
1 **Chemical properties, sources and size-resolved**
2 **hygroscopicity of submicron black carbon-containing**
3 **aerosols in urban Shanghai**

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18 *For Atmospheric Chemistry and Physics*

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20 **Abstract.** Refractory black carbon (*r*BC) aerosols play an important role in air quality
21 and climate change, yet high time-resolved and detailed investigation on the
22 physicochemical properties of *r*BC and its associated coating is still scarce. In this work,
23 we used a laser-only Aerodyne soot particle aerosol mass spectrometer (SP-AMS) to
24 exclusively measure the *r*BC-containing (*r*BCc) particles, and compared their
25 properties with the total non-refractory submicron particles (NR-PM₁) measured in
26 parallel by a high-resolution AMS (HR-AMS) in Shanghai. The observation shows that
27 *r*BC was overall thickly coated with an average mass ratio of coating to *r*BC core (R_{BC})
28 of $\sim 5.0 (\pm 1.7)$. However, mass of *r*BC coating species only occupied $19.1 (\pm 4.9)\%$ of
29 those in NR-PM₁; sulfate tended to condense preferentially on non-*r*BC particles
30 therefore its portion on *r*BC was only $7.4 (\pm 2.2)\%$, while the majority of primary
31 organic aerosols (POA) were associated with *r*BC ($72.7 \pm 21.0\%$). Positive matrix
32 factorization reveals that cooking emitted organics did not coat on *r*BC, and a portion
33 of organics coated on *r*BC was from biomass burning which was unidentifiable in NR-
34 PM₁. Small *r*BCc particles were predominantly from traffic, while large-sized ones
35 were often mixed with secondary components and typically had thick coating. Sulfate
36 and secondary organic aerosol (SOA) species were generated mainly through daytime
37 photochemical oxidation (SOA formation likely involved with in-situ chemical
38 conversion of traffic-related POA to SOA), while nocturnal heterogeneous formation
39 was dominant for nitrate; we also estimated the average times of 5~19 hours for those
40 secondary species to coat on *r*BC. Particles during a short period that was affected by
41 ship emissions, were characterized with a high vanadium concentration (on average 6.3
42 $\pm 3.1 \text{ ng m}^{-3}$) and a mean vanadium/nickel mass ratio of $2.0 (\pm 0.6)$. Furthermore, the
43 size-resolved hygroscopicity parameter (κ_{rBCc}) of *r*BCc particles was obtained based on
44 its fully chemical characterization, and was parameterized as $\kappa_{rBCc}(x) = 0.29 - 0.14 \times$
45 $\exp(-0.006 \times x)$ (x is from 150 to 1000 nm). Under critical supersaturations (SS_C) of
46 0.1% and 0.2%, the D_{50} values were $166 (\pm 16)$ and $110 (\pm 5)$ nm, respectively, and
47 with $16 (\pm 3)\%$ and $59 (\pm 4)\%$ of *r*BCc in number could be activated into cloud
48 condensation nuclei (CCN). Our findings are valuable to advance the understanding of
49 BC chemistry as well as the effective control of atmospheric BC pollution.

50 **1 Introduction**

51 Refractory black carbon (*rBC*) aerosols can directly absorb solar radiation,
52 indirectly change the nature of cloud and alter the albedo of snow or glaciers (Jacobi et
53 al., 2015), resulting in a positive radiative forcing that is second only to carbon dioxide
54 on both regional and global scales (Ramanathan and Carmichael, 2008; Bond et al.,
55 2013). The fresh *rBC* particles produced by incomplete combustion of biomass and
56 fossil fuel tend to be fractal in morphology and can mix with many other components
57 (Peng et al., 2016; Li et al., 2021). After entering into the atmosphere, fresh *rBC* can
58 further externally or internally mix with organic/inorganic species which are primarily
59 emitted or secondarily formed, and such aged *rBC*-containing (*rBCc*) particles (Chen
60 et al., 2017; Lee et al., 2017) might have contrasting chemical properties and
61 morphologies (or mixing states) (Liu et al., 2017a; Lee et al., 2019; Xie et al., 2019). In
62 addition, when *rBC* mixes with hydrophilic materials, its hygroscopicity, cloud
63 condensation nuclei (CCN) activity and size distribution, etc., can be significantly
64 changed, which subsequently affect its atmospheric behavior, impact and lifecycle (Liu
65 et al., 2013; Lambe et al., 2015). Therefore, it is necessary to elucidate the
66 physicochemical characteristics and sources of *rBC* cores and associated coating
67 materials, so as to better understand their influences on climate and air quality.

68 Chemical composition of ambient *rBCc* particles is largely dependent upon
69 atmospheric conditions and emission sources. In general, the thickness of coating, mass
70 contribution of secondary components (such as sulfate, nitrate and secondary organic
71 aerosol (SOA) species) and oxidation degree of the coated organics of *rBCc* particles,
72 increase with the aging time or oxidation capacity of ambient environment (Cappa et
73 al., 2012; Liu et al., 2015; Wang et al., 2017; Collier et al., 2018; Wang et al., 2019),
74 except in some specific cases that thickly coated *rBCc* might be dominated by primarily
75 emitted particles (such as from biomass burning (Wang et al., 2017)). Recent field
76 observations report that SOA species coated on *rBC* cores could account for 35% and
77 41% of the total SOA mass near traffic emission sources and in a polluted offshore
78 environment, respectively (Massoli et al., 2012; Massoli et al., 2015). A study of *rBCc*
79 particles in Singapore finds that over 90% of *rBC* derived from local combustion

80 sources (mainly traffic), while 30% of *r*BC was associated with fresh SOA generated
81 under the influences of daytime shipping and industrial emissions (Rivellini et al.,
82 2020). The SOA material concentrated on the surface of *r*BC was found to be chemically
83 different from the SOA that was externally mixed with *r*BC (Lee et al., 2017) in Fontana,
84 California, and another study in Shenzhen, China, reveals that more oxidized SOA
85 preferred to mix with *r*BC due to that abundant transition metals detected on *r*BC cores
86 might act as catalysts to convert less oxidized SOA to more oxidized SOA in aerosol
87 aqueous phase (Cao et al., 2022); the *r*BC could catalyze SO₂ to form sulfate as well,
88 as observed in Beijing (Zhang et al., 2020) and Guangzhou (Zhang et al., 2021), China.
89 Besides SOA, cooking-related OA is found to be externally mixed with *r*BC (Lee et al.,
90 2017; Wang et al., 2019), and a unique biomass burning related OA factor was identified
91 and was only present in *r*BCc rather than non-*r*BC particles during summertime in
92 Beijing (Wang et al., 2020a).

93 Size distribution of *r*BCc particles is also modulated greatly by their original
94 sources and ageing processes. For example, a study in Shanghai shows a bimodal size
95 distribution of *r*BCc, with a condensation mode dominated by traffic emissions (small
96 core size, thin coating) and a droplet mode including highly aged biomass burning
97 particles (large core size and thick coating) and highly aged traffic particles (small core
98 size and very thick coating) (Gong et al., 2016). Another study in Beijing (Liu et al.,
99 2019) further resolves four size modes of *r*BCc, relevant with traffic (small core, thin
100 coating), coal or biomass burning (moderate coating, both small and large cores), coal
101 combustion (large core, thick coating) and secondary process (thick coating, both small
102 and large cores).

103 Moreover, water uptake and CCN activity of *r*BCc particles can increase with the
104 encapsulation of water-soluble substances such as sulfate, nitrate, and SOA (Liu et al.,
105 2013; Wu et al., 2019). Based on the measured chemical composition of *r*BCc, our
106 previous work has established a method for calculating size-resolved hygroscopicity
107 parameters of *r*BCc (κ_{rBCc}), and determined the CCN activation diameters of *r*BCc
108 particles for given critical supersaturation (SS_C) values (Wu et al., 2019).

109 Highly time-resolved chemical characterization of *r*BCc particles were seldom

110 reported in China and is still lacking in Shanghai. In this study, we utilized an Aerodyne
111 soot particle aerosol mass spectrometer (SP-AMS) to determine the concentration,
112 composition and size distribution of *rBCc* particles exclusively (technical details in
113 Section 2.1) in urban Shanghai for the first time. We also compared the SP-AMS
114 measurement results with those from a co-located Aerodyne high-resolution time-of-
115 flight aerosol mass spectrometer (HR-AMS), to comprehensively investigate the
116 characteristics of *rBCc* particles. We analyzed κ_{rBCc} and estimated the proportions of
117 activated *rBCc* numbers at given SS_C as well.

118

119 **2 Experimental methods**

120 **2.1 Sampling site and instrumentation**

121 The field measurement was conducted from October 31 to December 2, 2018,
122 during which the instruments were deployed on 8th floor of the building of Shanghai
123 Academy of Environmental Sciences (SAES) (31°10'33.348" N, 121°26'10.978" E).
124 Shanghai can be a representative of the densely populated megacity across the world;
125 the measurement period also belonged to the cold season when sources of *rBCc* might
126 be complex, and can offer rich information about the *rBCc*; of course, future
127 measurements in other seasons are still essential to achieve a complete understanding
128 of *rBCc*. East of the sampling site is a large commercial shopping center, and the site
129 is surrounded by residential areas with two busy arterial roads directly to the east (~450
130 m) and south (~150 m), respectively (Figure S1 in the supporting information). In
131 addition, the adjacent areas are densely populated with roadside residents, office
132 workers, and market traders, as well as crowds in and out of the Caobao Road Metro
133 station (~100 m). The measurement period was dominated by northeasterly winds,
134 while many international freight companies located on northeastern side of the site, and
135 many freighters were reposing on the Huangpu River. Overall, the sampling site was
136 probably influenced by vehicular emissions, residential activities and the northeast
137 Cargo ship emission plumes, etc.

138 An Aerodyne SP-AMS and an HR-AMS were operated in parallel during the
139 campaign. The two AMSs shared a same sampling line with a PM_{2.5} cyclone (Model

140 URG-2000-30EN) in front to remove coarse particles. Ambient air pulled through the
141 sampling line was dried using a diffusion dryer filled with silicon gel and was
142 subsequently drawn into both instruments. Due to the transmission efficiency of the
143 inlet lens, both AMSs measured mainly particles of 30-1200 nm (denoted as PM_{10}).

144 The working principle of SP-AMS has been described in detail previously (Onasch
145 et al., 2012). However, in this work, we used only the intracavity infrared laser
146 vaporizer to selectively measure *r*BCc particles (*r*BC cores and associated coating
147 materials), as *r*BC can absorb 1064 nm laser light. The thermal tungsten vaporizer had
148 to be physically detached otherwise non-*r*BC particles can still be detected as the
149 filament can heat the vaporizer to ~ 200 °C even if it was turned off. Before sampling,
150 the SP-AMS was tuned and calibrated following the steps described previously (Lee et
151 al., 2015; Willis et al., 2016; Wang et al., 2017). During sampling, due to relatively low
152 *r*BCc mass loadings, the SP-AMS was operated with two mass sensitive V modes (2.5
153 minutes per cycle), one with a particle time-of-flight (PToF) mode (30 s) and another
154 one (120 s) with a mass spectral mode with mass-to-charge (m/z) ratio up to 500.
155 Filtered air was also measured in the middle of campaign (for 60 min) to determine the
156 limits of detection (LOD, three times the standard deviation) of various aerosol species
157 and to adjust the air-influenced mass spectral signals (Zhang et al., 2005).

158 Before removal of the tungsten heater, the calibrations of ionization efficiency (IE)
159 for nitrate and relative ionization efficiency (RIE) of sulfate were performed by using
160 pure ammonium nitrate and ammonium sulfate particles (Jayne et al., 2000), and the
161 values were assumed to be unchanged throughout the whole campaign (Willis et al.,
162 2016). RIE of *r*BC to nitrate was calibrated by using size-selected (300 nm) BC
163 particles (REGAL 400R pigment black, Cabot Corp.) (Onasch et al., 2012), and the
164 average ratio of C_1^+ to C_3^+ was calculated to be 0.584 to correct the interference on C_1^+
165 from other organics. RIEs of ammonium, nitrate, sulfate and *r*BC were determined to
166 be 4.53, 1.10, 1.01, and 0.17, respectively, and RIE of organics used the default value
167 of 1.4 (Canagaratna et al., 2007). The size was calibrated by Polystyrene latex (PSL)
168 spheres (100–700 nm) (Duke Scientific Corp., Palo Alto, CA) before the measurement.
169 This study applied a collection efficiency (CE) of 0.5 for SP-AMS.

170 The co-located HR-AMS (DeCarlo et al., 2006) was used to measure all PM₁
171 including both *r*BC_c and non-*r*BC particles, but it detected only non-refractory species
172 (NR-PM₁ species) as its 600 °C thermal heater is unable to vaporize *r*BC and other
173 refractory components. In addition, mass concentrations of gaseous pollutants, carbon
174 monoxide (CO), ozone (O₃), nitrogen dioxide (NO₂), and sulfur dioxide (SO₂) were
175 measured by the Thermo Scientific analyzers provided by SAES. Meteorological
176 parameters including air temperature (T), relative humidity (RH), wind speed (WS),
177 wind direction (WD) and precipitation, were obtained from Xujiahui Environmental
178 Monitoring Station of Shanghai (31°11'49.1424"N, 121°26'34.44" E)(~2400 m away
179 from the site). The concentrations of particle-phase vanadium (V) and nickel (Ni) that
180 were used to investigate ship influence were measured independently by an
181 Atmospheric heavy metal analyzer (XHAM-2000A, SAIL HERO., China).

182

183 **2.2 Data analysis**

184 The AMS data (both SP-AMS and HR-AMS) were analyzed using standard ToF-
185 AMS data analysis tool (Squirrel version 1.59D and Pika version 1.19D), based on Igor
186 Pro 6.37 (Wavemetrics, Lake Oswego, OR, USA). The mass concentrations and high
187 resolution mass spectra (HRMS) of *r*BC and coating species (*r*BC_{CT}) were calculated
188 from high-resolution (HR) fitting of V-mode data. Size distributions of *r*BC_c
189 components were determined by the PToF data with unit mass resolution and were
190 scaled to their mass concentrations obtained above. In particular, size distribution of
191 *r*BC was scaled to that of *m/z* 24 (C₂⁺) (the scaling factor is *r*BC mass concentration to
192 that of calculated based on its size distribution), because *m/z* 24 as a *r*BC fragment, has
193 least interference from other organic or inorganic species; such treatment was adopted
194 in earlier studies too (Collier et al., 2018; Wang et al., 2019; Wang et al., 2016).

195 The HR ion fitting of AMS data is able to distinguish various ions and isotopic
196 ions and calculate elemental ratios of organics such as oxygen-to-carbon (O/C),
197 hydrogen-to-carbon (H/C), nitrogen-to-carbon (N/C), and organic mass to organic
198 carbon (OM/OC) ratios, via the original Aiken-ambient (A-A) method (Aiken et al.,
199 2008) and the improved method (I-A) (Canagaratna et al., 2015b). Outcomes of both

200 methods correlated well. Average O/C, H/C, and OM/OC ratios from the I-A method
201 used in this work were 24.9%, 7.3%, and 5.6%, respectively, higher than those from the
202 A-A method.

