1	Siberian Arctic black carbon: gas flaring and wildfire impact
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18 Abstract

19 As explained in the latest Arctic Monitoring and Assessment Programme (AMAP) report released in early 2021, the Arctic has warmed three times more quickly than the planet as a whole, 20 and faster than previously thought. The Siberian Arctic is of great interest largely because 21 22 observations are sparse or largely lacking. A research aerosol station has been developed on the Bely Island (Kara Sea) in Western Siberia. Measurements of equivalent black carbon (EBC) 23 concentrations were carried out at the "Island Bely" station continuously from August 2019 to 24 November 2020. The source origin of the measured EBC, and the main contributing sources were 25 assessed using atmospheric transport modelling coupled with the most updated emission 26 inventories for anthropogenic and biomass burning sources of BC. 27

The obtained climatology for BC during the period of measurements showed an apparent 28 seasonal variation with the highest concentrations between December and April (60±92 ng/m3) 29 and the lowest between June and September (18 ± 72 ng/m³), typical of the Arctic Haze seasonality 30 31 reported elsewhere. When air masses arrived at the station through the biggest oil and gas extraction regions of Kazakhstan, Volga-Ural, Komi, Nenets and Western Siberia, BC contribution 32 from gas flaring dominated over domestic, industrial, and traffic sectors, ranging from 47 to 68%, 33 with a maximum contribution in January. When air was transported from Europe during the cold 34 season, emissions from transportation were more important. Accordingly, shipping emissions 35 increased due to the touristic cruise activities and the ice retreat in summertime. Biomass burning 36 (BB) played the biggest role between April and October, contributing 81% at maximum in July. 37 Long-range transport of BB aerosols appeared to induce large variability to the Absorption 38 Ångström Exponent (AAE) with values ranging >1.0 (excluding outliers). As regards to the 39 continental contribution to surface BC at the "Island Bely" station, Russian emissions dominated 40 41 during the whole year, while European and Asian ones contributed up to 20% in the cold period. Quantification of several pollution episodes showed an increasing trend in surface concentrations 42 43 and frequency during the cold period, as the station is directly in the Siberian gateway of the highest anthropogenic pollution sources to the Russian Arctic. 44

46 **1** Introduction

47 Global carbon pollution is annually produced by burning of fossil fuel and biomass. Combustion emissions are increasingly recognized as an important source of chemically active 48 aerosols. Black carbon (BC) originates from the incomplete combustion of fossil fuels and biomass 49 burning; it is a short-lived climate forcer and absorbs incoming solar radiation and, therefore, is of 50 high significance for the Arctic climate (Wang et al., 2011). The combined total effects of BC and 51 sulfates cause an Arctic surface warming of +0.29°K explaining approximately 20% of the 52 observed Arctic warming since the early 1980s (Ren et al., 2020). BC resides in the lowest 53 atmospheric layer, affects aerosol-cloud interactions (Yun et al., 2013) and has a cloud and sea-54 ice feedback when deposited (Flanner, 2013), thus accelerating melting (Quinn et al., 2008). 55

Long-range transport to the Arctic carries, among other aerosol constituents, many tracers 56 of anthropogenic and wildfire origin (Chang et al., 2011). (Winiger et al., 2016) showed that BC 57 in Arctic Scandinavia is predominantly linked to emissions in Europe. Over the whole Arctic 58 region (north of 66°N), Russia contributes 62% to surface BC (Zhu et al., 2020). Industrial and 59 residential sources are responsible for the highest measured BC concentrations at the Tiksi station 60 (Siberian Arctic) (Popovicheva et al., 2019b). (Stathopoulos et al., 2021) have demonstrated that 61 the long term impact of light absorbing carbon in the high Arctic is three times higher in the cold 62 period of the year compared to the warm period. There, fossil sources mostly prevail during winter-63 spring season, while biomass burning sources dominate during low BC concentration periods in 64 summer (Winiger et al., 2017). Although BC dominates light absorption by atmospheric aerosols, 65 other carbonaceous aerosol species (brown carbon, BrC) represent an important fraction of light 66 absorption in the UV and near-UV spectrum, thus having an important role in the assessment of 67 radiative forcing in the Arctic climate. Spectral dependence of the light absorption is generally 68 described by the Absorption Ångström Exponent (AAE), which is typically used to differentiate 69 between aerosol types (BC, BrC) and sources of BC (Sandradewi et al., 2008; Helin et al., 2021; 70 Zotter et al., 2017). 71

Quantification of the particulate Arctic pollution is a serious problem worldwide; reliable source emission inventories are challenged, and regional contributions of BC sources in the Arctic are still inconclusive (Zhu et al., 2020). Global anthropogenic emission dataset ECLIPSEv6 (Evaluating the Climate and Air Quality Impacts of Short-lived Pollutants) using the GAINS model (Klimont et al., 2017) includes all major economic sectors, such as energy and industrial production, transport, residential combustion, agriculture, and waste, distinguishing between sector-fuel-technology, fuels, and emission control options. The model predictions for European gateway to the Arctic were greatly improved, when the emission inventory from anthropogenic
sources was updated by estimates of European BC emissions (Winiger et al., 2016).

Due to large size and continuous production, gas flaring of oil industry is one of the highest 81 BC emission sources (Ismail and Umukoro, 2012) with a strong environmental and climatic impact 82 for the Arctic (Cho et al., 2019). Flaring in ECLIPSEv6 dominates BC emissions in the Arctic; 83 models have found that flaring contributes 42% to the annual mean BC surface concentrations in 84 the Arctic (Stohl et al., 2013). However, because flares are difficult to measure, their particulate 85 emissions and physicochemical properties are still underestimated (Conrad and Johnson, 2017; 86 Popovicheva et al., 2019a). Currently, models are struggling to reproduce BC concentrations 87 88 largely due to emission-related uncertainties in the Arctic region (Schacht et al., 2019). The observed annual mean contribution of fossil fuel combustion to the Arctic concentrations agrees 89 within a factor of two (Qi and Wang, 2019). 90

High latitude flaring emissions mainly originate from the North Sea, Norwegian Sea, the 91 92 northeastern part of European Russia (Komi Republic) and Western Siberia. The largest oil and 93 gas producing regions of northwestern Siberia are located along the main low-level pathway of air masses entering the Arctic and thus have a disproportionally large contribution to the Arctic lower 94 troposphere (Stohl, 2006). (Eleftheriadis et al., 2009) and (Tunved et al., 2013) identified these 95 regions as a key source for the highest measured BC concentrations and sub-micrometer aerosol 96 mass concentrations, respectively, at the Zeppelin station. The impact of BC long-range transport 97 98 from northwestern Siberia was also observed at the Ice Base Cape Baranova station located on Severnaya Zemlya archipelago (Eastern Siberia) (Manousakas et al., 2020). Accordingly, possible 99 gas flaring impact was observed at the Tiksi station (northeastern Siberia), despite the large 100 101 distance of the station from the largest oil and gas producing regions (Winiger et al., 2017). To 102 better understand and quantify the contribution of gas flaring to the Arctic environment, targeted aerosol and atmospheric composition measurements at the closest distance from the flaring 103 facilities are needed. The present operating Eurasian Arctic stations are all too far away to allow 104 assessing how air masses are affected by gas flares or what the contribution from different source 105 categories is (Stohl et al., 2013). Simulations combined with observations of BC at the proximity 106 of the source regions (e.g., the plumes from gas flaring regions over the Kara Sea) provide a better 107 constraint (Popovicheva et al., 2017). In addition, measurements of BC coupled by conditional 108 probability simulations performed inside the oil and gas producing region of the northwestern 109 Siberia have successfully distinguished between multiple industrial and urban sources 110 (Popovicheva et al., 2020). 111

112 Recent efforts have sought to develop a new Russian BC emission inventory (BCRUS) for the Siberian Arctic, based on activity data from local information, improved spatial distribution of 113 BC emissions and updated emission factors for oil and gas fields in northwestern Siberia (Huang 114 115 et al., 2015). According to this, it was found that BC emissions from gas flaring account for 36% 116 of the total anthropogenic BC emissions over Russia. Residential BC emissions, transportation, industry, and power plants contribute 25%, 20%, 13%, and 5.4%, respectively. The emissions from 117 gas flaring in BCRUS show a discrepancy of more than 40% higher than ECLIPSEv5. Using 118 BCRUS, modelled surface BC at Zeppelin, Barrow, and Alert stations were basically improved 119 (Huang et al., 2015). The contribution of anthropogenic emissions in Russia to the annual total 120 Arctic surface BC were calculated to be 56%, with gas flaring from the Yamalo-Nenets 121 Autonomous Okrug (YNAO), Khanty-Mansiysk Autonomous Okrug (KMAO), and Komi 122 Republic to be the main source (31% of Arctic surface BC) (Zhu et al., 2020). However, due to 123 the absence of BC inventories for industrial emissions and a denser observational network in the 124 125 western Siberian High Arctic, the spatial distribution of BC sources is still associated with large 126 uncertainties.

Agricultural fires in East Europe and North America are a major source of biomass burning 127 128 in the Eurasian Arctic (Treffeisen et al., 2007; Stohl et al., 2007; Stohl et al., 2006). Springtime fires in Siberia can double the North American Arctic background (Warneke et al., 2010). Long-129 130 term airborne observations of BC in Northern Siberia have revealed a strong impact from forest fires in summer (Kozlov et al., 2016; Paris et al., 2009). Particulate BrC emitted by intensive 131 wildfires has been measured in plumes transported for over two days (Forrister et al., 2015). In 132 summer 2019, wildfire activity in Central and Eastern Siberia occurred along the trans-Arctic 133 134 transport pathway of Siberian biomass burning emissions resulting in enhanced aerosol lamina observed in Western Canada (Johnson et al., 2021). 135

In 2019, a new aerosol station was developed by Moscow State University on the Bely Island 136 located in the Kara Sea (Western Siberian Arctic) (https://peexhq.home.blog/2019/12/11/new-137 research-aerosol-stations-in-the-russian-arctic), (Figure 1). The region was chosen because it is 138 close to the air pathway of large-scale emission plumes from populated industrial regions of 139 140 Eurasia and Siberian wildfires to the Arctic. We present here the ground-based continuous BC (equivalent BC, EBC) measurements from August 2019 until November 2020 at the "Island Bely" 141 station for the first time. The Arctic annual trends of BC are assessed, while the geospatial source 142 origin of the air arriving at the station is identified using a Lagrangian particle dispersion model. 143 Furthermore, the anthropogenic and biomass burning contributions to the modelled surface 144 concentrations of BC are evaluated against measured BC concentrations at the station. 145

146 Characterization of the pollution events in cold and warm periods separates the impact of gas 147 flaring versus biomass burning. In addition, the spectrally resolved absorption measurements 148 provide an opportunity for the characterization of BC sources. The present study assesses long 149 range transport of BC to the Western Siberian Arctic from the main large-scale emission regions 150 of the Eurasian continent using Lagrangian modelling coupled with continuous observations.

