



The relationship between PM_{2.5} and anti-cyclone wave activity during summer over the United States

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Abstract. To better understand the role of atmospheric dynamics in modulating surface concentrations of fine particulate matter (PM_{2.5}), we relate the anti-cyclone wave activity (AWA) metric and PM_{2.5} data from the Interagency Monitoring of Protected Visual Environment (IMPROVE) data for the period of 1988-2014 over the US. The observational results are compared with hindcast simulations over the past two decades using the National Center for Atmospheric Research-Community Earth System Model (NCAR CESM). We find that PM_{2.5} is positively correlated (up to R=0.65) with AWA changes close to the observing sites using regression analysis. The composite AWA for high aerosol days (all daily PM_{2.5} above the 90th percentile) shows a similarly strong correlation between PM_{2.5} and AWA. The most prominent correlation occurs in the Midwestern US. Furthermore, the higher quantiles of PM_{2.5} levels are more sensitive to the changes in AWA. For example, we find the averaged sensitivity of the 90th percentile PM_{2.5} to changes in AWA is approximately three times as strong as the sensitivity of 10th percentile PM_{2.5} at one site (Arendtsville, Pennsylvania; 39.92 °N, 77.31 °W). The higher values of the 90th percentile compared to the 50th percentile in quantile regression slopes are most prominent over the northeastern US. In addition, future changes in US PM_{2.5} based only on changes in climate are estimated to increase PM_{2.5} concentrations due to increased AWA in summer over areas where PM_{2.5} variations are dominated by meteorological changes, especially over the western US. Changes between current and future climates in AWA can explain up to 75% of PM_{2.5} variability using a linear regression model. Our analysis indicates that higher PM_{2.5} concentrations occur when a positive AWA anomaly is prominent, which could be critical for understanding how pollutants respond to changing atmospheric circulation, as well as developing robust pollution projections.

1 Introduction

Particulate matter less than 2.5 μm in diameter (PM_{2.5}) poses a considerable air quality concern due to its impacts on human health (Liu et al., 2020). PM_{2.5} has been linked to increased possibility of mortality (Krewski et al., 2009). Continuing exposure to PM_{2.5} can exacerbate existing cardiovascular and respiratory problems, and cause lung damage (Bernard et al., 2001). It can also alter the body's defense system against foreign materials, and even lead to premature death (Kappos et al., 2004).



Furthermore, PM_{2.5} could contribute to the degradation of visibility (Hand et al., 2011; Ashley et al., 2015) and the alteration of the hydrological cycle through changing rainfall formation mechanisms in clouds (Rosenfeld et al., 2008). Once deposited onto snow PM_{2.5} can cause an increase in melting of snow and ice (Painter et al., 2007) or modify land or ocean biogeochemistry (Mahowald et al., 2011). It also influences Earth's energy balance directly by scattering solar incoming radiation back into space (Charlson et al., 1992) or indirectly by altering cloud albedo and lifetime (Albrecht, 1989; Arimoto, 2001). Aerosols can also alter the large-scale atmospheric circulation and lead to historical changes to wintertime midlatitude winter extremes over northern Eurasia (Wang et al., 2020). Thus, understanding and predicting potential changes in PM_{2.5} is crucial to both human health concerns and current environmental issues.

Climate change and meteorological factors may have substantial impacts on surface concentrations of PM_{2.5}. A number of studies have demonstrated the influences of weather and climate change on PM_{2.5} variability (Brasseur et al., 2006; Liao et al., 2006; Dawson et al., 2007; Jacob and Winner, 2009; Thishan Dharshana et al., 2010; Tai et al., 2012; Liu et al., 2017; Chen et al., 2018; Wang et al., 2019; Wang et al., 2020), showing connections between specific meteorological conditions and PM_{2.5} concentration responses. Daily variation in temperature, rainfall, moisture, and circulation can explain up to 50% of the variability in PM_{2.5} (Tai et al., 2010). Meanwhile, Aw and Kleeman (2003) calculated a reduction in PM_{2.5} concentrations caused by rising temperature over southern California in a modeling study. Furthermore, Tai et al. (2012) identified cyclone passage with the associated cold front as a mechanism for temperature being strongly correlated with inter-annual variability of PM_{2.5} in the Midwestern US. Similarly, increased atmospheric stagnation may also have the potential to aggravate PM_{2.5} air quality in the future climate (Liao et al., 2006; Leibensperger et al., 2008). Mickley et al. (2004) also proposed an increased severity of summertime PM episodes due to a warmer future climate in the Northeastern and Midwestern US. In addition, a significantly negative correlation caused by the longer lifespan of gas-phase-produced sulfate was found between aerosol sulfate and cloud cover compared to aqueous-phase-produced sulfate (Koch et al., 2003). These studies suggest that changes in meteorology linked to climate modification may cause variations of PM_{2.5} levels and exposure risks. However, questions remain regarding whether and how PM_{2.5} concentrations directly related to atmospheric general circulation. Understanding the relationship between meteorology and PM_{2.5} levels will be critical to the understanding of pollutant response to a changing circulation due to climate change, as well as the development of robust pollution projections.

Many studies have focused on relating wave activity with extreme weather and climate events. Weather extremes are strongly influenced by the natural variability of the atmosphere at synoptic, intra-seasonal, and inter-annual time scales. For example, persistent high pressure blocking systems can cause extreme cold winter temperatures in Europe (Woolings et al., 2008) or summer heat waves (Coumou et al., 2015). Mid-latitude weather extremes can be influenced by the major modes of climate variability such as Arctic Oscillation (AO) (Michel and Riviere, 2011) or El Niño–Southern Oscillation (ENSO) (Ryoo et al., 2013). Furthermore, blocking anticyclones can decay by releasing accumulated wave activity as a stationary Rossby wave train (Takaya and Nakamura, 2001). Meanwhile, blocking highs are connected with high-amplitude, quasi-stationary anticyclonic anomalies that result in protracted and unusual weather events (Nakamura et al., 1997). However, no previous studies have investigated how wave activity can lead to changes in PM_{2.5} concentrations, although studies have shown a robust correlation



between surface ozone concentration and a measure of wave activity over the US using a linear regression model (Sun et al., 2019).

Here we propose the use of AWA, the anticyclonic part of local finite-amplitude wave activity (LWA) with quantile regressions, to be an effective method for diagnosing the tendency and sensitivities in the transport of PM_{2.5} concentrations, using a similar approach to a previous study focused on ozone (Sun et al., 2019). LWA assesses longitude-by-longitude anomalies of the finite-amplitude Rossby wave activity deviating from the circle of equivalent latitude using the meridional displacement of potential vorticity (PV) (Huang and Nakamura, 2016). As an example, the prominent Northern Hemisphere blocking episode that occurred in late October 2012 was well explained using LWA (Huang and Nakamura, 2016). Furthermore, recent modeling studies have illuminated the value of quantile regression in examining the influences of global changes on local air quality. For example, Porter et al. (2015) employ quantile regression to diagnose the meteorological sensitivities of higher ozone and PM_{2.5} levels by using observed and reanalysis meteorological data in the US over the past decade.

While meteorology sets the stage for the occurrence of dangerous pollutant levels in the present and future climates, the connection between the meteorological events and extreme pollution events is still not sufficiently understood (Dawson et al., 2014). In the present study, we apply a univariate linear regression model to analyze daily PM_{2.5}-AWA relationship from observations and simulations in the US, using a similar methodology as that used to examine LWA and ozone relationships (Sun et al., 2019). We also use quantile regression to calculate PM_{2.5} sensitivities to AWA across quantiles from 10% to 90%. In addition, the coefficients of the slope for the linear regression model between PM_{2.5} and AWA in the current climate are evaluated to see how much they can be used to predict future PM_{2.5} changes. Such an exploration of PM_{2.5} from atmospheric motion can provide new insights into understanding the mechanism of PM_{2.5} changes, as well as the sensitivity of higher PM_{2.5} concentrations to climate change.

2 Methodology

In order to quantify the relationship between AWA and surface PM_{2.5} concentrations during the summer months, five different combinations of datasets, one observational and four model output, are used in this study. Each of the five cases use surface PM_{2.5} concentrations and geopotential heights (for calculating a modified AWA, as explained below). Each modeling dataset uses 20 years for the analysis, while the observations are analyzed for the years between 1988 and 2014. Both PM_{2.5} and AWA time series are detrended and deseasonalized by eliminating the 21-day smoothed seasonal cycle from the original data. We use the resulting stationary residuals to focus on the synoptic-scale variability, minimizing aliasing from regular seasonal variations or long-term tendencies. This study is a followup study of Sun et al. (2019), using the same approach used there to analyze ozone, and here we focus on analyzing PM_{2.5}. The details of the methods are discussed in more detail in that paper, and summarized here.



