



1           **Formation of Organic Sulfur Compounds through SO<sub>2</sub> Initiated**  
2           **Photochemistry of PAHs and DMSO at the Air-Water Interface**

3  
4  
5  
6  
7

Haoyu Jiang<sup>1,2,3</sup>, Yingyao He<sup>4</sup>, Yiqun Wang<sup>1,2</sup>, Sheng Li<sup>5</sup>, Bin Jiang<sup>1,2</sup>, Luca Carena<sup>6</sup>, Xue  
Li<sup>7</sup>, Lihua Yang<sup>4</sup>, Tiangang Luan<sup>4</sup>, Davide Vione<sup>6</sup>, Sasho Gligorovski<sup>\*1,2,3</sup>

8           <sup>1</sup>State Key Laboratory of Organic Geochemistry and Guangdong Provincial Key Laboratory of  
9           Environmental Protection and Resources Utilization, Guangzhou Institute of Geochemistry,  
10           Chinese Academy of Sciences, Guangzhou 510 640, China

11           <sup>2</sup>Guangdong-Hong Kong-Macao Joint Laboratory for Environmental Pollution and Control,  
12           Guangzhou Institute of Geochemistry, Chinese Academy of Science, Guangzhou 510640,  
13           China

14           <sup>3</sup>Chinese Academy of Science, Center for Excellence in Deep Earth Science, Guangzhou,  
15           510640

16           <sup>4</sup>School of Marine Sciences, Sun Yat-sen University, Guangzhou 510006, China

17           <sup>5</sup>Hunan Research Academy of Environmental sciences, Changsha, 410004, China

18           <sup>6</sup>Dipartimento di Chimica, Università di Torino, Via Pietro Giuria 5, 10125 Torino, Italy

19           <sup>7</sup>Institute of Mass Spectrometry and Atmospheric Environment, Jinan University, Guangzhou  
20           510632, China

21  
22

23

24

25           \*Corresponding author:

26           Sasho Gligorovski

27           [gligorovski@gig.ac.cn](mailto:gligorovski@gig.ac.cn)

28



29 **ABSTRACT**

30 The presence of organic sulfur compounds (OSs) at the water surface, acting as organic  
31 surfactants, may influence the air-water interaction and contribute to new particle formation in  
32 the atmosphere. However, the impact of ubiquitous anthropogenic pollutant emissions, such as  
33 SO<sub>2</sub> and polycyclic aromatic hydrocarbons (PAHs) on the formation of OSs at the air-water  
34 interface still remains unknown. Here, we observe large amounts of OSs formation in presence  
35 of SO<sub>2</sub>, upon irradiation of aqueous solutions containing typical PAHs such as pyrene (PYR),  
36 fluoranthene (FLA), and phenanthrene (PHE), as well as dimethylsulfoxide (DMSO). We  
37 observe rapid formation of several gaseous OSs from light-induced heterogeneous reactions of  
38 SO<sub>2</sub> with either DMSO or a mixture of PAHs/DMSO, and some of these OSs (e.g.  
39 methanesulfonic acid) are well established secondary organic aerosol (SOA) precursors. A  
40 myriad of OSs and unsaturated compounds are produced and detected in the aqueous phase.  
41 The tentative reaction pathways are supported by theoretical calculations of the reaction Gibbs  
42 energies. Our findings provide new insights into potential sources and formation pathways of  
43 OSs occurring at the water (sea, lake, river) surface, that should be considered in future model  
44 studies to better represent the air-water interaction and SOA formation processes.

45

46 **Keywords:** PAHs, SO<sub>2</sub>, SOA, aromatic organosulfates, water surface, photochemical reactions,  
47 SPI-TOF-MS, FT-ICR-MS

48



## 49 1. INTRODUCTION

50 Organic sulfur compounds (OSs) are ubiquitous in atmospheric aerosols, and organosulfates  
51 are considered as important tracers of secondary organic aerosols (SOA). Based on the  
52 occurrence of hydrophilic and hydrophobic moieties in the same molecule, OSs are surface-  
53 active compounds that cause reductions of surface tension and enhance the formation potential  
54 of cloud condensation nuclei (CCN) in aerosol particles (Bruggemann et al., 2020).

55 Both biogenic and anthropogenic sources such as biomass and fossil fuel burning release OSs  
56 into the atmosphere (Bruggemann et al., 2020). The contribution of aromatic organosulfates  
57 could be up to two-thirds of the sum of the identified OSs (Riva et al., 2015; Riva et al., 2016),  
58 and their input is highest during winter (Ma et al., 2014). Although some aromatic  
59 organosulfates have similar chemical structures as those of potential aromatic VOC precursors,  
60 such as toluene and xylene (Kundu et al., 2013), monocyclic aromatics were not regarded as  
61 aromatic organosulfate precursors because they are more inclined to oxidize into ring-opening  
62 products (Staudt et al., 2014; Kamens et al., 2011). Polycyclic aromatic hydrocarbons (PAHs),  
63 such as naphthalene (NAP) and 2-methylnaphthalene (2-MeNAP) have rather been postulated  
64 as the precursors of aromatic OSs (Riva et al., 2015; Staudt et al., 2014).

65 Among all aromatic compounds, PAHs are ubiquitous organics enriched at both the sea and  
66 freshwater (lakes, river) surface (Cincinelli et al., 2001; Vácha et al., 2006; Chen et al., 2006;  
67 Lohmann et al., 2009; Seidel et al., 2017), where they can reach even 200–400 times higher  
68 concentrations compared to the water bulk (Cincinelli et al., 2001; Vácha et al., 2006; Chen et  
69 al., 2006; Lohmann et al., 2009; Seidel et al., 2017; Hardy et al., 1990). The origin of PAHs  
70 accumulated at the water surface stems from combustion processes such as biomass burning,  
71 and coal- and petroleum-based combustion (Lammel, 2015). At the surface of freshwater and  
72 seawater the PAHs concentrations vary, respectively, from 11.84 to 393.12 ng L<sup>-1</sup> (Li et al.,  
73 2017b), and from 5 to ~1900 ng L<sup>-1</sup> (Valavanidis et al., 2008; Otto et al., 2015; González-Gaya



74 et al., 2019; Pérez-Carrera et al., 2007; Ma et al., 2013). Phenanthrene (PHE), fluoranthene  
75 (FLA) and pyrene (PYR) are the most commonly detected PAHs in the coastal surface micro-  
76 layer (SML) (Guitart et al., 2007; Stortini et al., 2009), and they accounted for 92 to 96% of the  
77 total PAHs amount in the coastal surface waters of Nigeria (Benson et al., 2014).

78 Dimethylsulfoxide (DMSO) is an ubiquitous OS compound at the sea surface (Lee et al., 1999),  
79 which derives from degradation of phytoplankton (Andreae, 1980), photodegradation of  
80 dimethyl sulfide (DMS) (Barnes et al., 2006; Brimblecombe and Shooter, 1986), and microbial  
81 oxidation of DMS (Zhang et al., 1991). Because of high Henry's Law coefficient ( $\approx 10^7 \text{ M atm}^{-1}$ )  
82 and mass accommodation coefficient (0.10), gaseous DMSO can be deposited on the sea- and  
83 fresh-water surface where one finds enhanced DMSO concentrations (Legrand et al., 2001;  
84 Davidovits et al., 2006; González-Gaya et al., 2016). Highest levels of DMSO are detected in  
85 the ocean ( $1.5$  to  $532 \text{ nmol L}^{-1}$ ) (Hatton et al., 1996; Lee and De Mora, 1996; Andreae, 1980;  
86 Asher et al., 2017), followed by detected levels in rivers ( $<2.5$ – $210 \text{ nmol L}^{-1}$ ) (Andreae, 1980),  
87 lakes (up to  $180 \text{ nmol L}^{-1}$ ) (Richards et al., 1994), rainwater ( $2$ – $4 \text{ nmol L}^{-1}$ ) (Ridgeway et al.,  
88 1992), and aerosols ( $69$ – $125 \text{ pmol m}^{-3}$ ) (Harvey and Lang, 1986). One of the main degradation  
89 pathways of DMSO is the reaction with hydroxyl radicals (OH) (Barnes et al., 2006), yielding  
90 methanesulfinic and methanesulfonic acids (Librando et al., 2004). Because the concentrations  
91 of DMSO are 1–2 orders of magnitude higher than those of DMS, the photochemical oxidation  
92 of DMSO may be a relatively more important process than the photo-oxidation of DMS (Lee  
93 et al., 1999).

94 Sulfur dioxide ( $\text{SO}_2$ ) is directly emitted into the atmosphere by anthropogenic sources such as  
95 fossil fuel combustion, coal, oil, and industrial processes (Smith et al., 2001). In addition, it can  
96 also be formed during the oxidation of DMS (Hoffmann et al., 2016; Chen et al., 2018). As a  
97 key contributor to aerosol nucleation, the role of  $\text{SO}_2$  at the air-water interface is also recognized  
98 as an efficient precursor of HOSO and OH radicals, following light absorption below 340 nm



99 and its excitation to a very reactive triplet state ( $^3\text{SO}_2^*$ ) (Martins-Costa et al., 2018; Kroll et al.,  
100 2018). A majority of studies employed sulfate-containing seed particles to explore the  
101 formation pathway of OSs, while only a few were focused on the organosulfur formation  
102 pathway that is unique to  $\text{SO}_2$  chemistry (Blair et al., 2017; Shang et al., 2016; Passananti et al.,  
103 2016). Some previous studies have found that the heterogeneous reaction of  $\text{SO}_2$  with  
104 unsaturated acids can lead to the formation of OSs in the atmosphere (Shang et al., 2016;  
105 Passananti et al., 2016). However, it has also been found that monocyclic compounds such as  
106 terephthalic acid are not reactive *via* direct  $\text{SO}_2$  addition (Passananti et al., 2016), and there is  
107 still a knowledge gap concerning other OSs formation pathways involving heterogeneous  $\text{SO}_2$   
108 oxidation. Indeed, air quality models cannot explain the quick increase of sulfate amount in  
109 going from clean air to hazy events, when applying only the gas-phase and aqueous-phase  
110 chemistry of  $\text{SO}_2$  (Wang et al., 2014; Li et al., 2017a).

111 Several studies have assessed the heterogeneous chemistry of atmospherically relevant oxidants  
112 with PAHs (Donaldson et al., 2009; Monge et al., 2010; Styler et al., 2011; Zhou et al., 2019).  
113 Recently, Mekic et al (2020) and Jiang et al., (2021) have shown that the photosensitized  
114 degradation of DMSO by excited triplet state of typical PAH compounds (fluorene ( $^3\text{FL}^*$ ),  
115  $^3\text{PHE}^*$ ,  $^3\text{FLA}^*$  and  $^3\text{PYR}^*$ ) leads to the formation of OSs compounds, among the others, in  
116 both the gas- and aqueous- phase.

117 In this study we investigated the formation of OSs from aqueous DMSO and/or PAHs/DMSO,  
118 initiated by gaseous  $\text{SO}_2$  in the dark and in the presence of simulated sunlight irradiation  
119 ( $300\text{nm} < \lambda < 700\text{nm}$ ). The gaseous OS products were assessed by membrane inlet single photon  
120 ionization-time of flight-mass spectrometry (MI-SPI-TOFMS) (Zhang et al., 2019; Mekic et al.,  
121 2020a; Jiang et al., 2021). The formed aqueous-phase products were evaluated by means of  
122 ultrahigh resolution electrospray ionization Fourier-transform ion cyclotron resonance mass  
123 spectrometry (FT-ICR-MS) (Jiang et al., 2016). The tentative reaction pathways of the formed



124 OSs during the heterogeneous SO<sub>2</sub> oxidation of PAHs/DMSO under actinic illumination are  
125 supported by theoretical calculations of the reaction Gibbs energies. We show that oxidation by  
126 SO<sub>2</sub> of PAHs/DMSO can release gaseous OSs, such as methanesulfonic acid (MSA), which are  
127 known precursors of secondary organic aerosols (SOA) in the atmosphere. The formation of an  
128 important number of OSs and unsaturated compounds, among the others, was observed in the  
129 aqueous phase. We highlight the large amounts of generated linear and aromatic OSs, with  
130 potential to greatly influence the air-water exchange of organic compounds.

