

1 Potential environmental impact of bromoform from
2 Asparagopsis farming in Australia

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20 **Abstract**

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22 To mitigate the rumen enteric methane (CH₄) produced by ruminant livestock, *Asparagopsis*
23 *taxiformis* is proposed as an additive to ruminant feed. During the cultivation of *Asparagopsis*
24 *taxiformis* in the sea or in terrestrial based systems, this macroalgae, like most seaweeds and
25 phytoplankton, produces a large amount of bromoform (CHBr₃), which contributes to ozone
26 depletion once released into the atmosphere. In this study, we focus on the impact of CHBr₃ on
27 the stratospheric ozone layer resulting from potential emissions from proposed *Asparagopsis*
28 cultivation in Australia. The impact is assessed by weighting the emissions of CHBr₃ with its ozone
29 depletion potential (ODP), which is traditionally defined for long-lived halocarbons but has been
30 also applied to very short-lived substances (VSLs). An annual yield of $\sim 3.5 \times 10^4$ Mg dry weight
31 is required to meet the needs of 50% of the beef feedlot and dairy cattle in Australia. Our study
32 shows that the intensity and impact of CHBr₃ emissions varies, depending on location and
33 cultivation scenarios. Of the proposed locations, tropical farms near the Darwin region are
34 associated with largest CHBr₃ ODP values. However, farming of *Asparagopsis* using either ocean
35 or terrestrial cultivation systems at any of the proposed locations does not have potential to
36 significantly impact the ozone layer. Even if all *Asparagopsis* farming was performed in Darwin,
37 the CHBr₃ emitted into the atmosphere would amount to less than 0.02% of the global ODP-
38 weighted emissions. The impact of remaining farming scenarios is also relatively small even if
39 the intended annual yield in Darwin is scaled by a factor 30 to meet the global requirements, which
40 will increase the global ODP-weighted emissions up to $\sim 0.5\%$

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1. Introduction

Livestock is responsible for about 15% total anthropogenic greenhouse gas (GHG) emissions weighted by radiative forcing (Gerber et al., 2013), ranking it amongst the main contributors to climate change. The global demand for red meat and dairy is expected to increase >50% by 2050 compared to 2010 level, thus mitigation measures to reduce the GHG emissions from the global livestock industry are in high demand (Beauchemin et al., 2020). Total GHG emissions (e.g., CH₄) from ruminant livestock contribute about 18% of the total global carbon dioxide equivalent (CO₂-eq) inventory (Herrero and Thornton, 2013). With a global warming potential ~30 times higher than carbon dioxide (CO₂) and a much shorter lifetime (~10 years, IPCC, 2021), ruminant enteric CH₄ is an attractive and feasible target for global warming mitigation.

Enteric CH₄ from ruminant livestock is produced and released into the atmosphere through rumen microbial methanogenesis (Morgavi et al., 2010). Methanogenic archaea (methanogens) intercept substrate CO₂ and H₂ liberated during bacterial fermentation of feed materials (Kamra, 2005), and during this inefficient digestion process (Herrero and Thornton, 2013; Patra, 2012), methanogen metabolism leads to reductive CH₄ production and loss of feed energy as CH₄ emissions. To abate enteric methanogenesis, different strategies such as feeding management and antimethanogenic feed ingredients, have been proposed and assessed (e.g., Moate et al., 2016; Mayberry et al., 2019; Beauchemin et al., 2020). Some types of macroalgae have been demonstrated to mitigate production of CH₄ during *in vitro* and *in vivo* rumen fermentation significantly (Machado et al., 2014; Kinley and Fredeen 2015; Li et al., 2018; Kinley et al., 2020; Abbott et al., 2020). Among the different macroalgae species, Kinley et al. (2016a) concluded that the red algae *Asparagopsis* spp. showed the most potential for reducing CH₄ production. Kinley et al. (2016b) further demonstrated that forage with the addition of 2% *Asparagopsis taxiformis* could eliminate CH₄ production *in vitro* without negative effects on forage digestibility. In recent animal experiments, reduction of enteric CH₄ production by more than 98% was achieved with only 0.2% addition of freeze-dried and milled *Asparagopsis taxiformis* to the to the organic matter (OM) content of feedlot cattle feed (Kinley et al., 2020).

Halogenated, biologically active secondary metabolites are pivotal in the reduction of CH₄ induced by *Asparagopsis* (Abbott et al., 2020). Most of the reduction is ascribed to bromoform (CHBr₃) inhibition of the CH₄ biosynthetic pathway within methanogens (Machado et al., 2016). CHBr₃ as

74 a natural halogenated volatile organic compound originates from chemical and biological sources
75 including marine phytoplankton and macroalgae (Carpenter and Liss, 2000; Quack and Wallace,
76 2003). When emitted to the atmosphere, CHBr_3 has an atmospheric lifetime shorter than six months
77 and is often referred to as a very short-lived substance (VSLs). Once released into the atmosphere,
78 degraded halogenated VSLs can catalytically destroy ozone in the troposphere and stratosphere,
79 thus drawing them considerable interest (Engel and Rigby et al., 2018; Zhang et al., 2020).
80 Bromoform is the dominant compound among bromine-containing VSLs emissions, resulting
81 mostly from natural sources (Quack and Wallace, 2003) and to a lesser degree from anthropogenic
82 production (Maas et al., 2019; 2021). With an atmospheric lifetime of about 17 days (Carpenter
83 and Reimann et al., 2014), CHBr_3 can deliver bromine to the stratosphere under appropriate
84 conditions of emission strength and vertical transport (e.g., Aschmann et al., 2009; Liang et al.,
85 2010; Tegtmeier et al., 2015, 2020) and thus contribute to ozone depletion in the lower and middle
86 stratosphere (e.g., Yang et al., 2014; Sinnhuber and Meul, 2015). Global research on enabling
87 large-scale seaweed *Asparagopsis* farming is increasing (Black et al., 2021) as it appears to be one
88 of the most promising options as an antimethanogenic feed ingredient to achieve carbon neutrality
89 in the livestock sector within the next decade (Kinley et al., 2020; Roque et al., 2021). In
90 consequence, the environmental impact of CHBr_3 due to *Asparagopsis* farming also needs to be
91 explored and elucidated. In this study, only the impact on the stratosphere is considered.

92 The hypothesis was that large scale cultivation of *Asparagopsis* would not contribute significantly
93 to depletion of the ozone layer. The aim of this study was to assess the impact of anthropogenic
94 and natural processes that may contribute to CHBr_3 emissions inherent in large scale production
95 of *Asparagopsis spp.* and the subsequent impact of CHBr_3 release to the atmosphere by using
96 cultivation in Australia as the model. Specific objectives were to inform the industry, policy
97 makers, as well as the scientific community on: (i) the potential impact of CHBr_3 associated with
98 mass production of *Asparagopsis* on atmospheric halogen budgets and ozone depletion; (ii)
99 potential impacts relative to variability in regional climate, atmospheric conditions, and convection
100 trends with different potentials for transport of CHBr_3 to stratospheric ozone; (iii) the combined
101 CHBr_3 emissions potential of ocean and terrestrial based cultivation of *Asparagopsis* to supply
102 sufficient biomass for up to 50% of beef feedlot and dairy cattle in Australia; and (iv) extrapolation
103 of the impacts of production to requirements on a global scale.

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2. Data and Method

The potential impact of CHBr_3 on the atmospheric bromine budget and stratospheric ozone depletion, associated with *Asparagopsis spp.* mass production was assessed for assumed annual yields and particular production scenarios of macroalgae in Australia. Terrestrial systems cultivation and open ocean cultivation under different harvest conditions, variations of seaweed yield and growth rates for various scenarios and locations were tested as described in the following subsections.

2.1 Cultivation Scenarios

The cultivation scenarios in this study assume that sufficient seaweed is grown to supply *Asparagopsis spp.* to 50% of the Australian herds of beef cattle in feedlots (100%: $\sim 1.0 \times 10^6$) and dairy cows (100%: $\sim 1.5 \times 10^6$). For an effective reduction of CH_4 production from ruminants, a $\sim 0.4\%$ addition of freeze-dried and milled *Asparagopsis taxiformis* to the daily feed dry matter intake (DMI) is required (Kinley et al., 2020). This results in daily feed additions of 38 g dry weight (DW) *Asparagopsis* per head of feedlot cattle and 94 g DW *Asparagopsis* per head of dairy cows. In total, the required annual yield amounts to $\sim 3.5 \times 10^4$ Mg DW *Asparagopsis* to supplement the feed of roughly 50% of the Australian feedlot cattle and Australian dairy cows. Assuming that fresh weight (FW) has a DW content of 15%, a total of $\sim 2.3 \times 10^5$ Mg FW *Asparagopsis* needs to be harvested every year.

For a global scenario, we make the functional assumptions that: (i) there would be adoption of 30% of the global feed base to be supplemented with *Asparagopsis* farmed in Australia to reduce ruminant CH_4 production worldwide; (ii) *Asparagopsis* would be adopted by 50% of Australia's feedlot and dairy industries; and (iii) this is approximately equivalent to 1% of the global feedlot and dairy herds for the purpose of both assumed magnitude of production and adoption relevant for calculations of supply and emissions. This export scenario requires for 30 times increased production compared to the Australian scenario if all the required *Asparagopsis* was to be cultivated in Australia and an annual harvest of ~ 1 Tg DW *Asparagopsis* would be needed from Australian waters.

