

1 Evolution of source attributed organic aerosols and gases in a 2 megacity of central China

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18 **Abstract.** The secondary production of the oxygenated organic aerosol (OOA) impacts air quality, climate and human
19 health. The importance of various sources in contributing to the OOA loading and associated different ageing mechanisms
20 remains to be elucidated. Here we present concurrent observation and factorization analysis on the mass spectra of organic
21 aerosol (OA) by a high-resolution aerosol mass spectrometer and volatile organic compounds (VOCs) by a proton-transfer-
22 reaction mass spectrometer in Wuhan, a megacity in central China during autumn. The full mass spectra of organics
23 with two principle anthropogenic sources were identified as the traffic and cooking sources, for their primary emission
24 profiles in aerosol and gas phases, the evolutions, and their respective roles in producing OOA and secondary VOCs.
25 Primary emissions in gas and aerosol phases both contributed to the production of OOA. The photooxidation of traffic
26 sources from morning rush-hour caused 2.5-folds increase of OOA mass in a higher oxidation state (oxygen-to-carbon, O/C
27 = 0.72), coproducing gas-phase carboxylic acids; while at night, cooking aerosols and VOCs (particularly acrolein and
28 hexanal) importantly caused the nocturnal formation of oxygenated intermediate VOCs, increasing OOA mass by a factor of
29 1.7 (O/C = 0.42). The daytime and nighttime formation of secondary aerosols as contributed by different sources, was found
30 to be modulated by solar radiation and air moisture respectively. The environmental policy should therefore consider the
31 primary emissions and their respective ageing mechanisms influenced by meteorological conditions.

32 **1 Introduction**

33 The transformation between gas and aerosol phase of organic species produces secondary organic aerosols (SOA)
34 (Kroll and Seinfeld, 2008; Hallquist et al., 2009; Slowik et al., 2010), which forms important global budget of aerosol
35 loadings (Heald et al., 2005; Zhang et al., 2007), exerting climate (Poschl, 2005; Seinfeld et al., 2016) and environmental
36 impacts (Von Schneidemesser et al., 2011; Huang et al., 2014). Given the large complexities of organic species, the yields
37 (Goldstein and Galbally, 2007; Ortiz-Montalvo et al., 2014) and production rates (Sareen et al., 2013; Jokinen et al., 2015) of
38 SOA from gas precursors are influenced by the diversities of source profiles (emission mass percentages among species)
39 (Lin et al., 2012; Shrivastava et al., 2015) and environmental factors (such as radiation, air moisture and ambient
40 temperature) (Clark et al., 2016; Li et al., 2021b; Wang et al., 2021). This raises challenges for source-oriented
41 environmental policy making, until explicit understanding on the formation mechanism of SOA from different sources.

42 The application of detailed mass spectra of organic aerosols allows the online source attribution of organic aerosol (OA)
43 (Canagaratna et al., 2007; Ng et al., 2011a), based on the factorization analysis on the mass spectra of OA which groups the
44 covaried species from certain sources (Ulbrich et al., 2009). This factorization technique allows identification of primary
45 sources and aged secondary sources at a receptor measurement site. The complex aerosol sources in urban environment have
46 been commonly recognized as traffic sources mainly from vehicles on road (Zhu et al., 2021), biomass burning from open or
47 closed combustion (Adler et al., 2011), coal combustion normally in cold season for heating (Xu et al., 2019) and more
48 localized cooking sources (Allan et al., 2010; Zhang et al., 2020). The secondary sources are oxygenated OA which may be
49 from oxidizing volatile organic compounds (VOCs) following condensation (Donahue et al., 2006), or heterogenous
50 oxidation occurring on particle phase (Claeys et al., 2004). Among these sources, SOA has been found to be the main
51 contributor of OA mass loading (41–69 %) in urban environment of East Asia (Sun et al., 2014; Hu et al., 2017), especially
52 in warm season when primary emissions were low, along with high ambient temperature and more intensive chemical
53 reactions (Hu et al., 2016). The formation of SOA from VOCs may experience a few reaction generations (Kroll and
54 Seinfeld, 2008) and could interact with other sources of species during the process (Shrivastava et al., 2017), hereby
55 complicating the goal of identifying the key precursors in contributing to the consequent SOA. In addition, some primary
56 gases already have somehow low volatility and may not require a long reaction chain to become condensable, such as some
57 primarily emitted intermediate volatility organic compounds (IVOCs) may substantially contribute to the SOA (Robinson et
58 al., 2007; Huang et al., 2021). An understanding of source profiles of primary emissions in both gas and aerosol phases is
59 therefore important to rule out the source-dependant production of SOA. The above necessitates the concurrent investigation
60 on the compositions of gases and aerosols at a receptor site, along with their evolution and interaction, in order to elucidate
61 the role of each source in contributing to SOA.

62 In this study, we performed online continuous measurements on the detailed mass spectra of organics concurrently on
63 aerosol and gas phases, in a typical anthropogenically polluted region in central China, where such data had been rarely

64 reported. Factorization analysis is performed on the mass spectra of organics in both aerosol and gas phases, to investigate
65 the source-oriented gas-aerosol evolution and SOA formation in this region.

66 **2 Materials and methods**

67 **2.1 Sampling site**

68 The field experiment was performed in the campus of the China University of Geosciences (Wuhan) (114.40° E, 30.52°
69 N) during Oct.–Nov. 2019 (Hu et al., 2021). The site represents a typical region of mixed residential/traffic sources (Fig. S1).
70 Due to the preparation and hosting of the 7th CISM Military World Games during the experimental period, the government
71 implemented strict emission reduction measures, particularly for the industrial sources and heavy vehicle emissions in the
72 main roads. The more localized pollution sources, such as traffic emission from smaller sizes of vehicles and cooking
73 sources dominated the pollution in this study, besides those transported from surrounding regions during some heavy
74 pollution events (Zheng et al., 2019; Zheng et al., 2020). The HYSPLIT model (Draxler and Hess, 1997) with a 3-hourly 1°
75 × 1° wind field from the GDAS reanalysis products was used to obtain the 36-h backward trajectories initializing from the
76 location of the experiment site. The cluster analysis was performed to categorize the trajectories into three groups by
77 minimizing the differences in each group and maximizing the differences among groups (Moody and Galloway, 1988).

78 **2.2 Instrumentation**

79 **2.2.1 Measurements of mass spectra of aerosols and gases**

80 The mass concentration and chemical composition of NR-PM₁ were measured by a HR-ToF-AMS (Aerodyne Research
81 Inc., USA), including organics, nitrate (NO₃⁻), sulphate (SO₄²⁻), chloride (Cl⁻), and ammonium (NH₄⁺). The detailed can be
82 found elsewhere (Decarlo et al., 2006). Briefly, the aerosols are dried using a diffusion dryer before entering the AMS and
83 through a critical orifice into a narrow beam via an aerodynamic lens. The aerosol size is determined using the flight time of
84 particles to the thermal vaporization and ionization chamber. Then the aerosols are successively vaporized by a heated
85 surface (~600 °C), ionized by electron ionization (EI, 70 eV), and detected by a mass spectrometer detector. During this field
86 observation, the HR-ToF-AMS was operated under V mode with high sensitivity. The composition-dependent collection
87 efficiencies were applied (Middlebrook et al., 2012), and the ionization efficiency was calibrated using 300 nm pure
88 ammonium nitrate (Jayne et al., 2000). Elemental analysis (EA) was also executed using the “Improved Ambient” method
89 (Canagaratna et al., 2015) to obtain the hydrogen-to-carbon ratio (H/C), oxygen-to-carbon ratio (O/C), and nitrogen-to-
90 carbon ratio (N/C).

