

# COVID-19 lockdown ~~NO<sub>x</sub>~~ emission reductions can explain ~~more~~ **half** of the coincident increase in global atmospheric methane

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**Abstract.** Compared to 2019, measurements of the global growth rate of background (marine air) atmospheric methane rose by about 50% 5.5 ppb/yr in 2020, reaching 15.2 ppb/yr. Models of gGlobal atmospheric chemistry models have previously shown that reductions in nitrogen oxide (NO<sub>x</sub>) emissions reduce levels of the hydroxyl radical (OH), and lengthen the methane lifetime. Acting in the opposite sense, reductions of carbon monoxide (CO) and non-methane volatile organic compound (NMVOC) emissions increase OH and shorten methane's lifetime. Using estimates of NO<sub>x</sub>, CO and NMVOC emission reductions associated with COVID-19 lockdowns around the world in 2020, together with model-derived regional and sectoral aviation sensitivities of methane to these NO<sub>x</sub> emissions, we find that NO<sub>x</sub> emissions reductions led to a 4.3 (3.6 to 5.0) ppb/yr increase can fully explain the observed surge in the global methane growth rate. Reductions in CO and NMVOC emissions partly counteracted this, changing (reducing) the methane growth rate by -1.1 (-0.5 to -1.5) ppb/yr (CO) and -0.1 (0.0 to -0.3) ppb/yr (NMVOC). Uncertainties refer to ±1 standard deviation model ranges in sensitivities. Whilst changes in NO<sub>x</sub> anthropogenic emissions related to COVID-19 lockdowns are probably not the only important factor that has influenced methane since the beginning of 2020 during 2020, it is clear these results indicate that they have had a large impact are a key factor that will need to be included within any attribution study, and that the net effect of NO<sub>x</sub>, CO and NMVOC emissions changes may well be the dominant driver can explain over half of the observed 2020 se-recent methane changes. Large uncertainties remain in both emissions changes during the lockdowns and methane's response to them; nevertheless, this analysis suggests that further research into how atmospheric composition changed over the lockdown periods will help us to interpret past methane changes and to constrain future methane projections. The major global-scale changes in composition of the Earth's atmosphere measured during lockdown provide unprecedented constraints on the sensitivity of the atmospheric chemical system to changes in emissions, and are of great utility for evaluating policy-relevant models.

## 1 Introduction

Methane is a powerful greenhouse gas and important precursor of tropospheric ozone; both are key air pollutants and short-lived climate forcers (SLCFs). The 2013 Intergovernmental Panel on Climate Change assessment estimated methane's Global

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Warming Potential (GWP) over a 100-year time horizon to be 28 (Myhre et al., 2013); updates to its short-wave radiative forcing have increased this value by 14% (Etminan et al., 2016). Advances in our understanding of how GWP metrics relate to climate change indicate that it is not only the magnitude of emissions, but also the rates of change of SLCFs like methane that strongly influence near-term global temperature changes (Allen et al., 2018; Cain et al., 2019). Several factors in addition to rising anthropogenic methane emissions have influenced the evolution of atmospheric methane from its pre-industrial level of ~700 ppb to its present-day value of over 1900 ppb. The Intergovernmental Panel on Climate Change's Sixth Assessment Report (Szopa et al., 2021) assessed how changes in emissions of NO<sub>x</sub>, CO, and NMVOCs have contributed to historical changes in methane, through their impacts on OH, the main sink for methane. A range of modelling studies have explored these indirect impacts on methane (e.g., Shindell et al., 2005, 2009; Stevenson et al. 2013; Thornhill et al., 2021). For example, the Atmospheric Chemistry and Climate Model Intercomparison Project found that 1850-2000 increases in anthropogenic NO<sub>x</sub> emissions had reduced year 2000 methane levels by 955 ppb, whilst growing emissions of CO and NMVOCs had increased methane by 150 ppb and 59 ppb, respectively (Table 7 of Stevenson et al., 2013). These results have quite large uncertainties (at least ±10%, based on the model range in Stevenson et al. 2013), but indicate that non-methane (especially NO<sub>x</sub>) emissions have had very significant impacts on methane. These post-2013 updates increase the importance of better understanding of what controls methane and its evolution is vital for progress towards in the context of the Paris Climate Agreement target that seeks to limit warming to 1.5°C above pre-industrial levels.

Following the onset of the COVID-19 pandemic in early 2020, the trace gas composition of the global atmosphere ~~changed~~ has shown substantial ~~ly~~ changes. Atmospheric nitrogen oxide (NO<sub>x</sub>) levels have reduced as surface and aviation NO<sub>x</sub> emissions fell (Bauwens et al., 2020; Cooper et al., 2022), whilst the measured growth rate of methane (CH<sub>4</sub>) ~~rose~~ has risen sharply in 2020 (Laughner et al., 2021). The observed NO<sub>x</sub> changes are clearly linked to falls in emissions resulting from lockdowns, but the driver of the methane increases is less clear, with some studies discussing causes related to decreases in OH (e.g., Weber et al., 2020; Laughner et al., 2021) while others suggest rises in sources (e.g., Feng et al., 2022) remain unexplained (e.g., Vaughan, 2021). Methane, and NO<sub>x</sub>, CO and NMVOCs are linked through via the oxidising capacity of the atmosphere, specifically by the abundance of the hydroxyl (OH) radical. The response of global atmospheric chemistry to the large lockdown perturbation since early 2020 provides an opportunity to explore the sensitivity of the NO<sub>x</sub>-CO-NMVOC-OH-CH<sub>4</sub> system, and compare models and observations. Here we use model-derived sensitivities of global methane to NO<sub>x</sub>, CO and NMVOC emissions, together with estimated changes in anthropogenic NO<sub>x</sub> emissions of these species related to the COVID-19 lockdowns, to calculate ~~estimated~~ the impacts from lockdown emissions changes on the growth rate of global methane, and compare this to methane observations.

