



# 1 **Aerosol formation and growth rates from chamber experiments** 2 **using Kalman smoothing**

3 Matthew Ozon<sup>1</sup>, Dominik Stolzenburg<sup>2</sup>, Lubna Dada<sup>2</sup>, Aku Seppänen<sup>1</sup> and Kari E.J. Lehtinen<sup>1,3</sup>

4 <sup>1</sup> Department of Applied Physics, University of Eastern Finland, Kuopio, Finland

5 <sup>2</sup> Institute for Atmospheric and Earth System Research/ Physics, University of Helsinki, 00014 Helsinki, Finland

6 <sup>3</sup> Atmospheric Research Centre of Eastern Finland, Finnish Meteorological Institute, Kuopio, Finland

7  
8 *Correspondence to:* Kari E.J. Lehtinen ([kari.lehtinen@uef.fi](mailto:kari.lehtinen@uef.fi))

9 **Keywords.** Particle number size distribution, nucleation, growth, chamber experiments, Bayesian state estimation,  
10 Kalman smoother, DMA train

## 11 **Abstract.**

12 Bayesian state estimation in the form of Kalman smoothing was applied to Differential Mobility Analyser Train (DMA-  
13 train) measurements of aerosol size distribution dynamics. Four experiments were analysed in order to estimate the aerosol  
14 size distribution, formation rate and size-dependent growth rate, as functions of time. The first analysed case was a  
15 synthetic one, generated by a detailed aerosol dynamics model, and the other three chamber experiments performed at the  
16 CERN CLOUD facility. The estimated formation and growth rates were compared with other methods used earlier for  
17 the CLOUD data and with the true values for the computer-generated synthetic experiment. The agreement in the growth  
18 rates was remarkably good for all studied cases. The formation rates matched also well, especially considering the fact  
19 that they were estimated from data given by two different instruments, the other being the Particle Size magnifier (PSM).  
20 The presented Fixed Interval Kalman Smoother (FIKS) method has clear advantages compared with earlier methods that  
21 have been applied to this kind of data. First, FIKS can reconstruct the size distribution between possible size gaps in the  
22 measurement in such a way that it is consistent with aerosol size distribution dynamics theory, and second, the method  
23 gives rise to direct and reliable estimation of size distribution and process rate uncertainties if the uncertainties in the  
24 kernel functions and numerical models are known.

## 25 **1 Introduction**

26 Atmospheric new particle formation and growth are important phenomena when considering global aerosol  
27 concentrations. Their concentration together with their size distribution and chemical composition determines how  
28 aerosols affect visibility, health and climate (Albrecht, 1989; Appel et al., 1985; Daellenbach et al., 2020; Pope and  
29 Dockery, 2006; Twomey, 1974). These are determined by atmospheric dynamics and aerosol dynamics such as new  
30 particle formation and growth as well as removal rates. Nieminen et al. (2018) reviewed the existing literature on the  
31 formation and growth rates ranging from polar sites, with very small aerosol concentrations to polluted urban sites with  
32 extremely high concentrations. The rates have been typically estimated using the methodology reviewed in Kulmala et  
33 al. (2012), based on rather simple regression or balance equation approaches, suffering from potentially crude  
34 approximations and permitting no proper estimation of the uncertainties. At the same time, however, instrument  
35 development, especially advances in particle detection efficiency and mass spectrometry, has developed rapidly



36 (Kangasluoma et al., 2020). Potentially superior advanced data analysis methods have not been used, and, it is likely that  
37 there are significant inaccuracies in the estimated particle formation and growth rates estimated previously (Kürten et al.,  
38 2018).

39 There have been some attempts to estimate aerosol formation and growth rates with different inverse methods (Henze et  
40 al., 2004; Kuang et al., 2012; Lehtinen et al., 2004; Sandu et al., 2005; Verheggen and Mozurkewich, 2006; Viskari et al.,  
41 2012). We are, however, not aware of any of the above-mentioned methodology being used widely. The most promising  
42 ones in our view, that include also estimations of uncertainties, have been the ones by Kupiainen-Määttä (2016) and  
43 Shcherbacheva et al. (2020), who used Markov Chain Monte Carlo methodology to estimate evaporation rates as well as  
44 their uncertainties from synthetic cluster dynamics data. In addition, the INSIDE-method by Pichelstorfer et al. (2018),  
45 which is based on numerical solution of the aerosol general dynamic equation and matching the solution optimally to  
46 integrated measured concentrations of selected size intervals, has been successfully applied to determining growth rates  
47 at the CLOUD (Cosmics Leaving Outdoor Droplets) experiments at CERN (European Organization for Nuclear  
48 Research) (Stolzenburg et al., 2020). Furthermore, the results in very recent manuscript by McGuffin et al. (2020), in  
49 which nucleation, growth and emission rates of are estimated using techniques from the field of nonlinear process control,  
50 seem promising.

51 In a recent paper, Ozon et al. (2020) presented BAYROSOL, a Julia software package that combines a finite difference  
52 solution to the general dynamic equation for aerosols (GDE; Seinfeld and Pandis, 2016) to Bayesian state estimation in  
53 order to estimate unknown size dependent process rates (nucleation, condensation, losses) from known time evolution of  
54 the aerosol size distribution. Using synthetic SMPS measurements Ozon et al., (2020) showed that the Fixed Interval  
55 Kalman Smoother (FIKS; Kaipio and Somersalo, 2005) performed very well in estimating the process rates of the GDE  
56 in two distinct cases. First in a case in which continuous nucleation, growth and losses lead to a nearly steady state size  
57 distribution and second also in a case in which there is a growing nucleation mode after a nucleation burst. In the method,  
58 the unknowns (such as the discretized particle size distribution) are modeled as random variables, and their prior  
59 probability distributions are incorporated in the solution of the inverse problem. One important key feature of the Kalman  
60 smoother method is that it estimates also the error covariance matrices of the process rates that is their uncertainties, if  
61 the uncertainties of the measurement device are known.

62 In this manuscript, we show results of applying BAYROSOL, for the first time to real experimental data. We use  
63 experiments performed with the differential mobility analyser-train (DMA-train; Stolzenburg et al., 2017) measuring  
64 new particle formation and growth at the CERN CLOUD chamber (Duplissy et al., 2016; Kirkby et al., 2011). In addition  
65 to testing the method with synthetic DMA-train data (in which the ‘correct’ results are known), we estimate formation  
66 and growth rates from three different formation and growth experiments: sulphuric-ammonia (Stolzenburg et al., 2020),  
67 alpha-pinene ozonolysis (Heinritzi et al., 2020) and iodic acid (He et al., 2021; Sipilä et al., 2016). We compare the  
68 formation rates with results obtained by using the methodology by Dada et al. (2020) based on Particle Size Magnifier  
69 (PSM) measurements and the growth rates with the results obtained by the INSIDE method (Pichelstorfer et al., 2018).  
70 We chose the DMA-train measurements for three main reasons: the high time resolution makes it an ideal instrument for  
71 nucleation studies, the collection efficiencies of the channels have been carefully characterized (Stolzenburg et al., 2017;  
72 Wlasits et al., 2020) and finally, the data from this instrument has not been used to estimate directly particle formation  
73 rates so far.



## 74 2 Bayesian framework for parameter estimation

### 75 2.1 Aerosol measurement and evolution models

76 The quantities of interest in chamber experiments studying new particle formation are the particle number size-  
77 distribution, the formation rate at the critical cluster size and the growth rate of the freshly formed particles. The available  
78 measurements to infer these quantities usually size classify the aerosol and measure the size classified number  
79 concentrations. While the retrieval of the particle number size distribution from such measurements is a classical inverse  
80 problem (Chambolle and Pock, 2011; Kandlikar and Ramachandran, 1999; Wolfenbarger and Seinfeld, 1990), the  
81 estimation of the process parameters (formation and growth rate) is often done by analysis of the time evolution of the  
82 retrieved particle size distribution. Here, we focus on the formulation of the problem within a statistical Bayesian  
83 framework, where the state parameters described by a measurement model and an evolution model are treated as  
84 multivariate random processes and are estimated from measurements using a FIKS (Ozon et al., 2020).

85 A measurement can be described by a vector  $y^k$  representing  $m$  indirect observations (channels of the instrument) of the  
86 particle size distribution  $n(d_p, t_k)$ . The observations are linked to the size distribution by the so-called instrument transfer  
87 (or kernel) functions  $\mathcal{H}$  such that:

$$88 y^k = \int \mathcal{H}(d_p) n(d_p, t_k) dd_p \quad (1)$$

89 The transfer functions  $\mathcal{H}$  can be inferred from calibration experiments and instrument design considerations. Direct  
90 inversion of Eq. (1) for every time instant  $t_k$  is typically an underdetermined and ill-posed inverse problem and requires  
91 some additional assumptions in order to avoid oscillatory and unstable solutions (Kandlikar and Ramachandran, 1999).  
92 At the same time, the time evolution of the particle size-distribution  $n(d_p, t)$  can be described by the aerosol general  
93 dynamic equation (GDE):

$$94 \frac{\partial n}{\partial t}(d_p, t) = -\frac{\partial g(d_p, t)n(d_p, t)}{\partial d_p} - \lambda(d_p, t)n(d_p, t) - \text{CoagSink}(\beta, d_p, t) + \text{CoagSrc}(\beta, d_p, t) \quad (2)$$

95 Here,  $g(d_p, t)$  is the condensation growth/evaporation rate,  $\lambda(d_p, t)$  is the particle loss by deposition or dilution and  
96 CoagSink and CoagSrc are the sink and source rates due to particle coagulation within the size distribution with the  
97 coagulation coefficients  $\beta$ . An exact expression of all terms can be found in e.g. Ozon et al. (2020) and Seinfeld and  
98 Pandis (2016). The boundary conditions of Eq. (2) are given by the apparent formation rate  $J_{d_{min}}(t) =$   
99  $g(d_{min}, t)n(d_{min}, t)$  of newly formed particles at the minimum detectable size  $d_{min}$  and a zero numerical flux condition  
100  $g(d_{\infty}, t)n(d_{\infty}, t) = 0$  at very large sizes. Altogether, the process parameters  $g(d_p, t), J(t), \lambda(d_p, t)$  and  $\beta(d_i, d_j)$  as well  
101 as the initial and boundary conditions determine completely the evolution of the size distribution, but especially  $g(d_p, t)$   
102 and  $J(t)$  are usually not known. The coagulation coefficients  $\beta(d_i, d_j)$  can often be obtained from theory (and coagulation  
103 can be even neglected in many applications with low particle concentrations) and the loss parameters  $\lambda(d_p, t)$  are well  
104 quantified for controlled aerosol chamber experiments.

105 A single measurement of the size distribution  $y^k$  does not depend explicitly on the process parameters, but as  $g(d_p, t)$   
106 and  $J(t)$  determine the temporal evolution of  $n(d_p, t_k)$  the estimation of the process parameters is feasible from a  
107 sequence of  $l$  measurements  $y^k$  at several time instances.