91 The Proton-Transfer-Reactor Time-of-Flight Mass Spectrometer (PTR-ToF-MS 8000, Ionicon Analytik GmbH
92 Innsbruck, Austria) was deployed to quantify VOCs in this research. The operating and calibration of the PTR followed the
93 routine described previously (Cappellin et al., 2012; Bruns et al., 2016). Briefly, the PTR was operated with hydronium ions

94 (H_3O^+) as the reagent and with a drift tube pressure of 2.2 mbar, voltage of 600 V and temperature of 60 °C. The ratio of the
95 electric field (E) and the density of the buffer gas (N) in the drift tube, which dictates the ion drift velocity in the drift tube, is
96 135 Td. MS transmission function was performed using a mixture of VOCs (formaldehyde, methanol, acetonitrile,
97 acetaldehyde, acetone, isoprene, methyl ethyl ketone, benzene, toluene, styrene, benzaldehyde, ethylbenzene, 1,3,5-
98 trimethylbenzene). Mass calibration was done using H_3O^+ (m/z 21.0226), $\text{CH}_3\text{COCH}_4^+$ (m/z 59.0490) and monoterpenes (m/z
99 137.1290). The shift in m/z is minor which ensures that the mass calibration was sufficient for all compounds. The
100 background measurements were performed using a dry zero air cylinder. The measurement error is described in Text S2. A
101 separate reaction rate constant is applied to convert the ion signal into concentration (Bruns et al., 2016), or the default
102 reaction rate constant $2 \times 10^{-9} \text{ cm}^3 \text{ s}^{-1}$ can be applied to all other ions (Wang et al., 2020b). The vapor saturation concentration
103 (equilibrium vapor pressures) (C^*) of each VOC compound at 25 °C is estimated using the parameterization based on
104 elemental ratio and molecular weight (Fig. S2) (Li et al., 2016). A $\log C^*$ (in $\mu\text{g m}^{-3}$) lower than 6.5 is deemed to be
105 intermediate VOC (IVOC). According to Fig. S2, VOC species (with $m/z > 120$) are mainly identified to be IVOCs in this
106 study, thus the fraction of these larger molecular weight (MW) VOCs ($m/z > 120$) is used to evaluate the potential influence
107 of IVOCs.

108 2.2.2 PMF analysis on the mass spectra of OA and VOCs

109 Positive matrix factorization (PMF) (Paatero and Tapper, 1994) was performed on the high-resolution organic mass
110 spectral matrix of OA (Ulbrich et al., 2009; Decarlo et al., 2010). In this work, $m/z > 120$ and isotopic ions were excluded in
111 PMF analysis due to the limited mass resolution and low contributions to OA loading (~5 %). After a careful evaluation of
112 the mass spectral profiles and correlations with time series of tracers, diurnal variations, four OA factors from total OA were
113 identified with $f_{\text{peak}} = 0$, including hydrocarbon-related OA (HOA), cooking OA (COA), low-oxidized oxygenated OA
114 (LO-OOA, OOA1), and more-oxidized oxygenated OA (MO-OOA, OOA2). The detail of PMF diagnostic was summarized
115 in Text S1 and Fig. S3.

116 The EPA PMF 5.0 model (Paatero and Tapper, 1994) was used for the source apportionment of VOC species. The 109
117 VOCs were used for the PMF analysis. The uncertainties from each sample were determined according to the method
118 detection limit (MDL) and the error fraction (%). The detail of PMF diagnostic was summarized in Text S2 and Fig. S4. Five
119 factors were ultimately selected, and the Q/Q_{exp} ratio was 0.96 on average (Fig. S4). The rotation ambiguity was explored by
120 varying the f_{peak} values from -3 to +3, and the results with $f_{\text{peak}} = 0$ were selected for the lowest dQ (robust), indicating the
121 stability of the PMF solution (Zhou et al., 2019). Most of the residuals are distributed normally, ranging from -3 and +3,
122 suggesting the model fits the data well.

123 2.2.3 Other measurements

124 The concentration of BC particles was measured by a single particle soot photometer (SP2, DMT Inc.). The operation
125 and data analysis procedures of the SP2 have been described elsewhere (Schwarz et al., 2008; Liu et al., 2010). The SP2

126 incandescence signal was calibrated for refractory BC (rBC) mass using the Aquadag black carbon particle standards
127 (Acheson Inc., USA) and corrected for ambient rBC with a factor of 0.75 (Laborde et al., 2012).

128 The size distribution and number concentration of aerosols with an electrical mobility diameter from 12 to 552 nm were
129 also measured by a scanning mobility particle sizer (SMPS, Model 3081, impactor 50 % cut off at 0.677 μm ; CPC model
130 3775 at a flow rate 0.3 L min^{-1}). PM_{10} mass concentration calculated based on the volume concentration measured by the
131 SMPS agreed well with that from the sum of compositions by the HR-ToF-AMS and SP2 ($r = 0.71$, slope = 0.90).

132 The gas-phase species including NO_x , O_3 , and CO were measured in real time by a series of Wuhan Tianhong analyzers
133 (TH-2001H, TH-2003H, TH-2004H, respectively). These instruments were calibrated periodically with the corresponding
134 standard gas to ensure the accuracy of the observation data. In addition, the meteorological parameters including temperature
135 (T), relative humidity (RH), and wind speed and direction were recorded by an automatic weather station.

