Estimating Brewer-Dobson circulation trends from changes in stratospheric water vapour and methane

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Abstract. The stratospheric meridional overturning circulation, also referred to as the Brewer-Dobson circulation (BDC), controls the composition of the stratosphere, which, in turn, affects radiation and climate. As the BDC cannot be directly measured, one has to infer its strength and trends indirectly. For instance, trace gas measurements allow the calculation of average transit times.

Satellite measurements provide information on the distributions of trace gases for the entire stratosphere, with measurements of particularly long temporal and dense spatial coverage available for stratospheric water vapour (H_2O). Although chemical processes and boundary conditions confound interpretation, the influence of methane (CH_4) oxidation on H_2O in the stratosphere is relatively straightforward, and thus H_2O is an appealing tracer for transport analysis despite these caveats. In this work, we explore how mean age of air trends can be estimated from the combination of stratospheric H_2O and CH_4 data, by carrying out a proof-of-concept within the model environment of the Chemical Lagrangian Model of the Stratosphere (CLaMS). In particular, we assess the methodological uncertainties related to the two commonly-used approximations of (i) instantaneous stratospheric entry mixing ratio propagation, and (ii) constant correlation between mean age and the fractional release factor of CH_4. Performing various sensitivity studies with CLaMS, we test different methods of the mean age of air trend estimation, and we aim to provide a simple and practical advice on the adjustment of the used approximations for obtaining more reliable mean age of air trend from the measurements of H_2O and CH_4.

Our results show that the estimated mean age of air trends from the combination of observed stratospheric H_2O and CH_4 changes may be significantly affected by the assumed approximations. Depending on the investigated stratospheric region and the considered period, the error in estimated mean age of air decadal trends can be large, especially in the lower stratosphere. For particular periods, the errors from the two approximations can lead to opposite effects, which may even cancel out. Finally, for a more reliable estimate of the mean age of air trends, we propose to adjust the approximation method by using an idealised

age spectrum to propagate stratospheric entry mixing ratios. The findings of this work can be used for assessing the uncertainties in stratospheric BDC trend estimation from global satellite measurements.

1 Introduction

The stratospheric Brewer-Dobson circulation (BDC) affects the atmospheric distributions of radiatively active trace gases and is an important element in the climate system. This global-scale circulation transports air masses upwards in the tropics, polewards, and downwards in middle and high latitudes (e.g., Holton et al., 1995). A particularly important greenhouse gas affected by the BDC is stratospheric water vapour (H_2O), which induces cooling of the stratosphere and warming of the troposphere (e.g. Solomon et al., 2010; Maycock et al., 2011; Riese et al., 2012). The reliability of climate model predictions is significantly affected by the representation of the processes controlling the distribution of stratospheric H_2O. Although the BDC is such a crucial factor influencing stratospheric H_2O and Earth's climate, its long-term trends and associated effects on transport and dynamics are not well understood. In particular, climate models predict a strengthening BDC in a future climate with increasing greenhouse gas levels (e.g., Butchart, 2014), whereas trace gas observations show only insignificant changes (Engel et al., 2017; Fritsch et al., 2020).

Because of the slowness of BDC transport (global transit times on the order of years) and its zonal mean character, direct measurements of related circulation velocities are not possible and the circulation must be inferred from temperature or trace gas observations. A commonly used diagnostic for BDC transport is the mean age of air (AoA), the average time scale for transport through the stratosphere. Garcia et al. (2011) pointed out that the AoA is difficult to estimate from sparse stratospheric observations. However, under some conditions, it is possible to infer AoA from trace gas concentrations (e.g., Hall and Plumb, 1994; Strunk et al., 2000; Engel et al., 2009). Suitable species are so-called "clock tracers" – trace gases with a linearly increasing source, which can provide the first moment of the age spectrum, namely AoA (Waugh and Hall, 2002; Schoeberl et al., 2005). Two examples of such tracers are SF_6 or CO_2, and these are frequently used to infer stratospheric AoA (e.g., Stiller et al., 2017; Engel et al., 2017). However, the availability of suitable observations with global coverage and extending over sufficiently long time periods necessary for estimating trends is very limited. Also, there are other complications, for instance SF_6 has a strong mesospheric sink, and CO_2 has a seasonal cycle causing problems with inferring AoA. Hence, several recent and ongoing research activities focus on trace gas species other than SF_6 and CO_2, to infer stratospheric AoA and information on the BDC (e.g., Linz et al., 2017; Leedham Elvidge et al., 2018).

In particular, trace gas species with chemical sinks in the stratosphere provide information on the stratospheric circulation as the transport through the sink regions depends on the strength and depth of the circulation. For such long-lived species with stratospheric sinks, like N_2O, CH_4, or the chlorofluorocarbons (CFCs), the chemical loss can be described by a fractional release factor (FRF)

$$\alpha = 1 - \frac{\chi}{\chi_{[entry]}} , \tag{1}$$

where χ is the observed mixing ratio in the stratosphere and $\chi_{[entry]}$ the mixing ratio at the tropical tropopause where the air enters the stratosphere. Consequently, changes in the strength and pattern of the stratospheric circulation cause changes in α and, in general, FRF highly correlate with AoA. On the one hand, it is more complicated for chemically active species to disentangle the effects of chemistry and transport. On the other hand, atmospheric measurements may be of higher quality, more frequent, provide denser sampling, and cover longer time periods compared to the canonical species SF_6 and CO_2, which are the ones commonly used to investigate the BDC.

As many long-lived species are only sparsely measured, stratospheric H_2O is particularly appealing as a tracer for estimating long-term trends, with a suite of measurements covering the past decades. The longest continuous in situ time series of stratospheric H_2O (starting in 1980) comes from the frost point hygrometer observations in Boulder, Colorado located at 40.0°N, 105.2°W. In addition, stratospheric H_2O observations from different satellite platforms exist since the mid-1980s, such as SAGE II (Satellite Aerosol and Gas Experiment, covering the period of 1984-2005; e.g., McCormick, 1987; Chu et al., 1993; Rind et al., 1993; Thomason et al., 1997), HALOE (Halogen Occultation Experiment, 1991-2005; e.g., Russell III et al., 1993), MIPAS Envisat (Michelson Interferometer for Passive Atmospheric Sounding, 2002-2012; e.g., von Clarmann and Stiller, 2003; Raspollini et al., 2006; Fischer et al., 2008), ACE-FTS (Atmospheric Chemistry Experiment-Fourier Transform Spectrometer, 2004-2012; e.g., Bernath et al., 2005; Bernath, 2017) and Aura MLS (Microwave Limb Sounder, 2004-present; e.g., Waters et al., 1999, 2004, 2006). The different satellite observations are merged into homogeneous global data sets of high value for analysing stratospheric variability and trends (e.g. Hegglin et al., 2014; Davis et al., 2017). To this end, Hegglin et al. (2014) estimated trends of AoA from a novel merged satellite H_2O data record, based on the conservation property of total water in the stratosphere (mainly the sum of H_2O and two times CH_4; for details see Sect. 2.3). They showed that decreasing H_2O mixing ratios in the lower stratosphere, below about 10 hPa, from the mid 1980s to 2010 and increasing H_2O mixing ratios above are related to an accelerating shallow branch of the BDC (decreasing AoA below about 10 hPa) and to a decelerating deep branch of the BDC (increasing AoA above). It is, however, not straightforward to accurately determine the AoA from stratospheric H_2O distributions due to complex processes involved.

We consider that at a particular time and location of the stratosphere, H_2O mixing ratios are determined by the value of the stratospheric entry mixing ratio, the propagation of this entry value into the stratosphere, and the chemical source of stratospheric H_2O. This assumption does not hold in the lowermost stratosphere as convection and isentropic transport can cause multiple entry mixing ratios. The only significant source of H_2O considered in this work is CH_4 oxidation. The chemical source of stratospheric H_2O from the oxidation of CH_4 is done by $O(^1D)$, OH, and Cl radicals (e.g. Röckmann et al., 2004). The strength of the chemical source of H_2O depends on the transit time and the transit path of air since entering the stratosphere and is, thus, related to AoA. The full complexity of these processes is very challenging to represent in the analysis of stratospheric H_2O. Consequently, stratospheric H_2O and CH_4 are used in combination for AoA estimation, and the assumptions of (i) an instantaneous propagation of stratospheric entry mixing ratios, and (ii) stationarity of the correlation between AoA and the fractional release factor of CH_4, are frequently assumed approximations (e.g. Schoeberl et al., 2000, 2005; Hegglin et al., 2014). Hegglin et al. (2014) particularly point out that the assumption of stationarity is only valid due to the lack of a trend in tropospheric CH_4 over the specific period they considered.

In this paper, we investigate the methodology to estimate AoA trends from stratospheric H_2O and CH_4. In particular, we address in detail the impact of the two frequently employed approximations described above, of (i) instantaneous entry mixing ratio propagation, and (ii) constant AoA-FRF correlation. For this purpose, we employ a closed model environment (the "model world") in which the mean stratospheric AoA and its trend is known, which is not the case when atmospheric measurements are analysed. In this way, the effect of each approximation on the calculated AoA trend and the associated temporal development of H_2O can be quantified ("proof of concept"). We aim for simple and practical advice on the approximate methods to estimate AoA trends from stratospheric measurements of H_2O and CH_4. Our results highlight the importance of assessing the robustness of observation-based methods against uncertainties in the underlying assumptions to test their validity and general applicability.

Section 2 introduces the methods, describing the Chemical Lagrangian Model of the Stratosphere (CLaMS, McKenna et al., 2002a; Pommrich et al., 2014) which is employed here as the model framework, as well as the methods for estimating AoA trends from H_2O and CH_4. Section 3 presents the main results on the impact of the used approximations on the resulting AoA trend, as well as introduces the adjustment to the approximations for an improved AoA trend estimation. Sect. 4 contains the discussion, and Sect. 5 the overall conclusions.

