



1 **Secondary Organic Aerosols from OH Oxidation of Cyclic Volatile Methyl Siloxanes as an Important**
2 **Si Source in the Atmosphere**

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12 **Short summary:** We presented yields and compositions of Si-containing SOA generated from the reaction
13 of cVMS (D3-D6) with OH radicals. NO_x played negative roles on cVMS SOA formation, while ammonium
14 sulfate seeds enhanced D3-D5 SOA yields at short photochemical ages under high-NO_x conditions. The
15 aerosol mass spectra confirmed that the components of cVMS SOA significantly relied on OH exposure. A
16 global cVMS-derived SOA source strength was estimated to understand SOA formation potentials of cVMS.

17

18 **Abstract**

19 Cyclic volatile methyl siloxanes (cVMS) are active ingredients in widely used consumer products, which
20 can volatilize into the atmosphere, thus attracting much attention due to their potential environmental risks.
21 While in the atmosphere the cVMS undergo oxidation yielding both gaseous and particulate products. The
22 aerosol yields and compositions from the OH oxidation of four cVMS (D3-D6) were determined under low



23 and high-NO_x conditions in an oxidation flow reactor. The aerosol yields progressively increased from D3
24 to D6, consistent with the volatilities and molecule weights of these cVMS. NO_x can restrict the formation
25 of SOA, leading to lower SOA yields under high-NO_x conditions than under low-NO_x conditions, with a
26 yield decrease between 0.05-0.30 depending on the cVMS. Ammonium sulfate seeds exhibited minor
27 impacts on SOA yields under low-NO_x conditions, but significantly increased the SOA yields in the oxidation
28 of D3-D5 at short photochemical ages under high-NO_x conditions. The mass spectra of the SOA showed a
29 dependence of its chemical compositions on OH exposure. At high exposures, equivalent to photochemical
30 ages of >6 days in the atmosphere, D4-D6 SOA mainly consisted of C_xH_y and C_xH_yO_zSi_n under low-NO_x
31 conditions, whereas they primarily contained N_mO_z, C_xH_y, C_xH_yO₁, C_xH_yO_{>1} and C_xH_yO_zSi_n under high-NO_x
32 conditions. Using the yield data from the present study and reported cVMS annual production, a global
33 cVMS-derived SOA source strength is estimated to be 0.16 Tg yr⁻¹, distributed over major urban centers.

34

35 **1 Introduction**

36 Secondary organic aerosols (SOA), which contribute 50-85% to the mass of atmospheric organic aerosols
37 (OA) (Glasius and Goldstein, 2016), are mainly formed via the partitioning of low volatility products from
38 oxidation of volatile organic compounds (VOCs), semi- and intermediate volatile organic
39 compounds(S/IVOCs) (Riipinen et al., 2012). SOA has attracted significant attention due to their important
40 impacts on climate, ecosystems and human health (Berndt et al., 2016). Global budgets of SOA remain an
41 unresolved issue despite extensive research, largely due to uncertainties associated with aerosol yields and
42 the presence of unconsidered SOA precursors.

43 As one type of anthropogenic VOC and potential SOA precursors, cyclic volatile methyl siloxanes (cVMS)
44 are widely used in industrial applications and personal care products (Genualdi et al., 2011; Krogseth et al.,



45 2013a). cVMS have been classified as high-volume chemicals with an annual production of millions of tons
46 globally (Rücker and Kümmerer, 2015; Ahrens et al., 2014). Studies of cVMS in the environment have
47 focused on investigating health and environmental impacts particularly due to their potential persistence,
48 bioaccumulation and toxicity (Guo et al., 2019; Liu et al., 2018; Farasani and Darbre, 2017; Xu et al., 2019;
49 Kim et al., 2018; Coggon et al., 2018). As a result, the European Council has proposed a restriction on the
50 octamethylcyclotetrasiloxane (D4) and decamethylcyclopentasiloxane (D5) content in wash-off personal
51 care products to a limit of 0.1 mass% by 2020. The legislative actions notwithstanding, knowledge of
52 environmental behavior of cVMS remains surprisingly scarce as compared to their applications and
53 economic significance (Rücker and Kümmerer, 2015).

54 It has been estimated that approximately 90% of cVMS are emitted into the atmosphere due to their high
55 saturation vapor pressures (Allen et al., 1997). Gas-phase cVMS have been observed in both indoor and
56 outdoor air. Tang et al. (2015) reported that cVMS accounted for about one third of total VOC mass
57 concentration in a classroom. Outdoor air concentrations of cVMS have also been measured at different sites
58 worldwide (Li et al., 2020; Wang et al., 2018; Rauert et al., 2018), increasing from rural to urban sites and
59 consistent with increasing population density (Rücker and Kümmerer, 2015). For example, at a rural site in
60 Sweden, the concentration of hexamethylcyclotrisiloxane (D3), D4, D5 and dodecamethylcyclohexasiloxane
61 (D6) were 0.94, 3.5, 13 and 1 ng/m³, respectively (Kierkegaard and McLachlan, 2013), while they were 18,
62 55, 172 and 14 ng/m³ in urban areas of Toronto in Canada, respectively (Genualdi et al., 2011; Rauert et al.,
63 2018). cVMS have also been detected in the remote Arctic atmosphere, confirming their long-range transport
64 (Genualdi et al., 2011; Krogseth et al., 2013b). Atmospheric half-lives of cVMS are on the order of 30, 15,
65 10 and 1.6 days for D3-D6, respectively, which allow cVMS to exhibit a hemispherical distribution in the
66 atmosphere (Whelan et al., 2004; Allen et al., 1997; Yucuis et al., 2013; Xiao et al., 2015; Macleod et al.,



67 2013). These lifetimes are driven mostly by reactions with the OH radicals (Xiao et al., 2015; Wang et al.,
68 2013), which generate silanols and dimeric products that can be partitioned to condensed phases (Coggon et
69 al., 2018; Sommerlade, 1993; Wu and Johnston, 2016). The loss of cVMS in the atmosphere is slight through
70 O₃ and NO₃ due to their small reaction rates, and by Cl atoms on account of its low concentration (Atkinson,
71 1991; Alton and Browne, 2020).

