

We thank both the reviewers for their careful consideration of our manuscript and their helpful comments. We have reproduced the reviewer comments in full below, where we address each question and concern individually. Reviewer comments are written in black, our responses are in blue, and new text added to the manuscript is italicized.

REVIEWER 1:

Bates et al. have developed a new gas-phase chemical mechanism to simulate the oxidation chemistry for benzene, toluene, and xylenes in box and 3D models. They evaluated their model against chamber measurements as well as benchmarked it against other competing mechanisms. They find that their reduced mechanism performs well while noting that their mechanism alone models radical recycling from phenoxy-phenylperoxy that has modest effects on concentrations of smaller oxygenated products and oxidants. They also present interesting results from simulations performed separately with a boundary layer and 3D model.

Aromatics are an important class of organic compounds relevant for both anthropogenic (e.g., traffic) and biogenic (e.g., biomass burning) sources. Hence, the focus on aromatics and the development of a compact mechanism to simulate its atmospheric chemistry is well motivated. The manuscript was very easy to read and follow and I commend the authors for putting together a superb paper. I am less conversant with the chemical reactions described in the methods section so I am going to let the other reviewer(s) and editor to directly assess that section of the manuscript. However, I was able to follow the dominant chemical pathways and oxidation products and how they helped explain the results presented later. I recommend this paper for publication to ACP noting that the authors should consider my comments below before final submission.

1. *Beyond BTX: My sense is that the impacts described here are likely to be biased (mostly underestimated) for the following reasons. First, how are aromatic species other than benzene, toluene, and xylenes accounted for in the modeling? My sense is that C8 and higher aromatics are lumped into xylenes. Or are they parsed by reactivity with OH? What about the treatment of multi-ring aromatics (e.g., naphthalene)? **This attribution (or lumping) of non-BTX species to BTX surrogates needs to be made clear throughout this work, in addition to how this attribution is justified and/or adds to the uncertainty.** Furthermore, what would be good to consider here would be to document the relative magnitude of non-BTX species to BTX. Also, are the authors concerned that certain species (e.g., furan in SAPRC) that are assigned to an aromatic surrogate (in the case of furan and SAPRC, to xylene) for computational reasons skew the model findings presented in this work? (It should not be a concern for the box modeling, correct?) Second, recent combustion emissions work has pointed out that gasoline exhaust contains a wide variety of single-ring aromatic hydrocarbons separate from BTX species (e.g., Zhao et al., ES&T, 2016; Drozd et al., ES&T, 2019; Lu et al., ACP, 2018). This seems to be less true for diesel exhaust. These non-BTX species have been shown to be very important for SOA formation and, in some models, have been accounted for as intermediate volatility organic compounds (IVOC). Third, biomass burning has been shown to be an important source of oxygenated aromatics including phenols and methoxyphenols (e.g., Koss et al, ACP, 2018; Hatch

et al., ACP, 2018). Since some of these biomass burning aromatics feature in the chemical mechanism presented in this work as intermediate oxidation products (e.g., phenol, catechol), I believe that they are likely to have similar effects as the explored BTX species. It's likely that the emissions of these oxygenated aromatics are considered in the GEOS-Chem inventories but it has not been made clear. All 3 factors described above will tend to generally increase (?) the effect size for aromatics.

First, to clarify, neither the box nor global model simulations lump larger aromatics or non-aromatics into the BTX chemistry described here. C₉+ aromatics are simply not included in the chemistry of the box and global models, so their contributions are not artificially boosting the simulated importance of BTX. Multi-ring aromatics are only included in GEOS-Chem insofar as they contribute to SOA via emission of a naphthalene-like IVOC surrogate (see Pai et al., 2020, and Pye et al., 2010), and do not participate in GC13 gas-phase chemistry. As the reviewer notes, in some implementations of SAPRC furans are lumped with xylenes, but that is not the case here. Thus, we are not concerned that GC13 chemistry of non-BTX species affects the chemistry described in this manuscript. To clarify this, we have added the statement "No emissions of other gas-phase aromatics are included" to section 5.1

We agree, though, that the oxidation of larger aromatic species can be an important atmospheric contributor to OVOC budgets, oxidant and radical chemistry, and SOA production. Including the emissions and chemistry of these species (and emissions of oxygenated aromatics already in GC13) would likely, as the reviewer notes, increase the effect sizes shown in Figures 10-14. We initially considered such species beyond the scope of this manuscript, but based on the reviewer's interest have decided to add a small additional section (6) describing the potential importance of some larger aromatics and providing recommendations for their implementation in GEOS-Chem and other models with a heavily simplified mechanism:

" Our GC13 mechanism and its implementation in GEOS-Chem focused on the effect of BTX emissions. Taraborrelli et al. (2021) estimated that non-BTX aromatics contribute 54% of global aromatic emissions by carbon mass with C₆-C₈ oxygenates, C₈-C₉ aromatic hydrocarbons, and higher aromatics each contributing approximately one thirds of this non-BTX fraction. The contributions from C₆-C₈ oxygenate emissions including benzaldehyde, phenol, cresols, and catechols can be readily accounted for in GC13 since they are already included explicitly as secondary species. C₈-C₉ aromatic hydrocarbons including ethylbenzene, styrene, and trimethylbenzenes can be added as independent species using simplified initial oxidation reactions to convert them directly to products already included in GC13. We have provided such reactions, simplified from MCM (Jenkin et al., 2003), in Section S1 of the SI, and this would add 3 species and 7 reactions to GC13. Higher aromatics such as polycyclic aromatic hydrocarbons (PAHs) are semi-volatile and their oxidation products partition heavily into the aerosol phase (Chan et al., 2009; Chen et al., 2016), so their simulation is more relevant to SOA formation than to oxidant chemistry."

2. *Vapor wall losses: Page 10, line 4: Really? Recent work has argued that the timescales for vapor losses to the chamber walls are on the order of 5 to 15 minutes for a chamber volume

of $\sim 10 \text{ m}^3$. Since these losses have been shown to be important for SOA formation (Zhang et al., PNAS, 2014; Cappa et al., ACP, 2016; Akherati et al., ACP, 2019), I would expect them to have a similar effect on gas-phase chemistry and radical and oxidant concentrations. Further, Zhang et al. (2014) show that the vapor wall loss effect on SOA formation is different under different NO_x conditions. Perhaps this mechanism can shed light on why that might be? Vapor wall losses can be modeled relatively easily following the work of Zhang et al. (PNAS, 2014) and Krechmer et al. (ES&T, 2016).

Vapor wall losses can indeed be critically important for measured SOA yields, but GC13 does not seek to represent SOA formation nor to explicitly represent the wall-loss-prone compounds that contribute directly to SOA. While such semivolatile intermediates may be sources of OVOCs, even without wall losses their influence is likely to be small early in the chamber oxidation process, when most reactivity is attributable to the parent hydrocarbon. We purposely isolate early product yields from both simulations and experiments to avoid the longer-term, multi-generational chemistry that may be affected by the loss of intermediate products to the walls. Effects of semivolatile intermediates on oxidant concentrations and radical dynamics are also expected to be minor early in the experiment, and we do not use chamber simulations to constrain oxidant effects in the mechanism. Shedding light on the role of vapor wall losses in the NO_x -dependence of SOA yields is beyond the scope of this work and beyond the focus of the highly condensed GC13 mechanism; an improved representation of SOA from aromatics in GEOS-Chem is the subject of ongoing work independent of GC13. To clarify these points, we have expanded our discussion of vapor wall losses in section 4.1: *“Wall losses of semivolatile gases are not represented in these simulations, as we do not seek to model SOA formation or the yields of direct SOA precursors. An improved representation of SOA from aromatics in GEOS-Chem will be the subject of future work. While it is possible that wall losses of semivolatile intermediates affect experimental yields of OVOCs, this multigenerational chemistry is expected to play only a minor role at the short time scales isolated here.”*

3. *Chamber simulation results and comparisons (Figures 4-6): While I understand the rationale behind the study design to simulate a representative chamber experiment, this work should also explore sensitivity in model predictions (limited to GC13 perhaps) to a broader range of input conditions observed across the chamber experiments used to evaluate the mechanism. For instance, are the model predictions sensitive to whether one assumes a constant OH profile chamber experiment (relevant for OH produced from photolysis of H_2O_2) or an OH profile that decays with time (relevant for OH produced from photolysis of HONO)? Are the model predictions sensitive to the chamber lights (e.g., intensity, spectrum) that should control the time-varying concentrations of NO and NO_2 ? How do the model predictions change if the results were presented for a different time point (instead of 20 minutes after lights on)? How about the influence of chamber size (5 m^3 vs. 30 m^3) and mode of operation (batch mode vs. steady state model)? This mechanism sensitivity will be useful in interpreting the comparison with observations.

We were also curious about these various sensitivities and conducted a large number of such simulations to test our box model setup before arriving at the “default” simulations shown

in Figures 4-6. However, the focus of this work is on the mechanisms themselves (and sensitivities therein) rather than the parameters by which chamber experiments are conducted, so we did not opt to show these sensitivities in the main manuscript, but we did include results from many such sensitivity simulations in the SI. For the most part, initial yields reported in section 4.2 exhibit only minor sensitivities to the parameters we varied, which is why we did not elaborate on them in the main text. The most interesting sensitivities we found were to experiment duration (Figs. S3-5), temperature (Figs. S9-14), and initial VOC concentration (Figs. S15-S20), most of which were minor, although GC13 does exhibit some sensitivity of glyoxal and formic acid yields from benzene and of formaldehyde yields from toluene/xylene to initial VOC concentration. We have expanded this section of the SI by adding Figures S25-26 showing sensitivities to light intensity and oxidant concentration, and by providing greater descriptive detail in the accompanying text of Section S4. Furthermore, we have elaborated on these sensitivities in the main text and provided additional references to the SI: *“Initial’ yields shown in Figures 4-6 are after 10 min of photooxidation; additional results showing long-term yields after 24 h of oxidation are provided in Section S2 of the SI. Sensitivities to temperature and initial VOC concentrations are generally small, but additional results showing the effects of these parameters, as well as the effects of light intensity and oxidant source, can be found in the SI.” ... “We also find in GC13 simulations that the glyoxal yield from benzene is sensitive to the initial benzene concentration in chamber experiments (see Section S4 of the SI), which may further explain the spread of measured yields.” ... “long-term formaldehyde yields exhibit a temperature sensitivity of up to +3% K⁻¹”*. Finally, we have added a caveat to the beginning of Section 4.2 to draw readers’ attention

4. ***Boundary layer and 3D model evaluation:** The authors should discuss if and how the mechanism changes presented in this work have the potential to improve predictions of oxidants and gas-phase species that are likely to be biased based on past literature.

We hope that a major benefit of GC13 and its documentation in this work will be its potential to improve such predictions in future applications, and we believe that the detailed descriptions of the mechanism, its sensitivities, and its comparison to previously used mechanisms already go some way to showing how GC13 can improve biases. We believe that extensive evaluation of the ambient box modeling and global simulations against atmospheric data are beyond the scope of this work, and a detailed assessment of past literature would be so specific to the conditions (*e.g.*, mechanisms and models used) of each study as to lose general interest. However, we have elaborated throughout the manuscript on points of comparison with previous studies, in hopes of drawing attention to the ways in which GC13 might change model outcomes, improve predictions, and alleviate biases. Added passages include:

On ozone in box models: *“GC13 and MECCA have less ozone production than MCM, RACM, and SAPRC, all of which are known to overestimate peak ozone from chamber experiments (Bloss et al. 2005a, Goliff et al. 2013, Carter et al. 2013). ... The reduced ozone formation in GC13 and MECCA relative to other mechanisms may improve model biases relative to chamber experiments (Bloss et al. 2005a, Goliff et al. 2013, Carter et al. 2013), and would*

likely reduce the high simulated contribution of aromatics to ambient ozone formation in box model analyses of polluted environments (Luecken et al., 2018; Oak et al., 2019; Schroeder et al., 2020)."

On HO_x in box models: *"It has been reported that other mechanisms tend to underpredict HO_x concentrations in simulations of chamber experiments (Chen, 2008), and both Bloss et al. (2005a) and Carter et al. (2013) comment on the need to increase HO_x recycling from aromatic oxidation; GC13 thus brings HO_x concentrations into improved alignment with chamber results."*

On global glyoxal: *"Our simulated contribution of aromatic oxidation to the budgets of glyoxal and methylglyoxal are substantially larger than in previous GEOS-Chem studies using simplified mechanisms with lower yields (Fu et al., 2008; Silva et al., 2018), but are more consistent with results from the detailed mechanisms in Taraborrelli et al. (2021) and Yan et al. (2019). The increased glyoxal yields in GC13 align with general model findings of negative biases relative to satellite observations of glyoxal columns in regions with strong anthropogenic influence (Chan Miller et al., 2016; Liu et al., 2012; Myriokefalitakis et al., 2008; Silva et al., 2018; Stavrakou et al., 2011)."*

5. Page 10, line 30: I have always wondered what the NO_x concentrations are in 'low NO_x' chamber experiments. Literature from the Caltech group led by Prof. John Seinfeld has reported low NO_x experiments to have NO_x concentrations under 2 ppbv, reflecting the limits of quantification for their NO_x analyzer. Assuming other chamber groups have encountered the same quantification problem with NO_x, I wonder if a '10 pptv' assumption is justified here.

We agree with the reviewer that this can frequently be an issue in chamber experiments, and that assumptions of "NO_x-free" conditions may not always be verifiable with limited instrumentation. We have included error bars on reported [NO_x]₀ values where instrumental uncertainties are given, and have added a brief discussion of this point to the text in section 4.2: *"Experiments conducted with no added NO_x are shown at [NO_x]₀ = 10 ppt (below which modeled yields are invariant with [NO_x]₀) regardless of whether NO_x concentrations were monitored during the experiment. In some cases, NO_x may off-gas from chamber walls (Carter et al., 2005). When reported, instrumental uncertainties on measured NO_x mixing ratios are shown as horizontal error bars on experimental points."*

6. Page 11, line 6: Here and elsewhere, I would encourage the authors to be quantitative on what 'low-to-moderate' NO/NO_x and 'high' NO/NO_x means.