2.1 Observational Data

We use PM_{2.5} measurements from the Interagency Monitoring of Protected Visual Environment (IMPROVE) monitoring sites, which are located in National Parks and Class I Wilderness areas across the US (<http://vista.cira.colostate.edu/improve/>).
90 PM_{2.5} concentrations are reported every 3 days as mass per air volume at local temperature and pressure. Hand et al. (2011; 2012) have described the details regarding IMPROVE site locations, sampling, analysis approach, and detailed information of network operations. All stations with at least 1000 valid PM_{2.5} values between 1988 and 2014 are collected in this study, totaling 150 stations for summer (June, July, August, JJA hereafter). For the PM_{2.5} observations (called Obs case), the full names, states, latitudes and longitudes, as well as short names, are listed in Table S1 and locations are shown in Figure 1.
95 We chose three representative stations in different parts of the country to investigate the relation between AWA and PM_{2.5} in detail. The IMPROVE station names are AREN1 (Arendtsville, Pennsylvania; 39.92 °N, 77.31 °W; in northeast), SIPS1 (Sipsey Wilderness, Alabama; 34.34°N, 87.34°W; in northwest) and LAVO1 (Lassen Volcanic NP, California; 40.54°N, 121.58°W; in west), which are shown with the red dots in Figure 1.

In order to compare these in situ observations to the large scale meteorology, the 500hPa geopotential height (m) from the
100 European Reanalysis Interim version is used (ERA-Interim) for the time period of January 1991 to December 2010 (Dee et al., 2011). The ERA-Interim is a global reanalysis of recorded meteorological data over the past 3.5 decades and was undertaken by the European Centre for Medium-Range Weather Forecasts. This gridded dataset is created at approximately 0.7° spatial resolution with 37 vertical levels.

2.2 Model output

105 We compare the IMPROVE data with output from Community Earth System Model (CESM) as simulated for the Chemistry-Climate Model Initiative (CCMI) (Eyring et al., 2013). As a state-of-the-art Earth system modeling framework coordinated by the National Center for Atmospheric Research (NCAR), the CESM employed here is configured to fully couple the Community Atmosphere Model version 4 (CAM4) (Neale et al., 2010), the Community Land model version 4.0 (CLM4.0) (Oleson, 2010), the Parallel Ocean Program version 2 (POP2) (Smith et al., 2010) and the Los Alamos sea ice model (CICE version 4) (Hunke
110 and Lipscomb, 2008). All simulations are performed under current land cover conditions. The top of the simulated atmosphere reaches to 40 km with a horizontal resolution of 2.5° longitude by 1.9° latitude. The model has been widely used for modeling the Earth's past, present and future climate states (Neale et al., 2010; Hurrell et al., 2013).

To simulate the atmospheric chemistry, we include the CAM4-Chem module, which has been widely studied on its representation of atmospheric chemistry in the atmosphere (Aghedo et al., 2011; Lamarque et al., 2011a, b, 2012; Lamarque and
115 Solomon, 2010). CAM4-Chem employs the bulk aerosol model (BAM) with externally mixed aerosols considering black carbon, organic carbon, sulfate, sea spray and desert dust, and which simulates coarse mode aerosols in 4 bins for the latter two. Details of this implementation are discussed in Lamarque et al. (2012). CAM4 uses the Zhang–McFarlane deep convection scheme (Zhang and McFarlane, 1995), Hack shallow cumulus scheme (Hack et al., 2006), Holtslag and Boville's (1993) planetary boundary layer process, and the parameterization of cloud microphysics and macrophysics by Rasch and Kristjánsson



120 (1998) and Zhang et al. (2003). Additionally, the convective momentum transport is utilized to parameterize deep convection (Richter and Rasch, 2008). These two major revisions caused improvements in such aspects as the Madden-Julian oscillation and ENSO (Neale et al., 2008). The improved trend and magnitude of surface PM_{2.5} using this free-running model has been evaluated elsewhere (Tilmes et al., 2016).

The chemical emissions and forcing details for each of the model simulations are listed in Table 1. The simulation using specified dynamics (REFC1SD) for current levels of PM_{2.5} from 1991 to 2010 is driven by analyzed meteorological data from Modern-Era Retrospective Analysis for Research and Applications (MERRA) (see Tilmes et al., 2016). This simulation follows the conventions of the CCMI (Eyring et al., 2013). For the AWA analysis for this case, we use the 500-hPa geo-potential height from MERRA, which should be very similar to that from ERA-Interim, since they use largely the same observations. To compare the relationship between AWA and PM_{2.5} concentrations in online simulations, three simulations forced by trace gas projections and an interactively coupled ocean are employed. The GCM2000 and GCM2100 simulations are 25-year runs branched from the CCMI reference simulations in the year 2000 and the year 2100, respectively. Simulations over the first five years are discarded as spin-up, and results from the latter 20 years are discussed here (2006–2025 for GCM2000 and 2106–2125 for GCM2100). Note that while the CO₂ concentrations in the GCM2000 and GCM2100 simulations are kept at the year 2000 and 2100 level respectively, the concentrations of all other greenhouse gases including methane remain constant at the year 2000. In particular, the emissions and prescribed chemical species for longer-lived substances follow the protocol defined by CCMI hindcast simulations for the year 2000 (Eyring et al., 2013), which are repeated for all the simulated model years. Another future run (REFC2) is forced by future climate combined with future emissions following the REFC2 CCMI modeling protocol. In this run greenhouse gas forcing and emissions following the RCP6 scenario. The relationship between ozone and AWA has been examined in the GCM2000, GCM2100 and REFC2 simulations in Sun et al. (2019). Characteristics of the REFC1SD simulation are given in Phalitnonkiat et al (2018). Note that our REFC2 set-up covers volcanic eruptions in the past, but possible volcanic eruptions in the future are not included (Eyring et al., 2013).

2.3 AWA calculation

Studies on finite amplitude wave activity (FAWA) have identified the link between the pattern of atmospheric circulation and large-scale wave dynamics (Nakamura and Solomon, 2011; Methven, 2013; Chen and Plumb, 2014; Lu et al., 2015). LWA adds the longitude dimension to the zonally average quantity FAWA and is calculated from the meridional displacement of quasigeostrophic potential vorticity (PV) from zonal symmetry (Nakamura and Zhu, 2010). LWA helps differentiate longitudinally isolated events and describe extreme weather events at the local scales (Huang and Nakamura, 2016, Chen et al., 2015). Chen et al. (2015) used local finite-amplitude wave activity based on the 500-hPa geo-potential height for characterizing mid-latitude weather events. The total wave activity is composed of the cyclonic wave activity residing to the south of the equivalent latitude and the anticyclonic wave activity (AWA) to the north (see Fig. 1 in Sun et al, 2019). In this study, we focus on AWA to characterize its connection with changes in PM_{2.5} concentrations. Over the US in summer LWA is dominated by its anticyclonic component (Sun et al., 2019). Sun et al. (2019) also used AWA to characterize ozone variability.



2.4 Quantile regression

Quantile regression is used to estimate the slopes for several conditional quantile functions (Koenker and Bassett, 1978). It characterizes the connection between a range of predictor variables and specified percentiles (or quantiles) of the response variable. For example, Porter et al. (2015) analyzed the sensitivities of ozone and PM_{2.5} concentrations for response quantiles ranging from 2 to 98%. The parameters of quantile regression models evaluate the change in a specific quantile of the response variable caused by a one-unit change in the predictor variable. This permits us to measure how some percentiles of the PM_{2.5} may be more influenced by AWA than others, and this is indicated by changes in the regression coefficient. In order to illustrate the sensitivity of the PM_{2.5} concentration at different quantiles, we apply linear quantile regression for percentiles from 10th to 90th at the AREN1 site. And then we compare the 90th percentile quantile regression coefficient with 50th percentile quantile regression coefficient at each station.