For the future farm distributions in Australia, we assume that *Asparagopsis* will be cultivated in open ocean systems and terrestrial confinement systems (that may include, but not limited to, tanks,

136 raceways, and ponds) located near Geraldton, Triabunna, and Yamba (Figure 1). We assume that
137 one third of the required annual yield ($\sim 1.2 \times 10^4$ Mg DW) is grown near Triabunna (T), with 60%
138 in terrestrial systems and 40% in open ocean farms, one third is grown in terrestrial systems at
139 Yamba (Y), and the last third is grown in the open ocean in Geraldton (G). For comparison of the
140 environmental impact, we also adopt a tropical scenario where all farms with their total annual
141 yield of $\sim 3.5 \times 10^4$ Mg DW are assumed to be situated near Darwin.

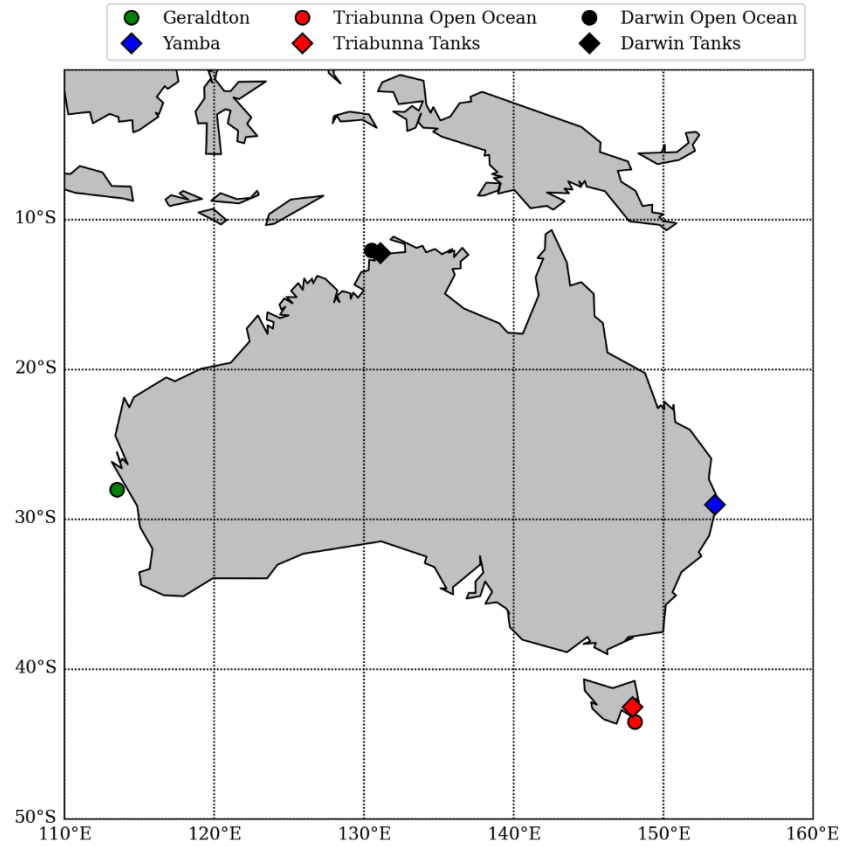
142 The emissions of CHBr_3 from the macroalgae farms can be derived based on estimates of the
143 standing stock biomass. For any given farming scenario, the standing stock biomass B_f (g DW) is
144 a function of time t and can be calculated from the initial biomass B_i (g DW) and the specific
145 growth rate GR (%/day) according to Hung et al. (2009):

$$146 \quad B_f(t) = B_i \cdot (1 + GR/100)^t \quad (1)$$

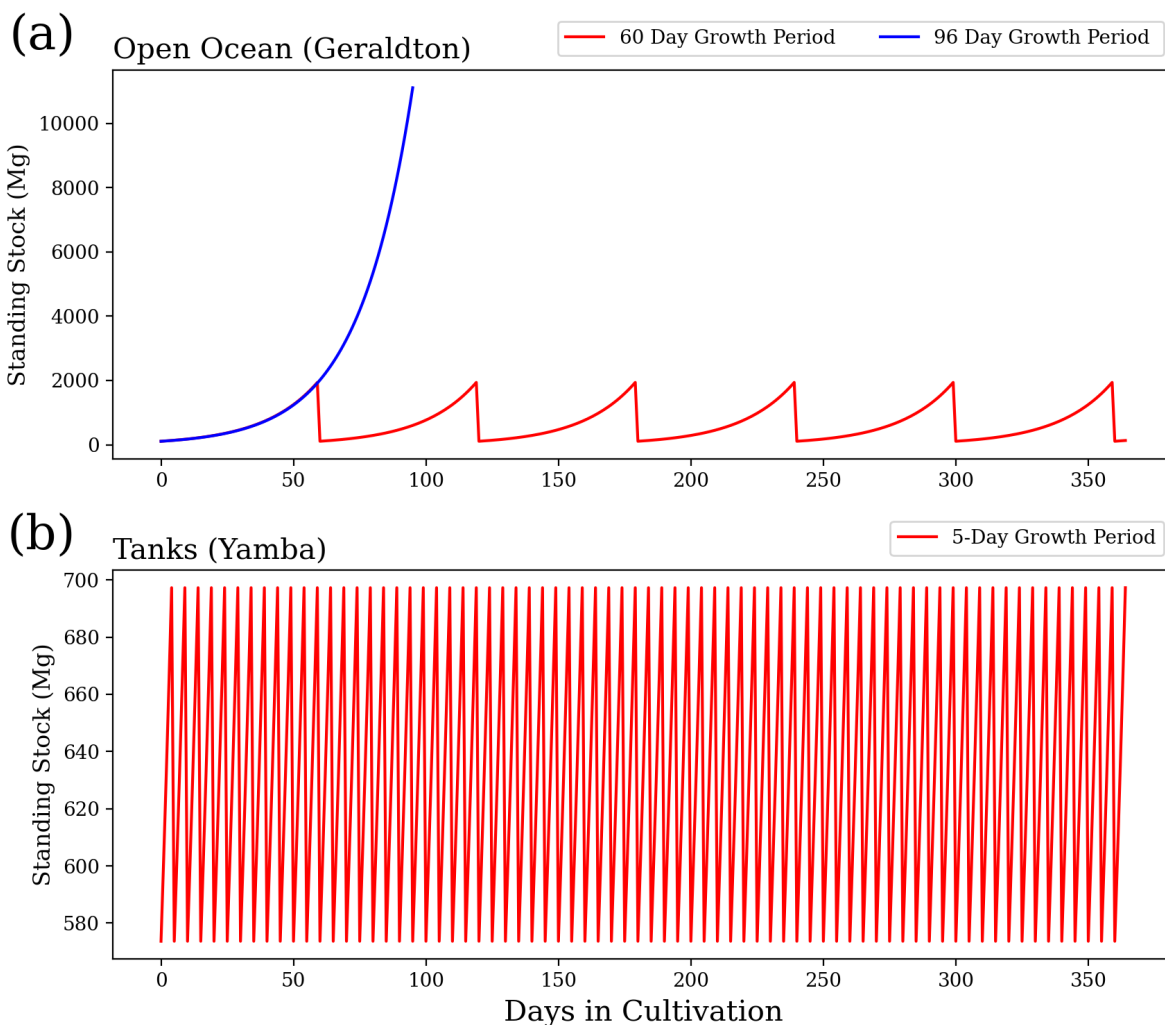
147 Terrestrial systems and open ocean cultivation scenarios are assuming a fixed targeted annual
148 yield. For a given initial biomass and growth rate, the length and frequency of the growth periods
149 per year need to be chosen accordingly, to achieve the required final yield. Yong et al. (2013)
150 checked the reliability of different equations for seaweed growth rate determination by comparing
151 the daily seaweed weight cultivated under optimized growth condition, and the most reliable
152 relationship between initial and final weight leads to the form of Eq (1). We also applied several
153 growth rates from 1 to 10% to show the possible influence of this parameter on the overall
154 emissions of the algae. Average growth rates of *Asparagopsis* ranged from 7 to 13 %/day in
155 samples from tropical and sub-tropical Australia during short-term experiments (Mata et al., 2017).
156 We used a lower growth rate of 5% for our scenario to provide an upper estimate of potential
157 CHBr_3 emissions. Note that emissions decrease by 27% when using a growth rate of 7% as
158 demonstrated in section 3.1.

159 Figure 2 provides an example of the variations of standing stock of *Asparagopsis* for the farms of
160 Geraldton (all open ocean) and Yamba (all terrestrial systems) with a growth rate of 5% per day.
161 For the open ocean cultures, we assume a scenario of six harvests per year and 60 day growth
162 periods to obtain the annual yield (Battaglia, 2020; Elsom, 2020). For a sensitivity study, we
163 assume an alternative scenario based on the same initial biomass, but only one harvest per year.
164 As evident from Figure 2, the same annual yield can be achieved with one harvest per year if

165 applying an extended growth period of 96 days. For the tank cultures, a harvest every 5 days (73
166 harvests per year) is assumed as a realistic scenario (Battaglia, 2020; Elsom, 2020).



167
168 **Figure 1.** Locations of actual *Asparagopsis* farms in Geraldton, Triabunna, Yamba, and theoretical
169 farms in Darwin.



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 171 **Figure 2.** Standing stock biomass of *Asparagopsis* cultivation a) in the open ocean for a 60-day
 172 growth period and 96 day growth period and b) in terrestrial systems culture for a 5 day growth
 173 period. Each of the three scenarios will achieve an annual yield of $\sim 1.6 \times 10^4$ Mg DW.