131



## 132        **2. EXPERIMENTAL**

### 133        **2.1. Photoreactor**

134        A double-wall rectangular ( $5 \times 5 \times 2$  cm) photoreactor was used to assess the reaction of gaseous  
135        SO<sub>2</sub> with a water/organic film containing DMSO or DMSO/PAHs. (Mekic et al., 2020a; Mekic  
136        et al., 2020b) The photoreactor was thermostated at ambient temperature ( $T = 293$  K) by  
137        thermostated bath (LAUDA ECO RE 630 GECCO, Germany).

138        A SO<sub>2</sub> flow of  $150 \text{ mL min}^{-1}$  mixed with an air flow of  $750 \text{ mL min}^{-1}$  ( $0\text{-}1 \text{ L min}^{-1}$  HORIBA  
139        METRON mass flow controller; accuracy,  $\pm 1\%$ ) allowed for dilution of SO<sub>2</sub> from a standard  
140        gas cylinder with 5 ppm concentration to a mixing ratio of  $\sim 800$  ppb that was flowing through  
141        the photoreactor. The applied mixing ratio of 800 ppb is higher compared to the atmospheric  
142        background mixing ratio that ranges from 1 to 70 ppb, but it is smaller than the SO<sub>2</sub> mixing  
143        ratio in dilute volcanic plumes (for instance, a mixing ratio of ca. 10 ppm is observed about 10  
144        km away from volcanic sources) (Oppenheimer et al., 1998), and it is also smaller compared to  
145        some previous laboratory studies that used 7.7-500 ppm of SO<sub>2</sub> (Librando et al., 2014). The  
146        applied mixing ratio of 800 ppb would probably amplify the intensity of the detected product  
147        compounds, but the formation profiles would still remain the same as in the case of smaller SO<sub>2</sub>  
148        mixing ratios.

149        The concentration of PHE, FLA, and PYR (Sigma-Aldrich) used separately was  $1 \times 10^{-4} \text{ mol L}^{-1}$ .

150        The three compounds were dissolved individually in a mixture of DMSO and ultrapure water  
151        with proportion 10:90 v/v (Mekic et al., 2020b). Such high DMSO concentration was necessary  
152        to dissolve the poorly water soluble PHE, FLA, and PYR (Mekic et al., 2020a; Mekic et al.,  
153        2020b). Several previous studies also used high concentrations of the organic co-solvent to  
154        assess the co-solvent effect on PAHs photolysis (Donaldson et al., 2009; Librando et al., 2014;  
155        Grossman et al., 2016). The reactor was filled with 10 mL of freshly prepared DMSO/PAH  
156        solution and irradiated with a Xenon lamp (Xe, 500 W,  $300\text{nm} < \lambda < 700\text{nm}$ ) during the SO<sub>2</sub>



157 oxidation of PAHs/DMSO. The spectral irradiance of the Xe lamp was measured with a  
158 calibrated spectroradiometer (Ocean Optics, USA) equipped with a linear-array CCD detector,  
159 and compared to the sunlight radiation (Mekic et al., 2020a; Mekic et al., 2020b). The  
160 irradiation time of the aqueous solution was 2 hours.

161 Blank experiments of SO<sub>2</sub> reaction were carried out with aqueous DMSO in all experimental  
162 conditions and with ultrapure water in the dark. All experiments were performed at least twice.

### 163 **2.2. Single Photon Ionization-Time of Flight-Mass Spectrometry (SPI-ToF-MS)**

164 A SPI-ToF-MS instrument (SPIMS 3000, Guangzhou Hexin Instrument Co., Ltd., China) was  
165 used to detect the gas-phase compounds formed during the light-induced heterogeneous  
166 reaction of SO<sub>2</sub> with PAHs/DMSO. The SPI-ToF-MS instrument was explained in details in  
167 our previous studies (Deng et al., 2021; Mekic et al., 2020a), so here only brief description is  
168 given in the supporting information (SI).

### 169 **2.3. Fourier - Transform Ion Cyclotron Resonance Mass Spectrometry (FT-ICR-MS)**

170 As FT-ICR-MS was described in our previous paper, a brief description only is given in the SI.  
171 An in-house software was used to calculate all mathematically possible formulae for all ions  
172 with a signal-to-noise ratio above 10, using a mass tolerance of ±0.2 ppm. Data of blank samples  
173 were subtracted from those of all samples according to the same possible formulae.

174 The double bond equivalent (DBE) of a chemical formula C<sub>c</sub>H<sub>h</sub>O<sub>o</sub>N<sub>n</sub>S<sub>s</sub> is calculated using the  
175 following equation (Yassine et al., 2014):

$$176 \quad DBE = \frac{2c+2-h+n}{2} \quad (\text{Eq-1})$$

177 The aromaticity equivalent ( $X_c$ ) is calculated using the following equation:

$$178 \quad X_c = \frac{3[DBE - (mO + nS)] - 2}{DBE - (mO + nS)} \quad (\text{Eq-2})$$

179 where  $X_c$  is introduced to improve the identification and characterization of aromatic and  
180 condensed aromatic compounds, while  $m$  and  $n$  are the respective fractions of oxygen and sulfur



181 atoms that are involved in the  $\pi$ -bonds of a molecular structure (Yassine et al., 2014). According  
182 to the model structural classes, the values of  $m$  and  $n$  were all set to 0. Threshold values of  $X_c$   
183 between 2.5 and 2.7 ( $2.5 \leq X_c < 2.7$ ) and equal or greater than 2.7 ( $X_c \geq 2.7$ ) were set as  
184 minimum criteria for the presence of aromatics or condensed aromatic compounds in an  
185 identified molecule (Blair et al., 2017; Jiang et al., 2016). It ought to be highlighted that electron  
186 spray ionization (ESI) and desorption electrospray ionization (DESI) of OSs is especially  
187 favorable in the negative-ion mode, making the observed fraction of OSs further amplified  
188 (Blair et al., 2017).

#### 189 **2.4. Theoretical Calculations**

190 The proposed structures of all identified tentative OSs are based on the reasonable inferred  
191 elemental compositions for a single mass, following speculation through NIST Chemistry  
192 WebBook (<https://webbook.nist.gov/chemistry/mw-ser/>) and database of MI-SPI-ToF-MS or  
193 FT-ICR-MS. Considering that PHE, PYR, FLA and SO<sub>2</sub> can all absorb lamp radiation,  
194 theoretical calculations were carried out to obtain some insights into the degradation pathways  
195 of DMSO initiated by the excited triplet states <sup>3</sup>PHE\*, <sup>3</sup>PYR\*, <sup>3</sup>FLA\* and <sup>3</sup>SO<sub>2</sub>\* in the aqueous-  
196 and gaseous- phase, as well as the interaction between light-excited PAHs and SO<sub>2</sub>.  
197 All calculations shown here were assessed by Gaussian 16W package (Frisch et al., 2016). The  
198 level of theory B3LYP/6-311G(d,p) was applied for geometry optimizations and frequency  
199 calculations for all molecules depicted in the reaction scheme (Mclean and Chandler, 1980;  
200 Binning Jr and Curtiss, 1990). There were no imaginary frequencies for all molecules optimized.  
201 Single-point energy calculations were performed at a more expensive level, i.e., M06-2X/Def2-  
202 TZVP level (Zhao and Truhlar, 2008; Weigend and Ahlrichs, 2005; Weigend, 2006). The  
203 existence of possible geometric isomers and conformers for each species were considered and  
204 investigated, and those with lowest calculated Gibbs free energies were selected. Molecular  
205 oxygen (O<sub>2</sub>), carbon dioxide (CO<sub>2</sub>) and water molecules (H<sub>2</sub>O) were placed as reactants or



206 products (if needed) to balance atoms in the schemes. Detailed Gibbs free energies for all  
207 molecules are presented in Table S1 and S2, and the corresponding reaction Gibbs energies are  
208 shown in Scheme 1.

209

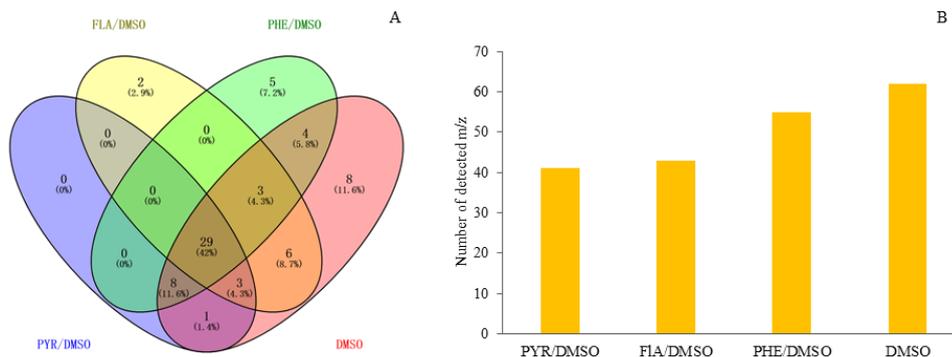


## 210 **3. RESULTS AND DISCUSSION**

### 211 **3.1. Gaseous OSs Detected by MI-SPI-ToF-MS**

212 To detect the gas-phase compounds formed by heterogeneous SO<sub>2</sub> oxidation of PAHs/DMSO  
213 in the dark and under light irradiation, we applied MI-SPI-ToF-MS as a novel promising  
214 technology for the real-time monitoring of VOCs (Zhang et al., 2019; Mekic et al., 2020a; Deng  
215 et al., 2021).

216 Figure 1A shows the Venn diagrams of the observed number of *m/z* signals corresponding to  
217 the gas-phase products of the light-induced heterogeneous SO<sub>2</sub> oxidation of PAHs/DMSO. The  
218 Venn diagrams showing the comparison of the gaseous products formed under different  
219 conditions are presented in Figure S1. The biggest contributor to the total number of secondarily  
220 formed products is the light-induced SO<sub>2</sub> oxidation of DMSO (Figure 1B). Among all detected  
221 *m/z* signals (Table S4), we tentatively identified a number of unsaturated multifunctional  
222 molecules and OSs released in the gas phase from the reaction of SO<sub>2</sub> with either DMSO or  
223 PAHs/DMSO, which are summarized in Table S5. It should be noted, the possibility for the  
224 existence of isomers and of different molecular formulas for the compounds with the same  
225 molecular weight (Nizkorodov et al., 2011). We emphasize the formation of gaseous OSs, and  
226 especially of those that are known to be SOA precursors. For example, the *m/z* signals of 80,  
227 94, 96, 112, 124, 126 were inferred as methanesulfinic acid (CH<sub>3</sub>SO<sub>2</sub>H, MSIA),  
228 methylsulfonylmethane ((CH<sub>3</sub>)<sub>2</sub>SO<sub>2</sub>, MSM), methanesulfonic acid (CH<sub>3</sub>SO<sub>3</sub>H, MSA),  
229 hydroxymethanesulfonic acid (CH<sub>4</sub>O<sub>4</sub>S, MSAOH), ethyl methanesulfonate (CH<sub>3</sub>SO<sub>3</sub>C<sub>2</sub>H<sub>5</sub>,  
230 EMS) and 2-hydroxyethenesulfonic acid (C<sub>2</sub>H<sub>5</sub>O<sub>4</sub>SH, ESAOH) (Berresheim et al., 1993;  
231 Berresheim and Eisele, 1998; Karl et al., 2007; Hopkins et al., 2008; Gaston et al., 2010; Ning  
232 et al., 2020; Dawson et al., 2012).



