



Exceptional middle latitude electron precipitation detected by balloon observations: implications for atmospheric composition

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Abstract. Energetic particle precipitation leads to ionization in the Earth's atmosphere, initiating the formation of active chemical species which destroy ozone and have the potential to impact atmospheric composition and dynamics down to the troposphere. We report on one exceptionally strong high-energy electron precipitation event detected by balloon measurements in middle latitudes on 14 December 2009 with ionization rates locally comparable to strong solar proton events. This electron

- 5 precipitation was likely caused by wave-particle interactions in the slot region between the inner and outer radiation belts, connected with still not well understood natural phenomena in the magnetosphere. Satellite observations of odd nitrogen and nitric acid are consistent with wide-spread electron precipitation into magnetic midlatitudes. Simulations with a 3D chemistry-climate model indicate almost complete destruction of ozone in the upper mesosphere over the region where high-energy electron precipitation occurred. Such an extraordinary type of energetic particle precipitation can have major implications for
- 10 the atmosphere, and their frequency and strength should be carefully studied.

1 Introduction

Energetic particle precipitation into the atmosphere initiates a chain of reactions starting with atmospheric ionization, leading to large changes in middle atmosphere composition, including the formation of hydrogen and nitric oxides followed by ozone loss in the stratosphere and mesosphere over \sim 30-80 km, and with potential relevance even for tropospheric weather

15 systems and regional climate (e.g. Seppälä et al., 2009; Mironova et al., 2015; Arsenovic et al., 2016; Sinnhuber and Funke, 2019). Permanent sources of atmospheric ionization are galactic cosmic rays and solar UV radiation, but the flux of energetic particles can increase by orders of magnitude through episodic precipitation of solar or magnetospheric energetic particles. The precipitation of electrons from the outer radiation belt is a consequence of the violation of the adiabatic motion of the trapped electrons, mostly as a result of the wave-particle interactions. Precipitation mainly occurs at high latitudes, in the zone





20 of the auroral oval corresponding to geomagnetic latitudes of $\sim 65-70^{\circ}$ or a McIlwaine parameter of L $\sim 5-6$. Comprehensive measurements of midlatitude electron precipitation from a slot between the outer and inner radiation belts at L $\sim 2-4$ have been made on the Van Allen Probes (e.g. Su et al., 2017; Foster et al., 2016).

Energetic electron precipitation (EEP) leads to the enhancement of odd nitrogen NO_x and odd hydrogen HO_x , which play a key role in the ozone balance of the middle atmosphere (e.g. Sinnhuber et al., 2012). The effect of high latitude EEP on atmospheric composition and ozone is confirmed by various observations (e.g. Newnham et al., 2011; Andersson et al., 2014; Newnham et al., 2013; Sinnhuber et al., 2016; Randall et al., 2006) and 3D chemistry-climate models (e.g. Rozanov et al., 2012; Arsenovic et al., 2016; Verronen et al., 2016; Sinnhuber et al., 2018) that account for EEP induced ionization.

- 30 Here, we present an exceptional case of high energy electron precipitation (with stratospheric and mesospheric ionization rates locally exceeding those of large solar proton events) from the slot region (2 < L < 4) observed over Moscow (55.96°N, 37.51°E, geomagnetic latitude ~52°N, L=2.7) on 14 December 2009. To confirm the balloon observations, and to bring those essentially point measurements into a broader context, Polar-orbiting Operational Environmental Satellites and VLF observations are studied as well. The EEP induced ionization and consequent enhancement of NO_x (N, NO) and HNO₃ are confirmed
- 35 by chemical composition observations from MLS (Waters et al., 2006) and MIPAS (Funke et al., 2014). Model studies with the 1D atmospheric chemistry model ExoTIC (Herbst et al., 2019) and the 3D chemistry-climate model HAMMONIA (Schmidt et al., 2006; Meraner et al., 2016) using the ionization rates derived from the balloon observations demonstrate formation and loss rates of a wide range of neutral species and ozone in the upper mesosphere.

40 2 Observations of middle latitude electron precipitation

2.1 Balloon observations

Balloon observations of energetic particle precipitation in the atmosphere are an important independent source of information for the evaluation of satellite-observed particle flux and energy used in chemistry-climate models, extending the useful energy range from hundreds of keV to several MeV.

