laser-only SP-AMS together with a HR-AMS. The campaign-average *r*BCc was 4.6 μg m<sup>-3</sup>, occupying ~19.1% mass of the total NR-PM<sub>1</sub>. The average mass ratio of coating to *r*BC cores (*R*<sub>BC</sub>) was ~5.0, indicating an overall thick coating, compared with the *r*BCc near combustion source. Sulfate was found to preferentially condense on non-*r*BC particles thus led to a low fraction of *r*BCc sulfate to that in NR-PM<sub>1</sub> (7.4%), while distribution of nitrate between *r*BCc and non-*r*BC particles showed no obvious difference. PMF analysis on *r*BCc and NR-PM<sub>1</sub> OA reveals that cooking-related organics were externally mixed with *r*BC, and a small portion of organics from biomass burning was only present in *r*BCc; the traffic-related OA species, however, was in a large part internally mixed with *r*BC.

A regression algorithm was applied to deconvolute the size distributions of individual *r*BCc OA factors, and results show that small *r*BCc particles were predominantly generated from traffic, and such particles could grow bigger because of condensation of secondary inorganic and organic components, resulting in thick coating. Investigation on diurnal patterns of the *r*BCc species reveals that sulfate and two SOA factors (LO-OOA<sub>*r*BC</sub> and MO-OOA<sub>*r*BC</sub>) were generated mainly through daytime photochemical oxidation; nitrate, on the other hand, was governed mainly by the nocturnal N<sub>2</sub>O<sub>5</sub> hydrolysis. Partial SOA was found to be produced from in-situ photochemical conversion from traffic-related POA. An average coating time (ACT) was proposed to roughly estimate the timescales for the secondary species to coat on *r*BC, and the ACT of sulfate, LO-OOA<sub>*r*BC</sub>, MO-OOA<sub>*r*BC</sub> and nitrate were approximately



800 5, 5, 7 and 19 hours, respectively.

801 Moreover, the size-resolved hygroscopicity parameters of *r*BCc particles ( $\kappa_{rBCc}$ )  
 802 and the coating material ( $\kappa_{CT}$ ) were obtained based on the elucidated composition of  
 803 *r*BCc particles. The fitted equations are:  $\kappa_{rBCc}(x) = 0.29 - 0.14 \times \exp(-0.006 \times x)$  and  
 804  $\kappa_{CT}(x) = 0.35 - 0.09 \times \exp(-0.003 \times x)$  ( $x$  is  $D_{rBCc}$ ,  $150 < x < 1000$  nm). The minimums  
 805 of both  $\kappa_{rBCc}$  and  $\kappa_{CT}$  were at  $\sim 150$  nm due to the abundances of hydrophobic *r*BC cores  
 806 and traffic-related HOA at this size. Under critical supersaturations ( $SS_c$ ) of 0.1% and  
 807 0.2%, the  $D_{50}$  values were estimated to be  $166 \pm 16$  and  $110 \pm 5$  nm, and the activated  
 808 number fractions of *r*BCc particles were  $16 \pm 3\%$  and  $59 \pm 4\%$ , respectively. Finally, a  
 809 typical case with influence from ship emissions was investigated. During this period,  
 810 the *r*BCc particles were enriched in V (on average  $5.8 \text{ ng m}^{-3}$ ), with a V/Ni ratio of 2.0,  
 811 and various secondary formation processes affect the ship-emitted particles at different  
 812 times of the day.

813 In summary, the findings from this comprehensive study on *r*BCc provide rich  
 814 information regarding the various primary sources and secondary formation pathways  
 815 of species coated on *r*BC, as well as the features of distributions of those species  
 816 between *r*BC and non-*r*BC particles. In particular, different types of *r*BCc particles  
 817 from diesel and gasoline vehicle emissions were resolved and elucidation of their  
 818 properties are useful to their future effective control. Understanding of the formation  
 819 processes and coating timescales of secondary components is helpful to understand the  
 820 impact of ambient BC particles too. At last, the parameterized relationship of  
 821 hygroscopicity with size distribution is useful for modelling the climate effect of *r*BC  
 822 (alteration of cloud properties).

823

824 *Data availability.* The data in this study are available from the authors upon request  
 825 (caxinra@163.com).

826

827 *Supplement.* The supplement related to this article is available online at: XXX

828

829 *Author contributions.* SJC, DDH, YZW, JFW, FZS, and XLG conducted the field



measurement. SJC, DDH, YZW, JFW, and JKX analyzed the data. YJZ, HLW, CH and  
HL reviewed the paper and provide useful suggestions. SJC and XLG wrote the paper.

832

*Competing interests.* The contact author has declared that neither they nor their co-  
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846

## 847 **References**

Aiken, A. C., DeCarlo, P. F., Kroll, J. H., Worsnop, D. R., Huffman, J. A., Docherty,  
K. S., Ulbrich, I. M., Mohr, C., Kimmel, J. R., Sueper, D., Sun, Y., Zhang, Q., Trimborn,  
A., Northway, M., Ziemann, P. J., Canagaratna, M. R., Onasch, T. B., Alfarra, M. R.,  
Prevot, A. S. H., Dommen, J., Duplissy, J., Metzger, A., Baltensperger, U., and Jimenez,  
J. L.: O/C and OM/OC ratios of primary, secondary, and ambient organic aerosols with  
high-resolution time-of-flight aerosol mass spectrometry, *Environ. Sci. Technol.*, 42,  
4478-4485, <https://doi.org/10.1021/es703009q>, 2008.

Ault, A. P., Moore, M. J., Furutani, H., and Prather, K. A.: Impact of emissions  
from the Los Angeles port region on San Diego air quality during regional transport  
events, *Environ. Sci. Technol.*, 43, 3500-3506, <https://doi.org/10.1021/es8018918>,  
2009.

Ault, A. P., Gaston, C. I., Ying, W., Gerardo, D., Thiemens, M. H., and Prather, K.



- 
- 860 A.: Characterization of the single particle mixing state of individual ship plume events  
 861 measured at the Port of Los Angeles, *Environ. Sci. Technol.*, 44, 1954-1961,  
 862 <https://doi.org/10.1021/es902985h>, 2010.
- 863 Becagli, S., Anello, F., Bommarito, C., Cassola, F., Calzolari, G., Di Iorio, T., di  
 864 Sarra, A., Gómez-Amo, J.-L., Lucarelli, F., Marconi, M., Meloni, D., Monteleone, F.,  
 865 Nava, S., Pace, G., Severi, M., Sferlazzo, D. M., Traversi, R., and Udisti, R.:  
 866 Constraining the ship contribution to the aerosol of the central Mediterranean, *Atmos.*  
 867 *Chem. Phys.*, 17, 2067-2084, <https://doi.org/10.5194/acp-17-2067-2017>, 2017.
- 868 Bond, T. C., Doherty, S. J., Fahey, D. W., Forster, P. M., Berntsen, T., DeAngelo,  
 869 B. J., Flanner, M. G., Ghan, S., Kärcher, B., Koch, D., Kinne, S., Kondo, Y., Quinn, P.  
 870 K., Sarofim, M. C., Schultz, M. G., Schulz, M., Venkataraman, C., Zhang, H., Zhang,  
 871 S., Bellouin, N., Guttikunda, S. K., Hopke, P. K., Jacobson, M. Z., Kaiser, J. W.,  
 872 Klimont, Z., Lohmann, U., Schwarz, J. P., Shindell, D., Storelvmo, T., Warren, S. G.,  
 873 and Zender, C. S.: Bounding the role of black carbon in the climate system: A scientific  
 874 assessment, *J. Geophys. Res.: Atmos.*, 118, 5380-5552,  
 875 <https://doi.org/10.1002/jgrd.50171>, 2013.
- 876 Canagaratna, M. R., Jayne, J. T., Jimenez, J. L., Allan, J. D., Alfarra, M. R., Zhang,  
 877 Q., Onasch, T. B., Drewnick, F., Coe, H., Middlebrook, A., Delia, A., Williams, L. R.,  
 878 Trimborn, A. M., Northway, M. J., DeCarlo, P. F., Kolb, C. E., Davidovits, P., and  
 879 Worsnop, D. R.: Chemical and microphysical characterization of ambient aerosols with  
 880 the aerodyne aerosol mass spectrometer, *Mass Spectrom. Rev.*, 26, 185-222,  
 881 <https://doi.org/10.1002/mas.20115>, 2007.
- 882 Canagaratna, M. R., Jimenez, J. L., Kroll, J. H., Chen, Q., Kessler, S. H., Massoli,  
 883 P., Hildebrandt Ruiz, L., Fortner, E., Williams, L. R., Wilson, K. R., Surratt, J. D.,  
 884 Donahue, N. M., Jayne, J. T., and Worsnop, D. R.: Elemental ratio measurements of  
 885 organic compounds using aerosol mass spectrometry: characterization, improved  
 886 calibration, and implications, *Atmos. Chem. Phys.*, 15, 253-272,  
 887 <https://doi.org/10.5194/acp-15-253-2015>, 2015a.
- 888 Canagaratna, M. R., Massoli, P., Browne, E. C., Franklin, J. P., Wilson, K. R.,  
 889 Onasch, T. B., Kirchstetter, T. W., Fortner, E. C., Kolb, C. E., Jayne, J. T., Kroll, J. H.,