108 **2.2 State estimation with Kalman smoothing**

109 Following Ozon et al. (2020) we formulate the problem as a Bayesian state estimation problem. After discretization of  
 110 the problem in size space ( $i = 1, \dots, q$ ) and time ( $k = 1, \dots, l$ ), we can define the state variable  $X^k = [N_i^k \ g_i^k \ \lambda_i^k \ J^k]$   
 111 for each time step  $k$  with the particle concentrations  $N_i^k$  per size discretization bin  $i$ , the condensation and loss terms  $g_i^k$   
 112 and  $\lambda_i^k$ , respectively, for each size discretization bin  $i$  and the nucleation rate  $J^k$ . The problem can then be formulated as:

113 
$$X^{k+1} = F(X^k) + w^k \quad (3)$$

114 
$$y^k = HX^k + v^k \quad (4)$$

115 Eq. (3) represents the discretized non-linear evolution model, which is based on the general dynamic equation for  $N^k$ , on  
 116 second order processes for  $g^k$  and  $J^k$ , and a random walk evolution for  $\lambda^k$  (see Section 2.3). Eq. (4) represents the  
 117 discretized linear observation model. The incorporation of positivity constraints for the process rates and the  
 118 aforementioned second order models require minor modifications in the definition of the state variable  $X^k$ . This, however,  
 119 does not cause any changes in the form of the state-space model (3)-(4). For the details on the above modifications as  
 120 well as discretization of the GDE model, we refer to Ozon et al. (2020). The terms  $w^k$  and  $v^k$  are the error terms, which  
 121 are approximated as normally distributed  $\mathcal{N}(0, \Gamma_w^k)$  and  $\mathcal{N}(0, \Gamma_v^k)$  with the covariance matrices  $\Gamma_w^k, \Gamma_v^k$ , which not only  
 122 include stochastic noise, but also errors due discretization, model and parameter uncertainties. The GDE, i.e. the non-  
 123 linear evolution model for  $N^k$  (Eq. (2)), is similar to an advection equation. Therefore, its numerical discretization  
 124 schemes are often unstable and must be treated carefully to avoid oscillation and divergence or to minimize numerical  
 125 diffusion (Shen et al., 2007; Smolarkiewicz, 1984). Thus, we show detailed considerations on the magnitude of the  
 126 different error terms in the Supplement.

127

**Algorithm 1** Extended Kalman Filter (EKF)

**Algorithm 2** Fixed Interval Kalman Smoother (FIKS)

**Initial state:** Expectation  $X^{0|0}$  and covariance  $\Gamma^{0|0}$

**Initialization:** Run Algorithm 1, store all variables

**for**  $k = 1, \dots, l$

**for**  $k = l - 1, \dots, 1$

Prediction: expectation and covariance

Backward gain matrix

$$X^{k|k-1} = F(X^{k-1|k-1})$$

$$A^k = \Gamma^{k|k} (\partial F)^T (\Gamma^{k+1|k})^{-1}$$

$$\Gamma^{k|k-1} = \partial F^{k-1} \Gamma^{k-1|k-1} (\partial F^{k-1})^T + \Gamma_w^{k-1}$$

Smoother expectation and covariance

Kalman gain matrix:

$$X^{k|K} = X^{k|k} + A^k (X^{k+1|K} - X^{k+1|k})$$

$$K^k = \Gamma^{k|k-1} (H^k)^T (H^k \Gamma^{k|k-1} (H^k)^T + \Gamma_v^k)^{-1}$$

$$\Gamma^{k|K} = \Gamma^{k|k} + A^k (\Gamma^{k+1|K} - \Gamma^{k+1|k}) (A^k)^T$$

Measurement update: filter expectation and covariance

**end**

$$X^{k|k} = X^{k|k-1} + K^k (Y^k - H^k X^{k|k-1})$$

$$\Gamma^{k|k} = (I - K^k H^k) \Gamma^{k|k-1}$$

**end**

128 **Table 1:** Extended Kalman Filter and Fixed Interval Kalman Smoother algorithms for estimation of the state variables and their  
 129 variances  $X^k$  and  $\Gamma^k$ .



130 Considering this structure of the problem, a non-linear extension to the Kalman Filter (Extended Kalman Filter; EKF) is  
131 a well suited algorithm for solving the unknown size-distribution and process parameters (Gelb, 1974; Kaipio and  
132 Somersalo, 2005). It is a two stage recursive procedure, where in the first stage the future state and propagation of error  
133 is predicted based on the state evolution model (Eq. (3)). In the second stage, the state variable and its covariance are  
134 estimated by updating the predicted state variable and covariance. This so-called measurement update accounts for the  
135 discrepancy between the realized measurements at time  $t_k$  and modelled measurements corresponding to the predicted  
136 state variable. This procedure is repeated until the final measurement  $k = l$ . After finishing the EKF recursions, we utilize  
137 a Fixed Interval Kalman Smoother (FIKS), which consists of a backward recursion from a backward gain matrix and  
138 smooths the results by backwards recursion from  $l$  to 1. The workflow of the EKF and FIKS are illustrated in Table 1 and  
139 more details on this algorithm can also be found in Ozon et al. (2020).

### 140 2.3 Adaption to chamber experiments

141 The state space model has been adjusted to represent best the evolution of an aerosol system during new particle formation  
142 experiments in an atmospheric simulation chamber like CLOUD. For the time evolution of the process parameters, we  
143 assume a rather smooth evolution for the nucleation and growth rates, approximated by a second order process (Ozon et  
144 al., 2020):

$$145 J^{k+1} = (r_1 + r_2)J^k - r_1 r_2 J^{k-1} + w_J^k \quad (5)$$

$$146 g^{k+1} = (r_1 + r_2)g^k - r_1 r_2 g^{k-1} + w_g^k \quad (6)$$

147 The constants  $r_1, r_2$  depend on the characteristic time of change, discretization time and a dampening factor and their  
148 definition can be taken from (Ozon et al., 2020) and the corresponding values for our experiments are listed in Table S1  
149 in the Supplement. The loss rates do not depend on time, in the chamber experiments of interest, and are well characterized  
150 by dedicated wall loss experiments (Stolzenburg et al., 2020). The dilution time of the chamber  $\tau_{dil} = Q_{tot}/V_{chamber}$ .  
151 The time evolution is described by a random walk with a small stochastic noise term  $w_\lambda^k$ , and the expectation of the initial  
152 state is set to the experimentally determined value with a standard deviation of  $\pm 10\%$ :

$$153 \lambda^{k+1} = \lambda^k + w_\lambda^k \quad (7)$$

154 For fully defining the problem (Eq. (3)-(4)), an estimate of the covariance matrices corresponding to the error terms  $v^k$   
155 and  $w^k$  is needed. The definition of the covariance matrices corresponding to the state noise  $w^k$  on the size-distribution  
156 evolution  $w_N^k$ , the growth rate  $w_g^k$ , the formation rate  $w_f^k$  and the wall loss rate  $w_\lambda^k$  follow the consideration of Ozon et al.  
157 (2020). The covariances of the wall losses, the growth rate and the size-distribution are dominantly diagonal with some  
158 additional off-diagonal terms in order to account for a correlation in size. The formulation given by Ozon et al. (2020)  
159 was slightly altered to give a stronger correlation between the closest size bins (values for the different experiments are  
160 given in Table S1 in the Supplement):

$$161 \Gamma_{N/g/\lambda}^k(i, j) = \sigma_{i, N/\lambda/g} \sigma_{j, N/\lambda/g} \exp\left(-\left(\frac{i-j}{\delta_{N/\lambda/g}}\right)^{\alpha_{N/\lambda/g}}\right) \quad (8)$$

162 For the size-distribution evolution, we find that  $\sigma_{i, N}^2 = (\delta^k)^2 \text{Var}(W_i^k)$ , with  $\delta^k$  the discretization time-step and  $W_i^k$  the  
163 error of the discretization of the size-distribution evolution. A detailed derivation of  $W_i^k$  is given in the Supplement.

164 The modelling error of the observation model and the measurement noise both contributing to  $v^k$  are assumed to be  
165 mutually independent. For this reason, the covariance of the error term  $\Gamma_v^k$  in the measurement model is written as a sum  
166 of the covariances of these two random variables, i.e.  $\Gamma_v^k = \Gamma_{mod}^k + \Gamma_y^k$ . For a detailed derivation we refer to the  
167 Supplement, where we also show that the discretization error is negligible compared to model and measurement error if



168 a fine enough size discretization is chosen. We approximate  $\Gamma_y^k$  with uncorrelated processes, and hence the covariance  
169 matrices are of the diagonal form. For the measurement error, the variance is given by Poisson counting statistics in the  
170 case of a single-particle counting detector such as a condensational particle counter (CPC):

$$171 \Gamma_y^k(i, i) = y^k(i) \quad (9)$$

172 For the model uncertainty, we assume the variance of the kernel  $\text{Var}(\Delta H_{i,j})$  is composed of an uncertainty proportional  
173 to  $H_{i,j}$  (for example due to an offset in the absolute calibration of the instrument) and a shifting size information error (for  
174 example discrepancy between set and actual classified size in a mobility spectrometer). It can then be formulated as  
175 (detailed values for the experiments under investigation can be found in Table S1 in the Supplement):

$$176 \Gamma_{mod}^k(i, i) = \sum_{j=1}^q (n(d_j) \Delta_j)^2 \text{Var}(\Delta H_{i,j}) \quad (10)$$

### 177 3 Experimental methods

178 We use experimental data from the CERN CLOUD experiment (Duplissy et al., 2016; Kirkby et al., 2011) where we  
179 measured particle size-distributions in the sub-10 nm range with a DMA-train (Stolzenburg et al., 2017). The raw data  
180 obtained from the DMA-train is used as input for the analysis of three different sets of experiments performed in the  
181 atmospheric simulation chamber. The DMA-train instrument kernels are also used for modelling an instrument response  
182 to simulated size-distribution data in order to verify the general performance of the FIKS to DMA-train like data.

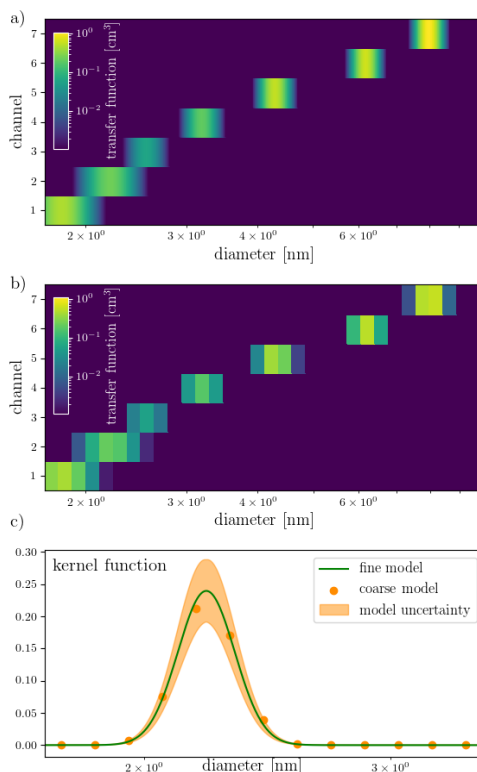
#### 183 3.1 DMA-train

184 The DMA-train is an electrical mobility spectrometer, specifically designed to measure sub-10 nm size-distributions  
185 (Stolzenburg et al., 2017). Six identical DMAs are applied in parallel i.e. they sample through the same inlet. They are  
186 set to six distinct but fixed voltages and hence classified sizes. The charging state of the aerosol is pre-conditioned in two  
187 TSI Inc. Advanced Aerosol Neutralizers (Model 3088), each supplying three DMAs at 5.5 litre per minute (lpm) total  
188 flow. We use the Wiedensohler approximation (Wiedensohler, 1988) to describe the steady-state charge distribution at  
189 the DMA inlets. Kallinger and Szymanski (2015) showed that for the used neutralizers the steady-state charge distribution  
190 is still achieved for flow rates up to 5 lpm. After size classification, the aerosol is detected in condensation particle  
191 counters. Four channels are equipped with TSI Inc. Model 3776 ultrafine CPCs for detection of aerosols down to 2.5 nm.  
192 They were operated at reduced temperature settings in order to increase the detection efficiency of the smallest particles,  
193 achieving a 50% detection efficiency for particles as small as 2 nm (Wlasits et al., 2020). Two channels of the DMA-train  
194 were operated with particle counters specifically designed for sub-2 nm particle detection using diethylene glycol (DEG),  
195 an Airmodus Ltd. particle size magnifier (Model A10, PSM) and a TSI Inc. nano-Enhancer (Model 3777). Both are used  
196 as boosters to activate the particles, which are subsequently detected by either an Airmodus Ltd. CPC (Model A20, for  
197 the PSM) or a TSI Inc. CPC (Model 3772, for the nano-Enhancer). Both channels have a higher aerosol flow rate of 2.5  
198 lpm resulting in a broader transfer function and higher transmission at the DMA compared to the 1.5 lpm sample flow in  
199 the other four channels. The sheath flow at the DMAs is kept constant at 15 lpm for all six channels.

