An assessment of macrophysical and microphysical cloud properties driving radiative forcing of shallow trade-wind clouds

Anna E. Luebke[1], André Ehrlich[1], Michael Schäfer[1], Kevin Wolf[1,2], and Manfred Wendisch[1]

[1]Leipzig Institute for Meteorology, University of Leipzig, Leipzig, Germany
[2]now: Institut Pierre-Simon Laplace, Sorbonne Université, Paris, France

Correspondence: Anna E. Luebke (anna.luebke@uni-leipzig.de)

Abstract. The clouds in the Atlantic trade-wind region are known to have an important ~~role in~~ impact on the global climate system. Acquiring a comprehensive characterization of these clouds based on observations is a challenge, but it is ~~a necessary piece of information~~ necessary for the evaluation of their representation in models. An exploration of how the macrophysical and microphysical cloud properties and organization of the cloud field impact the large-scale cloud radiative forcing is presented here. ~~Direct~~ In-situ measurements of the cloud radiative effects ~~from~~ based on the Broadband AirCrAft RaDiometer Instrumentation (BACARDI) on board the High Altitude and LOng Range Research Aircraft (HALO) and cloud observations from the GOES-16 satellite collected during the Elucidating the role of clouds-circulation coupling in climate (EUREC[4]A) campaign ~~provide evidence to~~ demonstrate what drives the cloud radiative effects in shallow trade-wind clouds. We find that the solar and terrestrial radiative effects of these clouds are largely driven by their macrophysical properties (cloud fraction and a scene-averaged liquid water path). ~~However, we~~ We also conclude that the microphysical properties, cloud top height and the organization of the cloud field ~~demonstrate an increasing relevance in determining~~ increasingly determine the cloud radiative effects as the cloud fraction increases.

1 Introduction

Shallow, marine trade-wind cumuli in the tropics have been established as a key component in influencing the radiative energy budget of the Earth's atmosphere in response to a changing climate (Bony et al., 2017; Klein et al., 2017; Stevens et al., 2021). Through various research efforts, the microphysical, macrophysical and radiative properties of these clouds have been mapped out, for example with the aid of high resolution satellite observations (Zhao and Di Girolamo, 2007; Mieslinger et al., 2019) and through dedicated research campaigns such as the Cloud, Aerosol, Radiation and tuRbulence in the trade regime over Barbados (CARRIBA; Siebert et al., 2013) and the Next-Generation Aircraft Remote Sensing for Validation field studies (NARVAL I and NARVAL II; Stevens et al., 2019). Such observational data sets are ideal for the purposes of process studies and the validation and improvement of numerical weather prediction models (NWPs) and global climate models (GCMs). However, constraining a comprehensive characterization of shallow trade-wind clouds based on their macrophysical and microphysical properties and assessing the associated dynamics remains difficult. An often cited reason for this is the lack of representative observations (e.g. Zhang et al., 2013; Herwehe et al., 2014; Khain et al., 2015; Bony et al., 2017). Specifically, though model studies are

25 able to quantify the radiative budget, it is challenging to obtain the suitable microphysical and radiative measurements for verification. Thus, the impact of shallow trade-wind clouds on the atmospheric radiative energy budget in a warming world remains a critical topic.

Despite this challenge, the influence of cloud properties on their radiative effects is already well known. For solar (shortwave) radiative effects, cloud fraction has been established as a useful parameter for estimating the radiative properties or effects of

30 the clouds for a given scene (e.g., Chen et al., 2000), particularly by using the relation between albedo, which is the combination of cloudy- and cloud-free-sky albedo, and cloud fraction (George and Wood, 2010; Bender et al., 2011, 2016; Engström et al., 2015). However, this relation can be further influenced by variations in liquid water path (LWP) when cloud fraction is held constant (Bender et al., 2016). Cloud microphysical characteristics have been found to have less of an impact, but George and Wood (2010) suggest that these impacts may be masked by the complex interactions between macrophysical and meterological

35 variability. Assessments of the cloud albedo demonstrate how changes in aerosol concentrations (Peng et al., 2002; Diamond et al., 2020), and thus droplet size via the Twomey effect (Twomey, 1977; Werner et al., 2014), enhances cloud albedo and cloud optical thickness due to a decrease in droplet effective radius (r_eff r_eff). However, this relation between r_eff r_eff and cloud albedo differs between optically thick and thin clouds (Lohmann et al., 2000), likely stemming from a dependence on whether the cloud is precipitating, which impacts LWP and the droplet size distribution. In optically thick clouds, cloud albedo increases

40 with decreasing r_eff r_eff, while cloud albedo decreases with decreasing r_eff r_eff in optically thin clouds. The spatial distribution or organization of the cloud field, which is strongly related to cloud fraction and LWP, also impacts the resulting albedo (Wood and Hartmann, 2006; McCoy et al., 2017).

With regard to terrestrial (longwave) cloud radiative effects, macrophysical properties are known to have the greatest influence. Cloud top height (or temperature) has been found to have the most impact, but cloud amount is also important (Ardanuy

45 et al., 1988; Chen et al., 2000). As described by Chen et al. (2000), because cloud top height is so critical to terrestrial cloud radiative effects, the cloud type is also known be an important factor. Among shallow clouds, stratocumulus along with stratus and altostratus, have the greatest impact on surface terrestrial effects. The organization of the cloud field also plays a role. For example, Tobin et al. (2012) found that the outgoing terrestrial radiation of deep convective clouds decreased as the cloud field became more aggregated. Whether this holds true for shallow trade-wind clouds is explored in this study.

50 Thus far, the observational research in this area has largely relied on satellite-based measurements of the cloud radiative effects, which has the benefit of providing long-term, global observations, but often with the trade-off of a coarser spatial and temporal resolution. Furthermore, the radiative flux densities are usually derived from combinations of radiance observations and radiative transfer models or the Clouds and the Earth's Radiant Energy System (CERES) instruments. ~~This means that~~ Thus, in these cases, irradiance is not measured directly, ~~which is one reason why airborne observations of this quantity are so~~

55 ~~important. The~~ and uncertainties are introduced due to the assumptions made in the calculations of irradiance. However, the recent ElUcidating the RolE of Cloud-Circulation Coupling in ClimAte (EUREC[4]A) campaign in 2020 used multiple platforms (e.g. airborne and satellite) and instrumentation methods to assess the cloud field from multiple angles, both literally and with regard to research topics, and with a strong degree of collocation (Stevens et al., 2021). Among these measurements are in-situ

airborne observations of solar and terrestrial broadband irradiance by broadband radiometers, which directly measure irradiance
60   without the constraints of satellite-based measurements.

With these new radiation observations, the major objective of this study is to understand how and to what degree the macro-physical and microphysical cloud properties influence the cloud radiative effects of shallow trade-wind clouds. Furthermore, we seek to understand their relative importance in driving the cloud radiative effects and how different properties might work together or against each other to result in a specific cloud radiative effect. In Sect. 2, the relevant instrumentation, data products,
65   and analysis methods for this study are discussed. The results of the analysis are presented in Sect. 3 with different approaches to exploring the data and the relations between the cloud properties and the solar and terrestrial cloud radiative effects. Section 4 concludes by summarizing and discussing the results and their implications as well as providing the context for future work.

2   Methodology

The analysis presented here is based on data taken during the EUREC^{4}A campaign (Bony et al., 2017; Stevens et al., 2021),
70   which took place in the winter of 2020. During this season, the trade-wind region is characterized by more cloud cover and stratiform structures in comparison to the rather convective, wet summer season (Nuijens et al., 2014). The general aim of this campaign was to sample the clouds and large-scale dynamics in the trade-wind region of the Atlantic Ocean, just off the coast of Barbados. With the participation of four research vessels, five aircraft, three remote sensing stations on Barbados and many unmanned research crafts, the synergistic approach of this campaign offered the opportunity to observe the trade-wind region
75   from several perspectives. Specifically, data from the High Altitude and LOng Range Research Aircraft (HALO) operated by the DLR (Deutsches Zentrum für Luft- und Raumfahrt, e.V.) and the Geostationary Operational Environmental Satellite 16 (GOES-16) operated by the National Oceanic and Atmospheric Administration (NOAA) are utilized in this study.

2.1   Airborne observations using HALO

As in NARVAL I and NARVAL II, which previously took place in this region, an instrument payload known as the cloud-
80   observatory configuration was installed on HALO (Stevens et al., 2019). This payload consists primarily of remote sensing instruments, comparable to the instrumentation of the current A-Train satellite constellation and the upcoming Earth Cloud Aerosol and Radiation Explorer (EarthCARE) satellite. For EUREC^{4}A, the instrumentation was extended with the Broadband AirCrAft RaDiometer Instrumentation (BACARDI; Konow et al., 2021) and the Video airbornE Longwave Observations within siX channels (VELOX; Schäfer et al., 2021) system. This instrument payload was designed to observe the cloud population as
85   well as the surrounding environment (i.e. dynamic and thermodynamic conditions) simultaneously. Active and passive sensors provide a characterization of the cloud properties – microphyiscal, macrophysical and radiative – in vertical and horizontal dimensions, while methodically released dropsondes along the flight path obtain ~~in situ~~ in-situ measurements of the vertical atmospheric profile. See Konow et al. (2021) for further information concerning the instrumentation and flights of HALO during EUREC^{4}A.

For the purposes of the campaign, specifically as a strategy for properly capturing the scale of the dynamics influencing cloud development, HALO primarily flew in a pattern defined by clockwise circles repeated in the same location for most flights (centered at 13.3° N, 57.717° W), approximately 220 km in diameter and requiring an hour to complete. As the synoptic situation is largely stable in this region with regularly occurring clouds (Bony et al., 2017), the chosen flight strategy for the campaign was independent of the meteorological situation and instead focused on achieving flights at different times of day to fully capture the diurnal cycle in the region (Vial et al., 2019; Konow et al., 2021). Other flight patterns were also observed for additional purposes, but the work here focuses on the circles so as to capture and constrain the observations used in this analysis to a single location.

The radiative component of this analysis is estimated from observations from BACARDI (Konow et al., 2021). This new radiometer package comprises two sets of Kipp and Zonen pyranometers (CMP22) and pyrgeometers (CGR4) mounted to the aircraft fuselage. Together, they provide measurements of the upward and downward solar (sensitive to wavelengths between $0.2-3.6\,\mu m$) and terrestrial ($4.5-42\,\mu m$) broadband irradiance at flight level. The respective uncertainties of the measurements from these sensors are 1 % and 4 %, which are calculated during calibration procedures before and after the campaign (Kipp and Zonen B.V, 2014, 2016). The data used here has been corrected in post-processing for temperature dependence and response time of the sensors as well as aircraft attitude, similar to the methods described by Ehrlich and Wendisch (2015).

2.2 Satellite observations from GOES-16

The cloud properties in this analysis come from the data products of the GOES-16 geostationary satellite Advanced Baseline Imager (ABI) (Schmit et al., 2017, 2018). GOES-16 is part of the GOES-R series of satellites. Throughout the duration of the campaign, GOES-16 was operated to produce images of the EUREC^{4}A research domain at a temporal resolution of 1 minute, when possible, which turns out to be a strong benefit for analyzing the evolution of the clouds. Specifically, we utilize the cloud mask product (Heidinger and Straka, 2012) as well as retrieved LWP, r_{eff} (Walther et al., 2013) and cloud top height (z_{ct}; Heidinger, 2012) products. The spatial resolution for the data products used here is 2x2 km at nadir. The cloud mask shows good agreement with the Cloud-Aerosol Lidar and Infrared Pathfinder Satellite Observation (CALIPSO) over the Contiguous United States, but low clouds (cloud tops below 2 km) were the most commonly missed cloud type (Jiménez, 2020). Also, further evaluation shows that the increased pixel resolution relative to its predecessor, GOES-13, results in an improved r_{eff} retrieval in comparisons with airborne measurements (Painemal et al., 2021). It is therefore expected that the LWP retrieval is also improved, however, both products continue to have a high bias.

