

Response to Reviewer #2:

Review of:

“Impacts of ice-nucleating particles from marine aerosols on mixed-phase orographic clouds during 2015 ACAPEX field campaign”

Authors: Lin et al.

Recommend major revisions.

General comment:

Overall, I find the paper and topic interesting and relevant to the microphysics modeling community. The potential impacts of marine sea-salt particles as INP is particularly interesting in, that historically, they have not been considered efficient INP. So, exploring their impact within a controlled microphysics modeling environment is quite important. As you will see in my specific comments below, I think the authors need to give a fairer assessment of the INP impacts on modeled fields and not attempt to overemphasize effects that appear to be rather minor when viewing the figures. Please show and explain the results in a balanced manner.

We thank the reviewer for your time and helpful comments. Our point-by-point response is enclosed. We would like to point out that this is about the potential impact of sea-spray aerosols, not sea salt aerosols. Sea spray particles that can contain sea salt, but also organics importantly. It is just in the model we used sea salt as a surrogate for sea spray particles given that most marine organic aerosols exist with coating on the surface of sea salt particles in the size range that dominates surface area.

Specific comments:

1. Abstract line 36-37: What is the difference between “post-AR” and “after AR”?

“After AR landfall” (these three words need be read together) is a stage after AR made landfall with large increase of IVT and precipitation. “Post-AR” begins from the point when AR started to retreat. The different AR stages are defined in Section 3.

2. Introduction line 44: Please be more specific regarding AR impacts on the “western” U.S., specifically when you state that it accounts for 30-50% of the precipitation. By western, do you mean Pacific Coast states?

We meant California. The sentence has been revised as “On a long-term average, AR storms contribute to 20–50% of California’s precipitation totals (Dettinger et al., 2011)” (Lines 44-45).

3. Section 2: Will you please provide the hydrometeor fall speed power law coefficients used in this version of HUCM-SBM for each ice species? I have found the power law coefficients to be quite important in such studies and would like to know what was used in this study.

The fall speed power law relationships are shown in a previous study <https://doi.org/10.1175/MWR-D-16-0385.1> (see Table 1 and Fig. 1 of this study). We have added a sentence to cite this study, i.e., “The fall speed power law relationships for ice/snow and graupel are depicted in Xue et al. (2017)” (lines 140-141)

4.Line 260: The 40% reduction in aerosol number is quite a lot. How much closer to observations do you get with use of the CARB2015 dataset over NEI2011?

We have calculated the mean number concentrations of aerosols below 2 km from observations and simulations using CARB2015 and NEI2015 and added the discussion, i.e., “The aerosol concentration averaged over 1-2 km altitudes is about 160 cm^{-3} with CARB2015 and 317 cm^{-3} with NEI2015, which is 26% lower and 47% higher than aircraft observations (215 cm^{-3}), respectively. Thus, the simulated aerosol concentrations with CARB2015 are in better agreement with observations” (Lines 251-254).

5.Section 2.2: How do you get realistic dust transport over the Pacific into California in this scenario with limited spin-up time?

Our chemical and aerosol initial and boundary conditions are from the global WRF-Chem simulations, not from the original WRF-Chem setup which requires a week to spin up. That means the chemistry and aerosol fields used for our model simulations already have steady-state values. Therefore, 2 days spin-up for our simulations are good enough. Since our model simulations overestimates aerosol concentrations in general, there is no indication of insufficient spin-up time.

6.Line 297: Why is only larger dust present? Wouldn't the larger dust tend to settle out before the smaller dust particles?

We meant to say the dust number concentration is dominated by small particles but it was said oppositely. Thanks for capturing the mistake. It is now corrected, and numbers are provided, i.e., “Although the simulated dust mass fraction is $\sim 14\%$, the derived number concentration for sizes larger than $0.5 \mu\text{m}$ is very low (less than 0.02 cm^{-3} , shown in a later figure). This is because the dust number concentration is dominated by small particles (14.71 cm^{-3} for the sizes smaller than $0.5 \mu\text{m}$)” (Lines 344-346).