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2.2 *Asparagopsis* CHBr₃ release rates

177 Rates of the CHBr₃ content in *Asparagopsis* given in the literature range between 3.4 to 43 mg
 178 CHBr₃/g DW, with values around 10 mg CHBr₃/g DW appearing to be realistic in current
 179 cultivation (Burreson and Moore, 1976; Mata et al., 2012, 2017; Paul et al., 2006; Vucko et al.,
 180 2017). We assume that *Asparagopsis* strain selection cultivated for feed supplements will lead to
 181 high yielding CHBr₃ varieties thus we assume augmented CHBr₃ production with a mean content

182 of 21.7 mg CHBr₃/g DW (Magnusson et al., 2020) for this study.

183 Very few values on the CHBr₃ release from *Asparagopsis* have been reported in the literature. A
 184 constant release of 1100 ng CHBr₃/g DW hr⁻¹ was measured for *Asparagopsis armata*
 185 *tetrasporophyte*, which has a CHBr₃ content of 14.5 mg CHBr₃/g DW (Paul et al., 2006). We
 186 assume a linear scaling between the CHBr₃ release rates and the content. Thus, a cultivated
 187 *Asparagopsis* for which we assume 21.7 mg CHBr₃/g DW should release around 1646 ng CHBr₃/g
 188 DW hr⁻¹, a rate which has been confirmed by Marshall et al. (1999). Therefore, for our calculations,
 189 we assume a constant release of 1600 ng CHBr₃/g DW hr⁻¹ for farmed *Asparagopsis* with a CHBr₃
 190 content of 21.7 mg CHBr₃/g DW. These content and release rates are higher than those for wild
 191 stock algae (Leedham et al., 2013; Nightingale et al., 1995) as the farming aims at algae varieties
 192 with high CHBr₃ yield. As available information on this topic is very sparse no variations of the
 193 release rate with life-cycle stages, season, location, or other environmental parameters were used
 194 in this study. Also, the two species *Asparagopsis armata* and *Asparagopsis taxiformis* were treated
 195 the same way as *Asparagopsis spp.*, as variations in CHBr₃ content and release within or between
 196 species are currently unknown (Mata et al., 2017) and more research on this topic is needed.

197 2.3 Parameterization of CHBr₃ Emission

198 The emissions of CHBr₃ from farmed macroalgae are a function of the standing stock biomass (in
 199 g DW) and can be calculated with the constant release rate (R_{CHBr_3}) of 1600 ng CHBr₃/g DW hr⁻¹
 200 multiplied with the standing stock. The total release of CHBr₃ (E_{CHBr_3}) over the complete growth
 201 period of T days is given by the integral over the daily emissions from day 1 to day T:

$$204 \quad E_{CHBr_3} = \int_0^T 24 \cdot B_i \cdot (1 + GR)^t \cdot R_{CHBr_3} dt = 24 \cdot B_i \cdot R_{CHBr_3} \cdot \frac{[(1+GR)^T - 1]}{\ln(1+GR)} \quad (2)$$

205 For our atmospheric impact studies we assume, that all CHBr₃ released from the algae is emitted
 206 into the atmosphere at its location of production. An increasing seawater concentration of CHBr₃
 207 shifts the equilibrium conditions between seawater and air towards the atmosphere, as CHBr₃
 208 easily volatilizes to the atmosphere. Consequently, air-sea exchange acts as a relatively fast loss
 209 process for CHBr₃ in surface water. Oceanic sinks can also impact CHBr₃, but act on relatively
 210 long timescales. Degradation through halide substitution and hydrolysis results in the ocean sink
 211 CHBr₃ half-life of 4.37 years (Hense and Quack, 2009). Thus, most of the CHBr₃ contained in

212 surface seawater is instantly outgassed into the atmosphere without oceanic loss processes, playing
213 a role as confirmed by the modelling study of Maas et al. (2021).

214 The air-sea exchange of CHBr_3 is expressed as the product of its transfer coefficient (k_w) and the
215 concentration gradient (Δc) (Eq. (3)). The gradient is computed between the water concentration
216 (c_w) and theoretical equilibrium water concentration (c_{atm}/H), where c_{atm} is the atmospheric
217 concentration and H is Henry's law constant (Moore et al., 1995a; Moore et al., 1995b).

$$218 \quad F = k_w \cdot \Delta c = k_w \cdot \left(c_w - \frac{c_{\text{atm}}}{H} \right) \quad (3)$$

219 The compound-specific transfer coefficient (k_w) is determined using the air-sea gas exchange
220 parameterization of Nightingale et al. (2000) (Eq. (4))

$$221 \quad k_w = k \cdot \sqrt{Sc} / 660 \quad (4)$$

222 The transfer coefficient k is a function of the wind speed at 10 m height (u_{10}): $k = 0.2u_{10}^2 +$
223 $0.3u_{10}$, and the Schmidt number (Sc) is a function of sea surface temperature (SST) from Quack
224 and Wallace (2003), which is expressed as $Sc = 4662.8 - 319.45 \cdot SST + 9.9012 \cdot SST^2 +$
225 $0.1159 \cdot SST^3$.

226 In this study, we use the CHBr_3 sea-to-air flux climatology from Ziska et al. (2013) as marine
227 background emissions. The global emission scenario from Ziska et al. (2013) is a bottom-up
228 estimate of the oceanic CHBr_3 fluxes, generated from atmospheric and oceanic surface ship-borne
229 *in situ* measurements between 1979 to 2013. Due to the paucity of data, the 35 year mean gridded
230 data set was filled by interpolating and extrapolating the *in situ* measurement data. The oceanic
231 emissions were calculated with the transfer coefficient parameterization of Nightingale et al.
232 (2000) and 6-hourly meteorological data, which allow a temporal emission variability related to
233 wind and temperature.

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235 **2.4 Emission Scenarios for FLEXPART Simulations**

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237 To quantify the atmospheric impact of CHBr_3 emissions from macroalgae farming, the Lagrangian
238 particle dispersion model FLEXPART (Pisso et al., 2019) is used. FLEXPART has been evaluated
239 extensively in previous studies (e.g., Stohl et al., 1998; Stohl and Trickl, 1999). The model includes
240 moist convection and turbulence parameterizations in the atmospheric boundary layer and free
241 troposphere (Forster et al., 2007; Stohl and Thomson, 1999). The European Centre for Medium-
242 Range Weather Forecasts (ECMWF) reanalysis product ERA-Interim (Dee et al., 2011) with a

243 horizontal resolution of $1^\circ \times 1^\circ$ and 60 vertical model levels is used for the meteorological input
244 fields, providing air temperature, winds, boundary layer height, specific humidity, as well as
245 convective and large-scale precipitation with a 3-hour temporal resolution.

246 We conduct FLEXPART simulations for year 2018 with different emission scenarios as explained
247 in the following and summarized in Table 1:

248 1.) Australian scenarios: CHBr_3 emissions from the *Asparagopsis* farming in Geraldton,
249 Triabunna, and Yamba are calculated for an overall annual yield of 34674 Mg DW according to
250 Equation 2. For the terrestrial systems, 5 day growth periods are assumed resulting in 73 harvests
251 per year. For the open ocean, the assumption of different growth periods results in three sub-
252 scenarios a) 6 times 60 day growth periods with the first period starting on January 1st (referred to
253 as GTY_O60), b) one 96 day growth period starting on January 1st (GTY_O96_Jan), and c) and
254 another starting on July 1st (GTY_O96_Jul).

255 For the last Australian scenario, we assume that all farms are located around Darwin in the
256 Northern Territory tropics with 6 times 60 day growth periods in the open ocean and 73 times 5
257 day growth periods in the terrestrial systems (Darwin_O60). While this is an unlikely scenario
258 according to current plans, it is useful to demonstrate the influence of potential farming locations
259 on their environmental impact.

260 2.) Global scenarios: Emissions from *Asparagopsis* farming in Geraldton, Triabunna, and Yamba
261 are estimated according to the annual yield, upscaled by a factor of 30 to global requirements.
262 amounting to 1.04×10^6 Mg DW. Growth periods and harvesting frequencies are set up in the same
263 way as for the Australian scenarios. Short names of the global scenarios are the same as for the
264 Australian scenarios with the additional label 30x.

265 3) Background scenario: Emission from Ziska et al. (2013) for the entire coastal region around
266 Australia defined as all $1^\circ \times 1^\circ$ grid cells directly neighbouring the coastline (Ziska_Coast).

267 4.) Extreme event scenarios: We assume extreme conditions where a hypothetical tropical cyclone
268 causes implausible release of all CHBr_3 from the macroalgae farm and water into the atmosphere.
269 We focus on the case study of Geraldton and the tropical cyclone Joyce, which occurred from 6-
270 13 January 2018 around western Australia. We base the amount of available macroalgae biomass
271 on the Australian scenario and assume that the entire CHBr_3 content of all *Asparagopsis* at this
272 location is released at once. The two scenarios defined here assume that the tropical cyclone occurs
273 at the end of the 60 day growth period (Geraldton_Ex60) resulting in the release of 41.8 Mg CHBr_3

274 (21.7 mg CHBr₃/g DW * 1926 Mg DW) or at the end of the 96 day growth period
275 (Geraldton_Ex96) resulting in the release of 250.8 Mg CHBr₃ (21.7 mg CHBr₃ /g DW* 11558 Mg
276 DW).