151 **2 Experimental**

152 **2.1 Aerosol station "Island Bely".**

Western Siberia is the world's largest gas flaring region with a leading oil and gas production 153 industry (Figure 1). YNAO is located north of the West Siberian Plain and covers a vast area of 154 769 thousand km². More than 94% of the region's economy is associated with industrial 155 applications related to the extraction of fuels, their processing, and transportation. Specifically, 156 157 YNAO has the largest reserves of Russia's natural gas and oil; YNAO emissions of BC are the largest in the Russian territory (Vinogradova, 2015). The relative contributions from gas flaring to 158 159 annual mean BC surface concentrations from all emission sources (surface transportation, industry, residential, biomass burning) exceed 70% (Stohl et al., 2013). 160

161 The Bely Island is located in the Kara Sea, north of the YNAO (Figure 1). For the purpose of atmospheric composition observations and sampling at the "Island Bely" station, the aerosol 162 pavilion has been built approximately half a km to the southeast of the Roshydromet 163 164 meteorological station continuously operating at the island (Figure 1). There are no other anthropogenic constructions on the island. Thus, the major advantage of a new-developed research 165 station is its long distance from any local anthropogenic sources. Previous research at the Tiksi 166 station has shown significant aerosol pollution from local sources (Popovicheva et al., 2019), 167 which is not the case in the "Island Bely" station. An aerosol sampling system was installed at the 168 aerosol pavilion in May 2019. Three total suspended particle (TSP) inlets were installed 169 approximately 1.5 m above the roof and 4 m above the ground. One is used for the real-time BC 170 171 monitoring with air flow 5 L/min, and two for aerosol chemical characterization operating with 2.3 m³/h flow. The TSP inlet is equipped with an electric heating wire to prevent rimming and ice 172 blocking of the system. 173

An Aethalometer model AE33 (Magee Scientific,Aerosol d.o.o.) was used to measure the light attenuation caused by particles depositing on two filter spots at different flow rates (Drinovec et al., 2015) and at seven wavelengths from ultraviolet (370 nm) to infrared (950 nm). The "dual spot" technique is applied for real-time loading effect compensation. The light-absorbing content of carbonaceous aerosol at 880 nm is reported as equivalent black carbon concentration (EBC), which is determined for each time interval from the change in the light attenuation at a wavelength
of 880, using a mass absorption cross-section of 7.7 m²/g and filter multiple scattering parameter
C of 1.57. Light absorbing organic components (BrC) absorb light at shorter wavelengths more
effectively than at 880 nm, which is observed as an increased AAE (Sandradewi et al., 2008;
Grange et al., 2020; Helin et al., 2021). AAE was calculated using Eq. 1 for 470 nm and 950 nm
wavelengths:

$$AAE = \frac{\ln (babs(470)/babs(950))}{\ln (950/470)}$$
(1)

185

186 where b_{abs} stands for the absorption coefficient at 470 nm and 950 nm. In order to avoid 187 instrumental noise when calculating the AAE, the following data processing was implemented. 188 One-minute absorption coefficients for the whole period were averaged to 1 hour. The dataset was 189 filtered to periods when EBC exceeded 20 ng/m³ (sensitivity level at 1 hour time resolution), and 190 then the AAE was averaged to 3 hours.

191 The Aethalometer model (Sandradewi et al., 2008) is typically used for the source apportionment of EBC, when measurements of absorption coefficient are performed by filter 192 193 photometers. The model uses an apriori assumed pair of AAE for traffic (AAETR) and biomass burning (AAEBB) to determine the contribution of both sources. Although the Aethalometer 194 195 model is an efficient tool for source apportionment of EBC in a well-mixed urban atmosphere, where two sources with distinct aerosol optical properties prevail (fossil fuel from traffic and fresh 196 197 biomass burning), the results can be affected when the characteristic optical properties of a specific source change over time. This is usually the case for wildfires, where different burning modes 198 (flaming or smoldering) and different types of wood can significantly influence BrC emissions and 199 its chemical composition (Kalogridis et al., 2018b). Furthermore, chemical evolution after 200 201 emissions and atmospheric aging (i.e. aerosol mixing state, particle morphology and size distribution) additionally influence aerosol absorption, which can be noticed especially for long-202 range transported air masses (Cappa et al., 2016; Saleh et al., 2013; Romshoo et al., 2021). 203 204 (Forrister et al., 2015) have shown that BrC emitted from wildfires was highly unstable, with 6% 205 of BrC remaining above background levels after two days.

BC measurements at the "Island Bely" station were performed from 10 August 2019 to 30 November 2020, with a time resolution of 1 min. Basic meteorological parameters, such as temperature, wind speed and direction were obtained every 3 hours from a meteorological station located 500 m from the "Island Bely" station. Cleaning of 1-min time resolved BC data was based on the definition of what can be considered as a peak: strong fast increase of BC value a few times higher than previous value and then a similar fast decrease. Analysis of meteorological parameters

was complementary to check whether the wind originated from the sector corresponding to 212 locations of diesel generators at the Roshydromet meteorological station (240-250 degrees from 213 the "Bely Island" station). In such cases, strong peaks of BC were removed from further analysis. 214 The total duration of the peaks under the influence of local contamination varies from several 215 216 minutes to one-two hours per day and monthly. For instance, in January 2020, when wind blew from 240-250 degrees for 24 hours in total, large peaks corresponding to a fraction of 13% of the 217 data were removed. In July 2020, when wind originated from the same location for a total of 45 218 hours, measurements corresponding to a fraction of 22% of the data were removed. Lack of windy 219 weather at Bely Island is a very rare event, only 0.7% of the observation time; during such weather, 220 peaks of BC were never observed. 221

222 **2.2 Atmospheric transport model coupling with emissions**

To investigate the possible origin of BC, the Lagrangian particle dispersion model 223 FLEXPART (FLEXible PARTicle dispersion model) version 10.4 was used (Pisso et al., 2019). 224 225 The model was driven by 3-h operational meteorological fields from the European Centre for Medium-Range Weather Forecasts (ECMWF) with 137 vertical levels and a horizontal resolution 226 of 1°×1°. In FLEXPART, computational particles were released at height 0 - 100 m from the 227 receptor ("Island Bely" station) and were tracked backward in time in FLEXPART's "retroplume" 228 229 mode. Simulations extended over 30 days backward in time, sufficient to include most BC emissions arriving at the station, given a typical BC lifetime of 1 week (Bond et al., 2013). 230

The tracking includes gravitational settling for spherical particles of the size observed. FLEXPART differs from trajectory models due to its ability to simulate dry and wet deposition of gases or aerosols (Grythe et al., 2017), turbulence (Cassiani et al., 2015), unresolved mesoscale motions (Stohl et al., 2005), while it includes a deep convection scheme (Forster et al., 2007). For our simulations, we assumed that BC has a density of 1500 kg m⁻³ and follows a logarithmic size distribution with an aerodynamic mean diameter of 0.25 μ m and a logarithmic standard deviation of 0.3 (Long et al., 2013).

FLEXPART simulations were performed every 3 hours during the studied period. The FLEXPART output consists of a footprint emission sensitivity which results in a modelled concentration at the receptor, when coupled with gridded emissions from an emission inventory. The emission sensitivity expresses the probability of any release occurring in each grid-cell to reach the receptor. The source contributions to receptor BC were derived by combining each gridded emission sector (e.g. gas flaring, transportation, waste management etc...) from an emission inventory with the footprint emission sensitivity. Calculations for anthropogenic sources (emission sectors are described below) and open biomass burning were performed separately. This
enabled identification of the exact origin of BC and allowed for quantification of its source
contribution. The modelled concentrations can also be displayed as a function of the time elapsed
since the emission has occurred (i.e., "age"), which can be shown as "age spectrum".

249 In this study, anthropogenic emission fluxes were adopted from the lastest version (v6b) of the ECLIPSE (Evaluating the CLimate and Air Quality ImPacts of ShortlivEd Pollutants) dataset, 250 an upgraded version of the previous version (Klimont et al., 2017). The inventory includes 251 252 industrial combustion (IND) emissions from industrial boilers and industrial production processes. Energy production sector (ENE) includes emissions from combustion processes in power plants 253 and generators. Residential and commercial sector (DOM) includes emissions from combustion in 254 heating and cooking stoves and boilers in households and public and commercial buildings. Waste 255 256 treatment and disposal sector (WST) resembles emissions from waste incineration and treatment. Transport sector (TRA) includes emissions from all land-based transport of goods, animals and 257 258 persons on road networks and off-road activities. Emissions from shipping activities in in-land waters (SHP) is included as a separate sector. Gas flaring (FLR) sector includes emissions from 259 260 oil and gas facilities. The methodology for obtaining emissions from FLR specifically over Russia has been improved in ECLIPSEv6 (Böttcher et al., 2021). Updates were based on new field-type 261 262 specific emission factors that were applied to VIIRS observations of the flared gas volume at individual flaring locations. For comparison, BCRUS emissions for the FLR sector (Huang et al., 263 264 2015) were also used in this study.