- The balloon measurements are performed by the radiosonde lifted up to the heights of 30-35 km and returning information on the ionizing particle fluxes at different levels of the atmosphere. The radiosonde sensor consists of two Geiger-Müller tubes arranged as a telescope with a 2 mm Al interlayer between the tubes (Fig. 1b). The device returns the count rates of the upper single tube and the telescope. The single tube is sensitive to X-rays and charged particles (electrons, protons, and muons), while a telescope measures only energetic charged particles but does not respond to the X-ray flux by the atmosphere. During quiet
- 50 conditions, the radiosonde records the fluxes of secondary cosmic rays. Precipitating electrons are absorbed at altitudes above 50 km, but they generate X-rays via bremsstrahlung, which penetrates into the atmosphere down to altitudes of \sim 20 km and can be recorded only by single tube. Intrusion into the atmosphere of solar particles causes count rate enhancement both of a





single tube and the telescope, which enables us to distinguish between solar proton and magnetospheric electron precipitation.
In the case of smooth growth of the Geiger-Müller tube count rates with altitude, we assume that it is caused by X-ray absorption in air rather than by temporal variations of X-ray flux. Taking the data of a previous balloon flight in quiet conditions as the background and subtracting it from the data of the flight that observed precipitating electrons, we get the X-ray flux vs. atmospheric pressure. The method of evaluation of the energy spectrum of electrons impinging on the atmosphere from the X-ray flux absorption in air was developed on the basis of the GEANT 4 simulation (Makhmutov et al., 2016).

In this study we use observations from the balloon experiment performed by the Lebedev Physical Institute (LPI) every few days since 1957 (Stozhkov et al., 2009), which has so far recorded 589 EEP events at polar latitudes, L=~5.5, over 1961-2019 (Makhmutov et al., 2016; Mironova et al., 2019a, b; Bazilevskaya et al., 2020), and is complemented by regular balloon launches at middle latitudes. Observations of EEP events in middle latitudes are very rare. However, several candidates have been found since the beginning of the 2000s which have not been studied properly yet. Here we present the most outstanding

EEP event recorded in the Moscow region so far.

Data from the LPI balloon observation at 13:26-13:45 UT on 14 December 2009 (Fig. 1c, curve 1) demonstrate a substantial enhancement in the count rate of the single Geiger-Müller tube above \sim 20 km (residual pressure \sim 55 hPa). This count rate increase of the single Geiger-Müller tube was due to X-ray bremsstrahlung generated by precipitating electrons in the

70 atmosphere at altitudes above 50 km. A typical quiet-day result derived on 11 December shows only background in the count rate of the single Geiger-Müller tube (Fig. 1 c, curve 2). The telescopes (Fig. 1c, curves 3 and 4), which are not sensitive to X-rays, recorded the background due to secondary cosmic rays, confirming particle precipitation as the source of the single tube count rate increase. At the polar station Apatity (67.57N, 33.56E, L=5.3), a radiosonde was aloft ~5 hours before the Moscow observation and did not observe enhanced electron precipitation.

75 2.2 POES observations

NOAA's Polar-orbiting Operational Environmental Satellites (POES) carry a suite of instruments that measure the flux of energetic protons and electrons at the altitude of the satellite. The Medium Energy Proton and Electron Detector (MEPED) onboard POES consists of telescopes pointing close to zenith (0°) and in the horizontal plane (90°). We have examined the POES data around the Moscow and Apatity regions for December 2009 with the following limitations: (i) McIllwain parameter L = 2-3,

80 foot-of-field-line latitudes Flat = 52-60°N, foot-of-field-line longitudes Flon = $30-55^\circ$, and (ii) L < 8, Flat = $60-70^\circ$ N, and Flon = $30-55^\circ$. The >30 keV, >100 keV, and >300 keV electron channels as well as both the horizontal (90°) and vertical (0°) telescopes were checked.