203 Furthermore, we performed Positive matrix factorization (PMF) (Paatero and
204 Tapper, 1994) analysis on the HRMS of organics measured by the SP-AMS via the PMF
205 Evaluation Tool (Ulbrich et al., 2009). The PMF solutions were thoroughly evaluated
206 following the protocols documented in Zhang et al. (2011). Finally, a 6-factor solution
207 was chosen as the optimal one. The final result included four primary OA (POA) factors,
208 namely hydrocarbon-like species enriched OA (HOA-rich), *r*BC-enriched OA (*r*BC-
209 rich), biomass burning OA (BBOA), water-soluble hydrocarbon-like species enriched
210 OA (WS-HOA), and two secondary OA (SOA) factors including a less oxidized
211 oxygenated OA (LO-OOA_{*r*BC}), and a more oxidized oxygenated OA (MO-OOA_{*r*BC})
212 (Key diagnostic plots in Fig. S2). As a comparison, the mass spectra of 5-factor and 7-
213 factor solutions are illustrated in Fig. S3 (for the 5-factor solution, factor 3 is a mixed
214 factor of BBOA and WS-HOA, while for the 7-factor solution, MO-OOA clearly
215 splits into factor 5 and factor 2). The HR-AMS data were processed in a similar way,
216 and four factors of NR-PM₁ organics were resolved, including hydrocarbon-related OA
217 (HOA_{NR-PM1}), cooking OA (COA_{NR-PM1}), less oxidized oxygenated OA (LO-OOA_{NR-PM1}-
218 PM₁), and more oxidized oxygenated OA (MO-OOA_{NR-PM1}) (Mass spectra and time
219 series shown in Fig. S4).

220

221 **2.3 Calculation of size-resolved hygroscopicity of *r*BCc**

222 The hygroscopicity parameter κ is a single parameter representing the
223 hygroscopicity of particles, calculated based on essentially the chemical composition
224 (Petters and Kreidenweis, 2007). SP-AMS measured size-resolved chemical
225 compositions of *r*BCc can thus lead to size-resolved hygroscopicity of *r*BCc (κ_{rBCc})
226 (Wu et al., 2019; Hu et al., 2021). This study applied a similar method proposed by Wu
227 et al. (2019) and the critical parameters involved in calculation are detailed in Table S1.
228 The procedures are briefly described here: (1) obtain the matrix of size-resolved mass
229 concentrations of each *r*BCc component from SP-AMS analysis; (2) convert the size-

230 resolved concentration matrix of inorganic ions (SO_4^{2-} , NO_3^- and NH_4^+) to
231 corresponding matrix of inorganic salts (NH_4NO_3 , NH_4HSO_4 , $(\text{NH}_4)_2\text{SO}_4$) using a
232 simplified solution of ion pairs (Gysel et al., 2007); (3) convert the size-resolved mass
233 matrix of all components (inorganic salts, OA and $r\text{BC}$) to those of size-resolved
234 volume fractions (Gysel et al., 2007; Chang et al., 2010; Wu et al., 2016); (4) combine
235 with previously reported hygroscopic parameters (Gysel et al., 2007; Chang et al.,
236 2010; Wu et al., 2016) to obtain the volumetric contributions of each component to the
237 hygroscopicity of $r\text{BCc}$; (5) use the Zdanovskii-Stokes-Robinson (ZSR) rule to derive
238 the size-resolved hygroscopicity of $r\text{BCc}$ ($\kappa_{r\text{BCc}}$) (Topping et al., 2005b, a).

239

240 **3 Results and discussion**

241 **3.1 Overview of chemical characteristics of $r\text{BCc}$ and NR- PM_{10} particles**

242 Figure 1 presents the time series of meteorological parameters, concentrations of
243 gaseous pollutants (CO , NO_2 , O_3 , and SO_2), ship emission tracers (vanadium and
244 nickel), $r\text{BC}$ and $r\text{BC}_{\text{CT}}$ species and their mass percentages to the total $r\text{BCc}$ mass,
245 PMF-resolved OA factors and their corresponding fractional contributions. The
246 sampling period was featured by relatively moderate temperatures and very stagnant
247 conditions with average ($\pm 1\sigma$) temperature of $15.3 (\pm 2.9) ^\circ\text{C}$ and wind speed (WS) of
248 $0.16 (\pm 0.29) \text{ m s}^{-1}$. The wind with speeds $< 0.5 \text{ m s}^{-1}$ dominated most of the sampling
249 days and 42% of sampling time was with near-zero wind, therefore overall influence of
250 WS and WD on surface mass loadings of $r\text{BCc}$ was insignificant. Yet one should keep
251 in mind that WD can affect $r\text{BCc}$ sources, and WS in higher altitude might be stronger
252 therefore long-range transport of air pollutants was still possible. The average
253 concentrations of CO , NO_2 , O_3 , SO_2 , V, and Ni were determined to be $0.60 (\pm 0.26)$
254 ppm, $29.20 (\pm 17.07)$ ppt, $27.10 (\pm 15.97)$ ppt, $1.27 (\pm 0.72)$ ppt, $4.05 (\pm 3.19) \text{ ng m}^{-3}$,
255 and $3.06 (\pm 2.45) \text{ ng m}^{-3}$, respectively.

256 The mass loadings of $r\text{BC}$ and $r\text{BC}_{\text{CT}}$ ranged from 0.04 to $11.00 \text{ } \mu\text{g m}^{-3}$ and 0.37
257 to $30.47 \text{ } \mu\text{g m}^{-3}$ with campaign-mean values ($\pm 1\sigma$) of $0.92 (\pm 0.81) \text{ } \mu\text{g m}^{-3}$ and $4.55 (\pm$
258 $4.40) \text{ } \mu\text{g m}^{-3}$. The coating materials accounted for $81.8 (\pm 5.3)\%$ of the total $r\text{BCc}$ mass,
259 of which organics was the most abundant species ($2.54 \pm 2.52 \text{ } \mu\text{g m}^{-3}$, $46.6 \pm 6.4\%$),

260 followed by nitrate ($1.20 \pm 1.30 \mu\text{g m}^{-3}$, $18.5 \pm 8.0\%$), ammonium ($0.44 \pm 0.40 \mu\text{g m}^{-3}$,
261 $8.0 \pm 2.0\%$), sulfate ($0.30 \pm 0.19 \mu\text{g m}^{-3}$, $7.1 \pm 3.6\%$), and chloride ($0.07 \pm 0.05 \mu\text{g m}^{-3}$,
262 $1.7 \pm 1.0\%$). The mass ratio of $r\text{BC}_{\text{CT}}$ to $r\text{BC}$ (R_{BC}) ranged from 2.2 to 9.0, with an
263 average of ~ 5.0 (± 1.7). The average R_{BC} was higher than that in California ($R_{\text{BC}} = 2.3$)
264 (Collier et al., 2018) and in Shenzhen ($R_{\text{BC}} = 2.5$) (Cao et al., 2022), lower than that in
265 Tibetan Plateau ($R_{\text{BC}} = 7.7$) (Wang et al., 2017) and similar to that in Beijing ($R_{\text{BC}} = 5.0$)
266 (Wang et al., 2019), suggesting $r\text{BC}$ was relatively thickly coated throughout the
267 campaign. Correlation between $r\text{BC}$ and $r\text{BC}_{\text{CT}}$ was moderate (Pearson's $r = 0.76$).
268 Correlation coefficients (r) of chloride, nitrate, sulfate and organics with $r\text{BC}$ were 0.72,
269 0.87, 0.71 and 0.73, respectively, suggesting variability of sources among different
270 coating components.

271 Figure 2 compares the campaign-averaged diurnal patterns of $r\text{BC}_c$ and NR- PM_{10}
272 species, chemical compositions of $r\text{BC}_c$ and NR- PM_{10} , and mass ratios of the species
273 coated on $r\text{BC}$ to those of NR- PM_{10} . We found that the diurnal variations of nitrate,
274 sulfate, ammonium, chloride were very similar ($r > 0.92$) between the two particle
275 groups, while apparent difference was found for $r\text{BC}_c$ organics with a much obvious
276 morning rush hour peak (6:00 - 9:00). The results indicate that the formation processes
277 of inorganic salts coated on $r\text{BC}$ were similar to those uncoated on $r\text{BC}$, but there were
278 large difference regarding sources/processes existed for organics. For $r\text{BC}$ itself, the
279 diurnal cycle presented clearly a morning peak and an evening peak, likely relevant
280 with rush hour traffic emissions (CO showed similar pattern). On the contrary, R_{BC} had
281 a minimum in the morning and dropped to a low level in later afternoon, probably due
282 to influence from traffic-emitted fresh and barely coated $r\text{BC}$ particles (details in
283 Section 3.2.1).

284 Distributions of species between $r\text{BC}_c$ and non- $r\text{BC}$ particles were also different,
285 leading to different chemical compositions (Figs. 2g and 2h). Sulfate tended to
286 preferentially condense on non- $r\text{BC}$ particles, as its mass contribution to total $r\text{BC}_c$
287 mass was only 6.5 (± 4.2)%, while its contribution to total NR- PM_{10} was 17.6 (± 8.3)%.
288 Apportionment of nitrate between $r\text{BC}_c$ and non- $r\text{BC}$ particles was about even as it
289 both occupied $\sim 26\%$ of the total $r\text{BC}_c$ and NR- PM_{10} masses. Organics occupied 55.9 (\pm

290 9.2)% of *rBCc* mass, larger than it in NR-PM₁ ($43.7 \pm 11.7\%$), due to that primary OA
291 species preferentially associated with *rBC*. Such result is similar to that observed in
292 winter in Beijing but contrary to the result that SOA was more abundant in *rBCc* in
293 Shenzhen (Cao et al., 2022).

294 On average, *rBCc* accounted for 19.1% of the total NR-PM₁ mass loading (21.61
295 $\pm 15.80 \mu\text{g m}^{-3}$)(Fig. 2i), comparable to that in Fontana, California (Lee et al., 2017).
296 The finding reveals that significant fractions of aerosol species were externally mixed
297 with *rBC*. Individually, sulfate captured by *rBC* only represented $7.4 (\pm 2.2)\%$ of NR-
298 PM₁ sulfate, similar to the earlier results (Lee et al., 2017; Wang et al., 2020a; Cao et al.,
299 2022; Ma et al., 2020); mass fractions of *rBCc* nitrate ($20.1 \pm 5.2\%$) and chloride (20.4
300 $\pm 5.5\%$) in NR-PM₁ were similar to the portion of total *rBCc* ($19.1 \pm 4.9\%$), while the
301 fraction of organics was higher ($26.1 \pm 7.3\%$). The lower fraction of sulfate in *rBCc*
302 than nitrate was likely due to a few reasons. One probable reason is traffic was a
303 dominant source of *rBC* (see Section 3.2.1 for details) and NO₂ is known to be mainly
304 from traffic as well, therefore secondarily formed nitrate was easy to condense on co-
305 emitted *rBC*, however SO₂ is mainly from other sources rather than traffic. Another
306 possible cause is that *rBC* concentration was relatively high during nighttime, and
307 nighttime formation of nitrate was significant; Sulfate, on the other hand, was mostly
308 formed in the afternoon due to photochemical oxidation in this study (see Section 3.3.2
309 for details), whereas afternoon *rBC* concentration was low. The relatively high ratio of
310 organics was attributed to the fact that majority of POA species were coated on *rBC*
311 (average ratio of $72.7 \pm 21.0\%$), while *rBC*-related SOA was $21.8 (\pm 7.7)\%$ of the total.
312 Note the *rBCc* POA here included all four POA factors, and COA_{NR-PM₁} did not coat on
313 *rBC* thus was not included in the calculation.

314

315 **3.2 Distinctive sources of OA in *rBCc* and in NR-PM₁**

316 As shown previously, source apportionment results of OA in *rBCc* and NR-PM₁
317 were different. This section discusses in details the characteristics of OA sources in
318 *rBCc* and in bulk NR-PM₁.

3.2.1 Source apportionment of OA in *r*BCc

Figure 3 shows the HRMS and temporal variations of the six OA factors resolved from PMF analysis of *r*BCc organics. Note the PMF analysis included *r*BC signals (i.e., C_x^+ ions) to aid identification of different factors, yet calculations of elemental ratios of these OA factors did not include C_x^+ ions in order to explicitly explore the properties of organic coating. The HRMS of HOA-rich and *r*BC-rich were similar to the OA previously reported in urban environments near traffic emissions and/or in gasoline/diesel vehicle exhaust (Massoli et al., 2012; Lee et al., 2015; Enroth et al., 2016; Saarikoski et al., 2016; Willis et al., 2016; Lee et al., 2017), therefore both factors were traffic-related. The HOA-rich mass spectrum was featured by intense hydrocarbon ion series (i.e., $C_nH_{2n+1}^+$ and $C_nH_{2n-1}^+$ ions in Fig. 3c), and a lowest O/C ratio of 0.07. Mass fraction of *r*BC signals (i.e., C_n^+ ions, such as m/z 12, 24, 36, 48, and 60, etc.) in HOA-rich was 8.1%. Mass spectrum of *r*BC-rich factor had remarkable contribution from *r*BC (24.2%). Beside C_n^+ ions, the *r*BC-rich factor contained more oxygenated organic fragments and presented a higher O/C ratio of 0.21 than that of HOA-rich, similar to previous studies (Willis et al., 2016; Lee et al., 2017). This result is reasonable as previous studies (Corbin et al., 2014; Malmberg et al., 2017) found that refractory organics could generate oxygenated ion fragments (such as CO^+ and CO_2^+ derived from oxygenated species on soot surface or inside soot nanostructure). HOA-rich factor correlated very well with the common AMS tracer of vehicular OA, $C_4H_9^+$ ($r=0.95$, Fig. 3i), while *r*BC-rich factor did match the variation of *r*BC well ($r=0.90$, Fig. 3g). Since diesel combustion often releases more *r*BC than that of gasoline, it is likely that the *r*BC-rich factor is representative of diesel vehicle exhaust while HOA-rich factor represents gasoline combustion emissions. This result demonstrates that laser-only SP-AMS is capable of distinguishing diesel and gasoline burning particles which typically cannot be separated by other AMS measurements. Further verification should be subject of future work.

In this work, a multiple linear regression for the three-dimension size-resolved mass spectra according to the method provided in Ulbrich et al. (2012) was used to resolve the average size distributions of six OA factors. The diagnostic plots are shown

349 in Fig. 4. Overall, the lumped size distribution of six OA factor could reproduce well
350 that of total OA (except for a few size bins, most deviations are within 10%).
351 Correlation between measured and reconstructed OA of all size bins was very tight with
352 r of 0.99 and a slope of 0.97, indicating the robustness of the regressed size distributions
353 of all OA factors. The results together with size distributions of r BCc components, and
354 corresponding mass fractional contributions of different components in all size bins are
355 illustrated in Fig. 5. The average HOA-rich size distribution peaked around 150 nm
356 (D_{va} , vacuum aerodynamic diameter), generally matching with previously reported size
357 distribution of HOA (Sun et al., 2012;Ulbrich et al., 2012;Zhou et al., 2016).
358 Interestingly, size distribution of r BC-rich factor presented two modes, with one
359 peaking \sim 260 nm, and a more pronounced one peaking \sim 580 nm (Fig. 5a).