## 60 2 Measurements of atmospheric methane and nitrogen oxides

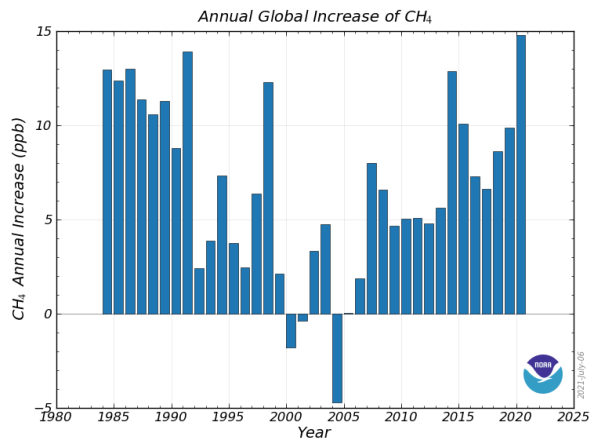
Recent methane measurements from the US National Oceanographic and Atmospheric Administration (NOAA) show that the atmospheric (marine air background) methane growth rate rose sharply from 9.79 ppb/yr in 2019 to 15.244.8 ppb/yr in 2020,

the highest than any preceding annual value in the 37-year NOAA record, that started in 1984 (Figure 1; Dlugokencky, 2022). Many of the earlier large year-to-year jumps in methane's growth rate relate in part to variability in climate and emissions associated with El Niño Southern Oscillation (ENSO), and in particular because of modulation of methane's main sink, oxidation by OH (Turner et al., 2018; Zhao et al., 2020). ENSO indices have not shown strong variations over 2019-2020 (WMO, 2021). The start of 2020 marked the onset of a La Niña that has persisted into 2022. Past La Niña's have not always shown clear links with methane's growth rate, and the influence of the current ENSO phase on methane is uncertain.

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Measurements of nitrogen dioxide (NO<sub>2</sub>) from satellite instruments and nitrogen monoxide (NO) and NO<sub>2</sub> from surface sites show that levels of atmospheric NO<sub>x</sub> (NO + NO<sub>2</sub>) dramatically fell globally during 2020 (Bauwens et al., 2020; Laughner et al., 2021; Cooper et al., 2022). This was driven by as COVID-19 lockdowns around the world that reduced emissions, mainly from transportation (Venter et al., 2020; Lamboll et al., 2021; Doumbia et al., 2021).



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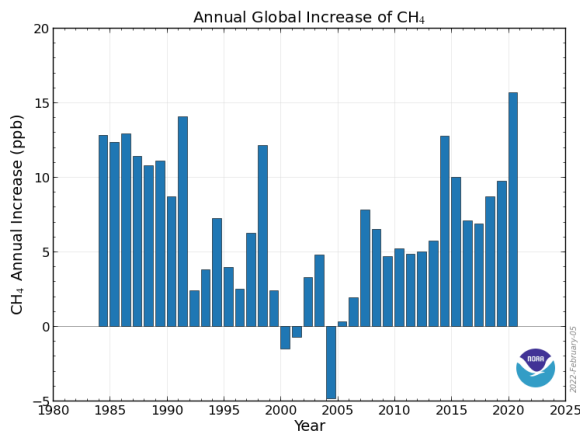


Figure 1: Global annual changes in surface atmospheric methane mole fraction (ppb) 1984-2020 (Dlugokencky, 2022).

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### 3 Sensitivity of global methane to NO<sub>x</sub>, CO and NMVOC emissions

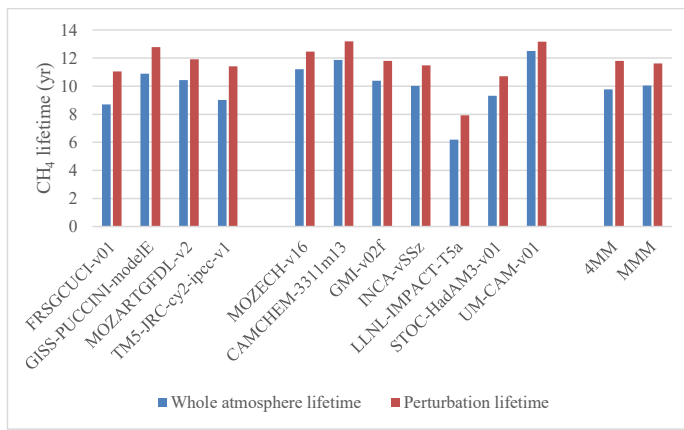
Global atmospheric chemistry model simulations indicate that decreases in NO<sub>x</sub> emissions lead to reductions in OH and increases in the global methane lifetime (Prather, 1994; Derwent et al., 2001; Wild et al., 2001; Stevenson et al., 2004; Weber et al., 2020). Similarly, decreases in CO and NMVOC emissions lead to increases in OH and decreases in methane lifetime (Derwent et al., 2001; Wild et al., 2001). Multi-model studies have calculated methane effects for NO<sub>x</sub> emissions from specific world regions (Fry et al., 2012) and the aviation sector (Lee et al., 2021). Although methane has an atmospheric lifetime of about 10 years, the models show that its peak response occurs within a few months of the cessation of a sudden short-lived (month- or year-long) pulse of extra-NO<sub>x</sub> emissions (Derwent et al., 2001; Wild et al., 2001; Stevenson et al., 2004). This indicates that the impacts on methane from the sudden changes in emissions associated with lockdowns will have had rapid impacts on methane's growth rate. Multi-model studies have calculated methane effects for NO<sub>x</sub> emissions from specific world regions (Fry et al., 2012) and the aviation sector (Lee et al., 2021).

