1   Seasonal Analysis of Reduced and Oxidized Nitrogen-Containing

2   Organic Compounds at a Coastal Site

3   Jenna C. Ditto[1,†], Jo Machesky[1], and Drew R. Gentner[1,*]

4   [1] Department of Chemical and Environmental Engineering, Yale University,

5   New Haven, CT, 06511, USA

6   [†] Now at: Department of Chemical Engineering and Applied Chemistry, University of Toronto,

7   Toronto, ON, M5S 3E5, Canada

8   [*] Correspondence to: drew.gentner@yale.edu

10   **Abstract**

11   Nitrogen-containing organic compounds, which may be directly emitted to the

12   atmosphere or may form via reactions with prevalent reactive nitrogen species (e.g. NH_3, NO_x,

13   NO_3), have important but uncertain effects on climate and human health. Using gas and liquid

14   chromatography with soft ionization and high-resolution mass spectrometry, we performed a

15   molecular-level speciation of functionalized organic compounds at a coastal site on the Long

16   Island Sound in summer (during the LISTOS 2018 campaign) and winter. This region often

17   experiences poor air quality due to the emissions of reactive anthropogenic, biogenic, and

18   marine-derived compounds and their chemical transformation products. Indeed, we observed a

19   range of functionalized compounds containing oxygen, nitrogen, and/or sulfur atoms resulting

20   from a mix of direct emissions and chemical transformations, including photochemical

21   processing in summer and aqueous-phase processing in winter. In both summer and winter,

22   nitrogen-containing organic aerosols dominated the observed distribution of functionalized

23   particle-phase species ionized by our analytical techniques, with 85% and 68% of measured

24 compound abundance containing a nitrogen atom, respectively. Nitrogen-containing particles

25 included reduced nitrogen functional groups (e.g. amines, imines, azoles) and common NO_z

26 contributors (e.g. organonitrates). The prevalence of reduced nitrogen functional groups

27 observed in the particle-phase, while frequently paired with oxygen-containing groups elsewhere

28 on the molecule, often rivaled that of oxidized nitrogen groups detected by our methods.

29 Supplemental gas-phase measurements, collected on adsorptive samplers and analyzed with a

30 novel liquid chromatography-based method, suggest that gas-phase reduced nitrogen compounds

31 are possible contributing precursors to the observed nitrogen-containing particles. Altogether,

32 this work highlights the prevalence of reduced nitrogen-containing compounds in the less-

33 studied Northeastern U.S., and potentially in other regions with similar anthropogenic, biogenic,

34 and marine source signatures.

36 **1. Introduction**

37   Coastal regions near the Long Island Sound often experience poor air quality due to a

38 combination of biogenic and anthropogenic emissions from upwind metropolitan areas along the

39 East Coast of the U.S. It is well established that these emissions undergo chemical

40 transformations to form secondary pollutants during hours to days of over-water transport to

41 downwind locations, including the states of Connecticut, Rhode Island, and Massachusetts (e.g.

42 Cleveland et al., 1976). Emissions of gas-phase organic compounds (e.g. volatile, intermediate,

43 and semi-volatile organic compounds (VOCs, IVOCs, SVOCs)) and primary organic aerosols

44 (POA) are oxidized via numerous pathways in the atmosphere to yield ozone (O_3) and secondary

45 organic aerosol (SOA) (Hallquist et al., 2009). SOA constitutes a variable but significant fraction

46 of particulate matter with a diameter of 2.5 μm or less (i.e. $PM_{2.5}$). Both O_3 and $PM_{2.5}$ are of

47   particular concern for human health and climate; O_3 is known to cause an increase in respiratory-

48   related illnesses (Di et al., 2017; Jerrett et al., 2009), while $PM_{2.5}$ is known to cause adverse

49   cardiovascular, respiratory, and cognitive effects and to impact climate forcings (Di et al., 2017;

50   Hallquist et al., 2009; Kilian and Kitazawa, 2018; Pope and Dockery, 2006). Coupled with local

51   emissions and chemistry, these incoming aged air parcels from coastal metropolitan areas

52   contribute to the Long Island Sound region often entering nonattainment for O_3 (United States

53   Environmental Protection Agency, 2020), especially in the summer.

54   The chemistry and composition of organic compound emissions and secondary

55   transformation products in the Long Island Sound area is historically understudied, though some

56   past work has advanced our understanding of important sources and chemical pathways in the

57   region. For example, VOC and sub-micron particulate matter composition were investigated

58   during the 2002 New England Air Quality Study (de Gouw et al., 2005), with further VOC

59   speciation in 2004 during the New England Air Quality Study – Intercontinental Transport and

60   Chemical Transformation campaign (Warneke et al., 2007). More recently, a 2015 aircraft

61   campaign in the Northeast U.S. called WINTER characterized wintertime chemistry in the region

62   and also investigated organic aerosol formation via aerosol mass spectrometry (Schroder et al.,

63   2018). Finally, the LISTOS campaign in 2018 focused on measuring and modeling O_3 mixing

64   ratios over the Sound to investigate the dynamics of O_3 formation linked to large metropolitan

65   areas along the coast and associated downwind impacts (Zhang et al., 2020).

66   However, little is known about the molecular-level chemical composition of the gas-

67   phase I/SVOCs and functionalized organic aerosol formed in the Northeastern U.S. This

68   molecular-level speciation is key to understanding the physical/chemical properties of these

69   compounds in the atmosphere and their chemical transformations, especially for classes of

70   compounds containing reduced and oxidized nitrogen functional groups, whose emissions,

71   lifetime, and ultimate impacts are generally poorly understood. For example, nitrogen-containing

72   compounds that serve as reservoir species for nitrogen oxides may increase the overall lifetime

73   of nitrogen oxides in the atmosphere via renoxification mechanisms (e.g. the photolysis of

74   particulate nitrates, which has been studied in the marine boundary layer (Ye et al., 2016)); some

75   may act as light absorbing chromophores (e.g. the brown carbon studied from a methylglyoxal

76   and ammonium sulfate system, which yielded mostly N-containing chromophores (Lin et al.,

77   2015)); and some may have adverse, but uncertain, effects on human health (e.g. impacts on

78   immune response to allergens (Ng et al., 2017)).

79   With the overall goal of examining the composition and contribution of nitrogen-

80   containing organic compounds from mixed anthropogenic, biogenic, and marine sources, we

81   collected measurements of organic gases and aerosols on the coast of the Long Island Sound in

82   Guilford, Connecticut. We note that we used this site as a case study, but our observations of

83   emissions and chemistry at this site are likely informative for other coastal urban and downwind

84   regions due to the ubiquity of nitrogen-containing emissions from varied sources.

85   Samples were collected during the summer and winter, and analyzed via high resolution

86   mass spectrometry to speciate the complex mixture of emissions and chemical transformation

87   products. These samples were taken alongside several targeted pollutant measurements including

88   O_3, nitrogen oxides (NO_x), particulate matter with a diameter of ≤ 2.5 μm ($PM_{2.5}$), and black

89   carbon (BC), all to inform our chemically-speciated analyses, and to contribute to a longer-term

90   characterization of this coastal area.

91   The objectives of this study were to: (1) investigate compositional trends and possible

92   chemical pathways contributing to measured summer and winter functionalized organic aerosols

93   at this site; (2) examine the relative contributions of reduced and oxidized nitrogen groups to

94   functionalized organic aerosol; and (3) use a novel sampling and liquid chromatography-based

95   analytical approach to probe the molecular-level composition of functionalized gas-phase

96   organic compounds and investigate possible nitrogen-containing gas-phase precursors to the

97   observed reduced nitrogen-containing particles.

98

99   **2.   Materials and Methods**

100   We collected measurements at the Yale Coastal Field Station (YCFS) in Guilford,

101   Connecticut (41.26°N, 72.73°W) (Rogers et al., 2020). Inlets were positioned facing the Long

102   Island Sound (i.e. South-Southeast) to capture onshore flow. The YCFS often received aged

103   urban incoming air from East Coast metropolitan areas, similar to known common air parcel

104   trajectories in the region (Figure S1). However, due to extensive mixing in the Northeast corridor

105   and over the Long Island Sound, along with extended collection times for offline gas- and

106   particle-phase samples, we also observed considerable biogenic and anthropogenic influence

107   from other areas of the Northeastern U.S.

108   2.1. Offline Samples of Organic Particles and Gases Analyzed via Liquid and Gas

109   Chromatography with Mass Spectrometry

110   We discuss three types of sampling and quadrupole time-of-flight mass spectrometry-

111   based analyses here: particles collected on Teflon filters and analyzed using liquid

112   chromatography with electrospray ionization, gases collected on packed adsorbent tubes and

113   analyzed using gas chromatography with atmospheric pressure chemical ionization, and

114   functionalized gases collected on cooled polyether ether ketone (PEEK) samplers and analyzed

115   using liquid chromatography with electrospray ionization. Teflon filter and adsorbent tube

116   measurements were collected at the YCFS during the summer as part of the 2018 Long Island

117   Sound Tropospheric Ozone Study (LISTOS), from July 9 to August 29, 2018. Additional filter

118   and adsorbent tube samples were collected during the following winter from February 25 to

119   March 5, 2019. Supplemental wintertime gas-phase samples on PEEK tubing were collected

120   briefly from March 5-6 2020, prior to the COVID-19 shutdown.

121   For filter and adsorbent tube collection, we used a short inlet (0.9 m long, 5/8" OD

122   stainless steel tubing, positioned 2.5 m above the ground) upstream of the sampler to allow the

123   sampling media to be housed in an air-conditioned trailer. A stainless steel mesh screen (84

124   mesh) was used at the opening of the inlet to limit particle size to ~PM$_{10}$ and to prevent large

125   particles from entering the sampler (Ditto et al., 2018). A custom filter and adsorbent tube

126   housing was connected to the inlet to simultaneously collect particle- and gas-phase organic

127   compounds, respectively (Sheu et al., 2018). The filter was positioned immediately upstream of

128   the adsorbent tube to collect particles for analysis and to prevent particles from reaching the gas-

129   phase sample. The housing was designed to minimize spacing between the filter and adsorbent

130   tube to reduce gas-phase loses to upstream surfaces, and was built out of a modified passivated

131   stainless steel filter holder and an aluminum block with sealed 6.34 mm (1/4") holes for

132   adsorbent tubes (Sheu et al., 2018).

133   Teflon filters (47 mm, 2.0 μm pores, Tisch Scientific) were used for particle-phase

134   sampling. Filters were collected at 20 L/min for 8 hours, during the day (9:00am-5:00pm) and at

135   night (9:00pm-5:00am). Samples were extracted in methanol and analyzed via liquid

136   chromatography (LC) using an Agilent 1260 Infinity LC and an Agilent Poroshell 120 SB-Aq

137   reverse phase column (2.1 x 50 mm, 2.7 μm particle size). The LC was coupled to an

138   electrospray ionization (ESI) source and a high-resolution mass spectrometer (Agilent 6550 Q-

139   TOF), and operated following previously described methods (Ditto et al., 2018, 2020). Filter

140   extracts were run with MS (i.e. TOF-only, to identify molecular formulas) and MS/MS (i.e.

141   tandem mass spectrometry, to identify functional groups) data acquisition, using both positive

142   and negative mode electrospray ionization. These methods are hereafter referred to as "LC-ESI-

143   MS" and "LC-ESI-MS/MS", respectively. Acquisition and non-targeted analysis methods,

144   including data QA/QC, are discussed in past work (Ditto et al., 2018, 2020). After stringent

145   QA/QC for peak shape and accurate molecular formula determination, non-targeted compound

146   identification from LC-ESI-MS examined an average of 200±56 and 167±47 compounds per

147   sample analyzed in summer and winter, respectively, across 34 samples in summer and 15 in

148   winter. For MS/MS, we used SIRIUS with CSI:FingerID for functional group identification with

149   a subset of compounds from MS analysis (Dührkop et al., 2015, 2019), as detailed in past work

150   (Ditto et al., 2020). We note that filter ion abundance data is presented as combined positive and

151   negative ionization mode data, showing all compounds as having equal ionization efficiency.

152   While exact ionization efficiency differences between compound types exist, their effects for

153   multifunctional compounds present in a complex mixture are uncertain. Thus, similar to other

154   studies and to our past work, we treat the intercomparison across compounds with equal

155   ionization efficiencies (Ditto et al., 2018).