2 Methods

2.1 The CLaMS model

This paper is based on a study performed within the "model world" using the Chemical Lagrangian Model of the Stratosphere, CLaMS (McKenna et al., 2002a,b). The model set-up is described in detail elsewhere (e.g., Poshyvailo et al., 2018). Briefly, CLaMS is a modular Lagrangian chemistry transport model based on 3D-forward trajectories with parameterisation of small-scale mixing. CLaMS consists of several modules, such as Lagrangian advection (TRAJ), stratospheric mixing (MIX), sinks of H_2O (CIRRUS), stratospheric chemistry, and several other modules responsible for simulation of various physical and chemical processes. The modules act successively at each time step of 24 h. The boundary conditions at the surface are prescribed based on ground-based measurements in the lowest model level (below $\approx$ 4 km). CH_4 mixing ratios are taken from the zonally-symmetric NOAA/ESRL dataset (e.g., Masarie et al., 1991) from 1990 to 2011, and from zonally-resolved AIRS data (e.g., Xiong et al., 2008, 2013) for 2011-2017. The CH_4 boundary condition has been switched in 2011 to take advantage of the better sampling of AIRS data in comparison to NOAA, although accepting the apparent discontinuity. Because the results of the CLaMS model are internally consistent, the discontinuity has only negligible effects.

The CLaMS trajectory module TRAJ performs fully Lagrangian 3-dimensional advection of an ensemble of approximately 2 million air parcels. The position of each air parcel is defined in hybrid isentropic coordinates (Pommrich et al., 2014) and longitude-latitude space. Horizontal resolution is about 100 km, while the vertical resolution is defined via a critical aspect ratio, of about 250 (Haynes and Anglade, 1997). The CLaMS simulations cover the atmosphere from the surface to the stratopause (2500 K or $\approx$ 60 km). The advection of forward trajectories in CLaMS is calculated using 6-hourly wind fields and total diabatic heating rates from meteorological ERA-Interim reanalysis (Dee et al., 2011). Wind fields are linearly interpolated

from the adjacent grid points to the locations of the air parcels. The trajectory calculation advects air parcels to new positions within one model time step.

The parameterisation of small-scale mixing in the CLaMS mixing module is based on the deformation rate in the large-scale flow: air parcels may be merged, or new air parcels may be inserted at each time step, depending on the critical distances between the air parcels (McKenna et al., 2002a; Konopka et al., 2004). Note, that the mixing parameterisation affects both horizontal (associated with deformation in the horizontal flow) and vertical (related to the vertical shear) diffusivity (Konopka et al., 2004, 2005).

Dehydration in CLaMS is performed with the CIRRUS module (e.g., Poshyvailo et al., 2018). The calculation includes freeze-drying in regions of cold temperatures, mainly occurring around the tropical tropopause and in the Southern polar vortex. These cold temperatures cause formation and sedimentation of ice particles. If saturation along a CLaMS air parcel trajectory exceeds a critical saturation, then the H_2O amount in excess is instantaneously transformed to the ice phase and sediments out. The parameterisation of sedimentation is based on a mean ice particle radius, the characteristic sedimentation length and the corresponding fall speed. The fall distance of the ice particles is calculated from the fall speed and the computation time step. After comparison of the fallen path with a characteristic sedimentation length of the vertical grid size, a respective fraction of ice will be removed. If the parcel is sub-saturated and ice exists, the ice will be instantaneously evaporated to maintain saturation. For further details see Poshyvailo et al. (2018). CH_4 oxidation is included in CLaMS as a source of H_2O in the middle and upper stratosphere (for details see Pommrich et al., 2014). Note, that due to the simple parameterisation of ice microphysics and the omission of a parameterisation of convective processes, the simulated H_2O results are meaningful only above the tropopause.

2.2 Age spectrum calculated with CLaMS

In general, the mixing ratio of any long-lived trace gas $\chi(r,\, t)$ at a specific time and specific location in the stratosphere, with assumed absence of integrated loss, can be expressed as the following integral over all past times (Hall and Plumb, 1994; Waugh and Hall, 2002)

$$\chi(r,\, t) \;=\; \int_0^\infty \chi(r_0,\, t-t')\; G(r,\, t \mid r_0,\, t-t')\; dt'. \tag{2}$$

where t is the field time when the volume is sampled, t' is the transit time; $G(r,\, t \mid r_0,\, t-t')$ is the boundary propagator or Green's function of the transport operator. Here, G is interpreted as a transit time distribution (the "age spectrum"), and is the probability that the transit time of the air parcel travelling from the source r_0 to the sample point r is in the range between t' and $t'+dt'$. The first momentum of the age spectrum is the mean age of air (AoA). In this way, the stratospheric tracer distribution can be described through the contributions of the tropospheric evolution and transport.

In our study, the age spectrum is computed with CLaMS driven by the ERA-Interim reanalysis, using the "Boundary Impulse (time-)Evolving" (BIER) method based on multiple tracer pulses (Ploeger and Birner, 2016). For the inert tracer χ with a pulse at the location r_0, the field time t_0, and the source time t_0^*, the time evolution of the source can be described with a δ-distribution

as $\chi(r_0,\, t_0) \;=\; \delta(t_0 - t_0^*)$. Thus, Eq. 2 can be transformed to

$$\chi(r,\, t' + t_0^*) \;=\; G(r,\, t' + t_0^* |\, r_0,\, t_0^*), \tag{3}$$

where $t' = t_0 - t_0^*$ is a transit time, $G(r,\, t' + t_0^* |\, r_0,\, t_0^*)$ is the boundary impulse response at location r to a δ-boundary condition at the location r_0 at source time t_0^*. Having N different tracers χ_i $(i = 1, ..., N)$ with pulses at the source location r_0 at times $t_{0[i]}$ provides the field time dependence of the propagator G. The age spectrum, may be constructed at each field time t and location r as $G(r,\, t |\, r_0,\, t - t'_{[i]}) \;=\; \chi_i(r,\, t)$. Hence, the N different tracers provide N pieces of information for the age spectrum at the discrete transit times $t'_{[i]} \;=\; t - t_{0[i]}$.

Here, $N = 60$ different boundary pulse tracers were used. These pulses were released directly at the tropical tropopause between 30° S-30° N. Precisely, the source region covers the potential temperature layer from 10 K below to 10 K above the WMO (lapse rate) tropopause. The particular tracer mixing ratio is set to 1 for each pulse for a period of 30 days at the location r_0, and it is set to 0 in r_0 at other times. Pulses are launched every month. Consequently, to build the age spectrum for January 1990, the most recent tracer pulse has source times in January 1990, the second tracer pulse in December 1989, and so on. In our study, the original length of each age spectrum is 10 years (threshold transit time). The tail of the age spectrum is approximated with an exponential function when transit time exceeds 10 years (e.g., Ploeger and Birner, 2016). We used the exponential correction for the tail back to January 1979 for each age spectrum.

2.3 Contributions to H_2O changes

In this paper, we investigate the methodology to estimate AoA trends within a closed model environment of CLaMS in which the mean stratospheric H_2O and CH_4 and their trends are known. Trends in AoA can be calculated from trends in stratospheric H_2O mixing ratios by using the conservation property of total hydrogen in the stratosphere, namely that the sum of H_2O and two times CH_4 mixing ratios is approximately constant (e.g., Le Texier et al., 1988; Dessler et al., 1994). This conservation property implies for the H_2O mixing ratio at a given location r and time t in the stratosphere

$$\mathrm{H_2O}(r,\, t) = \mathrm{H_2O_{[entry]}}(r,\, t) + 2\alpha(r,\, t)\; \mathrm{CH_{4\,[entry]}}(r,\, t), \tag{4}$$

where $\mathrm{H_2O_{[entry]}}(r,\, t)$ and $\mathrm{CH_{4\,[entry]}}(r,\, t)$ are H_2O and CH_4 mixing ratios respectively at the specific location and time in the stratosphere, being transported from their entry location at the tropical tropopause without any chemical effects.

The time series of CLaMS H_2O and CH_4 at the entry to the stratosphere, averaged over the tropics at 30°S-30°N and in the potential temperature layer 390-400 K (approximately 80 hPa), chosen just above the cold tropical tropopause region, are shown in Fig. 1. The H_2O time series is highly variable (Fig. 1a), while the CH_4 time series shows a clear positive trend (Fig. 1b). Besides freeze-drying at the tropical tropopause and in the Antarctic polar vortex, stratospheric H_2O is controlled mainly by CH4 oxidation in the middle and high stratosphere (Fig. 1c, d). Consequently, CH_4 mixing ratios generally decrease with increasing altitude, as it is gradually chemically transformed into H_2O.