233

234 **Figure 1:** Venn Diagrams of the detected  $m/z$  signals in the gas-phase for the heterogeneous  
235 reaction of  $\text{SO}_2$  with DMSO and PAHs/DMSO under light irradiation ( $300 \text{ nm} < \lambda < 700 \text{ nm}$ )  
236 (Panel A); The total number of identified  $m/z$  signals between the heterogeneous reactions of  
237  $\text{SO}_2$  with DMSO and PAHs/DMSO, under light irradiation (Panel B).

238

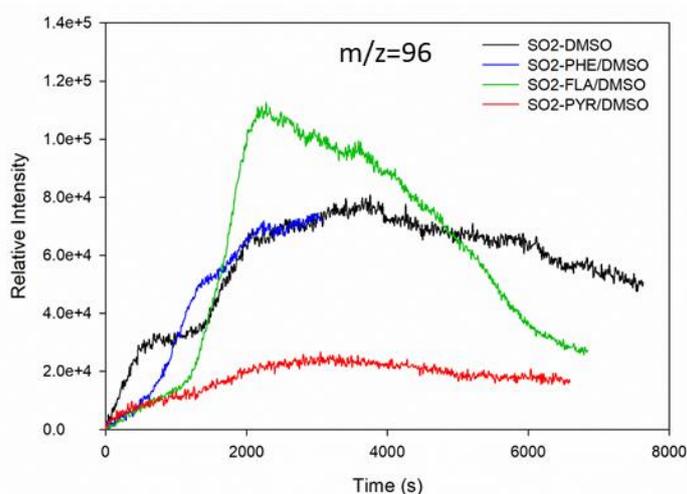
### 239 3.2. Formation Profiles of Gaseous OSs

240 In this study, we observed rapid formation of MSA, MSIA, MSM, EMS, MSAOH, and ESAOH  
241 (Figure 2 and Figure S2). The reaction of DMSO with OH leads to formation of MSIA (Barnes  
242 et al., 2006), which in turn may react again with OH radicals to form MSA (Librando et al.,  
243 2004; Barnes et al., 2006; Rosati et al., 2021). Gaseous MSA is a particularly important  
244 compound that can participate in the initial nucleation and growth step of particles, known as  
245 new particle formation (NPF) process (Schobesberger et al., 2013; Chen et al., 2016; Zhao et  
246 al., 2017; Perraud et al., 2015). Declining gaseous MSA concentrations were actually reported  
247 during marine NPF events, thereby suggesting that MSA may enter the aerosol particles at the  
248 earliest possible stage and significantly assist in cluster formation (Dall'osto et al., 2012; Bork  
249 et al., 2014). MSA is the simplest organosulfate compound, which can be mainly formed during  
250 the OH oxidation of DMS (Barnes et al., 2006; Rosati et al., 2021). Gas-phase MSA above the  
251 oceans and in coastal areas represents about 10 to 100% of the gas-phase sulfuric acid (SA)



252 concentration (Berresheim et al., 2002). It has been suggested that MSA can enhance particle  
253 formation from SA by 15-300%, if equal quantities of SA and MSA are present (Bork et al.,  
254 2014). There is a discrepancy between the modeled and measured concentration profiles of  
255 MSA, which indicates a missing MSA source. This source seems to be much stronger than the  
256 estimated production stemming from OH oxidation of DMS (Zhang et al., 2014a). Zhang et al.  
257 (2014) suggested a strong daytime source of MSA, the precursor of which may be DMSO  
258 (Zhang et al., 2014b). Their model estimations indicated that higher DMSO concentrations  
259 would lead to enhanced chemical production of MSA to reach  $4.9 \times 10^7$  molecules  $\text{cm}^{-2} \text{s}^{-1}$ ,  
260 which is similar to the strength of the missing source of MSA.

261 Here, we show that during daytime the reactions of light-excited  $\text{SO}_2$  and aqueous DMSO or  
262 DMSO/PAHs could represent an important source of gaseous MSA in the atmosphere near the  
263 water (ocean, lake and river) surface. Figure 2 shows typical time evolution profiles of MSA  
264 formed by the reaction of  $\text{SO}_2$  with PAHs/DMSO under light irradiation ( $300\text{nm} < \lambda < 700\text{nm}$ ).  
265 The formation profiles of MSM, EMS, MSIA, MSAOH and ESAOH are shown in Figure S2.



266

267 **Figure 2:** Formation profiles of  $m/z = 96$  (methanesulfonic acid, MSA) upon light-induced  
268 heterogeneous reactions of  $\text{SO}_2$  with DMSO and PAHs/DMSO.

269



270

271 The SO<sub>2</sub> oxidation of FLA/DMSO leads to MSA formation, the signal of which increases  
272 during the first 1 hour and then slowly decreases (Figure 2). The intensities of the product  
273 compounds (Figure 2 and Figure S2) decrease after one hour most probably due to their reaction  
274 with SO<sub>2</sub> and/or their photodegradation. The signal profile of MSA due to reaction of SO<sub>2</sub> with  
275 PHE/DMSO is shorter than the others due to a technical problem of the instrument.

276 Intriguingly, the light-induced heterogeneous reaction of SO<sub>2</sub> with DMSO, PYR/DMSO, and  
277 PHE/DMSO produces MSA that remains approximately stable over the course of the reaction.  
278 These results indicate that the suggested reaction between SO<sub>2</sub> and DMSO or DMSO/PAHs  
279 under sunlight irradiation can produce MSA, which can be persistent enough to affect the NPF  
280 process in the atmosphere. In section 3.4. we suggest a tentative reaction pathway for the  
281 production of MSA and other OSs that is supported by theoretical calculations of Gibbs free  
282 energies.

283

### 284 **3.3. Aqueous OSs Detected by FT-ICR-MS**

285 Numerous unsaturated multifunctional molecules and OSs were identified in the liquid phase  
286 during the reaction of SO<sub>2</sub> with either DMSO or PAHs/DMSO by using FT-ICR-MS. The  
287 number of detected product compounds in the aqueous phase was significantly higher compared  
288 to those detected in the gas phase, due to very high sensitivity of FT-ICR-MS. Figure S3 shows  
289 all the formulae detected upon the reaction of SO<sub>2</sub> with PAHs/DMSO under light irradiation.  
290 The shared formulae of the compounds detected upon reactions of SO<sub>2</sub> with PYR/DMSO and  
291 FLA/DMSO or PHE/DMSO were more abundant than those individually formed by the  
292 reaction of SO<sub>2</sub> with FLA/DMSO or PHE/DMSO. The number of detected compounds formed  
293 by reaction of SO<sub>2</sub> with PYR/DMSO was dominant among all the products of SO<sub>2</sub> oxidation of  
294 PAHs/DMSO, both in the dark and under light irradiation (Figure S3B and S4). Interestingly,



295 the light-induced heterogeneous reaction of SO<sub>2</sub> with PYR/DMSO released a small number of  
296 gaseous products, but the highest number of aqueous-phase products among all the studied  
297 PAHs/DMSO. In general, more C<sub>c</sub>H<sub>h</sub>O<sub>o</sub> (CHO) than C<sub>c</sub>H<sub>h</sub>O<sub>o</sub>S<sub>s</sub> (CHOS) products were formed  
298 in the aqueous phase with SO<sub>2</sub> + PAHs/DMSO under irradiation (Figure S5), with the only  
299 exception of PHE/DMSO. Even when subtracting the formulae detected upon SO<sub>2</sub> oxidation of  
300 DMSO in the dark from those detected in the corresponding light-induced heterogeneous  
301 reaction, more compounds were produced under irradiation than in the dark.

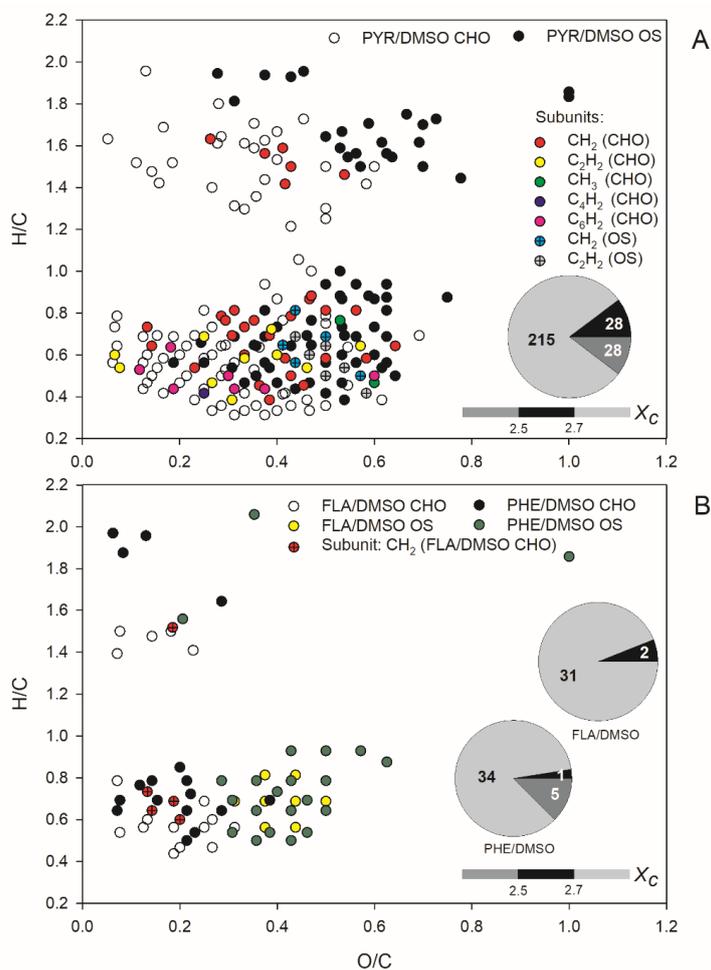
302 The analysis of iso-abundance plots of DBE vs. carbon numbers for the detected CHO and  
303 CHOS formulae is shown in text section S1 and Figures S6-S7. The two-dimensional van  
304 Krevelen (VK) plots for all CHOs and CHOSs formed during light-induced SO<sub>2</sub> oxidation of  
305 PAHs/DMSO are shown in Figure 3.

306 The CHO and CHOS compounds formed during the reaction of SO<sub>2</sub> with DMSO in the dark  
307 and under irradiation are shown in the VK plot depicted in Figure S8. The same classes of  
308 compounds, detected upon heterogeneous reactions of SO<sub>2</sub> with PAHs/DMSO under irradiation  
309 are illustrated in the VK plots depicted in Figure S9 and Figure S10, respectively. Most of the  
310 CHO and CHOS products identified with DMSO had high H/C ratios (1.5-2.1) but O/C ratios  
311 lower than 0.4, suggesting the formation of aliphatic compounds (Mekic et al., 2020a). However,  
312 the presence of aromatics could be additionally highlighted according to the analysis of a  
313 mathematical parameter, the aromaticity equivalent ( $X_c$ ) (Yassine et al., 2014). About one half  
314 of the observed product compounds and especially CHOSs exhibited  $X_c \geq 2.5$ , further  
315 indicating the formation of unsaturated compounds during the oxidation of DMSO by SO<sub>2</sub> in  
316 the presence of light (Figure S8).

317 The CHO and CHOS products displayed in the VK diagrams of Figure 3 could be all separated  
318 into two different clusters. The O/C ratios of CHOs were smaller than 0.7 for SO<sub>2</sub> reaction with  
319 PYR/DMSO, and smaller than 0.4 for SO<sub>2</sub> reactions with FLA/DMSO and PHE/DMSO.



320 Nonetheless, a cluster with relatively few observed products was located in the upper part of  
321 the VK diagrams with high H/C ratios (1.2-2.0), suggesting the formation of saturated aliphatic  
322 CHO compounds.