72 It has been demonstrated that elemental Si is a frequent constituent of nanoparticles in rural and urban
73 areas (Phares et al., 2003; Rhoads, 2003; Bein, 2005; Bzdek et al., 2014) and in remote regions (Li and
74 Winchester, 1990; Li and Winchester, 1993). These Si-containing nanoparticles have previously been
75 attributed to ore smelting processes, but recent studies have shown that Si-containing species are one of the
76 main components in cVMS SOA, suggesting that the oxidation of cVMS may be an important source of Si
77 in atmospheric aerosols (Wu and Johnston, 2016, 2017). In a modeling study, the oxidation products of
78 cVMS (D4, D5 and D6) were considered to quantify the maximum potential for aerosol formation through
79 reactions with the OH radicals (Janecek et al., 2017). Chandramouli and Kamens (2001) demonstrated the
80 gas-particle partitioning of silanols from D5 oxidation by the OH radicals. Wu and Johnston (2016, 2017)
81 analyzed the chemical composition of secondary aerosols from OH oxidation of D4 and D5, showing a large
82 number of monomeric and dimeric products. Janecek et al. (2017, 2019) reported physical properties of
83 SOA generated by OH oxidation of D5, including hygroscopicity, cloud seeding potential and volatility.
84 Charan et al. (2021) measured SOA yields of D5 using chambers and flow tube reactors, indicating necessary
85 conditions when extrapolating SOA yields to the ambient atmosphere. These studies mainly focused on D5
86 and occasionally D4 but rarely others. To better understand the SOA-forming potentials of typical cVMS in
87 the atmosphere, accurate yields and molecular compositions of SOA from the oxidation of cVMS under
88 various atmospheric conditions are needed.



89 In this work, the formation of SOA from the oxidation of four cVMS (D3-D6) by OH radical was
90 investigated in an oxidation flow reactor (OFR). Under various combinations of NO_x and ammonium sulfate
91 seed concentrations, the yields and compositions of SOA formed from the oxidation were measured using a
92 suite of instruments including a scanning mobility particle sizer (SMPS), a proton transfer reaction time of
93 flight mass spectrometer (PTR-ToF-MS) and an aerosol mass spectrometer (AMS). Based on these SOA
94 yields, the contribution of cVMS to SOA in the global atmosphere was estimated using reported cVMS
95 concentrations. The results obtained here can largely improve our understanding of the contribution and
96 composition of SOA from cVMS.

97 **2 Experiments and methods**

98 **2.1 Photo-oxidation experiments**

99 The reactions of cVMS with OH radicals were controlled at a constant temperature (21±1°C) and relative
100 humidity (35%±2%) in a custom-made oxidation flow reactor (the Environment and Climate Change Canada
101 oxidation flow reactor, ECCC-OFR), which is shown in Fig. S1 of the Supplement and has been described
102 in detail previously (Li et al., 2019a). Briefly, the ECCC-OFR is a fused quartz cylinder (length: 50.8 cm,
103 inner diameter: 20.3 cm) equipped with a conical inlet and 7 outlets. Wall losses of particles and gases in the
104 ECCC-OFR have been shown to be lower than in other OFRs (Huang et al., 2017; Lambe et al., 2011;
105 Simonen et al., 2017; Li et al., 2019a). The length and full angle of cone inlet are 35.6 cm and 30°,
106 respectively, designed to minimize the formation of jetting and recirculation in the OFR. The outlet at the
107 reactor center is a stainless-steel sampling port (inner diameter: 0.18 in) extending 12.7 cm long into the
108 ECCC-OFR. This sampling inlet reduces the impact of potential turbulent eddies caused by the back end of
109 the reactor. The remaining 6 outlets around the perimeter are designed to allow side flows to pass through
110 the OFR as a sheath flow, indirectly reducing wall losses of gases and particles inside the OFR upon sampling.



111 Ozone-free mercury UV lamps for generating OH radicals are housed in small quartz tubes around and in
112 parallel to the quartz reaction cylinder, and a large flow of air through each of these smaller quartz tubes is
113 used to remove the heat produced by the lamps. The relative humidity was adjusted by controlling the ratio
114 of dry air to wet air into the reactor, and was measured using a humidity sensor (Vaisala) at one of the sheath
115 flow outlets (side flows) of the reactor. The volume of the entire ECCC-OFR is about 16 L and the total flow
116 rate is 8 L min⁻¹, leading to a residence time of 2 min in the OFR.

117 OH radicals were produced through the reaction of water vapor with O(¹D) formed from O₃ photolysis at
118 254 nm. The OH concentration in the ECCC-OFR was regulated by controlling the input voltage and the
119 number of UV lamps. Methanol vapor, introduced into the ECCC-OFR through a bubbler containing
120 methanol solution, was used to determine the OH exposure (i.e., photochemical age) by tracking its decay
121 in the reactor from the reaction with the OH. The decay, or fractional loss, of gas-phase methanol,
122 [MeOH]/[MeOH]₀ was measured with the PTR-ToF-MS, and was used to calculate the OH concentration
123 via Equation 1,

$$124 \quad [\text{OH}] = -\frac{1}{k_{\text{MeOH}}} \ln \frac{[\text{MeOH}]}{[\text{MeOH}]_0} \quad (1)$$

125 where k_{MeOH} is the second-order rate constant of methanol reaction with OH at 298 K (9.4×10^{-13} cm³
126 molecule⁻¹ s⁻¹). It was noted that the OH exposure measurement was offline, because methanol can affect the
127 OH reactivity with cVMS. Under low and high-NO_x conditions described below, the OH exposure varied in
128 the range of 5.5×10^{10} - 1.8×10^{12} and 8.2×10^{10} - 1.1×10^{12} molecules cm⁻³ s, respectively. They correspond
129 to 0.4-14.2 and 0.6-8.5 equivalent days (photochemical age), respectively, assuming that an average OH
130 concentration in air is 1.5×10^6 molecules cm⁻³ (Mao et al., 2009).

