# A Novel Pathway of Atmospheric Sulfate Formation Through Carbonate Radical

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**Abstract.** Carbon dioxide is considered an inert gas that rarely participates in atmospheric chemical reactions. However, we show here that  $CO_2$  is involved in some important photo-oxidation reactions in the atmosphere through the formation of carbonate radicals ( $CO_3^{-1}$ ). This potentially active intermediate  $CO_3^{-1}$  is routinely overlooked in atmospheric chemistry regarding its effect on sulfate formation. Present work demonstrates that  $SO_2$  uptake coefficient is enhanced by 17 times on mineral dust particles driven by  $CO_3^{-1}$ . It can be produced through two routes over mineral dust surfaces: i) hydroxyl radical  $+ CO_3^{-2}$ ; ii) holes  $(h^+) + CO_3^{-2}$ . Importantly, upon irradiation mineral dust particles are able to produce gas-phase carbonate radical ions when the atmospherically relevant concentration of  $CO_2$  presents, therefore potentially promoting external sulfate aerosol formation and oxidative potential in the atmosphere. Employing a suite of laboratory investigations of sulfate formation in the presence of carbonate radical on the model and authentic dust particles, ground-based field measurements of sulfate and (bi)carbonate ions within ambient PM, together with density functional theory (DFT) calculations for single electron transfer processes in terms of  $CO_3^{-1}$ -initiated S(IV) oxidation, a novel role of carbonate radical in atmospheric chemistry is elucidated.

#### 1. Introduction

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Atmospheric composition changes are subjected to highly reactive light-induced radicals, such as hydroxyl ( $\cdot$ OH), hydroperoxyl ( $\text{HO}_2\cdot$ ), or nitrate radicals ( $\text{NO}_3\cdot$ ), which are able to alter not only compositions but also physical and chemical properties of particulate matter (Davis and Francisco, 2011; Platt et al., 1990; Prinn et al., 2001; Thompson, 1992). However, when atmospheric chemical reactions occur over nanometer-sized particles at ambient conditions, which creates a locally enriched aqueous medium of unique chemical activity, other radicals might likewise gain importance. The carbonate radical

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(CO<sub>3</sub>) is typically such an active radical. The lifetime of CO<sub>3</sub> ranges from a microsecond to even a few milliseconds and its concentration can be two orders of magnitude higher than that of hydroxyl radicals over the water surface (Chandrasekaran and Thomas, 1983; Goldstein et al., 2001; Shafirovich et al., 2001; Sulzberger et al., 1997). In addition, the one-electron reduction potential of  $E^0(CO_3^-/CO_3^-)$  couple is 1.78 V vs. NHE at neutral pH, leaving  $CO_3^-$  a strong oxidant in aquatic chemistry (Bisby et al., 1998; Cope et al., 1973; Merouani et al., 2010). Previous studies concerning carbonate radical in aqueous media demonstrate that it reacts rapidly with some organic compounds with higher selectivity (Merouani et al., 2010), especially for those electron-rich compounds amines (Stenman et al., 2003; Yan et al., 2019). Also, it has been pointed out that the scavenging of hydroxyl radicals by (bi)carbonate species leads to the formation of CO<sub>3</sub> ions (Graedel and Weschler, 1981), which promotes the degradation of phenol (Xiong et al., 2016). Besides, a higher second order of rate constant, lying at 109 M<sup>-1</sup> s<sup>-1</sup>, has been reported for the reaction of CO<sub>3</sub> with porphyrins (Ferrer-Sueta et al., 2003), indicating that this radical ion has great oxidation capability that may trigger atmospherically relevant chemical reactions. However, it is only regarded as a marginal intermediate in tropospheric anion chemistry so far (Beig and Brasseur, 2000: Dotan et al., 1977; Graedel and Weschler, 1981; Lehtipalo et al., 2016) and its underlying role as an active oxidant for heterogeneous reaction in the atmosphere is barely explored. Very recently, our group observed the promotional effect of CO<sub>3</sub> on atmospheric nitrate formation (Fang et al., 2021). Motivated by this finding, attempts were made to further explore its role in other important atmospherically-relevant reactions.

It is well documented that sulfate (SO<sub>4</sub><sup>2-</sup>) is also a key constituent of aerosols in the atmosphere (Huang et al., 2015; Su et al., 2016). It is able to serve as the precursors of efficient cloud condensation nuclei, with optical properties leading to a cooling effect (Wang et al., 2011). As a consequence, the mechanism aspect of secondary sulfate formation was the focus of numerous studies over the past decades (Hung et al., 2018; Stone, 2002; Zhang et al., 2015b). There is a consensus that high-valence sulfur (VI), produced from the oxidation of anthropogenic SO<sub>2</sub>, is the dominant source for atmospheric secondary sulfate. However, a remarkable missing sulfate budget emerges for the atmospheric modeling, which underpredicts SO<sub>4</sub><sup>2-</sup> by over 50 % (normalized mean bias) with respect to observational results when heterogeneous aerosol chemistry is not considered (Zheng et al., 2015). This indicates that the heterogeneous sulfate production pathway is a crucial process and exploring the unconsidered heterogeneous mechanism is very likely to narrow the gap between observations in lab studies, field measurements, and numerical modelings. However, due to the missing chemical mechanism that initiated fast SO<sub>2</sub> oxidation, atmospheric models fail to capture the key feature of atmospheric observations of high sulfate production (Wang et al., 2020). Consequently, there are unknown heterogeneous reaction pathways of significance and previously unconsidered promoters that have great potential to accelerate sulfate formation.