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We first illustrate the basis of our approach by describing the model experiments performed by Derwent et al. (2001), who conducted a series of experiments-simulations with the global tropospheric chemistry model STOCHEM to quantify the impact of NO<sub>x</sub> emissions on methane. They compared a 4-year long base simulation with a perturbation simulation that was identical

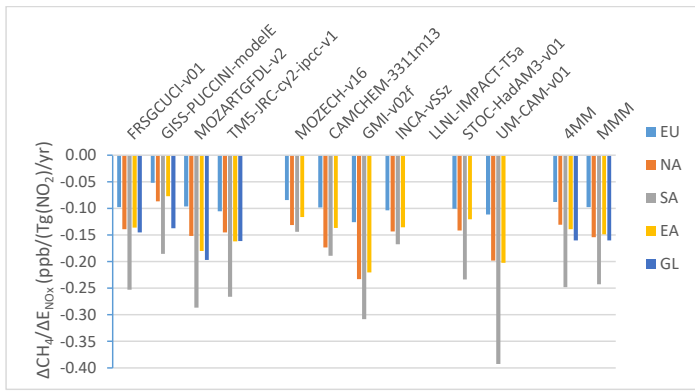
95 apart from an enhancement in NO<sub>x</sub> emissions of magnitude 1 Tg(NO<sub>2</sub>), added during the first month with the Northern Hemisphere surface anthropogenic NO<sub>x</sub> emissions distribution. The extra NO<sub>x</sub> produced a short-lived increase in OH, and this led to a rapid depletion of global methane, which peaked ~~with a magnitude of~~ around 0.39 Tg(CH<sub>4</sub>) after about six months. The methane deficit then exponentially decayed with an e-folding timescale of about 12 years (the methane perturbation lifetime,  $\tau$ ), with methane levels returning towards their base values. Wild et al. (2001) conducted similar  
100 experiments, with year-long emissions perturbations using a different model (UCI CTM), and found very similar behaviour but with slightly larger sensitivities: 1 Tg(NO<sub>2</sub>) from global fossil fuel sources yielded a 0.55 Tg depletion of CH<sub>4</sub>. These studies also investigated the impact of CO and NMVOC emissions. Changes in global methane burden (Tg) are converted to changes in tropospheric mole fraction (ppb) using the total atmosphere mass of  $5.113 \times 10^9$  Tg and a fill factor of 0.973 for conversion of a total atmosphere abundance to a tropospheric abundance (Prather et al., 2012). We assume the troposphere is  
105 well mixed, so surface changes will be the same as whole troposphere changes.  
More recently, Fry et al. (2012) analysed results from multi global models that took part in experiments from the Hemispheric Transport of Air Pollutants (HTAP) study that in order to isolated the impacts on methane of surface NO<sub>x</sub>, CO and NMVOC emissions from Europe (EU), North America (NA), and South Asia (SA) and East Asia (EA). We utilise that ensemble of model results here; these models are descriptions are given in Fiore et al. (2009). Models performed a base simulation, and a series of further repeat simulations with 20% lower anthropogenic emissions (For each region and each species for each region, Fry et al. (2012) compared each model performed a base simulation and to others one with 20% lower regional anthropogenic NO<sub>x</sub> emissions from that region. In addition to the 20% regional emission reduction experiments, some models also performed global 20% emission reduction experiments (Wild et al., 2012). Four models include results from all the regional and global perturbation simulations: FRSGUCI-v01, GISS-PUCCINI-modelE, MOZARTGFDL-v2, and TM5-JRC-cy2-ipcc-v1. We calculate a 'four model mean' (4MM) based on these model results. We also show results from the other models to illustrate the range of model behaviour, and show 'multi-model mean' (MMM) results from all available simulations. In the HTAP simulations, had methane was fixed as a prescribed boundary condition, precluding direct diagnosis of changes in methane. However, methane changes can be but diagnosed indirectly, by analysing the change in methane lifetime associated with the changes in tropospheric OH sink in each run. We convert these to whole atmosphere lifetimes by assuming fixed lifetimes for methane loss to soils (160 yr) and in the stratosphere (120 yr) (Prather et al., 2012). The HTAP experiments also included a global methane perturbation simulation – allowing the methane feedback factor and perturbation lifetime to be calculated (Prather, 1994; Holmes, 2018). Figure 1 shows whole atmosphere and perturbation methane lifetimes for the six HTAP models, with typical values of around 10 years and 12 years, respectively.

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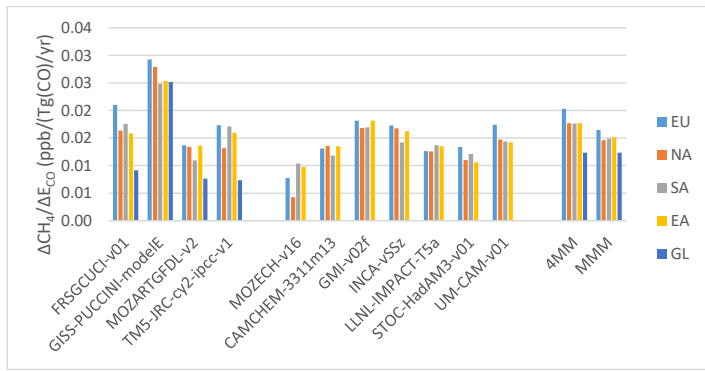
**Figure 1.** Whole atmosphere and perturbation methane lifetimes (years) for the HTAP models, together with the four model mean (4MM) of the core models (four models on the left), and the multi-model mean (MMM).

Differences between simulations yielded the change in methane lifetimes due to changes in regional emissions. From these changes in methane lifetime, the equilibrium change in methane was calculated; that is the change in methane that would have been achieved if methane levels had been free to respond (e.g., see Stevenson et al., 2013). In model simulations where methane is not prescribed, methane adjusts towards equilibrium with an e-folding timescale given by the methane perturbation lifetime ( $\tau$ ) (Derwent et al., 2001; Wild et al., 2001; Holmes, 2018). We convert equilibrium methane changes derived from sustained changes in emissions to the equivalent methane response for a pulse of emissions for each experiment. We use each model's perturbation lifetime to calculate the fraction of the equilibrium response that would have been reached after one year; e.g., for the multi-model mean (MMM) a methane perturbation lifetime of 11.2 years (Holmes, 2018) this fraction is  $(1 - e^{-1/\tau}) = 8.72\%$ . This method is appropriate because we compare to changes in the observed annual growth rate, and is justified by the rapid response of global methane seen in transient model simulations where methane is free to respond, and because the largest lockdown emissions' perturbations occurred in the first half of 2020. We normalise results to produce global methane sensitivities per Tg of gas emitted for each HTAP region and globally for each model. Figures 2, 3 and 4 show global methane sensitivities for NO<sub>x</sub>, CO and NMVOC emissions are shown in Figures 2, 3 and 4, respectively.