- 890 and Worsnop, D. R.: Chemical compositions of black carbon particle cores and coatings  
891 via soot particle aerosol mass spectrometry with photoionization and electron ionization,  
892 *J. Phys. Chem. A*, 119, 4589-4599, <https://doi.org/10.1021/jp510711u>, 2015b.
- 893 Cao, L.-M., Wei, J., He, L.-Y., Zeng, H., Li, M.-L., Zhu, Q., Yu, G.-H., and Huang,  
894 X.-F.: Aqueous aging of secondary organic aerosol coating onto black carbon: Insights  
895 from simultaneous L-ToF-AMS and SP-AMS measurements at an urban site in  
896 southern China, *J. Clean Prod.*, 330, 129888,  
897 <https://doi.org/10.1016/j.jclepro.2021.129888>, 2022.
- 898 Cappa, C. D., Onasch, T. B., Massoli, P., Worsnop, D. R., Bates, T. S., Cross, E.  
899 S., Davidovits, P., Hakala, J., Hayden, K. L., Jobson, B. T., Kolesar, K. R., Lack, D. A.,  
900 Lerner, B. M., Li, S.-M., Mellon, D., Nuaaman, I., Olfert, J. S., Petäjä, T., Quinn, P. K.,  
901 Song, C., Subramanian, R., Williams, E. J., and Zaveri, R. A.: Radiative Absorption  
902 Enhancements Due to the Mixing State of Atmospheric Black Carbon, *Science*, 337,  
903 1078-1081, <https://doi.org/10.1126/science.1230260>, 2012.
- 904 Chang, R. Y. W., Slowik, J. G., Shantz, N. C., Vlasenko, A., Liggio, J., Sjostedt, S.  
905 J., Leaitch, W. R., and Abbatt, J. P. D.: The hygroscopicity parameter ( $\kappa$ ) of ambient  
906 organic aerosol at a field site subject to biogenic and anthropogenic influences:  
907 relationship to degree of aerosol oxidation, *Atmos. Chem. Phys.*, 10, 5047-5064,  
908 <https://doi.org/10.5194/acp-10-5047-2010>, 2010.
- 909 Chen, B., Bai, Z., Cui, X., Chen, J., Andersson, A., and Gustafsson, O.: Light  
910 absorption enhancement of black carbon from urban haze in Northern China winter,  
911 *Environ. Pollut.*, 221, 418-426, <https://doi.org/10.1016/j.envpol.2016.12.004>, 2017.
- 912 Chen, C., Fan, X., Shaltout, T., Qiu, C., Ma, Y., Goldman, A., and Khalizov, A. F.:  
913 An unexpected restructuring of combustion soot aggregates by subnanometer coatings  
914 of polycyclic aromatic hydrocarbons, *Geophys. Res. Lett.*, 43, 11,080-011,088,  
915 <https://doi.org/10.1002/2016GL070877>, 2016.
- 916 Chen, C., Enekwizu, O. Y., Fan, X., Dobrzanski, C. D., Ivanova, E. V., Ma, Y., Gor,  
917 G. Y., and Khalizov, A. F.: Single Parameter for Predicting the Morphology of  
918 Atmospheric Black Carbon, *Environ. Sci. Technol.*, 52, 14169-14179,  
919 <https://doi.org/10.1021/acs.est.8b04201>, 2018.



- 920       Chen, D., Tian, X., Lang, J., Zhou, Y., Li, Y., Guo, X., Wang, W., and Liu, B.: The  
921       impact of ship emissions on PM<sub>2.5</sub> and the deposition of nitrogen and sulfur in Yangtze  
922       River Delta, China, *Sci. Total Environ.*, 649, 1609-1619,  
923       <https://doi.org/10.1016/j.scitotenv.2018.08.313>, 2019.
- 924       Clegg, S. L., Brimblecombe, P., and Wexler, A. S.: Thermodynamic model of the  
925       system  $\text{H}^+ \text{NH}_4^+ \text{SO}_4^{2-} \text{NO}_3^- \text{H}_2\text{O}$  at tropospheric temperatures, 102, 2155-2171,  
926       <https://doi.org/10.1021/jp973043j>, 1998.
- 927       Collier, S., Williams, L. R., Onasch, T. B., Cappa, C. D., Zhang, X., Russell, L.  
928       M., Chen, C.-L., Sanchez, K. J., Worsnop, D. R., and Zhang, Q.: Influence of Emissions  
929       and Aqueous Processing on Particles Containing Black Carbon in a Polluted Urban  
930       Environment: Insights From a Soot Particle-Aerosol Mass Spectrometer, *J. Geophys.*  
931       *Res.: Atmos.*, 123, 6648-6666, <https://doi.org/10.1002/2017jd027851>, 2018.
- 932       Corbin, J. C., Sierau, B., Gysel, M., Laborde, M., Keller, A., Kim, J., Petzold, A.,  
933       Onasch, T. B., Lohmann, U., and Mensah, A. A.: Mass spectrometry of refractory black  
934       carbon particles from six sources: carbon-cluster and oxygenated ions, *Atmos. Chem.*  
935       *Phys.*, 14, 2591-2603, <https://doi.org/10.5194/acp-14-2591-2014>, 2014.
- 936       Cubison, M. J., Ortega, A. M., Hayes, P. L., Farmer, D. K., Day, D., Lechner, M.  
937       J., Brune, W. H., Apel, E., Diskin, G. S., Fisher, J. A., Fuelberg, H. E., Hecobian, A.,  
938       Knapp, D. J., Mikoviny, T., Rierner, D., Sachse, G. W., Sessions, W., Weber, R. J.,  
939       Weinheimer, A. J., Wisthaler, A., and Jimenez, J. L.: Effects of aging on organic aerosol  
940       from open biomass burning smoke in aircraft and laboratory studies, *Atmos. Chem.*  
941       *Phys.*, 11, 12049-12064, <https://doi.org/10.5194/acp-11-12049-2011>, 2011.
- 942       DeCarlo, P. F., Kimmel, J. R., Trimborn, A., Northway, M. J., Jayne, J. T., Aiken,  
943       A. C., Gonin, M., Fuhrer, K., Horvath, T., Docherty, K. S., Worsnop, D. R., and Jimenez,  
944       J. L.: Field-Deployable, High-Resolution, Time-of-Flight Aerosol Mass Spectrometer,  
945       *Anal. Chem.*, 78, 8281-8289, <https://doi.org/10.1021/ac061249n>, 2006.
- 946       Enroth, J., Saarikoski, S., Niemi, J., Kousa, A., Ježek, I., Močnik, G., Carbone, S.,  
947       Kuuluvainen, H., Rönkkö, T., Hillamo, R., and Pirjola, L.: Chemical and physical  
948       characterization of traffic particles in four different highway environments in the





- 
- 949 Helsinki metropolitan area, *Atmos. Chem. Phys.*, 16, 5497-5512,  
 950 <https://doi.org/10.5194/acp-16-5497-2016>, 2016.
- 951 Fan, Q., Zhang, Y., Ma, W., Ma, H., Feng, J., Yu, Q., Yang, X., Ng, S. K., Fu, Q.,  
 952 and Chen, L.: Spatial and Seasonal Dynamics of Ship Emissions over the Yangtze River  
 953 Delta and East China Sea and Their Potential Environmental Influence, *Environ. Sci.*  
 954 *Technol.*, 50, 1322-1329, <https://doi.org/10.1021/acs.est.5b03965>, 2016.
- 955 Gao, X., Yang, L., Cheng, S., Gao, R., Zhou, Y., Xue, L., Shou, Y., Wang, J., Wang,  
 956 X., Nie, W., Xu, P., and Wang, W.: Semi-continuous measurement of water-soluble ions  
 957 in PM<sub>2.5</sub> in Jinan, China: Temporal variations and source apportionments, *Atmos.*  
 958 *Environ.*, 45, 6048-6056, <https://doi.org/10.1016/j.atmosenv.2011.07.041>, 2011.
- 959 Ge, X., Zhang, Q., Sun, Y., Ruehl, C. R., and Setyan, A.: Effect of aqueous-phase  
 960 processing on aerosol chemistry and size distributions in Fresno, California, during  
 961 wintertime, *Environ. Chem.*, 9, 221-235, <https://doi.org/10.1071/EN11168>, 2012.
- 962 Gilardoni, S., Massoli, P., Paglione, M., Giulianelli, L., Carbone, C., Rinaldi, M.,  
 963 Decesari, S., Sandrini, S., Costabile, F., Gobbi, G. P., Pietrogrande, M. C., Visentin, M.,  
 964 Scotto, F., Fuzzi, S., and Facchini, M. C.: Direct observation of aqueous secondary  
 965 organic aerosol from biomass-burning emissions, *Proc. Natl. Acad. Sci. U. S. A.*, 113,  
 966 10013-10018, <https://doi.org/10.1073/pnas.1602212113>, 2016.
- 967 Gong, X., Zhang, C., Chen, H., Nizkorodov, S. A., Chen, J., and Yang, X.: Size  
 968 distribution and mixing state of black carbon particles during a heavy air pollution  
 969 episode in Shanghai, *Atmos. Chem. Phys.*, 16, 5399-5411, [https://doi.org/10.5194/acp-](https://doi.org/10.5194/acp-16-5399-2016)  
 970 [16-5399-2016](https://doi.org/10.5194/acp-16-5399-2016), 2016.
- 971 Gysel, M., Crosier, J., Topping, D. O., Whitehead, J. D., Bower, K. N., Cubison,  
 972 M. J., Williams, P. I., Flynn, M. J., McFiggans, G. B., and Coe, H.: Closure study  
 973 between chemical composition and hygroscopic growth of aerosol particles during  
 974 TORCH2, *Atmos. Chem. Phys.*, 7, 6131-6144, [https://doi.org/10.5194/acp-7-6131-](https://doi.org/10.5194/acp-7-6131-2007)  
 975 [2007](https://doi.org/10.5194/acp-7-6131-2007), 2007.
- 976 Heald, C. L., Kroll, J. H., Jimenez, J. L., Docherty, K. S., DeCarlo, P. F., Aiken, A.  
 977 C., Chen, Q., Martin, S. T., Farmer, D. K., and Artaxo, P.: A simplified description of  
 978 the evolution of organic aerosol composition in the atmosphere, *Geophys. Res. Lett.*,



- 79 37, <https://doi.org/10.1029/2010gl042737>, 2010.
- 80 Hu, D., Liu, D., Kong, S., Zhao, D., Wu, Y., Li, S., Ding, S., Zheng, S., Cheng, Y.,  
81 Hu, K., Deng, Z., Wu, Y., Tian, P., Liu, Q., Huang, M., and Ding, D.: Direct  
82 Quantification of Droplet Activation of Ambient Black Carbon Under Water  
83 Supersaturation, *J. Geophys. Res.: Atmos.*, 126, e2021JD034649,  
84 <https://doi.org/10.1029/2021jd034649>, 2021.
- 85 Jacobi, H. W., Lim, S., Ménégot, M., Ginot, P., Laj, P., Bonasoni, P., Stocchi, P.,  
86 Marinoni, A., and Arnaud, Y.: Black carbon in snow in the upper Himalayan Khumbu  
87 Valley, Nepal: observations and modeling of the impact on snow albedo, melting, and  
88 radiative forcing, *The Cryosphere*, 9, 1685-1699, [https://doi.org/10.5194/tc-9-1685-](https://doi.org/10.5194/tc-9-1685-2015)  
89 [2015](https://doi.org/10.5194/tc-9-1685-2015), 2015.
- 90 Jang, H.-N., Seo, Y.-C., Lee, J.-H., Hwang, K.-W., Yoo, J.-I., Sok, C.-H., and Kim,  
91 S.-H.: Formation of fine particles enriched by V and Ni from heavy oil combustion:  
92 anthropogenic sources and drop-tube furnace experiments, *Atmos. Environ.*, 41, 1053-  
93 1063, <https://doi.org/10.1016/j.atmosenv.2006.09.011>, 2007.
- 94 Jayne, J. T., Leard, D. C., Zhang, X., Davidovits, P., Smith, K. A., Kolb, C. E., and  
95 Worsnop, D. R.: Development of an aerosol mass spectrometer for size and composition  
96 analysis of submicron particles, *Aerosol Sci. Technol.*, 33, 49-70,  
97 <https://doi.org/10.1080/027868200410840>, 2000.
- 98 Kim, K.-H., Shon, Z.-H., Mauulida, P. T., and Song, S.-K.: Long-term monitoring  
99 of airborne nickel (Ni) pollution in association with some potential source processes in  
1000 the urban environment, *Chemosphere*, 111, 312-319,  
1001 <https://doi.org/10.1016/j.chemosphere.2014.03.138>, 2014.
- 1002 Kroll, J. H., Donahue, N. M., Jimenez, J. L., Kessler, S. H., Canagaratna, M. R.,  
1003 Wilson, K. R., Altieri, K. E., Mazzoleni, L. R., Wozniak, A. S., Bluhm, H., Mysak, E.  
1004 R., Smith, J. D., Kolb, C. E., and Worsnop, D. R.: Carbon oxidation state as a metric  
1005 for describing the chemistry of atmospheric organic aerosol, *Nat. Chem.*, 3, 133-139,  
1006 <https://doi.org/10.1038/nchem.948>, 2011.
- 1007 Lambe, A. T., Ahern, A. T., Wright, J. P., Croasdale, D. R., Davidovits, P., and  
1008 Onasch, T. B.: Oxidative aging and cloud condensation nuclei activation of laboratory