On 14 December, the precipitation of >30 keV and >100 keV electrons was observed at polar latitudes close to the Apatity region at 05:34-05:35 (POES 16), 07:25-07:26 UT (POES 17) and 13:07-13:08 UT (POES 15). The precipitation was recorded also on 16 December (Fig. 2a). Although there was no strict coincidence with the balloon measurements, POES 16 occupied





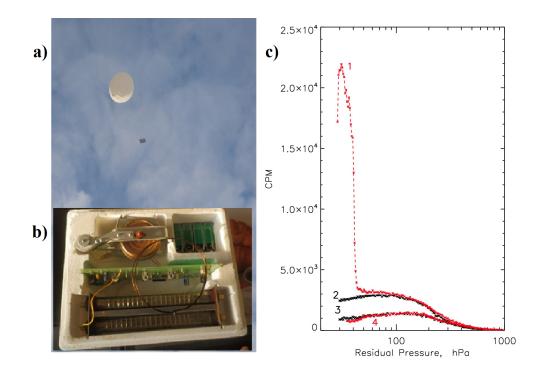


Figure 1. Left panel shows a balloon in flight a), and b) a radiosonde layout showing two Geiger-Müller tubes (a top counter referred to as "a single tube" and "telescope" arranged as a telescope detecting particles passing through both tubes and the filter between). Right panel c) shows the results (CPM – count rates of two Geiger-Müller tubes) of observations in the Moscow region on 14 December 2009 (curves 1, 4) and on 11 December 2009 (curves 2, 3). Curves 1 and 2 are the single Geiger-Müller tube count rates, sensitive to X rays. The telescope (sensitive to charged particles) count rate (curves 3 and 4) is multiplied by 3.

the closest position to Moscow at 13:48-13:49 UT on 14 December 2009. At the middle latitudes, there was no electron flux enhancement in the POES 16 vertical telescope data at this time, but the horizontal telescope channels of >30 keV and >100 keV electrons show increased particle flux on this day (Fig. 2b) as well as on 6 and 23 December.

90 2.3 VLF observations

Man-made, narrow-band radiowaves are transmitted in the Very Low Frequency (VLF) range from several, mainly midlatitude, locations around the world, particularly in the Northern Hemisphere. The radiowaves can propagate very long distances subionospherically, and are systematically recorded using a network of VLF receivers known as the Antarctic-Arctic Radiation-belt (Dynamic) Deposition - VLF Atmospheric Research Konsortium (AARDDVARK). Each receiver site is able to

95 log the amplitude and phase of 10 or more transmitter signals, with time resolution of typically 0.1-1 s. Perturbations to the phase and amplitude of the signals are caused by changes in ionization levels at altitudes close to the lower boundary of the D-region (50-85 km). Such perturbations can be caused by energetic particle precipitation (electron or proton), as well as solar flares. Determination of energetic particle precipitation characteristics from VLF perturbation levels requires knowledge of





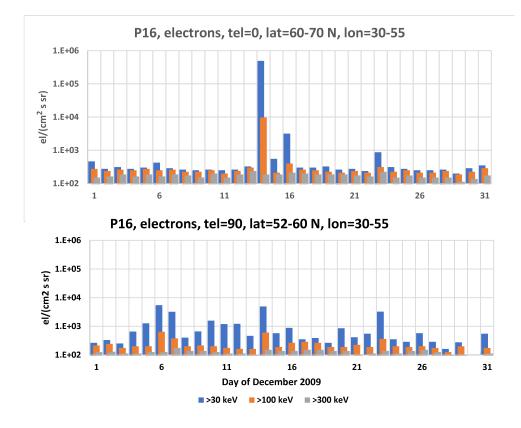


Figure 2. Daily data of POES 16 in December 2009. a) Data on electrons recorded by the vertical telescope at polar latitudes. b) The same but for the horizontal telescope at middle latitudes section.

either the flux of particles or the energy spectrum involved. Knowing one of these parameters allows the other to be calculated.See Rodger et al. (2012) for a detailed description of the calculations required and VLF perturbation responses that are likely to arise.

Subionospheric VLF propagation measurements from several receivers from the AARDDVARK located in the region of Scandinavia (Clilverd et al., 2009) showed a clear burst of precipitation from 13:30-15:00 UT, although the field of view did not include the region around Moscow. The propagation paths impacted by the precipitation spanned the 3<L<8 range, near to the geomagnetic latitude of Moscow.





3 Atmospheric response during the disturbed period

Geomagnetic disturbances 3.1

The main driver of energetic electron injection in the Earth environment is the solar wind interaction with Earth's magne-110 tosphere and related geomagnetic disturbances. Electron precipitation at polar latitudes is usually accompanied by enhanced auroral activity indicated by the auroral electrojet AE index, a substantial variation of the disturbance storm time index Dst, and the southward excursion of the Bz component of the interplanetary magnetic field. For this reason we took into account the behavior of AE, Dst, and Bz during December 2009. All these hourly-averaged parameters used in our study are collected by the Low-Resolution OMNI data set (King and Papitashvili, 2005).