277 The daily model output is recorded for all simulations. For the extreme event, which assumes the
278 destruction of a farm (Geraldton-Ex), the 3 hourly output is recorded. For all simulations, except
279 the background scenario and extreme scenario, trajectories are released from four regions of the
280 size of: a) Geraldton (open ocean, 11558 ha): 0.1°x0.1°; b) Triabunna (open ocean, 4623 ha):
281 0.06°x0.06°; c) Triabunna (terrestrial systems, 126 ha): 0.01°x0.01°; and d) Yamba (terrestrial
282 systems, 210 ha): 0.01°x0.01°. For the tropical and extreme scenarios, trajectories are released
283 from the Darwin and Geraldton farms, respectively. For the background scenario Ziska_Coast,
284 trajectories are released from the 1.0°x1.0° grid along the Australian coastline. Note that it is not
285 reasonable to compute the Ziska emission on the locations of farming as some farms are terrestrial.
286 However, if we assume all the farms are Geraldton-like (i.e., all grown in the open ocean), the
287 Ziska emission in Geraldton, Yamba and Triabunna will be 843 Mg, 295 Mg, and 676 Mg,
288 respectively. The amount of released CHBr₃ is evenly distributed among the trajectories and is
289 depleted during the Lagrangian simulations according to the atmospheric half-life of 17 days (*e*-
290 folding lifetime of 24 days) (Hossaini et al., 2010; Montzka and Reimann et al., 2010; Engel and
291 Rigby et al., 2018).

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297 **Table 1.** Detailed information on the scenarios set up for the atmospheric transport simulations
 298 with FLEXPART (Geraldton, Triubanna, and Yamba: GTY)

Name		Total Yield (Mg DW)	CHBr ₃ Emissions (Mg)	Notes	Simulation Period
Australian Scenarios	GTY_O60	Total: 34674	Total: 27.3	6 harvests (every 60 days) in the open ocean; 73 harvests (every 5 days) in the terrestrial systems.	01.01.2018 - 31.12.2018 2-month spin-up
	GTY_O96_Jan	Open Ocean: G: 11558 T: 4623 Y: -	Open Ocean: G: 9.10 T: 3.64	1 harvest (after 96 days) in open ocean; 73 harvests (every 5 days) in terrestrial systems. Growth in open ocean starts from 01.01.2018.	
	GTY_O96_Jul	Terrestrial systems: G: - T: 6935 Y: 11558	Terrestrial systems: T: 5.46 Y: 9.10	Same as GTY_O96_Jan but with growth in open ocean starting from 01.07.2018.	
	Darwin_O60	Total: 34674 Open Ocean: Darwin: 16181 Terrestrial systems: Darwin: 18493	Total: 27.3 Open Ocean: Darwin: 12.7 Terrestrial systems: Darwin: 14.6	Same as GTY_O60 but with farms near Darwin	
Global Scenarios	GTY_O60_30x	Total: 1040220	Total: 819	Same as GTY_O60 but with initial biomass and areas 30 times larger.	
	GTY_O96_Jan_30x	Open Ocean G: 346740 T: 138690 Y: -	Open Ocean G: 273 T: 109.2	Same as GTY_O96_Jan but with initial biomass and areas 30 times larger.	
	GTY_O96_Jul_30x	Terrestrial systems: G: - T: 208050 Y: 346740	Terrestrial systems: T: 163.8 Y: 273	Same as GTY_O96_Jul but with initial biomass and areas 30 times larger.	
	Darwin_O60_30x	Total: 1040220 Open Ocean Darwin: 485430	Total: 819 Open Ocean Darwin: 381 Terrestrial systems: Darwin: 438	Same as Darwin_O60 but with initial biomass and areas 30 times larger.	

		Terrestrial systems: Darwin: 554790			
Background Scenario	Ziska_Coast	-	3109	CHBr ₃ emission of the coastal region of Australia from Ziska et al. (2013)	
Extreme Scenarios	Geraldton_Ex60	Open Ocean: G: 1926	Open Ocean: G: 41.8	Extreme event: CHBr ₃ in Geraldton surface water before harvest is released due to tropical cyclone Joyce (07.01.2018 – 15.01.2018). Harvest period: 60 days.	9.01.2018 – 9.02.2018 No spin-up
	Geraldton_Ex96	Open Ocean: G: 11558	Open Ocean: G: 250.8	Same as Geraldton_Ex60 but with harvest period of 96 days	

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2.5 Ozone Depletion Potential (ODP)

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303 The ozone depletion potential (ODP) is defined as the time-integrated potential destructive effect
304 of a substance to the ozone layer relative to that of the reference substance CFC-11 (CCl₃F) on a
305 mass emitted basis (Wuebbles, 1983). The ODP is a well-established and extensively used concept
306 traditionally defined for anthropogenic long-lived halocarbons. However, the concept has been
307 also applied to VSLs (Brioude et al., 2010; Pisso et al., 2010): unlike the ODP for long-lived
308 halocarbons, which is one constant number, the ODP of a VSL is a function of time and location
309 of the emissions. This variable number still describes the time-integrated ozone depletion resulting
310 from a CHBr₃ unit mass emission relative to the ozone depletion resulting from the same unit mass
311 emission of CFC-11. The ODP for VSLs can be derived from chemistry-climate or chemistry
312 transport models simulating the changes of ozone due to certain compound (Claxton et al., 2019;
313 Zhang et al., 2020). The trajectory-derived ODP of VSLs such as CHBr₃ is calculated as a
314 function of location and time of the potential emissions (Brioude et al., 2010; Pisso et al., 2010).
315 As for the traditional ODP concept, the time and space dependent ODP describes only the potential
316 of a compound but not its actual damaging effect to the ozone layer and is independent of the total
317 emissions. It is noteworthy that many VSLs including CHBr₃ can impact ozone in the troposphere
318 and stratosphere. As ODPs are used to assess stratospheric ozone depletion only, the contribution
319 of VSLs to tropospheric ozone destruction needs to be excluded when calculating their ODP
320 (Pisso et al., 2010; Zhang et al., 2020). The trajectory based ODP from Pisso et al. (2010) used in
321 this study considers only the impact of CHBr₃ on the stratospheric ozone instead of the ozone

322 column. The fraction of originally emitted VSLs reaching the stratosphere depends strongly on
323 the meteorological conditions. In particular, it shows a pronounced seasonality. Here we apply
324 ODP values adapted from Pisso et al. (2010), originally calculated for a VSL with a lifetime of
325 20 days, which is very similar to that of CHBr_3 . ODPs for VSLs are calculated by means of
326 combining two sources of information: one corresponding to the slow stratospheric branch and the
327 other to the fast tropospheric branch of transport. The former is uniform for all species modelled
328 and is based on the calculation of the expected stratospheric residence time of a Lagrangian particle
329 entering the stratosphere. The latter is based on the probability of stratospheric injection of a given
330 unit emission of the tracer at the ground. The probability of injection depends not only on the
331 fraction of air reaching the tropopause but also on the time the air mass takes from the ground to
332 the tropopause. This is because during the transit of the air mass through the troposphere, the
333 precursor is chemically degraded, and the solubility of the products leads to mass loss due to wet
334 deposition.

335 In this study, we present the ODP-weighted emissions, which combine the information of the ODP
336 and surface emissions and are calculated by multiplying the CHBr_3 emissions with the trajectory-
337 derived ODP at each grid point. The ODP-weighted emissions provide insight into key factors of
338 CHBr_3 emission (i.e., where and when CHBr_3 is emitted) that impact stratospheric ozone
339 (Tegtmeier et al., 2015). The absolute values are subject to relatively large uncertainties arising
340 from uncertainties in the parameterization of the convective transport. Furthermore, the ODP
341 values applied here do not consider product gas entrainment and provide therefore a lower limit of
342 the impact of CHBr_3 on stratospheric ozone. Taking into account product gas entrainment can lead
343 to 30% higher ODP values (Engel and Rigby et al., 2018; Tegtmeier et al., 2020), but has no large
344 impact on the comparison between global ODP-weighted CHBr_3 emissions and farm-based ODP-
345 weighted CHBr_3 emissions presented here.

346

347 **3. CHBr_3 Emission and Atmospheric Mixing Ratio**

348

349 **3.1 CHBr_3 Emissions**

350

351 As shown in Eq. (2), the total CHBr_3 emissions are determined by the growth rate, growth period
352 and initial biomass. For our scenarios based on selected fixed growth rates, the growth periods are
353 adjusted so that the intended annual yield ($\sim 3.5 \times 10^4 \text{ Mg DW}$) is achieved. We conduct a sensitivity