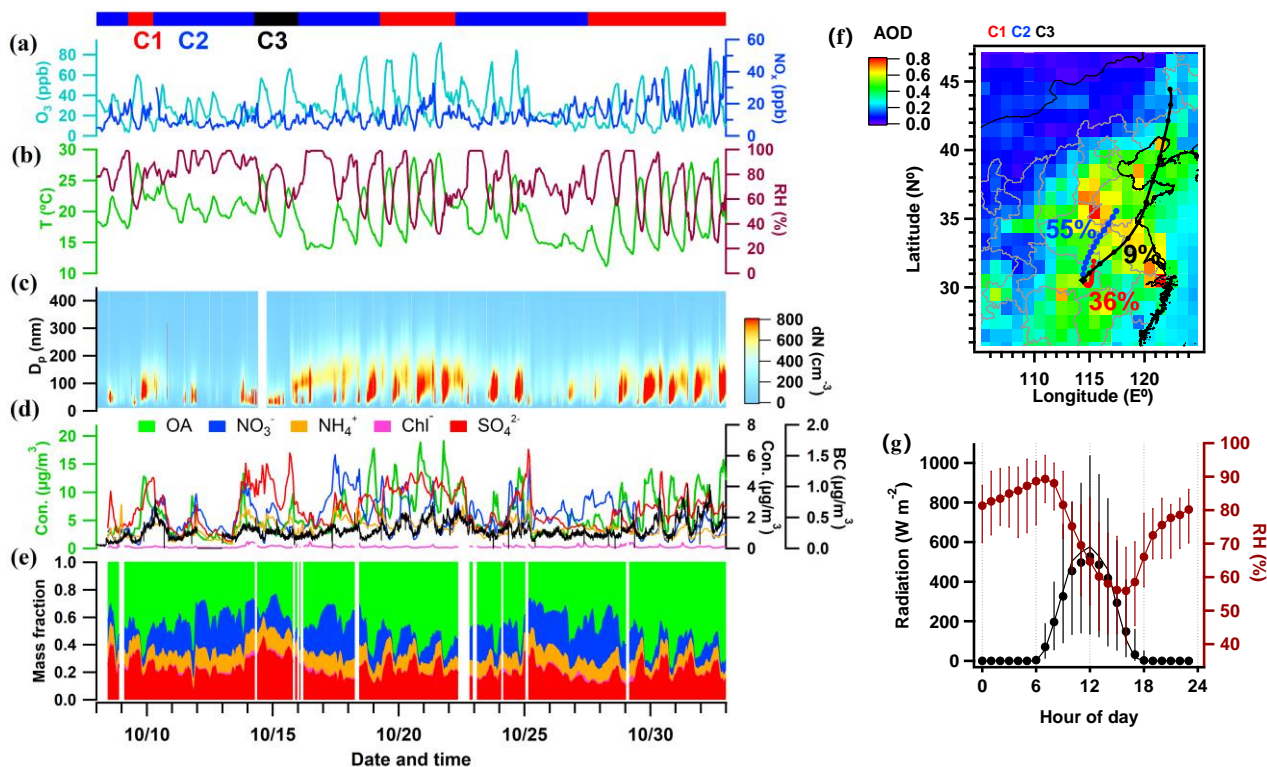
136 **3 Results and discussion**

137 **3.1 Chemical compositions of PM_{10}**

138 The time series of mass concentrations of BC, non-refractory- PM_{10} (NR- PM_{10}) species (i.e., OA, SO_4^{2-} , NO_3^- , NH_4^+ , and
139 Cl^-) and their relative contributions are summarized in Fig. 1. During this field observation, the mass concentrations of NR-
140 PM_{10} were in the range of 2.5–44.8 $\mu\text{g m}^{-3}$, with an average of $12.7 \pm 5.7 \mu\text{g m}^{-3}$, which is close to those observed in the
141 North China Plain in autumn 2019 (15.1 $\mu\text{g m}^{-3}$) (Li et al., 2021a), and was much lower than that ($41.3 \pm 42.7 \mu\text{g m}^{-3}$)
142 observed in Beijing in autumn 2012 (Hu et al., 2017). The mean mass concentration of BC was $0.29 \pm 0.17 \mu\text{g m}^{-3}$ with the
143 range of 0.1–1.0 $\mu\text{g m}^{-3}$. Among all species in NR- PM_{10} , OA contributed the major (49.2 %), indicating the dominant role of
144 OA in autumn time of PM_{10} pollution in this region. Inorganic aerosol accounted for 50.8 % of NR- PM_{10} in which sulphate
145 was the largest contributor (21.5 %), followed by nitrate (18.5 %), ammonium (10.7 %), and chloride (0.1 %), which has
146 similar relative contributions to those observed in field (Chen et al., 2021). The number concentration of all particles peaked
147 at $95 \pm 38 \text{ nm}$ in with regular growth during every field day. Fig. 1 also presents the temporal evolution variation in
148 meteorological parameters. During the field observation period, the average temperature and relative humidity (RH) were
149 $19.3 \pm 3.6^\circ\text{C}$ (11.2–29.4 $^\circ\text{C}$) and $75.0 \% \pm 17.5 \%$ (25.0–99.0 %), respectively. The O_3 concentration with an average of 30.7
150 $\pm 13.0 \text{ ppb}$, was highest 54.5 ppb at 15:00, likely due to the high temperature and enhanced photochemical processing
151 (Zheng et al., 2021).

152 Fig. 1f shows the spatial distribution of aerosol optical depth (AOD) and 36 h backward trajectories during the
153 campaign. Cluster-1 (C1) shows the circulated air mass with shortest transport distance for the past 36 h (36 % fraction); C2
154 is the northerly transported air mass (55 % fraction), over regions with higher AOD; C3 is the most rapid transport through

155 the northeast China and some coastal areas. It therefore separates the periods with more local air mass or influences from
 156 regional transport. Fig. 1g shows the diurnal variations of solar radiation and RH, peaking in the day and night respectively.



157

158 **Figure 1:** Time series of (a) mass concentrations of O₃ and NO_x; (b) ambient temperature (T) and relative humidity (RH); (c)
 159 number size distribution measured by the SMPS; (d) mass concentrations of key aerosol species (the green left y-axis represents
 160 OA, the right y-axis represents other aerosol species); (e) mass fractions of chemical species in non-refractory PM₁; (f) spatial
 161 distribution of mean aerosol optical depth (AOD) during the experiment and the clustered 36 h backward trajectories; (g) diurnal
 162 profiles of direct solar radiation and RH.