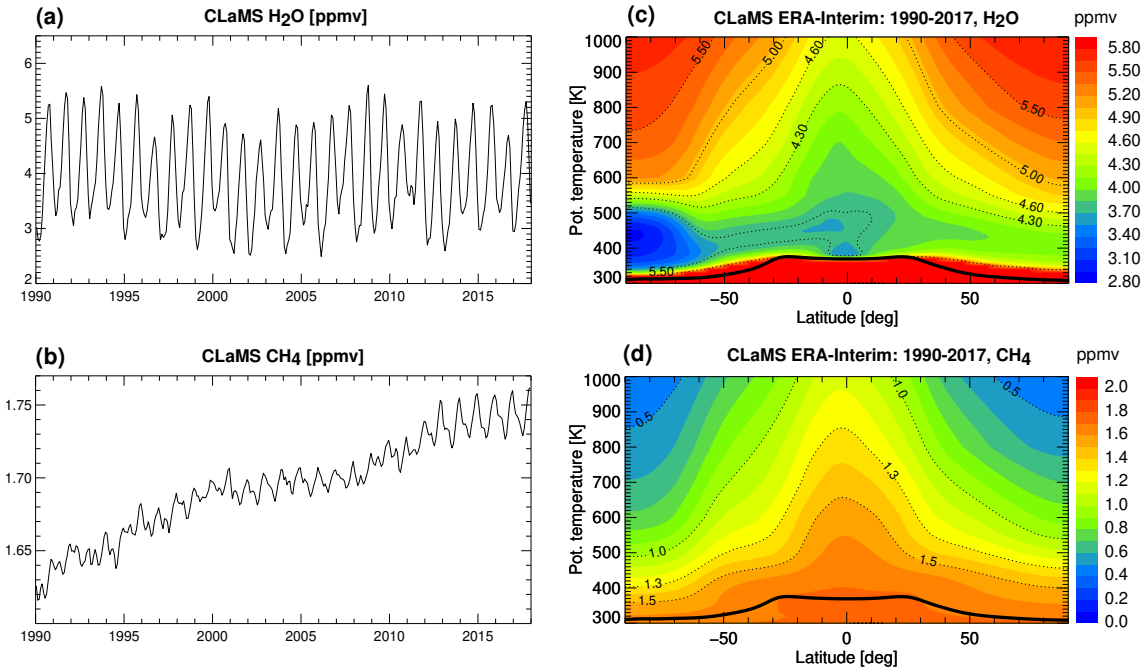

Figure 1. Time series of H_2O (a) and CH_4 (b) mixing ratios (in ppmv) for 1990-2017 averaged over the tropics at 30°S-30°N and in the potential temperature layer 390-400 K. Annual zonal mean H_2O distributions for H_2O (c) and CH_4 (d). Data shown is from the CLaMS simulation driven by ERA-Interim reanalysis.

The fractional release factor (FRF), α, describes the fraction of CH_4 which has been dissociated in the stratosphere (Solomon and Albritton, 1992), and it is determined as

$$\alpha(r,\,t) \;=\; 1 - \mathrm{CH_4}(r,\,t)/\mathrm{CH_{4\,[entry]}}(r,\,t). \tag{5}$$

The FRF is strongly affected by the vertical transport of the BDC. Hence, information on circulation trends (in particular on AoA) can be deduced from trends in FRF (Hegglin et al., 2014). Assuming long-term trends as small perturbations to the basic state, Eq. 4 can be rewritten as an equation for the linear trend in stratospheric H_2O over a given time period

$$\Delta\mathrm{H_2O}(r,\,t) = \Delta\mathrm{H_2O_{[entry]}}(r,\,t) + 2\alpha(r,\,t)\;\Delta\mathrm{CH_{4\,[entry]}}(r,\,t) + 2\mathrm{CH_{4\,[entry]}}(r,\,t)\;\Delta\alpha(r,\,t). \tag{6}$$

Here, $\Delta\mathrm{H_2O_{[entry]}}(r,\,t)$ and $\Delta\mathrm{CH_{4\,[entry]}}(r,\,t)$ are the trends in water vapour and methane transported from their stratospheric entry location at the TTL (neglecting chemistry), and $\Delta\alpha(r,\,t)$ is the trend in FRF. Note that the sum of H_2O and two times CH_4 mixing ratios is not conserved around and below the tropopause and in the polar vortex, where dehydration causes loss of H_2O. Hence, the presented analysis does not apply at the lowermost stratosphere, below the tropopause and in the polar vortex or near its edges.

Using the above approach, it is possible to investigate the different contributions to stratospheric H_2O changes related to changes in stratospheric H_2O entry mixing ratio, changes in CH_4 entry mixing ratio, and changes in the FRF, which are affected by stratospheric circulation changes and can be converted to an AoA trend.

2.4 Methods of assessing AoA trends from H_2O changes

There are four methods of AoA trend calculation used in this study, depending on the assumed approximations of (i) an instantaneous propagation of stratospheric entry mixing ratios, and (ii) stationarity of the correlation between AoA and the FRF. The summary of the different terms needed for AoA estimation with respect to the used method is shown in Table 1.

The full reconstruction method (FULL) includes the most detailed representation of the true atmospheric processes. In the FULL method, H_2O and CH_4 entry mixing ratios are propagated through the convolution of the TTL mixing ratios with the modelled age spectrum. The monthly varying AoA-FRF correlations are used for translating FRF into AoA changes, to include effects of the non-stationarity of the correlation. None of the approximations (i) or (ii) is used in the FULL method. Note, that for estimating AoA trend with the FULL method, the propagated entry H_2O mixing ratios are not needed; the AoA trend is deduced from the FRF, which requires only propagated entry CH_4 mixing ratios.

The constant correlation method (C-CORR) includes the propagation of entry H_2O and CH_4 mixing ratios by the modelled age spectrum (same as for FULL), but the stationary relationship between FRF and modelled AoA is used (ii approximation). The difference between C-CORR and FULL is in the correlation between AoA-FRF and the method of calculating FRF and its trend.

The method based on both approximations (i) and (ii) is named "approximation method" in the following, and is abbreviated APPROX. In fact, APPROX method is used in the literature, and we evaluate it in this paper in detail. Finally, we introduce an improvement to the approximation method: instead of using approximation (i), stratospheric entry H_2O and CH_4 mixing ratios are propagated with the parameterised idealised age spectrum.

3 Results

In the following we consider the consequences of the two major approximations (i) instantaneous propagation of stratospheric entry mixing ratios, and (ii) constant correlation (stationary relationship) between FRF and AoA. We evaluate the effects of these approximations on the AoA trends inferred from H_2O and CH_4 changes through comparison of the "true" AoA trend (actual modeled with CLaMS) and the AoA trends estimated with the different methods. First, we consider the extended 1990-2017 period and thereafter 1990-2006. The results of this work provide an estimate of the reliability of the approximation method to deduce circulation trends from observed stratospheric H_2O and CH_4 mixing ratios.

Table 1. Specification of the terms (first column of the table) required for AoA trends estimation from H_2O changes in different methods: full reconstruction (FULL), constant correlation (C-CORR), approximation (APPROX) and improved approximation (APPROX-improved) methods.

Term	**FULL**	**C-CORR**	**APPROX**	**APPROX-improved**
$\Delta H_2O_{[\text{entry}]}(r, t)$	Trend of H_2O propagated by monthly age spectrum, see Eq. 2	Trend of H_2O propagated by monthly age spectrum, see Eq. 2	Trend from H_2O time-series averaged over 390-400 K and 30° S-30° N	Trend of H_2O monthly propagated by parameterised idealised age spectrum
$2\alpha(r, t)\,\Delta CH_{4[\text{entry}]}(r, t)$	Trend of CH_4 propagated by monthly age spectrum, see Eq. 2; FRF is defined from Eq. 5	Trend of CH_4 propagated by monthly age spectrum, see Eq. 2; FRF is defined from Eq. 5	Trend from CH_4 time-series averaged over 390-400 K and 30° S-30° N; FRF is defined from approximated Eq. 5	Trend of CH_4 monthly propagated by parameterised idealised age spectrum, see Eq. 2; FRF is defined from Eq. 5
$2CH_{4[\text{entry}]}(r, t)\,\Delta\alpha(r, t)$	Trend of FRF defined from Eq. 5; used the climatological mean of CH_4 propagated by monthly age spectrum	Trend of FRF is residual from Eq. 6; used the climatological mean of CH_4 propagated by monthly age spectrum and trend of CLaMS H_2O	Trend of FRF is residual from Eq. 6; used CH_4 averaged over 390-400 K and 30° S-30° N and trend of CLaMS H_2O	Trend of FRF is residual from Eq. 6; used the climatological mean of CH_4 monthly propagated by parameterised idealised age spectrum and trend of CLaMS H_2O
ΔAoA	AoA trend is from recalculated AoA from FULL FRF: defined by the monthly varying correlation function f between FULL FRF and CLaMS AoA, where AoA = $f(\alpha)$	AoA trend is defined by the third-order polynomial constant in time empirical correlation function f, $\Delta\text{AoA} = f(\alpha + \Delta\alpha) - f(\alpha)$; note, that α is taken from APPROX for compatibility between the methods	AoA trend is defined by the third-order polynomial constant in time empirical correlation function f, $\Delta\text{AoA} = f(\alpha+\Delta\alpha)-f(\alpha)$, used α and $\Delta\alpha$ are from this method	AoA trend is defined by the d-order polynomial constant in time empirical correlation function f, $\Delta\text{AoA} = f(\alpha+\Delta\alpha)-f(\alpha)$; note, that α is taken from APPROX for compatibility between the methods

3.1 Contributions to stratospheric H_2O trend

AoA trends can be estimated from the separation of H_2O changes into three different contributions (see Eq. 6). In a first step, we estimate here each term of Eq. 6 using the full reconstruction method FULL through propagation of H_2O and CH_4 entry mixing ratios by the modelled age spectrum (for details see Table 1).

We calculate the first term of Eq. 6 by the convolution of H_2O mixing ratio at the tropical tropopause with the transport operator's Green's function using Eq. 2, The Green's function, or stratospheric age spectrum, has been simulated by CLaMS and is known over the considered period (for details see Sec. 2.2). Hence, the propagation of boundary mixing ratios to each grid point in the stratosphere provides the full reconstructed stratospheric tracer field. The analogous calculation is applied to derive CH_4 mixing ratios in the stratosphere, transported without including chemical effects.

In our study, the entry time series of H_2O and CH_4 are taken from zonally averaged monthly mean data simulated with CLaMS driven by ERA-Interim reanalysis (see Fig. 1a, b). The location of entry to the stratosphere is approximated as the 390-400 K layer between 30°S-30°N, which is located just above the cold point tropopause to avoid complications related to H_2O dehydration processes at the tropopause. The propagation procedure yields zonally averaged stratospheric entry H_2O and CH_4 distributions with a monthly resolution on the latitude-potential temperature grid.