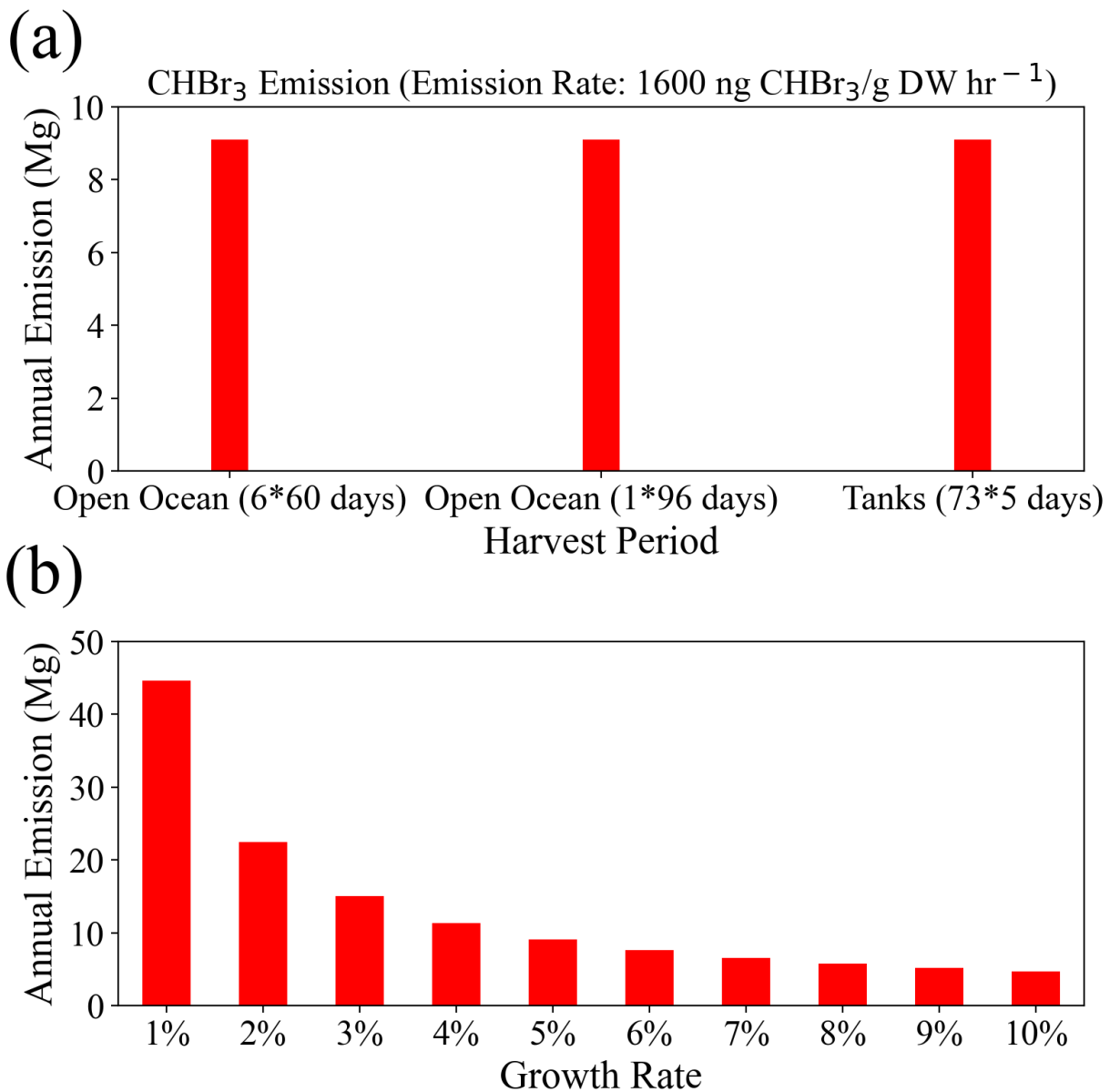
354 study to analyze how much the total emissions change for variations of the length and number of
355 the growth periods for a fixed annual yield. For this purpose, we compare Geraldton farming for
356 GTY_O60 (open ocean, six 60 day growth periods) with Geraldton farming for GTY_O96 (open
357 ocean, one 96 day growth period) and Yamba farming for GTY_O60 (terrestrial systems, 73
358 growth periods of 5 days). Our estimates show that the annual release of CHBr₃ from *Asparagopsis*
359 is the same for all three case studies (Fig. 3a), confirming that for a fixed annual yield and growth
360 rate, the culture conditions of open ocean and tank farming are not important for VSLs emissions.
361 A second sensitivity study investigates the variations of CHBr₃ emissions for different growth rates
362 and the same fixed annual yield. For this purpose, we compare Geraldton farming (open ocean,
363 with an intended annual yield of $\sim 1.1 \times 10^4$ Mg DW) for different growth rates varying between 1%
364 and 10%. The scenario with a 5% growth rate corresponds to Geraldton farming for GTY_O60
365 (open ocean, six 60 day growth periods), while for the other growth rates the growth periods have
366 been adjusted to achieve the same annual yield.

367 The CHBr₃ emissions depend strongly on the growth rates (Fig 3b), with emission calculated for
368 a 1% growth rate being almost 10 times higher than the emissions calculated for a 10% growth
369 rate. For a lower growth rate, the initial biomass needs to be higher to achieve the targeted seaweed
370 yield ($\sim 1.1 \times 10^4$ Mg) after one year and/or the growth period needs to be longer, thus resulting in
371 larger amounts of biomass in the ocean and higher annual CHBr₃ emissions. Vice versa, for higher
372 growth rates, the annual oceanic biomass is smaller and total emissions are lower.

373 The overall emissions from the intended Australian seaweed farming of $\sim 3.5 \times 10^4$ Mg DW range
374 from 13.5 Mg (0.05 Mmol) for a 10% growth rate to 134 Mg (0.5 Mmol) per year for a 1% growth
375 rate. For the growth rates higher than 5%, the differences of CHBr₃ emissions are less significant
376 than those derived for the lower growth rates. In our study, we choose 5% growth rate as
377 representative, which leads to emissions of ~ 27 Mg (0.1 Mmol) CHBr₃ per year for the targeted
378 final yield. For the global scenario with an annual yield of $\sim 1.0 \times 10^6$ Mg DW (30 times of the
379 Australian target), the emissions would range from 412 Mg (1.6 Mmol) to 4014 Mg (16 Mmol)
380 per year, with the annual emission of 810 Mg (3.2 Mmol) for 5% growth rates.

381 Interestingly, the potential local emissions for all the farming scenarios are generally 3 to 6 orders
382 of magnitude higher than the background coastal emissions. The maximum climatological
383 emissions derived from available observations (Ziska_Coast) are around 2000 pmol m⁻² hr⁻¹ for
384 the coastal waters of Australia, while the emissions from an *Asparagopsis* farm can reach more

385 than 2.0×10^6 pmol ($2 \mu\text{mol}$) $\text{m}^{-2} \text{hr}^{-1}$ from a terrestrial system and more than 5.0×10^5 pmol $\text{m}^{-2} \text{hr}^{-1}$
 386 1 from the open ocean. These differences are to a large degree related to the fact that the
 387 Ziska_Coast is given on a $1.0^\circ \times 1.0^\circ$ grid, with high coastal values averaging out over the relatively
 388 wide grid cells, while the values derived for the farms apply to much smaller areas. Tank emission
 389 rates ($0.01^\circ \times 0.01^\circ$) and open ocean farming emission rates ($0.1^\circ \times 0.1^\circ$) averaged over a $1^\circ \times 1^\circ$ grid
 390 cell result in 200 pmol $\text{CHBr}_3 \text{ m}^{-2} \text{hr}^{-1}$ and 5000 pmol $\text{CHBr}_3 \text{ m}^{-2} \text{hr}^{-1}$, respectively, thus being very
 391 similar to the Ziska emissions.



392
 393

394 **Figure 3.** The annual release of CHBr_3 (Mg yr^{-1}) from: a) same growth rate (5%) for different
395 growth periods; and b) under different growth rates but with same initial biomass, both a) and b)
396 are obtained with a total annual yield of 11558 Mg DW.

397

398

399

400

3.2 Atmospheric CHBr_3 mixing ratio

401 We use the CHBr_3 emissions calculated in section 3.1 to simulate the enhanced atmospheric CHBr_3
402 mixing ratios (above natural background) for each *Asparagopsis* farming scenario. Background
403 CHBr_3 levels are calculated based on the Ziska et al. (2013) Australian coastal emissions
404 (Ziska_Coast). The temporal evolution of CHBr_3 mixing ratio with height shows that the CHBr_3
405 resulting from the Australian farming scenarios are negligible (see Figure S1) compared to the
406 coastal background emissions of Australia (Ziska_Coast).

407 For the global scenarios (Figure 4), atmospheric CHBr_3 is comparable to CHBr_3 resulting from
408 Australian coastal background emissions, especially near the end of the growth period in the open
409 ocean. For almost all scenarios (except for GTY_O96_Jul_30x), the emissions generally reach
410 higher into the atmosphere in the first three months of the year with enhanced values around 15
411 km, reflecting the stronger convection during austral summer. For open ocean emissions occurring
412 during late austral winter (GTY_O96_Jul_30x, Figure 4c), high CHBr_3 mixing ratios are found
413 around September, however at a lower altitude range compared to the equivalent scenario with
414 open ocean emission occurring during late austral summer (GTY_O96_Jan_30x; Figure 3b).

415 The spatial distribution of annual mean CHBr_3 at 1 km (Figure 5) further confirms the
416 insignificance of the signals from the Australian farming scenarios compared with the background
417 CHBr_3 values. For the global scenarios, localized regions of high mixing ratios are found near the
418 locations of the farms due to the stronger emission. For Darwin_O60_30x, the belt of high mixing
419 ratios is extending northwestward, due to the prevailing easterlies in the tropics. At higher altitudes
420 (e.g., 5 km and 15 km; Figure S2-S3), localized high CHBr_3 is only found near Darwin for the
421 Darwin_O60_30x scenario, reflecting that strong tropical convection is needed to transport short
422 lived gases to such altitudes.