We thank the reviewer for pointing this out; references to "high" and "low" NO/NO_x can often be vague or have differing cutoffs depending on context. We have scoured the manuscript for mentions of high and low NO_x and, where possible and appropriate, clarified by adding quantitative thresholds. The passage singled out by the reviewer has been amended to read: *"fixed yields of 53-57% fit most data under low-to-moderate NO conditions, but when ~~at~~*

high $[NO_x]_0$ exceeds 100 ppb the observed yield declines,” and numerous other statements have been edited similarly (please see the track-changes document).

REVIEWER 2:

Bates et al. present work on the development of an updated chemical mechanism for aromatic chemistry for use in computational models of atmospheric chemistry. They demonstrate the mechanism performs well as assessed against state-of-the-art laboratory measurements and compare the simulated abundances of compounds like glyoxal, OH, and ozone across different implementations of aromatic chemistry. This work is scientifically sound, and the manuscript is generally well written. Barring one major accessibility issue, I recommend that this paper is accepted for publication following a few minor comments below.

Colour blindness and Figures 4-9: Figures 4-9 are not intelligible to colour blind readers. I strongly suggest the authors remake these figures to make them more accessible. There are a range of tools available online to explore if a given colour palette or figure is “colour blind friendly”.

We thank the reviewer for bringing this important point to our attention. Figures 4-9 have been re-rendered with a more accessible Matlab color palette that is distinguishable to readers with deuteranopia, protanopia, or tritanopia (<http://mkweb.bcgsc.ca/biovis2012/color-blindness-palette.png>).

Computational Impact: A major component of this work relates to the development and justification of the simplified mechanism, which is well justified in the manuscript given the described issues associated with computational limitations. It is thus surprising that the authors do not quantify the impact of the mechanism on global model simulation times. While this work likely improves the representation of chemistry in GEOS-Chem, what is the computational cost associated with the increased transport and chemistry demands? If I were a model developer, what are the costs of using this mechanism?

We agree with the reviewer that this is an important point to address, but it can be difficult to quantify in a meaningful way. The computational cost of the added aromatic chemistry will vary between users with a variety of simulation parameters (e.g., duration, spatial and temporal resolutions, mixing schemes, emission fields, and other customizable settings) and hardware parameters (since the CPU time of various components of the model scale differently with the number of cores used). However, we have added a statement in section 5.1 describing the computational cost of the aromatic chemistry: “The added aromatic chemistry in the GC13 simulation increases overall CPU time by an average of 1.7% relative to the base simulation, attributable predominantly to gas-phase chemistry (64%) and transport (26%).” We have also added a similar statement to the conclusions: “GC13 increases the computational cost of an annual simulation by 1.7% relative to a simulation without aromatic chemistry.”

P20 L5: Why are these results different from previous glyoxal simulations in GEOS-Chem (Silva et al 2018), which do not show a large change in glyoxal from aromatic oxidation over the Middle East?

Multiple factors may contribute to the different findings between Silva *et al.* (2018) and this manuscript. First, GC13 produces more glyoxal from aromatic oxidation than the previously implemented GEOS-Chem mechanism; this can be seen in greater detail in Figure S31. Second, the figures in Silva *et al.* (2018) show contributions of aromatics and other precursors to total column burdens of glyoxal, while this manuscript focuses on surface glyoxal. The short lifetimes of toluene and xylene confine their influence closer to the surface, while C₂H₂, which is simulated by Silva *et al.* (2018) to be the dominant contributor to total column glyoxal over much of the Middle East, has a longer lifetime and thus exerts more influence over free tropospheric glyoxal burdens. Third, the two studies use different anthropogenic emission inventories (EDGAR/RETRO with regional modifications in Silva *et al.*, CEDS here) in different years (2005 in Silva *et al.*, 2015-16 here), which changes the distribution of aromatic emissions as well as those of other glyoxal precursors. Finally, changes to oxidation mechanisms, SOA uptake, and other model mechanics between the GEOS-Chem versions used (11-02 in Silva *et al.*, 12.3 here) may affect simulated glyoxal. We have opted not to complicate the manuscript with a detailed analysis of these differences, but have added a reference to Silva *et al.* to the discussion of simulated glyoxal from the implementation of GC13 in GEOS-Chem: “*Our simulated contribution of aromatic oxidation to the budgets of glyoxal and methylglyoxal are substantially larger than in previous GEOS-Chem studies using simplified mechanisms with lower yields (Fu et al., 2008; Silva et al., 2018), but are more consistent with results from the detailed mechanisms in Taraborrelli et al. (2021) and Yan et al. (2019)*”.

P9 L27 & P10 L5: “Environmental chamber simulations” and “Continental boundary layer simulations” seem to be subtitles but aren’t formatted as such.

Rather than format these as headers of sub-sub-sections, we have chosen to keep them as undifferentiated paragraphs within sub-section 4.1, much like the paragraphs describing each mechanism in section 3.

P10 L30: [NO_x]₀ is not defined in the manuscript.

We have added a definition in the first paragraph of Section 4.2: “...we present our results as a function of *initial NO_x mixing ratios ([NO_x]₀)*.”

P18 L7: Does Kwon et al. 2021 differ from earlier implementations of C₂H_x chemistry into GEOS-Chem (e.g. Safieddine et al. 2017)?

The implementation of C₂H_x chemistry in GEOS-Chem by Safieddine et al. (2017) was drawn from the mechanism in CAM-Chem as detailed in Lamarque *et al.* (2012). The C₂H₂ chemistry included in Kwon *et al.* (2021) and here is identical to that in Lamarque *et al.* (2012), while our C₂H₄ chemistry is slightly more complex (with the addition of a hydroxynitrate and an

alkoxy radical), following MCM (Saunders *et al.*, 2003) and IUPAC (Atkinson *et al.*, 2004) recommendations. Because the mechanisms are provided in detail in Kwon *et al.* (2021), we choose not to describe them further here.

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Development and evaluation of a new compact mechanism for aromatic oxidation in atmospheric models

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ABSTRACT

Aromatic hydrocarbons (~~mainly~~ including benzene, toluene, and xylenes) play an important role in atmospheric chemistry but the associated chemical mechanisms are complex and uncertain. Sparse representation of this chemistry in models is needed for computational tractability. Here we develop a new compact mechanism for aromatic chemistry (GC13) that captures current knowledge from laboratory and computational studies with only 17 unique species and 44 reactions. We compare GC13 to six other currently used mechanisms of varying complexity in box model simulations of environmental chamber data and diurnal boundary layer chemistry, and show that GC13 provides results consistent with or better than more complex mechanisms for oxygenated products (alcohols, carbonyls, dicarbonyls), ozone, and hydrogen oxide ($\text{HO}_x \equiv \text{OH} + \text{HO}_2$) radicals. GC13 features in particular increased radical recycling and increased ozone destruction from phenoxy-phenylperoxy radical cycling relative to other mechanisms. We implement GC13 into the GEOS-Chem global chemical transport model and find higher glyoxal yields and net ozone loss from aromatic chemistry compared to other mechanisms. Aromatic oxidation in the model contributes 23%, 5%, and 8% of global glyoxal, methylglyoxal, and formic acid production respectively, and has mixed effects on formaldehyde. It drives small decreases in global tropospheric OH (-2.2%), NO_x (-3.7%) and ozone (-0.8%), but a large increase in NO_3 (+22%) from phenoxy-phenylperoxy radical cycling. Regional effects in polluted environments can be substantially larger, especially from photolysis of carbonyls produced by aromatic oxidation, which drives large wintertime increases in OH and ozone concentrations.

1. INTRODUCTION

Aromatic hydrocarbons are a major class of volatile organic compounds (VOCs) emitted to the atmosphere, with important effects on oxidant chemistry and secondary organic aerosol (SOA) formation. They are emitted from anthropogenic sources, including incomplete combustion, industrial processes, solvent evaporation, and fuel leakage (Na et al. 2004; Koppmann et al. 2007), and have additional sources from biomass burning and vegetation (Misztal et al. 2015; Cabrera-Perez et al. 2016). Aromatics compounds in the gas phase can be directly harmful to human health (Duarte-Davidson et al. 2001; Sarigiannis et al. 2008; Manuela et al. 2012) and their secondary chemistry can play a dominant role in ozone and SOA production in urban air (Barletta et al. 2005; Tan et al. 2012; Khan et al. 2018; Nault et al. 2018; Oak et al. 2019; Schroeder et al. 2020).

The simplest emitted aromatic hydrocarbons – benzene (C_6H_6), toluene (C_7H_8), and xylenes (C_8H_{10}), referred to as BTX – together make up over 20% of global anthropogenic non-methane VOC emissions on a carbon basis (Yan et al. 2019) and up to 60% of urban emissions (Lee et al. 2002; Ran et al. 2009). The oxidative chemistry of BTX, initiated by reaction with the hydroxyl radical (OH), is unique among VOCs due to the stability of the aromatic ring and to their low H/C ratios. Standard chemical mechanisms for the fates of the peroxy radicals (RO_2) and alkoxy radicals (RO) produced in VOC oxidation may thus not apply to aromatics (Vereecken, 2018; Xu et al. 2020). Low-volatility oxygenated organics produced from aromatic oxidation contribute to SOA formation (Ng et al. 2007; Hildebrandt et al. 2009; Schwantes et al. 2017) and may lead to new particle formation (Wang et al. 2017; Molteni et al. 2018; Garmash et al. 2020; Wang et al. 2020). Efficient production of peroxyacylnitrates (PANs) from BTX oxidation provides a reservoir of nitrogen oxide radicals ($NO_x \equiv NO + NO_2$), increasing ozone and OH on a global scale (Fischer et al., 2014). Interest in using satellite observations of formaldehyde (CH_2O) and glyoxal ($C_2H_2O_2$) as proxies of VOC emissions has further motivated the need to quantify yields of these species from aromatics (Chance et al. 2000; Wittrock et al. 2006; Liu et al. 2012; Chan Miller et al. 2016). Aromatics are particularly important for glyoxal and methylglyoxal production, providing another avenue for SOA formation (Fu et al. 2008; Lin et al. 2012).

The importance of aromatics for ozone formation spurred the initial development of BTX oxidation mechanisms for air quality models (Carter 1990; Stockwell et al. 1997). The Master Chemical Mechanism (MCM) gives a quasi-explicit representation of BTX atmospheric chemistry with thousands of reactions (Jenkin et al. 2003) but is computationally intractable for 3-D models. A range of simplified mechanisms are presently used in models (Jenkin et al. 2008; Goliff et al. 2013; Carter and Heo 2013; Emmons et al. 2020) but can differ greatly in their results. Two recent studies implementing BTX chemistry into global models found opposite effects on global tropospheric ozone (Yan et al. 2019; Taraborrelli et al. 2021). Differences between mechanisms reflect evolution of knowledge as well as remaining uncertainties and parameterization choices (Schwantes et al. 2017, Xu et al. 2020). Because of the high computational cost of chemical evolution and transport in 3-D models (Nielsen et al. 2017; Hu et al. 2018), it is imperative to minimize the number of species uniquely needed to describe aromatic chemistry (Stockwell et al. 2012; Brown-Steiner et al. 2018; Shen et al. 2020).

Here we present a new compact mechanism for BTX oxidation, GEOS-Chem version 13 (GC13), that is sufficiently simple for use in 3-D models but retains the accuracy of far more complex mechanisms and successfully fits laboratory data for BTX oxidation products. GC13 incorporates new knowledge on phenoxy-phenylperoxy radical cycling (Taraborrelli et al. 2021), later-generation chemistry of hydroxylated aromatics (Schwantes et al. 2017) and fragmentation products (Newland et al. 2019; Wang et al. 2020), and increased radical cycling in the reactions of first-generation aromatic peroxy radicals (Xu et al. 2020). We evaluate GC13 in box model simulations of laboratory chamber experiments and the continental boundary layer and intercompare it to six other mechanisms used in atmospheric models (CRI, MCM, MECCA, MOZART, RACM2, SAPRC). We implement GC13 into GEOS-Chem, a global chemical transport model (CTM), to diagnose the effects of aromatic chemistry in the troposphere on oxygenated organics and oxidant chemistry.

2. GC13: A NEW COMPACT AROMATIC MECHANISM FOR ATMOSPHERIC MODELS

GC13 includes 17 unique species and 44 unique reactions to describe BTX chemistry. These are listed in Section S1 of the Supplemental Information. Figures 1-3 show the dominant routes of BTX oxidation. Initial branching ratios in GC13 are shown in red, and species treated explicitly in blue. Starting from current knowledge on reaction pathways, we reduce the mechanism to be as simple as possible for atmospheric modeling while accurately representing important outcomes, including (a) ozone formation, (b) yields of major first-generation products, (c) short- and long-term yields of formaldehyde, glyoxal, and methylglyoxal, (d) effects on hydrogen oxide ($\text{HO}_x \equiv \text{OH} + \text{peroxy}$) radical budgets, and (e) closure of the total carbon budget. While we do not yet represent SOA formation with GC13, we include intermediates and pathways by which SOA formation is known to occur, facilitating future implementation of an SOA module. We find that separately representing individual xylene isomers provides negligible benefits, and therefore represent them as a single lumped species. We also combine many xylene oxidation products with toluene products, scaling product yields to retain mass balance.

The first steps of BTX oxidation are summarized in Figure 1. OH is the only significant oxidant; lifetimes with respect to oxidation at 298 K and $[\text{OH}] = 1 \times 10^6 \text{ molecules cm}^{-3}$ are 9.6 d for benzene, 2.1 d for toluene, and 16 h for xylene (Mellouki et al. 2020). Reactions of BTX with NO_3 radicals are at least a factor of 10^5 slower than their reactions with OH (Atkinson et al., 1984).

BTX oxidation by OH can proceed via either of two pathways: hydrogen abstraction from an alkyl substituent (**i**, only available to toluene and xylene), or OH addition to the aromatic ring (**ii**). Route **i** leads eventually to benzaldehyde (from toluene) or tolualdehyde (from xylene) (**1a**); as described in Section 4.1, we find that experimental benzaldehyde yields are best fit by skipping the intermediate peroxy radical and proceeding directly to aldehyde formation. Route **ii** can either be followed by H-abstraction (**iii**), leading to a stable hydroxylated compound (**1b**, e.g. phenol), or by O_2 addition (**iv**), leading to the formation of a bridged bicyclic peroxy radical (**1c**). (For a more detailed description of the dynamic system of reversible O_2 addition, see Xu et al. (2020)). While other mechanisms and past studies have suggested the intermediate formation of other products (pathway **v**) preceding **1c**, such as C_6 epoxides, these remain speculative or observed only under high-NO or low-pressure conditions unrepresentative of ambient BTX

oxidation (Yu and Jeffries 1997, Berndt and Boge 2001, Birdsall et al. 2010). Recent experimental and theoretical evidence suggests they do not form under ambient conditions (Wang et al. 2013, Zaytsev et al. 2019, Xu et al. 2020), so we assume that **1c** is the only product from route **iv**.