For the application of the cloud mask product from GOES-16, pixels labeled as "probably cloudy" and "cloudy" are accepted as cloud pixels in the calculation of the cloud fraction. The cloud mask is also used to calculate the organization of the cloud field, which is described by the organization index, I_{org} (Weger et al., 1992; Tompkins and Semie, 2017). This index classifies a given cloud scene as regular, random, or clustered. The classification is based on statistics of the nearest neighbor (NN) distances between clouds in comparison to a theoretical scene with a random distribution, which can be described as a Poisson point process. If the NN distances are on average larger than they would be if the clouds were randomly distributed

(I_{org} < 0.5), then the scene is regular. If they are smaller (I_{org} > 0.5), the scene is classified as clustered. An I_{org} value of 0.5 indicates that the scene is random. More detailed explanations can be found in Tompkins and Semie (2017) and Mieslinger et al. (2019).

2.3 Deriving cloud radiative effects

Cloud radiative forcing (ΔF) is defined as the difference between the observed net (downward - upward) irradiance ($F^\downarrow$ - $F^\uparrow$) at an altitude (z, in this case flight altitude) and the net irrandiance in cloud-free conditions, as shown in Eq. 1 (Stapf et al., 2021):

$$\Delta F(z) = \left[F^\downarrow(z) - F^\uparrow(z) \right]_{\text{cloud}} - \left[F^\downarrow(z) - F^\uparrow(z) \right]_{\text{cloud-free}}. \tag{1}$$

ΔF can also be further divided into its terrestrial (ΔF_{ter}) and solar (ΔF_{sol}) components, which are defined in Eqs. 2 and 3, respectively:

$$\Delta F_{\text{ter}}(z) = \left[F^\downarrow_{\text{ter}}(z) - F^\uparrow_{\text{ter}}(z) \right]_{\text{cloud}} - \left[F^\downarrow_{\text{ter}}(z) - F^\uparrow_{\text{ter}}(z) \right]_{\text{cloud-free}}, \tag{2}$$

$$\Delta F_{\text{sol}}(z) = \left[F^\downarrow_{\text{sol}}(z) - F^\uparrow_{\text{sol}}(z) \right]_{\text{cloud}} - \left[F^\downarrow_{\text{sol}}(z) - F^\uparrow_{\text{sol}}(z) \right]_{\text{cloud-free}}. \tag{3}$$

For ΔF_{sol}, Eq. 3 can be further simplified with the assumption that $F^\downarrow_{\text{sol}}$ is identical in cloudy and cloud-free cases while flying above cloud:

$$\Delta F_{\text{sol}}(z) = - \left[F^\uparrow_{\text{sol,cloud}}(z) \right] - \left[-F^\uparrow_{\text{sol,cloud-free}}(z) \right]. \tag{4}$$

However, ΔF_{sol}, is a function of solar zenith angle (SZA). For this reason, albedo (Eq. 5) is a more useful parameter to describe solar radiative fluxes because it expresses the relative difference between $F^\downarrow_{\text{sol}}$ and $F^\uparrow_{\text{sol}}$, thereby avoiding the SZA dependence:

$$\alpha = \frac{F^\uparrow_{\text{sol}}}{F^\downarrow_{\text{sol}}}. \tag{5}$$

With this in mind, the dependence of ΔF_{sol} on SZA can also be addressed. The same strategy applied to Eq. 3, wherein the cloud-free fluxes are subtracted from the total to isolate the cloud effect, can also be used with albedo (Eq. 5) to isolate the cloud effect on the observed albedo (α_{ce}). Simply, the cloudy and cloud-free $F^\uparrow_{\text{sol}}$ in Eq. 4 are normalized with $F^\downarrow_{\text{sol}}$, which removes the SZA dependence, thus making observations from different SZAs comparable:

$$\alpha_{\text{ce}}(z) = - \left[\frac{F^\uparrow_{\text{sol,cloud}}(z)}{F^\downarrow_{\text{sol}}(z)} \right] - \left[-\frac{F^\uparrow_{\text{sol,cloud-free}}(z)}{F^\downarrow_{\text{sol}}(z)} \right]. \tag{6}$$

Since measurements in cloud-free conditions cannot be obtained simultaneously or in similar conditions to those observed during each flight, simulated cloud-free cases of each flight are calculated using the libRadtran software package (Emde et al.,

Table 1. Estimations of the uncertainty of the simulated cloud-free conditions based on different possible sources.

Source of Uncertainty	Wavelength Range	Parameter	Uncertainty
Atmospheric Profile (dropsonde)	Solar	$F^{\downarrow}$	$< 1\,\%$
		$F^{\uparrow}$	$3\,\%$
	Terrestrial	$F^{\downarrow}$	$6\,\%$
		$F^{\uparrow}$	$8\,\%$
Surface Temperature ($\pm 3\,\mathrm{K}$)		$F^{\downarrow}$	$< 1\,\%$
		$F^{\uparrow}$	$2.5\,\%$

2016). Dropsondes along the flight track provide measured profiles of thermodynamic atmospheric properties (humidity, tem-
150 perature) used in the simulations (George et al., 2021). Estimations of the uncertainties of these simulations are given in Table
1. The largest source of uncertainty is found to be the atmospheric data, particularly for the cloud-free $F^{\uparrow}_{\mathrm{ter}}$. Due to the lim-
ited spatial resolution of the dropsondes, this method cannot always replicate an observed cloud-free situation. Additionally,
because the sondes may pass through clouds on their way down, the humidity and temperature profiles used for the simulation
could be biased. In order to minimize these effects, only the simulations at the dropsonde locations are used to approximate a
155 cloud-free $F^{\uparrow}_{\mathrm{ter}}$, which is then interpolated for the rest of the flight track. It should also be noted that aerosol particle properties
were not included in the simulations, as their effect is anticipated to be minimal relative to clouds, but their influence may still
be included in the cloud effects calculated here.

2.4 Combining satellite and aircraft observations

In the case of the analysis presented here, only a subset of seven flights are used, which are described in Table 2. Those
160 times when data from both BACARDI and GOES-16 were available are given there. Unfortunately, this also means that only
partial flights are usable in some cases. Additional filters were applied regarding the SZA and the presence of clouds above the
aircraft (e.g. cirrus). To prevent complications in the case of low sun conditions and to avoid the need to consider cirrus in this
study, the data were limited to SZAs less than 70° and to flights with cloud-free conditions above the flight altitude of HALO.
Furthermore, the maximum mean z_{ct} for an observed scene is limited to $4\,\mathrm{km}$ to avoid the inclusion of deeper convective
165 clouds, which have properties and associated relationships that should be considered separately. Scenes without clouds below
HALO are also excluded.

To enable the combined use of BACARDI and GOES-16 data, spatial collocation needs to be assured with both observations
covering the same footprint at the same time. BACARDI has a 180° field of view (FOV), but radiation from different directions
is weighted with the viewing zenith angle. Based on geometry, 95 % of the upward irradiance is determined by radiances within
170 a FOV of approximately ~~77~~154°. However, such a footprint may not necessarily be compatible with the radiation measured
by GOES-16. ~~Therefore, as a quality check , we compared the calculated α_{ce} from BACARDI with the mean reflectance from~~

Table 2. The flights from EUREC⁴A that are used in this analysis are given here with their respective flight IDs (HALO-mmdd) along with the times when HALO was in the circling flight pattern and GOES-16 data were available. The amounts of time and number of GOES images are for cloudy and cloud-free sky, but a total amount of time and number of GOES images for cloudy sky only is given as well. The general descriptions of the flights come from notes in the flight reports.

Flight ID	Time (UTC)	Duration (hh:mm)	GOES Images	General Description
HALO-0124	11:52–16:56	5:05	305	Shallow Cu, Sc, deeper convection
HALO-0128	15:18–18:38; 19:53–20:21	3:40	220	Shallow Cu clusters and lines
HALO-0131	15:28–18:50; 19:56–20:23	3:51	231	Dust, suppressed shallow Cu, "fish" clusters, Sc
HALO-0202	18:29–19:34	1:06	66	Large "flowers", some dissipating
HALO-0207	15:00–15:34; 17:14–20:23	3:45	225	Shallow Cu, "flowers", precipitation
HALO-0209	11:47–12:56; 13:54–17:24	4:41	281	Shallow and deeper Cu, cold pools, precipitation
HALO-0211	12:57–18:55	5:58	358	Convection with precipitation, Sc, shallow Cu
Total		**28:06**	**1686**	
Total (cloudy scenes only)		**14:02**	**842**	

the 0.65 μm channel for footprints cut out from A quality check was performed to ensure the compatibility of BACARDI and GOES-16 of varying sizes. Because we are comparing different quantities, they will not match exactly, but we assume that the footprint size at which they show the best correlation represents a footprint size where GOES-16 observations are comparable with BACARDI data. A FOV of 51° was found to provide the best correlation (R = 0.968). An example of this comparison from the flight on 24 January 2020 for , and it was found that a BACARDI footprint corresponding to a FOV of 51 102° is shown in Fig. A1. Based on the selected FOV and appropriate to use with observations from GOES-16. At a typical flight altitude of 10 km, the most appropriate footprint size has a FOV of 102° corresponds to a footprint with a radius of approximately 13 km. This quality check is described in more detail in Appendix A.

Comparison of α_{ce} from BACARDI and the mean reflectance from the 0.65 μm channel of GOES-16 for a FOV of 51° from the flight on 24 January 2020.

Additionally, the cloud scenes observed over the course of each minute are oblong instead of perfectly circular due to the movement of the aircraft. Over the course of a minute, the aircraft travels approximately 12 km with an average ground speed of $200\,\mathrm{m\,s^{-1}}$, so the area of each one minute scene is approximately $820\,\mathrm{km^2}$. Examples of the GOES-16 cloud mask product and the corresponding footprint cutout from the flight on 24 January 2020 are shown in Fig. 1.

Mean values are used to describe the cloud properties of each scene. The individual quantities are weighted with the cosine of the viewing angle across the footprint of BACARDI before calculating a mean value. This strategy is used to capture an expression of the properties that is compatible with the radiative view that BACARDI has (i.e. clouds on the periphery make less of an impact than clouds in the nadir view of the instrument). Furthermore, two representations of the LWP are used – mean LWP calculated for the cloud pixels only ($\mathrm{LWP_{cloud}}$) and a mean LWP calculation that includes both the cloudy and
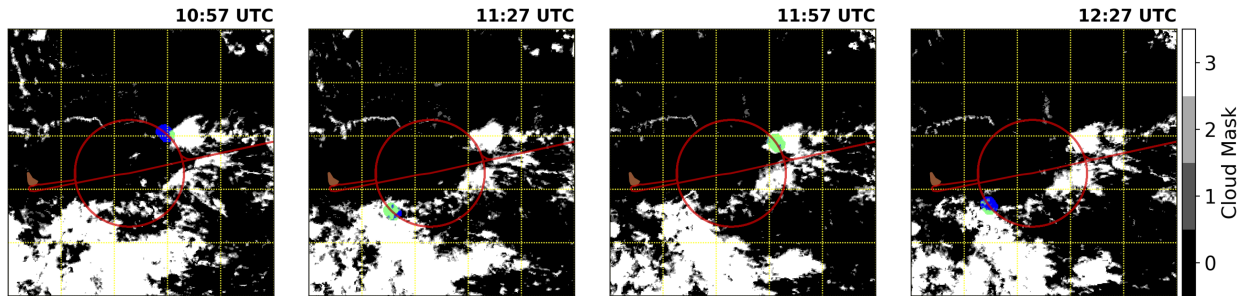
Figure 1. Four example scenes from the flight on 24 January 2020. The cloud mask product from GOES is displayed, with the blue (cloud free) and green (cloud) colors indicating the collocated ~~51~~102° footprint of BACARDI. The solid red line indicates the flight track. The island of Barbados is shown in brown for reference.

cloud-free pixels within the footprint-sized cloud scene (LWP_{scene}). LWP_{scene} considers LWP in a more macrophysical sense because of its relationship to cloud fraction and the fact that it represents the amount of water distributed over the scene, while LWP_{cloud} considers LWP as microphysical in nature.

The size of the footprint cut out from the full GOES scene is not sufficient for a meaningful quantification of the cloud organization. Trade-wind cumuli typically have sizes (measured as maximum diameter) from a few hundred meters to less than 2 km (Schnitt et al., 2017), so only a few clouds are present within the GOES cutouts. However, assuming that the cloud organization is driven on larger scales, we used a larger domain size of 5°x5° for this assessment of organization. Given the coarse resolution of GOES-16 this consideration seems appropriate. The full example scenes in Fig. 1 are the larger domain size, but for the analysis, the center of each larger domain would be aligned with the location of HALO at that time.