360 The BBOA was identified owing to its obviously higher signals of $C_2H_4O_2^+$ (2.03%)
361 and $C_3H_5O_2^+$ (1.62%) than those in other factors, as these two ions are well-known AMS
362 fragments of the biomass burning tracer, levoglucosan (Mohr et al., 2009;Cubison et
363 al., 2011). The time series of BBOA correlated particularly tightly with both marker
364 ions (r of 0.86 and 0.80, respectively); it in fact also correlated well with K^+ ($r=0.79$),
365 another tracer of biomass burning emission. The O/C and H/C ratios of BBOA were
366 0.12 and 1.78, and C_n^+ ions accounted for 9.1% of BBOA, all suggesting that the factor
367 was fresh and might be co-emitted with r BC. A relatively high N/C ratio (0.033) was
368 found for BBOA, which could be attributed to the large amounts of nitrogen-containing
369 organic species enriched in biomass burning OA (Laskin et al., 2009). In addition, the
370 size distribution of BBOA (Fig. 5a) (peak size \sim 500 nm) was similar to that of biomass
371 burning BC-containing particles obtained using single particle mass spectrometry in
372 Shanghai (Gong et al., 2016).

373 The PMF analysis deconvoluted a unique OA factor coated on r BC, namely WS-
374 HOA. Firstly, the WS-HOA mass spectrum had a series of hydrocarbon ion fragments
375 and its time series correlated well with them (e.g., r of 0.90 and 0.92 for WS-HOA vs.
376 $C_4H_7^+$ and $C_4H_9^+$, respectively). Secondly, this factor correlated the best ($r=0.57$) with
377 aerosol liquid water content (ALWC) (Fig. 3j) compared with the other five OA factors
378 (all $r<0.2$). Thirdly, a previous study (Ye et al., 2017) that investigated specially the

379 water-soluble fraction of OA via PMF analysis, separated also a HOA factor that
380 contained significant nitrogen-containing organic fragments, with a highest N/C ratio
381 among all other factors, and correlated well those nitrogenated fragments. The WS-
382 HOA defined here showed similar characteristics with the highest N/C (0.037) among
383 all factors and tight correlations with nitrogen-containing fragments ($r > 0.80$). At last,
384 although WS-HOA had a relatively high O/C (0.31) with remarkable contributions from
385 $C_2H_3O^+$ and CO_2^+ , yet its correlations with these two ions were in fact not strong (r of
386 0.46 and 0.44, respectively); and WS-HOA had the least fraction of rBC fragments
387 (0.9%) (note rBC is hydrophobic), even less than the two SOAs (Fig. 3d). Both results
388 suggest that this factor is a collection of water-soluble primary OA species. The peak
389 of WS-HOA size distribution was ~ 150 nm, close to that of HOA-rich (Figs. 5a).
390 Aqueous-phase processed SOA (aqSOA) were typically with very high O/C ratio (Xu
391 et al., 2019), and size distribution of aqSOA often presented a droplet mode, peaking in
392 relatively large sizes (such as 500~600 nm)(Gilardoni et al., 2016; Wang et al., 2021; Ge
393 et al., 2012). Therefore, the moderate O/C (0.31) and small mode size of WS-HOA
394 again manifest it was not from aqueous-phase reactions but more likely the water-
395 soluble fraction of POA.

396 The PMF analyses separated two SOA factors, LO- OOA_{rBC} and MO- OOA_{rBC} .
397 Mass spectral features of the two SOAs were consistent with previous studies: The LO-
398 OOA_{rBC} was rich in $C_xH_yO_1^+$ ions (38.7%) (such as $C_2H_3O^+$ at m/z 43) but with less
399 contribution from $C_xH_yO_2^+$ ions (7.6%) and an overall moderate O/C (0.25), while MO-
400 OOA_{rBC} had much more contribution from $C_xH_yO_2^+$ ion family (22.7%) (such as CO_2^+
401 at m/z 44) and $C_xH_yO_1^+$ ions (44.7%), with a high O/C ratio (0.56). In addition, LO-
402 OOA_{rBC} correlated better with nitrate ($r=0.83$) than it with sulfate ($r=0.69$), while the
403 correlation between MO- OOA_{rBC} and sulfate ($r=0.84$) is better than it with nitrate
404 ($r=0.76$). Size distributions of the two SOAs were also in accord with their secondary
405 behaviors, both accumulating at larger sizes (~ 450 nm for LO- OOA_{rBC} , and a bit larger
406 mode size of ~ 500 nm for MO- OOA_{rBC}), in agreement with previous observations (Sun
407 et al., 2012; Ulbrich et al., 2012; Zhou et al., 2016).

408 In total, traffic-related POA (sum of HOA-rich, rBC -rich and WS-HOA) was the

409 most abundant source of $rBCc$ organics (39.1%); BBOA occupied \sim 18.4%; the
410 contributions of two SOAs were on par with each other (20.2% for LO- OOA_{rBC} , and
411 22.3% for MO- OOA_{rBC}) (Fig. 2g). Among traffic POA, gasoline derived HOA-rich
412 factor outweighed the diesel derived rBC -rich factor (11.4% vs. 6.3% of the total $rBCc$).
413 Contributions of different factors varied greatly for different sizes of $rBCc$ particles
414 (Fig. 5b). Small-sized particles were overwhelmingly dominated by traffic POA; SOA
415 contributions increased with increase of size, and dominated over POA for 300-800 nm
416 ones; contribution of BBOA was also relative larger for 300-800 nm than for other-
417 sized ones; the very large particles (800-1000 nm) were found to be affected mainly by
418 traffic POA in this work. Correspondingly, for the total $rBCc$ particles, rBC cores
419 peaked at \sim 170 nm, while other secondary inorganic components, behaving like SOA
420 factors, all peaked at a big size (\sim 550 nm) (Fig. 5c) and their mass percentages were
421 also large for large-sized particles (Fig. 5d).

422 Figure 6a further demonstrates the changes of mass fractions of each $rBCc$
423 component as a function of R_{BC} . R_{BC} is a proxy of coating thickness. It was found that
424 the thinly coated $rBCc$ ($R_{BC}<3$) were dominated (up to \sim 80%) by traffic-related POA.
425 With the increase of R_{BC} , contribution of secondary components increased gradually,
426 especially the two SOAs and nitrate (sulfate showed little changes across the whole R_{BC}
427 range); SOA and nitrate contributions reached 40.2% and 31.3% at $R_{BC} = 8$, respectively.
428 Accordingly, the oxidation degree ($OSc = 2 * O / C - H / C$) (Kroll et al., 2011) of coated
429 organics increased with R_{BC} . In Fig. 6b, we presented the mass contributions of OA
430 factors to the rBC mass at different R_{BC} values. The rBC was as expected,
431 predominantly associated with POA (from 94% at $R_{BC}<2$ to 66% at $R_{BC}>8$), similar to
432 those from Fontana (Lee et al., 2017). Contribution of rBC -rich factor decreased
433 obviously, and those of SOA factors increased with R_{BC} . The contributions of HOA-
434 rich, WS-HOA and BBOA factors changed little.

435

436 3.2.2 Comparisons with NR- PM_{10} organics

437 As shown in Fig. S4, PMF analysis separated four OA factor for NR- PM_{10} organics.
438 Two SOA factors (LO- OOA and MO- OOA) were resolved for both $rBCc$ and NR- PM_{10} ,

439 and their contributions to them were also close (Figs. 2g and 2h). Correlations of time
440 series between the two LO-OOA factors and between the two MO-OOA factors were
441 also tight (r of 0.94 and 0.90, respectively), indicating similar formation processes for
442 each SOA. But, of course, the SOAs from $rBCc$ and from $NR-PM_1$ were not entirely
443 the same, as later ones had higher O/C ratios (0.52 of $LO-OOA_{NR-PM_1}$ and 0.62 of $MO-$
444 OOA_{NR-PM_1}). On average, the portion of LO-OOA coated on rBC took up 21.6% mass
445 of LO-OOA in total $NR-PM_1$, and the portion was 26.0% for MO-OOA. This result
446 suggests that there were some differences regarding the partitioning of LO-OOA and
447 MO-OOA onto $rBCc$ and non- rBC particles.

448 Compared with SOAs, source apportionment results of POA were quite distinct.
449 Firstly, there was only one HOA factor resolved for $NR-PM_1$, while three HOA factors
450 were separated for $rBCc$. Note the rBC -rich and WS-HOA factors occupied merely
451 3.1% and 2.1% of $NR-PM_1$ OA mass, respectively, probably the cause that they were
452 not identified in $NR-PM_1$ OA. Nevertheless, mass loadings of the sum of HOA-rich,
453 rBC -rich and WS-HOA factors (termed as HOA_{rBC}) agreed quite well with that of
454 HOA_{NR-PM_1} ($r=0.95$) (Fig. 7a), verifying both source apportionment results. And,
455 HOA_{rBC} took up 63.7% of HOA_{NR-PM_1} , while previous studies reported that 81%
456 (Massoli et al., 2012) and 87 % (Massoli et al., 2015) of HOA were associated with
457 rBC . These results imply that HOA species in $NR-PM_1$ were largely internally mixed
458 with rBC affected by vehicular emissions. Secondly, CO_{NR-PM_1} was only identified in
459 $NR-PM_1$ OA. AMS-resolved CO_{NR-PM_1} is mainly from cooking oil and food
460 ingredients, therefore it negligibly internally mixes with rBC . This result is consistent
461 with previous observations (Lee et al., 2015; Willis et al., 2016; Lee et al., 2017; Collier
462 et al., 2018). At last, BBOA was identified in $rBCc$ OA but not in $NR-PM_1$ OA. One
463 plausible reason was that the BBOA mass contribution was minor (equivalent to <5%
464 of $NR-PM_1$ OA mass) therefore was not able to be separated from other OA factors;
465 another speculation is that laser only SP-AMS can detect refractory species that HR-
466 AMS cannot, and a portion of these refractory species are likely rich in biomass burning
467 OA. Identification of BBOA in $rBCc$ rather than in simultaneously measured total $NR-$
468 PM_1 was also found in Tibet (Wang et al., 2017) and Beijing (Wang et al., 2020a), role

469 of such BBOA and its interplay with *r*BC core remain a subject of future work.

470 Diurnal cycles of the POA and SOA factors from both PMF analyses are compared
471 in Figs. 7b and 7c. The diurnal pattern of stacked HOA_{*r*BC} indeed agreed with that of
472 HOA_{NR-PM₁}, both with two peaks in the morning and evening rush hours, and, the
473 patterns of *r*BC-rich, HOA-rich, and WS-HOA factors showed similar behaviors
474 individually (Fig. 7b). The diurnal variation of COA_{NR-PM₁} had pronounced peaks
475 during lunch and dinner times, and its percentage in NR-PM₁ OA reached 54% at night.
476 Diurnal patterns of two LO-OOA factors were somewhat different ($r=0.35$). LO-
477 OOA_{*r*BC} has a major peak in the afternoon, while though LO-OOA_{NR-PM₁} concentration
478 rose in the afternoon too but peaked in early evening (~20:00). The daily variations of
479 two MO-OOA factors were similar ($r=0.83$), both peaking at 16:00. The afternoon
480 increases of both SOAs indicate an important role of photochemical reactions, yet
481 differences in formation mechanisms, volatilities and partitioning behaviors of SOA
482 products could lead to diversities of their diurnal patterns and HRMS.

483

484 **3.3 Evolution and formation *r*BCc components**

485 **3.3.1 Behaviors of *r*BC cores**

486 Size distribution of *r*BC cores shown in Fig. 5c was relatively wide. Beside the
487 main peak at ~170 nm, it extended significantly into large sizes and had a small peak at
488 ~550 nm, which was close to the peak of secondary components. With results shown in
489 Fig. 5a, we found that small *r*BC cores were often thinly coated, while thickly coated
490 *r*BCc particle were often highly aged and a portion of them also had large-sized *r*BC
491 cores. This result is likely owing to that oxidation of *r*BC-bounded organics and/or
492 condensation of secondary species onto *r*BC surface can induce restructuring of soot
493 aggregates to form compact and large cores (Chen et al., 2018;Chen et al., 2016). Such
494 phenomenon is in line with earlier studies (Liu et al., 2019;Gong et al., 2016). We
495 further show the image plot of size distributions of *r*BC at different R_{BC} in Fig. S6a. It
496 can be found that the *r*BC mass in a large part concentrated in particles with R_{BC} of 5-
497 8, and there was indeed a significant portion of *r*BC appearing in large size range (400-
498 800 nm) with very thick coating (R_{BC} of 8-9).

499

500 **3.3.2 Formation of inorganic salts**

501 Sulfate and nitrate both peaked at a big size ~ 550 nm (Fig. 5c) and were mainly
502 associated with thickly coated $rBCc$ ($R_{BC} > 6$, Figs. S6d and S6e). To investigate the
503 impacts of photochemistry and aqueous/heterogeneous chemistry on the formation of
504 rBC_{CT} species, we plotted the image plots of size distributions of nitrate, sulfate and
505 organics versus O_X ($O_3 + NO_2$) and relative humidity (RH) in Fig. 8. Here O_X is used
506 as a proxy of photochemical activity (Xu et al., 2017), and RH is an indicator of aqueous
507 reactions (Wu et al., 2018). Nitrate significantly concentrated in 65-100 ppb O_X range
508 but there was a weak accumulation in low O_X as well (Fig. 8a), while in Fig. 8d, nitrate
509 had a prominent hotspot in $RH > 85\%$. Generally, both strong photochemical activity
510 and high RH could promote nitrate formation. For sulfate, although the distribution was
511 scattered due to low level of mass loadings, high O_X level seemed to favor sulfate
512 formation (Fig. 8b); sulfate was scattered in the whole RH range and there were some
513 enhancements at high RH ($> 80\%$) but was much less clear-cut (Fig. 8e). Therefore
514 aqueous-phase production of sulfate was not important in this campaign.

515 We further calculated the sulfur oxidation ratio (SOR) and nitrogen oxidation ratio
516 (NOR) (Xu et al., 2014), and plotted their variations against O_X and RH in Figs. 9a and
517 9e, respectively. The NOR rose substantially at $O_X > 60$ ppb but showed no increase at
518 $O_X < 60$ ppb, while it increased continuously with RH. Mass ratio of nitrate to rBC
519 stayed at a high level during nighttime when RH was high as well (overall diurnal trend
520 of NO_3^-/rBC was similar to that of RH, see Figs. S7a and S7d). This result indicates a
521 likely more important role of nocturnal nitrate formation (N_2O_5 hydrolysis) (Pathak et
522 al., 2011) (Sun et al., 2011) than photochemical production of nitrate during this
523 campaign; moreover, low temperature and high RH favor nitrate partitioning into the
524 particle phase during nighttime too (Gao et al., 2011). For sulfate, The SOR increased
525 with O_X while it increased with RH under relatively dry conditions ($< 60\%$) but
526 decreased with RH when $RH > 60\%$. This result, on the other hand, highlights that
527 photochemical production is more important than aqueous/heterogenous formation for
528 sulfate. In addition, mass ratios of sulfate to rBC were enhanced remarkably during

529 daytime and peaked in the afternoon (Fig. S7e), supporting the key role of
530 photochemical formation of sulfate. Sulfate precursor SO₂ was at a high level during
531 daytime too. The main formation pathway of sulfate is strikingly different from that
532 observed in winter Nanjing (Wu et al., 2018), suggesting significant seasonal variability
533 of sulfate formation even in the same region.

534

535 3.3.3 Evolution of organics

536 Organics had a broad average size distribution (Fig. 5c), but unlike *r*BC, its main
537 peak appeared at 500~600 nm. Figure S6b shows that the majority of organics
538 partitioned in *r*BCc with *R*_{BC} of 5.0-9.0 and wide size coverage (300-800 nm).
539 Regarding its dependences on O_X and RH, it mainly accumulated at O_X>70 ppb (Fig.
540 8c) and very high RH (~90%) (Fig. 8f). Consistently, O/C ratio and OSc both peaked
541 in the afternoon (Fig. S7b), all suggesting a critical role of photochemistry in affecting
542 the behavior of organics.

