Contrasting characteristics of open- and closed- cellular stratocumulus cloud in the Eastern North Atlantic

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Abstract. Extensive regions of marine boundary layer cloud impact the radiative balance through their significant shortwave albedo while having little impact on outgoing longwave radiation. Despite this importance, these cloud systems remain poorly represented in large-scale models due to difficulty in representing the processes that drive their lifecycle and coverage. In particular, the mesoscale organization, and cellular structure of marine boundary clouds has important implications for the subsequent cloud feedbacks. In this study, we use long-term (2013-2018) observations from the Atmospheric Radiation Measurement (ARM) Facility's Eastern North Atlantic (ENA) site on Graciosa Island, Azores, Portugal to identify cloud cases with open- or closed-cellular organization. More than 500 hours of each organization type are identified. The ARM observations are combined with reanalysis and satellite products to quantify the cloud, precipitation, aerosol, thermodynamic and large-scale synoptic characteristics associated with these cloud types. Our analysis shows that both cloud organization populations occur during similar sea surface temperature conditions, but the open-cell cases are distinguished by stronger cold-air advection and large-scale subsidence compared to the closed-cell cases, consistent with their formation during cold-air outbreaks. We also find that the open-cell cases were associated with deeper boundary layers, stronger low-level winds, and higher-rain rates compared to their closed-cell counterparts. Finally, raindrops with diameters larger than one millimeter were routinely recorded at the surface during both populations, with a higher number of large drops during the open-cellular cases. The similarities and differences noted herein provide important insights into the environmental and cloud characteristics during varying marine boundary layer cloud mesoscale organization and will be useful for the evaluation of model simulations for ENA marine clouds.

1 Introduction

It is well established that a small increase in the global coverage of marine boundary layer (MBL) stratocumulus clouds could offset warming associated with a doubling of CO_2 (Hartmann and Short, 1980; Randall et al., 1984; Slingo, 1990). This is because the albedo of MBL clouds is much larger than that of the underlying ocean, generally causing a significant

decrease in the amount of solar radiation absorbed in the ocean's mixed layer, with a minimal difference in thermal radiation emitted to space. This large radiative impact coupled with their large areal coverage makes marine stratocumulus clouds an important component of the global energy balance. Bony and Dufresne (2005) have shown that the simulation and response to

35 the changing climate of MBL stratocumulus clouds represents the main source of uncertainty in cloud feedbacks simulated by Earth System Models used for predicting the future climate.

MBL stratocumulus clouds are intimately coupled to the turbulence in the boundary layer that is modulated primarily by the cloud top radiative cooling, entrainment, precipitation and surface turbulent fluxes (Wood, 2012). These clouds are known to occur in two distinct mesoscale (20-200 km) organizations known as closed cellular (unbroken) and open cellular

40 (broken) stratocumulus (Wood and Hartmann, 2006). These differing modes organize the internal diabatic forcings within the MBL, impacting the low-level cloud fraction, shortwave albedo and liquid water path driving the localized contribution to the radiative energy balance and water cycle (e.g., Rossow et al. 2002; Savic-Jovcic and Stevens 2008; Wood et al. 2008; Wood et al., 2016). The key processes responsible for these organizing states have been the subject of much research over the past several decades with the interaction among precipitation and boundary layer dynamics identified as a main driver (e.g.

45 Bretherton et al. 2004; Comstock et al. 2005, 2007; Sharon et al. 2006; Stevens et al. 2005; Savic-Jovcic and Stevens 2008; Wang and Feingold, 2009; Feingold et al. 2010). Aerosols, through their influences on the formation and suppression of drizzle have also been found to have important implications on MBL cloud organization (e.g. van Zanten and Stevens, 2005; Petters et al., 2006; Sharon et al., 2006; Wood et al., 2008; Xue et al., 2008; Wang and Feingold, 2009; Bretherton et al., 2010; Wood et al., 2011). The interplay of these different mechanisms and their relative importance under different regimes remains and

50 area of active research and a needed target for improved understanding and representation in large-scale atmospheric models (Wood et al., 2016; Jensen et al., 2016).

A number of studies based on observational analyses of field campaign data have investigated the characteristics of cellular MBL cloud fields aiming to define the processes responsible for the organization. Using aircraft observations off the coast of California, Sharon et al. (2006) and Stevens et al. (2005), observed much higher drizzle rates in pockets of open-

55 cellular clouds compared to nearby closed-cellular clouds hypothesizing that this precipitation was a driving force of organization. Comstock et al. (2005; 2007), using shipboard observations from the East Pacific Investigation of Climate (EPIC) 2001 field campaign (Bretherton et al., 2004), found significantly higher drizzle rates in open-cellular clouds, but also found that open-cell stratocumulus were associated with deeper, thermodynamically decoupled boundary layers compared to closed-cell stratocumulus. Several studies using observations from the VAMOS Ocean-Cloud-Atmosphere-Land Study (VOCALS)

60 field campaign (Wood et al., 2011) reported that cold pools were common in both open- and closed-cell stratocumulus cloud fields, however, drizzle was stronger and accumulation-mode aerosol concentrations were much lower during open-cellular conditions (Ghate et al., 2013; Wood et al., 2011; Terai and Wood, 2013; Wilbanks et al., 2015). However, Wood et al. (2011) and Terai et al. (2013) found that drizzle rates are not significantly different between open and closed cells, concluding that drizzle, and its associated thermodynamic feedbacks, are not the only factor causing the transition between mesoscale

65 organizations.

The driving mechanisms for changes in the organization of MBL stratocumulus clouds have also been the focus for a number of modelling studies over recent decades. Shao and Randall (1996) perform simulations incorporating increasing complexity of the model analysis for closed cellular MBL clouds. They found that cloud-top radiative cooling is a primary driver of the boundary layer dynamics and cloud organization. Feingold et al. (2010) combine satellite observations with numerical simulations to show the role of collisions between precipitation generated cold pools in driving the characteristics of open-cellular cloud organizations and their oscillation between different, weakly stable states. While many studies have noted the importance of drizzle in the transition between organizational cloud states, Yamaguchi and Feingold (2014), through idealized three-dimensional simulations, highlight the importance of the spatial distribution of the precipitation. Feingold et al. (2015) perform a series of idealized cloud-resolving model simulations to investigate the two-way transitions between open- and closed-cellular cloud populations. Their findings reiterate the importance of precipitation in the transition from the closed- to open-cellular states and emphasize that stabilization of the boundary layer due to this precipitation and increased longwave cooling acts as a barrier to cloud formation and recovery to closed-cellular cloud organization.

Within the large decks of marine stratocumulus clouds observed over the Eastern subtropical oceans, the transition from closed- to open-cellular is routinely observed with little diurnal variability (Burleysen and Yuter, 2015) or changes in the large-scale (inversion strength, subsidence, sea surface temperature [SST], etc.) conditions (Muhlbaur et al., 2014). Although the Eastern North Atlantic (ENA) broadly falls under the subsiding branch of the Hadley circulation, forcing a boundary layer inversion over a warmer sea surface that leads to shallow convection, it also routinely experiences cold air outbreaks (McCoy et al., 2017; Lamraoui et al., 2019, Ghate et al. 2020). These cold air outbreaks also contain open-cellular marine stratocumulus with the transition from the closed- to open-cellular cloud organization happening farther north. Although stratocumulus organized in closed- and open-cellular patterns are routinely observed, the differences between their cloud and environmental characteristics have not been quantified from long-term observations.