For deducing FRF, the relation from Eq. 5 is used, where $CH_4(r,\ t)$ is monthly mean CH_4 simulated with CLaMS driven by ERA-Interim reanalysis. The $CH_{4[\mathrm{entry}]}(r,\ t)$ is calculated as described above, through the convolution of the tropical entry mixing ratio time series with the age spectrum. Consequently, the resulting FRF has a monthly resolution.

Thus, we can estimate each contribution to the stratospheric H_2O mixing ratio. All trends are calculated through a linear regression which minimizes the standard deviation at each latitude-potential temperature grid. The different contributions to the total stratospheric H_2O change for 1990-2017 are shown in Fig. 2. Note that in this study we applied the propagation procedure on the period 1990-2017 because of the availability of the CLaMS data and the necessary age spectrum length.

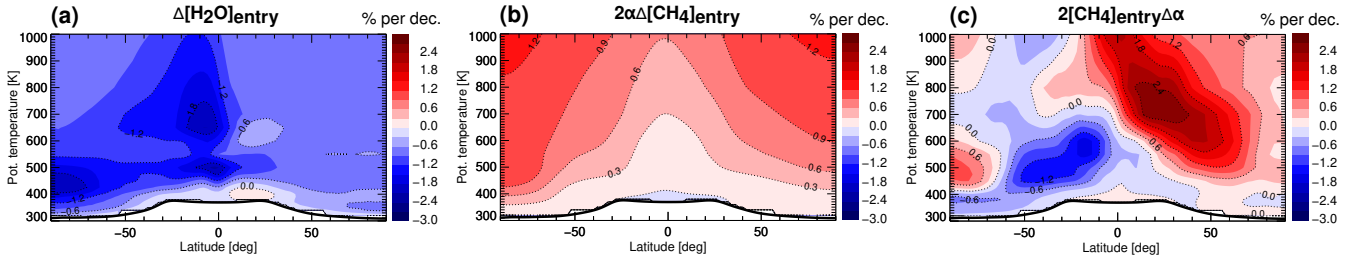

Figure 2. Contributions to stratospheric H_2O trends for 1990-2017 from (a) stratospheric H_2O entry mixing ratio changes, (b) stratospheric CH_4 entry mixing ratio changes, and (c) circulation changes. Stratospheric entry H_2O and CH_4 are derived through propagation of their stratospheric entry mixing ratios by convolution with the CLaMS modelled age spectrum. The data is from CLaMS simulations driven by ERA-Interim reanalysis, and is presented in percentage per decade, with relation to the climatological 1990-2017 stratospheric H_2O mixing ratios. The black line is the (lapse rate) tropopause calculated from ERA-Interim using the WMO definition (WMO, 1957). The white region below tropopause denotes the region, where the reconstruction method can not be applied.

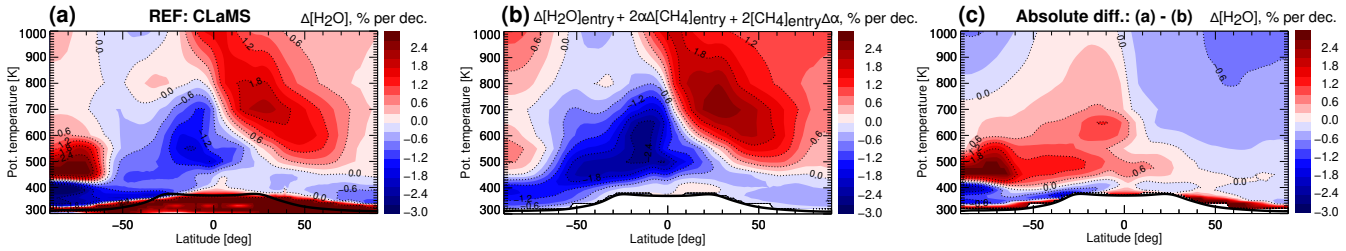

Figure 3. H_2O trends during 1990-2017 period and their reconstruction, shown in percentage per decade with respect to the 1990-2017 climatology. Sub-figure (a) represents the reference (REF) "true" stratospheric H_2O trend calculated from the CLaMS simulation driven by ERA-Interim reanalysis, (b) shows the reconstruction as sum of the different contributions from Fig. 2, (c) shows the absolute difference between (a) and (b). The black line is the tropopause calculated from ERA-Interim. Note, that due to the simple parameterisation of ice microphysics and the omission of a parameterisation of convective processes in CLaMS, the simulated H_2O shown in (a) is meaningful only above the tropopause.

Figures 2a, b represent the first two terms of Eq. 6, related to the entry H_2O and CH_4 mixing ratio trends. Figure 2c shows the impact from circulation changes, in terms of the change in FRF. In general, the different contributions affect the stratospheric H_2O changes differently in different regions, consistent with the findings of Hegglin et al. (2014). The strongest regional pattern is apparent in the contribution related to the stratospheric circulation change (Fig. 2c).

For assessing the reliability of the method applied to estimate the different contributions, we compare the stratospheric H_2O trend reconstructed as the sum of the calculated contributions (in Fig. 2) with the actual trend of CLaMS simulated H_2O. Figure 3a shows the "true" H_2O trend from the CLaMS simulation, while Fig. 3b shows the sum of the three terms from Fig. 2. Figure 3c shows the absolute difference between the true and reconstructed trends, indicating clear quantitative differences. Particularly large differences occur in the Antarctic region. This is expected due to the strong local dehydration occurring in that region, and the related failure of the total hydrogen conservation. Hence, the results of the reconstruction method should be interpreted with caution in the Southern polar region. Further disagreement between Fig. 3a and Fig. 3b is partly related to inaccuracies in the modelled age spectrum (monthly pulsing, limited spectrum length), and to inaccuracies in the boundary time-series (averaging in the layer of 390-400 K potential temperature and 30° S-30° N). Note, that through the convolution of the age spectrum and the stratospheric entry time series it is possible to reconstruct mixing ratios only above the tropopause (or the level of the boundary time-series, if chosen differently). Outside of the Southern high-latitude regions, the overall differences shown in Fig. 3c are small, and the propagation procedure provides a good estimate of stratospheric H_2O change and its contributions, at least regarding the large-scale patterns. Note that neither stratospheric entry H_2O (Fig. 2a) nor CH_4 changes (Fig. 2b) explain the pattern of the "true" H_2O trend (Fig. 3a). Instead, the circulation change term (Fig. 2c) includes the regional characteristics of the actual stratospheric H_2O trend.

3.2 AoA trends estimation using monthly AoA-FRF correlation in the full reconstruction method (FULL)

The changes in FRF can be translated into changes in AoA using the correlation between estimated FRF and aforehand known AoA, following the procedure described by Hegglin et al. (2014). In the full reconstruction method FULL (for details see Table 1), a monthly varying correlation between estimated FRF and CLaMS AoA is used, where AoA = $f(\alpha)$, and f is an empirically determined correlation function (Fig. 4).

In an example for January, April and July of 1995 shown in Fig. 4a, the AoA-FRF correlation functions are unique for each month, because the differences in magnitude of the coefficients are greater than the standard error's range. Moreover, monthly AoA-FRF correlation functions have a very small difference for relatively young air (< 4 years) and low FRF (< 0.4), but there are visible differences towards older AoA.

It is worth mentioning that the monthly AoA-FRF correlation functions are still a simplification and might introduce some bias in the reconstruction. In general, an accurate AoA-FRF correlation function depends not only on the considered time but also on longitude, latitude and altitude. An example of the relationship between AoA and FRF with regard to different latitude ranges (30° S-40° S, 10° S-10° N, 30° N-40° N) is shown in Fig. 4b for January, 2000. The AoA-FRF correlation functions are unique for each latitude range, because the differences in the magnitude of coefficients are out of the standard error's range. And, for instance, at the same FRF level of 0.3, the air at the Northern tropics (30° N-40° N) is younger than at the Southern tropics (30° S-40° S) by almost half a year.

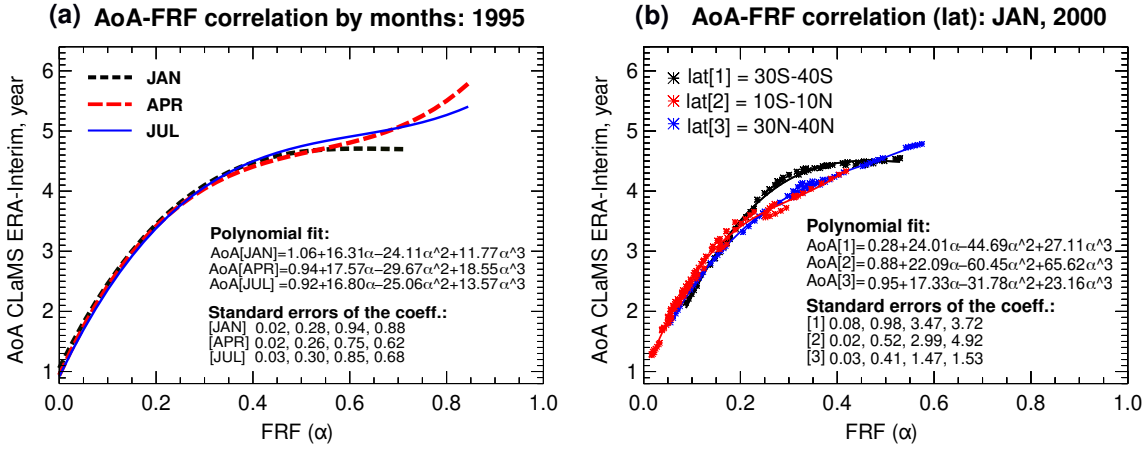

Figure 4. An example of relationship between estimated FRF (with FULL method) and CLaMS AoA. On sub-figure (a) are shown correlation functions for three months of 1995: January (black dashed line), April (red dashed line) and July (blue solid line); FRF and CLaMS AoA are taken at the whole range of latitudes (90° S-90° N). On sub-figure (b) are shown the relationships between estimated FRF and CLaMS AoA for January 2000 for three intervals of latitudes: between 30° S and 40° S (black line and black star-dots), 10° S-10° N (red line and red star-dots), 30° N-40° N (blue line and blue star-dots). The estimated FRF and CLaMS AoA driven by ERA-Interim (zonally averaged) are taken between 450 K and 1000 K, where FRF has positive values, in both sub-figures. The AoA-FRF relationships are shown by monthly fitting third-order polynomial function; in (a) the actual AoA-FRF distribution is not shown to avoid an overcrowded plot.