Due to the high stability of CO<sub>2</sub> under ambient conditions (Hossain et al., 2020), there are rare studies concerning the influence of CO<sub>2</sub> in atmospheric chemical processes (Deng et al., 2020; Liu et al., 2020a; Xia et al., 2021). CO<sub>2</sub> is demonstrated to form (bi)carbonate species over humidified dust particles (Baltrusaitis et al., 2011; Nanayakkara et al., 2014) and reduced to CO under solar illumination (Deng et al., 2020). However, its impact on atmospheric heterogeneous reactions remains poorly characterized. Our earlier laboratory study shows that CO<sub>2</sub> decreases the sulfate formation on aluminum

oxide particles in the dark (Liu et al., 2020b) while upon solar illumination its role in SO<sub>2</sub> oxidation over mineral dust surfaces is still an open question. In addition, carbonate salt is enriched in authentic dust aerosol (Cao et al., 2005) and reported to reach over 10 % wt. of Asian dust particles (McNaughton et al., 2009). It is generally accepted that CO<sub>3</sub><sup>2-</sup> affects atmospheric chemistry and aerosol characteristics mainly through its intrinsic alkalinity, which buffers aerosol acidity and favors the sulfate formation (Bao et al., 2010; Kerminen et al., 2001; Yu et al., 2018). In fact, either CO<sub>2</sub> or carbonate salt is able to produce the active CO<sub>3</sub><sup>-</sup> under the ambient circumstance and increase the oxidative capacity in the atmosphere. Combined with our previous investigation of CO<sub>3</sub><sup>-</sup> (Fang et al., 2021), this radical ion is likely to be a driving force for fast SO<sub>2</sub> oxidation. However, to the best of our knowledge, no work has considered how and to what extent the carbonate radical influences SO<sub>2</sub> heterogeneous oxidation in the atmosphere.

In the current study, through laboratory studies, we presented that carbon dioxide and calcium carbonate, working as the precursor of carbonate radicals, extend their ability to accelerate sulfate formation over authentic particles in the atmosphere. Together with quantum chemistry calculations, a detailed molecular mechanism regarding a single electron transfer (SET) process between carbonate radical and sulfite ions is elucidated. Furthermore, ground-based observations validate some findings from the laboratory-based simulations.

# 80 2. Experimental methods

## 2.1 Laboratory Studies

A series of characterizations were initially performed to investigate the mineral dust of concern by using X-ray diffraction (XRD) and Raman technique. The heterogeneous reaction of SO<sub>2</sub> on mineral dust particles in the presence of CO<sub>2</sub> and carbonate species were then investigated by the *in situ* Fourier transform spectrum (DRIFTS), ion chromatography (IC), Raman, electron spin resonance (ESR), and nanosecond transient absorption spectroscopy (NTAS). Furthermore, we employed three types of authentic dust particles and four kinds of synthesized authentic simulants to probe the proposed scheme. Besides, the steady concentration of CO<sub>3</sub><sup>-1</sup> ions in each reaction system was determined by High-Performance Liquid Chromatography (HPLC) using probe molecule aniline. All corresponding configuration setup, characterizations, and methodologies can be found in Supplement text 1, 2, 8-18, and Tables S1-S3, and Fig. S1-S12 and Fig. S15-S17. In terms of oxygen isotope experiments, more detailed procedures are available in Supplement, text 21.

## 2.2 Quantum Chemical Calculation

We employed density functional theory (DFT) calculations in the term of the single electron transfer (SET) process using Gaussian 09 package to investigate this novel route, detailed in Supplement text 19-20 and Fig. S13-S14.

## 2.3 Field Observations

Atmospheric aerosols were collected on the roof of our department using an 8-stage non-viable-cascade-impactor type sampler (TISCH TE Inc., USA), with sampling details shown in supplement Text 22 and daily variations of wind scales shown in Fig. S18. The concentration of water-soluble sulfate ion was by measured IC analysis while that of (bi)carbonate ions was determined by the ionization balance approach (Supplement text 23-24). The relationship between (bi)carbonate ions and sulfate ions during the daytime and nighttime hours were then determined.

## 3. Results and discussion

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#### 3.2.1 Accelerated sulfate production in the presence of carbonate.