131 Pure D3-D6 cVMS compounds (solid D3 and liquid D4, D5 and D6) were placed in a glass U-type tube
132 and maintained at the room temperature. Vapors from these compounds (Table S1 of the Supplement) were



133 separately introduced into the ECCC-OFR by a small flow of zero air (1-18 mL min⁻¹) passing over the
134 headspace of the U-tube containing the pure compounds. The concentrations of D3-D6 in the ECCC-OFR
135 ranged from 20 to 40 ppb, depending on their volatilities. To study the influence of existing particles on the
136 SOA formation, ammonium sulfate (AS) seed particles were produced using an atomizer, dried by a diffusion
137 dryer and neutralized by a neutralizer and injected into the reactor without size selection. The mass
138 concentration of AS seed particles was approximately 30 µg m⁻³ for all experiments.

139 N₂O was used as a source of NO to achieve high-NO_x conditions (Lambe et al., 2017). NO_x conditions
140 were defined by the fate of peroxy radicals (RO₂), which was described by the reaction rate ratio (R_{NO}) of
141 RO₂ + NO and RO₂ + HO₂ (Peng et al., 2017). The R_{NO} ratio increases with increasing OH exposures at a
142 constant concentration of N₂O (Li et al., 2019b). To achieve a constant branching ratio during the high-NO_x
143 experiments, the initial N₂O concentration in the OFR was varied (1.6%-8.0%) to maintain an R_{NO} value of
144 20 (Li et al., 2019b), as calculated using a model (OFR Exposures Estimator v3.1) (Peng et al., 2017). A
145 ratio of $R_{NO}=20$ indicates that 95% of RO₂ reacts with NO, ensuring the dominance of RO₂ + NO, which
146 represents conditions that are relevant for urban atmosphere (Peng et al., 2019). Under low-NO_x conditions,
147 N₂O was not introduced into the OFR, where the RO₂ radicals dominantly interacted with HO₂ radicals,
148 representing atmospheric scenarios with few NO_x sources.

149 2.2 Characterization and analysis

150 The concentrations of cVMS in the OFR were measured online with a proton transfer reaction time of
151 flight mass spectrometer (PTR-ToF-MS, Ionicon Analytik GmbH) (Liggio et al., 2016). The number and
152 mass size distribution of aerosols was monitored using a scanning mobility particle sizer (SMPS, TSI). The
153 mass spectra and elemental composition of aerosols was determined with a high-resolution time-of-flight
154 aerosol mass spectrometer (HR-ToF-AMS, Aerodyne) and analyzed with the AMS analysis software Squirrel



155 (version 1.62G) and Pika (version 1.22G).

156 SOA mass yields (Y) were calculated via Equation 2,

$$157 \quad Y = \frac{\Delta C_{\text{SOA}}}{\Delta C_{\text{cVMS}}} \quad (2)$$

158 where ΔC_{SOA} and ΔC_{cVMS} are the mass concentrations of SOA formed and cVMS lost, respectively. The mass

159 concentration of SOA was determined by multiplying the effective aerosol density by the integrated SOA

160 volume concentration from the SMPS, subtracting the AS seed volume for experiments with AS seeds. The

161 effective aerosol density (ρ) was calculated for unseeded experiments through the following Equation 3

162 (Lambe et al., 2015),

$$163 \quad \rho = \frac{D_{va}}{D_m} \quad (3)$$

164 where D_{va} is the vacuum aerodynamic diameter obtained from the HR-ToF-AMS, and D_m is the electric

165 mobility diameter measured by the SMPS. The ρ varied in the range of 1.6-1.8 depending on the cVMS. The

166 same ρ value was also used in the seeded experiments.

167 **3 Results and discussion**

168 **3.1 SOA yields**

169 Figure 1 shows the SOA yields from the photooxidation of the D3-D6 cVMS under low and high-NO_x

170 conditions as a function of photochemical age (PA), i.e., time-integrated OH exposure, with and without AS

171 seed particles. SOA yields have been widely used to estimate the potential of precursors to produce aerosol

172 mass (McFiggans et al., 2019; Li et al., 2019a; Bruns et al., 2015; Lambe et al., 2015). As shown in Fig. 1,

173 the cVMS SOA yields exhibited an overall increasing trend with PA, expressed in equivalent photochemical

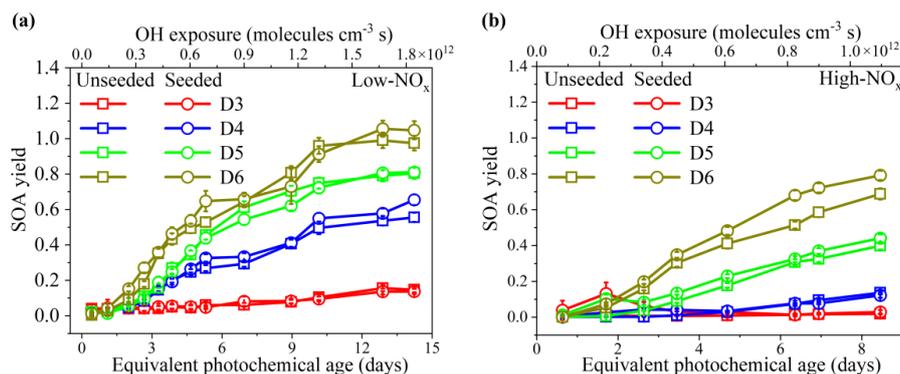
174 days, which agreed with the trend of D5 SOA yields reported by Janecek et al. (2019). Under low-NO_x

175 conditions (Fig. 1a), SOA yields exhibited a slow growth, reaching a plateau after 10 equivalent days. This

176 may be due to increased gas-phase fragmentation of cVMS to generate some higher volatility products,



177 leading to a small increasing amplitude of partition ratio of species into SOA at longer photochemical ages.