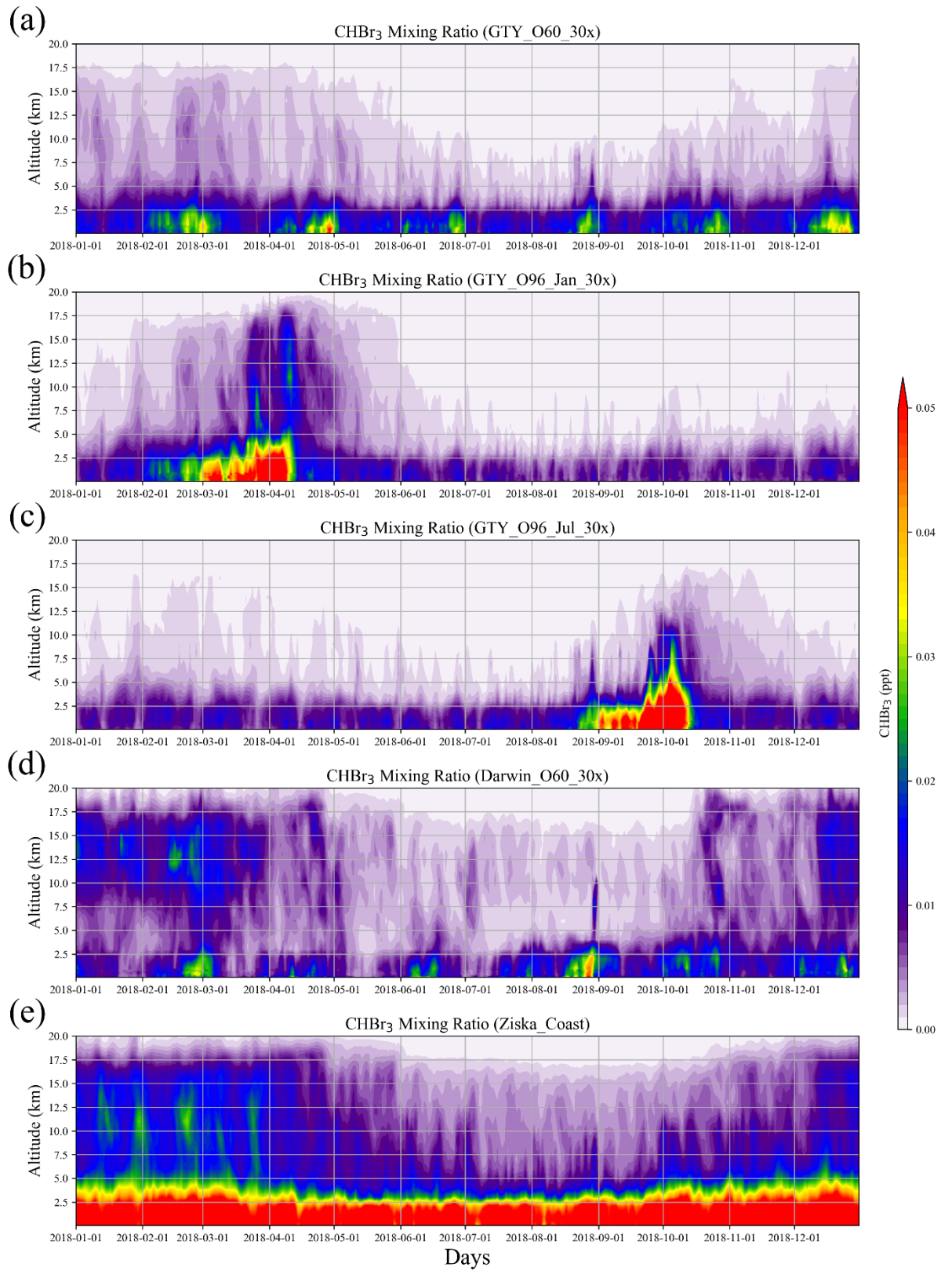
423 The results above suggest that in the boundary layer, global scenarios and extreme events could
424 lead to CHBr_3 comparable mixing ratios as those from the background scenario. Only in the global
425 tropical scenario (Darwin_O60_30x), CHBr_3 mixing ratios, which are larger than the background
426 values, can be found at high altitudes (Figure 4).

427 Simulations of the two extreme scenarios (Geraldton_Ex) for 60 and 96 day growth periods are
428 shown in Figure 6. For the Geraldton_Ex simulations, we assume the implausible scenario that the
429 farm could be totally damaged by cyclone Joyce on the day of harvest in January and the total
430 CHBr₃ content of the *Asparagopsis* stock was simultaneously released to the atmosphere during
431 the event. Both scenarios lead to significant CHBr₃ mixing ratios in the atmosphere, especially at
432 altitudes below 5 km. Among the two scenarios, the Geraldton_Ex96 contributes the larger amount
433 of CHBr₃ emission, as the macroalgae experienced a longer growth period, so the biomass was
434 higher and had accumulated more CHBr₃. When averaged over the same period (Jan 9-Jan 26,
435 2018), the CHBr₃ mixing ratios from Geraldton_Ex96 are much larger than those from
436 Ziska_Coast (Figure 6) below 5 km, and signals with comparable magnitudes, though with smaller
437 coverage, are found at 15 km.

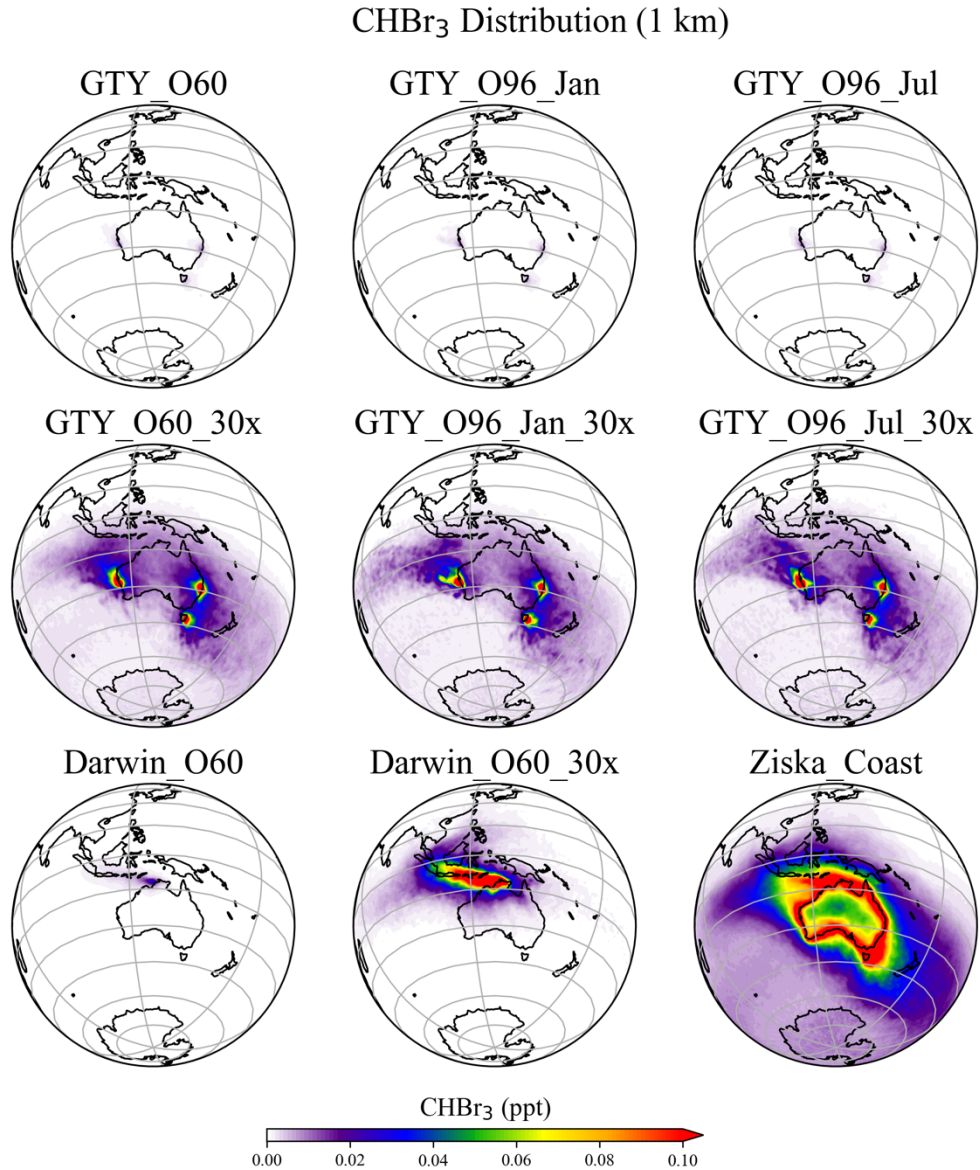
438 As mentioned in section 3.1, the local CHBr₃ emissions due to the seaweed cultivation are
439 generally higher than coastal emission given on 1.0x1.0 grid. However, due to the relatively small
440 spatial extent of the farms, the emissions quickly dilute in the atmosphere, and the magnitude of
441 the mixing ratios decline rapidly off the coast and vertically.