2.5 The univariate linear regression model

To help explore and measure the likely relationship between AWA and PM_{2.5} levels, we use the univariate linear regression model, similar to a previous study focused on ozone (Sun et al., 2019). Here the slope of PM_{2.5} with respect to wave activity ($S_{i_0, j_0}(i, j)$) on the daily time scale is used to show the linear association between changes (in time) of the normalized PM_{2.5} at a point (i_0, j_0) and the normalized wave activity at another point. We use the projection of PM_{2.5} onto AWA to reveal how closely the AWA anomaly field resembles the spatial pattern that enhances PM_{2.5} on the daily time scale during the summer. The projection of AWA (p_{i_0, j_0}) at all points in the domain onto S_{i_0, j_0} is defined according to the following equation:

170

$$p_{i_0, j_0} = AWA \cdot S_{i_0, j_0} = \sum_i \sum_j AWA(i, j) \times S_{i_0, j_0}(i, j) \quad (1)$$

The similarity between AWA spatial pattern and the PM_{2.5}-AWA regression coefficients' spatial structure is estimated by the projection value. The interannual change in PM_{2.5} due to changes in AWA is predicted based on a linear regression model as following equation (see Sun et al., 2019 for more discussion):

$$175 \quad PM_{2.5} = \beta p + \alpha \quad (2)$$

The change of PM_{2.5} (denoted by $\Delta PM_{2.5}$) in the future due to the change in AWA is calculated using the following equation, that is the climatological difference in future climate and the present climate ($\overline{AWA}_f - \overline{AWA}_p$), project it onto S, and multiply it by the slope β :

$$\Delta PM_{2.5} = \beta[(\overline{AWA}_f - \overline{AWA}_p) \cdot S] \quad (3)$$

180 Here, S is calculated from the values for the present climate. The projected value can measure the similarity between the AWA change and the PM_{2.5}'s trend with AWA by compressing the information of the AWA field into a single variable. This variable incorporates the non-local effect of AWA on PM_{2.5}'s variability.



2.6 The composite methodology

We use a composite methodology which is based around the most polluted ($>90^{th}$ percentile) daily PM_{2.5} and the correspond-
185 ing anomalies at every station. Composite 500-hPa geopotential height and AWA for daily values of PM_{2.5} larger than 90^{th}
percentile are produced by separately averaging all daily anomaly values of the corresponding 500-hPa geopotential height and
AWA.

3 Results and discussion

The monthly mean PM_{2.5} surface concentrations with standard deviation for different scenarios at three representative sites
190 are shown in Figure 2. PM_{2.5} concentrations are largest during summer at the three sites. The climatological average for
PM_{2.5} is greater in the Eastern than in the Western sites. A statistically significant correlation ($r>0.80$, $p<0.01$; Figure 2)
for current PM_{2.5} is found between observations and simulations of monthly mean climatological averages (REFC1SD and
GCM2000) at three representative sites. The highest correlation coefficients between model and observations (0.93) are seen
at SIPS1 perhaps due to the large seasonal variation in PM_{2.5} concentrations (Figure 2b). The future PM_{2.5} concentrations are
195 increased in GCM2100 under current emissions compared with current climate PM_{2.5} simulations. There is a strong decrease in
climatological mean for future PM_{2.5} at AREN1 under future emissions and meteorology (REFC2), while the climatological
average for future PM_{2.5} has no significant change under future emissions at SIPS1 and LAVO1. Such differences in the
monthly mean averages for PM_{2.5} suggests that emission changes are more important than climate changes at AREN1, but it
is not clear which is more important at SIPS1 or LAVO1.

200 Focusing on the spatial distribution, the highest PM_{2.5} concentrations over the 20-year average in summer occur in the
south-central US (Figure 3a; green lines; GCM2000 case). The 20-year averaged AWA on summer days exhibits a maximum
over the Southwestern US (Figure 3a; shading). The difference between two current climate simulations (REFC1SD minus
GCM2000) for summertime AWA is shown in Figure 3b. The reduced AWA in the reanalysis forced simulation is found across
most of the US, with the largest reduction in the Southwestern US as previously shown (Sun et al., 2019). In contrast, the AWA
205 in summer is higher over the Northeastern US in the forced simulations. The corresponding changes in summertime PM_{2.5}
concentration caused by a combination of different emissions and possibly changes in AWA is similar to changes in AWA,
although the reduction is largely over South-central US. The difference between two future scenarios (GCM2100; REFC2)
and current climate scenario (GCM2000) has a similar pattern (illustrated in Figure 3c and d), which shows a large increase in
AWA in the Southwestern US, but there is a difference in the amplitude of these changes (contrast Figure 3c vs. 3d) (Sun et
210 al., 2019). There is an increase in PM_{2.5} concentration for the future scenario with current emissions (GCM2100), while there
is a decrease in PM_{2.5} concentrations when future emissions are used (REFC2), showing the importance of future potential
decreases in emissions.



3.1 Relationship between PM_{2.5} concentrations and AWA at specific stations

For the three observation sites highlighted here (AREN1, SIPS1 and LAVO1), the PM_{2.5} concentrations are positively correlated with AWA in the areas close to the sites where presumably at least some of emissions of PM_{2.5} are located (Figure 4). The relationships between daily PM_{2.5} concentrations and AWA using CESM simulations presented here offer a test of the consistency between observational and model relationships in characterizing the response of PM_{2.5} to AWA. The highest regression coefficient occurs in the observational (Obs) and the reanalysis driven simulated cases (REFC1SD), as opposed to the case coupled meteorology (GCM2000) (Figure 4a, b, c (top row) and Figure 4d, e, and f (middle row) in contrast to Figure 4g, h and I (bottom row)). The highest spatial regression coefficients for site AREN1 and SIPS1 are located southward of the sites, while they are located to the northwest at LAVO1. Overall the model simulates similar spatial patterns to the observations for the case of the reanalysis driven simulations (GCM2000), but do less well for the coupled model simulations (REFC1SD). Of course, model predictions are not perfect, and include uncertain emissions and boundary conditions, as well as errors in model physical and chemical processes, which may be driving these inconsistencies between modeled and observed relationships. In addition Sun et al. (2019) shows considerable variation in the AWA pattern between different ensemble simulations suggesting even on the timescale of 20 years there is considerable internal variability between model ensemble members.

To obtain a more in-depth understanding of the physical mechanisms behind the relationship between PM_{2.5} concentration and AWA, we consider the composite AWA for high PM_{2.5} days at AREN1, SIPS1 and LAVO1 (Figure 5). The composite AWA calculated by averaging together all AWA corresponding to daily PM_{2.5} above the 90th quantile shows a similar strong correlation between PM_{2.5} and AWA as that seen for the average (Figure 4 vs. 5). The composite AWA are extremely pronounced over the areas where PM_{2.5} data originates. The geopotential composites for the highest pollution days generally showed spatial distributions which are similar to the regression coefficient distributions (Figure 5 vs. 4) at the three representative sites. The areas with the largest values of the composite AWA are located southward of the AREN1 and SIPS1 sites. But at LAVO1 the maximum is located to the northwest for the observational and reanalysis driven cases (Obs and REFC1SD), and eastward for the coupled model case (GCM2000). Overall, the composite AWA for PM_{2.5} also shows that the daily PM_{2.5} above its 90th quantile correlates strongly with AWA during summer. Note that there is a spatial displacement between maximum of geopotential height and the maximum of AWA, since wave activity fundamentally measures the waviness of atmospheric general circulation, rather than the magnitude.

3.2 Relationship of AWA and PM_{2.5} regionally

Next we consider how the local relationship between PM_{2.5} and AWA changes in space. To simplify the visualization of the spatial variability in the local relationship, we use the result from the previous section that the maximum regression coefficients between PM_{2.5} and AWA are usually close to the site where PM_{2.5} is measured (Figure 5). The highest composite AWA anywhere in the domain and the highest regression coefficient with AWA are shown at each gridpoint in Figure 6. If we look at the relationship between the PM_{2.5} concentration and the wave activity at each location, it can be seen that PM_{2.5} concentrations are positively correlated with AWA throughout the US, but with varying strengths (Figure 6). A roughly similar



spatial distribution is obtained when either the composite AWA for high PM_{2.5} (left-hand side: Figure 6a, c, e, g and i) and for the regression coefficients between the PM_{2.5} and AWA (right-hand side: Figure 6b, d, f, h and j) for the different observational or model combinations, showing the consistency of the approaches. The pattern of the distribution is consistent for observational (Figure 6a and 6b) and simulated (Figure 6c and 6d) data when forced with reanalyses winds, with the
250 largest values in the upper mid-west into the mid-Atlantic states. The climate model simulations tend to have a stronger correlation in the western/mountain regions, than seen in the reanalyses winds (Figure 6c vs. 6e), and there are some hints of this in the observations in Arizona, for example (Figure 6a). The correlations and composites tend to become stronger in the western/mountain regions in the future model simulations (Figure 6g and 6i vs. Figure 6e).

The consistency in the composite AWA under high PM_{2.5} conditions and the correlations argues that either of these metrics
255 can be useful tools to identify PM_{2.5} and AWA relationships. The anti-cyclonic condition is usually characterized by, low-level divergence, subsiding air, light wind, no rainfall, and high surface pressure. Taken together, the results above demonstrate the positive connection of PM_{2.5} concentrations with anti-cyclonic conditions everywhere across the US, which is likely accounted for arid weather and sinking inversions. This is consistent with the PM_{2.5} sensitivity study of Tai et al. (2010). They illustrated strong linkage between high PM_{2.5} concentrations, high 850-hPa geopotential height and stagnation which is characterized by
260 an anti-cyclonic condition. The significant association of PM_{2.5} with stagnation is also demonstrated by Cheng et al. (2007) in their examination of four Canadian cities. Furthermore, similar results were reported that a greatly strengthening ozone is associated with increased stagnation (Wu et al., 2008; Sun et al., 2014).