3 Sensitivity of cloud radiative effects to cloud properties

3.1 Overview of cloud data

The radiative data in this study was recorded by the BACARDI instrument and the cloud properties were obtained from GOES-16 satellite data products (see Sects. 2.1 and 2.2). A set of histograms depicting the distributions of the data is given in Fig. 2. The mean α_{ce} is 0.086, while the mean ΔF_{ter} is $6.35\,\mathrm{W m^{-1}}$. The mean ΔF_{sol} is $-77.64\,\mathrm{W m^{-1}}$, which demonstrates that the solar part of the cloud radiative forcing of shallow trade-wind cumuli is significantly stronger than the forcing from the terrestrial part. However, due to the influence of SZA on ΔF_{sol}, the mean value is provided only as an estimate for the purposes of comparison. The mean cosine weighted cloud fraction is 0.204, which is larger than the value of 0.087 reported by Mieslinger et al. (2019) from Advanced Spaceborne Thermal Emission and Reflection Radiomater (ASTER) imagery of trade-wind cumuli. Differences are likely to stem from the exclusion of non-cloudy scenes in this study as well as the small footprint and coarseness of the GOES satellite resolution. The mean z_{ct} of 1.9 km found here is slightly higher than the 1.3 km reported in their study, and the cloud field is found to be less clustered on average (~~I_{org}~~ I_{org} of 0.74 versus 0.89). The overall

8

observed values are within similar ranges, so the smaller sample size within a single season may also be the cause of differences to previous studies. The mean $\text{LWP}_{\text{scene}}$ and mean $\text{LWP}_{\text{cloud}}$ are $19\,\text{gm}^{-2}$ and $98\,\text{gm}^{-2}$, respectively. For comparison, during the NARVAL campaigns, retrievals of LWP from remote sensing instrumentation reported a mean LWP from the sampled clouds of about $63\,\text{gm}^{-2}$ during the dry, winter season (Jacob et al., 2019). The fact that thicker clouds ($\text{LWP} > 50\,\text{gm}^{-2}$) were also more frequent during this season was also noted. Mean r_{eff} in this study is $16\,\mu\text{m}$, which is comparable to previous observations in the vicinity of Barbados (10–17 µm in pristine cases; Werner et al., 2014).

3.2 Observed relations between cloud properties and cloud radiative effects

The impact of cloud properties on the radiation budget was analyzed separately for the solar and terrestrial components of the observed radiation because these two components are known to interact differently with cloud properties. A series of scatter plots in Figs. 3 and 4 demonstrate the relationships of α_{ce} and ΔF_{ter}, respectively, to the various cloud properties analyzed here.

In Fig. 3, it becomes apparent that the macrophysical properties (cloud fraction and $\text{LWP}_{\text{scene}}$) demonstrate the most obvious relationship to α_{ce} – as α_{ce} increases, the cloud fraction and $\text{LWP}_{\text{scene}}$, panel (a) and (b), also increase, as expected. The relation between α_{ce} and z_{ct} (Fig. 3(f)) is more weakly correlated, but generally as z_{ct} increases, so does α_{ce}. In the case of $\text{LWP}_{\text{cloud}}$ (Fig. 3(c)), two populations emerge – one where LWP and α_{ce} are strongly positively correlated and one where an increase in LWP does not appear to impact α_{ce}. This second branch may represent cases where changes in microphysical properties are less relevant to the radiative impact of the clouds. From the perspective of a scatter plot, the relationships of α_{ce} to r_{eff} and I_{org} are unclear given the very weak correlations (Fig. 3(d) and Fig. 3(e)).

Based on these results, we would expect to see the same patterns to be observed for individual flights as well. The different colored points and bars in Fig. 3 depict the median values (see also Table 3) of each parameter for each of the seven investigated flights and the range from the 10th to 90th percentiles of the different distributions, respectively. The clouds observed during the flight on 28 January 2020 (red star) have the lowest α_{ce}, while the highest median α_{ce} was observed during the flight on 24 January (black circle). The expectation that follows is that the cloud properties will also display a pattern according to their relationship to α_{ce}. For example, properties like cloud fraction and $\text{LWP}_{\text{scene}}$ are also generally increasing in a similar flight order as the α_{ce}, with 28 and 31 January showing some of the lowest cloud fraction and $\text{LWP}_{\text{scene}}$ values and 24 January showing the highest. However, the observations on 7 February (green plus) do not follow this trend for both cloud fraction and LWP. The remaining properties ($\text{LWP}_{\text{cloud}}$, r_{eff}, I_{org} and z_{ct}) appear not to follow any pattern given the order of the flights with increasing α_{ce}. This demonstrates again that the macrophysical properties are more closely linked to the observed α_{ce} than the microphysical properties.

From this series of panels in Fig. 3, it is also possible to pick out details about each flight and identify relationships between cloud properties and α_{ce}. For example, the α_{ce} values measured during the flight on 31 January (orange triangle) fall on the relatively lower end of observed α_{ce} as well as similarly low cloud fraction and $\text{LWP}_{\text{scene}}$. The mean $\text{LWP}_{\text{cloud}}$ is relatively high and the r_{eff} relatively low. On this day, the cloud field was characterized as quite clustered with lower cloud tops. During this flight, the cloud field was populated with suppressed shallow cumulus, so-called "fish" structures (Schulz et al., 2021) and
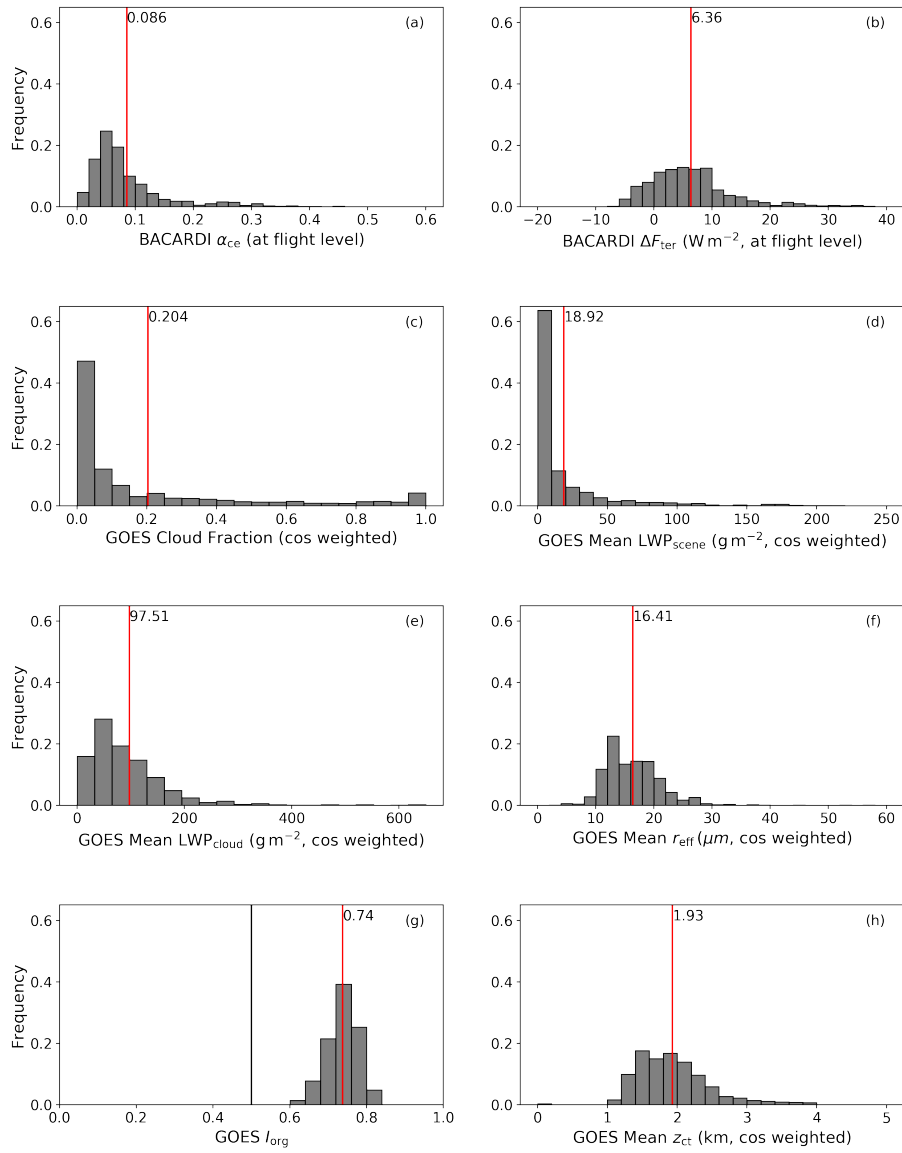
Figure 2. Histograms of all cloud properties for the subset of EUREC[4]A flights and data used in the analysis. Labeled red bars indicates the mean value. (a) Effect of clouds on albedo (b) Terrestrial cloud radiative forcing (c) Cloud fraction (d) Mean liquid water path for the scene (e) Mean liquid water path for the clouds (f) Mean effective radius (g) Organization index; black bar at 0.5 separates normal and clustered ~~I_{org}~~ I_{org} values (h) Mean cloud top height.
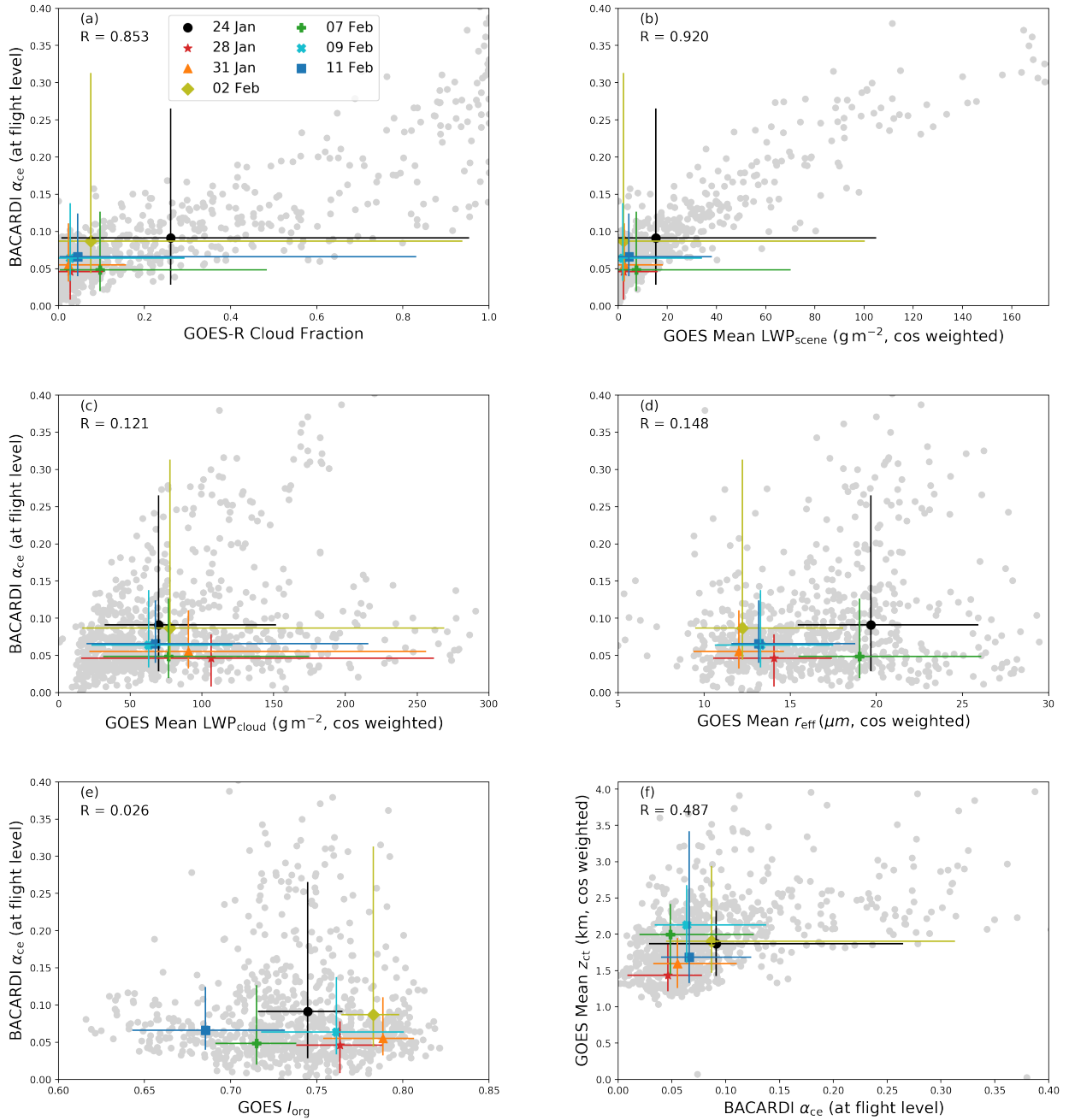
Figure 3. (a) - (f) Scatter plots (gray) of different cloud properties as a function of α_{ce} for the subset of EUREC[4]A flights used in the analysis. Different colored shapes denote the median observed values from individual flights, while the associated bars indicate the range of values between the 10th and 90th percentiles. Plots are zoomed in to include at least the 10th–90th percentile ranges for each variable.