178

179 Figure 1. SOA yields from unseeded and seeded ($30 \mu\text{g m}^{-3}$) photooxidation of cVMS by OH radicals. **(a)**

180

low-NO_x experiments; **(b)** high-NO_x experiments.

181

182 For the unseeded and low-NO_x experiments in Fig. 1a, SOA yields of four cVMS exhibited significant

183 differences in values over the same number of equivalent days. The SOA yields successively increased from

184 D3 to D6, consistent with the volatilities and molecular masses of the cVMS as well as their reaction rate

185 constants with the OH radical (Alton and Browne, 2020; Kim and Xu, 2017; Safron et al., 2015). The

186 maximum SOA yields of D3-D6 were (0.16 ± 0.02) , (0.56 ± 0.03) , (0.80 ± 0.03) and (0.99 ± 0.04) , respectively,

187 occurring after a PA of 12 equivalent days. It has been reported that D5 SOA yields varied in the range of

188 0.08-0.50 (Janecek et al., 2019; Wu and Johnston, 2017). Under a similar OH exposure ($1.6\text{-}1.7 \times 10^{12}$

189 molecules $\text{cm}^{-3} \text{ s}$), the D5 SOA yield (0.79 ± 0.03) obtained here is considerably higher than that (0.22)

190 measured by Janecek et al. (2019), which may be attributed to differences in experimental conditions such

191 as differences in wall losses, SOA measurement methods, RHs, and D5 concentrations (Table S2). Although

192 the amount of cVMS lost was variable, cVMS SOA yields positively depended on SOA mass concentrations

193 (Fig. S2), and this trend was observed in previous D5 SOA experiments with OH oxidation (Wu and Johnston,

194 2017).



195 As shown in Fig. 1b, the order of SOA yields from the four cVMS under high-NO_x conditions was the
196 same as that under low-NO_x conditions. However, the SOA yields under high-NO_x conditions were generally
197 smaller than the corresponding yields at similar OH exposures under low-NO_x conditions, with a decrease
198 of 0.05-0.30 depending on the cVMS. Such a reduction suggests that NO_x can restrict the formation of cVMS
199 SOA. NO_x has been shown to reduce SOA yields for some anthropogenic alkanes (Li et al., 2019b), aromatics
200 (Ng et al., 2007a; Chan et al., 2009; Zhou et al., 2019), monoterpenes (Zhao et al., 2018) and other terpenes
201 (Ng et al., 2007b), attributable to the formation of higher volatility products (e.g., organic nitrates) generated
202 by RO₂ + NO compared to RO₂ + HO₂ (Presto et al., 2005; Li et al., 2019b), which is also likely the case
203 here. The higher volatility products favor partitioning in the gas phase, thus reducing the potential for
204 forming SOA (Zhou et al., 2019). Moreover, high NO_x levels can suppress the formation of products for
205 nucleation, thereby reducing aerosol surface as a condensational sink and increasing the wall loss of
206 condensable species in an OFR under high-NO_x conditions (Zhao et al., 2018; Sarrafzadeh et al., 2016; Wildt
207 et al., 2014). Figure S3 indicates that the difference between SOA yields with and without NO_x decreased
208 with increasing silicon atoms within individual cVMS, indicating a less restricting effect of NO_x on the SOA
209 formation for larger cVMS. This means that high NO_x levels play a lessor role in the SOA yields of lower
210 volatility precursors.

211 SOA yields in the AS-seeded experiments under low and high-NO_x conditions are also shown in Figs. 1
212 and S3, indicating minimal impacts of the AS seed particles on SOA yields. A yield enhancement ratio
213 ($R_E = Y_{\text{seeded}}/Y_{\text{unseeded}}$, Fig. S4) was used to show the seed impacts more clearly. Under low-NO_x conditions,
214 the R_E values for all cVMS were close to 1.0 (Fig. S4a), suggesting negligible impact of AS seed particles
215 on SOA yields; however, under high-NO_x conditions, R_E was much larger (17.81, 13.18 and 15.51 for D3-
216 D5, respectively) at short PA but gradually decreased to 1.0 with increasing PA for D3-D5, while it was



217 always close to 1.0 for D6 regardless of PA (Fig. S4b). R_E values greater than 1.0 suggest that AS seed
218 particles played an enhancement role in the cVMS SOA formation, as similarly reported in SOA formation
219 from hydrocarbons (Sarrafzadeh et al., 2016; Lamkaddam et al., 2017; Li et al., 2019b). Under low- NO_x
220 conditions, the general lack of impact on cVMS SOA yields by the AS seed particles suggests that
221 condensation was not the main process driving SOA formation in cVMS oxidation. For the few cases of high
222 NO_x level at low PA, where R_E was >1 for D3-D5, it is possible that their early generations of oxidation
223 products were more volatile than successive generations of products and hence more prone to condensation
224 enhanced by AS seeds. As PA increased, further reactions of these early generation oxidation products with
225 OH radicals resulted in further generation products that were likely less volatile, thereby weakening the
226 enhancing role of AS seeds at high OH exposure. Such effect was less pronounced for D6, likely because its
227 oxidation products at different PA had similar volatilities. Figure S4b shows that the effect of AS seed
228 particles on SOA yields negatively correlated with the number of silicon atoms in the cVMS. Lower volatility
229 precursors (D5 and D6) formed lower volatile products (Alton and Browne, 2020), resulting in SOA yields
230 less sensitive to the pre-existing seeds. In fact, the oxidation products of D5 have been shown to be nearly
231 non-volatile (Janecek et al., 2019; Wu and Johnston, 2017).