323

324

325 **Figure 3:** The van Krevelen (VK) graph and aromaticity equivalent (grey with  $X_c < 2.5$ , black  
326 with  $2.5 \leq X_c < 2.7$ , and silver with  $X_c \geq 2.7$ ) for the detected CHO and CHOS compounds in  
327 ESI<sup>-</sup> mode, formed upon light-induced heterogeneous reaction of SO<sub>2</sub> with PAHs/DMSO. The



328  $X_c$  value is illustrated by the color bar of each VK diagram, while the pie chart shows the number  
329 in different  $X_c$  intervals during these reactions.  
330



331 The other cluster that includes most of the detected products is located in the lower part of the  
332 VK diagrams, exhibiting low H/C ratios (0.4-0.8) and indicating a degree of unsaturation  
333 (Mekic et al., 2020a; Lin et al., 2012). Most of the CHO compounds are probably condensed  
334 aromatic compounds as suggested by their  $X_c$  values higher than 2.5, especially in the case of  
335 products observed upon reaction of SO<sub>2</sub> with FLA/DMSO.

336 The observed CHOS products emerging from the reaction of SO<sub>2</sub> with PHE/DMSO, also  
337 depicted in Figure 3 could be divided as well into two groups based on their distribution of H/C  
338 and O/C ratios. Generally, the formed CHOS products that are located in the upper part of the  
339 VK diagrams (Figure 3) have broad range of O/C ratios spanning from 0.2 to 0.8, H/C ratios in  
340 a range of 1.4-2.0, and low DBE (1-4). This observation implies the formation of long-chain  
341 aliphatic-like CHOSs during the light-assisted oxidation of PHE/DMSO by SO<sub>2</sub>. In addition, a  
342 big fraction of compounds detected during the reaction of SO<sub>2</sub> with PHE/DMSO exhibits  $X_c <$   
343 2.5, and is thus consistent with the formation of long-chain aliphatic-like CHOSs (Wang et al.,  
344 2021). The CHOS products located in the lower part of the VK diagrams have H/C ratios of  
345 0.4-1.0, which are similar to those of the observed CHO products. However, the CHOS  
346 compounds exhibit higher O/C ratios that range between 0.2 and 0.8 for the reaction of SO<sub>2</sub>  
347 with PYR/DMSO, and 0.2-0.6 for the reaction of SO<sub>2</sub> with FLA/DMSO and PHE/DMSO. A  
348 possible reason is the occurrence of sulfates with R-OSO<sub>3</sub><sup>-</sup> groups, sulfonates with R-SO<sub>3</sub><sup>-</sup>  
349 groups and sulfones with R-SO<sub>2</sub>-R' groups (Bruggemann et al., 2020). These CHOS products  
350 partly overlap with the oxidized aromatic hydrocarbons (Kourtchev et al., 2014). Moreover, the  
351 majority of the CHOS products exhibit a condensed aromatic structure as indicated by their  $X_c$   
352 values higher than 2.7.

353 Based on the VK plots, here we used the different spacing patterns for Kendrick mass defect  
354 (KMD) analysis (Lin et al., 2012). The subunits including CH<sub>2</sub>, C<sub>2</sub>H<sub>2</sub>, CH<sub>3</sub>, C<sub>4</sub>H<sub>2</sub>, C<sub>6</sub>H<sub>2</sub> for the  
355 same CHO homologous series, and CH<sub>2</sub>, C<sub>2</sub>H<sub>2</sub> for CHOS homologues were identified during



356 the light-induced SO<sub>2</sub> oxidation of PAHs/DMSO. More subunits were found for CHOS  
357 homologues during the SO<sub>2</sub> oxidation of PAHs/DMSO under all experimental conditions  
358 (Figure S11). Literature suggests that subunits of CH<sub>2</sub> and C<sub>2</sub>H<sub>2</sub> were the most repeating mass  
359 building increments for either CHO or CHOS oxidized aromatic compounds (Lin et al., 2012).  
360 These identified subunits were responsible for the increase in the molecular mass of the detected  
361 compounds (Altieri et al., 2006), presumably leading to straight-chain alkanes and olefins.

### 362 **3.4. Tentative Reaction Mechanisms and the Production of OSs**

363 The proposed general pathway suggested in the literature is the direct addition of SO<sub>2</sub> to a  
364 double bond or the separate addition of SO<sub>2</sub> to the cleavage of a double bond. This may also  
365 apply to PAHs/DMSO, because of the chromophoric nature of the SO<sub>2</sub> adduct with C=C bonds  
366 (Passananti et al., 2016). Herein, we show that this reaction pathway of SO<sub>2</sub> may indeed proceed  
367 on PAHs and is supported by theoretical calculations of the reaction Gibbs energies, based on  
368 the information obtained from the detected tentative products.

369 The reaction pathway suggests that photoexcitation of PAHs and SO<sub>2</sub> proceeds through the  
370  $\pi \rightarrow \pi^*$  electronic transition, followed by intersystem crossing to produce the corresponding  
371 excited triplet states (<sup>3</sup>PAHs\* and <sup>3</sup>SO<sub>2</sub>\*) (Mekic et al., 2020a), which most likely play an  
372 important role during the oxidation of PAHs/DMSO. Previous work concluded that the  
373 photochemical reaction of <sup>3</sup>PAHs\* with DMSO in water would lead to the formation of singlet  
374 oxygen (<sup>1</sup>O<sub>2</sub>) *via* energy transfer reaction with ground triplet-state oxygen (<sup>3</sup>O<sub>2</sub>) (Wilkinson et  
375 al., 1995), the formation of hydroxyl radical upon water oxidation by other triplets states  
376 (Brigante et al., 2010), and the formation of more radicals through electron transfer. Many  
377 triplet states work in the presence of oxygen as O<sub>2</sub> quenches most, but not all of them. The  
378 excited triplet state of SO<sub>2</sub> (<sup>3</sup>SO<sub>2</sub>\*) formed upon light irradiation reacts with water molecules  
379 yielding OH radicals as follows (Martins-Costa et al., 2018; Kroll et al., 2018):



R-1



381 The OH radical attack on PAHs/DMSO yields carbon-centered radicals or alkyl radical which  
382 further react with O<sub>2</sub> yielding peroxy (RO<sub>2</sub>) and hydroxyperoxy radicals (HO<sub>2</sub>) (Von Sonntag  
383 et al., 1997).

384 Although RO<sub>2</sub> could directly transform into organic sulfates by SO<sub>2</sub>, it has been shown that SO<sub>2</sub>  
385 could accelerate heterogeneous OH oxidation rates by 10 to 20 times, originating from the  
386 radical chain reactions propagated by alkoxy radicals formed by the reaction of peroxy radicals  
387 with SO<sub>2</sub> (Richards-Henderson et al., 2016). Peroxy radicals undergoes self-reactions leading  
388 to the formation of stable products (ketone or alcohol) or alkoxy radicals. (Richards-Henderson  
389 et al., 2016) In the presence of <sup>3</sup>SO<sub>2</sub><sup>\*</sup>, the OH production rate increases by several orders of  
390 magnitude at the air-water interface, thus resulting the increase of radical chain length by alkoxy  
391 radicals (Martins-Costa et al., 2018).

392 In this study, the comprehensive reaction schemes fully explain the detected sulfur-containing  
393 unsaturated multifunctional compounds emerging from light-induced SO<sub>2</sub> oxidation of PAHs  
394 and DMSO at the air-water interface. The tentative reaction pathways describing the formation  
395 of aqueous-phase products including their description are given in the text section S2 and  
396 Scheme S1. The suggested reaction mechanism for the formation of gas-phase product  
397 compounds is shown in Scheme 1 in the following section.

### 398 **3.5. Reaction Mechanism of the Gaseous Compounds**

399 A detailed mechanism for the <sup>3</sup>SO<sub>2</sub><sup>\*</sup> oxidation of PAHs/DMSO is presented in Scheme 1, which  
400 could be divided into two proposed general pathways including 1) self-oxidation of <sup>3</sup>SO<sub>2</sub><sup>\*</sup> and  
401 oxidation of DMSO initiated by <sup>3</sup>SO<sub>2</sub><sup>\*</sup> and <sup>3</sup>PAH<sup>\*</sup>, and 2) photodegradation of sulfur-  
402 containing PAH compounds that were initially formed from PAHs by <sup>3</sup>SO<sub>2</sub><sup>\*</sup>.

403 *Pathway A:* In this pathway, we emphasize the formation and transformation of MSA that plays  
404 a key role during the formation of the detected OSs (Scheme 1A).



405 The self-oxidation of  $^3\text{SO}_2^*$  could yield sulfuric acid ( $\text{H}_2\text{SO}_4$ ) (1). Meanwhile, DMSO can be  
406 oxidized by oxygen and the OH radical formed by  $^3\text{SO}_2^*$  and  $^3\text{PAH}^*$ , yielding MSIA ( $\text{CH}_4\text{O}_2\text{S}$ )  
407 (2) and MSM ( $\text{C}_2\text{H}_6\text{O}_2\text{S}$ ) (3). Dehydration of MSIA (2) (Urbanski et al., 1998; Kukui et al.,  
408 2003; Allen et al., 1999; Arsene et al., 2002) and sulfuric acid (1) would lead to the production  
409 of MSA ( $\text{CH}_4\text{SO}_3$ ) (4). More radical species such as the methyl radical ( $\text{CH}_3$ ), would be formed  
410 during the OH-addition route in the gas-phase atmospheric oxidation of DMSO, resulting in the  
411 formation of  $\text{SO}_2$  that would participate in the subsequent oxidation reactions (Falbe-Hansen et  
412 al., 2000; González-García et al., 2006; Arsene et al., 2002; Urbanski et al., 1998).

413 Further transformation would occur upon oxidation of MSA (4) or self-oxidation of  $^3\text{SO}_2^*$  to  
414 yield MSAOH ( $\text{CH}_4\text{O}_4\text{S}$ ) (5) via  $\text{OH}_2$  radical. The OH and  $\text{CH}_2\text{OH}$  radicals could oxidize MSA  
415 (4) into ESAOH ( $\text{C}_2\text{H}_6\text{O}_4\text{S}$ ) (6) and  $\text{C}_2\text{H}_6\text{O}_3\text{S}$  (7).

416 *Pathway B:* Although linear OS products would also be generated upon addition of  $^3\text{SO}_2^*$  to  
417 double bonds, formed after ring-opening of the PAH molecules, we stress that the occurrence  
418 of gaseous OSs mostly emerges from the photodegradation of aromatic OS products. Under our  
419 experimental conditions,  $^3\text{SO}_2^*$  oxidation of PAHs would prevail over PAH photodegradation  
420 and would lead to sulfur-containing PAHs. Here we only present the general degradation  
421 process and one proposed pathway for the given species. The details of intermediate products  
422 were not shown in this scheme.

423 For example,  $\text{C}_9\text{H}_8\text{O}_4\text{S}$  (9),  $\text{C}_{10}\text{H}_6\text{O}_3\text{S}$  (10), and  $\text{C}_{10}\text{H}_6\text{O}_4\text{S}$  (11) (Scheme 1B) appear as the  
424 photodegradation products of  $\text{C}_{16}\text{H}_{10}\text{O}_3\text{S}$  or  $\text{C}_{16}\text{H}_{12}\text{O}_3\text{S}$ ,  $\text{C}_{14}\text{H}_{10}\text{O}_3\text{S}$  and  $\text{C}_{14}\text{H}_{10}\text{O}_4\text{S}$  (Scheme  
425 S1). These compounds were further transformed upon oxygen attack, then followed by cleavage  
426 of the phenyl ring through a highly exoergonic process. Cleavage of the five-member ring  
427 would yield  $\text{C}_{10}\text{H}_8\text{O}_3\text{S}$  (12) and  $\text{C}_{10}\text{H}_8\text{O}_4\text{S}$  (13), followed by the yield of phenyl hydrogen  
428 sulfate ( $\text{C}_6\text{H}_6\text{O}_4\text{S}$ ) (14) via phenyl-ring cleavage. Meanwhile, phenyl-ring cleavage of  
429 condensed aromatics would also yield  $\text{C}_5\text{H}_6\text{O}_4\text{S}$  (15) and 1,3,2-dioxathiole 2,2-dioxide



430 (C<sub>2</sub>H<sub>4</sub>O<sub>4</sub>S) (16). Subsequently, the attack of oxygen or radicals initiates the exoergonic opening  
431 of a five-member ring or a phenyl ring, resulting in linear OSs under a rapid oxidative scission.