Table 3. Median values shown in Figs.3 and 4. The highest and lowest values for each parameter are marked in bold.

	24 Jan	28 Jan	31 Jan	02 Feb	07 Feb	09 Feb	11 Feb
α_{ce}	**0.091**	**0.046**	0.055	0.087	0.048	0.064	0.066
ΔF_{ter} (Wm^{-1})	6.38	3.73	2.89	**-0.81**	5.36	1.94	**9.47**
Cloud fraction	**0.261**	0.027	**0.023**	0.075	0.096	0.028	0.045
LWP$_{scene}$ (gm^{-2})	**15.34**	2.19	2.44	2.24	7.39	**1.92**	4.32
LWP$_{cloud}$ (gm^{-2})	69.73	**106.43**	90.63	77.77	76.81	**62.82**	67.64
r_{eff} (µm)	**19.68**	14.06	**12.02**	12.22	19.03	13.26	13.18
I_{org}	0.74	0.76	**0.79**	0.78	0.72	0.76	**0.69**
z_{ct} (km)	1.87	**1.43**	1.60	1.91	2.00	**2.13**	1.68

lots of dust. The dust could be an explanation for why the r_{eff} is quite small, and the notion of suppressed shallow cumulus fits to the lower cloud top heights, but confirmation of this assertion would require further information and analysis. While small droplets typically indicate higher α_{ce}, the characteristics of the macrophysical properties seem to be more dominant in this case. In contrast, the flight on 7 February (green plus) has a lower α_{ce} relative to other flights, but has a higher cloud

250 fraction and LWP$_{scene}$. The LWP$_{cloud}$ is also relatively lower, the r_{eff} is relatively larger, the cloud tops are higher and the cloud field is less clustered than most flights. This flight was dominated by cloud structures known as "flowers" (Schulz et al., 2021) and precipitation was noted. Thus, it could be possible that despite the higher cloud fraction and therefore higher amount of liquid water distributed in the scene, which can be attributed to the large "flower" clouds, the clouds themselves had a lower LWP and droplets with a higher r_{eff}. This could demonstrate a case wherein the macrophysical and microphysical

255 properties counteract each other and raises new questions to be answered. For example, does this always happen or are there specific requirements for this? In the case of the 31 January flight, do the microphysical properties make a contribution to α_{ce} at all? If, for example, the r_{eff} had been larger, would the α_{ce} have been noticeably lower?

Another feature of the measurements collected on 7 February that is worth noting in Fig. 3 is that the percentile line for cloud fraction reveals that a relatively large range of cloud fractions were sampled throughout the flight. This is true also for

260 the flights on 24 January, 2 February and 11 February. The implication here is that a more diverse range of clouds and/or dynamic situations were encountered and sampled in a single flight, unlike 31 January, for example. Therefore, Fig. 3 may be an incomplete picture as the inclusion of multiple cloud situations in this statistical view could certainly make the interpretation more difficult.

The terrestrial component of the cloud radiative effect has a different relationship to the same cloud properties. Figure

265 4 contains a series of scatter plots showing the general relationship of different cloud properties to the ΔF_{ter} calculated using Eq. 1. The majority of the calculated ΔF_{ter} values are positive, indicating a weak warming effect above the clouds, with some indication of weak cooling for the lowest clouds (below 2.5 km). This weak warming is typical for low-level clouds given the

low temperature contrast to the warm surface (Chen et al., 2000). The cooling, however, might fit into the uncertainties of the observations and simulations, as mentioned in Sect. 2.3. Similar to α_{ce}, the cloud fraction (Fig. 4(a)) appears to be strongly tied to the resulting ΔF_{ter}. The relationship to z_{ct} also appears more pronounced, showing that z_{ct} is linked to a decrease in cloud top temperature as expected (Fig. 4(f)). Consequently, the cloud warming effect increases in magnitude. In contrast to what is shown in Fig. 3, both the LWP$_{scene}$ and LWP$_{cloud}$ (Fig. 4(b) and (c)) show a difficult to interpret relationship to ΔF_{ter}. Also, the relation between ~~r_{eff} and I_{org}~~ r_{eff} and I_{org} to ΔF_{ter} (panels (d) and (e)) is potentially masked by the dominating effects of cloud fraction and z_{ct}. Additional methods are required to untangle these relationships.

In Fig. 4, the lowest median ΔF_{ter} was observed during the flight conducted on 2 February (yellow diamond), while the highest was on 11 February (blue square). The order of flights by increasing median ΔF_{ter} is different than that of increasing α_{ce} in Fig. 3. In contrast to α_{ce}, the pattern of increasing median ΔF_{ter} with increasing cloud fraction is not as clear. The z_{ct} relationship, which is more clear in the scatter plot in Fig. 4 is not clearly captured in the pattern of the individual flights. This is perhaps due to the fact that between days, other properties like the atmospheric profile are changing, which could mask the z_{ct} effect. For a constant temperature and humidity profile between days, we would expect the relation between ΔF_{ter} and z_{ct} to be more clear among the individual flights. Although it is poorly correlated in the subset of data as a whole, the most defined relationship found here between individual flights is with the organization of the cloud field, ~~I_{org}~~ I_{org}. As the cloud field becomes more clustered, ΔF_{ter} decreases. This is in contrast to the results of Tobin et al. (2012), but their results pertain to deep convection, and here shallow clouds are the focus.

Looking again at some flights individually, we can deduce why the observations from different flights may look the way they do. If we look at the flight on 24 January (black circle), this flight stands out in Fig. 4(a) and (b) as having the highest median cloud fraction and LWP$_{scene}$, but not the highest ΔF_{ter}. The LWP$_{cloud}$ is relatively low and the R$_{reff}$ relatively large, and the cloud field has a middle range clustering compared to the other flights. The z_{ct} is also in a middle range. Using cloud fraction alone, we would expect a higher ΔF_{ter} relative to other clouds, but based on the results in this figure, the ~~I_{org}~~ I_{org} or z_{ct} could be factors dampening or overtaking the impact of cloud fraction. The high z_{ct} on the 9 February flight may also be having a similar impact since the median ΔF_{ter} is high considering the low median cloud fraction and LWP$_{scene}$ also observed during that flight.

3.3 Standardized regression analysis

To better understand the relative influence that different properties or the organization of the cloud field have, we use a parameter known as a beta coefficient (e.g., Neter et al., 1983; Bring, 1994). A beta coefficient represents the slope of the linear regression calculated for variables that have been standardized, so that they have a mean of 0 and a standard deviation of 1. Standardization is simply carried out with

$$Z = \frac{x - \mu}{\sigma}, \tag{7}$$

where Z is the standardized version of variable x, μ is the mean and σ is the standard deviation. The benefit of using this method is that the variables are independent of their units, and the slopes of the linear regression of each cloud parameter with
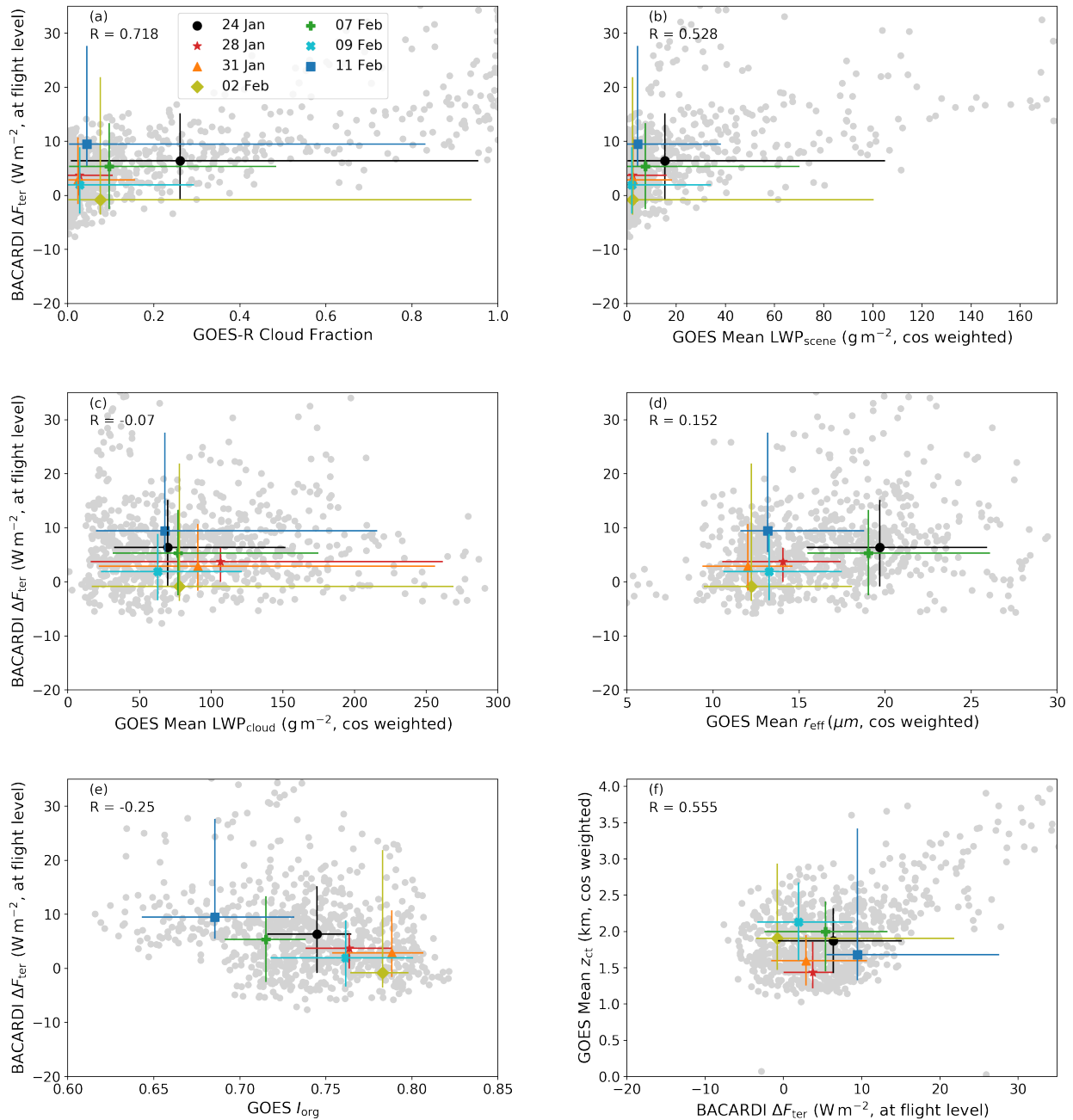
13

Figure 4. Same as Fig. 3 but for ΔF_{ter}.