543 Figure 9 illustrates the mass ratios of each OA factor to *r*BC varying with O_X and
544 RH. Mass ratios of all four POA factors generally presented decreasing trends (despite
545 some fluctuations) against O_X (Fig. 9b), and the total POA_{*r*BC} showed evident decrease
546 with increase of O_X (Fig. 9d). Instead, both LO-OOA_{*r*BC} and MO-OOA_{*r*BC}, as well as
547 their sum (SOA_{*r*BC}) increased continuously with O_X (Figs. 9c and 9d). This result proves
548 that photochemical oxidation contributed significantly to both LO-OOA_{*r*BC} and MO-
549 OOA_{*r*BC} formations. Comparatively, decreases of POA_{*r*BC} perhaps point to its reaction
550 loss upon photochemical oxidation. With regard to RH, besides WS-HOA, the other
551 three POA_{*r*BC} factors showed almost no dependences on RH (Fig. 9f); note the increase
552 of WS-HOA with RH did not indicate the aqueous production of WS-HOA (see
553 discussion in Section 3.2.1), but a result of enhanced dissolution with increase of
554 moisture. Overall small increase of POA_{*r*BC} (Fig. 9h) with RH then mainly attributed to
555 WS-HOA. Contrary to the trends with O_X, mass ratios of two SOA factors as well as
556 the total SOA to *r*BC went down with increase of RH (Figs. 9g and 9h), indicating a
557 trivial role of aqueous/heterogenous oxidation for the SOA coated on *r*BC observed
558 during this campaign.

559 The aging of OA is generally characterized by the increase of O/C and decrease of
560 H/C (Ng et al., 2011;Zhao et al., 2019).The different aging pathways of OA follow
561 different slopes in the Van Krevelen (VK) diagram (Heald et al., 2010). For example,
562 addition of only one oxygen atom to the carbon skeleton results in a slope equal to 0,
563 while replacement of the hydrogen atom with a carboxylic acid group ($-\text{COOH}$) results
564 in a slope of -1 without fragmentation (C-C bond breaking), and -0.5 with
565 fragmentation (Heald et al., 2010;Ng et al., 2011;Zhao et al., 2019). As presented in Fig.
566 10a, fitting of all OA data yielded a slope of -0.96 , very close to -1 , suggesting that
567 OA ageing process resembled the hydrogen substitution with a $-\text{COOH}$ group
568 (carboxylation). Interestingly, the four OA factors (HOA-rich, WS-HOA, LO- OOA_{rBC}
569 and MO- OOA_{rBC}) aligned almost in a straight line with a slope of -0.77 (Fig. 10a), also
570 close to -1 . Figure S8 further reveals that there was a strong anti-correlation between
571 mass fractions of sum of HOA-rich and WS-HOA and sum of LO- OOA_{rBC} and MO-
572 OOA_{rBC} ($r=-0.97$); the slope of fitted line was -0.86 . All these results suggest that OA
573 evolution may contain a channel of photochemical transformations from HOA-rich and
574 WS-HOA to LO- OOA_{rBC} and then to MO- OOA_{rBC} . This result is also in line with the
575 observed decrease of POA_{rBC} and increase of SOA_{rBC} against O_x .

576 CHO^+ , CHO_2^+ and $\text{C}_2\text{H}_2\text{O}_2^+$ are the AMS tracer ions for carbonyl, carboxylic acid
577 and glyoxal, respectively (Wang et al., 2020b;Canagaratna et al., 2015a;Yu et al., 2014).
578 Mass fractions of these three fragment ions presented decreasing trends (or no clear
579 trends) against RH (Fig. S9), suggesting again that aqueous processing is not an
580 important pathway in OA evolution during this campaign. Conversely, fractional
581 contributions of these three ions presented increasing trends versus O_x , supporting the
582 dominance of photochemical oxidation pathway (Figs. 10b-d). Figures 10e-g show the
583 scatter plots of H/C versus O/C at different O_x concentrations. The regressed slope was
584 -1.03 for low O_x (<60 ppb) conditions (Fig. 10e), indicating that the OA aging at low
585 O_x level is mainly analogue to the carboxylation process. This result corresponds
586 precisely to the fact that mass fractions of CHO_2^+ and $\text{C}_2\text{H}_2\text{O}_2^+$ increased gradually with
587 O_x at low O_x , whereas that of CHO^+ remained essentially unchanged (Figs. 10e-g). The
588 fitted slope was -1.14 for moderate O_x conditions (60-80 ppb), and it changed to -0.43

589 for high O_x level (>80 ppb) but correlation became weaker ($r=-0.57$). This result
590 implies that the OA evolution under high O_x conditions might include oxidations by
591 the additions of alcohols, peroxides and carboxylation. In all, the evolution of *rBCc* OA
592 in Shanghai during this campaign is governed by photochemistry rather than aqueous
593 chemistry, but with different oxidation pathways at different O_x levels.

594 **3.3.4 Coating time of secondary species onto *rBC***

595 Although the *rBCc* organics was dominated by primary sources (Fig. 2g), the
596 diurnal variations of OSc, O/C and H/C of the total organics, were controlled
597 predominantly by the two SOA factors. Correlations between the diurnal cycles of MO-
598 OOA_{*rBC*}/*rBC* and OSc were extremely well ($r=0.97$ with OSc and $r=0.98$ with O/C),
599 and those of LO-OOA_{*rBC*} were also tight ($r=0.91$ with OSc and $r=0.92$ with O/C).
600 The correlations with LO-OOA_{*rBC*} were a bit weaker than those of MO-OOA_{*rBC*},
601 indicating that MO-OOA_{*rBC*} was probably the final products and was more important
602 in governing the overall oxidation level of organic coating. Figure 11a depicts the
603 diurnal variations of SOA_{*rBC*}/*rBC* and POA_{*rBC*}/*rBC*. Diurnal variations of POA_{*rBC*}/*rBC*
604 and *rBC* were overall similar, while the daily pattern of SOA_{*rBC*}/*rBC* was almost
605 opposite to that of *rBC*. This result likely indicates that most POA_{*rBC*} species were co-
606 emitted and coated on *rBC* cores originally, therefore the coating process during *rBCc*
607 lifecycle was mainly relevant with SOA species rather than POA species. This process
608 is mainly through photochemical reactions, including in-situ oxidation of originally
609 coated POA species (for example, oxidation of HOA-rich plus WS-HOA to LO-
610 OOA_{*rBC*}, then to MO-OOA_{*rBC*}), and partitioning of secondary species formed in gas-
611 phase reactions onto *rBC* surface. In addition, sulfate and nitrate were both secondarily
612 formed, but sulfate was dominated by photochemical production while nitrate was
613 governed by nocturnal heterogenous formation (as discussed in Section 3.3.2); different
614 diurnal patterns of them (Fig. S7) point to different coating processes too.

615 In this regard, we hereby propose a concept of average coating time (ACT), which
616 is used to roughly estimate the timescales required for secondary components coated
617 onto *rBC*. The specific method is listed as follows:

618 1. Move forward the diurnal variation of SA_{*rBC*}/*rBC* (SA represents a secondary

619 aerosol species) for n hours to get a new SA_{rBC}/rBC diurnal pattern, labelled as
620 “ $SA_{rBC}/rBC(r-nh\text{-ahead})$ ”. Here, the r value is the linear correlation coefficient between
621 the new SA_{rBC}/rBC diurnal pattern with that of rBC .

622 2. Choose the best correlation coefficient ($r-nh\text{-ahead}$), and nh corresponds to the
623 ACT for this secondary component.

624 Diurnal patterns of $LO\text{-}OOA_{rBC}/rBC$ and $MO\text{-}OOA_{rBC}/rBC$ are depicted in Fig.
625 11b. They were both opposite to the trend of rBC , and they were similar to each other,
626 except that $MO\text{-}OOA_{rBC}/rBC$ peaked hours later in the afternoon, signifying that the
627 $MO\text{-}OOA_{rBC}$ needs longer time to coated on rBC than $LO\text{-}OOA_{rBC}$ does, consistent
628 with the fact that $MO\text{-}OOA_{rBC}$ was “more aged”. Correspondingly, we obtained an ACT
629 of 7 hours for $MO\text{-}OOA_{rBC}$ ($0.35\text{-}7\text{h}\text{-ahead}$) and 5 hours for $LO\text{-}OOA_{rBC}$ ($0.57\text{-}5\text{h}\text{-}$
630 ahead) (Fig. 11c) using the method described above. The ACT of sulfate ($0.65\text{-}7\text{h}\text{-ahead}$)
631 and nitrate ($0.30\text{-}19\text{h}\text{-ahead}$) were 7 and 19 hours, respectively (Fig. 11d). The results
632 suggest that the rBC emitted mainly in the morning rush hours requires a few hours to
633 be adequately coated by $LO\text{-}OOA_{rBC}$, $MO\text{-}OOA_{rBC}$ and sulfate, as these three species
634 are mainly produced in the afternoon by photochemical reactions; while photochemical
635 production of nitrate is insignificant, thus rBC was coated by nitrate until later night
636 when nitrate was formed efficiently by heterogenous N_2O_5 hydrolysis. Note the best r
637 values were not high (for example, 0.35 for $MO\text{-}OOA_{rBC}$ and 0.30 for nitrate), yet the
638 adjusted trends of all secondary components (Figs. 11c and 11d) matched that of rBC
639 quite well during 3:00~12:00 (r of 0.90, 0.91, 0.84 and 0.84 for $MO\text{-}OOA_{rBC}$, $LO\text{-}$
640 OOA_{rBC} , sulfate and nitrate, respectively), likely meaning that daytime variations of
641 two SOAs and sulfate were indeed controlled by the coating process while governing
642 factors of their nighttime variations might be complex, and *vice versa* for nitrate.

643

644 **3.4 Size-resolved hygroscopicity of $rBCc$**

645 By using the method in Section 2.3, we calculated the size-resolved hygroscopicity
646 parameters for the total $rBCc$ (κ_{rBCc}) and the coatings materials (κ_{CT}) across the whole
647 campaign. The image plots are illustrated in Figs. 12a and b. Generally, large κ_{rBCc} and
648 κ_{CT} values occurred at big particle sizes, and this result was overall similar to that in

649 Nanjing during winter (Wu et al., 2019). We further illustrate the size-resolved κ_{rBCc} as
650 a function of R_{BC} in Fig. 12c. The figure shows that κ_{rBCc} overall became larger with
651 increasing particle size regardless of the coating thickness. However, there were some
652 (though not significant) relatively high κ_{rBCc} values in the range of 80-150 nm (bottom
653 left and bottom right in Fig. 12c).

654 In Figs. 13a and b, we further determined the average size-resolved κ_{rBCc} and κ_{CT}
655 as a function of coated diameter (D_{rBCc}). Both κ_{rBCc} and κ_{CT} values were relatively high
656 at $D_{rBCc} < 100$ nm and presented slight decreases from 100 to 150 nm. This is distinctive
657 from those observed in Nanjing, where κ_{CT} increased with D_{rBCc} from 50 nm (Wu et al.,
658 2019). From Figs. 5b and 5d, it can be seen the $rBCc$ with $D_{rBCc} < 150$ nm was dominated
659 by a portion of ammonium and sulfate (8-10%) and organics (~60%), of which organics
660 was dominated by WS-HOA. Such composition explains the relatively high
661 hygroscopicity at $D_{rBCc} < 150$ nm as both ammonium sulfate and WS-HOA are
662 hydrophilic; a slight decrease of the hygroscopicity from 100 to 150 nm D_{rBCc} was also
663 a response of decreased mass contributions of ammonium sulfate and WS-HOA (and
664 increase of hydrophobic HOA-rich contribution).

665 Figures 13a and b provide the fitted exponential functions for the mean κ_{rBCc} and
666 κ_{CT} with D_{rBCc} . The equations are: $\kappa_{rBCc}(x) = 0.29 - 0.14 \times \exp(-0.006 \times x)$ and $\kappa_{CT}(x) =$
667 $0.35 - 0.09 \times \exp(-0.003 \times x)$ (x is D_{rBCc} , $150 < x < 1000$ nm). Here, 0.29 and 0.35 are
668 the upper limits of κ_{rBCc} and κ_{CT} , higher than those reported in Nanjing (0.28 and 0.30
669 for κ_{rBCc} and κ_{CT}); yet the increasing rates of κ_{rBCc} and κ_{CT} with D_{rBCc} are 0.14 and 0.09,
670 respectively, which are much lower than those from Nanjing (0.35 and 0.27 for κ_{rBCc}
671 and κ_{CT}) (Wu et al., 2019). Smaller increased contributions of hydrophilic secondary
672 species from 150 to 1000 nm in Shanghai than those from 100 to 1000 nm in Nanjing
673 are likely the cause of smaller increasing rates of hygroscopicity parameters. In addition,
674 it should be noted that the hygroscopicity is not only determined by the bulk
675 composition, but also affected by the phase state of particles. For instance, a recent
676 study reveals that the hygroscopic growth of phase-separated particles (with ammonium
677 sulfate as cores) can be reduced by the secondary organic shells and is dependent on
678 the thickness of organic coating (Li et al., 2021).

679 The critical supersaturation (SS_C) for a selected dry diameter (D_{rBCc} measured by
680 SP-AMS) of a $rBCc$ particle with a hygroscopicity parameter κ_{rBCc} , can be calculated
681 by the “ κ -Kohler theory” equation (Petters and Kreidenweis, 2007). Based on the size-
682 resolved κ_{rBCc} , the CCN activation diameter (D_{50}) of particles at a given critical SS_C can
683 be calculated (Wu et al., 2019). Then, by combining the measured $rBCc$ number size
684 distribution and the D_{50} value, activation fraction (f_{AC}) of $rBCc$ number population (i.e.,
685 the fraction greater than the D_{50}) can be obtained. Figure 13c shows the SS_C as a
686 function of D_{rBCc} for the entire sampling period to obtain the D_{50} at a specific SS_C . The
687 D_{50} values of the $rBCc$ particles were determined to be $166 (\pm 16)$ nm and $110 (\pm 5)$ nm
688 for SS_C of 0.1% and 0.2%, respectively. The two D_{50} values are both smaller than those
689 determined for $rBCc$ particle in Nanjing (Wu et al., 2019), likely owing to the presence
690 of WS-HOA in Shanghai. Figure 13d shows the f_{AC} at SS_C of 0.1% (D_{50} of 166 nm) was
691 $16 (\pm 3)\%$, and the f_{AC} at SS_C of 0.2% (D_{50} of 110 nm) was $59 (\pm 4)\%$.