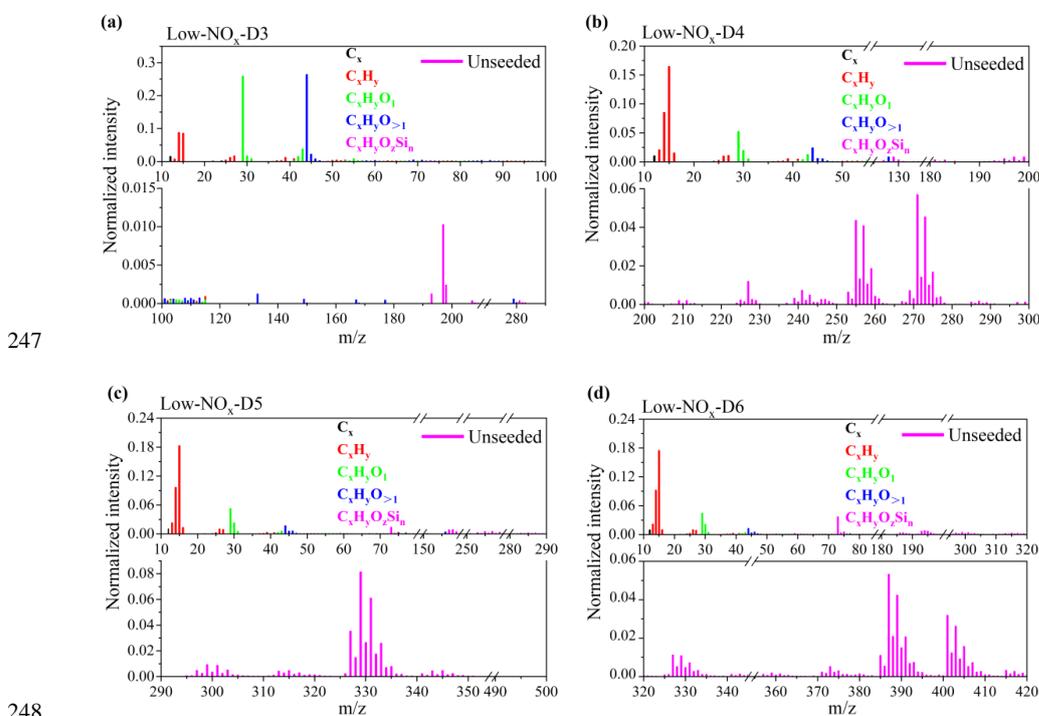
232 3.2 Aerosol compositions

233 3.2.1 Compositions of SOA under low- NO_x conditions.

234 Figures 2 and S5 show the normalized HR-ToF-AMS mass spectra of cVMS SOA from unseeded
235 experiments under low- NO_x conditions at OH exposures of 9.0×10^{11} molecules cm^{-3} s. The mass spectral
236 signals can be identified as fragments with a formula of $\text{C}_x\text{H}_y\text{O}_z\text{Si}_n$. For D3 SOA, the most prominent peaks
237 were at m/z 44 and 29, dominated by CO_2^+ and CHO^+ , which were tracers for organic acids (Ng et al., 2010),
238 alcohols and aldehydes (Lee et al., 2012), respectively. They may result from the oxidation of the methyl



239 groups in D3 by OH radicals. For the mass spectra of D4-D6 SOA, the two highest peaks at m/z 14 and 15
240 were CH_2^+ and CH_3^+ , respectively. In addition, there were several dominant $\text{C}_x\text{H}_y\text{O}_z\text{Si}_n$ peaks, which were
241 fragments of silicon-containing products. For the $\text{C}_x\text{H}_y\text{O}_z\text{Si}_n$ group in D4 SOA, there were four typical peaks
242 at m/z 255, 257, 271 and 273, with formulae of $\text{C}_4\text{H}_{11}\text{O}_7\text{Si}_3^+$, $\text{C}_3\text{H}_9\text{O}_8\text{Si}_3^+$, $\text{C}_3\text{H}_7\text{O}_9\text{Si}_3^+$ and $\text{C}_3\text{H}_9\text{O}_9\text{Si}_3^+$,
243 respectively. The $\text{C}_x\text{H}_y\text{O}_z\text{Si}_n$ fragment group containing Si of D5 SOA had three dominant peaks at m/z 327,
244 329 and 331, corresponding to $\text{C}_{12}\text{H}_{11}\text{O}_2\text{Si}_5^+$, $\text{C}_9\text{H}_9\text{O}_8\text{Si}_3^+$ and $\text{C}_5\text{H}_{15}\text{O}_9\text{Si}_4^+$, respectively. For the $\text{C}_x\text{H}_y\text{O}_z\text{Si}_n$
245 group containing Si in D6 SOA, there were five main peaks at m/z 73 ($\text{C}_3\text{H}_9\text{Si}$), 387 ($\text{C}_8\text{H}_{23}\text{O}_8\text{Si}_5^+$), 389
246 ($\text{C}_8\text{H}_9\text{O}_9\text{Si}_5^+$), 401 ($\text{C}_9\text{H}_{21}\text{O}_{10}\text{Si}_4^+$) and 403 ($\text{C}_7\text{H}_{15}\text{O}_{12}\text{Si}_4^+$).