The bridged bicyclic peroxy radical **1c** would traditionally be treated as a mechanistic branching point, leading potentially to the formation of a hydroperoxide, an organonitrate, an alcohol, or an alkoxy radical depending on the bimolecular reaction partner. However, recent work by Xu et al. (2020) found that nitrate and hydroperoxide yields from **1c** never exceeded 0.06% and 1% respectively for benzene. Instead, they suggest that reaction with either NO or HO₂ leads exclusively to alkoxy radical formation (**1d**, with NO₂ and OH as respective coproducts) and subsequent fragmentation, and that this pattern should hold for larger aromatics as well. These higher-than-expected radical recycling rates are consistent with other recent results and hypotheses (Orlando et al. 2012, Praske et al. 2015, Zaytsev et al. 2019), and help to bring OH radical concentrations in line with values inferred from BTX decay rates in chamber experiments (Bloss et al. 2005a). In our mechanism, this means that we can bypass representation of the individual peroxy radical species and proceed straight to ring-opening products, using a single generic peroxy radical species to accomplish conversion of bimolecular reaction partners (e.g., NO to NO₂).

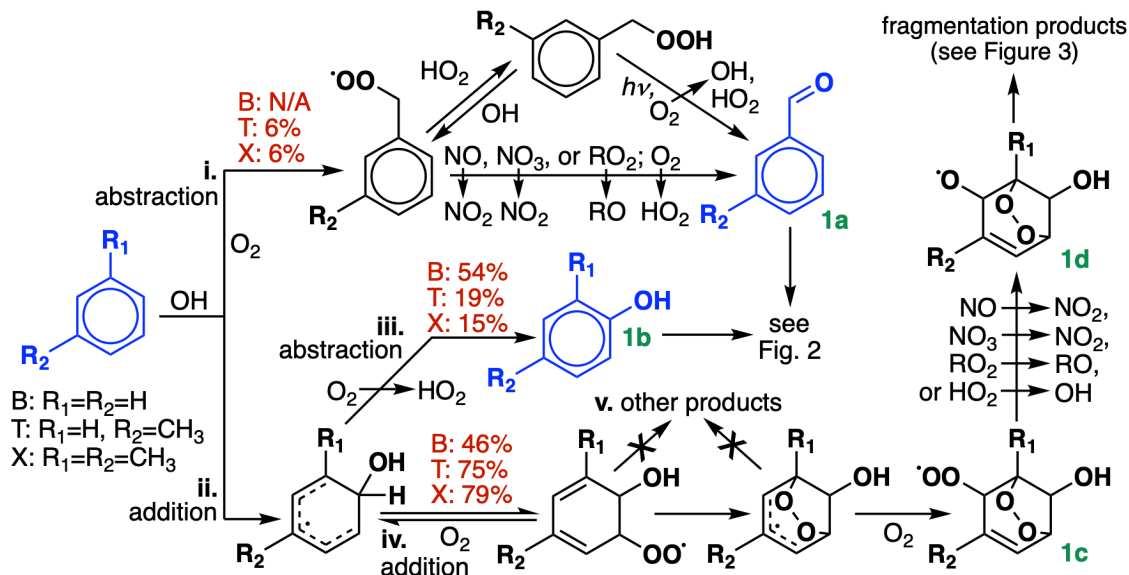


Figure 1. Initial oxidation chemistry of benzene, toluene, and xylenes. R₁ and R₂ denote either H or CH₃ depending on the species. For simplicity, only *m*-xylene is shown; in GC13, the xylene isomers are lumped together. Initial branching ratios in GC13 are shown in red, and species treated explicitly and uniquely in the aromatic oxidation mechanism are highlighted in blue.

The secondary chemistry of the ring-retaining aldehyde (**1a**) and alcohol (**1b**) products is shown in Figure 2. We lump benzaldehyde with tolualdehyde, given their similar chemistries. Benzaldehyde oxidation by either OH or NO₃ generates a peroxyacyl radical (**2a**) that can react with NO₂ to form peroxybenzoyl nitrate (**2b**), a PAN-like species we represent explicitly for its ability to sequester and transport NO_x. Reaction of **2a** with HO₂ yields perbenzoic acid (**2c**), benzoic acid and ozone, or a benzoyloxy radical and OH, in a 65:15:20 ratio (Roth et al. 2010);

subsequent chemistry leads eventually to formation of the phenylperoxy radical (**2d**). Benzaldehyde photolysis is also implemented with recently reevaluated rates (Mellouki et al., 2020), producing OH, CO₂, and the phenylperoxy radical (**2d**).

Rates and branching ratios for phenol and cresol oxidation by OH and NO₃ are taken directly from MCM. We lump xylenols with cresols, given their similar reactivity, and lump all products of cresol/xylenol oxidation with those of phenol. Both OH and NO₃ can abstract hydrogens, leading either to bridged peroxy radicals (**1c**) or to phenoxy radicals (**2e**), or undergo addition. NO₃ addition leads to ring-retaining organonitrates, which we lump with nitrophenols (**2f**), while OH addition forms catechols (**2g**). While other mechanisms including MCM typically assume that the reaction of catechols with OH proceeds primarily by abstraction from the hydroxyl group, leading to functionalized phenoxy radicals, Schwantes et al. (2017) found that addition pathways dominate, leading to a proliferation of oxidized species that may act as SOA precursors. We cannot represent these species in detail in a condensed mechanism; future work will focus on parameterizing SOA yields from these pathways. Instead, we adopt branching ratios from Schwantes et al. (2017), lump polyhydroxylated aromatics with cresols and other oxidized products with analogs, and adjust the cresol + OH reaction rate such that it represents only the abstraction pathways. H-abstraction from cresols would make functionalized phenoxy radicals, which we combine with the phenoxy radical in GC13. We also include the reactions of catechols with NO₃ and O₃, with rates from MCM and products lumped in with those from direct phenol and cresol oxidation.

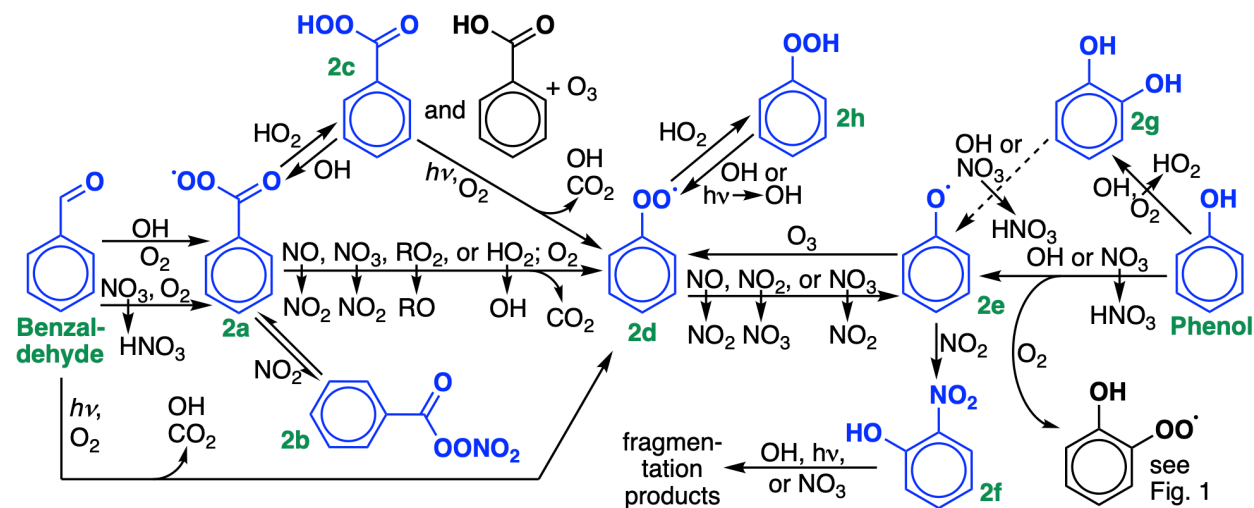


Figure 2. Oxidation chemistry of stable ring-retaining aldehydes and alcohols from benzene, toluene, and xylene. For simplicity, only benzaldehyde and phenol are shown; cresol chemistry is not shown, but is similar to phenol. GC13 lumps tolualdehyde with benzaldehyde, xylenols with cresol, and methylcatechols with catechol. Species treated explicitly and uniquely in the GC13 mechanism are highlighted in blue. The reaction of catechols (**2g**) with OH produces substituted phenoxy radicals, lumped here with the phenoxy radical (**2e**).

Both benzaldehyde and phenol/cresol oxidation lead eventually to the formation of phenylperoxy and/or phenoxy radicals (**2d** & **2e**), the chemistry of which can have important implications for radical and ozone budgets via their cycling and formation of nitrophenols (**2f**; see Section 4.3). Due to its aromatic stability, the phenoxy radical is far more stable than a

typical alkoxy radical, and rather than fragmenting, tends to react either with NO₂ (forming nitrophenols) or ozone (forming the phenylperoxy radical). The phenoxy and phenylperoxy radicals thus cycle until interrupted by phenyl hydroperoxide (**2h**) formation from phenylperoxy + HO₂, which is only a temporary radical reservoir, or by nitrophenol formation from phenoxy + NO₂, which provides a terminal sink and the only route to fragmentation products from this radical system. Here, we treat the phenoxy-phenylperoxy system similarly to their representation in recent work by Taraborrelli et al. (2021), including the phenylperoxy + NO₂ reaction (Jagiella and Zabel 2007), the phenoxy + O₃ reaction (Tao and Li, 1999), explicit treatment of phenyl hydroperoxide and nitrophenols (lumped with nitrocatechols), and HONO formation from nitrophenol photolysis (Bejan et al. 2006, Chen et al. 2011).

Finally, Figure 3 shows the chemical pathways of the bridged alkoxy radical **1d** produced following OH addition to BTX (Figure 1). These pathways remain uncertain, particularly for toluene and xylene, but represent the main source of glyoxal and methylglyoxal from BTX oxidation. A theoretical study by Wang et al. (2013) suggested ring-closure (**vi**) as the dominant fate of the bicyclic alkoxy radical from benzene, forming an epoxide (**3a**). However, follow-up calculations suggested instead that ring-breaking (**vii**) dominates (98%+; Wang et al., 2020), resulting in an acyclic alkoxy radical (**3b**), which was confirmed experimentally by Xu et al. (2020). The acyclic alkoxy radical can then decompose directly, forming a conjugated γ -dicarbonyl (**3c**) and glyoxal or methylglyoxal. Conjugated γ -dicarbonyls are highly photolabile, with typical daytime lifetimes of 10-15 minutes (Newland et al., 2019); however, their photolysis does not result in substantial HO_x production, instead leading primarily to ketene-enol formation (**3e**), with a minor route decomposing to CO and an acrolein derivative (Newland et al., 2019).

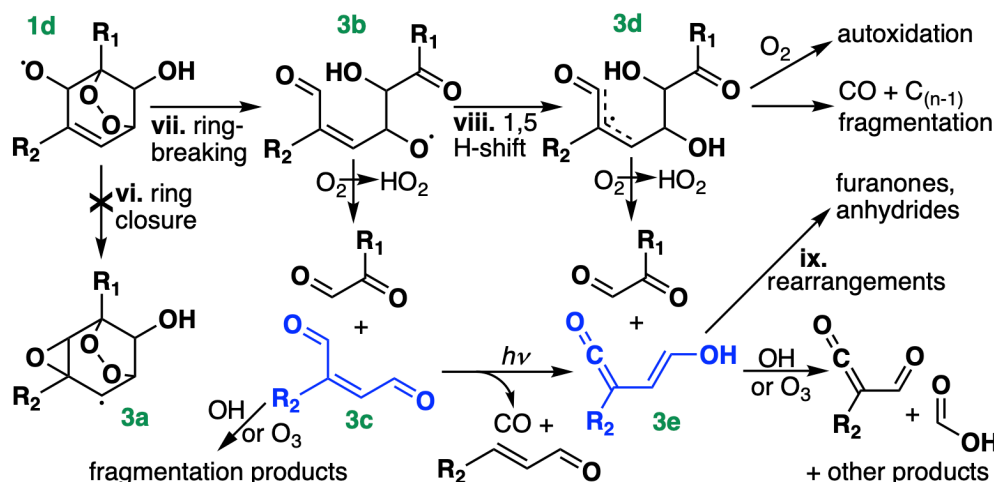


Figure 3. Chemical pathways following OH addition to BTX, continued from Figure 1. R₁ and R₂ denote either H or CH₃ depending on the species. For toluene and xylene, only OH addition *ortho*- to R₁ is shown, and for xylene, only the *meta* isomer is shown; isomeric differences could lead to differing placement of R₁ and R₂ on the stable products, including formation of acetic rather than formic acid (bottom right). Compounds in blue (**3c** and **3e**) are treated as lumped stable intermediates in GC13.

Recent theoretical (Wang et al., 2020) and experimental (Xu et al., 2020) studies have shown that instead of directly decomposing, the acyclic alkoxy radical (**3b**) can undergo a 1,5

aldehydic H-shift (**viii**), yielding a resonance-stabilized allyl radical (**3d**). This allyl radical represents a mechanistic branching point with many possible fates, dominated by decomposition to an α -dicarbonyl and a ketene-enol (**3e**). Between this direct production and secondary formation from γ -dicarbonyl photolysis, Wang et al. (2020) and Xu et al. (2020) suggest that ketene-enols are the primary product of the ring-opening pathway; however, Xu et al. (2020) note that they do not observe unity yields from benzene, and suggest that other reactive pathways from **3d** may account for the missing carbon. Such pathways include CO elimination and O₂ addition, the latter of which could lead to autoxidation and the formation of highly oxidized SOA precursors from aromatics (Wang et al. 2017, Molteni et al. 2018, Garmash et al. 2020).

