Technical note: Entrainment-limited kinetics of bimolecular reactions in clouds

Christopher D. Holmes[1]

5   [1] Earth, Ocean, and Atmospheric Science, Florida State University, Tallahassee, FL 32306, USA

Correspondence to: Christopher D. Holmes (cdholmes@fsu.edu)

Abstract. The method of entrainment-limited kinetics enables atmospheric chemistry models that do not resolve clouds to simulate heterogeneous (surface and multiphase) cloud chemistry more accurately

10   and efficiently than previous numerical methods. The method, which was previously described for reactions with first-order kinetics in clouds, incorporates cloud entrainment into the kinetic rate coefficient. This technical note shows how bimolecular reactions with second-order kinetics in clouds can also be treated with entrainment-limited kinetics, enabling efficient simulations of a wider range of cloud chemistry reactions. Accuracy is demonstrated using oxidation of SO_2 to $S(VI)$—a key step in

15   formation of acid rain—as an example. Over a large range of reaction rates, cloud fractions, and initial reactant concentrations, the numerical errors in the entrainment-limited bimolecular reaction rates are typically $\ll 1\%$ and always $< 4\%$, which is far smaller than the errors found in several commonly used methods of simulating cloud chemistry with fractional cloud cover.

1 Introduction

20   Aqueous reactions in clouds play an important role in atmospheric chemistry, production of acid rain from SO_2 being a prominent example (Seinfeld and Pandis, 2016). Rapid heterogeneous (surface and multiphase) reactions can consume reactants within clouds, making the overall reaction rate dependent on entrainment to supply additional reactants from the surrounding air. Since clouds are sub-grid-scale features in many large-scale regional and global atmospheric models, accounting for these processes in

25   chemical transport models is challenging. To address these challenges, Holmes et al. (2019) introduced entrainment-limited uptake, an algorithm to accurately and efficiently account for cloud chemistry occurring in just a fraction of a grid cell. The method incorporates cloud fraction and entrainment into

the kinetic rate expression, enabling calculation of concentrations in a partly cloudy model grid cell with very little computational effort. The original paper applied entrainment-limited uptake to first-

30 order loss of nitrogen oxide compounds (NO_2, NO_3, N_2O_5) and showed that clouds are a globally significant sink for these gases (Holmes et al., 2019). The method has since been applied to nitrogen oxide isotopes (Alexander et al., 2020), nitrate in urban haze (Chen et al., 2021), dimethyl sulfide oxidation (Novak et al., 2020), and reactive halogens (Wang et al., 2021), all of which also involved first-order loss reactions in clouds. This note derives entrainment-limited reaction kinetics for

35 bimolecular reactions with second-order kinetics so that the entrainment-limited method can be applied to a wider range of chemical systems that are important in the atmosphere.

2 Derivation

The computational challenge of cloud chemistry in a fractionally cloudy grid cell is that explicitly calculating reactant concentrations in the cloudy and clear fractions would increase the model's

40 variables and computational effort. For cloud reactions with first-order kinetics, however, Holmes et al. (2019) showed that explicitly calculating concentrations within clouds can be avoided. For a reaction with loss frequency k_i ins clouds, the reaction rate in a partly cloudy grid cell is

$$R_1 = k_1[A] \qquad \text{1a}$$
$$k_1 = k_i \left(\frac{x}{1+x} \right) \qquad \text{1b}$$

where [A] is the reactant concentration in the grid cell (averaged over cloudy and clear fractions),

45 $x/(1+x)$ is the fraction of reactant inside cloud, and

$$x = \frac{1}{2}(f' - k' - 1) + \frac{1}{2}(1 + k'^2 + f'^2 + 2k' + 2f' - 2k'f')^{1/2}, k' \equiv \frac{k_i}{k_c}, f' \equiv \frac{f_c}{1-f_c}. \qquad 2$$

The cloud fraction is f_c and $1/k_c$ is the mean residence time of air in clouds. The expression is exact for steady decay in which concentrations in and out of clouds decline at the same fractional rate. The overall idea is that kinetics governing grid-cell concentration follows the usual first-order form (Eq. 1a)

50 with rate coefficients that depend on entrainment as well as chemical kinetics. We will follow a similar approach for bimolecular reactions.

