

1 Long Range Prediction and the Stratosphere

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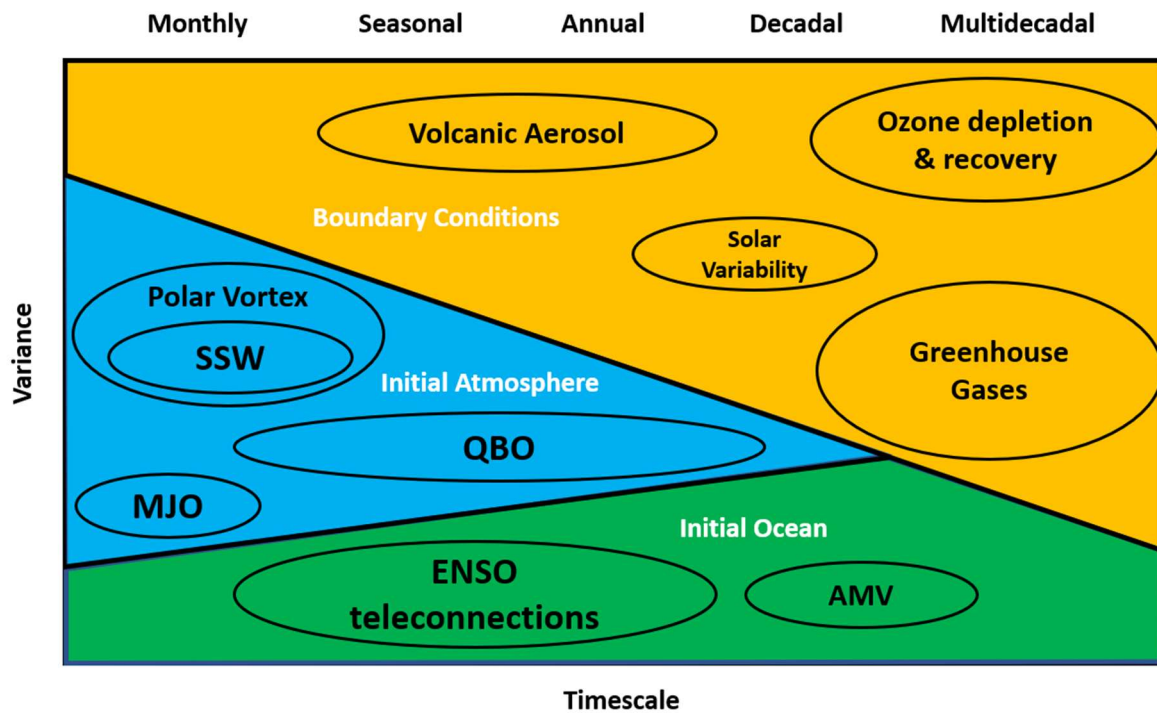
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37 **Abstract.** Over recent years there have been ~~parallel-concomitant~~ advances in the development of
38 stratosphere resolving numerical models, our understanding of stratosphere-troposphere interaction and
39 the extension of long-range forecasts to explicitly include the stratosphere. These advances are now
40 allowing new and improved capability in long range prediction. We present an overview of this
41 development and show how the inclusion of the stratosphere in forecast systems aids monthly, seasonal,
42 ~~annual to~~ decadal climate predictions and multidecadal projections. We end with an outlook towards
43 the future ~~of climate forecasts~~ and identify areas for improvement that could further benefit these rapidly
44 evolving predictions.

46 1 Introduction

47 ~~The climate system contains significant unpredictable variance and for daily weather fluctuations~~
48 ~~at least it is~~are thought to have a deterministic predictability horizon of around two weeks due to the
49 sensitivity of the evolution of the atmospheric state to small errors in initial conditions (Lorenz 1969) -
50 the so-called ‘butterfly effect’. Recent estimates (Leung et al., 2020; Domeisen et al., 2018) as well as
51 tests of the predictability of midlatitude *daily* weather using the latest global prediction models (Zhang
52 et al., 2019; Son et al., 2020) produce similar estimates for this predictability limit. However, this does
53 not preclude skilful forecasts of the *statistics* (most notably the average) of conditions at long range
54 beyond this timescale (e.g. Shukla 1981). This predictability owes its existence to slowly varying
55 predictable components of the climate system in the ocean, and in some cases the atmosphere, as well
56 as externally forced changes such as volcanic or solar variability effects (e.g. Kushnir et al., 2019).
57 Some of the more prominent examples of stratospheric variability such as sudden stratospheric
58 warmings and their subsequent impact on the stratosphere and interaction with the troposphere
59 (Baldwin et al., 2021), ~~or~~and the quasi-biennial oscillation and its associated teleconnections (Scaife et
60 al., 2014a), have been shown to fall into this predictable category be predictable out to timescales well
61 beyond the traditional two-week predictability horizon from initial tropospheric conditions alone. Other
62 examples involve stratospheric pathways for teleconnections originating in the troposphere or ocean
63 (e.g., Schwartz and Garfinkel 2017, Byrne et al. 2019). On longer timescales, boundary forcing, for
64 example from composition changes such as ozone depletion and recovery allows the stratosphere to
65 thereby provideing relatively slowly varying conditions to guide the turbulent troposphere and hence
66 provide long range predictability (e.g., Thompson et al. 2011) ~~of conditions beyond the two-week limit.~~
67 The relative importance of stratospheric initial conditions to boundary conditions decreases with lead
68 time as shown in the schematic in Figure 1.



69

70 **Figure 1: Schematic representation of the role of the stratosphere in long range prediction showing the**
 71 **transition from initial condition predictability in the atmosphere (blue) and the ocean (green), to boundary**
 72 **condition predictability at longer timescales (orange). Individual mechanisms involving the stratosphere**
 73 **are labelled in black showing their approximate contributions to forecast variance at different lead times.**

74 The extension of long-range prediction systems to explicitly include representation of the stratosphere
 75 follows many years of development of stratosphere resolving general circulation models (GCMs). By
 76 the late 20th century many leading centres for climate research had started to include the stratosphere in
 77 versions of their GCMs (Pawson et al., 2000; Gerber et al., 2012). Much of the early model development
 78 was motivated by the discovery of the ozone hole in the 1980s (Farman et al., 1985) and the need for
 79 simulations of ozone depletion and potential recovery of the ozone hole following the 1987 Montreal
 80 Protocol, which required atmospheric models that represented both the atmospheric dynamics and
 81 chemistry of stratospheric ozone depletion (Molina and Rowland 1974; Crutzen 1974). In most cases
 82 this was achieved by adding further quasi-horizontal layers to the domain of existing climate models to
 83 extend their representation of the atmosphere to the stratopause or beyond (e.g. Rind et al 1988; Beagley
 84 et al., 1997; Swinbank et al., 1998; Sassi et al., 2002), while also incorporating key radiative (e.g. Fels
 85 et al., 1985), chemical (e.g. Steil et al., 1998) and dynamical (e.g. Scaife et al., 2000) processes.