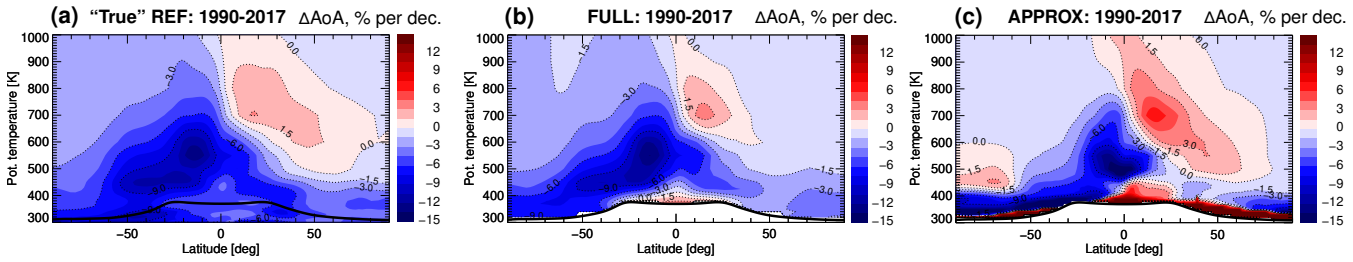

Figure 5. Decadal changes of AoA over 1990-2017: (a) "true" AoA changes from the CLaMS simulation driven by ERA-Interim reanalysis, (b) using CH_4 propagation by CLaMS age spectrum and monthly varying AoA-FRF correlation (FULL), (c) AoA changes estimated with the approximation method (APPROX). The changes are presented in percentage per decade relatively to the AoA 1990-2017 climatologies in each used method of AoA calculation. The black line is the tropopause calculated from ERA-Interim.

AoA trend is calculated from the resulting AoA applying a linear fit (minimizing the standard deviation) at each latitude-potential temperature grid point. The reference "true" AoA trend calculated directly from CLaMS simulated AoA is shown in Fig. 5a, whereas the resulting AoA trend calculated in FULL method is shown in Fig. 5b. The estimated AoA trend with FULL is qualitatively and quantitatively highly reliable, when comparing to the reference CLaMS AoA trend. Note, that for estimating AoA trends with the FULL method, the propagated entry H_2O mixing ratios are not used; the AoA trend is deduced from the FRF, which requires only propagated entry CH_4 mixing ratios. Visible differences between Fig. 4a and Fig. 4b are related to approximations in the AoA-FRF correlation: monthly AoA-FRF correlation functions are still a simplification and might introduce some bias in the reconstruction. The averaged (between 390-400 K potential temperature, from 30° S to 30° N) CH_4 boundary mixing ratios, which are used for the reconstruction, could induce biases as well.

3.3 AoA trends estimation in the approximation method (APPROX)

Accurate estimation of AoA from observed trace gas distributions is a complicated task. Even though, it is desirable to have a complete age spectrum for AoA calculations, it is very difficult to obtain it from measurements. Consequently, different approximations are often applied when deriving AoA from trace gases observation with non-linear increase, as well as assumptions about the age spectrum and its shape (e.g., Schoeberl et al., 2000, 2005; Ehhalt et al., 2007; Hegglin et al., 2014). In the following, we further investigate the method of AoA trend estimation from combination of H_2O and CH_4 changes, applying the two major approximations introduced above: (i) instantaneous propagation of stratospheric entry mixing ratios, and (ii) constant correlation (stationary relationship) between FRF and modelled AoA.

In the approximation method APPROX, each term of Eq. 6 is approximated. The actual trend of stratospheric H_2O (left side of the Eq. 6) is assumed to be known aforehand. In our case it is the CLaMS simulated H_2O change over the considered period; if the method is applied to observations it would be the observed H_2O change. The needed variables for the first and the second term of Eq. 6 are obtained as a linear trend of the model entry mixing ratio time series, averaged over the 390-400 K potential temperature layer and 30° S-30° N, from the considered period (see Fig. 1). The FRF required for the second term is derived from Eq. 5. Therefore, following Hegglin et al. (2014) we use zonal mean CLaMS simulated $CH_4(r,\, t)$ mixing ratios averaged

over 2005-2006, and we calculate $CH_{4[entry]}$ as a mean mixing ratio between 390-400 K potential temperatures and 30° S-30° N over 2002-2006. The third term is calculated as a residual between the actual CLaMS H_2O trend over the considered period (left side of Eq. 6) and the other two terms. Dividing the third term by $2CH_{4[entry]}$ yields the FRF changes.

In order to estimate the AoA trend induced by the changes in stratospheric H_2O, we define the relationship between FRF from APPROX and CLaMS modelled AoA (taken as 2005-2006 climatology inthe APPROX method). The correlation function is derived by fitting a third-order polynomial, as suggested by Hegglin et al. (2014). In our study, the empirical relationship between APPROX FRF and CLaMS AoA is described by the function $f(\alpha) = 0.85 + 16.49\alpha - 25.30\alpha^2 + 13.77\alpha^3$ (see Appendix A). The approximation method assumes that the AoA-FRF relationship is stable in time (stationary). Note that by applying a constant AoA-FRF correlation function some atmospheric variability can be lost. Using the correlation function $f(\alpha)$ we obtain the AoA trends from previously estimated FRF and its changes, as $\Delta\text{AoA} = f(\alpha + \Delta\alpha) - f(\alpha)$. A summary on all terms defined within the APPROX method is provided in Table 1.

The resulting decadal AoA trend for 1990-2017 estimated with the approximation method is shown in Fig. 5c, and can be compared to the "true" AoA trend from the reference CLaMS simulation, shown in Fig. 5a. There are visible quantitative differences between the "true" and the estimated trend of the AoA, especially in the Antarctic region. These differences are related to the dehydration processes occurring in that region. Also, the approximation method overestimates the AoA trend in the Northern hemispheric subtropical middle stratosphere. In the extratropical lowermost stratosphere (below about 380 K), AoA trends calculated with the approximation method are even opposite compared to the true trends, likely related to significant transport into this region across the subtropical tropopause (e.g., Hauck et al., 2019) which is not represented in the APPROX. But overall, both APPROX and the "true" AoA trend show good agreement: decreasing AoA in the lower stratosphere, and increasing AoA in the Northern hemispheric middle stratosphere. Interestingly, for the 1990-2017 period AoA trend show clear differences between the Northern and Southern hemispheres. The hemispheric differences might be related to the effect of mixing (e.g., Ploeger et al., 2015) and shifting stratospheric circulation patterns (Stiller et al., 2017), previously found in long-term AoA trend derived from the observed stratospheric CH_4 (Remsberg, 2015). Overall, APPROX provides a good estimate of the AoA trend for 1990-2017, corroborating the validity of the applicability of this method to H_2O observations over similar time periods (e.g., Hegglin et al., 2014).

To assess the general applicability of the approximation method APPROX, we consider another period, 1990-2006. The "true" AoA trend for this period from the CLaMS simulation is presented in Fig. 6a, and the result from APPROX is shown in Fig. 6c. In this case, the APPROX AoA trend disagrees substantially when compared with the "true" CLaMS trend. Differences occur even in the sign of the AoA trend. Particularly clear differences occur in the strength of the AoA trend and its detailed pattern. Thus, the accuracy of the estimated AoA changes from APPROX largely depends on the considered period. In the following section, we further investigate the effects of the applied approximations and discuss their impact on the quality of the estimated AoA trend.