**Figure 2.** Sensitivity of global methane (ppb) to changes in surface anthropogenic NO<sub>x</sub> emissions (Tg(NO<sub>2</sub>)/yr) for the HTAP models for four regions (Europe, EU; North America, NA; South Asia, SA; and East Asia (EA), and globally (GL). Global results are refer just only available for the four core models, shown on the left of the figure. Also shown are the 4MM and MMM. There are no results for the LLNL-IMPACT-T5a model for NO<sub>x</sub>; it is included to maintain consistency with Figures 3 and 4.

Figure 2 shows relatively consistent responses to NO<sub>x</sub> emissions, with all models least sensitive to EU NO<sub>x</sub> emissions and most sensitive to SA, with NA and EA in between. The 4MM sensitivities are slightly lower than the MMM.



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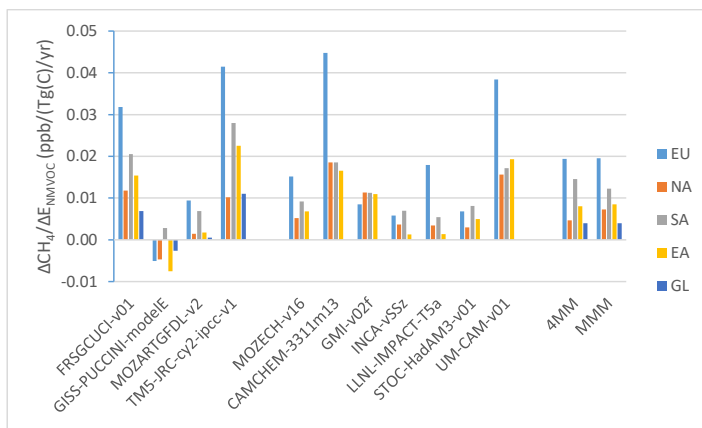
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**Figure 3.** As Figure 2, but methane sensitivities for changes in surface anthropogenic CO emissions (Tg(CO)/yr).

As for NOx, Figure 3 shows relatively consistent behaviour across the models for CO, with less variation between regions, reflecting the longer lifetime of CO, which makes the location of emissions less important. The 4MM sensitivities for CO are slightly larger than the MMM values.



**Figure 4.** As Figure 2, but methane sensitivities for changes in surface anthropogenic NMVOC emissions (Tg(C)/yr).

Figure 4 shows more divergence in model response to NMVOC emissions, with one model (GISS-PUCCINI-modelE) displaying an opposite sensitivity to the other models (apart from for SA emissions), and some models showing quite large sensitivities, whilst others are small. This probably reflects differing methods of representing NMVOCs in each model, in terms of both the number of species, grouping together of species, and the sophistication of their oxidation chemistry. Somewhat fortuitously, the 4MM and MMM are similar.

The HTAP experiments used 2001 as their base year, prescribing global methane to be 1760 ppb<sub>v</sub>, and each model used their own best estimates of global 2001 emissions. In 2020, surface level background global mean methane was ~1870 ppb<sub>v</sub>, and emissions of NOx, CO and NMVOCs had changed relative to 2001. Sensitivities of methane to emissions derived from the HTAP results will differ somewhat from those that would be found if 2020 conditions were used, and this represents an important caveat to our results. However, these differences are unlikely to be substantial, and no more up-to-date multi-model study of the impacts of regional NOx, CO and NMVOC emissions on methane has been published to date, so it represents our best source of information in the literature.



180 \*\*\*Revised to here 13/07/22\*\*\*

and for NO<sub>x</sub> from aviation (Lee et al., 2021)

185 Similar model simulations have calculated the sensitivity of methane to aviation NO<sub>x</sub> emissions. Wild et al. (2001) and Stevenson et al. (2004) conducted pulse experiments adding NO<sub>x</sub> using the global aviation NO<sub>x</sub> emissions distribution, and found a peak impact on global methane of about 2.5-2.6 Tg (equivalent to mole fractions of 0.88-0.92 ppb) for a 1 Tg(NO<sub>2</sub>) emission perturbation. Stevenson and Derwent (2009) also found spatial variation in sensitivity for aviation NO<sub>x</sub>, with the more sensitive regions tending to have lower background NO<sub>x</sub> levels. The most up-to-date study of aviation NO<sub>x</sub> is Lee et al. (2021), who assessed multi-model results for aviation NO<sub>x</sub> emissions using sustained emissions changes, similarly to the HTAP study. Lee et al. (2021) report (their Table 3) a methane radiative forcing sensitivity to aviation NO<sub>x</sub> emissions of -15.8 mW m<sup>-2</sup> (Tg(N) yr<sup>-1</sup>)<sup>-1</sup>. We convert this to a methane mole fraction sensitivity to NO<sub>x</sub> emissions using the relationship between changes in mole fraction and radiative forcing given by Myhre et al. (1998), and then, which we convert these, using a similar methodology to that described above, to the equivalent response for a pulse of emissions. This yields a sensitivity of methane to a pulse change in aviation NO<sub>x</sub> emissions of 1.12 ppb (CH<sub>4</sub>)/Tg(NO<sub>2</sub>) yr<sup>-1</sup>, and find a similar to, but slightly higher than sensitivity results from the earlier studies. Lee et al. (2021) also report a 95% likelihood range on the radiative forcing sensitivity, which translates to a standard deviation of 0.21 ppb (CH<sub>4</sub>)/Tg(NO<sub>2</sub>) yr<sup>-1</sup>, which we take to be a representative uncertainty for the mole fraction sensitivity to aviation NO<sub>x</sub> emissions.