200 The constant sampling at fixed sizes allows for either a higher time-resolution at large aerosol number concentrations or  
201 a higher sensitivity towards low number concentrations due to longer signal averaging times compared to a scanning or  
202 stepping differential mobility spectrometer. To increase the number of measured particle sizes, one DMA is still operated  
203 in an alternating mode, switching between 6.2 and 8 nm every ten seconds. The other DMAs are set to classify particles  
204 of 4.3, 3.2, 2.55, 2.2 and 1.8 nm. The instrument kernels are obtained from calibration experiments, where we use the



205 DMA transfer function and sampling loss characterization from Stolzenburg et al. (2017), the CPC activation efficiencies  
206 from Wlasits et al. (2020) and the charging efficiency was tested to follow the Wiedensohler approximation in Tauber et  
207 al. (2020). The kernel functions for all seven classified sizes are shown in Fig. 1 for an instrument averaging time of 120  
208 seconds, sulphuric acid-like test particles (using the Ammonium Sulfate detection efficiencies from Wlasits et al. (2020))  
209 and including the detector flow rates of each condensation particle counter. Therefore, the kernels can be used to convert  
210 raw particle counts at the detecting CPCs into a particle size-distribution (within an inverse problem) and vice-versa. Note  
211 that for different chemical composition of the input particles, the CPC response might be different. Therefore, the kernels  
212 used for analysing experiments where particles were formed from oxidized organics (Kirkby et al., 2016) or from iodic  
213 acid (He et al., 2021) are different and approximated by the calibration curves for oxidized beta-caryophyllene and sodium  
214 chloride from Wlasits et al. (2020).



215  
216 **Figure 1:** Kernel functions of the DMA-train when classifying sulphuric acid particles at a signal averaging time of 120 seconds. (a)  
217 shows the continuous form of the transfer function (color code) of the seven DMA-train channels (y-axis), inferred from instrument  
218 calibrations (Stolzenburg et al., 2017). (b) shows the discretization into 32 size bins from 1.7 to 10 nm used in the Kalman smoother  
219 (c) shows the comparison between continuous form and used discretization for channel 2 with centroid diameter at 2.2 nm, together  
220 with the model uncertainty, which is used for the error estimate (see Table S1 in the Supplement).

### 221 3.2 CLOUD

222 We use experimental data from the DMA-train measuring new particle formation in nucleation experiments at the CERN  
223 CLOUD chamber. The 26.1 m<sup>3</sup> stainless-steel chamber provides a high-purity, temperature-controlled environment in  
224 order to perform experiments under atmospherically relevant conditions, where trace gases can be added precisely at pptv  
225 (parts per trillion by volume) level and sunlight can be simulated by UV-illumination of the chamber. We use three



226 different sets of experiments, of varying chemical composition, in order to demonstrate the performance of the FIKS in  
227 reconstructing formation and growth rates. See also Kirkby et al. (2011) and Duplissy et al. (2016).  
228 First, a nucleation experiment using sulphuric acid and ammonia was performed at 5°C by adding SO<sub>2</sub> and O<sub>3</sub> to the  
229 chamber and through the photo-dissociation of ozone, the formation of OH radicals and sulphuric acid was induced which  
230 resulted in strong new particle formation (Stolzenburg et al., 2020). Second, we performed nucleation and growth  
231 experiments at 5°C using oxidized organics from dark (i.e. no UV-illumination) ozonolysis of alpha-pinene (Kirkby et al.,  
232 2016; Stolzenburg et al., 2018). Both experiments resulted in moderate new particle formation rate and thus, in the Kalman  
233 smoother, we used DMA-train data that was averaged over 120 seconds time intervals. Third, we studied nucleation from  
234 iodine oxides at 10°C (He et al., 2021), which resulted in high particle formation rates and fast growth. For the third  
235 experiment, we reduced the DMA-train averaging time down to 20 seconds, while keeping high counting statistics over  
236 the averaging interval.

### 237 3.3 PSM derived formation rates

238 Particle formation rates ( $J_{dp}$ ) are calculated from the time derivative of the total particle number concentration larger than  
239 1.7 nm following the method introduced in Dada et al. (2020). The particle number size distribution is measured with the  
240 particle size magnifier (PSM) coupled with a condensation particle counter (1.5 – 2.5 nm), a TSI nano-SMPS (3 – 65 nm)  
241 and home-built long-SMPS (10 – 800 nm). The formation rates are corrected for the size dependent wall and coagulation  
242 losses. Additionally, since the chamber is operated in continuous mode, the particle concentrations are corrected for  
243 dilution losses. For more information on the PSM derived formation rates, see Dada et al. (2020). The uncertainty on  
244  $J_{1.7}$  was assumed to be 30% for the CLOUD chamber calculated using propagation of error in the concentration  
245 measurement, dilution, coagulation and wall losses as well as the error on the time-derivative of the total particle  
246 concentration.

### 247 3.4 Growth rates using INSIDE

248 In order to compare the particle growth rates derived by Kalman smoothing, we use the size- and time-dependent growth  
249 rate analysis tool INSIDE (Pichelstorfer et al., 2018). It uses input particle size-distributions at time  $t_1$  in order to simulate  
250 the known aerosol dynamics (coagulation, wall losses and dilution) until a time  $t_2$  (typically separated by one  
251 measurement cycle of an instrument, i.e. the 120 seconds averaging time mentioned above). At  $t_2$ , the simulated aerosol  
252 size-distribution is compared to the measured size-distribution and by evaluating the general dynamics equation above a

253 certain diameter  $d_{\text{eval}}$  the growth term  $\left. \frac{dd_p}{dt} \right|_{d_{\text{eval}}} = \frac{\left. \frac{dN_{\text{exp}}}{dt} \right|_{d_{\text{eval}}}^{\infty} - \left. \frac{dN_{\text{sim}}}{dt} \right|_{d_{\text{eval}}}^{\infty}}{n(d_p, t)|_{d_{\text{eval}}}}$  can be computed. Evaluating this equation for

254 several evaluation diameters  $d_{\text{eval}}$  and at all measurement times  $t_i$ , this results in time- and size-resolved growth rates  
255 from a size-distribution measurement, which is distinct from most others, integrative growth rate analysis approaches  
256 (Dada et al., 2020). However, compared to the Kalman smoothing, each time step is analysed individually and the analysis  
257 framework relies on already inverted size distributions, where a point-by-point inversion procedure is used for the DMA-  
258 train data of this work (Stolzenburg and McMurry, 2008). Moreover, the INSIDE method does not provide an uncertainty  
259 estimate on the growth rate calculation and hence a comparison with other growth rate calculations is beneficial to verify  
260 the results.



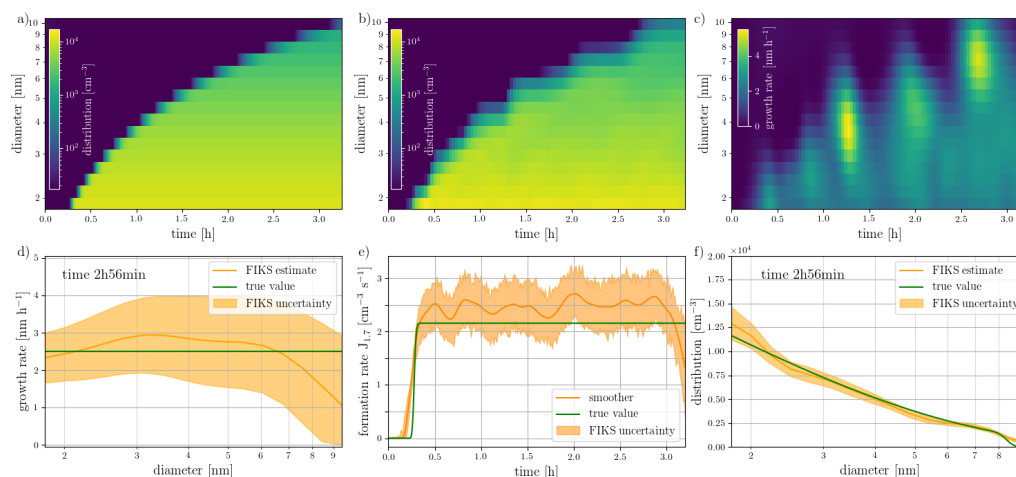


## 261 4 Results: Simulation and Experimental

### 262 4.1 Numerical simulation test

263 First, we modelled the DMA-train instrument numerically and simulated a data set representing a typical nucleation  
264 experiment performed in an atmospheric simulation chamber like CLOUD and applied the FIKS to this synthetic dataset.  
265 We used the same framework as in Ozon et al. (2020) to simulate a nucleation experiment with formation rate at 1 nm  
266  $J_{1.0} = 5.25 \text{ cm}^{-3} \text{ s}^{-1}$ , and size-independent growth rate  $GR = 2.5 \text{ nm h}^{-1}$  and the loss rates equal to the CERN CLOUD  
267 experiment (Stolzenburg et al., 2020). The evolution of the simulated size-distribution is shown in Fig. 2a. The  
268 measurement data  $y^k$  are then simulated using the kernel functions from Fig. 1a and altered with a Poisson-distributed  
269 random counting error. The FIKS is then applied to the measurement data with the input parameters given in Table S1 in  
270 the Supplement.

271 Figure 2 also shows the Kalman smoother estimates for the size-distribution, growth rate and formation rate at 1.7 nm.  
272 The reconstructed size-distribution (Fig. 2b) is very similar to the true size-distribution (Fig. 2a), especially taking into  
273 account the sparser discretization of the former. Also the estimated growth and nucleation rates agree well with true values  
274 of the respective process rates specified in the simulation. Moreover, the uncertainty estimates are feasible: For both  
275 quantities, the true values are within the uncertainty limits given by FIKS. The reconstructed size-distribution and  
276 especially, growth rate (Fig. 2 b and c, respectively) show some temporal oscillations, which are related to periods for the  
277 particle population to grow from the size range visible for one DMA-train channel to the next one. That is, the oscillation  
278 is a result of an insufficient coverage of the size-range by the DMA-train kernels and is discussed in more detail in Section  
279 4.5. Nevertheless, the overall good retrieval of the simulated size distribution and process rates demonstrates that Kalman  
280 smoothing approach is well suited for analysing DMA-train data from chamber nucleation experiments.