Subsequent chemistry of the ketene-enols (**3e**) was studied by Newland et al. (2019) and Wang et al. (2020), and includes unimolecular cyclization to furanones as well as direct formation of formic and acetic acids via reaction with OH and O₃. Newland et al. (2019) also observed tautomerization of the ketene-enols to ketene-carbonyls, possibly mediated by chamber surfaces, followed by cyclization to form anhydrides. Further studies are needed to constrain the contributions of these pathways, particularly for ring-opening products from toluene and xylene.

Because the branching ratios of the ring-opening pathways in Figure 3 remain poorly constrained, particularly for toluene and xylene, we do not speciate them in detail in GC13. Instead, we tune the yields of glyoxal and methylglyoxal (coproduced with **3c** and **3e**) to match observed yields from chamber experiments, and lump larger products into two representative intermediates, which stand in for the C₄₊ stable products in Figure 3 (primarily **3c** and **3e**). One, treated as a C₄ compound and produced from all three BTX precursors, yields glyoxal and other compounds lacking methyl groups upon its subsequent oxidation, while the other, treated as a C₅ compound and produced only from toluene and xylene, also yields methylglyoxal and other methylated products. The subsequent chemistry of these representative C₄-C₅ intermediates is a weighted combination of the reactive fates of the conjugated dicarbonyls (**3c**) and the ketene-enols (**3e**), drawn from Newland et al. (2019) and MCM for **3c** and from Newland et al. (2019) and Wang et al. (2020) for **3e**, adjusted slightly to tune the later-generation yields of glyoxal and methylglyoxal. We do not explicitly track the furanones and anhydrides produced in unimolecular rearrangements of the ketene-enols (**ix**), allocating this carbon instead to other oxygenated intermediates with similar functionalities and lifetimes while maintaining carbon balance. We include a route to direct CO formation from **3d**, using a branching ratio from Xu et al. (2020) for benzene, but do not include the minor pathway to highly oxidized molecule (HOM) formation; subsequent updates to the mechanism focused on SOA formation could represent these pathways explicitly.

3. PREVIOUS AROMATIC MECHANISMS

In the following sections, we will compare GC13 to a suite of commonly used aromatic mechanisms. Here, we briefly describe each mechanism, in order from most to least complex. The number of species and reactions in each mechanism, excluding inorganic species and generic C₁-C₃ compounds, are given in Table 1. The mechanisms span a wide range of complexity, with two orders of magnitude separating the numbers of species and reactions in the largest and smallest mechanisms.

Table 1. Aromatic (BTX) chemistry mechanisms ^a

Mechanism	Number of:		Reference
	Species	Reactions	
MCM v3.1	1271	3788	Bloss et al. (2005b)
MECCA	229	666	Taraborrelli et al. (2021)
SAPRC-11	55	374	Carter and Heo (2013)
CRI v2-R5	56	128	Watson et al. (2008)
RACM2	34	115	Goliff et al. (2013)
MOZART-T1	33	56	Emmons et al. (2020)
MOZART-GC	13	43	Porter et al. (2017)
GC13	17	44	This work

^a Mechanisms used in intercomparison to GC13. Species count includes both stable and radical species, but does not include inorganic reactants and common/generic C₁-C₄ species. Reaction count does not include reactions that only contain these excluded species.

MCM. The Master Chemical Mechanism (MCM) is a near-explicit mechanism that treats the full oxidative degradation of benzene, toluene, and each xylene isomer. The mechanism was developed by Jenkin et al. (2003) and updated to version 3.1 in Bloss et al. (2005b) based on an assessment of experimental work. Bloss et al. (2005a, 2005b) compared the mechanism to a series of chamber experiments and noted that, while ozone was well-simulated in benzene oxidation, simulations typically overestimated ozone formation from the larger aromatics while underestimating OH concentrations in all experiments.

MECCA. The Module Efficiently Calculating the Chemistry of the Atmosphere (MECCA; Sander et al. 2011) includes detailed aromatic chemistry described by Cabrera-Perez et al. (2016). The chemistry of toluene and benzene is taken from MCM without simplification, but with some important updates (notably to phenoxy-phenylperoxy radical cycling) described in detail by Cabrera-Perez et al. (2016) and Taraborrelli et al. (2021). MECCA lumps the xylene isomers and sets their chemistry identical to that of toluene, aside from the branching ratios in their initial reaction with OH. Despite its large number of species and reactions, MECCA has been used in a global 3-D model study by Taraborrelli et al. (2021).

SAPRC. The Statewide Air Pollution Research Center (SAPRC) mechanisms are a family of moderately reduced mechanisms widely used in airshed models for the prediction of ozone formation and representation of organic pollutants. Here we use SAPRC-11 (Carter and Heo, 2013), which was specifically designed to optimize aromatic oxidation and ozone formation rates in comparison to environmental chamber experiments. SAPRC-11 represents the xylene isomers individually and includes many early-generation BTX oxidation products, including phenol, cresol, xylenols, catechol, dicarbonyl compounds, nitrophenols, and benzaldehyde. SAPRC-11 was previously implemented in GEOS-Chem (v9-02) by Yan et al. (2019). Versions of SAPRC have also been implemented in the Weather Research and Forecasting chemistry model (WRF-Chem), the Community Multiscale Air Quality (CMAQ) model, and other regional 3D models (Yu et al. 2010, Cai et al. 2011, Zhang et al. 2012, Kitayama et al. 2019, Shareef et al. 2019).

CRI. The Common Representative Intermediates (CRI) mechanism, first developed by Jenkin et al. (2002), is a reduced mechanism based on MCM that uses a series of generalized

intermediates, rather than explicitly simulating multigenerational chemistry, to simulate ozone formation rates. Optimization and comparisons to MCM for the most recent version, v2, are described by Jenkin et al. (2008). Here, we use CRI v2-R5, a further reduction that lumps together the xylene isomers (Watson et al. 2008). CRI v2-R5 has been widely implemented in regional and global models, including WRF-Chem (Archer-Nicholls et al. 2014), STOCHEM (Utembe et al. 2010, Khan et al. 2015), and UKCA (Archer-Nicholls et al. 2020).

RACM. The Regional Atmospheric Chemistry Mechanism (RACM) is a reduced mechanism first presented by Stockwell et al. (1997) and intended for use in modeling gas-phase chemistry in a wide range of ambient conditions. Here we use RACM2 as described by Goliff et al. (2013). RACM2 is based on MCM for benzene and on Calvert et al. (2002) for toluene and the individual xylene isomers, and represents major first-generation products of BTX oxidation, including benzaldehyde, phenol, cresol, an epoxide, and photolabile dicarbonyls. RACM is implemented in both WRF-Chem and CMAQ (Sarwar et al. 2013, Kitayama et al. 2019, Shareef et al. 2019).

MOZART. The Model for Ozone and Related chemical Tracers (MOZART) chemical mechanisms are designed for implementation in the NCAR Community Earth System Model (CESM). The most recent iteration, MOZART-T1 (Emmons et al. 2020), is the first to differentiate between the BTX compounds. With an oxidation scheme based on MCM, it represents some major early-generation products explicitly, including benzaldehyde, photolabile dicarbonyls, and peroxybenzoyl nitrate, but ignores some later-generation products assumed to undergo efficient deposition or aerosol uptake. An earlier version of the aromatic chemistry in MOZART-T1, first described by Knote et al. (2014), was implemented into GEOS-Chem by Porter et al. (2017); for the sake of comparison and contextualization of past studies, we also implement the Porter et al. (2017) mechanism here, which we label “MOZART-GC”.

4. MECHANISM EVALUATION AND INTERCOMPARISON

4.1. METHODS

We implemented GC13 and the other mechanisms of Section 3 into box model simulations for comparisons to environmental chamber data and for mechanism intercomparisons under a range of boundary layer conditions. The simulations use a 4th-order Rosenbrock kinetic solver implemented with the Kinetic PreProcessor tool (KPP; Damian et al. 2002, Daescu et al 2003, Sandu et al 2003). We standardize the inorganic and C₁-C₃ chemistry to that of MCM in all mechanisms so that the only differences between the mechanisms are in their BTX oxidation chemistry. For mechanisms with speciated xylenes, we assume equal contributions from the three isomers, and comparisons of xylene product yields with literature values are only conducted for experimental studies that targeted all three isomers.

Environmental chamber simulations. To quantify product yields and compare them to environmental chamber data, we simulate BTX chemistry in a box model representative of chamber experimental conditions. For each mechanism, we initialize the box model with fixed mixing ratios of one aromatic precursor, NO, and H₂O₂ as a photolytic OH source, then run the simulation forward in time with a fixed light intensity and temperature until the aromatic

precursor is 99% depleted. We then vary all initial settings (temperature, light intensity, and each reactant concentration) individually and rerun the simulation to sample the full range of possible experimental conditions. Results shown in Sections 4.2-4.3 are for simulations with initial $[\text{NO}]_0 = 5 \text{ ppt} - 2.5 \text{ ppm}$, $P = 1 \text{ atm}$, $T = 298 \text{ K}$, $[\text{VOC}]_0 = 100 \text{ ppb}$, $[\text{H}_2\text{O}_2]_0 = 2.5 \text{ ppm}$, and an NO_2 photolysis rate (j_{NO_2}) of $8 \times 10^{-3} \text{ s}^{-1}$ (other photolysis rates, given in Section S1 of the SI, are scaled to j_{NO_2}). “Initial” yields shown in Figures 4-6 are after 10 min of photooxidation; additional results showing long-term yields after 24 h of oxidation are provided in Section S2 of the SI. Sensitivities to temperature and initial VOC concentrations are generally small, but additional results showing the effects of these parameters, as well as the effects of light intensity and oxidant source, can be found in the SI. Wall losses of semivolatile gases are not represented in these simulations, as we do not seek to model SOA formation or the yields of direct SOA precursors. An improved representation of SOA from aromatics in GEOS-Chem will be the subject of future work. While it is possible that wall losses of semivolatile intermediates affect experimental yields of OVOCs, this multigenerational chemistry is expected to play only a minor role at in the short-term chemistry described time scales isolated here.

Continental boundary layer simulations. To examine longer-term product yields and effects of BTX oxidation on the ambient atmosphere, the same box model described above is also run under conditions meant to simulate a continental boundary layer like that of the heavily studied Seoul Metropolitan Area (Oak et al. 2019, Schroeder et al. 2020) with constant NO and aromatic emissions. The well-mixed boundary layer exchanges with the background free troposphere with a fixed ventilation timescale of 1 d for all species. Simulations are initialized with 75 ppb O_3 , 1.8 ppm CH_4 , 200 ppb CO , 300 ppt CH_2O , and 1% H_2O , and these species are also present in the same concentrations in the background free troposphere with which the boundary layer box exchanges. Photolysis rates follow a clear-sky diurnal profile at 45° latitude at the summer solstice with an ozone column of 350 DU, while temperature varies sinusoidally with an amplitude of 4°C , centered at 25°C , peaking at 13:00 solar time, and a period of 1 d. Results shown in Sections 4.2-4.3 are for a total aromatic VOC emission rate of 120 ppt h^{-1} , distributed between benzene, toluene, and xylene in a 2:2:1 molar ratio or for a single precursor, and for fixed NO emission rates between 1 ppt h^{-1} and 10 ppb h^{-1} . Additional results showing sensitivities to VOC emission rates can be found in the SI. The model does not represent deposition or aerosol uptake processes, except to impose a 1 h loss rate on N_2O_5 for conversion to HNO_3 . We apply 7 days of initialization to reach diurnal steady state; the results shown in Section 4.3 are from the 8th simulated day.

4.2. OXYGENATED VOC YIELDS

We use environmental chamber simulations to determine prompt product yields from each mechanism and compare them to experimental data, and use both chamber and boundary layer simulations to investigate the differences in long-term yields of later- and multi-generational products between mechanisms. In both cases we present our results as a function of initial NO_x concentrations mixing ratio ($[\text{NO}_x]_0$) imposed in the model either directly (chamber simulations) or through emissions (boundary layer simulations). In some cases, yields from organic peroxy radical reactions may depend on the branching between bimolecular and unimolecular reactions in addition to the branching between reaction with NO and other reactions, and comparisons would be more appropriately made as a function of peroxy radical

lifetime against all bimolecular reactions rather than just NO (see, e.g., Xu et al. 2020); however, it is difficult to approximate these bimolecular lifetime conditions for many past experimental results, so we opt for the more straightforward comparisons as a function of initial NO_x .