Bimolecular reactions, $A + B \rightarrow$ products, typically follow second-order kinetic rate expressions of the form $R = k_{AB}[A][B]$, where k_{AB} is the rate coefficient. For reactions within clouds, the rate depends on

55   reactant concentrations within clouds, designated $[A]_i$ and $[B]_i$. These concentrations are related to the grid-average concentration via $[A]_i/[A] = x_A/f_c(1 + x_A)$, where x_A is defined by Eq. 2 using the loss frequency for A within cloud. $[B]_i/[B]$ and x_B are defined similarly. The loss frequency for A within cloud is the pseudo-first order rate $k_{i,A} = k_{AB}[B]_i$ and $k_{i,B} = k_{AB}[A]_i$ is the analogous loss for B. This forms a system of equations that collectively define in-cloud reaction rates for bimolecular reactions:

60
$$k_{i,A} = k_{AB}[B]\left(\frac{x_B}{f_c(1 + x_B)}\right) \qquad \text{3a}$$

$$k_{i,B} = k_{AB}[A]\left(\frac{x_A}{f_c(1 + x_A)}\right) \qquad \text{3b}$$

The system of equations 2 and 3 can be solved by root finding methods or fixed-point iteration. After evaluating x_A and x_B, the overall reaction rate in a partly cloudy grid cell is found by substituting Eq. 3a into Eq. 1:

65
$$R_2 = k_2[A][B] \qquad \text{4a}$$

$$k_2 = \frac{k_{AB}x_A x_B}{f_c(1 + x_A)(1 + x_B)}. \qquad \text{4b}$$

Equation 4b is the exact form of entrainment-limited bimolecular reaction rate coefficient.

We can also derive an approximation to the entrainment-limited bimolecular rate coefficient that does

70   not require iteration to solve. In the limit where the in-cloud reaction is much faster than entrainment $(k_{i,A} \gg k_c$ or $k_{i,B} \gg k_c)$, the grid-scale losses of A and B are determined by the rate at which the limiting reactant is entrained into clouds:

$$R_2 \approx f' k_c \min([A], [B]). \qquad \text{5}$$

In the limit where in-cloud reactions are slow $(k_{i,A} \ll k_c$ and $k_{i,B} \ll k_c)$, the losses follow second-order

75   kinetics determined by the grid-scale mean concentrations:

$$R_2 \approx f_c k_{AB}[A][B] \qquad \text{6}$$

Combining these limits gives an approximation of the entrainment-limited bimolecular loss rates, expressed as a grid-scale 2nd order rate coefficient

$$k_2 \approx \left(\left(\frac{f' k_c \min([A],[B])}{[A][B]} \right)^{-1} + (f_c k_{AB})^{-1} \right)^{-1} \qquad 7$$

80 This approximate entrainment-limited bimolecular reaction rate coefficient can be used in Eq. 4a.

3 Evaluation

The accuracy of entrainment-limited bimolecular reaction rates will now be demonstrated using oxidation of S(IV) by aqueous H_2O_2, which is a prominent step in the formation of S(VI) and acid rain,

85 as an example (Chameides, 1984). One key aqueous reaction is $HSO_3^- + H_2O_2 + H^+ \rightarrow SO_4^{2-} + 2H^+ + H_2O$, where the reactants are dissolved gasses SO_2 and H_2O_2. While the reaction occurs in cloud droplets, the reaction rate can be expressed in terms of the gas-phase concentrations of SO_2 and H_2O_2 by incorporating the solubility and dissociation equilibria, cloud liquid water content, and aqueous kinetics into the effective rate coefficient (e.g., Park et al., 2004). For a cloud with 1 g m^{-3} liquid water at pH 5,

90 284 K, and 800 hPa, the effective, gas-phase bimolecular rate coefficient is $k_{eff} = 3.7 \times 10^{-14}$ cm^3 molecule^{-1} s^{-1}, which will be used in examples below. A similar approach can be applied to other bimolecular aqueous reactions.

Figure 1 shows that the exact entrainment-limited algorithm (Eq. 4) is nearly identical to a reference

95 solution in a two-box model that explicitly represents concentrations inside clouds and entrainment mixing with clear air. The approximate entrainment-limited solution (Eq. 7) also resembles the exact entrainment-limited and reference solutions, but remaining reactant concentrations diverge by 3 % after 1 hour and 10 % after 4 hours. Two other cloud chemistry methods that are used in current atmospheric chemistry models are also shown in Figure 1: the thin-cloud approximation, in which loss is computed

100 for the entire grid cell using grid-average liquid water content, and the cloud partitioning method, in which only reactants within the cloudy fraction can react, but the concentrations are homogenized across cloudy and clear regions each time step of the chemical solver. Holmes et al. (2019) describe these other methods in greater detail. Both of the other methods diverge from the reference solution and entrainment-limited method by large amounts.

105

Figure 2 shows accumulated error in the entrainment-limited kinetics over a wide range of initial reactant concentrations and cloud fractions. Results are presented as the error in total product formed, relative to the reference two-box model, after one hour of integration. Over most of the parameter space, the errors in the entrainment-limited calculations are much less than 1 %. The largest errors occur over a
110 narrow range of $k_{\mathrm{AB}}[\mathrm{B}]/k_c$ values in regions that are about half cloudy and these errors do not exceed 4 %. By the same metric, the approximate entrainment-limited bimolecular algorithm has 10-30 % error (Figure 2) and the thin-cloud and partitioning methods have several hundred percent errors (not shown).