163 3.2 Source attributed organic aerosol

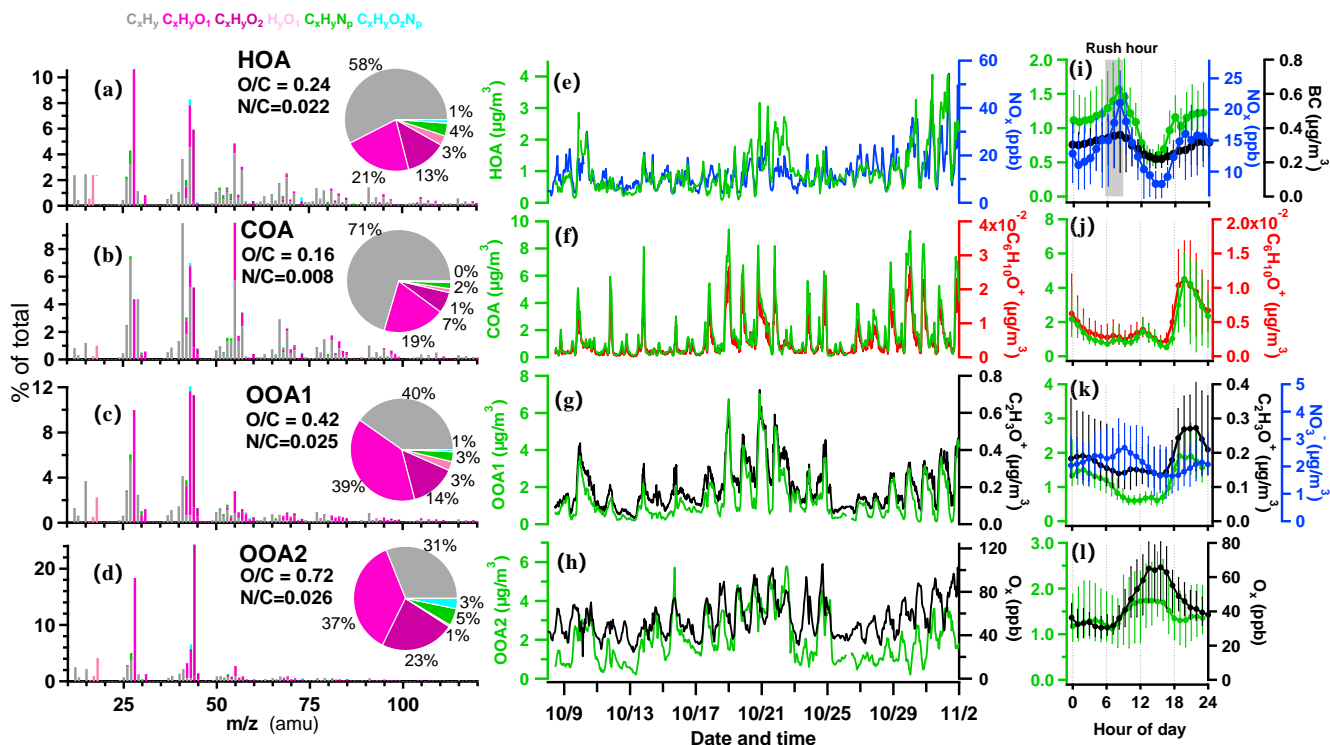
164 The resolved OA factors by the PMF analysis on the AMS measurements are shown in Fig. 2. The first factor is
 165 characterized by prominent hydrocarbon ion series of C_nH_{2n-1}⁺ and C_nH_{2n+1}⁺ (Fig. 2a), with a low O/C ratio (0.24), which is
 166 generally considered to be related to the emissions of fossil fuel combustion and vehicle emissions (Huang et al., 2010;
 167 Morgan et al., 2010; Ng et al., 2011b). The time series of the factor correlated well with NO_x (r = 0.71) and BC (r = 0.83)
 168 (Fig. 2e and Table S1). The diurnal pattern of this factor showed peaks in the morning and afternoon rush-hour (Fig. 2i),
 169 with a major increase from 5:30, reaching a peak value of 1.5 μg m⁻³ at 8:30. The concentration gradually decreased around
 170 noontime due to boundary layer dilution until 15:00 and reached a minimum of 0.6 μg m⁻³. This spectrum also contained
 171 some fragment markers for possible coal combustion OA (CCOA), i.e. C₉H₇⁺ (m/z 115, r = 0.73) (Hu et al., 2013). This

172 factor was not distinctly resolved in this dataset maybe due to the urban nature of the site, where the traffic source may have
173 overwhelmed, due to the less significant coal combustion pollutants during the sampling period.

174 The second factor was characterized by m/z 55 ($C_4H_7^+$, $C_3H_3O^+$) and 57 ($C_4H_9^+$, $C_3H_5O^+$), accounting for 10 % and 3.5
175 % of the total spectrum, respectively (Fig. 2b), with the lowest O/C ratio among factors (0.16). This factor has a
176 $C_3H_3O^+/C_3H_5O^+$ of 3.0 and $C_4H_7^+/C_4H_9^+$ of 2.0 (1.0 is usually for HOA), indicating it as cooking source rather than HOA
177 (Mohr et al., 2012). The correlation coefficient of the COA factor and marker fragment $C_6H_{10}O^+$ was 0.91, which is also
178 similar to a previous study (Sun et al., 2011). The diurnal pattern of this factor showed a major peak during 18:00–20:00,
179 reaching up to $4.0 \mu\text{g m}^{-3}$ on average, in addition to a smaller peak at lunch time, which corroborated the diurnal cooking
180 activities. Previous studies reported COA accounted for 6.5–30 % of the total OA in urban (Rogge et al., 1991; Lanz et al.,
181 2008; Allan et al., 2010; Xu et al., 2014). Here the concentrations of COA in OA were in the range of 0.5–4.5 $\mu\text{g m}^{-3}$ and
182 accounted for 23 % of OA mass fraction on average. The cooking emission in the studying region was likely from charcoal-
183 grilling activities, which was popular in the surrounding areas.

184 Besides the two primary OA, additional two oxygenated OA (OOA) factors are identified. According to the oxidation
185 state, OOA was separated into lower (OOA1) and more oxidized (OOA2) factors. OOA1 factor contained abundant oxygen-
186 containing fragments, accounting for more than 50 % of the total mass spectrum, with an O/C of 0.42. In particular, the
187 oxygenated fragment containing one oxygen accounted for 39 %. The $C_2H_3O^+$ ion (m/z 43) is an important component of less
188 oxidized SOA (Wang et al., 2020a), which was highly correlated with OOA1 ($r = 0.94$). OOA1 showed lower concentration
189 during daytime but higher concentration during nighttime (Fig. 2k). Such diurnal variation was similar to that of RH ($r =$
190 0.45), but opposite to that of solar radiation, which was similar to a previous report (Sun et al., 2014). OOA1 may thus be
191 associated with aqueous reactions when high RH. The co-occurrence of nighttime peak of OOA1 and COA suggested the
192 primary source of cooking emission may have considerably contributed to the production of OOA1.

193 The factor OOA2 had the highest O/C of 0.72 and contained 61 % oxygen-containing fragments (Fig. 2d), which is very
194 similar to the spectra of OOA factor resolved in other cities (0.60 and 67 % in Mexico City, 0.80 and 66 % in Pasadena)
195 (Aiken et al., 2009; Hayes et al., 2013). This factor was correlated strongly with the fragment of CO_2^+ ($r = 0.79$). Compared
196 to OOA1, this factor showed obvious diurnal variations with a major enhancement at around noontime (10:00–15:00), up to
197 $2.9 \mu\text{g m}^{-3}$, indicating the photochemical production of SOA during daytime. The variation of OOA2 correlated with odd
198 oxygen ($O_x = O_3 + NO_2$, $r = 0.70$), agreeing with previous observations that oxidized OA had strong correlation with O_x
199 (Wood et al., 2010; Hu et al., 2016). Considering the high RH during the experiment ($> 60 \%$, Fig. 1g), OOA2 factor may
200 have also experienced aqueous chemistry and showed good correlation with sulphate ($r = 0.82$).



201
 202 **Figure 2: Results of source-attributed organic aerosols by the PMF analysis on the OA mass spectra measured by the HR-TOF-**
 203 **AMS. Mass spectra of PMF factors for (a) hydrocarbon-like OA (HOA); (b) cooking OA (COA); (c) oxygenated OA1 (OOA1); (d)**
 204 **oxygenated OA2 (OOA2). The pies show the relative contributions of the six ion categories to each factor. (e-h) Temporal**
 205 **variations of four OA factors (HOA, COA, OOA1 and OOA2) with each correlated species as NO_x , $\text{C}_6\text{H}_{10}\text{O}^+$, $\text{C}_2\text{H}_3\text{O}^+$ and O_x . (i-l)**
 206 **Diurnal profiles of NO_x , BC, $\text{C}_6\text{H}_{10}\text{O}^+$, $\text{C}_2\text{H}_3\text{O}^+$, NO_3^- and O_x . The lines and whiskers denote the median, the 25th and 75th**
 207 **percentiles at each hour, respectively.**

208 3.3 Source attributed VOCs

209 Fig. 3 and Fig. S5 summarize the source attributed of VOCs and key indicators. Two primary and three secondary VOC
 210 factors were identified. Here 109 species of VOCs from the PTR-TOF-MS were used, which are mostly oxygenated or
 211 contain cyclic functional groups, in contrast with the GC-MS measurement which contains many aliphatic hydrocarbon
 212 compounds and usually only decades of species were used for the PMF analysis (Zheng et al., 2021). These may lead to
 213 some unsolved primary VOC sources for the current analysis using the PTR-measured VOCs, if the source contained crucial
 214 markers of aliphatic species, but this method may have great advantage in comprehensively resolving factors of secondary
 215 VOCs. This may explain the less primary but more secondary VOC factors compared to conventional PMF-resolved VOC
 216 sources (Cai et al., 2010).