In comparisons with experimental yields (Figures 4-6), simulated yields are shown after 20 min of oxidation, while experimental yields are the earliest reported value. **Because the experimental yields shown here come from a range of chamber studies with differing setups (e.g., chamber size, oxidant source, light spectrum and intensity, duration, initial concentrations, etc.), care should be taken when comparing observed yields as a function of $[\text{NO}_x]_0$ to each other or to simulated yields.** Experiments conducted with no added NO_x are shown at $[\text{NO}_x]_0 = 10$ ppt₅ (below which modeled yields are invariant with $[\text{NO}_x]_0$) **regardless of whether NO_x concentrations were monitored during the experiment. In some cases, NO_x may off-gas from chamber walls (Carter et al., 2005). When reported, instrumental uncertainties on measured NO_x mixing ratios are shown as horizontal error bars on experimental points.**

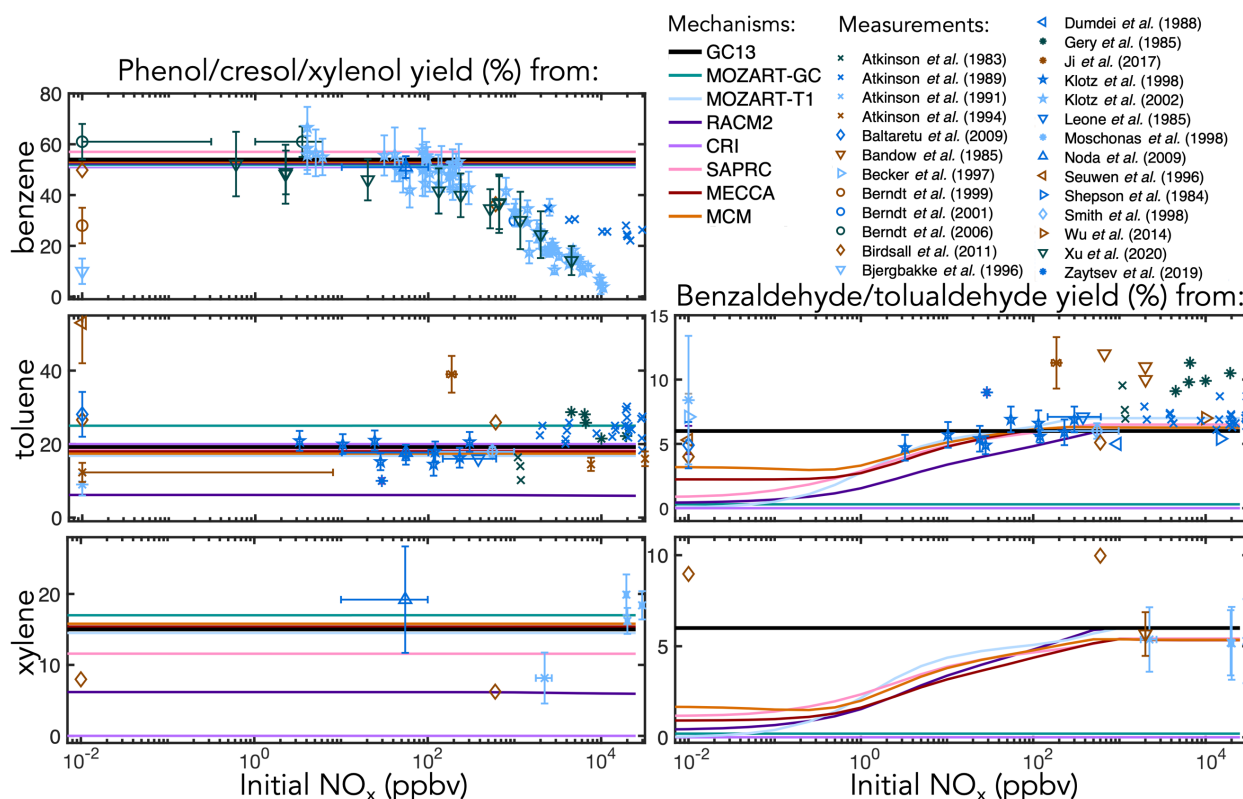


Figure 4: Experimental (points) and simulated (lines) prompt molar yields of aromatic alcohols and aldehydes from BTX oxidation in environmental chambers as a function of initial NO_x mixing ratio. Overlapping simulation lines are offset slightly for visibility.

Ring-retaining products. Figure 4 shows environmental chamber molar yields of ring-retaining alcohols and aldehydes from BTX oxidation in the mechanisms and in the experimental literature as functions of initial NO_x **mixing ratio concentrations**. For the ring-retaining alcohols formed via OH addition followed by H-abstraction (Figure 1iii), all mechanisms implement fixed direct yields from the BTX precursors, which therefore do not vary with NO. GC13 uses a phenol yield from benzene of 54%, derived from an error-weighted average of the yields

measured by Klotz et al. 2002 ($53 \pm 7\%$), Berndt and Boge 2006 ($61 \pm 7\%$), and Xu et al. 2020 ($49 \pm 13\%$) at atmospherically relevant NO_x levels. Among the other mechanisms implemented here, SAPRC uses a fixed phenol yield of 57% from benzene, while all others use 53%. Fixed cresol and xylenol yields of 19% and 15% in GC13, from toluene and xylene respectively, are taken from MECCA and are consistent with observations.

For phenol, the mechanisms' fixed yields of 53-57% fit most data under low-to-moderate NO conditions, but **when at high $[\text{NO}_x]_0$ exceeds 100 ppb** the observed yield declines as bimolecular reactions become too fast to allow equilibration to occur. This behavior is not exhibited by any of the mechanisms studied here; however, such high NO concentrations are rarely seen in ambient environments, and the fixed yields are therefore suitable for atmospheric simulations. Experimental yields of cresol (from toluene) and xylenol (from xylene) exhibit little correlation with NO_x , and are therefore adequately represented by fixed branching ratios, although the ratios implemented vary widely between mechanisms, with low yields in RACM2 (6% from toluene and xylene) and no representation of xylenols in CRI v2-R5.

The ring-retaining aldehydes benzaldehyde and tolualdehyde are formed following H-abstraction from the methyl groups of toluene and xylene respectively (Figure 1i). Most mechanisms that include benzaldehyde chemistry (CRI and MOZART-GC do not) explicitly treat the benzyl peroxy intermediate in this reaction pathway, resulting in peak aldehyde yields of $\sim 6\%$ **when at high $[\text{NO}_x]_0$ exceeds 100 ppb**, dropping to 1-3% at **low- $[\text{NO}_x]_0 < 1$ ppb** due to competing formation of benzyl hydroperoxide. However, this results in lower benzaldehyde yields than observed under **low- NO_x -free** conditions. The lack of observed NO_x -dependence in the benzaldehyde yield may reflect a short lifetime of benzyl hydroperoxide, a high incidence of $\text{RO}_2 + \text{RO}_2$ reactions (Moschonas et al. 1998), a non-hydroperoxide-forming channel in the benzyl peroxy + HO_2 reaction (Baltaretu et al. 2009), or another pathway altogether (Salta et al. 2020). Regardless, to fit observations and further simplify the mechanism, we bypass the benzyl peroxy intermediate and form benzaldehyde directly from the reaction of toluene and xylene with OH with a fixed yield of 6%, consistent with most experimental results at both high and low NO_x .

C1-C3 carbonyl products. Figure 5 shows prompt environmental chamber yields of glyoxal from BTX oxidation in both the mechanisms and the experimental literature as functions of initial NO_x **mixing ratio concentrations** (initialized as NO in simulations and in most chamber experiments, with some NO_2 in chamber experiments). Initial glyoxal yields from BTX oxidation generally range between 10% and 40%, with the highest yields from benzene and the lowest from xylenes. With a fixed NO_x -independent first-generation glyoxal yield and secondary yields from the representative C_4 - C_5 intermediates (Figure 3), GC13 accurately simulates experimentally measured glyoxal yields across a wide range of $[\text{NO}_x]_0$ for all three aromatic precursors, while most other mechanisms exhibit excessive glyoxal formation **at under-high- $[\text{NO}_x]_0 > 100$ ppb conditions** and insufficient production **at under-low- $[\text{NO}_x]_0 < 1$ ppb conditions**. Due to the long lifetime of benzene relative to its intermediate oxidation products, it is difficult to isolate prompt vs. multigenerational glyoxal yields, which **partially** explains the range of experimental results; the prompt yield we implement in GC13 (18%) is able to match the lower limit of observed yields at both high and low NO_x . **We also find in GC13 simulations that the**

glyoxal yield from benzene is sensitive to the initial benzene concentration in chamber experiments (see Section S4 of the SI), which may further explain the spread of measured yields.

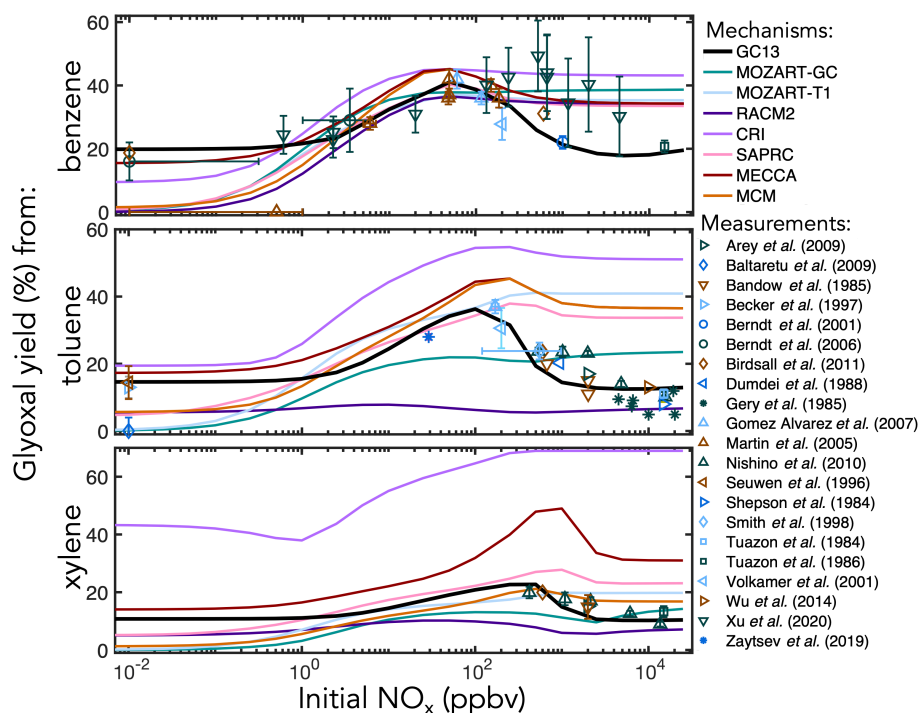


Figure 5: Experimental (points) and simulated (lines) prompt molar yields of glyoxal from BTX oxidation in environmental chambers as a function of initial NO_x mixing ratio.

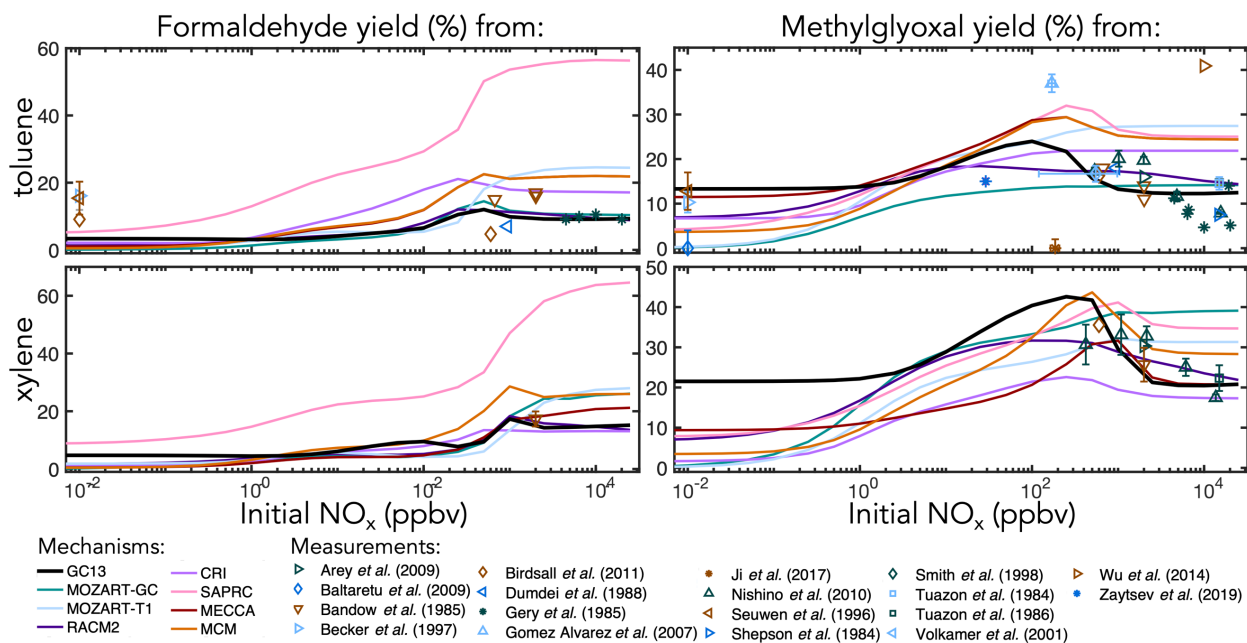


Figure 6: Experimental (points) and simulated (lines) prompt molar yields of methylglyoxal and formaldehyde from toluene and xylene oxidation in environmental chambers as a function of initial NO_x mixing ratio.

Figure 6 shows simulated and observed environmental chamber yields of formaldehyde and methylglyoxal from toluene and xylene oxidation as functions of initial NO_x mixing ratio concentrations. Because it lacks alkyl substituents, benzene is unable to produce either formaldehyde or methylglyoxal. As with glyoxal, GC13 is consistent with experimental yields across the full spectrum of $[\text{NO}_x]_0$ conditions, although some observed methylglyoxal yields from toluene deviate considerably from the general trend. Most other mechanisms exhibit lower prompt formaldehyde and methylglyoxal yields at under-low- $[\text{NO}_x]_0 < 1$ ppb conditions and higher yields at under-high- $[\text{NO}_x]_0 > 100$ ppb conditions than GC13 or observed values, although observational evidence is sparse for xylene. In particular, SAPRC and MCM exhibit high prompt formaldehyde and methylglyoxal yields, respectively, at under-high- $[\text{NO}_x]_0 > 100$ ppb conditions.

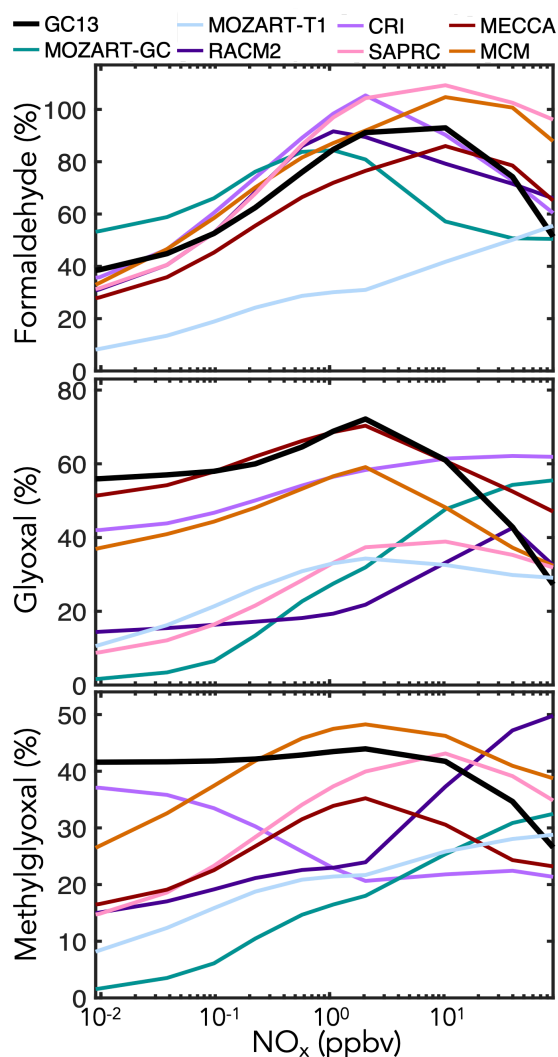


Figure 7: Simulated molar yields of formaldehyde, glyoxal, and methylglyoxal from BTX oxidation in a continental boundary layer as a function of midday NO_x mixing ratio. Model setup is described in Section 4.1. Molar yields are averaged over the 8th simulated day.

