



1 **Mixing state of refractory black carbon at different atmospheres in China**

2 Gang Zhao¹, Tianyi Tan¹, Shuya Hu¹, Zhuofei Du¹, Dongjie Shang¹, Zhijun Wu^{1,2},
3 Song Guo^{1,2}, Jing Zheng¹, Wenfei Zhu¹, Mengren Li¹, Limin Zeng¹, Min Hu^{1,2*}

4 1 State Key Joint Laboratory of Environmental Simulation and Pollution Control,
5 International Joint Laboratory for Regional Pollution Control, Ministry of Education,
6 College of Environmental Sciences and Engineering, Peking University, Beijing,
7 100871, China

8 2 Collaborative Innovation Center of Atmospheric Environment and Equipment
9 Technology, Nanjing University of Information Science & Technology, Nanjing,
10 China

11 ***Correspondence author:** Min Hu (minhu@pku.edu.cn)

12 **Abstract**

13 Black carbon (BC) particles exert a significant influence on the earth's climate
14 system. However, large uncertainties remain when estimating the radiative forcing by
15 BC because the corresponding microphysical properties have not been well addressed.
16 Knowledge of the BC mixing states of different aging degree can help better
17 characterise the corresponding environmental and climate effects. In this study, the BC
18 size distributions were studied based on three different field campaigns at an urban site,
19 a suburban site, and a background site in China using a single particle soot photometer
20 (SP2) in tandem with a differential mobility diameter. Measurements from the SP2
21 indicates that the BC particles were composed of either fresh or aged aerosols. The
22 mean number fractions of the fresh BC aerosols were 51%, 67%, and 21% for the



23 urban, suburban, and background sites, respectively. The corresponding mobility
24 diameters of these aged (fresh) BC-containing aerosols were 294 nm (193 nm), 244
25 nm (161 nm), and 257 nm (162 nm). The measured aged (fresh) BC core number
26 median diameters were 115 nm (114 nm), 107 nm (95 nm), and 127 nm (111 nm) for
27 urban, suburban, and background sites, respectively. The corresponding aged (fresh)
28 core mass median diameters were 187 nm (154 nm), 182 nm (146 nm), and 238 nm
29 (163 nm) respectively. The mean diameter of the aged BC-containing aerosols was
30 larger than that of the fresh BC-containing aerosols, while the mean BC core diameter
31 of the aged BC-containing aerosols was smaller than that of the fresh BC-containing
32 aerosols. About 10% of the BC-containing aerosols with the BC core were attached to
33 the other non-BC components, which were mainly generated by coagulation between
34 the BC and non-BC components. The measurement results in our study can help better
35 understand the BC size distributions and mixing status in the different atmospheres in
36 China and can be further used in modeling studies to help constrain the uncertainties
37 of the BC radiative effects.

38 **Introduction**

39 Black carbon (BC) plays an important role in the climate system by absorbing
40 solar radiation (Ramanathan et al., 2008), interacting with the cloud (Roberts et al.,
41 2008), and changing the albedo of the snow (Menon et al., 2002). It is the second most
42 important aerosol component after carbon dioxide, contributing to global warming
43 (Bond et al., 2013). The solar absorption of BC has a significant influence on the
44 development of the boundary layer and then aggravates the air pollution (Ding et al.,
45 2016). The turbulence in the atmospheric boundary layer can be suppressed due to the
46 existence of BC (Wilcox et al., 2016). The BC also plays a remarkable role in driving
47 the formation and trend of regional haze (Zhang et al., 2020).



48 BC is mainly generated by the incomplete combustion of biofuels and fossil fuels
49 (Bond et al., 2006). After emission, the morphology of BC transforms from fractal to
50 spherical and subsequently grows to a fully compact particle with other chemical
51 components coating on it (Peng et al., 2016). During the aging process, the BC optical
52 properties change significantly up to a factor of 3 and then the corresponding
53 magnitude of climate forcing contributed by BC is increased by up to a factor of 2
54 (Zhang et al., 2008; Cappa et al., 2012). Large uncertainties remain in estimating the
55 BC radiative effects due to the large variation in BC microphysical properties, such as
56 size distributions and mixing states during the aging process (Moffet et al., 2016;
57 Matsui et al., 2018; Zhao et al., 2019). Therefore, characterizing the differences in size
58 distributions and mixing states between the fresh and aged BC particles can help better
59 constrain the uncertainties of BC aerosol radiative effects. To our best understanding,
60 few studies have specified the mixing states and size distributions of both the fresh and
61 aged BC aerosols.

62 The BC-containing particles can also be classified into two morphological types:
63 bare BC on the surface of non-BC particles (attached type) and BC embedded within
64 or coated by non-BC components (coated type). With the same amount of non-BC
65 components, the mass absorption cross-sections of BC by the attached type are much
66 smaller than those by the coated type (Moteki et al., 2008; Moteki et al., 2010a;
67 Moteki et al., 2014). Therefore, the impact of BC on climate can be better estimated
68 when accurately identifying the two types of ambient BC-containing particles.
69 Observations are required to constrain the spatial and temporal microphysical
70 properties of the atmospheric BC.

71 The single-particle soot photometer (SP2) is always used to measure the mixing
72 states and size distributions of ambient BC particles. The measured signals from SP2
73 can be used to distinguish the BC-containing aerosols as fresh thinly and aged thickly



74 coated ones. The measured results can also be employed to distinguish the
75 BC-containing particles between attached and coated types, which were described in
76 detail in the methodology part.

77 In this study, the tandem SP2 and differential mobility analyzer (DMA) was
78 employed at an urban site, a suburban site, and a background site in China to
79 investigate the microphysical properties of the BC particles. The size distributions and
80 mixing states of both the fresh and aged BC aerosols at different atmospheres were
81 characterized. We also investigated the corresponding morphology properties of the
82 BC-containing aerosols. The measured microphysical properties provide the basis for
83 future modeling studies of the BC radiative effects of different environment in China.

84 **2 Methodology**

85 **2.1 Measurement sites**

86 The measurements were conducted at three different atmospheric sites in China,
87 namely the urban site of Peking University Urban Atmosphere Environment
88 Monitoring Station (PKU, 39.9°N, 116.1°E, 58m a.s.l) in Beijing between 20 January
89 and 4 February 2016, the suburban site of Changping (CP, 40.3°N, 116.2°E, 70m a.s.l))
90 in Beijing between 15 May and 5 June 2016, and the background site of Lijiang (LJ,
91 27.2°N, 100.2°E, 3410 m a.s.l) in Yunnan Province between 22 March and 4 April
92 2015. The PKU site is located in the northwest of Beijing. This site could
93 characterize the air pollution of the urban Beijing (Hu et al., 2017; Hu et al., 2021).
94 The CP site locates at the northwest of the Beijing urban area, representing a regional
95 atmosphere (Wang et al., 2019b; Zhao et al., 2021). The LJ site represents the
96 background areas, located in the Mountain Yulong, in the Yunan Province of China
97 (Zheng et al., 2017; Shang et al., 2018; Wang et al., 2019a). The aerosol optical depth
98 at the wavelength of 550 nm during the year 2020 indicated that the LJ site was very



99 clean and the PKU and CP sites were more polluted as shown in Fig. S1 in the
100 supplement.

101 **2.2 Instruments**

102 **2.2.1 DMA-SP2 system**

103 As for the SP2, the continuous Nd: YAG laser beam with the wavelength of 1064
104 nm is generated intensively in the instrument chamber. When the BC-containing
105 particles pass through the laser beam, they absorb the radiation and then are heated to
106 around 3500-5000 K. The intensity of the emitted incandescent light from the heated
107 BC particle is then transformed to the BC mass concentration. The scattering signals
108 of the BC particle are recorded to estimate the BC particle mixing state.

109 In this study, the SP2 (Droplet Measurement Technology, Inc., USA) was placed
110 after the DMA (Model 3081, TSI, USA) to measure the size-resolved BC mixing
111 states, and the instrument setup is schematically shown in Fig. S2. The DMA was set
112 to scan the aerosol over the size range between 12.3 and 697 nm every five minutes.
113 The flow rate leading to the SP2 and the condensation particle counter (CPC, Model
114 3776, TSI, USA) were 0.12 and 0.28 L/min, respectively. The sheath flow of the DMA
115 was 4 L/min.