- 
- 1009 combustion soot, J. Aerosol. Sci., 79, 31-39,  
 1010 <https://doi.org/10.1016/j.jaerosci.2014.10.001>, 2015.
- 1011 Laskin, A., Smith, J. S., and Laskin, J.: Molecular characterization of nitrogen-  
 1012 containing organic compounds in biomass burning aerosols using high-resolution mass  
 1013 spectrometry, Environ. Sci. Technol., 43, 3764-3771,  
 1014 <https://doi.org/10.1021/es803456n>, 2009.
- 1015 Lee, A. K. Y., Willis, M. D., Healy, R. M., Onasch, T. B., and Abbatt, J. P. D.:  
 1016 Mixing state of carbonaceous aerosol in an urban environment: single particle  
 1017 characterization using the soot particle aerosol mass spectrometer (SP-AMS), Atmos.  
 1018 Chem. Phys., 15, 1823-1841, <https://doi.org/10.5194/acp-15-1823-2015>, 2015.
- 1019 Lee, A. K. Y., Chen, C.-L., Liu, J., Price, D. J., Betha, R., Russell, L. M., Zhang,  
 1020 X., and Cappa, C. D.: Formation of secondary organic aerosol coating on black carbon  
 1021 particles near vehicular emissions, Atmos. Chem. Phys., 17, 15055-15067,  
 1022 <https://doi.org/10.5194/acp-17-15055-2017>, 2017.
- 1023 Lee, A. K. Y., Rivellini, L. H., Chen, C. L., Liu, J., Price, D. J., Betha, R., Russell,  
 1024 L. M., Zhang, X., and Cappa, C. D.: Influences of primary emission and secondary  
 1025 coating formation on the particle diversity and mixing state of black carbon particles,  
 1026 Environ. Sci. Technol., 53, 9429-9438, <https://doi.org/10.1021/acs.est.9b03064>, 2019.
- 1027 Li, W., Teng, X., Chen, X., Liu, L., Xu, L., Zhang, J., Wang, Y., Zhang, Y., and Shi,  
 1028 Z.: Organic coating reduces hygroscopic growth of phase-separated aerosol particles,  
 1029 Environ. Sci. Technol., 55, 16339-16346, <https://doi.org/10.1021/acs.est.1c05901>,  
 1030 2021.
- 1031 Liu, D., Allan, J., Whitehead, J., Young, D., Flynn, M., Coe, H., McFiggans, G.,  
 1032 Fleming, Z. L., and Bandy, B.: Ambient black carbon particle hygroscopic properties  
 1033 controlled by mixing state and composition, Atmos. Chem. Phys., 13, 2015-2029,  
 1034 <https://doi.org/10.5194/acp-13-2015-2013>, 2013.
- 1035 Liu, D., Whitehead, J., Alfarra, M. R., Reyes-Villegas, E., Spracklen, Dominick V.,  
 1036 Reddington, Carly L., Kong, S., Williams, Paul I., Ting, Y.-C., Haslett, S., Taylor,  
 1037 Jonathan W., Flynn, Michael J., Morgan, William T., McFiggans, G., Coe, H., and  
 1038 Allan, James D.: Black-carbon absorption enhancement in the atmosphere determined



- 1039 by particle mixing state, *Nat. Geosci.*, 10, 184-188, <https://doi.org/10.1038/ngeo2901>,  
1040 2017a.
- 1041 Liu, D., Joshi, R., Wang, J., Yu, C., Allan, J. D., Coe, H., Flynn, M. J., Xie, C., Lee,  
1042 J., Squires, F., Kotthaus, S., Grimmond, S., Ge, X., Sun, Y., and Fu, P.: Contrasting  
1043 physical properties of black carbon in urban Beijing between winter and summer,  
1044 *Atmos. Chem. Phys.*, 19, 6749-6769, <https://doi.org/10.5194/acp-19-6749-2019>, 2019.
- 1045 Liu, S., Aiken, A. C., Gorkowski, K., Dubey, M. K., Cappa, C. D., Williams, L. R.,  
1046 Herndon, S. C., Massoli, P., Fortner, E. C., Chhabra, P. S., Brooks, W. A., Onasch, T.  
1047 B., Jayne, J. T., Worsnop, D. R., China, S., Sharma, N., Mazzoleni, C., Xu, L., Ng, N.  
1048 L., Liu, D., Allan, J. D., Lee, J. D., Fleming, Z. L., Mohr, C., Zotter, P., Szidat, S., and  
1049 Prevot, A. S. H.: Enhanced light absorption by mixed source black and brown carbon  
1050 particles in UK winter, *Nat. Commun.*, 6, 8435, <https://doi.org/10.1038/ncomms9435>,  
1051 2015.
- 1052 Liu, Z., Lu, X., Feng, J., Fan, Q., Zhang, Y., and Yang, X.: Influence of ship  
1053 emissions on urban air quality: a Comprehensive study using highly time-resolved  
1054 online measurements and numerical simulation in Shanghai, *Environ. Sci. Technol.*, 51,  
1055 202-211, <https://doi.org/10.1021/acs.est.6b03834>, 2017b.
- 1056 Malmberg, V. B., Eriksson, A. C., Shen, M., Nilsson, P., Gallo, Y., Waldheim, B.,  
1057 Martinsson, J., Andersson, O., and Pagels, J.: Evolution of In-Cylinder Diesel Engine  
1058 Soot and Emission Characteristics Investigated with Online Aerosol Mass Spectrometry,  
1059 *Environ. Sci. Technol.*, 51, 1876-1885, <https://doi.org/10.1021/acs.est.6b03391>, 2017.
- 1060 Mar, V., Fulvio, A., Andrés, A., Xavier, Q., Teresa, M., Saúl García, D. S., María  
1061 Dolores, H., and Rosalía, F. P.: Chemical tracers of particulate emissions from  
1062 commercial shipping, *Environ. Sci. Technol.*, 43, 7472-7477,  
1063 <https://doi.org/10.1021/es901558t>, 2009.
- 1064 Massoli, P., Fortner, E. C., Canagaratna, M. R., Williams, L. R., Zhang, Q., Sun,  
1065 Y., Schwab, J. J., Trimborn, A., Onasch, T. B., Demerjian, K. L., Kolb, C. E., Worsnop,  
1066 D. R., and Jayne, J. T.: Pollution Gradients and Chemical Characterization  
1067 of Particulate Matter from Vehicular Traffic near Major Roadways: Results from the  
1068 2009 Queens College Air Quality Study in NYC, *Aerosol Sci. Technol.*, 46, 1201-1218,



- 
- 1069 <https://doi.org/10.1080/02786826.2012.701784>, 2012.
- 1070 Massoli, P., Onasch, T. B., Cappa, C. D., Nuamaan, I., Hakala, J., Hayden, K., Li,  
 1071 S.-M., Sueper, D. T., Bates, T. S., Quinn, P. K., Jayne, J. T., and Worsnop, D. R.:  
 1072 Characterization of black carbon-containing particles from soot particle aerosol mass  
 1073 spectrometer measurements on the R/VAtlantis during CalNex 2010, J. Geophys. Res.:  
 1074 Atmos., 120, 2575-2593, <https://doi.org/10.1002/2014jd022834>, 2015.
- 1075 Mazzei, F., D'Alessandro, A., Lucarelli, F., Nava, S., Prati, P., Valli, G., and Vecchi,  
 1076 R.: Characterization of particulate matter sources in an urban environment, Sci. Total  
 1077 Environ., 401, 81-89, <https://doi.org/10.1016/j.scitotenv.2008.03.008>, 2008.
- 1078 Mohr, C., Huffman, J. A., Cubison, M. J., Aiken, A. C., Docherty, K. S., Kimmel,  
 1079 J. R., Ulbrich, I. M., Hannigan, M., and Jimenez, J. L.: Characterization of Primary  
 1080 Organic Aerosol Emissions from Meat Cooking, Trash Burning, and Motor Vehicles  
 1081 with High-Resolution Aerosol Mass Spectrometry and Comparison with Ambient and  
 1082 Chamber Observations, Environ. Sci. Technol., 43, 2443-2449,  
 1083 <https://doi.org/10.1021/es8011518>, 2009.
- 1084 Moldanová, J., Fridell, E., Popovicheva, O., Demirdjian, B., Tishkova, V.,  
 1085 Faccinotto, A., and Focsa, C.: Characterisation of particulate matter and gaseous  
 1086 emissions from a large ship diesel engine, Atmos. Environ., 43, 2632-2641,  
 1087 <https://doi.org/10.1016/j.atmosenv.2009.02.008>, 2009.
- 1088 Ng, N. L., Canagaratna, M. R., Jimenez, J. L., Chhabra, P. S., Seinfeld, J. H., and  
 1089 Worsnop, D. R.: Changes in organic aerosol composition with aging inferred from  
 1090 aerosol mass spectra, Atmos. Chem. Phys., 11, 6465-6474, [https://doi.org/10.5194/acp-](https://doi.org/10.5194/acp-11-6465-2011)  
 1091 [11-6465-2011](https://doi.org/10.5194/acp-11-6465-2011), 2011.
- 1092 Onasch, T. B., Trimborn, A., Fortner, E. C., Jayne, J. T., Kok, G. L., Williams, L.  
 1093 R., Davidovits, P., and Worsnop, D. R.: Soot Particle Aerosol Mass Spectrometer:  
 1094 Development, Validation, and Initial Application, Aerosol Sci. Technol., 46, 804-817,  
 1095 <https://doi.org/10.1080/02786826.2012.663948>, 2012.
- 1096 Paatero, P., and Tapper, U.: Positive matrix factorization: A non-negative factor  
 1097 model with optimal utilization of error estimates of data values, Environmetrics, 5, 111-  
 1098 126, <https://doi.org/10.1002/env.3170050203>, 1994.