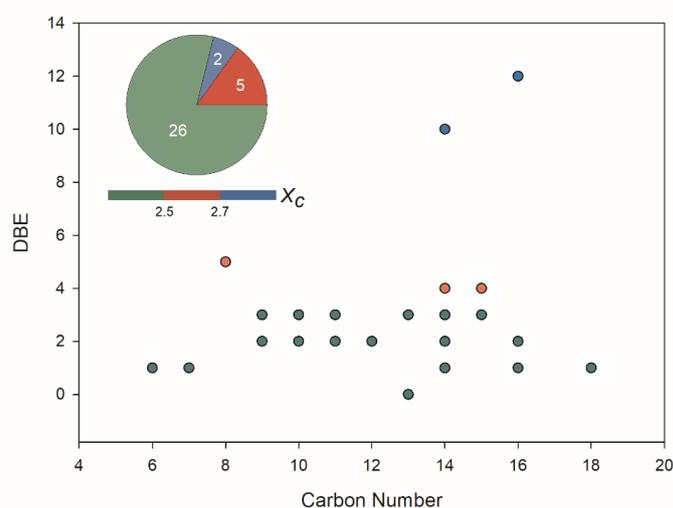
### 400        **3.6. Comparison with OSs Identified Within Atmospheric Aerosols**

401 Table S6 shows the intercomparison of the OSs detected here, mainly formed upon light-  
402 induced reactions of SO<sub>2</sub> with PAHs/DMSO, with OSs identified during various field  
403 campaigns. It should be stressed that the agreement between the chemical formulae of the OSs  
404 detected in this study and those from field campaigns does not necessarily imply the same  
405 molecular structure, because multiple structural isomers are plausible for each formula (Mekic  
406 et al., 2020a; Nizkorodov et al., 2011). A total of 81 tentative OSs from this study overlapped  
407 with those identified in ambient aerosol samples, wherein 33 OS formulae were detected in the  
408 aqueous phase. Only 4 liquid-phase OS formulae were detected following the reactions of SO<sub>2</sub>  
409 with PAH/DMSO in the dark. The blue-shaded OSs are generated exclusively by the reaction  
410 of SO<sub>2</sub> with DMSO. The group of compounds with yellow-shaded area are produced by the  
411 reaction of SO<sub>2</sub> with PAHs/DMSO. Finally, the green-shaded area summarizes the OSs  
412 detected with both SO<sub>2</sub> + DMSO and SO<sub>2</sub> + PAHs/DMSO. Most of the liquid-phase OS  
413 formulae (chemical formulae in bold) were formed individually upon SO<sub>2</sub> oxidation of  
414 PYR/DMSO and DMSO. The product compounds individually formed by the reaction of SO<sub>2</sub>  
415 with PHE/DMSO and the shared formulae generated by the reaction of SO<sub>2</sub> with DMSO and  
416 PAHs/DMSO make the biggest contribution to the gas-phase OSs. These observations highlight  
417 the importance of the SO<sub>2</sub> oxidation reactions of DMSO and/or PAHs/DMSO at the freshwater  
418 and sea surface, or in the liquid films of the aerosol particles, which would represent an  
419 important source of OSs.

420 Tentatively identified VOC precursors are also highlighted in Table S6. The tentative  
421 precursors of more than half of the identified OSs in the atmosphere still remain unexplained  
422 and unknown. Most of the identified overlapped OSs are probably long alkyl chain OSs, thus  
423 their tentative precursors were designated as alkyl-containing OSs from anthropogenic sources  
424 (Tao et al., 2014). There are three overlapped compounds which were assumed as aromatic OSs,



425 thus their precursors were considered as 2-MeNAP and methylbenzyl sulfate (Staudt et al., 2014;  
426 Riva et al., 2015).  
427 Figure 4 shows the iso-abundance plot of DBE vs. carbon numbers according to the  
428 corresponding  $X_c$  of liquid-phase OS formulae, intuitively indicating that most of the identified  
429 OS formulae with low DBE are presumably long-chain aliphatic-like compounds. Moreover,  
430 there were seven detected OSs with aromatic structure that were consistent with those identified  
431 in aerosol samples from field studies. Among those aromatic OSs, only  $C_8H_8O_5S$  has its  
432 tentatively identified VOC precursor, 2-MeNAP (Riva et al., 2015). It has been shown that  
433 aromatic OSs are produced from the interaction of aromatics with sulfur-containing species.  
434 For example, the gas-phase photo-oxidation of PAHs in the presence of sulfate aerosol, and the  
435 aqueous-phase reactions of several monocyclic aromatics with sulfite in the presence of  $Fe^{3+}$   
436 mediated by sulfoxy radical anions can generate aromatic OSs (Riva et al., 2015; Huang et al.,  
437 2020). The reaction of SOA particles formed from photo-oxidation of biodiesel and diesel fuel  
438 with  $SO_2$  was demonstrated to yield a large number of aromatic OS species (Blair et al., 2017).



439

440 **Figure 4:** Iso-abundance plot of DBE vs carbon numbers according to the corresponding  
441 aromaticity equivalent (green with  $X_c < 2.5$ , red with  $2.5 \leq X_c < 2.7$ , and blue with  $X_c \geq 2.7$ )



442 for the detected organic sulfur species in the liquid phase, which had also been identified in  
443 ambient aerosol samples.

444 In our previous study we have identified the OSs formed in the absence of SO<sub>2</sub> as a precursor,  
445 during the photodegradation of DMSO initiated by the excited triplet state of fluorene in both  
446 liquid and gas phase (Mekic et al., 2020a; Mekic et al., 2020b). It is evident that the inclusion  
447 of SO<sub>2</sub> as a precursor in the reaction system enhances the formation of OSs, including those  
448 having similar structures as those detected in field studies. It can be considered that the light-  
449 induced reaction of SO<sub>2</sub> with DMSO or a mixture of DMSO and PAHs is a previously unknown,  
450 additional atmospheric source of OSs (Berresheim et al., 1993; Berresheim and Eisele, 1998;  
451 Karl et al., 2007; Hopkins et al., 2008; Gaston et al., 2010; Ning et al., 2020; Dawson et al.,  
452 2012).

453 Here we show that light-induced heterogeneous processing of SO<sub>2</sub> with PAHs/DMSO  
454 represents an additional source of organic sulphur compounds in the atmosphere and in the  
455 freshwater or seawater. In particular, the reaction of SO<sub>2</sub> with DMSO leads to enhanced  
456 formation of organic sulphur compounds compared to the photosensitized degradation of  
457 DMSO initiated by the excited triplet states of PAHs compounds (Figure S13). This research  
458 results point out to the complexity of the chemical processes responsible for the formation of  
459 organic sulphur compounds. Considering the abundance of SO<sub>2</sub> in the atmosphere and the  
460 omnipresence of water adsorbed PAHs and DMSO compounds, the future modelling studies  
461 should consider both pathways, 1) photosensitized degradation of DMSO initiated by the  
462 excited triplet states of PAHs (Zhang et al., 2019; Jiang et al., 2021), and 2) heterogeneous  
463 chemistry of SO<sub>2</sub> with PAHs/DMSO and DMSO, as alternative formation pathways of organic  
464 sulphur compounds in the atmosphere. In this study, many detected aromatic and linear OSs  
465 formed during the light-induced SO<sub>2</sub> oxidation of PAHs/DMSO are reported for the first time.  
466 An important number of detected compounds overlapped with those of ambient OSs identified



467 during field measurements. We suggest that a plausible mechanism for OSs formation *via* direct  
468 addition of SO<sub>2</sub> to the C=C double bond is not only limited to unsaturated acids but is also valid  
469 for anthropogenic precursors such as PAHs. A large amount of organosulfates and especially  
470 aromatic organosulfates could be formed through this pathway and released into water and  
471 ambient atmosphere. These OSs can form surfactant films on aerosol particles in the boundary  
472 layer, by which means they influence the surface tension and hygroscopicity of particles  
473 (Decesari et al., 2011; Tao et al., 2014). Indeed, the OSs formation pathway in the  
474 heterogeneous reaction between SO<sub>2</sub> and PAHs/DMSO is of great significance because OSs  
475 generation at the water surface would influence the air-water exchange, and enhanced gaseous  
476 OSs would result in the formation of SOA. Moreover, the OS products formation from PAHs  
477 initiated by SO<sub>2</sub> is also of importance in urban areas where PAH concentrations are usually  
478 high in ambient air (Zhu et al., 2019; Cai et al., 2020). Finally, aromatic OSs occurring in urban  
479 aerosols represent a still unrecognized source of toxic products (Riva et al., 2015).

480



481 **Supporting Information**

482 Additional 16 figures, 6 tables and 1 reaction scheme. Short description of MI-SPI-TOF-MS  
483 and FT-ICR-MS. Analysis of FT - ICR - MS aqueous phase products based on DBE vs carbon  
484 number iso-abundance plot. Reaction mechanism of the aqueous phase product compounds.

485

486 **Acknowledgements**

487 This study was financially supported by National Natural Science Foundation of China (N<sup>o</sup>:  
488 42007200, N<sup>o</sup>: 41773131, N<sup>o</sup>: 41977187, and N<sup>o</sup>: 42177087), Guangdong Foundation for  
489 Program of Science and Technology Research (N<sup>o</sup>: 2021A1515011555), and China  
490 Postdoctoral Science Foundation (N<sup>o</sup>: 2019M653105). This study was funded by Institute  
491 Director Foundation of GIG (N<sup>o</sup>: 2019SZJJ-10), International Cooperation Grant of Chinese  
492 Academy of Science (N<sup>o</sup>: 132744KYSB20190007), State Key Laboratory of Organic  
493 Geochemistry, Guangzhou Institute of Geochemistry (SKLOG2020-5, and KTZ\_17101).

494 The authors thank Dr. Jiangping Liu, Huifan Deng, Wentao Zhou and Gwendal Loisel for their  
495 assistance in the lab analysis. All data in this manuscript are freely available upon request  
496 through the corresponding author (gligorovski@gig.ac.cn). This is a contribution of the  
497 GIGCAS.

498

499 **Author Contributions**

500 H.J. and S.G. wrote the paper; S.G. designed the research; H.J. and Y.W. performed the  
501 laboratory experiments; H.J. interpreted the MI-SPI-TOF-MS and FT-ICR-MS data and  
502 relevant discussion; H.J. Y.H and S.L speculated reaction pathways; Y.H. carried out the Gibbs  
503 energies theoretical calculation of the reaction initiated by <sup>3</sup>PAH\* and <sup>3</sup>SO<sub>2</sub>\*; B.J. contributed  
504 to the FT-ICR-MS data analysis. X.L. contributed to the MI-SPI-TOF-MS data analysis, L.Y.  
505 and T.L. contributed to the relevant discussion of reaction mechanisms; L.C., and D.V.



506 contributed to the discussion of the results and implications. All the authors contributed to  
507 revise the manuscript and approved the final version.