Emissions from biomass burning (BB) were adopted from Copernicus Atmosphere 265 Monitoring Services (CAMS) Global Fire Assimilation System (GFAS). CAMS GFAS assimilates 266 fire radiative power (FRP) observations from satellite-based sensors converting the energy 267 released during fire combustion into gases and aerosol daily fluxes (Di Giuseppe et al., 2016; 268 269 Kaiser et al., 2012). Data are available globally on a regular grid with a horizontal resolution of 0.1 degrees from 2003 to the present. FRP observations assimilated in GFAS are the NASA Terra 270 MODIS and Aqua MODIS active fire products (http://modis-fire.umd.edu/, (Kaufman et al., 271 272 2003). FRP measures the heat power emitted by fires, as a result of the combustion process and is 273 directly related to the total biomass combusted (Wooster et al., 2005). Using land-use-dependent conversion factors, GFAS converts FRP into emission estimates for 44 smoke constituents (Kaiser 274 et al., 2012), one of which is BC. 275

Biomass burning emissions were also adopted from the Global Fire Emission Dataset version 4.1 (GFEDv4.1). The product combines satellite information on fire activity and vegetation productivity to estimate gridded monthly burned area and fire emissions, as well as 279 scalars that can be used to calculate higher temporal resolution emissions. All data are publicly 280 available for use in large-scale atmospheric and biogeochemical studies and include (i) burned area (Giglio et al., 2013), (ii) burned area from "small" fires based on active fire detections outside 281 the burned area maps detailed in (Randerson et al., 2012) and updates in (Werf et al., 2017), (iii) 282 283 carbon and dry matter emissions from van der Werf et al. (2017), (iv) fractional contributions of various fire types to total emissions and (v) list of emission factors to compute trace gas and aerosol 284 emissions based on (Akagi et al., 2011) and (Andreae and Merlet, 2001). The current version (v4) 285 has a spatial resolution of 0.25 degrees and is available from 1997 onwards. 286

In the present paper, several different configurations were used to calculated modelled surface BC concentrations at "Island Bely" station, namely ECLIPSEv6 with GFED4 (ECLIPSEv6-GFED4), and ECLIPSEv6 with CAMS (ECLIPSEv6-CAMS). The same two configurations were also used after substituting the FLR emissions in ECLIPSEv6 with those from BCRUS (Huang et al., 2015).

292 **3 Results and discussion**

3.1 Monthly climatology of black carbon

294 The climate at the Bely Island is characterized by an average annual temperature of -8°C, precipitation of 450 mm, and stable snow coverage from October to May. Meteorology displays a 295 296 large annual variability determined by alternating periods of the polar night and midnight sun. 297 Median temperatures stay above 0°C for 4 months each year between June and September. This period is also characterized by the highest-frequency occurrence of ocean air masses and the most 298 stable wind speeds. A shift occurred in October with decreased solar insolation resulting in a 299 temperature shift to below 0°C. The cold month winds were primarily continental, with a low-300 frequency occurrence of ocean air masses. 301

The cycle of temperature and wind speed variations observed during the study period is 302 shown in Figure 2a,b. The period from 1 November 2019 to 1 April 2020, when temperature 303 304 dropped below -10 °C, as well as November 2020, is denoted as the "cold period". The remaining period of our study, from 10 August to 31 October 2019 as well as from April to 1 November 305 2020, is considered as the "warm period". Figure 2c illustrates the long-term time series of 24h 306 median EBC concentrations measured at wavelength of 880 nm (EBC(880)) during the study 307 period, with median value 37 ± 64 ng/m³ (maximum: 520 ng/m³, minimum: 6 ng/m³). The polar 308 frequency plot of wind speed/direction shows that the maximum number of hours the wind was 309 from north-east and south-west directions with around 5 m/s (Figure 3a). BC concentration roses 310

in Figure 3 indicate the sources of the highest concentrations, which originated from the continentin both cold and warm periods.

Figure 4 illustrates a long-term time series of monthly EBC concentrations at the "Island 313 Bely" station during the period from August 2019 to November 2020. The highest concentrations 314 were observed from November to April and the lowest ones from June to August, in agreement 315 with the typical seasonal trend of the Arctic aerosol concentrations (Stone et al., 2014). EBC 316 monthly climatology during the study period is shown in Figure 4a in terms of the median, and 317 upper and lower quartiles. For winter months, the maximum median EBC concentration was 165 318 ng/m³ observed in December 2019. The increase of the Arctic concentrations in winter, known as 319 the Arctic Haze, was more pronounced in November-December 2019 and January-March 2020. 320 On average, concentrations in summer were about 10 times lower than those in winter, with a 321 minimum median value of 30 ng/m³ in July 2020. Observations at the "Island Bely" station for the 322 second year started from August 2020 and lasted to November 2020 to confirm the general annual 323 324 trend of low summer and high winter BC concentrations. However, monthly median EBC in September 2020 demonstrated a value of 30.7 ng/m³. 325

326 Similar annual trend was recorded in 2015-2016 at the Tiksi station (coast of Laptev sea), with high concentrations reaching 130 ng/m³ during winter-spring and low concentrations of about 327 328 20 ng/m³ observed from May to October (Popovicheva et al., 2019b). As shown by earlier studies 329 at various polar stations, such as in Ny-Ålesund, Alert, and Barrow, aerosols display Arctic Haze peak concentrations during winter and early spring months (Stone et al., 2014). EBC during Arctic 330 Haze at both "Island Bely" and Tiksi stations are typically higher as compared to those observed 331 at Alert (100±65 ng/m³), a station that has shown the largest concentrations among all Polar 332 stations (Sharma et al., 2004). The latter confirms previous findings from (Eckhardt et al., 2015) 333 and (Winiger et al., 2017) that the Siberian Arctic is mainly polluted as a result of the influence 334 335 from emissions occurring in the Eurasian continent.

Near-surface measurements allow for evaluation of the capability of transport models to 336 337 reproduce the distribution of BC in the Arctic based on different emission datasets (Schacht et al., 2019; Zhu et al., 2020). Figure 4a and Supplentary Table S 2 show observed and modelled BC 338 339 monthly median mass concentrations at the "Island Bely" station. Use of ECLIPSEv6 emissions caused overestimations of modelled BC concentrations of up to 46% (February 2020). All 340 simulated BC concentrations were found in the range between the 25th and 75th percentiles of 341 measured EBC. Modelled BC is underestimated in March-May 2020, being 29 ng/m³ below the 342 25th percentile of EBC in April 2020. When FLR emissions in ECLIPSEv6 were substituted by 343 BCRUS FLR, similar modelled BC monthly median concentrations were calculated, thus 344

indicating that other sectorial emissions might have a large contribution to surface BC at thestation.

Figure 4b shows the so-called "age spectrum" of modelled BC for the "Island Bely" 347 station. In the cold period of high EBC concentrations, the longest age of more than 19 days back, 348 affects up to 60% of the surface concentrations. In this time, due to the geographical proximity, 349 Russia dominates. However, both Europe and Asia contribute around 20% to the monthly averaged 350 surface BC, with the largest contribution to be in February 2019 and November 2020, (Figure 351 352 4b,d). The most aged air masses (from 28 to 30 days back) contributed up to 50%, arriving at the "Island Bely" station in December 2019, which is the month of the highest observed EBC 353 concentrations during the study period. The impact of the closest regions with age between 7 and 354 9 days is more significant in the winter months, while in the warm period, such short-term 355 356 contributions become negligible. The calculated age and continental spectrum of BC obtained for the "Island Bely" station mainly denote the variability of airmass transport patterns in different 357 358 seasons. In the cold season, the Siberian Arctic tends to force the air from south to north into the Arctic (Stohl, 2006), thus bringing more anthropogenic BC from highly populated regions. 359

360 Monthly averaged BC contributions from different sources simulated by FLEXPART using ECLIPSEv6 emissions are shown in Figure 4c and Supplementary Table S 3. From 361 362 November 2019 to March 2020 the FLR sector contributed 47%–68% (maximum in January 2020) 363 to surface BC, when air masses arrived at Bely through oil and gas extraction sites. February and November 2020 demonstrated the biggest non-gas flaring impact. More specifically, February 364 2020 coincides with the largest model overestimation (Figure 4a) implying a likely misestimation 365 of non-gas flaring emissions in ECLIPSEv6. From April 2020 the impact of FLR dropped 366 significantly (Supplementary Table S 4), with minimum of 12% in June. Starting from April to 367 October 2020, BB emissions played the biggest role in surface BC contributing 81% in July 2020. 368 It is noteworthy that the impact of SHP emissions became quite perceptible in the warm period, 369 when the oceanic ice is absent in the Arctic and touristic cruises peak. 370

371 Emission sensitivities of surface BC presented over the whole Arctic (north of 66°N) have been also simulated using the same model (Zhu et al., 2020). Anthropogenic sources contributed 372 373 82% of the annual BC, as estimated from BCRUS emission dataset. Arctic BC originated 374 predominantly from anthropogenic emissions in Russia (56%), mainly FLR from YNAO, Khanty-375 Mansiysk Autonomous Okrug (KMAO), and Komi Republic (31% of surface Arctic BC). In summer (July-August), open BB in Siberia, Alaska, and Canada contributed 75%. At Zeppelin, 376 377 modelled BC (39.1 ng/m³ for annual mean) was reported to be 85% higher than the observed value (21.1 ng/m³ for annual mean) (Zhu et al., 2020). At Tiksi, modelled BC was underestimated (74.4 378

 ng/m^3 for annual mean) by 40% compared with observations (104.2 ng/m³ for annual mean) (Zhu 379 et al., 2020). Annual (from September 2019 to August 2020) median modelled concentrations of 380 BC using ECLIPSEv6, BCRUS, and CAMS for the "Island Bely" station are shown in 381 Supplementary Table S 2. We find that modelled BC (78.4 ng/m³ for annual mean) is 26% higher 382 383 than the observed value (61.8 ng/m^3 for annual mean); the overestimation is much smaller than observed for other remote stations. Annual averaged contributions of anthropogenic emissions by 384 ECLIPSEv6 and ECLIPSEv6 with flaring from BCRUS were equal to 76% and 80%, respectively, 385 due to the difference in FLR emissions in the two datasets (Supplementary Table S 4). 386

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3.2 Cold season pollution

388 Figure 5a shows EBC concentrations measured at the "Island Bely" station during the cold period, from November 2019 to April 2020 and from 1st to 30th November 2020. Time series 389 indicates that EBC undergoes the Arctic-typical seasonal trend with higher concentrations in 390 winter and early spring and lower in summer. Background pollutant concentrations in Arctic 391 392 stations are generally very low without any detectable influence from local or regional sources (Eleftheriadis et al., 2004; Popovicheva et al., 2019b). We relate the Arctic background to the 393 lowest 20th percentile of EBC data (10 ng/m³). Long-term pollution episodes were assumed to be 394 repeated events of high EBC concentration above the 80th percentiles (90 ng/m³) that are clearly 395 396 distinguishable from the background (Figure 5a).