- 1099 Pathak, R. K., Wang, T., and Wu, W. S.: Nighttime enhancement of PM<sub>2.5</sub> nitrate  
1100 in ammonia-poor atmospheric conditions in Beijing and Shanghai: Plausible  
1101 contributions of heterogeneous hydrolysis of N<sub>2</sub>O<sub>5</sub> and HNO<sub>3</sub> partitioning, *Atmos.*  
1102 *Environ.*, 45, 1183-1191, <https://doi.org/10.1016/j.atmosenv.2010.09.003>, 2011.
- 1103 Peng, J., Hu, M., Guo, S., Du, Z., Zheng, J., Shang, D., Levy, Z. M., Zeng, L.,  
1104 Shao, M., and Wu, Y. S.: Markedly enhanced absorption and direct radiative forcing of  
1105 black carbon under polluted urban environments, *Proc. Natl. Acad. Sci. U. S. A.*, 113,  
1106 4266-4271, <https://doi.org/10.1073/pnas.1602310113>, 2016.
- 1107 Petit, J. E., Favez, O., Albinet, A., and Canonaco, F.: A user-friendly tool for  
1108 comprehensive evaluation of the geographical origins of atmospheric pollution: Wind  
1109 and trajectory analyses, 88, 183-187, <https://doi.org/10.1016/j.envsoft.2016.11.022>,  
1110 2017.
- 1111 Petters, M. D., and Kreidenweis, S. M.: A single parameter representation of  
1112 hygroscopic growth and cloud condensation nucleus activity, *Atmos. Chem. Phys.*, 7,  
1113 1961-1971, <https://doi.org/10.5194/acp-7-1961-2007>, 2007.
- 1114 Ramanathan, V., and Carmichael, G.: Global and regional climate changes due to  
1115 black carbon, *Nat. Geosci.*, 1, 221-227, <https://doi.org/10.1038/ngeo156>, 2008.
- 1116 Rivellini, L.-H., Adam, M. G., Kasthuriarachchi, N., and Lee, A. K. Y.:  
1117 Characterization of carbonaceous aerosols in Singapore: insight from black carbon  
1118 fragments and trace metal ions detected by a soot particle aerosol mass spectrometer,  
1119 *Atmos. Chem. Phys.*, 20, 5977-5993, <https://doi.org/10.5194/acp-20-5977-2020>, 2020.
- 1120 Saarikoski, S., Timonen, H., Carbone, S., Kuuluvainen, H., Niemi, J. V., Kousa,  
1121 A., Rönkkö, T., Worsnop, D., Hillamo, R., and Pirjola, L.: Investigating the chemical  
1122 species in submicron particles emitted by city buses, *Aerosol Sci. Technol.*, 51, 317-  
1123 329, <https://doi.org/10.1080/02786826.2016.1261992>, 2016.
- 1124 Sun, Y. L., Zhang, Q., Schwab, J. J., Demerjian, K. L., Chen, W. N., Bae, M. S.,  
1125 Hung, H. M., Hogrefe, O., Frank, B., Rattigan, O. V., and Lin, Y. C.: Characterization  
1126 of the sources and processes of organic and inorganic aerosols in New York city with a  
1127 high-resolution time-of-flight aerosol mass spectrometer, *Atmos. Chem. Phys.*, 11,  
1128 1581-1602, <https://doi.org/10.5194/acp-11-1581-2011>, 2011.



- 1129 Sun, Y. L., Zhang, Q., Schwab, J. J., Yang, T., Ng, N. L., and Demerjian, K. L.:  
1130 Factor analysis of combined organic and inorganic aerosol mass spectra from high  
1131 resolution aerosol mass spectrometer measurements, *Atmos. Chem. Phys.*, 12, 8537-  
1132 8551, <https://doi.org/10.5194/acp-12-8537-2012>, 2012.
- 1133 Topping, D. O., McFiggans, G. B., and Coe, H.: A curved multi-component aerosol  
1134 hygroscopicity model framework: Part 1 – Inorganic compounds, *Atmos. Chem. Phys.*,  
1135 5, 1205-1222, <https://doi.org/10.5194/acp-5-1205-2005>, 2005a.
- 1136 Topping, D. O., McFiggans, G. B., and Coe, H.: A curved multi-component aerosol  
1137 hygroscopicity model framework: Part 2 – Including organic compounds, *Atmos. Chem.*  
1138 *Phys.*, 5, 1223-1242, <https://doi.org/10.5194/acp-5-1223-2005>, 2005b.
- 1139 Ulbrich, I. M., Canagaratna, M. R., Zhang, Q., Worsnop, D. R., and Jimenez, J. L.:  
1140 Interpretation of organic components from Positive Matrix Factorization of aerosol  
1141 mass spectrometric data, *Atmos. Chem. Phys.*, 9, 2891-2918,  
1142 <https://doi.org/10.5194/acp-9-2891-2009>, 2009.
- 1143 Ulbrich, I. M., Canagaratna, M. R., Cubison, M. J., Zhang, Q., Ng, N. L., Aiken,  
1144 A. C., and Jimenez, J. L.: Three-dimensional factorization of size-resolved organic  
1145 aerosol mass spectra from Mexico City, *Atmos. Meas. Tech.*, 5, 195-224,  
1146 <https://doi.org/10.5194/amt-5-195-2012>, 2012.
- 1147 Wang, J., Ge, X., Chen, Y., Shen, Y., Zhang, Q., Sun, Y., Xu, J., Ge, S., Yu, H., and  
1148 Chen, M.: Highly time-resolved urban aerosol characteristics during springtime in  
1149 Yangtze River Delta, China: insights from soot particle aerosol mass spectrometry,  
1150 *Atmos. Chem. Phys.*, 16, 9109-9127, <https://doi.org/10.5194/acp-16-9109-2016>, 2016.
- 1151 Wang, J., Zhang, Q., Chen, M., Collier, S., Zhou, S., Ge, X., Xu, J., Shi, J., Xie,  
1152 C., Hu, J., Ge, S., Sun, Y., and Coe, H.: First Chemical Characterization of Refractory  
1153 Black Carbon Aerosols and Associated Coatings over the Tibetan Plateau (4730 m a.s.l.),  
1154 *Environ. Sci. Technol.*, 51, 14072-14082, <https://doi.org/10.1021/acs.est.7b03973>,  
1155 2017.
- 1156 Wang, J., Liu, D., Ge, X., Wu, Y., Shen, F., Chen, M., Zhao, J., Xie, C., Wang, Q.,  
1157 Xu, W., Zhang, J., Hu, J., Allan, J., Joshi, R., Fu, P., Coe, H., and Sun, Y.:  
1158 Characterization of black carbon-containing fine particles in Beijing during wintertime,



- 
- 1159 Atmos. Chem. Phys., 19, 447-458, <https://doi.org/10.5194/acp-19-447-2019>, 2019.
- 1160 Wang, J., Ye, J., Liu, D., Wu, Y., Zhao, J., Xu, W., Xie, C., Shen, F., Zhang, J.,  
 1161 Ohno, P. E., Qin, Y., Zhao, X., Martin, S. T., Lee, A. K. Y., Fu, P., Jacob, D. J., Zhang,  
 1162 Q., Sun, Y., Chen, M., and Ge, X.: Characterization of submicron organic particles in  
 1163 Beijing during summertime: comparison between SP-AMS and HR-AMS, Atmos.  
 1164 Chem. Phys., 20, 14091-14102, <https://doi.org/10.5194/acp-20-14091-2020>, 2020a.
- 1165 Wang, J., Ye, J., Zhang, Q., Zhao, J., Wu, Y., Li, J., Liu, D., Li, W., Zhang, Y., Wu,  
 1166 C., Xie, C., Qin, Y., Lei, Y., Huang, X., Guo, J., Liu, P., Fu, P., Li, Y., Lee, H. C., Choi,  
 1167 H., Zhang, J., Liao, H., Chen, M., Sun, Y., Ge, X., Martin, S. T., and Jacob, D. J.:  
 1168 Aqueous production of secondary organic aerosol from fossil-fuel emissions in winter  
 1169 Beijing haze, Proc. Natl. Acad. Sci. U. S. A., 118, e2022179118,  
 1170 <https://doi.org/10.1073/pnas.2022179118>, 2021.
- 1171 Wang, S., Newland, M. J., Deng, W., Rickard, A. R., Hamilton, J. F., Muñoz, A.,  
 1172 Ródenas, M., Vázquez, M. M., Wang, L., and Wang, X.: Aromatic Photo-oxidation, A  
 1173 New Source of Atmospheric Acidity, Environ. Sci. Technol., 54, 7798-7806,  
 1174 <https://doi.org/10.1021/acs.est.0c00526>, 2020b.
- 1175 Willis, M. D., Healy, R. M., Riemer, N., West, M., Wang, J. M., Jeong, C.-H.,  
 1176 Wenger, J. C., Evans, G. J., Abbatt, J. P. D., and Lee, A. K. Y.: Quantification of black  
 1177 carbon mixing state from traffic: implications for aerosol optical properties, Atmos.  
 1178 Chem. Phys., 16, 4693-4706, <https://doi.org/10.5194/acp-16-4693-2016>, 2016.
- 1179 Wu, Y., Ge, X., Wang, J., Shen, Y., Ye, Z., Ge, S., Wu, Y., Yu, H., and Chen, M.:  
 1180 Responses of secondary aerosols to relative humidity and photochemical activities in  
 1181 an industrialized environment during late winter, Atmos. Environ., 193, 66-78,  
 1182 <https://doi.org/10.1016/j.atmosenv.2018.09.008>, 2018.
- 1183 Wu, Y., Liu, D., Wang, J., Shen, F., Chen, Y., Cui, S., Ge, S., Wu, Y., Chen, M.,  
 1184 and Ge, X.: Characterization of Size-Resolved Hygroscopicity of Black Carbon-  
 1185 Containing Particle in Urban Environment, Environ. Sci. Technol., 53, 14212-14221,  
 1186 <https://doi.org/10.1021/acs.est.9b05546>, 2019.
- 1187 Wu, Y., Liu, D., Wang, X., Li, S., Zhang, J., Qiu, H., Ding, S., Hu, K., Li, W., Tian,  
 1188 P., Liu, Q., Zhao, D., Ma, E., Chen, M., Xu, H., Ouyang, B., Chen, Y., Kong, S., Ge, X.,





