1	Responses to the comments by Referee #1
2	
3	
4	
5	Manuscript number: acp-2021-647
6	
7	
8	
9	Title: Enhanced upward motion through the troposphere over the tropical
10	western Pacific and its implications for the transport of trace gases from the
11	troposphere to the stratosphere
12	
13	
14	
15	
16	Author(s): Kai Qie, Wuke Wang, Wenshou Tian [*] , Rui Huang, Mian Xu, Tao Wang,
17	Yifeng Peng
18	
19	
20	
21	December 2021
22	

This is an interesting and useful study. However the scientific content, the quality of the study and its presentation should be improved. In particular, the text is in some parts very descriptive and technical. I suggest some major revisions before publication by ACP.

Re: Thank you very much for your helpful suggestions which help us improve our manuscript substantially. We have modified our manuscript according to the comments. Our point-to-point responses to the reviewer's comments are below:

30

31 General comments:

1) In general in the manuscript it is very often written 'we found a positive or negative 32 trend'. Please specify here your message by adding some numbers in the text (a trend 33 of xxx per year or a change of xxx within 60 years from 1958 to 2017). It would be 34 35 also very helpful to give the reader an impression whether these trends are of minor or major importance by adding some numbers from the literature for comparison. In 36 general, I am wondering that the results are not discussed more quantitatively (see 37 specific comments below). Further, please explain in detail how the trends 38 39 are calculated and how the El Niño Southern Oscillation (ENSO) is considered in calculating the trends. 40

Re: We thank the reviewer for the constructive comments. The quantitative results are added to the revised manuscript according to the referee's specific comments below. The methods of how the trends are calculated and how the impact of ENSO is evaluated are also described in the revised manuscript. The details are shown in the responses to the referee's specific comments below.

46

47 2) Figures: In general, the font size of the labels is very small and should be enlarged.
48 Further, the text in the figure captions is very similar to each other. Please give here
49 the reader more information which data or model simulations are shown and add
50 some explanation what is important or what is the main message of the figure.

51 Re: Thanks for the suggestion. The font sizes of the labels in each figure are 52 enlarged, and the figure captions are rephrased.

53

3) In Section 2 the used data sets and model simulations are described. However, I am
missing a bit more motivation for the reader to understand why these data sets and
model simulations are used. A bit more explanation would be helpful.

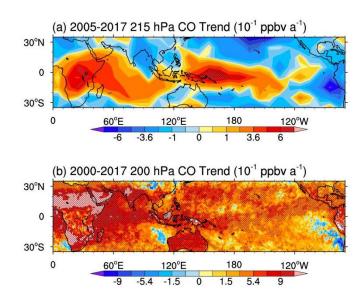
57 Re: Thanks for the comment. We have added some text to explain why the 58 datasets and model simulations are used in this study, and the descriptions about 59 the reanalysis datasets and model simulations are rephrased according to the 60 referee's specific comments.

61

3) The use of observations such as CO satellite measurements would strengthen the
main message of the manuscript. Therefore, I recommend to add some satellite data
(e.g. MLS CO https://mls.jpl.nasa.gov/eos-aura-mls/data-products/co)

65 Re: We thank the reviewer's good suggestion. An extra figure showing the trends of CO observed by MOPITT and MLS at near 200 hPa during 2000-2017 and 66 2005-2017 is added in the revised manuscript. The CO shows significantly 67 increasing trends over the TWP in NDJFM using MOPITT (at 200 hPa during 68 69 2000-2017) and MLS data (at 215 hPa during 2005-2017). The MLS CO data show that the area-averaged CO increased approximately 2.0±3.7 ppbv decade⁻¹ 70 over the TWP, while the CO increased 5.0±3.1 ppbv decade⁻¹ near the equator, 71 150°E at 215 hPa in NDJFM during 2005-2017 (Fig. R1). The area-averaged 72 MOPITT CO data increased at a rate of 5.0±3.1 ppbv decade⁻¹ at 200 hPa over 73 the TWP in NDJFM during 2000-2017. It should be pointed out that the linear 74 trends of CO are calculated based on the satellite data which only cover 14 or 18 75 years due to the data limitation here. Hence, the linear trends of CO may have 76 77 uncertainties particularly in the regions with large interannual variations in CO. To partially overcome this shortage, the trends of MLS CO at 215 hPa during 78

time periods of 2005-2016, 2006-2016, 2006-2017, and 2007-2016 and the trends 79 of MOPITT CO at 200 hPa during time periods of 2000-2016, 2001-2016, 80 2001-2017, and 2002-2016 are shown in Fig. R2 (Supplementary Fig. 6). It could 81 be found that the CO near 200 hPa shows robustly increasing trends over the 82 TWP in satellite data (both of MLS and MOPITT). Overall, though the observed 83 CO only covers less than 20 years, the results from the satellite data may provide 84 extra evidence for the impact of the positive trends of upward motion over the 85 86 TWP on the trace gases in the upper troposphere. The above discussion is added to the revised manuscript. We hope these results may further support our main 87 88 conclusions in this study.



89

Fig. R1. The trends of CO derived from the MLS and MOPITT data. (a) The trends of CO (10⁻¹ ppbv a⁻¹) at 215 hPa using MLS data in NDJFM during 2005-2017. (b) The trends of CO (10⁻¹ ppbv a⁻¹) at 200 hPa using MOPITT data in NDJFM during 2000-2017. The trends of CO over the dotted region are statistically significant at the 90% confidence level.

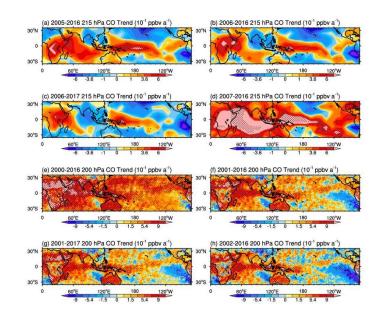


Fig. R2. The trends of CO derived from the MLS and MOPITT data. (a)-(d) The
trends of CO (10⁻¹ ppbv a⁻¹) at 215 hPa using MLS data in NDJFM during
periods of (a) 2005-2016; (b) 2006-2016; (c) 2006-2017; and (d) 2007-2016. (e)-(h)
The trends of CO (10⁻¹ ppbv a⁻¹) at 200 hPa using MOPITT data in NDJFM
during periods of (e) 2000-2016; (f) 2001-2016; (g) 2001-2017; and (h) 2002-2016.
The trends of CO over the dotted region are statistically significant at the 90%
confidence level.

103

95

104 Specific Comments:

P2 L2: 'A significantly intensified upward motion through the troposphere over the
TWP in the boreal wintertime (November to March of the next year) has been
detected.' Please make this statement more quantitative.

108 Re: Corrected. The phrase is rewritten as: "A significantly intensified upward 109 motion through the troposphere over the TWP in the boreal wintertime 110 (November to March of the next year, NDJFM) has been detected using multiple 111 reanalysis datasets. The upward motion over the TWP is intensified at rates of 112 $8\pm3.1\%$ decade⁻¹ and $3.6\pm3.3\%$ decade⁻¹ in NDJFM at 150 hPa from 1958 to 2017

- 113 using JRA55 and ERA5 reanalysis datasets, while the MERRA2 reanalysis data
- show a 7.5±7.1% decade⁻¹ intensified upward motion for the period 1980-2017."
- 115 P2 L18: Please specify here which reanalyses are used.
- 116 **Re: Added.**
- 117 P2 L23: 'numerical simulation' --> 'simulation with WACCM4'?
- 118 Re: Updated.