195 Table 1 summarises results from all these existing studies. For surface NO<sub>x</sub> emissions reductions of 1 Tg(NO<sub>2</sub>) over one year, models find peak increases in global methane burden of about 0.5 Tg, with about five times higher sensitivities for reductions in aviation NO<sub>x</sub>. There is significant variation in sensitivity between regions for surface NO<sub>x</sub> emissions, with South Asia about twice as sensitive as North America and East Asia, and three times as sensitive as Europe (Fry et al., 2012). Stevenson and Derwent (2009) also found spatial variation in sensitivity for aviation NO<sub>x</sub>, with the more sensitive regions tending to have lower background NO<sub>x</sub> levels.

NO <sub>x</sub> emission region/sector	ACH <sub>4</sub> /AE <sub>NO<sub>x</sub></sub> Tg(CH <sub>4</sub> )/Tg(NO <sub>2</sub> ) yr <sup>-1</sup>	2020-2019 AE <sub>NO<sub>x</sub></sub> Tg(NO <sub>2</sub> ) yr <sup>-1</sup>	ACH <sub>4</sub> Tg(CH <sub>4</sub> )
Surface emissions			
Global	-0.55 <sup>a</sup>	-19.38	
N. Hemisphere	-0.39 <sup>b</sup>	-16.72	
S. Hemisphere	-1.1 <sup>b</sup>	-2.66	2.93
Europe	-0.28 <sup>c</sup>	-2.65	0.74

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N. America	-0.47 <sup>e</sup>	-2.55	1.20
E. Asia	-0.44 <sup>e</sup>	-4.40	1.94
S. Asia	-0.88 <sup>e</sup>	-3.78	3.33
NH minus 4 regions	-0.39 <sup>b</sup>	-3.34	1.30
<i>Aviation emissions</i>			
Global	-2.6 <sup>a</sup>	-0.83	
Global	-2.5 <sup>d</sup>	-0.83	
Global	-2.3 <sup>e</sup>	-0.83	1.94
Various 10° x 10° model grid boxes	-1.9 to -15 <sup>f</sup>		

205 **Table 1: Sensitivity of changes in the global methane burden ( $\Delta CH_4$ ; units  $Tg(CH_4)$ ) to changes in NO<sub>x</sub> emissions ( $\Delta E_{NO_x}$ ; units  $Tg(NO_2) \cdot yr^{-1}$ ) from several modelling studies, calculated for a variety of surface and aviation emissions from different regions. Also shown are COVID-19 lockdown impacts on NO<sub>x</sub> emissions ( $Tg(NO_2) \cdot yr^{-1}$ ) between 2019 and 2020 from Lamboll et al. (2021) for regions and global aviation, and the corresponding contributions to the change in global methane burden ( $Tg(CH_4)$ ; values only given for contributions used in Section 4). References for the sensitivity values: (a) Wild et al. (2001); (b) Derwent et al. (2001); (c) Fry et al. (2012); (d) Stevenson et al. (2004); (e) Lee et al. (2021); and (f) Stevenson and Derwent (2009).**

#### 4 COVID-19 lockdown impacts on NO<sub>x</sub> emissions

Lamboll et al. (2021) compiled estimates of the impact of COVID-19 lockdowns on global anthropogenic NO<sub>x</sub>, CO and NMVOC emissions, as monthly mean time series, with spatial resolution 0.5° latitude by 0.5° longitude. We use these data to calculate the difference in surface NO<sub>x</sub> emissions between 2019 (pre-lockdown) and 2020 for the four HTAP regions, as well as globally, and hence for the 'Rest of the World' (ROW) region (i.e. everywhere beyond the four HTAP regions) for the Northern and Southern Hemispheres. The annual reduction in global surface NO<sub>x</sub> emissions from 2019 to 2020 was about 19.38  $Tg(NO_2)$ , or 15%. Lamboll et al. (2021) also compiled data on aviation emissions, estimating a global reduction of about 0.83  $Tg(NO_2)$ , or 25%. Global and regional annual changes in NO<sub>x</sub>, CO and NMVOC emissions are summarised in Table 1.

	NO <sub>x</sub>	CO	NMVOC
	$Tg(NO_2)$	$Tg(CO)$	$Tg(C)$

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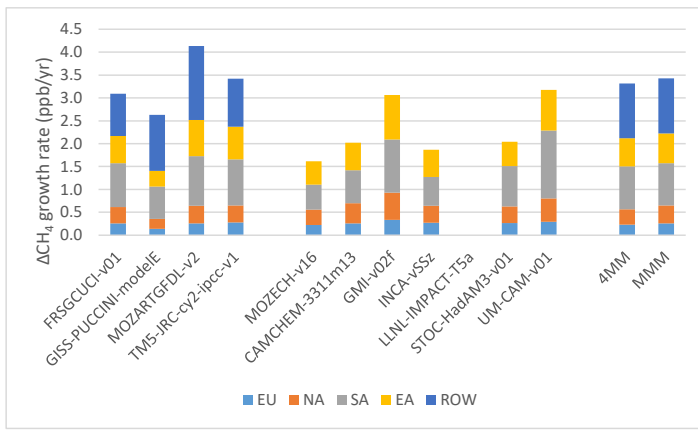
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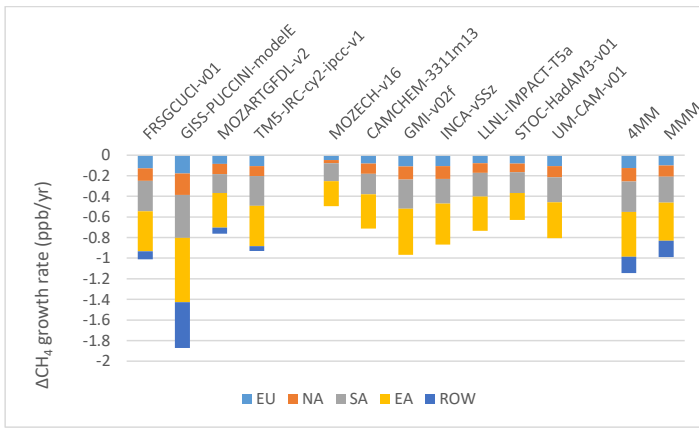
Table 1 shows calculated regional and aviation components of the methane change. We calculate a total methane burden change of 13.36 Tg(CH<sub>4</sub>), comprising 11.44 Tg(CH<sub>4</sub>) from surface NO<sub>x</sub> changes, with a further 1.92 Tg(CH<sub>4</sub>) from aircraft. The more sensitive regions (South Asia, the Southern Hemisphere) and aviation make proportionally larger contributions to the total methane change. We convert the overall change in global methane burden (Tg) to a change in tropospheric mole fraction (ppb) using the total atmosphere mass of 5.113 x 10<sup>27</sup> Tg and a fill factor of 0.973 for conversion of a total atmosphere abundance to a tropospheric abundance (Prather et al., 2012). This yields a global mean increase in tropospheric methane mole fraction of 4.9 ppb associated with the NO<sub>x</sub> reductions. Since the troposphere is well mixed, this is also the change at the surface.