- 1189 and Liu, H.: Ambient marine shipping emissions determined by vessel operation mode  
1190 along the East China Sea, *Sci. Total Environ.*, 769, 144713,  
1191 <https://doi.org/10.1016/j.scitotenv.2020.144713>, 2021.
- 1192 Wu, Z. J., Zheng, J., Shang, D. J., Du, Z. F., Wu, Y. S., Zeng, L. M., Wiedensohler,  
1193 A., and Hu, M.: Particle hygroscopicity and its link to chemical composition in the  
1194 urban atmosphere of Beijing, China, during summertime, *Atmos. Chem. Phys.*, 16,  
1195 1123-1138, <https://doi.org/10.5194/acp-16-1123-2016>, 2016.
- 1196 Xie, C., Xu, W., Wang, J., Liu, D., Ge, X., Zhang, Q., Wang, Q., Du, W., Zhao, J.,  
1197 Zhou, W., Li, J., Fu, P., Wang, Z., Worsnop, D., and Sun, Y.: Light absorption  
1198 enhancement of black carbon in urban Beijing in summer, *Atmos. Environ.*, 213, 499-  
1199 504, <https://doi.org/10.1016/j.atmosenv.2019.06.041>, 2019.
- 1200 Xu, J., Zhang, Q., Chen, M., Ge, X., Ren, J., and Qin, D.: Chemical composition,  
1201 sources, and processes of urban aerosols during summertime in northwest China:  
1202 insights from high-resolution aerosol mass spectrometry, *Atmos. Chem. Phys.*, 14,  
1203 12593-12611, <https://doi.org/10.5194/acp-14-12593-2014>, 2014.
- 1204 Xu, W., Han, T., Du, W., Wang, Q., Chen, C., Zhao, J., Zhang, Y., Li, J., Fu, P.,  
1205 Wang, Z., Worsnop, D. R., and Sun, Y.: Effects of Aqueous-Phase and Photochemical  
1206 Processing on Secondary Organic Aerosol Formation and Evolution in Beijing, China,  
1207 *Environ. Sci. Technol.*, 51, 762-770, <https://doi.org/10.1021/acs.est.6b04498>, 2017.
- 1208 Xu, W., Sun, Y., Wang, Q., Zhao, J., Wang, J., Ge, X., Xie, C., Zhou, W., Du, W.,  
1209 Li, J., Fu, P., Wang, Z., Worsnop, D. R., and Coe, H.: Changes in Aerosol Chemistry  
1210 From 2014 to 2016 in Winter in Beijing: Insights From High-Resolution Aerosol Mass  
1211 Spectrometry, *J. Geophys. Res.: Atmos.*, 124, 1132-1147,  
1212 <https://doi.org/10.1029/2018jd029245>, 2019.
- 1213 Ye, Z., Liu, J., Gu, A., Feng, F., Liu, Y., Bi, C., Xu, J., Li, L., Chen, H., Chen, Y.,  
1214 Dai, L., Zhou, Q., and Ge, X.: Chemical characterization of fine particulate matter in  
1215 Changzhou, China, and source apportionment with offline aerosol mass spectrometry,  
1216 *Atmos. Chem. Phys.*, 17, 2573-2592, <https://doi.org/10.5194/acp-17-2573-2017>, 2017.
- 1217 Yu, G., Zhang, Y., Yang, F., He, B., Zhang, C., Zou, Z., Yang, X., Li, N., and Chen,  
1218 J.: Dynamic Ni/V Ratio in the Ship-Emitted Particles Driven by Multiphase Fuel Oil



- 
- 1219 Regulations in Coastal China, *Environ. Sci. Technol.*, 55, 15031-15039,  
 1220 <https://doi.org/10.1021/acs.est.1c02612>, 2021.
- 1221 Yu, L., Smith, J., Laskin, A., Anastasio, C., Laskin, J., and Zhang, Q.: Chemical  
 1222 characterization of SOA formed from aqueous-phase reactions of phenols with the  
 1223 triplet excited state of carbonyl and hydroxyl radical, *Atmos. Chem. Phys.*, 14, 13801-  
 1224 13816, <https://doi.org/10.5194/acp-14-13801-2014>, 2014.
- 1225 Zhang, F., Wang, Y., Peng, J., Chen, L., Sun, Y., Duan, L., Ge, X., Li, Y., Zhao, J.,  
 1226 Liu, C., Zhang, X., Zhang, G., Pan, Y., Wang, Y., Zhang Annie, L., Ji, Y., Wang, G., Hu,  
 1227 M., Molina Mario, J., and Zhang, R.: An unexpected catalyst dominates formation and  
 1228 radiative forcing of regional haze, *Proc. Natl. Acad. Sci. U. S. A.*, 117, 3960-3966,  
 1229 <https://doi.org/10.1073/pnas.1919343117>, 2020.
- 1230 Zhang, G., Fu, Y., Peng, X., Sun, W., Shi, Z., Song, W., Hu, W., Chen, D., Lian,  
 1231 X., Li, L., Tang, M., Wang, X., and Bi, X.: Black Carbon Involved Photochemistry  
 1232 Enhances the Formation of Sulfate in the Ambient Atmosphere: Evidence From In Situ  
 1233 Individual Particle Investigation, *J. Geophys. Res.: Atmos.*, 126, e2021JD035226,  
 1234 <https://doi.org/10.1029/2021JD035226>, 2021.
- 1235 Zhang, Q., Canagaratna, M. R., Jayne, J. T., Worsnop, D. R., and Jimenez, J.-L.:  
 1236 Time- and size-resolved chemical composition of submicron particles in Pittsburgh:  
 1237 Implications for aerosol sources and processes, *J. Geophys. Res.*, 110,  
 1238 <https://doi.org/10.1029/2004JD004649>, 2005.
- 1239 Zhang, Q., Jimenez, J. L., Canagaratna, M. R., Ulbrich, I. M., Ng, N. L., Worsnop,  
 1240 D. R., and Sun, Y.: Understanding atmospheric organic aerosols via factor analysis of  
 1241 aerosol mass spectrometry: a review, *Anal. Bioanal. Chem.*, 401, 3045-3067,  
 1242 <https://doi.org/10.1007/s00216-011-5355-y>, 2011.
- 1243 Zhao, J., Qiu, Y., Zhou, W., Xu, W., Wang, J., Zhang, Y., Li, L., Xie, C., Wang, Q.,  
 1244 Du, W., Worsnop, D. R., Canagaratna, M. R., Zhou, L., Ge, X., Fu, P., Li, J., Wang, Z.,  
 1245 Donahue, N. M., and Sun, Y.: Organic Aerosol Processing During Winter Severe Haze  
 1246 Episodes in Beijing, *J. Geophys. Res.: Atmos.*, 124, 10248-10263,  
 1247 <https://doi.org/10.1029/2019jd030832>, 2019.
- 1248 Zhao, M., Zhang, Y., Ma, W., Fu, Q., Yang, X., Li, C., Zhou, B., Yu, Q., and Chen,



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1249 L.: Characteristics and ship traffic source identification of air pollutants in China's  
1250 largest port, *Atmos. Environ.*, **64**, 277-286,  
1251 <https://doi.org/10.1016/j.atmosenv.2012.10.007>, 2013.

1252 Zhou, S., Collier, S., Xu, J., Mei, F., Wang, J., Lee, Y.-N., Sedlacek, A. J.,  
1253 Springston, S. R., Sun, Y., and Zhang, Q.: Influences of upwind emission sources and  
1254 atmospheric processing on aerosol chemistry and properties at a rural location in the  
1255 Northeastern U.S, *J. Geophys. Res.: Atmos.*, **121**, 6049-6065,  
1256 <https://doi.org/10.1002/2015jd024568>, 2016.

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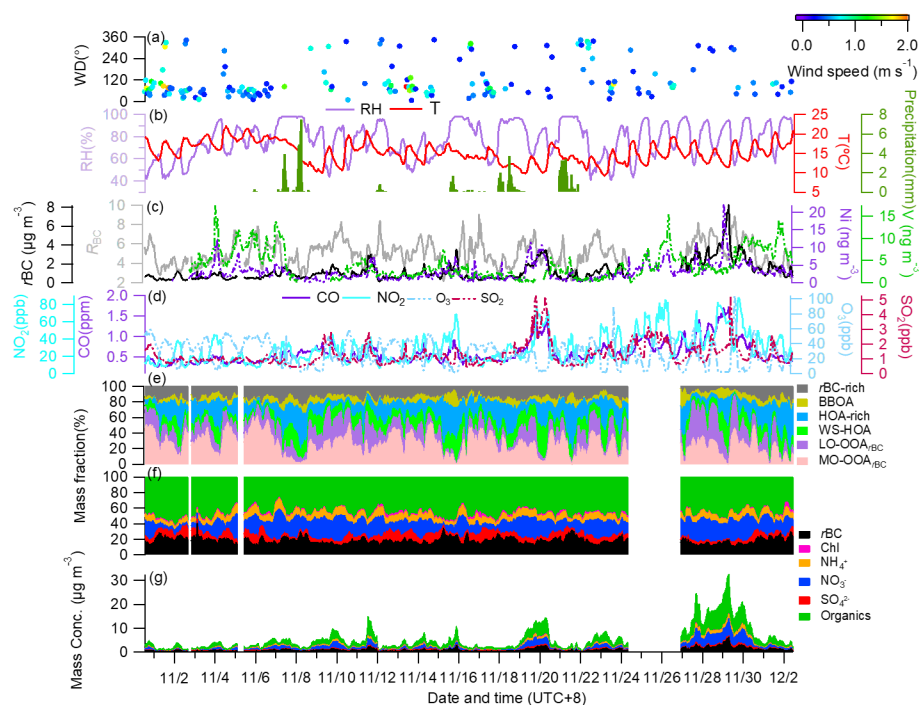


Figure 1. Time series of (a) wind direction (WD) colored by wind speed (WS), (b) air temperature (T), relative humidity (RH) and precipitation, (c) mass concentrations of *r*BC, Ni, V, and *R*<sub>BC</sub> (mass ratio of all coating species to *r*BC), (d) mass concentrations of gas pollutants of CO, NO<sub>2</sub>, O<sub>3</sub> and SO<sub>2</sub>, (e) mass fractions (%) of different OA factors to the total *r*BCc OA, (f) mass fractions (%) of different components to the total *r*BCc mass, and (g) mass concentrations of stacked *r*BCc components.