217 The first factor was dominated by aromatic compounds, such as C_6H_6 (m/z 79.054) and C_7H_8 (m/z 93.070) in Fig. 3a.
 218 They were well established markers for vehicle emissions (Gkatzelis et al., 2021) and had a good correlation with this factor
 219 ($r = 0.97$ and 0.63 , respectively Fig. S6). This factor showed peaks in the morning and afternoon rush-hour (Fig. 3k) and was

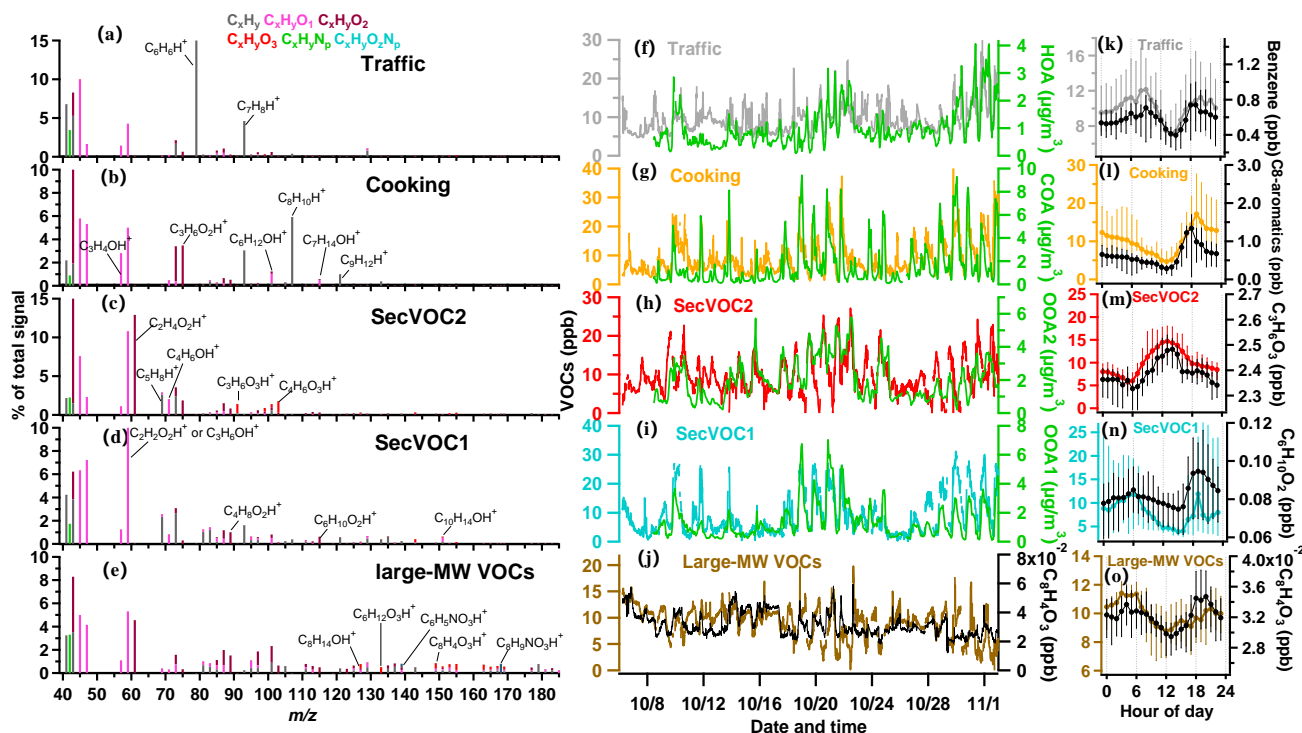
220 well correlated with HOA and NO_x ($r = 0.63$ and 0.53 Table S1). The concentration of this factor had a major increase
221 during the early morning, reaching a peak value of 12.1 ± 1.3 ppb at 08:30, further corroborating the traffic source of this
222 factor. The concentration decreased around noontime until 15:00 because of the dilution by well-developed boundary layer
223 and in addition to its consumption through photochemical reaction due to intense solar radiation. The diurnal pattern was
224 also high at night, although the peak was lower than early morning, the average concentration was higher by a factor of 2
225 than daytime. This suggests that traffic VOCs is more likely to participate in photochemical reaction, and other primary
226 emissions may be precursors for nocturnal chemistry.

227 The second factor contained abundant aldehydes such as C₃H₄O (acrolein), C₃H₆O₂ (hydroxyacetone), C₆H₁₂O
228 (hexanal), and C₇H₁₄O (heptanal) at m/z 57.069, 75.044, 101.096, and 115.112, respectively, as well as C₈H₁₀ (C8-aromatics)
229 and C₉H₁₂ (C9-aromatics) at m/z 107.086 and 121.101 in VOC mass spectra (Fig. 3b), which are footprint VOCs identified
230 from primary cooking emissions during the charbroiling and frying (Klein et al., 2016). This factor had similar time series
231 (Fig. 3g) with COA and had high correlation ($r = 0.67$, Table S1). The concentration of this factor decreased during the
232 daytime and increased after 18:00 with a peak value at 19:00 (17.2 ± 3.0 ppb). As shown in Fig. 3l, the diurnal concentration
233 decreased strongly after emission and continued to decline throughout the night, suggesting that cooking VOCs may be
234 major precursors and were consumed during night.

235 Besides the two primary VOCs, three secondary oxygenated VOCs (SecVOCs) factors are identified. These three
236 factors are not well correlated with any primary factors attributed by the OA, are thus considered to be mainly contributed by
237 secondary production. The daytime biogenic emission, e.g. isoprene, may also contribute to the SecVOC2 formation by
238 interacting with OH in the presence of NO_x (Lin et al., 2013), producing methacrylic acid epoxide (MAE) and methacrolein
239 (MACR) as intermediately involved in SOA formation (Gu et al., 2021). This factor was thus also contributed by this
240 primary source. SecVOC2 contributed a large fraction of C₂H₄O₂, C₃H₆O₃ and C₄H₆O₃ (Fig. 3c), which were gas-phase
241 carboxylic acids, the oxidation products from photochemical processes (Hartikainen et al., 2018; Li et al., 2021b). SecVOC2
242 had one clear peak at 13:00 (14.7 ± 2.8 ppb Fig. 3m). The time series of SecVOC2 correlated well with OOA2 ($r = 0.67$),
243 following the variation of solar radiation. It had a rapid enhancement starting from 07:00 to 12:00 and declined continuously
244 after 13:00. This indicated that many of these species can be formed rapidly during daytime and may have a short lifetime
245 owing to the partitioning to the condensed phase and forming SOA.