P2 L24: 'show that more CO could be elevated to the tropical tropopause layer (TTL)'
Please make this statement more quantitative.

Re: Rephrased as: "Using CO as a tropospheric tracer, the WACCM4 simulations show that an increase of CO at a rate of 0.4 ppbv decade⁻¹ at the layer 150-70 hPa in the tropics is mainly resulted from the global SST warming and the subsequent enhanced upward motion over the TWP in the troposphere and strengthened tropical upwelling of Brewer-Dobson (BD) circulation in the lower stratosphere."

P2 L27: Why is aerosol explicitly emphasized here. Please clarify (e.g. outflow frompolluted air from South Asia?)

Re: We thank the reviewer's comment. This sentence has been rewritten as: "This implies that more tropospheric trace gases and aerosols from both natural maritime source and outflow from polluted air from South Asia may enter the stratosphere through the TWP region and affect the stratospheric chemistry and climate."

P3 L42: Please add possible sources of ozone-depleting halogen-containing
substances in TWP (outflow from anthropogenic emissions from South Asia, natural
maritime bromine-containing substances?).

Re: We thank the reviewer's comment. This sentence has been rewritten as: 137 "Through the TWP region, tropospheric trace gases, e.g., the natural maritime 138 bromine-containing substances and outflow from anthropogenic emissions from 139 140 South Asia, are lifted to the upper troposphere by the strong upward motion and 141 the deep convection and subsequently into the stratosphere by the large-scale upwelling (e.g., Levine et al., 2007, 2008; Navarro et al., 2015), which affect the 142 ozone concentration and other chemical processes in the stratosphere (e.g., Feng 143 144 et al., 2007; Sinnhuber et al., 2009)."

145 P4 L45: (Saiz-Lopez and von Glasow, 2012; Wang et al., 2015). -> (e.g.
146 Saiz-Lopez ...).

147 **Re: Corrected.**

148 P4 L46: 'the coldest tropopause' of what? Please specify.

Re: Here we mean that the TWP region has the lowest tropopause temperature over the globe. Corrected as "At the same time, the TWP region has the lowest cold-point tropopause temperature (CPTT) over the globe and plays an important role in controlling the water vapor concentration in the stratosphere."

P4 L49: 'an important region for troposphere-to-stratosphere transport' Please addsome references.

155 Re: Added.

P4 L50: Is the TWP more important for stratospheric chemistry as other regions in theatmosphere? Please clarify?

158 Re: We thank for the reviewer's comment. Here we want to summarize the 159 importance of the TWP region. The sentence was modified as "The TWP is an 160 important region for tropospheric trace gases being transported from the 161 troposphere to the stratosphere, and therefore influencing the stratospheric 162 chemistry (e.g., Fueglistaler et al., 2004; Levine et al., 2007; Krüger et al., 2008;
163 Pan et al., 2016)."

164

P4 L66-70: The impact of ozone-depleting halogen-containing substances is already
mentioned on P3 L42. I propose to combine these two sentences in one paragraph.

Re: These sentences are combined in the first paragraph of Introduction section 167 as: "Through the TWP region, tropospheric trace gases, e.g., the natural 168 169 maritime bromine-containing substances and outflow from anthropogenic 170 emissions from South Asia, are lifted to the upper troposphere by the strong upward motion and the deep convection and subsequently into the stratosphere 171 by the large-scale upwelling (e.g., Levine et al., 2007, 2008; Navarro et al., 2015), 172 which affects the ozone concentration and other chemical processes in the 173 174 stratosphere (e.g., Feng et al., 2007; Sinnhuber et al., 2009)."

175

P4 L71: 'Based on a trajectory model, Fueglistaler et al. (2004) pointed out that the TWP region is a primary source of the tropospheric air entering the stratosphere and approximately 80% of the trajectories ascending into the stratosphere enter the TTL from the TWP'. However, in L63 it is written: 'the TWP is not the dominant entry of trace gases transported from the troposphere into the lower stratosphere'. Please rephrase this statement more carefully.

Re: Thanks for the comment. The statement is rephrased as: "Though the 182 vertical transport from TTL to the lower stratosphere is dominated by the BD 183 circulation, numerous studies confirmed that the TWP region is an important 184 185 pathway of the surface air entering the TTL (Fueglistaler et al., 2004; Levine et al., 2007; Krüger et al., 2008; Haines and Esler, 2014). Based on a trajectory 186 model, Fueglistaler et al. (2004) pointed out that approximately 80% of the 187 trajectories ascending into the stratosphere from the TTL are originated from 188 the TWP region." 189

1	00	
1	90	

191

192

193 2.' 194 **Re: Corrected.** 195 P6 L102: 'is also discussed.' --> ' will be discussed in Sect. 3' 196 **Re: Corrected.** 197 198 P6 L110: Please add the horizontal resolution of ERA5 data ($0.3^{\circ} \times 0.3^{\circ}$), which is 199 much higher as in JRA55 and MERRA2. What about differences in vertical and 200 temporal resolution. Please specify. 201 Re: Thanks for the comment. The description of the JRA55, ERA5 and 202 MERRA2 datasets are rephrased in Section 2, and the information about the 203 204 vertical, horizontal, and temporal resolution are added. 205 P6 L124: 'UTLS' is not yet introduced in the text. 206 **Re: Corrected.** 207 P6 L125: 'even though there are still large biases in the reanalysis datasets' What are 208 the differences between the three different reanalyses (JRA55, ERA5 and MERRA2) 209 used here? Please specify. 210 211 Re: According to the results of Uma et al. (2021), the description is added to the 212 manuscript as: "the updrafts from the JRA55 data in the UTLS are stronger than those from ERA5 and MERRA2 data." It should be mentioned that Uma et 213 al. (2021) did not give quantitative differences between them. 214 215

P6 L100: 'using reanalysis datasets and model simulations' --> 'using JRA55, ERA5

and MERRA2 reanalysis and different WACCAM4 simulations as described in Sect.

P8 L145: 'except that the global SSTs are fixed to the climatological mean values during 1955-2018 (long-term mean for each calendar month during 1955-2018.' Why are the SST not fixed to a value representative for the beginning of the 60-year period?

220 Re: The Control and Fixsst simulations are designed to investigate the impact of SST changes on the intensified upward motion over the TWP. For this purpose, 221 using the SST climatology representative for the beginning of the 60-year period 222 to force the simulation should also be proper. Since we compare the trends 223 224 between the Control (transient) and the Fixsst (constant) simulations, the state of the Fixsst simulation should not influence the results. The SSTs are fixed to the 225 mean of 1958-2017 rather than 1960s to make the mean state of the two 226 simulations more consistent with each other. 227

228

P8 L146 Please explain the added-value of a time-slice experiment compared to thehindcast simulation.