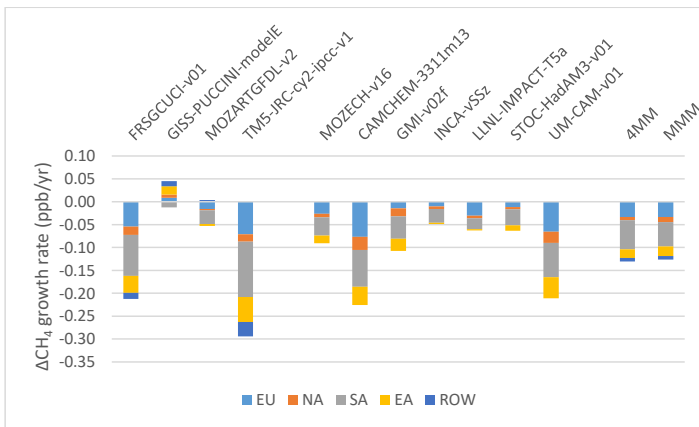
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**Figure 5.** Calculated changes in global methane growth rate from changes in surface NO<sub>x</sub> emissions during the 2020 lockdown, for each of the HTAP models. Also shown are values for the mean of the four core models (shown on left) (4MM) that reported results for all simulations, together with multi-model mean (MMM) results based on all available models.

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**Figure 6.** As Figure 5, but for CO emissions.



**Figure 7.** As Figure 5, but for NMVOC emissions.

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295 [al., 2021; Feng et al., 2022](#)) that will yield further information. There are undoubtedly several other factors, in addition to  
changes in [anthropogenic NOx, CO and NMVOC emissions](#) that influenced methane during 2020. Nevertheless, it seems  
likely that the dramatic reductions in [theseNOx emissions, especially NOx](#), brought about by the COVID-19 lockdowns can  
explain a large component of the surge in methane growth rate seen [duringsince early 2020. These influences on methane](#)  
[related to changes in OH need to be carefully accounted for in any attribution study that attempts to explain the recent observed](#)  
300 [dramatic changes in methane.](#)

#### Author contributions

DSS wrote the text and performed the main analysis. OW and WJC performed additional analysis and commented on the text.  
RGD commented on the text.

#### Competing interests

305 The authors declare that they have no conflict of interest.

#### Code/data availability

Original data used here are all freely available in the cited references.

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#### References

Allen, M. R., Shine, K. P., Fuglested, J. S., Millar, R. J., Cain, M., Frame, D. J., and Macey, A.H.: A solution to the  
misrepresentations of CO<sub>2</sub>-equivalent emissions of short lived climate pollutants under ambitious mitigation. *npj Clim. Atmos.*  
315 *Sci.* 1, 16, <https://doi.org/10.1038/s41612-018-0026-8>, 2018.  
Bauwens, M., Compennolle, S., Stavrakou, T., Müller, J.-F., van Gent, J., Eskes, H., et al.: Impact of coronavirus outbreak on  
NO<sub>2</sub> pollution assessed using TROPOMI and OMI observations. *Geophys. Res. Lett.*, 47, e2020GL087978,  
<https://doi.org/10.1029/2020GL087978>, 2020.

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320 [Cain, M., Lynch, J., Allen, M.R. et al.: Improved calculation of warming equivalent emissions for short lived climate pollutants. npj Clim. Atmos. Sci. 2, 29, <https://doi.org/10.1038/s41612-019-0086-4>, 2019.](#) [Cooper, M.J., Martin, R.V., Hammer, M.S. et al.: Global fine-scale changes in ambient NO<sub>2</sub> during COVID-19 lockdowns. Nature 601, 380–387, <https://doi.org/10.1038/s41586-021-04229-0>, 2022.](#)

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Derwent, R.G., Collins, W.J., Johnson, C.E., and Stevenson, D.S.: Transient Behaviour of Tropospheric Ozone Precursors in a Global 3-D CTM and Their Indirect Greenhouse Effects, Climatic Change 49, 463–487, <https://doi.org/10.1023/A:1010648913655>, 2001.

Field Code Changed

325 [Dlugokencky, E.: NOAA/GML Trends in Atmospheric Methane, \[https://gml.noaa.gov/ccgg/trends\\\_ch4/\]\(https://gml.noaa.gov/ccgg/trends\_ch4/\), accessed 0325/086/20224, 20224.](#)

Field Code Changed

[Dombia, T., C. Granier, Elguindi N., Bouarar I., Darras S., Brasseur G., Gaubert B., Liu Y., Shi X., Stavrou J., Tilmes S., Lacey F., Deroubaix A., Wang T. 2021 Changes in global air pollutant emissions during the Covid-19 pandemic: a dataset for atmospheric chemistry modeling Earth Syst. Sci. Data 4191–4206 <https://doi.org/10.5194/essd-13-4191-2021>, 2021.](#)