508

509 **Author information**

510 Corresponding Author

511 \* Phone: +86 2085291497; Email: gligorovski@gig.ac.cn

512

513 **Notes**

514 The authors declare no competing financial interest.

515

516



## 517 References

- 518 Allen, H. C., Gragson, D., and Richmond, G.: Molecular structure and adsorption of dimethyl sulfoxide  
519 at the surface of aqueous solutions, *J. Phys. Chem. B*, 103, 660-666, 1999.
- 520 Altieri, K. E., Carlton, A. G., Lim, H.-J., Turpin, B. J., and Seitzinger, S. P.: Evidence for oligomer formation  
521 in clouds: Reactions of isoprene oxidation products, *Environ. Sci. Technol.*, 40, 4956-4960,  
522 10.1021/es052170n, 2006.
- 523 Andreae, M. O. J. L.: Dimethylsulfoxide in marine and freshwaters, *Limnol. Oceanogr.*, 25, 1054-1063,  
524 1980.
- 525 Arsene, C., Barnes, I., Becker, K. H., Schneider, W. F., Wallington, T. T., Mihalopoulos, N., and Patroescu-  
526 Klotz, I. V. J. E. s.: Formation of methane sulfinic acid in the gas-phase OH-radical initiated oxidation of  
527 dimethyl sulfoxide, *Environ. Sci. Technol.*, 36, 5155-5163, 2002.
- 528 Asher, E., Dacey, J. W., Ianson, D., Peña, A., and Tortell, P. D.: Concentrations and cycling of DMS, DMSP,  
529 and DMSO in coastal and offshore waters of the Subarctic Pacific during summer, 2010 - 2011, *J.*  
530 *Geophys. Res. Oceans*, 122, 3269-3286, 2017.
- 531 Barnes, I., Hjorth, J., and Mihalopoulos, N.: Dimethyl sulfide and dimethyl sulfoxide and their oxidation  
532 in the atmosphere, *Chem. Rev.*, 106, 940-975, 10.1021/cr020529+, 2006.
- 533 Benson, N. U., Essien, J. P., Asuquo, F. E., and Eritobor, A. L.: Occurrence and distribution of polycyclic  
534 aromatic hydrocarbons in surface microlayer and subsurface seawater of Lagos Lagoon, Nigeria,  
535 *Environ. Monit. Assess.*, 186, 5519-5529, 2014.
- 536 Berresheim, H. and Eisele, F.: Sulfur chemistry in the Antarctic Troposphere Experiment: An overview  
537 of project SCATE, *J. Geophys. Res-Atmos*, 103, 1619-1627, 1998.
- 538 Berresheim, H., Eisele, F., Tanner, D., McInnes, L., Ramsey - Bell, D., and Covert, D.: Atmospheric sulfur  
539 chemistry and cloud condensation nuclei (CCN) concentrations over the northeastern Pacific coast, *J.*  
540 *Geophys. Res-Atmos*, 98, 12701-12711, 1993.
- 541 Berresheim, H., Elste, T., Tremmel, H. G., Allen, A. G., Hansson, H. C., Rosman, K., Dal Maso, M., Makela,  
542 J. M., Kulmala, M., and O'Dowd, C. D.: Gas-aerosol relationships of H<sub>2</sub>SO<sub>4</sub>, MSA, and OH: Observations  
543 in the coastal marine boundary layer at Mace Head, Ireland, *J. Geophys. Res-Atmos*, 107,  
544 10.1029/2000jd000229, 2002.
- 545 Binning Jr, R. and Curtiss, L.: Compact contracted basis sets for third - row atoms: Ga - Kr, *J. Comput.*  
546 *Chem.*, 11, 1206-1216, 1990.
- 547 Blair, S. L., MacMillan, A. C., Drozd, G. T., Goldstein, A. H., Chu, R. K., Paša-Tolić, L., Shaw, J. B., Tolić, N.,  
548 Lin, P., and Laskin, J.: Molecular characterization of organosulfur compounds in biodiesel and diesel  
549 fuel secondary organic aerosol, *Environ. Sci. Technol.*, 51, 119-127, 2017.
- 550 Bork, N., Elm, J., Olenius, T., and Vehkamäki, H.: Methane sulfonic acid-enhanced formation of  
551 molecular clusters of sulfuric acid and dimethyl amine, *Atmospheric Chem. Phys.*, 14, 12023-12030,  
552 10.5194/acp-14-12023-2014, 2014.
- 553 Brigante, M., Charbouillot, T., Vione, D., and Mailhot, G.: Photochemistry of 1-nitronaphthalene: A  
554 potential source of singlet oxygen and radical species in atmospheric waters, *J. Phys. Chem. A*, 114,  
555 2830-2836, 2010.
- 556 Brimblecombe, P. and Shooter, D.: Photo-oxidation of dimethylsulphide in aqueous solution, *Mar.*  
557 *Chem.*, 19, 343-353, 1986.
- 558 Bruggemann, M., Xu, R., Tilgner, A., Kwong, K. C., Mutzel, A., Poon, H. Y., Otto, T., Schaefer, T., Poulain,  
559 L., Chan, M. N., and Herrmann, H.: Organosulfates in Ambient Aerosol: State of Knowledge and Future  
560 Research Directions on Formation, Abundance, Fate, and Importance, *Environ. Sci. Technol.*, 54, 3767-  
561 3782, 10.1021/acs.est.9b06751, 2020.
- 562 Cai, D., Wang, X., Chen, J., and Li, X.: Molecular Characterization of Organosulfates in Highly Polluted  
563 Atmosphere Using Ultra-High-Resolution Mass Spectrometry, *J. Geophys. Res-Atmos*, 125,  
564 10.1029/2019jd032253, 2020.



- 565 Chen, H., Varner, M. E., Gerber, R. B., and Finlayson-Pitts, B. J.: Reactions of Methanesulfonic Acid with  
566 Amines and Ammonia as a Source of New Particles in Air, *J. Phys. Chem. B*, **120**, 1526-1536,  
567 10.1021/acs.jpcc.5b07433, 2016.
- 568 Chen, J., Ehrenhauser, F. S., Valsaraj, K. T., and Wornat, M. J.: Uptake and UV-photooxidation of gas-  
569 phase PAHs on the surface of atmospheric water films. 1. Naphthalene, *J. Phys. Chem. A*, **110**, 9161-  
570 9168, 2006.
- 571 Chen, Q., Sherwen, T., Evans, M., and Alexander, B.: DMS oxidation and sulfur aerosol formation in the  
572 marine troposphere: a focus on reactive halogen and multiphase chemistry, *Atmospheric Chem. Phys.*,  
573 **18**, 13617-13637, 10.5194/acp-18-13617-2018, 2018.
- 574 Cincinelli, A., Stortini, A. M., Perugini, M., Checchini, L., and Lepri, L.: Organic pollutants in sea-surface  
575 microlayer and aerosol in the coastal environment of Leghorn - (Tyrrhenian Sea), *Mar. Chem.*, **76**, 77-  
576 98, 10.1016/s0304-4203(01)00049-4, 2001.
- 577 Dall'Osto, M., Ceburnis, D., Monahan, C., Worsnop, D. R., Bialek, J., Kulmala, M., Kurten, T., Ehn, M.,  
578 Wenger, J., Sodeau, J., Healy, R., and O'Dowd, C.: Nitrogenated and aliphatic organic vapors as possible  
579 drivers for marine secondary organic aerosol growth, *J. Geophys. Res-Atmos*, **117**,  
580 10.1029/2012jd017522, 2012.
- 581 Davidovits, P., Kolb, C. E., Williams, L. R., Jayne, J. T., and Worsnop, D. R.: Mass accommodation and  
582 chemical reactions at gas-liquid interfaces, *Chem. Rev.*, **106**, 1323-1354, 2006.
- 583 Dawson, M. L., Varner, M. E., Perraud, V., Ezell, M. J., Gerber, R. B., and Finlayson-Pitts, B. J.: Simplified  
584 mechanism for new particle formation from methanesulfonic acid, amines, and water via experiments  
585 and ab initio calculations, *Proc. Natl. Acad. Sci. U.S.A.*, **109**, 18719-18724, 2012.
- 586 Decesari, S., Finessi, E., Rinaldi, M., Paglione, M., Fuzzi, S., Stephanou, E., Tziaras, T., Spyros, A.,  
587 Ceburnis, D., and O'Dowd, C.: Primary and secondary marine organic aerosols over the North Atlantic  
588 Ocean during the MAP experiment, *J. Geophys. Res-Atmos*, **116**, 2011.
- 589 Deng, H., Liu, J., Wang, Y., Song, W., Wang, X., Li, X., Vione, D., and Gligorovski, S.: Effect of Inorganic  
590 Salts on N-Containing Organic Compounds Formed by Heterogeneous Reaction of NO<sub>2</sub> with Oleic Acid,  
591 *Environ. Sci. Technol.*, **55**, 7831-7840, 10.1021/acs.est.1c01043, 2021.
- 592 Donaldson, D., Kahan, T., Kwamena, N., Handley, S., and Barbier, C.: Atmospheric chemistry of urban  
593 surface films, in: *Atmospheric Aerosols Characterization, Chemistry, Modeling, and Climate*, ACS  
594 Publications, 79-89, 2009.
- 595 Falbe-Hansen, H., Sørensen, S., Jensen, N., Pedersen, T., and Hjorth, J.: Atmospheric gas-phase  
596 reactions of dimethylsulphoxide and dimethylsulphone with OH and NO<sub>3</sub> radicals, Cl atoms and ozone,  
597 *Atmos. Environ.*, **34**, 1543-1551, 2000.
- 598 Frisch, M., Trucks, G., Schlegel, H., Scuseria, G., Robb, M., Cheeseman, J., Scalmani, G., Barone, V.,  
599 Petersson, G., and Nakatsuji, H.: Gaussian 16, revision C. 01, 2016.
- 600 Gaston, C. J., Pratt, K. A., Qin, X., and Prather, K. A.: Real-time detection and mixing state of  
601 methanesulfonate in single particles at an inland urban location during a phytoplankton bloom,  
602 *Environ. Sci. Technol.*, **44**, 1566-1572, 2010.
- 603 González-García, N., González-Lafont, À., and Lluch, J. M.: Variational transition-state theory study of  
604 the dimethyl sulfoxide (DMSO) and OH reaction, *J. Phys. Chem. A*, **110**, 798-808, 2006.
- 605 González-Gaya, B., Fernández-Pinos, M.-C., Morales, L., Méjanelle, L., Abad, E., Piña, B., Duarte, C. M.,  
606 Jiménez, B., and Dachs, J.: High atmosphere-ocean exchange of semivolatile aromatic hydrocarbons,  
607 *Nat. Geosci.*, **9**, 438-442, 10.1038/ngeo2714, 2016.
- 608 González-Gaya, B., Martínez-Varela, A., Vila-Costa, M., Casal, P., Cerro-Gálvez, E., Berrojalbiz, N.,  
609 Lundin, D., Vidal, M., Mompeán, C., and Bode, A.: Biodegradation as an important sink of aromatic  
610 hydrocarbons in the oceans, *Nat. Geosci.*, **12**, 119-125, 2019.
- 611 Grossman, J. N., Stern, A. P., Kirich, M. L., and Kahan, T. F.: Anthracene and pyrene photolysis kinetics  
612 in aqueous, organic, and mixed aqueous-organic phases, *Atmos. Environ.*, **128**, 158-164, 2016.
- 613 Guitart, C., Garcia-Flor, N., Bayona, J. M., and Albaiges, J.: Occurrence and fate of polycyclic aromatic  
614 hydrocarbons in the coastal surface microlayer, *Mar. Pollut. Bull.*, **54**, 186-194,  
615 10.1016/j.marpolbul.2006.10.008, 2007.