397 The aerosol optical properties with respect to absorption, presented as daily median AAE is shown in Figure 2d. The AAE for highly aged aerosols measured during periods of low BC was 398 lower than 1 (reaching values as low as 0.2) and is mostly related to the aerosol size distribution 399 400 (large particles) and internally mixed BC particles (Cappa et al., 2016). As shown by modeling studies (Virkkula, 2021), pure BC particles surrounded by non-absorbing coatings can have AAE 401 in the range from <1 to 1.7, also depending on the morphology of the fractal aggregates (Romshoo 402 et al., 2021). The AAE increased in periods of higher aerosol concentration levels in the cold 403 period ranging from 0.6 to 1.35. 404

In many cases, when AAE exceeded 1 in the cold period, the pollution episodes could be identified as influenced by BB. However, due to the mixing with background aerosol and aging processes, a large variability in AAE values might be observed at receptors of long range transported pollution and AAE may not be representative of BB sources. Nevertheless, it can still be used as a qualitative parameter, when extra information is available. Such events of increased AAE were rarely observed in our study, and the most prominent BB impact occurred during the pollution episodes C4, C7, and C8 when impact of domestic sources was significant (**Figure 5**a). In general, FLEXPART coupled with ECLIPSEv6-CAMS emissions captures periods with both high and low concentrations relatively well, (**Figure 2**c). A good correlation between measurements and simulations, with a Pearson coefficient R of 0.7 and the root mean squared error (RMSE) of 89.2 ng/m³ wasobtained for the cold period (**Figure 6**a). According to monthly median contributions to BC concentrations in the cold period, the impact of anthropogenic sources, namely FLR, DOM and TRA dominated surface BC by 97.7% (**Figure 4**c).

418 Looking closely to specific episodes, during pollution episode C1, three events of high 419 EBC concentrations were observed, (Figure 5a). On 5 November 2019, measured EBC reached 180 ng/m³, while FLEXPART simulated similarly high BC values. Footprint emission sensitivities 420 at this time showed that airmasses originated from East and North Europe, passed south of 421 European Russia, then turned the direction straight through the West Siberia approaching the Bely 422 423 Island from the southeast (Figure 7). The same airmassed moved towards the large Russian FLR sources of YNAO, KMAO, and Krasnoyarsky Krai (see Figure 1) causing up to 71% contribution 424 425 to surface BC (Supplementary Table S 5).

426 On 12 November 2019, airmasses arrived at the Bely Island through the Yamal peninsula 427 after passing from the ocean (Supplementary Figure S 1). Model strongly underestimated 428 measured EBC concentrations by about 10 times (Figure 7). We fail to provide a concrete 429 explanation for this; a simplified hypothesis is that a number of flaring sites located at the Yamal 430 peninsula might have not been included into the emission database, but this certainly needs further research. In contrast, the model strongly overestimated measured EBC concentrations on 16 431 November 2019. At that time, airmasses passed through remote regions of Eastern Siberia and 432 arrived through the gas flaring sites of Krasnoyarsky Krai at the station (Figure 7) causing an FLR 433 434 contribution of 98.6% to surface BC (Supplementary Table S 5). The reason might be use of incorrect emission factors for BC at the FLR facilities of Krasnovarsky Krai in the adopted 435 emissions, because direct transport from this region was observed. During 12 and 16 November 436 2019 the AAE was in the range from 0.7 to 1, which agrees with the expected optical properties 437 438 for the FLR sources.

Pollution episode C2 in December 2019 gave the highest EBC concentrations observed during the whole cold period, (Figure 5a). On 4 December, EBC approched 400 ng/m³, when airmasses originated from Kazakhstan and Russian gas flaring regions of KMAO and YNAO (Figure 8). The maximum EBC concentration of approximately 500 ng/m³ with AAE 1.05 was observed on 19 December, when air came from Europe, initially through the Russian oil and gas basins of Volga-Ural at the south of European Russia and then through KMAO and YNAO in

Western Siberia (Figure 8). During the December pollution events, FLR contribution dominated, 445 reaching 73% on 19 December (Supplementary Table S 5). 446

The highest FLR contributions were observed during the pollution episodes C3-C6 447 (Supplementary Table S 5). Similar airmass transportation through either gas and oil fields of 448 YNAO and KMAO in Western Siberia or Komi and Nenets regions north of European Russia 449 occurred in all of the events, (Supplementary Figure S 1). 450

451 In contrast to the aforementioned events, the pollution episode C7 was unrelated to FLR, as airmasses did not cross flaring regions (Figure 8). On 16 November 2020, retroplumes confirm 452 origin of surface BC from Central and Eastern Europe and the Kola Peninsula (Figure 8). DOM 453 454 and TRA hold the largest share of the source contribution with 73% and 20%, respectively 455 (Supplementary Table S 5), while the model overestimated measured EBC. Episode C8 gave the 456 biggest EBC (370) concentration which reached 346 ng/m³ and exceeded EBC (880) (133 ng/m³) 457 on 24 November 2020 (Supplementary Table S 5). At this time, airmasses came to the Bely Island 458 directly from the most populated region of European Russia (Figure 8). The contribution of DOM and TRA was 34% and 23%, respectively. AAE approached the highest value observed (1.3) 459 460 during the study period. This might show a detectable impact of biomass burning in the classified DOM emissions. BC from wood burning contributes around 61% of the total residential emissions, 461 462 especially in areas where there is limited use of natural gas (Kalogridis et al., 2018a), and in forest 463 regions (Huang et al., 2015). Note that the impact of IND emissions was the largest in episodes C7 and C8 as compared to the whole cold period (Supplementary Table S 5), due to industrial 464 emissions from sites in Central European Russia. 465

466

3.3 Warm season pollution

467 Figure 5b shows EBC concentrations measured at the "Island Bely" station during the warm period, from 10 August to 31 October 2019 and from 1 April to 1 November 2020. It is 468 469 immediately seen that BC in the warm period was mainly affected by Russian emissions (90%), and only in October 2020 and August 2019 partly (~20%) from Europe and North America, 470 (Figure 4). EBC concentrations rarely exceeded the 80th percentile that was set as the pollution 471 criterion, while the duration of the warm period episodes was shorter. 472

473 Due to the mixing with background aerosol and aging processes, airmasses influenced by BB events should be expected to have increased AAE as compared to the BC produced by fossil 474 475 fuel. However, aging processes may induce a high variability in AAE in areas affected by long-476 range transport, and hence AAE may not be representative of a BB source. Pollution events were 477 rarely observed in this season, and the most sufficient BB impact occurred during the pollution478 episodes W4, W7, and W8.

However, events characterised by higher AAE were observed more often indicating that 479 BB impact was more significant during the warm period, mainly during spring and summer 480 (episodes W3, W4, and W6). Comparison between measured and modelled concentrations showed 481 poor correlation (R of 0.41 and RMSE of 121 ng/m³) (Figure 6). According to monthly median 482 483 contributions to surface BC concentrations in the warm period, the impact of BB emissions was 484 as high as 50% (Figure 4c). SHP emissions contributed about 1%, as a result of the increase of touristic activity in the Arctic and the more active use of the Northern Sea Route due to the Arctic 485 ice retreat. 486

From the beginning of the study period in August 2019, large wildfires were observed in Siberia (Voronova et al., 2020). The latter resulted in a strong BB impact at the "Island Bely" station (**Figure 5**b). However, during episode W1, EBC concentrations reached approximately 200 ng/m³ not caused by wildfire plumes (**Figure 5**). During this time, airmasses were transported from Northern Europe (Supplementary Figure S 1), and the main contribution to surface BC at the "Island Bely" station was due to TRA emissions (36%, Supplementary Table S 5).

Episode W2 during October 2019 (Figure 5) was characterised by EBC of 119 ng/m³, 493 while modelled BC was strongly overestimated (Supplementary Figure S 1). The calculated BB 494 contribution to station's surface BC was 64% (Supplementary Table S 5) and the hotspot BB 495 sources were near the Pur River (YNAO), as recorded by CAMS (Supplementary Figure S 1). The 496 measured AAE does not indicate any contribution from BrC, as it would be expected for BB 497 498 sources and observed AAE values were lower than 1 (Supplementary Table S 5). Note that the 499 FIRMS active fire data analyses (https://firms.modaps.eosdis.nasa.gov/) indicate that the fire spots 500 were in the same grid-cell as industrial facilities of oil extraction field in the Purovski region (YNAO). Thus, it might be that thermal anomalies from flaring facilities were mistakenly related 501 to fires in CAMS. This hypothesis is reinforced by the fact that no wildfires were recorded by the 502 local forest fire service (https://aviales.ru) during October 2020 in Western Siberia and 503 504 Krasnoyarsky Krai.

Pollution episode W3 is related to a strong springtime wildfire activity that occurred in
Southern Siberia. The retroplumes on 18 and 23 April 2020 showed that the air originated from
Central Asia, a large territory of Southern Siberia and Krasnoyarsk Kray, and arrived at the Bely
Island through the Western Siberia from the southeast (Supplementary Figure S 1, Figure 9). High
footprint emission sensitivities coincided with the location of large wildfires resulting in BB

contribution to surface BC at the station equal to 28% (18 April 2020) and 19% (23 April 2020),
respectively. The most significant impact of wildfires was observed on 23 April 2020, when 6-h
median EBC concentration reached 700 ng/m3 with AAE ranging from 1.3 to 1.5, clearly
indicating an elevated contribution of BrC (Supplementary Table S 5).

Wildfires occurred in northern Krasnoyarsk Krai and Sakha Republic, Central Siberia between April and November 2020 (<u>https://aviales.ru/popup.aspx?news=6286</u>) that burned around seven million hectares of forest. The pollution episode W4 on 7 July 2020 recorded a 6-h median EBC of 150 ng/m³ and an AAE of around 1.4 clearly indicating BB impact. The model captures this event well, providing the highest BB contribution exactly when observed, equal to 90% (Supplementary Table S 5). Airmasses arrived from the east and passed north of Krasnoyarsk Kray where the large wildfires occurred (**Figure 9**).

521 Unprecedented high wildfire-related BC concentration was observed in September 2020 (pollution episode W6). EBC concentrations exceeded 5 and 20 times the 80th percentile of the 522 measurements. Maximum 6-h median EBC reached 534 ng/m³ on 1st September 2020 and it was 523 even higher than the biggest Arctic haze concentration observed in December 2019, 524 525 (Supplementary Table S 5). Increased AAE of around 1.4 revealed strong BB impact. This event resulted from long-range transport of BC from the Eurasian continent during the intensive 526 527 wildfires in Western Siberia (Krasnoyarsk Kray and Yakutia) (Figure 9), where around one 528 million hectares of forest were burned in August 2020. The contribution of BB to surface BC at the "Island Bely" station was as high as 95%. 529

Despite the exclusive BB origin of the light absorbing carbon measured at the "Island Bely" 530 station, the AAE was much lower than the established value for fresh BB (close to 2) (Sandradewi 531 et al., 2008) likely due to aging. This apparent reduction of the BrC contribution to absorption is 532 in agreement with (Forrister et al., 2015), who examined BrC concentrations and AAE from 533 Western U.S. forest fires as a function of aging. Their results show that most of the BrC (~94%) 534 emitted from wildfires was lost within a day. Similar observations have been reported for long-535 536 range transported North American smoke over the Northeastern Atlantic (Zheng et al., 2020) and 537 for transported Russian smoke over the Mediterranean (Diapouli et al., 2014).