116 The Aquadag was used to calibrate the measured incandescence signal of the SP2
117 using the DMA-SP2 system. The formula from Gysel et al. (2011) was used to convert
118 the mobility diameter into the mass of Aquadag. A correction factor of 0.75 was
119 applied to account for the different response sensitivity of SP2 to Aquadag and
120 ambient BC (Moteki et al., 2010b).



121 In this study, the coating thickness of the BC-containing aerosols was calculated
122 by the difference between the total mobility diameter measured by the DMA and the
123 mass equivalent diameters of the BC core with the assumption that the density of the
124 BC-core is 1.8 g/cm^3 .

125 **2.2.2 Other instruments**

126 The submicron particles (PM_{10}) chemical compositions were measured using a
127 high-resolution time-of-flight aerosol mass spectrometer (AMS; Aerodyne Research
128 Inc., Billerica, MA, USA). The data processing software PIKA (version 1.16) was
129 used for data analysis. The positive matrix factorization (PMF) analysis was conducted
130 for source appointment of the organic aerosols (Ulbrich et al., 2009). More details of
131 the measurement of the aerosol chemical compositions and data processing can be
132 found in Zheng et al. (2017).

133 The mass concentrations of O_3 were measured using UV absorption (model 49i,
134 Thermo Fischer Inc. USA) with a time resolution of 1 minute. The mass
135 concentrations of NO and NO_2 were measured using the chemiluminescence technique
136 ($\text{NO-NO}_2\text{-NO}_x$ Analyzer, Model 42i, Thermo Scientific, USA). The mass
137 concentrations of SO_2 were measured using the ultraviolet fluorescence method (SO_2
138 analyzer, model 43i-TLE, Thermo Scientific, USA). The temperature (T), relative
139 humidity (RH), wind speed (WS), and wind direction (WD) were monitored
140 continuously during these campaigns.

141 **2.3 Methodology**

142 For the BC-containing aerosol, there is a lag between the peak time of the
143 scattering and the incandescence signal (Metcalf et al., 2012). The lag time between
144 the peak scattering signal and the peak incandescence signal can be employed to



145 describe the coating thickness (Schwarz et al., 2006; Moteki et al., 2007) and further
146 used to distinguish the BC-containing aerosols as fresh thinly and aged thickly coated
147 ones.

148 The measured scattering and incandescence signal can also be employed to
149 distinguish the BC-containing particles as attached and coated types (Moteki et al.,
150 2014) by calculating the time-dependent scattering cross-sections of BC-containing
151 particles (Moteki et al., 2007). For the coated type, all of the coating material will
152 evaporate and the scattering cross-sections will decrease to zero after passing through
153 the laser beam, while the scattering cross-section of the attached BC-containing
154 aerosol will not decrease to zero (Moteki et al., 2008). The method adopted by
155 Dahlkötter et al. (2014) was employed here to characterize the morphology of the
156 BC-containing aerosols.

157 **3 Results and discussions**

158 **3.1 Overview of the measurement results at different atmospheres**

159 The time series of the measurement results are shown in Fig. S3, Fig. S4, and Fig.
160 S5 for the PKU, CP, and LJ sites, respectively.

161 As for the PKU site, the wind was mainly from the north and the wind speed was
162 low with a mean value of 2.2 m/s. The ambient atmosphere was very dry with a mean
163 RH of 27.6%, with minimum and maximum values of 5.8% and 72.6%, respectively.
164 The temperature in the winter of Beijing had a mean value of 0.8 °C between -5.9 °C
165 and 9.2 °C. The mean mass concentration of PM_{2.5} was 49.3 ± 55.4 µg/m³. The
166 concentration of SO₂ and NO_x (NO_x=NO + NO₂) had the same trends as PM_{2.5}, with
167 mean values of 16.3 ± 11.9 ppb and 68.2 ± 63.4 ppb, respectively. The O₃



168 concentration is anti-correlated with $PM_{2.5}$. The measurement site experienced four
169 main pollution periods between 20, January and 4, February, with each period lasting
170 2~4 days. The four pollution periods happened from 21 January to 24 January, from
171 24 January to 26 January, from 28 January to 29 January, and from 31 January to 3
172 February. The $PM_{2.5}$ peaked at the first pollution period, with $272.8 \mu\text{g}/\text{m}^3$. For each
173 period, the high RH and low wind speed favored the development of pollution. At the
174 end of each pollution period, the $PM_{2.5}$ dropped dramatically with the increment of the
175 wind speed and the change of the wind direction. The environment in the winter of
176 Beijing was polluted, which was highly influenced by both primary particle emissions
177 and secondary formation influenced by the meteorology conditions.

178 For the suburban site CP, the wind showed obvious diurnal cycles with high-speed
179 west wind during the day and low-speed east wind during the night. The mean wind
180 speed was $2.4 \pm 1.6 \text{ m/s}$. The RH during the campaign was $38.8 \pm 16.0\%$, with a
181 maximum value of 80.5%. The temperature during the campaign was $21.8 \pm 5.2 \text{ }^\circ\text{C}$
182 with a maximum value of $33.2 \text{ }^\circ\text{C}$. As for the NO_x , the mean concentration was $21.4 \pm$
183 17.7 ppb . The concentration of NO_x experienced high value during the early morning,
184 and fluctuated dramatically, which is highly related to the anthropogenic activities.
185 The mean concentration of SO_2 was $2.89 \pm 1.10 \text{ ppb}$. The measured SO_2 concentration
186 values during the day were higher than those at night. There was no obvious diurnal
187 cycle for the SO_2 concentration, and dramatic fluctuation was not observed, which
188 indicates that the SO_2 was mainly from transportation. The measured mean O_3
189 concentration was $54.5 \pm 38.8 \text{ ppb}$. The mean $PM_{2.5}$ concentration was 22.6 ± 16.8
190 $\mu\text{g}/\text{m}^3$, with a maximum value of $71.8 \mu\text{g}/\text{m}^3$.



191 As for the background LJ site, The mean value of the wind speed, RH, and T were
192 3.13 m/s, 50.23%, and 6.5 °C, respectively. The mean PM_{2.5} mass concentration was
193 $6.2 \pm 5.7 \mu\text{g}/\text{m}^3$. The mean NO_x and SO₂ concentrations were 0.05 ppb and 0.97 ppb
194 respectively.

195 The characteristics of the measurement sites are summarized and shown in Fig. 1.
196 The differences in the temperature and RH among these sites were mainly resulted
197 from the that the measurements were conducted in different seasons. The
198 concentrations of SO₂, NO_x, and PM_{2.5} indicated that the urban site PKU was most
199 polluted. The suburban site CP was slightly polluted and the background LJ was the
200 cleanest.

201 **3.2 Mixing states of the fresh and aged BC-containing aerosols**

202 The measured lag time probability distributions for the PKU, CP and LJ sites are
203 shown in Fig. 2 (a), (b), and (c), respectively. The lag time had two modes for each
204 measurement site. In this study, the BC-containing aerosols with a lag time larger than
205 1.4 μs were classified as aged thickly coated particles. The other BC-containing
206 aerosols were classified as fresh thinly coated particles. Our critical lag time of 1.4 μs
207 is smaller than the previous studies that distinguished the BC-containing aerosols
208 between fresh BC and aged BC with a lag time of 2 μs (Moteki et al., 2007; Metcalf et
209 al., 2012) or 1.8 μs (Metcalf et al., 2012), which was determined by the internal setup
210 up of the SP2.

211 For each type of BC-containing aerosols, we calculated the coating thickness
212 probabilities and the results are shown in Fig. (d), (e) and (f) for the PKU, CP, and LJ
213 sites, respectively. Results showed that the BC-containing aerosols were mainly



214 composed of thickly coated aged BC aerosols and thinly coated fresh BC aerosols. The
215 coating thickness for the fresh BC-containing aerosol was smaller than that of the aged
216 BC-containing aerosols. However, the coating thickness of the aged BC-containing
217 aerosols spread wider than that of the fresh ones.