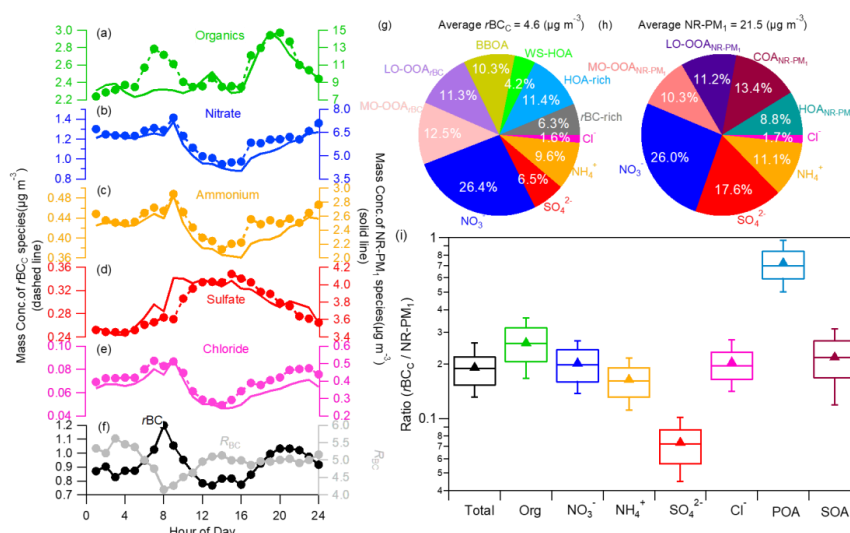
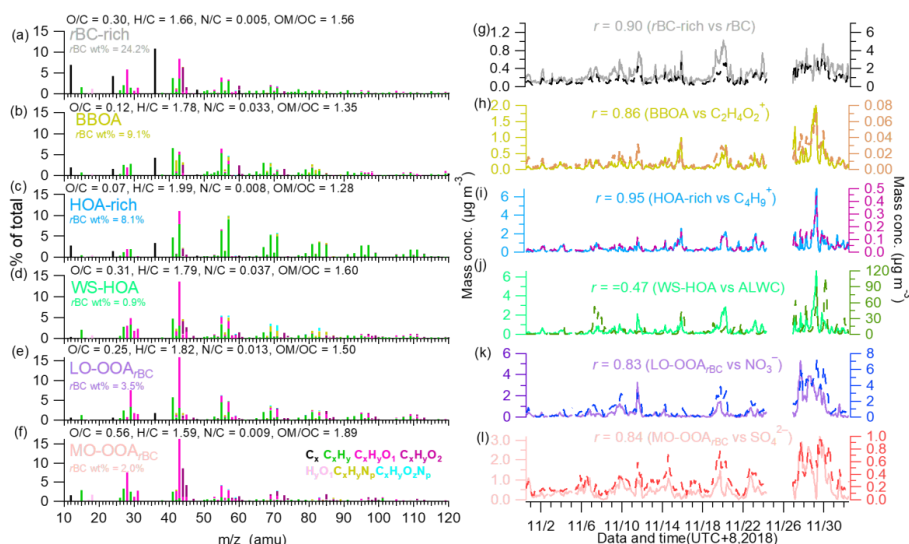


Figure 2. Diurnal cycles of mass concentrations of (a–e)  $rBC_c$  and NR-PM<sub>1</sub> species (organics, nitrate, ammonium, sulfate, and chloride), and (f)  $rBC$  and  $R_{BC}$ . Campaign-average chemical composition of  $rBC_c$  (g) and NR-PM<sub>1</sub> (h). (i) Mass ratios of species in  $rBC_c$  to those in NR-PM<sub>1</sub> (the whiskers above and below the boxes mark the 90% and 10% percentiles, respectively; the upper and lower edge of the boxes represent the 75% and 25% percentiles, respectively; and the lines and triangles inside the boxes denote the median and mean values, respectively; SOA represents  $([LO-OOA_{rBC}] + [MO-OOA_{rBC}]) / ([LO-OOA_{NR-PM1}] + [LV-OOA_{NR-PM1}])$ , and POA represents  $([rBC-rich] + [HOA-rich] + [BBOA] + [WS-HOA]) / [HOA_{NR-PM1}]$ ).

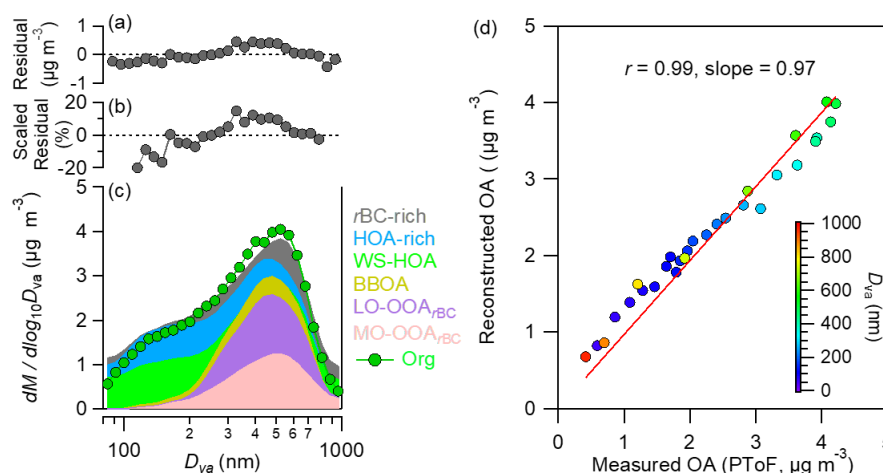


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1276 Figure 3. High resolution mass spectra of (a) *r*BC-rich, (b) BBOA, (c) HOA-rich, (d)  
 1277 WS-HOA, (e) LO-OOA<sub>rBC</sub>, and (f) MO-OOA<sub>rBC</sub>. (g-l) Time series of corresponding  
 1278 factors, their tracers (*r*BC, C<sub>2</sub>H<sub>4</sub>O<sub>2</sub><sup>+</sup>, C<sub>4</sub>H<sub>9</sub><sup>+</sup>, ALWC, nitrate and sulfate) as well as the  
 1279 correlation coefficients (ALWC refers to aerosol liquid water content, which was  
 1280 estimated by using the extended aerosol inorganic model (Clegg et al., 1998).  
 1281 Calculated ALWC at different RH values is shown in Fig. S4)  
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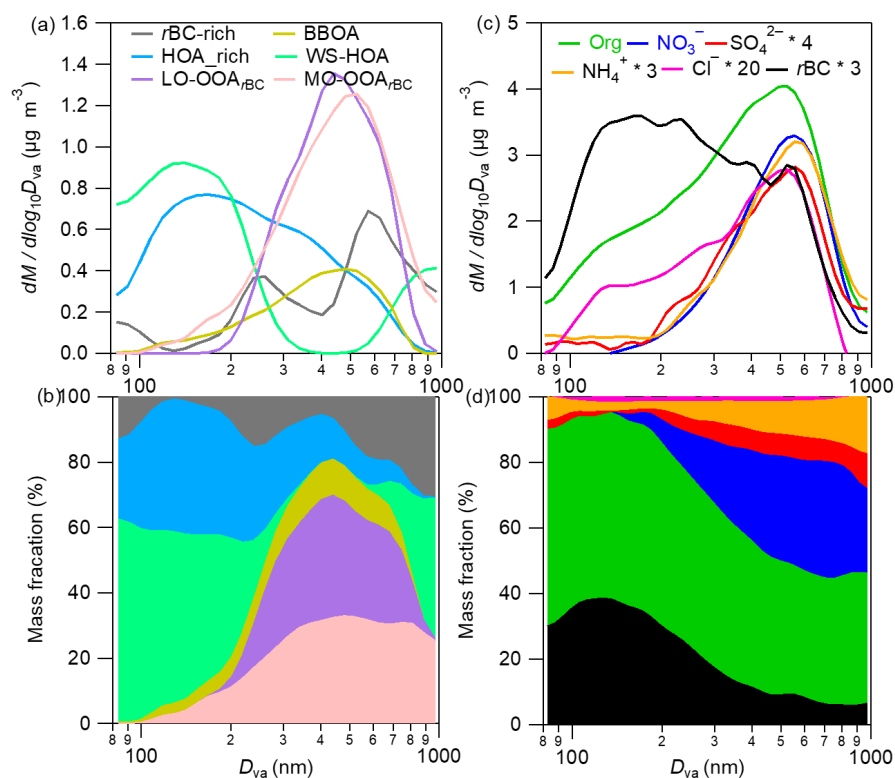
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1285 Figure 4. Summary of key diagnostic plots of derivation of size distributions of  
 1286 individual *rBCc* OA factors. (a) Absolute and (b) relative residuals between the  
 1287 reconstructed and measured OA mass concentrations in different size bins. (c) Stacked  
 1288 size distributions of the six OA factors compared to the size distributions of total OA.  
 1289 (d) Reconstructed OA mass concentrations compared to the measured values for  
 1290 different size bins (80-1000 nm).

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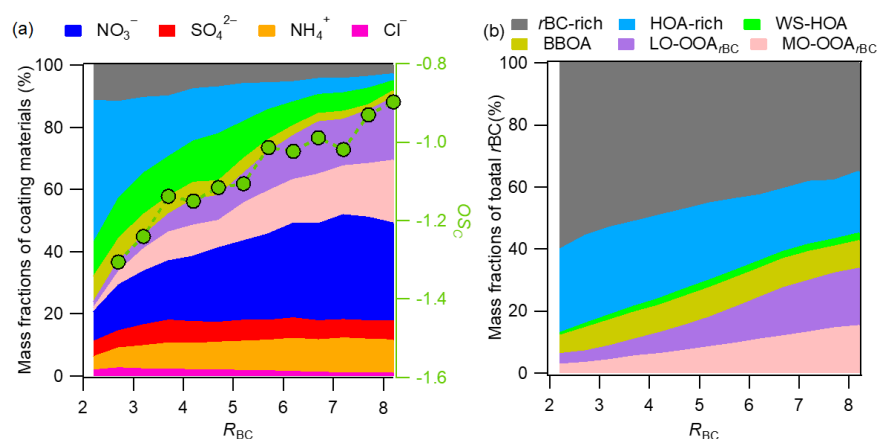


1292  
 1293 Figure 5. Campaign-average size distributions of six *r*BCc OA factors (a) and individual  
 1294 *r*BCc components (b), and corresponding mass contributions of the six factors to the  
 1295 total *r*BCc OA (c), and the major components to the total *r*BCc (d) at different sizes  
 1296 (80-1000 nm).  
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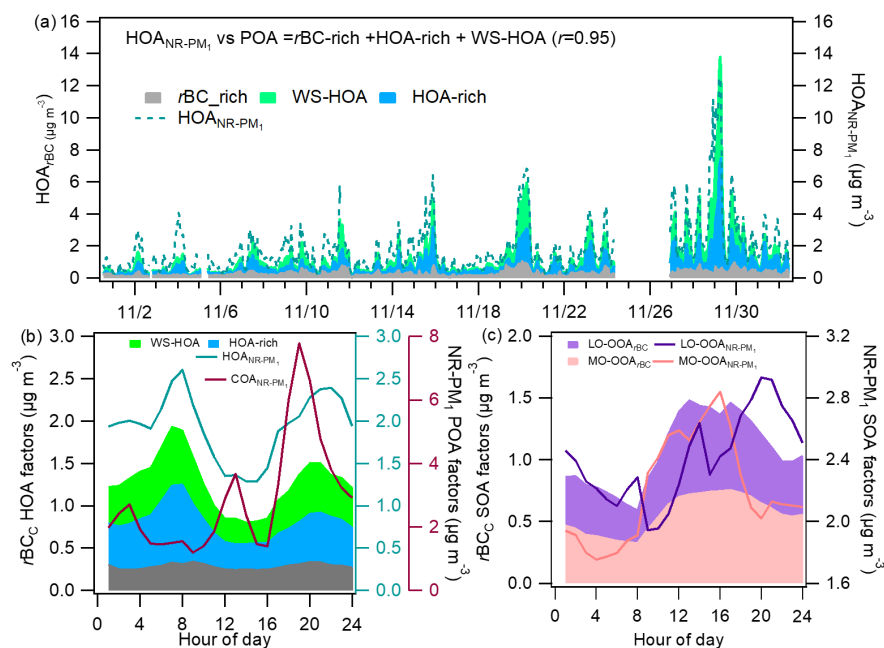


1299

1300 Figure 6. (a) Variations of mass fractions of the major  $rBCc$  components against  $R_{BC}$ .

1301 (b) Variations of mass contributions of individual  $rBCc$  OA factors to  $rBC$  against  $R_{BC}$ .