The composite AWA is for PM_{2.5} that larger than the 90th quantile, while the regression coefficient is for all PM_{2.5}. The composite AWA and regression coefficient have similar spatial distributions suggesting the positive connection between daily
265 PM_{2.5} and AWA is mainly produced by high PM_{2.5} concentration above its 90th quantile. Overall, the relationship between behavior of AWA and extreme PM_{2.5} concentration is generally consistent with the existing meteorological studies (Woolings et al., 2008; Coumou et al., 2015; Michel and Riviere, 2011; Ryoo et al., 2013).

3.3 The sensitivity of quantiles in PM_{2.5} concentrations to AWA

To examine the sensitivity of different levels of PM_{2.5} concentrations to AWA, we fit the linear regression and quantile regres-
270 sion for AWA and daily PM_{2.5} for summers between 1988 to 2014 from IMPROVE monitoring sites for different percentiles (10th to 90th percentiles) using an “impact region” of AWA at AREN1 site. Here the averaged AWA over the “impact region” is defined as an elliptic area bounded by the maximum and minimum longitude and latitude of the maximum composite AWA for PM_{2.5} larger than 90th percentile minus the 0.05 contour line (blue elliptic circle in Figure 5a). One can clearly see that the higher percentiles of PM_{2.5} are more sensitive to the change in the averaged AWA over the impact region, e.g., the 90th
275 percentile of PM_{2.5} is approximately three times more sensitive to the averaged AWA over the impact region when compared with the 10th percentile of PM_{2.5} (Figure 7a).

In order to examine whether the relationship that high quantile of PM_{2.5} is more sensitive to the AWA than the low quantile of PM_{2.5} applies to the other sites, we calculate the difference of 90th percentile quantile regression coefficient (slope) from 50th percentile quantile regression coefficient across all sites (90th percentile quantile regression coefficient (slope) minus



280 50th percentile quantile regression coefficient, shown in Figure 7b). Out of the 150 sites, 145 sites show that 90th percentile
PM2.5 increases more than the 50th percentile of PM2.5 with the enhancement of the AWA. In the Northeast region (north and
east of New York state with New York state included), this relationship is the most pronounced. This difference in response
between the highest and median PM2.5 values indicates the different sensitivities within various percentiles of the PM2.5 levels.
These results are to some extent consistent with those from Porter et al. (2015), which addressed that averaged sensitivity of
285 95th percentile summertime ozone to changes in highest daily temperature was larger than the sensitivity of 50th percentile
summertime ozone.

3.4 Projected PM2.5 concentrations due to changes in future AWA

The strong association between PM2.5 concentrations and AWA in the current climate prompts us to investigate the extent to
which we can utilize a linear regression model to predict changes in PM2.5 concentrations from AWA change in future climate.
290 Employing daily present-day summertime concentrations of PM2.5 and AWA for current climate from the coupled model
simulation (GCM2000) and equation (1)-(3), we derive that how much of PM2.5's interannual variance can be explained by the
projection of JJA AWA anomalies onto the daily PM2.5-AWA regression coefficients pattern. The coefficient of determination
(R^2) of the linear regression model using simulated PM2.5 and AWA for the present climate varies from 0 to 0.75 depending
on gridbox (Figure 8). This means that the projected value (using only AWA changes) captures up to 75% of the interannual
295 variability in PM2.5 over Great Plains and West. Wise and Comrie (2005) similarly determined R^2 values of 0.1-0.5 for
associations of PM with atmospheric variables across sites in the Southwest. Because of the high correlation coefficients
(75%) this suggests that the regression results reveal the broad population instead of a small number of influential outliers
(Cook, 1979). The R^2 measures the part of variance of PM2.5 that can be explained by the linear regression model (Kutner et
al., 2004).

300 Next we explore how much of the future change in PM2.5 concentrations can be predicted just on the basis of changes in
AWA. Using PM2.5-AWA relationships determined from current coupled model output (GCM2000), future PM2.5 changes
can be estimated by using the linear relationship fitted with the current data and projected change of AWA in the future (as
shown in equation (1)-(3)). Here we assume that the linear relationship between the predictors and PM2.5 do not change very
much in the future compared to the present to extrapolate the current linear relationship between PM2.5 and AWA to the future.
305 Future climate change is simulated to cause an increase in PM2.5 concentrations over most of the US if there are no changes
in emissions (GCM2100; Figure 9a). Using CAM4-chem fitted slope pattern and regression coefficient, enhanced PM2.5 is
found over most of the western US and a small area in the Northeast (Figure 9b). The prediction suggests that the significant
increase in future PM2.5 resulted from AWA changes arises over the western US, which is up to $0.92 \mu\text{g m}^{-3}$. Great Plains,
South-central, Midwest and Southwest show a small decrease of PM2.5 in the future, where a negative value of the projection
310 value occurs when projecting the positive anomaly onto it. The projection of PM2.5 change in most parts of the South is less
reliable, because of the low R^2 in this region in the interannual variability metrics (Figure 8).

In order to investigate how much change in PM2.5 can be caused by the change of AWA, the fraction of predicted JJA
PM2.5 change in total JJA PM2.5 change from simulation is calculated in Figure 10. Over all, the maximum values are found



in the western US. We infer that AWA can be generally utilized in PM_{2.5} predictions during summer in the US, especially over
315 the western US. The impact of AWA change alone on summer PM_{2.5} concentrations is likely to be quite significant (above
50%) in the western US, suggesting that AWA is a suitable tool for air quality predictions for most regions where meteorology
dominates. These results are somewhat consistent with Tagaris et al. (2007) who showed that the Midwest was modelled to
have the larger daily average PM_{2.5} levels in the future, but our signal is in the west. In addition, studies suggest increases in
PM_{2.5} over polluted regions in the future climate caused by intensified stagnation (Liao et al., 2006; Jacob and Winner, 2009),
320 which is consistent with the anti-cyclonic condition seen with AWA in this study.

Increased AWA over most parts of US in the future climate are projected to increase PM_{2.5} levels over western regions where
meteorology dominates fluctuations in PM_{2.5}. This is consistent with some studies reporting an increased PM_{2.5} concentra-
tions due to more common and extended stagnation periods across northern mid-latitudes in the future climate if anthropogenic
emissions remain constant (Mickley et al., 2004; Liao et al., 2006; Leibensperger et al., 2008; Jacob and Winner, 2009). Sim-
325 ilarly in simulations, a mean increase of $0.24 \mu\text{g m}^{-3}$ in summer average PM_{2.5} levels with a largest growth of $0.93 \mu\text{g m}^{-3}$
in the Midwest has been shown (Tai et al., 2010). An increase in PM_{2.5} despite globally increasing precipitation is also
obtained by using coupled chemistry climate models, revealing a decreased precipitation on a large scale across polluted re-
gions and seasons (Fang et al., 2011; Kloster et al., 2010). In contrast, Avise et al. (2009) found that changes in meteorology
tend to reduce summertime PM_{2.5} concentrations (approximately $-1 \mu\text{g m}^{-3}$) in most regions with the maximum reductions
330 over the Southeastern US caused by intensified wet deposition. Furthermore, some models suggest there could be a regional
increase in summertime PM_{2.5} over the eastern US due to lower precipitation (Racherla and Adams, 2006), while Tagaris et al.
(2007) found that climate change, alone, with no emissions increase or controls affects the US PM_{2.5} concentrations slightly.
The discrepancy between studies mentioned above and the results given here is most likely attributable to differences in model
formulation. Although earlier studies predict important changes in PM_{2.5} levels in a warming climate throughout the US, there
335 is no consistency across studies (Jacob and Winner, 2009). The unpredictable sensitivity of PM_{2.5} levels to climate change
could be explained by the complication of the reliance of different PM_{2.5} elements on climatic variables, and the uncertainties
in regional boundary layer ventilation and precipitation, as well as the diversity of PM_{2.5} components (Racherla and Adams,
2006; Pye et al., 2009).

It should be noted that the change in surface PM_{2.5} predicted by the future change of AWA using the univariate linear
340 regression models is different from the simulated future change of PM_{2.5} from the constant emissions run, which simulate
the most significant change in the eastern US (Figure 9a), consistent with the other model projections (Tai et al., 2012). This
discrepancy mostly results from the distribution of projected AWA change (Figure 3). In addition, Figure 9b only includes the
change caused by the change of AWA. These studies did not account for the change of PM_{2.5} separately due to stagnation,
temperature or other meteorological conditions, which could also act a significant part in the PM_{2.5} changes. In addition, the
345 predicting capability of the linear models is limited, and the model only looks at linear relationships between PM_{2.5} levels and
AWA. Also, it only looks at the mean of PM_{2.5} levels.