Re: Thanks for the comment. The SSTs in the hindcast simulation are prescribed as the observed SSTs, with changes of SSTs over the globe. SSTs in the time-slice simulations are only modified in the eastern maritime continent and the tropical western Pacific (20°S-20°N, 120°E-160°E), which emphasizes the importance of the SSTs over these areas. The descriptions are clarified in the revised manuscript.

237

P8 L150: For better motivation, please explain in more detail why this set up is usedfor the two time-slice simulations.

Re: Thanks for the suggestion. Some explanations are added to the manuscript as: "To figure out the impact of the warming SST over the TWP region on the intensifying trend of the upward motion over the TWP region, a couple of time-slice simulations (R1 and R2) are also integrated for 33 years... Since the SSTs over the TWP show significantly warming trends, the SSTs during
1998-2017 are higher than the SSTs during 1958-1977. Hence, the difference
between R1 and R2 reflect the impact of the warmed SSTs over the TWP on the
atmospheric circulation."

248

P9 L171: 'the climatological distribution of the vertical velocity at 150 hPa for each
month of the year.' --> Mean values of the vertical velocity at 150 hPa for each month
averaged over 60 years from 1958 to 2017. Yes?

252 **Re: Yes. The statement is corrected correspondingly.**

Why is JRA55 and not ERA5 or MERRA2 selected for Fig.1? What are the difference between JRA55 and ERA5/MERRA2?

Re: The pattern of the 150 hPa vertical velocity from JRA55 data shown in Fig. 1 is similar to the patterns of the 150 hPa vertical velocity from ERA5 and MERRA2 datasets. To avoid repetition, only the result from JRA55 data is shown in Fig. 1. According to the referee's comment, the climatological mean vertical velocity in NDJFM in ERA5 and MERRA2 is added to the supplementary material. The vertical velocity differences between JRA55 and the ERA5 and MEERA2 data are further discussed in the revised manuscript.

P9 L180: please add text within ++: 'which is more important to the transport of air
over the TWP from the lower troposphere to the TTL +compared to the summer
months (as shown in Fig. 1) + and subsequently to the lower stratosphere.

265 **Re: Corrected.**

P9 L182: 'Notably, the 150 hPa w shows no subsidence over the maritime continent,
while there is descending motion over the maritime continent at 100 hPa (not shown),
which is referred to the "stratospheric drain" (Gettleman et al., 2000; Sherwood,
2000).' The 100 hPa values should be shown in an electronic supplement.

270 Re: The 100 hPa *w* values using JRA55, ERA5 and MERRA2 are shown in
271 Supplementary Fig. 2.

272 P10 L186: Please explain in detail how the trend is calculated.

273 Re: We thank for the reviewer's suggestion. The description about the trend and
274 the significance test is added to Section 2 as:

²⁷⁵ "Linear trends and the significance test. The linear trends are estimated ²⁷⁶ using a simple least square regression method. The significances of the ²⁷⁷ correlation coefficients, mean differences, and trends are determined via a ²⁷⁸ two-tail Student's t-test. The confidence interval of trend is calculated using the

following equation (Shirley et al., 2004):
$$\left(b - t_{1-\frac{\alpha}{2}}(n-2)\sigma_b, b + t_{1-\frac{\alpha}{2}}(n-2)\sigma\right)$$

where b is the estimated slope, σ denotes the standard error of the slope, and $t_{1-\frac{\alpha}{2}}(n-2)$ represents the value of t-distribution with the degree of freedom equal to *n*-2. α is the two-tailed confidence level. σ is calculated as:

283
$$\sigma = b \sqrt{\frac{\frac{1}{r^2} - 1}{n - 2}}.$$

P10 L187: 'using reanalysis datasets' -> 'using JARA55, ERA5 and MERRA2
reanalyses.'

286 **Re: Corrected.**

287 P10 L191: ->'is intensifying through the troposphere from 1958 to 2017.'

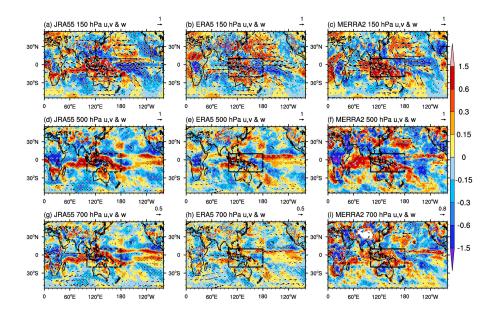
288 Re: Corrected.

289 P10 L193 : add 'used here' or 'used in this study'

Figure 2: In MERRA2 the horizontal winds seems to be much stronger compared to 291 JARA55 and ERA5. Could you make a comment on this. Please discuss the 292 293 similarities and differences of the three reanalyses in more detail. Maybe you could 294 show an additional figure showing the differences of ERA5 and MERRA2 compared to JARA55. ERA5 has much higher spacial and temporal resolution as JRA55 and 295 MERRA2, therefore I would expect pronounced differences to JARA55 and 296 MERRA2, in particular convection is much improved compared to the previous 297 ECMWF reanalysis ERA-Interim. 298

299 Re: Thanks for the comment. In Fig. 2, the trends of the horizontal winds seem to be much stronger in MERRA2 compared to JRA55 and ERA5. It should be 300 noted that the wind trends in JRA55 and ERA5 are calculated during the period 301 1958-2017, however, the wind trends of in MERRA2 are calculated during the 302 303 period 1980-2017. To further figure out whether there are large differences between the trends of the winds between JRA55, ERA5 and MERRA2, the 304 trends of winds during 1980-2017 in NDJFM derived from JRA55, ERA5 and 305 MERRA2 are shown here (and in the supplementary material). It could be seen 306 that the trends of horizontal winds in Figs. R3a and R3b are larger than the 307 trends of horizontal winds in Figs. 2a and 2b (in manuscript). And there are 308 insignificant differences between the trends of horizontal winds in JRA55, ERA5, 309 and MERRA2. Hence, the differences of the trends of the horizontal winds in Fig. 310 311 2 are mainly due to the different time periods which are used to calculate the trends. The trend patterns of the winds in JRA55, ERA5, and MERRA2 are 312 similar. However, there are also some differences between the trends of vertical 313 velocity in JRA55, ERA5, and MERRA2. There are significantly positive trends 314 over the TWP regions in JRA55, ERA5, and MERRA2, while the positive trends 315 of vertical velocity over the TWP in ERA5 seem to be weaker than those in 316 JRA55 and MERRA2. Comparing to the negative trends of the vertical velocity 317

over the central Pacific in JRA55 and ERA5, the negative trends of the vertical
velocity over the central Pacific in MERRA2 extend more northward. The above
discussion is added to the corresponding paragraph in the revised manuscript.



321

322

Fig. R3. The trends of the vertical velocity and horizontal winds in NDJFM using JRA55 (a, d, g), ERA5(b, e, h) and MERRA2(c, f, i) data during 1980-2017 at different levels. (a)-(c) are the trends of winds at 150 hPa. (d)-(f) are the trends of winds at 500 hPa. (g)-(i) are the trends of winds at 700 hPa. The trends of vertical velocity over the dotted region are statistically significant at the 90% confidence level.

329

Figure 3: Please Explain how 'standardized intensity' is calculated. What is the reason
for the extreme minima (1981, 1991, 1999)? El Niño Southern Oscillation (ENSO)?