246 SecVOC1 factor featured with some less-oxygenated VOCs e.g., C₂H₂O₂, C₆H₁₀O₂ and C₁₀H₁₄O (Fig. 3d). This factor
247 had a peak at 19:00 which was consistent with the primary cooking VOCs factor, but also increased throughout the night,
248 peaking at midnight. This factor had a similar temporal trend with OOA1 ($r = 0.76$, Fig. 3i), which was less oxygenated than
249 photochemistry dominated OOA2. Combining the features above, SecVOC1 tended to be contributed by some immediately
250 reacted species from emissions in the late afternoon and early night. A particular factor (Fig. 3e) is significantly composed of
251 large molecular weight (large-MW) oxidized VOCs, i.e. the mean contribution of ionic compounds with $m/z > 120$ was
252 above 50 % in this factor (Fig. S5e), which was much higher than that in SecVOC2 (Fig. S5c). Fig. 3e shows its signature

253 compounds of $C_8H_{14}O$, $C_6H_{12}O_3$ and $C_8H_4O_3$, and some are nitrogen-containing VOCs, such as $C_6H_5NO_3$ and $C_8H_9NO_3$.
 254 These VOCs with $m/z > 120$ tend to be intermediate-volatility organic compounds (IVOCs) as the estimated vapor saturation
 255 concentration ($\log_{10}C^*$, 298 K) is less than $6.5 \mu\text{g m}^{-3}$ (Fig. S2). This factor is hereby termed as large-MW VOCs to indicate
 256 the fraction of IVOCs, which only require few oxidation steps to become semi-volatile (Robinson et al., 2007). Fig. 3o
 257 shows an increase of this factor at mid-night, later than the peak of SecVOC1, which may imply the ageing process in
 258 producing these VOCs.
 259



260
 261 **Figure 3.** Source attributed VOCs by the PMF analysis measured by the PTR-TOF-MS. (a-e) Mass spectra of the five factors
 262 (traffic VOCs, cooking VOCs, secondary VOCs (SecVOC2, SecVOC1) and large molecular weight (MW) VOCs (large-MW
 263 VOCs), with major fingerprint peaks labelled in the mass spectra. (f-j) Temporal variations of the five VOCs factors and their
 264 respective correlated OA species. (k-o) Diurnal profiles of the VOC factors and their respective signature species. The lines and
 265 whiskers denote the median, the 25th and 75th percentiles at each hour, respectively.

266 3.4 Oxidation process of organics in the day and night

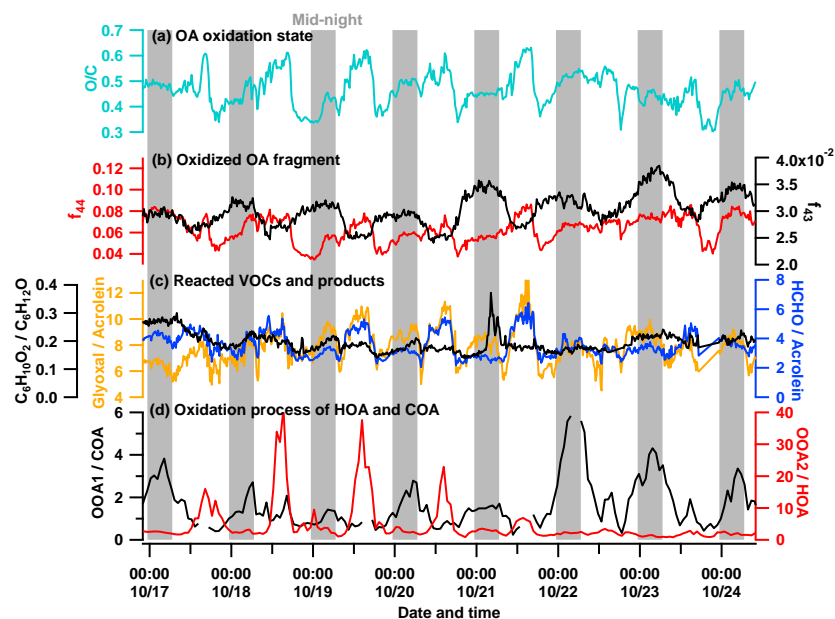
267 After source attribution of organics in aerosol and gas phase, we are able to identify the emission structure of primary
 268 sources and the consequent evolution. Given the identified primary traffic and cooking sources were emitted in the day and
 269 late afternoon respectively, this provides the potential opportunity to study the evolution of different primary sources in the
 270 potentially contrasting ageing mechanisms in the day and night.

271 Fig. 4a-d shows the temporal evolution of the key factors in a week to reflect the day and night ageing. The O/C ratio in
272 the day was higher than that at night by 0.2 (Fig. 4a), corresponding with the increase of highly oxygenated fragments of
273 AMS f_{44} (CO_2^+) in the day (Fig. 4b). During night, the O/C ratio increased with the increase moderately oxygenated fragment
274 f_{43} ($\text{C}_2\text{H}_3\text{O}^+$). The diurnal variation of atomic ratio O/C of OA (Fig. 4e) showed a slight decrease at morning rush-hour and
275 sharp drop at 18:00 (by 0.2) due to the significant contribution of the traffic and cooking sources respectively, corresponding
276 with the two steps of increase at each time. This variation of O/C clearly showed the daytime and nighttime ageing processes
277 of OA for different primary sources. Notably, the O/C showed peak value of 0.54 at 15:00, implying the importance of
278 photooxidation in producing highly oxygenated OA.

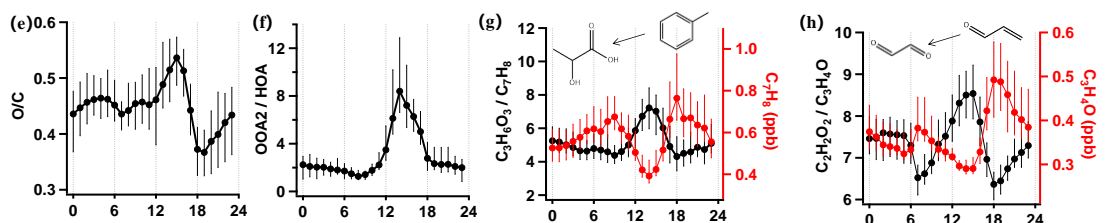
279 Considering the diurnal pattern of anthropogenic activities, the traffic emission at rush-hour is deemed to be the major
280 source in contributing to the daytime production of OOA. The ratio OOA2/HOA is thus used to indicate the daytime
281 oxidation of OOA. As Fig. 4d and 4f showed, OOA2/HOA had a clear peak during daytime, and increased after 8:00 and
282 peaked at 14:00. After 15:00, the ratio gradually decreased to minimum at 20:00 and maintained to be low throughout the
283 night. This clearly demonstrated the photooxidation in producing OOA2 from oxidizing HOA. Fig. 4c and 4g give a few
284 examples of photooxidation in gas phase: the toluene showed a production of $\text{C}_3\text{H}_6\text{O}_3$ (hydroxypropionic acid) and $\text{C}_4\text{H}_6\text{O}_3$
285 (acetic anhydride) by a factor of 2 in 3 hours (Fig. S7c); the $\text{C}_3\text{H}_4\text{O}$ (acrolein) produced the oxidized product $\text{C}_2\text{H}_2\text{O}_2$
286 (glyoxal) by a factor of 1.3 (Fig. 4h). All of these reacted species are from traffic VOCs and the corresponding products are
287 from SecVOC2. The daytime biogenic emission, e.g. isoprene, may also contribute to the SecVOC2 formation by interacting
288 with OH in the presence of NO_x (Lin et al., 2013), producing methacrylic acid epoxide (MAE) as intermediately involved in
289 SOA formation, where a considerable nitrogen content was also found in the OOA2 factor ($\text{N}/\text{C} = 0.026$). Overall, the highly
290 oxidized OOA2 ($\text{O}/\text{C} = 0.72$) is considered to be mainly contributed by the traffic source, via the oxidation of VOCs and
291 partitioning to condensed phase, direct oxidation on HOA through heterogenous oxidation (Guo et al., 2020), or VOCs
292 evaporated from HOA and further condensation after oxidation (Zhao et al., 2015). All factors may have contributed to the
293 daytime production of OOA2.