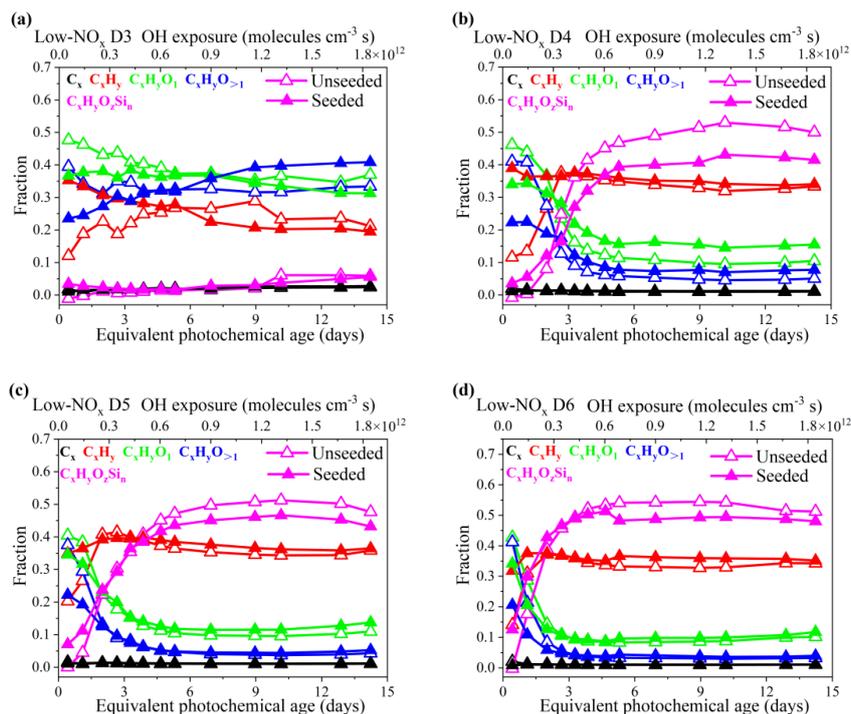


249 Figure 2. HR-ToF-AMS mass spectra of cVMS SOA at OH exposure of 9.0×10^{11} molecules cm^{-3} s under
250 low- NO_x conditions in unseeded experiments. **a-d** represent the mass spectra of D3-D6 SOA, respectively.

251



252 Figure 3 shows the evolution of different groups of ions in the HR-ToF-AMS spectra of the cVMS SOA
253 as a function of PA in equivalent days. For D3-D6 under unseeded conditions, $C_xH_yO_1$ and $C_xH_yO_{>1}$ ions
254 significantly decreased within 0-4 equivalent days of PA, but remained essentially unchanged when PA
255 increased to 7-15 equivalent days. The C_xH_y ion increased to its peak value at about 9 equivalent days of PA
256 for D3 and 2-3 equivalent days of PA for D4-D6, and then gradually decreased with further PA increases.
257 The $C_xH_yO_zSi_n$ group of ions maintained an increasing trend until 9-10 days of PA, thereafter it decreased
258 slightly for D4-D6 SOA. The weighted values of the atomic number ratios Si/C (n/x) and Si/O (n/z) for the
259 $C_xH_yO_zSi_n$ groups in D5 and D6 SOA at different PA are plotted in Fig. S6, which can be used to indicate
260 the changes in the Si element of SOA. The n/x ratio at initial SOA formation stage was close to that (0.50)
261 in D5 and D6 molecules, and then increased continuously with increasing PA. The n/z ratio kept increasing
262 from 0.53 to 1.15 for D5 SOA and from 0.72 to 1.32 for D6 SOA, but varying around 1.0 that was the Si/O
263 ratio in D5 and D6. While it is difficult to separate the effect of fragmentation due to the AMS ionization
264 process, the relative changes of group intensities and the evolution of n/x and n/z in $C_xH_yO_zSi_n$ over different
265 PA may be attributed to the initial substitution of methyl groups on the -Si-O- ring of the cVMS by OH or
266 CH_2OH when oxidized by the OH radicals (Wu and Johnston, 2016; Alton and Browne, 2020). Si-containing
267 products such as silanols may partition into SOA and result in an appearance of $C_xH_yO_zSi_n$ ions in the AMS
268 mass spectra. It is possible that in this process, the Si-O bonds may continue to be cleaved from OH radical
269 attack, followed by further deposition of less volatile and nonvolatile products containing $C_xH_yO_zSi_n$ on the
270 SOA.



271

272

273 Figure 3. Fraction of C_x, C_xH_y, C_xH_yO₁, C_xH_yO_{>1} and C_xH_yO₂Si_n ion groups for SOA derived from the
274 oxidation of cVMS (a-d) by OH radicals at different photochemical ages under low-NO_x conditions.

275 Empty and solid triangles represent experimental data under unseeded and seeded conditions, respectively.

276

277 As shown in Fig. 3, the presence of seeds led to some changes in the evolution trends of ion groups in the
278 AMS spectra. For instance, the initial fraction of C_xH_y in seeded experiments was larger than that in unseeded
279 experiments, whereas C_xH_yO₁ and C_xH_yO_{>1} exhibited opposite changes. The presence of seeds led to larger
280 initial and smaller steady-state C_xH_yO₂Si_n fractions than those in unseeded experiments. Regardless of the
281 presence of seeds, C_xH_y, C_xH_yO₁ and C_xH_yO_{>1} mainly contributed to the composition of all cVMS SOA at
282 initial OH radicals oxidation, but D4-D6 SOA primarily consisted of C_xH_y and C_xH_yO₂Si_n after 6 equivalent
283 days.



284 **3.2.2 Compositions of SOA under high-NO_x conditions.**

285 Figure S7 shows the HR-ToF-AMS mass spectra of cVMS SOA under high-NO_x conditions at OH
286 exposures of 9.0×10^{11} molecules cm⁻³ s (-6.9 d). Compared to that under low-NO_x conditions (Figs. 2 and
287 S5), there was one additional N-containing group (N_mO_z) in the SOA mass spectra under high-NO_x
288 conditions, which accounted for 16%-31%. For the mass spectra of D3-D6 SOA originating from unseeded
289 experiments under high-NO_x conditions in Fig. S7, the dominating peaks of the N_mO_z family were m/z 30
290 (NO⁺) and m/z 46 (NO₂⁺). The common main peaks were located at m/z 30 (NO⁺) for cVMS SOA, m/z 44
291 (CO₂⁺) for D3-D4 SOA, and m/z 46 (NO₂⁺) for D4-D6 SOA. In addition, there were other primary peaks at
292 m/z 29 (CHO⁺) for D4 SOA, while m/z 15 (CH₃⁺) and m/z 28 (CO⁺) for D5-D6 SOA. The m/z 28 (CO⁺),
293 similar with m/z 44 (CO₂⁺), is considered as a tracer for organic acids (Ng et al., 2010). In the mass spectra
294 for D3-D6 SOA under high-NO_x conditions, the presence of NO⁺ (m/z 30) and NO₂⁺ (m/z 46) illustrated the
295 formation of nitrates in SOA (Ng et al., 2007b).