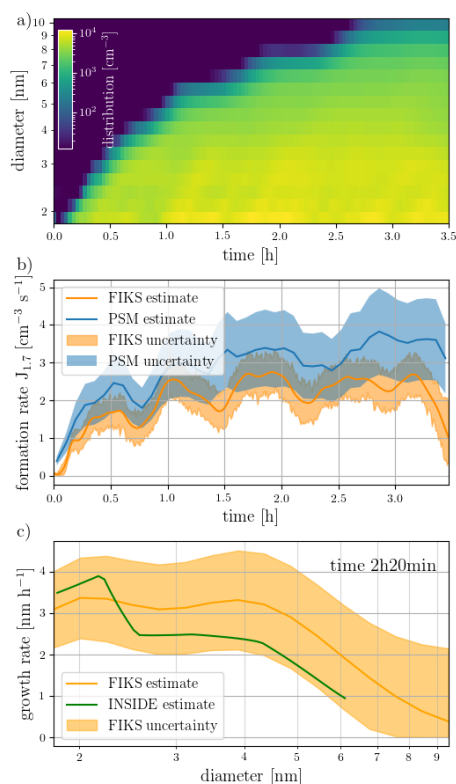


281 **Figure 2:** Results of the simulated chamber nucleation experiment: (a) The simulated, temporally evolving size distribution of aerosols.  
282 (b) The FIKS estimate for the size distribution. This estimate was computed based on the numerically simulated DMA-train data  
283 corresponding to the synthetic size distribution shown in (a). (c) The FIKS estimate for the growth rate of particles. (d) The growth rate  
284 corresponding to a single instant of time (2h 56min); here, the FIKS estimate and the associated uncertainty limits are plotted together  
285 with the true growth rate. (e) The FIKS estimate for the formation rate at 1.7 nm and its uncertainty, and the true formation rate. (f)  
286 FIKS estimate, its uncertainty and the true value of the particle size distribution corresponding to a single instant of time (2h 56min).  
287 In subfigures d-f, the orange lines represent the FIKS estimates (posterior expectations) and the orange shaded areas illustrate the  
288 uncertainties of the associated variables (more specifically, their approximate posterior standard deviation limits). The true values of  
289 the quantities are plotted in green.  
290



## 291 4.2 Sulphuric Acid-Ammonia experiment

292 We applied the FIKS to experimental data from a sulphuric acid-ammonia nucleation and growth experiment performed  
293 at 5°C at the CERN CLOUD chamber. The raw data measured with the DMA-train were averaged in 120-second time  
294 intervals in order to increase the counting statistics per channel and then used as input for BAYROSOL. The details of  
295 parameter choices in FIKS are given again in Table S1 in the Supplement. The results of applying the Kalman smoothing  
296 to this experimental data are shown in Fig. 3.



297

298 **Figure 3:** Results of an experiment performed at the CERN CLOUD chamber: The DMA-train data was acquired during sulphuric  
299 acid-ammonia nucleation and growth processes. Experimental conditions: 5°C, 60% RH, 5 ppb SO<sub>2</sub>, 120 ppb O<sub>3</sub> and 40 ppt NH<sub>3</sub>. (a)  
300 The FIKS estimate for the size distribution. (b) The evolution of the formation rate during the experiment. The orange line and shaded  
301 area represent the FIKS estimate and its uncertainty, respectively, while the solid blue line and blue area show the nucleation rate  
302 estimate and the associated 30% uncertainty based on a different set of instruments (particle size magnifier, PSM) and the standard  
303 approach for inferring nucleation rates at 1.7 nm (Dada et al., 2020). (c) The growth rate corresponding to time 2h19 min after the start  
304 of the experiment. The FIKS estimate and the associated uncertainty are marked with orange line and shaded area, respectively. The  
305 green line represents the growth rate estimate, which is obtained from DMA-train data by traditional inversion (INSIDE method).

306 The size- and time-dependence of the FIKS estimate for the true size-distribution (Fig. 3a) is very smooth, and it is also  
307 able to bridge the information gaps between the largest size-distribution channels. This is a significant improvement from  
308 the traditional point-by-point inversion, where the data of each DMA-train channel is inverted independently: When point-  
309 by-point inversion was applied to the same dataset, those information gaps caused severe reconstruction errors  
310 (discontinuities) in the size-distribution, see (Stolzenburg et al., 2020). The size-distribution shows also small temporal  
311 fluctuations. The fluctuations are even more clearly visible in the reconstructed evolution of the formation rate (Fig. 3b).  
312 Remarkably, the same fluctuations are also recovered when the formation rate is inferred from the PSM and nano-SMPS,

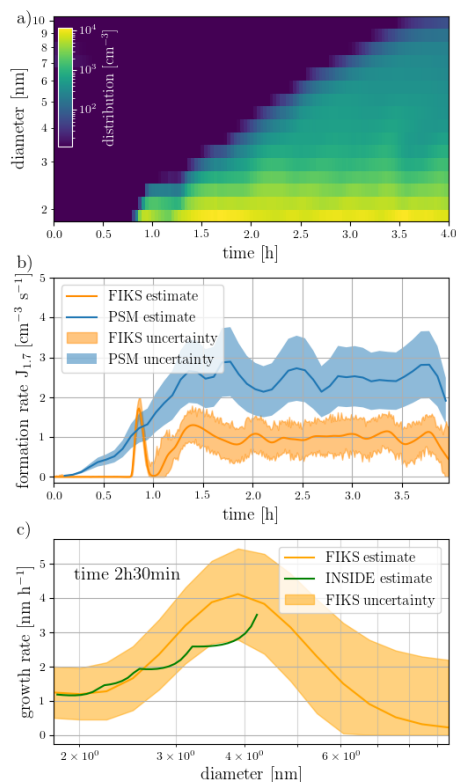


313 i.e. an entirely different set of instrumentation and different type of data-analysis. This result strongly suggests that the  
314 fluctuation of the particle formation truly occurs physically in the experiment and is not a reconstruction error caused by  
315 instrument noise or a bias caused by the inversion method. The absolute values of the two independent formation rate  
316 estimates agree upon a factor of 1.5. Furthermore, large portions of the uncertainty intervals of these two estimates overlap  
317 with each other, which is another indicator of the feasibility of FIKS to analysing DMA-train data. In addition, also the  
318 inferred growth rates agree within the systematic uncertainties for both approaches (Fig. 3c for one time instant). It is  
319 worth noting, however, that both growth rate estimates rely on data from the same instrument and are hence more  
320 interdependent than the formation rate estimates. Nevertheless, the good agreement between them corroborates the  
321 feasibility of Kalman smoothing to reconstructing nanoparticle growth rates from experimental data.

#### 322 4.3 Alpha-pinene ozonolysis experiment

323 Third, we applied the FIKS to experimental data obtained by the DMA-train from an alpha-pinene ozonolysis experiment.  
324 Besides the different chemical composition of the growing particles (resulting in different assumptions on the DMA-train  
325 transfer functions) in comparison to the experimental results used in section 4.2, this experiment is characterized by a  
326 different size-dependency of the growth rate, due to the increased condensation of low- and semi-volatile organics with  
327 increasing particle size (Simon et al., 2020; Stolzenburg et al., 2018; Tröstl et al., 2016). The formation rate, however,  
328 remains rather similar to the sulphuric acid-ammonia experiment under these specific experimental conditions, but the  
329 slower initial growth rates result in generally lower produced particle concentrations. The results of applying FIKS to the  
330 data from this experiment are shown in Fig. 4. Again, the size-distribution given by FIKS is much smoother than that  
331 obtained with a standard inversion procedure (see Heinritzi et al. (2020) for the same dataset inverted by the standard  
332 approach).

333 Because in alpha-pinene ozonolysis experiment the nucleation and growth rates were known to be similar to those in the  
334 sulphuric acid-ammonia experiment, the parameters of their evolution models were selected as in Section 4.2 (Table S1  
335 in the Supplement). The formation rate is again slightly lower than the one obtained from the PSM. Considering the fact  
336 that inter-instrument deviations in sub-10 nm size-distribution measurements can be as large as one order of magnitude  
337 (Kangasluoma et al., 2020), the achieved agreement is remarkable, especially as some fluctuations within the chamber  
338 can again be reconstructed in both approaches. The retrieved growth rates from the FIKS estimate and the INSIDE method  
339 agree remarkably well and both show the increasing growth rates with increasing particle size up to 4.3 nm, an indication  
340 for a strong Kelvin-effect in organic condensation (Stolzenburg et al., 2018; Tröstl et al., 2016). The decreasing Kalman  
341 smoother estimate above 4.3 nm is related to the fact that the FIKS searches a smooth estimate of the growth rate (Eq.  
342 (6)), but at that point (2h30min after experiment start), practically no information from the size-distribution above 4.3 nm  
343 is available, resulting in a slow decrease towards zero. This is due to the strong smoothness a priori used in the FIKS  
344 algorithm (Table S1 in the Supplement). The INSIDE method does not report growth rates from regions with no  
345 information (Pichelstorfer et al., 2018) from the size-distribution and hence the estimate stops at 4.3 nm.



346

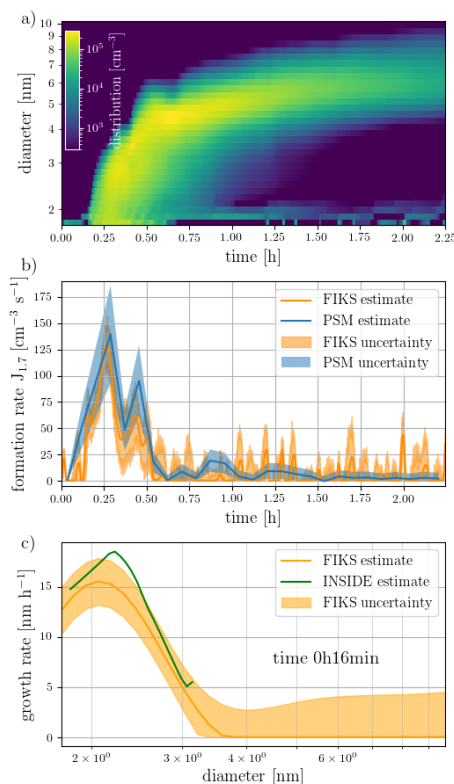
347 **Figure 4:** Results of an experiment performed at the CERN CLOUD chamber: The DMA-train data was acquired during an alpha-  
348 pinene ozonolysis. Experimental conditions: Experimental conditions: 5°C, 40% RH, 300 ppt alpha-pinene, 40 ppb O<sub>3</sub>. (a) The FIKS  
349 estimate for the size distribution. (b) The evolution of the formation rate during the experiment. The orange line and shaded area  
350 represent the FIKS estimate and its uncertainty, respectively, while the solid blue line and blue area show the nucleation rate estimate  
351 and the associated 30% uncertainty based on a different set of instruments (particle size magnifier, PSM) and the standard approach  
352 for inferring nucleation rates at 1.7 nm (Dada et al., 2020). (c) The growth rate corresponding to time 2h30 min after the start of the  
353 experiment. The FIKS estimate and the associated uncertainty are marked with orange line and shaded area, respectively. The green  
354 line represents the growth rate estimate, which is obtained from DMA-train data by traditional inversion (INSIDE method).