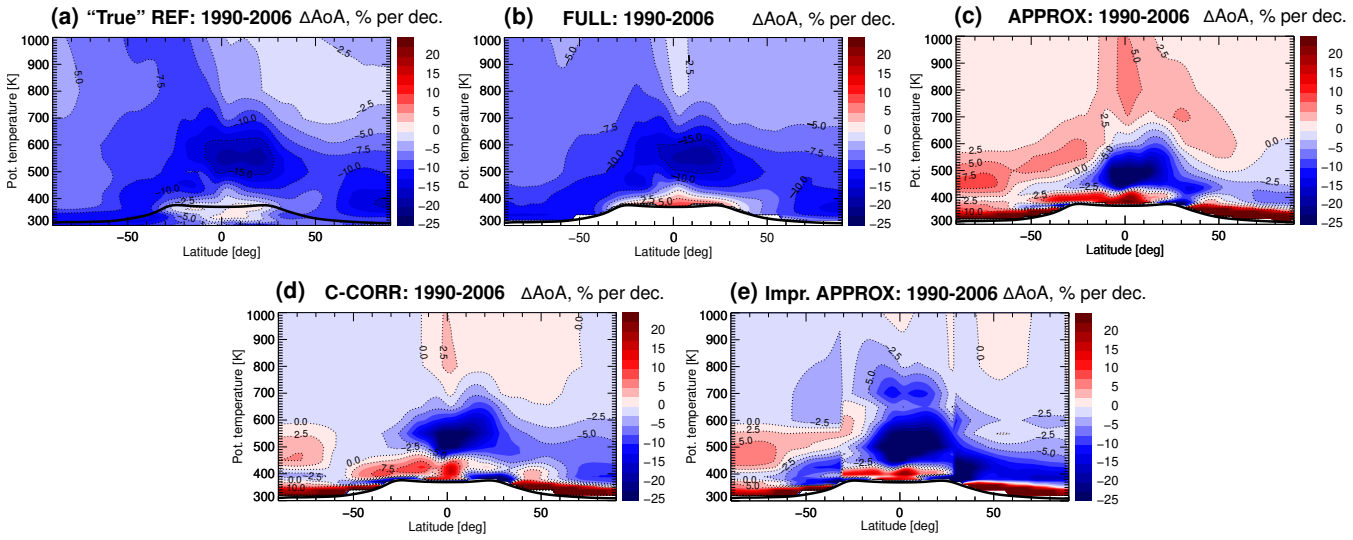

Figure 6. Comparison of decadal AoA trends for the 1990-2006 period, estimated using different methods: (a) "true" AoA trend from CLaMS simulation driven by ERA-Interim reanalysis, (b) FULL, (c) APPROX, (d) C-CORR (stratospheric entry H_2O and CH_4 propagation by CLaMS age spectrum and a constant AoA-FRF correlation), (e) improved approximation method (stratospheric entry H_2O and CH_4 propagation by the parameterised age-spectrum and a constant AoA-FRF correlation). The AoA trends are presented in percentage per decade with respect to the 1990-2006 AoA climatologies from the respective method. The black line is the climatological 1990-2006 tropopause calculated from ERA-Interim.

3.4 Effect of the approximations: entry mixing ratio propagation and constant AoA-FRF correlation

Firstly, we evaluate the effect of approximation (i) of the instantaneous entry mixing ratio propagation. For this purpose, we perform an additional sensitivity study, where the stationary relationship between FRF and modelled AoA is kept, but entry H_2O and CH_4 mixing ratios are propagated through the convolution of the CLaMS mixing ratios with the modelled age spectrum. This method is termed as constant correlation method, C-CORR, in the following (see Table 1 for details). Note that in both methods, C-CORR and APPROX, the same FRF distribution is used, but the changes in circulation (namely $\Delta\alpha$) are different.

The estimated AoA trend from C-CORR for 1990-2006 is shown in Fig. 6d. The approximation of the instantaneous entry propagation largely affects the AoA trend, as evident from comparison of the resulting AoA trends from C-CORR, APPROX, and the CLaMS reference AoA trend (Fig. 6a, c, d). Including the entry H_2O and CH_4 propagation by the age spectrum in the method clearly improves the estimated AoA trend. When comparing C-CORR to APPROX the general trend patterns stay similar, but improvements are visible in the extratropical lower stratosphere and above about 600 K.

For evaluating the effect of the approximation (ii) of a constant correlation between FRF and AoA, we compare the resulting AoA trends from C-CORR and FULL methods. The difference between the AoA trend results from C-CORR and FULL

methods stems from the differences in the AoA-FRF correlations used in each method, and the explicit FRF trend calculation (see Table 1).

In C-CORR, the AoA trend is estimated from the residual circulation contribution (see Table 1). Consequently, the monthly varying AoA-FRF correlation improves the accuracy of the estimated AoA trend both qualitatively and quantitatively, when compared to the stationary AoA-FRF correlation (Fig. 6a, b, d). It was mentioned earlier in the paper that stratospheric H_2O is a highly variable tracer, and can lead to difficulties in estimating AoA trends. Hence, the good performance of the FULL method can be related to the fact that stratospheric entry H_2O mixing ratios are not influencing the calculation.

For a more precise assessment of the effects of the two approximations, the differences among AoA trends estimated with different methods (APPROX, C-CORR, FULL) and for two different periods (1990-2017 and 1990-2006) are analysed. The difference between AoA trends from APPROX and C-CORR gives an estimate of the effect of the approximation (i) assuming instantaneous entry mixing ratio propagation, whereas the difference between C-CORR and FULL gives an estimate of the effect of the approximation (ii) using a constant AoA-FRF correlation (Fig. 7).

Figure 7a, b shows the differences in AoA trends for 1990-2017 between C-CORR and APPROX, and FULL and C-CORR. The differences are less than 5% per decade above 600 K. Below 600 K, the differences in AoA trends are higher, with the maximum at approximately 480 K. The larger differences in the lowest stratosphere directly above the tropopause should not be over-interpreted as the transit time resolution of used age spectra of 1 month is too coarse for a reliable reconstruction there. Interestingly, the effects of the first and second approximations are opposite in sign for 1990-2017 (Fig. 7a, b). Consequently,

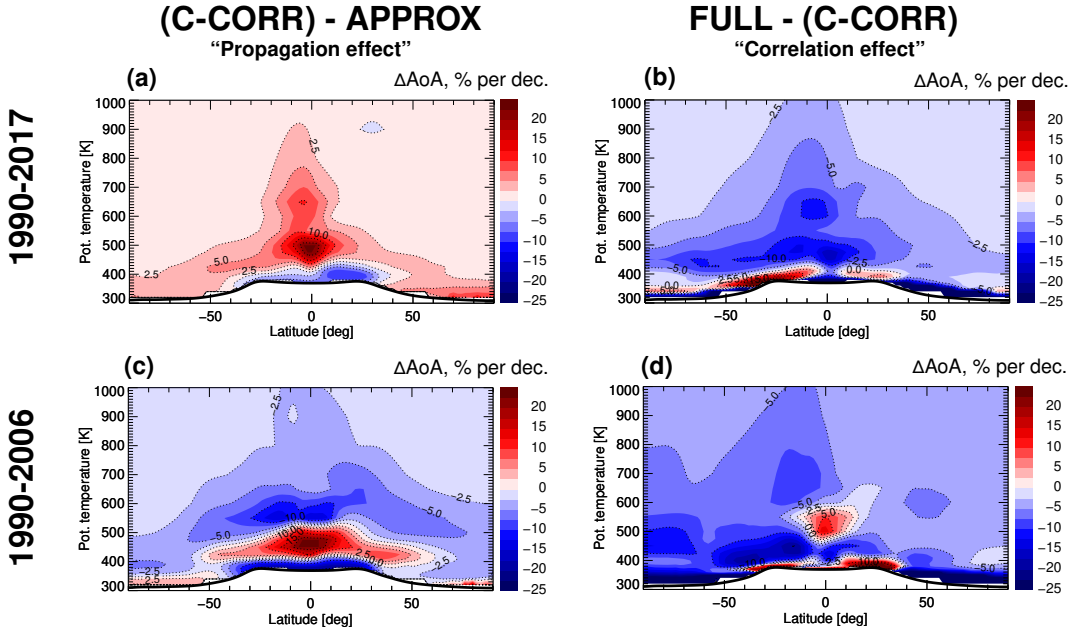

Figure 7. Differences in AoA trends estimated with three methods, APPROX, C-CORR, and FULL, for the periods of 1990-2017 (a, b) and 1990-2006 (c, d). The black line is the climatological tropopause calculated from ERA-Interim for the considered period.

the effects from both approximations cancel out to some extent, such that the APPROX method yields results remarkably close to the CLaMS reverence AoA trend (see Fig. 5). However, in general such cancellation can not be expected.

Figure 7c, d shows the difference in AoA trends for 1990-2006. The difference above 600 K is around 5% per decade or less. A more complex structure is found below 600 K. Generally, the instantaneous entry mixing ratio propagation causes an error in the estimated stratospheric entry H_2O and CH_4 contributions (see Sect. 4). This, in turn, causes an error in the derived residual circulation impact in C-CORR, which is further translated into the estimated AoA trend error.

The above analysis shows that the effects of both approximations (instantaneous propagation of stratospheric H_2O and CH_4 entry mixing ratios, and a constant correlation between FRF and AoA) on the estimated AoA trend are comparable in magnitude, however they depend on the exact period considered. Interestingly, the effects of the approximations can be opposite in sign cancelling out each other to some extent. Consequently, the approximation method can lead to a reliable estimation of AoA trends for certain periods, but this should not always be expected. A further improvement of the approximation method is proposed in the following Sect. 3.5.

3.5 Improved AoA trend estimation using parameterised age spectra

For estimating AoA trends from H_2O and CH_4 observations, the approximations (i) instantaneous entry mixing ratio propagation, and (ii) constant AoA-FRF correlation are necessary. Nevertheless, the higher reliability of AoA trends can be achieved if the used approximations are adjusted. As a simple and practical improvement, we propose to use an analytical, parameterised age spectrum for propagating stratospheric entry H_2O and CH_4 mixing ratios. Note that for a further improvement of AoA trend estimates, a non-stationary AoA-FRF relationship in principle would be needed as well. But due to the sparseness of available stratospheric CH_4 measurements deducing such a relationship from observations is challenging, and we refrain from including it in the methodological improvement.