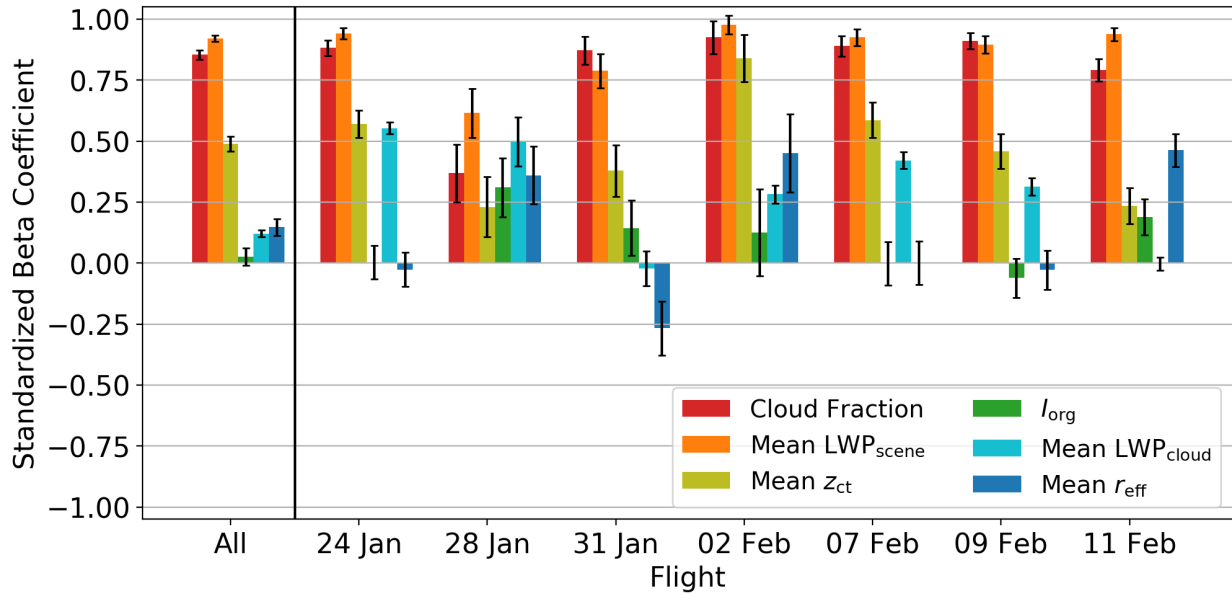**Figure 5.** Beta coefficients demonstrating the relative correlation of different cloud properties and the organization of the cloud field to α_{ce} for the subset of flights (all) and individual flights.

α_{ce} and ΔF_{ter} can be compared, thus indicating how much a change in any one variable leads to a change in α_{ce} or ΔF_{ter}, respectively.

Figure 5 shows the beta coefficients for each variable for the subset of flights as a whole as well as for individual flights. Cloud fraction and LWP$_{\mathrm{scene}}$ show the highest correlation to α_{ce} for the subset of flights anaylzed here (all flights) and for the individual flights. z_{ct} is also relatively high for all flights and some individual flights, whereas the correlations of the remaining parameters are low in most cases. It becomes clear that cloud fraction and LWP$_{\mathrm{scene}}$ are the main drivers of α_{ce} for both the campaign data set as a whole and for individual flights. The microphysical properties, on the other hand, are less straightforward to interpret. For all flights, neither LWP$_{\mathrm{cloud}}$ nor ~~r_{eff}~~ r_{eff} demonstrate a strong correlation to α_{ce}. The organization of the cloud field produces a similarly weak correlation.

Looking at individual flights, the overall correlation of LWP$_{\mathrm{cloud}}$ is weakly positive, but for flights on the 24 and 28 January the LWP$_{\mathrm{cloud}}$ demonstrates a stronger influence on the α_{ce}. ~~r_{eff}~~ r_{eff} is also an interesting parameter due to the fact that overall and in some individual flight cases, it has a positive correlation to α_{ce}, while in other cases, like 31 January, the correlation is negative. Also, the importance of z_{ct} varies between flights, while the importance of ~~I_{org}~~ I_{org} maintains a low level of correlation to α_{ce}.

In the case of ΔF_{ter}, for all flights and most of the individual flights, it is clear that the macrophysical properties are also the most highly correlated to ΔF_{ter}, although not as strongly as with α_{ce}. The 28 January and 9 February show the ~~r_{eff} and I_{org}~~ r_{eff} and I_{org} to have the largest impact on ΔF_{ter}, and with a magnitude stronger than what was observed for α_{ce}. z_{ct} is also
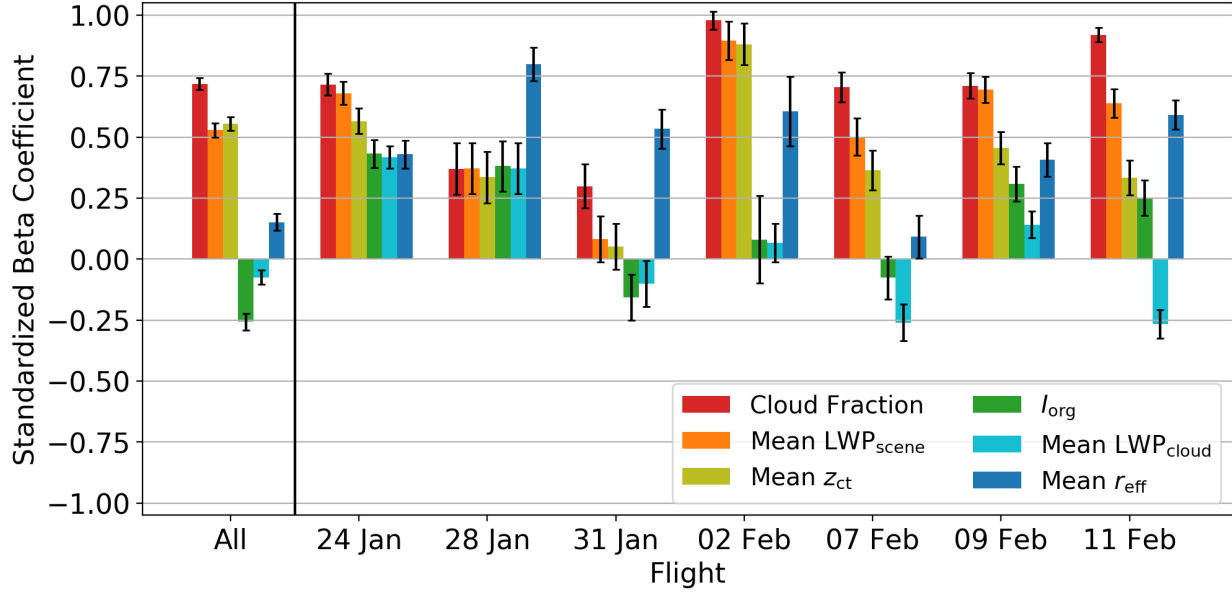
Figure 6. Same as Fig. 5 but for ΔF_{ter}.

generally strongly positively correlated with ΔF_{ter}. ~~I_{org}~~ I_{org} often demonstrates a stronger correlation to ΔF_{ter} than α_{ce}, but the positive or negative sign of that correlation changes from flight to flight, which suggests that the effects of this parameter

320 could be linked to other properties or processes. Looking at all flights shows that correlation to be negative. $\text{LWP}_{\text{cloud}}$ shows a mostly minor correlation with ΔF_{ter}.

It should be noted that while the beta coefficients are a useful tool for pointing out the correlation of the different variables to α_{ce} and ΔF_{ter}, they are not necessarily good indicators of how the different variables work in concert with each other. Furthermore, the differences between flights rely very strongly on the individual distribution of the data for each flight. Because

325 of the limited amount of data from each flight, outliers can significantly influence the results of a linear regression approach. In some cases it is not possible or is unreliable as evidenced by the large error bars in Figs. 5 and 6. Thus, such results should be interpreted with caution.

3.4 Separation into cloud fraction regimes

The analysis has shown that macrophysical properties, cloud fraction and $\text{LWP}_{\text{scene}}$, are the main drivers of the α_{ce}. To extract

330 the impact of the microphysical properties, the data is divided into different cloud fraction classes – low cloud fraction for values below 33 %, middle cloud fraction for values between 33 % and 66 % and high cloud fraction for values greater than 66 %. This approach also allows for the exploration of those cases shown in Figs. 3 and 4 where the percentile range lines reveal that a large range of cloud fraction values were sampled in a single flight. The data is further categorized into different $\text{LWP}_{\text{cloud}}$, ~~$r_{\text{eff}}, I_{\text{org}}$~~ $r_{\text{eff}}, I_{\text{org}}$ and z_{ct} ranges, and the results are shown in Fig. 7. It should be noted that due to the limited
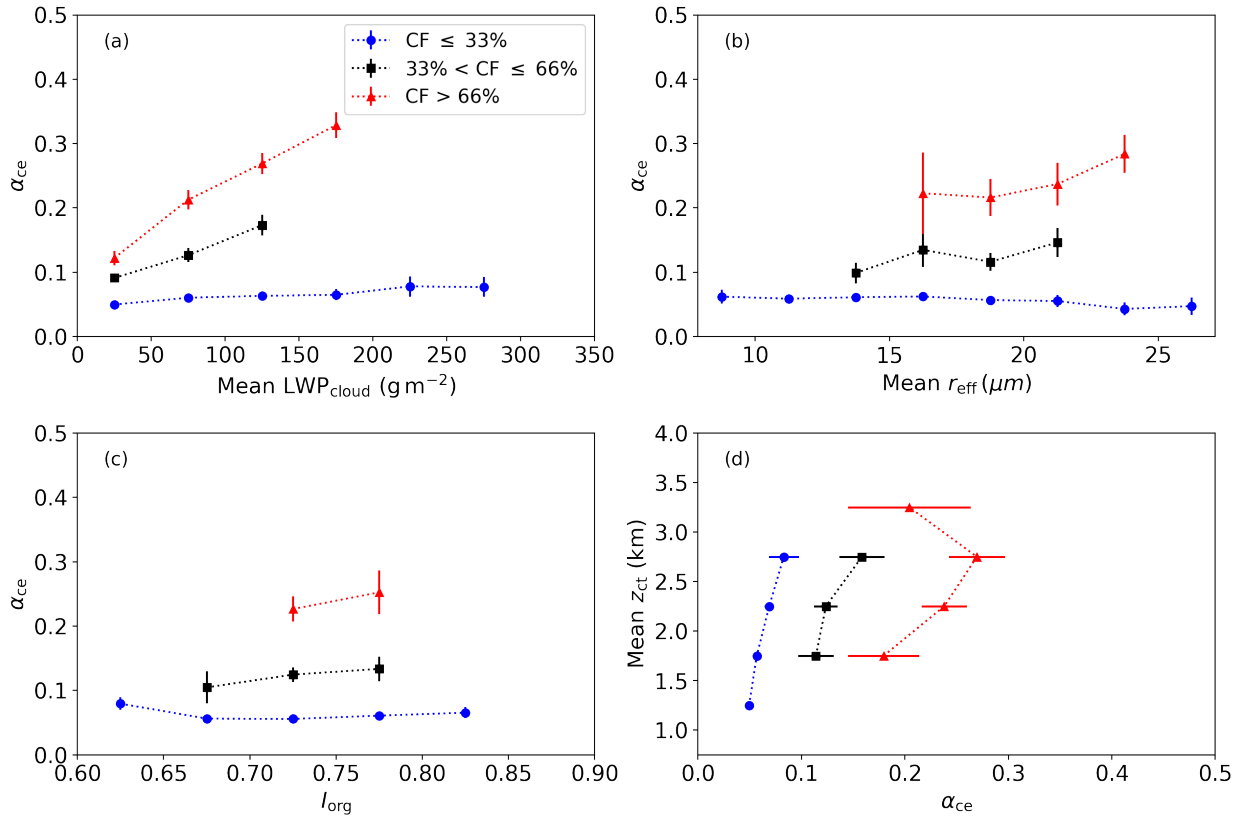
Figure 7. The relationship between α_{ce} and different cloud properties also as a function of different cloud fraction classifications. (a) LWP$_{cloud}$, (b) ~~r$_{eff}$~~ r_{eff}, (c) ~~I$_{org}$~~ I_{org} and (d) z_{ct}. The bars indicate 95% confidence intervals.

amount of data, this evaluation is not carried out for individual flights, where we might expect to see certain details expressed differently for different situations.

The range of values observed in the data is of course much more diverse than what is shown here, but only small amounts of data are found at the extremes of those ranges. Thus, in an effort to preserve the statistical representativeness of the data, only cloud property bins with at least 10 data points are shown. This explains why the lowest cloud fraction bin has the largest ranges among the different cloud properties; the majority of the data observed during this campaign had low cloud fractions.

