Evolution of source attributed organic aerosols and gases in a megacity of central China

3 Siyuan Li¹, Dantong Liu^{1,*}, Shaofei Kong^{2,*}, Yangzhou Wu¹, Kang Hu¹, Huang Zheng², Yi Cheng

⁴², Shurui Zheng ², Xiaotong Jiang ¹, Shuo Ding ¹, Dawei Hu ³, Quan Liu ⁴, Ping Tian ⁵, Delong Zhao ⁵,

5 Jiujiang Sheng ⁵

6 ¹ Department of Atmospheric Sciences, School of Earth Sciences, Zhejiang University, Hangzhou 310027, China

- ² Department of Atmospheric Sciences, School of Environmental Studies, China University of Geosciences, Wuhan, 430074,
 China
- ³ Centre for Atmospheric Sciences, School of Earth and Environmental Sciences, University of Manchester, Manchester M13
 9PL, UK

⁴ State Key Laboratory of Severe Weather & Key Laboratory of Atmospheric Chemistry of CMA, Chinese Academy of
 Meteorological Sciences, Beijing 100081, China

13 ⁵ Beijing Key Laboratory of Cloud, Precipitation and Atmospheric Water Resources, Beijing Meteorological Service, Beijing

- 14 100089, China; Field Experiment Base of Cloud and Precipitation Research in North China, China Meteorological
- 15 Administration, Beijing 100089, China.
- 16

17 Correspondence to: Dantong Liu (dantongliu@zju.edu.cn); Shaofei Kong (kongshaofei@cug.edu.cn)

18 Abstract. The secondary production of the oxygenated organic aerosol (OOA) impacts air quality, climate and human 19 health. The importance of various sources in contributing to the OOA loading and associated different aging mechanisms 20 remains to be elucidated. Here we present concurrent observation and factorization analysis on the mass spectra of organic 21 aerosol (OA) by a high-resolution aerosol mass spectrometer and volatile organic compounds (VOCs) by a proton-transfer-22 reaction mass spectrometer in Wuhan, a megacity in central China during autumntime. The full mass spectra of organics 23 with two principle anthropogenic sources were identified as the traffic and cooking sources, for their primary emission 24 profiles in aerosol and gas phases, the evolutions, and their respective roles in producing OOA and secondary VOCs. 25 Primary emissions in gas and aerosol phases both contributed to the production of OOA. The photooxidation of traffic 26 sources from morning rush-hour caused 2.5-folds increase of OOA mass in a higher oxidation state (oxygen-to-carbon, O/C 27 = 0.72), coproducing gas-phase carboxylic acids; while at night, cooking aerosols and VOCs (particularly acrolein and 28 hexanal) importantly caused the nocturnal formation of oxygenated intermediate VOCs, increasing OOA mass by a factor of 29 1.7 (O/C = 0.42). The daytime and nighttime formation of secondary aerosols as contributed by different sources, was found 30 to be modulated by solar radiation and air moisture respectively. The environmental policy should therefore consider the 31 primary emissions and their respective ageing mechanisms influenced by meteorological conditions.

32 1 Introduction

33 The transformation between gas and aerosol phase of organic species produces secondary organic aerosols (SOA) 34 (Kroll and Seinfeld, 2008; Hallquist et al., 2009; Slowik et al., 2010), which forms important global budget of aerosol 35 loadings (Heald et al., 2005; Zhang et al., 2007), exerting climate (Poschl, 2005; Seinfeld et al., 2016) and environmental 36 impacts (Von Schneidemesser et al., 2011; Huang et al., 2014). Given the large complexities of organic species, the yields 37 (Goldstein and Galbally, 2007; Ortiz-Montalvo et al., 2014) and production rates (Sareen et al., 2013; Jokinen et al., 2015) of 38 SOA from gas precursors are influenced by the diversities of source profiles (emission mass percentages among species) 39 (Lin et al., 2012; Shrivastava et al., 2015) and environmental factors (such as radiation, air moisture and ambient 40 temperature) (Clark et al., 2016; Li et al., 2021b; Wang et al., 2021). This raises challenges for source-oriented environmental policy making, until explicit understanding on the formation mechanism of SOA from different sources. 41

42 The application of detailed mass spectra of organic aerosols allows the online source attribution of organic aerosol (OA) 43 (Canagaratna et al., 2007; Ng et al., 2011a), based on the factorization analysis on the mass spectra of OA which groups the 44 covaried species from certain sources (Ulbrich et al., 2009). This factorization technique allows identification of primary 45 sources and aged secondary sources at a receptor measurement site. The complex aerosol sources in urban environment have been commonly recognized as traffic sources mainly from vehicles on road (Zhu et al., 2021), biomass burning from open or 46 47 closed combustion (Adler et al., 2011), coal combustion normally in cold season for heating (Xu et al., 2019) and more 48 localized cooking sources (Allan et al., 2010; Zhang et al., 2020). The secondary sources are oxygenated OA which may be 49 from oxidizing volatile organic compounds (VOCs) following condensation (Donahue et al., 2006), or heterogenous oxidation occurring on particle phase (Claevs et al., 2004). Among these sources, SOA has been found to be the main 50 51 contributor of OA mass loading (41-69 %) in urban environment of East Asia (Sun et al., 2014; Hu et al., 2017), especially 52 in warm season when primary emissions were low, along with high ambient temperature and more intensive chemical 53 reactions (Hu et al., 2016). The formation of SOA from VOCs may experience a few reaction generations (Kroll and 54 Seinfeld, 2008) and could interact with other sources of species during the process (Shrivastava et al., 2017), hereby 55 complicating the goal of identifying the key precursors in contributing to the consequent SOA. In addition, some primary 56 gases already have somehow low volatility and may not require a long reaction chain to become condensable, such as some 57 primarily emitted intermediate volatility organic compounds (IVOCs) may substantially contribute to the SOA (Robinson et 58 al., 2007; Huang et al., 2021). An understanding of source profiles of primary emissions in both gas and aerosol phases is 59 therefore important to rule out the source-dependant production of SOA. The above necessitates the concurrent investigation 60 on the compositions of gases and aerosols at a receptor site, along with their evolution and interaction, in order to elucidate the role of each source in contributing to SOA. 61

In this study, we performed online continuous measurements on the detailed mass spectra of organics concurrently on aerosol and gas phases, in a typical anthropogenically polluted region in central China, where such data had been rarely 64 reported. Factorization analysis is performed on the mass spectra of organics in both aerosol and gas phases, to investigate

65 the source-oriented gas-aerosol evolution and SOA formation in this region.

66 2 Materials and methods

67 2.1 Sampling site

The field experiment was performed in the campus of the China University of Geosciences (Wuhan) (114.40° E. 30.52° 68 N) during Oct.–Nov. 2019 (Hu et al., 2021). The site represents a typical region of mixed residential/traffic sources (Fig. S1). 69 Due to the preparation and hosting of the 7th CISM Military World Games during the experimental period, the government 70 71 implemented strict emission reduction measures, particularly for the industrial sources and heavy vehicle emissions in the 72 main roads. The more localized pollution sources, such as traffic emission from smaller sizes of vehicles and cooking sources dominated the pollution in this study, besides those transported from surrounding regions during some heavy 73 74 pollution events (Zheng et al., 2019; Zheng et al., 2020). The HYSPLIT model (Draxler and Hess, 1997) with a 3-hourly 1° 75 \times 1° wind field from the GDAS reanalysis products was used to obtain the 36-h backward trajectories initializing from the location of the experiment site. The cluster analysis was performed to categorize the trajectories into three groups by 76 77 minimizing the differences in each group and maximizing the differences among groups (Moody and Galloway, 1988).

78 2.2 Instrumentation

79 2.2.1 Measurements of mass spectra of aerosols and gases

80 The mass concentration and chemical composition of NR-PM₁ were measured by a HR-ToF-AMS (Aerodyne Research 81 Inc., USA), including organics, nitrate (NO₃⁻), sulphate (SO₄^{2–}), chloride (Chl⁻), and ammonium (NH₄⁺). The detailed can be 82 found elsewhere (Decarlo et al., 2006). Briefly, the aerosols are dried using a diffusion dryer before entering the AMS and 83 through a critical orifice into a narrow beam via an aerodynamic lens. The aerosol size is determined using the flight time of 84 particles to the thermal vaporization and ionization chamber. Then the aerosols are successively vaporized by a heated 85 surface (~600 °C), ionized by electron ionization (EI, 70 eV), and detected by a mass spectrometer detector. During this field observation, the HR-ToF-AMS was operated under V mode with high sensitivity. The composition-dependent collection 86 efficiencies were applied (Middlebrook et al., 2012), and the ionization efficiency was calibrated using 300 nm pure 87 88 ammonium nitrate (Jayne et al., 2000). Elemental analysis (EA) was also executed using the "Improved Ambient" method 89 (Canagaratna et al., 2015) to obtain the hydrogen-to-carbon ratio (H/C), oxygen-to-carbon ratio (O/C), and nitrogen-to-90 carbon ratio (N/C).

91 The Proton-Transfer-Reactor Time-of-Flight Mass Spectrometer (PTR-ToF-MS 8000, Ionicon Analytik GmbH 92 Innsbruck, Austria) was deployed to quantify VOCs in this research. The operating and calibration of the PTR followed the 93 routine described previously (Cappellin et al., 2012; Bruns et al., 2016). Briefly, the PTR was operated with hydronium ions

 (H_3O^+) as the reagent and with a drift tube pressure of 2.2 mbar, voltage of 600 V and temperature of 60 °C. The ratio of the 94 95 electric field (E) and the density of the buffer gas (N) in the drift tube, which dictates the ion drift velocity in the drift tube, is 96 135 Td. MS transmission function was performed using a mixture of VOCs (formaldehyde, methanol, acetonitrile, 97 acetaldehyde, acetone, isoprene, methyl ethyl ketone, benzene, toluene, styrene, benzaldehyde, ethylbenzene, 1,3,5-98 trimethylbenzene). Mass calibration was done using H₃O⁺ (m/z 21.0226), CH₃COCH₄⁺ (m/z 59.0490) and monoterpenes (m/z99 137.1290). The shift in m/z is minor which ensures that the mass calibration was sufficient for all compounds. The 100 background measurements were performed using a dry zero air cylinder. The measurement error is described in Text S2. A 101 separate reaction rate constant is applied to convert the ion signal into concentration (Bruns et al., 2016), or the default reaction rate constant 2×10^{-9} cm³ s⁻¹ can be applied to all other ions (Wang et al., 2020b). The vapor saturation concentration 102 (equilibrium vapor pressures) (C^*) of each VOC compound at 25 °C is estimated using the parameterization based on 103 elemental ratio and molecular weight (Fig. S2) (Li et al., 2016). A $\log C^*$ (in µg m⁻³) lower than 6.5 is deemed to be 104 105 intermediate VOC (IVOC). According to Fig. S2, VOC species (with m/z > 120) are mainly identified to be IVOCs in this 106 study, thus the fraction of these larger molecular weight (MW) VOCs (m/z > 120) is used to evaluate the potential influence 107 of IVOCs.