#### 355 4.4 Iodic Acid experiment

356 Finally, we analysed a more dynamic experiment of nucleation and growth from iodic acid (He et al., 2020, 2021). The  
357 experiment is characterized by extremely high nucleation rates, which are two orders of magnitude higher than in the  
358 sulphuric acid-ammonia and alpha-pinene ozonolysis experiments. However, the growth rate is only half an order of  
359 magnitude higher compared to the other two example cases. Figure 5 shows that, in spite of the highly dynamic  
360 experiment, the formation rate recovered by Kalman smoothing agrees with the estimate obtained from the PSM (Fig.  
361 5b). The usage of the DMA-train data with a time resolution of 20 seconds allows for the precise recovery of the spike in  
362 formation rate in the beginning of the experiment. This causes the build-up of a high condensation sink and vapour/cluster  
363 depletion almost shutting off any further nucleation during the continuation of the experiment. The four-minute time  
364 resolution of the data for the calculation of the nucleation rate from the PSM is limited in that respect. The reconstructed  
365 growth rates (Fig. 5c) show again agreement between the FIKS estimate and the INSIDE method, indicating a clear  
366 decreasing trend with size, which is expected for condensation at the kinetic limit if the vapour molecular size is taken  
367 into account (He et al., 2021; Lehtinen and Kulmala, 2003; Nieminen et al., 2010; Stolzenburg et al., 2020). The lower



368 values towards 1.8 nm could be caused by a biased estimate of the PSM detection efficiency, because neither a calibration  
369 for iodic acid clusters nor for sodium chloride particles (which was used for the other detectors in the DMA-train) was  
370 available. The instabilities in the size-distribution at the smallest sizes and fluctuations of the formation rate are expected  
371 considering the highly dynamic process of this experiment Overall, the good agreement for the inferred process  
372 parameters of the aerosol general dynamics equation, i.e. the formation and growth rates, is still remarkable. However,  
373 the reconstructed size-distribution from the FIKS estimate shows some discontinuities, especially during the growth above  
374 3 nm. This is because the available instrument information from the DMA-train starts to get very sparse in that size range  
375 given the dynamic processes involved in the iodic acid nucleation and growth. More available size channels (hence more  
376 DMAs in the case of the DMA-train) would help to resolve such discontinuities. We will therefore provide an instrument  
377 design recommendation based on simulated data in the next section.



378

379 **Figure 5:** Results of an experiment performed at the CERN CLOUD chamber: The DMA-train data was acquired during iodic acid  
380 nucleation and growth processes. Experimental conditions: Experimental conditions: 10°C, 80% RH, 100 ppt I<sub>2</sub>, 40 ppb O<sub>3</sub>. (a) The  
381 FIKS estimate for the size distribution. (b) The evolution of the formation rate during the experiment. The orange line and shaded area  
382 represent the FIKS estimate and its uncertainty, respectively, while the solid blue line and blue area show the nucleation rate estimate  
383 and the associated 30% uncertainty based on a different set of instruments (particle size magnifier, PSM) and the standard approach  
384 for inferring nucleation rates at 1.7 nm (Dada et al., 2020). (c) The growth rate corresponding to time 0h16min after the start of  
385 the experiment. The FIKS estimate and the associated uncertainty are marked with orange line and shaded area, respectively. The green  
386 line represents the growth rate estimate, which is obtained from DMA-train data by traditional inversion (INSIDE method).

#### 387 4.5 Instrument design recommendation from a signal processing point of view

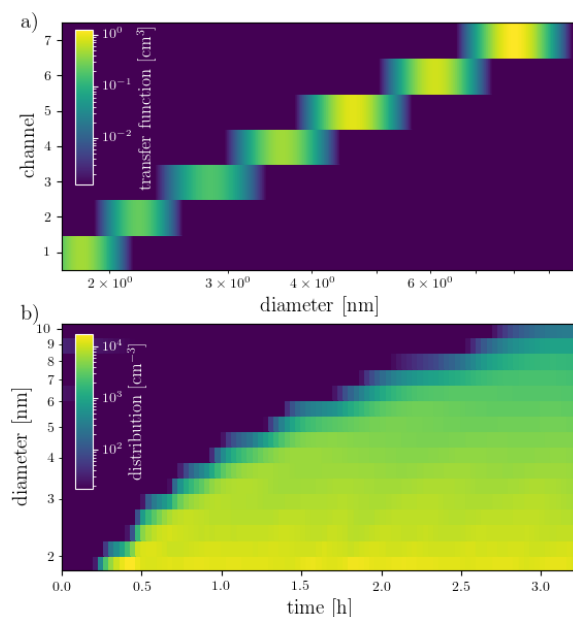
388 The size range covered by the seven DMA-train channels was chosen semi-arbitrarily based on some external constraints:  
389 the lowest measured centroid diameter was supposed to be as close as possible to 1.7 nm where the formation rate is



390 typically measured for experiments performed at the CLOUD chamber. In order to cover the sub-10 nm range the largest  
391 channels was set to 8 nm. The other channel diameters were chosen to yield sufficient coverage of the sub-3 nm range.  
392 The width of the transfer functions was fixed by the detector sample flow rates and the 15 lpm critical orifices provided  
393 by Grimm Aerosol for the DMA sheath air supply (also the standard flow rate used by the manufacturer for this type of  
394 DMA). However, the chosen centroid diameters (and hence fixed voltages at the DMAs) and selected sheath flow rates  
395 could easily be altered.

396 In order to study numerically the effect of choice for the DMA-train channels, we constructed a model corresponding to  
397 a channel choice different from that in the previous sections: The kernel for the DMA-train with seven channels having  
398 centroid diameters of 1.8, 2.2, 2.8, 3.6, 4.6, 6.1 and 8.0 nm is illustrated in Fig. 6a. Moreover, channels 3 to 7 have been  
399 altered to use a 2.5 lpm sample flow rate (could be achieved by a make-up flow at each CPC) and a reduced sheath flow  
400 rate of 10 lpm, only providing a resolution of theoretical non-diffusive resolution of  $\sim 4$ , which is significantly lower.  
401 However, this permits covering the entire size-range between 1.8 and 8 nm by overlapping channels. We revisited the  
402 numerical simulation described in Section 4.1 using DMA-train model corresponding to this configuration. The resulting  
403 size-distribution given by FIKS is shown in Fig. 6b. Comparison between the two reconstructions in Fig. 2b and Fig. 6b  
404 reveals that the new choice of DMA-channels stabilizes the size-distribution estimate significantly. While in Fig. 2b, the  
405 size distribution evolves in step-wise manner when the growing particle mode reaches larger size-channels, the size  
406 distribution in Fig. 6b is very smooth, also temporally. The latter estimate also resembles the true (simulated) size-  
407 distribution (Fig. 2a) more than the former one.

408 The result of this additional numerical study thus demonstrates that in the chosen conditions of the simulation, the  
409 reconstruction quality improves when the resolution of individual channels is lowered. This seemingly counterintuitive  
410 effect stems from the fact that FIKS estimates do not rely only on the measurements, but are also advised by the GDE  
411 model, which makes the problem of optimizing the measurement design a somewhat cumbersome task. A rigorous  
412 investigation of the optimal experimental design is out of the scope of this paper, but the above observation is worth  
413 noticing -- especially because a lot of recent experimental effort in the sub-10 nm range has been devoted to improving  
414 the instrument resolution (Kangasluoma et al., 2020). While on the individual channel level this might reduce systematic  
415 uncertainties, as discussed in Kangasluoma et al. (2020), signal processing rather requires a broad coverage of the size  
416 distribution than high resolution. However, the ideal instrument would combine both, full coverage of the size distribution  
417 but achieved with more, high-resolution channels. For the DMA-train principle, this would require new ideas in  
418 instrument design in order to incorporate more DMAs without the instrument becoming impractically bulky.



419

420 **Figure 6:** Adjusted DMA-train kernel for better signal processing (a), which could be achieved by choosing suitable centroid voltages  
421 and altering the resolution of the DMAs through sheath-flow and sample-flow adjustments. (b) shows the reconstructed simulated size-  
422 distribution of Fig. 2 using the adjusted kernel.

## 423 5 Conclusion

424 A recently developed methodology (BAYROSOL) applying the Fixed Interval Kalman Smoother (FIKS) to a finite  
425 difference solution to the aerosol GDE was used to analyze DMA-train measurements of aerosol dynamics at the CLOUD  
426 chamber facility at CERN. The overall aim of this methodology is to estimate unknown aerosol microphysical process  
427 rates as well as their uncertainties from size-distribution evolution measurements. In a previous paper, the methodology  
428 was shown to be able to predict new particle formation, growth and loss rates from synthetic computer-generated aerosol  
429 size distribution evolution data, while here the method has been applied to real experimental data for the first time. Four  
430 experimental cases with particle formation and growth were tested: 1) a computer-generated synthetic case, 2) sulphuric  
431 acid-ammonia, 3) alpha-pinene ozonolysis, and 4) iodic acid.

432 The DMA-train was selected for two main reasons: first, the instrument kernel functions are well characterized giving  
433 rise to reliable estimation of the instrument uncertainties and second, new particle formation rates have not been estimated  
434 directly from DMA-train measurements before. In addition, as the current version of the DMA-train is designed in such  
435 a way that the individual DMAs have a rather narrow collection kernels for the channels, with significant gaps between  
436 some of the channels, the FIKS can reconstruct the size distributions from the measured signals in such a way that the  
437 distributions are rather smooth and consistent with the GDE.

438 We compared the growth rates, which with BAYROSOL can be estimated as functions of both size and time, with those  
439 obtained by INSIDE, a method applied earlier to CLOUD data, and the agreement was remarkably good for all studied  
440 cases. INSIDE is also based on matching the GDE solution to measured size distribution dynamics, however without the  
441 capability of estimating uncertainties of the estimations. The FIKS-based estimates for the particle formation rates were  
442 compared with those estimated from data obtained by a separate instrument, the Particle Size Magnifier (PSM), based on



443 the rate of change of the total number concentration measured by the instrument corrected by coagulation and wall-losses.  
444 Again, the agreement was very good, especially considering the fact that instrument uncertainties are large at the very  
445 lowest end of the measured size spectrum. For the iodine-acid case, the FIKS estimate of the formation rate was even able  
446 to capture rapidly changing dynamics of the experiment. It was remarkable for all cases that some fluctuations in the  
447 formation rates were recovered by both methods independently, indicating that these are physical variations during the  
448 experiment  
449 Finally, we utilized the FIKS from an instrument development point of view. Typically, an as-high-as-possible resolution  
450 for the different measurement channels has been the aim when measuring nanometer-sized particles. This aim, however,  
451 gives rise to gaps in the measured size range, as is the case in the DMA-train studied. Thus we studied whether a better  
452 coverage of the size spectrum, but lower resolution of the individual channels would be advantageous for size distribution  
453 estimation.  
454 Summarizing, we believe that Bayesian state estimation methods such as FIKS can be very useful in the field of aerosol  
455 science in many aspects. As mentioned above, they can be used to fill gaps in measurements in such a way that not only  
456 the obtained size distributions but also unknown process rates are consistent with theory describing aerosol size  
457 distribution dynamics. In addition, the methodology provides estimations of uncertainties both for size distributions as  
458 well as process rates based on uncertainty estimations in the measurements and the used models, which is unfortunately  
459 not common when reporting results of aerosol measurements. Finally, it is conceivable that the methodology presented  
460 here will be superior to several previous approaches when combining measurement data obtained with several different  
461 instruments that operate at different size ranges. This will be a topic of our forthcoming studies.