In this study, we use observations collected at the Atmospheric Radiation Measurement's (ARM's) ENA site to characterize the aerosol, cloud, precipitation, and thermodynamic conditions during closed- and open-cellular stratocumulus conditions observed in the region. The data and instrumentation at the ENA site are described in section 2. Section 3 describes the methodology for selecting closed- and open-cellular cloud cases. In section 4, we compare and contrast the cloud and environmental characteristics associated with each cloud population. The article is concluded with a summary and discussion in section 5.

2 Data and Instrumentation

The U. S. Department of Energy's ARM (https://www.arm.gov) User Facility operates a number of surface-based measurement sites around the globe (e.g., Mather and Voyles, 2013). Each of these sites include a comprehensive suite of instrumentation for continuous remote sensing and in situ measurements of cloud, aerosol, radiation, atmospheric state and precipitation. Following a successful 2-year deployment of an ARM Mobile Facility (Miller et al., 2016) at Graciosa Island,

Azores, Portugal (Wood et al., 2015), the ARM ENA long-term fixed site was established at this location (Fig. 1, 39.09^o N, 28.03^o W, 15 m) in October 2013.

110   For this study, we use active and passive remote sensing and *in situ* observations from the ARM site, combined with satellite observations and reanalysis products to quantify the cloud and environmental characteristics associated with both closed- and open-cellular cloud organization. The Ka-band ARM Zenith Radar (KAZR, Kollias et al., 2007, 2016; Widener et al., 2012) collects profiles of raw Doppler spectra and its first three moments at 2 second and 30 m resolution in co- and cross-polarization modes. Collocated with the KAZR is a ceilometer that records the first three optical cloud base heights at 15

115   second and 30 m resolution, a Micro-Pulse Lidar (MPL) that records raw backscatter at a 7 second and 30 m resolution, and a microwave radiometer (MWR; Morris 2019; Cadeddu et al., 2013) that records brightness temperatures at 23 GHz, 32 GHz and 90 GHz frequencies. The data from the KAZR, MPL, MWR and ceilometer is combined in the Active Remote Sensing of CLouds (ARSCL) value-added product (ARM, 2015) to produce profiles of noise-filtered moments of Doppler spectra and cloud boundaries (Clothiaux et al. 2001, 2011; Kollias et al. 2005, 2016). The data from the MWR was also used to retrieve

120   column integrated values of Liquid Water Path (LWP) and Precipitable Water Vapor (PWV) at 20 second temporal resolution (ARM, 2014a; Gaustad et al., 2011; Turner et al., 2007). MWR measurements that are flagged as being influenced by rain or PWV measurements that lie beyond 2 standard deviations of the daily mean are removed. Balloon born radiosondes are launched twice daily at the site (00 and 12 UTC) using a Vaisala RS-92, or more recently RS-41 radiosonde (Holdridge et al., 2011; Jensen et al., 2016), and report profiles of temperature, humidity, pressure and winds (ARM, 2013). The lower

125   tropospheric stability (LTS), quantified as the difference between the potential temperature at 700 hPa and that at the surface (Klein and Hartmann, 1993), is also calculated using the radiosonde data. Surface raindrop size distributions (DSDs) are measured by a two-dimensional video disdrometer (2DVD, e.g., Kruger and Krajewski, 2002; Bartholomew 2017; ARM, 2014b). For improved data quality, we apply a fall speed correction and drop size and number thresholds. The DSD parameters are calculated using open source PyDSD code (Hardin and Guy, 2017). More detailed information regarding the disdrometer

130   processing is described in Wang et al. (2018) and Giangrande et al. (2019). Also present at the site is a surface meteorological station that reports temperature, pressure, humidity, and winds at a 1-minute temporal resolution. From the plethora of aerosol instrumentation present at the ENA site, we used the data collected by the Condensation Particle Counter (CPC) that measures the total concentration of aerosol with diameters between 3 to 3000 nanometers. Visible imagery (0.65 micrometer) and associated data products available at 30-60 minute temporal and 9 km spatial resolution from the Spinning Enhanced Visible

135   and Infrared Imager (SEVIRI) onboard the Meteosat-11 satellite were used to characterize cloud conditions around the site. These data were obtained from the Satellite ClOud and Radiation Property (SatCORPS) team and NASA Langley Research Center (NASA 2019a; NASA 2019b).

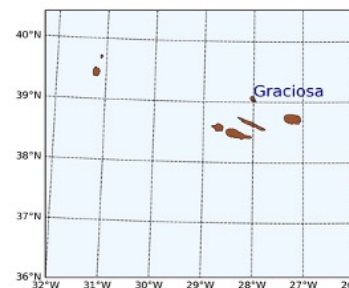
The data from the National Center for Environmental Prediction (NCEP) reanalysis available at 2 degree spatial and 6-hour temporal resolution was used to characterize the large-scale environmental conditions at the site (Kalnay et al., 1996).

140   Primarily the SST, horizontal wind speed and direction at the surface, and large-scale subsidence as reported by NCEP reanalysis were retrieved. The reanalysis data were also used to force the Hybrid Single Particle Lagrangian Integrated

Figure 1 (top left) Map showing the location of the Azores and Graciosa Island, (top right) elevation map of Graciosa Island showing the location of the ARM ENA site on the northern coastline, and (bottom) photo of the ARM ENA site courtesy of the U. S. Department of Energy ARM User Facility.

Trajectory (HYSPLIT; Stein et al., 2015; Rolph et al. 2017) model to deduce the location of air-parcels before reaching the site. Due to high uncertainty in the reanalysis-reported large-scale vertical air motion, the HYSPLIT simulations were performed at 500 m height assuming isobaric vertical air motion.