Re: "The intensity of the upward motion over the TWP is simply defined as the area-averaged upward mass flux at a specific level. And the standardized intensity is the intensity divided by the standard deviation of the intensity at the corresponding level." The explanation of the standardized intensity is added to

the manuscript. The extreme minima (actually, the years are 1982, 1991, and 336 1997) are mainly due to the ENSO events (El Niño), which may result in a weak 337 upward motion over the TWP (e.g., Levine et al., 2008; Hosking et al., 2012; Hu 338 et al., 2016). To figure out the influence of the El Niño events (1982, 1991, 1997), 339 the time series of the standardized intensity of the upward motion over the TWP 340 341 in NDJFM after removing the ENSO signal using the linear regression method (Hu et al., 2018) in JRA55, ERA5, and MERRA2 are shown here (Fig. R4 and 342 343 Supplementary Fig. 5). It could be seen that the extreme minima become much weaker after removing the ENSO signal using the linear regression method. This 344 result suggests that the El Niño events could affect the upward motion over the 345 TWP and to a large extent result in the extreme minima (1982, 1991, and 1997). 346 Notably, the upward motions over the TWP at 150 hPa, 500 hPa, and 700 hPa in 347 NDJFM in JRA55, ERA5, and MERRA2 still show statistically significant 348 intensifying trends after removing the ENSO signal in Supplementary Fig. 5, 349 which suggests that ENSO events exert limited impacts on the trends of the 350 351 upward motion over the TWP in NDJFM during 1958-2017. Some of above discussions are added to the revised manuscript. 352

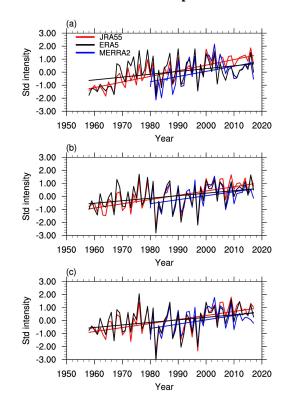


Fig. R4. The time series of the standardized intensity of the upward motion over the tropical western Pacific (20°S-10°N, 100°E-180°E) at (a) 150 hPa; (b) 500 hPa; and (c) 700 hPa extracted from JRA55 (red), ERA5 (black) and MERRA2 (blue) datasets after removing the ENSO signal using linear regression method. The straight lines in each figure indicate the linear trends. The linear trends of the upward motion intensity over the TWP at 150 hPa, 500 hPa, and 700 hPa from three datasets are statistically significant at the 95% confidence level.

361

P10 L201: 'This suggests a comprehensive enhancement of vertical velocity though
the whole troposphere, which is evident from the surface to 100 hPa (not shown).'
Figures demonstrating this could be shown in an electronic supplement.

Re: The trends of vertical velocity from the surface to 100 hPa in NDJFM
derived from JRA55, ERA5, and MERRA2 are added in the supplementary
material (Supplementary Fig. 4)

368

P10 L205 :'Due to the data limitation, it is not possible to show the corresponding changes of trace gases by observations.' I agree that it is difficult to find observation from 1958 to 2017. However satellite measurements from shorter time period could be used (e.g. MLS CO available since August 2004; https://mls.jpl.nasa.gov).

Re: We thank for the referee's comment. An extra figure showing the trends of CO observed by MOPITT and MLS at around 200 hPa during 2000-2017 and 2005-2017 is added in the revised manuscript. The details could be found in the responses to the major comments above.

377

378 P11 L210: 'of observed OLR' --> 'of observed OLR provided by NOAA (see Sect.
379 2)'

380 **Re: Corrected.**

P11 L222: 'CPTT' is not yet introduced. Fig. 4b is not referred to in the text --> '.. the cold-point tropopause temperature (CPTT; see Fig. 4b) shows significantly decreasing trends over the TWP in NDJFM during 1958-2017,... However negative trends are also found in other regions in low and mid-altitudes, except in the Pacific.'

385 386

Re: CPTT is introduced in the revised manuscript *Line 55*. The statement is corrected.

P12 L242: 'The SSTs over the TWP are positively correlated with the upward motion 387 intensity over the TWP, while the SSTs over tropical central, eastern Pacific, and 388 Indian Ocean show negative correlations.' I am wondering that the positive correlation 389 390 pattern is somewhat shifted to the east, then the western part of the maritime continent (100°E-120°E) is also negative correlated. However, in the western part of the 391 maritime continent (100°E-120°E) the trends of horizontal winds (Fig. 2) are large. 392 Maybe, it is useful to avoid misunderstandings to mark the region of the TWP 393 394 somehow (e.g. by a box).

Re: We are sorry for the possible confusion. The TWP is marked by a box in the
figures of the revised manuscript, and the corresponding statement is corrected
to avoid the confusion.

398

P13 L253: 'a couple of model simulations' --> 'a couple of model simulations with
WACCAM4'

401 **Re: Corrected.**

P14 L277: 'a couple of time-slice runs (R1 and R2) are performed (more details are
given in the section 2).' --> It is maybe a matter of taste, but I would prefer in
general to say 'simulations instead of 'run'. Please repeat the main features of R1 and
R2 as a reminder for the reader.

406 Re: Corrected. And the main features of R1 and R2 are added to the
407 corresponding paragraph.

408

409 P14 L289: 'The changes in the OLR' --> 'The changes in the OLR simulated in
410 WACCAM4'

411 **Re: Corrected.**

412 P15 L300: 'We now discuss about the relationship between the trends of the upward motion over the TWP and the changes of the trace gases in the lower stratosphere.' 413 -->'The relationship between the trends of the upward motion over the TWP and the 414 415 change of CO and water vapor in the lower stratosphere simulated with WACCAM4 will be analyzed. It is expected, that a positive trend in the upward motion over the 416 TWP yield higher CO in the lower stratosphere caused be enhanced vertical upward 417 transport. However, water vapor mixing ratios in the lower stratosphere depends in 418 419 addition from the temperature in the UTLS' Is that what you would like to discuss here? 420

421 **Re: Yes. The corresponding phrases are corrected.**

422 Section 3.3 is written somewhat confusing, therefore I propose to write a short 423 introduction of Sect. 3.3 summarizing previous results from the literature and 424 subsequent the new results of Qie et al.

Re: Thanks for the comment. A short introduction of Section 3.3 is added to the manuscript according to the comments of the referee and the literature.

427 "Previous studies showed that the enhanced deep convection and upward motion
428 could lead to increased CO in the UTLS (e.g., Duncan et al., 2007; Livesey et al.,
429 2013). At the same time, water vapor mixing ratios in the UTLS may increase
430 due to the enhanced upward motion which could bring more wet air from low
431 altitude to high altitude (e.g., Rosenlof, 2003; Lu et al., 2020). However, the water

vapor mixing ratios in the lower stratosphere also depend on the tropopause
temperature (e.g., Highwood and Hoskins, 1998; Garfinkel et al., 2018; Pan et al.,
2019). Hence, the relationship between the intensity of upward motion and the
water vapor concentration in the UTLS is complex. Here, the relationship
between the trends of the upward motion over the TWP and the changes in CO
and water vapor in the ULTS simulated with WACCM4 are analyzed."

- 438
- P15 L303: 'in different simulations are displayed' --> 'are shown based on the Control
 and the Fixsst simulation as well as using their difference..'