#### 462 **Data availability**

463 The version of the implementation of the estimation method (BAYROSOL1.1)[<https://doi.org/10.5281/zenodo.4450492>]  
464 is available under the MIT Expat License; it is the version used to generate the results discussed in this paper. The package  
465 also contains the code used to obtain the results discussed in Ozon et al. 2020. It is possible to generate the simulated data  
466 described in section 4.1. The experimental data (DMA-train kernels) and raw data, are also included in the repository.

#### 467 **Competing interest**

468 The authors declare no competing interests.

#### 469 **Acknowledgements**

470 This publication has been produced within the framework of the EMME-CARE project, which has received funding from  
471 the European Union's Horizon 2020 Research and Innovation Programme (under grant agreement no. 856612) and the  
472 Government of Cyprus. This research has also received funding from the European Union's Horizon 2020 research and  
473 innovation programme under the Marie Skłodowska-Curie grant agreement no. 895875 ("NPF-PANDA") as well as  
474 Academy of Finland Project #325647. The sole responsibility of this publication lies with the author. The European Union  
475 is not responsible for any use that may be made of the information contained therein.





476 **Author contributions**

477 MO, DS, LD analysed the experimental data, MO developed the software, all authors were involved in the scientific  
478 discussion and in the writing of the manuscript.

479 **References**

- 480 Albrecht, B. A.: Aerosols, Cloud Microphysics, and Fractional Cloudiness, *Science*, 245(4923), 1227–1230,  
481 doi:10.1126/science.245.4923.1227, 1989.
- 482 Appel, B. R., Tokiwa, Y., Hsu, J., Kothny, E. L. and Hahn, E.: Visibility as related to atmospheric aerosol constituents,  
483 *Atmos. Environ.*, 19(9), 1525–1534, doi:https://doi.org/10.1016/0004-6981(85)90290-2, 1985.
- 484 Chambolle, A. and Pock, T.: A First-Order Primal-Dual Algorithm for Convex Problems with Applications to Imaging,  
485 *J. Math. Imaging Vis.*, 40(1), 120–145, doi:10.1007/s10851-010-0251-1, 2011.
- 486 Dada, L., Lehtipalo, K., Kontkanen, J., Nieminen, T., Baalbaki, R., Ahonen, L., Duplissy, J., Yan, C., Chu, B., Petäjä, T.,  
487 Lehtinen, K., Kerminen, V.-M., Kulmala, M. and Kangasluoma, J.: Formation and growth of sub-3-nm aerosol particles  
488 in experimental chambers, *Nat. Protoc.*, doi:10.1038/s41596-019-0274-z, 2020.
- 489 Daellenbach, K. R., Uzu, G., Jiang, J., Cassagnes, L.-E., Leni, Z., Vlachou, A., Stefenelli, G., Canonaco, F., Weber, S.,  
490 Segers, A., Kuenen, J. J. P., Schaap, M., Favez, O., Albinet, A., Aksoyoglu, S., Dommen, J., Baltensperger, U., Geiser,  
491 M., El Haddad, I., Jaffrezo, J.-L. and Prévôt, A. S. H.: Sources of particulate-matter air pollution and its oxidative potential  
492 in Europe, *Nature*, 587(7834), 414–419, doi:10.1038/s41586-020-2902-8, 2020.
- 493 Duplissy, J., Merikanto, J., Franchin, A., Tsagkogeorgas, G., Kangasluoma, J., Wimmer, D., Vuollekoski, H.,  
494 Schobesberger, S., Lehtipalo, K., Flagan, R. C., Brus, D., Donahue, N. M., Vehkamäki, H., Almeida, J., Amorim, A.,  
495 Barmet, P., Bianchi, F., Breitenlechner, M., Dunne, E. M., Guida, R., Henschel, H., Junninen, H., Kirkby, J., Kürten, A.,  
496 Kupc, A., Määttä, A., Makhmutov, V., Mathot, S., Nieminen, T., Onnela, A., Praplan, A. P., Riccobono, F., Rondo,  
497 L., Steiner, G., Tome, A., Walther, H., Baltensperger, U., Carslaw, K. S., Dommen, J., Hansel, A., Petäjä, T., Sipilä, M.,  
498 Stratmann, F., Vrtala, A., Wagner, P. E., Worsnop, D. R., Curtius, J. and Kulmala, M.: Effect of ions on sulfuric acid-  
499 water binary particle formation: 2. Experimental data and comparison with QC-normalized classical nucleation theory, *J.*  
500 *Geophys. Res.-Atmos.*, 121(4), 1752–1775, doi:10.1002/2015JD023539, 2016.
- 501 Gelb, A.: Applied optimal estimation, MIT Press, Cambridge, Massachusetts., 1974.
- 502 He, X.-C., Iyer, S., Sipilä, M., Ylisirniö, A., Peltola, M., Kontkanen, J., Baalbaki, R., Simon, M., Kürten, A., Tham, Y.  
503 J., Pesonen, J., Ahonen, L. R., Amanatidis, S., Amorim, A., Baccharini, A., Beck, L., Bianchi, F., Brilke, S., Chen, D.,  
504 Chiu, R., Curtius, J., Dada, L., Dias, A., Dommen, J., Donahue, N. M., Duplissy, J., El Haddad, I., Finkenzeller, H.,  
505 Fischer, L., Heinritzi, M., Hofbauer, V., Kangasluoma, J., Kim, C., Koenig, T. K., Kubečka, J., Kvashnin, A.,  
506 Lamkaddam, H., Lee, C. P., Leiminger, M., Li, Z., Makhmutov, V., Xiao, M., Marten, R., Nie, W., Onnela, A., Partoll,  
507 E., Petäjä, T., Salo, V.-T., Schuchmann, S., Steiner, G., Stolzenburg, D., Stozhkov, Y., Tauber, C., Tomé, A., Väisänen,  
508 O., Vazquez-Pufleau, M., Volkamer, R., Wagner, A. C., Wang, M., Wang, Y., Wimmer, D., Winkler, P. M., Worsnop,  
509 D. R., Wu, Y., Yan, C., Ye, Q., Lehtinen, K., Nieminen, T., Manninen, H. E., Rissanen, M., Schobesberger, S., Lehtipalo,  
510 K., Baltensperger, U., Hansel, A., Kerminen, V.-M., Flagan, R. C., Kirkby, J., Kurtén, T. and Kulmala, M.: Determination  
511 of the collision rate coefficient between charged iodine acid clusters and iodine acid using the appearance time method,  
512 *Aerosol Sci. Tech.*, 1–12, doi:10.1080/02786826.2020.1839013, 2020.
- 513 He, X.-C., Tham, Y. J., Dada, L., Wang, M., Finkenzeller, H., Stolzenburg, D., Iyer, S., Simon, M., Kürten, A., Shen, J.,



514 Rörup, B., Rissanen, M., Schobesberger, S., Baalbaki, R., Wang, D. S., Koenig, T. K., Jokinen, T., Sarnela, N., Beck, L.,  
515 Almeida, J., Amanatidis, S., Amorim, A., Ataei, F., Baccarini, A., Bertozzi, B., Bianchi, F., Brilke, S., Caudillo, L., Chen,  
516 D., Chiu, R., Chu, B., Dias, A., Ding, A., Dommen, J., Duplissy, J., El-Haddad, I., Gonzalez Carracedo, L., Granzin, M.,  
517 Hansel, A., Heinritzi, M., Hofbauer, V., Junninen, H., Kangasluoma, J., Kempainen, D., Kim, C., Kong, W., Krechmer,  
518 J. E., Kvashnin, A., Laitinen, T., Lamkaddam, H., Lee, C. P., Lehtipalo, K., Leiminger, M., Li, Z., Makhmutov, V.,  
519 Manninen, H. E., Marie, G., Marten, R., Mathot, S., Mauldin, R. L., Mentler, B., Möhler, O., Müller, T., Nie, W., Onnela,  
520 A., Petäjä, T., Pfeifer, J., Philippov, M., Ranjithkumar, A., Saiz-Lopez, A., Salma, I., Scholz, W., Schuchmann, S.,  
521 Schulze, B., Steiner, G., Stozhkov, Y., Tauber, C., Tomé, A., Thakur, R. C., Väisänen, O., Vazquez-Pufleau, M., Wagner,  
522 A. C., Wang, Y., Weber, S. K., Winkler, P. M., Wu, Y., Xiao, M., Yan, C., Ye, Q., Ylisirniö, A., Zauner-Wieczorek, M.,  
523 Zha, Q., Zhou, P., Flagan, R. C., Curtius, J., Baltensperger, U., Kulmala, M., Kerminen, V.-M., Kurten, T., et al.: Role of  
524 iodine oxoacids in atmospheric aerosol nucleation, *Science*, 371(6529), 589–595, doi:10.1126/science.abe0298, 2021.

525 Heinritzi, M., Dada, L., Simon, M., Stolzenburg, D., Wagner, A. C., Fischer, L., Ahonen, L. R., Amanatidis, S., Baalbaki,  
526 R., Baccarini, A., Bauer, P. S., Baumgartner, B., Bianchi, F., Brilke, S., Chen, D., Chiu, R., Dias, A., Dommen, J.,  
527 Duplissy, J., Finkenzeller, H., Frege, C., Fuchs, C., Garmash, O., Gordon, H., Granzin, M., El Haddad, I., He, X., Helm,  
528 J., Hofbauer, V., Hoyle, C. R., Kangasluoma, J., Keber, T., Kim, C., Kürten, A., Lamkaddam, H., Laurila, T. M.,  
529 Lampilahti, J., Lee, C. P., Lehtipalo, K., Leiminger, M., Mai, H., Makhmutov, V., Manninen, H. E., Marten, R., Mathot,  
530 S., Mauldin, R. L., Mentler, B., Molteni, U., Müller, T., Nie, W., Nieminen, T., Onnela, A., Partoll, E., Passananti, M.,  
531 Petäjä, T., Pfeifer, J., Pospisilova, V., Quéléver, L. L. J., Rissanen, M. P., Rose, C., Schobesberger, S., Scholz, W.,  
532 Scholze, K., Sipilä, M., Steiner, G., Stozhkov, Y., Tauber, C., Tham, Y. J., Vazquez-Pufleau, M., Virtanen, A., Vogel, A.  
533 L., Volkamer, R., Wagner, R., Wang, M., Weitz, L., Wimmer, D., Xiao, M., Yan, C., Ye, P., Zha, Q., Zhou, X., Amorim,  
534 A., Baltensperger, U., Hansel, A., Kulmala, M., Tomé, A., Winkler, P. M., Worsnop, D. R., Donahue, N. M., Kirkby, J.  
535 and Curtius, J.: Molecular understanding of the suppression of new-particle formation by isoprene, *Atmos. Chem. Phys.*,  
536 20(20), 11809–11821, doi:10.5194/acp-20-11809-2020, 2020.