218 The number fractions of the aged BC-containing aerosols were significantly
219 different for different atmospheres as shown in Fig. 2 (g), (h), and (i). At the polluted
220 urban site, the number concentration of the aged BC-containing aerosols was
221 comparable to that of the fresh BC-containing aerosols with the number fractions of
222 51% and 49% for the fresh and aged BC particles, respectively. The number fraction
223 of the aged BC aerosols at the CP site was 67 %. However, the BC-containing aerosols
224 at the background LJ site were dominated by aged ones with a number fraction of
225 79%.

226 The difference in the number fraction of the aged BC particles was synthetically
227 influenced by the ambient pollution levels and the sources of the BC aerosols. The
228 suburban site CP had the largest number fraction of the fresh BC particles. The CP site
229 is not far from the urban, and thus the fresh BC particles from the traffic contribute a
230 large amount of the total ones. The urban site PKU had a larger number fraction of the
231 aged BC than that of the CP site. This might be resulted from that the PKU site being
232 more polluted than the CP site and then the aging processing at the PKU site was
233 faster than that at the CP site. The LJ site is far from the traffic sources. The measured
234 BC particles at the LJ site were mainly from long-range transportation and
235 experienced a long time of aging process than that at the CP and PKU sites. Therefore,
236 the BC-containing aerosols were dominated by the aged ones at the LJ sites.

237 We compared the number fraction of the aged BC at different measurement sites
238 from literature (Shiraiwa et al., 2007; Schwarz et al., 2008a; Schwarz et al., 2008b;



239 Subramanian et al., 2010; Huang et al., 2012; McMeeking et al., 2012; Metcalf et al.,
240 2012; Holder et al., 2014; Wang et al., 2014; Ueda et al., 2016; Wang et al., 2016;
241 Wang et al., 2017a; Wang et al., 2017b; Wang et al., 2017c; Wu et al., 2017;
242 Krasowsky et al., 2018; Saha et al., 2018) and the results are shown in Fig. 3. The
243 number fraction values were divided into three different kinds of groups, namely the
244 roadside, urban or suburban, and background. Results from Fig. 3 show that the
245 number fractions at the roadside tend to be the lowest. These sites were close to the
246 traffic sources and the measured BC-containing aerosols were mainly from the traffic.
247 The left part of the green circles correspond to the relatively clean urban or suburban
248 sites with the number fractions of the aged BC around 30%. However, the number
249 fractions of the relative polluted urban or suburban sites had a larger number fraction
250 of the aged BC around 50%. The number fractions of the aged BC at the background
251 sites were the largest. Therefore, the number fractions of the aged BC-containing
252 aerosols were synthetically influenced by the distance from the primary source and the
253 pollution levels of the ambient atmosphere. The number fraction of the aged
254 BC-containing aerosols increased with the distance from the primary emission sources
255 and the pollution levels. Our results were consistent with the aerial measurement by
256 Metcalf et al. (2012), who found that the number fraction of the aged BC was
257 29%–41% at the top of the Los Angeles city and 47%–54% for the out plume of this
258 city.

259 For a better understanding of the source of the fresh and aged BC, we compared
260 the number concentrations of the BC-containing aerosols with the source
261 apportionment results from the AMS data. Among the PMF results, the factor of
262 hydrocarbon-like organic aerosol (HOA) is mainly composed of long-chain
263 hydrocarbon, and oxygenated organic aerosol (OOA) is mainly from the secondary
264 formation. HOA is mainly from the diesel exhaust, gasoline exhaust, and lubricating



265 oil emission. From Fig. 4(a), the number concentration of the fresh BC and mass
266 concentration of HOA showed good consistency, with R^2 equaling 0.69 as shown in
267 Fig. S6, which further proved the evidence that the fresh BC-containing aerosols were
268 from the traffic sources. The time series of the aged BC and OOA showed good
269 consistency as shown in Fig. 4 (b), with R^2 equaling 0.87. Therefore, the aging
270 processing of the ambient BC was accompanied by the ambient OA. The mass
271 concentration of OOA and number concentration of aged BC can be used as good
272 indicators for each other.

273 **3.3 Size distributions of the fresh and aged BC-containing aerosols**

274 The size distributions of the BC-containing aerosols exert significant influence on
275 their corresponding radiative effects (Matsui et al., 2018; Zhao et al., 2019). We
276 calculated the number size distribution (NSD) of BC-containing aerosols for the fresh
277 and aged ones at different sites, and the results are shown in Fig. 5. It should be noted
278 that the D_p in Fig. 5 corresponds to the mobility diameter from the DMA. The
279 BC-containing aerosol NSD was further fit using the log-normal distribution.

280 As for the fresh BC-containing aerosols, the geometric mean diameters (D_m) were
281 193, 161, and 162 nm for the PKU, CP, and LJ sites, respectively. The geometric
282 standard deviations (GSD) of the BC-containing aerosol NSD were 1.50, 1.63, 1.91
283 for the PKU, CP, and LJ sites, respectively. The GSD to some extent reflects the
284 diversity of the BC sources. The LJ site had the largest GSD, which indicated multiple
285 sources of fresh BC-containing aerosols. The LJ site was highly influenced by
286 atmospheric transportation, due to the high altitude of this location (Zheng et al., 2017;
287 Tan et al., 2021). Therefore, the fresh BC-containing aerosols could be originated from
288 different orientations. As for the urban site PKU, the fresh BC aerosols were mainly
289 from urban lifestyle emissions. Therefore, the fresh BC aerosols at the PKU site had



290 the lowest value of the GSD. However, the fresh BC aerosols at the suburban site CP
291 were influenced synthetically by urban lifestyle sources and some other sources from
292 suburban, and thus had a larger value of GSD than that of PKU.

293 As for the aged BC, it is obvious that they had larger diameters than those of the
294 fresh BC due to the coating of other non-BC components. The D_m values of the aged
295 BC were 294, 244, and 257 nm for the PKU, CP, and LJ sites, respectively. The
296 corresponding GSD values were 1.37, 1.41, and 1.46.

297 Based on the above results, the D_m values of the aged BC aerosols were larger
298 than that of the fresh BC aerosols by 52%, 52%, and 59% for the PKU, CP, and LJ
299 sites, respectively. The GSD values of the aged BC were consistent with that of the
300 fresh BC with the lowest value at the PKU site and highest value at the LJ site, which
301 is consistent with the diversity of the sources of BC-containing aerosols. For each site,
302 the GSD values of the aged BC aerosols were smaller than that of the fresh ones. The
303 GSD of BC-containing aerosols tend to be smaller during the aging processing
304 because the increment of the diameter should decrease with the diameter.

305 **3.4 Size distribution of the fresh and aged BC core**

306 The number and mass concentrations of the BC core under different mass
307 equivalent diameters were calculated with the assumption that the BC core has an
308 effective density of 1.8 g/cm^3 and the results are shown in Fig. 6 and Table 1. From
309 Fig. 6, the geometric mean mass equivalent diameter (D_{me}) of the fresh BC particles
310 were 115, 107, and 127 nm, for the PKU, CP, and LJ sites respectively. The
311 corresponding GSD values are 1.58 1.53 and 1.68, respectively. The D_{me} for the aged
312 BC particles were 114, 95, and 111 nm for the PKU, CP, and LJ sites respectively and
313 the corresponding GSD values were 1.40, 1.45, and 1.43, respectively. Both the GSD



314 and the D_{me} of the aged BC were smaller than that of the fresh BC. This might be
315 resulted from the fact that the small BC particles have a longer life than that of the
316 large BC particles.

317 **3.5 Morphology of the BC-containing aerosols**

318 The time series of the number fractions of the attached BC-containing aerosols to
319 the total BC-containing aerosols ($f_{attached}$) are shown in Fig. 7. From Fig. 7, the $f_{attached}$
320 ranged between 0 and 0.21 with a mean value of $7.2 \pm 3.7\%$, $11.0 \pm 3.7\%$, and $10.1 \pm$
321 4.1% . Moteki et al. (2014) found that the $f_{attached}$ was generally less than 0.1 in Tokyo.
322 The $f_{attached}$ ranged between 3% and 16% in suburban London (Liu et al., 2015). A
323 mean value of 12% was found for biomass burning particles using electron
324 microscopy (China et al., 2013). Our measurement results were consistent with the
325 previous studies. The $f_{attached}$ tend to increase with the $PM_{2.5}$ for different sites, which
326 may indicate that the attached BC-containing aerosols were generated from the
327 coagulation of BC and non-BC aerosols.