441 **Re: Corrected.**

442 P15 L303: --> 'in Fig. 7d-i'

443 **Re: Corrected.**

444

P16 L328: 'As mentioned above in section 3.1, the observed tracer gases (e.g., CO)
have very limited data record and may be affected by a mixture of anthropogenic and
natural (e.g., biomass burning) emissions and the ENSO events (e.g., Duncan et al.,
2007; Logan et al., 2008). It is therefore very hard to identify the relative contribution
of single factors.' This sentence is here not very helpful, please remove it.

450 **Re: Removed.**

P16 L332: 'We utilize the numeric simulations' --> 'We use the Control and the Fixsst
simulation with WACCAM4 ..'

- 453 **Re: Corrected.**
- 454

455 P17 L344: 'increasing trends over the TWP' How much is the increase in CO within 456 60 years? Please add some numbers in the text. $(4*10^{-4})$ ppm per year -> 0.024

457 ppm change in CO in 60 years; that seems not to be much.)

Give some reference about CO values and variability of CO in this region frommeasurements to assess the trend in CO over TWP.

Re: Thanks for the suggestion. We show the climatological mean CO values at 460 461 215 hPa in NDJFM from MLS observations during 2005-2017 and at 200 hPa in 462 NDJFM from MOPITT observations during 2000-2017. The concentration of MLS CO over the TWP is approximately 80 ppbv at 215 hPa and MOPITT CO 463 is 70 ppbv at 200 hPa, which is consistent with previous study (e.g., Huang et al., 464 2016). The increasing trends of CO at 150 hPa over the TWP in the Control and 465 Fixsst simulations are approximately 3.4 ppbv decade⁻¹ (20.4 ppbv within 60 466 years) and 3.2 ppbv decade⁻¹ (19.2 ppbv within 60 years). The CO at 150 hPa 467 over the TWP derived from the difference between the Control and Fixsst 468 increased 0.2 ppbv decade⁻¹ (1.2 ppbv within 60 years), which suggests that the 469 470 enhanced deep convection and intensified upward motion could lead to an extra 6% increasing trend of CO at 150 hPa over the TWP. It should be mentioned 471 that the changes in the CO at 150 hPa caused by the intensified upward motion 472 over the TWP not only depend on the vertical transport but also on the gradient 473 474 of CO concentration at around 150 hPa (Garfinkel et al., 2013). This may be the reason why the intensifying upward motion over the TWP only contribute to an 475 extra 6% increasing trend of CO at 150 hPa in NDJFM during 1958-2017. For 476 example, CO derived from the difference between the Control and Fixsst 477 478 simulations shows higher increasing trends in the layer 150-70 hPa (0.4 ppbv decade⁻¹) than those at 150 hPa (0.2 ppbv decade⁻¹), which is due to the greater 479 CO gradient in the UTLS comparing to the CO gradient in the upper 480 troposphere. 481

482

P17 L354: 'This is consistent with our results which show intensified northerlies over
the subtropical Indian Ocean and strengthened westerlies over the subtropical Indian
Ocean and western Pacific'

Please add some numbers in the text: how much is the strengthening. Is it a large orweak change. Please give the reader some numbers to assess this change.

Re: Thanks for the suggestion. The trends of the northerlies over the subtropical
Indian Ocean (15°S-25°S, 60°E-100°E) are approximately 0.2 m s⁻¹ decade⁻¹ and
the trends of westerlies over the subtropical Indian Ocean and western Pacific
(20°N-35°N, 60°E-160°E) are approximately 0.3 m s⁻¹ decade⁻¹ (Figs. 5c and f).
The discussion is added to the revised manuscript.

493

P18 L377: 'In summary, the increase of CO as shown in Figs. 8a-8b is mainly caused
by surface emissions.' My understanding is that the surface emissions are the same in
the Control and Fixsst simulation and that the increase of UTLS CO is caused by
stronger upwelling. Please clarify.

Re: We are sorry for the confusion. The surface emissions are the same in the 498 Control and Fixsst simulations, which are increasing in NDJFM during 499 1958-2017. Hence, the trends of CO in Fig. 9a (in the revised manuscript) contain 500 the CO trends induced both by the increased surface emissions and the enhanced 501 upward motion. The trends of CO over the TWP in Fig. 9b (in the revised 502 manuscript) only include the CO trends induced by the increased surface 503 emissions since the upward motion over the TWP in the Fixsst simulation shows 504 weak trends. Furthermore, the CO increased through the troposphere over the 505 TWP using the difference between the Control and Fixsst simulations, which 506 507 suggests that the increase of CO in the upper troposphere in Fig. 9c (in the revised manuscript) is caused by the intensified upward motion over the TWP. 508 509 Some discussions are added to the text.

510

511 Figure 11: '(a) Control run; (b) Fixsst run; (c) difference between the Control run and

512 the Fixsst run; and (d) JRA55.' --> labels a,b,c,d are not consistent to Fig.11.

513 **Re: We are sorry for the mistake. The figure caption is corrected.**

514

515 Why is MERRA2 and ERA5 not shown. How is the trend of the BD circulation 516 calculated? Are zonal mean values shown? Please clarify.

Re: Thanks for the suggestion. We have added the trends of the BDC derived from ERA5 and MERRA2 to the supplementary material. The trend of the BDC is calculated using the simple least square regression. The w^* used in the manuscript is calculated using the TEM formula and w^* denotes the monthly zonal mean of the vertical component of the BDC. To avoid confusion, the \overline{w}^* and \overline{v}^* in the equation mentioned in the original manuscript are corrected as w^* and v^* in the revised manuscript.

524

P19 L384: 'The tropical upwelling of BDC (w*) are significantly increased in the
lower stratosphere over past decades as seen in both reanalysis data and the control
run (Figs. 11a and b).' --> 'in JARA55 and control simulation'

528 Re: Corrected.

529

530 Please indicate that the TEM is used to calculate w*. Please specify 'significantly 531 increased' with some numbers. Please compare the increase with numbers from other 532 references.

533 Re: We thank the referee's comment. The manuscript is revised correspondingly.

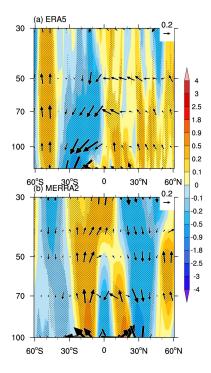
534 The quantitative results and the comparison with other references are added.