- 616 Hardy, J. T., Crecelius, E. A., Antrim, L. D., Kiesser, S. L., Broadhurst, V. L., Boehm, P. D., Steinhauer, W.  
617 G., and Coogan, T. H.: Aquatic surface microlayer contamination in Chesapeake Bay, *Mar. Chem.*, **28**,  
618 333-351, 1990.
- 619 Harvey, G. R. and Lang, R. F.: Dimethylsulfoxide and dimethylsulfone in the marine atmosphere,  
620 *Geophys. Res. Lett.*, **13**, 49-51, 1986.
- 621 Hatton, A., Malin, G., Turner, S., and Liss, P.: DMSO, A Significant Compound in the Biogeochemical  
622 Cycle of DMS., in: *Biological and Environmental Chemistry of DMSP and Related Sulfonium Compounds*,  
623 Springer, 405-412, 1996.
- 624 Hoffmann, E. H., Tilgner, A., Schrodner, R., Brauer, P., Wolke, R., and Herrmann, H.: An advanced  
625 modeling study on the impacts and atmospheric implications of multiphase dimethyl sulfide chemistry,  
626 *Proc Natl Acad Sci U S A*, **113**, 11776-11781, 10.1073/pnas.1606320113, 2016.
- 627 Hopkins, R. J., Desyatnik, Y., Tivanski, A. V., Zaveri, R. A., Berkowitz, C. M., Tyliszczak, T., Gilles, M. K.,  
628 and Laskin, A.: Chemical speciation of sulfur in marine cloud droplets and particles: Analysis of  
629 individual particles from the marine boundary layer over the California current, *J. Geophys. Res.-Atmos.*,  
630 **113**, 2008.
- 631 Huang, L., Liu, T., and Grassian, V. H.: Radical-Initiated Formation of Aromatic Organosulfates and  
632 Sulfonates in the Aqueous Phase, *Environ. Sci. Technol.*, **54**, 11857-11864, 10.1021/acs.est.0c05644,  
633 2020.
- 634 Jiang, B., Kuang, B. Y., Liang, Y., Zhang, J., Huang, X. H. H., Xu, C., Yu, J. Z., and Shi, Q.: Molecular  
635 composition of urban organic aerosols on clear and hazy days in Beijing: a comparative study using FT-  
636 ICR MS, *Environ. Chem.*, **13**, 888-901, 10.1071/en15230, 2016.
- 637 Jiang, H., Carena, L., He, Y., Wang, Y., Zhou, W., Yang, L., Luan, T., Li, X., Brigante, M., Vione, D., and  
638 Gligorovski, S.: Photosensitized degradation of DMSO initiated by PAHs at the air-water interface, as  
639 an alternative source of organic sulfur compounds to the atmosphere, *J. Geophys. Res.-Atmos.*, **126**,  
640 10.1029/2021JD035346, 2021.
- 641 Kamens, R. M., Zhang, H., Chen, E. H., Zhou, Y., Parikh, H. M., Wilson, R. L., Galloway, K. E., and Rosen,  
642 E. P.: Secondary organic aerosol formation from toluene in an atmospheric hydrocarbon mixture:  
643 Water and particle seed effects, *Atmos. Environ.*, **45**, 2324-2334, 10.1016/j.atmosenv.2010.11.007,  
644 2011.
- 645 Karl, M., Gross, A., Leck, C., and Pirjola, L.: Intercomparison of dimethylsulfide oxidation mechanisms  
646 for the marine boundary layer: Gaseous and particulate sulfur constituents, *J. Geophys. Res.-Atmos.*,  
647 **112**, 10.1029/2006jd007914, 2007.
- 648 Kourtchev, I., O'Connor, I. P., Giorio, C., Fuller, S. J., Kristensen, K., Maenhaut, W., Wenger, J. C., Sodeau,  
649 J. R., Glasius, M., and Kalberer, M.: Effects of anthropogenic emissions on the molecular composition  
650 of urban organic aerosols: An ultrahigh resolution mass spectrometry study, *Atmos. Environ.*, **89**, 525-  
651 532, 10.1016/j.atmosenv.2014.02.051, 2014.
- 652 Kroll, J. A., Frandsen, B. N., Kjaergaard, H. G., and Vaida, V.: Atmospheric hydroxyl radical source:  
653 Reaction of triplet SO<sub>2</sub> and water, *J. Phys. Chem.A*, **122**, 4465-4469, 2018.
- 654 Kukui, A., Borissenko, D., Laverdet, G., and Le Bras, G.: Gas-phase reactions of OH radicals with  
655 dimethyl sulfoxide and methane sulfinic acid using turbulent flow reactor and chemical ionization mass  
656 spectrometry, *J. Phys. Chem.A*, **107**, 5732-5742, 2003.
- 657 Kundu, S., Quraishi, T. A., Yu, G., Suarez, C., Keutsch, F. N., and Stone, E. A.: Evidence and quantitation  
658 of aromatic organosulfates in ambient aerosols in Lahore, Pakistan, *Atmospheric Chem. Phys.*, **13**,  
659 4865-4875, 10.5194/acp-13-4865-2013, 2013.
- 660 Lammel, G.: Polycyclic aromatic compounds in the atmosphere—a review identifying research needs,  
661 *Polycyclic Aromat. Compd.*, **35**, 316-329, 2015.
- 662 Lee, P. and De Mora, S.: DMSP, DMS and DMSO concentrations and temporal trends in marine surface  
663 waters at Leigh, New Zealand, in: *Biological and Environmental Chemistry of DMSP and Related*  
664 *Sulfonium Compounds*, Springer, 391-404, 1996.
- 665 Lee, P. A., de Mora, S. J., and Lévassieur, M.: A review of dimethylsulfoxide in aquatic environments,  
666 *Atmosphere-Ocean*, **37**, 439-456, 10.1080/07055900.1999.9649635, 1999.



- 667 Legrand, M., Sciare, J., Jourdain, B., and Genthon, C.: Subdaily variations of atmospheric  
668 dimethylsulfide, dimethylsulfoxide, methanesulfonate, and non - sea - salt sulfate aerosols in the  
669 atmospheric boundary layer at Dumont d'Urville (coastal Antarctica) during summer, *J. Geophys. Res-*  
670 *Atmos*, 106, 14409-14422, 2001.
- 671 Li, G., Bei, N., Cao, J., Huang, R., Wu, J., Feng, T., Wang, Y., Liu, S., Zhang, Q., Tie, X., and Molina, L. T.:  
672 A possible pathway for rapid growth of sulfate during haze days in China, *Atmospheric Chem. Phys.*,  
673 17, 3301-3316, 10.5194/acp-17-3301-2017, 2017a.
- 674 Li, J., Li, F., and Liu, Q.: PAHs behavior in surface water and groundwater of the Yellow River estuary:  
675 evidence from isotopes and hydrochemistry, *Chemosphere*, 178, 143-153, 2017b.
- 676 Librando, V., Tringali, G., Hjorth, J., and Coluccia, S.: OH-initiated oxidation of DMS/DMSO: reaction  
677 products at high NO<sub>x</sub> levels, *Environ. Pollut.*, 127, 403-410, 2004.
- 678 Librando, V., Bracchitta, G., de Guidi, G., Minniti, Z., Perrini, G., and Catalfo, A.: Photodegradation of  
679 anthracene and benzo [a] anthracene in polar and apolar media: new pathways of photodegradation,  
680 *Polycyclic Aromat. Compd.*, 34, 263-279, 2014.
- 681 Lin, P., Rincon, A. G., Kalberer, M., and Yu, J. Z.: Elemental composition of HULIS in the Pearl River Delta  
682 Region, China: results inferred from positive and negative electrospray high resolution mass  
683 spectrometric data, *Environ. Sci. Technol.*, 46, 7454-7462, 10.1021/es300285d, 2012.
- 684 Lohmann, R., Gioia, R., Jones, K. C., Nizzetto, L., Temme, C., Xie, Z., Schulz-Bull, D., Hand, I., Morgan, E.,  
685 and Jantunen, L. J. E. s.: Organochlorine pesticides and PAHs in the surface water and atmosphere of  
686 the North Atlantic and Arctic Ocean, *Environ. Sci. Technol.*, 43, 5633-5639, 2009.
- 687 Ma, Y., Xu, X., Song, W., Geng, F., and Wang, L.: Seasonal and diurnal variations of particulate  
688 organosulfates in urban Shanghai, China, *Atmos. Environ.*, 85, 152-160,  
689 10.1016/j.atmosenv.2013.12.017, 2014.
- 690 Ma, Y., Xie, Z., Yang, H., Möller, A., Halsall, C., Cai, M., Sturm, R., and Ebinghaus, R.: Deposition of  
691 polycyclic aromatic hydrocarbons in the North Pacific and the Arctic, *J. Geophys. Res-Atmos*, 118, 5822-  
692 5829, 2013.
- 693 Martins-Costa, M. T., Anglada, J. M., Francisco, J. S., and Ruiz-López, M. F.: Photochemistry of SO<sub>2</sub> at  
694 the air–water interface: a source of OH and HOSO radicals, *J. Am. Chem. Soc.*, 140, 12341-12344, 2018.
- 695 McLean, A. and Chandler, G.: Contracted Gaussian basis sets for molecular calculations. I. Second row  
696 atoms, Z= 11–18, *J. Chem. Phys.*, 72, 5639-5648, 1980.
- 697 Mekic, M., Zeng, J., Jiang, B., Li, X., Lazarou, Y. G., Brigante, M., Herrmann, H., and Gligorovski, S.:  
698 Formation of Toxic Unsaturated Multifunctional and Organosulfur Compounds From the  
699 Photosensitized Processing of Fluorene and DMSO at the Air - Water Interface, *J. Geophys. Res-Atmos*,  
700 125, 10.1029/2019jd031839, 2020a.
- 701 Mekic, M., Zeng, J., Zhou, W., Loisel, G., Jin, B., Li, X., Vione, D., and Gligorovski, S.: Ionic Strength Effect  
702 on Photochemistry of Fluorene and Dimethylsulfoxide at the Air–Sea Interface: Alternative Formation  
703 Pathway of Organic Sulfur Compounds in a Marine Atmosphere *ACS Earth Space Chem.*, 4, 1029-1038,  
704 10.1021/acsearthspacechem.0c00059, 2020b.
- 705 Monge, M. E., George, C., D'Anna, B., Doussin, J.-F., Jammoul, A., Wang, J., Eyglunet, G., Solignac, G.,  
706 Daele, V., and Mellouki, A.: Ozone formation from illuminated titanium dioxide surfaces, *J. Am. Chem.*  
707 *Soc.*, 132, 8234-8235, 2010.
- 708 Ning, A., Zhang, H., Zhang, X., Li, Z., Zhang, Y., Xu, Y., and Ge, M.: A molecular-scale study on the role  
709 of methanesulfinic acid in marine new particle formation, *Atmos. Environ.*, 227,  
710 10.1016/j.atmosenv.2020.117378, 2020.
- 711 Nizkorodov, S. A., Laskin, J., and Laskin, A.: Molecular chemistry of organic aerosols through the  
712 application of high resolution mass spectrometry, *PCCP*, 13, 3612-3629, 2011.
- 713 Oppenheimer, C., Francis, P., Burton, M., Maciejewski, A., and Boardman, L.: Remote measurement of  
714 volcanic gases by Fourier transform infrared spectroscopy, *Appl. Phys. B: Lasers Opt.*, 67, 1998.
- 715 Otto, S., Streibel, T., Erdmann, S., Klingbeil, S., Schulz-Bull, D., and Zimmermann, R.: Pyrolysis–gas  
716 chromatography–mass spectrometry with electron-ionization or resonance-enhanced-multi-photon-