328 We calculated the $f_{attached}$ under different aerosol diameters and the results are
329 shown in Fig. 8. There were few attached BC-containing aerosols when the diameter
330 was smaller than 250 nm with $f_{attached}$ lowing than 2%. The $f_{attached}$ increased with the
331 diameter for all of the measurement sites. It could reach 30% for the LJ sites. Based on
332 the results from the electron microscopy, the BC volume fractions are smaller than
333 those of the non-BC volume fractions in the attached BC aerosols (Moteki et al., 2014).
334 Our results further indicate that the attached BC aerosols were formed from
335 coagulation, as the coagulation efficiency of the two particles increased with the
336 difference between their sizes (Kim et al., 2016; Cai et al., 2017; Mahfouz et al.,
337 2021).



338 The f_{attached} under different aerosol number concentrations (N) and different ratios
339 of the BC-free aerosol number concentrations to the BC-containing aerosol number
340 concentrations are shown in Fig. 9. Results showed that the f_{attached} increased with the
341 above two factors. The results were consistent with the fact that the coagulation
342 between BC and non-BC components is more likely to happen with the increment of
343 the BC-free aerosol number concentrations. Based on the analysis above, we
344 concluded that the attached BC- containing aerosols are mainly formed through
345 coagulation.

346 **4 Conclusions**

347 In this study, the BC microphysical properties were studied based on field
348 measurement using the DMA-SP2 system at the urban site PKU, suburban site CP and
349 a background site LJ.

350 The BC-containing aerosols were sorted as aged thickly coated BC and fresh thinly
351 coated BC based on the lag time between the peak position of the light scattering
352 signals and the incandescence signals. The number fractions of the aged
353 BC-containing aerosols were 49%, 33%, and 79% for the PKU, CP, and LJ sites
354 respectively. The mass concentrations of the fresh BC-containing aerosols showed
355 good consistency with that of HOA, which indicated that the fresh BC-containing
356 aerosols were mainly generated from the emission of vehicles. The aged
357 BC-containing aerosols are highly correlated with the OOA.

358 The geometric diameter of the fresh BC-containing aerosols ranged between 160
359 nm and 200 nm, while the corresponding range was 240~300 nm for the aged
360 BC-containing aerosols. The GSD of the BC-containing aerosols decreased during the
361 aging process. The corresponding mobility diameters of these aged (fresh)
362 BC-containing aerosols were 294 (193), 244 (161), and 257 (162) nm. The measured



363 aged (fresh) BC core number median diameters were 115 (114), 107 (95), and 127
364 (111) nm for the urban, suburban, and background sites, respectively. The
365 corresponding aged (fresh) core mass median diameters were 187 (154), 182 (146),
366 and 238 (163) nm respectively. The mean diameter of the aged BC-containing aerosols
367 was larger than that of the fresh BC-containing aerosols, while the mean BC core
368 diameter of the aged BC-containing aerosols was smaller than that of the fresh
369 BC-containing aerosols.

370 The BC-containing aerosols were sorted as the coated type when the scattering
371 cross-section decreased to zeros, while the BC-containing aerosols were sorted as the
372 attached type when the scattering cross-section was still larger than a critical point
373 after passing through the SP2 laser beam. There are about 10% of the BC-containing
374 aerosols with the BC core attached to the other non-BC components. We concluded
375 that the attached BC-containing aerosols were mainly generated by coagulation
376 between the BC and non-BC components.

377 **Data availability.** The data is available at <https://doi.org/10.5281/zenodo.5816310>.

378 **Author contributions.** **Gang Zhao:** Conceptualization, Writing - Original Draft,
379 Visualization, Software, **Tianyi Tan:** Data Curation, Conceptualization, Visualization,
380 **Shuya Hu:** Data Curation, Conceptualization, **Zhuofei Du:** Data Curation, **Dongjie**
381 **Shang:** Data Curation, **Zhijun Wu:** Data Curation, Conceptualization, **Song Guo:**
382 Data Curation, Conceptualization, **Jing Zheng:** Data Curation, Conceptualization,
383 **Wenfei Zhu:** Data Curation, Conceptualization, **Mengren Li:** Data Curation,
384 Conceptualization, **Limin Zeng:** Data Curation, Conceptualization, **Min Hu:**
385 Resources, Supervision, Data Curation, Conceptualization, Revision.

386 **Competing interests.** The authors declare that they have no conflict of interest.



387 **Acknowledgments.** This work is supported by the China Postdoctoral Science
388 Foundation (2021M700192) and National Natural Science Foundation of China
389 (91844301).

390

391 **References**

- 392 Bond, T. C., and Bergstrom, R. W.: Light Absorption by Carbonaceous Particles: An
393 Investigative Review, *Aerosol Sci. Technol.*, 40, 27-67, 10.1080/02786820500421521,
394 2006.
- 395 Bond, T. C., Doherty, S. J., Fahey, D. W., Forster, P. M., Berntsen, T., DeAngelo, B.
396 J., Flanner, M. G., Ghan, S., Karcher, B., Koch, D., Kinne, S., Kondo, Y., Quinn, P. K.,
397 Sarofim, M. C., Schultz, M. G., Schulz, M., Venkataraman, C., Zhang, H., Zhang, S.,
398 Bellouin, N., Guttikunda, S. K., Hopke, P. K., Jacobson, M. Z., Kaiser, J. W., Klimont,
399 Z., Lohmann, U., Schwarz, J. P., Shindell, D., Storelvmo, T., Warren, S. G., and
400 Zender, C. S.: Bounding the role of black carbon in the climate system: A scientific
401 assessment, *J Geophys Res-Atmos*, 118, 5380-5552, 10.1002/jgrd.50171, 2013.
- 402 Cai, R., and Jiang, J.: A new balance formula to estimate new particle formation rate:
403 reevaluating the effect of coagulation scavenging, *Atmospheric Chemistry and Physics*,
404 17, 12659-12675, 10.5194/acp-17-12659-2017, 2017.
- 405 Cappa, C. D., Onasch, T. B., Massoli, P., Worsnop, D. R., Bates, T. S., Cross, E. S.,
406 Davidovits, P., Hakala, J., Hayden, K. L., Jobson, B. T., Kolesar, K. R., Lack, D. A.,
407 Lerner, B. M., Li, S. M., Mellon, D., Nuaaman, I., Olfert, J. S., Petaja, T., Quinn, P. K.,
408 Song, C., Subramanian, R., Williams, E. J., and Zaveri, R. A.: Radiative Absorption
409 Enhancements Due to the Mixing State of Atmospheric Black Carbon, *Science*, 337,
410 1078-1081, 10.1126/science.1223447, 2012.