In the following, we discuss the results of an additional sensitivity study with CLaMS stratospheric entry H_2O and CH_4 mixing ratios propagated by the parameterised idealised age spectrum, and using a constant AoA-FRF correlation. This method is hereinafter referred to as the "improved approximation method" (see Table 1 for details). In this method we use an inverse Gaussian distribution (e.g., Newman et al., 2007; Bönisch et al., 2009; Hauck et al., 2019) as a parameterised age spectrum

$$G(t,\Gamma) = \sqrt{\frac{\Gamma^3}{4\pi\Delta^2 t^3}} \cdot \exp\left(\frac{\Gamma\cdot(t-\Gamma)^2}{4\Delta^2 t}\right), \tag{7}$$

where Γ is the mean AoA and Δ is the width of the age spectrum. Here, we parameterise AoA in different zones or "regions" depending on the considered latitude, longitude and height. The finer the separation into different regions, the less pronounced the discontinuities at the edges of the regions are. For a simple and practical method without assuming a priori knowledge of model age of air, we propose to divide the stratosphere into seven regions, prescribing one mean value of AoA for each region (see Appendix B). We apply the empirical relation between the age spectrum width and AoA proposed by Hall and Plumb (1994) and used in several other studies (e.g., Engel et al., 2002; Bönisch et al., 2009)

$$\frac{\Delta^2}{\Gamma} = C, \tag{8}$$

with the constant $C = 0.7$ years, although we note that recent work of Hauck et al. (2019) suggests a larger value (2.0 years) for the lower stratosphere.

The resulting AoA trend estimated with the improved approximation method for 1990-2006 is shown in Fig. 6e. There is a clear improvement in the AoA trend estimation when compared to the pure approximation method APPROX (Fig. 6c). Note, that in both methods the same FRF distribution is used, but the FRF changes are different depending on the propagation method of stratospheric entry H_2O and CH_4. The discrepancies in estimated AoA trends between APPROX and the improved approximation method stem from the residual calculation of the third term of Eq. 6, $CH_{4\,[\mathrm{entry}]}(r,\,t)\;\Delta\alpha(r,\,t)$, with the major impact of the stratospheric entry H_2O trend (see Sect. 4). The propagation of stratospheric entry H_2O and CH_4 by the proposed parameterised idealised age spectrum results in AoA trends close to those from the C-CORR method (Fig. 6d, e), which is the best estimate possible for the improved approximation method due to the usage of constant AoA-FRF correlation.

Hence, we encourage the usage of the improved approximation method when estimating AoA from the combination of H_2O and CH_4 observational data. Stratospheric entry mixing ratio time series for H_2O and CH_4 can be deduced from satellite measurements, such as ACE-FTS, HALOE, MIPAS, or SCIAMACHY (e.g., Bernath, 2017; Russell III et al., 1993; Nassar et al., 2005; Raspollini et al., 2006; Scherer et al., 2008; Müller et al., 2016; Lossow et al., 2017; Noël et al., 2018). Due to the limited stratospheric CH_4 observations, the stratospheric entry CH_4 mixing ratio contribution can be kept as in APPROX (instantaneous propagation), because this term has only little effect on the resulting AoA trend (see Sect. 4). Our study shows that the usage of the parameterised idealised age spectrum clearly improves the representation of the stratospheric entry H_2O term, and hence the final AoA trend estimate.

4 Discussion

The contributions of stratospheric entry H_2O mixing ratio trends ($\Delta H_2O_{[\mathrm{entry}]}(r,\,t)$ in Eq. 6) calculated with APPROX, improved approximation and FULL methods, as well as the "true" stratospheric H_2O trend from CLaMS simulation are shown in Fig. 8a-d. Note that the stratospheric entry H_2O mixing ratio propagated by the CLaMS age spectrum is shown here only for comparison and is not used for AoA trend calculations in FULL method. Comparison of Fig. 8b and Fig. 8d shows that APPROX overestimates the stratospheric entry H_2O mixing ratio trend by approximately 3% per decade, especially in the middle and upper stratosphere. The propagation of stratospheric entry H_2O by the parameterised idealised age spectrum improves the representation of stratospheric entry H_2O trend (see Fig. 8c, d), and thus leads to a more reliable AoA trend estimation.

The contributions from stratospheric entry CH_4 mixing ratio trends ($2\alpha(r,\,t)\;\Delta CH_{4\,[\mathrm{entry}]}(r,\,t)$ in Eq. 6) are presented in Fig. 8e-g. The differences between the various stratospheric entry CH_4 trends are below 0.5% per decade. Although, the pattern of the stratospheric entry CH_4 mixing ratio trends is slightly improved when including propagation with the parameterised idealised age spectrum (Fig. 8f, g), overall, this term has a weak effect on the circulation contribution and the resulting AoA trend.

The circulation contributions ($2CH_{4\,[\mathrm{entry}]}(r,\,t)\;\Delta\alpha(r,\,t)$ in Eq. 6) in the APPROX and improved approximation methods are calculated as the residual between CLaMS H_2O trend (see Fig. 8a) and the other two components. As FULL is the most

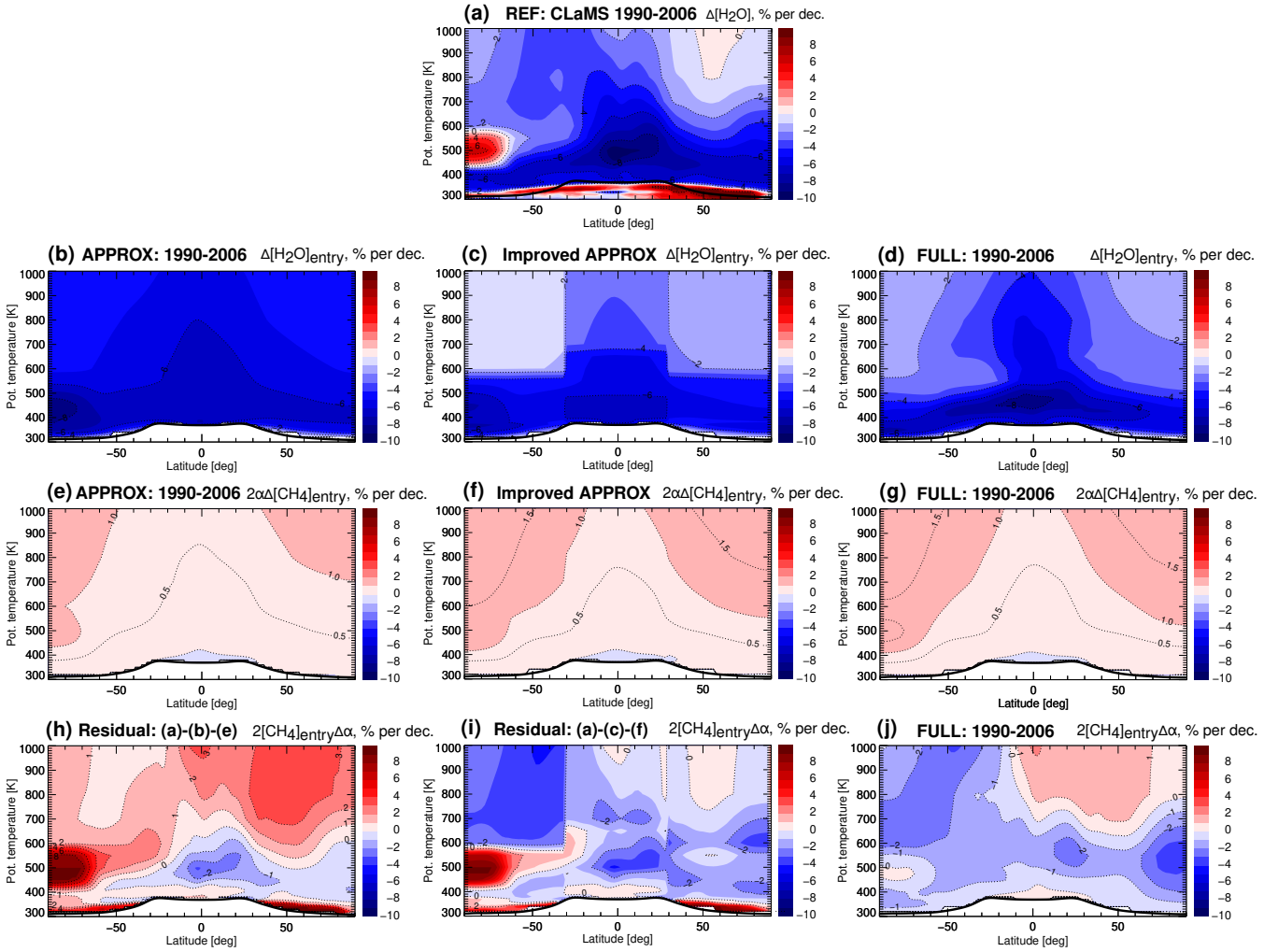

Figure 8. The "true" stratospheric H_2O 1990-2006 trend calculated from CLaMS simulations driven by ERA-Interim reanalysis is shown in (a). Further plots represent contributions to H_2O trend calculated through different methods: APPROX (b, e, h), the improved approximation (c, f, i), and FULL (d, g, j) methods. The contributions from circulation changes are calculated as the residual in APPROX and improved approximation methods (h, i), and as a linear trend in FULL (j). Note, that all sub-figures are presented in percentage per decade, with relation to the climatological mean 1990-2006 CLaMS stratospheric H_2O mixing ratio. The black line is the tropopause calculated from ERA-Interim.

exact method, the circulation terms from the approximate methods are evaluated against it. Figure 8h, j shows that the error in the estimated circulation contribution in APPROX is large, and the sign of the circulation contributions is even opposite in particular regions. Propagating stratospheric entry H_2O and CH_4 by the parameterised idealised age spectrum (Fig. 8i) improves the representation of circulation change significantly. From the comparison of Fig. 8i, j, we conclude that calculating the
440 circulation impact as a residual yields a reliable representation of the circulation contribution when an idealised parameterised

age spectrum is used. Large discrepancies still occur in the Antarctic region where the reconstruction method is expected to fail because of local dehydration processes. The polar dehydration in the Antarctic region has a substantial drying effect, reaching 1 ppmv and even more below 600 K potential temperature in the Southern hemisphere (Poshyvailo, 2020). This dehydration effect induces discrepancies in methods with the residual circulation term. Furthermore, the circulation contribution calculated as a residual depends also on the accuracy of the used stratospheric H_2O trend, shown in Fig. 8a.