108 2.2.2 PMF analysis on the mass spectra of OA and VOCs

Positive matrix factorization (PMF) (Paatero and Tapper, 1994) was performed on the high-resolution organic mass spectral matrix of OA (Ulbrich et al., 2009; Decarlo et al., 2010). In this work, m/z > 120 and isotopic ions were excluded in PMF analysis due to the limited mass resolution and low contributions to OA loading (~5 %). After a careful evaluation of the mass spectral profiles and correlations with time series of tracers, diurnal variations, four OA factors from total OA were identified with *f*peak = 0, including hydrocarbon-related OA (HOA), cooking OA (COA), low-oxidized oxygenated OA (LO-OOA, OOA1), and more-oxidized oxygenated OA (MO-OOA, OOA2). The detail of PMF diagnostic was summarized in Text S1 and Fig. S3.

The EPA PMF 5.0 model (Paatero and Tapper, 1994) was used for the source apportionment of VOC species. The 109 VOCs were used for the PMF analysis. The uncertainties from each sample were determined according to the method detection limit (MDL) and the error fraction (%). The detail of PMF diagnostic was summarized in Text S2 and Fig. S4. Five factors were ultimately selected, and the Q/Q_{exp} ratio was 0.96 on average (Fig. S4). The rotation ambiguity was explored by varying the *f*peak values from -3 to +3, and the results with *f*peak = 0 were selected for the lowest dQ (robust), indicating the stability of the PMF solution (Zhou et al., 2019). Most of the residuals are distributed normally, ranging from -3 and +3, suggesting the model fits the data well.

123 **2.2.3 Other measurements**

124 The concentration of BC particles was measured by a single particle soot photometer (SP2, DMT Inc.). The operation 125 and data analysis procedures of the SP2 have been described elsewhere (Schwarz et al., 2008; Liu et al., 2010). The SP2 incandescence signal was calibrated for refractory BC (rBC) mass using the Aquadag black carbon particle standards
(Acheson Inc., USA) and corrected for ambient rBC with a factor of 0.75 (Laborde et al., 2012).

The size distribution and number concentration of aerosols with an electrical mobility diameter from 12 to 552 nm were also measured by a scanning mobility particle sizer (SMPS, Model 3081, impactor 50 % cut off at 0.677 μ m; CPC model 3775 at a flow rate 0.3 L min⁻¹). PM₁ mass concentration calculated based on the volume concentration measured by the SMPS agreed well with that from the sum of compositions by the HR-ToF-AMS and SP2 (r = 0.71, slope = 0.90).

The gas-phase species including NO_x, O₃, and CO were measured in real time by a series of Wuhan Tianhong analyzers (TH-2001H, TH-2003H, TH-2004H, respectively). These instruments were calibrated periodically with the corresponding standard gas to ensure the accuracy of the observation data. In addition, the meteorological parameters including temperature (T), relative humidity (RH), and wind speed and direction were recorded by an automatic weather station.

136 3 Results and discussion

137 **3.1 Chemical compositions of PM**₁

The time series of mass concentrations of BC, non-refractory-PM₁ (NR-PM₁) species (i.e., OA, SO_4^{2-} , NO_3^{-} , NH_4^+ , and 138 139 Cl⁻) and their relative contributions are summarized in Fig. 1. During this field observation, the mass concentrations of NR-PM₁ were in the range of 2.5–44.8 μ g m⁻³, with an average of 12.7 ± 5.7 μ g m⁻³, which is close to those observed in the 140 North China Plain in autumn 2019 (15.1 μ g m⁻³) (Li et al., 2021a), and was much lower than that (41.3 ± 42.7 μ g m⁻³) 141 observed in Beijing in autumn 2012 (Hu et al., 2017). The mean mass concentration of BC was $0.29 \pm 0.17 \ \mu g \ m^{-3}$ with the 142 range of 0.1–1.0 µg m⁻³. Among all species in NR-PM₁, OA contributed the major (49.2 %), indicating the dominant role of 143 144 OA in autumntime of PM₁ pollution in this region. Inorganic aerosol accounted for 50.8 % of NR-PM₁ in which sulfate was 145 the largest contributor (21.5 %), followed by nitrate (18.5 %), ammonium (10.7 %), and chloride (0.1 %), which has similar 146 relative contributions to those observed in field (Chen et al., 2021). The number concentration of all particles peaked at $95 \pm$ 38 nm in with regular growth during every field day. Fig. 1 also presents the temporal evolution variation in meteorological 147 148 parameters. During the field observation period, the average temperature and relative humidity (RH) were $19.3 \pm 3.6^{\circ}$ C 149 (11.2-29.4 °C) and $75.0 \% \pm 17.5 \% (25.0-99.0 \%)$, respectively. The O₃ concentration with an average of $30.7 \pm 13.0 \text{ ppb}$, 150 was highest 54.5 ppb at 15:00, likely due to the high temperature and enhanced photochemical processing (Zheng et al., 151 2021).

Fig. 1f shows the spatial distribution of aerosol optical depth (AOD) and 36 h backward trajectories during the campaign. Cluster-1 (C1) shows the circulated air mass with shortest transport distance for the past 36 h (36 % fraction); C2 is the northerly transported air mass (55 % fraction), over regions with higher AOD; C3 is the most rapid transport through

155 the northeast China and some coastal areas. It therefore separates the periods with more local air mass or influences from

156 regional transport. Fig. 1g shows the diurnal variations of solar radiation and RH, peaking in the day and night respectively.



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Figure 1: Time series of (a) mass concentrations of O₃ and NO_x; (b) ambient temperature (T) and relative humidity (RH); (c) number size distribution measured by the SMPS; (d) mass concentrations of key aerosol species (the green left y-axis represents OA, the right y-axis represents other aerosol species); (e) mass fractions of chemical species in non-refractory PM₁; (f) spatial distribution of mean aerosol optical depth (AOD) during the experiment and the clustered 36 h backward trajectories; (g) diurnal profiles of direct solar radiation and RH.

163 **3.2 Source attributed organic aerosol**

164 The resolved OA factors by the PMF analysis on the AMS measurements are shown in Fig. 2. The first factor is characterized by prominent hydrocarbon ion series of $C_nH_{2n-1}^+$ and $C_nH_{2n+1}^+$ (Fig. 2a), with a low O/C ratio (0.24), which is 165 generally considered to be related to the emissions of fossil fuel combustion and vehicle emissions (Huang et al., 2010; 166 Morgan et al., 2010; Ng et al., 2011b). The time series of the factor correlated well with NO_x (r = 0.71) and BC (r = 0.83) 167 (Fig. 2e and Table S1). The diurnal pattern of this factor showed peaks in the morning and afternoon rush-hour (Fig. 2i), 168 169 with a major increase from 5:30, reaching a peak value of 1.5 μ g m⁻³ at 8:30. The concentration gradually decreased around noontime due to boundary layer dilution until 15:00 and reached a minimum of 0.6 µg m⁻³. This spectrum also contained 170 some fragment markers for possible coal combustion OA (CCOA), i.e. $C_9H_7^+$ (m/z 115, r = 0.73) (Hu et al., 2013). This 171

172 factor was not distinctly resolved in this dataset maybe due to the urban nature of the site, where the traffic source may have 173 overwhelmed, due to the less significant coal combustion pollutants during the sampling period.

174 The second factor was characterized by m/z 55 (C₄H₇⁺, C₃H₃O⁺) and 57 (C₄H₉⁺, C₃H₅O⁺), accounting for 10 % and 3.5 175 % of the total spectrum, respectively (Fig. 2b), with the lowest O/C ratio among factors (0.16). This factor has a $C_{3}H_{3}O^{+}/C_{3}H_{5}O^{+}$ of 3.0 and $C_{4}H_{7}^{+}/C_{4}H_{9}^{+}$ of 2.0 (1.0 is usually for HOA), indicating it as cooking source rather than HOA 176 177 (Mohr et al., 2012). The correlation coefficient of the COA factor and marker fragment $C_6H_{10}O^+$ was 0.91, which is also similar to a previous study (Sun et al., 2011). The diurnal pattern of this factor showed a major peak during 18:00–20:00, 178 reaching up to 4.0 µg m⁻³ on average, in addition to a smaller peak at lunch time, which corroborated the diurnal cooking 179 180 activities. Previous studies reported COA accounted for 6.5–30 % of the total OA in urban (Rogge et al., 1991; Lanz et al., 181 2008; Allan et al., 2010; Xu et al., 2014). Here the concentrations of COA in OA were in the range of $0.5-4.5 \,\mu g \, m^{-3}$ and 182 accounted for 23 % of OA mass fraction on average. The cooking emission in the studying region was likely from charcoal-183 grilling activities, which was popular in the surrounding areas.