3 Identification of Closed- and Open-Cellular Cloud Organization Cases

The mesoscale cellularity of marine stratocumulus clouds has been previously identified using data from polar orbiting (e.g., Wood and Hartmann, 2006; Jensen et al., 2008; Zhang et al., 2010) and geostationary satellites (Burleysen et al., 2015), either through analysis of retrieved estimates of LWP or cloud top temperatures. As the goal of this study is not to quantify the cellularity, but rather to determine the differences between the cloud and environmental characteristics associated with the two mesoscale organizations, we use a more traditional approach of combining data from ground-based and satellite

data to select closed- and open-cellular cloud organization cases via visual inspection. To minimize the potential island influences on the observed boundary layer and cloud properties, since the site is located at the northern edge of the island, we focus on cases with dominant boundary layers wind directions from the northern half of the wind rose (e.g., Miller et al. 1998, Rémillard and Tselioudis 2015). Days that were mostly overcast (minimum 12 hours), had relatively uniform (within 500 m) cloud top heights, and had only a single cloud layer below 4 km were selected as closed-cellular cloud organization cases. The SEVIRI images were used to ensure uniform cloud-cover within 2° north of the ENA site. Days that had relatively uniform cloud top heights of heavily precipitating stratocumulus clouds, had shallow cumulus clouds underneath the precipitating stratocumulus, and had noticeable cellular structure encompassing clear areas (identified from the satellite imagery) with stratocumulus cloud cover lower than 100% were chosen for open-cellular cloud organization cases. Additionally, cases with any mid-level clouds (4-7 km), that lasted less than 12 hours, or were missing any of the ground-based measurements listed above, were discarded. In Figures 2 and 3, we show illustrative examples of the time-height KAZR reflectivity observations and the SEVIRI visible satellite imagery for a closed-cellular and an open cellular cloud organization case, respectively. Using these criteria, 26 closed-cellular cases (588 hours) and 24 open-cellular cases (536 hours) were identified from the data collected from October 2015 to December 2018. The list of cases for both open and closed cellular stratocumulus clouds is included in Table 1.

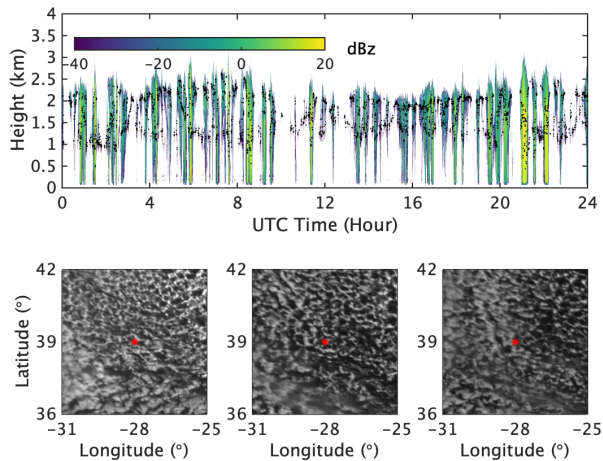



Figure 2 Example of open cells from 22 October 2016. (top) Time-height profile of the best estimate radar reflectivity from the ARSCL value-added product. (bottom) reflectance at 0.65 microns wavelength from Meteosat observations.

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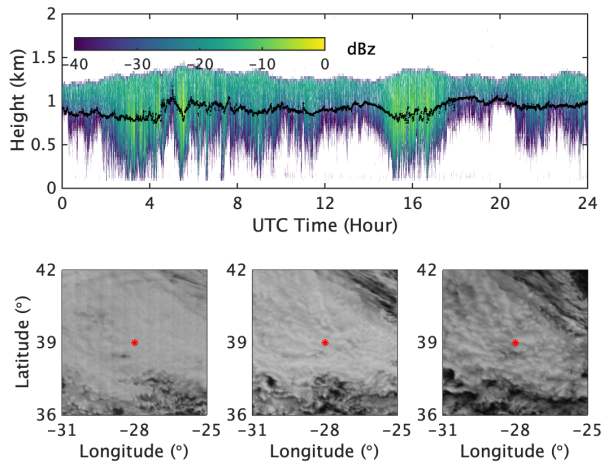
Figure 3 Example of closed cells from 04 March 2016. (top) Time-height profile of the best estimate radar reflectivity from the ARSCL value-added product. (bottom) reflectance at 0.65 microns wavelength from Meteosat observations.

Cloud Organization	Dates [YYYYMMDD], Hours (if not 0-24)
Closed-cellular mesoscale organization Total 26 days	20151019, 20151020, 20151021, 20151213 [0-12], 20160301, 20160303, 20160304, 20160524, 20161015, 20161030, 20161031, 20161101 [0-12], 20161102, 20161116, 20161117, 20170226 [0-12], 20170527, 20180204, 20180205, 20180502, 20180503, 20180512, 20180513, 20180805, 20181028, 20181029
Open-cellular mesoscale organization Total 24 days	20151207, 20151226, 20160305, 20160326, 20160329 [6-24], 20160509, 20161022, 20161023, 20161104, 20161128, 20170209 [0-16], 20170302, 20170329 [0-18], 20170409 [0-16], 20180115, 20180216 [0-12], 20180401, 20181010, 20181012, 20181106, 20181107, 20181111, 20181120, 20181121

190 **Table 1 List of days and hours if not the whole day, with closed and open cellular mesoscale organization at the ARM ENA site.**

The annual cycle of the synoptic and cloud conditions at the Azores is driven by the interplay of the midlatitude storm track and the subtropical Azorian high pressure system. During the boreal winter months (December-January-February), the southern displacement of the midlatitude storm track drives an increase of deeper, storm-related cloud systems and subsidence
195 associated with cold-air outbreaks drives the formation of boundary layer cloudiness (Rémillard and Tselioudis 2015). During the boreal summer months (June-July-August) subsidence associated with the subtropical Azores high pressure system drives a maximum in stratocumulus cloud occurrence (Rémillard and Tselioudis, 2015). Most of the cases for both cloud organizations were observed during the transition months, 12 (11) and 9 (8) closed- (open-) cellular cloud organization during September-October-November and March-April-May respectively (Table 2). There was only one case of closed-cellular and
200 none of open-cellular cloud organization during the summer season. The abundance of open-cellular cases during the transition months is consistent with the climatology presented by McCoy et al. (2017), however the lack of closed-cellular cases during summer months was unexpected. Further scrutiny of the observations during the summer months shows that stratocumulus cloud cases often do not persist through the afternoon hours, lasting for less than 12 hours, and are often complicated by

embedded cumulus and mid- and multi-level clouds which do not meet our selection criteria for closed-cellular cloud organization cases.

Season	Closed Cellular mesoscale organization	Open Cellular mesoscale organization
MAM	9	8
JJA	1	0
SON	12	11
DJF	4	5
Total	26	24

Table 2 Summary of open/closed cell cases used in this study for each season (MAM = March/April/May, JJA = June/July/August, SON = September/October/November, DJF = December/January/February)

4 Composite Characteristics for Cloud Organization Populations

To quantify, compare and contrast the cloud characteristics and associated environment for closed- and open-cellular cloud organizations at the ENA site, we determine the mean and standard deviation of a number of important large-scale meteorological, environmental thermodynamic, and cloud, precipitation and aerosol variables over the defined populations of cloud organization cases. From these statistical measures we present composite visualizations of the relevant cloud and explanatory environmental parameters.

Beginning at the largest scale, we first composite the large-scale meteorological variables (SST, winds, subsidence) from NCEP reanalysis output for each group of cloud organization cases. Figure 4 combines these composited variables with 24-hour back trajectories, originating at a height of 500 m, for all the cloud cases in each population. Both closed- and open-cellular cloud organization cases show a similar SST at the ENA site and, as expected based on our selection criteria, a significant northerly component to the surface wind direction but with stronger winds for the open-cellular cases. The closed-cellular cases are essentially occurring under a classic Azorian surface high pressure system with weaker winds, resulting in noticeable variability in the originating locations of the 24-hour back trajectories. On the other hand, the open-cell cases form under relatively stronger cold-air advection (see also Table 3), i.e., cold air outbreaks, with the large majority of the back trajectories originating from the Greenland region.