The tropical upwelling of BDC (w^*) calculated using the TEM formula increased significantly in the lower stratosphere over past decades as seen in the JRA55 data and the Control simulation (Figs. 12a and 12b). We found that the 70 hPa upward mass flux in NDJFM in the tropics (15°S-15°N) increased 2.8±1.9% decade⁻¹ (significant at the 95% confidence level) in the JRA55 data from 1958

to 2017 (Fig. 12a) and 4.6±4.3% decade⁻¹ (significant at the 95% confidence level) 540 in the MERRA2 data from 1980 to 2017 (Supplementary Fig. 7b). From the 541 ERA5 data, the 70 hPa upward mass flux in NDJFM increased in the north 542 hemisphere (0-15°N) at a rate of 5±2.8% decade-1 (significant at the 95% 543 confidence level), but decreased significantly in the south hemisphere (0-15°S) 544 during 1958-2017 (Supplementary Fig. 7a). On average, the trend of the 70 hPa 545 upward mass flux in NDJFM in the tropics (15°S-15°N) is insignificant in ERA5. 546 547 In fact, many previous studies have investigated the trends of BDC. For example, Abalos et al. (2015) investigated the trends of BDC using JRA55, MERRA, and 548 ERA-Interim data during 1979-2012 and suggested that the BDC in JRA55 and 549 MERRA significantly strengthened throughout the layer 100-10 hPa with a rate 550 of 2-5% decade⁻¹, while the BDC in ERA-Interim shows weakening trends. Diallo 551 et al. (2021) compared the trends of the BDC in the ERA5 and ERA-Interim 552 during 1979-2018 and pointed out that the BDC in the ERA-Interim shows 553 weakening trend and the BDC in the ERA5 strengthened at a rate of 1.5% 554 555 decade⁻¹ which is more consistent with other studies. In the present study, we only focus on the trend of the BDC in the wintertime (NDJFM) in the tropics 556 (15°S-15°N) during 1958-2017, which may lead to some differences between our 557 result and the previous studies. Overall, the trends of the tropical upwelling of 558 BDC using JRA55, MERRA2 data and the Control simulation are similar to the 559 previous studies using both reanalysis datasets and model results (e.g., Butchart 560 et al., 2010; Abalos et al., 2015; Fu et al., 2019; Rao et al., 2019; Diallo et al., 561 2021). However, the tropical upwelling of the BDC decreased using ERA5 data 562 563 in the tropics (15°S-15°N), which are different from the results in JRA55 and MERRA2. In summary, the tropical upwelling of the BD circulation is likely 564 strengthened as shown in JRA55 and MERRA2 reanalyses as well as model 565 566 simulations, although there are some uncertainties since the ERA5 data show a 567 negative trend. This may contribute to the transport of the tropospheric trace gases from the TTL to a higher level. The increased concentration of CO in the 568 UTLS in Fig. 9c and 10f may be due to a combined effect of the strengthened 569

570 tropical upwelling of the BD circulation and the enhanced upward motion over

571 the TWP.



572

Fig. R5. The trends of the BD circulation (vectors) calculated using the TEM formula using ERA5 and MERRA2 data. (a) The trends of w* (10^{-5} m s⁻¹ a⁻¹) and v* (10^{-2} m s⁻¹ a⁻¹) in NDJFM during 1958-2017 using ERA5 data. (b) The trends of w* (10^{-5} m s⁻¹ a⁻¹) and v* (10^{-2} m s⁻¹ a⁻¹) in NDJFM during 1980-2017 using MERRA2 data. The shadings are the trends of the vertical velocities (10^{-5} m s⁻¹ a⁻¹). The trends of the vertical velocity over the dotted regions are statistically significant at the 90% confidence level.

580

P19 L400: 'The recent trends of the upward motion from the lower to the upper troposphere in boreal winter over the TWP is investigated for the first time based on the reanalysis datasets and model simulations.' Specify which reanalysis and which model runs are used.

585 **Re: Corrected.**

586

P19 L405: 'Warmer SSTs over the TWP lead to a strengthened Pacific Walker circulation, enhanced deep convection and stronger upward motion over the TWP.' Please make this statement more quantitative. From the analysis it is not clear for me what is enhanced: convection or subsequent upward motion over the TWP by diabatic heating or both.

Re: Thanks for the suggestion. The statement is rephrased. Both of the deep convection and the subsequent upward motion over the TWP by diabatic heating are enhanced. We are sorry for the confusion.

595

How is downward transport over TWP by the Pacific Walker circulation during ElNiño considered within the analysis? Please clarify?

Re: Thanks for the comment. The impact of ENSO events on the upward motion
over the TWP is discussed in the revised manuscript according to the referee's
suggestion. Some discussions are also added in the Summary and Discussion.

P20 L410:' Model simulations indicate that the CO concentration increases
significantly from the surface to the stratosphere with increased surface emissions.'
Please make the statement more quantitative.

Re: Thanks for the comment. The statement is rephrased as: "Results from the 604 Control simulation indicate that the CO concentration increased significantly 605 from the surface to the stratosphere over the TWP. The CO at 150 hPa increased 606 at a rate of approximately 3.4 ppby decade⁻¹ with increased surface emissions 607 and the enhanced upward motion over the TWP. Specifically, an enhancement of 608 tropospheric upward motion and subsequent upward transport of trace gases 609 over the TWP lead to an extra 6% increasing trend of CO concentrations in the 610 upper troposphere. Furthermore, the upward mass fluxes at 70 hPa in the 611 tropics (15°S-15°N) show strengthening trends at rates of 2.8±1.9% decade⁻¹ and 612 4.6±4.3% decade⁻¹ in JRA55 data (during 1958-2017) and MERRA2 data (during 613

- 614 1980-2017), respectively, which is consistent with previous studies (e.g., Butchart
 615 et al., 2010; Fu et al., 2019; Rao et al., 2019)."
- 616

P20 L417: 'Trace gases and aerosols in the stratosphere have important impacts on the
stratospheric processes, and hence influence the troposphere weather and climate
through their radiative and dynamical feedback'. This statement is very general.
Please be more specific here.

Re: We thank the referee's comment. The statement is rephrased as: "Trace 621 gases and aerosols entering the stratosphere from the troposphere have 622 important impacts on the stratospheric processes. For example, ozone-depleting 623 substances, CH₄ and N₂O could influence on the stratospheric ozone significantly 624 (e.g., Shindell et al., 2013; Wang et al., 2014; WMO, 2018), which also modify the 625 626 temperature in the stratosphere significantly through their strong radiative effects. Water vapor in the lower stratosphere, in particular, has a significant 627 warming effect on the surface climate (Solomon et al., 2010). Therefore, changes 628 of trace gases in the UTLS have important impacts on both tropospheric and 629 630 stratospheric climate."

My impression is that the conclusion section should be revised to summarize theresults of Qie et al in a much more quantitative way.

Re: Thanks for the referee's suggestion. The conclusion section is revised
according to the quantitative results in the revised manuscript.