184 Besides the two primary OA, additional two oxygenated OA (OOA) factors are identified. According to the oxidation 185 state, OOA was separated into lower (OOA1) and more oxidized (OOA2) factors. OOA1 factor contained abundant oxygen-186 containing fragments, accounting for more than 50 % of the total mass spectrum, with an O/C of 0.42. In particular, the 187 oxygenated fragment containing one oxygen accounted for 39 %. The $C_2H_3O^+$ ion (m/z 43) is an important component of less 188 oxidized SOA (Wang et al., 2020a), which was highly correlated with OOA1 (r = 0.94). OOA1 showed lower concentration 189 during daytime but higher concentration during nighttime (Fig. 2k). Such diurnal variation was similar to that of RH ($\mathbf{r} =$ 190 0.45), but opposite to that of solar radiation, which was similar to a previous report (Sun et al., 2014). OOA1 may thus be 191 associated with aqueous reactions when high RH. The co-occurrence of nighttime peak of OOA1 and COA suggested the 192 primary source of cooking emission may have considerably contributed to the production of OOA1.

193 The factor OOA2 had the highest O/C of 0.72 and contained 61 % oxygen-containing fragments (Fig. 2d), which is very 194 similar to the spectra of OOA factor resolved in other cities (0.60 and 67 % in Mexico City, 0.80 and 66 % in Pasadena) 195 (Aiken et al., 2009; Hayes et al., 2013). This factor was correlated strongly with the fragment of CO_2^+ (r = 0.79). Compared to OOA1, this factor showed obvious diurnal variations with a major enhancement at around noontime (10:00–15:00), up to 196 197 2.9 µg m⁻³, indicating the photochemical production of SOA during daytime. The variation of OOA2 correlated with odd 198 oxygen ($O_x = O_3 + NO_2$, r = 0.70), agreeing with previous observations that oxidized OA had strong correlation with O_x 199 (Wood et al., 2010; Hu et al., 2016). Considering the high RH during the experiment (> 60 %, Fig. 1g), OOA2 factor may 200 have also experienced aqueous chemistry and showed good correlation with sulfate (r = 0.82).

C_xH_y C_xH_yO₁ C_xH_yO₂ H_yO₁ C_xH_yN_p C_xH_yO_zN_p Rush hour HOA 2.0 58% 10 O/C = 0.24(a) 8 HOA (µg/m N/C=0.022 1% 6. 4% 2 3% 4 -21% 13% 2 0.0 0.0 0 10 8 2.0x10 71% (i) COA (f) 8 8 O/C = 0.166 (b) (m/gu) 0% 6 N/C=0.008 6 ò 2% 4 1% COA of total E 7% 2 19% 0 0 0.8 8 ×12 4 (k) **00A1** 40% (g) С²Н 6 (c) O/C = 0.4200A1 (µg/m[~] 2 8 Ő. N/C=0.025 1% 3% 4 3% 39% 0 1 4% n 0.0 n 0.0 3.0 80 31% 120 (1) **OOA2** (h) 20-(d) O/C = 0.72OA2 (µg/m) 60 ò o 2.0 N/C=0.026 3% (ppb) 40 € 5% 37% 10 ġ 20 23% 0 0.0 25 50 75 100 0 6 12 18 24 10/9 10/13 10/17 10/21 10/25 10/29 11/2 m/z (amu) Hour of day Date and time



201 202 Figure 2: Results of source-attributed organic aerosols by the PMF analysis on the OA mass spectra measured by the HR-TOF-203 AMS. Mass spectra of PMF factors for (a) hydrocarbon-like OA (HOA); (b) cooking OA (COA); (c) oxygenated OA1 (OOA1); (d) 204 oxygenated OA2 (OOA2). The pies show the relative contributions of the six ion categories to each factor. (e-h) Temporal 205 variations of four OA factors (HOA, COA, OOA1 and OOA2) with each correlated species as NO_x, C₆H₁₀O⁺, C₂H₃O⁺ and O_x. (i-l) 206 Diurnal profiles of NO_x, BC, C₆H₁₀O⁺, C₂H₃O⁺, NO₃⁻ and O_x. The lines and whiskers denote the median, the 25th and 75th 207 percentiles at each hour, respectively.

208 3.3 Source attributed VOCs

209 Fig. 3 and Fig. S5 summarize the source attributed of VOCs and key indicators. Two primary and three secondary VOC 210 factors were identified. Here 109 species of VOCs from the PTR-TOF-MS were used, which are mostly oxygenated or 211 contain cyclic functional groups, in contrast with the GC-MS measurement which contains many aliphatic hydrocarbon 212 compounds and usually only decades of species were used for the PMF analysis (Zheng et al., 2021). These may lead to some unsolved primary VOC sources for the current analysis using the PTR-measured VOCs, if the source contained crucial 213 markers of aliphatic species, but this method may have great advantage in comprehensively resolving factors of secondary 214 VOCs. This may explain the less primary but more secondary VOC factors compared to conventional PMF-resolved VOC 215 sources (Cai et al., 2010). 216

The first factor was dominated by aromatic compounds, such as $C_{6}H_{6}$ (m/z 79.054) and $C_{7}H_{8}$ (m/z 93.070) in Fig. 3a. 217 They were well established markers for vehicle emissions (Gkatzelis et al., 2021) and had a good correlation with this factor 218

(r = 0.97 and 0.63, respectively Fig. S6). This factor showed peaks in the morning and afternoon rush-hour (Fig. 3k) and was 219

well correlated with HOA and NO_x (r = 0.63 and 0.53 Table S1). The concentration of this factor had a major increase during the early morning, reaching a peak value of 12.1 ± 1.3 ppb at 08:30, further corroborating the traffic source of this factor. The concentration decreased around noontime until 15:00 because of the dilution by well-developed boundary layer and in addition to its consumption through photochemical reaction due to intense solar radiation. The diurnal pattern was also high at night, although the peak was lower than early morning, the average concentration was higher by a factor of 2 than daytime. This suggests that traffic VOCs is more likely to participate in photochemical reaction, and other primary emissions may be precursors for nocturnal chemistry.

227 The second factor contained abundant aldehvdes such as C_3H_4O (acrolein), $C_3H_6O_2$ (hydroxyacetone), $C_6H_{12}O_2$ 228 (hexanal), and $C_7H_{14}O$ (heptanal) at m/z 57.069, 75.044, 101.096, and 115.112, respectively, as well as C_8H_{10} (C8-aromatics) 229 and C₉H₁₂ (C9-aromatics) at m/z 107.086 and 121.101 in VOC mass spectra (Fig. 3b), which are footprint VOCs identified 230 from primary cooking emissions during the charbroiling and frying (Klein et al., 2016). This factor had similar time series (Fig. 3g) with COA and had high correlation (r = 0.67, Table S1). The concentration of this factor decreased during the 231 232 daytime and increased after 18:00 with a peak value at 19:00 (17.2 ± 3.0 ppb). As shown in Fig. 31, the diurnal concentration 233 decreased strongly after emission and continued to decline throughout the night, suggesting that cooking VOCs may be major precursors and were consumed during night. 234

235 Besides the two primary VOCs, three secondary oxygenated VOCs (SecVOCs) factors are identified. These three 236 factors are not well correlated with any primary factors attributed by the OA, are thus considered to be mainly contributed by 237 secondary production. The daytime biogenic emission, e.g. isoprene, may also contribute to the SecVOC2 formation by 238 interacting with OH in the presence of NO_x (Lin et al., 2013), producing methacrylic acid epoxide (MAE) and methacrolein 239 (MACR) as intermediately involved in SOA formation (Gu et al., 2021). This factor was thus also contributed by this 240 primary source. SecVOC2 contributed a large fraction of $C_2H_4O_2$, $C_3H_6O_3$ and $C_4H_6O_3$ (Fig. 3c), which were gas-phase 241 carboxylic acids, the oxidation products from photochemical processes (Hartikainen et al., 2018; Li et al., 2021b). SecVOC2 242 had one clear peak at 13:00 (14.7 \pm 2.8 ppb Fig. 3m). The time series of SecVOC2 correlated well with OOA2 (r = 0.67), 243 following the variation of solar radiation. It had a rapid enhancement starting from 07:00 to 12:00 and declined continuously 244 after 13:00. This indicated that many of these species can be formed rapidly during daytime and may have a short lifetime 245 owing to the partitioning to the condensed phase and forming SOA.

SecVOC1 factor featured with some less-oxygenated VOCs e.g., $C_2H_2O_2$, $C_6H_{10}O_2$ and $C_{10}H_{14}O$ (Fig. 3d). This factor had a peak at 19:00 which was consistent with the primary cooking VOCs factor, but also increased throughout the night, peaking at midnight. This factor had a similar temporal trend with OOA1 (r = 0.76, Fig. 3i), which was less oxygenated than photochemistry dominated OOA2. Combining the features above, SecVOC1 tended to be contributed by some immediately reacted species from emissions in the late afternoon and early night. A particular factor (Fig. 3e) is significantly composed of large molecular weight (large-MW) oxidized VOCs, i.e. the mean contribution of ionic compounds with m/z > 120 was above 50 % in this factor (Fig. S5e), which was much higher than that in SecVOC2 (Fig. S5c). Fig. 3e shows its signature compounds of $C_8H_{14}O$, $C_6H_{12}O_3$ and $C_8H_4O_3$, and some are nitrogen-containing VOCs, such as $C_6H_5NO_3$ and $C_8H_9NO_3$. These VOCs with m/z > 120 tend to be intermediate-volatility organic compounds (IVOCs) as the estimated vapor saturation concentration ($log_{10}C^*$, 298 K) is less than 6.5 µg m⁻³ (Fig. S2). This factor is hereby termed as large-MW VOCs to indicate the fraction of IVOCs, which only require few oxidation steps to become semi-volatile (Robinson et al., 2007). Fig. 3o shows an increase of this factor at mid-night, later than the peak of SecVOC1, which may imply the ageing process in producing these VOCs.

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Figure 3. Source attributed VOCs by the PMF analysis measured by the PTR-TOF-MS. (a-e) Mass spectra of the five factors (traffic VOCs, cooking VOCs, secondary VOCs (SecVOC2, SecVOC1) and large molecular weight (MW) VOCs (large-MW VOCs), with major fingerprint peaks labelled in the mass spectra. (f-j) Temporal variations of the five VOCs factors and their respective correlated OA species. (k-o) Diurnal profiles of the VOC factors and their respective signature species. The lines and whiskers denote the median, the 25th and 75th percentiles at each hour, respectively.