- 411 China, S., Mazzoleni, C., Gorkowski, K., Aiken, A. C., and Dubey, M. K.:
412 Morphology and mixing state of individual freshly emitted wildfire carbonaceous
413 particles, *Nature communications*, 4, 2122, 10.1038/ncomms3122, 2013.
- 414 Dahlkötter, F., Gysel, M., Sauer, D., Minikin, A., Baumann, R., Seifert, P., Ansmann,
415 A., Fromm, M., Voigt, C., and Weinzierl, B.: The Pagami Creek smoke plume after
416 long-range transport to the upper troposphere over Europe – aerosol properties
417 and black carbon mixing state, *Atmospheric Chemistry and Physics*, 14, 6111-6137,
418 10.5194/acp-14-6111-2014, 2014.
- 419 Ding, A., Huang, X., Nie, W., Sun, J., Kerminen, V.-M., Petäjä, T., Su, H., Cheng, Y.,
420 Yang, X.-Q., Wang, M., Chi, X., Wang, J., Virkkula, A., Guo, W., Yuan, J., Wang, S.
421 R., Zhang, R., Wu, Y., Song, Y. C., Zhu, T., Zilitinkevich, S., Kulmala, M., and Fu, C.:
422 Enhanced haze pollution by black carbon in megacities in China, *Geophys. Res. Lett.*,
423 43, 2873-2879, 2016.
- 424 Gysel, M., Laborde, M., Olfert, J. S., Subramanian, R., and Gröhn, A. J.: Effective
425 density of Aquadag and fullerene soot black carbon reference materials used for SP2
426 calibration, *Atmospheric Measurement Techniques*, 4, 2851-2858,
427 10.5194/amt-4-2851-2011, 2011.
- 428 Holder, A. L., Hagler, G. S. W., Yelverton, T. L. B., and Hays, M. D.: On-road black
429 carbon instrument intercomparison and aerosol characteristics by driving environment,
430 *Atmospheric Environment*, 88, 183-191,
431 <https://doi.org/10.1016/j.atmosenv.2014.01.021>, 2014.
- 432 Hu, S., Zhao, G., Tan, T., Li, C., Zong, T., Xu, N., Zhu, W., and Hu, M.: Current
433 challenges of improving visibility due to increasing nitrate fraction in PM_{2.5} during
434 the haze days in Beijing, China, *Environmental Pollution*, 290, 118032,
435 10.1016/j.envpol.2021.118032, 2021.
- 436 Hu, W., Hu, M., Hu, W.-W., Zheng, J., Chen, C., Wu, Y., and Guo, S.: Seasonal
437 variations in high time-resolved chemical compositions, sources, and evolution of



- 438 atmospheric submicron aerosols in the megacity Beijing, *Atmospheric Chemistry and*
439 *Physics*, 17, 9979-10000, 10.5194/acp-17-9979-2017, 2017.
- 440 Huang, X.-F., Sun, T.-L., Zeng, L.-W., Yu, G.-H., and Luan, S.-J.: Black carbon
441 aerosol characterization in a coastal city in South China using a single particle soot
442 photometer, *Atmospheric Environment*, 51, 21-28,
443 <https://doi.org/10.1016/j.atmosenv.2012.01.056>, 2012.
- 444 Kim, Y.-h., Yiacoumi, S., Nenes, A., and Tsouris, C.: Charging and coagulation of
445 radioactive and nonradioactive particles in the atmosphere, *Atmospheric Chemistry*
446 *and Physics*, 16, 3449-3462, 10.5194/acp-16-3449-2016, 2016.
- 447 Krasowsky, T. S., McMeeking, G. R., Sioutas, C., and Ban-Weiss, G.: Characterizing
448 the evolution of physical properties and mixing state of black carbon particles: from
449 near a major highway to the broader urban plume in Los Angeles, *Atmos. Chem. Phys.*,
450 18, 11991-12010, 10.5194/acp-18-11991-2018, 2018.
- 451 Liu, S., Aiken, A. C., Gorkowski, K., Dubey, M. K., Cappa, C. D., Williams, L. R.,
452 Herndon, S. C., Massoli, P., Fortner, E. C., Chhabra, P. S., Brooks, W. A., Onasch, T.
453 B., Jayne, J. T., Worsnop, D. R., China, S., Sharma, N., Mazzoleni, C., Xu, L., Ng, N.
454 L., Liu, D., Allan, J. D., Lee, J. D., Fleming, Z. L., Mohr, C., Zotter, P., Szidat, S., and
455 Prevot, A. S.: Enhanced light absorption by mixed source black and brown carbon
456 particles in UK winter, *Nature communications*, 6, 8435, 10.1038/ncomms9435, 2015.
- 457 Mahfouz, N. G. A., and Donahue, N. M.: Technical note: The enhancement limit of
458 coagulation scavenging of small charged particles, *Atmos. Chem. Phys.*, 21,
459 3827-3832, 10.5194/acp-21-3827-2021, 2021.
- 460 Matsui, H., Hamilton, D. S., and Mahowald, N. M.: Black carbon radiative effects
461 highly sensitive to emitted particle size when resolving mixing-state diversity, *Nature*
462 *communications*, 9, 3446, 10.1038/s41467-018-05635-1, 2018.
- 463 McMeeking, G. R., Bart, M., Chazette, P., Haywood, J. M., Hopkins, J. R., McQuaid,
464 J. B., Morgan, W. T., Raut, J. C., Ryder, C. L., Savage, N., Turnbull, K., and Coe, H.:



465 Airborne measurements of trace gases and aerosols over the London metropolitan
466 region, *Atmos. Chem. Phys.*, 12, 5163-5187, 10.5194/acp-12-5163-2012, 2012.

467 Menon, S., Hansen, J., Nazarenko, L., and Luo, Y.: Climate effects of black carbon
468 aerosols in China and India, *Science*, 297, 2250-2253, 10.1126/science.1075159,
469 2002.

470 Metcalf, A. R., Craven, J. S., Ensberg, J. J., Brioude, J., Angevine, W., Sorooshian, A.,
471 Duong, H. T., Jonsson, H. H., Flagan, R. C., and Seinfeld, J. H.: Black carbon aerosol
472 over the Los Angeles Basin during CalNex, *Journal of Geophysical Research:*
473 *Atmospheres*, 117, <https://doi.org/10.1029/2011JD017255>, 2012.

474 Moffet, R. C., and Brien, R. E., Alpert, P. A., Kelly, S. T., Pham, D. Q., Gilles,
475 M. K., Knopf, D. A., and Laskin, A.: Morphology and mixing of black carbon
476 particles collected in central California during the CARES field study, *Atmospheric*
477 *Chemistry and Physics*, 16, 14515-14525, 10.5194/acp-16-14515-2016, 2016.

478 Moteki, N., and Kondo, Y.: Effects of Mixing State on Black Carbon Measurements
479 by Laser-Induced Incandescence, *Aerosol Sci. Technol.*, 41, 398-417,
480 10.1080/02786820701199728, 2007.

481 Moteki, N., and Kondo, Y.: Method to measure time-dependent scattering cross
482 sections of particles evaporating in a laser beam, *Journal of Aerosol Science*, 39,
483 348-364, 10.1016/j.jaerosci.2007.12.002, 2008.

484 Moteki, N., and Kondo, Y.: Dependence of Laser-Induced Incandescence on Physical
485 Properties of Black Carbon Aerosols: Measurements and Theoretical Interpretation,
486 *Aerosol Sci. Technol.*, 44, 663-675, Pii 924375405
487 10.1080/02786826.2010.484450, 2010a.

488 Moteki, N., Kondo, Y., and Nakamura, S.-i.: Method to measure refractive indices of
489 small nonspherical particles: Application to black carbon particles, *Journal of Aerosol*
490 *Science*, 41, 513-521, <https://doi.org/10.1016/j.jaerosci.2010.02.013>, 2010b.



491 Moteki, N., Kondo, Y., and Adachi, K.: Identification by single-particle soot
492 photometer of black carbon particles attached to other particles: Laboratory
493 experiments and ground observations in Tokyo, *Journal of Geophysical Research:*
494 *Atmospheres*, 119, 1031-1043, <https://doi.org/10.1002/2013JD020655>, 2014.

495 Peng, J., Hu, M., Guo, S., Du, Z., Zheng, J., Shang, D., Levy Zamora, M., Zeng, L.,
496 Shao, M., Wu, Y.-S., Zheng, J., Wang, Y., Glen, C. R., Collins, D. R., Molina, M. J.,
497 and Zhang, R.: Markedly enhanced absorption and direct radiative forcing of black
498 carbon under polluted urban environments, *Proceedings of the National Academy of*
499 *Sciences*, 201602310, 10.1073/pnas.1602310113, 2016.

500 Ramanathan, V., and Carmichael, G.: Global and regional climate changes due to
501 black carbon, *Nature Geoscience*, 1, 221-227, 10.1038/ngeo156, 2008.

502 Roberts, G. C., Ramana, M. V., Corrigan, C., Kim, D., and Ramanathan, V.:
503 Simultaneous observations of aerosol-cloud-albedo interactions with three stacked
504 unmanned aerial vehicles, *Proceedings of the National Academy of Sciences of the*
505 *United States of America*, 105, 7370-7375, 10.1073/pnas.0710308105, 2008.