442

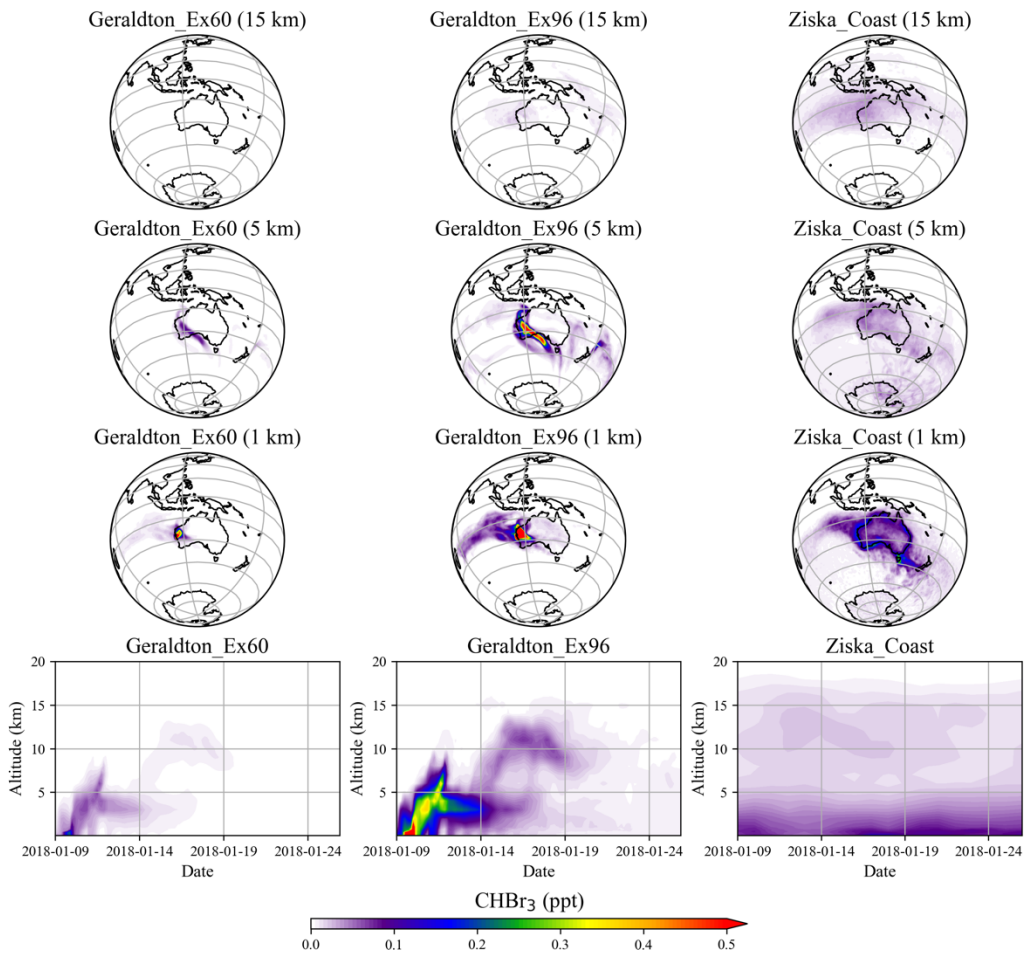
443



445 **Figure 4.** Altitude-time cross-sections of CHBr₃ mixing ratio averaged over [10°-45° S, 105°-
 446 165° E] from Global Scenarios: a) GTY_O60_30x, b) GTY_O96_Jan_30x, c)
 447 GTY_O96_Jul_30x, d) Darwin_O60_30x, and Background Scenario: e) Ziska_Coast.
 448



449 **Figure 5.** Annual mean CHBr₃ spatial distribution from GTY_O60, GTY_O60_30x,
 450 GTY_O96_Jan, GTY_O96_Jan_30x, GTY_O96_Jul, GTY_O96_Jul_30x, Darwin_O60,
 451 Darwin_O60_30x, and Ziska_Coast at 1 km altitude.
 452



453
 454 **Figure 6.** 17-day average of spatial distribution and altitude-time cross-sections of CHBr₃ mixing
 455 ratio averaged over [10°-45° S, 105°-165° E] for Geraldton_Ex60, Geraldton_Ex96, and
 456 Ziska_Coast.

457
 458
 459 **4. Ozone depletion potential for CHBr₃**

460
 461 The ODP distribution from Pisso et al. (2010) for the region around Australia, South-East Asia,
 462 and the Indian Ocean for the Southern Hemisphere (SH) summer and winter is shown in Figure 7.
 463 The ODP distribution changes strongly with season as the transport of short-lived halogenated
 464 substances such as CHBr₃ depends on the seasonal variations of the location of the Intertropical
 465 Convergence Zone (ITCZ). Highest ODP values of 0.5, which imply that any amount (per mass)
 466 of CHBr₃ released from the specific location will destroy half as much stratospheric ozone as the
 467 same amount of CFC-11 released from this location, are found during July over the Maritime
 468 continent and during January over the West Pacific south of the equator. The northern Australian

469 coastline shows highest ODP values during January when the thermal equator and the ITCZ are
470 shifted southwards and ODP values for Yamba and Darwin are 0.26 and 0.29, respectively. The
471 other two locations as well as all four locations during SH winter, show ODP values of only up to
472 0.1.

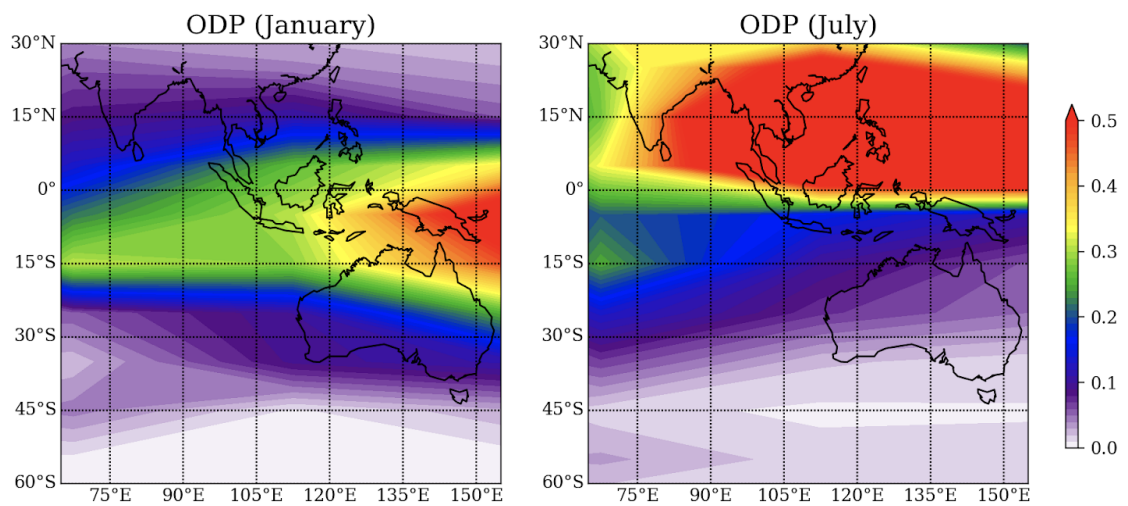
473 As demonstrated in section 3, the total annual CHBr_3 emissions from any location are independent
474 of the details of the farming practice, however, the ODP-weighted emissions change for the
475 different scenarios as the growth periods fall into different seasons with varying ODP values. In
476 general, the scenario of one harvest period in SH summer leads to larger ODP-weighted emissions
477 when compared to the same biomass harvested throughout the year. In addition to the harvesting
478 practice, the locations of the farms have a large impact on the efficiency of the CHBr_3 transport to
479 the stratosphere and thus on the ODP-weighted emissions.

480 The ODP-weighted emissions of CHBr_3 for different emission scenarios are shown in Figure 8.
481 *Asparagopsis* farming at GTY (GTY_O60) leads to additional CHBr_3 emissions of up to 2.53 Mg
482 per year. If all farming ($\sim 3.5 \times 10^4$ Mg DW *Asparagopsis*) occurs in Darwin (Darwin_O60), ODP-
483 weighted emissions would increase to 6.48 Mg CHBr_3 per year. In comparison, all naturally
484 occurring emissions around the Australian coastline (Ziska_coast) lead to ODP-weighted CHBr_3
485 emissions of 221.52 Mg per year. In consequence, *Asparagopsis* farming in the three locations
486 Geraldton, Triabunna and Yamba would lead to an increase of the ODP-weighted emissions from
487 Australian coastal emissions of 1.14%. If all farming would take place in Darwin, ODP-weighted
488 CHBr_3 emissions would increase by 2.93%.