4 Conclusions

We employed a univariate linear regression model to determine the correlation of PM_{2.5} levels and AWA on synoptic-scales over the US. This analysis demonstrates that PM_{2.5} is positively linked to the local anticyclonic finite-amplitude wave activity over the past two decades during JJA, and the high PM_{2.5} concentrations are more sensitive to the AWA than those low ones. The relationship between AWA and PM_{2.5} in model-simulated and observational data agrees in their general pattern and amplitude. These results provide insights into the drivers behind high PM_{2.5} pollution episodes in the observed record, emphasizing the significance of atmospheric circulation to the pollutant accumulation.

We found that AWA is positively correlated with PM_{2.5} at every available station in the summer using regression analysis in this study. The most prominent relationship between PM_{2.5} and AWA occurs in the Midwestern US for current climate, while it moves westward in the future climates. The composite AWA for PM_{2.5} larger than its 90th percentile can also demonstrate the positive relationship between PM_{2.5} and AWA. Climate change in the future is likely to cause a response in regional PM_{2.5}. The sensitivities of PM_{2.5} levels to changes in AWA are, generally, more robust for higher percentiles through quantile regression, which is most prominent in the northeastern US. It means that changes to AWA are likely to influence the extent of PM_{2.5} extremes more strongly than they influence moderate PM_{2.5} levels. This study presents new perspectives to explore both the observed and simulated PM_{2.5} responses to climate change. Furthermore, the contrast of observed sensitivities to those simulated by CESM could determine essential model biases relating to the prediction of future PM_{2.5}, potentially offering perceptions into the fundamental mechanistic reasons behind those biases.

The coefficient of determination of the linear regression model using simulated PM_{2.5} and AWA for the present climate is up to 75% over Great Plains and West, which shows that the daily variation in AWA can project up to 75% of interannual PM_{2.5} variability across the US. These effects suggest that AWA could have significant impacts on the PM_{2.5} levels. Significant regional variation is found in these results, indicating that while the positive association between PM_{2.5} and AWA are generally consistent, the extent to which AWA influences PM_{2.5} is local.

Data availability. PM_{2.5} data are available at <http://vista.cira.colostate.edu/improve>. ERA-Interim data can be accessed at <http://www.ecmwf.int/en/research/climate-reanalysis/era-interim>. The CCMI output data can be downloaded from the Centre for Environmental Data Analysis (<http://data.ceda.ac.uk/badc/wcrp-ccmi/data/CCMI-1/output/>).

Author contributions. YW analysed the observational data and visualized the model output. YW and NM wrote the manuscript. PH and WS provided critical feedback, and helped shape the research and analysis. GC provided the expertise on local finite-amplitude wave activity. All authors reviewed the manuscript.

Competing interests. The authors declare that they have no conflict of interest.



Acknowledgements. We acknowledge high-performance computing support from Cheyenne (doi:10.5065/D6RX99HX) provided by NCAR's Computational and Information Systems Laboratory, sponsored by the National Science Foundation. We also thank Simone Tilmes.

Financial support. This research has been supported by NUAA grants (NS2019037) and the National Science Foundation (NSF AGS-1608775).



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Table 1. The model case names used in this study. Designated model descriptions for the four attribution cases. Online means that the model is calculating the meteorology prognostically online. Case names are based on names described previously (Eyring et al., 2013; Sun et al., 2019).

Simulation name	Model Case Name	Period	<i>GHG</i> ¹ forcing	<i>SST</i> ² and sea ice	Emissions	Meteorological driver
REFC1SD	f.e11.TSREFC1SD.f19. f19.ccmi23.001	1991-2010	<i>CMIP5</i> ³	<i>HadISST2</i> ⁴	Anthropogenic and biomass burning emission: <i>MACCity</i> ⁵ Biogenic emissions: <i>MEGAN2</i> ⁶	<i>MERRA</i> ⁷
GCM2000	b.e11.TSREFC2.femis 2000.y2000.f19.f19. ccmi23.001	2006-2025	<i>CO</i> ₂ = 369 ppm	<i>Online</i> ⁸	Anthropogenic and biomass burning from <i>AR5</i> ⁹ Biogenic emissions: Monthly values from <i>MEGAN2</i> for 2000	Online
GCM2100	b.e11.TSREFC2.femis 2000.y2100.f19.f19. ccmi23.001	2106-2125	<i>CO</i> ₂ = 669 ppm	<i>Online</i> ⁸	Same as GCM2000	<i>Online</i> ⁸
REFC2	b.e11.TSREFC2.f19.g1 6.ccmi23.001.cam.h7	2080-2099	A1B scenario	Online	A1B scenario	Online

¹Greenhouse gas. ²Sea surface temperature. ³Coupled model intercomparison project. ⁴Hadley Center Sea Ice and Sea surface temperature data set (Titchner and Rayner (2014)). ⁵Granier et al. (2011). ⁶Guenther et al. (2012). ⁷Modern era retrospective-analysis for research and applications (Rienecker et al. (2011)). ⁸Tilmes et al. (2016). ⁹Assessment report 5 (Eyring et al. (2013)).

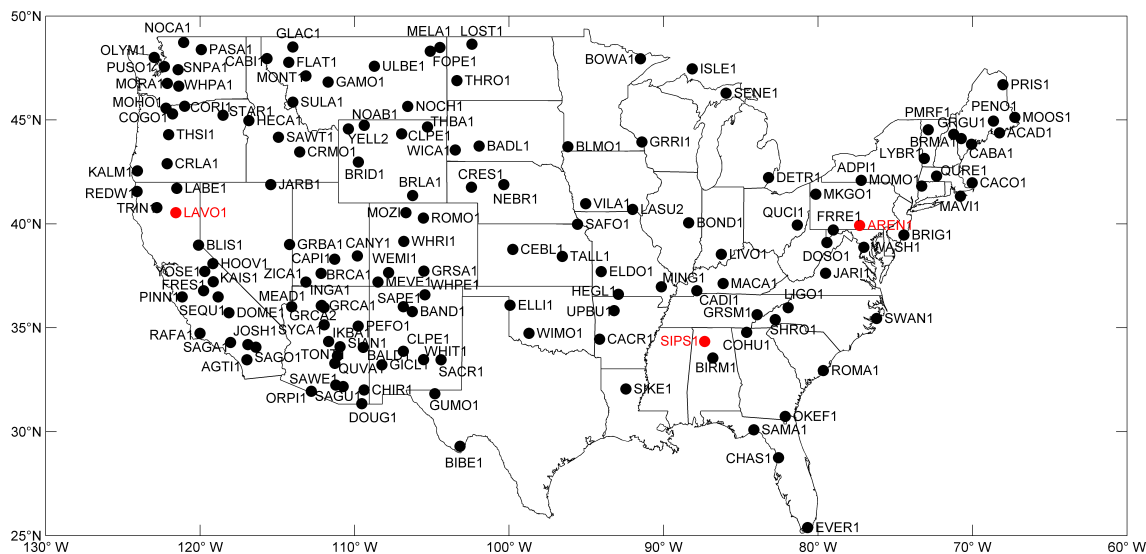


Figure 1. Geographical location of IMPROVE sites for PM_{2.5}. Red dots are three representative sites.

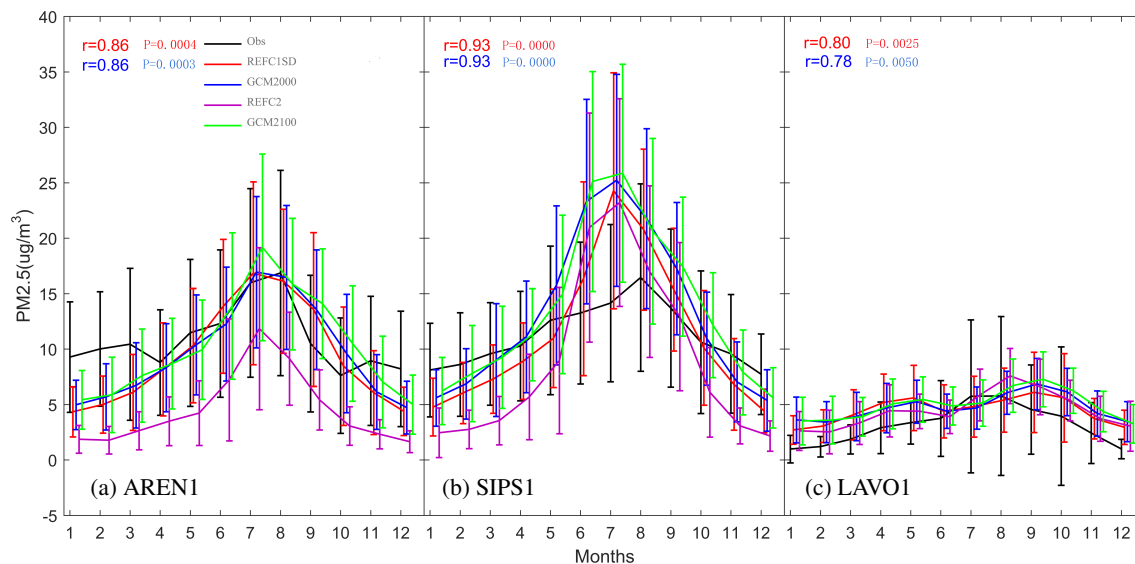


Figure 2. Climatological monthly mean average with standard deviation for PM_{2.5} used in this study at three sites (AREN1 (a), SIPS1 (b) and LAVO1(c)) for five scenarios used in this study. Red (r) represents the correlation coefficient between observation (Obs) and simulation for current from the REFC1SD simulation. Blue r represents the correlation coefficient between observation (Obs) and simulation for current from GCM2000 simulation. The p-values are included.