86 The early development of so called ‘high top’ climate models, which represent the whole depth of the
 87 stratosphere, in general preceded the discovery of the main body of evidence that the variability of the
 88 stratosphere is not only affected by, but also interacts with the lower atmosphere and surface climate.
 89 Pioneering early studies suggested that the stratosphere might have direct effects on the troposphere
 90 and surface climate (e.g. Labitzke 1965; Boville 1984; Kodera et al., 1990, 1995; Haynes et al., 1991;
 91 Perlwitz and Graf 1995). In subsequent years, as reliable observational records lengthened and large
 92 enough samples of stratospheric variability were amassed it was unequivocally demonstrated that
 93 stratospheric variability precedes important tropospheric changes in the extratropics (Baldwin and
 94 Dunkerton 1999, 2001). There was debate about causality and whether the stratosphere really does
 95 affect the atmosphere below (e.g. Plumb and Semeniuk 2003). However, experiments where the
 96 stratosphere is perturbed in numerical models show changes in surface climate and reproduce similar
 97 patterns of response at the surface to those found in real world observations (e.g. Polvani and Kushner
 98 2002; Norton et al., 2003; Scaife et al., 2006; Joshi et al., 2006; Scaife and Knight 2008; Hitchcock and
 99 Haynes 2016, White et al., 2020). These involve changes to planetary scale waves and also baroclinic

100 eddies in the troposphere that are consistent with changes in baroclinicity near the tropopause (Kushner
101 and Polvani 2004; Song and Robinson 2004; Wittman et al., 2004, 2007; Scaife et al., 2012; Domeisen
102 et al., 2013; Hitchcock and Simpson 2014; White et al., 2020). Importantly, as we discuss below, the
103 same mechanisms also appear to be at work across a broad range of timescales (Kidston et al., 2015).

104 In recent years, motivated by the evidence of surface effects of stratospheric variability in the mid-
105 latitudes, the high-top model configurations used for stratospheric research were incorporated into
106 leading ~~long-range~~ prediction systems. Improved vertical resolution was already known to improve the
107 atmospheric data assimilation of satellite instrument observations whose sensitivity was often heavily
108 weighted towards stratospheric altitudes. This also provided initial stratospheric conditions for sets
109 of ~~This was initially done in retrospective forecast test experiments~~, some of which were internationally
110 coordinated (e.g. Butler et al., 2016; Tompkins et al., 2017). ~~A~~ However, a growing number of
111 operational systems are now producing regular ensembles of predictions at lead times of months or
112 years with coupled ocean-atmosphere models that extend to the stratopause or beyond; for example at
113 Environment Canada (Merryfield et al., 2013), the Met Office in the UK (MacLachlan et al., 2014), the
114 German Weather Service DWD (Baehr et al., 2015), the Japan Meteorological Agency (Takaya et al.,
115 2017) and the European Centre for Medium Range Weather Forecasts (Johnson et al., 2019). In the
116 following sections we document the emerging impacts and benefits of this new capability for surface
117 climate predictions at monthly, seasonal, and annual to decadal lead times starting with the shorter-
118 range, initial condition cases and ending with the longer-range boundary-condition cases.

119

120 **2 The stratosphere and monthly prediction**

121 The best-established phenomenon that gives rise to predictability of surface climate from the
122 stratosphere is the tropospheric circulation changes that follow strong and weak conditions in the
123 stratospheric polar vortex (Baldwin and Dunkerton 1999, 2001). For example, weak vortex conditions
124 such as those found in a sudden stratospheric warming (SSW, Baldwin et al., 2021) are typically
125 followed by a weakening and southward shift of the tropospheric mid-latitude jet stream (see e.g.
126 Kidston et al., 2015 and references therein) and thus the negative polarity of the North Atlantic
127 Oscillation (NAO), Arctic Oscillation (AO) and Northern Annular Mode (NAM). These fluctuations
128 also show a tendency to vacillate between strong westerly and weak (SSW) states on subseasonal
129 timescales (Kuroda and Kodera 2001; Hardiman et al., 2020a). The changes in the troposphere persist
130 roughly as long as those in the lower stratosphere, and last for around two months (Baldwin and
131 Dunkerton 2001; Baldwin et al., 2003; Hitchcock et al., 2013; Son et al., 2020; Domeisen 2019). The
132 impacts on surface climate also include changes in the frequency of extremes of temperature and
133 rainfall (Scaife et al., 2008; King et al., 2019; Cai et al., 2016; Domeisen et al., 2020b).

134 Although *major* SSW events, involving a complete reversal of the zonal flow in the mid stratosphere,
135 are rare in the southern hemisphere (Wang et al., 2020; Jucker et al., 2021), variations of the Antarctic
136 polar vortex are likewise followed by similar signatures in the underlying tropospheric flow, in this case
137 via the Southern Annular Mode (SAM). Weakening of the vortex is typically followed by a negative
138 shift in the SAM and associated changes in rainfall and near surface temperature (Thompson et al.,
139 2005; E. Lim-E. et al., 2018, 2019, 2021, Rao et al., 2020e). These changes in Southern Hemisphere
140 circulation typically take longer to reach the surface than their Northern Hemisphere counterparts
141 (Graverson and Christiansen 2003), perhaps due to the stronger stratospheric polar vortex and weaker
142 wave driving in the southern hemisphere, but they are nonetheless better predicted by improving
143 stratospheric resolution of forecast models (Roff et al., 2011). The timescale of weeks for the
144 predictability of sudden warmings is limited by the predictability of weather patterns in the troposphere
145 which might trigger SSW events (e.g. Mukougawa et al., 2005; Taguchi 2016; Garfinkel and Schwarz
146 2017; Jucker and Reichler 2018; Lee et al., 2020a). However, if we add this timescale to the timescale
147 of a month or more for the persistence of lower stratospheric anomalies and their surface effects (e.g.

148 Baldwin et al., 2003; Butler et al., 2019), we arrive at the conclusion that on these occasions at least,
149 initial conditions in the atmosphere can provide predictability well beyond the usual two-week horizon
150 for daily weather in either hemisphere.

151 Predictability of the atmosphere at monthly lead times is also known to originate in part from the
152 Madden Julian Oscillation (MJO) in the troposphere and its teleconnection to the extratropics (e.g.
153 Vitart 2017). The circulation pattern associated with the MJO resembles a poleward and eastward
154 propagating Rossby wave with centres of action over the Pacific and extending into the Atlantic sector
155 where it also maps strongly onto the North Atlantic Oscillation. The lead time of around 10 days for the
156 impact of a change in the MJO to appear in the extratropical flow (e.g. Cassou 2008; Lin et al., 2009)
157 is also consistent with the timescale for poleward propagation of Rossby waves (e.g. Scaife et al., 2017).
158 It turns out that this tropospheric MJO teleconnection on monthly timescales also interacts with the
159 stratosphere (Garfinkel and Schwartz 2017). The MJO teleconnection to the North Pacific affects the
160 region most strongly associated with tropospheric precursors to SSW events, and consistent with this,
161 SSWs in the observational record have tended to follow certain MJO phases. The subsequent weak
162 vortex anomaly then propagates down to the troposphere (Garfinkel et al 2012), where it may strengthen
163 and prolong any existing negative NAO signal that is directly linked to the MJO via the troposphere
164 (Schwartz and Garfinkel 2017, 2020; Barnes et al., 2019).