To summarise, the differences between the residual circulation components of APPROX and improved approximation methods are caused by the discrepancies in the stratospheric entry H_2O and CH_4 mixing ratio contributions, with the major impact from the stratospheric entry H_2O trend. Consequently, the correct representation of the stratospheric entry H_2O mixing ratio trend is crucial for a reliable estimate of the AoA trend.

We showed that biases in AoA trends estimated from stratospheric H_2O can be large, questioning the usefulness of this approach. On the other hand, AoA estimates based on other trace gas species, like SF_6 or CO_2, show substantial uncertainties as well (e.g., Engel et al., 2009; Fritsch et al., 2020). An advantage of using H_2O is the existence of several long and homogenized records of satellite measurements of comparatively high quality. Furthermore, the significant bias reduction in estimated AoA trends, related to the relatively simple methodological improvement by using a parameterised idealised age spectrum for entry mixing ratio propagation, seems very promising. Further improvements could be realized by including a chemistry-dependent propagator instead of the idealised age spectrum (see Ostermöller et al., 2017), or by using an inversion algorithm for fitting the parameterised age spectrum (e.g., Hauck et al., 2019).

5 Conclusions

We investigated the effects of two commonly used approximations, (i) instantaneous stratospheric entry mixing ratio propagation and (ii) stationary AoA-FRF correlation, to estimate long-term BDC changes from the combination of stratospheric H_2O and CH_4 by deducing AoA trend. We carried out different sensitivity experiments within the CLaMS "model world": including both above mentioned approximations (APPROX), only the constant correlation approximation but representing the correct entry mixing ratio propagation (C-CORR), and representing the non-stationarity of the correlation as well as the entry mixing ratio propagation (FULL). Estimated AoA trends were compared to the actual CLaMS AoA trend. We considered as case studies the 1990-2006 and 1990-2017 periods.

The results show that both approximations have an important effect on the calculated AoA trend, leading to noticeable differences in the derived mean AoA trend compared to the "true" CLaMS AoA trend. The discrepancies in the AoA trends are up to 5 % per decade above 600 K, and more than 10 % below 600 K due to the applied approximations. Depending on the considered period, the effects from both approximations can also be opposite, and may even cancel out to some extent, producing, incidentally, an estimation of AoA trend remarkably close to the "true".

In order to increase the reliability of the derived AoA trend, we propose a simple and practical adjustment of the approximation (i) by propagating the stratospheric entry H_2O and CH_4 mixing ratios using a parameterised idealised age spectrum instead of instantaneous propagation. This refinement of the method considerably improves the reconstructed stratospheric entry H_2O

contribution, and, consequently, the derived AoA trend. The results of this article are of particular relevance for assessing the uncertainty in estimates of stratospheric circulation and BDC changes from global satellite measurements of stratospheric H_2O.

Appendix A: Stationary correlation function between FRF and modelled AoA

In order to estimate the AoA trend induced by the changes in stratospheric H_2O, we define a stationary relationship between zonally averaged FRF and AoA from CLaMS simulations driven by ERA-Interim reanalysis. We use CH_4 mixing ratios averaged over 2005-2006 for FRF calculations (see Eq. 5), following Hegglin et al. (2014). This specific period is characterised by relatively constant tropospheric CH_4 values. Accordingly, the dependence of stratospheric CH_4 entry mixing ratios on transit time can be neglected, and CH_4 can be assumed constant over these years. We assume that the stratospheric CH_4 during 2005-2006 was originated in the tropical tropopause layer during approximately 2002-2006. Thus, stratospheric entry CH_4 can be approximated as the mean CH_4 mixing ratio over 390-400 K and 30° S-30° N during 2002-2006 period. Consequently, FRF is stable in time and does not depend on the chosen period.

The correlation function of the CLaMS modelled AoA and estimated FRF function is presented in Fig. A1. It is derived by fitting a third-order polynomial to the AoA-FRF distribution, as suggested by Hegglin et al. (2014). The empirical relationship

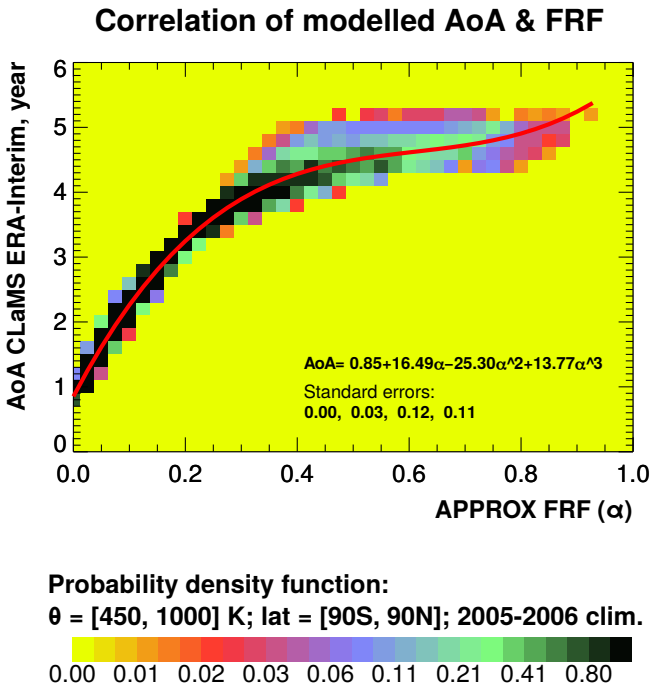

Figure A1. Relationship between CLaMS FRF and AoA for the methods of AoA trend estimation using stationary approximation. FRF is defined from the climatological 2005-2006 CLaMS simulated CH_4, AoA is the climatological 2005-2006 from CLaMS simulations. The considered correlated region is between 450-1000 K and 90° S-90° N. The colour bar represents the probability density function. The red line is a third-order polynomial fitting function.

between the CLaMS AoA and the FRF is $f(\alpha) = 0.85 + 16.49\alpha - 25.30\alpha^2 + 13.77\alpha^3$ (see Fig. A1). The same relationship is applied for any investigated period, if the method implies a stationary AoA-FRF correlation function.

Appendix B: Parameterisation of AoA for idealised age spectrum calculations

In the improved approximation method, we use parameterised idealised age spectrum to propagate stratospheric entry H_2O and CH_4 mixing ratios for the AoA trend estimation. The idealised age spectrum requires the width of the age spectrum that is defined from the empirical relation Eq. 8. Here, we make a simple parameterisation of AoA (Γ) required for the idealised age spectrum. We separate the stratosphere (up to 1000 K or approximately 37 km) in seven zones, see Fig. B1. However, further improvement by other zone divisions is, in principle, possible. A better spatial resolution of propagated H_2O, CH_4 by idealised age spectrum, and consequently estimated AoA could be gained by using more zones or even by assigning different shape for zones (e.g. triangles).

In the latitudinal directions, the middle zone between 30° S to 30° N is associated with the tropical pipe region. In the height-direction at the units of potential temperature, the lowest AoA value of 1 year is located just above the tropical tropopause layer between 400-500 K (or about 15-20 km). The mean value of AoA equaled to 2.5 years, we assign to three regions: in both hemispheres between 380-600 K (approx. 12-24 km), and between 500-700 K (approx. 20-28 km) at the tropics. The averaged AoA value of 3.5 years is at the tropical middle stratosphere from 700 to 1000 K (approx. 28-37 km). The averaged AoA value is different for both hemispheres at the region between 600-1000 K (approx. 24-37 km): 4.5 years for the Southern and 4.0 years for the Northern hemispheres. Such AoA asymmetry exists, since the deep branch of the circulation is stronger in the Northern hemisphere and thus causes younger AoA (Butchart, 2014; Konopka et al., 2015).

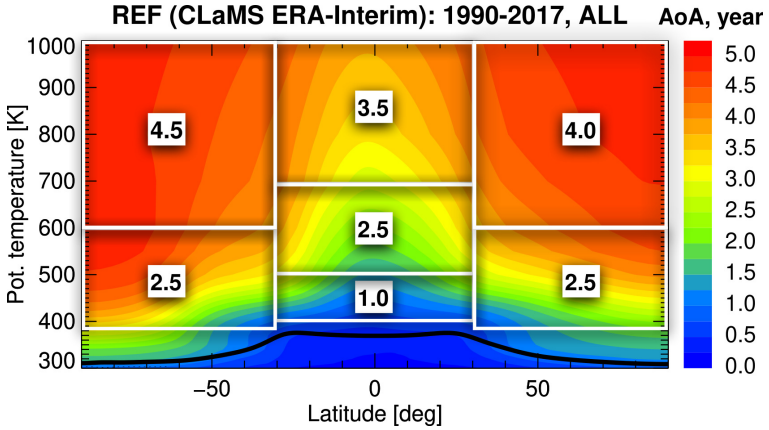

Figure B1. Zonal mean of AoA from CLaMS simulation driven by ERA-Interim reanalysis. Data shown are climatology for 1990-2017 years. Regions defined with the white boxes are used for the idealised age spectrum calculations in the improved approximation method. Values denoted in each box is an approximate averaged AoA value of the region bounded by white box.

Code and data availability. The CLaMS code used in this article is available on theGitLab server: https://jugit.fz-juelich.de/clams/CLaMS.

Author contributions. The study was designed by F.P. and L.P-S. with contributions by S.F. and R.M.. L.P.-S. and F.P. conducted the model runs and analysed the results. Partially, the work is based on the original concept from the study of M.H.. L.P.-S. wrote the paper with further inputs from F.P. and R.M. All co-authors contributed to the interpretation of the results, active discussions, and the revision of the paper.

Competing interests. R.M. is an editor of ACP, otherwise there are no competing interests.

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