266 **3.4 Oxidation process of organics in the day and night**

After source attribution of organics in aerosol and gas phase, we are able to identify the emission structure of primary sources and the consequent evolution. Given the identified primary traffic and cooking sources were emitted in the day and late afternoon respectively, this provides the potential opportunity to study the evolution of different primary sources in the potentially contrasting ageing mechanisms in the day and night. 271 Fig. 4a-d shows the temporal evolution of the key factors in a week to reflect the day and night ageing. The O/C ratio in 272 the day was higher than that at night by 0.2 (Fig. 4a), corresponding with the increase of highly oxygenated fragments of 273 AMS f_{44} (CO₂⁺) in the day (Fig. 4b). During night, the O/C ratio increased with the increase moderately oxygenated fragment 274 f_{43} (C₂H₃O⁺). The diurnal variation of atomic ratio O/C of OA (Fig. 4e) showed a slight decrease at morning rush-hour and 275 sharp drop at 18:00 (by 0.2) due to the significant contribution of the traffic and cooking sources respectively, corresponding 276 with the two steps of increase at each time. This variation of O/C clearly showed the daytime and nighttime ageing processes of OA for different primary sources. Notably, the O/C showed peak value of 0.54 at 15:00, implying the importance of 277 278 photooxidation in producing highly oxygenated OA.

279 Considering the diurnal pattern of anthropogenic activities, the traffic emission at rush-hour is deemed to be the major 280 source in contributing to the daytime production of OOA. The ratio OOA2/HOA is thus used to indicate the daytime 281 oxidation of OOA. As Fig. 4d and 4f showed, OOA2/HOA had a clear peak during daytime, and increased after 8:00 and 282 peaked at 14:00. After 15:00, the ratio gradually decreased to minimum at 20:00 and maintained to be low throughout the 283 night. This clearly demonstrated the photooxidation in producing OOA2 from oxidizing HOA. Fig. 4c and 4g give a few 284 examples of photooxidation in gas phase: the toluene showed a production of $C_3H_6O_3$ (hydroxypropionic acid) and $C_4H_6O_3$ 285 (acetic anhydride) by a factor of 2 in 3 hours (Fig. S7c); the C_3H_4O (acrolein) produced the oxidized product $C_2H_2O_2$ 286 (glyoxal) by a factor of 1.3 (Fig. 4h). All of these reacted species are from traffic VOCs and the corresponding products are 287 from SecVOC2. The daytime biogenic emission, e.g. isoprene, may also contribute to the SecVOC2 formation by interacting with OH in the presence of NO_x (Lin et al., 2013), producing methacrylic acid epoxide (MAE) as intermediately involved in 288 289 SOA formation, where a considerable nitrogen content was also found in the OOA2 factor (N/C = 0.026). Overall, the highly 290 oxidized OOA2 (O/C = 0.72) is considered to be mainly contributed by the traffic source, via the oxidation of VOCs and 291 partitioning to condensed phase, direct oxidation on HOA through heterogenous oxidation (Guo et al., 2020), or VOCs 292 evaporated from HOA and further condensation after oxidation (Zhao et al., 2015). All factors may have contributed to the davtime production of OOA2. 293

294 The nocturnal oxidation is mainly contributed by the sources emitting from late afternoon throughout the midnight. 295 Both traffic and cooking sources contributed to the emission since late afternoon, with cooking source as the predominant 296 contributor in both aerosol (Fig. 2j) and gas (Fig. 3l) phases. The ratio between the nighttime OOA1 and cooking aerosol 297 (OOA1/COA) is therefore used to indicate the nocturnal oxidation of SOA (Fig. 4d and j). There was a sudden increase of 298 OOA1/COA during the daytime because COA was consumed rapidly in the afternoon after a small amount of emission at 299 noon (Fig. 2j). The lowest OOA1/COA at 0.5 corresponded with the fresh cooking emission at 18:00, and kept increasing 300 until peaking in the early morning at 6:00 up to 3, which was an increase by a factor of 6 compared to the minima (Fig. 4j). In addition to evidence for daytime reaction, Fig. 4h and Fig. S7d also gave evidence for certain reacted species during the 301 302 ageing at nighttime. The first-generation oxidation products from acrolein (ACR $C_3H_4OH^+$, m/z 57.033) are glyoxal 303 $(C_2H_2O_2, m/z 59.036)$ and formaldehyde (HCHOH⁺, m/z 31.018) according to the database in Master Chemical Mechanism

- 304 (MCM). This night oxidation is also evidenced by the formation of some nitrated organic compounds and ketone 305 compounds, such as $C_6H_5NO_3$ (nitrophenol, NP, Fig. 41) and $C_6H_{10}O_2$ (hexanedione, Fig. 4k) produced from phenol (C_6H_6O) 306 and hexanal ($C_6H_{12}O$), respectively. Notably, an important fraction of large molecular weight-VOCs (which are mostly 307 intermediate VOCs, Fig. S2) peaked at 3:00-4:00 (Fig. 3o), consistent with the variation of OOA1/COA during night. Some 308 of these nitrated and oxygenated IVOCs may have been further oxidized and partitioned to aerosol phase, contributing to the 309
- OOA1. Given the larger molecular usually has a lower O/C ratio (Hatch et al., 2017) (because of a higher content of carbon),
- 310 this may explain the lower O/C observed for nighttime formed OOA1 (O/C = 0.42), than OOA2 produced by daytime
- 311 photooxidation (O/C = 0.72). Notably, nighttime SOA had a high N/C (0.025), implying the NO₃·-initiated from cooking
- 312 emissions oxidation, which was different from the organic nitrate formation mechanism during the daytime.



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314 Figure 4. Time series showing the aging of aerosols and gases, with grey vertical bars denoting nighttime (00:00–06:00), (a) the 315 oxygen to carbon ratio (O/C); (b) the fragment fraction at m/z 44 (f_{44}), m/z 43 (f_{43}); (c) concentration ratio of C₆H₁₀O₂ to C₆H₁₂O (hexanedione to hexanal), glyoxal to acrolein, and formaldehyde to acrolein; (d) concentration ratio of OOA1 to COA and OOA2 316 317 to HOA. Diurnal variations of key species showing the daytime photooxidation and nighttime oxidation, (e) ratio of O/C; (f) ratio 318 of OOA2 to HOA; (g) concentration of C₇H₈ and ratio of C₃H₆O₃ to C₇H₈ (hydroxypropionic acid to toluene); (h) concentration of 319 C₃H₄O and ratio of C₂H₂O₂ to C₃H₄O (glyoxal to acrolein); (i) particle diameter (D_p); (j) ratio of OOA1 to COA; (k) concentration 320 of C₆H₁₂O and ratio of C₆H₁₀O₂ to C₆H₁₂O (hexanedione to hexanal); (I) concentration of C₆H₆O and ratio of C₆H₅NO₃ to C₆H₆O 321 (nitrophenol to phenol).

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Fig. 5 summarizes the key indicators for the contrasting daytime and nighttime oxidation process. The concentrations are normalized by ΔCO , where ΔCO is the total CO concentration subtracted by the background concentration (1th) 325 percentile of the dataset), to indicate the variation of species regardless of the boundary layer evolution, wind speed, wind 326 direction and temperature (Gouw and Jimenez, 2009). The odd oxygen O_x (O₃ + NO₂) has been widely used to generally 327 indicate the activity of daytime photochemistry (Hu et al., 2017), and the enhanced moisture is the main driving factor for 328 nighttime chemistry. The O_x concentration and RH are therefore used as references with which the variations of species were 329 correlated in the day and night, respectively. As Fig. 5a showed, the traffic primary emissions in both gas (traffic VOCs) and 330 aerosol phase (HOA) declined with increased O_x by 60 % and 40 % respectively, suggesting their roles as precursors in the 331 daytime reaction. The produced species are oxygenated SecVOC2 and OOA2, showing enhancement with O_x, peaking at 332 midday-afternoon. This process was rapid as the SOA production by a factor of 2.5 and the oxygenated VOCs production by 333 a factor of 1.7 within 6 hours. The traffic VOCs are widely observed to contribute to SOA production, with aromatic 334 compounds serving as key precursors (Fang et al., 2021). The semi-volatile nature of HOA means it could be evaporated to 335 gas phase and further oxidized to recondense as SOA (Robinson et al., 2007). The decrease rate of HOA with increased 336 photochemical age was also found in urban environment (Zhu et al., 2021), generally consistent with the reaction rate in this 337 study. Here we linked the declining rate as a function of photochemical activities for both reacted aerosol and gas phases for 338 traffic sources. The gases evaporated from aerosol phase (especially under higher temperature when increased saturation 339 pressure for semi-volatile or intermediate volatile species) and primary VOCs may be simultaneously involved in the 340 photooxidation, further contributing to the SOA formation.

341 For nocturnal oxidation shown in Fig. 5b, the reacted species are cooking VOCs and COA (decrease by 35 % and 77 %, respectively), producing SecVOC1 and large-MW oxygenated VOCs, with an increase of night SOA formation by a factor 342 343 of 1.7. The nocturnal processes may have largely involved aqueous reactions, because the variations of reacted or produced 344 species were highly correlated with RH (r > 0.95, Fig. 5b). The large-MW VOCs (mostly IVOCs) increased by 50 % and 345 reached maxima when highest RH. This suggests the moisture may have been involved in converting some primary VOCs to IVOCs, which further contribute to the SOA production during nighttime. Previous studies also found the oxidation of 346 IVOCs from cooking sources can be an important source of SOA (Zhang et al., 2020). The evidence is given here that 347 348 organic aerosols and gases from cooking emission had been reacted and contributed to SOA. An approximate production rate 349 of 0.2 μ g m⁻³ h⁻¹ of OOA from cooking source can be obtained by considering the ageing time of ~10 h (from COA peak 350 18:00 to OOA1 peak 4:00, which was generally consistent with previous laboratory work using gas precursors from cooking 351 sources $(0.07-0.5 \,\mu g \,\mathrm{m}^{-3} \,\mathrm{h}^{-1})$ (Liu et al., 2017). Notably, daytime SOA had a higher oxidation state, implying the importance 352 of photooxidation in producing highly oxidized OA. This may be because of the high temperature at daytime and a species 353 may require a lower volatility (hereby more oxygenated) to be in condensed phase than at night.