In addition to these prompt yields, later-generation formation of C_1 - C_3 carbonyl species can be important for HO_x radical and carbon budgets. Data on long-term yields are sparse, so here we rely mostly on model intercomparisons of boundary layer simulations with MCM taken

as a reference. Results are shown in Figure 7 for mixed aromatic emissions, and in Figure S6 in the Supporting Information for individual aromatic precursors. The maxima in yields at intermediate NO_x , particularly evident for formaldehyde, reflect the corresponding maxima of OH concentrations. Generally, MCM is able to produce higher late-generation yields of the C_1 - C_3 carbonyls than more reduced mechanisms, but GC13 exhibits high yields similar to MCM. GC13 also simulates similar midday glyoxal-to-formaldehyde concentration ratios from aromatic oxidation to MCM, ranging from 0.3 ~~when under high- NO_x conditions exceeds 1 ppb~~ to > 0.7 at ~~low~~ $\text{NO}_x < 10 \text{ ppt}$ (Figure S8). Results speciated by BTX precursor reveal a similar trend. MOZART-GC and MOZART-T1, the most reduced of previous mechanisms, tend to simulate the lowest long-term carbonyl yields. MECCA is generally able to reproduce MCM's high long-term carbonyl yields except from xylene, which MECCA lumps as toluene; this results in an overprediction of glyoxal yields and underprediction of formaldehyde and methylglyoxal yields relative to MCM. Long-term yields from simulated chamber experiments (Figures S3-S4) show results similar to the boundary layer simulations, and long-term formaldehyde yields exhibit a temperature sensitivity of up to $+3\% \text{ K}^{-1}$.

Other VOC products. Some experimental and theoretical evidence exists for formation pathways of fumaraldehydic, formic, and acetic acids from BTX oxidation (Berndt et al. 1999, Dumdei et al. 1988, Wang et al. 2020, Xu et al. 2020), but this is generally not included in mechanisms. Global models tend to underestimate ambient concentrations of these compounds, and additional formation pathways could help alleviate this discrepancy (Millet et al. 2015, Khan et al. 2018). Here, we include the formation of formic and acetic acids as described in Newland et al. (2019) and Wang et al. (2020) via the ketene-enols, represented as part of the lumped C_4 and C_5 products. Few chamber data are available for comparison; Berndt et al. (1999) measured a 13% yield of formic acid from benzene under ~~nominally low- NO_x -free~~ conditions, while Dumdei et al. (1988) measured a 6% yield of acetic acid from toluene under high- NO (940 ppb) conditions. Both are consistent with prompt yields in GC13. Long-term yields of formic acid in the mechanism can reach 32% from benzene, 28% from toluene, and 16% from xylene, while acetic acid yields can reach 12-13% from toluene and xylene (See Figures S2 and S5).

Most mechanisms also represent the formation of larger (C_{4+}) dicarbonyls, including biacetyl, photolabile conjugated dialdehydes, and less reactive ketones, with varying degrees of complexity. In GC13, these compounds are not treated explicitly, and are instead grouped into the two representative C_4 - C_5 reactive intermediates. For this reason, we do not optimize any branching ratios in GC13 with environmental chamber yields of C_{4+} dicarbonyls, but Figure S1 in the Supporting Information provides a comparison between measured and simulated yields of photolabile dicarbonyls in each mechanism.

Finally, as described in Section 2, most mechanisms include the direct formation of a ring-opened C_{6+} epoxide from BTX oxidation. These yields are NO_x -independent and span a wide range, from 0% from all precursors in GC13 to $>70\%$ from xylene in RACM. While little evidence for this pathway exists under atmospherically relevant conditions, it may be useful as an intermediate, particularly given that other known pathways cannot achieve carbon closure in many experimental studies (e.g. Xu et al. 2020). Although these epoxides are not included in GC13, Figure S1 provides a comparison between measured yields in environmental chambers and simulated yields in other mechanisms.

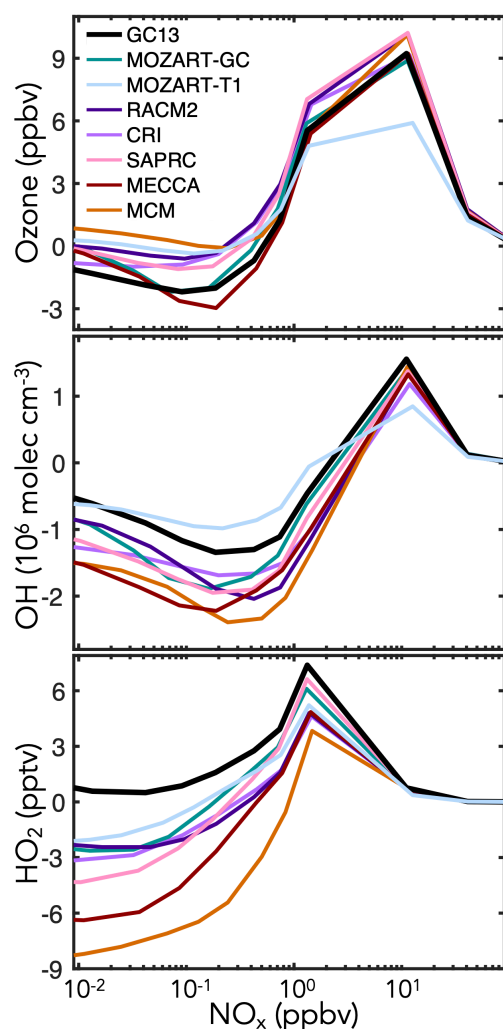


Figure 8: Changes in midday ozone, OH, and HO₂ concentrations due to aromatic chemistry in box model simulations of the continental boundary layer with different aromatic oxidation mechanisms. Changes in species concentrations are calculated by subtracting their midday (10:00-14:00) mean values in a simulation without aromatic emissions from those in an equivalent simulation with aromatic emissions, and are plotted against midday NO_x mixing ratio. Aromatic emissions are 120 ppt h⁻¹ with 2:2:1 molar ratios for benzene:toluene:xylenes. Other model setup is described in Section 4.1.

4.3. EFFECTS ON OXIDANTS

Figure 8 shows the effects of BTX oxidation on daytime abundances of ozone, OH, and HO₂ in the continental boundary layer simulations. Results speciated by aromatic precursor are shown in Figure S7. GC13 and MECCA have less ozone production than MCM, RACM, and SAPRC, all of which are known to overestimate peak ozone from chamber experiments (Bloss et al. 2005a, Goliff et al. 2013, Carter et al. 2013). This is primarily due to phenoxy-phenylperoxy radical cycling chemistry, represented only in GC13 and MECCA (Sander et al. 2019, Taraborrelli et al. 2021). The dominant phenoxy-phenylperoxy cycle converts NO₂ to NO₃ and consumes one ozone molecule; during the day, this cycle largely balances via NO₃ photolysis (although the minor channel to NO + O₂ represents an odd oxygen sink), but at night, conversion

of NO_3 to N_2O_5 and on to HNO_3 amplifies ozone loss. MECCA and GC13 both exhibit a surge of NO_3 (and corresponding reduction in NO_2) at sunset, which propagates to lower NO_2 and ozone levels throughout the following day (Figure S32). Flux through this phenoxy-phenylperoxy system is highest from benzene because of the high phenol yields (Figure 1), so the ozone differences between mechanisms are strongest for benzene (see Figure S7). **The reduced ozone formation in GC13 and MECCA relative to other mechanisms may improve model biases relative to chamber experiments (Bloss et al. 2005a, Goliff et al. 2013, Carter et al. 2013), and would likely reduce the high simulated contribution of aromatics to ambient ozone formation in box model analyses of polluted environments (Luecken et al., 2018; Oak et al., 2019; Schroeder et al., 2020).**

GC13 also simulates higher HO_x concentrations than other mechanisms, especially under low- NO_x (< 1 ppb) conditions. This effect is due primarily to increased radical propagation from the bimolecular reactions of the bridged bicyclic peroxy radicals (**1c**), which do not form radical-terminating hydroperoxides or organonitrates in GC13. Higher radical recycling from subsequent reactions of the representative C_4 and C_5 intermediates also contributes. It has been reported that other mechanisms tend to underpredict HO_x concentrations in simulations of chamber experiments (Chen, 2008), and both Bloss et al. (2005a) and Carter et al. (2013) comment on the need to increase HO_x recycling from aromatic oxidation; GC13 thus brings HO_x concentrations into improved alignment with chamber results. These effects are stronger for toluene and xylene because of their higher bridged bicyclic peroxy radical and C_4 - C_5 intermediate yields relative to benzene (see Figure S7).

To test the sensitivities of these outcomes to specific aspects of the GC13 mechanism, we conduct additional simulations with individually perturbed reaction rate constants and yields. Results from the sensitivity simulations with the most prominent changes are shown in Figure 9. We find that ozone is most sensitive to changes in the rates of the key reactions in the phenoxy-phenylperoxy system. These rates remain uncertain; Tao and Li (1999) measured a rate **constant** of $2.86 (\pm 0.35) \times 10^{-13} \text{ cm}^3 \text{ molecule}^{-1} \text{ s}^{-1}$ at 298 K for the $\text{C}_6\text{H}_5\text{O} + \text{O}_3$ reaction, which we use in GC13, but noted that it might be a lower limit, while Jagiella and Zabel (2007) did not set uncertainty bounds on their best fit rate constant of $7 \times 10^{-12} \text{ cm}^3 \text{ molecule}^{-1} \text{ s}^{-1}$ for $\text{C}_6\text{H}_5\text{O}_2 + \text{NO}_2$ (used in GC13), instead only specifying a minimum of $1 \times 10^{-12} \text{ cm}^3 \text{ molecule}^{-1} \text{ s}^{-1}$ consistent with their results. Increasing either rate by a factor of ten substantially increases ozone loss due to phenoxy-phenylperoxy cycling, highlighting the importance of better constraints on these rates.

Ozone and OH are both sensitive to changes in the fates of catechols and methylcatechols. While most mechanisms assume that the reactions of catechols and methylcatechols with OH proceed by abstraction to form functionalized phenoxy radicals, Schwantes et al. (2017) showed that addition pathways dominate, leading to heavily substituted low-volatility products that may contribute to SOA formation. In a sensitivity simulation with the product channels from catechols + OH turned off (representing complete loss of products to aerosols), the effects of aromatic oxidation on ozone production and OH are strongly diminished, as this represents a major loss of later-generation gas-phase products such as the C_1 - C_3 carbonyls and C_4 - C_5 intermediates. For both catechol and phenoxy-phenylperoxy perturbations, benzene is the most sensitive of the primary aromatics, due to its higher yields of the phenolic pathway than toluene or xylene (Figure S27).

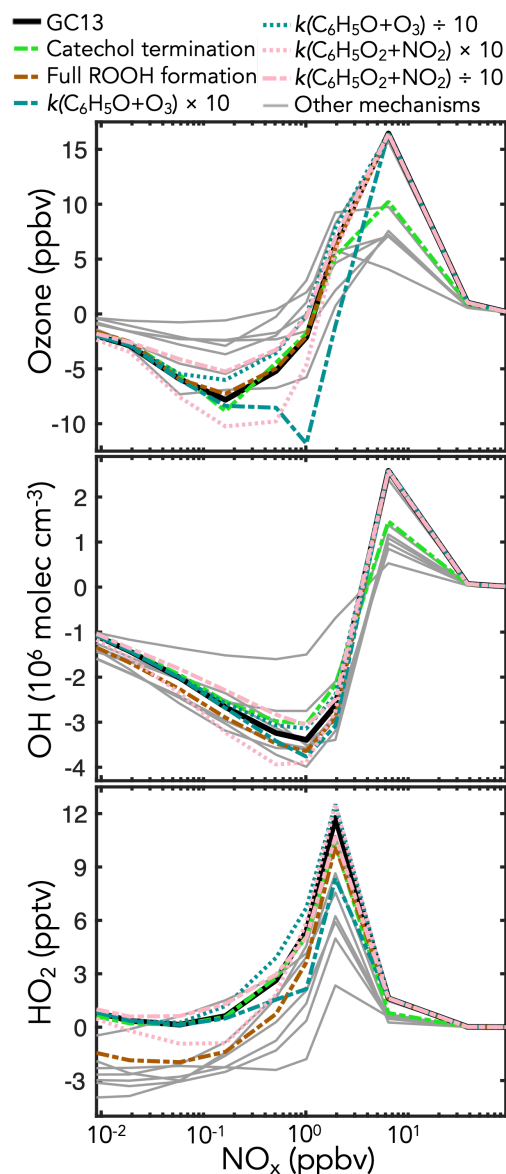


Figure 9: Sensitivity of midday ozone, OH, and HO₂ concentrations in the GC13 aromatic mechanism to changes in individual reaction rate constants and yields. The black lines show the standard GC13 box model simulation for the effect of mixed aromatic emissions in the continental boundary layer at 10:00-14:00 local time, relative to a simulation with no aromatic emission. The colored and patterned lines show the effects of individual changes in the GC13 mechanism. Gray lines show the results from the other mechanisms in Table 1. Results shown for the GC13 standard simulation and other mechanisms are the same as for Figure 8 except that BTX emissions were doubled to better show the changes from the individual reactions.

HO_x concentrations are also highly sensitive to the relative contributions of the radical propagation and termination pathways from the reactions of HO₂ with the initial bridged bicyclic peroxy radicals from BTX + OH under low-NO_x (< 1 ppb) conditions. A perturbation simulation with 100% radical termination (i.e. hydroperoxide formation), as in most mechanisms, reduces HO₂ in GC13 to levels similar to the other mechanisms, and causes a smaller reduction in OH,

highlighting the importance of this branching ratio. However, Xu et al. (2020) showed that hydroperoxide formation is minimal from the benzene-derived bridged bicyclic peroxy radical, motivating our treatment in GC13.