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Energetic electron precipitation on 14 December occurred during a period of enhanced auroral activity as indicated by values of AE index larger than 200 nT, but with moderate levels of geomagnetic activity, indicated by negative values -12 nT of Dst, see panel a) of Figure 3. The southward orientation of the interplanetary magnetic field in the near-Earth space, indicated by negative Bz, emphasizes that electron precipitation was possible at the time. A similar combination of negative Bz, moderately negative Dst and enhanced auroral activity occurred during 5-6 and 23-27 December. However, on 14 December, the period

of southward oriented interplanetary magnetic field directly related to electron injection was strongest and most prolonged (Fig. 3a).

3.2 Satellite observations of trace gases (MLS and MIPAS)

Energetic electrons precipitating into the atmosphere decelerate by collisions with the most abundant species. In the middle 125 atmosphere below ~ 90 km, these are N₂ and O₂, which are either ionized or dissociated, starting a complex ion-chemistry reaction chain which ultimately leads to the formation of nitric oxide NO and nitric acid HNO₃, see Sinnhuber and Funke (2019) for a recent review.

The Earth Observing System (EOS) Microwave Limb Sounder (MLS) (Waters et al., 2006) is an instrument on NASA's Aura satellite, launched in July 2004. MLS observes millimeter and submillimeter-wavelength thermal emission, vertically scanning 130 Earth's limb in the orbit plane from the ground to about 90 km to give daily near-global (82°S-82°N latitude) coverage with \sim 15 orbits per day, making measurements during both day and night. Aura is in a sun-synchronous orbit with an ascending (north-going) equator-crossing time of 13:45 local time; it therefore passes the latitude of Moscow shortly after noon local time. One of the important MLS retrieval products that controls stratospheric ozone depletion is nitric acid (HNO₃). However, due to the relatively poor precision of HNO_3 in the upper stratosphere and lower mesosphere, only zonal average data can be 135

used. Here we use the latest version 5 MLS HNO₃ measurements (Livesey, N. J., Read, W. G., Wagner, P. A., Froidevaux, L., Lambert, A., Manney, G. L., et al., 2020).





The Michelson Interferometer for Passive Atmospheric Sounding (MIPAS) on board ENVISAT measured mid-infrared emis-140 sion spectra in the middle and upper atmosphere during 2002-2012, enabling the retrieval of temperature and a large number of trace species with daily global coverage (Fischer et al., 2008). In this study we use IMK/IAA MIPAS NO_x (N, NO) data (version V5R NO 220) (Funke et al., 2014). ENVISAT was in a sun-synchronous orbit with an equator crossing time of 10/22 hours local solar time; it therefore passed over Moscow several hours before and several hours after the balloon observation of the electron precipitation on 14 December. Because of the fast horizontal transport in the mesosphere as shown in Figure 5d, a direct observation of the impact of a localized, short-lived event is unlikely. 145

We analyzed MIPAS NO_x and MLS HNO₃ at 68 km altitude at high latitudes (50°-81°N) and at geomagnetic midlatitudes $(10^{\circ}-55^{\circ})$ geomag. lat.), see Figure 3b and 3c. Selection of these latitudinal-longitudinal regions as well as the altitude of the observations was based on balloon energetic electron precipitation observations and HAMMONIA chemistry-climate model results.

Daily mean values at geomagnetic midlatitudes are highest for both species on 14 December. Given that detection of a localized transient event is unlikely, that an increase in NO_x was observed by MIPAS (and an increase in average HNO₃ by MLS) thus suggests either a number of events in different locations on this day, or an event covering a larger area not observable 155 by balloon observations alone. For NO_x , maximal values were also much higher than on any other day in December 2009; for HNO₃, a very noisy observation with a large spread, maximal values were not conclusive (not shown). Slightly higher values than on average in either NO_x or HNO_3 (or both) were also observed during the periods of negative Bz and substorm activity on 5-6 and 23-27 December. While these enhancements could not be attributed clearly to the location of Moscow and were not statistically significant, the coincidence in both species with negative Bz and substorm activity suggests electron precipitation 160 into geomagnetic midlatitudes on these days that was strongest on 14 December. A closer view of the distribution of enhanced

- NO_x values on 6, 14 and 24 December (Fig. 3d) shows enhanced values mostly within the auroral oval (over North America) on 6 December, as expected from a period of auroral substorm activity; on 14 and 24 December, the enhanced values occurred mostly southward of the auroral oval in an area reaching from North America over the Atlantic to Northern Europe, with a spread indicating either sporadic precipitation hot-spots at very low latitudes, or fast horizontal transport. 11 December is 165

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shown as a reference for a "quiet" day without precipitation.