156   Gas-phase samples were collected on glass adsorbent tubes (6.35 mm OD, 88.9 mm long)

157   packed with quartz wool, glass beads, Tenax TA, and Carbopack X (Sheu et al., 2018). Samples

158   were collected at 200 mL/min for 2 hours, during the day (2:00-4:00pm) and at night (2:00-

159   4:00am). Adsorbent tubes were analyzed using a GERSTEL TD3.5+ thermal desorption unit and

160   an Agilent 7890B gas chromatograph (GC) with a DB5-MS UI column (30 m x 320 μm x 0.25

161   μm). The GC was coupled to an atmospheric pressure chemical ionization (APCI) source and the

162   same Q-TOF as above, operated with MS (i.e. TOF-only) data acquisition and positive ionization

163   mode only. These methods are hereafter called "GC-APCI-MS" and acquisition and analysis

164   methods are discussed in past work (Ditto et al., 2021; Khare et al., 2019). After QA/QC, this

165   non-targeted analysis yielded an average of 388±201 and 612±133 compounds per sample in

166   summer and winter, respectively, across 34 samples in summer and 14 samples in winter.

167   Finally, as a supplemental analysis to probe the composition of functionalized gases that

168   were not GC amenable and thus not measured using the adsorbent tube and thermal desorption-

169   gas chromatography techniques mentioned above, we used PEEK-based sample collectors and

170   liquid chromatography to trap and speciate oxygen-, nitrogen-, and/or sulfur-containing gases

171   without thermal desorption. This method was designed to target functionalized gases, which

172   represent important precursors, intermediates, and by-products in the atmospheric processing of

173   emitted organic compounds but are often challenging to speciate with traditional GC techniques

174   due to their chemical functionality, reactivity, and/or thermal lability. Additionally, in many gas-

175   phase measurement systems, primary emissions (i.e. hydrocarbons) can overwhelm the signal of

176   more functionalized analytes, adding to the challenge of speciating these lower abundance

177   compounds.

178   Thus, to probe the chemical composition of these functionalized gases, we used a

179   sampling approach, desorption method, separation method, and ionization technique that

180   leveraged their relatively lower volatility and higher polarity. This included adsorptive sampling

181   onto cooled PEEK tubing followed by direct inline desorption into the LC mobile phase for LC-

182   ESI-MS analysis. ESI was specifically chosen here because it is sensitive to functionalized

183   compounds. Testing was performed in positive and negative ionization mode, but field samples

184   were run in positive mode only. While further details and discussion can be found in Section S1

185   and Tables S1-3, briefly, PEEK tubing was cooled to 2°C and used as an adsorptive collector,

186   with a Teflon filter positioned upstream of the PEEK tubing to remove particles. PEEK was

187   selected due to its inert behavior, thus reducing the possibility for surface-analyte interactions

188   that might inhibit effective inline solvent desorption and dissolution. PEEK is also compatible

189   with the solvents used in the LC system, and is frequently used in LC instruments. Field samples

190   were collected on cooled PEEK tubing during the subsequent winter (March 5-6, 2020), for 2

191   hours each between 8:00am-2:00pm. For these 2 hour (~2.6 L) field samples, functionalized

192   gases in a typical 100-250 g/mol molecular weight range were resolvable at ~25-60 ppt in the

193   atmosphere, based on instrument detection limits (Ditto et al., 2018). For analysis, each PEEK

194   collector was installed in the LC system flow path, and analytes were directly desorbed using the

195   LC mobile phase solvents then trapped and focused on the LC column for 20 minutes, before

196   being analyzed using the same LC-ESI-MS system in positive ionization mode (Figure 1). This

197   inline mobile phase desorption step gently mobilized potentially fragile analytes from the PEEK

198   collector and trapped and focused them on the LC column prior to chromatographic separation

199   and mass spectral analysis. Additionally, this preconcentration step allows for the detection and

200   characterization of lower concentration species.

201   We note that there are other existing approaches for offline collection of highly-

202   functionalized organic gases and particles that are compatible with LC analysis such as spray

203   chambers, particle into liquid samplers, coated denuders, PUF sampling, and more. This PEEK

204   sampling method with inline desorption into the LC mobile phase was pursued to reduce sample

205   preparation steps and thus possibilities for losses (e.g. during solvent extraction or evaporative

206   preconcentration), as well as for its direct similarity to the filter-based particle-phase LC-ESI-MS

207   analysis.

208   We also note that for all LC analyses, it is possible that some functional groups of interest

209   (e.g. organonitrates) may have undergone hydrolysis in the LC mobile phase, which was

210   primarily water at the beginning of the LC solvent gradient. If so, some of the observed

211   compounds could be byproducts of other functionalized species, though this would require

212   hydrolysis to occur on the order of a few minutes. While we did not observe any of our nitrogen-

213   containing test standards to hydrolyze over these timescales, standards were not available to

214   reflect every functional group observed in these datasets.

215   2.2. Supporting Measurements

216   O_3, NO_x, $PM_{2.5}$, and BC concentrations were recorded concurrently during both summer

217   and winter sampling periods. O_3 was measured with a 2B Tech Model 202 Ozone Monitor, NO_x

218   with a Thermo Scientific Model 42i-TL Analyzer, $PM_{2.5}$ with a MetOne BAM-1020 instrument,

219   and BC with a Magee Scientific AE33 Aethalometer. O_3 and NO_x inlets were constructed of FEP

220   tubing (1/4" OD), with a Teflon filter housed in a PFA filter holder upstream to remove particles.

221   The $PM_{2.5}$ inlet was made of stainless steel tubing (1 ¼" OD) and the BC inlet was made of

222   copper tubing (3/8" OD). Both particle inlets were outfitted with a $PM_{2.5}$ cyclone to limit particle

223   size to 2.5 μm and below.

224   All inlets were mounted 3 m above the ground. Instrument flow rates were calibrated

225   with an external mass flow controller. O_3 and NO_x monitors were zeroed with laboratory-

226   generated zero air. The O_3 monitor was calibrated against Connecticut Department of Energy and

227   Environmental Protection instrumentation and further confirmed with an O_3 generator in the lab.

228   The NO_x monitor was calibrated using a NO standard (AirGas, 2 ppm NO in nitrogen, ± 5%)

229   diluted to 25 ppb with laboratory-generated nitrogen gas. The BC instrument was programmed to

230   conduct an automatic performance check using particle-free air and the $PM_{2.5}$ instrument was

231   zeroed following MetOne protocols with particle-free air. O_3 and NO_x data were collected at 1-

232   second intervals, BC data were collected at 1-minute intervals, and $PM_{2.5}$ data were collected at

233   1-hour intervals. BC data were saved directly from the instrument, while O_3, NO_x, and $PM_{2.5}$

234   data were recorded with a LabJack T7 datalogger and custom LabView code. In addition, hourly

235   weather data (temperature, relative humidity, wind speed, wind direction) were collected with a

236   WeatherHawk weather station on top of the 3 m tower.

237   During the summer, we also collected a small number of size-resolved particle samples

238   on quartz filters using an eight-stage cascade impactor (Thermo Scientific). Quartz filters were

239   extracted and analyzed following the same procedure as the Teflon filters discussed above, with

240   the addition of a syringe filtration step to remove insoluble fibers. The cascade impactor was

241   positioned on the roof of the trailer and pulled 28.3 L/min through the inlet for periods of 8 hours

242   during the day and at night (same timing as above) or for 24-hour periods.

243   Finally, we computed 48-hour backward trajectories for every hour during each offline

244   sample collection period with the HYSPLIT Backward Trajectory Model (accessed online at

245   https://www.ready.noaa.gov/HYSPLIT.php), using GDAS1.0 meteorological data, the field site's

246   coordinates as each trajectory's end point, and a final trajectory height of 50 m above the ground.

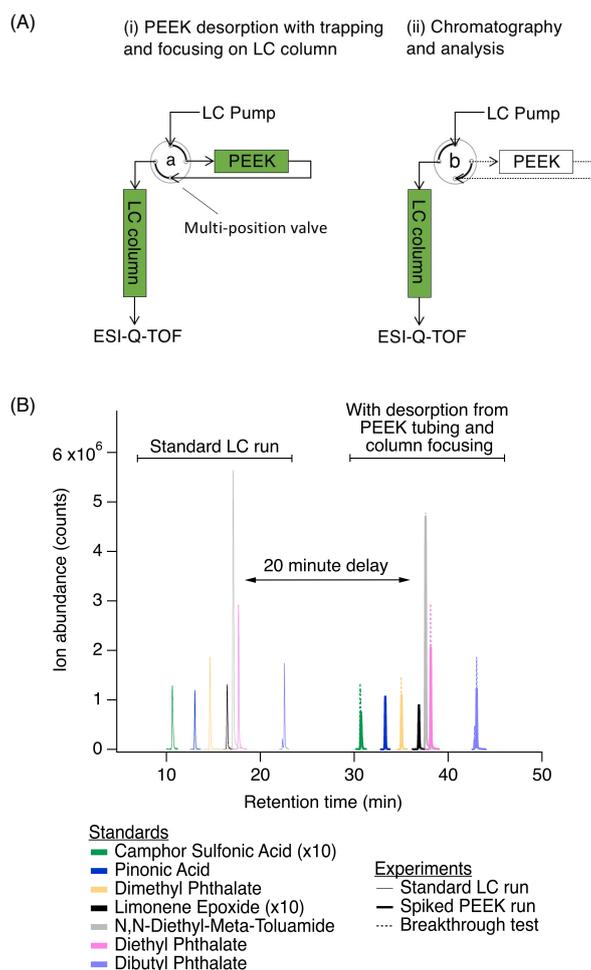



Figure 1. (A) Simplified analytical system setup for functionalized gas-phase compounds, showing (i) desorption from the PEEK collector and trapping on the LC column in order to focus analytes prior to chromatographic separation, and (ii) subsequent chromatographic separation and analysis. Green shading indicates active solvent flow through PEEK collector and/or LC column. A multi-position valve was switched from position "a" (panel (i)) to position "b" (panel (ii)) to remove the PEEK collector from the flow path for chromatography and analysis. Table S1 describes the flow rates and solvents used in each of these steps. (B) Comparison of select peaks from a typical LC run (solid traces from 10-23 min) to that from a PEEK collector spiked with a standard (bold traces from 30-43 min) demonstrates desorption, trapping/focusing, and similar chromatography. Comparable results from a 2-hour breakthrough test at 2°C with 22 mL/min air flow are also shown (overlayed dotted traces from 30-43 min). Spiked PEEK and breakthrough tests were performed to validate this sampling and analysis methods, and are discussed further in Section S1. Test analytes were used across a range of functionality, with examples shown here and the full list in Table S2.

3. Results and Discussion

3.1. Characteristics of the Urban Regional Site

Backward trajectories for summertime and wintertime samples showed a strong urban influence. Summertime trajectories ranged from the northwest, west, and especially the southwest (i.e. New York City and other coastal metropolitan areas, similar to well-established and expected air flow patterns near the Long Island Sound). In contrast, trajectories were almost exclusively from the northwest in the winter (Figure S1). These air parcels brought a range of compounds from a mixture of anthropogenic, biogenic, and marine sources to the site, all with differences in gas- and particle-phase source profiles. However, due to the varied backward trajectories, dynamic variations in wind direction over the long duration filter samples (Figure S2), and a high degree of mixing over the Sound, our 8-hour samples are representative of mixed regional conditions in summer and winter, and are thus discussed in this context. Further detailed site characterization can be found in Section S2 and Figures S1-S3.