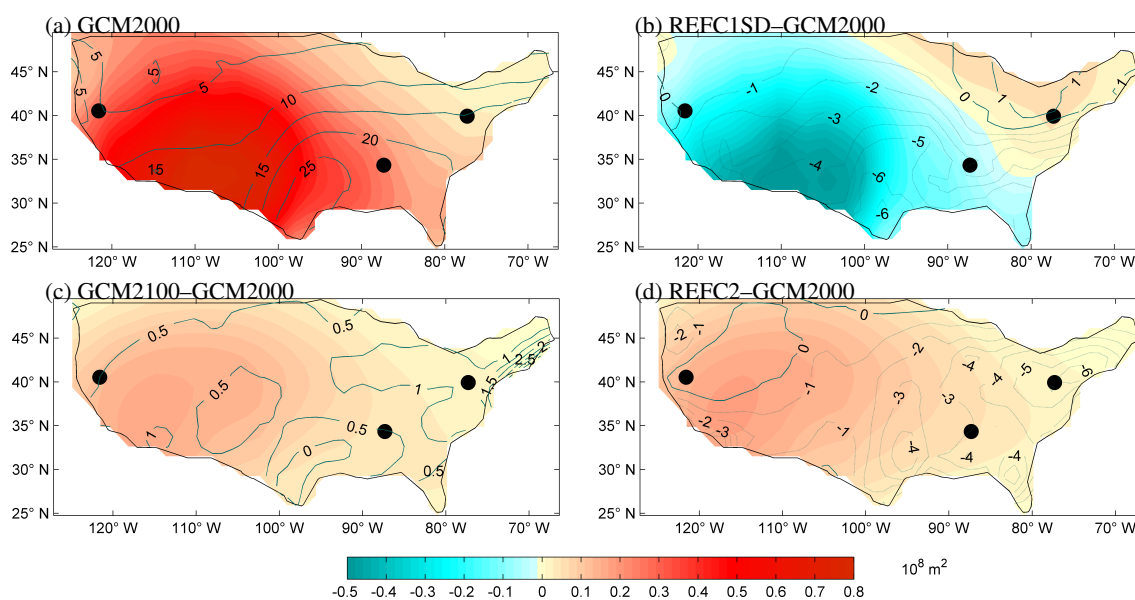


Figure 3. Wave activity (AWA: shading using the legend in $10^8 m^2$) and PM_{2.5} concentrations (green contour lines in $\mu g m^{-3}$) for (a) the current climate (GCM2000, 2006-2025 summer days' average) (b) reanalysis driven case (REFC1SD, 1991-2010 summer days' average) minus the current climate online case (GCM2000, 2006-2025 summer days' average); (c) Future climate with current emission (GCM2100, 2106-2125 summer days' average) minus the current climate (GCM2000, 2006-2025 summer days' average); (d) Future climate with future emission (REFC2, 2080-2099 summer days' average) minus current climate (GCM2000, 2006-2025 summer days' average). Three black dots are representative stations (AREN1, SIPS1 and LAVO1).

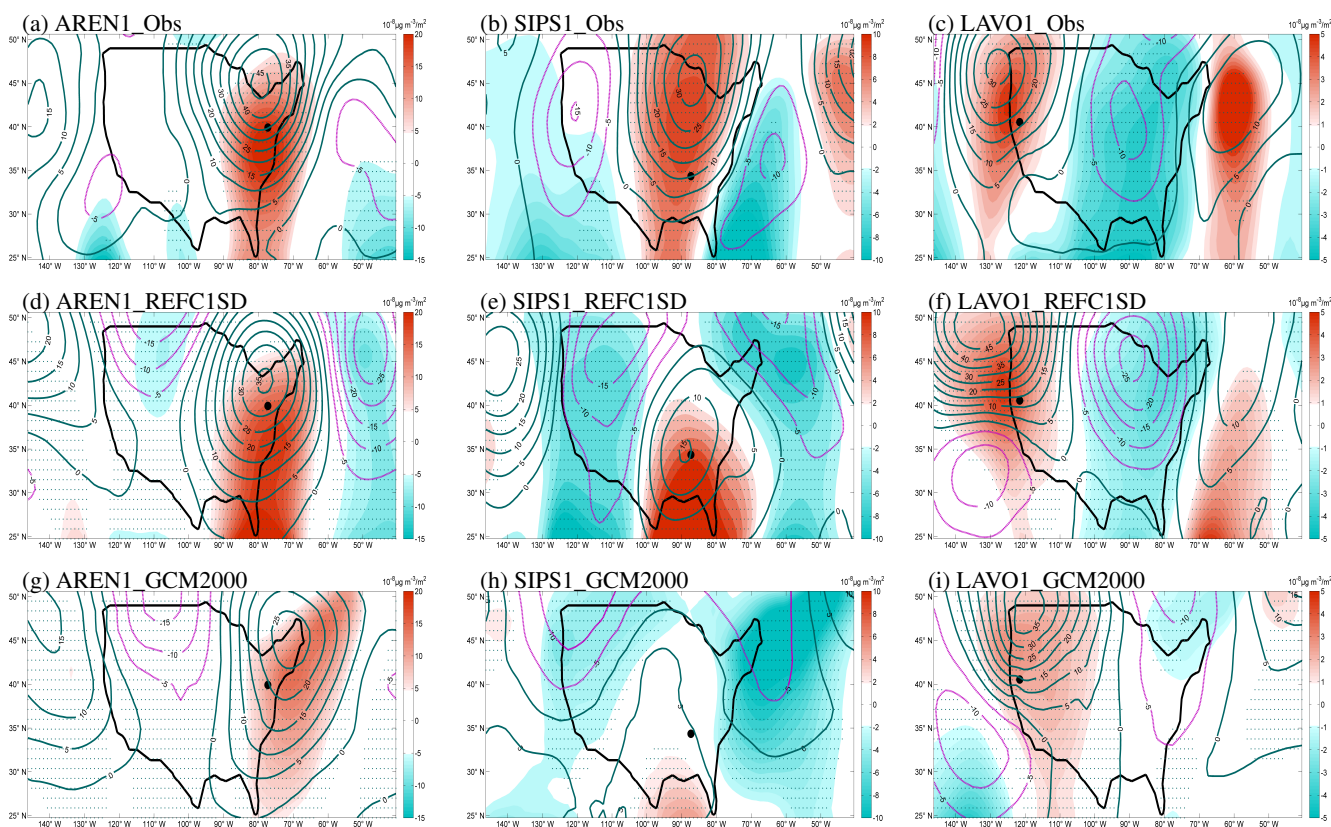


Figure 4. (Shaded) regression coefficients between daily AWA and PM2.5 at site (denoted by the black dots) (a, d, g) AREN1, (b, e, h) SIPS1 and (c, f, i) LAVO1 in the study domain for daily JJA time series of current climates. The top row are results using IMPROVE PM2.5 and reanalysis AWA, the middle row uses the reanalysis driven simulated PM2.5 (REFC1SD) and reanalysis AWA, and the bottom row uses current climate simulated PM2.5 and AWA (GCM2000). Stippling indicates the regions that are statistically significant at the 95% confidence level. Unit: $10^{-8} \mu\text{g m}^{-3} / \text{m}^2$ for regression coefficients.