The physico-chemical properties of employed mineral dust proxies were first characterized (Fig. S1), consistent with earlier studies (Balachandran and Eror, 1982; Shang et al., 2010; Su et al., 2008), and spectral irradiance of the solar simulator applied in the present study is well covered by natural sunlight (Fig. S2), as much as possible having experimental results from the lab simulate the real atmosphere. Upon irradiation, the sulfate yield on TiO<sub>2</sub>-CaCO<sub>3</sub> mixture particles (50 wt. % CaCO<sub>3</sub>), measured by IC, is significantly enhanced by 7 times and 23 times compared to that of pristine TiO<sub>2</sub> and CaCO<sub>3</sub> (Fig. 1a), respectively. In stark contrast, there is a negligible increase of sulfate production detected on the TiO<sub>2</sub>-CaCO<sub>3</sub> mixture relative to that of pristine CaCO<sub>3</sub> and TiO<sub>2</sub> in dark experiments (Fig. 1a). While great discrepancies in sulfate yield between dark and irradiation experiments, it remains unclear for the role of carbonate salt in promoting sulfate formation. There is a prevailing view that neutralization of H<sub>2</sub>SO<sub>4</sub> accounts for rapid SO<sub>2</sub> oxidation over carbonate salt particles, which needs careful consideration. Following this speculation, two types of mixtures TiO<sub>2</sub>-CaCO<sub>3</sub> and TiO<sub>2</sub>-CaO were employed. In the dark condition (Fig. S3), both TiO<sub>2</sub>-CaO and TiO<sub>2</sub>-CaCO<sub>3</sub> almost yield an identical concentration of sulfite and sulfate as they are likely to present similar physical and chemical properties, e.g. surface pH and neutralization capability. Once irradiated, TiO<sub>2</sub>-CaCO<sub>3</sub> particles produce nearly two times of sulfate than TiO<sub>2</sub>-CaO particles, along with a sharp decrease of S(IV) species on the surface of TiO<sub>2</sub>-CaCO<sub>3</sub> surfaces (see additional discussion in the supplementary text 2). These results allow us to assert that the carbonate-containing system contains another important mechanism for sulfate generation beyond the production of an alkaline environment. Fig. 1b and 1c show the *in situ* diffuse reflectance infrared DRIFTS features of S(IV) and S(VI) species formed on theoretical and experimental TiO<sub>2</sub>-CaCO<sub>3</sub> mixtures (wt./wt. = 50/50) upon irradiation for 90 min, respectively. The "theoretical" here is calculated based on the in situ DRIFTS of pristine TiO<sub>2</sub> and CaCO<sub>3</sub> through a simple linear superposition. These results suggest a synergistic effect presented in this mixture for sulfate formation under solar irradiation. In addition, a more evident fingerprint SO<sub>4</sub><sup>2</sup>- feature (Dong et al., 2009; Yann Batonneau et al., 2008) monitored by Raman spectroscopy appears over TiO<sub>2</sub>-CaCO<sub>3</sub> particles compared to pristine TiO<sub>2</sub> particles (Fig. S4), in good agreement with the IC analyses and in situ DRIFTS measurements. Combining DRIFTS experiments with the obtained calibration curve (Fig. S5), we estimated that the uptake coefficient of TiO<sub>2</sub>-CaCO<sub>3</sub> mixture (50 wt. % CaCO<sub>3</sub>) is increased by about 17 times as compared to that of pure CaCO<sub>3</sub> or TiO<sub>2</sub> (Table S3). More importantly, upon irradiation SO<sub>2</sub> uptake coefficients for these dust proxies lie at the order of magnitudes of 10<sup>-4</sup>, indicating that the photochemical pathway associated with carbonate species is likely a potential driving force to trigger fast SO<sub>2</sub> oxidation in the atmosphere.

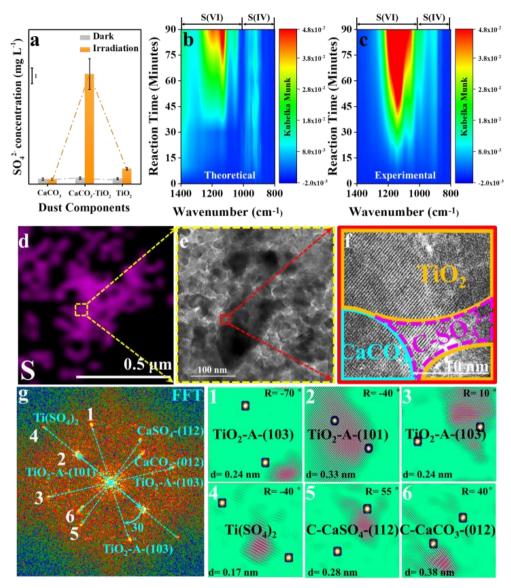


Figure 1. (a) Sulfate concentration quantified by IC on mineral dust particles after exposure to gaseous SO<sub>2</sub> under irradiation or dark for 30 min. *In situ* DRIFTS of S(IV) and S(VI) species yield on theoretical (b) and (c) experimental TiO<sub>2</sub>-CaCO<sub>3</sub> mixtures (wt./wt. = 50/50) upon irradiation for 90 min. Reaction conditions: RH = 30 %, Light intensity (I) = 30 mW cm<sup>-2</sup>, Total flow rate = 52.5 mL min<sup>-1</sup> and SO<sub>2</sub> = 2.21×10<sup>14</sup> molecules cm<sup>-3</sup>. All spectra were processed by the Kubelka-Munk (K-M) algorithm. Noting that the production of sulfur species in theoretical TiO<sub>2</sub>+CaCO<sub>3</sub> mixtures refer to 0.5 × K-M bands of sulfur species of TiO<sub>2</sub> + 0.5 × K-M bands of sulfur species of CaCO<sub>3</sub> while that for experimental TiO<sub>2</sub>+CaCO<sub>3</sub> mixtures refer to 1 × K-M bands of sulfur species of TiO<sub>2</sub>+CaCO<sub>3</sub> mixtures (wt./wt. = 50/50). (d)

Energy Dispersive Spectroscopy (EDS) mapping of sulfur. (e) Selected HRTEM region containing a high density of sulfur for further observation and the red rectangle refers to the region shown in panel f. (f) The HRTEM image in high resolution with lattice fringes and (g) corresponding FFT power spectra, lattice indexing, and (1-6) inverse FFT analysis of lattice signal shown in panel g. In panel f, the term

 $C-SO_4^{2-}$  stands for crystalline  $SO_4^{2-}$ , i.e.  $CaSO_4$  and  $Ti(SO_4)_2$ . Particles for the HRTEM measurement refer to  $TiO_2-CaCO_3$  mixture particles upon exposure to the  $4.42\times10^{14}$  molecules cm<sup>-3</sup>  $SO_2/N_2+O_2$  for 60 min while other reaction conditions are as same as that of above sulfate quantification experiments.