3.2. Seasonal Comparisons of Functionalized Organic Aerosols

3.2.1. Summertime composition and the influence of photochemistry and NO_x

During this period of active photochemistry, the observed distribution of particle-phase compounds in summertime samples spanned across the intermediate volatility (IVOC) to ultralow volatility organic compound (ULVOC) range, with a predominance of semivolatile (SVOC), low volatility (LVOC), and extremely low volatility organic compounds (ELVOC) (Figure 2A). These compounds existed across a range of functionality. Due to elevated summertime O_3 mixing ratios at the site (shown in Figure S2, 8-hour maximum mixing ratio in summer: 57 ± 20 ppb, vs. winter: 46 ± 5 ppb, $p < 0.05$, including day and night sampling periods), O_3 may have influenced the photochemical processing of emitted volatile species,

284   especially unsaturated biogenic VOCs which readily undergo ozonolysis due to their chemical

285   structure. However, we did not observe a correlation between 8-hour maximum (or 8-hour

286   average) O_3 mixing ratios with average particle-phase volatility, carbon number, or O/C (nor did

287   we observe such relationships for gas-phase organic compounds). There were, however, weak

288   relationships between NO_x mixing ratios and each of these particle-phase characteristics in the

289   summer. While average NO_x mixing ratios were slightly lower during the summer (as shown in

290   Figure S2, 2.3 ± 1.5 ppb in summer vs. 3.7 ± 2.7 ppb in winter, $p < 0.05$), NO_x mixing ratios

291   trended weakly with particle-phase O/C ($r \sim 0.45$), volatility (as saturation mass concentration,

292   log(C_0), $r \sim 0.49$), and inversely with carbon number ($r \sim -0.56$) in summer.

293   While our correlations and conclusions are somewhat limited by the 8-hour filter

294   sampling duration and the resulting highly regionally-mixed samples, one possible hypothesis is

295   that the presence of NO_x could have promoted more fragmentation reactions in the gas-phase

296   (Loza et al., 2014) that decreased average carbon number, and correspondingly increased

297   volatility and O/C. In fact, we observed highly oxidized C_3-C_6 compounds in the gas-phase (from

298   adsorbent tube measurements with GC-APCI (Section S2)) that were possibly products of these

299   fragmentation reactions of larger compounds. These trends of NO_x mixing ratios with O/C,

300   volatility, and carbon number were not apparent for the observed complex mixture of gas-phase

301   organic compounds. However, these highly oxidized gases may not have persisted in the gas-

302   phase and could have been taken up by the condensed/aqueous phase due to their water

303   solubility, where they would have instead contributed to the observed trends of NO_x with carbon

304   number, volatility, and O/C in the particle-phase. We note that if there was significant uptake of

305   gas-phase NO_z to the particle-phase, this may have in part contributed to the particle-phase

306   correlations with NO_x given that the chemiluminescence NO_x analyzer used in this study is

307   known to also respond to gas-phase NO_z (Dunlea et al., 2007).

308   Additionally, NO_x could have been involved in heterogeneous chemistry, promoting

309   oxidation and/or nitrogen addition reactions, such as interaction with $NO_3^•$ to yield organonitrates

310   (Lim et al., 2016), formation and interaction with HONO to yield nitrophenols (Vidović et al.,

311   2018), or other pathways.

312   *3.2.2. Comparison to wintertime composition and the role of aqueous-phase chemistry*

313   In the winter, these same relationships between NO_x and particle-phase characteristics

314   were not observed. This is possibly due to the decreased role of photochemistry in the winter and

315   the increased role of other competing physical and chemical processes, such as aqueous-phase

316   chemistry. The observed chemical composition of the particle-phase in the winter was indeed

317   indicative of increased aqueous-phase processing, with a shifted compound distribution that

318   included lower molecular weight and higher volatility particle-phase species compared to

319   summer (shown by saturation mass concentration in Figure 2A-B). This is consistent with

320   another past study that suggested that aged fog-water samples contained organic compounds with

321   smaller carbon backbone structures than aged non-aqueous particles, and linked this difference to

322   aqueous-phase fragmentation reactions, the uptake of smaller water-soluble gases to the aqueous-

323   phase, and/or less oligomerization (Brege et al., 2018). In contrast, there was a larger proportion

324   of LVOCs, ELVOCs, and ULVOCs in summer, which decreased the average summertime

325   volatility of the functionalized organic aerosol components. This was likely driven by increased

326   photochemical processing in summer, leading to lower saturation mass concentrations (summer:

327    $\log(C_0) = -4.1 \pm 5.0$ $\mu g/m^3$, vs. winter: $\log(C_0) = -1.8 \pm 4.8$ $\mu g/m^3$, $p < 0.05$) and higher O/C

328   (summer: O/C = 0.5 ± 0.4 vs. winter: O/C = 0.4 ± 0.4, $p < 0.05$). To assess the potential

329   contribution of aqueous-phase chemistry, we also estimated aerosol liquid water concentrations

330   based on available data in Section S2.1.

331   Furthermore, from MS/MS analysis, we observed functional groups that were possible

332   indicators of aqueous-phase processing, including the presence of nitrophenols during the winter,

333   which may have formed via dark aqueous-phase reactions with HONO (Vidović et al., 2018),

334   and relatively low contributions from carbonyls across seasons, possibly linked to carbonyl

335   hydrolysis (Ditto et al., 2020). Based on laboratory studies, the presence of azole functional

336   groups and other heterocyclic nitrogen species could also indicate aqueous phase processing, and

337   may be formed from small carbonyl precursors such as glyoxal (DeHaan et al., 2009; Grace et

338   al., 2019) and biacetyl (Grace et al., 2020) reacting with atmospheric ammonia or small amines.

339   Many of the N-only containing azoles observed here had similar substructures to those formed in

340   the aqueous-phase reactions of small carbonyls with ammonia/amines (DeHaan et al., 2009;

341   Grace et al., 2019). In addition, as discussed above, we observed many small gas-phase C_3-C_6

342   compounds at the site in the summer, which likely included multifunctional isoprene oxidation

343   products (e.g. glycoaldehyde, hydroxyacetone, and isomers); these potential precursors could

344   have reacted with atmospheric ammonia or species containing amino groups to form the

345   observed azole-containing reaction products. We observed more azoles during the summer (Ditto

346   et al., 2020), perhaps due to the increased prevalence of the C_3-C_6 precursors and overall

347   prominence of aerosol liquid water.

348   Lastly, the role of aqueous-phase chemistry in the region is further supported by prior

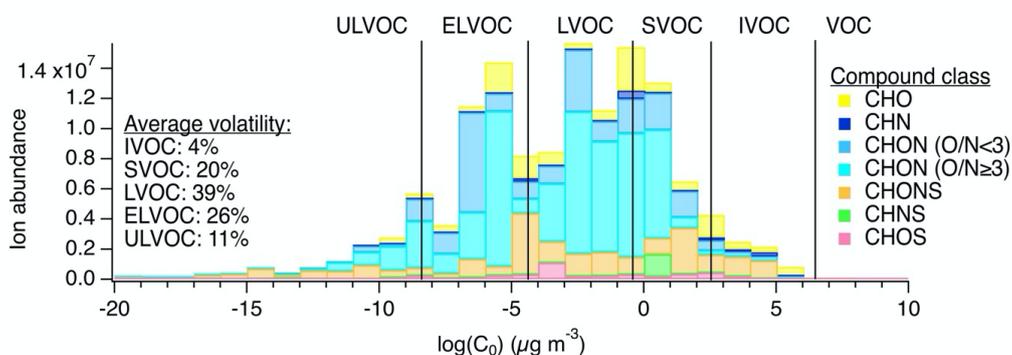
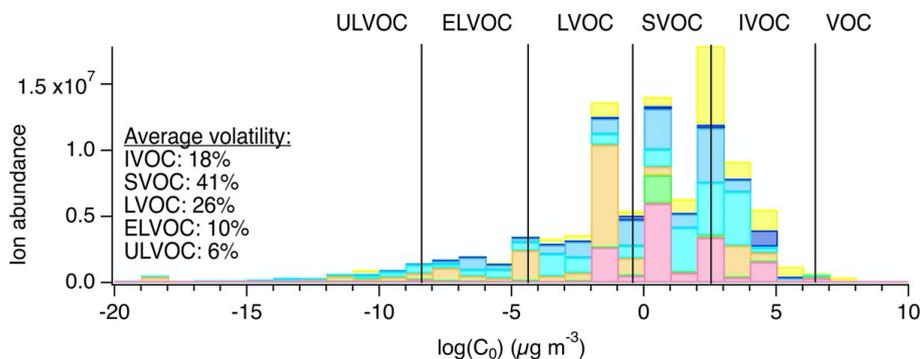
349   summertime observations at Brookhaven National Laboratory (on the opposite side of the Long

350   Island Sound), which examined a low-volatility oxygenated organic aerosol factor in the source

351   apportionment of aerosol mass spectrometry measurements, and showed a strong contribution

352   from carboxylic acids and other ELVOCs that were attributed to aqueous-phase processing

353   (Zhou et al., 2016).

(A) Particle-phase summer volatility distribution (LC-ESI-MS)

(B) Particle-phase winter volatility distribution (LC-ESI-MS)

354   **Figure 2.** Chemical composition of particle-phase organic compound mixtures at the YCFS from
355   LC-ESI-MS measurements. (A) and (B) show particle-phase volatility distributions by
356   compound class in the summer (N=34) and winter (N=15), respectively. Compound volatility
357   was estimated and grouped following approaches in the literature (Li et al., 2016; Schervish and
358   Donahue, 2020). For direct comparison, volatility bins were defined for the same reference
359   temperature in (A) and (B) (i.e. 300 K, the average summertime sampling period temperature),
360   though wintertime saturation mass concentrations for the observed compounds would shift
361   approximately 2 orders of magnitude lower due to lower temperatures (i.e. 270K). Thus, in the
362   winter, a larger fraction of I/SVOCs would have partitioned to the particle phase, though this
363   effect was likely more pronounced for SVOCs than IVOCs (see Table S4).

364   *3.2.3. Comparison to other sites*

365   The distribution of compound classes observed at the YCFS was significantly different

366   from observations at a range of field sites discussed in past studies (Figure 3), including a remote

367   forested site (i.e. the PROPHET site in Northern Michigan), an urban inland site (i.e. near

368   downtown Atlanta) across two seasons, and in New York City (Ditto et al., 2018, 2019). While a

369   detailed site-to-site comparison is outside the scope of this work, the proximity of the YCFS to

370   the ocean and thus the impact of marine emissions and over-water chemistry likely contributed to

371   the differences between the YCFS and inland locations.

372   In particular, at the YCFS, we observed notably smaller relative contributions from

373   compounds containing carbon, hydrogen, and oxygen (i.e. CHO, 11-16% of observed

374   functionalized compounds here vs. 34-50% at other sites), and the contributions from nitrogen-

375   containing particle-phase compounds at the YCFS were in stark contrast to other sites. Here,

376   85% of compounds in summer and 68% of compounds in winter contained at least one nitrogen

377   atom, compared to 38-51% at the other previously studied sites (Figure 3). These nitrogen-

378   containing species were comprised of compounds with various reduced and oxidized nitrogen-

379   containing functional groups with varying oxygen-to-nitrogen ratios (O/N), which are broadly

380   classified and discussed below as compounds containing carbon, hydrogen, and nitrogen (i.e.

381   CHN), and compounds containing carbon, hydrogen, oxygen, and nitrogen (i.e. CHON (O/N

382   ratio < 3), and CHON (O/N ratio ≥ 3)). There were notably greater contributions at the site from

383   nitrogen-containing compounds that also contained at least one oxygen atom, including CHON

384   compounds with O/N < 3 (19-20% here vs. 10-15% at other sites), CHON compounds with O/N

385   ≥ 3 (24-44% here vs. 14-19% at other sites), as well as compounds containing oxygen, nitrogen,

386   and sulfur (i.e. CHONS, 20-21% here vs. 9-10% at other sites) (Ditto et al., 2018). See Section

387   S3 and Figure S4 for discussion of additional compound classes.

388   We note that while these were PM_{10} measurements, the observations of high nitrogen

389   content were not biased by the inclusion of larger, possibly primary particles. Quartz filter

390   samples collected with a cascade impactor at the site during the summer and analyzed with the

391   same LC-ESI-MS methods did not show any statistically significant differences between any of

392   these nitrogen-containing compound classes as a function of particle size across particles ranging

393   from 0.4 to 10 μm. This is consistent with past studies which have demonstrated that amines, as

394   an example of reduced nitrogen, are ubiquitous in size-resolved aerosol samples in urban and

395   rural locations (VandenBoer et al., 2011).

396   The prevalence of nitrogen-containing species at the YCFS is consistent with the study at

397   Brookhaven National Laboratory discussed above, where a dedicated nitrogen-enriched aerosol

398   mass spectrometry factor was identified, and contained prevalent signal from aliphatic amines

399   and amides. However, in the Brookhaven study, the nitrogen-enriched factor was associated with

400   industrial amine emissions that were enhanced during periods of south/southwestern backward

401   trajectory influence, and that had correlations with tracers linked to industrial processes. In our

402   study, there was no correlation between backward trajectory direction and the contribution of

403   nitrogen-containing species. Also, wintertime air parcels arrived predominantly from directions

404   other than south/southwest, suggesting that the nitrogen-containing species observed in our study

405   were the result of mixed anthropogenic, biogenic, and marine precursors and their transformation

406   products.