294 The nocturnal oxidation is mainly contributed by the sources emitting from late afternoon throughout the midnight.
295 Both traffic and cooking sources contributed to the emission since late afternoon, with cooking source as the predominant
296 contributor in both aerosol (Fig. 2j) and gas (Fig. 3l) phases. The ratio between the nighttime OOA1 and cooking aerosol
297 (OOA1/COA) is therefore used to indicate the nocturnal oxidation of SOA (Fig. 4d and j). There was a sudden increase of
298 OOA1/COA during the daytime because COA was consumed rapidly in the afternoon after a small amount of emission at
299 noon (Fig. 2j). The lowest OOA1/COA at 0.5 corresponded with the fresh cooking emission at 18:00, and kept increasing
300 until peaking in the early morning at 6:00 up to 3, which was an increase by a factor of 6 compared to the minima (Fig. 4j).
301 In addition to evidence for daytime reaction, Fig. 4h and Fig. S7d also gave evidence for certain reacted species during the
302 ageing at nighttime. The first-generation oxidation products from acrolein (ACR $\text{C}_3\text{H}_4\text{OH}^+$, m/z 57.033) are glyoxal
303 ($\text{C}_2\text{H}_2\text{O}_2$, m/z 59.036) and formaldehyde (HCHOH^+ , m/z 31.018) according to the database in Master Chemical Mechanism

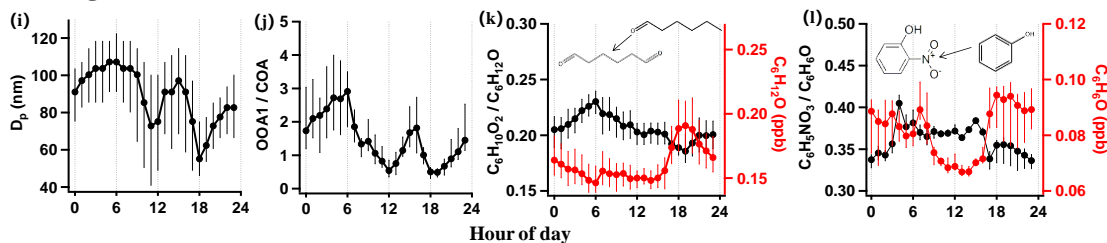
304 (MCM). This night oxidation is also evidenced by the formation of some nitrated organic compounds and ketone
305 compounds, such as $C_6H_5NO_3$ (nitrophenol, NP, Fig. 4l) and $C_6H_{10}O_2$ (hexanedione, Fig. 4k) produced from phenol (C_6H_6O)
306 and hexanal ($C_6H_{12}O$), respectively. Notably, an important fraction of large molecular weight-VOCs (which are mostly
307 intermediate VOCs, Fig. S2) peaked at 3:00–4:00 (Fig. 3o), consistent with the variation of OOA1/COA during night. Some
308 of these nitrated and oxygenated IVOCs may have been further oxidized and partitioned to aerosol phase, contributing to the
309 OOA1. Given the larger molecular usually has a lower O/C ratio (Hatch et al., 2017) (because of a higher content of carbon),
310 this may explain the lower O/C observed for nighttime formed OOA1 (O/C = 0.42), than OOA2 produced by daytime
311 photooxidation (O/C = 0.72). Notably, nighttime SOA had a high N/C (0.025), implying the NO_3^- -initiated from cooking
312 emissions oxidation, which was different from the organic nitrate formation mechanism during the daytime.



Daytime photo-oxidation



Nighttime oxidation



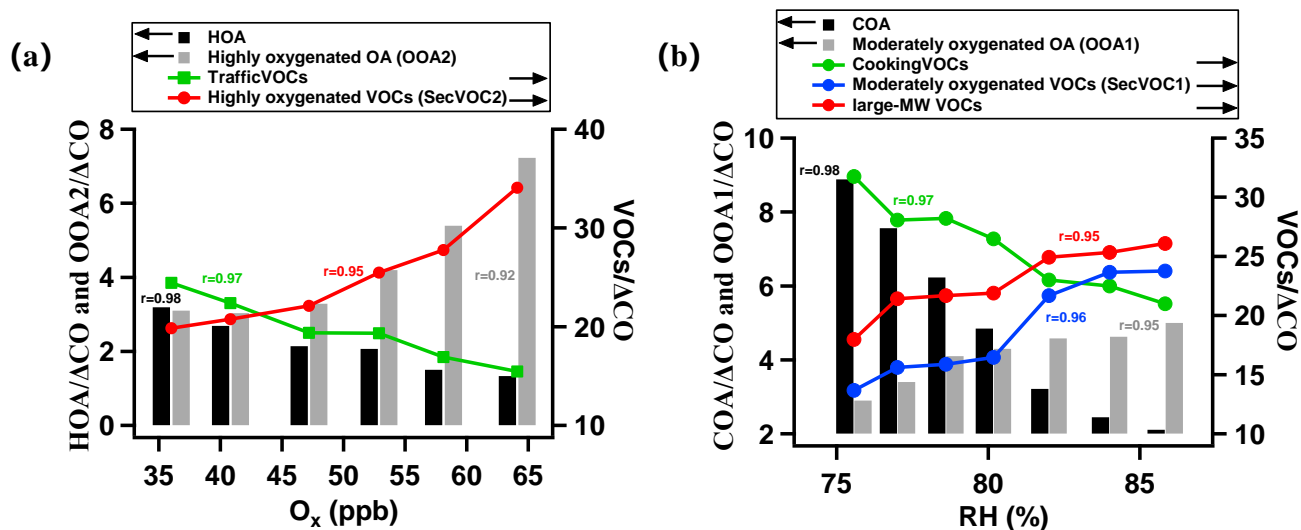
313
 314 **Figure 4.** Time series showing the ageing of aerosols and gases, with grey vertical bars denoting nighttime (00:00–06:00), (a) the
 315 oxygen to carbon ratio (O/C); (b) the fragment fraction at m/z 44 (f_{44}), m/z 43 (f_{43}); (c) concentration ratio of $C_6H_{10}O_2$ to $C_6H_{12}O$
 316 (hexanedione to hexanal), glyoxal to acrolein, and formaldehyde to acrolein; (d) concentration ratio of OOA1 to COA and OOA2
 317 to HOA. Diurnal variations of key species showing the daytime photooxidation and nighttime oxidation, (e) ratio of O/C; (f)
 318 ratio of OOA2 to HOA; (g) concentration of C_7H_8 and ratio of $C_3H_6O_3$ to C_7H_8 (hydroxypropionic acid to toluene); (h) concentration of
 319 C_3H_4O and ratio of $C_2H_2O_2$ to C_3H_4O (glyoxal to acrolein); (i) particle diameter (D_p); (j) ratio of OOA1 to COA; (k) concentration
 320 of $C_6H_{12}O$ and ratio of $C_6H_{10}O_2$ to $C_6H_{12}O$ (hexanedione to hexanal); (l) concentration of C_6H_6O and ratio of $C_6H_5NO_3$ to C_6H_6O
 321 (nitrophenol to phenol).