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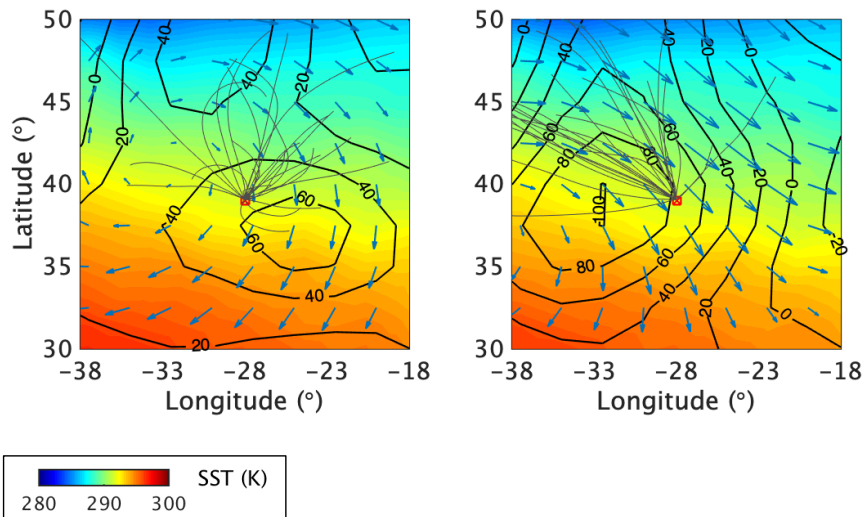
Figure 4: Averaged sea surface temperature [K] (colors), winds (arrows) and 700 hPa large scale subsidence [hPa day⁻¹] (contours) for closed cell (left) and open cell (right) cases from NCEP. The thin gray lines indicate the location of parcels 24 hours prior to reaching the site. The ARM ENA site is indicated by the red square.

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Next, we investigate the differences in the boundary layer thermodynamic structure between the groups of open- and closed-cellular cloud cases. Using twice-daily (nominally 1100 and 2300 UTC) radiosonde profiles, we calculate the lower tropospheric profiles of potential temperature, water vapor mixing ratio, and the zonal and meridional components of the horizontal wind. These profiles are used to compile composites for the closed- and open-cellular cloud cases (Fig. 5). For these composite profiles, the height coordinate is normalized by the height of the base of the subsidence inversion to avoid smoothing of the boundary layer thermodynamic structure (Augstein et al. 1974; Mahrt, 1976; Albrecht et al., 1995). We note a distinct difference showing a higher inversion-base height (approx. 2 km compared to approx. 1 km), identified both from the potential temperature and water vapor mixing ratio profiles, for the open-cellular cases which manifests in a higher cloud-top height (approx. 1.9 km compared to 1.4 km), see Table 3. The composite sounding shows that the closed-cellular cases tend to occur in a more well-mixed, moister boundary layer compared to open-cellular cases. This is consistent with our expectations as the organization in closed-cellular cases are driven by relatively stronger cloud-top radiative cooling (e.g., Shao and Randall, 1996) and occur under shallow boundary layer depths with less precipitation. Consistent with Figure 4, the radiosonde profiles also show composite winds near the surface from the north-northwest, as expected from our selection criteria, for both groups

of cloud organization cases, with significantly stronger wind speeds for the open-cellular cases. When the composites are broken down by the radiosonde launch time (Fig. 6), a difference in the time evolution becomes apparent with open-cellular cases showing greater variability as a function of time with stabilization in the boundary layer (i.e., a greater increase in potential temperature with height) and a moistening of the free troposphere (above 2 km). x This time evolution of the thermodynamic structure is due to a combination of the diurnal cycle and advection. Given that the mean advective tendencies show cold- and dry-air advection for both cloud organization populations, we would expect advection to drive cooling and drying with time for each cloud organization population. Figure 6 shows small to little increase in the temperature and a moistening of the boundary layer with time for both open- and closed-cellular cases. This suggests that the local diurnal cycle plays the more significant role in the evolution of boundary layer thermodynamics. The difference in the mean sensible and latent heat fluxes between open- and closed-cellular cases is consistent with the changes in thermodynamic structure showing larger fluxes, and so larger increases in temperature and moisture for the open-cellular cases.

Both open- and closed-cellular cases show a moistening of the boundary layer with time, with the change being more pronounced for the open-cellular cases.

The most obvious signature of MBL cloud mesoscale organization are the horizontal and vertical distributions of cloud occurrence. Figure 7 shows a two-dimensional histogram for KAZR radar reflectivity and the vertical profile of cloud occurrence (Xie et al., 2010) for each cloud population. The closed-cellular cases have shallower cloud layers with lower radar reflectivity while the open-cellular cases show smaller cloud frequency of occurrence in the main cloud layer, but deeper cloud layers with overall higher values of radar reflectivity. This is qualitatively consistent with single (or a few) case radar observations of transitions between open- and closed cellular cases from previous studies (e.g. Stevens et al. 2005, Comstock et al. 2007, Wood et al. 2011) and provide confidence in the visual inspection selection procedure that we used.

Given this confirmation of our selection criteria, we calculate the mean and standard deviation for a number of descriptive and potentially explanatory variables for the closed- and open-cell cloud populations (Table 3). There are several noteworthy differences between the populations:

- Regarding the thermodynamic structure of the MBL, closed-cellular cases occur during times of stronger mean inversions (regardless of the thermodynamic variable used to quantify the inversion strength), but with a similar physical inversion depth. The lower tropospheric stability (LTS; Klein and Hartmann, 1993) is significantly stronger for the closed-cellular cases despite a similar mean SST.

- The mean subsidence rate (quantified as the value of the large-scale vertical velocity, omega, at the 700 hPa pressure level) is larger, by a factor of 50 percent for the open-cellular cases. This greater subsidence rate is likely the result of post-frontal circulations during cold-air outbreak conditions. The deeper boundary layers for the open-cellular cases must be maintained by a combination of increased turbulence (Ghate et al., 2019, 2020) generated by the stronger wind speeds and significantly larger surface sensible and latent hear fluxes compared to closed-cell cases (Kazil et al., 2014; Wang et al., 2010), a consequence of the smaller cloud fractions and thus greater surface downwelling radiative fluxes.

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- When precipitation is reaching the surface, and thus measured by the surface disdrometer, the mean surface rainfall rate is significantly higher and mean raindrop sizes are significantly larger for the open-cellular cases.

- Mean aerosol concentrations, as estimate from the CPC, are significantly larger for the closed-cell cases This is consistent with the larger rainfall rates and related increase in precipitation scavenging (Zheng et al., 2019).

- The mean cloud-top height and cloud-base height are lower, with an accompanying smaller cloud thickness for the closed-cellular cases compared to the open-cellular cases. Correspondingly, mean LWP, when clouds are present, is twice as large for the open-cellular cases compared to the closed cellular cases. Also, the PWV is identical for the two organizations with slightly more variation for the open-cellular cases.