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Figure 5. Daytime and nighttime evolution of key species against O_x and RH respectively. (a) Daytime oxidation showing the primary emissions of traffic source in gas and aerosol phases including HOA, traffic VOCs; and secondary products including SecVOC2 and OOA2, with concentrations all normalized by Δ CO. (b) Nighttime oxidation showing the primary emissions of cooking source in gas and aerosol phases including COA and cooking VOCs; and secondary products including SecVOC1, large-MW VOC and OOA1, with concentrations all normalized by Δ CO.

361 4 Conclusion

362 In this study, organic gases and aerosols were concurrently characterized through online mass spectrometers at a megacity in central China. Through the factorization analysis on the organic mass spectra, two principal sources-the traffic 363 364 and cooking sources were identified for both aerosol and gas phases, hereby the reacted and produced species between phases were interlinked. We observed clear evidence of daytime and nighttime oxidation of source-attributed OA and VOCs. 365 Daytime photooxidation caused 60 % decrease of primary aerosol and 40 % of primary VOCs reduction for traffic sources, 366 367 producing oxygenated SecVOCs by a factor of 1.7 and OOA by a factor of 2.5, in a 6 hours photochemical ageing. Nocturnal ageing caused a reduction of primary OA (by 77 %) and primary VOCs (by 35 %) from cooking sources, 368 369 producing oxygenated VOCs and OOA by a factor of 1.4 and 1.7, respectively. In particular, larger molecular IVOCs 370 produced (by a factor of 1.7) at night may importantly contribute to the OOA. This implies primary species in aerosol and gas phases both contribute to the production of OOA. A higher oxidation state of OOA from daytime photooxidation was 371 372 found than nighttime, suggesting different compositions of produced OOA modulated by solar radiation and moisture, 373 respectively. Vehicle and cooking emissions are the major contributors of organic aerosols in urban areas, especially in 374 megacities. These results provide direct observations about the reaction rate for primary precursors and production rate for 375 secondary aerosols, as influenced by primary sources and meteorological conditions. The environmental policy making 376 should therefore consider the respective primary sources and ageing mechanisms for local and regional atmospheric 377 environmental problems. The particular regulation should be placed to avoid the formation of daytime highly-oxidized

378 species when high solar radiation, which may contribute to the reactive oxygen species and exert adverse health impacts

379 (Tao et al., 2003; Verma et al., 2009).

380 Data availability

381 The data in this study are available from the corresponding author upon request.

382 Author contribution

383 DL, SK, and SL led and designed the study. SL, YW, HZ, YC, SZ and DH set up and conducted the experiment. SL,

DL, YW, KH, XJ and SD contributed to the data analysis. QL, DZ, JS provided technical support and assistance. SL and DL
 wrote the manuscript. All authors read and approved the final manuscript.

386 Competing interests

387 The authors declare no competing interests.

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392 References

- Adler, G., Flores, J. M., Abo Riziq, A., Borrmann, S., and Rudich, Y.: Chemical, physical, and optical evolution of biomass burning
 aerosols: a case study, Atmos. Chem. Phys., 11, 1491-1503, <u>https://doi.org/10.5194/acp-11-1491-2011</u>, 2011.
- 395 Aiken, A., Salcedo, D., Cubison, M., Huffman, J., DeCarlo, P., Ulbrich, I., Docherty, K., Sueper, D., Kimmel, J., Worsnop, D., A, T., M,
- N., Stone, E., Schauer, J., Volkamer, R., Fortner, E., de Foy, B., Wang, J., Laskin, A., and Jimenez, J.: Mexico City Aerosol Analysis
 During Milagro Using High Resolution Aerosol Mass Spectrometry at the Urban Supersite (T0) Part 1: Fine Particle Composition
 and Organic Source Apportionment, Atmos. Chem. Phys., 9, 6633-6653, https://doi.org/10.5194/acpd-9-8377-2009, 2009.
- Allan, J. D., Williams, P., Morgan, W., Martin, C., Flynn, M., Lee, J., Nemitz, E., Phillips, G., Gallagher, M., and Coe, H.: Contributions
 from transport, solid fuel burning and cooking to primary organic aerosols in two UK cities, Atmos. Chem. Phys., 10, 647-668,
 https://doi.org/10.5194/acp-10-647-2010, 2010.
- Bruns, E. A., El Haddad, I., Slowik, J. G., Kilic, D., Klein, F., Baltensperger, U., and Prévôt, A. S. H.: Identification of significant
 precursor gases of secondary organic aerosols from residential wood combustion, Sci. Rep., 6, 1-9, <u>https://doi.org/10.1038/srep27881</u>,
 2016.
- Cai, C., Geng, F., Tie, X., Yu, Q., and An, J.: Characteristics and source apportionment of VOCs measured in Shanghai, China, Atmos.
 Environ., 44, 5005-5014, https://doi.org/10.1016/j.atmosenv.2010.07.059, 2010.

- 407 Canagaratna, M. R., Jayne, J. T., Jimenez, J. L., Allan, J. D., Alfarra, M. R., Zhang, Q., Onasch, T. B., Drewnick, F., Coe, H.,
 408 Middlebrook, A., Delia, A., Williams, L. R., Trimborn, A. M., Northway, M. J., DeCarlo, P. F., Kolb, C. E., Davidovits, P., and
- 409 Worsnop, D. R.: Chemical and microphysical characterization of ambient aerosols with the aerodyne aerosol mass spectrometer,
- 410 Mass Spectrom Rev., 26, 185-222, <u>https://doi.org/10.1002/mas.20115</u>, 2007.
- 411 Canagaratna, M. R., Jimenez, J. L., Kroll, J. H., Chen, Q., Kessler, S. H., Massoli, P., Hildebrandt Ruiz, L., Fortner, E., Williams, L. R.,
- 412 Wilson, K. R., Surratt, J. D., Donahue, N. M., Javne, J. T., and Worsnop, D. R.: Elemental ratio measurements of organic compounds
- using aerosol mass spectrometry: characterization, improved calibration, and implications, Atmos. Chem. Phys., 15, 253-272,
 https://doi.org/10.5194/acp-15-253-2015, 2015.
- Cappellin, L., Karl, T., Probst, M., Ismailova, O., Winkler, P. M., Soukoulis, C., Aprea, E., Mark, T. D., Gasperi, F., and Biasioli, F.: On
 quantitative determination of volatile organic compound concentrations using proton transfer reaction time-of-flight mass
 spectrometry, Environ. Sci. Technol., 46, 2283-2290, https://doi.org/10.1021/es203985t, 2012.
- Chen, T., Liu, J., Ma, Q., Chu, B., Zhang, P., Ma, J., Liu, Y., Zhong, C., Liu, P., Wang, Y., Mu, Y., and He, H.: Measurement report:
 Effects of photochemical aging on the formation and evolution of summertime secondary aerosol in Beijing, Atmos. Chem. Phys., 21, 1341-1356, https://doi.org/10.5194/acp-21-1341-2021, 2021.
- 421 Claeys, M., Graham, B., Vas, G., Wu, W., Vermeylen, R., Pashynska, V., Cafmeyer, J., Cuyon, P., Andreae, M. O., and Artaxo, P.:
 422 Formation of secondary organic aerosols through photooxidation of isoprene, Science, 303, 1173-1176, 423 https://doi.org/10.1126/science.1092805, 2004.
- 424 Clark, C. H., Kacarab, M., Nakao, S., Asa-Awuku, A., Sato, K., and Cocker III, D. R.: Temperature effects on secondary organic aerosol
 425 (SOA) from the dark ozonolysis and photo-oxidation of isoprene, Environ. Sci. Technol., 50, 5564-5571,
 426 https://doi.org/10.1021/acs.est.5b05524, 2016.
- 427 Decarlo, P. F., Kimmel, J. R., Trimborn, A., Northway, M. J., Jayne, J. T., Aiken, A. C., Gonin, M., Fuhrer, K., Horvath, T., and Docherty,
 428 K. S.: Field-deployable, high-resolution, time-of-flight aerosol mass spectrometer, Anal. Chem., 78, 8281-8289,
 429 https://doi.org/10.1021/ac061249n, 2006.
- 430 Decarlo, P. F., Ulbrich, I. M., Crounse, J., Foy, B. D., Dunlea, E. J., Aiken, A. C., Knapp, D., Weinheimer, A. J., Campos, T., and
 431 Wennberg, P. O.: Investigation of the sources and processing of organic aerosol over the Central Mexican Plateau from aircraft
 432 measurements during MILAGRO, Atmos. Chem. Phys., 10, 5257–5280, https://doi.org/10.5194/acp-10-5257-2010, 2010.
- Donahue, N. M., Robinson, A. L., Stanier, C. O., and Pandis, S. N.: Coupled partitioning, dilution, and chemical aging of semivolatile
 organics, Environ. Sci. Technol., 40, 2635-2643, <u>https://doi.org/10.1021/es052297c</u>, 2006.
- 435 Draxler, R. and Hess, G.: Description of the HYSPLIT_4 modelling system, NOAA Tech. Mem. ERL ARL-224, 1997.
- 436 Fang, H., Huang, X., Zhang, Y., Pei, C., Huang, Z., Wang, Y., Chen, Y., Yan, J., Zeng, J., Xiao, S., Luo, S., Li, S., Wang, J., Zhu, M., Fu,
- X., Wu, Z., Zhang, R., Song, W., Zhang, G., Hu, W., Tang, M., Ding, X., Bi, X., and Wang, X.: Measurement report: Emissions of
 intermediate-volatility organic compounds from vehicles under real-world driving conditions in an urban tunnel, Atmos. Chem.
- 439 Phys., 21, 10005-10013, https://doi.org/10.5194/acp-21-10005-2021, 2021.
- 440 Gkatzelis, G. I., Coggon, M. M., McDonald, B. C., Peischl, J., Gilman, J. B., Aikin, K. C., Robinson, M. A., Canonaco, F., Prevot, A. S.
- H., Trainer, M., and Warneke, C.: Observations confirm that volatile chemical products are a major source of petrochemical
 emissions in US cities, Environ. Sci. Technol., 55, 4332-4343, <u>https://doi.org/10.1021/acs.est.0c05471</u>, 2021.
- 443 Goldstein, A. H. and Galbally, I. E.: Known and Unexplored Organic Constituents in the Earth's Atmosphere, Environ. Sci. Technol., 41,
- 444 1514-1521, <u>https://doi.org/10.1021/es072476p</u>, 2007.