506 Saha, P. K., Khlystov, A., and Grieshop, A. P.: Downwind evolution of the volatility
507 and mixing state of near-road aerosols near a US interstate highway, *Atmos. Chem.*
508 *Phys.*, 18, 2139-2154, 10.5194/acp-18-2139-2018, 2018.

509 Schwarz, J. P., Gao, R. S., Fahey, D. W., Thomson, D. S., Watts, L. A., Wilson, J. C.,
510 Reeves, J. M., Darbeheshti, M., Baumgardner, D. G., Kok, G. L., Chung, S. H., Schulz,
511 M., Hendricks, J., Lauer, A., Kärcher, B., Slowik, J. G., Rosenlof, K. H., Thompson, T.
512 L., Langford, A. O., Loewenstein, M., and Aikin, K. C.: Single-particle measurements
513 of midlatitude black carbon and light-scattering aerosols from the boundary layer to
514 the lower stratosphere, *Journal of Geophysical Research*, 111, 10.1029/2006jd007076,
515 2006.

516 Schwarz, J. P., Gao, R. S., Spackman, J. R., Watts, L. A., Thomson, D. S., Fahey, D.
517 W., Ryerson, T. B., Peischl, J., Holloway, J. S., Trainer, M., Frost, G. J., Baynard, T.,



- 518 Lack, D. A., de Gouw, J. A., Warneke, C., and Del Negro, L. A.: Measurement of the
519 mixing state, mass, and optical size of individual black carbon particles in urban and
520 biomass burning emissions, *Geophys. Res. Lett.*, 35,
521 <https://doi.org/10.1029/2008GL033968>, 2008a.
- 522 Schwarz, J. P., Spackman, J. R., Fahey, D. W., Gao, R. S., Lohmann, U., Stier, P.,
523 Watts, L. A., Thomson, D. S., Lack, D. A., Pfister, L., Mahoney, M. J., Baumgardner,
524 D., Wilson, J. C., and Reeves, J. M.: Coatings and their enhancement of black carbon
525 light absorption in the tropical atmosphere, *J Geophys Res-Atmos*, 113,
526 10.1029/2007jd009042, 2008b.
- 527 Shang, D., Hu, M., Zheng, J., Qin, Y., Du, Z., Li, M., Fang, J., Peng, J., Wu, Y., Lu, S.,
528 and Guo, S.: Particle number size distribution and new particle formation under the
529 influence of biomass burning at a high altitude background site at Mt. Yulong
530 (3410 m), China, *Atmospheric Chemistry and Physics*, 18, 15687-15703,
531 10.5194/acp-18-15687-2018, 2018.
- 532 Shiraiwa, M., Kondo, Y., Moteki, N., Takegawa, N., Miyazaki, Y., and Blake, D. R.:
533 Evolution of mixing state of black carbon in polluted air from Tokyo, *Geophys. Res.*
534 *Lett.*, 34, <https://doi.org/10.1029/2007GL029819>, 2007.
- 535 Subramanian, R., Kok, G. L., Baumgardner, D., Clarke, A., Shinozuka, Y., Campos, T.
536 L., Heizer, C. G., Stephens, B. B., de Foy, B., Voss, P. B., and Zaveri, R. A.: Black
537 carbon over Mexico: the effect of atmospheric transport on mixing state, mass
538 absorption cross-section, and BC/CO ratios, *Atmos. Chem. Phys.*, 10, 219-237,
539 10.5194/acp-10-219-2010, 2010.
- 540 Tan, T., Hu, M., Du, Z., Zhao, G., Shang, D., Zheng, J., Qin, Y., Li, M., Wu, Y., Zeng,
541 L., Guo, S., and Wu, Z.: Measurement report: Strong light absorption induced by aged
542 biomass burning black carbon over the southeastern Tibetan Plateau in pre-monsoon
543 season, *Atmospheric Chemistry and Physics*, 21, 8499-8510,
544 10.5194/acp-21-8499-2021, 2021.



- 545 Ueda, S., Nakayama, T., Taketani, F., Adachi, K., Matsuki, A., Iwamoto, Y., Sadanaga,
546 Y., and Matsumi, Y.: Light absorption and morphological properties of
547 soot-containing aerosols observed at an East Asian outflow site, Noto Peninsula, Japan,
548 *Atmos. Chem. Phys.*, 16, 2525-2541, 10.5194/acp-16-2525-2016, 2016.
- 549 Ulbrich, I. M., Canagaratna, M. R., Zhang, Q., Worsnop, D. R., and Jimenez, J. L.:
550 Interpretation of organic components from Positive Matrix Factorization of aerosol
551 mass spectrometric data, *Atmos. Chem. Phys.*, 9, 2891-2918,
552 10.5194/acp-9-2891-2009, 2009.
- 553 Wang, J., Zhang, Q., Chen, M., Collier, S., Zhou, S., Ge, X., Xu, J., Shi, J., Xie, C.,
554 Hu, J., Ge, S., Sun, Y., and Coe, H.: First Chemical Characterization of Refractory
555 Black Carbon Aerosols and Associated Coatings over the Tibetan Plateau (4730 m
556 a.s.l), *Environmental Science & Technology*, 51, 14072-14082,
557 10.1021/acs.est.7b03973, 2017a.
- 558 Wang, Q., Huang, R. J., Cao, J., Han, Y., Wang, G., Li, G., Wang, Y., Dai, W., Zhang,
559 R., and Zhou, Y.: Mixing State of Black Carbon Aerosol in a Heavily Polluted Urban
560 Area of China: Implications for Light Absorption Enhancement, *Aerosol Sci. Technol.*,
561 48, 689-697, 10.1080/02786826.2014.917758, 2014.
- 562 Wang, Q., Huang, R.-J., Zhao, Z., Zhang, N., Wang, Y., Ni, H., Tie, X., Han, Y.,
563 Zhuang, M., Wang, M., Zhang, J., Zhang, X., Dusek, U., and Cao, J.: Size distribution
564 and mixing state of refractory black carbon aerosol from a coastal city in South China,
565 *Atmospheric Research*, 181, 163-171, <https://doi.org/10.1016/j.atmosres.2016.06.022>,
566 2016.
- 567 Wang, Q., Huang, R., Zhao, Z., Cao, J., Ni, H., Tie, X., Zhu, C., Shen, Z., Wang, M.,
568 Dai, W., Han, Y., Zhang, N., and Prevot, A. S. H.: Effects of photochemical oxidation
569 on the mixing state and light absorption of black carbon in the urban atmosphere of
570 China, *Environmental Research Letters*, 12, 10.1088/1748-9326/aa64ea, 2017b.



- 571 Wang, Y., Liu, F., He, C., Bi, L., Cheng, T., Wang, Z., Zhang, H., Zhang, X., Shi, Z.,
572 and Li, W.: Fractal Dimensions and Mixing Structures of Soot Particles during
573 Atmospheric Processing, *Environmental Science & Technology Letters*, 4, 487-493,
574 10.1021/acs.estlett.7b00418, 2017c.
- 575 Wang, Y., Hu, M., Lin, P., Tan, T., Li, M., Xu, N., Zheng, J., Du, Z., Qin, Y., Wu, Y.,
576 Lu, S., Song, Y., Wu, Z., Guo, S., Zeng, L., Huang, X., and He, L.: Enhancement in
577 Particulate Organic Nitrogen and Light Absorption of Humic-Like Substances over
578 Tibetan Plateau Due to Long-Range Transported Biomass Burning Emissions, *Environ
579 Sci Technol*, 53, 14222-14232, 10.1021/acs.est.9b06152, 2019a.
- 580 Wang, Y., Hu, M., Wang, Y., Zheng, J., Shang, D., Yang, Y., Liu, Y., Li, X., Tang, R.,
581 Zhu, W., Du, Z., Wu, Y., Guo, S., Wu, Z., Lou, S., Hallquist, M., and Yu, J. Z.: The
582 formation of nitro-aromatic compounds under high NO_x and anthropogenic VOC
583 conditions in urban Beijing, China, *Atmospheric Chemistry and Physics*, 19,
584 7649-7665, 10.5194/acp-19-7649-2019, 2019b.
- 585 Wilcox, E. M., Thomas, R. M., Praveen, P. S., Pistone, K., Bender, F. A. M., and
586 Ramanathan, V.: Black carbon solar absorption suppresses turbulence in the
587 atmospheric boundary layer, *Proceedings of the National Academy of Sciences*, 113,
588 11794-11799, 10.1073/pnas.1525746113, 2016.
- 589 Wu, Y., Wang, X., Tao, J., Huang, R., Tian, P., Cao, J., Zhang, L., Ho, K. F., Han, Z.,
590 and Zhang, R.: Size distribution and source of black carbon aerosol in urban Beijing
591 during winter haze episodes, *Atmos. Chem. Phys.*, 17, 7965-7975,
592 10.5194/acp-17-7965-2017, 2017.
- 593 Zhang, F., Wang, Y., Peng, J., Chen, L., Sun, Y., Duan, L., Ge, X., Li, Y., Zhao, J.,
594 Liu, C., Zhang, X., Zhang, G., Pan, Y., Wang, Y., Zhang, A. L., Ji, Y., Wang, G., Hu,
595 M., Molina, M. J., and Zhang, R.: An unexpected catalyst dominates formation and
596 radiative forcing of regional haze, *Proceedings of the National Academy of Sciences*,
597 117, 3960-3966, 10.1073/pnas.1919343117, 2020.