- The cloud fraction, defined as the frequency of occurrence of cloud below 4 km in the column above the ceilometer (Wu et al., 2014), is greater for the closed-cellular cases compared to the open-cellular cases.

- Both open- and closed-cellular populations show mean cold- and dry-air advection with the open-cellular cases showing stronger cold-air advection, and the closed-cellular cases showing stronger dry-air advection.

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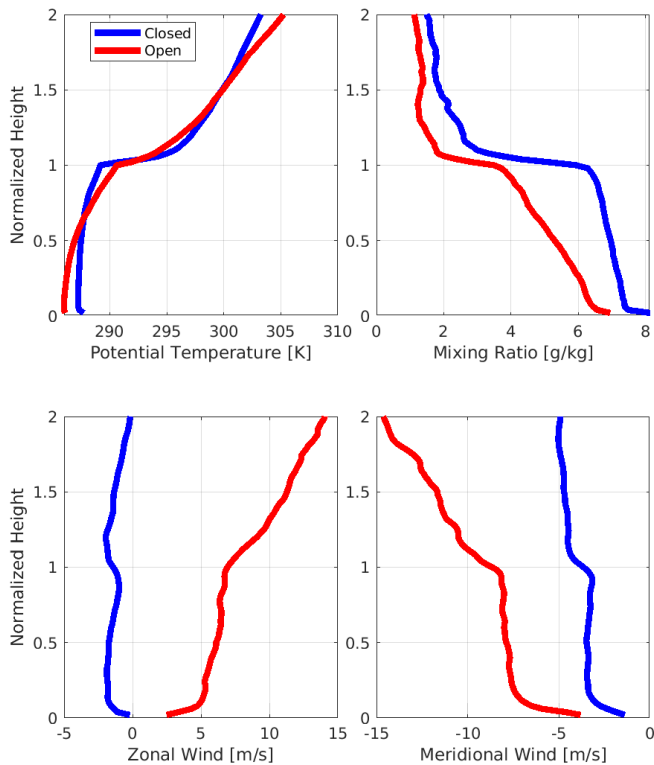
Figure 5: Composite profiles of radiosonde reported (a) potential temperature [K], (b) water vapor mixing ratio [g/kg], and (c) zonal and (d) meridional wind speed [m/s] for closed- (blue) and open-cellular (red) cloud populations. The height coordinate is normalized by the boundary layer depth.

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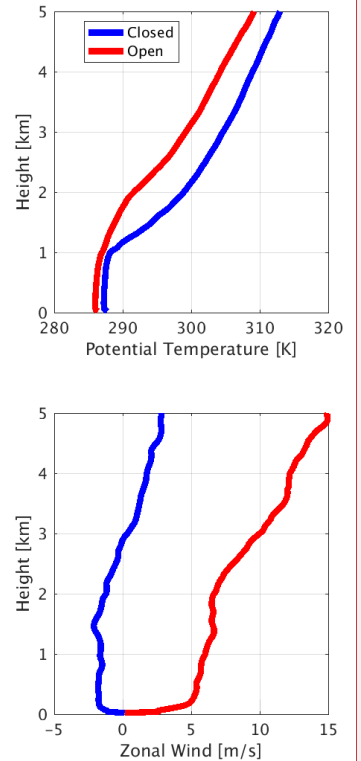
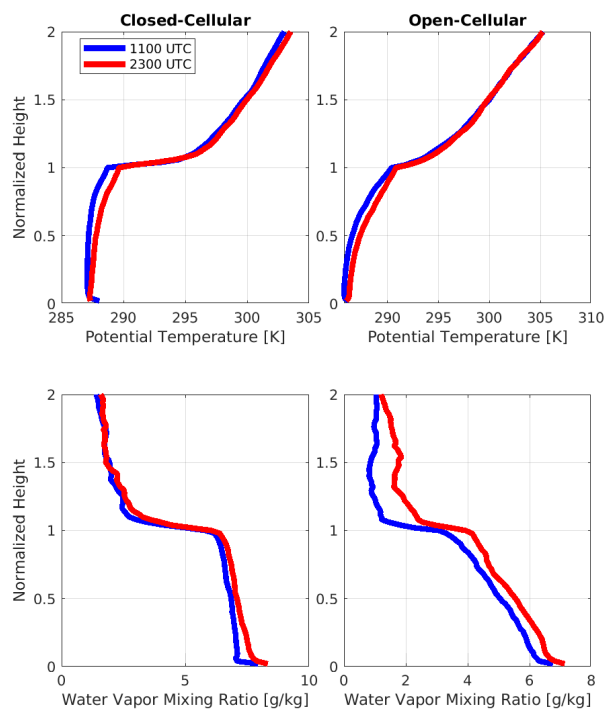
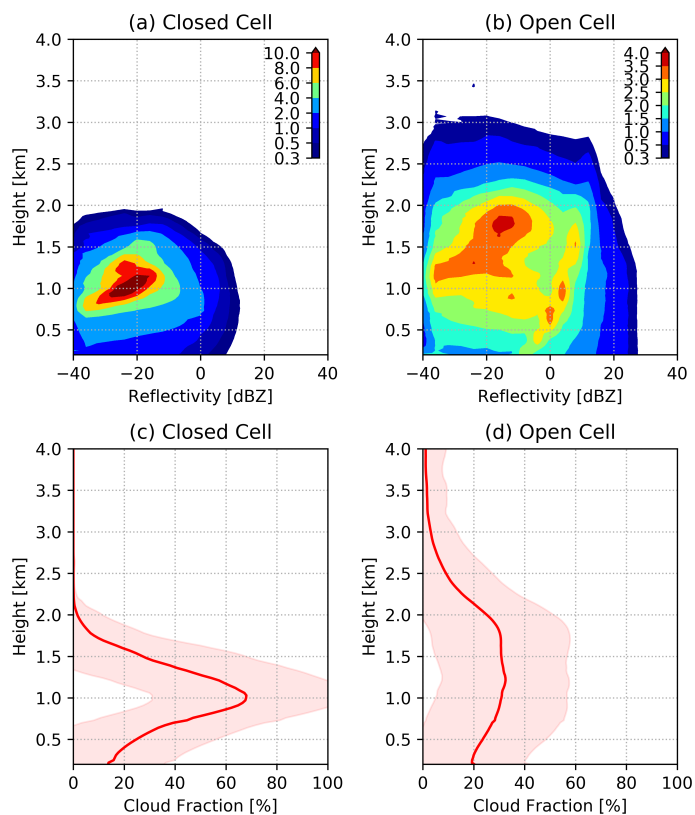


Figure 6: Comparison of composite profiles of radiosonde reported (top) potential temperature [K] and (bottom) water vapor mixing ratio [g/kg] for closed-cell (left column) and open-cell (right-column) at 11 UTC (blue) and 23 UTC (red). The height coordinate is normalized by the boundary layer depth.