537 Henze, D. K., Seinfeld, J. H., Liao, W., Sandu, A. and Carmichael, G. R.: Inverse modeling of aerosol dynamics:  
538 Condensational growth, *J. Geophys. Res.-Atmos.*, 109(D14), doi:https://doi.org/10.1029/2004JD004593, 2004.

539 Kaipio, J. and Somersalo, E.: *Statistical and Computational Inverse Problems*, Springer-Verlag, New York., 2005.

540 Kallinger, P. and Szymanski, W. W.: Experimental determination of the steady-state charging probabilities and particle  
541 size conservation in non-radioactive and radioactive bipolar aerosol chargers in the size range of 5–40 nm, *J. Nanopart.*  
542 *Res.*, 17(4), 171, doi:10.1007/s11051-015-2981-x, 2015.

543 Kandlikar, M. and Ramachandran, G.: Inverse methods for analysing aerosol spectrometer measurements: a critical  
544 review, *J. Aerosol Sci.*, 30(4), 413–437, doi:https://doi.org/10.1016/S0021-8502(98)00066-4, 1999.

545 Kangasluoma, J., Cai, R., Jiang, J., Deng, C., Stolzenburg, D., Ahonen, L. R., Chan, T., Fu, Y., Kim, C., Laurila, T. M.,  
546 Zhou, Y., Dada, L., Sulo, J., Flagan, R. C., Kulmala, M., Petäjä, T. and Lehtipalo, K.: Overview of measurements and  
547 current instrumentation for 1–10 nm aerosol particle number size distributions, *J. Aerosol Sci.*, 105584,  
548 doi:https://doi.org/10.1016/j.jaerosci.2020.105584, 2020.

549 Kirkby, J., Curtius, J., Almeida, J., Dunne, E., Duplissy, J., Ehrhart, S., Franchin, A., Gagné, S., Ickes, L., Kürten, A.,  
550 Kupc, A., Metzger, A., Riccobono, F., Rondo, L., Schobesberger, S., Tsagkogeorgas, G., Wimmer, D., Amorim, A.,  
551 Bianchi, F., Breitenlechner, M., David, A., Dommen, J., Downard, A., Ehn, M., Flagan, R. C., Haider, S., Hansel, A.,  
552 Hauser, D., Jud, W., Junninen, H., Kreissl, F., Kvashnin, A., Laaksonen, A., Lehtipalo, K., Lima, J., Lovejoy, E. R.,  
553 Makhmutov, V., Mathot, S., Mikkilä, J., Minginette, P., Mogo, S., Nieminen, T., Onnela, A., Pereira, P., Petäjä, T.,



- 554 Schnitzhofer, R., Seinfeld, J. H., Sipilä, M., Stozhkov, Y., Stratmann, F., Tomé, A., Vanhanen, J., Viisanen, Y., Vrtala,  
555 A., Wagner, P. E., Walther, H., Weingartner, E., Wex, H., Winkler, P. M., Carslaw, K. S., Worsnop, D. R., Baltensperger,  
556 U. and Kulmala, M.: Role of sulphuric acid, ammonia and galactic cosmic rays in atmospheric aerosol nucleation, *Nature*,  
557 476, 429–433, doi:10.1038/nature10343, 2011.
- 558 Kirkby, J., Duplissy, J., Sengupta, K., Frege, C., Gordon, H., Williamson, C., Heinritzi, M., Simon, M., Yan, C., Almeida,  
559 J., Tröstl, J., Nieminen, T., Ortega, I. K., Wagner, R., Adamov, A., Amorim, A., Bernhammer, A.-K., Bianchi, F.,  
560 Breitenlechner, M., Brilke, S., Chen, X., Craven, J., Dias, A., Ehrhart, S., Flagan, R. C., Franchin, A., Fuchs, C., Guida,  
561 R., Hakala, J., Hoyle, C. R., Jokinen, T., Junninen, H., Kangasluoma, J., Kim, J., Krapf, M., Kürten, A., Laaksonen, A.,  
562 Lehtipalo, K., Makhmutov, V., Mathot, S., Molteni, U., Onnela, A., Peräkylä, O., Piel, F., Petäjä, T., Praplan, A. P.,  
563 Pringle, K., Rap, A., Richards, N. A. D., Riipinen, I., Rissanen, M. P., Rondo, L., Sarnela, N., Schobesberger, S., Scott,  
564 C. E., Seinfeld, J. H., Sipilä, M., Steiner, G., Stozhkov, Y., Stratmann, F., Tomé, A., Virtanen, A., Vogel, A. L., Wagner,  
565 A. C., Wagner, P. E., Weingartner, E., Wimmer, D., Winkler, P. M., Ye, P., Zhang, X., Hansel, A., Dommen, J., Donahue,  
566 N. M., Worsnop, D. R., Baltensperger, U., Kulmala, M., Carslaw, K. S. and Curtius, J.: Ion-induced nucleation of pure  
567 biogenic particles, *Nature*, 533, 521–526, doi:10.1038/nature17953, 2016.
- 568 Kuang, C., Chen, M., Zhao, J., Smith, J., McMurry, P. H. and Wang, J.: Size and time-resolved growth rate measurements  
569 of 1 to 5 nm freshly formed atmospheric nuclei, *Atmos. Chem. Phys.*, 12(7), 3573–3589, doi:10.5194/acp-12-3573-2012,  
570 2012.
- 571 Kulmala, M., Petäjä, T., Nieminen, T., Sipilä, M., Manninen, H. E., Lehtipalo, K., Dal Maso, M., Aalto, P. P., Junninen,  
572 H., Paasonen, P., Riipinen, I., Lehtinen, K. E. J., Laaksonen, A. and Kerminen, V.-M.: Measurement of the nucleation of  
573 atmospheric aerosol particles, *Nat. Protoc.*, 7(9), 1651–1667, doi:10.1038/nprot.2012.091, 2012.
- 574 Kupiainen-Määttä, O.: A Monte Carlo approach for determining cluster evaporation rates from concentration  
575 measurements, *Atmos. Chem. Phys.*, 16(22), 14585–14598, doi:10.5194/acp-16-14585-2016, 2016.
- 576 Kürten, A., Li, C., Bianchi, F., Curtius, J., Dias, A., Donahue, N. M., Duplissy, J., Flagan, R. C., Hakala, J., Jokinen, T.,  
577 Kirkby, J., Kulmala, M., Laaksonen, A., Lehtipalo, K., Makhmutov, V., Onnela, A., Rissanen, M. P., Simon, M., Sipilä,  
578 M., Stozhkov, Y., Tröstl, J., Ye, P. and McMurry, P. H.: New particle formation in the sulfuric acid–dimethylamine–  
579 water system: reevaluation of CLOUD chamber measurements and comparison to an aerosol nucleation and growth  
580 model, *Atmos. Chem. Phys.*, 18(2), 845–863, doi:10.5194/acp-18-845-2018, 2018.
- 581 Lehtinen, K. E. J. and Kulmala, M.: A model for particle formation and growth in the atmosphere with molecular  
582 resolution in size, *Atmos. Chem. Phys.*, 3(1), 251–257, doi:10.5194/acp-3-251-2003, 2003.
- 583 Lehtinen, K. E. J., Rannik, Ü., Petäjä, T., Kulmala, M. and Hari, P.: Nucleation rate and vapor concentration estimations  
584 using a least squares aerosol dynamics method, *J. Geophys. Res.-Atmos.*, 109(D21), doi:10.1029/2004JD004893, 2004.
- 585 McGuffin, D. L., Huang, Y., Flagan, R. C., Petäjä, T., Ydstie, B. E. and Adams, P. J.: Novel estimation of aerosol  
586 processes with particle size distribution measurements: a case study with TOMAS algorithm, *Geosci. Model Dev.*  
587 *Discuss.*, 2020, 1–27, doi:10.5194/gmd-2020-281, 2020.
- 588 Nieminen, T., Lehtinen, K. E. J. and Kulmala, M.: Sub-10 nm particle growth by vapor condensation – effects of vapor  
589 molecule size and particle thermal speed, *Atmos. Chem. Phys.*, 10(20), 9773–9779, doi:10.5194/acp-10-9773-2010, 2010.
- 590 Nieminen, T., Kerminen, V.-M., Petäjä, T., Aalto, P. P., Arshinov, M., Asmi, E., Baltensperger, U., Beddows, D. C. S.,  
591 Beukes, J. P., Collins, D., Ding, A., Harrison, R. M., Henzing, B., Hooda, R., Hu, M., Hörrak, U., Kivekäs, N., Komsaare,  
592 K., Krejci, R., Kristensson, A., Laakso, L., Laaksonen, A., Leitch, W. R., Lihavainen, H., Mihalopoulos, N., Németh,  
593 Z., Nie, W., O’Dowd, C., Salma, I., Sellegri, K., Svenningsson, B., Swietlicki, E., Tunved, P., Ulevicius, V., Vakkari, V.,