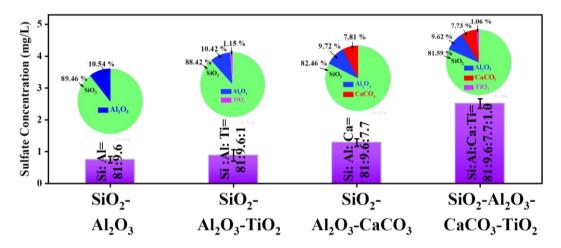
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High-resolution transmission electron microscopy (HRTEM) analysis of TiO<sub>2</sub>-CaCO<sub>3</sub> particles after reaction, in combination with energy dispersive spectrometer mapping measurements of sulfur component, were conducted to investigate the synergistic effect between TiO<sub>2</sub> and carbonate ions (Fig. 1 d-f and Fig. S6). A region with a relatively high density of sulfur species was selected for further observation and the distribution of each component (Fig. 1g) was determined by fast Fourier transformation (FFT) and inverse FFT analyses of the selected HRTEM image in high resolution with lattice fringes shown in Fig. 1f. Observation of crystalline Ti(SO<sub>4</sub>)<sub>2</sub> and CaSO<sub>4</sub> on the interface of TiO<sub>2</sub> and CaCO<sub>3</sub> components imply that the synergistic effect on sulfate production likely originates from interplays of those two types of components under solar illumination. We further assessed the importance of interfacial contact between TiO<sub>2</sub> and CaCO<sub>3</sub> in sulfate production by two synthesis approaches in which the interface abundance is modulated for comparison. Typically, a "grinding" method was used to make TiO<sub>2</sub>-CaCO<sub>3</sub> mixture with compact contact between those two components, thus leading to a strong interaction. Meanwhile, the "shaking" method is designed to create a TiO<sub>2</sub>-CaCO<sub>3</sub> mixture with weak interplay, leaving relatively fewer amounts of interfaces within the mixtures. The resulting mixing statuses of two samples meet our expectations, evidenced by the scanning electron microscope (SEM) technique (Fig. S7). IC quantification analysis suggests that particles with considerable junctions exhibit a more pronounced promotion for sulfate yield than those having relatively few junctions (Fig. S8). These results emphasize the importance of an indispensable interface connection between TiO<sub>2</sub> and CaCO<sub>3</sub> in fast production upon irradiation.



**Figure 2.** Sulfate concentration quantified by IC. Sulfate concentration was measured by IC on mineral dust simulants after exposure to gaseous SO<sub>2</sub> (2.46×10<sup>14</sup> molecule cm<sup>-3</sup>) under irradiation. Noting that SiO<sub>2</sub>: Al<sub>2</sub>O<sub>3</sub>: CaCO<sub>3</sub>: TiO<sub>2</sub> refers to the mass fraction ratios of the components in simulants. Experiments were all conducted at RH of 30 % and Light intensity (*I*) of 30 mW cm<sup>-2</sup>.

The rapid SO<sub>2</sub> oxidation pathway was further probed by employing mineral dust simulants where two dominant crust constituents SiO<sub>2</sub> and Al<sub>2</sub>O<sub>3</sub> were introduced into TiO<sub>2</sub>-CaCO<sub>3</sub> particles to mimic the authentic mineral dust particles in the atmosphere, with specific component and corresponding ratio information shown in Table S1. It is worth mentioning that the determination of the ratio of each component in the simulants relies on the EDS mapping results of ATD particles. In Fig. 2, the introduction of TiO<sub>2</sub> components ( $\approx 1$  % wt.) into SiO<sub>2</sub>-Al<sub>2</sub>O<sub>3</sub> leads to 81.6 % enhancement of sulfate production while merely 24.8 % wt. increase of sulfate yield was observed once  $\approx 8$  % wt. of CaCO<sub>3</sub> was incorporated into SiO<sub>2</sub>-Al<sub>2</sub>O<sub>3</sub> dust particles. Surprisingly, mixing of  $\approx 1$  % mass fraction of TiO<sub>2</sub> and  $\approx 8$  % wt. of CaCO<sub>3</sub> into SiO<sub>2</sub>-Al<sub>2</sub>O<sub>3</sub> gives rise to a 235 % increase of sulfate formation relative to that of SiO<sub>2</sub>-Al<sub>2</sub>O<sub>3</sub>. Hence, the observed synergistic effect on heterogeneous oxidation of SO<sub>2</sub> is likely to take effect in the atmosphere.

Fe<sub>2</sub>O<sub>3</sub> is also one of the crucial components found in authentic mineral dust (El Zein et al., 2013), and it has been reported to produce ROS under solar irradiation (Li et al., 2019), thus likely involving the reaction mechanism proposed in this work. Similar to experiments using TiO<sub>2</sub>+CaCO<sub>3</sub> mixture, alpha-Fe<sub>2</sub>O<sub>3</sub>+CaCO<sub>3</sub> are prepared by grinding alpha-Fe<sub>2</sub>O<sub>3</sub> and CaCO<sub>3</sub>. In Fig. S9a, our results show that alpha-Fe<sub>2</sub>O<sub>3</sub> can not trigger fast SO<sub>2</sub> oxidation in the presence of carbonate ions upon irradiation, which is distinguished from results we derived from TiO<sub>2</sub>+CaCO<sub>3</sub> mixture. This can be explained by the fact that Fe<sub>2</sub>O<sub>3</sub> shows a lower redox activity relative to TiO<sub>2</sub> (Fig. S9b), where its strong redox capability essentially enables photo-induced electrons and holes to produce O<sub>2</sub>- and ·OH radical ions. In stark contrast, the valence band and conduct band of Fe<sub>2</sub>O<sub>3</sub> lie at -0.18 and at 1.68 V vs. NHE (pH = 7), lower than the redox potential required for generating O<sub>2</sub>-, ·OH as well as CO<sub>3</sub>- (Li et al., 2016). Hence, no promoted sulfate production is seen for Fe<sub>2</sub>O<sub>3</sub>+CaCO<sub>3</sub> particles under irradiation. More discussion on the inconsistency between our study and the previous results regarding the response of SO<sub>2</sub> oxidation to solar irradiation can be found in supplementary text 3.