320 **Figure 7: Contoured frequency by altitude diagrams (CFAD) of radar reflectivity for (a) open cell and (b) closed cell cases. Vertical profiles of hourly cloud fraction from ARSCL for (c) open cell and closed cell cases. Lines are means; shadings are 1 sigma standard deviations.**

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Parameter	Closed Cellular mesoscale organization (Mean±Std)	Open Cellular mesoscale organization Mean±Std
Number of Hours	588	536
Ceilometer Cloud Fraction (%)	93±14	57±24
Fraction of Column Max Echoes > -20 dBz (%)	71±28	47±24
Cloud Top (km)	1.4±0.3	1.9±0.7
Cloud Base (km)	1.0±0.3	1.3±0.5
Thickness (km)	0.4±0.3	0.6 ± 0.6
LWP (g/m^2)	91±108	240±340
PWV (cm)	1.66±0.49	1.66±0.58
Inversion Strength (ΔT K)	7.31±2.02	4.19±2.89
Inversion Strength (Δr g/kg)	-3.02±1.82	-1.76±1.36
Inversion Depth (ΔZ m)	184.05±110.68	184.54±178.93
SST	291.67±1.95	291.50±2.21

	(K)		
Sensible Heat Flux	30.20±26.95	66.48±39.77
Latent Heat Flux	131.10±79.31	246.88±123.35
LTS (K)	22.41±0.35	13.31±0.91
Omega at 700 hPa (hPa/day)	50.50±72.53	74.23±112.76
Temperature Advection (K/day)	-1.79±2.39	-2.36±2.82
Moisture Advection (g/kg/day)	-1.4±1.87	-1.09±2.22
CPC Aerosol Concentration (cm^{-3})	328±250	273±179
Surface 5-min Rain Rate (mm/hour)	0.15±0.32	0.73±1.62
Surface D_0 (mm)	0.54±0.25	0.88±0.38

Table 3: Mean and standard deviations of important bulk parameters during closed- and open-cellular conditions.

To further explore the characteristics of the closed- and open-cellular cloud organization populations, the probability
340   density function (PDF) of LWP, PWV, cloud top height, cloud depth, cloud fraction and echo fraction are examined in Figure
8. Both the open- and closed-cellular cloud cases exhibit a bimodal distribution of LWP (Fig. 8a), with the local maximum
being more clearly separated for the open-cellular cases. This bimodal distribution is consistent with contributions from
drizzling and non-drizzling clouds and higher precipitation rates and associated LWP for the open-cellular cases. The two
cloud populations show very little difference in the PWV (Fig. 8b), with the closed-cellular cloud cases showing a more well-
345   defined peak at the mode values, consistent with the similarity in the bulk properties presented in Table 3. Both cloud
organization populations show a peak cloud-top height (Fig. 8c) at 1.2 km, however, the open-cellular cases show a stronger

Deleted:

Deleted: peak value for

Deleted: lightly lower than that for open-cellular cases

350 peak around 2 km, consistent with the bimodal structure in LWP (Fig. 8a) where higher cloud-top heights are associated with larger LWP and precipitating clouds. A comparison of the observed cloud depths (Fig. 8d) is consistent with the LWP and cloud-top height differences showing the open-cellular cloud organization with an extended tail towards larger cloud thicknesses compared to the closed-cellular cloud population. Finally, the cloud fraction (Fig. 8e, ceilometer observed frequency of occurrence) and echo fraction (Fig. 8f, includes observed drizzle but will exclude thinnest cloud) of open-cellular

355 cloud cases suggests large variability in cloud coverage with a mode around 50% compared to much greater coverage for closed-cellular cases. In summary, the open-cellular cloud organization cases have more varied cloud top heights, thickness and cloud fraction, while the closed-cellular cloud organization cases overcast with more uniform cloud cover and thickness.

The diurnal variation of composite cloud macro-physical properties within the two cloud populations is shown in Figure 9. As expected from the results shown in Table 3 and Figure 7, the composite total cloud fraction (Fig. 9a) for closed-cell cases

360 is greater than that for open-cell cases through the entire diurnal cycle. Closed-cell cases are nearly overcast with little variation in total cloud fraction during night-time and early morning hours (nominally 22 – 10 UTC) with a subsequent decrease, but greater variability, during the daytime. Similarly, open-cell cases show relatively smaller total cloud fraction during the afternoon-evening hours (nominally 16 – 23 UTC), compared to the overnight and morning hours, but with similar variability over the population across the diurnal cycle. These diurnal cycles in cloud fraction, with decreased low cloud coverage during

365 the afternoon hours, are broadly consistent with previous studies in the Azores region (Miller et al., 1998; Rémillard et al. 2012). Also consistent with Table 3, the mean cloud depth (or thickness) (Fig. 9b) for open-cell cases is larger than that for closed-cell cases, however the variability among the open-cell cases is much greater than the closed-cell cases. This is consistent with the regular occurrence of a mixture of deeper, precipitating cumulus and stratocumulus in the open-cell cases. For the LWP and PWV (Fig. 9c,d), the open-cell cases exhibit much greater variability, with the mean values generally being

370 slightly larger during the night-time and early morning hours. The hourly mean values, of both LWP and PWV, tend to converge during the day-time hours consistent with the decrease in physical cloud thickness (Fig. 9b).

Histograms of surface rainfall DSD characteristics show distinct differences between closed- and open-cellular cloud populations (Fig. 10). Open-cellular cloud cases tend to have larger median drop sizes, liquid water contents (LWCs) and rain rates compared to closed-cellular cases, while the closed-cellular cases have somewhat larger numbers of droplets based on

375 the intercept of the DSD. This result is consistent with the results in Fig. 8 and Table 3 showing that the open-cellular cloud organization cases tend to have larger cloud thickness, higher cloud-top heights, and larger LWPs compared to the closed-cellular cloud organization cases. It is also important to note that rain drops larger than 1 mm diameter are often observed in both closed- and open-cell cases. This is an important threshold for millimeter cloud radar observations (Lehrmitte 2002; Giangrande et al. 2010) as the Rayleigh approximation becomes invalid and non-Rayleigh scattering must be considered when

380 interpreting the observed radar reflectivity factor.
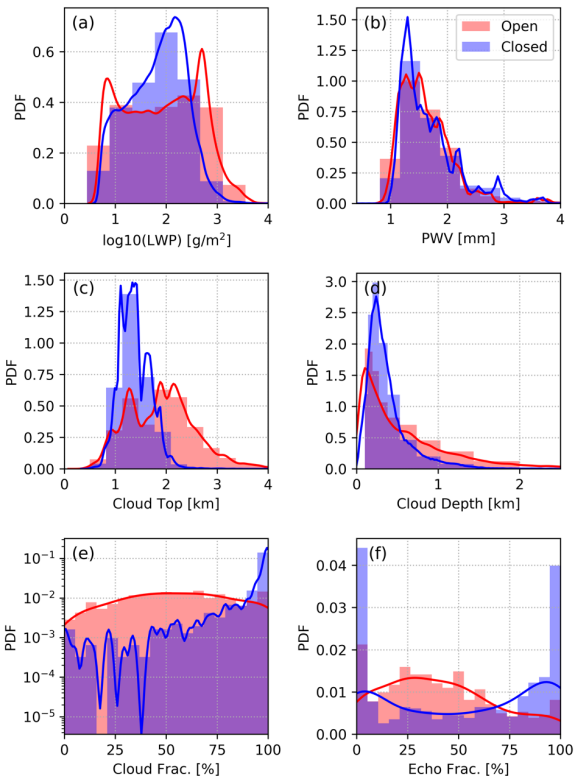
Figure 8: Histograms of (a) LWP, (b) PWV, (c) cloud top height, (d) cloud depth, (e) ceilometer cloud fraction [Note: logarithmic scale on vertical axis], and (f) KAZR echo fraction for open (red) and closed cellular (blue) mesoscale organizations. The lines are the probability density functions (PDFs) calculated using gaussian kernel density estimation.