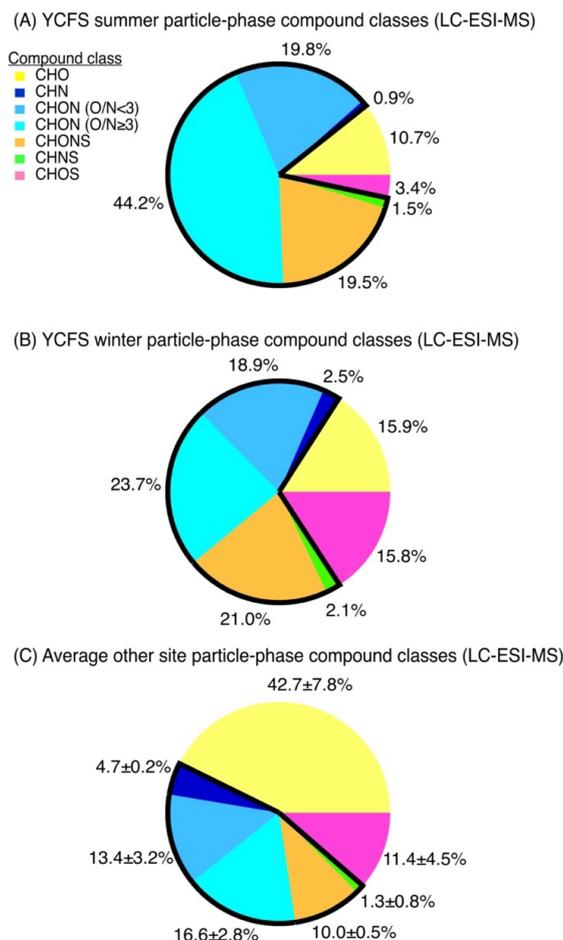



Figure 3. Enhancements in nitrogen-containing particle-phase compounds compared to other sites, from LC-ESI-MS measurements. Particle-phase compound class distributions in the summer (A) and winter (B) at the YCFS, weighted by compound abundance, in contrast with the average compound class distribution from previously studied forested, urban inland, and urban coastal sites (C). Nitrogen-containing compound class contributions are outlined in black.

3.3. Speciating Particle-Phase Multifunctional Nitrogen-Containing Compounds

The observed particle-phase species were highly functionalized, often multi-functional, and contained combinations of oxygen, nitrogen, and/or sulfur heteroatoms. Here, we discuss the functional groups present, broken up by the nitrogen-containing compound classes shown in Figures 2-3, with additional discussion of other relevant compound classes in Section S3.

418   *3.3.1. CHN compounds*

419   While nitrogen-containing compounds in general were very prominent at the site (Figure

420   3A-B), CHN compounds were relatively less abundant in these samples of functionalized

421   organic aerosol. Particle-phase CHN compounds represented just 1% and 3% of observed

422   functionalized organic aerosol abundance in summer and winter, respectively, which was similar

423   to observations at other ambient sites (~5% CHN) (Ditto et al., 2018).

424   In the summertime LC-ESI-MS/MS measurements, CHN particle-phase compounds were

425   comprised primarily of amines (72% of CHN species contained an amine group) and nitriles

426   (28% of CHN species contained a nitrile group), as shown in Figure 4. In the winter, these

427   compounds were nearly exclusively amines (present in 99% of CHN species). Amines have

428   many primary land-based sources (e.g. biogenic emissions (Kieloaho et al., 2013), agricultural

429   activity (Ge et al., 2011), emissions from decomposing organic matter (Ge et al., 2011;

430   Sintermann and Neftel, 2015), biomass burning (Ge et al., 2011), emissions from port activity

431   (Gaston et al., 2013), chemical products (Khare and Gentner, 2018), and vehicle exhaust

432   (Sodeman et al., 2005)), but their presence on the coast could also indicate marine contributions.

433   Amines have been detected both in bulk ocean water, the surface microlayer, and in sea spray

434   aerosol, and their emissions and chemical transformations in the marine environment have been

435   the topic of many recent studies (e.g. Brean et al., 2021; Dall'Osto et al., 2019; Decesari et al.,

436   2020; Di Lorenzo et al., 2018; van Pinxteren et al., 2012, 2019; Quinn et al., 2015; Wu et al.,

437   2020).

438   Recent studies have also evaluated amine phase partitioning or formation in cloud/flog

439   water (e.g. Chen et al., 2018; Youn et al., 2015), as well as condensed-phase or aqueous-phase

440   pathways that may transform emitted amines (e.g. Ge et al., 2016; Lim et al., 2019; Tao et al.,

441   2021). Interestingly, the observed amines at this site, as well as other reduced nitrogen groups

442   like nitriles, imines, and enamines, were not present exclusively in CHN species and thus were a

443   mix of both direct emissions and chemically processed compounds. Reduced nitrogen groups

444   were often paired with hydroxyl groups, carboxylic acids, carbonyls, ethers, and esters as part of

445   nitrogen and oxygen containing compounds with a range of O/N ratios. As such, we discuss

446   CHON species as a function of O/N ratio to focus on differences between less-oxygenated (O/N

447   < 3) and more-oxygenated (O/N ≥ 3, e.g. organonitrates) species, using a ratio of 3 to distinguish

448   between the two as informed by the O/N ratio of the organonitrate functional group.

449   *3.3.2. CHON (O/N < 3) compounds*

450   CHON (O/N < 3) compounds were notably more important at this site than other sites,

451   representing 20% and 19% of observed functionalized organic aerosol abundance in summer and

452   winter, respectively (Figure 3A-B), compared to ~13% at other sites (from predominantly

453   summer measurements). These CHON compounds included some functional groups that

454   contained both oxygen and nitrogen, such as amide groups (12% of this compound class's

455   nitrogen content in summer, vs. 1% in winter, Figure 4) and nitro groups (15% of this nitrogen

456   content in summer, vs. 6% in winter, Figure 4). However, most CHON (O/N < 3) compounds

457   were comprised of a combination of nitrogen- *or* oxygen-containing groups, rather than a

458   functional group containing both nitrogen and oxygen. This included large contributions from

459   hydroxyls and ethers across both seasons, as well as important contributions from amines,

460   isocyanates, and heterocyclic nitrogen (Ditto et al., 2020). The presence of these functional

461   groups in the winter could be indicative of wood burning emissions in the region, which has been

462   observed in the wintertime in past ambient sampling in the Northeast U.S. (Sullivan et al., 2019).

463   Isocyanates contributed notably to this compound class during the winter, which could similarly

464   be linked to burning wood, other biomass, building materials (Leslie et al., 2019; Priestley et al.,

465   2018; Roberts et al., 2014), or could be photochemically produced via the oxidation of amines

466   and amides (Borduas et al., 2015; Leslie et al., 2019). Importantly, levoglucosan, a common

467   biomass burning tracer, was observed across nearly all daytime and nighttime winter particle-

468   phase samples (verified with an authentic standard), supporting the influence of biomass burning

469   compounds at the site. Together, the overall high prevalence of reduced nitrogen at this site

470   could be influenced by the mixing of aged biomass burning plumes with marine air, which is

471   consistent with past observations of very high alkylamine concentrations in biomass burning

472   particles that mixed with marine air prior to sampling (Di Lorenzo et al., 2018).

473   *3.3.3. CHON (O/N ≥ 3) compounds*

474   CHON (O/N ≥ 3) compounds were the dominant compound class in the observed

475   summertime distribution and played an important role in the wintertime distribution as well,

476   comprising 44% of observed functionalized organic aerosol abundance in summer vs. 24% in

477   winter (Figure 3A-B). These contributions were far greater than the contributions of CHON (O/N

478   ≥ 3) species at other sites, which typically ranged from 14-19% (predominantly from

479   summertime measurement, Figure 3C).

480   Similar to CHON (O/N < 3), we observed some CHON (O/N ≥ 3) compounds with

481   functional groups containing 3 oxygen atoms and 1 nitrogen atom, e.g. nitrophenols and

482   organonitrates (Figure 4), but also contributions from nitrogen-only functional groups paired

483   with oxygen-containing groups. Notably, in the summer, there were important contributions from

484   amines (47% of this compound class's nitrogen content), imines (19%), organonitrates (10%),

485   and azoles (16%) (Figure 4). In contrast, in the winter, nitrogen content in the CHON (O/N ≥ 3)

486   compound class was dominated by I/SVOC nitrophenols, comprising 64% of the CHON (O/N ≥

487   3) ion abundance.

488   NO_x mixing ratios were typically low in both summer and winter (2.3 ± 1.5 ppb in

489   summer vs. 3.7 ± 2.7 ppb in winter), but were slightly higher during winter. In the winter, CHON

490   (O/N ≥ 3) compounds showed a weak positive relationship with NO_x mixing ratios (r ~ 0.58) and

491   a stronger correlation with NO mixing ratios (r ~ 0.81). This relationship between CHON (O/N ≥

492   3) and NO (and NO_x) suggests that many of these oxidized nitrogen species were products of

493   NO_x-related chemistry (i.e. NO_z compounds). The enhancements in nitrophenols serves as one

494   example of this, as NO mixing ratios also correlated with the contribution of nitrophenols in the

495   winter (r ~ 0.69).

496   In past work, we discussed nitrophenol nighttime enhancements during winter, and noted

497   their reported aqueous formation pathways mentioned in prior laboratory studies (Ditto et al.,

498   2020). Here, we demonstrate that nitrophenols were important contributors to the CHON (O/N ≥

499   3) compound class, and highlight their role as examples of NO_z due to their possible formation

500   via dark aqueous-phase nitration pathways of oxygenated aromatics with ambient nitrous acid

501   (HONO) (Vidović et al., 2018). While nitrophenols may have other sources (e.g. diesel exhaust),

502   our observations of a clear nighttime enhancement during the winter suggest that these functional

503   groups were most likely formed by secondary chemistry related to NO_x oxidation, as this field

504   site was removed from major roadways. Our wintertime observations suggest that HONO could

505   have been derived from local wood burning, and could have reacted away as the smoke plume

506   aged to form stable products like nitrophenols, similar to HONO transformation chemistry into

507   other forms of oxidized nitrogen (e.g. particulate nitrates, PANs, organic nitrates) that has

508   recently been observed in wildfire smoke (Juncosa Calahorrano et al., 2021).

509   Furthermore, the correlation between NO and CHON (O/N ≥ 3) could also be influenced

510   by the daytime formation of organonitrates via reaction with OH$^\bullet$ and NO (i.e. RO$_2^\bullet$ + NO)

511   (Liebmann et al., 2019; Ng et al., 2017; Perring et al., 2013; Takeuchi and Ng, 2018), though

512   organonitrates contributed to a smaller fraction of CHON (O/N ≥ 3) species (i.e. 10% of this

513   compound class's nitrogen content across seasons).

514   *3.3.4. Overall contributions of reduced and oxidized nitrogen groups*

515   In the summer and winter, contributions from reduced nitrogen groups (e.g. groups

516   shown in black/grey in Figure 4) rivaled that of oxidized nitrogen groups in CHON compounds

517   across a range of O/N ratios. In the summer, reduced nitrogen groups contributed to 50% of all

518   detected CHON (O/N < 3) compounds by ion abundance, while in the winter they contributed

519   47% (Figure 4). For CHON compounds with O/N ≥ 3, reduced nitrogen groups contributed to

520   68% of compound ion abundance in the summer, while in the winter they contributed just 13%.

521   Interestingly, 90% of the dominant reduced nitrogen functional groups observed (amines and

522   imines) were present in acyclic rather than cyclic structures, which may have been the result of

523   either direct emissions or formation via reactions with ammonia or other small amines.