- 594 Vana, M., Wiedensohler, A., Wu, Z., Virtanen, A. and Kulmala, M.: Global analysis of continental boundary layer new  
595 particle formation based on long-term measurements, *Atmos. Chem. Phys.*, 18(19), 14737–14756, doi:10.5194/acp-18-  
596 14737-2018, 2018.
- 597 Ozon, M., Seppänen, A., Kaipio, J. P. and Lehtinen, K. E. J.: Retrieval of process rate parameters in the general dynamic  
598 equation for aerosols using Bayesian state estimation, *Geosci. Model Dev. Discuss.*, 2020, 1–32, doi:10.5194/gmd-2020-  
599 236, 2020.
- 600 Pichelstorfer, L., Stolzenburg, D., Ortega, J., Karl, T., Kokkola, H., Laakso, A., Lehtinen, K. E. J., Smith, J. N., McMurry,  
601 P. H. and Winkler, P. M.: Resolving nanoparticle growth mechanisms from size- and time-dependent growth rate analysis,  
602 *Atmos. Chem. Phys.*, 18(2), 1307–1323, doi:10.5194/acp-18-1307-2018, 2018.
- 603 Pope, C. A. and Dockery, D. W.: Health Effects of Fine Particulate Air Pollution: Lines that Connect, *J. Air Waste*  
604 *Manag.*, 56(6), 709–742, doi:10.1080/10473289.2006.10464485, 2006.
- 605 Sandu, A., Liao, W., Carmichael, G. R., Henze, D. K. and Seinfeld, J. H.: Inverse Modeling of Aerosol Dynamics Using  
606 Adjoint: Theoretical and Numerical Considerations, *Aerosol Sci. Tech.*, 39(8), 677–694,  
607 doi:10.1080/02786820500182289, 2005.
- 608 Seinfeld, J. and Pandis, S.: *Atmospheric Chemistry and Physics: From Air Pollution to Climate Change*, 3rd edition, 3rd  
609 edition, Wiley., 2016.
- 610 Shcherbacheva, A., Balehowsky, T., Kubečka, J., Olenius, T., Helin, T., Haario, H., Laine, M., Kurtén, T. and Vehkamäki,  
611 H.: Identification of molecular cluster evaporation rates, cluster formation enthalpies and entropies by Monte Carlo  
612 method, *Atmos. Chem. Phys. Discuss.*, 2020, 1–58, doi:10.5194/acp-2019-1036, 2020.
- 613 Shen, J., Shu, C. and Zhang, M.: High Resolution Schemes for a Hierarchical Size-Structured Model, *SIAM J. Numer.*  
614 *Anal.*, 45(1), 352–370, doi:10.1137/050638126, 2007.
- 615 Simon, M., Dada, L., Heinritzi, M., Scholz, W., Stolzenburg, D., Fischer, L., Wagner, A. C., Kürten, A., Rörup, B., He,  
616 X.-C., Almeida, J., Baalbaki, R., Baccarini, A., Bauer, P. S., Beck, L., Bergen, A., Bianchi, F., Bräkling, S., Brilke, S.,  
617 Caudillo, L., Chen, D., Chu, B., Dias, A., Draper, D. C., Duplissy, J., El-Haddad, I., Finkenzeller, H., Frege, C., Gonzalez-  
618 Carracedo, L., Gordon, H., Granzin, M., Hakala, J., Hofbauer, V., Hoyle, C. R., Kim, C., Kong, W., Lamkaddam, H.,  
619 Lee, C. P., Lehtipalo, K., Leiminger, M., Mai, H., Manninen, H. E., Marie, G., Marten, R., Mentler, B., Molteni, U.,  
620 Nichman, L., Nie, W., Ojdanic, A., Onnela, A., Partoll, E., Petäjä, T., Pfeifer, J., Philippov, M., Quéléver, L. L. J.,  
621 Ranjithkumar, A., Rissanen, M. P., Schallhart, S., Schobesberger, S., Schuchmann, S., Shen, J., Sipilä, M., Steiner, G.,  
622 Stozhkov, Y., Tauber, C., Tham, Y. J., Tomé, A. R., Vazquez-Pufleau, M., Vogel, A. L., Wagner, R., Wang, M., Wang,  
623 D. S., Wang, Y., Weber, S. K., Wu, Y., Xiao, M., Yan, C., Ye, P., Ye, Q., Zauner-Wieczorek, M., Zhou, X., Baltensperger,  
624 U., Dommen, J., Flagan, R. C., Hansel, A., Kulmala, M., Volkamer, R., Winkler, P. M., Worsnop, D. R., Donahue, N.  
625 M., Kirkby, J. and Curtius, J.: Molecular understanding of new-particle formation from alpha-pinene between -50 and  
626 +25 °C, *Atmos. Chem. Phys.*, 20(15), 9183–9207, doi:10.5194/acp-20-9183-2020, 2020.
- 627 Sipilä, M., Sarnela, N., Jokinen, T., Henschel, H., Junninen, H., Kontkanen, J., Richters, S., Kangasluoma, J., Franchin,  
628 A., Peräkylä, O., Rissanen, M. P., Ehn, M., Vehkamäki, H., Kurten, T., Berndt, T., Petäjä, T., Worsnop, D., Ceburnis, D.,  
629 Kerminen, V.-M., Kulmala, M. and O’Dowd, C.: Molecular-scale evidence of aerosol particle formation via sequential  
630 addition of HIO<sub>3</sub>, *Nature*, 537, 532, doi:10.1038/nature19314, 2016.
- 631 Smolarkiewicz, P. K.: A fully multidimensional positive definite advection transport algorithm with small implicit  
632 diffusion, *J. Comput. Phys.*, 54(2), 325–362, doi:https://doi.org/10.1016/0021-9991(84)90121-9, 1984.
- 633 Stolzenburg, D., Steiner, G. and Winkler, P. M.: A DMA-train for precision measurement of sub-10 nm aerosol dynamics,



- 634 Atmos. Meas. Tech., 10(4), 1639–1651, doi:10.5194/amt-10-1639-2017, 2017.
- 635 Stolzenburg, D., Fischer, L., Vogel, A. L., Heinritzi, M., Schervish, M., Simon, M., Wagner, A. C., Dada, L., Ahonen, L.  
636 R., Amorim, A., Baccarini, A., Bauer, P. S., Baumgartner, B., Bergen, A., Bianchi, F., Breitenlechner, M., Brilke, S.,  
637 Buenrostro Mazon, S., Chen, D., Dias, A., Draper, D. C., Duplissy, J., El Haddad, I., Finkenzeller, H., Frege, C., Fuchs,  
638 C., Garmash, O., Gordon, H., He, X., Helm, J., Hofbauer, V., Hoyle, C. R., Kim, C., Kirkby, J., Kontkanen, J., Kürten,  
639 A., Lampilahti, J., Lawler, M., Lehtipalo, K., Leiminger, M., Mai, H., Mathot, S., Mentler, B., Molteni, U., Nie, W.,  
640 Nieminen, T., Nowak, J. B., Ojdanic, A., Onnela, A., Passananti, M., Petäjä, T., Quéléver, L. L. J., Rissanen, M. P.,  
641 Sarnela, N., Schallhart, S., Tauber, C., Tomé, A., Wagner, R., Wang, M., Weitz, L., Wimmer, D., Xiao, M., Yan, C., Ye,  
642 P., Zha, Q., Baltensperger, U., Curtius, J., Dommen, J., Flagan, R. C., Kulmala, M., Smith, J. N., Worsnop, D. R., Hansel,  
643 A., Donahue, N. M. and Winkler, P. M.: Rapid growth of organic aerosol nanoparticles over a wide tropospheric  
644 temperature range, P. Nat. Acad. Sci. USA, 115(37), 9122–9127, doi:10.1073/pnas.1807604115, 2018.
- 645 Stolzenburg, D., Simon, M., Ranjithkumar, A., Kürten, A., Lehtipalo, K., Gordon, H., Ehrhart, S., Finkenzeller, H.,  
646 Pichelstorfer, L., Nieminen, T., He, X.-C., Brilke, S., Xiao, M., Amorim, A., Baalbaki, R., Baccarini, A., Beck, L.,  
647 Bräkling, S., Caudillo Murillo, L., Chen, D., Chu, B., Dada, L., Dias, A., Dommen, J., Duplissy, J., El Haddad, I., Fischer,  
648 L., Gonzalez Carracedo, L., Heinritzi, M., Kim, C., Koenig, T. K., Kong, W., Lamkaddam, H., Lee, C. P., Leiminger, M.,  
649 Li, Z., Makhmutov, V., Manninen, H. E., Marie, G., Marten, R., Müller, T., Nie, W., Partoll, E., Petäjä, T., Pfeifer, J.,  
650 Philippov, M., Rissanen, M. P., Rörup, B., Schobesberger, S., Schuchmann, S., Shen, J., Sipilä, M., Steiner, G., Stozhkov,  
651 Y., Tauber, C., Tham, Y. J., Tomé, A., Vazquez-Pufleau, M., Wagner, A. C., Wang, M., Wang, Y., Weber, S. K., Wimmer,  
652 D., Wlasits, P. J., Wu, Y., Ye, Q., Zauner-Wieczorek, M., Baltensperger, U., Carslaw, K. S., Curtius, J., Donahue, N. M.,  
653 Flagan, R. C., Hansel, A., Kulmala, M., Lelieveld, J., Volkamer, R., Kirkby, J. and Winkler, P. M.: Enhanced growth rate  
654 of atmospheric particles from sulfuric acid, Atmos. Chem. Phys., 20(12), 7359–7372, doi:10.5194/acp-20-7359-2020,  
655 2020.
- 656 Stolzenburg, M. R. and McMurry, P. H.: Equations Governing Single and Tandem DMA Configurations and a New  
657 Lognormal Approximation to the Transfer Function, Aerosol Sci. Tech., 42, 421–432, doi:10.1080/02786820802157823,  
658 2008.
- 659 Tauber, C., Schmoll, D., Gruenwald, J., Brilke, S., Wlasits, P. J., Winkler, P. M. and Wimmer, D.: Characterization of a  
660 Non-Thermal Plasma Source for the Use as a Mass Spec Calibration Tool and Non-Radioactive Aerosol Charger, Atmos.  
661 Meas. Tech., 13(11), 5993–6006, doi:10.5194/amt-13-5993-2020, 2020.
- 662 Tröstl, J., Chuang, W. K., Gordon, H., Heinritzi, M., Yan, C., Molteni, U., Ahlm, L., Frege, C., Bianchi, F., Wagner, R.,  
663 Simon, M., Lehtipalo, K., Williamson, C., Craven, J. S., Duplissy, J., Adamov, A., Almeida, J., Bernhammer, A.-K.,  
664 Breitenlechner, M., Brilke, S., Dias, A., Ehrhart, S., Flagan, R. C., Franchin, A., Fuchs, C., Guida, R., Gysel, M., Hansel,  
665 A., Hoyle, C. R., Jokinen, T., Junninen, H., Kangasluoma, J., Keskinen, H., Kim, J., Krapf, M., Kürten, A., Laaksonen,  
666 A., Lawler, M., Leiminger, M., Mathot, S., Möhler, O., Nieminen, T., Onnela, A., Petäjä, T., Piel, F. M., Miettinen, P.,  
667 Rissanen, M. P., Rondo, L., Sarnela, N., Schobesberger, S., Sengupta, K., Sipilä, M., Smith, J. N., Steiner, G., Tomé, A.,  
668 Virtanen, A., Wagner, A. C., Weingartner, E., Wimmer, D., Winkler, P. M., Ye, P., Carslaw, K. S., Curtius, J., Dommen,  
669 J., Kirkby, J., Kulmala, M., Riipinen, I., Worsnop, D. R., Donahue, N. M. and Baltensperger, U.: The role of low-volatility  
670 organic compounds in initial particle growth in the atmosphere, Nature, 533, 527–531, doi:10.1038/nature18271, 2016.
- 671 Twomey, S.: Pollution and the planetary albedo, Atmos. Environ., 8(12), 1251–1256, doi:https://doi.org/10.1016/0004-  
672 6981(74)90004-3, 1974.
- 673 Verheggen, B. and Mozurkewich, M.: An inverse modeling procedure to determine particle growth and nucleation rates



674 from measured aerosol size distributions, *Atmos. Chem. Phys.*, 6(10), 2927–2942, doi:10.5194/acp-6-2927-2006, 2006.  
675 Viskari, T., Asmi, E., Kolmonen, P., Vuollekoski, H., Petäjä, T. and Järvinen, H.: Estimation of aerosol particle number  
676 distributions with Kalman Filtering – Part 1: Theory, general aspects and statistical validity, *Atmos. Chem. Phys.*, 12(24),  
677 11767–11779, doi:10.5194/acp-12-11767-2012, 2012.  
678 Wiedensohler, A.: An approximation of the bipolar charge distribution for particles in the submicron size range, *J. Aerosol*  
679 *Sci.*, 19(3), 387–389, doi:10.1016/0021-8502(88)90278-9, 1988.  
680 Wlasits, P. J., Stolzenburg, D., Tauber, C., Brilke, S., Schmitt, S. H., Winkler, P. M. and Wimmer, D.: Counting on  
681 chemistry: laboratory evaluation of seed-material-dependent detection efficiencies of ultrafine condensation particle  
682 counters, *Atmos. Meas. Tech.*, 13(7), 3787–3798, doi:10.5194/amt-13-3787-2020, 2020.  
683 Wolfenbarger, J. K. and Seinfeld, J. H.: Inversion of aerosol size distribution data, *J. Aerosol Sci.*, 21(2), 227–247,  
684 doi:[https://doi.org/10.1016/0021-8502\(90\)90007-K](https://doi.org/10.1016/0021-8502(90)90007-K), 1990.  
685