Potential impact of the mid-latitude electron precipitation event on the ozone layer 4

To estimate the potential impact of the Moscow event on atmospheric composition, we used the 1D atmospheric chemistry model ExoTIC (Herbst et al., 2019) and the 3D chemistry-climate model HAMMONIA (Schmidt et al., 2006; Meraner et al., 2016). Using the ionization rates derived from the balloon observations (Fig. 4a), a model experiment was carried out with the ExoTIC ion chemistry model for the Moscow region, considering ionization from 12-20 UT, to provide formation and loss

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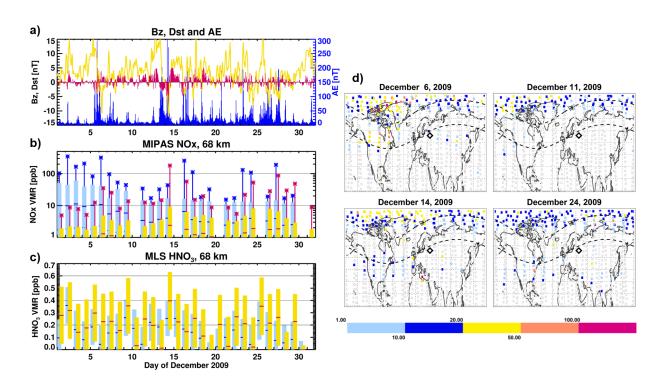


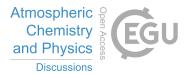
Figure 3. a) Bz (red), Dst index (yellow) and AE index (blue). b) MIPAS NO_x (NO+NO₂) at 68 km altitude at high latitudes ($50^{\circ}-81^{\circ}N$, blue) and at geomagnetic midlatitudes ($10^{\circ}-55^{\circ}$ geomag. lat., red); horizontal line segments mark the mean values, the error bars the 95% percentile, and the stars the maximal values of the day. c) MLS HNO₃ at 68 km altitude at high latitudes ($50^{\circ}-81^{\circ}N$, blue) and at geomagnetic midlatitudes ($10^{\circ}-55^{\circ}$ geomag. lat., red); horizontal line segments mark the mean values ($50^{\circ}-81^{\circ}N$, blue) and at geomagnetic midlatitudes ($10^{\circ}-55^{\circ}$ geomag. lat., red); horizontal line segments mark the mean values and the error bars the 2σ standard error of the mean. d) MIPAS NO_x on the satellite overpass footprints on four days (6, 11, 14 and 23 December) at 68 km altitude in the Northern Hemisphere. The dashed lines mark geomagnetic latitudes of 50° and 75° . The rhombi mark the position of Moscow. Colored symbols mark observations larger than the monthly mean plus one standard error; grey symbols are observations not statistically significant in this sense.

rates of a wide range of neutral species (Fig. 4b).

4.1 Ionization rates calculations

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Computation of ionization rate (*IR*) requires knowledge of the energy spectra and parameterization of ion production via ionization yield functions. The ionization yield function at the atmospheric depth is the number of ion pairs created by one precipitating electron with the initial energy E at the upper boundary of the atmosphere. The ionization rates (ion pairs $g^{-1}s^{-1}$) are computed as $IR(x) = \int Y(x, E) * F(E) dE$, where Y(x, E) are yield functions, F(E) is a flux of precipitating electrons at the top of atmosphere, x is atmospheric depth and E is energy of the considered particles. The limits of integration are defined by maximum and minimum energy of the considered electrons. During the EEP event observed over Moscow, ionization rates (*IR*)