- Gouw, J. D. and Jimenez, J. L.: Organic aerosols in the earth's atmosphere, Environ. Sci. Technol., 43, 7614-7618,
 https://doi.org/10.1021/es9006004, 2009.
- 447 Gu, C., Wang, S., Zhu, J., Wu, S., Duan, Y., Gao, S., and Zhou, B.: Investigation on the urban ambient isoprene and its oxidation 448 processes, Atmos. Environ., 270, 118870, https://doi.org/10.1016/j.atmosenv.2021.118870, 2021.
- 449 Guo, J., Zhou, S., Cai, M., Zhao, J., Song, W., Zhao, W., Hu, W., Sun, Y., He, Y., Yang, C., Xu, X., Zhang, Z., Cheng, P., Fan, Q., Hang,
- 450 J., Fan, S., Wang, X., and Wang, X.: Characterization of submicron particles by time-of-flight aerosol chemical speciation monitor
- (ToF-ACSM) during wintertime: aerosol composition, sources, and chemical processes in Guangzhou, China, Atmos. Chem. Phys.,
 20, 7595-7615, https://doi.org/10.5194/acp-20-7595-2020, 2020.
- Hallquist, M., Wenger, J. C., Baltensperger, U., Rudich, Y., Simpson, D., Claeys, M., Dommen, J., Donahue, N. M., George, C.,
 Goldstein, A. H., Hamilton, J. F., Herrmann, H., Hoffmann, T., Iinuma, Y., Jang, M., Jenkin, M. E., Jimenez, J. L., Kiendler-Scharr,
 A., Maenhaut, W., Mcfiggans, G., Mentel, T. F., Monod, A., Prevot, A. S. H., Seinfeld, J. H., Surratt, J. D., Szmigielski, R., and
 Wildt, J.: The formation, properties and impact of secondary organic aerosol: current and emerging issues, Atmos. Chem. Phys., 9,
- 457 5155-5236, https://doi.org/10.5194/acp-9-5155-2009, 2009.
- 458 Hartikainen, A., Yli-Pirilä, P., Tiitta, P., Leskinen, A., Kortelainen, M., Orasche, J., Schnelle-Kreis, J., Lehtinen, K., Zimmermann, R.,
- Jokiniemi, J., and Sippula, O.: Volatile organic compounds from logwood combustion: Emissions and transformation under dark and
 photochemical aging conditions in a smog chamber, Environ. Sci. Technol., 52, 4979–4988, <u>https://doi.org/10.1021/acs.est.7b06269</u>,
 2018.
- Hatch, L. E., Yokelson, R. J., Stockwell, C. E., Veres, P. R., Simpson, I. J., Blake, D. R., Orlando, J. J., and Barsanti, K. C.: Multiinstrument comparison and compilation of non-methane organic gas emissions from biomass burning and implications for smokederived secondary organic aerosol precursors, Atmos. Chem. Phys., 17, 1471-1489, https://doi.org/10.5194/acp-17-1471-2017, 2017.
- 465 Hayes, P. L., Ortega, A. M., Cubison, M. J., Froyd, K. D., Zhao, Y., Cliff, S. S., Hu, W. W., Toohey, D. W., Flynn, J. H., Lefer, B. L., 466 Grossberg, N., Alvarez, S., Rappenglück, B., Taylor, J. W., Allan, J. D., Holloway, J. S., Gilman, J. B., Kuster, W. C., de Gouw, J. 467 A., Massoli, P., Zhang, X., Liu, J., Weber, R. J., Corrigan, A. L., Russell, L. M., Isaacman, G., Worton, D. R., Kreisberg, N. M., 468 Goldstein, A. H., Thalman, R., Waxman, E. M., Volkamer, R., Lin, Y. H., Surratt, J. D., Kleindienst, T. E., Offenberg, J. H., 469 Dusanter, S., Griffith, S., Stevens, P. S., Brioude, J., Angevine, W. M., and Jimenez, J. L.: Organic aerosol composition and sources 470 Res.-Atmos., in Pasadena, California, during the 2010 CalNex campaign, J. Geophys. 118. 9233-9257. 471 https://doi.org/10.1002/jgrd.50530, 2013.
- Heald, C. L., Jacob, D. J., Park, R. J., Russell, L. M., Huebert, B. J., Seinfeld, J. H., Liao, H., and Weber, R. J.: A large organic aerosol
 source in the free troposphere missing from current models, Geophys. Res. Lett., 32, <u>https://doi.org/10.1029/2005gl023831</u>, 2005.
- Hu, D., Liu, D., Kong, S., Zhao, D., Wu, Y., Li, S., Ding, S., Zheng, S., Cheng, Y., Hu, K., Deng, Z., Wu, Y., Tian, P., Liu, Q., Huang, M.,
 and Ding, D.: Direct quantification of droplet activation of ambient black carbon under water supersaturation, J. Geophys. Res.-
- 476 Atmos., 126, <u>https://doi.org/10.1029/2021jd034649</u>, 2021.
- Hu, W., Hu, M., Hu, W.-W., Zheng, J., Chen, C., Wu, Y., and Guo, S.: Seasonal variations in high time-resolved chemical compositions,
 sources, and evolution of atmospheric submicron aerosols in the megacity Beijing, Atmos. Chem. Phys., 17, 9979-10000,
 https://doi.org/10.5194/acp-17-9979-2017, 2017.
- 480 Hu, W., Hu, M., Hu, W., Jimenez, J., Yuan, B., Chen, W., Wang, M., Wu, Y., Chen, C., Wang, Z., Peng, J., Zeng, L., and Shao, M.:
- 481 Chemical composition, sources, and aging process of submicron aerosols in Beijing: Contrast between summer and winter, Journal of
- 482 Geophysical Research: Atmospheres, 121, 1955-1977, <u>https://doi.org/10.1002/2015JD024020</u>, 2016.

- Hu, W. W., Hu, M., Yuan, B., Jimenez, J. L., Tang, Q., Peng, J. F., Hu, W., Shao, M., Wang, M., Zeng, L. M., Wu, Y. S., Gong, Z. H.,
 Huang, X. F., and He, L. Y.: Insights on organic aerosol aging and the influence of coal combustion at a regional receptor site of
 central eastern China, Atmos. Chem. Phys., 13, 10095-10112, https://doi.org/10.5194/acp-13-10095-2013, 2013.
- Huang, L., Wang, Q., Wang, Y., Emery, C., Zhu, A., Zhu, Y., Yin, S., Yarwood, G., Zhang, K., and Li, L.: Simulation of secondary
 organic aerosol over the Yangtze River Delta region: The impacts from the emissions of intermediate volatility organic compounds
 and the SOA modeling framework, Atmos. Environ., 246, 118079, https://doi.org/10.1016/j.atmosenv.2020.118079, 2021.
- 489 Huang, R. J., Zhang, Y., Bozzetti, C., Ho, K. F., Cao, J. J., Han, Y., Daellenbach, K. R., Slowik, J. G., Platt, S. M., Canonaco, F., Zotter,
- 490 P., Wolf, R., Pieber, S. M., Bruns, E. A., Crippa, M., Ciarelli, G., Piazzalunga, A., Schwikowski, M., Abbaszade, G., Schnelle-Kreis,
- J., Zimmermann, R., An, Z., Szidat, S., Baltensperger, U., El Haddad, I., and Prevot, A. S.: High secondary aerosol contribution to
 particulate pollution during haze events in China, Nature, 514, 218-222, https://doi.org/10.1038/nature13774, 2014.
- Huang, X. F., He, L. Y., Hu, M., Canagaratna, M. R., Sun, Y., Zhang, Q., Zhu, T., Xue, L., Zeng, L. W., Liu, X. G., Zhang, Y. H., Jayne,
 J. T., Ng, N. L., and Worsnop, D. R.: Highly time-resolved chemical characterization of atmospheric submicron particles during 2008
 Beijing Olympic Games using an Aerodyne High-Resolution Aerosol Mass Spectrometer, Atmos. Chem. Phys., 10, 8933-8945,
 https://doi.org/10.5194/acp-10-8933-2010, 2010.
- Jayne, J. T., Leard, D. C., Zhang, X., Davidovits, P., Smith, K. A., Kolb, C. E., and Worsnop, D. R.: Development of an aerosol mass
 spectrometer for size and composition analysis of submicron particles, Aerosol Sci. Tech., 33, 49-70,
 https://doi.org/10.1080/027868200410840, 2000.
- Jokinen, T., Berndt, T., Makkonen, R., Kerminen, V. M., and Sipil, M.: Production of extremely low volatile organic compounds from
 biogenic emissions: Measured yields and atmospheric implications, P. Natl. Acad. Sci. USA, 112, 7123-7128,
 https://doi.org/10.1073/pnas.1423977112, 2015.
- Klein, F., Platt, S. M., Farren, N. J., Detournay, A., Bruns, E. A., Bozzetti, C., Daellenbach, K. R., Kilic, D., Kumar, N. K., Pieber, S. M.,
 Slowik, J. G., Temime-Roussel, B., Marchand, N., Hamilton, J. F., Baltensperger, U., Prevot, A. S., and El Haddad, I.:
 Characterization of gas-phase organics using proton transfer reaction time-of-flight mass spectrometry: cooking emissions, Environ.
 Sci. Technol., 50, 1243-1250, https://doi.org/10.1021/acs.est.5b04618, 2016.
- Kroll, J. and Seinfeld, J.: Chemistry of secondary organic aerosol: Formation and evolution of low-volatility organics in the atmosphere,
 Atmos. Environ., 42, 3593-3624, https://doi.org/10.1016/j.atmosenv.2008.01.003, 2008.
- Laborde, M., Schnaiter, M., Linke, C., Saathoff, H., Naumann, K.-H., Möhler, O., Berlenz, S., Wagner, U., Taylor, J., Liu, D., Flynn, M.,
 Allan, J., Coe, H., Heimerl, K., Dahlkötter, F., Weinzierl, B., Wollny, A. G., Zanatta, M., Cozic, J., and Gysel, M.: Single Particle
- 511 Soot Photometer intercomparison at the AIDA chamber, Atmos. Meas. Tech., 5, 3077-3097, <u>https://doi.org/10.5194/amt-5-3077-</u> 512 2012, 2012.
- Lanz, V. A., Alfarra, M. R., Baltensperger, U., Buchmann, B., Hueglin, C., Szidat, S., Wehrli, M. N., Wacker, L., Weimer, S., and Caseiro,
 A.: Source attribution of submicron organic aerosols during wintertime inversions by advanced factor analysis of aerosol mass
 spectra, Environ. Sci. Technol., 42, 214-220, https://doi.org/10.1021/es0707207, 2008.
- 516 Li, J., Cao, L., Gao, W., He, L., Yan, Y., He, Y., Pan, Y., Ji, D., Liu, Z., and Wang, Y.: Seasonal variations in the highly time-resolved
- aerosol composition, sources and chemical processes of background submicron particles in the North China Plain, Atmos. Chem.
 Phys., 21, 4521-4539, https://doi.org/10.5194/acp-21-4521-2021, 2021a.
 - 19