From this figure, the effect that different cloud properties have on cloud scenes of similar cloud fractions can be ascertained. For all cloud properties, an increase or decrease in that property has a limited effect on the α_{ce} when the cloud fractions are low. Specifically, there is a weak increase in α_{ce} with increasing LWP$_{cloud}$ and z_{ct} and a weak decrease in α_{ce} with increasing ~~r$_{eff}$~~ r_{eff}. In the case of LWP$_{cloud}$, this low cloud fraction line coincides with the consistent low α_{ce} mode in Fig. 3(c). The impact of ~~I$_{org}$~~ I_{org} on α_{ce} for low cloud fractions is not clearly defined. When the cloud fractions are in the middle range, the impact of increasing LWP$_{cloud}$ and z_{ct} becomes stronger, resulting in higher α_{ce}. Additionally, both ~~r$_{eff}$ and I$_{org}$~~ r_{eff} and I_{org} show

17

a weak positive correlation with α_{ce}. This is the first indication from the analyses presented here that the organization of the cloud field, in terms of the degree of clustering, has an impact on the solar radiative effects of the trade-wind cumuli. For high cloud fraction cases, the impact of all properties shown in Fig. 7 increases as indicated by the stronger slopes. With respect to LWP$_{cloud}$, this characterizes the mode in Fig. 3(c), where an increase in LWP$_{cloud}$ coincides with an increase in α_{ce}. Notably, around 3 km, the z_{ct} reverses its correlation to α_{ce} from positive to negative. This may be due to the limited amount of data in this cloud fraction and z_{ct} combination, which subsequently limits the representation of this group.

This result leads to the conclusion that while macrophysical properties are the main driver of the α_{ce}, as cloud fraction increases, the microphysical properties as well as the horizontal cloud field organization and z_{ct} make a greater contribution to the resulting α_{ce}. When cloud fraction values are low, the impact of those properties, regardless of magnitude, is minor.

Figure 8 shows how the relationship of different cloud properties to ΔF_{ter} changes when the cloud fraction is held constant. In the low cloud fraction group, the microphysical properties LWP$_{cloud}$ and r_{eff} appear to have no impact on ΔF_{ter} regardless of their magnitude. The impact of increasing z_{ct} is also minimal. The organization of the cloud field shows a clearly negative impact on ΔF_{ter} for this cloud fraction class; as the clustering of the cloud field increases, the ΔF_{ter} decreases. For the middle cloud fraction range, the negative correlation to I_{org} continues to be present. The relationship of ΔF_{ter} to LWP$_{cloud}$ and r_{eff} also shows a negative correlation. In this cloud fraction range and the high cloud fraction range, the relationship of ΔF_{ter} to z_{ct} is even stronger. The impact of I_{org} continues to have a negative correlation, but does not appear to increase or decrease in strength relative to the other cloud fraction groups. The impacts of LWP$_{cloud}$ and r_{eff} are unclear for the high cloud fraction range. Given that they do not decrease with increasing ΔF_{ter}, like in the middle cloud fraction regime, this could suggest that competing mechanisms are represented here. For example, because ΔF_{ter} is so strongly tied to z_{ct} at high cloud fractions, the effects of LWP$_{cloud}$ could be masked by the combination of high and low clouds in the high cloud fraction class. Thus, it could be necessary to further categorize the data by cloud top height or other macrophysical characteristics before drawing any conclusions about the effects of microphysical cloud properties.

4 Discussion and Conclusions

Bringing together a comprehensive characterization of trade-wind clouds based on their macrophysical and microphysical properties and interactions is a challenge, one reason being the lack of representative observational data sets. The recent EUREC^{4}A campaign provides a solid database to resolve this issue. In this study, we assess how different cloud properties (cloud fraction, mean liquid water path including cloudy and cloud-free pixels in the scene (LWP$_{scene}$), mean liquid water path of only the cloudy pixels in the scene (LWP$_{cloud}$), droplet effective radius (r_{eff}), cloud top height (z_{ct}) and the organization of the cloud field (I_{org}) – affect the cloud radiative forcing of shallow trade-wind clouds and to what degree these parameters matter relative to each other. Using irradiance observations from the BACARDI broadband radiometer onboard HALO, we calculate the solar and terrestrial cloud radiative effects, here represented by α_{ce} and ΔF_{ter}. The cloud properties are obtained via one-minute collocated cloud products from GOES-16.
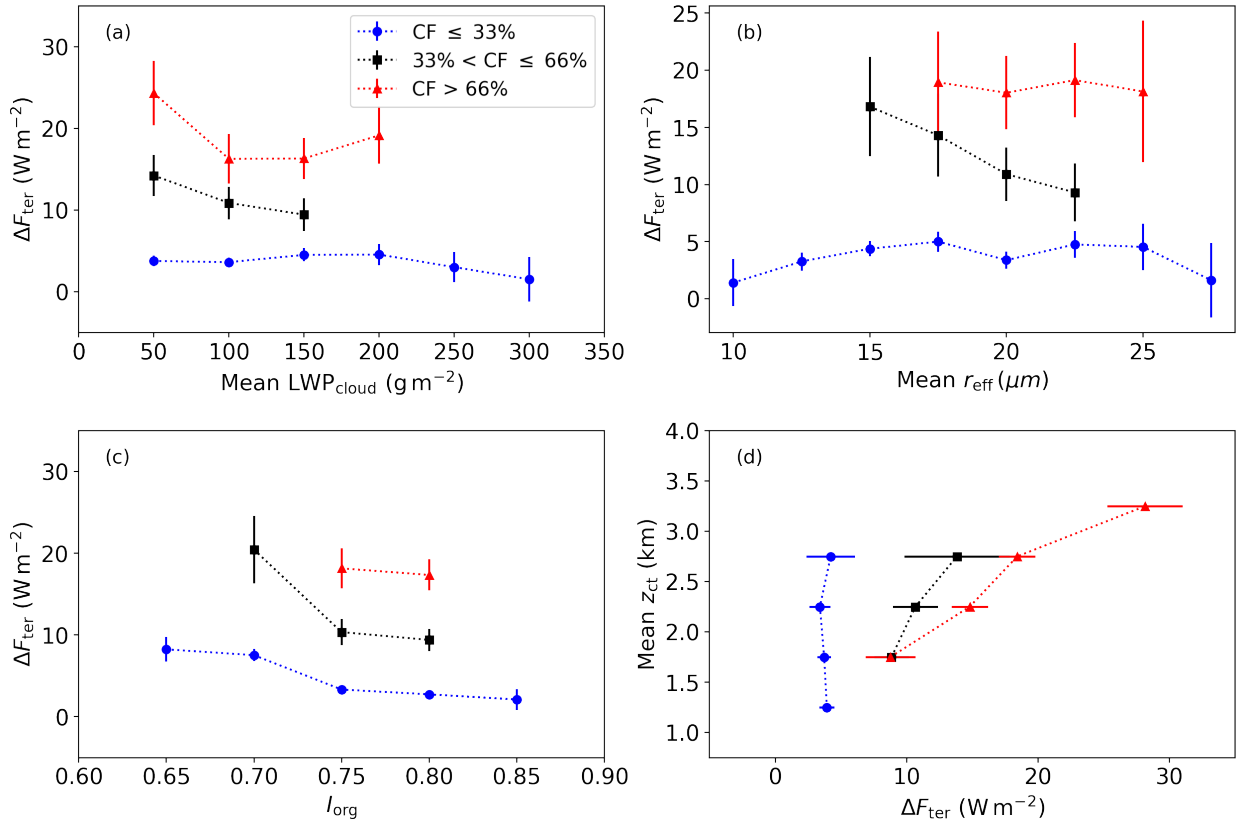
Figure 8. Same as Fig. 7 but for ΔF_{ter}.

The relationships between cloud radiative forcing and different cloud properties are complex and a challenge to disentangle. However, we have demonstrated that it is possible to use a combination of airborne irradiance measurements and satellite-based cloud property observations for this purpose. For both the solar and terrestrial component, we find that the macrophysical properties, cloud fraction and $\mathrm{LWP_{scene}}$, are the main drivers of changes in the radiative forcing of trade-wind clouds. For the solar component, in cases where the cloud fraction is below 33 %, we observe that cloud microphysical properties, z_{ct} and the organization of the cloud field are of minor importance for determining the cloud radiative forcing, demonstrating weak or no impact on the α_{ce}. As the cloud fraction increases, the relevance of those additional quantities also increases. For the terrestrial component, the impact of z_{ct} and ΔF_{ter} strengthens with increasing cloud fraction. ~~I_{org}~~ I_{org} also demonstrates a consistent negative impact on ΔF_{ter} at any cloud fraction, while the impacts of $\mathrm{LWP_{cloud}}$ and ~~r_{eff}~~ r_{eff} are more difficult to interpret.

In comparison to each other, the solar component is more strongly positively correlated to $\mathrm{LWP_{cloud}}$ and ~~r_{eff}~~ r_{eff}. The correlations to z_{ct} and ~~I_{org} are also~~ I_{org} are weakly positive for α_{ce}. This is in contrast to ΔF_{ter}, where the positive correlation to z_{ct} is much stronger, particularly as cloud fraction increases, and the correlation to ~~I_{org}~~ I_{org} is also strong, but negative. This last point indicates that as the cloud field becomes more clustered, the terrestrial cloud radiative effects decrease, while the α_{ce}

increases slightly. With these points in mind, it should be possible to discern a general idea of the cloud radiative forcing by looking only at the properties of a cloud field based on satellite measurements.

395   In those instances where certain relations were unclear, it could be possible that the addition of other information or ways of categorizing the data could be the key to extracting ~~more clear results~~even more clear conclusions. Here, the results should be constrained to specific cloud types observed during EUREC[4]A, which are largely a mixture of shallow cumulus and low stratiform clouds. This could be further extended into categorization based on the description of four cloud types given in Schulz et al. (2021). Also, using $\cancel{I_{org}}$ I_{org} as the parameter for defining the organization of the cloud field may do so correctly

400 in a macrophysical sense, but it fails to capture other important qualities about the clouds, such as the optical thickness, which other methods for describing cloud morphology capture more effectively (McCoy et al., 2017). Furthermore, the complexities of better understanding the relation of $\cancel{r_{eff}}$ r_{eff} to cloud radiative forcing may be found by separating precipitating versus non-precipitating clouds (Lohmann et al., 2000).

     Another point to consider is the fact that the cloud fraction ~~,~~or organization of the cloud field ~~, etc.~~ may not be accurately

405 depicted due to the coarse resolution of GOES-16. Coarse resolution leads to missing small clouds, cloud fractions that are overestimated and reflectances that are falsely attributed to cloudy and cloud-free pixels (Koren et al., 2008). However, because the cloud radiative forcing ~~in this study~~ is calculated from airborne observations in this study, we do not expect the irradiance observations to be falsely representative of cloudy sky conditions. Also, by adjusting the BACARDI footprint used for this study to be compatible with the reflectance observations of GOES, we are ensuring a fair comparison. Nevertheless, the

410 question remains concerning whether higher resolution observations of cloud properties are necessary for determining what cloud properties drive the cloud radiative forcing. To answer this, additional work is planned for a subsequent study including imaging remote sensing at a high spatial resolution (below 10 m), such as the VELOX thermal IR camera that was also present ~~onboard~~ on board HALO during EUREC[4]A.