The last pollution episode W7 was observed at the end of October 2020. Although it occurred in the warm period, it is rather related to Russian FLR and European TRA emissions (Supplementary Table S 5, **Figure 5**b). At the end of October 2020 airmasses came mainly from Europe, passing through the Yamal Peninsula.

542 **4** Conclusions

- The present paper aims at performing a quantitative analysis of the Arctic pollution via high-resolution measurements from a recently developed aerosol station at the Bely Island (Kara Sea) combined with Lagrangian modelling. A consequent goal is to examine the impact of anthropogenic and natural sources to the high Arctic as a result of long-range transport. The main results can be summarised as follows:
- EBC monthly climatology is following the typical Arctic aerosol seasonal variation
 characterised with higher EBC concentrations in winter and lower in summer.
- AAE for aged BC larger than 1 indicates wildfire impact in the warm period, but mixing with
 gas flaring emissions from nearby regions was also observed.
- The recently upgraded ECLIPSEv6 emissions and ECLIPSEv6 coupled with FLR from
 BCRUS represent measured EBC accurately in the cold period. Annual average contributions
 of anthropogenic emissions to surface BC were 76% and 80% (50% and 59% from gas flaring)
 for each dataset, respectively.
- The most significant model overestimation was observed in February 2020, when air masses
 passed through non gas flaring regions. The largest underestimation occurred in April 2020
 during the period of spring agriculture fires.
- Daily BB emissions from CAMS were more efficient in representing pollution episodes than
 monthly GFED4 emissions, and therefore they were mainly used here.
- Russian emissions dominate during the whole year, European and Asian ones contribute up
 to 20% in the cold period. Pollution episodes with EBC concentrations above 90 ng/m³ occur
 in 18.5% of observation time. Monthly average FLR emissions dominate (98%) any other
 emission sector.
- FLR and BB emissions contribute the largest share of EBC to the "Island Bely" station during
 the cold and warm period, respectively. This is consistent with previously-reported source
 contributions to the the Russian Arctic. When air is transported from Europe, other sources
 such as TRA become important. The same applies for SHP emissions that become important
 in summertime, because of cruise activities and ice retreat.
- Emissions from gas and oil fields in Western Siberia, and the Northern European Russia cause
 the vast majority of the pollution episodes in the Arctic.
- 572 15 pollution episodes with concentrations reaching close to 723 ng/m3 were detected. The
 573 duration of the cold pollution episodes is longer than of the warm period, and the median (up
 574 to 160 ng/m3) and maximum EBC (up to 450 ng/m3) concentrations higher.
- 575 In conclusion, the significance of high-quality measurements at the "Island Bely" station 576 is pronounced from the present study, because (i) the station is located along the main pathway of

airmasses entering the Arctic, and (ii) it is north of the world's largest gas flaring regions. The
operation of the "Island Bely" station is an asset in source emission optimisation, because EBC
measurements in the High Arctic are still rare.

580

581Data availability. All model data used in the present publication together with all figures of582footprint analysis and source contributions to surface BC are open through the websites583https://niflheim.nilu.no/NikolaosPY/Bely_2020_cams.py584https://niflheim.nilu.no/NikolaosPY/Bely_2020_huang_cams.py585obtained from the corresponding author upon request. The definitions of the regions and continents586used in the current analysis are based on regional masks that can be seen in Supplementary Figure587S 3.

588

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590

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- of the results. K.E. provided BB impact and AAE aging evaluation. A.G. performed AAE
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- 618

619 **References**

- 620
- Akagi, S., Yokelson, R. J., Wiedinmyer, C., Alvarado, M., Reid, J., Karl, T., Crounse, J., and
 Wennberg, P.: Emission factors for open and domestic biomass burning for use in atmospheric
 models, Atmospheric Chemistry and Physics, 11, 4039-4072, 2011.
- 624 Andreae, M. O. and Merlet, P.: Emission of trace gases and aerosols from biomass burning,
- Global biogeochemical cycles, 15, 955-966, 2001.
- Bond, T. C., Doherty, S. J., Fahey, D., Forster, P., Berntsen, T., DeAngelo, B., Flanner, M.,
- 627 Ghan, S., Kärcher, B., and Koch, D.: Bounding the role of black carbon in the climate system: A
- 628 scientific assessment, Journal of Geophysical Research: Atmospheres, 118, 5380-5552, 2013.
- Böttcher, K., Paunu, V.-V., Kupiainen, K., Zhizhin, M., Matveev, A., Savolahti, M., Klimont, Z.,
- 630 Väätäinen, S., Lamberg, H., and Karvosenoja, N.: Black carbon emissions from flaring in Russia
- 631 in the period 2012-2017, Atmospheric Environment, 118390, 2021.
- 632 Cappa, C. D., Kolesar, K. R., Zhang, X., Atkinson, D. B., Pekour, M. S., Zaveri, R. A.,
- 633
 Zelenyuk, A., and Zhang, Q.: Understanding the optical properties of ambient sub- and

 633
 Zelenyuk, A., and Zhang, Q.: Understanding the optical properties of ambient sub- and
- 634 supermicron particulate matter: results from the CARES 2010 field study in northern California,
- 635 Atmos. Chem. Phys., 16, 6511-6535, 10.5194/acp-16-6511-2016, 2016.
- 636 Cassiani, M., Stohl, A., and Brioude, J.: Lagrangian stochastic modelling of dispersion in the
- 637 convective boundary layer with skewed turbulence conditions and a vertical density gradient:
- Formulation and implementation in the FLEXPART model, Boundary-Layer Meteorology, 154,
 367-390, 2015.
- 640 Chang, R.-W., Leck, C., Graus, M., Müller, M., Paatero, J., Burkhart, J. F., Stohl, A., Orr, L.,
- Hayden, K., and Li, S.-M.: Aerosol composition and sources in the central Arctic Ocean during
 ASCOS, Atmospheric Chemistry and Physics, 2011.
- 643 Cho, M.-H., Park, R. J., Yoon, J., Choi, Y., Jeong, J. I., Labzovskii, L., Fu, J. S., Huang, K.,
- Jeong, S.-J., and Kim, B.-M.: A missing component of Arctic warming: black carbon from gas
 flaring, Environmental Research Letters, 14, 094011, 2019.
- 646 Conrad, B. M. and Johnson, M. R.: Field measurements of black carbon yields from gas flaring,
- Environmental science & technology, 51, 1893-1900, 2017.
- 648 Di Giuseppe, F., Remy, S., Pappenberger, F., and Wetterhall, F.: Improving CAMS biomass
- burning estimations by means of the Global ECMWF Fire Forecast system (GEFF), ECMWF
 Tech. Memo. 790, 18 pp., <u>https://www</u>. ecmwf. int/sites/default/files ..., 2016.
- Drinovec, L., Močnik, G., Zotter, P., Prévôt, A., Ruckstuhl, C., Coz, E., Rupakheti, M., Sciare,
- 52 J., Müller, T., and Wiedensohler, A.: The" dual-spot" Aethalometer: an improved measurement
- of aerosol black carbon with real-time loading compensation, Atmospheric Measurement
- 654 Techniques, 8, 1965-1979, 2015.
- 655 Eckhardt, S., Quennehen, B., Olivié, D. J. L., Berntsen, T. K., Cherian, R., Christensen, J. H.,
- 656 Collins, W., Crepinsek, S., Daskalakis, N., Flanner, M., Herber, A., Heyes, C., Hodnebrog, Ø.,
- Huang, L., Kanakidou, M., Klimont, Z., Langner, J., Law, K. S., Lund, M. T., Mahmood, R.,
- Massling, A., Myriokefalitakis, S., Nielsen, I. E., Nøjgaard, J. K., Quaas, J., Quinn, P. K., Raut,
- J. C., Rumbold, S. T., Schulz, M., Sharma, S., Skeie, R. B., Skov, H., Uttal, T., von Salzen, K.,
- and Stohl, A.: Current model capabilities for simulating black carbon and sulfate concentrations

- 661 in the Arctic atmosphere: a multi-model evaluation using a comprehensive measurement data set,
- 662 Atmos. Chem. Phys., 15, 9413-9433, 10.5194/acp-15-9413-2015, 2015.
- 663 Eleftheriadis, K., Vratolis, S., and Nyeki, S.: Aerosol black carbon in the European Arctic:

Measurements at Zeppelin station, Ny-Ålesund, Svalbard from 1998–2007, Geophysical
 Research Letters, 36, n/a-n/a, 10.1029/2008GL035741, 2009.

- 666 Eleftheriadis, K., Nyeki, S., Psomiadou, C., and Colbeck, I.: Background aerosol properties in
- the European arctic, Water, Air and Soil Pollution: Focus, 4, 23-30, 2004.
- 668 Flanner, M. G.: Arctic climate sensitivity to local black carbon, Journal of Geophysical
- 669 Research: Atmospheres, 118, 1840-1851, 10.1002/jgrd.50176, 2013.
- 670 Forrister, H., Liu, J., Scheuer, E., Dibb, J., Ziemba, L., Thornhill, K. L., Anderson, B., Diskin,
- 671 G., Perring, A. E., and Schwarz, J. P.: Evolution of brown carbon in wildfire plumes,
- 672 Geophysical Research Letters, 42, 4623-4630, 2015.
- Forster, C., Stohl, A., and Seibert, P.: Parameterization of convective transport in a Lagrangian

particle dispersion model and its evaluation, Journal of applied meteorology and climatology, 46,403-422, 2007.