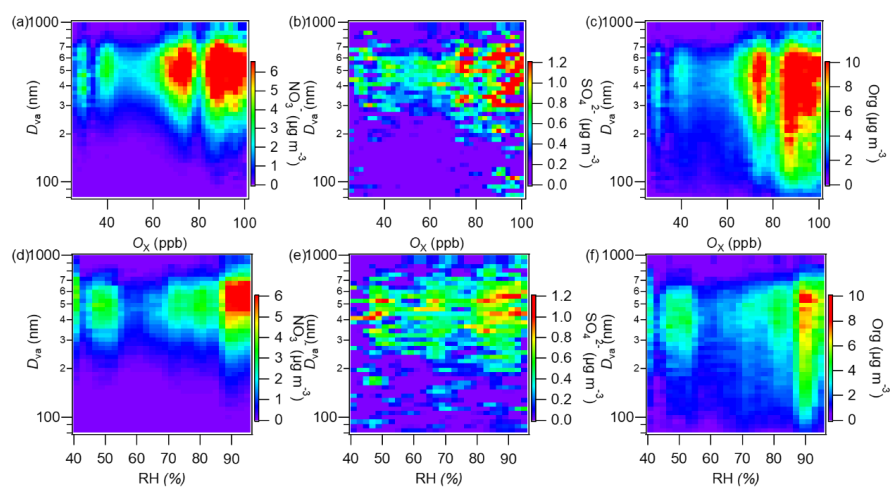
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 1304 Figure 7. (a) Time series of stacked three  $r\text{BC}_c$  POA factors (i.e.,  $r\text{BC-rich}$ , HOA-rich,  
 1305 and WS-HOA) and  $\text{HOA}_{\text{NR-PM}_1}$ . Comparisons of the diurnal patterns of different POA  
 1306 factors (b) and SOA factors (c) of  $r\text{BC}_c$  and NR-PM<sub>1</sub>.  
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1308



1309

1310 Figure 8. Image plots of size distributions of *r*BCc nitrate, sulfate, organics as a function  
 1311 of (a-c)  $O_X$  and (d-f) RH, respectively (color represents its concentration).  
 1312

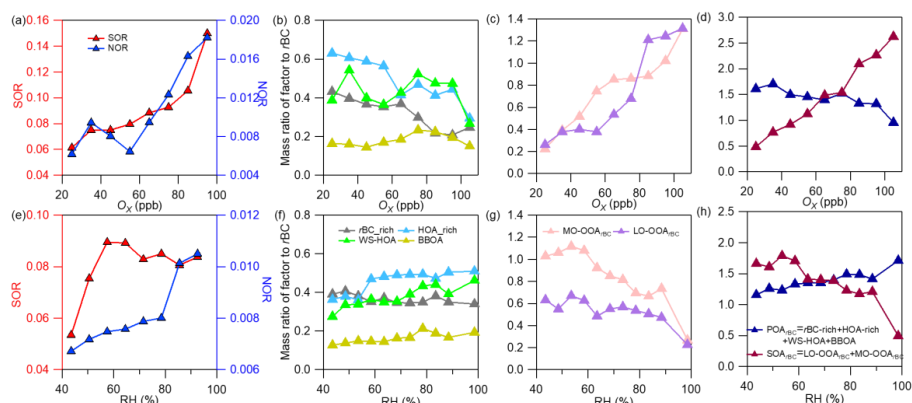
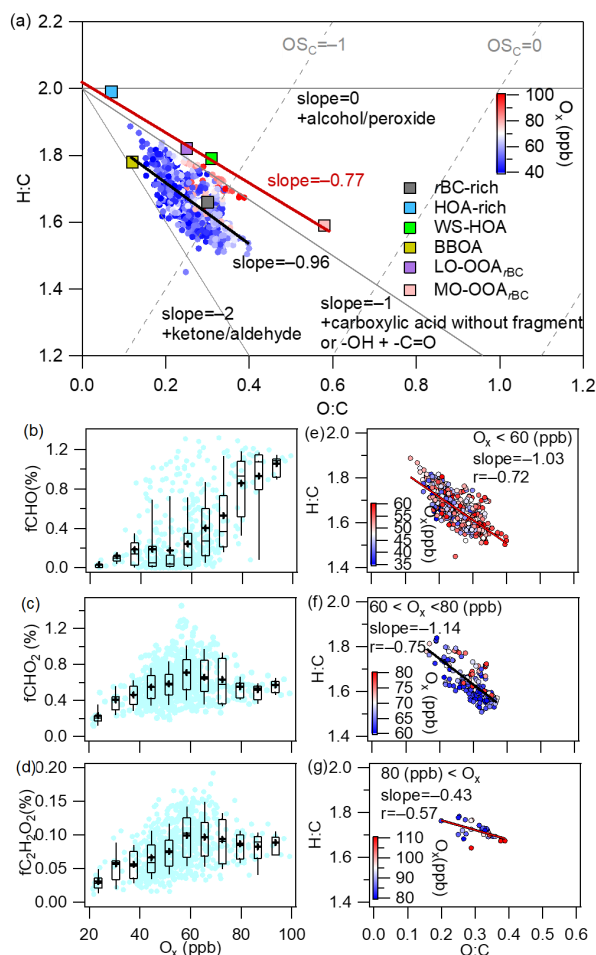
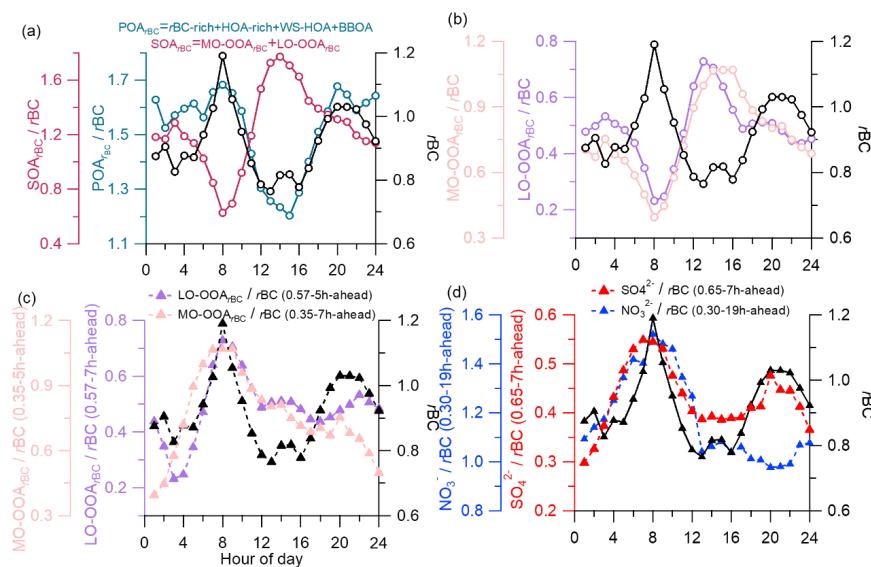


Figure 9. Variations of nitrogen oxidation ratio (NOR) and sulfur oxidation ratio, mass ratios of different POA factors, SOA factors and total POA and SOA to  $rBC$  against  $O_X$  (a-d) and RH (e-h) ( $NOR = nNO_3^- / (nNO_3^- + nNO_2 + nNO)$  and  $SOR = nSO_4^{2-} / (nSO_4^{2-} + nSO_2)$ , where  $nNO_3^-$ ,  $nSO_4^{2-}$ ,  $nNO_2$ ,  $nNO$  and  $nSO_2$  are the molar concentrations of particle-phase sulfate, nitrate, gaseous  $NO_2$ ,  $NO$  and  $SO_2$ , respectively).



1320

1321 Figure 10. (a) Van Krevelen diagram of H/C versus O/C ratios for all rBCC OA and the  
 1322 six factors colored by  $O_x$  concentrations (the black line represents the linearly fitted  
 1323 line of all OA data, and the red line is the fitted line of the four OA factors). (b-d) Mass  
 1324 fractions of selected oxygenated ion fragments as a function of  $O_x$  (meanings of the  
 1325 boxes are the same as those described in Fig. 2). (e-g) Scatter plots of H/C versus O/C  
 1326 ratios under different  $O_x$  levels (data are colored by  $O_x$  concentrations).