Figure 9: Diurnal cycle of the mean and standard deviation of the (a) 1-D cloud fraction [%], (b) cloud depth (or thickness) [km], (c) liquid water path [g/m2] and (d) precipitable water vapor [mm] for open-cell (red) and closed cell (blue) cases. The markers represent the mean values, the shadings represent the 1 sigma standard deviations.
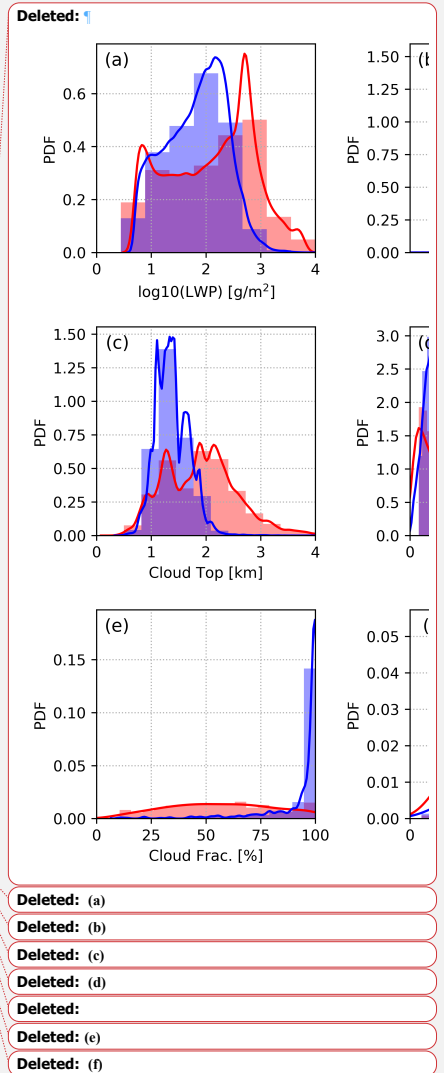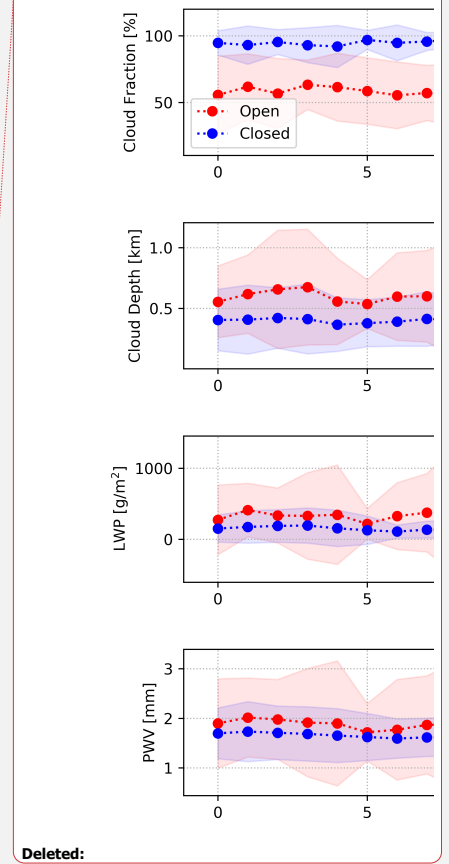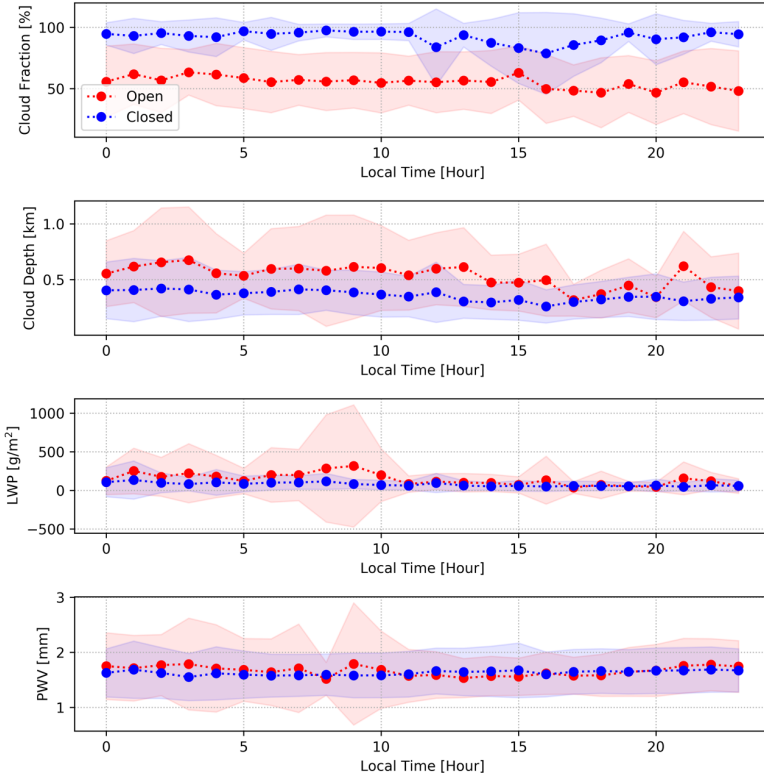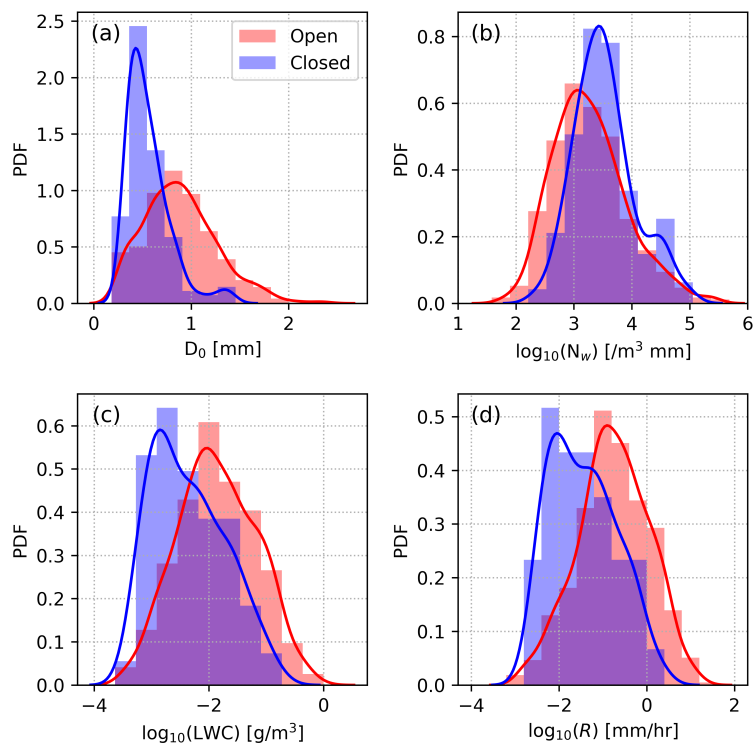
Figure 10: Comparison histograms of precipitation drop size distribution parameters: (a) median volume drop size (D_0), (b)DSD intercept parameter (N_w), (c) liquid water content, and (d) rain rate (RR) for open and closed cell cases. The lines are the probability density functions (PDFs) calculated using gaussian kernel density estimation.