- 717 ionization for characterization of polycyclic aromatic hydrocarbons in the Baltic Sea, *Mar. Pollut. Bull.*,  
718 99, 35-42, 2015.
- 719 Passananti, M., Kong, L., Shang, J., Dupart, Y., Perrier, S., Chen, J., Donaldson, D. J., and George, C.:  
720 Organosulfate Formation through the Heterogeneous Reaction of Sulfur Dioxide with Unsaturated  
721 Fatty Acids and Long - Chain Alkenes, *Angew. Chem. Int. Ed.*, 55, 10336-10339, 2016.
- 722 Pérez-Carrera, E., León, V. M. L., Parra, A. G., and González-Mazo, E.: Simultaneous determination of  
723 pesticides, polycyclic aromatic hydrocarbons and polychlorinated biphenyls in seawater and interstitial  
724 marine water samples, using stir bar sorptive extraction–thermal desorption–gas chromatography–  
725 mass spectrometry, *J. Chromatogr. A*, 1170, 82-90, 2007.
- 726 Perraud, V., Horne, J. R., Martinez, A. S., Kalinowski, J., Meinardi, S., Dawson, M. L., Wingen, L. M.,  
727 Dabdub, D., Blake, D. R., Gerber, R. B., and Finlayson-Pitts, B. J.: The future of airborne sulfur-containing  
728 particles in the absence of fossil fuel sulfur dioxide emissions, *Proc Natl Acad Sci U S A*, 112, 13514-  
729 13519, 10.1073/pnas.1510743112, 2015.
- 730 Richards-Henderson, N. K., Goldstein, A. H., Wilson, K. R. J. E. s., and technology: Sulfur dioxide  
731 accelerates the heterogeneous oxidation rate of organic aerosol by hydroxyl radicals, *Environ. Sci.*  
732 *Technol. Lett.*, 50, 3554-3561, 2016.
- 733 Richards, S., Rudd, J., and Kelly, C.: Organic volatile sulfur in lakes ranging in sulfate and dissolved salt  
734 concentration over five orders of magnitude, *Limnol. Oceanogr.*, 39, 562-572, 1994.
- 735 Ridgeway, R. G., Thornton, D. C., and Bandy, A. R.: Determination of trace aqueous dimethylsulfoxide  
736 concentrations by isotope dilution gas chromatography/mass spectrometry: Application to rain and  
737 sea water, *J. Atmos. Chem.*, 14, 53-60, 1992.
- 738 Riva, M., Da Silva Barbosa, T., Lin, Y.-H., Stone, E. A., Gold, A., and Surratt, J. D.: Chemical  
739 characterization of organosulfates in secondary organic aerosol derived from the photooxidation of  
740 alkanes, *Atmospheric Chem. Phys.*, 16, 11001-11018, 10.5194/acp-16-11001-2016, 2016.
- 741 Riva, M., Tomaz, S., Cui, T., Lin, Y. H., Perraudin, E., Gold, A., Stone, E. A., Villenave, E., and Surratt, J.  
742 D.: Evidence for an unrecognized secondary anthropogenic source of organosulfates and sulfonates:  
743 gas-phase oxidation of polycyclic aromatic hydrocarbons in the presence of sulfate aerosol, *Environ.*  
744 *Sci. Technol.*, 49, 6654-6664, 10.1021/acs.est.5b00836, 2015.
- 745 Rosati, B., Christiansen, S., de Jonge, R. W., Roldin, P., Jensen, M. M., Wang, K., Moosakutty, S. P.,  
746 Thomsen, D., Salomonsen, C., Hyttinen, N., Elm, J., Feilberg, A., Glasius, M., and Bilde, M.: New Particle  
747 Formation and Growth from Dimethyl Sulfide Oxidation by Hydroxyl Radicals, *ACS Earth Space Chem.*,  
748 5, 801-811, 10.1021/acsearthspacechem.0c00333, 2021.
- 749 Schobesberger, S., Junninen, H., Bianchi, F., Lonn, G., Ehn, M., Lehtipalo, K., Dommen, J., Ehrhart, S.,  
750 Ortega, I. K., Franchin, A., Nieminen, T., Riccobono, F., Hutterli, M., Duplissy, J., Almeida, J., Amorim,  
751 A., Breitenlechner, M., Downard, A. J., Dunne, E. M., Flagan, R. C., Kajos, M., Keskinen, H., Kirkby, J.,  
752 Kupc, A., Kurten, A., Kurten, T., Laaksonen, A., Mathot, S., Onnela, A., Praplan, A. P., Rondo, L., Santos,  
753 F. D., Schallhart, S., Schnitzhofer, R., Sipila, M., Tome, A., Tsagkogeorgas, G., Vehkamäki, H., Wimmer,  
754 D., Baltensperger, U., Carslaw, K. S., Curtius, J., Hansel, A., Petaja, T., Kulmala, M., Donahue, N. M., and  
755 Worsnop, D. R.: Molecular understanding of atmospheric particle formation from sulfuric acid and  
756 large oxidized organic molecules, *Proc Natl Acad Sci U S A*, 110, 17223-17228,  
757 10.1073/pnas.1306973110, 2013.
- 758 Seidel, M., Manecki, M., Herlemann, D. P., Deutsch, B., Schulz-Bull, D., Jürgens, K., and Dittmar, T.:  
759 Composition and transformation of dissolved organic matter in the Baltic Sea, *Front. Earth Sci.*, 5, 31,  
760 2017.
- 761 Shang, J., Passananti, M., Dupart, Y., Ciuraru, R., Tinel, L., Rossignol, S. p., Perrier, S. b., Zhu, T., and  
762 George, C.: SO<sub>2</sub> Uptake on oleic acid: A new formation pathway of organosulfur compounds in the  
763 atmosphere, *Environ. Sci. Technol. Lett.*, 3, 67-72, 2016.
- 764 Smith, S. J., Pitcher, H., and Wigley, T. M. L.: Global and regional anthropogenic sulfur dioxide emissions,  
765 *Global Planet. Change*, 29, 99-119, 10.1016/s0921-8181(00)00057-6, 2001.
- 766 Staudt, S., Kundu, S., Lehmler, H.-J., He, X., Cui, T., Lin, Y.-H., Kristensen, K., Glasius, M., Zhang, X.,  
767 Weber, R. J., Surratt, J. D., and Stone, E. A.: Aromatic organosulfates in atmospheric aerosols: Synthesis,



- 768 characterization, and abundance, *Atmos. Environ.*, 94, 366-373, 10.1016/j.atmosenv.2014.05.049,  
769 2014.
- 770 Stortini, A., Martellini, T., Del Bubba, M., Lepri, L., Capodaglio, G., and Cincinelli, A.: n-Alkanes, PAHs  
771 and surfactants in the sea surface microlayer and sea water samples of the Gerlache Inlet sea  
772 (Antarctica), *Microchem. J.*, 92, 37-43, 2009.
- 773 Styler, S., Loiseaux, M.-E., and Donaldson, D.: Substrate effects in the photoenhanced ozonation of  
774 pyrene, *Atmospheric Chem. Phys.*, 11, 1243-1253, 2011.
- 775 Tao, S., Lu, X., Levac, N., Bateman, A. P., Nguyen, T. B., Bones, D. L., Nizkorodov, S. A., Laskin, J., Laskin,  
776 A., and Yang, X.: Molecular characterization of organosulfates in organic aerosols from Shanghai and  
777 Los Angeles urban areas by nanospray-desorption electrospray ionization high-resolution mass  
778 spectrometry, *Environ. Sci. Technol.*, 48, 10993-11001, 2014.
- 779 Urbanski, S., Stickel, R., and Wine, P.: Mechanistic and kinetic study of the gas-phase reaction of  
780 hydroxyl radical with dimethyl sulfoxide, *J. Phys. Chem. A*, 102, 10522-10529, 1998.
- 781 Vácha, R., Jungwirth, P., Chen, J., and Valsaraj, K.: Adsorption of polycyclic aromatic hydrocarbons at  
782 the air–water interface: Molecular dynamics simulations and experimental atmospheric observations,  
783 *PCCP*, 8, 4461-4467, 2006.
- 784 Valavanidis, A., Vlachogianni, T., Triantafyllaki, S., Dassenakis, M., Androustos, F., and Scoullou, M.:  
785 Polycyclic aromatic hydrocarbons in surface seawater and in indigenous mussels (*Mytilus*  
786 *galloprovincialis*) from coastal areas of the Saronikos Gulf (Greece), *Estuar. Coast. Shelf Sci.*, 79, 733-  
787 739, 2008.
- 788 von Sonntag, C., Dowideit, P., Fang, X., Mertens, R., Pan, X., Schuchmann, M. N., Schuchmann, H.-P. J.  
789 W. S., and Technology: The fate of peroxy radicals in aqueous solution, *Water Sci. Technol.*, 35, 9-15,  
790 1997.
- 791 Wang, Y., Mekić, M., Li, P., Deng, H., Liu, S., Jiang, B., Jin, B., Vione, D., and Gligorovski, S.: Ionic Strength  
792 Effect Triggers Brown Carbon Formation through Heterogeneous Ozone Processing of Ortho-Vanillin,  
793 *Environ. Sci. Technol.*, 55, 4553-4564, 10.1021/acs.est.1c00874, 2021.
- 794 Wang, Y., Zhang, Q., Jiang, J., Zhou, W., Wang, B., He, K., Duan, F., Zhang, Q., Philip, S., and Xie, Y.:  
795 Enhanced sulfate formation during China's severe winter haze episode in January 2013 missing from  
796 current models, *J. Geophys. Res.-Atmos.*, 119, 10.1002/2013jd021426, 2014.
- 797 Weigend, F.: Accurate Coulomb-fitting basis sets for H to Rn, *PCCP*, 8, 1057-1065, 2006.
- 798 Weigend, F. and Ahlrichs, R.: Balanced basis sets of split valence, triple zeta valence and quadruple  
799 zeta valence quality for H to Rn: Design and assessment of accuracy, *PCCP*, 7, 3297-3305, 2005.
- 800 Wilkinson, F., Helman, W. P., and Ross, A. B.: Rate constants for the decay and reactions of the lowest  
801 electronically excited singlet state of molecular oxygen in solution. An expanded and revised  
802 compilation, *J. Phys. Chem. Ref. Data*, 24, 663-677, 1995.
- 803 Yassine, M. M., Harir, M., Dabek - Zlotorzynska, E., and Schmitt - Kopplin, P.: Structural  
804 characterization of organic aerosol using Fourier transform ion cyclotron resonance mass  
805 spectrometry: aromaticity equivalent approach, *Rapid Commun. Mass Spectrom.*, 28, 2445-2454, 2014.
- 806 Zhang, L., Kuniyoshi, I., Hirai, M., and Shoda, M.: Oxidation of dimethyl sulfide by *Pseudomonas*  
807 *acidovorans* DMR-11 isolated from peat biofilter, *Biotechnol. Lett.*, 13, 223-228, 1991.
- 808 Zhang, M., Gao, W., Yan, J., Wu, Y., Marandino, C. A., Park, K., Chen, L., Lin, Q., Tan, G., and Pan, M.:  
809 An integrated sampler for shipboard underway measurement of dimethyl sulfide in surface seawater  
810 and air, *Atmos. Environ.*, 209, 86-91, 2019.
- 811 Zhang, S. H., Yang, G. P., Zhang, H. H., and Yang, J.: Spatial variation of biogenic sulfur in the south  
812 Yellow Sea and the East China Sea during summer and its contribution to atmospheric sulfate aerosol,  
813 *Sci. Total Environ.*, 488-489, 157-167, 10.1016/j.scitotenv.2014.04.074, 2014a.
- 814 Zhang, Y., Wang, Y., Gray, B. A., Gu, D., Mauldin, L., Cantrell, C., and Bandy, A.: Surface and free  
815 tropospheric sources of methanesulfonic acid over the tropical Pacific Ocean, *Geophys. Res. Lett.*, 41,  
816 5239-5245, 10.1002/2014gl060934, 2014b.
- 817 Zhao, H., Jiang, X., and Du, L.: Contribution of methane sulfonic acid to new particle formation in the  
818 atmosphere, *Chemosphere*, 174, 689-699, 10.1016/j.chemosphere.2017.02.040, 2017.



819 Zhao, Y. and Truhlar, D. G.: The M06 suite of density functionals for main group thermochemistry,  
820 thermochemical kinetics, noncovalent interactions, excited states, and transition elements: two new  
821 functionals and systematic testing of four M06-class functionals and 12 other functionals, *Theor. Chem.*  
822 *Acc.*, 120, 215-241, 2008.

823 Zhou, S., Hwang, B. C. H., Lakey, P. S. J., Zuend, A., Abbatt, J. P. D., and Shiraiwa, M.: Multiphase  
824 reactivity of polycyclic aromatic hydrocarbons is driven by phase separation and diffusion limitations,  
825 *Proc. Natl. Acad. Sci. U. S. A.*, 116, 11658-11663, 10.1073/pnas.1902517116, 2019.

826 Zhu, M., Jiang, B., Li, S., Yu, Q., Yu, X., Zhang, Y., Bi, X., Yu, J., George, C., and Yu, Z.: Organosulfur  
827 Compounds Formed from Heterogeneous Reaction between SO<sub>2</sub> and Particulate-Bound Unsaturated  
828 Fatty Acids in Ambient Air, *Environ. Sci. Technol. Lett.*, 6, 318-322, 2019.

829