635 The conclusion section is rewritten as:

636 "The recent trends of the upward motion from the lower to the upper 637 troposphere in boreal winter over the TWP is investigated for the first time based 638 on the JRA55, ERA5, MERRA2 datasets and four WACCM4 simulations (more 639 details could be found in Section 2). The upward motion at 150 hPa over the 640 TWP in NDJFM increased 8±3.1% decade⁻¹ and 3.6±3.3% decade⁻¹ in NDJFM 641 from 1958 to 2017 in JRA55 and ERA5 reanalysis datasets, respectively. Despite

the possible discontinuities between the radiosonde era (after 1958) and the 642 satellite era (after 1979), the upward motion at 150 hPa over the TWP in NDJFM 643 increased 7.5±7.1% decade⁻¹ during 1980-2017 in MERRA2 data. Such 644 intensification of the upward motion over the TWP also exist in the middle- and 645 lower-troposphere in NDJFM in JRA55, ERA5, and MERRA2, which can be 646 confirmed by the WACCM4 model simulations. Comparing the results between 647 the Control and Fixsst simulations with WACCM4, it is found that the trend of 648 649 the upward motion over the TWP is closely related to the changes in global SSTs, especially the SST warming over the eastern maritime continent and tropical 650 western Pacific (see the results from the experiments R1 and R2 in Fig. 7). 651 Warmer SSTs over the eastern maritime continent and tropical western Pacific 652 (approximately 0.5 K) lead to a strengthened Pacific Walker circulation, 653 enhanced deep convection and approximately 27% intensified upward motion at 654 150 hPa over the TWP as shown by the results from the experiments R1 and R2. 655 The enhanced deep convection over the TWP could lead to a dryer lower 656 657 stratosphere over the TWP, as the strong upward motion and the Rossby-Kelvin wave responses induce a colder tropopause over the TWP. It should be pointed 658 out that the results in the present study are mainly based on the reanalyses data, 659 660 and some uncertainties may exist. More observational data are expected to be used to obtain a more robust result in the future. 661

Results from the Control simulation indicate that the CO concentrations increased significantly from the surface to the stratosphere over the TWP. The CO at 150 hPa increased at a rate of approximately 3.4 ppbv decade⁻¹ with increased surface emissions and the enhanced upward motion over the TWP. Specifically, an enhancement of tropospheric upward motion and subsequent upward transport of trace gases over the TWP lead to an extra 6% increasing trend of CO concentrations in the upper troposphere.

Furthermore, the upward mass fluxes at 70 hPa in the tropics (15°S-15°N)
show strengthening trends at rates of 2.8±1.9% decade⁻¹ and 4.6±4.3% decade⁻¹
using JRA55 data (during 1958-2017) and MERRA2 data (during 1980-2017) in

NDJFM, which is consistent with previous studies (e.g., Butchart et al., 2010; Fu 672 et al., 2019; Rao et al., 2019). However, such enhancement in tropical upward 673 674 mass flux at 70 hPa has large uncertainties since the ERA5 data show a negative and insignificant trend (Supplementary Fig. 7a). The results from the Control 675 and Fixsst simulations indicate that the elevated CO in the upper troposphere is 676 677 further uplifted to the lower stratosphere by the intensified tropical upwelling of the BD circulation due mainly to global SST warming and lead to an increase of 678 679 CO in the lower stratosphere. An extra 14% increasing trend of CO at the layer 150-70 hPa over the TWP is derived from the Control and Fixsst simulations..." 680

681

682 **References:**

Abalos, M., Legras, B., Ploeger, F., and Randel, W. J.: Evaluating the advective
Brewer-Dobson circulation in three reanalyses for the period 1979-2012, J.
Geophys. Res., 120,7534-7554, doi:10.1002/2015JD023182, 2015.

Butchart, N., Cionni, I., Eyring, V., Shepherd, T. G., Waugh, D. W., Akiyoshi, H.,
Austin, J., Brühl, C., Chipperfield, M. P., Cordero, E., Dameris, M., Deckert,
R., Dhomse, S., Frith, S. M., Garcia, R. R., Gettelman, A., Giorgetta, M. A.,
Kinnison, D. E., Li, F., Mancini, E., McLandress, C., Pawson, S., Pitari, G.,

- Plummer, D. A., Rozanov, E., Sassi, F., Scinocca, J. F., Shibata, K., Steil, B.,
 and Tian, W.: Chemistry–Climate Model simulations of twenty-first century
 stratospheric climate and circulation changes, J. Climate, 23, 5349–5374,
 https://doi.org/10.1175/2010JCLI3404.1, 2010.
- Deeter, M. N., Edwards, D. P., Francis, G. L., Gille, J. C., Mao, D., 694 Martinez-Alonso, S., Worden, H. M., Ziskin, D., and Andreae, M. O.: 695 Radiance-based retrieval bias mitigation for the MOPITT instrument: the 696 697 version 8 product, Atmos. Meas. Tech., 12, 4561-4580, 698 https://doi.org/10.5194/amt-12-4561-2019, 2019.
- Diallo, M., Ern, M., and Ploeger, F.: The advective Brewer–Dobson circulation
 in the ERA5 reanalysis: climatology, variability, and trends, Atmos. Chem.
 Phys., 21, 7515–7544, https://doi.org/10.5194/acp-21-7515-2021, 2021.

- Duncan, B. N., Logan J. A., Bey, I., Megretskaia, I. A., Yantosca, R. M., Novelli, P.
 C., Jones, N. B. and Rinsland, C. P.: Global budget of CO, 1988-1997:
 Source estimates and validation with a global model, J. Geophys. Res., 112,
 D22301, doi:10.1029/2007JD008459, 2007.
- Fu, Q., Solomon, S., Pahlavan, H. A., and Lin, P.: Observed changes in
 Brewer–Dobson circulation for 1980–2018, Environ. Res. Lett., 14, 114 026,
 https://doi.org/10.1088/1748-9326/ab4de7, 2019.
- Fueglistaler, S., Wernli, H., and Peter, T.: Tropical troposphere-to-stratosphere
 transport inferred from trajectory calculations, J. Geophys. Res., 109,
 D03108, doi:10.1029/2003JD004069, 2004.
- Garfinkel, C. I., Waugh, D. W., Oman, L. D., Wang, L., and Hurwitz, M. M.:
 Temperature trends in the tropical upper troposphere and lower
 stratosphere: Connections with sea surface temperatures and implications
 for water vapor and ozone, J. Geophys. Res., 118, 9658-9672,
 doi:10.1002/jgrd.50772, 2013.
- Garfinkel, C. O., Gordon, A., Oman, L. D., Li, F., Davis, S., and Pawson, S.:
 Nonlinear response of tropical lower-stratospheric temperature and water
 vapor to ENSO, Atmos. Chem. Phys., 18, 4597-4615,
 https://doi.org/10.5194/acp-18-4597-2018, 2018.
- Gettelman, A., Holton, J. R., and Douglass, A. R.: Simulations of water vapor in
 the lower stratosphere and upper troposphere, J. Geophys. Res., 105(D7),
 9003-9023, https://doi.org/10.1029/1999JD901133, 2000.
- Haines, P. E., and Esler, J. G.: Determination of the source regions for surface to
 stratosphere transport: An Eulerian backtracking approach, Geophys. Res.
 Lett., 41, 1343-1349, doi:10.1002/2013GL058757, 2014.
- Highwood, E. J., and Hoskins, B. J.: The tropical tropopause, Q. J. R. Meteorol.
 Soc., 124(549), 1579-1604, DOI: 10.1002/qj.49712454911, 1998.
- 729 Hosking, J. S., Russo, M. R., Braesicke, P., and Pyle, J. A.: Tropical convective
- 730 transport and the Walker circulation, Atmos. Chem. Phys., 12, 9791-9797,
- 731 doi:10.5194/acp-12-9791-2012, 2012