598 Zhang, R., Khalizov, A. F., Pagels, J., Zhang, D., Xue, H., and McMurry, P. H.:
599 Variability in morphology, hygroscopicity, and optical properties of soot aerosols
600 during atmospheric processing, *Proceedings of the National Academy of Sciences of*
601 *the United States of America*, 105, 10291-10296, 10.1073/pnas.0804860105, 2008.

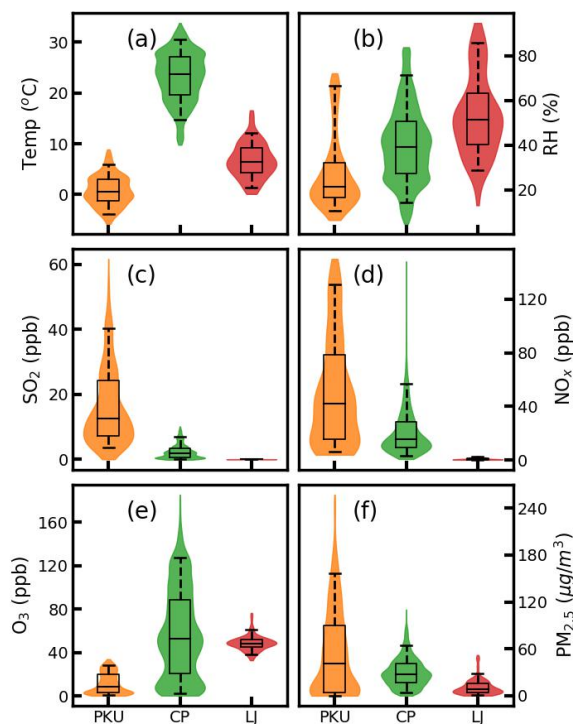
602 Zhao, G., Tao, J., Kuang, Y., Shen, C., Yu, Y., and Zhao, C.: Role of black carbon
603 mass size distribution in the direct aerosol radiative forcing, *Atmos. Chem. Phys.*, 19,
604 13175-13188, 10.5194/acp-19-13175-2019, 2019.

605 Zhao, G., Hu, M., Fang, X., Tan, T., Xiao, Y., Du, Z., Zheng, J., Shang, D., Wu, Z.,
606 Guo, S., and Zhao, C.: Larger than expected variation range in the real part of the
607 refractive index for ambient aerosols in China, *Science of The Total Environment*, 779,
608 146443, 10.1016/j.scitotenv.2021.146443, 2021.

609 Zheng, J., Hu, M., Du, Z., Shang, D., Gong, Z., Qin, Y., Fang, J., Gu, F., Li, M., Peng,
610 J., Li, J., Zhang, Y., Huang, X., He, L., Wu, Y., and Guo, S.: Influence of biomass
611 burning from South Asia at a high-altitude mountain receptor site in China,
612 *Atmospheric Chemistry and Physics*, 17, 6853-6864, 10.5194/acp-17-6853-2017,
613 2017.

614

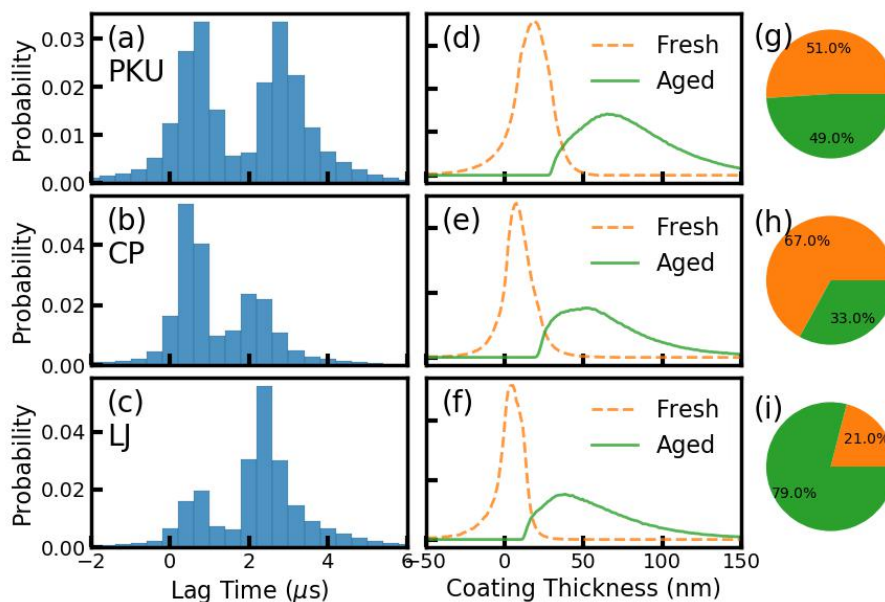
615



616

617 **Figure 1.** The measured distribution of (a) temperature, (b) RH, (c) SO₂, (d) NO_x, (e)
618 O₃ and (f) PM_{2.5} for PKU (orange), CP (green) and LJ (red) sites, respectively. The
619 box and whisker plots represent the 5th, 25th, 75th, and 95th percentiles. The width of
620 the filled colors represents the probability distributions of the corresponding measured
621 values.

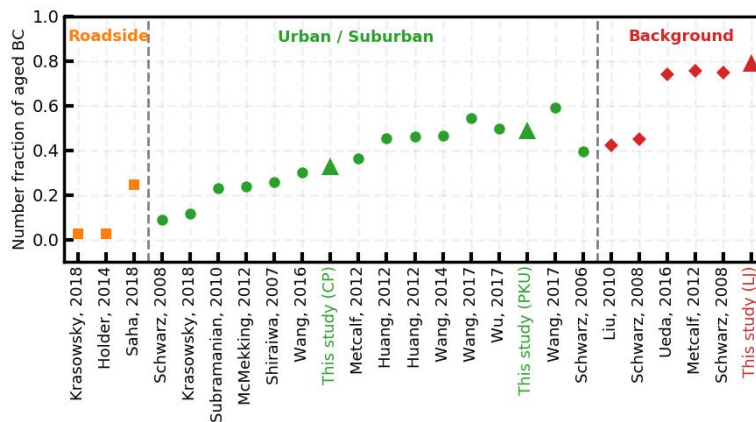
622



623

624 **Figure 2.** (a) The measured probability distribution of the lag time for the PKU site.
625 Panel (d) shows the corresponding coating thickness distributions of fresh (orange)
626 and aged (green) BC-containing aerosols. Panel (g) gives the number fraction of the
627 fresh (orange) and aged (green) BC-containing aerosols. Panel (b), (e), and (h) are the
628 corresponding values for the CP site. Panel (c), (f), and (g) give the results for LJ sites.