Appendix A: BACARDI Footprint Quality Check

420 As mentioned in Sect. 2.4, the irradiance measured by BACARDI may not necessarily be compatible with the radiation measured by GOES-16 within the full BACARDI footprint size. Therefore, as a quality check, we compared the calculated α_{ce} from BACARDI with the mean reflectance from the 0.65 μm channel for footprints of varying sizes cut out from GOES-16. Because we are comparing different quantities, they will not match exactly, but we assume that the footprint size at which they show the best correlation represents a footprint size where GOES-16 mean footprint reflectance is comparable with BACARDI
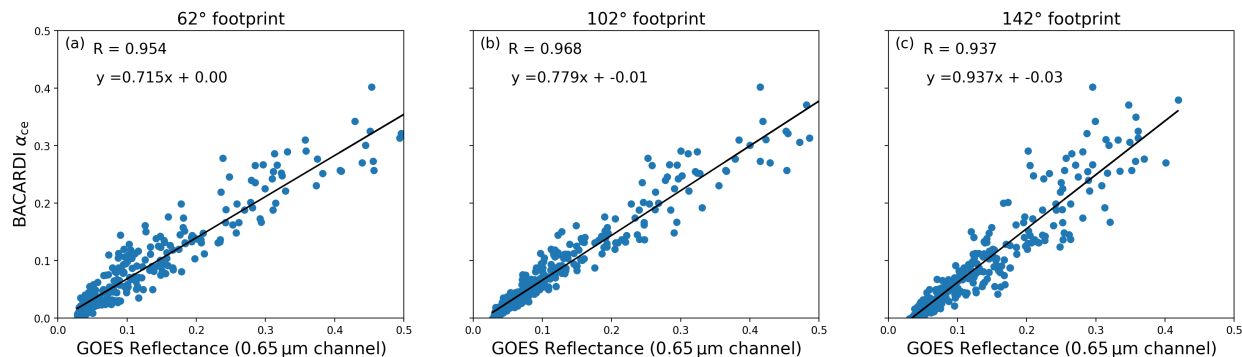
Figure A1. Comparison of α_{ce} from BACARDI and the mean reflectance from the 0.65 μm channel of GOES-16 for a FOV of (a) 62°, (b) 102° and (c) 142° from the flight on 24 January 2020. The data used here corresponds to the location of the circle flight pattern given in Sect. 2.1, at a typical flight altitude of approximately 10 km. The exact measurement times can be found in Table 2.

irradiance. A FOV of 102° was found to provide the best correlation (R = 0.968). An example of this comparison from the flight on 24 January 2020 for a FOV of 102° is shown in Fig. A1(b). The same comparison for a FOV of 62° and 142° are also shown in Fig. A1(a) and (c), respectively, to demonstrate that 102° is the best solution. Based on the selected FOV and a typical flight altitude of 10 km, the most appropriate footprint size has a radius of approximately 13 km.

References

Ardanuy, P. E., Stowe, L. L., Gruber, A., Weiss, M., and Long, C. S.: Longwave Cloud Radiative Forcing as Determined from Nimbus-7 Observations, J. Climate, 2, 766–799, https://doi.org/10.1175/1520-0442(1989)002<3C0766:LCRFAD>3E2.0.CO;2, 1988.

Bender, F. A.-M., Charlson, R. J., Ekman, A. M. L., and Leahy, L. V.: Quantification of Monthly Mean Regional-Scale Albedo of Marine Stratiform Clouds in Satellite Observations and GCMs, J. Appl. Meteorol. Clim., 50, 2139 – 2148, https://doi.org/10.1175/JAMC-D-11-049.1, 2011.

Bender, F. A.-M., Engström, A., and Karlsson, J.: Factors Controlling Cloud Albedo in Marine Subtropical Stratocumulus Regions in Climate Models and Satellite Observations, J. Climate, 29, 3559 – 3587, https://doi.org/10.1175/JCLI-D-15-0095.1, 2016.

Bony, S., Stevens, B., Ament, F., Bigorre, S., Chazette, P., Crewell, S., Delanoë, J., Emanuel, K., Farrell, D., Flamant, C., Gross, S., Hirsch, L., Karstensen, J., Mayer, B., Nuijens, L., Ruppert, J. H., Sandu, I., Siebesma, P., Speich, S., Szczap, F., Totems, J., Vogel, R., Wendisch, M., and Wirth, M.: EUREC4A: A Field Campaign to Elucidate the Couplings Between Clouds, Convection and Circulation, Surv. Geophys., 38, 1529–1568, https://doi.org/10.1007/s10712-017-9428-0, 2017.

Bring, J.: How to Standardize Regression Coefficients, Am. Stat., 48(3), 209–213, https://doi.org/10.2307/2684719, 1994.

Chen, T., Rossow, W. B., and Zhang, Y.: Radiative Effects of Cloud-Type Variations, J. Climate, 13, 264–286, https://doi.org/10.1175/1520-0442(2000)013<3C0264:REOCTV>3E2.0.CO;2, 2000.

Diamond, M. S., Director, H. M., Eastman, R., Possner, A., and Wood, R.: Substantial Cloud Brightening From Shipping in Subtropical Low Clouds, AGU Advances, 1, e2019AV000 111, https://doi.org/10.1029/2019AV000111, e2019AV000111 2019AV000111, 2020.

Ehrlich, A. and Wendisch, M.: Reconstruction of high-resolution time series from slow-response broadband terrestrial irradiance measurements by deconvolution, Atmos. Meas. Tech., 8, 3671–3684, https://doi.org/10.5194/amt-83671-2015, 2015.

Ehrlich, A., Wolf, K., Luebke, A., Zoeger, M., and Giez, A.: Broadband solar and terrestrial, upward and downward irradiance measured by BACARDI on HALO during the EUREC4A Field Campaign, AERIS [data set], https://doi.org/10.25326/160, 2021.

Emde, C., Buras-Schnell, R., Kylling, A., Mayer, B., Gasteiger, J., Hamann, U., Kylling, J., Richter, B., Pause, C., Dowling, T., and Bugliaro, L.: The libRadtran software package for radiative transfer calculations (version 2.0.1), Geosci. Model Dev., 9, 1647–1672, https://doi.org/10.5194/gmd-9-1647-2016, 2016.

Engström, A., Bender, F. A.-M., Charlson, R. J., and Wood, R.: The nonlinear relationship between albedo and cloud fraction on near-global, monthly mean scale in observations and in the CMIP5 model ensemble, Geophys. Res. Lett., 42, 9571–9578, https://doi.org/10.1002/2015GL066275, 2015.

George, G.: JOANNE : Joint dropsonde Observations of the Atmosphere in tropical North atlaNtic meso-scale Environments (v2.0.0), AERIS [data set], https://doi.org/10.25326/246, 2021.

George, G., Stevens, B., Bony, S., Pincus, R., Fairall, C., Schulz, H., Kölling, T., Kalen, Q. T., Klingebiel, M., Konow, H., Lundry, A., Prange, M., and Radtke, J.: JOANNE : Joint dropsonde Observations of the Atmosphere in tropical North atlaNtic meso-scale Environments, Earth Sys. Sci. Data, 13, 5253–5272, https://doi.org/essd-13-5253-2021, 2021.

George, R. C. and Wood, R.: Subseasonal variability of low cloud radiative properties over the southeast Pacific Ocean, Atmos. Chem. Phys., 10, 4047–4063, https://doi.org/10.5194/acp-10-4047-2010, 2010.

Heidinger, A.: Algorithm Theoretical Basis Document: ABI Cloud Height, Version 3.0, NOAA NESDIS Center for Satellite Applications and Research, College Park, MD, USA, https://www.star.nesdis.noaa.gov/goesr/documents/ATBDs/Baseline/ATBD_GOES-R_Cloud%20Height_v3.0_July%202012.pdf, 2012.

475   Heidinger, A. and Straka, W. C.: Algorithm Theoretical Basis Document: ABI Cloud Mask, Version 3.0, NOAA NESDIS Center for Satellite Applications and Research, College Park, MD, USA, https://www.star.nesdis.noaa.gov/goesr/documents/ATBDs/Baseline/ATBD_GOES-R_Cloud_Mask_v3.0_July%202012.pdf, 2012.

Herwehe, J. A., Alapaty, K., Spero, T. L., and Nolte, C. G.: Increasing the credibility of regional climate simulations by introducing subgrid-scale cloud-radiation interactions, J. Geophys. Res.-Atmos., 119, 5317–5330, https://doi.org/10.1002/2014JD021504, 2014.

480   Jacob, M., Ament, F., Gutleben, M., Konow, H., Mech, M., Wirth, M., and Crewell, S.: Investigating the liquid water path over the tropical Atlantic with synergistic airborne measurements, Atmos. Meas. Tech., 12, 3237–3254, https://doi.org/10.5194/amt-12-3237-2019, 2019.

Jiménez, P. A.: Assessment of the GOES-16 Clear Sky Mask Product over the Contiguous USA Using CALIPSO Retrievals, Remote Sens.-Basel, 12, https://doi.org/10.3390/rs12101630, 2020.

Khain, A. P., Beheng, K. D., Heymsfield, A., Korolev, A., Krichak, S. O., Levin, Z., Pinsky, M., Phillips, V., Prabhakaran, T., Teller, A.,

485   van den Heever, S. C., and Yano, J.-I.: Representation of microphysical processes in cloud-resolving models: Spectral (bin) microphysics versus bulk parameterization, Rev. Geophys., 53, 247–322, https://doi.org/10.1002/2014RG000468, 2015.

Kipp and Zonen B.V: Instruction Manual CG4 Pyrgeometer, Delft, Netherlands, https://www.kippzonen.com/Download/38/Manual-CGR4-Pyrgeometer, 2014.

Kipp and Zonen B.V: Instruction Manual CMP Series Pyranometer, Delft, Netherlands, https://www.kippzonen.com/Download/72/

490   Manual-Pyranometers-CMP-series-English, 2016.

Klein, S. A., Hall, A., Norris, J. R., and Pincus, R.: Low-Cloud Feedbacks from Cloud-Controlling Factors: A Review, Surv. Geophys., 38, 1307–1329, https://doi.org/10.1007/s10712-017-9433-3, 2017.

Konow, H., Ewald, F., George, G., Jacob, M., Klingebiel, M., Kölling, T., Luebke, A. E., Mieslinger, T., Pörtge, V., Radtke, J., Schäfer, M., Schulz, H., Vogel, R., Wirth, M., Bony, S., Crewell, S., Ehrlich, A., Forster, L., Giez, A., Groß, S., Gutleben, M., Hagen, M., Hirsch, L.,

495   Jansen, F., Lang, T., Mayer, B., Mech, M., Prange, M., Schnitt, S., Vial, J., Walbröl, A., Wendisch, M., Wolf, K., Zinner, T., Zöger, M., Ament, F., and Stevens, B.: EUREC[4]A's HALO, Earth Sys. Sci. Data, 13, 5545–5563, https://doi.org/10.5194/essd-13-5545-2021, 2021.

Koren, I., Oreopoulos, L., Feingold, G., Remer, L. A., and Altaratz, O.: How small is a small cloud?, Atmos. Chem. and Phys., 8, 3855–3864, https://doi.org/10.5194/acp-8-3855-2008, 2008.

Lohmann, U., Tselioudis, G., and Tyler, C.: Why is the cloud albedo — Particle size relationship different in optically thick and optically

500   thin clouds?, Geophys. Res. Lett., 27, 1099–1102, https://doi.org/10.1029/1999GL011098, 2000.

McCoy, I. L., Wood, R., and Fletcher, J. K.: Identifying Meteorological Controls on Open and Closed Mesoscale Cellular Convection Associated with Marine Cold Air Outbreaks, J. Geophys. Res.-Atmos., 122, 11,678–11,702, https://doi.org/10.1002/2017JD027031, 2017.

Mieslinger, T., Horváth, A., Buehler, S. A., and Sakradzija, M.: The Dependence of Shallow Cumulus Macrophysical Properties on Large-Scale Meteorology as Observed in ASTER Imagery, J. Geophys. Res.-Atmos., 124, 11 477–11 505,

505   https://doi.org/10.1029/2019JD030768, 2019.

Neter, J., Wasserman, W., and Kutner, M.: Applied Linear Regression Models, Richard D. Irwin, Inc., Homewood, IL, USA, 1st edn., 1983.

Nuijens, L., Serikov, I., Hirsch, L., Lonitz, K., and Stevens, B.: The distribution and variability of low-level cloud in the North Atlantic trades, Q. J. Roy. Meteor. Soc., 140, 2364–2374, https://doi.org/10.1002/qj.2307, 2014.

Painemal, D., Spangenberg, D., Smith Jr., W. L., Minnis, P., Cairns, B., Moore, R. H., Crosbie, E., Robinson, C., Thornhill, K. L., Winstead,

510   E. L., and Ziemba, L.: Evaluation of satellite retrievals of liquid clouds from the GOES-13 imager and MODIS over the midlatitude North Atlantic during the NAAMES campaign, Atmos. Meas. Tech., 14, 6633–6646, https://doi.org/10.5194/amt-14-6633-2021, 2021.

Peng, Y., Lohmann, U., Leaitch, R., Banic, C., and Couture, M.: The cloud albedo-cloud droplet effective radius relationship for clean and polluted clouds from RACE and FIRE.ACE, J. Geophys. Res.-Atmos., 107, AAC 1–1–AAC 1–6, https://doi.org/10.1029/2000JD000281, 2002.

515 Schmit, T., Lindstrom, S., Gerth, J., and Gunshor, M.: Applications of the 16 Spectral Bands on the Advanced Baseline Imager (ABI), J. Oper. Meteor., 6, 33–46, https://doi.org/10.15191/nwajom.2018.0604, 2018.