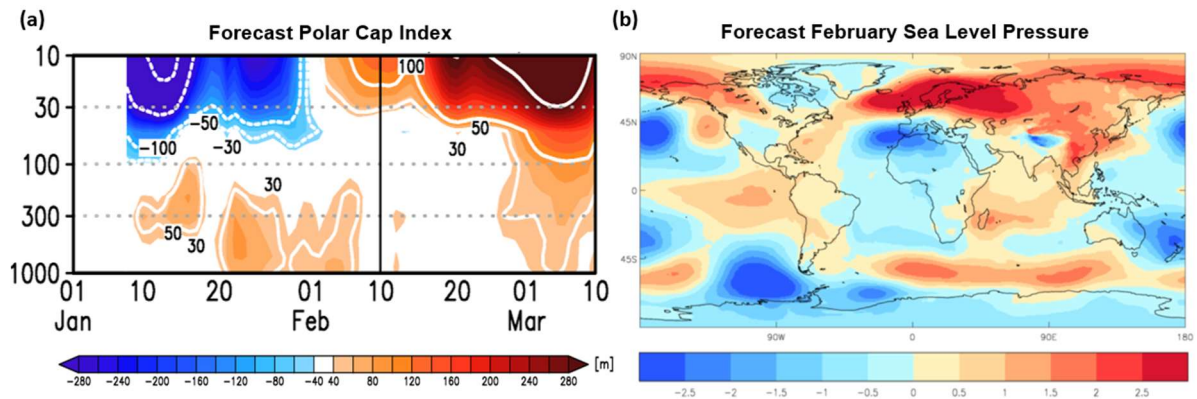
165 In addition to the interaction of the MJO with the extratropical stratosphere, a further, completely
166 different link between the stratosphere and the MJO has recently been uncovered which modulates MJO
167 amplitude and persistence in the troposphere via the phase of the Quasi-Biennial Oscillation (QBO) in
168 the tropical lower stratosphere (Liu et al., 2014; Yoo and Son 2016; Martin et al., 2021). In this case,
169 easterly phases of the QBO appear to energise the MJO compared to westerly QBO phases, likely due
170 to changes in temperature and hence static stability close to the tropopause (Hendon and Abhik 2018;
171 Martin et al., 2019) with a potential contribution of cloud-radiation feedbacks (Son et al., 2017, see
172 Martin et al., 2021 for a review). This modulation of the MJO is in turn important for predictability as
173 it gives rise to higher monthly prediction skill of the MJO and its surface teleconnections during the
174 easterly phase of the QBO (Marshall et al., 2017; Abhik and Hendon 2019; Y. Lim-Y. et al., 2019).

175 ~~Other mechanisms have also been found that can provide potentially predictable signals on the monthly~~
176 ~~timescale.~~ The traditional view of stratosphere-troposphere interaction involves upward propagation of
177 planetary scale Rossby waves (Charney and Drazin 1961), but this linear theory applies equally well to
178 downward propagation. Harnik and Lindzen (2001) and Perlwitz and Harnik (2003) identified a
179 possible source of downward propagating planetary waves in the form of reflecting surfaces in the
180 winter stratosphere. Examples of specific reflection events, showing upwards and then downward
181 propagation have since been observed (e.g. Kodera et al., 2008; Harnik 2009; Kodera and Mukougawa
182 2017; Mukougawa et al., 2017; Matthias and Kretschmer 2020). These results suggest that the details
183 of the stratospheric circulation such as regions of negative vertical wind shear could be important for
184 the formation of reflecting conditions (Perlwitz and Shaw 2013) and may yet provide a further
185 mechanism by which the stratosphere can affect the troposphere (Domeisen et al., 2019; Butler et al.,
186 2019).

187 Following studies demonstrating enhanced tropospheric predictability after SSW events in individual
188 climate models (e.g. Kuroda 2008; Mukougawa et al 2009; Marshall and Scaife 2010; Sigmond et al.
189 2013), subseasonal forecast systems which explicitly represent the stratosphere in the climate system
190 were developed and implemented at operational prediction centres worldwide. It is often difficult to
191 demonstrate significant increases in overall skill (e.g. Richter et al. 2020a) but routinely produced
192 ensembles of subseasonal predictions show that both stratospheric variability and its subsequent
193 tropospheric signature are predictable at monthly lead times (Domeisen et al. 2020a, 2020b). The
194 strongest surface impacts occur if the polar vortex in the lower stratosphere is in a weakened state at
195 the time of the SSW (Karpechko et al., 2017) and there appears to be a roughly linear relationship

196 between the strength of these lower stratospheric anomalies and the tropospheric response (e.g. Runde
 197 et al., 2016; White et al., 2020 and see Baldwin et al., 2019 for a review). We should note however that
 198 there is no one-to-one correspondence between stratospheric variability and tropospheric events, and
 199 some prominent examples of sudden stratospheric warmings are followed by differing tropospheric
 200 anomalies (e.g. Charlton-Perez et al., 2018; Knight et al., 2020; Butler et al., 2020; Rao et al., 2020a).
 201 Nevertheless, the canonical response is seen in the majority (~70%) of cases and periods of intense
 202 wintertime stratospheric variability are important windows of opportunity to provide skilful monthly
 203 forecasts (Mariotti et al., 2020; Tripathi et al., 2015a).

204



205 **Figure 21:** Monthly forecasts prior to the 2018 sudden stratospheric warming and severe cold event over
 206 northern Europe. Forecast Polar Cap Index (a) and February sea level pressure anomalies (b). **Forecasts**
 207 **ensemble mean anomalies are shown for the average of forecasts initialised between 8th and 22nd of were**
 208 **initialised in January 2018 relative to hindcasts over the 1993-2016 period (initialisation dates: 8th, 15th and**
 209 **22nd)** using the Met Office Hadley Centre GloSea prediction system (MacLachlan et al., 2015). Sea level
 210 pressure is measured in hPa and Polar Cap Index is the geopotential height anomaly (m) averaged over
 211 65N to the North Pole.

212 These forecast systems are now important tools for national meteorological and hydrological services
 213 to monitor impending stratospheric variability and associated surface impacts in real time. Recent
 214 extreme examples illustrate the importance of this activity. In February 2018 a major SSW occurred
 215 and was followed by a strong negative NAO-like pattern at the surface with easterly wind anomalies
 216 over Europe and multiple cold air outbreaks over the following weeks, including extreme snowfall
 217 across northern Europe (Figure 21, Karpechko et al., 2018; Knight et al., 2020; Rao et al., 2020a) and
 218 an abrupt end to Iberian drought in Southern Europe (Ayarzagueña et al., 2018). Studies of monthly
 219 ensemble predictions of this event with operational stratosphere resolving systems showed that the
 220 stratospheric event was predictable at least 2 weeks in advance (Figure 21) and that the ensemble
 221 forecasts indicated increased likelihood of cold surface conditions for several weeks after the event
 222 (Karpechko 2018; Butler et al., 2020; Statnaia et al., 2020; Rao et al., 2020a). Again, as in the analysis
 223 of previous events, there was also a strong association with the MJO entering Phase 7 with increased
 224 convection in the West Pacific (cf. Garfinkel and Schwartz 2017) in the 2018 event. Finally, we should
 225 also note that cases of monthly forecasts where the stratosphere plays an important role are not restricted
 226 to winters with sudden stratospheric warmings and periods when the stratospheric polar vortex is above
 227 normal strength also provide opportunities for skilful monthly forecasts (Tripathi et al., 2015b; Scaife
 228 et al., 2016). In this case an opposite but symmetric surface response results, with strong *positive* NAO.
 229 A very recent example occurred in February 2020, when, following an extremely strong polar vortex
 230 (Hardiman et al., 2020b; Lee et al., 2020b; Lawrence et al., 2020; Rao et al., 2021a), the tropospheric
 231 jet in the Atlantic sector strengthened, and the associated increased storminess and rainfall in this case
 232 resulted in UK monthly rainfall reaching a new record high (Davies et al., 2021).