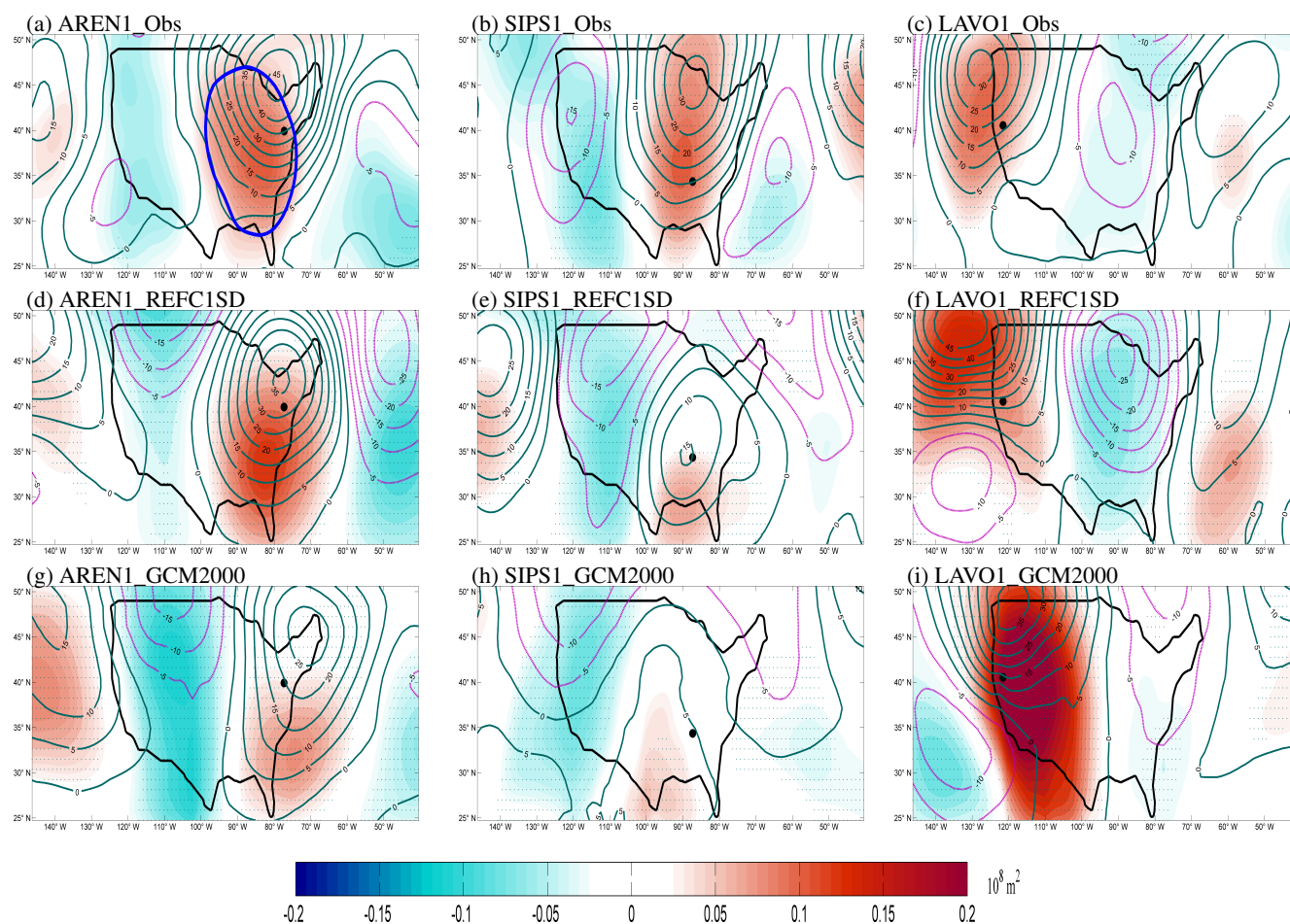


Figure 5. (Contour) composite 500-hPa geopotential height anomaly (positive values are represented by solid green lines and negative values by dashed purple lines) and (shading) corresponding AWA for PM_{2.5} larger than 90th quantile at site (denoted by the black dots) (a, d, g) AREN1, (b, e, h) SIPS1 and (c, f, i) LAVO1. The top row are results using IMPROVE PM_{2.5} and reanalysis AWA, the middle row uses the REFC1SD simulated PM_{2.5} (REFC1SD) and reanalysis AWA, and the bottom row uses REFC2 simulated PM_{2.5} and AWA (GCM2000). Stippling indicates the regions that are statistically significant at the 95% confidence level. Unit: m for 500-hPa geopotential height and 10^8 m^2 for AWA. The blue outlined area in a) is the impact region, which is defined as the region of the maximum regression coefficient minus 0.05.

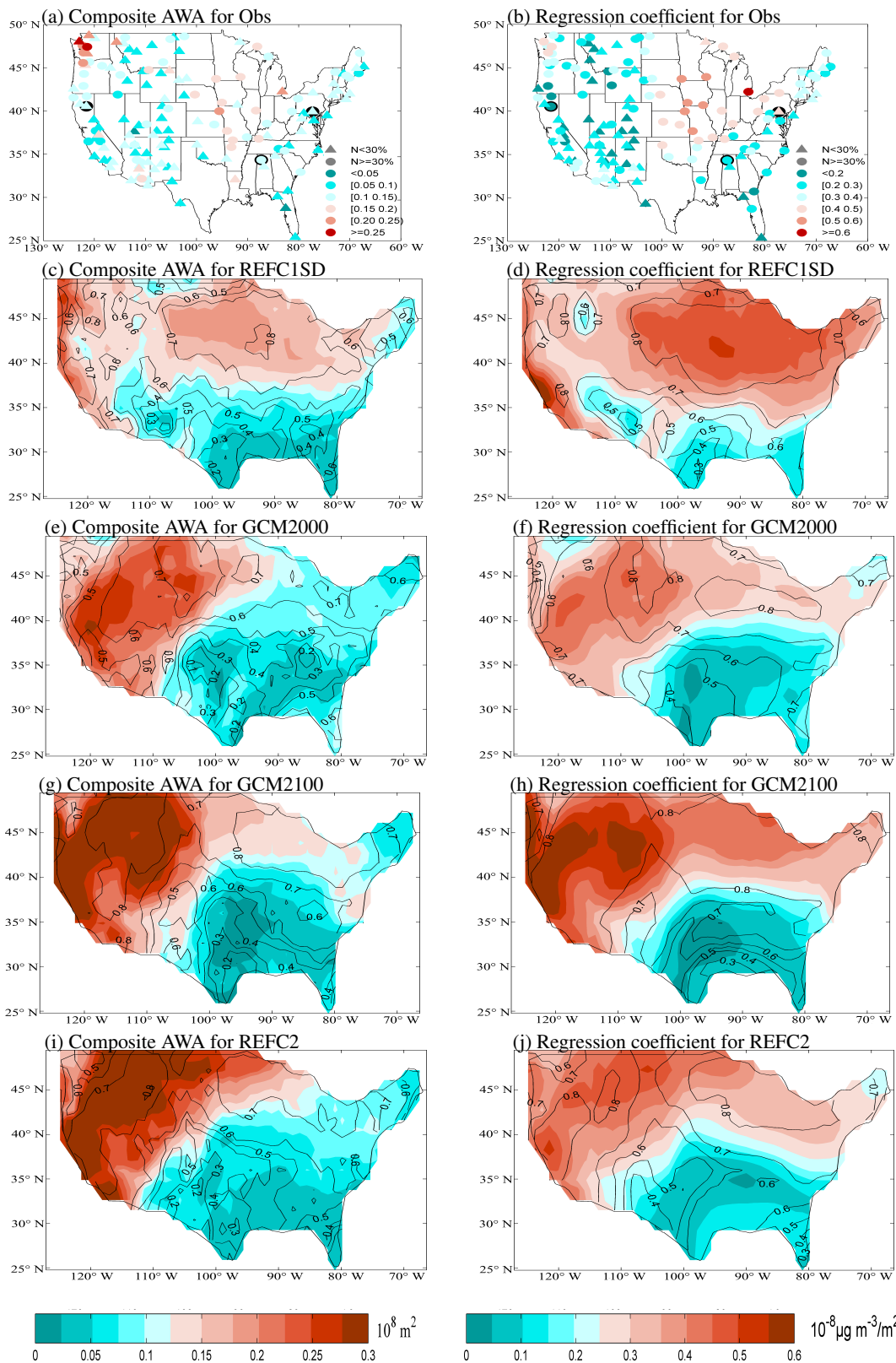




Figure 6. The maximum of the composite AWA distribution for PM_{2.5} larger than 90th quantile (Shading) (a, c, e, g, i), and the centers of the spatial regression coefficient distribution between PM_{2.5} and AWA (b, d, f, h, j): observations (Obs, first row), current climate from the reanalysis driven simulation (REFC1SD, second row), current climate from the coupled model simulation (GCM2000, third row), future climate with future emission (GCM2100, fourth row) and future climate with current emission (REFC2, bottom row). At each grid point, the highest composite AWA anywhere in the domain based on the PM_{2.5} larger than 90th quantile and the highest regression coefficient between AWA and PM_{2.5} are shown. In a) and b), the three representative sites are denoted by the black dots. In a) and b) the different shapes (circle or triangle) indicate the number of values for every grid that are statistically significant (at the 95% confidence level) is more than 30% or not. The different colors indicate different highest composite AWA and regression coefficients as indicated in the legend. In c) through j) the number of values for every grid that are statistically significant at the 95% confidence level are shown (in black contours).