296 For the C_xH_yO_zSi_n group in the D4 SOA mass spectrum under high-NO_x conditions, the dominating peaks
297 and their formulas were same as those under low-NO_x conditions. For the C_xH_yO_zSi_n group in D5 SOA, in
298 addition to two typical peaks at m/z 327 and 329 in the low-NO_x experiments, there was another prominent
299 peak at m/z 328, with a formula C₈H₁₂O₅Si₅. The C_xH_yO_zSi_n group in D6 SOA had three typical peaks at m/z
300 73 (C₃H₉Si), m/z 387 (C₈H₂₃O₈Si₅⁺) and m/z 401 (C₉H₂₁O₁₀Si₄⁺). For the C_xH_yO_zSi_n groups in cVMS SOA,
301 there was little difference in the x, y, z and n value assignment of C_xH_yO_zSi_n peaks in SOA generated under
302 low-NO_x and high-NO_x conditions, suggesting the formation of similar Si-containing oxidation products.

303 For cVMS SOA under high-NO_x conditions, the evolution of family groups as a function of OH exposure
304 was summarized in Fig. S8. The dominated composition at initial stage was C_xH_yO_{>1} groups for D3-D6 SOA.
305 At equivalent days larger than 6, D3 SOA primarily consisted of C_xH_yO_{>1}, N_mO_z and C_xH_yO₁ groups, while



306 D4-D6 SOA was mainly composed of $C_xH_yO_zSi_n$, N_mO_z , $C_xH_yO_1$, C_xH_y and $C_xH_yO_{>1}$ groups. Figure S8 also
307 shows influences of seeds on the evolution of family groups under high- NO_x conditions. It was observed
308 that all groups in D3-D6 SOA displayed similar change trends regardless of the existence of seeds. As shown
309 in Fig. S9, the trend of the weighted values of the atomic ratio n/z in the $C_xH_yO_zSi_n$ groups at different
310 photochemical ages under high- NO_x conditions was similar to that in low- NO_x experiments. However, the
311 n/x ratios remained almost unchanged, and were close to the initial value (0.5) in cVMS. This may be
312 attributed to possible suppression of cleavage of methyl groups from the -Si-O- ring of the cVMS under
313 high- NO_x conditions.

314 4 Conclusions and implications

315 The yields and compositions of SOA generated from the photooxidation of four cVMS (D3-D6) with OH
316 radicals were investigated using an oxidation flow reactor. cVMS SOA yields exhibited an overall increasing
317 trend with PA, and their values gradually increased with cVMS from D3 to D6. SOA formations depended
318 on NO_x , as shown by smaller SOA yields under high- NO_x conditions. Ammonium sulfate seeds significantly
319 enhanced SOA yields of D3-D5 at short PA under high- NO_x conditions. The SOA mass spectra showed that
320 Si-containing species were one of main chemical compositions at PA of >6 days.

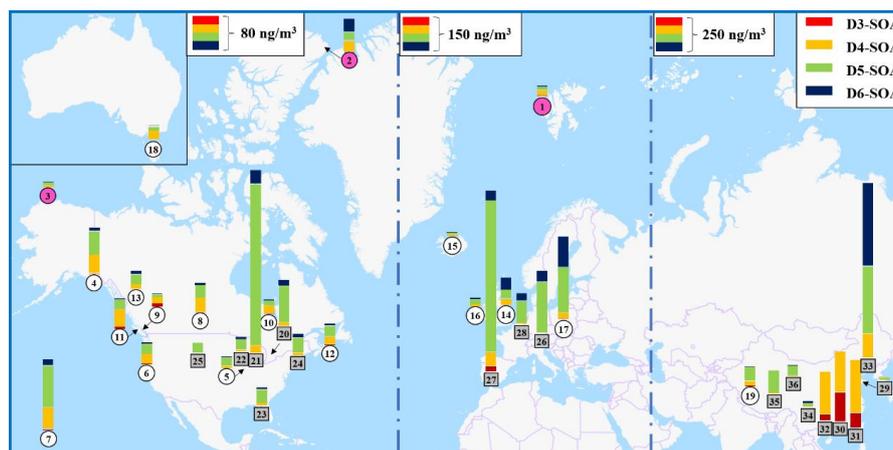
321 To evaluate the potential contributions of cVMS to SOA formation in the atmosphere, global SOA
322 concentrations produced from cVMS were estimated according to the cVMS SOA yields measured in this
323 work and using the cVMS concentrations reported from multiple studies, which were listed in Table S3.
324 Here, under the seeded conditions, the high- NO_x SOA yields at 8.5 equivalent days (D3: 0.028; D4: 0.122;
325 D5: 0.441; D6: 0.792) are employed in the calculation of cVMS SOA concentrations at urban sites, while
326 the low- NO_x SOA yields under the unseeded conditions at 14.2 equivalent days (D3: 0.148; D4: 0.556; D5:
327 0.805; D6: 0.975) are used to estimate cVMS SOA at background and polar sites. Figure 4 shows the global



328 concentration distribution of SOA from four cVMS (D3-D6) at 36 sites worldwide estimated by the Equation
329 4,

$$330 \quad C_{\text{cVMS-SOA}} = C_{\text{cVMS}} \times \frac{\Delta C_{\text{cVMS}}}{C_{\text{in-cVMS}}} \times Y \quad (4)$$