Results from additional sensitivity simulations are shown in Section S5 of the Supporting Information (Figures S27-S31). Changes to other reactions in the phenoxy-phenylperoxy system can be important – perturbations to the $C_6H_5O + NO_2$ rate have similar effects to those of $C_6H_5O + O_3$ because the two reactions are in direct competition, while increasing the $C_6H_5O_2 + NO_3$ rate increases ozone loss in a manner similar to increasing the $C_6H_5O_2 + NO_2$ rate, but peaking at slightly lower ambient NO_x concentrations. Mechanism outcomes are only mildly sensitive to changes in the reactions of nitrophenols, a pathway implemented in GC13 and MECCA. Other novel aspects of GC13, including the 0% yield of organonitrates from the reactions of NO with the initial bridged bicyclic peroxy radicals (following Xu et al. 2020) and increased HO_x recycling from the benzoylperoxy radical + HO_2 reaction, have minimal effects on ambient ozone and HO_x .

5. GLOBAL EFFECTS OF AROMATIC OXIDATION

5.1 METHODS

To investigate more deeply the effects of the BTX oxidation on atmospheric chemistry, we implement the GC13 mechanism into the GEOS-Chem CTM and compare it to selected other mechanisms in the GEOS-Chem environment. GEOS-Chem is driven by meteorology from the Modern-Era Retrospective analysis for Research and Applications, Version 2 (MERRA-2) assimilation product of the NASA Global Modeling and Assimilation Office (GMAO). We use GEOS-Chem version 12.3 (DOI 10.5281/zenodo.2658178) with added C_2H_4 and C_2H_2 chemistry (Kwon et al. 2021) as a base, which includes 196 species in its chemical mechanism (not including aromatic chemistry), of which 149 are **transported** ~~adveected~~. We run global simulations at $2^\circ \times 2.5^\circ$ horizontal resolution with 47 vertical layers. For each simulation, we perform an initial 8-month spinup (1 Mar – 1 Dec 2015), followed by one year of simulation from which seasonal and annual averages are output. We conduct one simulation with no aromatic emissions as a base case, and one simulation with GC13 for comparison. For simulations with GC13 chemistry, Henry's Law coefficients of newly included species (Table S1) are taken from Sander (2015) and Cabrera-Perez et al. (2016) for use in GEOS-Chem dry and wet deposition modules. **The added aromatic chemistry in the GC13 simulation increases overall CPU time by an average of 1.7% relative to the base simulation, attributable predominantly to gas-phase chemistry (64%) and transport (26%).**

Anthropogenic VOC emissions in our GEOS-Chem simulations are from the Community Emissions Data System (CEDS) (Hoesly et al., 2018), overwritten with the Multi-resolution Emission Inventory for China (MEIC; Zheng et al. 2018) and with the KORUS v5 inventory for the rest of East Asia (Woo et al. 2012, Jang et al. 2020). Biogenic emissions in GEOS-Chem are from the Model of Emissions of Gases and Aerosols from Nature (MEGAN) version 2.1 (Guenther et al., 2012), and open fire emissions are from the Global Fire Emissions Database (GFED) version 4 (van der Werf et al., 2010). Global annual emissions are 7.23 Tg, 10.42 Tg, and 7.30 Tg for benzene, toluene, and xylene from anthropogenic sources, and 1.67

Tg, 0.88 Tg, and 0.26 Tg respectively from ~~open fires biomass-burning~~. ~~No emissions of other gas-phase aromatics are included~~. Total BTX emissions are 60% higher than in the global model simulation of Taraborrelli et al. (2021), but only 4% higher (in carbon mass) than their total emissions of C₆-C₉ aromatics (including phenol, benzaldehyde, ethyl benzene, and lumped C₉ aromatics).

We also implement the two simplest alternative mechanisms, RACM2 and MOZART-T1, in the GEOS-Chem environment. MOZART-GC and SAPRC-11 were previously implemented in GEOS-Chem by Porter et al. (2017) and Yan et al. (2019) respectively, but neither were incorporated into the standard version of GEOS-Chem. Instead, aromatic chemistry previously implemented in the standard version of GEOS-Chem was simply parameterized to achieve reasonable glyoxal and methylglyoxal yields with fixed branching ratios for SOA and peroxyacetylnitrate (PAN) formation (Fu et al. 2008; Fischer et al. 2014). Comparison of GC13 to this parameterized GEOS-Chem aromatic chemistry and to MOZART-GC is shown in the SI (Section S7-S8) for reference to past GEOS-Chem studies.

5.2 EFFECTS ON OXYGENATED VOCs

Figure 10 shows the impact of GC13 aromatic chemistry on concentrations of glyoxal, methylglyoxal, and formic acid in the lowest 1 km of the atmosphere. Aromatic oxidation increases the tropospheric production of these three oxygenated VOCs by 30%, 5%, and 9% respectively. Although absolute changes are strongest in source regions, the relative contribution of aromatic chemistry to these concentrations extends globally, due both to later-generation production and to the longer lifetimes of aromatics relative to other precursors (e.g. anthropogenic alkenes and isoprene). Changes to gas-phase acetic acid are similar but smaller in magnitude to those of formic acid, with aromatic chemistry increasing production by 5%. Because these oxygenated VOCs are also formed in isoprene oxidation, the relative contribution of aromatics is much lower in high-isoprene regions, especially the tropics. In the Middle East, aromatics are responsible for >80% of glyoxal ~~because anthropogenic VOC emissions are high while biogenic emissions are low. which may play a critical role in aqueous SOA formation.~~

~~These findings for~~ Our simulated contributions of aromatic oxidation to the budgets of glyoxal and methylglyoxal are substantially larger than in previous GEOS-Chem studies using simplified mechanisms with lower yields (Fu et al., 2008; Silva et al., 2018), but are more consistent with results from the detailed mechanisms in Taraborrelli et al. (2021) and Yan et al. (2019). The increased glyoxal yields in GC13 align with general model findings of negative biases relative to satellite observations of glyoxal columns in regions with strong anthropogenic influence (Chan Miller et al., 2016; Liu et al., 2012; Myriokefalitakis et al., 2008; Silva et al., 2018; Stavrakou et al., 2011). Taraborrelli et al. (2021) ~~but generally smaller simulated even higher aromatic contributions to the global glyoxal budget~~, due in part to differences in the mechanisms (higher long-term yields of glyoxal and methylglyoxal from MECCA; see Figures S3-S4) and potentially to differences in non-aromatic glyoxal and methylglyoxal sources between the models. Neither Taraborrelli et al. (2021) nor Yan et al. (2019) discuss the contribution of aromatic chemistry to gas-phase formic or acetic acid budgets.

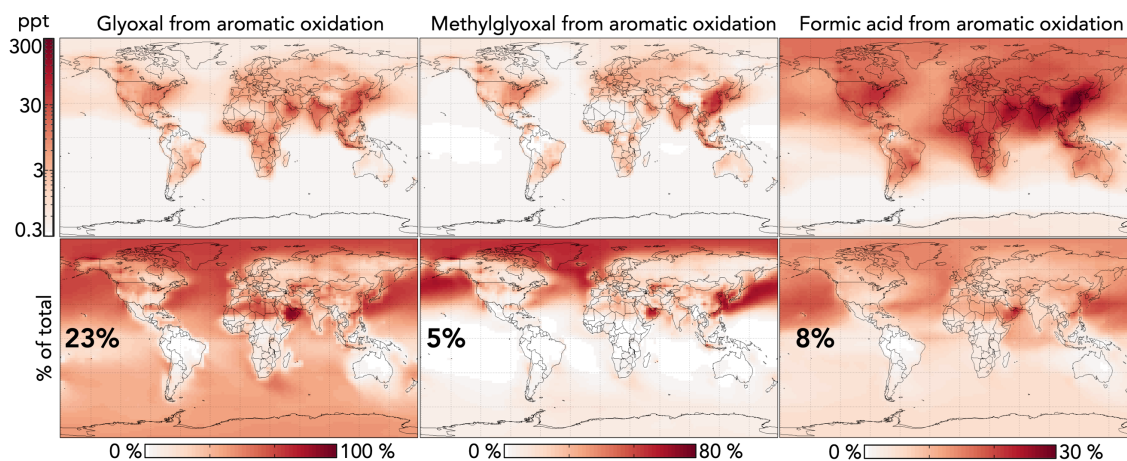


Figure 10: Effects of GC13 aromatic chemistry on glyoxal, methylglyoxal, and formic acid concentrations in GEOS-Chem. The top panels show absolute differences in annual mean concentrations below 1 km altitude between GEOS-Chem simulations with GC13 aromatic chemistry and with no aromatic chemistry. The bottom panels show the percentage contributions of aromatic chemistry to the total model concentrations of the three species. **Inset numbers are global percentages.** Color scales are logarithmic for top panels and linear for bottom panels.

Additional effects from aromatic chemistry on the global distribution of oxygenated VOCs are shown in Figure 11. PAN is subject to competing influences: methylglyoxal formation from aromatic oxidation increases the source strength of the acylperoxy radical, the organic precursor to PAN, while lower NO_2 due to phenoxy-phenylperoxy cycling tends to decrease PAN production. The former effect dominates in source regions; aromatic oxidation with the GC13 mechanism increases PAN mixing ratios over Northern Hemisphere continents by up to 40%. Downwind, particularly over oceans, PAN decreases due to phenoxy-phenylperoxy consumption of NO_2 .

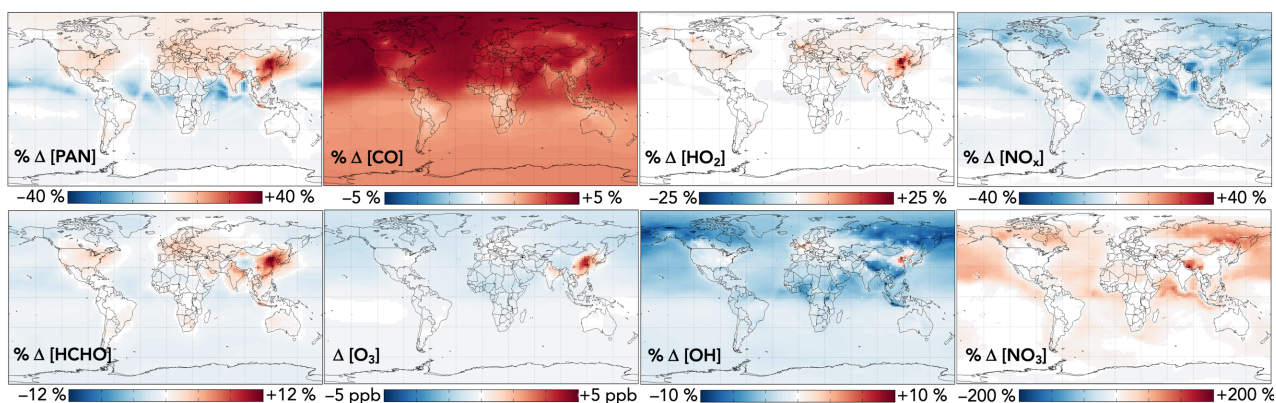


Figure 11: Effects of GC13 aromatic chemistry on PAN, formaldehyde, CO, ozone, HO_2 , OH, NO_x , and NO_3 concentrations. Panels show the absolute (for ozone) and relative (for others) differences in annual mean concentrations below 1 km altitude between GEOS-Chem simulations with GC13 aromatic chemistry and with no aromatic chemistry. Color scales are linear.

Formaldehyde exhibits a similar spatial pattern to PAN, with competing effects from its direct secondary production via aromatic oxidation, leading to locally increased mixing ratios of

up to 12% from aromatic oxidation, and indirect decreases due to reduced OH, which dominates downwind. Global formaldehyde production changes by just -0.1% from aromatic oxidation. GC13 increases the tropospheric CO burden relative to the simulation without aromatic chemistry by 3%, due both to direct production and the decreased OH sink, with less spatial heterogeneity than other effects. The changes to both formaldehyde and CO are spatially consistent with the findings of Taraborrelli et al. (2021) using MECCA, but are smaller in magnitude, driven primarily by the smaller change in tropospheric OH in GC13.

Figure 12 compares the effects of aromatic chemistry on oxygenated VOC concentrations in GEOS-Chem with GC13 to simulations with the MOZART-T1 and RACM2 mechanisms. The most prominent difference between the simulations is the higher overall yield of glyoxal from aromatic oxidation in GC13, especially in later-generation chemistry. This results in increases of up to 60% in surface glyoxal when switching from either MOZART-T1 or RACM2 to GC13, with the strongest effects over the Middle East (where a lack of biogenic emissions renders aromatics the dominant glyoxal source) and in remote areas where later-generation chemistry dominates and decreases in OH increase VOC lifetimes. Overall tropospheric glyoxal loadings are 10% lower in MOZART-T1 and 13% lower in RACM2 than in GC13, while tropospheric glyoxal production from aromatics is 38% lower in MOZART-T1 and 61% lower in RACM2 than in GC13.

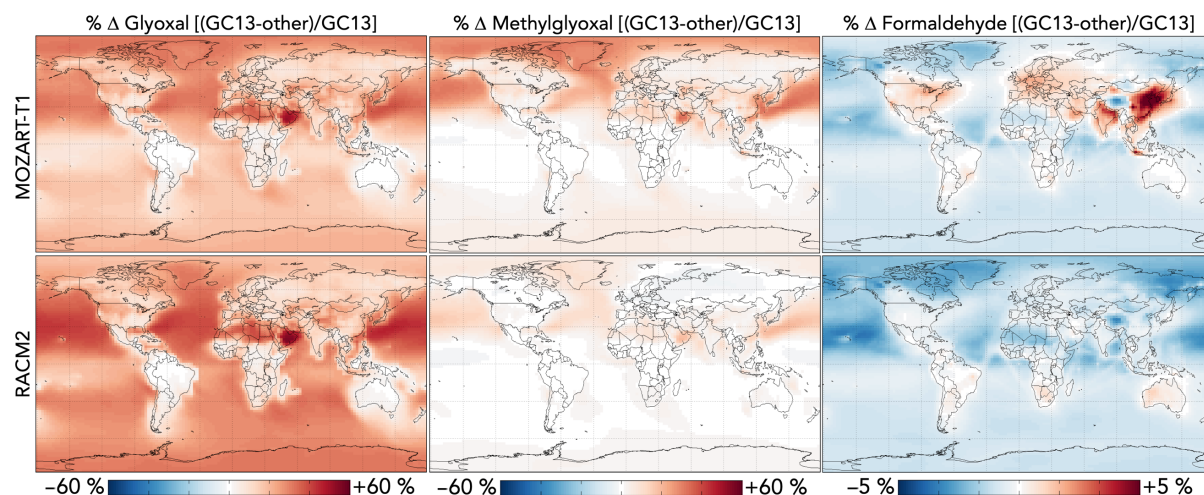


Figure 12: Differences in glyoxal, methylglyoxal, and formaldehyde concentrations between the GC13 mechanism and the MOZART-T1 and RACM2 mechanisms for aromatic chemistry. Values are percent differences in annual mean concentrations below 1 km altitude between GEOS-Chem simulations with the GC13 mechanism and with the MOZART-T1 or RACM2 mechanism. Color scales are linear.