692

693 **3.5 A case study influenced by ship emissions**

694 **3.5.1 Potential source areas of $rBCc$**

695 To explore the potential geographic origins of $rBCc$ at the receptor site, the hybrid
696 single-particle Lagrangian integrated trajectory (HYSPLIT) model (version 4.9) was
697 applied here. Figure 14a shows that the backward trajectories were classified into four
698 air mass clusters, including one long-range transport from northern sea (Cluster1), one
699 long-range transport from northeastern sea (Cluster2), a local one from eastern ports
700 (Cluster3), and one from northwestern inland region (Cluster4). The four clusters
701 occupied 23.8%, 33.8%, 37.3%, and 5.0% of the total trajectories, respectively. It is
702 clear that the sampling period was influenced by offshore air masses in most of the time
703 (95%). Cluster3 had the highest mean $rBCc$ concentrations ($13.2 \pm 10.9 \mu\text{g m}^{-3}$) while
704 the mean concentrations of the other three clusters were apparently lower ($4.3\sim 5.2 \mu\text{g}$
705 m^{-3}). This result is plausible as Cluster3 was the shortest in length therefore was least
706 diluted compared with other three clusters. Average chemical compositions of the $rBCc$
707 from four clusters (Fig. 14b) showed some differences as well: Cluster1 and Cluster2
708 had higher MO-OOA $_{rBC}$ contributions, possibly owing to the interceptions of more

709 aged SOA species during the long-range transports; While Cluster4 had less MO-
710 OOA_{rBC} but a bit more POA_{rBC} contributions, likely attributing to more primary species
711 emitted in inland regions (similarly, a higher fraction of nitrate was likely because of
712 enhanced NO₂ emissions).

713

714 **3.5.2 A typical case of *rBC*c affected by ship emissions**

715 Ship emissions are found to have an important impact on the air quality of
716 Shanghai and the Yangtze River Delta (Zhao et al., 2013; Fan et al., 2016; Liu et al.,
717 2017b; Chen et al., 2019). The ship engines usually burn heavy fuel oil (HFO), and
718 vanadium (V) and nickel (Ni) can be adopted as reliable tracers for the ship-emitted
719 particles (Ault et al., 2009; Moldanová et al., 2009; Ault et al., 2010). The long-term
720 variation of Ni/V ratio in ship-emitted particles in Shanghai has been recently reported
721 (Yu et al., 2021). The main conclusion is that Ni/V ratio was close to 0.4 in 2018, while
722 it became to be greater than 2.0 in 2020. Our measurement was carried out in 2018, and
723 we therefore chose a period from November 3rd to 5th as a typical case affected by ship
724 emissions (SEP period), since the average Ni/V ratio was ~0.50 and high concentrations
725 of both Ni and V were found. Figure S10 shows the concentration-weighted trajectories
726 (CWT) of ship emission tracers (V, Ni), *rBC* and *rBC*-rich factor during SEP, displaying
727 that these components were mainly from sea. During SEP, the correlation coefficient (*r*)
728 between V(Ni) and *rBC* was 0.69 (0.74), indicating the SEP was indeed impacted by
729 ship emissions.

730 Figure 15 displays the chemical characteristics of *rBC*c and NR-PM₁ components
731 (especially the OA factors), V and Ni, gaseous pollutants and the meteorological
732 parameters during SEP. As a comparison, we also selected a period with no ship
733 emissions with the same time span as SEP (from 0:00 on November 10 to 0:00 on
734 November 12, termed as non-SEP), and a similar plot during non-SEP is presented in
735 Fig. S11. During SEP, the wind was mainly from east, and the average wind speed was
736 ~0.5 (± 0.3) m s⁻¹ (Fig. 15a); Wusong, Luojing and Waigaoqiao ports located northeast
737 of the sampling site (Fig. S1). Instead, the wind was mainly from north during non-SEP
738 (Fig. S11a). During SEP, the average mass concentrations of V and Ni were 6.3 (± 3.1)

739 ng m^{-3} and $3.2 (\pm 1.4) \text{ ng m}^{-3}$, respectively, while those during non-SEP were only 2.9
740 $(\pm 1.4) \text{ ng m}^{-3}$ and $2.6 (\pm 1.6) \text{ ng m}^{-3}$. The average mass ratio of V/Ni during SEP was
741 2.0 in agreement with those reported in ship-influenced PM_1 (Mazzei et al., 2008;Mar
742 et al., 2009), and within the range of 1.9 to 3.5 for domestic HFO (Zhao et al., 2013),
743 while the average ratio of V/Ni (1.1 ± 0.8) during non-SEP was outside the range.
744 Moreover, the major air pollutants emitted from ships include nitrogen oxides (NO_x),
745 sulfur oxides (SO_2), carbon monoxide (CO), hydrocarbons and primary/secondary
746 particles (Becagli et al., 2017;Wu et al., 2021). As demonstrated in Fig. 15d and Fig.
747 S11d, SO_2 concentration was overall higher during SEP (10:00-16:00 on November 11
748 not included); mass loading of NO_2 was higher during the rush hours of non-SEP, but
749 was higher at night (no traffic) during SEP. Mass proportion of sulfate in NR- PM_1
750 during SEP was also higher than that during non-SEP ($22.7 \pm 8.1 \% \text{ vs. } 19.1 \pm 5.3 \%$),
751 but *vice versa* for nitrate ($17.8 (\pm 9.9)\%$ during SEP *vs.* $26.2 (\pm 9.1)\%$ during non-SEP).

752 We further investigated the dependences of $r\text{BCc}$ and NR- PM_1 species on V during
753 SEP and non-SEP, as shown in Fig. S12. Here we only used V since Ni level might be
754 influenced by other emission sources, such as refining industry (Jang et al., 2007;Kim
755 et al., 2014) in urban Shanghai, and during non-SEP, Ni still presented a good
756 correlation with $r\text{BC}$ ($r=0.80$). During SEP, V concentrations (most of them $>4 \text{ ng m}^{-3}$)
757 overall positively correlated with both sulfate and nitrate (except for $r\text{BCc}$ sulfate) (Fig.
758 S12a). Considering that V concentration was independently measured for all fine
759 particles, a better correlation with total NR- PM_1 sulfate than it with $r\text{BCc}$ sulfate is
760 reasonable. Similarly, V-rich particles had positive correlations with traffic-related OA
761 and LO-OOA no matter in $r\text{BCc}$ or in total NR- PM_1 (Fig. S12b). Conversely, during
762 non-SEP, particles with low-V content (most of them $<4 \text{ ng m}^{-3}$) had no clear links with
763 sulfate, nitrate, POA and SOA species (even anti-correlations for $V > 2.5 \text{ ng m}^{-3}$) (Figs.
764 S12c and S12d; a detailed comparison of the correlation coefficients of V with OA
765 factors of $r\text{BCc}$ and NR- PM_1 organics are provided in Table S2). The results above
766 demonstrate that during SEP, chemical properties of the particles (both fresh and aged
767 ones) were clearly linked with ship influences.

768 Previous studies (Ault et al., 2009;Ault et al., 2010;Liu et al., 2017b) have shown

769 that the fresh ship-emitted V-rich particles are typically accompanied by high sulfate
770 contribution, while those aged V-containing ship particles are on the other hand with
771 relatively high nitrate contribution. In order to further probe chemical characteristics
772 and evolution processes of particles influenced by ship emissions, we divided SEP into
773 three episodes (i.e., EP1-EP3) (marked in Fig. 15). During EP1, r_{BC} content was
774 highest (Fig. 15k) and coating was thinnest (Fig. S13i) indicating the particles were
775 relatively fresh, nevertheless the SOA contents were not low (Figs.15j and 15l), OS_C
776 was moderate (Fig. S13j), sulfate portions in NR- PM_{10} and r_{BCc} were both the highest
777 (26.5%), and nitrate portion was the lowest (9.8%) among three episodes (Figs.15i and
778 15k). Such composition is not common in other cases, demonstrating it was a specific
779 period impacted by fresh ship emissions. EP2 had the highest mass loadings of V, gas
780 pollutants as well as the lowest planetary boundary layer (PBL) height (~200 meters)
781 (Fig. S13). Mass contributions of nitrate increased and sulfate decreased, and r_{BC}
782 content decreased from those during EP1, signifying that the particle population likely
783 contained some aged ship-emitted particles. Surely, particle composition during EP2
784 was also influenced by the formation mechanisms of secondary species: EP2 was
785 mostly during nighttime, therefore photochemical formation of sulfate and SOA were
786 weak (as shown in Figs. 15i and 15l, sulfate contribution was only 15.1%, and SOA
787 contribution was only 33.7%), while nitrate formation was enhanced due to the
788 nocturnal process. During EP3, r_{BC} was the lowest, sulfate and V were moderate,
789 nitrate and SOA were highest in both r_{BCc} and NR- PM_{10} , OS_C and R_{BC} were highest in
790 r_{BCc} among the three episodes (Figs. 15i-l and Figs. S13i-k), therefore it was also a
791 period with influence from aged ship-emitted particles; the difference from EP2 is that
792 photochemically formed sulfate and SOA were still significant as EP3 was in the later
793 afternoon and early evening, though heterogeneously formed nitrate played a non-
794 negligible role too (see the increase of RH, increase of nitrate and decrease of O_3
795 concentrations from the beginning of EP3 in Fig. 15).

796

797 **4. Conclusions and implications**

798 We conducted a field measurement during November of 2018 in urban Shanghai,

799 China, focusing on the elucidation of physical and chemical properties of the ambient
800 particles containing *r*BC cores (*r*BCc) by using a laser-only SP-AMS together with a
801 HR-AMS. The campaign-average *r*BCc was $4.6 (\pm 4.4) \mu\text{g m}^{-3}$, occupying $\sim 19.1 (\pm$
802 $4.9) \%$ mass of the total NR-PM₁. The average mass ratio of coating to *r*BC cores (R_{BC})
803 was $\sim 5.0 (\pm 1.7)$, indicating an overall thick coating, compared with the *r*BCc near
804 combustion source. Sulfate was found to preferentially condense on non-*r*BC particles
805 thus led to a low fraction of *r*BCc sulfate to that in NR-PM₁ ($7.4\% \pm 2.2 \%$), while
806 distribution of nitrate between *r*BCc and non-*r*BC particles showed no obvious
807 difference. PMF analysis on *r*BCc and NR-PM₁ OA reveals that cooking-related
808 organics were externally mixed with *r*BC, and a small portion of organics from biomass
809 burning was only present in *r*BCc; the traffic-related OA species, however, was in a
810 large part internally mixed with *r*BC.

811 A regression algorithm was applied to deconvolute the size distributions of
812 individual *r*BCc OA factors, and results show that small *r*BCc particles were
813 predominantly generated from traffic, and such particles could grow bigger because of
814 condensation of secondary inorganic and organic components, resulting in thick coating.
815 Investigation on diurnal patterns of the *r*BCc species reveals that sulfate and two SOA
816 factors (LO-OOA_{*r*BC} and MO-OOA_{*r*BC}) were generated mainly through daytime
817 photochemical oxidation; nitrate, on the other hand, was governed mainly by the
818 nocturnal N₂O₅ hydrolysis. Partial SOA was found to be produced from in-situ
819 photochemical conversion from traffic-related POA. An average coating time (ACT)
820 was proposed to roughly estimate the timescales for the secondary species to coat on
821 *r*BC, and the ACT of sulfate, LO-OOA_{*r*BC}, MO-OOA_{*r*BC} and nitrate were approximately
822 7, 5, 7 and 19 hours, respectively.

823 Moreover, the size-resolved hygroscopicity parameters of *r*BCc particles ($\kappa_{r\text{BCc}}$)
824 and the coating material (κ_{CT}) were obtained based on the elucidated composition of
825 *r*BCc particles. The fitted equations are: $\kappa_{r\text{BCc}}(x) = 0.29 - 0.14 \times \exp(-0.006 \times x)$ and
826 $\kappa_{\text{CT}}(x) = 0.35 - 0.09 \times \exp(-0.003 \times x)$ (x is $D_{r\text{BCc}}$, $150 < x < 1000$ nm). The minimums
827 of both $\kappa_{r\text{BCc}}$ and κ_{CT} were at ~ 150 nm due to the abundances of hydrophobic *r*BC cores
828 and traffic-related HOA at this size. Under critical supersaturations (SS_{C}) of 0.1% and

829 0.2%, the D_{50} values were estimated to be 166 (± 16) and 110 (± 5) nm, and the
830 activated number fractions of $rBCc$ particles were 16 (± 3)% and 59 (± 4)%,
831 respectively. Finally, a typical case with influence from ship emissions was investigated.
832 During this period, the $rBCc$ particles were enriched in V (6.3 ± 3.1 ng m⁻³), with a
833 V/Ni ratio of 2.0 (± 0.6), and various secondary formation processes affect the ship-
834 emitted particles at different times of the day.

835 In summary, the findings from this comprehensive study on $rBCc$ provide rich
836 information regarding the various primary sources and secondary formation pathways
837 of species coated on rBC , as well as the features of distributions of those species
838 between rBC and non- rBC particles. In particular, different types of $rBCc$ particles
839 from diesel and gasoline vehicle emissions were resolved and elucidation of their
840 properties are useful to their future effective control. Understanding of the formation
841 processes and coating timescales of secondary components is helpful to understand the
842 impact of ambient BC particles too. At last, the parameterized relationship of
843 hygroscopicity with size distribution is useful for modelling the climate effect of rBC
844 (alternation of cloud properties).

845

846 *Data availability.* The data in this study are available from the authors upon request
847 (caxinra@163.com).

848

849 *Supplement.* The supplement related to this article is available online at: XXX

850

851 *Author contributions.* SJC, DDH, YZW, JFW, FZS, and XLG conducted the field
852 measurement. SJC, DDH, YZW, JFW, and JKX analyzed the data. YJZ, HLW, CH and
853 HL reviewed the paper and provide useful suggestions. SJC and XLG wrote the paper.

854

855 *Competing interests.* The contact author has declared that neither they nor their co-
856 authors have any competing interests.

857

858 *Disclaimer.* Publisher's note: Copernicus Publications remains neutral with regard to

859 jurisdictional claims in published maps and institutional affiliations.

860

861 *Acknowledgements.* We sincerely thank SAES to provide data of gaseous pollutants and
862 particulate V and Ni, and the logistic help during the campaign.

863

864 *Financial support.* This work has been supported by the National Natural Science
865 Foundation of China (42021004, 21976093 and 21777073).

866

867 *Review statement.* This paper was XXX.