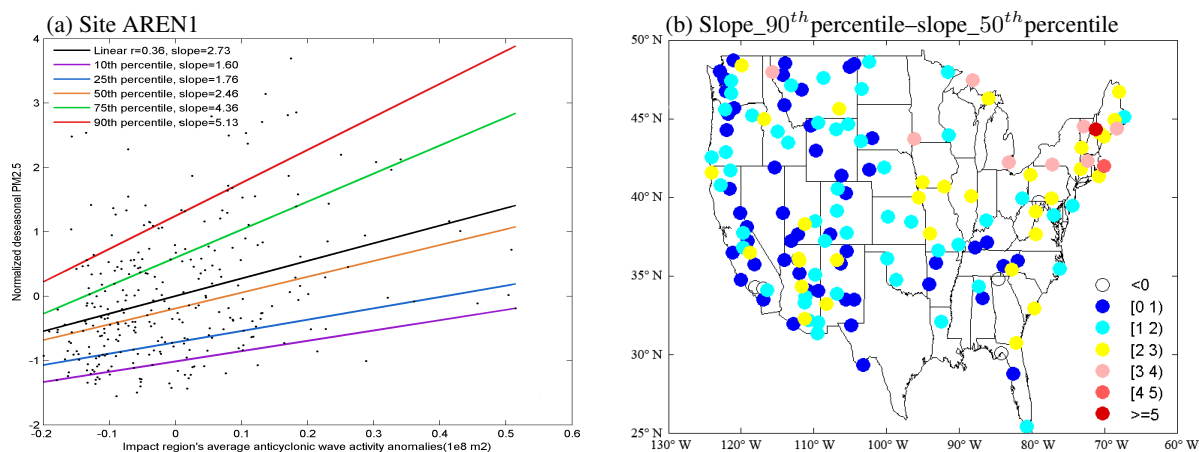


Figure 7. (a) Scatterplot for site AREN1 JJA deseasonalized PM_{2.5} and impact region's average AWA with linear and quantile regression results. (b) The subtraction of 50th percentile quantile regression slope from 90th percentile quantile regression slope between PM_{2.5} and impact region's average anti-cyclonic wave activity across all 150 sites in the U. S.

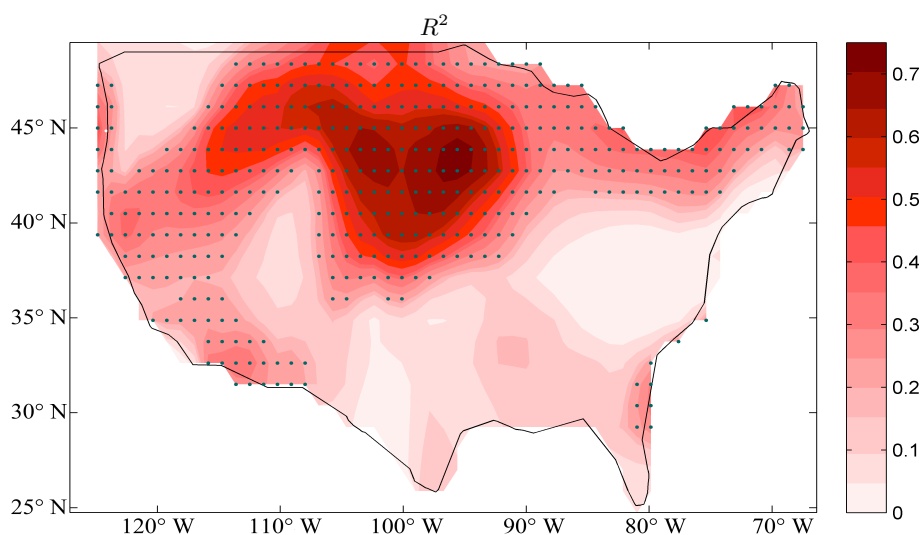


Figure 8. PM_{2.5}'s interannual variance explained (R^2) of the linear regression model using AWA projection value as the explanatory variable with modeled results. Stippling indicates where R^2 is significant (at 5% significance level) at model grids.

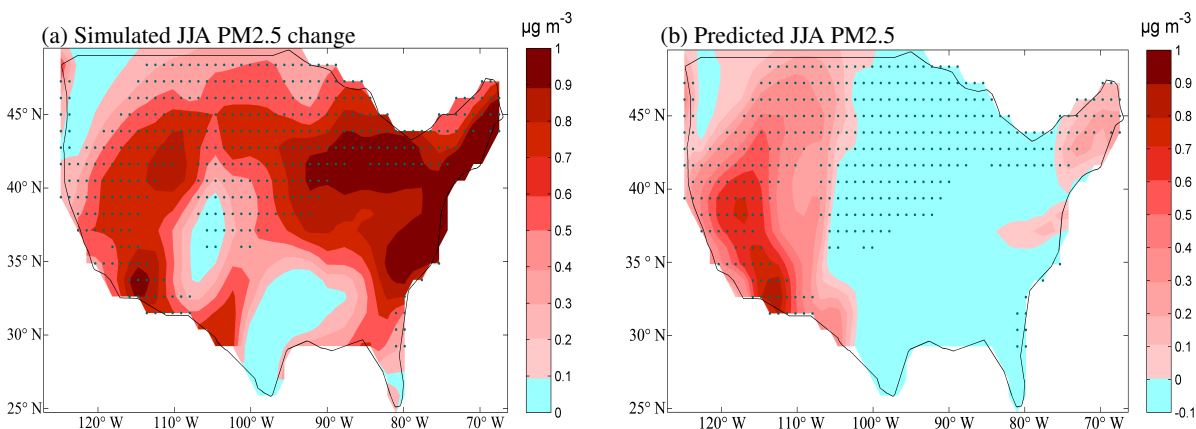


Figure 9. (a) Simulated change of JJA PM_{2.5} between simulated future (future climate with current emission, GCM2100, 2106-2125 mean) and current (current climate from the REFC2 simulation, GCM2000, 2006-2025 mean). (b) Predicted JJA PM_{2.5} change using the linear regression model fitted with simulated current PM_{2.5} (GCM2000, 2006-2025 mean). Stippling indicates where R^2 is significant (at 5% significance level) at model grids. Unit: $\mu\text{g m}^{-3}$ for PM_{2.5}.

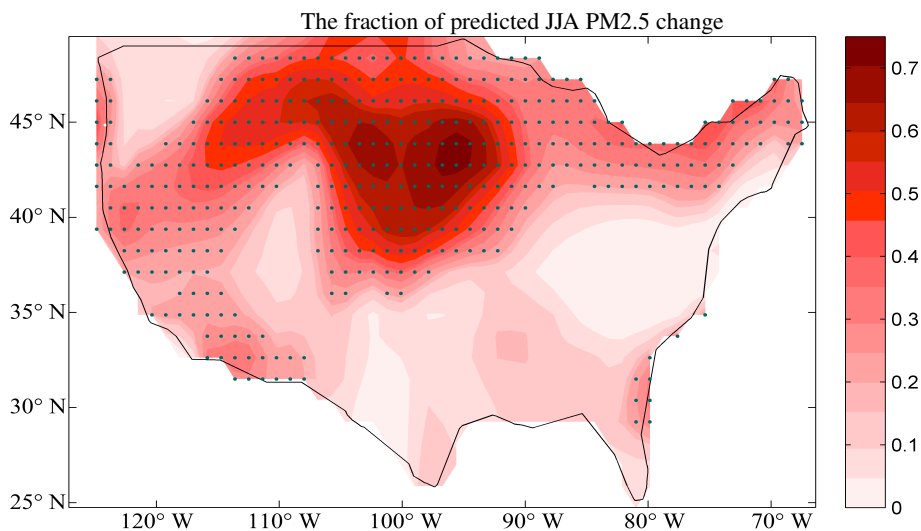


Figure 10. The fraction of predicted JJA PM_{2.5} change using simulated data (current climate from the REFC2 simulation, GCM2000, 2006-2025 mean): Figure 9a divided by Figure 9b) fitted model accounts for the total JJA PM_{2.5} change from simulations. The fraction that less than zero is regarded as zero.