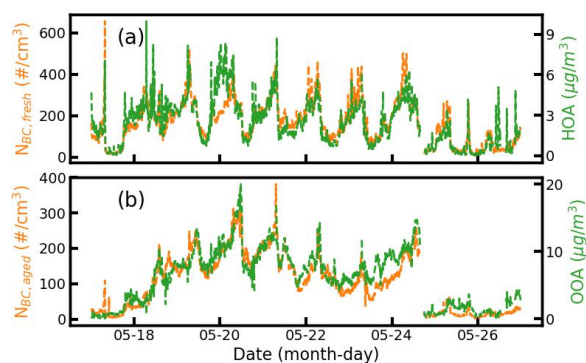
629



630

631 **Figure 3.** Measured number fraction of the aged BC under different atmospheric
632 environments based on literature. Our measured values are shown as triangles.

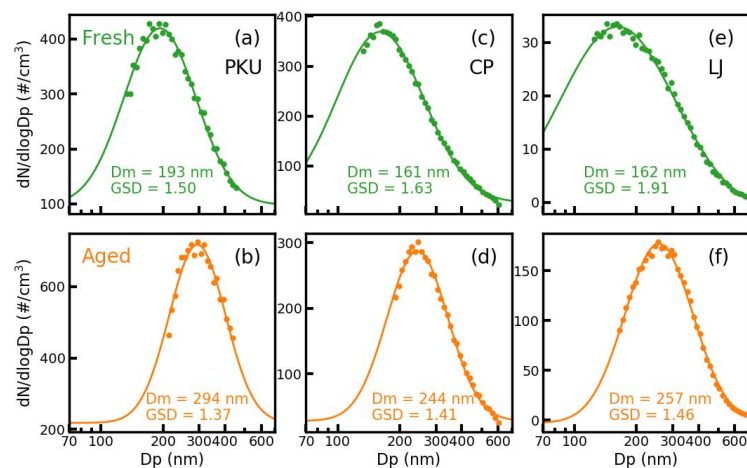
633



634

635 **Figure 4.** The time series of (a) the number concentration of the fresh BC (orange)
636 and the mass concentration of HOA (green), (b) the number concentration of aged BC
637 (orange), and the mass concentration of OOA (green).

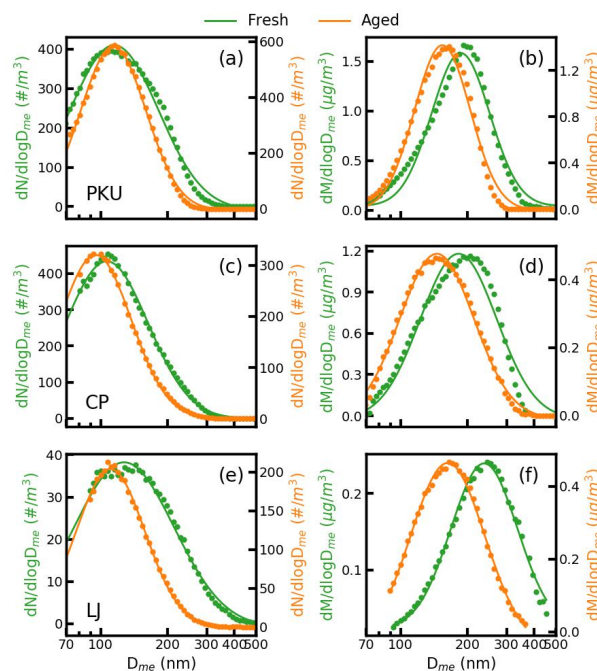
638



639

640 **Figure 5.** The number size distributions of the fresh BC-containing aerosols at (a)
641 PKU, (c) CP, and (e) LJ sites. Panels (b), (d), and (f) are the number size distributions
642 of the aged BC-containing aerosols for the PKU, CP, and LJ sites, respectively. The
643 dots in the figure are the measurement results and the lines are the corresponding fit
644 results with a log-normal distribution.

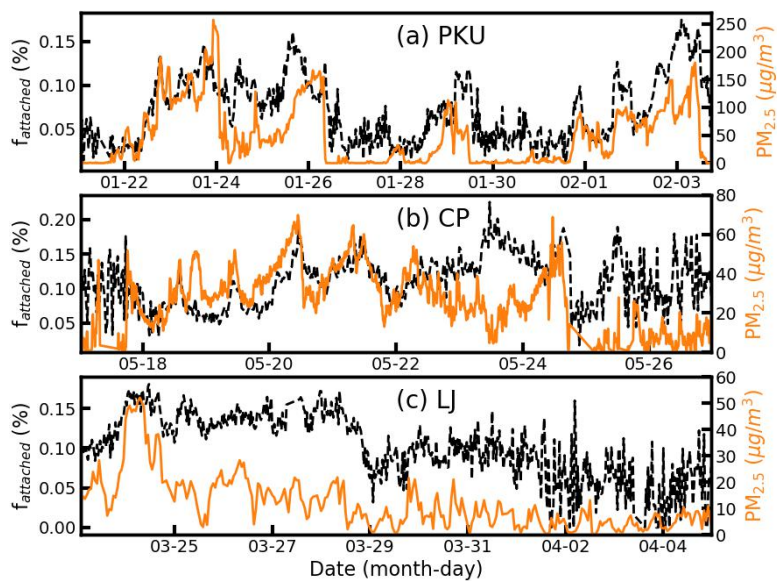
645



646

647 **Figure 6.** The BC core number size distributions of the fresh (green) and aged
648 (orange) BC aerosols for the (a) PKU, (c) CP, and (e) LJ sites. Panel (b), (d) (f) show
649 the BC core mass distributions of the fresh (green) and aged (orange) BC aerosols for
650 the PKU, CP, and LJ sites, respectively.

651



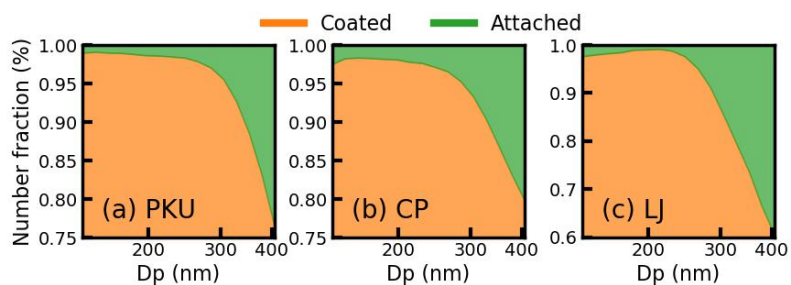
652

653 **Figure 7.** The time series of the number fractions of the attached BC (black) and
654 $PM_{2.5}$ mass concentrations (orange) for the (a) PKU, (b) CP, and (c) LJ sites.

655



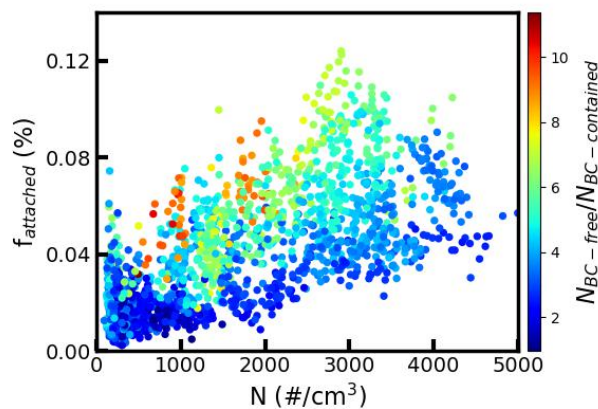
656



657

658 **Figure 8.** The number fractions of the coated and attached BC under different
659 diameters for the (a) PKU, (b) CP, and (C) LJ sites.

660



661

662 **Figure 9.** The number fractions of the attached BC aerosols under different total
663 aerosol number concentrations for the CP sites. The filled colors represent the ratios
664 between the BC-free aerosol number concentrations to the BC-containing aerosol
665 number concentrations.

666



667 **Table 1.** The D_{me} and GSD values of the BC core at different sites.

Site	Value	Number Distribution		Mass Distribution	
		Fresh	Aged	Fresh	Aged
CP	D_{me} (nm)	115	114	187	154
	GSD	1.58	1.40	1.35	1.34
PKU	D_{me} (nm)	107	95	182	146
	GSD	1.53	1.45	1.48	1.47
LJ	D_{me} (nm)	127	111	238	163
	GSD	1.68	1.43	1.47	1.41

668