868

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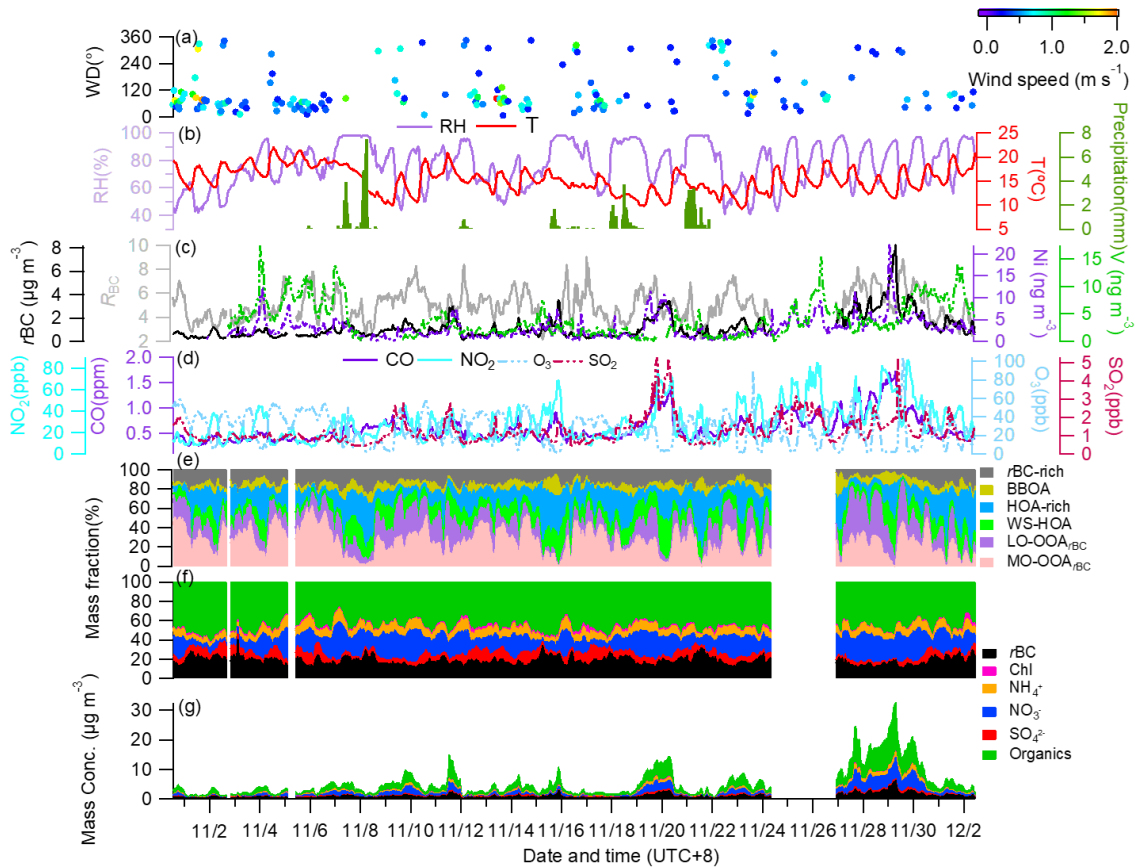
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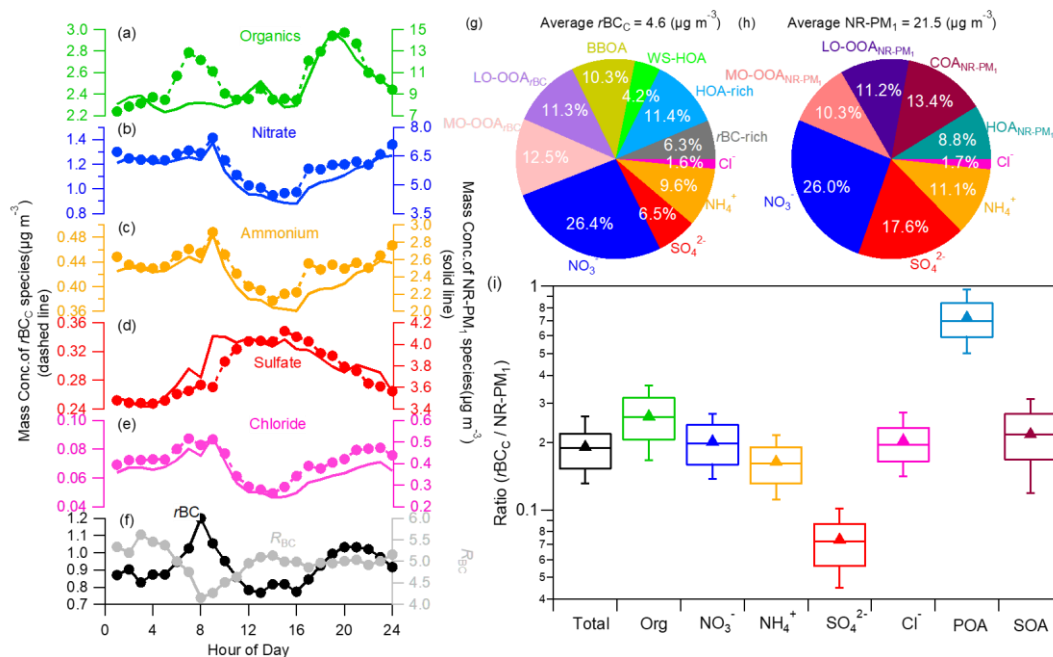
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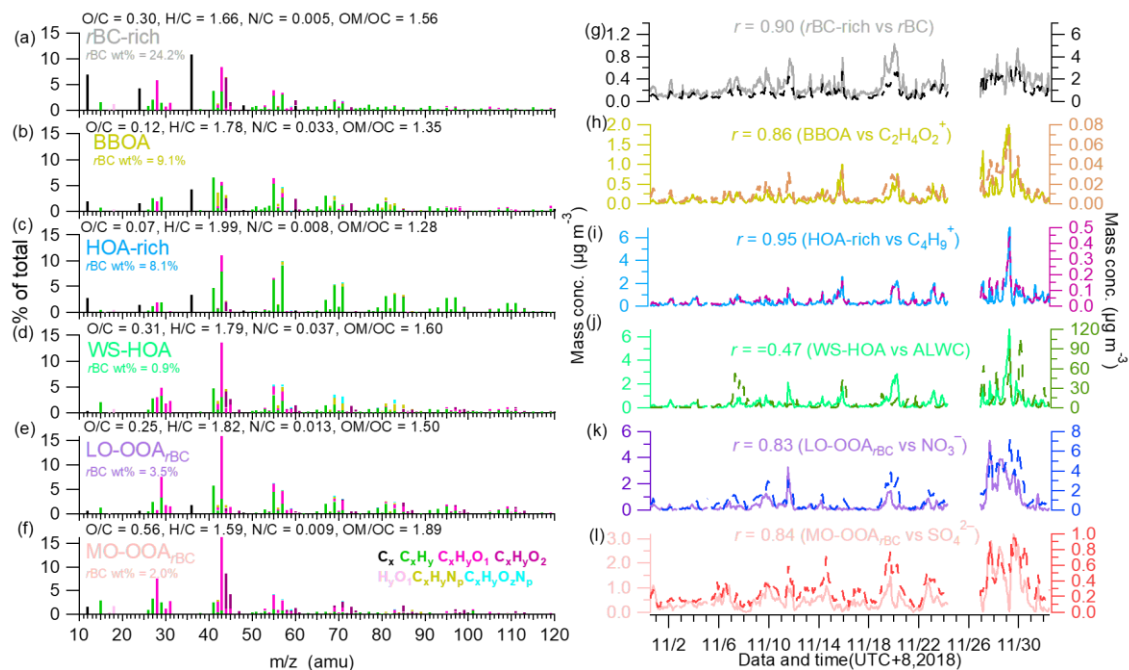
1284

1285 Figure 1. Time series of (a) wind direction (WD) colored by wind speed (WS), (b) air
 1286 temperature (T), relative humidity (RH) and precipitation, (c) mass concentrations of
 1287 *r*BC, Ni, V, and *R*_{BC} (mass ratio of all coating species to *r*BC), (d) mass concentrations
 1288 of gas pollutants of CO, NO₂, O₃ and SO₂, (e) mass fractions (%) of different OA factors
 1289 to the total *r*BCc OA, (f) mass fractions (%) of different components to the total *r*BCc
 1290 mass, and (g) mass concentrations of stacked *r*BCc components.



1291

1292 Figure 2. Diurnal cycles of mass concentrations of (a–e) r_{BCc} and NR-PM₁ species
 1293 (organics, nitrate, ammonium, sulfate, and chloride), and (f) r_{BC} and R_{BC} . Campaign-
 1294 average chemical composition of r_{BCc} (g) and NR-PM₁ (h). (i) Mass ratios of species
 1295 in r_{BCc} to those in NR-PM₁ (the whiskers above and below the boxes mark the 90%
 1296 and 10% percentiles, respectively; the upper and lower edge of the boxes represent the
 1297 75% and 25% percentiles, respectively; and the lines and triangles inside the boxes
 1298 denote the median and mean values, respectively; SOA represents $([LO-OOA_{r_{BCc}}] +$
 1299 $[MO-OOA_{r_{BCc}}])/([LO-OOA_{NR-PM_1}] + [LV-OOA_{NR-PM_1}])$, and POA represents $([r_{BC-rich}$
 1300 $+ HOA-rich + BBOA + WS-HOA])/HOA_{NR-PM_1}$).



1301

1302 Figure 3. High resolution mass spectra of (a) *r*BC-rich, (b) BBOA, (c) HOA-rich, (d)

1303 WS-HOA, (e) LO-OOA_{*r*BC}, and (f) MO-OOA_{*r*BC}. (g-l) Time series of corresponding

1304 factors, their tracers (*r*BC, C₂H₄O₂⁺, C₄H₉⁺, ALWC, nitrate and sulfate) as well as the

1305 correlation coefficients (ALWC refers to aerosol liquid water content, which was

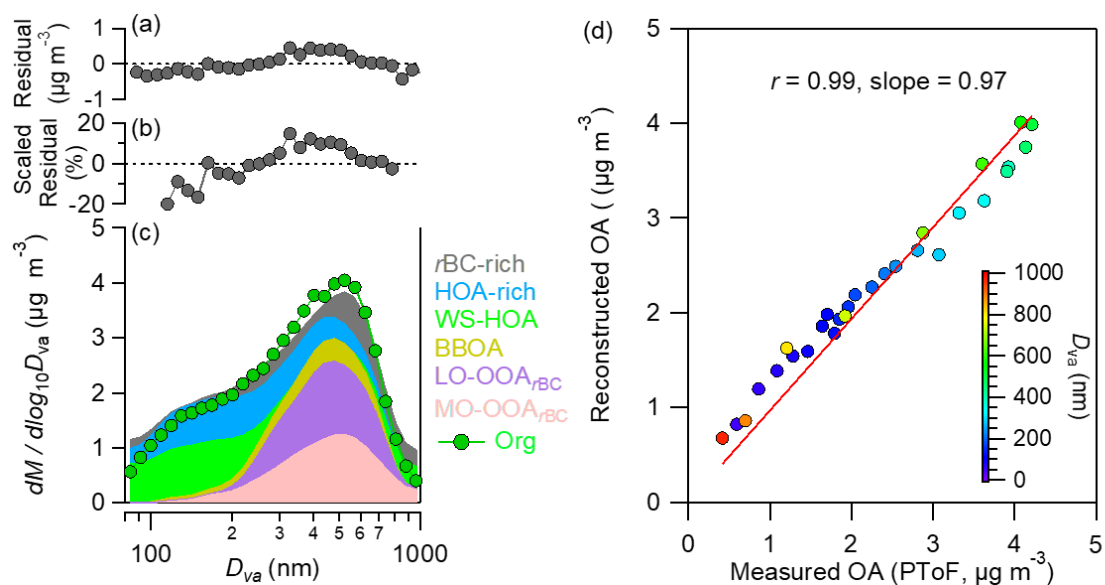
1306 estimated by using the model II of extended aerosol inorganic model (E-AIM II)(Clegg

1307 et al., 1998); calculation details and results at different RH values are described in the

1308 caption and shown in Fig. S5)

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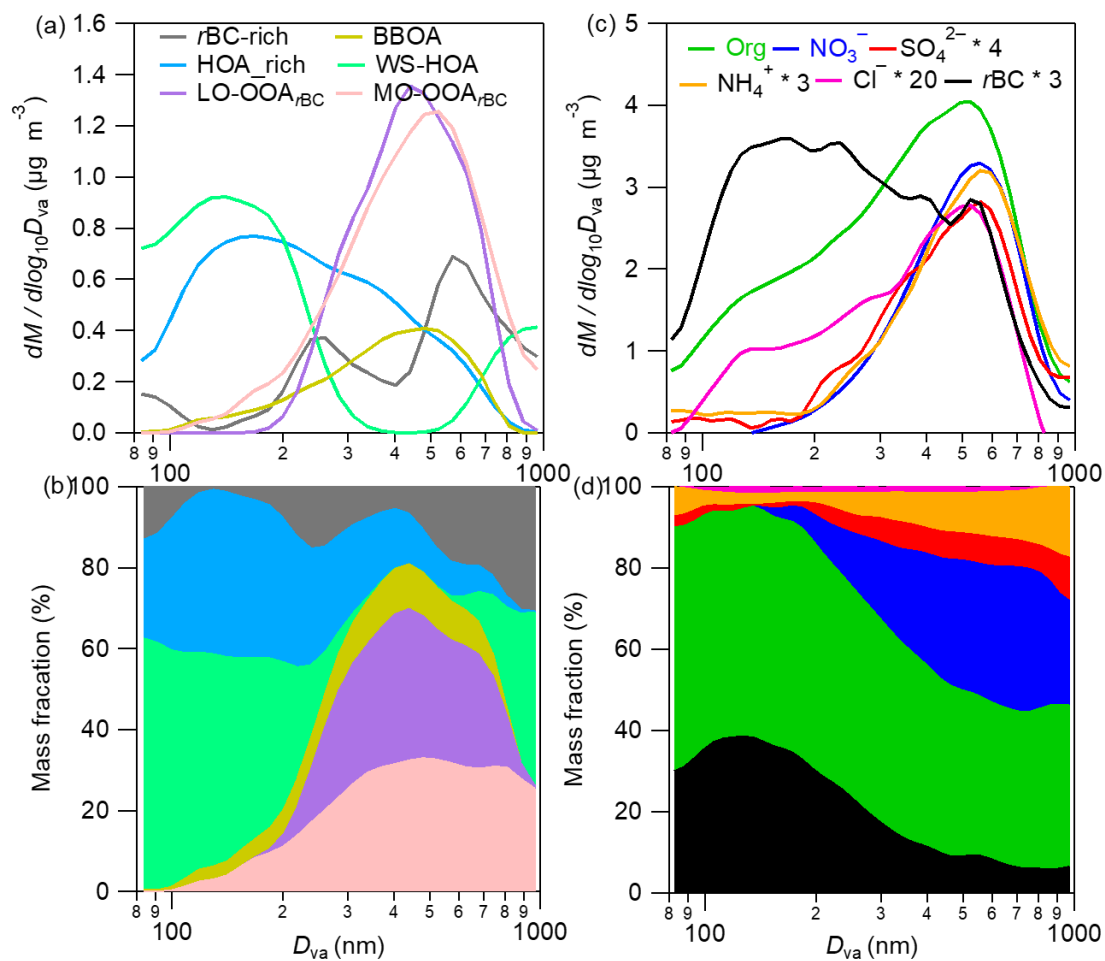
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1312 Figure 4. Summary of key diagnostic plots of derivation of size distributions of
1313 individual r BCc OA factors. (a) Absolute and (b) relative residuals between the
1314 reconstructed and measured OA mass concentrations in different size bins. (c)
1315 stacked size distributions of the six OA factors compared to the size distributions of total OA.
1316 (d) Reconstructed OA mass concentrations compared to the measured values for
1317 different size bins (80-1000 nm).

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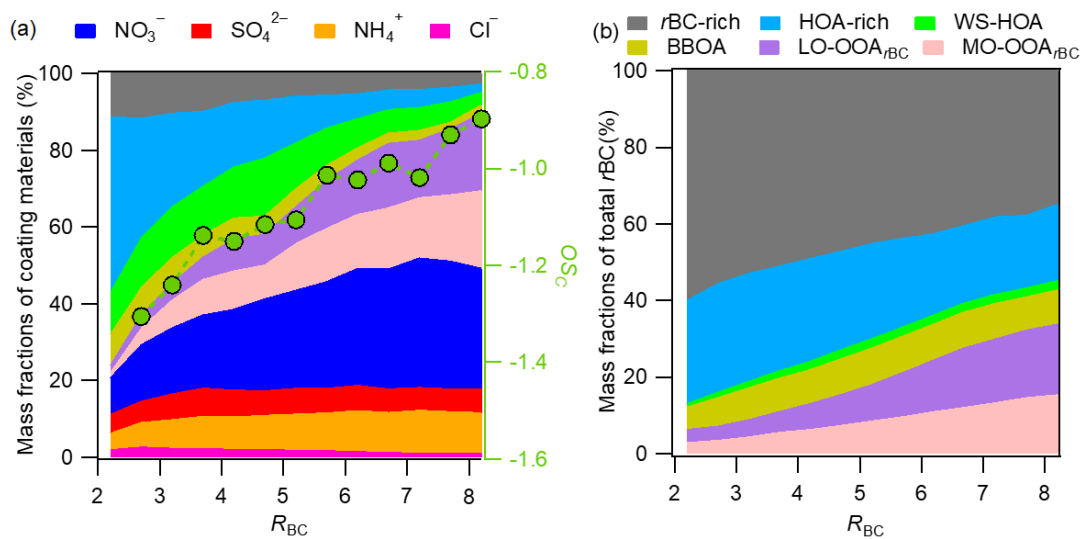


1319

1320 Figure 5. Campaign-average size distributions of six *r*BCc OA factors (a) and individual
 1321 *r*BCc components (b), and corresponding mass contributions of the six factors to the
 1322 total *r*BCc OA (c), and the major components to the total *r*BCc (d) at different sizes
 1323 (80-1000 nm).