Schmit, T. J., Griffith, P., Gunshor, M. M., Daniels, J. M., Goodman, S. J., and Lebair, W. J.: A Closer Look at the ABI on the GOES-R Series, B. Am. Meteorol. Soc., 98, 681 – 698, https://doi.org/10.1175/BAMS-D-15-00230.1, 2017.

Schnitt, S., Orlandi, E., Mech, M., Ehrlich, A., and Crewell, S.: Characterization of Water Vapor and Clouds During the
520 Next-Generation Aircraft Remote Sensing for Validation (NARVAL) South Studies, IEEE J. Sel. Top. Appl., 10, 3114–3124, https://doi.org/10.1109/JSTARS.2017.2687943, 2017.

Schulz, H., Eastman, R., and Stevens, B.: Characterization and Evolution of Organized Shallow Convection in the Downstream North Atlantic Trades, J. Geophys. Res.-Atmos., 126, e2021JD034 575, https://doi.org/10.1029/2021JD034575, e2021JD034575 2021JD034575, 2021.

Schäfer, M., Wolf, K., Ehrlich, A., Hallbauer, C., Jäkel, E., Jansen, F., Luebke, A. E., Müller, J., Thoböll, J. C., Röschenthaler, T., Stevens,
525 B., and Wendisch, M.: A new thermal infrared imager for airborne remote sensing of cloud and surface properties, Atmos. Meas. Tech. Discuss. [preprint], in review, https://doi.org/10.5194/amt-2021-341, 2021.

Siebert, H., Beals, M., Bethke, J., Bierwirth, E., Conrath, T., Dieckmann, K., Ditas, F., Ehrlich, A., Farrell, D., Hartmann, S., Izaguirre, M. A., Katzwinkel, J., Nuijens, L., Roberts, G., Schäfer, M., Shaw, R. A., Schmeissner, T., Serikov, I., Stevens, B., Stratmann, F., Wehner, B., Wendisch, M., Werner, F., and Wex, H.: The fine-scale structure of the trade wind cumuli over Barbados – an introduction to the
530 CARRIBA project, Atmos. Chem. Phys., 13, 10 061–10 077, https://doi.org/10.5194/acp-13-10061-2013, 2013.

Stapf, J., Ehrlich, A., and Wendisch, M.: Influence of Thermodynamic State Changes on Surface Cloud Radiative Forcing in the Arctic: A Comparison of Two Approaches Using Data From AFLUX and SHEBA, J. Geophys. Res.-Atmos., 126, e2020JD033 589, https://doi.org/10.1029/2020JD033589, 2021.

Stevens, B., Ament, F., Bony, S., Crewell, S., Ewald, F., Gross, S., Hansen, A., Hirsch, L., Jacob, M., Kölling, T., Konow, H., Mayer, B.,
535 Wendisch, M., Wirth, M., Wolf, K., Bakan, S., Bauer-Pfundstein, M., Brueck, M., Delanoë, J., Ehrlich, A., Farrell, D., Forde, M., Gödde, F., Grob, H., Hagen, M., Jäkel, E., Jansen, F., Klepp, C., Klingebiel, M., Mech, M., Peters, G., Rapp, M., Wing, A. A., and Zinner, T.: A High-Altitude Long-Range Aircraft Configured as a Cloud Observatory: The NARVAL Expeditions, B. Am. Meteorol. Soc., 100, 1061 – 1077, https://doi.org/10.1175/BAMS-D-18-0198.1, 2019.

Stevens, B., Bony, S., Farrell, D., Ament, F., Blyth, A., Fairall, C., Karstensen, J., Quinn, P. K., Speich, S., Acquistapace, C., Aemisegger, F.,
540 Albright, A. L., Bellenger, H., Bodenschatz, E., Caesar, K.-A., Chewitt-Lucas, R., de Boer, G., Delanoë, J., Denby, L., Ewald, F., Fildier, B., Forde, M., George, G., Gross, S., Hagen, M., Hausold, A., Heywood, K. J., Hirsch, L., Jacob, M., Jansen, F., Kinne, S., Klocke, D., Kölling, T., Konow, H., Lothon, M., Mohr, W., Naumann, A. K., Nuijens, L., Olivier, L., Pincus, R., Pöhlker, M., Reverdin, G., Roberts, G., Schnitt, S., Schulz, H., Siebesma, A. P., Stephan, C. C., Sullivan, P., Touzé-Peiffer, L., Vial, J., Vogel, R., Zuidema, P., Alexander, N., Alves, L., Arixi, S., Asmath, H., Bagheri, G., Baier, K., Bailey, A., Baranowski, D., Baron, A., Barrau, S., Barrett, P. A., Batier,
545 F., Behrendt, A., Bendinger, A., Beucher, F., Bigorre, S., Blades, E., Blossey, P., Bock, O., Böing, S., Bosser, P., Bourras, D., Bouruet-Aubertot, P., Bower, K., Branellec, P., Branger, H., Brennek, M., Brewer, A., Brilouet, P.-E., Brügmann, B., Buehler, S. A., Burke, E., Burton, R., Calmer, R., Canonici, J.-C., Carton, X., Cato Jr., G., Charles, J. A., Chazette, P., Chen, Y., Chilinski, M. T., Choularton, T., Chuang, P., Clarke, S., Coe, H., Cornet, C., Coutris, P., Couvreux, F., Crewell, S., Cronin, T., Cui, Z., Cuypers, Y., Daley, A., Damerell, G. M., Dauhut, T., Deneke, H., Desbios, J.-P., Dörner, S., Donner, S., Douet, V., Drushka, K., Dütsch, M., Ehrlich, A., Emanuel, K.,

550   Emmanouilidis, A., Etienne, J.-C., Etienne-Leblanc, S., Faure, G., Feingold, G., Ferrero, L., Fix, A., Flamant, C., Flatau, P. J., Foltz, G. R., Forster, L., Furtuna, I., Gadian, A., Galewsky, J., Gallagher, M., Gallimore, P., Gaston, C., Gentemann, C., Geyskens, N., Giez, A., Gollop, J., Gouirand, I., Gourbeyre, C., de Graaf, D., de Groot, G. E., Grosz, R., Güttler, J., Gutleben, M., Hall, K., Harris, G., Helfer, K. C., Henze, D., Herbert, C., Holanda, B., Ibanez-Landeta, A., Intrieri, J., Iyer, S., Julien, F., Kalesse, H., Kazil, J., Kellman, A., Kidane, A. T., Kirchner, U., Klingebiel, M., Körner, M., Kremper, L. A., Kretzschmar, J., Krüger, O., Kumala, W., Kurz, A., L'Hégaret, P., Labaste,

555   M., Lachlan-Cope, T., Laing, A., Landschützer, P., Lang, T., Lange, D., Lange, I., Laplace, C., Lavik, G., Laxenaire, R., Le Bihan, C., Leandro, M., Lefevre, N., Lena, M., Lenschow, D., Li, Q., Lloyd, G., Los, S., Losi, N., Lovell, O., Luneau, C., Makuch, P., Malinowski, S., Manta, G., Marinou, E., Marsden, N., Masson, S., Maury, N., Mayer, B., Mayers-Als, M., Mazel, C., McGeary, W., McWilliams, J. C., Mech, M., Mehlmann, M., Meroni, A. N., Mieslinger, T., Minikin, A., Minnett, P., Möller, G., Morfa Avalos, Y., Muller, C., Musat, I., Napoli, A., Neuberger, A., Noisel, C., Noone, D., Nordsiek, F., Nowak, J. L., Oswald, L., Parker, D. J., Peck, C., Person, R., Philippi, M.,

560   Plueddemann, A., Pöhlker, C., Pörtge, V., Pöschl, U., Pologne, L., Posyniak, M., Prange, M., Quiñones Meléndez, E., Radtke, J., Ramage, K., Reimann, J., Renault, L., Reus, K., Reyes, A., Ribbe, J., Ringel, M., Ritschel, M., Rocha, C. B., Rochetin, N., Röttenbacher, J., Rollo, C., Royer, H., Sadoulet, P., Saffin, L., Sandiford, S., Sandu, I., Schäfer, M., Schemann, V., Schirmacher, I., Schlenczek, O., Schmidt, J., Schröder, M., Schwarzenboeck, A., Sealy, A., Senff, C. J., Serikov, I., Shohan, S., Siddle, E., Smirnov, A., Späth, F., Spooner, B., Stolla, M. K., Szkółka, W., de Szoeke, S. P., Tarot, S., Tetoni, E., Thompson, E., Thomson, J., Tomassini, L., Totems, J., Ubele, A. A., Villiger,

565   L., von Arx, J., Wagner, T., Walther, A., Webber, B., Wendisch, M., Whitehall, S., Wiltshire, A., Wing, A. A., Wirth, M., Wiskandt, J., Wolf, K., Worbes, L., Wright, E., Wulfmeyer, V., Young, S., Zhang, C., Zhang, D., Ziemen, F., Zinner, T., and Zöger, M.: EUREC[4]A, Earth Syst. Sci. Data, 13, 4067–4119, https://doi.org/10.5194/essd-13-4067-2021, 2021.

Tobin, I., Bony, S., and Roca, R.: Observational Evidence for Relationships between the Degree of Aggregations of Deep Convection, Water Vapor, Surface Fluxes, and Radiation, J. Climate, 25, 6885–6904, https://doi.org/10.1175/JCLI-D-11-00258.1, 2012.

570   Tompkins, A. M. and Semie, A. G.: Organization of tropical convection in low vertical wind shears: Role of updraft entrainment, J. Adv. Model Earth Sy., 9, 1046–1068, https://doi.org/10.1002/2016MS000802, 2017.

Twomey, S.: The Influence of Pollution on the Shortwave Albedo of Clouds, J. Atmos. Sci., 34, 1149–1152, https://doi.org/10.1175/1520-0469(1977)034<1149:TIOPOT>2.0.CO;2, 1977.

Vial, J., Vogel, R., Bony, S., Stevens, B., Winker, D. M., Cai, X., Hohenegger, C., Naumann, A. K., and Brogniez, H.: A New Look at the
575   Daily Cycle of Trade Wind Cumuli, J. Adv. Model. Earth Sy., 11, 3148–3166, https://doi.org/10.1029/2019MS001746, 2019.

Walther, A., Straka, W., and Heidinger, A. K.: Algorithm Theoretical Basis Document: Daytime Cloud Optical and Microphysical Properties (DCOMP), Version 3.0, NOAA NESDIS Center for Satellite Applications and Research, College Park, MD, USA, https://www.star.nesdis. noaa.gov/goesr/documents/ATBDs/Baseline/ATBD_GOES-R_Cloud_DCOMP_v3.0_Jun2013.pdf, 2013.

Weger, R. C., Lee, J., Zhu, T., and Welch, R. M.: Clustering, randomness and regularity in cloud fields: 1. Theoretical considerations, J.
580   Geophys. Res.-Atmos., 97, 20 519–20 536, https://doi.org/10.1029/92JD02038, 1992.

Werner, F., Ditas, F., Siebert, H., Simmel, M., Wehner, B., Pilewskie, P., Schmeissner, T., Shaw, R. A., Hartmann, S., Wex, H., Roberts, G. C., and Wendisch, M.: Twomey effect observed from collocated microphysical and remote sensing measurements over shallow cumulus, J. Geophys. Res.-Atmos., 119, 1534–1545, https://doi.org/10.1002/2013JD020131, 2014.

Wood, R. and Hartmann, D. L.: Spatial Variability of Liquid Water Path in Marine Low Cloud: The Importance of Mesoscale Cellular
585   Convection, J. Climate, 19, 1748 – 1764, https://doi.org/10.1175/JCLI3702.1, 2006.

Zhang, F., Liang, X.-Z., Zeng, Q., Gu, Y., and Su, S.: Cloud-Aerosol-Radiation (CAR) ensemble modeling system: Overall accuracy and efficiency, Adv. Atmos. Sci., 30, 955–973, https://doi.org/10.1007/s00376-012-2171-z, 2013.

Zhao, G. and Di Girolamo, L.: Statistics on the macrophysical properties of trade wind cumuli over the tropical western Atlantic, J. Geophys. Res.-Atmos., 112, https://doi.org/10.1029/2006JD007371, 2007.