[Etminan, M., G. Myhre, E. J. Highwood, and K. P. Shine: Radiative forcing of carbon dioxide, methane, and nitrous oxide: A significant revision of the methane radiative forcing, Geophys. Res. Lett., 43, 12,614–12,623, doi:10.1002/2016GL071930, 2016.](#) [Feng, L., Palmer, P. I., Parker, R. J., Lunt, M. F., and Boesch, H.: Methane emissions responsible for record-breaking atmospheric methane growth rates in 2020 and 2021, Atmos. Chem. Phys. Discuss. \[preprint\], <https://doi.org/10.5194/acp-2022-425>, in review, 2022.](#)

[Fiore, A. M., et al.: Multimodel estimates of intercontinental source-receptor relationships for ozone pollution, J. Geophys. Res., 114, D04301, doi:10.1029/2008JD010816, 2009.](#)

340 [Fry, M. M., et al.: The influence of ozone precursor emissions from four world regions on tropospheric composition and radiative climate forcing, J. Geophys. Res., 117, D07306, doi:10.1029/2011JD017134, 2012.](#)

[Holmes, C. D.: Methane feedback on atmospheric chemistry: Methods, models, and mechanisms. Journal of Advances in Modeling Earth Systems, 10, 1087–1099, <https://doi.org/10.1002/2017MS001196>, 2018.](#)

Field Code Changed

345 [Lamboll, R. D., Jones, C. D., Skeie, R. B., Fiedler, S., Samsel, B. H., Gillett, N. P., Rogelj, J., and Forster, P. M.: Modifying emissions scenario projections to account for the effects of COVID-19: protocol for CovidMIP, Geosci. Model Dev., 14, 3683–3695, <https://doi.org/10.5194/gmd-14-3683-2021>, 2021.](#)

Field Code Changed

350 [Laughner, Joshua L., Jessica L. Neu, David Schimel, Paul O. Wennberg, Kelley Barsanti, Kevin W. Bowman, Abhishek Chatterjee, Bart E. Croes, Helen L. Fitzmaurice, Daven K. Henze, Jinsol Kim, Eric A. Kort, Zhu Liu, Kazuyuki Miyazaki, Alexander J. Turner, Susan Anenberg, Jeremy Avise, Hansen Cao, David Crisp, Joost de Gouw, Annmarie Eldering, John C. Fyfe, Daniel L. Goldberg, Kevin R. Gurney, Sina Hasheminassab, Francesca Hopkins, Cesunica E. Ivey, Dylan B. A. Jones, Junjie Liu, Nicole S. Lovenduski, Randall V. Martin, Galen A. McKinley, Lesley Ott, Benjamin Poulter, Muye Ru, Stanley P. Sander, Neil Swart, Yuk L. Yung, Zhao-Cheng Zeng, Societal shifts due to COVID-19 reveal large-scale complexities and](#)



[feedbacks between atmospheric chemistry and climate change. Proceedings of the National Academy of Sciences Nov 2021, 118 \(46\) e2109481118; DOI: 10.1073/pnas.2109481118, 2021.](#)

Lee, D.S., D.W. Fahey, A. Skowron, M.R. Allen, U. Burkhardt, Q. Chen, S.J. Doherty, S. Freeman, P.M. Forster, J. Fuglestedt, A. Gettelman, R.R. De León, L.L. Lim, M.T. Lund, R.J. Millar, B. Owen, J.E. Penner, G. Pitari, M.J. Prather, R. Sausen, and L.J. Wilcox: The contribution of global aviation to anthropogenic climate forcing for 2000 to 2018, Atmospheric Environment, Volume 244, 117834, <https://doi.org/10.1016/j.atmosenv.2020.117834>, 2021.

Miyazaki, K., K. Bowman, T. Sekiya, M. Takigawa, J. L. Neu, K. Sudo, G. Osterman, and H. Eskes: Global tropospheric ozone responses to reduced NOx emissions linked to the COVID-19 worldwide lockdowns, *Science Advances*, 7, 24, eabf7460 DOI:10.1126/sciadv.abf7460, 2021.

[Myhre, G., E. Highwood, K. Shine, and F. Stordal: New estimates of radiative forcing due to well mixed greenhouse gases. Geophys. Res. Lett., 25\(14\), 2715–2718, doi:10.1029/98GL01908, 1998.](#)

[Myhre, G., D. Shindell, F. M. Bréon, W. Collins, J. Fuglestedt, J. Huang, D. Koch, J. F. Lamarque, D. Lee, B. Mendoza, T. Nakajima, A. Robock, G. Stephens, T. Takemura and H. Zhang: Anthropogenic and Natural Radiative Forcing. In: Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change \[Stocker, T.F., D. Qin, G. K. Plattner, M. Tignor, S.K. Allen, J. Bosechung, A. Nauels, Y. Xia, V. Bex and P.M. Midgley \(eds.\)\]. Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA, 2013.](#)

Prather, M.J.: Lifetimes and Eigenstates in Atmospheric Chemistry, *Geophys. Res. Lett.*, 21, 801-804, 1994.

Prather, M. J., Holmes, C. D., and Hsu, J.: Reactive greenhouse gas scenarios: Systematic exploration of uncertainties and the role of atmospheric chemistry, *Geophys. Res. Lett.*, 39, L09803, doi:10.1029/2012GL051440, 2012.

[Shindell, D.T., G. Faluvegi, N. Bell, and G.A. Schmidt: An emissions-based view of climate forcing by methane and tropospheric ozone. Geophys. Res. Lett., 32, L04803, doi:10.1029/2004GL021900, 2005.](#)

[Shindell, D.T., G. Faluvegi, D.M. Koch, G.A. Schmidt, N. Unger, and S.E. Bauer: Improved attribution of climate forcing to emissions. Science, 326, 716-718, doi:10.1126/science.1174760, 2009.](#)

Stevenson, D.S. and R.G. Derwent: How does the location of aircraft nitrogen oxide emissions affect their climate impact? *Geophys. Res. Lett.*, 36, L17810, doi:10.1029/2009GL039422, 2009.

Stevenson, D. S., Doherty, R. M., Sanderson, M. G., Collins, W. J., Johnson, C. E., and Derwent, R. G.: Radiative forcing from aircraft NOx emissions: Mechanisms and seasonal dependence, *J. Geophys. Res.*, 109, D17307, doi:10.1029/2004JD004759, 2004.