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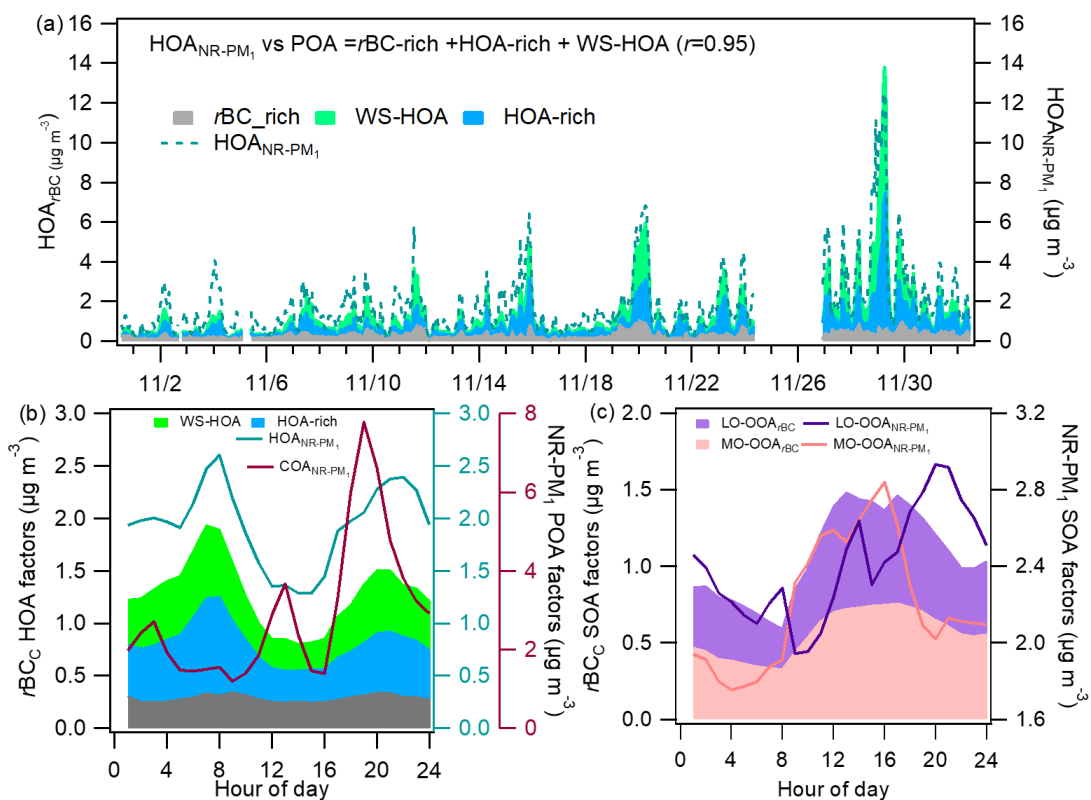


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1327 Figure 6. (a) Variations of mass fractions of the major $rBCc$ components against R_{BC} .

1328 (b) Variations of mass contributions of individual $rBCc$ OA factors to rBC against R_{BC} .

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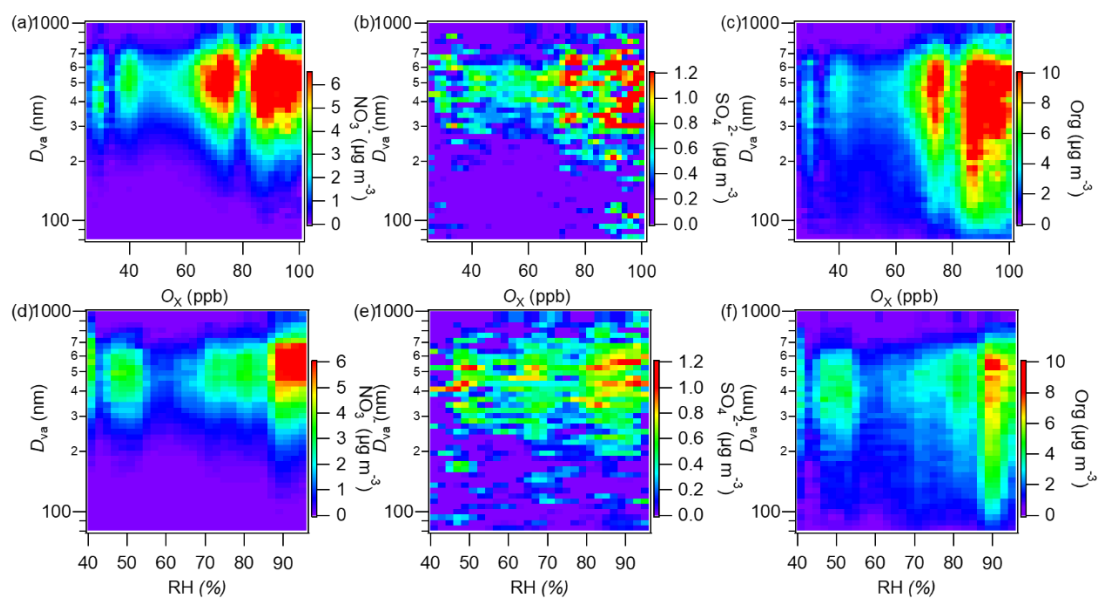


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1331 Figure 7. (a) Time series of stacked three rBC_c POA factors (i.e., rBC -rich, HOA-rich,
 1332 and WS-HOA) and HOA_{NR-PM_1} . Comparisons of the diurnal patterns of different POA
 1333 factors (b) and SOA factors (c) of rBC_c and $NR-PM_1$.

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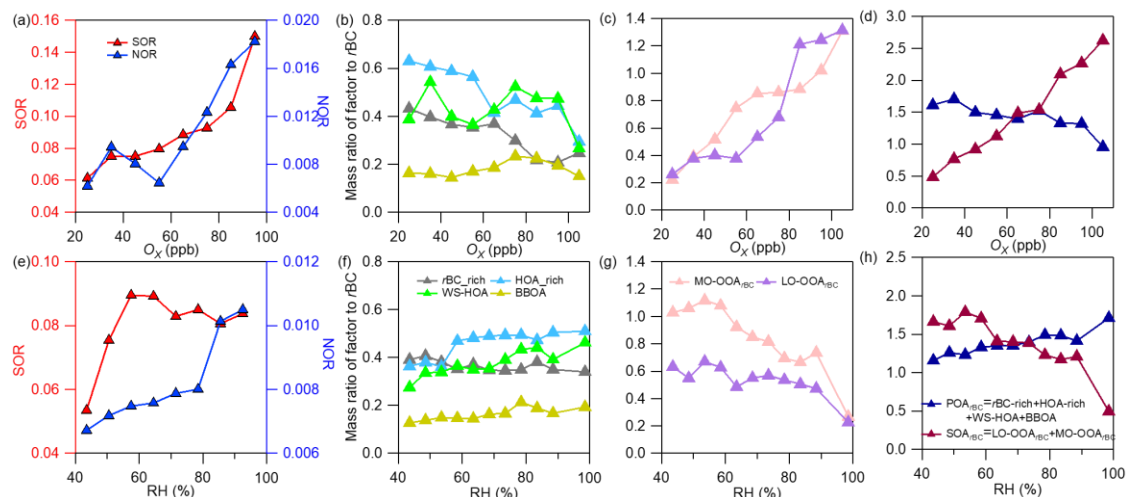
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1337 Figure 8. Image plots of size distributions of rBCc nitrate, sulfate, organics as a function
1338 of (a-c) O_x and (d-f) RH, respectively (color represents its concentration).

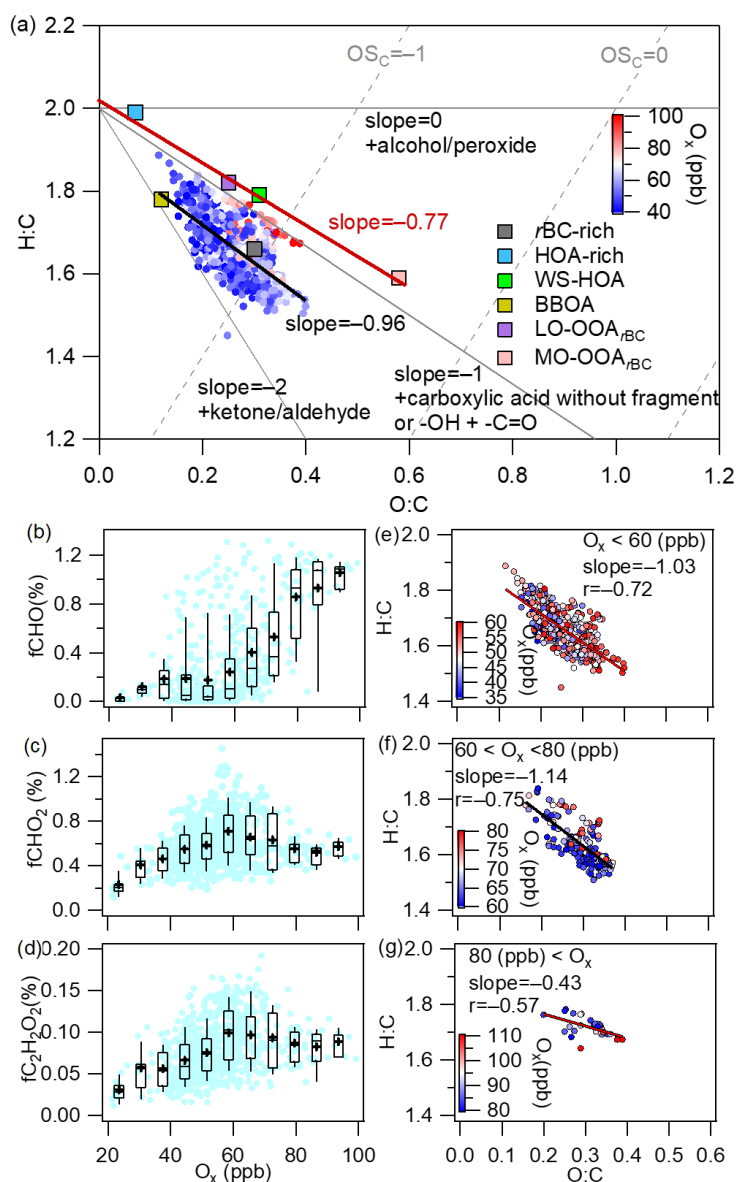
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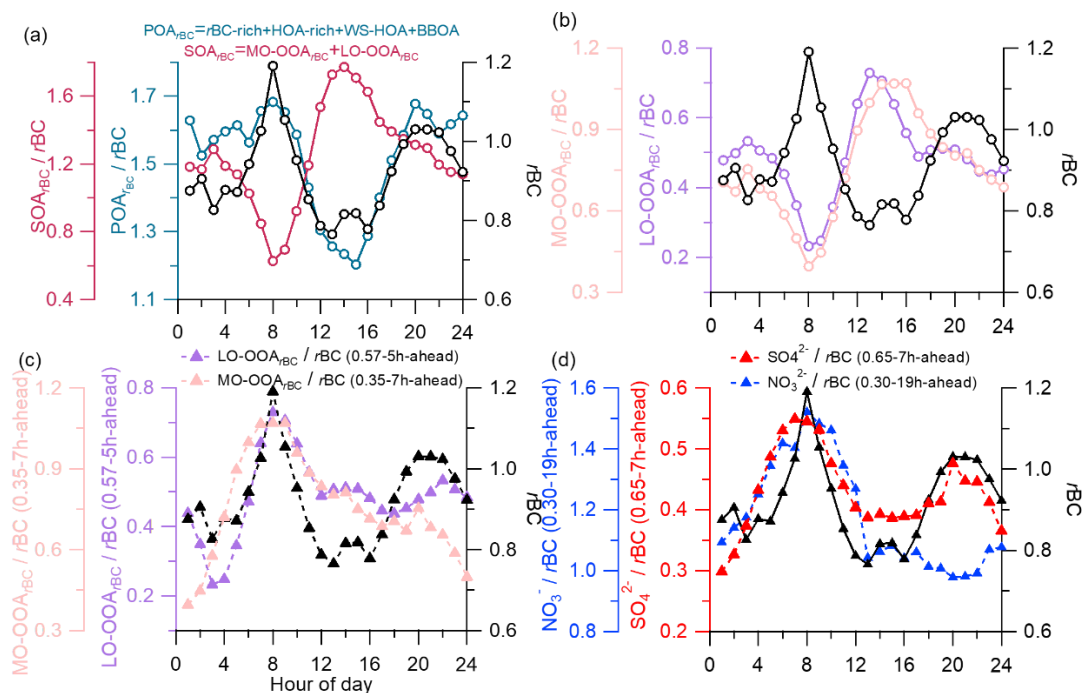
1341 Figure 9. Variations of nitrogen oxidation ratio (NOR) and sulfur oxidation ratio, mass
 1342 ratios of different POA factors, SOA factors and total POA and SOA to *rBC* against O_X
 1343 (a-d) and RH (e-h) ($NOR = \frac{nNO_3^-}{nNO_3^- + nNO_2 + nNO}$ and $SOR = \frac{nSO_4^{2-}}{nSO_4^{2-} + nSO_2}$,
 1344 where nNO_3^- , nSO_4^{2-} , nNO_2 , nNO and nSO_2 are the molar concentrations of particle-
 1345 phase sulfate, nitrate, gaseous NO_2 , NO and SO_2 , respectively).