- 676 Giglio, L., Randerson, J. T., and Van Der Werf, G. R.: Analysis of daily, monthly, and annual
- burned area using the fourth-generation global fire emissions database (GFED4), Journal of
 Geophysical Research: Biogeosciences, 118, 317-328, 2013.
- 679 Grange, S. K., Lötscher, H., Fischer, A., Emmenegger, L., and Hueglin, C.: Evaluation of
- 680 equivalent black carbon source apportionment using observations from Switzerland between 2008 and 2018. Atmospheric Massurement Tashnigues, 12, 1867, 1885, 2020
- 681 2008 and 2018, Atmospheric Measurement Techniques, 13, 1867-1885, 2020.
- Grythe, H., Kristiansen, N. I., Groot Zwaaftink, C. D., Eckhardt, S., Ström, J., Tunved, P.,
- 683 Krejci, R., and Stohl, A.: A new aerosol wet removal scheme for the Lagrangian particle model
- 684 FLEXPART v10, Geosci. Model Dev., 10, 1447-1466, 10.5194/gmd-10-1447-2017, 2017.
- Helin, A., Virkkula, A., Backman, J., Pirjola, L., Sippula, O., Aakko-Saksa, P., Väätäinen, S.,
- 686 Mylläri, F., Järvinen, A., and Bloss, M.: Variation of absorption Ångström exponent in aerosols 687 from different emission sources, Journal of Geophysical Research: Atmospheres, 126,
- from different emission sources, Journal of Geophysical Research: Atmospheres, 126
 e2020JD034094, 2021.
- 689 Huang, K., Fu, J. S., Prikhodko, V. Y., Storey, J. M., Romanov, A., Hodson, E. L., Cresko, J.,
- 690 Morozova, I., Ignatieva, Y., and Cabaniss, J.: Russian anthropogenic black carbon: Emission
- 691 reconstruction and Arctic black carbon simulation, Journal of Geophysical Research:
- 692 Atmospheres, 120, 11,306-311,333, 2015.
- Ismail, O. S. and Umukoro, G. E.: Global impact of gas flaring, Energy and Power Engineering,4, 290, 2012.
- Johnson, M. S., Strawbridge, K., Knowland, K. E., Keller, C., and Travis, M.: Long-range
- transport of Siberian biomass burning emissions to North America during FIREX-AQ,
- 697 Atmospheric Environment, 252, 118241, 2021.
- Kaiser, J., Heil, A., Andreae, M., Benedetti, A., Chubarova, N., Jones, L., Morcrette, J.-J.,
- 699 Razinger, M., Schultz, M., and Suttie, M.: Biomass burning emissions estimated with a global
- fire assimilation system based on observed fire radiative power, Biogeosciences, 9, 527-554,
- 701 2012.
- 702 Kalogridis, A.-C., Vratolis, S., Liakakou, E., Gerasopoulos, E., Mihalopoulos, N., and
- 703 Eleftheriadis, K.: Assessment of wood burning versus fossil fuel contribution to wintertime black
- carbon and carbon monoxide concentrations in Athens, Greece, Atmospheric Chemistry and
- 705 Physics, 18, 10219-10236, 2018a.
- 706 Kalogridis, A. C., Popovicheva, O. B., Engling, G., Diapouli, E., Kawamura, K., Tachibana, E.,
- 707 Ono, K., Kozlov, V. S., and Eleftheriadis, K.: Smoke aerosol chemistry and aging of Siberian
- biomass burning emissions in a large aerosol chamber, Atmospheric Environment, 185, 15-28,
 <u>https://doi.org/10.1016/j.atmosenv.2018.04.033</u>, 2018b.
- 710 Kaufman, Y., Ichoku, C., Giglio, L., Korontzi, S., Chu, D., Hao, W., Li, R.-R., and Justice, C.:
- 711 Fire and smoke observed from the Earth Observing System MODIS instrument--products,
- validation, and operational use, International Journal of Remote Sensing, 24, 1765-1781, 2003.

- 713 Klimont, Z., Kupiainen, K., Heyes, C., Purohit, P., Cofala, J., Rafaj, P., Borken-Kleefeld, J., and
- 714 Schöpp, W.: Global anthropogenic emissions of particulate matter including black carbon,
- Atmospheric Chemistry and Physics Discussions, 17, 8681-8723, 2017.
- 716 Kozlov, V. S., Panchenko, M. V., Shmargunov, V. P., Chernov, D. G., Yausheva, E. P., Pol'kin,
- 717 V. V., and Terpugova, S. A.: Long-term investigations of the spatiotemporal variability of black
- carbon and aerosol concentrations in the troposphere of West Siberia and Russian Subarctic,
- 719 Химия в интересах устойчивого развития, 24, 423-440, 2016.
- 720 Long, C. M., Nascarella, M. A., and Valberg, P. A.: Carbon black vs. black carbon and other
- airborne materials containing elemental carbon: Physical and chemical distinctions,
- 722 Environmental Pollution, 181, 271-286, 2013.
- 723 Manousakas, M., Popovicheva, O., Evangeliou, N., Diapouli, E., Sitnikov, N., Shonija, N., and
- 724 Eleftheriadis, K.: Aerosol carbonaceous, elemental and ionic composition variability and origin
- at the Siberian High Arctic, Cape Baranova, Tellus B: Chemical and Physical Meteorology, 72,
 1-14, 2020.
- 727 Paris, J.-D., Stohl, A., Nédélec, P., Arshinov, M. Y., Panchenko, M., Shmargunov, V., Law, K.
- S., Belan, B., and Ciais, P.: Wildfire smoke in the Siberian Arctic in summer: source
- characterization and plume evolution from airborne measurements, Atmospheric Chemistry andPhysics, 9, 9315-9327, 2009.
- 731 Pisso, I., Sollum, E., Grythe, H., Kristiansen, N. I., Cassiani, M., Eckhardt, S., Arnold, D.,
- 732 Morton, D., Thompson, R. L., and Groot Zwaaftink, C. D.: The Lagrangian particle dispersion
- model FLEXPART version 10.4, Geoscientific Model Development, 12, 4955-4997, 2019.
- 734 Popovicheva, O., Chichaeva, M., Kobelev, V., Sinitskiy, A., and Hansen, A.: Black Carbon in
- urban emissions on the Polar Circle, 26th International Symposium on Atmos. Ocean Optics,
 Proc. SPIE, 11560J,
- 737 Popovicheva, O., Timofeev, M., Persiantseva, N., Jefferson, M. A., Johnson, M., Rogak, S. N.,
- and Baldelli, A.: Microstructure and chemical composition of particles from small-scale gas
- flaring, Aerosol and Air Quality Research, 19, 2205-2221, 2019a.
- 740 Popovicheva, O., Diapouli, E., Makshtas, A., Shonija, N., Manousakas, M., Saraga, D., Uttal, T.,
- and Eleftheriadis, K.: East Siberian Arctic background and black carbon polluted aerosols at
 HMO Tiksi, Science of the Total Environment, 655, 924-938, 2019b.
- Popovicheva, O. B., Evangeliou, N., Eleftheriadis, K., Kalogridis, A. C., Sitnikov, N., Eckhardt,
- 744 S., and Stohl, A.: Black Carbon Sources Constrained by Observations in the Russian High
- 745 Arctic, Environmental Science & Technology, 51, 3871-3879, 10.1021/acs.est.6b05832, 2017.
- 746 Qi, L. and Wang, S.: Sources of black carbon in the atmosphere and in snow in the Arctic,
- 747 Science of The Total Environment, 691, 442-454, 2019.
- 748 Quinn, P. K., Bates, T. S., Baum, E., Doubleday, N., Fiore, A. M., Flanner, M., Fridlind, A.,
- Garrett, T. J., Koch, D., Menon, S., Shindell, D., Stohl, A., and Warren, S. G.: Short-lived
- 750 pollutants in the Arctic: their climate impact and possible mitigation strategies, Atmos. Chem.
- 751 Phys., 8, 1723-1735, 10.5194/acp-8-1723-2008, 2008.
- 752 Randerson, J., Chen, Y., Van Der Werf, G., Rogers, B., and Morton, D.: Global burned area and
- biomass burning emissions from small fires, Journal of Geophysical Research: Biogeosciences,117, 2012.
- Ren, L., Yang, Y., Wang, H., Zhang, R., Wang, P., and Liao, H.: Source attribution of Arctic
- black carbon and sulfate aerosols and associated Arctic surface warming during 1980–2018,
 Atmospheric Chemistry and Physics, 20, 9067-9085, 2020.
- 758 Romshoo, B., Müller, T., Pfeifer, S., Saturno, J., Nowak, A., Ciupek, K., Quincey, P., and
- 759 Wiedensohler, A.: Optical properties of coated black carbon aggregates: numerical simulations,
- radiative forcing estimates, and size-resolved parameterization scheme, Atmos. Chem. Phys., 21,
- 761 12989-13010, 10.5194/acp-21-12989-2021, 2021.
- Saleh, R., Hennigan, C., McMeeking, G., Chuang, W., Robinson, E., Coe, H., Donahue, N., and
- 763 Robinson, A.: Absorptivity of brown carbon in fresh and photo-chemically aged biomass-
- burning emissions, Atmospheric Chemistry and Physics, 13, 7683-7693, 2013.