524   In contrast, possible NO$_z$ products (e.g. groups shown in blue in Figure 4) were present in

525   18% and 7% of CHON (O/N < 3) compounds in the summer and winter, respectively. For

526   CHON (O/N ≥ 3) compounds, they were present in 18% and 86% in the summer and winter,

527   respectively, with the latter wintertime increase in oxidized N-groups largely driven by the

528   presence of nitrophenols at night (Ditto et al., 2020). The remaining fraction of nitrogen-

529   containing groups also contained oxygen, but with a reduced nitrogen atom (e.g. amide,

530   isocyanate, nitrogen/oxygen-containing azole, shown in brown in Figure 4). We note that

531   CHONS compounds also represented a sizable fraction of observed organic nitrogen (Figure 3),

532   and contained a mix of reduced and oxidized functional groups (Section S3 and Figure S4).

533   The importance of reduced nitrogen functional groups in CHON compounds highlights

534   that not all oxygen- and nitrogen-containing species in the CHON (O/N ≥ 3) compound class

535   were NO_z, despite their apparent molecular formulas and the observed correlation observed

536   between CHON (O/N ≥ 3) species with NO and NO_x mixing ratios. For instance, many of the

537   observed reduced nitrogen-containing functional groups co-occurred with several oxygen-

538   containing groups like hydroxyls, carboxylic acids, esters, ethers, and carbonyls, and thus had

539   molecular formulas with O/N ≥ 3, which could incorrectly be assumed to be an organonitrate or

540   similar structure based on molecular formula alone.

541   We note that the relative distribution of reduced and oxidized nitrogen-containing groups

542   shown here is subject to sampling and ionization conditions. While the electrospray ionization

543   source used for the particle-phase analysis discussed here effectively ionized these nitrogen-

544   containing groups, their relative sensitivity may differ because many of these functional groups

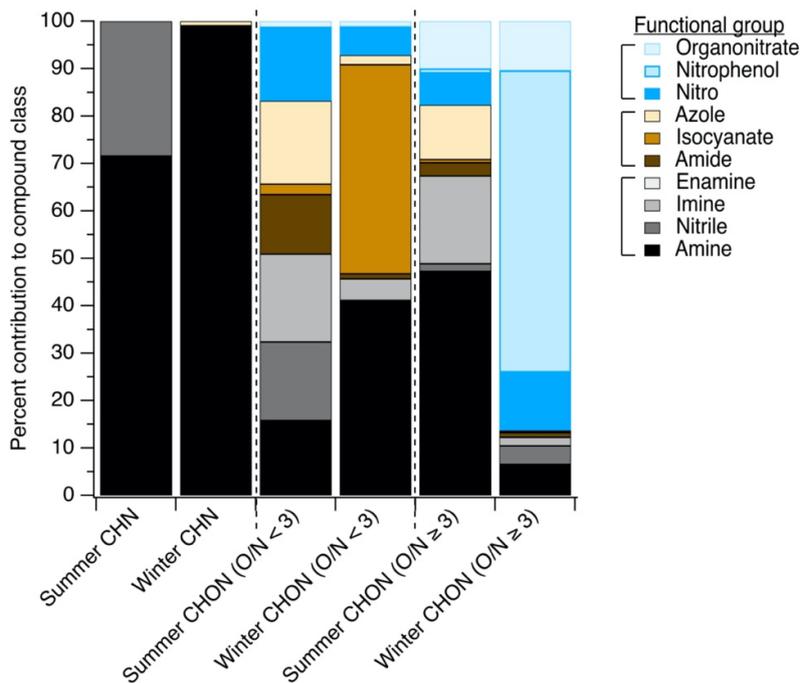
545   were present in multifunctional compounds whose other features may also contribute to

546   ionization behavior. Also, other aspects of the sample collection and extraction process could

547   cause variability in observed signal (e.g. PM size cut, organonitrate stability over long duration

548   samples). Thus, we emphasize that the observed relative abundances here are valuable because

549   they suggest that fully reduced nitrogen-containing groups are important contributors to

550   multifunctional CHON species, but their exact mass contributions remain uncertain.

551 **Figure 4.** The distribution of functional groups in particle-phase nitrogen-containing compounds
552 measured via LC-ESI-MS/MS. The breakdown of CHN, CHON (O/N < 3), and CHON (O/N ≥
553 3) compounds is shown as a function of contributions of each functional group to ion abundance,
554 with possible NO_z species shown in blue shades, fully-reduced nitrogen-containing groups
555 shown in black/grey shades, and groups containing both oxygen and nitrogen where the nitrogen
556 atom itself is not oxidized shown in brown shades.

557 3.4. Probing Possible Nitrogen-Containing Gas-Phase Precursors to Observed Nitrogen-

558 Containing Particles with Adsorptive Sampling and LC-ESI-MS

559   The particle-phase volatility distribution in the winter consisted of 18% IVOCs and 41%

560 SVOCs (Figure 2B). Of the observed compounds in winter, 68% contained nitrogen; these likely

561 included contributions from functionalized gas-phase precursors and likely were influenced by

562 the active multiphase partitioning of these precursors, and their gas- or particle-phase reaction

563 products, with changes in organic aerosol loading, atmospheric liquid water concentrations, and

564   temperature (Donahue et al., 2011; Ervens et al., 2011). This emphasizes the need to measure a

565   broader range of these functionalized gas-phase compounds, which have known limitations with

566   GC transmission, but represent uncertain and important-to-measure SOA precursors.

567   However, despite evidence of substantial contributions from particle-phase I/SVOCs with

568   diverse nitrogen-containing functionalities that could dynamically partition between phases

569   (Figure 2A-B), the observed compound class distribution from gas-phase adsorbent tube

570   measurements analyzed via GC-APCI-MS was dominated by hydrocarbons (i.e. CH, 24% of

571   detected ion abundance in summer vs. 18% in winter) and oxygenates (i.e. CHO, 66% in summer

572   vs. 69% in winter) (Figure S5). These gas-phase species appeared to be lightly functionalized

573   oxygenates (average O/C: 0.12 ± 0.13), showing minimal contributions from nitrogen (or sulfur)

574   heteroatoms; only 9% of detected ion abundance from gas-phase adsorbent tubes in summer and

575   11% in winter contained a nitrogen heteroatom. This is likely due to measurement limitations;

576   while GC-APCI techniques are extremely well-suited for the analysis of less functionalized

577   organic compounds from both instrument transmission and ionization efficiency perspectives,

578   these techniques are not as effective for more polar, more functionalized, more thermally-labile,

579   or otherwise less-GC-amenable species. Thus, to examine a broader range of functionalized gas-

580   phase compounds, we used an offline adsorptive sampling method on cooled PEEK tubing

581   collectors and inline mobile phase desorption for LC-ESI-MS analysis (Figure 1). CH and CHS

582   compound classes were excluded from this gas-phase LC-ESI-MS analysis due to their poor ESI

583   ionization efficiency.

584   Due to variations in trapping and desorption effectiveness (Section S1), this method was

585   not intended to be used as a quantitative measurement of concentration, but rather a relative

586   assessment of the distribution of nitrogen-containing gas-phase organic compounds. The

587   variation between analytes in breakthrough testing does not influence our conclusions about the

588   overall prevalence of observed gas-phase organic nitrogen. In laboratory tests, gas-phase sample

589   collection, inline desorption to the mobile phase, trapping on the LC column, and

590   chromatographic separation performed well. We observed limited breakthrough for most

591   analytes during sampling, effective focusing prior to LC analysis, and similar separations for

592   spiked collectors and breakthrough tests compared to standard LC runs (Figure 1B).

593   Results from the application of this new method at the YCFS revealed a wide range of

594   compounds with oxygen-, nitrogen-, and/or sulfur-containing functionality (Figure 5) that existed

595   at a lower average saturation mass concentration than the adsorbent tube methods during winter,

596   with a $\log(C_0)$ of 3.8 ± 2.6 $\mu g/m^3$ for adsorbent tubes analyzed with GC-APCI-MS compared to

597   0.6 ± 2.2 $\mu g/m^3$ for functionalized gases observed via LC-ESI-MS ($p < 0.05$). This decrease in

598   volatility corresponded to an increase in the average O/C ratio of these functionalized gases to

599   0.2 ± 0.2, which can partly be attributed to LC-ESI's poor ionization of CH compounds and to

600   the collection system's design (targeting heteroatom-containing species and not higher volatility

601   hydrocarbons). This may be a lower limit of O/C among functionalized compounds, as during

602   testing with a mixture of standards, we often observed poor retention of high O/C sugars like

603   xylitol and mannose on the LC analytical column (Table S2).

604   The gas-phase LC-ESI-MS data provide a valuable comparison to the wintertime

605   particle-phase samples analyzed using the same instrument. These particle-phase samples had

606   major contributions from CHO, CHON (O/N < 3 and O/N ≥ 3), CHONS, and CHOS compound

607   classes (Figure 3B). While not collected concurrently, the functionalized gas-phase samples in

608   winter had similar contributions from CHO (20%) and CHON (O/N ≥3) compounds (16%),

609   relatively more CHN (11%) and CHON (O/N < 3) (46%) compounds, and fewer CHONS (2.7%)

610   and CHOS (4.4%) compounds (Figure 5A). The prevalence of gas-phase CHN, CHON (O/N <

611   3), and CHON (O/N ≥ 3) is of particular interest given the abundance of CHON compounds

612   observed in the particle phase, and the potential of these gases to partition to the particle-phase

613   and/or act as reactive precursors to other oxidized nitrogen-containing species.

614   The presence of these nitrogen-containing compounds in the gas-phase also suggests that

615   these compound classes observed in the particle-phase at least partly originated in the gas-phase

616   and partitioned, rather than formed exclusively as a result of particle-phase chemistry. These

617   species could have also formed in the particle-phase and partitioned to the gas-phase with or

618   without condensed-phase fragmentation (discussed above). In either scenario, these nitrogen-

619   containing compounds likely actively partitioned between phases due to their volatility (e.g.

620   I/SVOCs shown in Figure 5A). Also, their polarity and high Henry's Law coefficients (relative

621   to non-functionalized hydrocarbons (Sander, 2015)) suggests that these compounds could have

622   been readily taken up by the aqueous phase. To check that these compounds were indeed gas-

623   phase species under ambient conditions, we predicted the saturation mass concentration for

624   individual compounds using individual ion formulas and estimated their gas-particle partitioning

625   to a pre-existing condensed phase. While the range of compounds in Figure 5A can be expected

626   to dynamically partition, the results confirm that the overall suite of observed compounds would

627   have predominately existed as gases, with on average ~80% of observed ion abundance predicted

628   to equilibrate to the gas-phase across compound classes (Figure S6).

629   Of all the gas-phase species observed with at least one nitrogen atom (i.e. CHN, CHON,

630   CHONS, CHNS), we note that 78% of these compounds had an O/N ratio of less than 3 (Figure

631   5B), indicating that most of these gas-phase species were not organonitrates, nitrophenols, or

632   other similar structures. This is similar to our particle-phase results, which showed important

633   contributions from reduced nitrogen-containing groups paired with oxygen-containing groups in

634   CHON (O/N < 3) compounds. Notably, we observed an 11% contribution of gas-phase CHN

635   species with this gas-phase LC-ESI-MS method (Figure 5A), in contrast to 2% CHN in the

636   wintertime particle-phase samples (Figure 3). In the winter particle-phase samples, most CHN

637   compounds contained amines (discussed above), and thus we postulate that these functionalized

638   gas-phase CHN species were possibly also amines that acted as precursors to observed nitrogen-

639   containing particle-phase compounds following oxidation and partitioning (or vice versa).

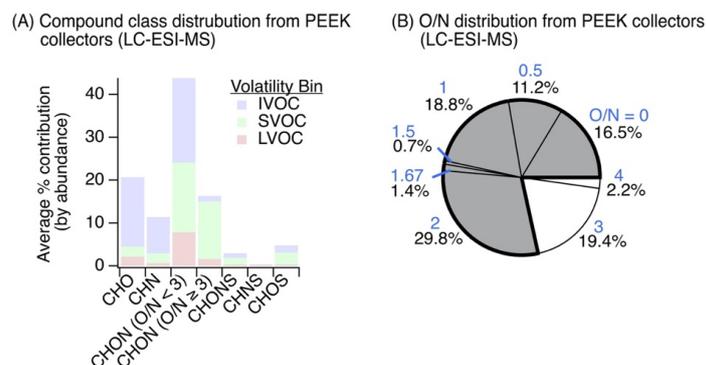
640   The substantial contribution from CHON with O/N < 3 (46%) to the functionalized gas-

641   phase samples could be linked to less photochemical processing of CHON compounds relative to

642   the particle phase and/or the emissions/oxidation of CHN or CHON compounds. Moreover, in

643   the particle-phase, we observed a weak negative relationship between CHN contribution and

644   hydroxyl group prevalence in summertime measurements (r~ -0.57), which may support the

645   transformation of CHN to CHON compounds via the formation of hydroxyl-containing species.