- Li, S., Liu, D., Hu, D., Kong, S., Wu, Y., Ding, S., Cheng, Y., Qiu, H., Zheng, S., Yan, Q., Zheng, H., Hu, K., Zhang, J., Zhao, D., Liu, Q.,
 Sheng, J., Ye, J., He, H., and Ding, D.: Evolution of organic aerosol from wood smoke influenced by burning phase and solar
 radiation, J. Geophys. Res.-Atmos., 126, https://doi.org/10.1029/2021JD034534, 2021b.
- Li, Y., Pöschl, U., and Shiraiwa, M.: Molecular corridors and parameterizations of volatility in the chemical evolution of organic aerosols,
 Atmos. Chem. Phys., 16, 3327-3344, https://doi.org/10.5194/acp-16-3327-2016, 2016.
- Lin, G., Penner, J. E., Sillman, S., Taraborrelli, D., and Lelieveld, J.: Global modeling of SOA formation from dicarbonyls, epoxides,
 organic nitrates and peroxides, Atmos. Chem. Phys., 12, 4743-4774, https://doi.org/10.5194/acp-12-4743-2012, 2012.
- 526 Lin, Y. H., Zhang, H., Pye, H. O., Zhang, Z., Marth, W. J., Park, S., Arashiro, M., Cui, T., Budisulistiorini, S. H., Sexton, K. G., Vizuete,
- W., Xie, Y., Luecken, D. J., Piletic, I. R., Edney, E. O., Bartolotti, L. J., Gold, A., and Surratt, J. D.: Epoxide as a precursor to
 secondary organic aerosol formation from isoprene photooxidation in the presence of nitrogen oxides, P. Natl. Acad. Sci. USA, 110,
 6718-6723, https://doi.org/10.1073/pnas.1221150110, 2013.
- Liu, D., Flynn, M., Gysel, M., Targino, A., Crawford, I., Bower, K., Choularton, T., Jurányi, Z., Steinbacher, M., Hüglin, C., Curtius, J.,
 Kampus, M., Petzold, A., Weingartner, E., and Baltensperger, U.: Single particle characterization of black carbon aerosols at a
 tropospheric alpine site in Switzerland, Atmos. Chem. Phys., 10, 7389-7407, https://doi.org/10.5194/acpd-10-8765-2010, 2010.
- Liu, T., Wang, Z., Huang, D. D., Wang, X., and Chan, C. K.: Significant production of secondary organic aerosol from emissions of
 heated cooking oils, Environ. Sci. Technol. Lett., 5, 32-37, <u>https://doi.org/10.1021/acs.estlett.7b00530</u>, 2017.
- 535 Middlebrook, A. M., Bahreini, R., Jimenez, J. L., and Canagaratna, M. R.: Evaluation of composition-dependent collection efficiencies for 536 the aerodyne aerosol spectrometer using field data. Aerosol Sci. Tech.. 46, 258-271, mass 537 https://doi.org/10.1080/02786826.2011.620041.2012.
- Mohr, C., DeCarlo, P. F., Heringa, M. F., Chirico, R., Slowik, J. G., Richter, R., Reche, C., Alastuey, A., Querol, X., Seco, R., Peñuelas,
 J., Jiménez, J. L., Crippa, M., Zimmermann, R., Baltensperger, U., and Prévôt, A. S. H.: Identification and quantification of organic
 aerosol from cooking and other sources in Barcelona using aerosol mass spectrometer data, Atmos. Chem. Phys., 12, 1649-1665,
 https://doi.org/10.5194/acp-12-1649-2012, 2012.
- 542 Moody, J. L. and Galloway, J. N.: Quantifying the relationship between atmospheric transport and the chemical composition of 543 precipitation on Bermuda, Tellus B, 40, 463-479, https://doi.org/10.1111/j.1600-0889.1988.tb00117.x, 1988.
- Morgan, W., Allan, J., Bower, K., Highwood, E., Liu, D., McMeeking, G., Northway, M., Williams, P., Krejci, R., and H, C.: Airborne
 measurements of the spatial distribution of aerosol chemical composition across Europe and evolution of the organic fraction, Atmos.
 Chem. Phys., 10, https://doi.org/10.5194/acpd-9-27215-2009, 2010.
- Ng, N., Jimenez, J., Chhabra, P., Seinfeld, J., and Worsnop, D.: Changes in organic aerosol composition with aging inferred from aerosol
 mass spectra, Atmos. Chem. Phys., 11, 6465-6474, https://doi.org/10.5194/acp-11-6465-2011, 2011a.
- Ng, N. L., Canagaratna, M. R., Jimenez, I. L., Zhang, Q., Ulbrich, I. M., and Worsnop, D. R.: Real-time methods for estimating organic
 component mass concentrations from aerosol mass spectrometer data, Environ. Sci. Technol., 45, 910-916,
 https://doi.org/10.1021/es102951k, 2011b.
- Ortiz-Montalvo, D. L., Hakkinen, S. A., Schwier, A. N., Lim, Y. B., McNeill, V. F., and Turpin, B. J.: Ammonium addition (and aerosol pH) has a dramatic impact on the volatility and yield of glyoxal secondary organic aerosol, Environ. Sci. Technol., 48, 255-262, https://doi.org/10.1021/es4035667, 2014.
- 555 Paatero, P. and Tapper, U.: Positive matrix factorization: A non-negative factor model with optimal utilization of error estimates of data
- 556 values, Environmetrics, 5, 111-126, <u>https://doi.org/10.1002/env.3170050203</u>, 1994.

- Poschl, U.: Atmospheric aerosols: composition, transformation, climate and health effects, Angew. Chem. Int. Edit., 44, 7520-7540,
 https://doi.org/10.1002/anie.200501122, 2005.
- Robinson, A. L., Donahue, N. M., Shrivastava, M. K., Weitkamp, E. A., Sage, A. M., Grieshop, A. P., Lane, T. E., Pierce, J. R., and
 Pandis, S. N.: Rethinking organic aerosols: semivolatile emissions and photochemical aging, Science, 315, 1259-1262,
 https://doi.org/10.1126/science.1133061, 2007.
- Rogge, W. F., Hildemann, L. M., Mazurek, M. A., Cass, G. R., and Simoneit, B.: Sources of fine organic aerosol. 1. Charbroilers and meat
 cooking operations, Environ. Sci. Technol., 25, 1112-1125, https://doi.org/10.1021/es00018a015, 1991.
- Sareen, N., Moussa, S. G., and McNeill, V. F.: Photochemical aging of light-absorbing secondary organic aerosol material, J. Phys. Chem.
 A., 117, 2987-2996, https://doi.org/10.1021/jp309413j, 2013.
- Schwarz, J., Gao, R., Spackman, J. R., Watts, L., Thomson, D. S., Fahey, D., Ryerson, T., Peischl, J., Holloway, J., Trainer, M., Frost, G.,
 Baynard, T., Lack, D., de Gouw, J., warneke, C., and Del Negro, L.: Measurement of the mixing state, mass, and optical size of
 individual black carbon particles in urban and biomass burning emissions, Geophys. Res. Lett., 35, L13810,
 <u>https://doi.org/10.1029/2008GL033968</u>, 2008.
- 570 Seinfeld, J., Bretherton, C., Carslaw, K., Coe, H., DeMott, P., Dunlea, E., Feingold, G., Ghan, S., Guenther, A., Kahn, R., Kraucunas, I.,
- 571 Kreidenweis, S., Molina, M., Nenes, A., Penner, J., Prather, K., Ramanathan, V., Ramaswamy, V., Rasch, P., and Wood, R.:
- 572 Improving our fundamental understanding of the role of aerosol-cloud interactions in the climate system, P. Natl. Acad. Sci. USA,
- 573 113, 5781-5790, <u>https://doi.org/10.1073/pnas.1514043113</u>, 2016.
- Shrivastava, M., Cappa, C. D., Fan, J., Goldstein, A. H., Guenther, A. B., Jimenez, J. L., Kuang, C., Laskin, A., Martin, S. T., Ng, N. L.,
 Petaja, T., Pierce, J. R., Rasch, P. J., Roldin, P., Seinfeld, J. H., Shilling, J., Smith, J. N., Thornton, J. A., Volkamer, R., Wang, J.,
 Worsnop, D. R., Zaveri, R. A., Zelenyuk, A., and Zhang, Q.: Recent advances in understanding secondary organic aerosol:
 Implications for global climate forcing, Rev. Geophy., 55, 509-559, https://doi.org/10.1002/2016rg000540, 2017.
- Shrivastava, M., Easter, R. C., Liu, X., Zelenyuk, A., Singh, B., Zhang, K., Ma, P.-L., Chand, D., Ghan, S., Jimenez, J. L., Zhang, Q., Fast,
 J., Rasch, P. J., and Tiitta, P.: Global transformation and fate of SOA: Implications of low-volatility SOA and gas-phase
 fragmentation reactions, J. Geophys. Res.-Atmos., 120, 4169-4195, https://doi.org/10.1002/2014jd022563, 2015.
- Slowik, J. G., Vlasenko, A., McGuire, M., Evans, G. J., and Abbatt, J. P. D.: Simultaneous factor analysis of organic particle and gas mass
 spectra: AMS and PTR-MS measurements at an urban site, Atmos. Chem. Phys., 10, 1969-1988, <u>https://doi.org/10.5194/acp-10-1969-2010</u>, 2010.
- Sun, Y., Jiang, Q., Wang, Z., Fu, P., Li, J., Yang, T., and Yin, Y.: Investigation of the sources and evolution processes of severe haze
 pollution in Beijing in January 2013, J. Geophys. Res.-Atmos., 119, 4380-4398, https://doi.org/10.1002/2014jd021641, 2014.
- 586 Sun, Y. L., Zhang, Q., Schwab, J. J., Demerjian, K. L., Chen, W. N., Bae, M. S., Hung, H. M., Hogrefe, O., Frank, B., Rattigan, O. V., and
- Lin, Y. C.: Characterization of the sources and processes of organic and inorganic aerosols in New York city with a high-resolution
 time-of-flight aerosol mass apectrometer, Atmos. Chem. Phys., 11, 1581-1602, https://doi.org/10.5194/acp-11-1581-2011, 2011.
- Tao, F., Gonzalez-Flecha, B., and Kobzik, L.: Reactive oxygen species in pulmonary inflammation by ambient particulates, Free Radical
 Bio. Med., 35, 327-340, https://doi.org/10.1016/S0891-5849(03)00280-6, 2003.
- 591 Ulbrich, I. M., Canagaratna, M. R., Zhang, Q., Worsnop, D. R., and Jimenez, J. L.: Interpretation of organic components from Positive
- 592 Matrix Factorization of aerosol mass spectrometric data, Atmos. Chem. Phys., 9, 2891-2918, <u>https://doi.org/10.5194/acp-9-2891-</u> 593 2009, 2009.