Overall, we show that upon irradiation atmospherically relevant content of TiO<sub>2</sub> (nearly 1 %) found in authentic dust simulants is able to interact with carbonate ions to launch a fast SO<sub>2</sub> oxidation channel, which is beyond the conventional regime of alkaline neutralization of H<sub>2</sub>SO<sub>4</sub>. Unlike TiO<sub>2</sub>, alpha- Fe<sub>2</sub>O<sub>3</sub> lacks the ability to initiate fast SO<sub>2</sub> oxidation by generating CO<sub>3</sub>. due to its limited photo-chemical activity although ferric chemistry is important in secondary sulfate formation in the atmosphere (Sullivan et al., 2007; Yermakov and Purmal, 2003).

### 3.2.2 Accelerated sulfate production in the presence of CO<sub>2</sub>.

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Atmospheric CO<sub>2</sub> is also an important source of (bi)carbonate. Its influence on photochemical SO<sub>2</sub> uptake on mineral dust was thus studied. In the presence of atmospherically relevant CO<sub>2</sub> (9.83×10<sup>15</sup> molecules cm<sup>-3</sup>), sulfate yield was increased under irradiation as compared to CO<sub>2</sub>-free case (Fig. 3a and b). We cautiously examined the buffering effect of formed (bi)carbonate on sulfate production by time-resolved DRIFTS spectra (Fig. 3c and d). CO<sub>2</sub> suppresses both S(IV) and S(VI) products under the dark. This observation implies that (bi)carbonate ions that evolve active intermediates upon irradiation

may be a plausible force to drive rapid sulfate formation rather than accumulating much more sulfur species through buffering effect. The reaction kinetics of SO<sub>2</sub> on mineral dust particles follows the pseudo-first-order, as evidenced by the SO<sub>2</sub> concentration dependence experiments (Fig. S10 a-f). Besides, a nearly 50 % increase of SO<sub>2</sub> uptake coefficient is observed for the mineral dust proxy TiO<sub>2</sub> after being exposed to 9.83×10<sup>15</sup> molecules cm<sup>-3</sup> (400 ppm) CO<sub>2</sub>+SO<sub>2</sub>/N<sub>2</sub>+O<sub>2</sub> mixture.

As another step toward a real scenario in the atmosphere, experimental trials employing authentic dust particles, i.e. Arizona test dust (ATD), clays IMt-2 (Illite, Mont., USA), and K-Ga-2 (Kaolin, Georgia, USA), were implemented (Table S2). In Fig. 4, K-Ga-2 clay exhibits the most marked promotional effect on sulfate yield (by nearly 100 % increased sulfate production in the CO<sub>2</sub>-involved case under irradiation). This correlates with its considerable TiO<sub>2</sub> contents (3.43 %) in the K-Ga-2 clay, in which active intermediates are readily evolved from TiO<sub>2</sub> and (bi)carbonate species upon irradiation. However, the promotional effect of CO<sub>2</sub> on sulfate production under irradiation is weak for IMt-2 (the content of TiO<sub>2</sub>  $\approx$  0.99 %) and ATD (the content of TiO<sub>2</sub>  $\approx$  0.46 %) as compared to K-Ga-2 particles. This may correlate to their higher mass fraction of alkaline earth metal oxide (denoted as A.E.), which enables dust particles to possess a large number of (bi)carbonate species in the natural environment where they have experienced long-term exposure to atmospheric CO<sub>2</sub> during the regional transport. Therefore, the aforementioned synergetic effect takes effect over IMt-2 and ATD particles even without exposure to CO<sub>2</sub> due to the presence of abundant carbonate formed, and a less evident increase of sulfate yield was thus observed.

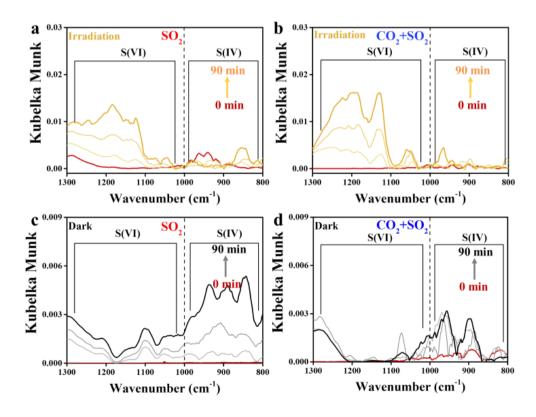
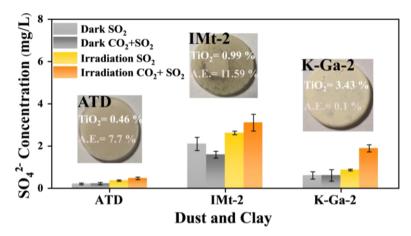


Figure 3. Time-resolved DRIFTS of S(IV) and S(VI) products over TiO<sub>2</sub> particles after exposure to  $SO_2/N_2+O_2$  in the absence and presence of  $CO_2$  upon irradiation (**a** and **b**) and those reactions under dark (**c** and **d**). Reaction conditions: RH = 30 %, Light intensity (I) = 30 mW cm<sup>-2</sup>, total flow rate = 52.5 mL min<sup>-1</sup> and  $SO_2 = 7.37 \times 10^{13}$  molecules cm<sup>-3</sup>.



**Figure 4.** Laboratory studies of sulfate production on authentic dust and clay membranes under the dark and irradiation (30 mW cm<sup>-2</sup>) upon exposure to  $4.91 \times 10^{14}$  molecules cm<sup>-3</sup> SO<sub>2</sub>/N<sub>2</sub>+O<sub>2</sub> and  $2.46 \times 10^{18}$  molecules cm<sup>-3</sup> CO<sub>2</sub>+  $4.91 \times 10^{14}$  molecules cm<sup>-3</sup> SO<sub>2</sub>/N<sub>2</sub>+O<sub>2</sub> at RH of 30 % (total flow rate = 100 mL min<sup>-1</sup>).