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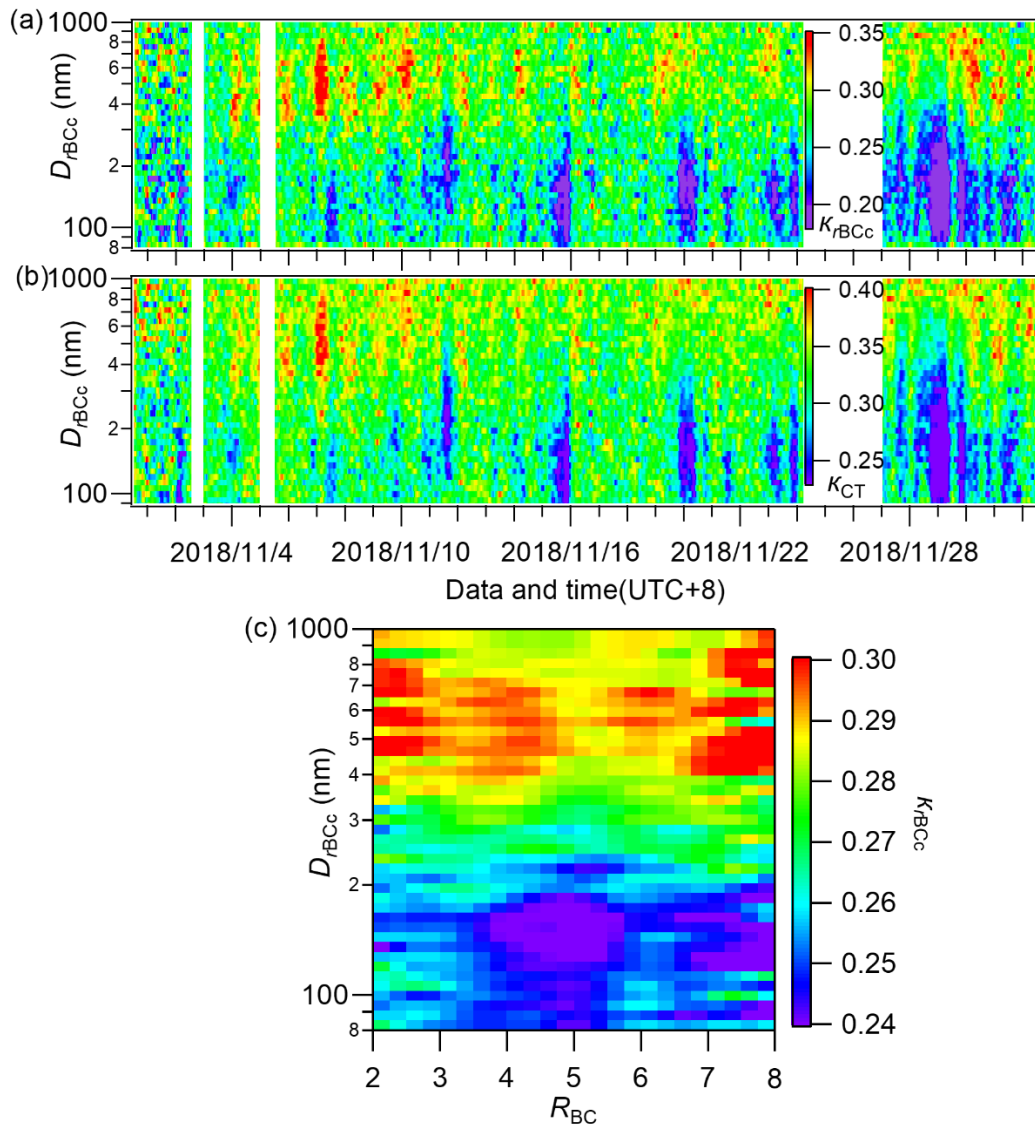
1348 Figure 10. (a) Van Krevelen diagram of H/C versus O/C ratios for all rBCc OA and the
 1349 six factors colored by O_x concentrations (the black line represents the linearly fitted
 1350 line of all OA data, and the red line is the fitted line of the four OA factors). (b-d) Mass
 1351 fractions of selected oxygenated ion fragments as a function of O_x (meanings of the
 1352 boxes are the same as those described in Fig. 2). (e-g) Scatter plots of H/C versus O/C
 1353 ratios under different O_x levels (data are colored by O_x concentrations).



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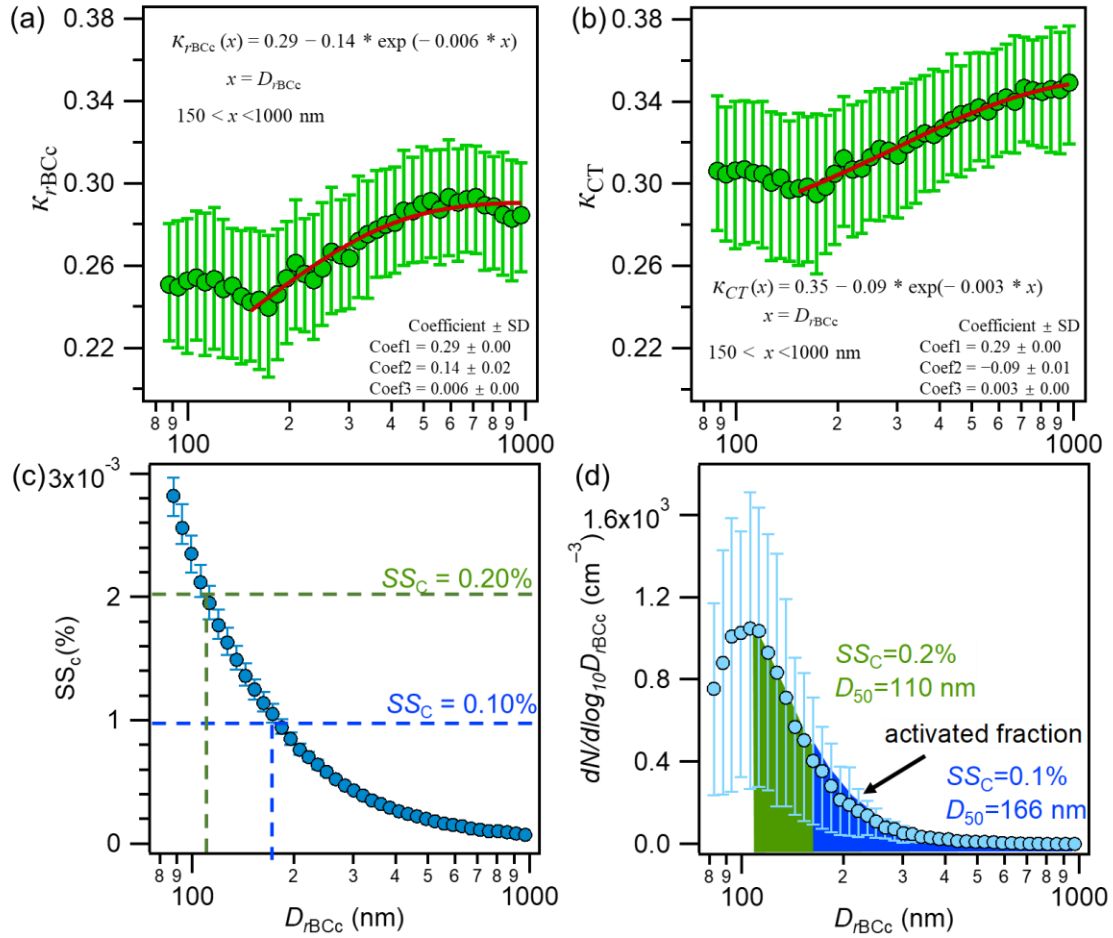
1355 Figure 11. Campaign-average diurnal patterns of (a) rBC , POA_{rBC}/rBC and
 1356 SOA_{rBC}/rBC , and (b) rBC , $MO\text{-}OOA_{rBC}/rBC$ and $LO\text{-}OOA_{rBC}/rBC$. Adjusted diurnal
 1357 patterns by the average coating time (ACT) for (c) $LO\text{-}OOA_{rBC}/rBC$, $MO\text{-}OOA_{rBC}/rBC$,
 1358 and (d) SO_4^{2-}/rBC , NO_3^-/rBC .

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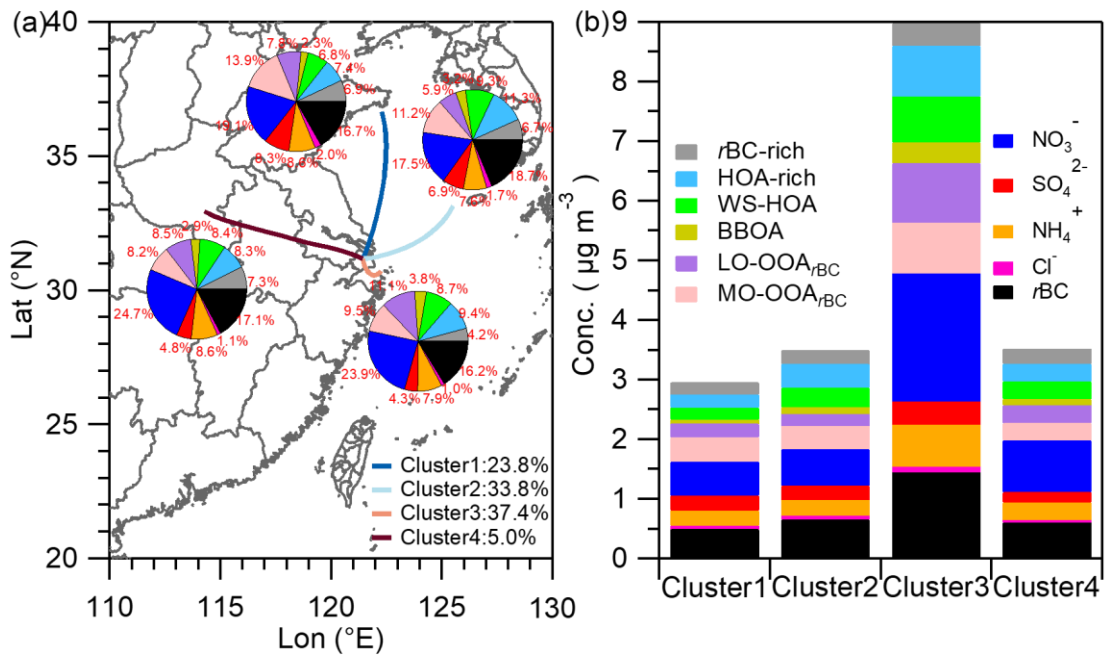
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1361 Figure 12. Image plots of size-resolved hygroscopicity parameters of (a) $rBCc$ (κ_{rBCc}),
 1362 (b) its coating materials (κ_{CT}) during the whole campaign, and (c) the campaign-average
 1363 size-resolved κ_{rBCc} at different R_{BC} .



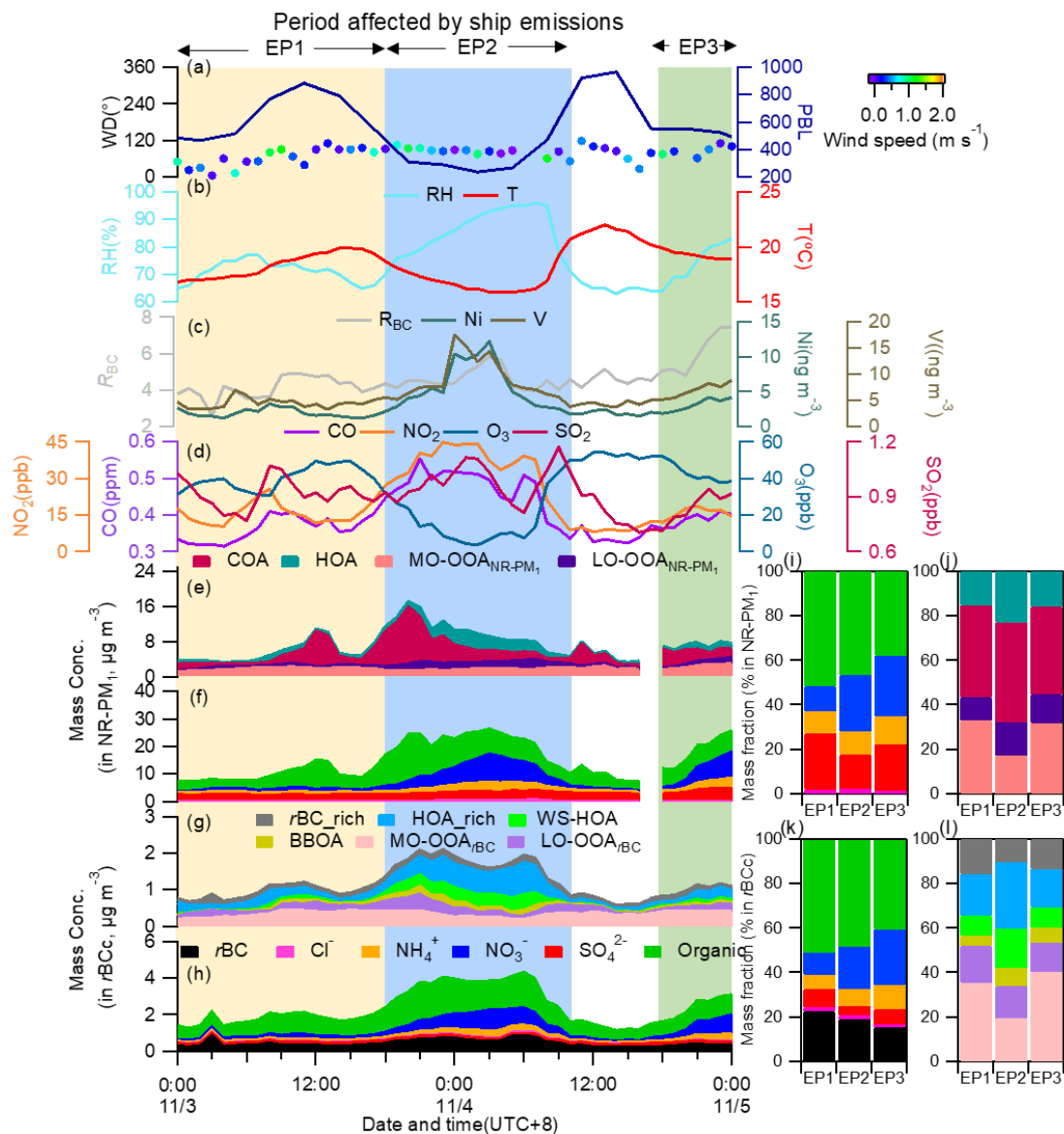
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1365 Figure 13. Campaign-average size-resolved hygroscopic parameters for $rBCc$ (κ_{rBCc})
 1366 and (a) for its coatings (κ_{CT}) (b) (the red lines are exponential fits of the data of 150-
 1367 1000 nm). (c) Campaign-average size-resolved critical supersaturation (SS_C), and (d)
 1368 the predicted activated fraction of $rBCc$ number concentration based on D_{50} at SS_C of
 1369 0.1% (166 nm) and 0.2% (110 nm) (the solid circles are mean values, the upper and
 1370 lower lines are the 75th and 25th percentiles, respectively).



1371

1372 Figure 14. (a) Four clusters of 24-h backward trajectories (at altitude of 500 m) analyzed
 1373 by NOAA HYSPLIT model (<http://www.arl.noaa.gov/ready/hysplit4.html>) embedded
 1374 in Zefir(Petit et al., 2017), with the pie chart showing the average rBC_c chemical
 1375 compositions in each cluster. (b) Stacked mass concentrations of the rBC_c components
 1376 of the four clusters.



1377
 1378 Figure 15. Time series of (a) wind direction (WD) colored by wind speed (WS),
 1379 planetary boundary layer (PBL) height, (b) relative humidity (RH) and temperature (T),
 1380 (c) mass concentrations of particle-phase Ni and V, and R_{BC} , (d) mass concentrations of
 1381 CO, NO₂, O₃, SO₂, stacked concentrations of (e) NR-PM₁ OA factors, (f) NR-PM₁
 1382 species, (g) $rBCc$ OA factors, and (h) $rBCc$ components during the ship emission period
 1383 (SEP). Mass contributions of different components to NR-PM₁ (i), different OA factors
 1384 to total NR-PM₁ OA (j), different components to $rBCc$ (k), and different OA factors to
 1385 total $rBCc$ OA (l) for the three episodes (EP1, EP2 and EP3).
 1386