The entrainment-limited approach is best suited for applications and models that do not require highly
115 detailed cloud and aqueous chemistry. For example, the derivation above assumes that reactants A and B are consumed in only one reaction. While additional in-cloud reactions and reactants can be incorporated into the pseudo-first order loss rates (Eq. 3), to account for their effects on x_A and x_B, solving the system becomes more computationally intensive as more reactants are involved. For cloud reactions that depend on [H$^+$], the pH must be assumed or calculated via another method because it is
120 infeasible to account for the relevant aqueous equilibria within the entrainment-limited equations. Overcoming these limitations, however, requires explicit representation of reactant concentrations and entrainment in the cloudy fraction of a grid cell, along with the extra computational burden that incurs. Despite the progression of atmospheric models to ever higher resolutions, fractional cloudiness is likely to remain a feature of many global and regional models for many years to come, necessitating some
125 means of accounting for its effect on chemistry.

4 Conclusion

The results here and in the earlier work of Holmes et al. (2019) show that the entrainment-limited reaction kinetics can provide an efficient and accurate means of representing heterogeneous cloud
130 chemistry in atmospheric models with fractional cloud cover. By incorporating cloud fraction and entrainment into the rate coefficient, the usual first- and second-order rate expressions are retained, allowing the entrainment-limited kinetics to be easily implemented in numerical codes. The entrainment-limited approach provides far greater accuracy than other methods currently in use; typical

errors for bimolecular reactions are $\ll 1\ \%$ error after 1 hour and always $< 4\ \%$. Entrainment-limited
kinetics have already been applied to numerous first-order reactions and the extension here to
bimolecular reactions should further expand its applicability and usefulness in atmospheric chemistry
modeling.

Code availability

Python code implementing the entrainment-limited bimolecular kinetics is provided in the supplement.

Competing interests

The author declares that he has no conflicts of interest.

Acknowledgments

I thank Daniel Jacob and Mike Long for helpful discussions. This work was supported by the NASA
New Investigator Program (grant NNX16AI57G).

References

Alexander, B., Sherwen, T., Holmes, C.D., Fisher, J.A., Chen, Q., Evans, M.J., and Kasibhatla, P.:
Global inorganic nitrate production mechanisms: comparison of a global model with nitrate isotope
observations, Atmos. Chem. Phys., 20, 3859-3877, https://doi.org/10.5194/acp-20-3859-2020, 2020.

Chameides, W.L: The photochemistry of a remote marine stratiform cloud, J. Geophys. Res., 89 (D3),
4739-4755, https://doi.org/10.1029/JD089iD03p04739, 1984.

Chan, Y-C, Evans, M.J., He, P., Holmes, C.D., Jaegle, L., Kasibhatla, P., Liu, X.-Y., Sherwen, T.,
Thornton, J.A., Wang, X., Xie, Z., Zhai, S., Alexander, B.: Heterogeneous nitrate production
mechanisms in intense haze events in North China, J. Geophys. Res., 126, e2021JD034688,
https://doi.org/10.1029/2021JD034688, 2021.

Holmes, C. D., Bertram, T. H., Confer, K. L., Graham, K. A., Ronan, A. C., Wirks, C. K., & Shah, V.:
The role of clouds in the tropospheric NO_X cycle: A new modeling approach for cloud chemistry

and its global implications, Geophys. Res. Lett., 46, 4980–4990, https://doi.org/10.1029/2019GL081990, 2019.

Novak, G.A., Veres, P.R., Fite, C.H., Neuman, J.A., et al.: Rapid cloud removal of dimethyl sulfide oxidation products short circuits new particle formation in the marine boundary layer, American Geophysical Union Fall Meeting 2020, 1-17 December 2020, Abstract A082-08, 2020.

Park, R. J., Jacob, D. J., Field, B. D., Yantosca, R. M., and Chin, M.: Natural and transboundary pollution influences on sulfate-nitrate-ammonium aerosols in the United States: implications for policy, J. Geophys. Res., 109, D15204, https://doi.org/10.1029/2003JD004473, 2004.

Seinfeld, J.H. and Pandis, S.N.: Atmospheric Chemistry and Physics: From Air Pollution to Climate Change (3rd Edition), Wiley, Hoboken, USA, 2016.

Wang, X., Jacob, D.J., Downs, W., Zhai, S., Zhu, L., Shah, V., Holmes, C.D., Sherwen, T., Alexander, B., Evans, M.J., Eastham, S.D., Neuman, J.A., Veres, P., Koenig, T.K., Volkamer, R., Huey, L.G., Bannan, T.J., Percival, C.J., Lee, B.H., and Thornton, J.A.: Global tropospheric halogen (Cl, Br, I) chemistry and its impact on oxidants, Atmos. Chem. Phys., https://doi.org/10.5194/acp-2021-441, 2021.