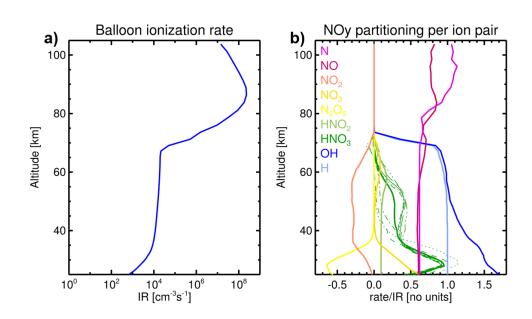


Figure 4. a) The atmospheric ionization profile derived from the balloon observations of 14 December 2009 over Moscow. b) Mean formation and loss rates of NO_y species due to ion chemistry calculated hourly from 12:00 - 20:00 UTC over Moscow with the ExoTIC ion chemistry model, showing formation of N, NO, H and OH as well as re-partitioning of NO_y species from NO_2 and N_2O_5 to NO_3 , HNO₂ and HNO₃. For HNO₃, individual hourly values are also shown in different line styles to highlight its strong diurnal variability, with distinctly higher values in the 35—60 km region during night-time.

180 (see Fig. 4a) are computed using a look-up table Y(x, E) with ion production for isotropic flux of precipitating monoenergetic electrons (Artamonov et al., 2017) and electron energy distribution F(E) proposed by balloon-borne observations (see Fig. 1c, curve 1). The background prescribed ionization rates used during December 2009 in the HAMMONIA model are based on the EEP spectra obtained by POES satellites and computed by Atmospheric Ionization Module Osnabruck (AIMOS v1.6, Wissing and Kallenrode (2009)).

185 4.2 ExoTIC ion chemistry model results

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The Exoplanetary Terrestrial Ion Chemistry model ExoTIC is a 1D stacked-box model of atmospheric neutral and ion composition. It is based on the UBIC model developed for the terrestrial middle atmosphere (Winkler et al., 2009), but has recently been generalized to planetary atmospheres with a wide range of orbital parameters, stellar systems, and base compositions (Herbst et al., 2019). Temperature, pressure, an initial atmospheric composition and particle impact ionization rates are prescribed externally. The particle energy is distributed to primary ions and excited species based on the atmospheric composition; 60 neutral and 120 charged species are considered, which interact due to neutral, neutral–ion, and ion–ion gas-phase reactions, as well as photolysis and photoelectron attachment and detachment reactions (Sinnhuber et al., 2012). The ion chemistry module called hourly from the base model uses an iterative chemical equilibrium approach and provides formation and loss rates of





all neutral species due to primary ionization, positive and negative ion chemistry which can be used as a parameterization for 195 global chemistry-climate models (Nieder et al., 2014).

The ExoTIC results indicate strong formation of NO_x (N, NO) throughout the middle atmosphere, formation of HO_x (H, OH), and re-partitioning of NO_u species in the altitude range where large positive and negative cluster ions form (below 75 km), with strong HNO₃ production in the upper stratosphere and lower mesosphere particularly during night-time, see Fig. 4b. These formation and loss rates were provided as input for the HAMMONIA global chemistry-climate model, which does not 200 include a detailed description of the ionospheric D-layer, thus allowing consideration of, e.g., the direct HNO₃ production from ion chemistry. As detailed information about the spatial and temporal characteristics of the event comes mainly from the balloon observations over Moscow, we limited the 3D model experiment to forcing by the available information; i.e., the ionization rates were applied only in the model profile above Moscow and only at the 6 hours prescribed by the balloon and POES observations.

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4.3 HAMMONIA chemistry-climate model results

The Hamburg Model for the Neutral and Ionized Atmosphere HAMMONIA is a revised version of the general atmospheric circulation model ECHAM5 (Roeckner et al., 2006), in which the upper boundary is raised to \sim 200-250 km or 1.7e⁻⁷ hPa. A detailed description of the model can be found in Schmidt et al. (2006) and Meraner et al. (2016). The system of hydrothermodynamic equations in the model is solved by the spectral method with triangular truncation T63, which approximately 210 corresponds to a horizontal resolution of $1.9^{\circ} \times 1.9^{\circ}$ in latitude and longitude. Vertical resolution is 119 levels. Here we use the model in specified dynamics mode, assimilating ECMWF ERA-Interim data up to 1 hPa. HAMMONIA includes the MOZART3 package to describe atmospheric chemistry (Kinnison et al., 2007). Background ionization rates from auroral and medium-energy electrons and solar protons as well as heavier ions are provided by the Atmospheric Ionization Module Osnabruck AIMOS v1.6 (Wissing and Kallenrode, 2009) with a two-hourly resolution. Ionization effects are described by the 215 5-ion chemistry scheme in the thermosphere (Kieser, 2011) and by the parameterization of NO_x and HO_x production by energetic particles in the middle atmosphere (Jackman et al., 2005) below ~90 km. Since HAMMONIA does not have a detailed ion chemistry treatment in the ionospheric D layer, the parameterization of Jackman et al. (2005) has been supplemented here by the formation and loss rates of neutral species estimated for the event from the 1D ExoTIC model. Two experiments were conducted with HAMMONIA: with just the background ionization rates from AIMOS and with the background plus the ion-220 ization rates estimated from the balloon observations over Moscow on 14 December 2009.