#### 3.2.3 Reaction Mechanism.

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The heterogeneous reaction of SO<sub>2</sub> on dust particles in the atmosphere is a complicated process, covering a series of reactions taking place through both homogeneous and heterogeneous ways. At a sufficiently low RH condition (normally below 10 % RH), water readily dissociates on the surface of metal oxide under ambient atmospheric conditions, where metal oxide surface is terminated by hydroxyl groups that hydrogen bond to adsorbed water molecules (Cwiertny et al., 2008). In this case, SO<sub>2</sub> oxidation over dust particles is dominated by the heterogenous pathway, where the resulting hydroxyl groups capture SO<sub>2</sub> in the gas phase first and then stabilize it as adsorbed S(IV)<sub>ad</sub>. Afterward, S(IV)<sub>ad</sub> will be oxidized by oxidants in the atmosphere or photo-induced active intermediates produced from the dust surface upon irradiation. As the RH increases beyond 10 % -15 %, multilayer water coverage occurs, reaching approximately two monolayers at RH of 30 % (Mogili et al., 2006). Under these circumstances, the amount of water adsorbed onto the surface of the dust particles is believed to be sufficiently large that it is liquid-like in its physical and chemical properties (Cwiertny et al., 2008) (Peters and Ewing, 1997). In this work, heterogenous SO<sub>2</sub> oxidation over mineral dust proxies proceeds at the RH of 30 %, and two water layers absorb on dust particles. Thus, radical ions are expected to play a key role in fast SO<sub>2</sub> oxidation and mechanism studies performed in solution phase are persuasive to some extent.

Our preliminary sulfate quantification results (Fig. 1-4) suggest that the presence of (bi)carbonate ions under solar light contributes to increased sulfate yield. In our carbonate-containing reaction system, a plausible intermediate is the active

carbonate radical. They are readily produced via the following two pathways. First of all, carbonate anion can be directly oxidized by produced photo-induced holes (Eqs 2 and 3), as the redox potential of  $CO_3^-/CO_3^{2-}$  is 1.78 V (vs NHE, at pH = 7), which is lower than the  $TiO_2$  valence band (VB) potential of 2.67 V (vs NHE, at pH = 7) (Li et al., 2016; Xiong et al., 2016):

Mineral dust 
$$+hv \rightarrow h^+ + e^-$$
 (Eq. 1)

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$$h^{+} + CO_3^{2-} \rightarrow CO_3^{-}$$
 (Eq. 2)

In the second pathway, carbonate radicals evolve through the reaction of (bi)carbonate anion with formed hydroxyl radicals ·OH over mineral dust surfaces (Zhang et al., 2015a) (Eqs 3 and 4).

$$h^+ + H_2O_{ad} \rightarrow \cdot OH + H^+$$
 (Eq. 3)

$$\cdot OH + HCO_3 \cdot /CO_3 \cdot - CO_3 \cdot + H_2O/OH \cdot$$
 (Eq. 4)

245 The above assumptions are supported by nanosecond transient absorption spectra (NTAS), in which signal ( $\Delta$ OD) of carbonate radical CO<sub>3</sub> at 600 nm (Bhattacharya et al., 1998) only emerged for dust suspension containing (bi)carbonate species (Fig. 5a). A promoted degradation of aniline in TiO<sub>2</sub> suspension due to the presence of carbonate ions presents additional evidence of the formation of active CO<sub>3</sub> ions and strengthened oxidation capability of TiO<sub>2</sub> (Fig. S11, see additional discussion in supplementary text 4). The CO<sub>3</sub> -induced chemistry was further evidenced by OH scavenging 250 experiments using tertiary Butyl Alcohol (TBA) and isopropanol (i-PrOH) as they show lower reaction rates with CO<sub>3</sub>  $(k_{\rm CO_3^{--},i\text{-}PrOH} < 4.0 \times 10^4 \ \rm M^{-1} \ s^{-1} \ and \ k_{\rm CO_3^{--},TBA} < 1.6 \times 10^2 \ \rm M^{-1} \ s^{-1}) \ relative \ to \ that \ with \ \cdot OH \ (k_{i\text{-}PrOH} \sim \rm OH} < 1.9 \times 10^9 \ \rm M^{-1} \ s^{-1} \ and \ k_{\rm CO_3^{--},TBA} < 1.6 \times 10^2 \ \rm M^{-1} \ s^{-1})$  $(k_{\text{-OH,TBA}} = 6 \times 10^8 \, \text{M}^{-1} \, \text{s}^{-1})$  (Buxton et al., 2009; Liu et al., 2015) (Liu et al., 2015) (Li et al., 2020). Tertiary Butyl Alcohol (TBA) sharply decreases the yield of sulfate on TiO<sub>2</sub> surface by nearly 70 %, with sulfite ions being the dominant sulfur species (Fig. S12). Meanwhile, a significant loss of sulfate yield when TiO<sub>2</sub> suspension was added with i-PrOH (Fig. 5b). 255 This is in strong contrast to the result of a carbonate-involved system where the reactivity is sustained, i.e., carbonate radicals offer an alternative reaction pathway for SO<sub>2</sub> oxidation. This is plausible since the carbonate ions are excellent OH scavenger, and CO<sub>3</sub> becomes predominant species at a relatively strong alkaline aqueous-like environment in the presence of carbonate salt. This is supported by the previous work (Sun et al., 2016), in which adding 0.1 M of NaHCO<sub>3</sub> into the  $UV/H_2O_2$  system ( $H_2O_2 = 0.3$  mM) were sufficient to suppress ·OH concentration to around  $10^{-15}$  M, creating a carbonate radical dominated reaction system ([CO<sub>3</sub>-] =  $8.64 \times 10^{-12}$  M). In our supplementary experiments (Fig. S11), 0.2 M of 260 carbonate salt was employed, and the reaction rate of CO<sub>3</sub><sup>2-</sup> with ·OH is nearly two orders of magnitude higher than that of HCO<sub>3</sub>, thus giving rise to carbonate radical being the substitute of hydroxyl radical in the reaction. The above results suggest that OH is a major contributor to sulfate yield on TiO<sub>2</sub> particles in the absence of carbonate ions while CO<sub>3</sub> ions dominate SO<sub>2</sub> oxidation over carbonate-containing dust particles upon irradiation. In addition to experimental investigations, the 265 carbonate radical formation process is proved to be thermodynamically favorable, supported by density functional theory (DFT) calculations (Fig. S13).