7.Lines 316-323: This discussion on accumulated precipitation hinges on small changes seen in figure 3. It is difficult to see the changes being discussed in this manner. Perhaps figure 3 should include difference plots so that we can more readily see the spillover effects being discussed. As figure 3 is currently presented, all the simulations look very similar with only minor differences in the details as one would expect when changing a parameterization.

For the spatial distribution plots, difference plots have a problem to clearly show the comparison with the observations at the stations. We have used white boxes to mark up the exact location where these differences are observed (Figure 4a) and calculated mean precipitation for the white box area. Also, for the time series plot, we have added a panel for the differences between the simulation and observations (Fig. 4c). All quantitative discussion has been added as “All three

simulations predict a narrower but higher peak precipitation compared with the observed wider peak with lower values (Fig. 4b). However, the overestimation of the peak value by DM15+MC18 is lower than the other two (30% vs. 45% for DM15 and 58% for Bigg; Fig. 4b-c). The accumulated precipitation in the southern mountain range (the south part of white boxes in Fig. 4a) is generally less than 100 mm in observations and less than 120 mm in DM15+MC18 but more than 140 mm in other two simulations. The mean precipitation over the white box accumulated over the AR period are 89, 128, 130, and 116 mm for observations, Bigg, DM15, and DM15+MC18, respectively. Again, although all three simulations overestimate the precipitation, DM15+MC18 simulates the lowest value and closer to observations” (Lines 358-367).

8.Lines 333-334: While long time integration could be impacting cell formation, there are many other model artifacts that could be hindering better prediction compared to the obs. Since you are using this statement to justify moving forward with the analysis, there needs to be better justification or explanation for why the predicted cells differ from the observations. We need to be convinced that the simulations are trustworthy.

It will need quite a lot of effort to figure out factors leading to the model biases to provide more specifics here, which would also be a large distraction of the focus. We agree that many other things in model could affect the biases, and we have revised the sentence to “The deviations of the simulation from observations for the postfrontal clouds could be because of various reasons such as (a) the long-time model integration time (the 4th day after model initiation) and (b) the spatial mismatch of simulated and observed clouds since those postfrontal clouds are small” (Lines 385-387).

9.Lines 353-363: In this section the authors seems to be focused on the improvement to the glaciation ratio in the simulation with sea salt INP while downplaying the large overestimation in water content in figure 5c. Should the simulation with MC18 be considered “better” than the others?

We have added text to discuss the large overestimation of LWC and IWC by MC18 because the post-frontal clouds are invigorated a lot (see Lines 403-407). We specifically state that MC18 is only better in simulating the cloud phase states. Since all simulations do not predict the post-frontal clouds well, we would recommend focus on how significant those clouds are changed by marine INPs instead of which simulation is better.

10.Lines 393: The comment on the 330% increase in precipitation is very misleading here. From figure 7a, it can be seen that the 330% increase occurs from a VERY small absolute increase in precipitation at a time when precipitation rate is very small. These sorts of statements regarding the analysis that over emphasizes a small impact should be clarified or not included in the discussion. The overall changes in precipitation rate due to sea salt INP is quite small as seen in figure 7a, with at most 0.1 mm/hr change.

Yes, we agree. We have changed to “over 50% in general” in Lines 431-432 and we have emphasized small precipitation before AR landfall by adding a sentence “Note the precipitation

is very small at some point before AR landfall so the large increases might not mean that much” (Lines 435-436).

11.Lines 423-425: While the rain vs snow argument is valid for hydrologic reasons, you have already shown that the accumulated precipitation differences are very small between simulations. So, does this really matter?

We would say the ratio of rain vs. snow matters more to people and society than the accumulated precipitation. Also, the spatial distribution of the precipitation matters more than the accumulated precipitation over a large region to our stakeholders.

12.In general, the discussion of the spillover effect is quite interesting and perhaps should be highlighted rather than over-emphasizing minor changes in precipitation rate. Further, I find figure 9, and the discussion on the glaciation of the cloud, the most fascinating part of this story thus far. It is my opinion that these features should be emphasized earlier in the paper and place it in the context of figure 5c. Does the MC18 simulation produce too much overall condensate while better predicting the relative proportions of water to ice?