Methylglyoxal exhibits similar though less pronounced differences between the mechanisms, confined mostly to the northern hemisphere where its production is greatest. The strongest differences are seen for the MOZART-T1 mechanism, which produces only 62% as much methylglyoxal from BTX as GC13 in global simulations, resulting in decreases in surface methylglyoxal mixing ratio of up to 30% (2% globally). The RACM2 mechanism produces more methylglyoxal than MOZART-T1 (74% as much as GC13 globally), and therefore exhibits smaller changes (up to 10% decreases locally, 1% globally), and even some local increases.

Differences in formaldehyde between the mechanisms are minor; global tropospheric formaldehyde is 0.4% higher with MOZART-T1 and 0.8% higher with RACM2 relative to GC13. The most prominent change is an increase of 5% in boundary layer formaldehyde over Northeast China with GC13 relative to MOZART-T1.

5.3 EFFECTS ON OXIDANTS

The effects of aromatic chemistry on radical and ozone budgets are shown in Figure 11. Impacts on HO_x and ozone are consistent with the results of the continental boundary layer simulations in Section 4.3. HO_2 is increased by aromatic chemistry by up to 20% annually averaged in high- NO_x aromatic source regions, but exhibits little change ($<0.1\%$) globally. OH and ozone both increase in high- NO_x aromatic source regions, by up to 6% and 5 ppb respectively on annual averages, but decrease elsewhere, largely due to phenoxy-phenylperoxy radical cycling. On a global scale, these decreases slightly dominate; aromatic chemistry reduces tropospheric OH and ozone by 2.2% and $<1\%$ (0.37 ppb) on annual average. The effect on OH has a strong seasonal variation (Figure 13), with increases in source regions in the NH winter – up to 24% over Northeast China – but small effects and even slight decreases in the summer. This is due to the importance of carbonyl photolysis as a wintertime OH source (Li et al. 2021).

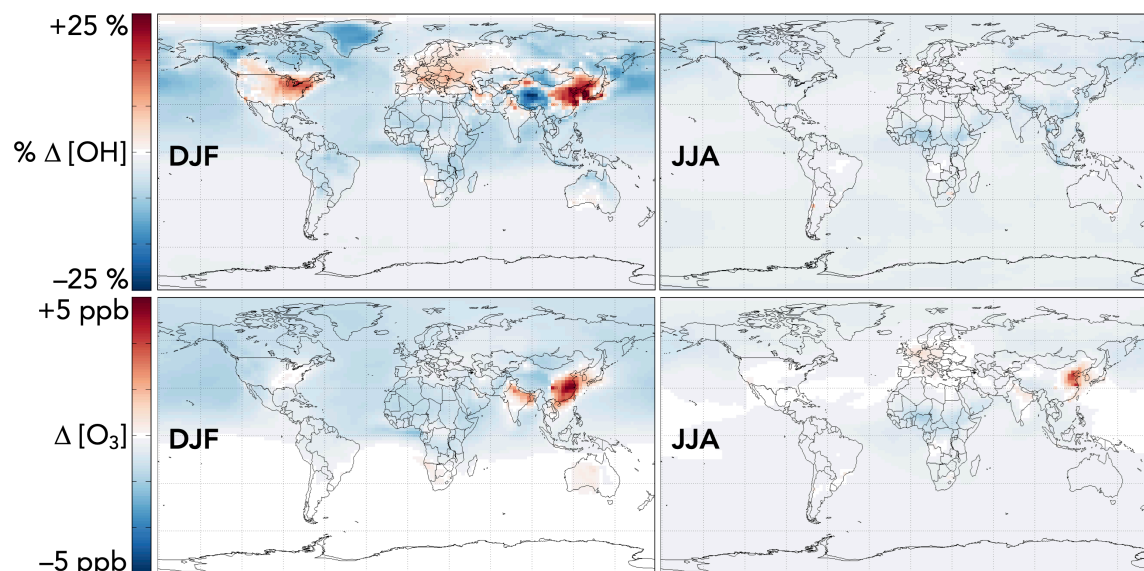


Figure 13: Seasonal effects of aromatic chemistry on OH and ozone concentrations. Panels show the relative (for OH) and absolute (for ozone) differences in mean concentrations below 1 km altitude between GEOS-Chem simulations with GC13 aromatic chemistry and with no aromatic chemistry for Dec. 1 2015 – Mar. 1 2016 (left) and Jun. 1 2016 – Sep. 1 2016 (right). Color scales are linear.

NO_x concentrations decrease everywhere as a result of aromatic chemistry, most notably in regions downwind of aromatic emissions in the Northern Hemisphere (Figure 11). Though PAN and peroxybenzoyl nitrate can act as NO_x reservoirs, releasing NO_x in remote air, there are additional NO_x sinks from phenoxy-phenylperoxy cycling and nitrophenol formation. The impacts of NO_3 production from phenylperoxy + NO_2 are particularly pronounced in downwind

regions with very low NO_x concentrations, where this pathway can increase annual average NO_3 concentrations by up to 200%.

Generally, these effects of aromatic chemistry on oxidants in GC13 are consistent with those from MECCA in Taraborrelli et al. (2021), which also showed global decreases in NO_x , OH, and ozone from aromatic chemistry, along with local increases and seasonal cycles for OH and ozone in areas of strong aromatic emissions, and strong increases in NO_3 . Global average changes tend to be stronger in Taraborrelli et al. (2021), especially for OH and ozone, consistent with the sharper decreases in these compounds due to aromatic chemistry in continental boundary layer simulations (Figure 8). By contrast, SAPRC implemented in GEOS-Chem by Yan et al. (2019) showed increases in ozone, OH, and NO_x concentrations as well as decreases in NO_3 . We attribute this primarily to the absence of phenoxy-phenylperoxy cycling in SAPRC.

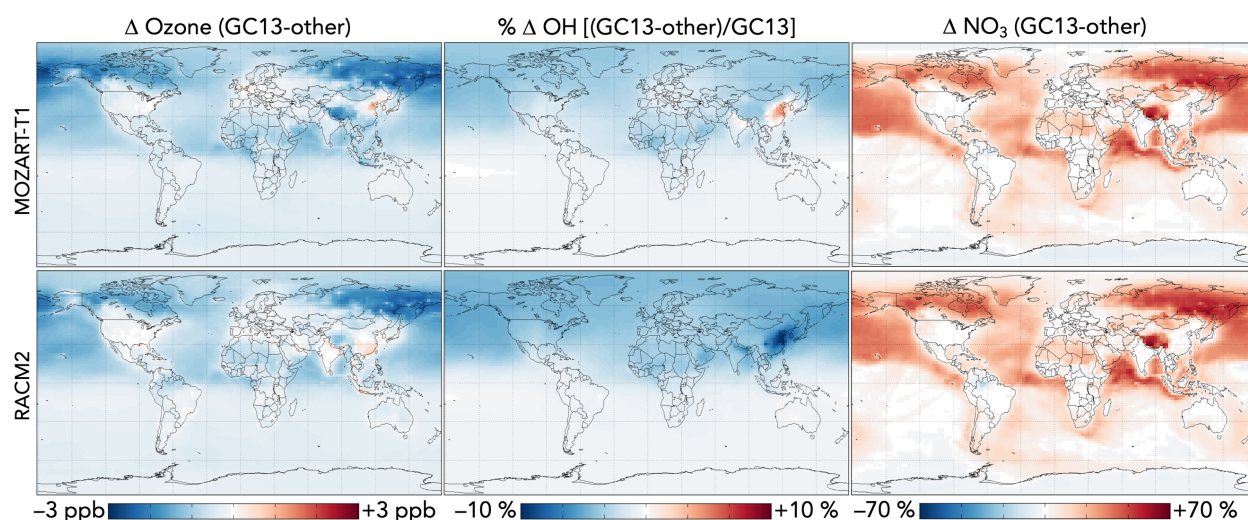


Figure 14: Differences in ozone, OH, and NO_3 concentrations between the GC13 mechanism and the MOZART-T1 and RACM2 mechanisms for aromatic chemistry. Values are absolute (for OH) or percent (for other species) differences in annual mean concentrations below 1 km altitude between GEOS-Chem simulations with the GC13 mechanism and with the MOZART-T1 or RACM2 mechanism. Color scales are linear.

Differences in oxidant concentrations between global simulations with GC13 and with the MOZART-T1 and RACM2 mechanisms are shown in Figure 14. Changes to ozone, HO_x , and NO_x between the mechanisms can be attributed primarily to the inclusion of phenoxy-phenylperoxy cycling and increased OH recycling in GC13. Surface ozone and OH are reduced in GC13 relative to MOZART-T1 and RACM2, consistent with findings in box model simulations (Section 4.3). Both MOZART-T1 and RACM2 cause increases in the tropospheric ozone burden relative to a simulation without aromatic chemistry, as Yan et al. (2019) also showed for SAPRC, while aromatic chemistry in GC13, like MECCA (Taraborrelli et al. 2021), causes a global ozone decrease. As an additional consequence of their lack of phenoxy-phenylperoxy cycling, the MOZART-T1 and RACM2 mechanisms simulate much lower tropospheric burdens of NO_3 (by 18% and 19% respectively) and slightly higher NO_x burdens (by 3% and 2% respectively) than GC13. Finally, higher OH recycling in GC13 leads to local increases in OH relative to MOZART-T1 (up to 4%) in source regions, but globally, differences in tropospheric and surface OH are <2%.

6. ACCOUNTING FOR PRIMARY NON-BTX AROMATICS

Our GC13 mechanism and its implementation in GEOS-Chem focused on the effect of BTX emissions. Taraborrelli et al. (2021) estimated that non-BTX aromatics contribute 54% of global aromatic emissions by carbon mass with C₆-C₈ oxygenates, C₈-C₉ aromatic hydrocarbons, and higher aromatics each contributing approximately one third of this non-BTX fraction. The contributions from C₆-C₈ oxygenate emissions including benzaldehyde, phenol, cresols, and catechols can be readily accounted for in GC13 since they are already included explicitly as secondary species. C₈-C₉ aromatic hydrocarbons including ethylbenzene, styrene, and trimethylbenzenes can be added as independent species using simplified initial oxidation reactions to convert them directly to products already included in GC13. We have provided such reactions, simplified from MCM (Jenkin et al., 2003), in Section S1 of the SI, and this would add 3 species and 7 reactions to GC13. Higher aromatics such as polycyclic aromatic hydrocarbons (PAHs) are semi-volatile and their oxidation products partition heavily into the aerosol phase (Chan et al., 2009; Chen et al., 2016), so their simulation is more relevant to SOA formation than to oxidant chemistry.

7. CONCLUSIONS

We developed a new compact mechanism (GC13) for fast and accurate simulation of benzene, toluene, and xylene (BTX) oxidation chemistry in atmospheric models. ~~Our mechanism~~ GC13 includes only 17 unique species to describe the aromatic oxidation chain and 44 unique reactions. It includes recent information from experimental and computational studies, and captures the important features of much more complex ~~and computationally costly~~ mechanisms. In particular, it incorporates recent evidence for efficient radical recycling from the bimolecular reactions of bridged bicyclic peroxy radicals, and explicitly treats phenoxy-phenylperoxy radical cycling as a sink for ozone and NO_x. ~~Ethylbenzene, styrene, and trimethylbenzene can be added to GC13 with just 3 additional species and 7 reactions. GC13 presently focuses on gas-phase chemistry, but future application to secondary organic aerosol (SOA) formation~~ ~~development of a chemically detailed aerosol module could build on the current treatment~~ ~~leverage the inclusion of methylcatechols and nitrophenols as explicit species in GC13 as key SOA precursors.~~

We compared outcomes of GC13 to other aromatic oxidation mechanisms in box model simulations of environmental chamber observations and of the continental boundary layer. Product yields from the mechanism exhibit good agreement with environmental chamber observations and result in increased glyoxal and methylglyoxal yields from aromatic oxidation relative to previous reduced mechanisms. Radical cycling in GC13 tends to increase simulated HO_x radical concentrations, which past mechanisms have tended to underestimate (Bloss et al. 2005a, Carter et al. 2013). Phenoxy-phenylperoxy radical cycling decreases ozone production, which past mechanisms have tended to overestimate (Bloss et al. 2005a). We find that the effects of aromatic chemistry on HO_x, NO_x, and ozone are strongly sensitive to uncertainties in the chemistry of the phenoxy-phenylperoxy system, with smaller but significant sensitivities to radical recycling from bridged bicyclic peroxy radicals and the oxidative fates of catechols.

We implemented the GC13 mechanism in the GEOS-Chem global atmospheric chemistry model and compared results to simulations without aromatic chemistry and with alternative

mechanisms (MOZART-T1, RACM2). GC13 increases the computational cost of an annual simulation by 1.7% relative to a simulation without aromatic chemistry. Aromatic oxidation plays a particularly important role in the tropospheric budgets of small dicarbonyl species, contributing 23% and 5% to the global production of glyoxal and methylglyoxal, respectively. These values are substantially higher than those simulated with the MOZART-T1 and RACM2 mechanisms. Formaldehyde concentrations over Northeast China increase by 12% due to aromatic chemistry. Formic and acetic acids increase globally by 9% and 5%, respectively, through oxidation of ketene-enols generated by aromatic ring-breaking. Aromatic chemistry decreases global tropospheric OH by 2.2% and ozone by less than 1% but increases them in polluted environments in winter, such as in Northeast China where wintertime OH increases by 24% and surface ozone increases by 5 ppb.

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