Since we are interested in determining the maximum potential atmospheric impact of the observed midlatitudinal EEP, we estimate the effects with the HAMMONIA model (see Figure 5), applying spatial extreme statistics (global or zonal maximum 225 and minimum values) instead of averaging globally or over a certain region. This approach is justified by the forcing localization and the 3D dynamics of the middle atmosphere, which quickly transports the anomalies induced in chemical species away from their source region. NO_x produced during the event is 1-2 orders of magnitude larger than the unperturbed maximum values





in the middle and upper mesosphere above ~ 60 km (Fig. 5a). The HNO₃ mixing ratio in the lower mesosphere reaches values of up to 2 ppbv during the event (Fig. 5b) at 55-68 km, 2-2.5 orders of magnitude larger than the unperturbed maximum 230 values at those altitudes. Modelled HNO₃ is additionally plotted at 68 km as maximum zonal values (Fig. 5d) to illustrate the meridional transport of the plume. Transport is mainly defined by the position and shape of the polar vortex, which is an oval with vertices extending to Europe and Alaska (not shown). The initial location of the plume is within the modeled vortex and it circled the globe in about 3 days. From Figures 5a and 5b, we can also see the downward propagation of the signal below 0.01 hPa and upward propagation above 0.01 hPa following the large-scale residual circulation. The downward propagation of the odd nitrogen produced by energetic particles is an important contributor to the stratospheric high-latitude ozone budget 235 (Sinnhuber and Funke, 2019). However, because the event is so localized in the model, this effect is indistinguishable in the global average. The modeled ozone response is caused, therefore, almost completely by the mesospheric HO_x enhancement, leading to the destruction of as much as 95% of the ozone in the upper mesosphere above 68 km over Moscow during the event (Fig. 5c).

Discussion and Conclusions 240 5

On 14 December 2009, surprisingly strong high-energy electron precipitation was observed clearly by midlatitude balloon measurements. Satellite POES data and VLF receivers confirm these electron precipitations and show that the EEP event extended over a larger area and continued for some time after the observed balloon event, moving northward. Midlatitude energetic electron precipitation can be triggered by wave-particle interactions in the slot region (2 < L < 4) between the inner and outer radiation belts. Inside the magnetosphere, the generation of waves called plasmaspheric hiss is especially intense 245 near the plasmasphere boundary (plasmapause). Here the electron scattering dominates the inward radial diffusion, resulting in an "impenetrable barrier" for electrons (Baker et al., 2014; Foster et al., 2016), which precipitate into the atmosphere. It is commonly accepted that the slot between the belts (L \sim 3) arises from electron scattering by the waves which can be of either natural or artificial origin initiated by VLF emission of man-made transmitters (e.g. Gombosi et al., 2017; Frolov et al., 250 2020; Zhao et al., 2019). However, natural or anthropogenic phenomena in the magnetosphere resulting in midlatitude electron precipitation still are not well understood.

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The energetic electron precipitation occurred during a period of mid-level geomagnetic activity with southward orientation of the interplanetary magnetic field in the near-Earth space. The atmospheric energy deposition during this event was much larger than expected for midlatitude precipitation due to, e.g., hiss forcing, and resembled in strength and altitude coverage large solar proton events. While the perturbations evident in the balloon observations are too short and localized to be directly detectable in the coarser resolution satellite measurements, the hemispheric response of atmospheric species like NO_x and HNO3 are in agreement with a midlatitude precipitation event on this day. Analysis of VLF subionospheric propagation perturbations shows evidence of precipitation during 04-16 UT on 14 December, with several bursts observed within 3<L<8,