322

323 Fig. 5 summarizes the key indicators for the contrasting daytime and nighttime oxidation process. The concentrations
 324 are normalized by ΔCO , where ΔCO is the total CO concentration subtracted by the background concentration (1th

325 percentile of the dataset), to indicate the variation of species regardless of the boundary layer evolution, wind speed, wind
326 direction and temperature (Gouw and Jimenez, 2009). The odd oxygen O_x ($O_3 + NO_2$) has been widely used to generally
327 indicate the activity of daytime photochemistry (Hu et al., 2017), and the enhanced moisture is the main driving factor for
328 nighttime chemistry. The O_x concentration and RH are therefore used as references with which the variations of species were
329 correlated in the day and night, respectively. As Fig. 5a showed, the traffic primary emissions in both gas (traffic VOCs) and
330 aerosol phase (HOA) declined with increased O_x by 60 % and 40 % respectively, suggesting their roles as precursors in the
331 daytime reaction. The produced species are oxygenated SecVOC2 and OOA2, showing enhancement with O_x , peaking at
332 midday-afternoon. This process was rapid as the SOA production by a factor of 2.5 and the oxygenated VOCs production by
333 a factor of 1.7 within 6 hours. The traffic VOCs are widely observed to contribute to SOA production, with aromatic
334 compounds serving as key precursors (Fang et al., 2021). The semi-volatile nature of HOA means it could be evaporated to
335 gas phase and further oxidized to recondense as SOA (Robinson et al., 2007). The decrease rate of HOA with increased
336 photochemical age was also found in urban environment (Zhu et al., 2021), generally consistent with the reaction rate in this
337 study. Here we linked the declining rate as a function of photochemical activities for both reacted aerosol and gas phases for
338 traffic sources. The gases evaporated from aerosol phase (especially under higher temperature when increased saturation
339 pressure for semi-volatile or intermediate volatile species) and primary VOCs may be simultaneously involved in the
340 photooxidation, further contributing to the SOA formation.

341 For nocturnal oxidation shown in Fig. 5b, the reacted species are cooking VOCs and COA (decrease by 35 % and 77 %,
342 respectively), producing SecVOC1 and large-MW oxygenated VOCs, with an increase of night SOA formation by a factor
343 of 1.7. The nocturnal processes may have largely involved aqueous reactions, because the variations of reacted or produced
344 species were highly correlated with RH ($r > 0.95$, Fig. 5b). The large-MW VOCs (mostly IVOCs) increased by 50 % and
345 reached maxima when highest RH. This suggests the moisture may have been involved in converting some primary VOCs to
346 IVOCs, which further contribute to the SOA production during nighttime. Previous studies also found the oxidation of
347 IVOCs from cooking sources can be an important source of SOA (Zhang et al., 2020). The evidence is given here that
348 organic aerosols and gases from cooking emission had been reacted and contributed to SOA. An approximate production rate
349 of $0.2 \mu\text{g m}^{-3} \text{h}^{-1}$ of OOA from cooking source can be obtained by considering the ageing time of ~ 10 h (from COA peak
350 18:00 to OOA1 peak 4:00), which was generally consistent with previous laboratory work using gas precursors from cooking
351 sources ($0.07\text{--}0.5 \mu\text{g m}^{-3} \text{h}^{-1}$) (Liu et al., 2017). Notably, daytime SOA had a higher oxidation state, implying the importance
352 of photooxidation in producing highly oxidized OA. This may be because of the high temperature at daytime and a species
353 may require a lower volatility (hereby more oxygenated) to be in condensed phase than at night.



355
 356 **Figure 5. Daytime and nighttime evolution of key species against O_x and RH respectively. (a) Daytime oxidation showing the**
 357 **primary emissions of traffic source in gas and aerosol phases including HOA, traffic VOCs; and secondary products including**
 358 **SecVOC2 and OOA2, with concentrations all normalized by Δ CO. (b) Nighttime oxidation showing the primary emissions of**
 359 **cooking source in gas and aerosol phases including COA and cooking VOCs; and secondary products including SecVOC1, large-**
 360 **MW VOC and OOA1, with concentrations all normalized by Δ CO.**

361 4 Conclusion

362 In this study, organic gases and aerosols were concurrently characterized through online mass spectrometers at a
 363 megacity in central China. Through the factorization analysis on the organic mass spectra, two principal sources—the traffic
 364 and cooking sources were identified for both aerosol and gas phases, hereby the reacted and produced species between
 365 phases were interlinked. We observed clear evidence of daytime and nighttime oxidation of source-attributed OA and VOCs.
 366 Daytime photooxidation caused 60 % decrease of primary aerosol and 40 % of primary VOCs reduction for traffic sources,
 367 producing oxygenated SecVOCs by a factor of 1.7 and OOA by a factor of 2.5, in a 6 hours photochemical ageing.
 368 Nocturnal ageing caused a reduction of primary OA (by 77 %) and primary VOCs (by 35 %) from cooking sources,
 369 producing oxygenated VOCs and OOA by a factor of 1.4 and 1.7, respectively. In particular, larger molecular IVOCs
 370 produced (by a factor of 1.7) at night may importantly contribute to the OOA. This implies primary species in aerosol and
 371 gas phases both contribute to the production of OOA. A higher oxidation state of OOA from daytime photooxidation was
 372 found than nighttime, suggesting different compositions of produced OOA modulated by solar radiation and moisture,
 373 respectively. Vehicle and cooking emissions are the major contributors of organic aerosols in urban areas, especially in
 374 megacities. These results provide direct observations about the reaction rate for primary precursors and production rate for
 375 secondary aerosols, as influenced by primary sources and meteorological conditions. The environmental policy making
 376 should therefore consider the respective primary sources and ageing mechanisms for local and regional atmospheric

377 environmental problems. The particular regulation should be placed to avoid the formation of daytime highly-oxidized
378 species when high solar radiation, which may contribute to the reactive oxygen species and exert adverse health impacts
379 (Tao et al., 2003; Verma et al., 2009).

380 **Data availability**

381 The data in this study are available from the corresponding author upon request.

382 **Author contribution**

383 DL, SK, and SL led and designed the study. SL, YW, HZ, YC, SZ and DH set up and conducted the experiment. SL,
384 DL, YW, KH, XJ and SD contributed to the data analysis. QL, DZ, JS provided technical support and assistance. SL and DL
385 wrote the manuscript. All authors read and approved the final manuscript.

386 **Competing interests**

387 The authors declare no competing interests.

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