The spillover effect was brought up earlier when discussing Fig. 4a and more detailed discussion is provided later with two figures (Figures 8 and 9). So it is well highlighted. Because section 4.1 focuses on comparisons with observations and the discussion about marine INP impacts starts from Figure 7, those features are already discussed in the earliest places they appear and the reasons leading to such features are elaborated in the section when marine INP effects are focused. The same for the overestimations of condensate content by MC18, which is described in Figure 5c but further discussed in section 4.3. We also added more pointers/hintes for readers to connect those discussions.

“Does the MC18 simulation produce too much overall condensate while better predicting the relative proportions of water to ice?”, Yes, and it is discussed with Fig. 6 (original Fig. 5) in section 4.1.

13.Lines 528-532: Here you are stating that competition for ice nucleation between dust and marine INP explains the differences in deep cloud occurrence and precipitation, yet you stated earlier that the ice nucleation for dust and marine INP occurs in two different temperature regimes and thus different vertical locations in the cloud. So, how do they compete in this scenario if they are activated in different locations?

In fact, the much higher ice nucleation rate from marine INPs in this case is only because of the large amount of marine INPs, not because their nucleation temperature regimes are different. Sorry about the misleading sentence (we forgot to change it to be consistent with the discussion of Figure 10 on Page 23). Now throughout the paper, we have made a consistent argument.

Here we have removed the argument. After we replotted Figure 10 c and d with the same log scale, we did not see a clear competition between dust and sea-spray nucleation.

14.Lines 538-542: Here you state that cloud dynamics (vertical velocity) is not changed much. Yet in line 524 you discuss invigoration of postfrontal cloud cells. This appear contradictory. Please clarify and discuss how invigoration works in this scenario. You mention the term invigoration several times but have not explained the process.

Here “invigoration” means increased condensates and cloud fraction, not stronger convection, which was already defined with “The mostly significantly invigorated postfrontal cloud cells by the marine INP effect (i.e., the increase in both LWC and IWC and a large increase in cloud fraction) might...” (now Lines 564-566). We have avoided using the term “invigoration” in other places in the revised manuscript. There was a paragraph in Section 4.3 (now the last paragraph P. 25) devoted to discussing why the postfrontal clouds can be invigorated. i.e., “As discussed earlier, the largest ice nucleation rates from marine aerosols at the post-AR stage explain the largest marine INP effects among the three stages. The factors contributing to the larger ice nucleation rates include the increased abundance of marine aerosols compared to the previous two stages (Fig. 10b). In addition, with the ~ 6 °C colder temperatures below 8-km altitudes during the post-AR stage compared to the other two stages, ice nucleation from marine aerosols becomes more efficient (Fig. 10d). The most significantly invigorated postfrontal cloud cells by the marine INP effect (i.e., the increase in both LWC and IWC and a large increase in cloud fraction) might also be related to small scale thermodynamic changes through the feedback of microphysical changes over the first two AR stages”.

15.Conclusion: The papers first referenced in the conclusion should be discussed earlier in the manuscript.

That is not necessary. Depending on how you summarize. We think our organization of results and summary makes sense.

Figure 2: Please show panel A on a linear scale rather than log scale. The log scale somewhat minimizes the larger differences in aerosol number between the obs and model. Near 2.8km altitude, I estimate the aerosol concentrations to be 40/cm³ (obs) and 200/cm³ (model).

Aerosol concentrations vary by a few magnitudes in the vertical profile so it is usually plotted with log scale. We have provided more quantitative discussions now, i.e., “The simulation overestimates the total aerosol number concentrations by ~ 2 -times averaged over the altitudes of 2.2-3.2 km. At 2.8 km, the difference between the simulation (219 cm⁻³) and observations (55 cm⁻³) is about 4 times” (Lines 334-336).

Figure 10: Why is there no homogeneous ice nucleation in panel D? Is this a contouring issue since the values in panel D are much larger than panel C? Perhaps you could plot panels C and D on a common log scale so that we can see the comparable differences. Also, the figure caption should say “The freezing rates in (c) and (d)”

Yes, it is because values in Panel d were plotted in three magnitude higher values. We have replotted c and d with the same log scale (also replotted a and b with the same log scale). The typo in the figure cation is also corrected. Thanks for the suggestion.