5 Summary and Conclusions

MBL stratocumulus clouds often exist in two distinct states of mesoscale organization: closed- and open-cellular. The different cloud organization populations differ significantly in their low-level cloud fraction, shortwave albedo and LWP. A number of studies over recent decades have identified the importance of aerosols, drizzle and their impacts on boundary layer

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dynamics in determining the cellular organization of clouds within the MBL. However, the interplay of these different mechanisms and their relative importance under different regimes remains and area of active research and advances in our understanding are necessary for improved representation in large-scale atmospheric models (Wood et al., 2016; Jensen et al., 2016). Using long-term observations from the ARM ENA site, we have compiled composite cloud, precipitation, and thermodynamic properties over a carefully selected population of open- and closed-cellular MBL cloud cases. Important highlights of this analysis include:

- Both open- and closed-cellular cloud organization cases occur during similar SST conditions, however, open-cellular cases are distinguished by stronger cold-air advection and subsidence compared to closed-cellular cases, consistent with the open-cellular cases forming during cold-air outbreaks.
- The open-cellular clouds were associated with deeper boundary layers, stronger winds, higher rain rates and stronger large-scale subsidence compared to their closed-cellular counterparts.
- Raindrops with diameter larger than 1 mm were routinely recorded at the surface during both organizations, with a higher number of large drops during open-cellular cloud cases.

To distinctly summarize the similarities and differences in the environmental and cloud properties during the two mesoscale cloud organization populations, we normalized the data using the average of the samples available for both closed- and open-cellular cloud cases. Figure 11 shows the mean of the normalized data for a number of important boundary layer and cloud properties. These results comparing the surface condensation nuclei (CN), PWV, and depth of the boundary layer inversion (Inv. ΔZ) show very little difference between the two populations (difference in normalized values less than 0.2). More significant differences are shown in all the other variables. The LWP and rain rate, which are expected to co-vary (e.g., Wood 2005; Zuidema et al., 2005; Geoffroy et al., 2008; Serpetzoglou et al. 2008; Kubar et al. 2009; Rémillard et al., 2012), exhibit the largest normalized differences of greater than 0.8. The boundary layer inversion strength, in terms of the jump in either potential temperature (Inv. $\Delta\theta$) or water vapor mixing ratio (Inv. Δr), was stronger during closed-cellular cases. Although the rain fraction was higher for the closed-cellular cloud cases, due to their greater coverage, they had significantly lower rain rates as compared to the open-cellular clouds cases. In addition to providing insights into the environmental and cloud characteristics during these two mesoscale cloud organization populations, the similarities and differences noted here can be used to evaluate model simulations of the two cloud organizational states.

Follow-on research activities are aimed at exploring differences among the internal properties of the marine boundary layer such as dynamics and microphysics for the closed- and open-cell cloud populations. In addition, another research avenue is the use of satellite and reanalysis data to investigate the evolution of the cloud and boundary layer properties along the parcel trajectories for both cloud populations.
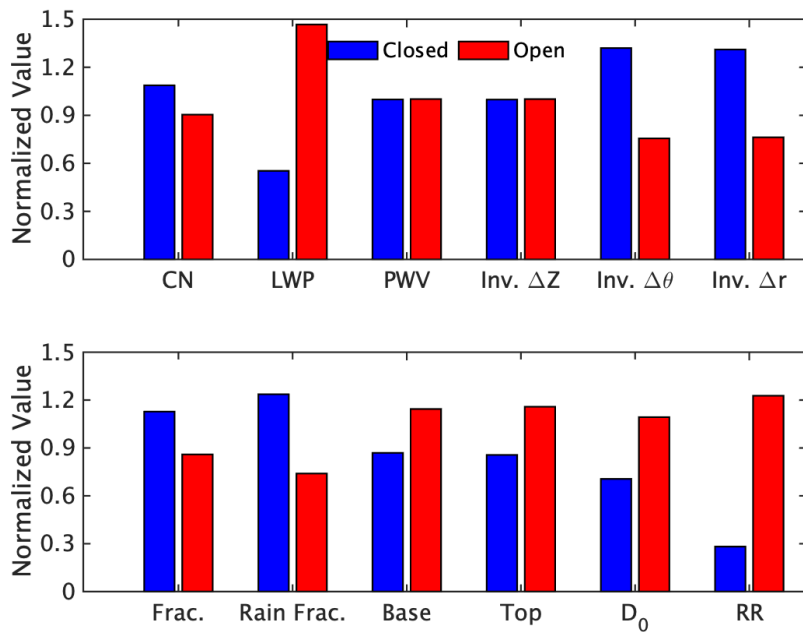
Figure 11: Normalized values of boundary layer and cloud properties for closed cellular (blue) and open cellular (red) organizations.

Data Availability

455 All ARM ENA datasets are available from the ARM data archive (https://www.arm.gov/data) (ARM 2013, 2014a, 2014b, 2015). Meteosat data were obtained from the Satellite ClOud and Radiation Property (SatCORPS) team at NASA Langley Research Center https://satcorps.larc.nasa.gov/prod/exp/amf_azores_reprocessed/visst-pixel-netcdf and https://search.earthdata.nasa.gov/search?q5CER_GEO_Ed4_MET10. (NASA 2019a, 2019b). One-degree global reanalysis data is available from http://www.esrl.noaa.gov/psd/data/gridded/data.ncep.reanalysis.html.

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Author Contributions

Michael Jensen, Virendra Ghate: project design and management, data analysis, interpretation, lead writing. Die Wang: data analysis and visualization, writing – review and editing, Diana Apoznanski: data analysis and visualization, Scott Giangrande, interpretation of results, writing – review and editing, Mary Jane Bartholomew, Karen Johnson, Mandana Thieman – data preparation and analysis, writing – review and editing

Competing Interests

The authors declare that they have no conflict of interest.

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