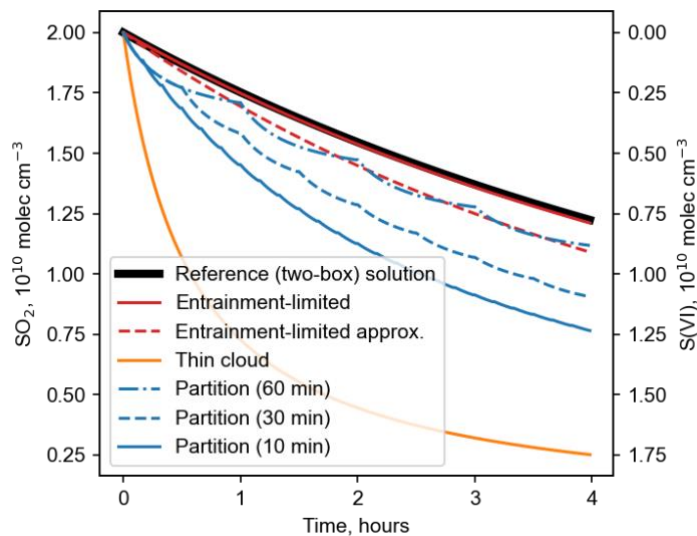



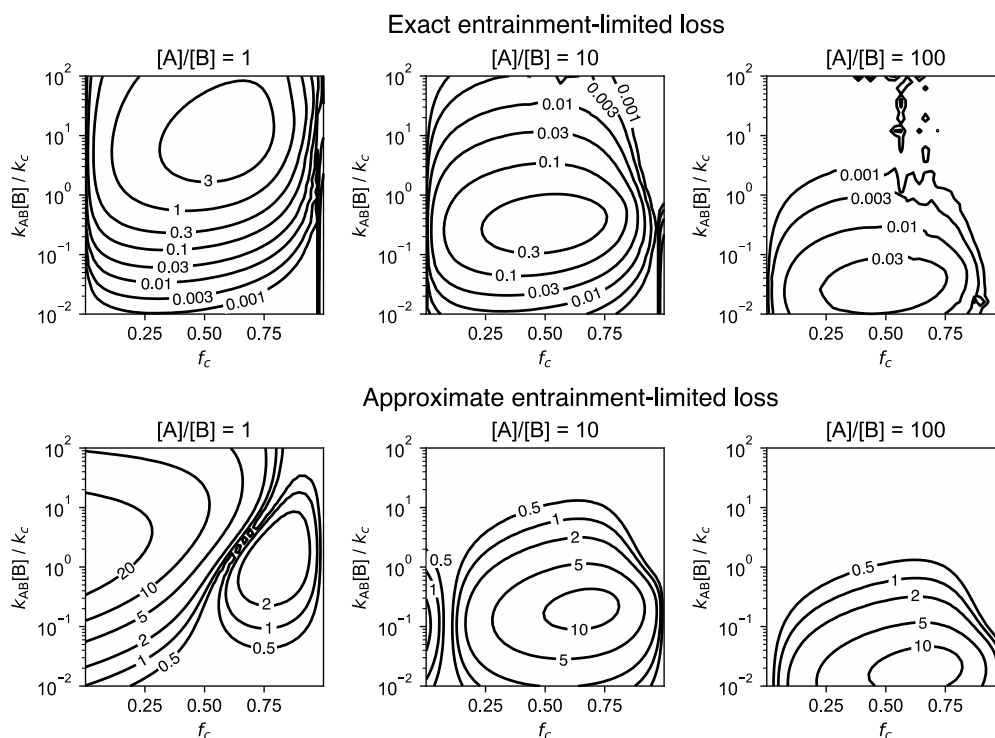
Figure 1: Comparison of numerical solutions for reaction of dissolved SO$_2$ with H$_2$O$_2$ in cloud water in a partly cloudy region. Calculations use conditions T = 284 K, p = 800 hPa, 1 g m^{-3} liquid water in cloud, pH = 5, f_c = 0.2, k_c = 1 h^{-1}, k_{eff} = 3.7×10^{-14} cm^3 molecule^{-1} s^{-1}, and initial concentrations [SO$_2$]$_0$ = [H$_2$O$_2$]$_0$ = 1 ppb. For the cloud partitioning method, the numbers in parentheses give the time step for homogenizing reactant concentrations.

185

Figure 2: Accuracy of exact entrainment-limited bimolecular kinetics (Eq. 4, top row) and approximate entrainment-limited kinetics (Eqs. 4a and 7, bottom row). Accuracy is shown as the percent difference (%) in the cumulative loss of reactants after 1 hour relative to a reference two-box model.