- 765 Sandradewi, J., Prévôt, A. S., Szidat, S., Perron, N., Alfarra, M. R., Lanz, V. A., Weingartner, E.,
- and Baltensperger, U.: Using aerosol light absorption measurements for the quantitativedetermination of wood burning and traffic emission contributions to particulate matter,
- 768 Environmental science & technology, 42, 3316-3323, 2008.
- 769 Schacht, J., Heinold, B., Quaas, J., Backman, J., Cherian, R., Ehrlich, A., Herber, A., Huang, W.
- T. K., Kondo, Y., Massling, A., Sinha, P. R., Weinzierl, B., Zanatta, M., and Tegen, I.: The
- importance of the representation of air pollution emissions for the modeled distribution and
- radiative effects of black carbon in the Arctic, Atmos. Chem. Phys., 19, 11159-11183,
- 773 10.5194/acp-19-11159-2019, 2019.
- Sharma, S., Lavoué, D., Cachier, H., Barrie, L., and Gong, S.: Long-term trends of the black
- carbon concentrations in the Canadian Arctic, Journal of Geophysical Research: Atmospheres,109, 2004.
- 577 Stathopoulos, V., Evangeliou, N., Stohl, A., Vratolis, S., Matsoukas, C., and Eleftheriadis, K.:
- Large circulation patterns strongly modulate long term variability of Arctic black carbon levels
 and areas of origin, Geophysical Research Letters, e2021GL092876, 2021.
- 780 Stohl, A.: Characteristics of atmospheric transport into the Arctic troposphere, Journal of
- 781 Geophysical Research: Atmospheres, 111, n/a-n/a, 10.1029/2005JD006888, 2006.
- 782 Stohl, A., Forster, C., Frank, A., Seibert, P., and Wotawa, G.: Technical note: The Lagrangian
- particle dispersion model FLEXPART version 6.2, Atmos. Chem. Phys., 5, 2461-2474,
- 784 10.5194/acp-5-2461-2005, 2005.
- 785 Stohl, A., Klimont, Z., Eckhardt, S., Kupiainen, K., Shevchenko, V. P., Kopeikin, V. M., and
- 786 Novigatsky, A. N.: Black carbon in the Arctic: the underestimated role of gas flaring and
- residential combustion emissions, Atmos. Chem. Phys., 13, 8833-8855, 10.5194/acp-13-88332013, 2013.
- 789 Stohl, A., Andrews, E., Burkhart, J., Forster, C., Herber, A., Hoch, S., Kowal, D., Lunder, C.,
- 790 Mefford, T., and Ogren, J.: Pan-Arctic enhancements of light absorbing aerosol concentrations
- due to North American boreal forest fires during summer 2004, Journal of Geophysical
- 792 Research: Atmospheres, 111, 2006.
- 793 Stohl, A., Berg, T., Burkhart, J., Fjáraa, A., Forster, C., Herber, A., Hov, Ø., Lunder, C.,
- 794 McMillan, W., and Oltmans, S.: Arctic smoke–record high air pollution levels in the European
- Arctic due to agricultural fires in Eastern Europe in spring 2006, Atmospheric Chemistry and
- 796 Physics, 7, 511-534, 2007.
- 797 Stone, R. S., Sharma, S., Herber, A., Eleftheriadis, K., and Nelson, D. W.: A characterization of
- Arctic aerosols on the basis of aerosol optical depth and black carbon measurements, Elem SciAnth, 2, 2014.
- 800 Treffeisen, R., Turnved, P., StrÖm, J., Herber, A., Bareiss, J., Helbig, A., Stone, R. S.,
- 801 Hoyningen-Huene, W., Krejci, R., Stohl, A., and Neuber, R.: Arctic smoke ? aerosol
- characteristics during a record air pollution event in the European Arctic and its radiative impact,
 Atmospheric Chemistry and Physics Discussions, 7, 2275-2324, 2007.
- 804 Tunved, P., Ström, J., and Krejci, R.: Arctic aerosol life cycle: linking aerosol size distributions
- observed between 2000 and 2010 with air mass transport and precipitation at Zeppelin station,
 Ny-Å lesund, Svalbard, Atmospheric Chemistry & Physics, 13, 3643-3660, 2013.
- Ny-A result, Svaluard, Atmospheric Chemistry & Physics, 15, 3043-3000, 2013.
 Vinogradova, A.: Anthropogenic Black Carbon emissions to the atmosphere: surface distribution
- winogradova, A.: Anthropogenic Black Carbon emissions to the atmosphere: surface distribut
 through Russian territory, Atmospheric and Oceanic Optics, 28, 158-164, 2015.
- 809 Virkkula, A.: Modeled source apportionment of black carbon particles coated with a light-
- scattering shell, Atmospheric Measurement Techniques, 14, 3707-3719, 2021.
- 811 Voronova, O., Zima, A., Kladov, V., and Cherepanova, E.: Anomalous Wildfires in Siberia in Summer 2010, Investigation Atmospheric and Oceanic Physics 56 (1042) 1052 (2020)
- 812 Summer 2019, Izvestiya, Atmospheric and Oceanic Physics, 56, 1042-1052, 2020.
- 813 Wang, Q., Jacob, D. J., Fisher, J. A., Mao, J., Leibensperger, E. M., Carouge, C. C., Le Sager, P.,
- 814 Kondo, Y., Jimenez, J. L., Cubison, M. J., and Doherty, S. J.: Sources of carbonaceous aerosols
- 815 and deposited black carbon in the Arctic in winter-spring: implications for radiative forcing,
- 816 Atmos. Chem. Phys., 11, 12453-12473, 10.5194/acp-11-12453-2011, 2011.

- 817 Warneke, C., Froyd, K., Brioude, J., Bahreini, R., Brock, C., Cozic, J., De Gouw, J., Fahey, D.,
- 818 Ferrare, R., and Holloway, J.: An important contribution to springtime Arctic aerosol from
- 819 biomass burning in Russia, Geophysical Research Letters, 37, 2010.
- 820 Werf, G. R., Randerson, J. T., Giglio, L., Leeuwen, T. T. v., Chen, Y., Rogers, B. M., Mu, M.,
- Van Marle, M. J., Morton, D. C., and Collatz, G. J.: Global fire emissions estimates during
 1997–2016, Earth System Science Data, 9, 697-720, 2017.
- Winiger, P., Andersson, A., Eckhardt, S., Stohl, A., and Gustafsson, Ö.: The sources of
- atmospheric black carbon at a European gateway to the Arctic, Nature communications, 7, 1-8,
 2016.
- 826 Winiger, P., Andersson, A., Eckhardt, S., Stohl, A., Semiletov, I. P., Dudarev, O. V., Charkin,
- A., Shakhova, N., Klimont, Z., Heyes, C., and Gustafsson, Ö.: Siberian Arctic black carbon
 sources constrained by model and observation, Proceedings of the National Academy of
- 829 Sciences, 114, E1054-E1061, 10.1073/pnas.1613401114, 2017.
- 830 Wooster, M. J., Roberts, G., Perry, G., and Kaufman, Y.: Retrieval of biomass combustion rates
- and totals from fire radiative power observations: FRP derivation and calibration relationships
- between biomass consumption and fire radiative energy release, Journal of GeophysicalResearch: Atmospheres, 110, 2005.
- 834 Yun, Y., Penner, J. E., and Popovicheva, O.: The effects of hygroscopicity on ice nucleation of
- fossil fuel combustion aerosols in mixed-phase clouds, Atmos. Chem. Phys., 13, 4339-4348,
- 836 10.5194/acp-13-4339-2013, 2013.
- 837 Zhu, C., Kanaya, Y., Takigawa, M., Ikeda, K., Tanimoto, H., Taketani, F., Miyakawa, T.,
- 838 Kobayashi, H., and Pisso, I.: FLEXPART v10. 1 simulation of source contributions to Arctic
- black carbon, Atmospheric Chemistry and Physics, 20, 1641-1656, 2020.
- 840 Zotter, P., Herich, H., Gysel, M., El-Haddad, I., Zhang, Y., Močnik, G., Hüglin, C.,
- 841 Baltensperger, U., Szidat, S., and Prévôt, A. S.: Evaluation of the absorption Ångström
- 842 exponents for traffic and wood burning in the Aethalometer-based source apportionment using
- radiocarbon measurements of ambient aerosol, Atmospheric Chemistry and Physics, 17, 4229-
- **4249**, 2017.

847 FIGURE LEGENDS

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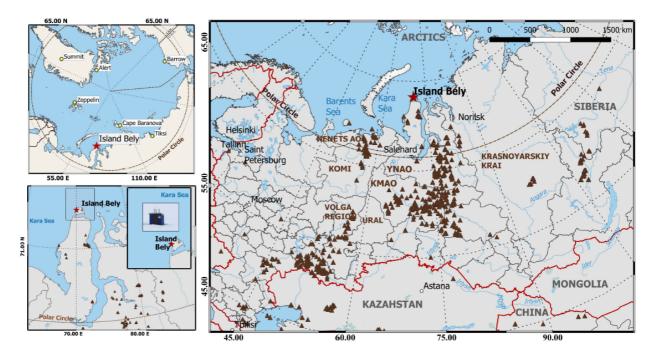




Figure 1. Top left map shows the location of the newly established "Island Bely" aerosol station
in contrast to other Arctic stations (Zeppelin, Alert, Barrow, Summit, Tiksi, and Cape Baranova).
Bottom left map shows a zoomed version of the location of the Bely Island in the Kara Sea, where
the new station was developed (73°20'7.57"N, 70°4'49.05"E). The map on the right shows the
"Island Bely" aerosol station in combination with the European part of Russia and Western Siberia
and the Yamalo-Nenets Autonomic Okrug (YNAO). Flares of oil and gas fields are shown for
2019 year in brown triangles (adopted from https://skytruth.org/).

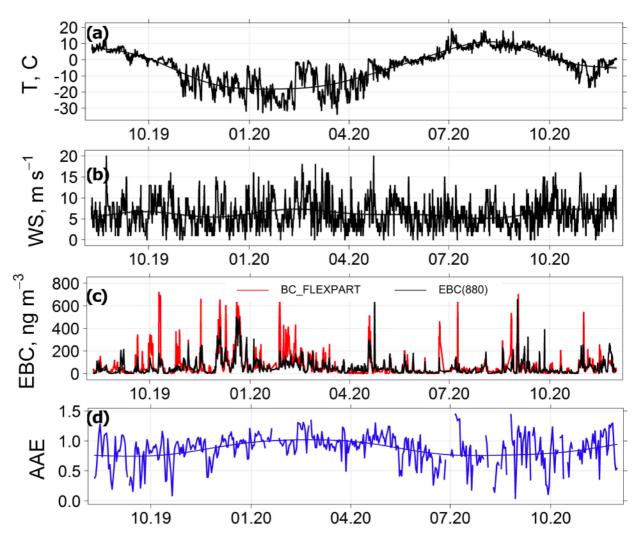
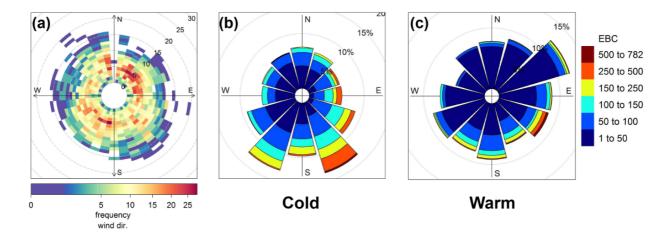


Figure 2. Meteorological conditions with respect to (a) mean temperature and (b) wind speed (data
were smoothed to show long-term variations), (c) time-series of 24-h median EBC (black) and
model BC using ECLIPSEv6 - CAMS emissions (red), and (d) 24-h average Absorption Ångström
Exponent (AAE) measured at "Island Bely" station from 10 August 2019 to 30 November 2020
(date format in mm.yy).



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Figure 3. (a) Polar frequency plots of wind speed and direction. Each cell gives the total number of hours the wind was originating from a certain wind direction. The dashed circular grey lines show the wind speed (in m s⁻¹). Rose diagrams show 3h EBC concentrations during the cold (b) and warm (c) periods.