646   **Figure 5.** Observations of gas-phase nitrogen-containing compounds. (A) The distribution of
647   functionalized gases observed via sampling on PEEK collectors (N = 6) and inline mobile phase
648   desorption with non-targeted LC-ESI-MS analysis contained a diversity of oxygen-, nitrogen-,
649   and/or sulfur-containing compounds in the IVOC-LVOC range (volatility assignment and
650   grouping was the same as discussed in Figure 2). While we cannot rule out gas-phase LVOC
651   contributions from evaporation off of the upstream particle filter, LVOC contributions were
652   limited (~12%). (B) Oxygen-to-nitrogen (O/N) ratio distribution of observed gas-phase nitrogen-
653   containing species where O/N ratios < 3 are colored grey and O/N ratios ≥ 3 are white (blue text
654   above each percentage signifies the O/N ratio).

4. Conclusions and Opportunities for Future Research

655

656   Together, these results suggest that a mix of direct emissions and chemical processes

657   during summer and winter in the Long Island Sound region resulted in a diverse mixture of

658   multifunctional gases and particles, where more than two-thirds of observed particle-phase

659   compounds contained at least one nitrogen atom.

660   The observed nitrogen-containing functional groups existed across a range of fully

661   reduced (e.g. amines, imines) and oxidized (e.g. nitro, organonitrate) structures. These fully

662   reduced nitrogen functional groups were prevalent across all nitrogen-containing compound

663   classes, including CHON species, and we highlight their importance as contributors to these

664   multifunctional compounds beyond typical NO_z-type compounds that are commonly studied

665   using online mass spectrometers and share similar CHON molecular formulas. For instance,

666   these gas- and particle-phase measurements of nitrogen-containing compounds are

667   complementary to the measurements of these species made by chemical ionization mass

668   spectrometers (CIMS) or by proton transfer reaction mass spectrometers (PTR-MS), whose

669   ionization mechanisms can be tuned for sensitivity towards functionalized compounds of interest

670   (Riva et al., 2019). While online mass spectrometers excel at high time resolution measurements

671   that capture dynamic chemical processes in the atmosphere, their mass resolution is typically

672   lower and they normally do not utilize separations, so they largely depend on parent ion mass-to-

673   charge ratios to assign molecular formulas without structural attribution. The offline methods

674   used here cannot match the time resolution of online techniques. However, the use of

675   chromatography to separate isomers, longer sampling times to increase sensitivity toward a

676   greater range of compounds, and the use of higher resolution mass spectrometers with MS/MS

677   capabilities allow for improved compound identification and determination of functional group

678 distribution at the molecular level. This enables us to distinguish between true NO_z species and

679 those that contain combinations of nitrogen and oxygen but are not NO_x oxidation products.

680 Thus, both these online and offline methods should be employed together to differentiate a wider

681 range of nitrogen-containing species and to achieve both temporal and chemical resolution.

682   As discussed throughout this work, the Long Island Sound region is affected by a mixture

683 of anthropogenic, biogenic, and marine sources, all of which contain known emitters of organic

684 nitrogen. Understanding the combined effect of these individual sources and their chemical

685 transformations will be important in regions like the Long Island Sound, where a significant

686 degree of mixing occurs over the Sound before air parcels arrive inland. For example, past work

687 has noted extremely high contributions from alkylamines in biomass burning-influenced air

688 mixed with marine air (Di Lorenzo et al., 2018). Similar enhancements could be expected when

689 mixing other prominent sources of amines with marine air, such as in the aging urban outflow

690 from the Central Atlantic and Northeast U.S., which may be transported up the coast and impact

691 states in the surrounding region.

692   As with any ambient site, these mixed emissions are chemically processed in the

693 atmosphere via a multitude of pathways. Here, we observed evidence of photochemical and

694 aqueous processes occurring in both seasons, but in the winter we observed various mixture-wide

695 trends that suggested an enhanced role for aqueous-phase processing. These observations

696 included higher overall particle-phase volatility and smaller carbon backbone sizes, which may

697 indicate a more important role for aqueous-phase fragmentation reactions or aqueous uptake of

698 water soluble gases (Brege et al., 2018). We also observed key marker functional groups that

699 may be formed via aqueous phase chemistry (e.g. nitrophenols, azoles). The role of aqueous-

700 phase chemistry and aqueous-phase uptake of gases is increasingly studied in laboratory and

701   ambient contexts (Herrmann et al., 2015), and such chemistry should further examined especially

702   in coastal and other humid regions.

703   For example, the aqueous-phase processing of atmospherically relevant nitrogen-

704   containing species is particularly important to understand in ambient air due to the potential of

705   brown carbon formation, which has significant impacts on climate forcing (Laskin et al., 2015).

706   The role of ammonia and amines reacting with carbonyls is of interest for this type of chemistry

707   (e.g. DeHaan et al., 2009; Grace et al., 2020; McNeill, 2015; Sareen et al., 2010) and should

708   continue to be explored, particularly in coastal settings where concentrations of small gas-phase

709   amines may be high due to their marine sources. As discussed above, our ambient observations

710   of azoles could be indicative of such chemistry, and should be explored in future comparisons of

711   ambient and laboratory-generated species. Also, we observed a significant contribution from

712   nitrophenols at our site, and while they are not formed by this same chemistry, they represent

713   another important form of light absorbing nitrogen-containing organic mass in the atmosphere

714   (Hems and Abbatt, 2018). Finally, many of the nitrogen-containing functional groups observed

715   in this work may be susceptible to hydrolysis, so the balance between hydrolysis and other

716   aqueous pathways is important to consider and understand for appropriate representation of

717   nitrogen-containing compounds in models for both aqueous aerosol and in-cloud/fog chemistry.

718   As another example, the greater prevalence of I/SVOCs observed in the winter particle-

719   phase samples suggested possible dynamic partitioning or aqueous uptake of lighter gas-phase

720   compounds; to explore the composition of these lighter gas-phase compounds that could exist as

721   I/SVOCs and thus participate in phase partitioning, we supplemented our particle-phase analyses

722   with a novel approach for investigating functionalized gases with LC-ESI-MS. Further

723   investigation of these nitrogen-containing gases will facilitate new understanding of their gas-

748   **Competing interests:** The authors declare that they have no substantive conflicts of interest but

749   acknowledge that D.R.G. is an associate editor with *Atmospheric Chemistry and Physics*.

750

751   **Acknowledgements:** We acknowledge support from the National Science Foundation

752   (AGS1764126) and the Yale Natural Lands Program. We also acknowledge GERSTEL for their

753   collaboration with the thermal desorption unit used here. We thank David Wheeler at the New

754   York Department of Environmental Conservation, Pete Babich and Adam Augustine at the

755   Connecticut Department of Energy and Environmental Protection, Luke Valin at the

756   Environmental Protection Agency, and Jordan Peccia at Yale for use of sampling equipment, as

757   well as Paul Miller (NESCAUM) for organizing the LISTOS project. We also thank Richard

758   Boardman and the Yale Peabody Museum for enabling us to set up and collect samples at the

759   YCFS site, and acknowledge the help of Yale Peabody Museum EVOLUTIONS interns: Amir

760   Bond, Ethan Weed, Paula Mock, and Aurea Orencia. We thank Trevor VandenBoer and Barbara

761   Ervens for helpful feedback on the draft manuscript. Finally, we acknowledge the NOAA Air

762   Resources Laboratory (ARL) for the provision of the HYSPLIT transport and dispersion model.

763

References

766   Borduas, N., Da Silva, G., Murphy, J. G. and Abbatt, J. P. D.: Experimental and theoretical understanding of the gas
767   phase oxidation of atmospheric amides with OH radicals: Kinetics, products, and mechanisms, J. Phys. Chem. A,
768   119(19), 4298–4308, doi:10.1021/jp503759f, 2015.

770   Brean, J., Dall'Osto, M., Simó, R., Shi, Z., Beddows, D. C. S. and Harrison, R. M.: Open ocean and coastal new
771   particle formation from sulfuric acid and amines around the Antarctic Peninsula, Nat. Geosci., 14(6), 383–388,
772   doi:10.1038/s41561-021-00751-y, 2021.

774   Brege, M., Paglione, M., Gilardoni, S., Decesari, S., Facchini, M. C. and Mazzoleni, L. R.: Molecular insights on
775   aging and aqueous-phase processing from ambient biomass burning emissions-influenced Po Valley fog and aerosol,
776   Atmos. Chem. Phys., 18(17), 13197–13214, doi:10.5194/acp-18-13197-2018, 2018.

778   Chen, C. L., Chen, S., Russell, L. M., Liu, J., Price, D. J., Betha, R., Sanchez, K. J., Lee, A. K. Y., Williams, L.,
779   Collier, S. C., Zhang, Q., Kumar, A., Kleeman, M. J., Zhang, X. and Cappa, C. D.: Organic Aerosol Particle
780   Chemical Properties Associated With Residential Burning and Fog in Wintertime San Joaquin Valley (Fresno) and

781 With Vehicle and Firework Emissions in Summertime South Coast Air Basin (Fontana), J. Geophys. Res. Atmos.,
782 123(18), 10707–10731, doi:10.1029/2018JD028374, 2018.
783
784 Cleveland, W. S., Kleiner, B., McRae, J. E. and Warner, J. L.: Photochemical air pollution: Transport from the New
785 York City area into Connecticut and Massachusetts, Science., 191(4223), 179–181, doi:10.1126/science.1246603,
786 1976.
787
788 Dall'Osto, M., Airs, R. L., Beale, R., Cree, C., Fitzsimons, M. F., Beddows, D., Harrison, R. M., Ceburnis, D.,
789 O'Dowd, C., Rinaldi, M., Paglione, M., Nenes, A., Decesari, S. and Simó, R.: Simultaneous Detection of
790 Alkylamines in the Surface Ocean and Atmosphere of the Antarctic Sympagic Environment, ACS Earth Sp. Chem.,
791 3(5), 854–862, doi:10.1021/acsearthspacechem.9b00028, 2019.
792
793 Decesari, S., Paglione, M., Rinaldi, M., Dall'osto, M., Simó, R., Zanca, N., Volpi, F., Cristina Facchini, M.,
794 Hoffmann, T., Götz, S., Johannes Kampf, C., O'Dowd, C., Ceburnis, D., Ovadnevaite, J. and Tagliavini, E.:
795 Shipborne measurements of Antarctic submicron organic aerosols: An NMR perspective linking multiple sources
796 and bioregions, Atmos. Chem. Phys., 20(7), 4193–4207, doi:10.5194/acp-20-4193-2020, 2020.
797
798 DeHaan, D. O., Corrigan, A. L., Smith, K. W., Stroik, D. R., Turley, J. J., Lee, F. E., Tolbert, M. A., Jimenez, J. L.,
799 Cordova, K. E., Ferrell, G. R., Haan, D. O. De, Corrigan, A. L., Smith, K. W., Stroik, D. R., Turley, J. J., Lee, F. E.,
800 Tolbert, M. A., Jimenez, J. L., Cordova, K. E. and Ferrell, G. R.: Secondary organic aerosol-forming reactions of
801 glyoxal with amino acids, Environ. Sci. Technol., 43(8), 2818–2824, doi:10.1021/es803534f, 2009.
802
803 Di, Q., Wang, Y., Zanobetti, A., Wang, Y., Koutrakis, P., Choirat, C., Dominici, F. and Schwartz, J. D.: Air
804 pollution and mortality in the medicare population, N. Engl. J. Med., 376(26), 2513–2522,
805 doi:10.1056/NEJMoa1702747, 2017.
806
807 Ditto, J. C., Barnes, E. B., Khare, P., Takeuchi, M., Joo, T., Bui, A. A. T., Lee-Taylor, J., Eris, G., Chen, Y.,
808 Aumont, B., Jimenez, J. L., Ng, N. L., Griffin, R. J. and Gentner, D. R.: An omnipresent diversity and variability in
809 the chemical composition of atmospheric functionalized organic aerosol, Commun. Chem., 1(1), 75,
810 doi:10.1038/s42004-018-0074-3, 2018.
811
812 Ditto, J. C., Joo, T., Khare, P., Sheu, R., Takeuchi, M., Chen, Y., Xu, W., Bui, A. A. T., Sun, Y., Ng, N. L. and
813 Gentner, D. R.: Effects of Molecular-Level Compositional Variability in Organic Aerosol on Phase State and
814 Thermodynamic Mixing Behavior, Environ. Sci. Technol., 53(22), 13009–13018, doi:10.1021/acs.est.9b02664,
815 2019.
816
817 Ditto, J. C., Joo, T., Slade, J. H., Shepson, P. B., Ng, N. L. and Gentner, D. R.: Nontargeted Tandem Mass
818 Spectrometry Analysis Reveals Diversity and Variability in Aerosol Functional Groups across Multiple Sites,
819 Seasons, and Times of Day, Environ. Sci. Technol. Lett., 7(2), 60–69, doi:10.1021/acs.estlett.9b00702, 2020.
820
821 Ditto, J. C., He, M., Hass-Mitchell, T. N., Moussa, S. G., Hayden, K., Li, S.-M., Liggio, J., Leithead, A., Lee, P.,
822 Wheeler, M. J., Wentzell, J. J. B. and Gentner, D. R.: Atmospheric Evolution of Emissions from a Boreal Forest
823 Fire: The Formation of Highly-Functionalized Oxygen-, Nitrogen-, and Sulfur-Containing Compounds, Atmos.
824 Chem. Phys., 21, 255–267, 2021.
825
826 Donahue, N. M., Epstein, S. A., Pandis, S. N. and Robinson, A. L.: A two-dimensional volatility basis set: 1.
827 organic-aerosol mixing thermodynamics, Atmos. Chem. Phys., 11(7), 3303–3318, doi:10.5194/acp-11-3303-2011,
828 2011.
829
830 Dührkop, K., Shen, H., Meusel, M., Rousu, J. and Böcker, S.: Searching molecular structure databases with tandem
831 mass spectra using CSI:FingerID, Proc. Natl. Acad. Sci., 112(41), 12580–12585, doi:10.1073/pnas.1509788112,
832 2015.
833
834 Dührkop, K., Fleischauer, M., Ludwig, M., Aksenov, A. A., Melnik, A. V., Meusel, M., Dorrestein, P. C., Rousu, J.
835 and Böcker, S.: SIRIUS 4: a rapid tool for turning tandem mass spectra into metabolite structure information, Nat.
836 Methods, 16(4), 299–302, doi:10.1038/s41592-019-0344-8, 2019.