331 where C_{cVMS} and $C_{\text{cVMS-SOA}}$ are the mass concentration of cVMS reported from literatures and cVMS SOA
332 estimated in global sites, respectively; $C_{\text{in-cVMS}}$ and ΔC_{cVMS} are the mass concentration of initial and lost
333 cVMS at the selected equivalent days during the experiments of this work, respectively; Y is the cVMS SOA
334 mass yields mentioned above. Table S3 summarizes the details regarding sites and concentrations of cVMS
335 SOA. The derived concentrations of cVMS SOA varies significantly among urban, background and polar
336 sites. The total cVMS SOA concentrations in urban areas are the highest, up to 1249 ng/m³. They are 14.5-
337 347 and 10.5-105.3 ng/m³ for background and Arctic sites, respectively. cVMS SOA concentrations in urban
338 regions of Asia (sites 29-36) and Europe (sites 26-28) are generally larger than that of North America (sites
339 20-25). In China, the total cVMS SOA concentrations in urban sites range from 14.7 to 1249 ng/m³. The
340 main precursors of cVMS SOA are different among Chinese cities. For three cities along the southeast coast
341 of China (Guangzhou, Macau and Foshan), the dominant precursors of cVMS SOA are D3 and D4, which
342 are related to industrial emissions of these two siloxanes in this region (Wang et al., 2001). For Dalian in
343 China, mainly D5 and D6 contribute to cVMS SOA, which have the highest concentrations among all sites.
344 This can be attributed to the industrial production of D5 and D6 and the use of personal care products in
345 Dalian (Li et al., 2020). In the other Chinese urban areas with reported cVMS concentrations (Lhasa, Golmud,
346 Kunming and Yantai), the total cVMS SOA concentrations are considerably smaller than those in the urban
347 areas above, with D5 acting as the main precursor, which may be ascribed to the relatively low population
348 densities in these cities (Wang et al., 2018).



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Figure 4. Global concentrations of cVMS SOA (ng/m^3) on the basis of the cVMS concentrations

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reported from multiple studies and the cVMS SOA yields measured in this work. The numbers of polar,

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background and urban sites are enclosed in pink, white circles and gray boxes, respectively. The details

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about cVMS and SOA concentrations at different sites were summarized in Table S3 of Supplement.

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At urban sites in Europe and North America, cVMS SOA concentrations are reported in the range of 131-

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773 ng/m^3 and 21.8–416 ng/m^3 , respectively. Among these cVMS, D5 is the main contributor to cVMS SOA

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at these locations, averaging 80.4% of total cVMS SOA. This contribution is higher than that (73.9%) at

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Chinese urban sites. For instance, D5 SOA is calculated to be 652 ng/m^3 in Catalan, Spain, 369 ng/m^3 in

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Chicago, USA, 218 ng/m^3 in Zurich, Switzerland and 94 ng/m^3 in Paris, France, where there are high levels

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of economic activities and high population densities. These results suggest that personal care products as a

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main source of D5 may be the most important anthropogenic origins of Si-containing SOA in Europe and

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North America.

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At background and Arctic sites, cVMS SOA are primarily derived from D4 and D5. The background sites

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are located in mountains, rural areas, forested areas, lakes and at high altitudes. The three highest cVMS



365 SOA concentrations at background sites are located at Kosetice in the Czech Republic (347 ng/m³), Hilo,
366 Hawaii, USA (157 ng/m³) and Tibetan Plateau in China (134 ng/m³), where the contribution percentages of
367 SOA from both D4 and D5 are 61.9%, 90.3% and 94.9%, respectively. The cVMS SOA concentrations at
368 the Little Fox Lake site in Yukon, Canada is the highest (105.3 ng/m³) among the four locations in the Arctic,
369 91% of which is accounted for by both D4 and D5 SOA. The dominance of D4 and D5 SOA in both
370 background and the Arctic regions highlights their persistence in atmosphere and the potential for long-range
371 atmospheric transport.

372 Furthermore, global concentration distribution of Si-containing SOA estimated for the four cVMS (D3-
373 D6) at 36 sites worldwide is also presented in Table S3, which clearly shows the global importance of Si in
374 SOA with the estimated percentages of cVMS SOA that contains the Si element. For example, up to 49.6%
375 and 31.2% of cVMS SOA contained Si elements in background and urban sites, respectively. These results
376 are similar to the summary observations that reported percentages of aerosols with a Si mole fraction >0.01
377 (%) at different sites (Bzdek et al., 2014). The above results demonstrated that Si is a frequent component of
378 SOA in background and urban areas.

379 The global annual production of D4, D5 and D6 is about 1, 0.1 and 0.01 Tg·yr⁻¹, respectively, and 90% of
380 these cVMS is eventually released into the atmosphere (Li et al., 2020; Genualdi et al., 2011; Wang et al.,
381 2013; Sakurai et al., 2019). Based on the results shown in Fig. 4, the annual production of cVMS (D4-D6)
382 SOA was estimated to be 0.16 Tg·yr⁻¹, which was about 5.5% of SOA (2.9 Tg·yr⁻¹) produced from mobile
383 source emissions in the USA and 5-8 times of SOA generated by Athabasca oil sands (0.02-0.03 Tg·yr⁻¹, one
384 of the largest sources of anthropogenic secondary organic aerosols in North America) (Tkacik et al., 2014;
385 Liggio et al., 2016). Moreover, it was also 0.8% and 2.3% of isoprene-SOA (20 Tg·yr⁻¹) and monoterpenes-
386 SOA (7 Tg·yr⁻¹) (typical biogenic SOA), respectively, indicating the potential importance of cVMS SOA



387 (Jokinen et al., 2015). While these cVMS SOA sources may seem small, they can make substantially higher
388 contributions to ambient air SOAs in population centers where cVMS compounds are primarily used.

389

390 **Author contributions**

391 CH designed and conducted all experiments; CH and HY analyzed the data and prepared the paper with
392 contributions from KL, PL, JL, AL and SML. SML supervised the project.

393

394 **Competing interests**

395 The authors declare that they have no conflict of interest.

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