233

234 3 The stratosphere and seasonal prediction

235

236 Prior to the advent of dynamical forecast systems which explicitly represent the stratosphere, seasonal
237 forecasts using empirical relationships and statistical methods were proposed. These relied on the prior
238 state of the polar vortex and other predictable factors such as the QBO that are known to have links to
239 surface climate (Thompson et al., 2002; Charlton et al., 2003; Christiansen et al., 2005; Boer and
240 Hamilton 2008). In some cases they indicated additional predictability that was absent in existing
241 operational forecast systems, providing further evidence of predictability involving the stratosphere and
242 further motivating the extension of dynamical forecast systems to properly represent the stratosphere.
243 Similar empirical forecast studies continue, and although they cannot provide evidence of predictability
244 that is as strong as from ~~forecasts using GCMs experiments~~ based on ~~fundamental~~ physical principles,
245 they do continue to be useful to indicate sources of predictability that need to be properly represented
246 in comprehensive forecast systems (e.g. Folland et al., 2012; Wang et al., 2017; Hall et al., 2017; Byrne
247 and Shepherd 2018).

248 Following the introduction of dynamical seasonal forecast systems with a good representation of the
249 stratosphere, clear links between successful seasonal prediction of the North Atlantic Oscillation, the
250 closely related Arctic Oscillation and the state of the stratospheric polar vortex have been identified in
251 forecast output (e.g. Scaife et al., 2014b; Stockdale et al., 2015; Jia et al., 2017; ~~Byrne et al., 2019~~).
252 Similar signals are also seen in the southern hemisphere in relation to predictability of the Southern
253 Annular Mode (Seviour et al., 2014; Byrne et al., 2019; Lim et al., 2021). Statistically significant
254 increases in overall skill directly attributable to the inclusion of the stratosphere in prediction systems
255 is sometimes difficult to demonstrate (e.g. Butler et al., 2016), especially given that other factors such
256 as horizontal resolution and physical parametrizations are often simultaneously changed. Nevertheless,
257 the body of evidence now weighs heavily in favour of predictability of the NAO and SAM from the
258 stratospheric polar vortex and from analyses showing reduced surface prediction skill in the absence of
259 stratospheric variability (e.g. Hardiman et al 2011; Sigmond et al., 2013; Scaife et al., 2016).

260 A second clear example of seasonal predictability originating in the stratosphere is the Quasi-Biennial
261 Oscillation (QBO). The QBO has such inherently long timescales that it persists for several months in
262 seasonal forecasts from initial atmospheric conditions alone and its regularity means that it can be
263 predicted from simple composites of earlier cycles. Nevertheless, a growing number of numerical
264 models used in seasonal forecast systems can now simulate and predict the oscillation within climate
265 forecasts (Garfinkel et al., 2018; Richter et al., 2020b; Stockdale et al., 2021) with the aid of forcing
266 from parametrized non-orographic gravity waves, and there is skill in predicting QBO phase changes
267 at lead times of a few months (e.g. Pohlman et al., 2013, Scaife et al., 2014a). ~~The~~A surface impact of
268 the QBO is also well established and has stood the test of time since it was first identified in the 1970s
269 (Ebdon 1975; Thompson et al., 2002; Anstey and Shepherd 2014; Gray et al 2018). Yet again this
270 response projects closely onto the North Atlantic Oscillation (and hence the Arctic Oscillation/Northern
271 Annular Mode) ~~in the northern hemisphere,~~ and the Southern Annular Mode ~~in the SH~~. The favoured
272 mechanism involves refraction of vertically propagating Rossby waves in the lower stratosphere
273 (Holton and Tan 1980), although other pathways may also be involved (e.g. Inoue et al., 2011;
274 Yamazaki et al., 2020; Rao et al., 2020**a**, 2021**b**). The observed magnitude of the QBO teleconnection
275 is also large enough to provide seasonal predictability of surface climate (Boer and Hamilton 2008) but
276 its modelled amplitude at the surface appears to be under-represented in current operational prediction
277 systems and models (Scaife et al., 2014b; Garfinkel et al., 2018; O'Reilly et al., 2019; Rao et al., 2020b;
278 Anstey et al. 2021).

279 In addition to the stratosphere acting as a source of predictability, other mechanisms by which the
280 stratosphere plays a role in seasonal predictions involve a pathway for global scale teleconnections.
281 These often originate in the tropics where the longer timescales of coupled ocean-atmosphere variability

282 such as the El Niño Southern Oscillation (ENSO, L’Heureux et al. 2020) provide a predictable source
283 of low frequency variability. Effects on the extratropics can occur by tropical excitation of anomalous
284 Rossby waves which propagate polewards but also upwards into the stratosphere, as in the case of
285 ENSO (Manzini et al., 2006; Domeisen et al., 2019), giving two pathways for extratropical influence
286 (Butler et al., 2014; [Kretschmer et al., 2021](#)). These highly predictable tropical sources of climate
287 variability alter the strength and position of the stratospheric polar vortex in the extratropics as well as
288 the frequency of SSWs (Polvani et al., 2017) and these are followed by changes in the seasonal westerly
289 jets in the troposphere and surface climate via the North Atlantic Oscillation (Ineson and Scaife 2009;
290 Cagnazzo and Manzini 2009) or the Southern Annular Mode (Byrne et al., 2019). As might be expected,
291 both the QBO and ENSO teleconnections are best represented in seasonal forecast systems which
292 contain a well resolved stratosphere (Butler et al., 2016). We note that new examples of the stratosphere
293 acting as a conduit for seasonal teleconnections are still being uncovered (Hurwitz et al., 2012, Woo et
294 al., 2015). For example, the Indian Ocean Dipole (IOD) received little attention in this context until the
295 recent record event of late 2019, when it appears to have driven an extreme winter strengthening of the
296 northern hemisphere stratospheric polar vortex. This strengthening took many weeks to decay, giving
297 rise to extreme yet highly predictable conditions in the stratosphere and around the Atlantic sector in
298 late boreal winter (Hardiman et al., 2020b; Lee et al., 2020b). The same event was also implicated in
299 extreme changes in the polar vortex and the near SSW in the southern hemisphere (Rao et al., 2020e);
300 an event that itself likely helped to drive the extreme summer conditions and wildfires over Australia
301 that year (Lim et al., 2021).

302 Apparent links between Arctic sea ice and seasonal winter climate in the mid latitudes have also been
303 suggested to be mediated by the stratosphere, with increased Rossby wave activity and a weakening of
304 the stratospheric polar vortex in response to reduced sea ice, especially in the Barents-Kara Sea (Jaiser
305 et al., 2013; Kim et al. 2014; King et al., 2016; Kretschmer et al., 2016). Some studies also reproduced
306 surface signals in response to sea ice anomalies in seasonal forecasts of particular years that are in
307 apparent agreement with observational estimates (e.g. Balmaseda et al., 2010; Orsolini et al., 2012).
308 However, recent updates to observational records show [a](#) weakening of these apparent effects
309 (Blackport and Screen 2020) and significant non-stationarity (Kolstad and Screen 2019). Subsequent
310 modelling studies with larger samples of simulations have provided mixed results (Zhang et al 2018;
311 Dai and Song 2020; Smith et al 2021), and some [have](#) argued that the atmospheric response to sea ice
312 is weak and that while the sensitivity to Barents-Kara sea ice may be stronger, the stratospheric response
313 in particular is highly variable (McKenna et al. 2017). While there may well be a longer-term effect via
314 the stratosphere from sea ice decline (Sun et al., 2015; Screen and Blackport 2019; Kretschmer et al.,
315 2020), sensitivity of the response to the background state complicates the issue (Labe et al., 2019; Smith
316 et al. 2017), as do possible confounding influences from the tropics (Warner et al. 2020) and to date
317 there is no clear consensus for strong enough year--to--year effects to provide significant seasonal
318 predictability.