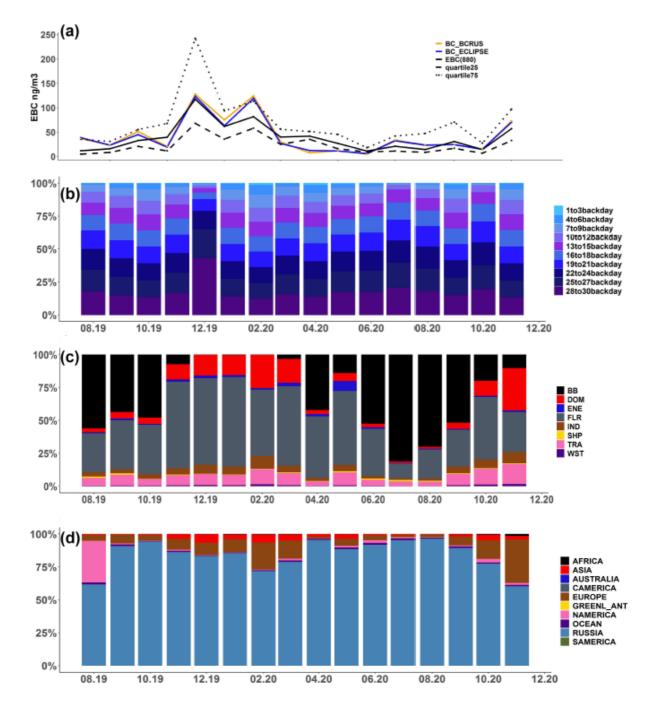


Figure 4. (a) Monthly climatology of EBC at the "Island Bely" station depicting medians, 25th and 872 75th percentiles (dashed lines). Near-surface monthly median BC concentrations simulated with 873 874 FLEXPART coupled to ECLIPSEv6-CAMS (steel blue) and ECLIPSEv6-BCRUS-CAMS (red) emissions are also shown. (b) Age spectrum of modelled BC from all possible sources showing 875 the contribution of emissions from each day back in time to the surface concentration of BC. (c) 876 Contribution from different emission source types to surface BC concentrations. The emission 877 sources of biomass burning (BB) adopted from GFEDv4.1, and residential and commercial 878 879 (DOM), power plants, energy conversion, and extraction (ENE), gas flaring (FLR), industrial combustion and processing (IND), shipping (SHP), and transportation (TRA) adopted from 880 ECLIPSEv6 were considered. (d) Continental spectrum showing the contribution from each 881 continent or region to surface BC concentrations . 10 regions were considered namely, Africa, 882 Asia, Australia, Central America, Europe, Greenland/Antarctica, North America, World Ocean, 883 Russia, and South America (see Supplementary Figure S 3). 884

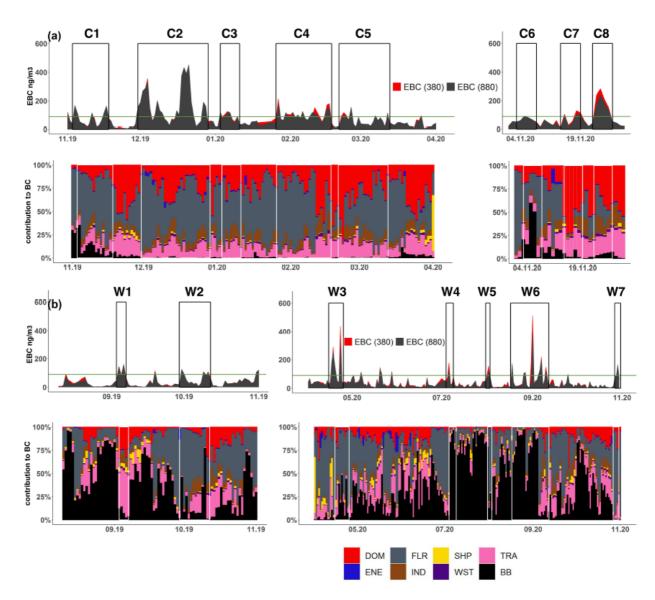


Figure 5. 24-h median EBC concentrations measured at 880 nm (black) and 370 nm (red), and
source contributions to surface BC from anthropogenic (DOM, ENE, FLR, IND, SHP, WST,
TRA) and BB sources for (a) the cold and (b) the warm period. Pollution episodes were composed
from the periodically repeated events of high EBC concentration. The green straight line indicates
the pollution level of the 80th percentile.

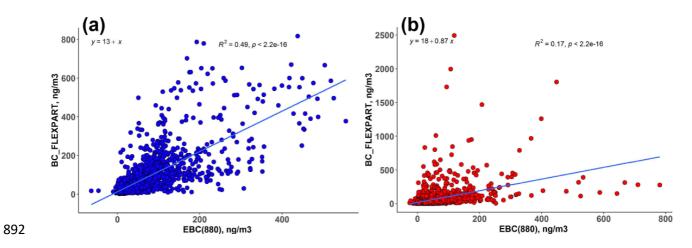


Figure 6. Scatter-plots of 3-h median measured EBC (880) against modelled BC from FLEXPART for the (a) cold and (b) warm period. Solid line is the linear regression fit of the comparison between modelled and observed values.



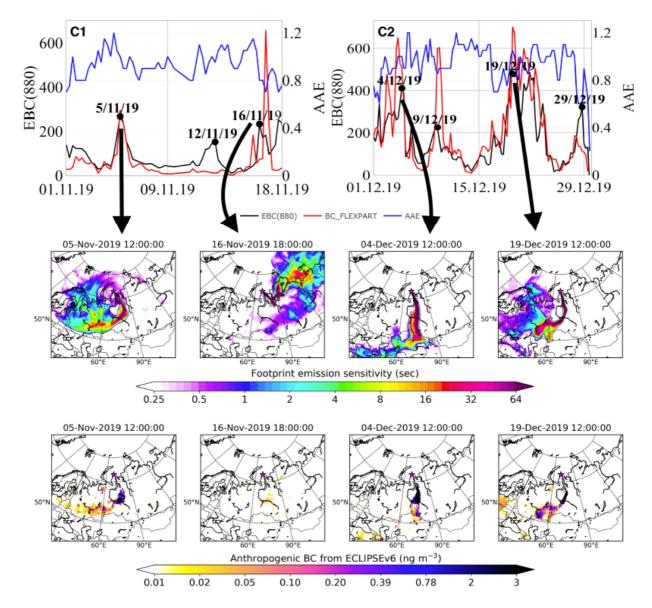


Figure 7. Examples of pollution episodes C1 and C2 observed in the cold period (see **Figure 5**a), where FLR contribution prevails. 6-h median EBC (880) (black line), BC simulated with FLEXPART (red line), AAE (blue line) (upper row). Footprint emissions sensitivities in seconds showing the largest probability of emission origin (middle row). Spatial distribution of anthropogenic contribution (in ng/m3) to surface BC at the "Island Bely" station (bottom row).

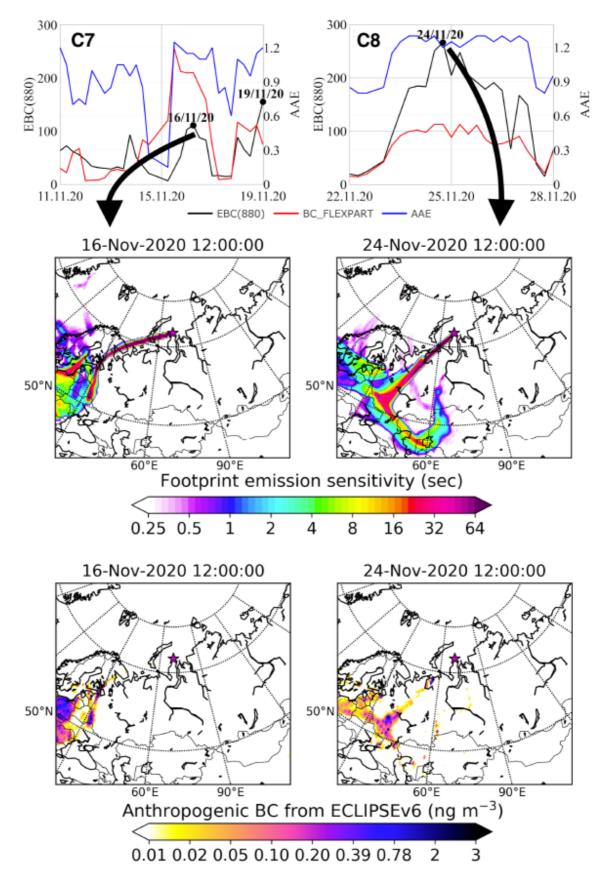


Figure 8. Examples of pollution episodes C7 and C8 observed in the cold period (see Figure 5a),
where DOM and TRA contribution prevails. Timeseries of measured EBC, modelled BC and AAE
(upper row), footprint emissions sensitivities (middle row) and anthropogenic contribution to
surface BC (bottom row) are shown.



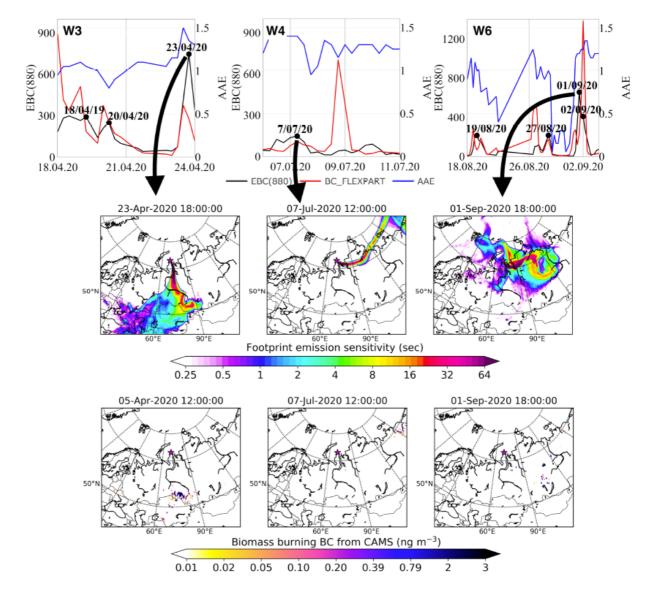


Figure 9. Examples of pollution episodes W3, W4 and W6 in the warm period (see Figure 5b),
where BB contribution prevails. The figure has been arranged similar to Figure 7 (timeseries of
measured EBC, modelled BC and AAE, footprint emissions sensitivities and BB contribution to
surface BC).