- Hu, D., Guan, Z., Tian, W., and Ren, R.: Recent strengthening of the
 stratospheric Arctic vortex response to warming in the central North Pacific,
 9, 1697, https://doi.org/10.1038/s41467-018-04138-3, 2018.
- Hu, D., Guo, Y., Wang, F., Xu, Q., Li, Y., Sang, W., Wang, X., and Liu, M.:
 Brewer-Dobson circulation: Recent-Past and near-future trends simulated
 by chemistry-climate models, Adv. Meteorol., 2017, 1-13,
 https://doi.org/10.1155/2017/2913895, 2017.
- Hu, D., Guo, Y., Wang, F., Xu, Q., Li, Y., Sang, W., Wang, X., and Liu, M.:
 Brewer-Dobson circulation: Recent-Past and near-future trends simulated
 by chemistry-climate models, Adv. Meteorol., 2017, 1-13,
 https://doi.org/10.1155/2017/2913895, 2017.
- Hu, D., Tian, W., Guan, Z., Guo, Y., and Dhomse, S.: Longitudinal asymmetric
 trends of tropical cold-point tropopause temperature and their link to
 strengthened Walker circulation, J. Climate, 29(21), 7755–7771,
 https://doi.org/10.1175/JCLI-D-15-0851.1, 2016.
- Huang, L., Jiang, J. H., Murray, L. T., Damon, M. R., Su, H., and Livesey, N.:
 Evaluation of UTLS carbon monoxide simulations in GMI and
 GEOS-Chem chemical transport models using Aura MLS observations,
 Atmos. Chem. Phys., 16, 5641-5663, doi:10.5194/acp-16-5641-2016, 2016.
- Krüger, K., Tegtmeier, S., and Rex, M.: Long-term climatology of air mass
 transport through the Tropical Tropopause Layer (TTL) during NH winter,
 Atmos. Chem. Phys., 8, 813–823, doi:10.5194/acpd-7-13989-2007, 2008.
- Levine, J. G., Braesicke, P., Harris, N. R. P., Pyle, J. A.: Seasonal and
 inter-annual variations in troposphere-to-stratosphere transport from the
 tropical tropopause layer, Atmos. Chem. Phys., 8, 3689-3703,
 DOI:10.5194/acpd-8-489-2008, 2008.
- Levine, J. G., Braesicke, P., Harris, N. R. P., Savage, N. H., and Pyle, J. A.:
 Pathways and timescales for troposphere-to-stratosphere transport via the
 tropical tropopause layer and their relevance for very short lived substances,
 J. Geophys. Res., 112, D04308, doi:10.1029/2005JD006940, 2007.

- Livesey, N. J., Logan, J. A., Santee, M. L., Waters, J. W., Doherty, R. M., Read,
 W. G., Froidevaux, L., and Jiang, J. H.: Interrelated variatinos of O3, CO
 and deep convection in the tropical/subtropical upper troposphere observed
 by the Aura Microwave Lim Sounder (MLS) during 2004-2011, Atmos.
 Chem. Phys, 13, 579-598, doi:10.5194/acp-13-579-2013, 2013.
- Livesey, N. J., Read,W. G., Wagner, P. A., Froidevaux, L., Lambert, A., Manney,
 G. L., Millán, L., Pumphrey, H. C., Santee, M. L., Schwartz, M. J., Wang, S.,
 Fuller, R. A., Jarnot, R. F., Knosp, B. W., and Martinez, E.: EOS MLS
 Version 4.2x Level 2 data quality and description document, Jet Propulsion
 Laboratory, California Institute of Technology, Pasadena, CA, 2015.
- Lu, J., Xie, F., Sun, C., Luo, J., Cai, Q., Zhang, J., Li, J., and Tian, H.: Analysis
 of factors influencing tropical lower stratospheric water vapor during
 1980–2017, npj Clim. Atmos. Sci., 3(1), 35,
 https://doi.org/10.1038/s41612-020-00138-7, 2020.
- Navarro, M. A., Atlas, E. L., Saiz-Lopez, A., Rodriguez-Lloveras, X., Kinnison, D.
 E., Lamarque, J., Tilmes, S., Filus, M., and Harris, N. R. P., et al.: Airborne
 measurements of organic bromine compounds in the Pacific tropical
 tropopause layer, P. Natl. Acad. Sci. USA, 112, 13789-13793,
 doi:10.1073/pnas.1511463112, 2015.
- Pan, L. L., Atlas, E. L., Salawitch, R. J., Honomichl, S. B., Bresch, J. F., and
 Randel, W. J., et al.: The Convective Transport of Active Species in the
 Tropics (CONTRAST) Experiment, B. Am. Meteorol. Soc., 98(1), 106-128,
 DOI: 10.1175/BAMS-D-14-00272.1, 2016.
- Pan, L. L., Honomichl, S. B., Thornberry, T., Rollins, A., Bui, T. P., Pfister, L.,
 and Jensen E. E.: Observational Evidence of Horizontal Transport-Driven
 Dehydration in the TTL, Geophys. Res. Lett., 46(13), 7848-7856,
 DOI: 10.1029/2019GL083647, 2019.
- Qie, K., Qie, X., and Tian, W.: Increasing trend of lightning activity in the South
 Asia region, Sci. Bull., 66, 78-84, https://doi.org/10.1016/j.scib.2020.08.033,
 2021.

- Rao, J., Yu, Y., Guo, D., Shi, C., Chen, D., and Hu, D.: Evaluating the
 Brewer-Dobson circulation and its responses to ENSO, QBO, and the solar
 cycle in different reanalyses, Earth Planet. Phys., 3(2), 1-16,
 http://doi.org/10.26464/epp2019012, 2019.
- Rex, M., Wohltmann, I., Ridder, T., Lehmann, R., Rosenlof, K., Wennberg, P.,
 Weisenstein, D., Notholt, J., Krüger, K., Mohr, V., and Tegtmeier, S.: A
 tropical West Pacific OH minimum and implications for stratospheric
 composition, Atmos. Chem. Phys., 14, 4827-4841,
 doi:10.5194/acp-14-4827-2014, 2014.
- Rosenlof, K. H. How water enters the stratosphere, Science, 302, 1691-1692,
 doi:10.1126/science.1092703, 2003.
- Ryu, J., and Lee, S.: Effect of tropical waves on the tropical tropopause
 transition layer upwelling, J. Atmos. Sci., 67(10), 3130-3148,
 DOI: 10.1175/2010JAS3434.1, 2010.
- 806 Sherwood, S. C.: A stratospheric "drain" over the maritime continent, Geophys.
 807 Res. Lett., 27(5), 677-680, https://doi.org/10.1029/1999GL010868, 2000.
- Shirley, D., Stanley, W., & Daniel, C.: Statistics for Research (Third Edition), (p.
 627), Hoboken, New Jersey: John Wiley & Sons Inc., 2004.
- Uma, K. N., Das, S. S., Ratnam, M. V., and Suneeth, K. V.: Assessment of vertical
 air motion among reanalyses and qualitative comparison with
 very-high-frequency radar measurements over two tropical stations, Atmos.
- 813 Chem. Phys., 21, 2083-2103, https://doi.org/10.5194/acp-21-2083-2021, 2021.
- Wang, S., Schmidt, J. A., Baidar, S., Coburn, S., Dix, B., and Koenig, T. K., et al.: 814 815 Active and wide-spread halogen chemistry in the tropical and subtropical Acad. USA, 816 free troposphere, P. Natl. Sci. 112, 9281-9286, DOI: 10.1073/pnas.1505142112, 2015. 817
- 818
- 819
- 820