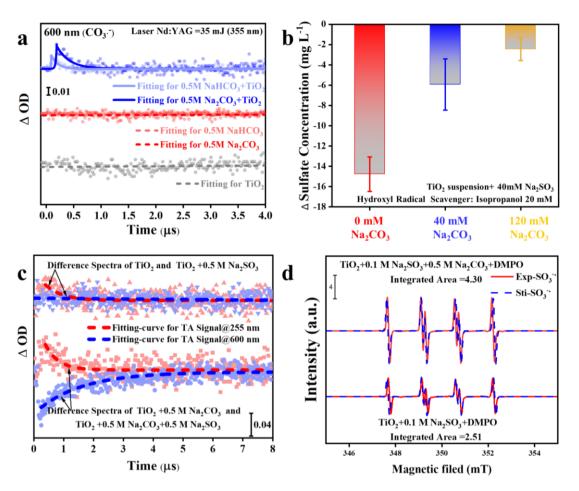


Figure 5. (a) Single-wavelength transient absorption spectra of various aqueous solutions. (b) Sulfate formation change Δ(SO<sub>4</sub><sup>2-</sup>) determined by different sulfate concentrations with and without the addition of isopropanol as hydroxyl radical scavenger. (c) The difference in transient absorption kinetics of sulfite radical and carbonate radical at the various aqueous solutions and their corresponding growth-decay fit curves. ΔA-signal was recorded at 255 and 600 nm after pulsed 355 nm laser excitation. (d) ESR spectrometry of [DMPO–SO<sub>3</sub><sup>-</sup>] intermediate formed in a solution of d TiO<sub>2</sub> (3mg~4 mL) + 0.1 M Na<sub>2</sub>SO<sub>3</sub> and TiO<sub>2</sub> (3mg~4 mL) + 0.5 M Na<sub>2</sub>CO<sub>3</sub> + 0.1 M Na<sub>2</sub>SO<sub>3</sub>. For clarity, the integrated areas of ESR profiles were also presented for direct comparison. Exp. and Sti. stand for experimental results and corresponding fitting results using software Isotropic Radicals.

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On the other hand, the previous studies (Chameides and Davis, 1982; Das, 2001; Neta and Huie, 1985) agree with the key role of sulfite radical (SO<sub>3</sub><sup>-</sup>) in rapid sulfate production in an aqueous medium, and the present reaction system creates a localized environment where SO<sub>3</sub><sup>-</sup> can be readily produced from the TiO<sub>2</sub> and S(IV) species upon solar illumination (Salama et al., 1995). Consequently, probe light of NTAS at wavelength 255 nm (ascribed to sulfite radical) and 600 nm (ascribed to carbonate radical) were simultaneously monitored (Ghalei et al., 2016; Goldstein et al., 2001; Hayon et al., 1972). A weak signal of sulfite radical was observed in the system of TiO<sub>2</sub>+Na<sub>2</sub>SO<sub>3</sub> suspension under irradiation (Fig. 5c). On the contrary, the sulfite radical signal is strengthened after the introduction of carbonate ions into the TiO<sub>2</sub>+Na<sub>2</sub>SO<sub>3</sub>

suspension, along with a significant decrease of signal for carbonate radical. ESR data (Fig. 5d) further confirms the increase of SO<sub>3</sub> after 2 min UV irradiation in the presence of carbonate ion. Based on the above results, one may deduce that the interplay between carbonate radical and sulfite ions is a crucial step giving rise to the increased SO<sub>3</sub> which is responsible for rapid SO<sub>2</sub> oxidation through chain propagation reactions (Deng et al., 2017). Nevertheless, there are two possibilities that might explain the aforementioned interaction. One is the oxygen transfer and the other route is electron transfer, which needs further clarification.

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We first examined the oxygen transfer path through <sup>18</sup>O isotope labeling experiments. TiO<sub>2</sub> particles were initially exposed to C<sup>16</sup>O<sub>2</sub>/N<sub>2</sub> and C<sup>18</sup>O<sub>2</sub>/N<sub>2</sub>, followed by the exposure of SO<sub>2</sub>/N<sub>2</sub>+O<sub>2</sub> under irradiation (Fig. 6a). Bidentate carbonate band centered at 1573 cm<sup>-1</sup> appears after the introduction of C<sup>16</sup>O<sub>2</sub>/N<sub>2</sub>, while this band shifts to 1558 cm<sup>-1</sup> when C<sup>18</sup>O<sub>2</sub>/N<sub>2</sub> is introduced, indicating the incorporation of <sup>18</sup>O into bidentate carbonate species, in good agreement with the previous report (Liao et al., 2002). However, no shift of IR features at 1269, 1219, and 1159 cm<sup>-1</sup>, assigned to (bi)sulfate species on TiO<sub>2</sub> particles, were observed throughout the reaction. This implies that the oxygen transfer path does not account for the rapid SO<sub>2</sub> oxidation on particles of concern.