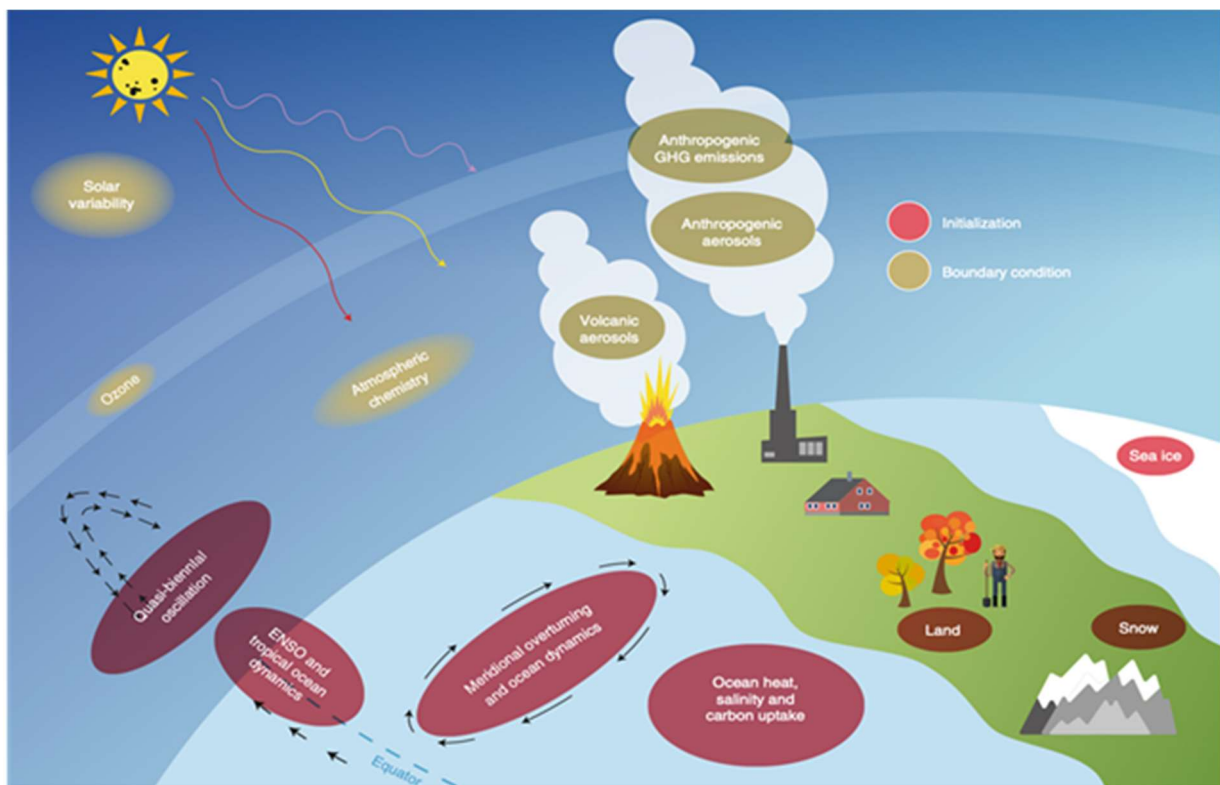
319 Other proposed teleconnections acting via the stratosphere have been found in observations but remain
320 to be confirmed with successful reproduction in physically based climate models. A prominent example
321 involves a proposed link between Eurasian snow amounts and the stratosphere, followed by a return
322 influence on the NAO and surface climate. In this case, enhanced snow cover or depth is associated
323 with high pressure over north Eurasia, an increase in the flux of Rossby wave activity into the
324 stratosphere and a subsequent weakening of the stratospheric polar vortex, followed by the expected
325 negative shift in the NAO and AO (Cohen and Entekhabi; 1999, Cohen and Jones 2011; Cohen et al.,
326 2014; Furtado et al., 2015). However, the strength of this link in climate models and seasonal predictions
327 is modest (Fletcher et al., 2009; Riddle et al., 2013; Tyrrell et al., 2018, 2019) and does not agree with
328 apparent links to the AO in observations (Kretschmer et al., 2016; Garfinkel et al., 2020) even when
329 model mean state biases are corrected (Tyrrell et al., 2020). It has also been suggested that
330 teleconnections to snow are non-stationary or non-causal and there is continued debate about its long-
331 term robustness (Peings et al., 2013; Henderson et al., 2018).

332 In summary, a number of mechanisms by which the stratosphere acts to provide seasonal predictability
 333 either by acting directly as a source of predictable variability (e.g. the QBO, SSWs), or as a conduit for
 334 teleconnections (e.g. ENSO, MJO, IOD) have now been established in observations and have been
 335 confirmed using climate model simulations based on first principles. These operate in seasonal forecast
 336 systems, albeit with remaining errors such as the weakness of the QBO connection to surface climate.
 337 Meanwhile, other mechanisms involving the stratosphere (for example the response to snow cover
 338 variations) have been proposed based on apparent observed relationships, but until we have agreement
 339 between these observations and theory (model simulations), scientists remain sceptical of whether they
 340 represent actual sources of seasonal predictability and these ~~topics~~ remain topics of current research.

341

342 4 The stratosphere and annual to decadal prediction

343 In recent years, initialised predictions on longer timescales were developed on the premise of multiyear
 344 memory in the ocean (e.g. Smith et al 2007), and following the development pathway mapped out by
 345 seasonal forecasts in the past, these are now being run operationally to produce real time multi-model
 346 forecasts (Smith et al., 2013). Kushnir et al., (2019) mapped out this operational development of annual
 347 to decadal predictions and highlighted a number of sources of predictability, some of which involve the
 348 stratosphere (Figure 32), but not all of which are fully represented in climate prediction systems.



349

350 **Figure 32. Sources of annual to decadal predictability, some of which involve the stratosphere through the**
 351 **response to external forcing, internal atmospheric dynamics, or ozone chemistry changes. After Kushnir et**
 352 **al., 2019.**

353 Despite common misconceptions, not all annual to decadal predictability stems from the ocean. Indeed,
 354 it has been clearly demonstrated that multiyear predictability of the QBO exists in current decadal
 355 predictions systems out to lead times of several years (Pohlman et al., 2013; Scaife et al., 2014^{ab}). This
 356 offers the prospect of a stratospheric contribution to multiyear predictability of the extratropics through
 357 the teleconnection with the Arctic Oscillation (Anstey and Shepherd 2014, Gray et al., 2018) and to

358 tropical predictability through links to the MJO (e.g. Martin et al., 2021) and wider tropical climate
359 variability (Haynes et al., 2021).

360 Although it is more important on multidecadal timescales (see below), external forcing of the
361 stratosphere can also act as a source of decadal predictability. Forced climate signals from changes in
362 greenhouse gases or stratospheric effects such as ozone depletion occur on a much longer timescale
363 than the lead time of decadal forecasts but their contribution to the skill of predictions is not trivial. For
364 example, it is not immediately obvious whether the slow changes from multidecadal forced signals
365 would simply be swamped by unpredictable internal variability on decadal timescales, rendering long
366 term external forcing changes useless for decadal predictions. However, this is not the case and long-
367 term forcing is now known to be an important source of decadal prediction skill (Smith et al., 2019,
368 2020).