Dunlea, E. J., Herndon, S. C., Nelson, D. D., Volkamer, R. M., San Martini, F., Sheehy, P. M., Zahniser, M. S., Shorter, J. H., Wormhoudt, J. C., Lamb, B. K., Allwine, E. J., Gaffney, J. S., Marley, N. A., Grutter, M., Marquez, C., Blanco, S., Cardenas, B., Retama, A., Ramos Villegas, C. R., Kolb, C. E., Molina, L. T. and Molina, M. J.: Evaluation of nitrogen dioxide chemiluminescence monitors in a polluted urban environment, Atmos. Chem. Phys., 7(10), 2691–2704, doi:10.5194/acp-7-2691-2007, 2007.

Ervens, B., Turpin, B. J. and Weber, R. J.: Secondary organic aerosol formation in cloud droplets and aqueous particles (aqSOA): A review of laboratory, field and model studies, Atmos. Chem. Phys., 11(21), 11069–11102, doi:10.5194/acp-11-11069-2011, 2011.

Gaston, C. J., Quinn, P. K., Bates, T. S., Gilman, J. B., Bon, D. M., Kuster, W. C. and Prather, K. A.: The impact of shipping, agricultural, and urban emissions on single particle chemistry observed aboard the R/V Atlantis during CalNex, J. Geophys. Res. Atmos., 118(10), 5003–5017, doi:10.1002/jgrd.50427, 2013.

Ge, X., Wexler, A. S. and Clegg, S. L.: Atmospheric amines - Part I. A review, Atmos. Environ., 45(3), 524–546, doi:10.1016/j.atmosenv.2010.10.012, 2011.

Ge, Y., Liu, Y., Chu, B., He, H., Chen, T., Wang, S., Wei, W. and Cheng, S.: Ozonolysis of Trimethylamine Exchanged with Typical Ammonium Salts in the Particle Phase, Environ. Sci. Technol., 50(20), 11076–11084, doi:10.1021/acs.est.6b04375, 2016.

de Gouw, J. A., Middlebrook, A. M., Warneke, C., Goldan, P. D., Kuster, W. C., Roberts, J. M., Fehsenfeld, F. C., Worsnop, D. R., Canagaratna, M. R., Pszenny, A. A. P., Keene, W. C., Marchewka, M., Bertman, S. B. and Bates, T. S.: Budget of organic carbon in a polluted atmosphere: Results from the New England Air Quality Study in 2002, J. Geophys. Res. D Atmos., 110(16), 1–22, doi:10.1029/2004JD005623, 2005.

Grace, D. N., Sharp, J. R., Holappa, R. E., Lugos, E. N., Sebold, M. B., Griffith, D. R., Hendrickson, H. P. and Galloway, M. M.: Heterocyclic Product Formation in Aqueous Brown Carbon Systems, ACS Earth Sp. Chem., 3(11), 2472–2481, doi:10.1021/acsearthspacechem.9b00235, 2019.

Grace, D. N., Lugos, E. N., Ma, S., Griffith, D. R., Hendrickson, H. P., Woo, J. L. and Galloway, M. M.: Brown Carbon Formation Potential of the Biacetyl-Ammonium Sulfate Reaction System, ACS Earth Sp. Chem., 4(7), 1104–1113, doi:10.1021/acsearthspacechem.0c00096, 2020.

Hallquist, M., Wenger, J. C., Baltensperger, U., Rudich, Y., Simpson, D., Claeys, M., Dommen, J., Donahue, N. M., George, C., Goldstein, A. H., Hamilton, J. F., Herrmann, H., Hoffmann, T., Iinuma, Y., Jang, M., Jenkin, M. E., Jimenez, J. L., Kiendler-Scharr, A., Maenhaut, W., McFiggans, G., Mentel, T. F., Monod, A., Prevot, A. S. H., Seinfeld, J. H., Surratt, J. D., Szmigielski, R. and Wildt, J.: The formation, properties and impact of secondary organic aerosol: current and emerging issues, Atmos. Chem. Phys., 9(14), 5155–5236, doi:10.5194/acp-9-5155-2009, 2009.

Hems, R. F. and Abbatt, J. P. D.: Aqueous Phase Photo-oxidation of Brown Carbon Nitrophenols: Reaction Kinetics, Mechanism, and Evolution of Light Absorption, ACS Earth Sp. Chem., 2(3), 225–234, doi:10.1021/acsearthspacechem.7b00123, 2018.

Herrmann, H., Schaefer, T., Tilgner, A., Styler, S. A., Weller, C., Teich, M. and Otto, T.: Tropospheric Aqueous-Phase Chemistry: Kinetics, Mechanisms, and Its Coupling to a Changing Gas Phase, Chem. Rev., 115(10), 4259–4334, doi:10.1021/cr500447k, 2015.

Jerrett, M., Burnett, R. T., Pope, C. A., Ito, K., Thurston, G., Krewski, D., Shi, Y., Calle, E. and Thun, M.: Long-term ozone exposure and mortality, N. Engl. J. Med., 360(11), 1085–1095, doi:10.1056/NEJMoa0803894, 2009.

Juncosa Calahorrano, J. F., Lindaas, J., O'Dell, K., Palm, B. B., Peng, Q., Flocke, F., Pollack, I. B., Garofalo, L. A., Farmer, D. K., Pierce, J. R., Collett, J. L., Weinheimer, A., Campos, T., Hornbrook, R. S., Hall, S. R., Ullmann, K., Pothier, M. A., Apel, E. C., Permar, W., Hu, L., Hills, A. J., Montzka, D., Tyndall, G., Thornton, J. A. and Fischer,

893   E. V.: Daytime Oxidized Reactive Nitrogen Partitioning in Western U.S. Wildfire Smoke Plumes, J. Geophys. Res.
894   Atmos., 126(4), 1–22, doi:10.1029/2020JD033484, 2021.
895
896   Khare, P. and Gentner, D. R.: Considering the future of anthropogenic gas-phase organic compound emissions and
897   the increasing influence of non-combustion sources on urban air quality, Atmos. Chem. Phys., 18(8), 5391–5413,
898   doi:10.5194/acp-18-5391-2018, 2018.
899
900   Khare, P., Marcotte, A., Sheu, R., Walsh, A. N., Ditto, J. C. and Gentner, D. R.: Advances in offline approaches for
901   trace measurements of complex organic compound mixtures via soft ionization and high-resolution tandem mass
902   spectrometry, J. Chromatogr. A, 1598, 163–174, doi:10.1016/j.chroma.2019.03.037, 2019.
903
904   Kieloaho, A. J., Hellén, H., Hakola, H., Manninen, H. E., Nieminen, T., Kulmala, M. and Pihlatie, M.: Gas-phase
905   alkylamines in a boreal Scots pine forest air, Atmos. Environ., 80, 369–377, doi:10.1016/j.atmosenv.2013.08.019,
906   2013.
907
908   Kilian, J. and Kitazawa, M.: The emerging risk of exposure to air pollution on cognitive decline and Alzheimer's
909   disease – Evidence from epidemiological and animal studies, Biomed. J., 41(3), 141–162,
910   doi:10.1016/j.bj.2018.06.001, 2018.
911
912   Laskin, A., Laskin, J. and Nizkorodov, S. A.: Chemistry of Atmospheric Brown Carbon, Chem. Rev., 115(10),
913   4335–4382, doi:10.1021/cr5006167, 2015.
914
915   Leslie, M. D., Ridoli, M., Murphy, J. G. and Borduas-Dedekind, N.: Isocyanic acid (HNCO) and its fate in the
916   atmosphere: A review, Environ. Sci. Process. Impacts, 21(5), 793–808, doi:10.1039/c9em00003h, 2019.
917
918   Li, Y., Pöschl, U. and Shiraiwa, M.: Molecular corridors and parameterizations of volatility in the chemical
919   evolution of organic aerosols, Atmos. Chem. Phys., 16(5), 3327–3344, doi:10.5194/acp-16-3327-2016, 2016.
920
921   Liebmann, J., Sobanski, N., Schuladen, J., Karu, E., Hellén, H., Hakola, H., Zha, Q., Ehn, M., Riva, M., Heikkinen,
922   L., Williams, J., Fischer, H., Lelieveld, J. and Crowley, J. N.: Alkyl nitrates in the boreal forest: formation via the
923   NO3-, OH- and O3-induced oxidation of biogenic volatile organic compounds and ambient lifetimes, Atmos. Chem.
924   Phys., 19, 10391–10403, doi:10.5194/acp-2019-463, 2019.
925
926   Lim, S., McArdell, C. S. and von Gunten, U.: Reactions of aliphatic amines with ozone: Kinetics and mechanisms,
927   Water Res., 157, 514–528, doi:10.1016/j.watres.2019.03.089, 2019.
928
929   Lim, Y. Bin, Kim, H., Kim, J. Y. and Turpin, B. J.: Photochemical organonitrate formation in wet aerosols, Atmos.
930   Chem. Phys., 16(19), 12631–12647, doi:10.5194/acp-16-12631-2016, 2016.
931
932   Lin, P., Laskin, J., Nizkorodov, S. A. and Laskin, A.: Revealing Brown Carbon Chromophores Produced in
933   Reactions of Methylglyoxal with Ammonium Sulfate, Environ. Sci. Technol., 49(24), 14257–14266,
934   doi:10.1021/acs.est.5b03608, 2015.
935
936   Di Lorenzo, R. A., Place, B. K., VandenBoer, T. C. and Young, C. J.: Composition of Size-Resolved Aged Boreal
937   Fire Aerosols: Brown Carbon, Biomass Burning Tracers, and Reduced Nitrogen, ACS Earth Sp. Chem., 2(3), 278–
938   285, doi:10.1021/acsearthspacechem.7b00137, 2018.
939
940   Loza, C. L., Craven, J. S., Yee, L. D., Coggon, M. M., Schwantes, R. H., Shiraiwa, M., Zhang, X., Schilling, K. A.,
941   Ng, N. L., Canagaratna, M. R., Ziemann, P. J., Flagan, R. C. and Seinfeld, J. H.: Secondary organic aerosol yields of
942   12-carbon alkanes, Atmos. Chem. Phys., 14(3), 1423–1439, doi:10.5194/acp-14-1423-2014, 2014.
943
944   McNeill, V. F.: Aqueous organic chemistry in the atmosphere: Sources and chemical processing of organic aerosols,
945   Environ. Sci. Technol., 49(3), 1237–1244, doi:10.1021/es5043707, 2015.
946
947   Ng, N. L., Brown, S. S., Archibald, A. T., Atlas, E., Cohen, R. C., Crowley, J. N., Day, D. A., Donahue, N. M., Fry,
948   J. L., Fuchs, H., Griffin, R. J., Guzman, M. I., Herrmann, H., Hodzic, A., Iinuma, Y., Kiendler-Scharr, A., Lee, B.