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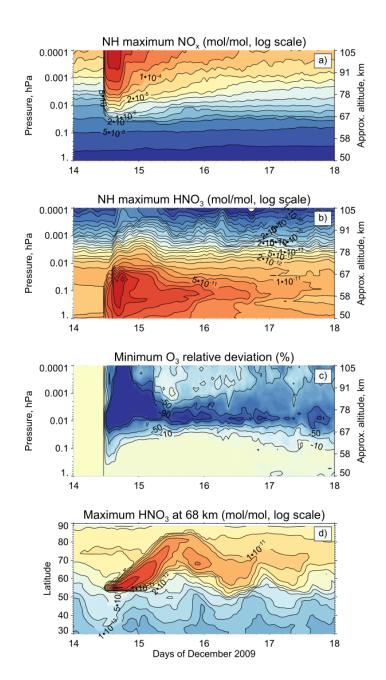


Figure 5. HAMMONIA results for a) Northern Hemispheric (NH) maximum value of NO_x volume mixing ratio (VMR). b) NH maximum value of HNO_3VMR . c) NH minimum value of the relative difference between runs with and without the event, and d) zonal maximum value of HNO_3VMR at 68 km.

260 including one at the time of the event observed over Moscow.





Both POES and VLF data on 14 December, as well as NO_x and HNO_3 observations throughout December, suggest that events indeed lasted for a few hours and covered extended areas during this time and that high-energy electron precipitation can occur even during relatively "quiet" periods without a geomagnetic storm.

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Satellite observations of NO_x and HNO₃ are consistent with precipitation into midlatitudes on several days in December 2009, with the strongest response on 14 December, highlighting that this was an exceptionally strong high-energy electron precipitation with ionization rates locally comparable to strong solar proton events. In the daily mean average (50°-81°N, 10-55° geomagnetic latitude) the increase is very small and not statistically significant considering the 95% percentile (NO_x) or 2σ standard error (HNO₃). However, averaging over large areas / large amounts of data (few hundreds to > 1000 profiles) mutes the maxima, thus increases in hotspots could be much larger. This is indicated in NO_x measurements, where maximal values of about 200 ppb are observed on 14 December compared to a mean value of 2-3 ppb, as well as by the results of the HAMMONIA model experiments. Complete destruction of ozone in the upper mesosphere over the region where high-energy electron precipitation occurred is also shown by HAMMONIA numerical experiments.

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The frequency, duration and spatial coverage of these newly discovered electron precipitation events are yet unknown, but results from first model simulations indicate a potentially large impact on atmospheric composition. If such EEP occur frequently and in a larger area over the middle latitudinal region, they could have an accumulated impact much larger than our model results (which assumed only one short, highly localized, event) suggest. Such midlatitudinal EEP events with ionization rates locally comparable to strong solar proton events could be recurrent and have major implications for the atmosphere. Thus,

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their frequency and strength should be carefully studied.

This conclusion inspires further studies involving a wider network of the balloon-based instruments.

285 Code availability. HAMMONIA chemistry-climate model: code and simulation results can be obtained by contacting TS (timofei.sukhodolov@pmodwrc.ch).
 ExoTIC ion chemistry model: code and simulation results can be obtained by contacting MS (miriam.sinnhuber@kit.edu).

 Data availability. Balloon observations: https://sites.lebedev.ru/ru/sites/DNS_FIAN/479.html POES: http://www.ngdc.noaa.gov/stp/satellite/poes
 OMNI: https://omniweb.gsfc.nasa.gov/form/dx1.html MLS: https://acdisc.gesdisc.eosdis.nasa.gov/data/Aura_MLS_Level2/ML2HNO3.004/ MIPAS: https://www.imk-asf.kit.edu/english/308.php AARDDVARK: http://psddb.nerc-bas.ac.uk/data/access/





Author contributions. IM, MS, GB, VM, ER, TS - discussed the idea and wrote the manuscript.

295 GB and VM -balloon measurements and spectra retrieval.

GB, IM and VM - EEP events selection.

IM - ionization rates calculation.

MS - ExoTIC ion chemistry model: code and simulation results.

ER and TS - HAMMONIA chemistry-climate model: code and simulation results.

300 Data analysis:

GB – preparation and analysis of POES data, IM and VM – analysis of geomagtetic indexes, MS and BF – preparation and analysis of MIPAS data, IM and MLS – preparation and analysis of MLS data, MC – preparation and analysis of VLF data. All authors discussed the results and commented on the manuscript.

Competing interests. The authors declare no competing interests.

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