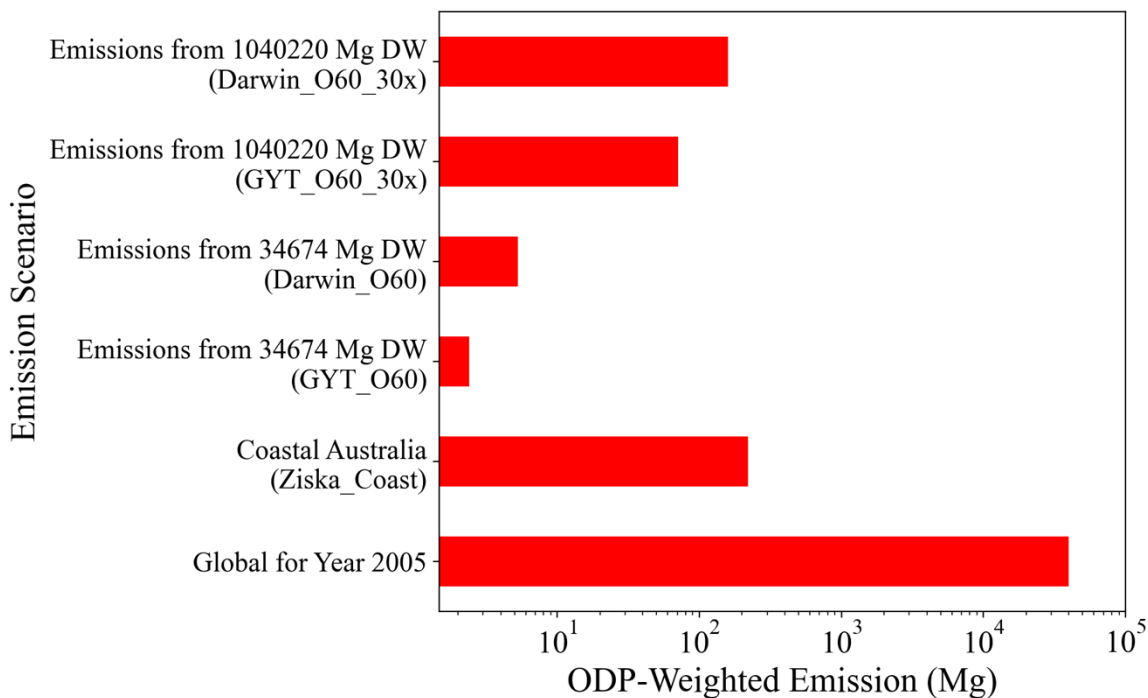
489 As the global ODP-weighted emissions were estimated to be around 4.0×10^4 Mg per year (bottom-
490 most bar in Figure 8., Tegtmeier et al., 2015), the additional contribution due to the Australian
491 farming scenarios in GTY or Darwin would be negligible increasing the contribution of CHBr_3
492 emissions to ozone depletion by 0.006% and 0.016%, respectively. Even if the farming would be
493 upscaled to cover the global needs ($\sim 1.0 \times 10^6$ Mg DW), the ODP-weighted CHBr_3 emissions would
494 only increase to 75 Mg and 195 Mg for farming in GTY (GTY_O60_30x) and Darwin
495 (Darwin_O60_30x), respectively. Thus produced CHBr_3 would increase the current contribution
496 of CHBr_3 to stratospheric ozone depletion by 0.19% and 0.48%, which is again a very small
497 contribution.

498 To assess the increase of the ODP-weighted CHBr_3 emissions under the most extreme and
499 implausible conditions, we envision the total harvest of one year, which contains 752 Mg (21.7

500 mg $\text{CHBr}_3/\text{g DW}$ *34674 Mg DW) CHBr_3 , stored in a warehouse of 50 x 25 x 5 m in either of the
501 four locations. We assume that the facility is destroyed, and all 750 Mg released to the atmosphere.
502 Then maximum ODP-weighted CHBr_3 emissions would occur for the release in Darwin during
503 January and amount to 215.9 Mg almost doubling the ODP-weighted coastal CHBr_3 emissions of
504 Australia. If the entire content of $\sim 1.0 \times 10^6$ Mg *Asparagopsis* DW (21.7 mg $\text{CHBr}_3/\text{g DW}$ *1040220
505 Mg DW=22573 Mg CHBr_3) would be released in Darwin, the additional contribution of CHBr_3 to
506 global ozone depletion could reach 16%.
507



508
509 **Figure 7.** Spatial distribution of the ODP in January and July from Pisso et al. (2010), plotted with
510 interval of 0.01.
511



512
 513
 514 **Figure 8.** The ODP-weighted emissions of CHBr_3 for Global Scenarios (GTY_O60_30x and
 515 Darwin_O60_30x), Australian Scenarios (GTY_O60 and Darwin_O60), Coastal Australian
 516 emission (Ziska_Coast), and global ODP-weighted emission for 2005 taken from Tegtmeier et al.
 517 (2015) as a reference, note that the x-axis is exponential.

518
 519
 520 **5. Summary and Conclusions**

521
 522 In this study, we assessed the potential risks of CHBr_3 released from *Asparagopsis* farming near
 523 Australia for the stratospheric ozone layer by analyzing different cultivation scenarios. We
 524 conclude that the intended operation of *Asparagopsis* seaweed cultivation farms with an annual
 525 yield of 34674 Mg DW in either open ocean or terrestrial cultures at the locations Triubanna,
 526 Yamba, Geraldton, and Darwin will not impact the ozone layer under normal operating conditions.
 527 For Australia scenarios with an annual yield of $\sim 3.5 \times 10^4$ Mg DW and algae growth rate of 5% per
 528 day, the expected annual CHBr_3 emission from the considered *Asparagopsis* farms into the
 529 atmosphere (~ 27 Mg, 0.11 Mmol) is less than 0.9% of the coastal Australian emissions (~ 3109
 530 Mg, 12.3 Mmol). This contribution is negligible from a global perspective by adding less than
 531 0.01% to the worldwide CHBr_3 emissions from natural and anthropogenic sources. The overall
 532 emissions from the farms would be even smaller with a faster growth rate for the same annual

533 yield. We have assumed a high CHBr_3 production of 21.7 mg/g DW from superior strains and
534 expected lower CHBr_3 production of 14 mg/g DW would likewise reduce emissions to the
535 atmosphere.

536 The CHBr_3 emissions from the localized *Asparagopsis* farms could be larger than emissions from
537 coastal Australia. However, the overall atmospheric impact of the *Asparagopsis* farms is
538 negligible, as the CHBr_3 dilutes rapidly and degrades in the atmosphere under normal weather
539 conditions. Mixing ratios of CHBr_3 are generally dominated by the coastal Australian emissions.
540 In global scenarios with annual yield $\sim 1.0 \times 10^6$ Mg DW, localized CHBr_3 mixing ratios
541 comparable to the background values can be found in the lower troposphere. In the upper
542 troposphere, on the other hand, mixing ratios larger than background values only appear in the
543 global tropical scenario (Darwin_O60_30x). The release of the complete CHBr_3 content from the
544 macroalgae to the environment on very short timescales (e.g., days) due to extreme weather
545 situations could contribute significant amounts to the atmosphere, especially during times when
546 the standing stock biomass is relatively large (Geraldton_Ex96). While such extreme scenarios
547 could lead to much larger mixing ratios than background values, such mass release events are
548 implausible because even if a farm was totally destroyed the seaweed stock could not
549 instantaneously release all the accumulated CHBr_3 . Such scenarios have been included here to
550 evaluate a catastrophic and likely impossible worst-case scenario.

551 The impact of CHBr_3 from the proposed seaweed farms on the stratospheric ozone layer is assessed
552 by weighting the emissions with the ozone depletion potential of CHBr_3 . In total, Australia
553 scenarios could lead to additional ODP-weighted CHBr_3 emissions of up to 2.53 Mg per year with
554 farms located in Geraldton, Triubana and Yamba. With all farming performed in Darwin
555 (Darwin_O60), the emitted CHBr_3 could reach the stratosphere on shorter time scales and ODP-
556 weighted emissions would increase to 6.48 Mg, which is less than 0.02% of the global ODP-
557 weighted emissions. For global tropical scenario (Darwin_O60_30x), the ODP-weighted
558 emissions amount to 175 Mg, increasing the global ozone depletion by 0.48%, resulting in a very
559 small contribution. The ODP used in this study, does not include the impact of VLSL product
560 gases. Previous modelling studies have highlighted the role of product gas treatment and their
561 impact on the stratospheric halogen budget (e.g., Fernandez et al., 2021). Including product gas
562 entrainment can lead to up to 30% larger ODP values for CHBr_3 (Engel and Rigby et al., 2018;
563 Tegtmeier et al., 2020), thus the ODP-weighted emissions presented here can be up to 30% larger.

564 However, this does not affect our assessment of the potential importance of cultivation induced
565 CHBr₃ as the ratios of the impact of each scenario compared with the global ODP-weighted
566 emission remain the same. Another point to note is that Ziska emission potentially underestimate
567 background levels of CHBr₃ in coastal regions (Ziska et al., 2013), e.g. CHBr₃ measurement in
568 Cape Grim, which is close to Triabunna, shows much larger CHBr₃ mixing ratio (~1.5 ppt, Dunse
569 et al., 2020). However, including such higher atmospheric background values in the Ziska
570 methodology would not really increase the fluxes as the flux is driven by air-sea gradient (Eq. 3).
571 Our conclusions will still hold if stronger coastal emission is applied, as it will increase the
572 background CHBr₃ mixing ratios.

573 We note that all data characterizing the potential systems for the production of *Asparagopsis* are
574 based on few available literature data, laboratory scale tests and relatively small-scale field trials.
575 This not only places limitations on the technological representativeness of a future system and the
576 temporal validity of the study, but also demonstrates importance for directed studies, especially on
577 the release of CHBr₃ from *Asparagopsis* during cultivation. As this understanding evolves so will
578 the cultivation and processing technologies engineered to conserve the antimethanogenic CHBr₃
579 in the seaweed biomass which is the primary value feature of *Asparagopsis*. These limitations are
580 largely mitigated in our study by evaluating various environmental and meteorological conditions
581 ranging from conservative to most extreme scenarios and by investigating different farming
582 practices based on various sensitivity studies.

583

584 **Data availability**

585 The CHBr₃ emission data and FLEXPART output can be obtained from the authors on request via
586 BQ (bquack@geomar.de), ST (susann.tegtmeier@usask.ca), or YJ (yue.jia@noaa.gov).

587

588 **Author Contributions**

589 BQ initialized the idea. YJ, BQ, and ST carried out the calculations and analysis. YJ performed
590 the FLEXPART simulations and produced the figures. YJ, BQ, and ST wrote the manuscript with
591 the contribution from other co-authors RK and IP. RK contributed to conceptualization, design,
592 writing, editing, procurement of funding. All the authors contributed to discussions and revisions
593 of the manuscript.

594

595 **Competing interests**

596 The authors declare that they have no conflict of interest.

597

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