- Verma, V., Ning, Z., Cho, A. K., Schauer, J. J., Shafer, M. M., and Sioutas, C.: Redox activity of urban quasi-ultrafine particles from
 primary and secondary sources, Atmos. Environ., 43, 6360-6368, <u>https://doi.org/10.1016/j.atmosenv.2009.09.019</u>, 2009.
- von Schneidemesser, E., Monks, P. S., Gros, V., Gauduin, J., and Sanchez, O.: How important is biogenic isoprene in an urban
 environment? A study in London and Paris, Geophys. Res. Lett., 38, <u>https://doi.org/10.1029/2011gl048647</u>, 2011.
- 598 Wang, J., Ye, J., Liu, D., Wu, Y., Zhao, J., Xu, W., Xie, C., Shen, F., Zhang, J., Ohno, P. E., Qin, Y., Zhao, X., Martin, S. T., Lee, A. K.
- Y., Fu, P., Jacob, D. J., Zhang, Q., Sun, Y., Chen, M., and Ge, X.: Characterization of submicron organic particles in Beijing during
 summertime: comparison between SP-AMS and HR-AMS, Atmos. Chem. Phys., 20, 14091-14102, <u>https://doi.org/10.5194/acp-20-</u>
 14091-2020, 2020a.
- Wang, J., Ye, J., Zhang, Q., Zhao, J., Wu, Y., Li, J., Liu, D., Li, W., Zhang, Y., Wu, C., Xie, C., Qin, Y., Lei, Y., Huang, X., Guo, J., Liu,
 P., Fu, P., Li, Y., Lee, H. C., Choi, H., Zhang, J., Liao, H., Chen, M., Sun, Y., Ge, X., Martin, S. T., and Jacob, D. J.: Aqueous
 production of secondary organic aerosol from fossil-fuel emissions in winter Beijing haze, P. Natl. Acad. Sci. USA, 118,
 https://doi.org/10.1073/pnas.2022179118, 2021.
- Wang, L., Slowik, J. G., Tripathi, N., Bhattu, D., and Prévt, A.: Source characterization of volatile organic compounds measured by
 proton-transfer-reaction time-of-flight mass spectrometers in Delhi, India, Atmos. Chem. Phys., 20, 9753-9770,
 https://doi.org/10.5194/acp-20-9753-2020, 2020b.
- 609 Wood, E. C., Canagaratna, M. R., Herndon, S. C., Onasch, T. B., Kolb, C. E., Worsnop, D. R., Kroll, J. H., Knighton, W. B., Seila, R.,
- 610 Zavala, M., Molina, L. T., DeCarlo, P. F., Jimenez, J. L., Weinheimer, A. J., Knapp, D. J., Jobson, B. T., Stutz, J., Kuster, W. C., and
- Williams, E. J.: Investigation of the correlation between odd oxygen and secondary organic aerosol in Mexico City and Houston,
 Atmos. Chem. Phys., 10, 8947-8968, <u>https://doi.org/10.5194/acp-10-8947-2010</u>, 2010.
- Ku, J., Zhang, Q., Chen, M., Ge, X., Ren, J., and Qin, D.: Chemical composition, sources, and processes of urban aerosols during
 summertime in northwest China: insights from high-resolution aerosol mass spectrometry, Atmos. Chem. Phys., 14, 12593-12611,
 https://doi.org/10.5194/acp-14-12593-2014, 2014.
- Ku, W., Sun, Y., Wang, Q., Zhao, J., Wang, J., Ge, X., Xie, C., Zhou, W., Du, W., Li, J., Fu, P., Wang, Z., Worsnop, D. R., and Coe, H.:
 Changes in aerosol chemistry from 2014 to 2016 in winter in Beijing: Insights from high-resolution aerosol mass spectrometry, J.
 Geophys. Res.-Atmos., 124, 1132-1147, https://doi.org/10.1029/2018jd029245, 2019.
- 619 Zhang, Q., Jimenez, J. L., Canagaratna, M. R., Allan, J. D., Coe, H., Ulbrich, I., Alfarra, M. R., Takami, A., Middlebrook, A. M., Sun, Y.
- 620 L., Dzepina, K., Dunlea, E., Docherty, K., DeCarlo, P. F., Salcedo, D., Onasch, T., Jayne, J. T., Miyoshi, T., Shimono, A.,
- 621 Hatakeyama, S., Takegawa, N., Kondo, Y., Schneider, J., Drewnick, F., Borrmann, S., Weimer, S., Demerjian, K., Williams, P.,
- 622 Bower, K., Bahreini, R., Cottrell, L., Griffin, R. J., Rautiainen, J., Sun, J. Y., Zhang, Y. M., and Worsnop, D. R.: Ubiquity and
- dominance of oxygenated species in organic aerosols in anthropogenically-influenced Northern Hemisphere midlatitudes, Geophys.
 Res. Lett., 34, https://doi.org/10.1029/2007gl029979, 2007.
- Zhang, Z., Zhu, W., Hu, M., Wang, H., Chen, Z., Shen, R., Yu, Y., Tan, R., and Guo, S.: Secondary organic aerosol from typical Chinese
 domestic cooking emissions, Environ. Sci. Technol. Lett., 8, 24-31, https://doi.org/10.1021/acs.estlett.0c00754, 2020.
- Zhao, Y., Nguyen, N. T., Presto, A. A., Hennigan, C. J., May, A. A., and Robinson, A. L.: Intermediate volatility organic compound
 emissions from on-road diesel vehicles: chemical composition, emission factors, and estimated secondary organic aerosol production,
- 629 Environ. Sci. Technol., 49, 11516-11526, https://doi.org/10.1021/acs.est.5b02841, 2015.

- Kong, S., Chen, N., Niu, Z., Zhang, Y., Jiang, S., Yan, Y., and Qi, S.: Source apportionment of volatile organic compounds:
 Implications to reactivity, ozone formation, and secondary organic aerosol potential, Atmos. Res., 249, 105344,
 https://doi.org/10.1016/j.atmosres.2020.105344, 2021.
- Zheng, H., Kong, S., Chen, N., Yan, Y., Liu, D., Zhu, B., Xu, K., Cao, W., Ding, Q., Lan, B., Zhang, Z., Zheng, M., Fan, Z., Cheng, Y.,
 Zheng, S., Yao, L., Bai, Y., Zhao, T., and Qi, S.: Significant changes in the chemical compositions and sources of PM_{2.5} in Wuhan
 since the city lockdown as COVID-19, Sci. Total. Environ., 739, 140000, https://doi.org/10.1016/j.scitotenv.2020.140000, 2020.
- 636 Zheng, H., Kong, S., Yan, Q., Wu, F., Cheng, Y., Zheng, S., Wu, J., Yang, G., Zheng, M., Tang, L., Yin, Y., Chen, K., Zhao, T., Liu, D.,
- 637 Li, S., Qi, S., Zhao, D., Zhang, T., Ruan, J., and Huang, M.: The impacts of pollution control measures on PM_{2.5} reduction: Insights of 638 197, chemical composition, source variation and health risk, Environ., 103-117, Atmos. https://doi.org/10.1016/j.atmosenv.2018.10.023, 2019. 639
- Kang, K., Li, Z., Zhang, T., Wang, F., Wang, F., Tao, Y., Zhang, X., Wang, F., and Huang, J.: Volatile organic compounds in a typical
 petrochemical industrialized valley city of northwest China based on high-resolution PTR-MS measurements: Characterization,
 sources and chemical effects, Sci. Total. Environ., 671, 883-896, https://doi.org/10.1016/j.scitotenv.2019.03.283, 2019.
- 643 Zhu, W., Guo, S., Zhang, Z., Wang, H., Yu, Y., Chen, Z., Shen, R., Tan, R., Song, K., Liu, K., Tang, R., Liu, Y., Lou, S., Li, Y., Zhang,
- W., Zhang, Z., Shuai, S., Xu, H., Li, S., Chen, Y., Hu, M., Canonaco, F., and Prévôt, A. S. H.: Mass spectral characterization of
 secondary organic aerosol from urban cooking and vehicular sources, Atmos. Chem. Phys., 21, 15065-15079,
 https://doi.org/10.5194/acp-21-15065-2021, 2021.

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