Stevenson, D. S., Young, P. J., Naik, V., Lamarque, J.-F., Shindell, D. T., Voulgarakis, A., Skeie, R. B., Dalsoren, S. B., Myhre, G., Bernsten, T. K., Folberth, G. A., Rumbold, S. T., Collins, W. J., MacKenzie, I. A., Doherty, R. M., Zeng, G., van Noije, T. P. C., Strunk, A., Bergmann, D., Cameron-Smith, P., Plummer, D. A., Strode, S. A., Horowitz, L., Lee, Y. H., Szopa, S., Sudo, K., Nagashima, T., Josse, B., Cionni, I., Righi, M., Eyring, V., Conley, A., Bowman, K. W., Wild, O., and Archibald, A.: Tropospheric ozone changes, radiative forcing and attribution to emissions in the Atmospheric Chemistry and Climate

Field Code Changed

Model Intercomparison Project (ACCMIP), *Atmos. Chem. Phys.*, 13, 3063–3085, <https://doi.org/10.5194/acp-13-3063-2013>, 2013.

[Szopa, S., V. Naik, B. Adhikary, P. Artaxo, T. Bernsten, W.D. Collins, S. Fuzzi, L. Gallardo, A. Kiendler-Scharr, Z. Klimont, H. Liao, N. Unger, and P. Zanis, 2021: Short-Lived Climate Forcers. In \*Climate Change 2021: The Physical Science Basis.\*](#)

390 [Contribution of Working Group I to the Sixth Assessment Report of the Intergovernmental Panel on Climate Change \[Masson-Delmotte, V., P. Zhai, A. Pirani, S.L. Connors, C. Péan, S. Berger, N. Caud, Y. Chen, L. Goldfarb, M.I. Gomis, M. Huang, K. Leitzell, E. Lonnoy, J.B.R. Matthews, T.K. Maycock, T. Waterfield, O. Yelekçi, R. Yu, and B. Zhou \(eds.\)\]. Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA, pp. 817–922, doi:10.1017/9781009157896.008.](#)

395 [Thornhill, G. D., Collins, W. J., Kramer, R. J., Olivić, D., Skeie, R. B., O'Connor, F. M., Abraham, N. L., Checa-Garcia, R., Bauer, S. E., Deushi, M., Emmons, L. K., Forster, P. M., Horowitz, L. W., Johnson, B., Keeble, J., Lamarque, J.-F., Michou, M., Mills, M. J., Mulcahy, J. P., Myhre, G., Nabat, P., Naik, V., Oshima, N., Schulz, M., Smith, C. J., Takemura, T., Tilmes, S., Wu, T., Zeng, G., and Zhang, J.: Effective radiative forcing from emissions of reactive gases and aerosols – a multi-model comparison, \*Atmos. Chem. Phys.\*, 21, 853–874, <https://doi.org/10.5194/acp-21-853-2021>, 2021.](#)

400 Turner, A.J., I. Fung, V. Naik, L. W. Horowitz, R. C. Cohen: Modulation of hydroxyl variability by ENSO in the absence of external forcing, *Proceedings of the National Academy of Sciences*, 115 (36) 8931-8936; DOI: 10.1073/pnas.1807532115, 2018.

~~Vaughan, A.: A mysterious rise in methane levels is sparking global warming fears, *New Scientist*, <https://www.newscientist.com/article/mg25033350-700-a-mysterious-rise-in-methane-levels-is-sparking-global-warming-fears/> [accessed 14<sup>th</sup> July 2021], 2021.~~ [Vasquez, K.: Cleaner pandemic air led to reduced lightning strikes worldwide, \*Eos\*, 103, <https://doi.org/10.1029/2022EO220048>, 2022.](#)

405 Venter, Z.S., K. Aunan, S. Chowdhury, J. Lelieveld: COVID-19 lockdowns cause global air pollution declines, *Proceedings of the National Academy of Sciences*, 117 (32) 18984-18990; DOI: 10.1073/pnas.2006853117, 2020.

Weber, J., Shin, Y. M., Staunton Sykes, J., Archer-Nicholls, S., Abraham, N. L., & Archibald, A. T.: Minimal climate impacts from short-lived climate forcers following emission reductions related to the COVID-19 pandemic. *Geophysical Research Letters*, 47, e2020GL090326, <https://doi.org/10.1029/2020GL090326>, 2020.

410 [Wild, O., Fiore, A. M., Shindell, D. T., Doherty, R. M., Collins, W. J., Dentener, F. J., Schultz, M. G., Gong, S., MacKenzie, I. A., Zeng, G., Hess, P., Duncan, B. N., Bergmann, D. J., Szopa, S., Jonson, J. E., Keating, T. J., and Zuber, A.: Modelling future changes in surface ozone: a parameterized approach, \*Atmos. Chem. Phys.\*, 12, 2037–2054, <https://doi.org/10.5194/acp-12-2037-2012>, 2012.](#)

415 Wild, O., Prather, M.J., and Akimoto, H.: Indirect long-term global radiative cooling from NO<sub>x</sub> emissions, *Geophys. Res. Lett.*, 28, 1719-1722, 2001.

World Meteorological Organization: State of the Global Climate 2020 (WMO-No. 1264), <https://public.wmo.int/en/our-mandate/climate/wmo-statement-state-of-global-climate>, accessed 14-07-21, 2021.

Field Code Changed

Field Code Changed

Field Code Changed

420 Zhao, Y., Sauniois, M., Bousquet, P., Lin, X., Berchet, A., Hegglin, M. I., Canadell, J. G., Jackson, R. B., Deushi, M., Jöckel, P., Kinnison, D., Kirner, O., Strode, S., Tilmes, S., Dlugokencky, E. J., and Zheng, B.: On the role of trend and variability in the hydroxyl radical (OH) in the global methane budget, *Atmos. Chem. Phys.*, 20, 13011–13022, <https://doi.org/10.5194/acp-20-13011-2020>, 2020.

Field Code Changed