369 External forcing involving the stratosphere on shorter timescales is also important for annual to decadal
370 predictions. The stratosphere has long been known to be influenced by volcanic eruptions, particularly
371 in the case of tropical volcanic eruptions which are powerful enough to inject significant quantities of
372 sulphur dioxide into the atmosphere. Here it reacts with water to form sulphuric acid and persists in
373 aerosol form, leading to predictable multiyear global surface cooling, tropical stratospheric warming
374 and an intensification of the westerly stratospheric polar vortex in the extratropics (Robock and Mao
375 1992). Although the sample of observed events is limited, modelling studies have reproduced an
376 observed post-eruption intensification of the westerly winds in the stratosphere and some impacts on
377 the surface Arctic Oscillation. However, generations of models have struggled to reproduce the two-
378 year persistence of volcanic effects seen in observations and the observed magnitude of the effect on
379 the winter AO (e.g. Stenchikov et al., 2006; Marshall et al., 2009; Charlton-Perez et al., 2013, Bittner
380 et al., 2016). In addition to these changes in the atmosphere, the intensification of stratospheric
381 westerlies and hence Arctic Oscillation also combines with surface cooling of the ocean to generate
382 predictable changes in the Atlantic meridional overturning circulation (Reichler et al., 2012) which can
383 extend the volcanic influence to decadal timescales (Swingedouw et al., 2015). Finally, although the
384 mechanism is debated, there is also evidence of a multiyear effect of tropical volcanic eruptions on
385 ENSO, presumably requiring the persistent radiative forcing that arises through the long residence time
386 of volcanic products, particularly sulphate aerosols, in the stratosphere. This reportedly increases the
387 frequency of El Nino events by a factor of two in the years following volcanic eruptions (Adams et al.,
388 2003), again suggesting an important source of multiannual predictability via the stratosphere.

389 A second source of multiannual predictability from external forcing originates from solar variability
390 and in particular the 11-year solar activity cycle. Although a number of alternative mechanisms have
391 been proposed (see Gray et al., 2010 for a review), the established mechanism for surface effects via
392 the stratosphere is the change in the polar vortex that results from changes in upper stratospheric heating
393 over the course of each cycle between solar minimum and solar maximum. Atmospheric wave-mean
394 flow interactions amplify the initial radiatively driven change and drive its descent to the troposphere
395 (Kodera and Kuroda 2002; Marsh et al 2007; Ineson et al., 2011; Givon et al., 2021), where changes in
396 the extratropical jets result in a negative (positive) Arctic Oscillation pattern following solar minimum
397 (maximum). There is also evidence that it contributes to interannual prediction skill (Dunstone et al.,
398 2016) and an interesting aspect that has emerged in recent years is the integrating effect of the ocean on
399 solar induced changes in the NAO via interannual persistence of ocean heat content anomalies which
400 leads to a lag of around 3 years ($\pi/2$ cycles) in the peak response, as would be expected if the ocean is
401 integrating the effects of a periodic solar forcing (Scaife et al., 2013; Gray et al., 2013; Andrews et al.,
402 2015; Thiéblemont et al., 2015). However, debate continues as to whether the solar signal is indeed
403 large enough to be detectable in observations in the presence of large internal tropospheric variability
404 (Chiodo et al., 2019).

405 Perhaps the longest known timescale for predictability from initial conditions, which also involves the
406 stratosphere, is the interaction of Atlantic Multidecadal Variability (AMV) with the stratospheric
407 circulation. The Atlantic has followed pronounced multidecadal variations over the last century (Mann
408 et al., 1995) and these variations are predictable out to years ahead (Hermanson et al., 2014). Some
409 studies link these variations to the stratosphere and the NAO/AO (Reichler et al., 2012; Omrani et al.,
410 2014). Indeed, the pronounced multidecadal increase in the surface NAO between the 1960s and 1990s
411 is strongly coupled to changes in the strength of the stratospheric polar night jet (Scaife et al., 2005).
412 Although current models simulate weak coupling between the AMV and the free atmosphere, this
413 coupling appears to increase with model resolution (Lai et al., 2021) suggesting that the links between
414 AMV, the stratosphere and the NAO offer potential for improved decadal scale prediction involving
415 the stratosphere.

416 The currently recognised role of the stratosphere in decadal forecasts of surface climate again appears
417 mainly via the impact on annular modes and, in the northern hemisphere, the North Atlantic Oscillation.
418 Indeed, while current decadal prediction systems are now able to produce skilful predictions of
419 variations in the NAO on multiyear lead times (Smith et al., 2019, 2020; Athanassiadis et al., 2020),
420 much work is still needed to attribute these variations to in these modes to external forcing or internal
421 variability ~~and to understand the interaction between boundary and initial conditions which blurs~~
422 ~~the simple distinction between the two, current decadal prediction systems are now able to produce~~
423 ~~skilful predictions of variations in the NAO on multiyear lead times (Smith et al., 2019, 2020;~~
424 ~~Athanassiadis et al., 2020).~~ These new results are important because they indicate new-found decadal
425 predictability of events like the high NAO of the 1990s which yielded a run of mild but wet and stormy
426 winters in northern Europe and the eastern USA. These winters are well known to have caused
427 significant impact for example on the insurance sector (Leckebusch et al., 2007) and coincided with the
428 longest observed absence of SSW events (Pawson and Naujokat 1999; Domeisen 2019). Given the
429 indications of coupled stratosphere-troposphere variations on decadal timescales (Scaife et al., 2005;
430 Omrani et al., 2014; Garfinkel et al., 2017; Woo et al., 2015), understanding the role of the stratosphere
431 in extratropical decadal predictions needs further investigation.

433 **5 The stratosphere and multidecadal projeediction**

434 The importance of the stratosphere for climate projeedictions on multidecadal timescales was generally
435 recognised before its role in predictions on shorter timescales. This is in part a legacy of the early
436 development of stratosphere-troposphere models for ozone depletion studies described in the
437 introduction for example. On these longer timescales, coupling between stratospheric composition,
438 thermal structure and atmospheric circulation gives rise to surface climate predictability.

439 Perhaps the best-known case for the stratosphere affecting multidecadal projeedictions of surface
440 climate is the influence of ozone depletion on the southern annular mode (SAM; Thompson and
441 Solomon 2002, 2005; McLandress et al., 2011; Polvani et al., 2011; Son et al., 2018) where decreasing
442 ozone in the late 20th century lead to a strengthened pole-to-equator temperature gradient, a stronger
443 stratospheric polar vortex and a shift to strong positive SAM phases at the surface. In this case, studies
444 again show the importance of stratospheric resolution to generate the full response, consistent with a
445 genuine downward influence (Karpechko et al., 2008). The associated poleward shift in the tropospheric
446 jet is connected to a delay in the spring breakdown of the stratospheric polar vortex (Byrne et al., 2017)
447 and delivered significant and prolonged changes in rainfall across many regions of the southern
448 hemisphere (Kang et al., 2011; Purich and Son 2012). Implementation of the Montreal Protocol in 1987
449 and subsequent reductions in the rate of ozone depletion mean that recovery of the ozone layer is now
450 expected over the coming decades and the reversible effects of this on the surface climate form an
451 important element of current multidecadal projeedictions (Thompson et al., 2011; Previdi and Polvani
452 2014; Solomon et al., 2016; Banarjee et al., 2020, Zambri et al., 2021) where they are expected to play

453 an important role alongside other changes in the southern stratosphere due to continuing increases in
454 greenhouse gases (Son et al., 2009; Barnes et al., 2012), some of which occur via the stratospheric polar
455 vortex in a similar way to those from ozone depletion and recovery (Ceppi and Shepherd, 2019).