H., Luecken, D. J., Mao, J., McLaren, R., Mutzel, A., Osthoff, H. D., Ouyang, B., Picquet-Varrault, B., Platt, U., Pye, H. O. T., Rudich, Y., Schwantes, R. H., Shiraiwa, M., Stutz, J., Thornton, J. A., Tilgner, A., Williams, B. J. and Zaveri, R. A.: Nitrate radicals and biogenic volatile organic compounds: oxidation, mechanisms, and organic aerosol, Atmos. Chem. Phys., 17(3), 2103–2162, doi:10.5194/acp-17-2103-2017, 2017.

Perring, A. E., Pusede, S. E. and Cohen, R. C.: An Observational Perspective on the Atmospheric Impacts of Alkyl and Multifunctional Nitrates on Ozone and Secondary Organic Aerosol, Chem. Rev., 113(8), 5848–5870, doi:10.1021/cr300520x, 2013.

van Pinxteren, M., Müller, C., Iinuma, Y., Stolle, C. and Herrmann, H.: Chemical Characterization of Dissolved Organic Compounds from Coastal Sea Surface Microlayers (Baltic Sea, Germany), Environ. Sci. Technol., 46(19), 10455–10462, doi:10.1021/es204492b, 2012.

van Pinxteren, M., Fomba, K. W., van Pinxteren, D., Triesch, N., Hoffmann, E. H., Cree, C. H. L., Fitzsimons, M. F., von Tümpling, W. and Herrmann, H.: Aliphatic amines at the Cape Verde Atmospheric Observatory: Abundance, origins and sea-air fluxes, Atmos. Environ., 203, 183–195, doi:10.1016/j.atmosenv.2019.02.011, 2019.

Pope, C. A. and Dockery, D. D.: Health Effects of Fine Particulate Air Pollution: Lines that Connect, J. Air Waste Manage. Assoc., 56, 709–742, doi:10.1080/10473289.2006.10464485, 2006.

Priestley, M., Le Breton, M., Bannan, T. J., Leather, K. E., Bacak, A., Reyes-Villegas, E., De Vocht, F., Shallcross, B. M. A., Brazier, T., Khan, M. A., Allan, J., Shallcross, D. E., Coe, H. and Percival, C. J.: Observations of Isocyanate, Amide, Nitrate, and Nitro Compounds From an Anthropogenic Biomass Burning Event Using a ToF-CIMS, J. Geophys. Res. Atmos., 123, 7687–7704, doi:10.1002/2017JD027316, 2018.

Quinn, P. K., Collins, D. B., Grassian, V. H., Prather, K. A. and Bates, T. S.: Chemistry and Related Properties of Freshly Emitted Sea Spray Aerosol, Chem. Rev., 115(10), 4383–4399, doi:10.1021/cr500713g, 2015.

Riva, M., Rantala, P., Krechmer, J. E., Peräkylä, O., Zhang, Y., Heikkinen, L., Garmash, O., Yan, C., Kulmala, M., Worsnop, D. and Ehn, M.: Evaluating the performance of five different chemical ionization techniques for detecting gaseous oxygenated organic species, Atmos. Meas. Tech., 12, 2403–2421, doi:https://doi.org/10.5194/amt-12-2403-2019, 2019.

Roberts, J. M., Veres, P. R., VandenBoer, T. C., Warneke, C., Graus, M., Williams, E. J., Lefer, B., Brock, C. A., Bahreini, R., Ozturk, F., Middlebrook, A. M., Wagner, N. L., Dube, W. P. and DeGouw, J. A.: New insights into atmospheric sources and sinksof isocyanic acid, HNCO, from recent urbanand regional observations, J. Geophys. Res. Atmos., 119, 1060–1072, doi:10.1002/2013JD019931, 2014.

Rogers, H. M., Ditto, J. C. and Gentner, D. R.: Evidence for impacts on surface-level air quality in the northeastern US from long-distance transport of smoke from North American fires during the Long Island Sound Tropospheric Ozone Study (LISTOS) 2018, Atmos. Chem. Phys., 20(2), 671–682, doi:10.5194/acp-20-671-2020, 2020.

Sander, R.: Compilation of Henry's law constants (version 4.0) for water as solvent, Atmos. Chem. Phys., 15(8), 4399–4981, doi:10.5194/acp-15-4399-2015, 2015.

Sareen, N., Schwier, A. N., Shapiro, E. L., Mitroo, D. and Mcneill, V. F.: Secondary organic material formed by methylglyoxal in aqueous aerosol mimics, Atmos. Chem. Phys., 10, 997–1016, doi:10.5194/acpd-9-15567-2009, 2010.

Schervish, M. and Donahue, N. M.: Peroxy radical chemistry and the volatility basis set, Atmos. Chem. Phys., 20(2), 1183–1199, doi:10.5194/acp-20-1183-2020, 2020.

Schroder, J. C., Campuzano-Jost, P., Day, D. A., Shah, V., Larson, K., Sommers, J. M., Sullivan, A. P., Campos, T., Reeves, J. M., Hills, A., Hornbrook, R. S., Blake, N. J., Scheuer, E., Guo, H., Fibiger, D. L., McDuffie, E. E., Hayes, P. L., Weber, R. J., Dibb, J. E., Apel, E. C., Jaeglé, L., Brown, S. S., Thornton, J. A. and Jimenez, J. L.: Sources and Secondary Production of Organic Aerosols in the Northeastern United States during WINTER, J. Geophys. Res.

1005 Atmos., 123(14), 7771–7796, doi:10.1029/2018JD028475, 2018.
1006
1007 Sheu, R., Marcotte, A., Khare, P., Charan, S., Ditto, J. C. and Gentner, D. R.: Advances in offline approaches for
1008 chemically speciated measurements of trace gas-phase organic compounds via adsorbent tubes in an integrated
1009 sampling-to-analysis system, J. Chromatogr. A, 1575, 80–90, doi:10.1016/j.chroma.2018.09.014, 2018.
1010 Sintermann, J. and Neftel, A.: Ideas and perspectives: On the emission of amines from terrestrial vegetation in the
1011 context of new atmospheric particle formation, Biogeosciences, 12(11), 3225–3240, doi:10.5194/bg-12-3225-2015,
1012 2015.
1013
1014 Sodeman, D. A., Toner, S. M. and Prather, K. A.: Determination of Single Particle Mass Spectral Signatures from
1015 Light-Duty Vehicle Emissions, Environ. Sci. Technol., 39(12), 4569–4580, doi:10.1016/j.atmosenv.2007.01.025,
1016 2005.
1017
1018 Sullivan, A. P., Guo, H., Schroder, J. C., Campuzano-Jost, P., Jimenez, J. L., Campos, T., Shah, V., Jaeglé, L., Lee,
1019 B. H., Lopez-Hilfiker, F. D., Thornton, J. A., Brown, S. S. and Weber, R. J.: Biomass Burning Markers and
1020 Residential Burning in the WINTER Aircraft Campaign, J. Geophys. Res. Atmos., 124(3), 1846–1861,
1021 doi:10.1029/2017JD028153, 2019.
1022
1023 Takeuchi, M. and Ng, N. L.: Organic nitrates and secondary organic aerosol (SOA) formation from oxidation of
1024 biogenic volatile organic compounds, ACS Symp. Ser. Multiph. Environ. Chem. Atmos., 1299, 105–125,
1025 doi:10.1021/bk-2018-1299.ch006, 2018.
1026
1027 Tao, Y., Liu, T., Yang, X. and Murphy, J. G.: Kinetics and Products of the Aqueous Phase Oxidation of
1028 Triethylamine by OH, ACS Earth Sp. Chem., 5(8), 1889–1895, doi:10.1021/acsearthspacechem.1c00162, 2021.
1029
1030 United States Environmental Protection Agency: Historical Exceedance Days in New England, Reg. 1 EPA New
1031 Engl. [online] Available from: https://www3.epa.gov/region1/airquality/standard.html (Accessed 14 August 2021),
1032 2020.
1033
1034 VandenBoer, T. C., Petroff, A., Markovic, M. Z. and Murphy, J. G.: Size distribution of alkyl amines in continental
1035 particulate matter and their online detection in the gas and particle phase, Atmos. Chem. Phys., 11(9), 4319–4332,
1036 doi:10.5194/acp-11-4319-2011, 2011.
1037
1038 Vidović, K., Lašič Jurković, D., Šala, M., Kroflič, A. and Grgić, I.: Nighttime Aqueous-Phase Formation of
1039 Nitrocatechols in the Atmospheric Condensed Phase, Environ. Sci. Technol., 52(17), 9722–9730,
1040 doi:10.1021/acs.est.8b01161, 2018.
1041
1042 Warneke, C., McKeen, S. A., de Gouw, J. A., Goldan, P. D., Kuster, W. C., Holloway, J. S., Williams, E. J., Lerner,
1043 B. M., Parrish, D. D., Trainer, M., Fehsenfeld, F. C., Kato, S., Atlas, E. L., Baker, A. and Blake, D. R.:
1044 Determination of urban volatile organic compound emission ratios and comparison with an emissions database, J.
1045 Geophys. Res. Atmos., 112(10), 1–13, doi:10.1029/2006JD007930, 2007.
1046
1047 Wu, C., Wen, Y., Hua, L., Jiang, J., Xie, Y., Cao, Y., Chai, S., Hou, K. and Li, H.: Rapid and highly sensitive
1048 measurement of trimethylamine in seawater using dynamic purge-release and dopant-assisted atmospheric pressure
1049 photoionization mass spectrometry, Anal. Chim. Acta, 1137, 56–63, doi:10.1016/j.aca.2020.08.060, 2020.
1050
1051 Ye, C., Zhou, X., Pu, D., Stutz, J., Festa, J., Spolaor, M., Tsai, C., Cantrell, C., Mauldin, R. L., Campos, T.,
1052 Weinheimer, A., Hornbrook, R. S., Apel, E. C., Guenther, A., Kaser, L., Yuan, B., Karl, T., Haggerty, J., Hall, S.,
1053 Ullmann, K., Smith, J. N., Ortega, J. and Knote, C.: Rapid cycling of reactive nitrogen in the marine boundary layer,
1054 Nature, 532, 489–491, doi:10.1038/nature17195, 2016.
1055
1056 Youn, J. S., Crosbie, E., Maudlin, L. C., Wang, Z. and Sorooshian, A.: Dimethylamine as a major alkyl amine
1057 species in particles and cloud water: Observations in semi-arid and coastal regions, Atmos. Environ., 122, 250–258,
1058 doi:10.1016/j.atmosenv.2015.09.061, 2015.
1059
1060 Zhang, J., Ninneman, M., Joseph, E., Schwab, M. J., Shrestha, B. and Schwab, J. J.: Mobile Laboratory

1061   Measurements of High Surface Ozone Levels and Spatial Heterogeneity During LISTOS 2018: Evidence for Sea
1062   Breeze Influence, J. Geophys. Res. Atmos., 125(11), 1–12, doi:10.1029/2019JD031961, 2020.
1063
1064   Zhou, S., Collier, S., Xu, J., Mei, F., Wang, J., Lee, Y.-N., Sedlacek, A. J., Springston, S. R., Sun, Y. and Zhang, Q.:
1065   Influences of upwind emission sources and atmospheric processing on aerosol chemistry and properties at a rural
1066   location in the Northeastern U.S., J. Geophys. Res. Atmos., 121(10), 6049–6065, doi:10.1002/2015JD024568, 2016.
1067