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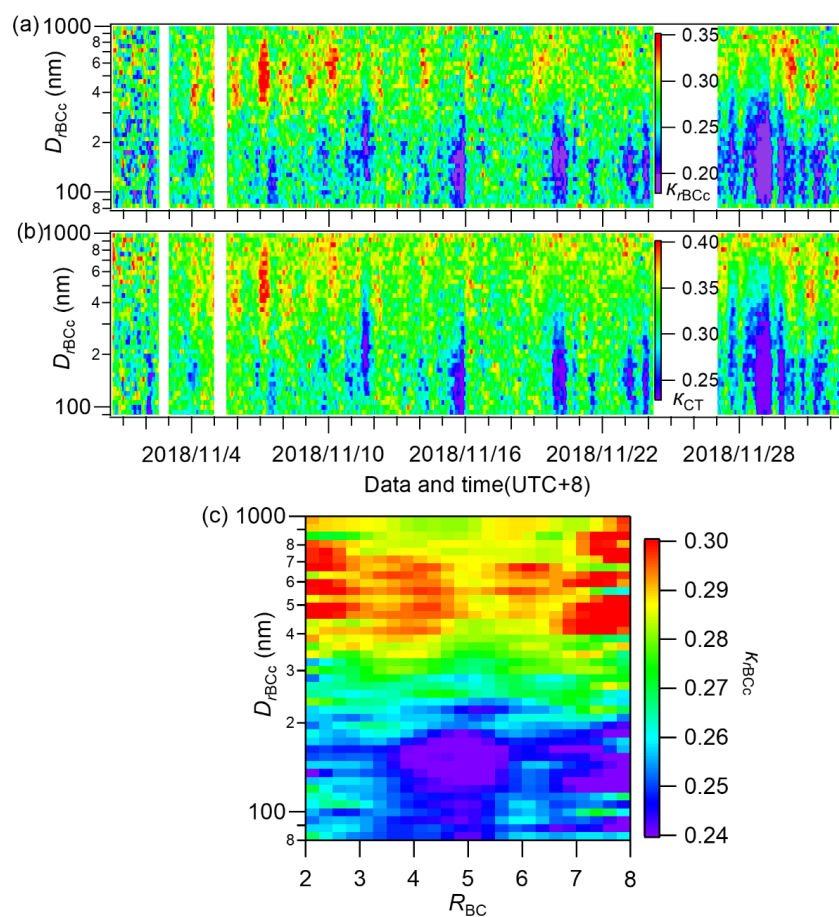
1328 Figure 11. Campaign-average diurnal patterns of (a)  $rBC$ ,  $POA_{rBC}/rBC$  and

1329  $SOA_{rBC}/rBC$ , and (b)  $rBC$ ,  $MO-OOA_{rBC}/rBC$  and  $LO-OOA_{rBC}/rBC$ . Adjusted diurnal

1330 patterns by the average coating time (ACT) for (c)  $LO-OOA_{rBC}/rBC$ ,  $MO-OOA_{rBC}/rBC$ ,

1331 and (d)  $SO_4^{2-}/rBC$ ,  $NO_3^{2-}/rBC$ .

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1334 Figure 12. Image plots of size-resolved hygroscopicity parameters of (a)  $rBCc$  ( $\kappa_{rBCc}$ ),  
 1335 its coating materials ( $\kappa_{CT}$ ) during the whole campaign, and (c) the campaign-average  
 1336 size-resolved  $\kappa_{rBCc}$  at different  $R_{BC}$ .

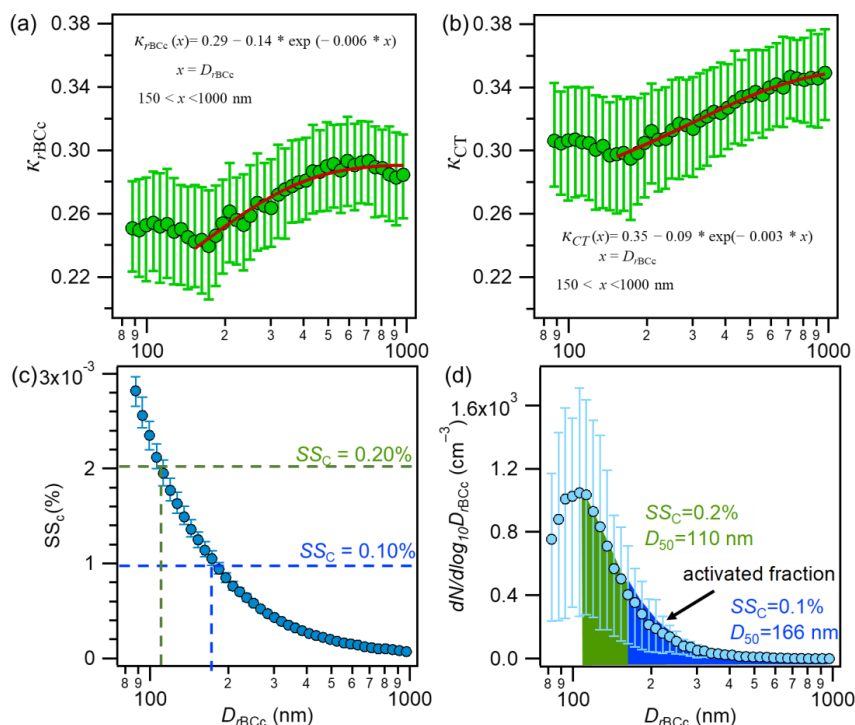
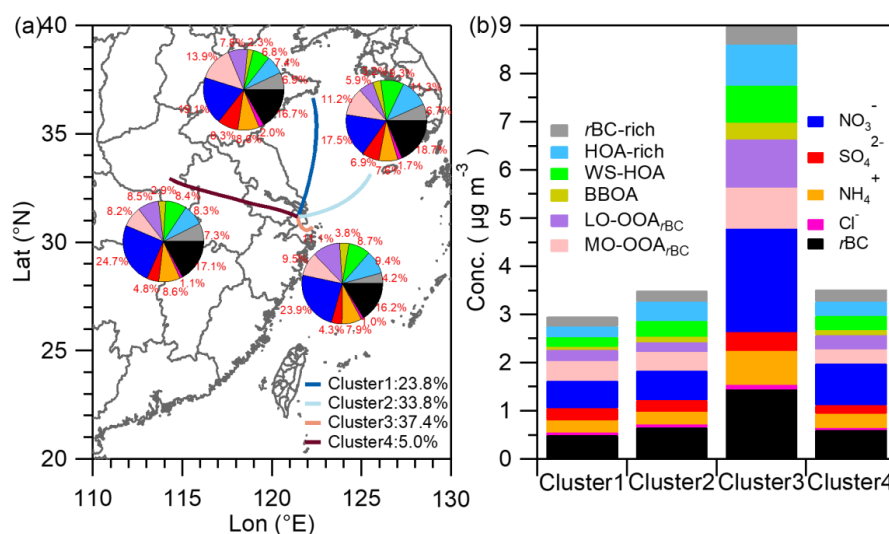


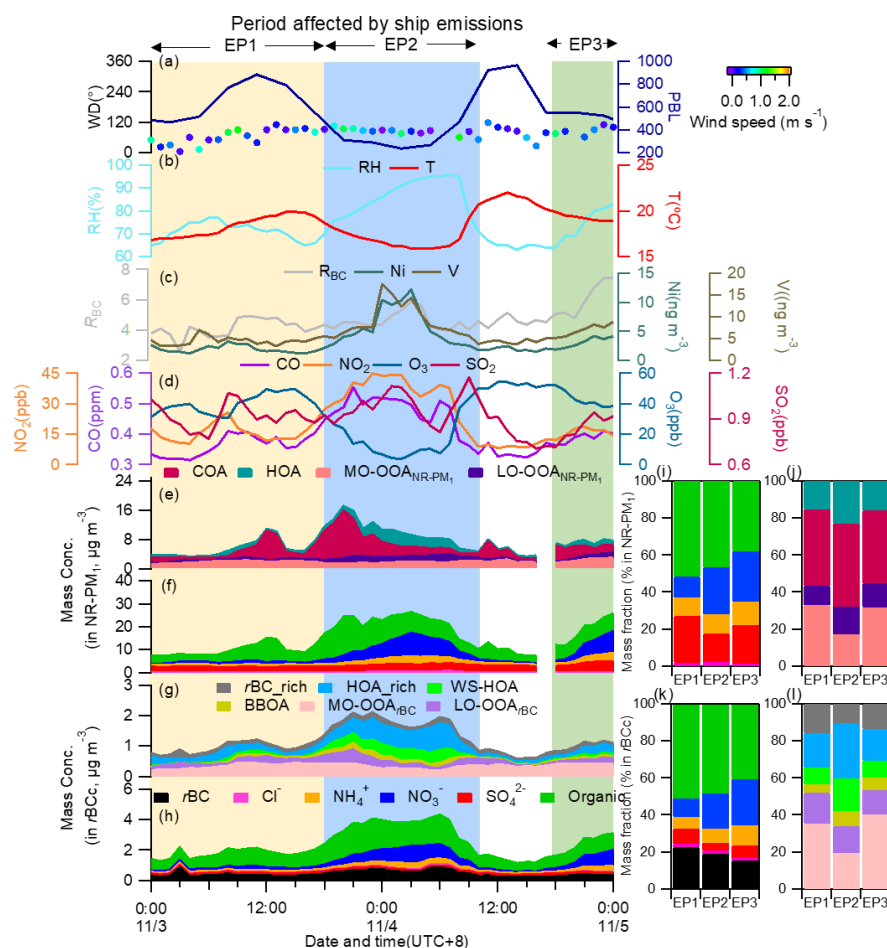
Figure 13. Campaign-average size-resolved hygroscopic parameters for  $rBCc$  ( $\kappa_{rBCc}$ ) and (a) for its coatings ( $\kappa_{CT}$ ) (b) (the red lines are exponential fits of the data of 150-1000 nm). (c) Campaign-average size-resolved critical supersaturation ( $SS_c$ ), and (d) the predicted activated fraction of  $rBCc$  number concentration based on  $D_{50}$  at  $SS_c$  of 0.1% (166 nm) and 0.2% (110 nm) (the solid circles are mean values, the upper and lower lines are the 75<sup>th</sup> and 25<sup>th</sup> percentiles, respectively).





1344

1345 Figure 14. (a) Four clusters of 24-h backward trajectories (at altitude of 500 m) analyzed  
 1346 by NOAA HYSPLIT model (<http://www.arl.noaa.gov/ready/hysplit4.html>) embedded  
 1347 in Zefir(Petit et al., 2017), with the pie chart showing the average *r*BCc chemical  
 1348 compositions in each cluster. (b) Stacked mass concentrations of the *r*BCc components  
 1349 of the four clusters.



1350  
 1351 Figure 15. Time series of (a) wind direction (WD) colored by wind speed (WS),  
 1352 planetary boundary layer (PBL) height, (b) relative humidity (RH) and temperature (T),  
 1353 (c) mass concentrations of particle-phase Ni and V, and  $R_{BC}$ , (d) mass concentrations of  
 1354 CO, NO<sub>2</sub>, O<sub>3</sub>, SO<sub>2</sub>, stacked concentrations of (e) NR-PM<sub>1</sub> OA factors, (f) NR-PM<sub>1</sub>  
 1355 species, (g)  $rBCc$  OA factors, and (h)  $rBCc$  components during the ship emission period  
 1356 (SEP). Mass contributions of different components to NR-PM<sub>1</sub> (i), different OA factors  
 1357 to total NR-PM<sub>1</sub> OA (j), different components to  $rBCc$  (k), and different OA factors to  
 1358 total  $rBCc$  OA (l) for the three episodes (EP1, EP2 and EP3).  
 1359