456 The more limited effects of ozone depletion in the northern hemisphere meant that the role of the
457 stratosphere in multidecadal projeredictions took longer to become established. Some early studies found
458 potential amplification of positive Arctic Oscillation trends under climate change when the stratosphere
459 was included (Shindell et al., 2001). However, this was not borne out in later studies as simulations
460 with other fully coupled ocean-troposphere-stratosphere models suggested weakening of the
461 stratospheric polar vortex (e.g. Huebener et al., 2007). Subsequent studies with multiple models also
462 indicated a southward shift in the polar night jet with weakening high latitude winds and strengthening
463 subtropical winds (Scaife et al., 2012; Manzini et al., 2014). These changes result from increased
464 atmospheric wave driving of the winds which can overwhelm the cooling effect of greenhouse gases
465 (Karpechko and Manzini 2012) and can lead to important differences in future surface climate, for
466 example in regional rainfall in areas typically affected by the stratosphere via the Arctic Oscillation and
467 NAO (Scaife et al., 2012). There is still significant uncertainty due to the diversity of modelled
468 stratospheric responses to greenhouse gas increases (Manzini et al., 2014, Simpson et al., 2018, Zappa
469 and Shepherd 2017), and it has proved difficult to identify any clear change in the frequency of sudden
470 stratospheric warmings (Ayarzagüena et al., 2018, 2020; Rao et al., 2020c). This is perhaps due to the
471 competition between strengthening latitudinal temperature gradients near the tropopause and enhanced
472 meridional overturning in the mid stratosphere. There is also strong inherent unpredictable variability
473 from decade to decade in the frequency of SSW occurrence (Butchart et al., 2000; McLandress and
474 Shepherd 2009).

475 Other aspects of future climate change where the stratosphere plays a role have also been identified, for
476 example, in the debate over the response to future levels of Arctic sea ice. In this case it seems that the
477 response of the mid-latitude circulation involves a negative shift in the Arctic Oscillation (Screen et al.,
478 2018; Zappa et al., 2018; McKenna et al., 2018). This could again be amplified by interaction with the
479 stratosphere as some studies suggest that the stratospheric response is necessary for a large surface
480 response (Kim et al., 2014), while others highlight that the stratospheric interaction is sensitive to the
481 regional pattern of sea ice decline (McKenna et al., 2018), and still others show evidence of non-linear
482 stratospheric, and stratosphere-mediated surface response (Manzini et al., 2018), coincident with the
483 time when the Barents and Kara seas become ice-free (Kretschmer et al., 2020). Furthermore, studies
484 also indicate that the surface climate response to sea ice decline depends systematically on the phase of
485 the stratospheric QBO (Labe et al., 2019).

486 Although it is much less certain than anthropogenic climate change, there have also been suggestions
487 of a multidecadal decline of external solar irradiance which can impact multidecadal climate
488 projeredictions via the stratosphere. Previous multidecadal solar minima, so called ‘grand minima’, have
489 occurred in sunspot records and have been connected to the Little Ice Age period around the end of the
490 17th century using proxy and other data (Owens et al., 2017). Given recent weak amplitude 11 year solar
491 cycles, there are now suggestions of a future solar ‘grand minimum’ where the 11 year cycle described
492 above could become muted or even absent for a prolonged period (Lockwood et al., 2010). In this case,
493 the upper stratospheric cooling in the tropics and summer hemisphere can change the meridional
494 temperature gradient in a similar fashion to the 11 year cycle (Maycock et al., 2015) and leads to a
495 negative shift in the AO and; the NAO, and hence affects regional climate (Ineson et al., 2015).
496 However, in this case it appears that while regional changes could be significant, they are generally
497 much smaller than the surface warming due to anticipated levels of anthropogenic greenhouse gases
498 (Anet et al., 2013; Ineson et al., 2015; Maycock et al., 2015).

499 Finally, we note that although low frequency variability in teleconnections is observed (e.g. Garfinkel
500 et al., 2019), it is often unclear whether this is a systematic variation or simply due to sampling

501 variability of an underlying stationary process (Jain et al., 2018). ~~Nevertheless, there is also~~ growing
502 evidence for systematic climate change in some of the teleconnections by which the stratosphere enables
503 surface predictability. Under future climate change it appears that some of the teleconnections discussed
504 above may *strengthen* in amplitude. For example, ~~the connection between ENSO and the extratropical~~
505 ~~Atlantic/European sector—the strength of ENSO induced anomalies in the extratropical~~
506 ~~Atlantic/European sector~~ increases in future climate projections (Müller and Roeckner 2006; Fereday
507 et al., 2020). Similarly, recent analyses suggest that the MJO teleconnection to the extratropics increases
508 *in amplitude* under climate change (Samarasinghe et al., 2020). The same is also true of the extratropical
509 effects of the stratospheric QBO, where in this case, the ~~strength–amplitude~~ of the teleconnection *in*
510 *composite anomalies* doubles under future climate change (Rao et al., 2020d) despite the QBO itself
511 becoming weaker (Richter et al., 2020c).

512 ~~Providing they are not swamped by changes in interannual variability, these strengthening~~
513 ~~teleconnections suggest a growing importance of the stratosphere in surface climate prediction over the~~
514 ~~coming decades.~~

517 6 Outlook

518 Long range prediction has evolved quickly in recent years (Merryfield et al., 2017; 2020, Butler et al.,
519 2019; Meehl et al., 2021) and this rapid development is due in part to the improved representation of
520 stratospheric processes and stratospheric initial conditions in ensemble prediction systems. The long-
521 range forecast community originally focused on predictability from initial ocean conditions and this
522 remains the primary source of long-range predictability, for example from ENSO, but some of these
523 long-range prediction systems contained poor representations of the stratosphere. In the meantime,
524 those working in parallel on climate modelling of the stratosphere were rarely involved in initialised
525 long-range prediction, instead being driven primarily by the ozone depletion problem. Knowledge
526 exchange across fields is important in science and precursors to a new paradigm often occur when a
527 topic is investigated ~~by~~ *from* researchers from outside the field (Kuhn 1970). The crossover and
528 collaboration between long range prediction and stratospheric research communities is no exception
529 and *the interaction between these communities* has yielded rapid progress and new insights. Examples
530 where initial atmospheric conditions can provide predictability beyond the usually assumed limit have
531 been demonstrated, particularly for the extratropics but also for the tropics, and we now know that in
532 some situations, *for example when sudden stratospheric warmings occur, the* initial conditions ~~in the~~
533 ~~ocean have less impact than initial conditions~~ in the stratosphere *can have more impact than initial*
534 *conditions in the ocean* (Thompson et al 2002; Scaife and Knight 2008; Polvani et al 2017). This
535 suggests that initial atmospheric conditions *in the stratosphere* are likely to be more important for long
536 range forecasts than previously assumed (Mukougawa et al., 2005, 2009; Stockdale et al., 2015;
537 Noguchi et al., 2016, 2020a; Choi and Son 2019; O’Reilly et al., 2019; Nie et al., 2019), not least
538 because the overturning and breaking of Rossby waves in the stratosphere is followed by long lived
539 atmospheric anomalies due to synoptic scale eddy feedbacks that prolong the effects in the troposphere
540 (Kunz and Greatbatch 2013) ~~and enhance long range predictability~~ (Kang et al., 2011; *White et al.,*
541 *2020*). *More research on the initial conditions in the stratosphere might therefore help to reveal potential*
542 *for further improvements in prediction skill.*

543 A notable simplification to understanding the role of the stratosphere, at least in extratropical long-
544 range predictions, is its apparently seamless mechanism across different timescales and different
545 phenomena. Following the early ground-breaking studies showing surface impacts of stratospheric
546 variability (*e.g. Labitzke 1965, Boville 1984*) and a multitude of studies on individual teleconnections
547 between the stratosphere and surface climate, the projection of stratospheric ~~impacts–variability~~ onto

548 the Arctic Oscillation/North Atlantic Oscillation/Annular Mode circulation patterns across timescales
549 and hemispheres is now well established (see the review by Kidston et al., 2015). This suggests that
550 similar coupling processes occur between the stratosphere and troposphere from months to decades and
551 these processes ~~are responsible for~~ lead to some of the most intense extratropical climate extremes, in
552 winter in the northern hemisphere NH and in late spring/early summer in the southern hemisphere
553 (Karpechko et al., 2018; Fereday et al., 2012, Kautz et al., 2019; Domeisen and Butler 2020). While
554 studies point to changes in upper tropospheric baroclinicity and tropospheric eddy feedbacks as crucial
555 in these teleconnections, a full mechanistic understanding of how this occurs is still lacking.

556 Some, but not all, leading forecast systems now include a well resolved stratosphere with reasonable
557 representation of relevant processes such as the body force from sub-grid orographic and non-
558 orographic gravity waves. However, many outstanding problems remain. Although their number is
559 increasing, only a subset of current GCMs have the ability to simulate a realistic QBO beyond its decay
560 from initial conditions and it seems that all GCMs have problems with the fidelity of modelled QBO
561 teleconnections, which are either too weak or absent altogether (Scaife et al., 2014a; Kim et al., 2020;
562 Anstey et al., 2021). Even the relatively well studied ENSO teleconnection via the stratosphere to the
563 extratropics still has outstanding questions, such as whether the northern hemisphere stratosphere
564 exhibits more SSW events during the La Niña phase (Butler and Polvani 2011; Song and Son 2012).
565 This is not generally reproduced in modelling systems (Garfinkel et al., 2012) but occurred ~~again~~ in the
566 ~~most~~ recent La Niña winter ~~at the time of writing in of 2020/-2020~~21. Similarly, while the increased
567 monthly predictability from the MJO during the easterly phase of the QBO has been detected in monthly
568 forecast experiments, the QBO-MJO connection does not persist in longer predictions and simulations
569 with current models (Kim et al., 2020). Research and model development on stratosphere-troposphere
570 interaction, including tropical effects (Noguchi et al., 2020b), will no doubt lead to further progress in
571 resolving this issue (Haynes et al., 2021).

572 ~~In addition to teleconnection errors,~~ Errors in the modelled climatological mean mean biases in
573 stratospheric climate are inevitably present to varying degrees in even the latest climate models. The
574 common protocol of running a set of retrospective predictions to allow these mean biases to be estimated
575 and hence subtracted from real time predictions may well correct for much of this error. However, the
576 degree to which biases have a nonlinear, state dependent impact on the predictions is not fully
577 understood. In some contexts, the nonlinear impacts of biases may be minimal (Karpechko et al., 2021)
578 while others show sensitivity (Sigmond et al., 2008, 2010) and increases of prediction skill occur under
579 certain background conditions, for example during Easterly QBO phases (Taguchi 2018). Other
580 processes generally omitted from long range predictions include interactive variations of ozone and
581 other trace gases. Although reports of impacts and benefits have varied, it is thought that surface signals
582 on interannual timescales come mainly from dynamical rather than chemical changes (Seviour et al.,
583 2014; Harari et al., 2019). Nevertheless, some studies suggest detectable effects from interannual
584 variability of ozone and it may be that ozone fluctuations could help to amplify surface signals
585 (Karpechko et al., 2014; Son et al., 2013; Smith and Polvani 2014; Oehrlein et al., 2020; Hendon et al.,
586 2020), providing a further area for future development. Given that the cost of full atmospheric chemistry
587 schemes remains computationally expensive, it seems likely that simple parametrizations of ozone
588 chemistry (e.g. Monge-Sanz et al., 2021) would be valuable in this context.

589 We end with a pointer to an issue that has now been found to affect ~~all~~ long-range predictions from
590 monthly to seasonal to decadal and multidecadal timescales, particularly in the extratropics. So called
591 ‘perfect model studies’, which test the ability of models to predict their own ensemble members, are
592 now known to *underestimate* the true predictability of climate in some regions, particularly around the
593 Atlantic basin, and so models are better at predicting real world variations than they are at predicting
594 themselves. ~~This—the~~ so-called ‘Signal to Noise Paradox’ (Scaife and Smith 2018). ~~This is~~ at first
595 surprising, because perfect model prediction scores are often assumed to represent an upper (rather than
596 lower) limit for prediction skill of the real world. The problem can be understood in terms of

597 ~~unrealistically weak ensemble mean predictions (e.g. Eade et al., 2014) but w-~~Whether the stratosphere
598 is involved directly in the cause of this problem remains to be seen (Saito et al., 2017; Stockdale et al.,
599 2015), as it initially appears ~~first~~ in the troposphere rather than the stratosphere in long range forecasts
600 (Domeisen et al., 2020a) ~~and studies are undecided whether predictions of the stratosphere exhibit the~~
601 ~~same issue (Saito et al., 2017; Stockdale et al., 2015).~~ Nevertheless, the unrealistically weak amplitude
602 of ensemble mean predictions~~the same signal to noise issues~~ may well have the same root cause as
603 ~~the reason for~~ the weaker than observed amplitude of ~~many~~ modelled ~~tropospheric~~ teleconnections
604 involving to the stratosphere discussed in this review, including, for example, the under-representation
605 of the surface impact of the QBO. Resolving this problem will therefore likely amplify these signals,
606 provide greater levels of prediction skill, and further strengthen the role of the stratosphere in long range
607 predictions of surface climate.

608 Author Contributions

609 AAS wrote the draft manuscript. All other co-authors contributed relevant references and input to
610 revisions and editing of the manuscript. SWS ~~helped produce~~provided Figure 1a.

611 Acknowledgements

612 AAS and SCH were supported by the Met Office Hadley Centre Climate Programme funded by BEIS
613 and Defra. MPB was supported by the Natural Environment Research Council (grant number
614 NE/M006123/1). JHR was supported by the Regional and Global Model Analysis (RGMA) component
615 of the Earth and Environmental System Modeling Program of the U.S. Department of Energy’s Office
616 of Biological & Environmental Research (BER) via NSF Interagency Agreement 1844590. SN was
617 supported by the Japan Society for the Promotion of Science (KAKENHI, Grant Number: 19K14798).
618 EPL was supported by the Australian government’s National Environmental Science Program Phase 2
619 and the Victorian Water and Climate Initiative Phase 2. SWS was supported by the National Research
620 Foundation of Korea (NRF) grant funded by the South Korean government (Ministry of Science and
621 ICT) (2017R1E1A1A01074889). DWJT is supported by the US National Science Foundation Climate
622 and Large-Scale Dynamics program. Support from the Swiss National Science Foundation through
623 projects PP00P2_170523 and PP00P2_198896 to D.D. is gratefully acknowledged. ~~Support from the~~
624 ~~Swiss National Science Foundation through project PP00P2_170523 to D.D. is gratefully~~
625 ~~acknowledged.~~ CIG was supported by an European Research Council starting Grant under the European
626 Unions Horizon 2020 research and innovation program (Grant agreement 677756).

627

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