In light of the above analysis, the electron transfer might be a plausible pathway to explain the fast oxidation within the reaction system. DFT calculations provide an accessible approach to study the electron transfer pathway. The result in Fig. 6b shows SO<sub>3</sub> formation is a SET process of CO<sub>3</sub> and SO<sub>3</sub><sup>2</sup>, where O atom in SO<sub>3</sub><sup>2</sup> transfers an electron to O atom in CO<sub>3</sub> to form SO<sub>3</sub> and CO<sub>3</sub><sup>2</sup>. This SET reaction is a thermodynamically favorable process, with the difference of Gibbs free energy between reactant and product lying at -24.09 kcal mol<sup>-1</sup>. However, we noted that the insufficient O<sub>2</sub> supply in aqueous media may be an underlying constraint to the proposed CO<sub>3</sub> initiated SO<sub>2</sub> oxidation pathway. Hence, we estimated both oxygen consumption and supply rates, and oxygen supply flux can be several orders of magnitude larger than corresponding consumption (see additional details in the supplementary text 5). Therefore, oxygen is sufficient in the reaction, allowing the considered chain reactions to continually proceed. Taken above results and discussions together, the following reactions are proposed accordingly (Eqs. 5-8):

$$CO_3^{-} + SO_3^{2-} \rightarrow CO_3^{2-} + SO_3^{-}$$
 (Eq. 5)

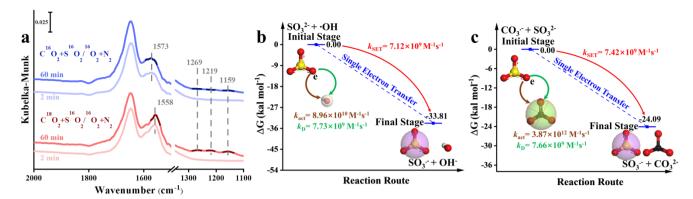
$$SO_3$$
 +  $O_2 \rightarrow SO_5$  (Eq. 6)

$$SO_5^{-} + SO_3^{2-} \rightarrow SO_4^{-} + SO_4^{2-}$$
 (Eq. 7)

$$SO_4$$
 +  $SO_3$  2-  $SO_4$  2-  $SO_3$  (Eq. 8)

Another important issue needs to be addressed as well. SO<sub>3</sub><sup>-</sup> can also be formed via the conventional reaction of ·OH and SO<sub>3</sub><sup>2</sup>- (Eqs. 9) and this process is also considered.

$$\cdot OH + SO_3^{2-} \rightarrow SO_3^{-+} OH^{-}$$
 (Eq. 9)



**Figure 6.** (a) *In situ* DRIFTS of heterogeneous reaction of SO<sub>2</sub> on the TiO<sub>2</sub> particles for 2 and 60 min after being exposed to  $C^{16(18)}O_2/N_2$  for 20 min under irradiation. (b) Reaction pathway of interaction between hydroxyl radical (·OH) and sulfite (SO<sub>3</sub><sup>2-</sup>) and (c) Interaction between carbonate radical (CO<sub>3</sub><sup>-</sup>) and sulfite (SO<sub>3</sub><sup>2-</sup>) through the SET process at the CCSD (T)-F12/cc-PVDZ-F12//M06-2X/6-311++G (3df, 3pd) level and  $\Delta G_0^{SET}$  represents the difference in Gibbs free energy between reactant and product. The white, black, yellow, and red spheres represent H, C, S, and O atoms, respectively. In order to visualize the variation of surface products in oxygen isotope experiments (panel a), DRIFTS features of these concerned species were plotted in dark colors. For interpretation of the references to color in the legends of panels b and c, the reader is referred to the Web version of this article.

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In this SET process, electron donor  $SO_3^{2-}$  reacts spontaneously with electron acceptor ·OH (Fig. 6c) and the calculated activation free energy barrier  $\Delta G^{\neq}_{SET}$  for this SET reaction is 2.50 kcal mol<sup>-1</sup>. Hence, the reaction process of ·OH with  $SO_3^{2-}$  is diffusion-controlled, and the total rate constant  $k_{SET-2}$  was calculated to be  $7.12 \times 10^9$  M<sup>-1</sup>s<sup>-1</sup>. In comparison, the rate constant  $k_{SET-1}$  of the diffusion-controlled SET process for  $CO_3^{--}$  and  $SO_3^{2-}$  was estimated to be  $7.42 \times 10^9$  M<sup>-1</sup>s<sup>-1</sup>. Despite a slight net increase of the rate, the distinguishable concentration of  $CO_3^{--}$  and ·OH should also be taken into account for the rate comparison in varied reaction paths. To visualize the difference, relative rates were calculated according to Eq. 10:

$$r = \frac{v_{\text{CO}_3}^{-} + \text{SO}_3^{2-}}{v_{\text{OH}^+ \text{SO}_3}^{2-}} = \frac{k_{\text{SET-1}}[\text{CO}_3^{-}][\text{SO}_3^{2-}]}{k_{\text{SET-2}}[\text{OH}][\text{SO}_3^{2-}]}$$
(Eq.10)

Where *r* is the ratio of two reaction rates, [CO<sub>3</sub><sup>-</sup>], [SO<sub>3</sub><sup>2</sup>-], and [·OH] refer to the concentration of corresponding reactants.

Previous literature suggests the concentration of carbonate radicals is able to show two orders of magnitude higher than that of hydroxyl radical at the surface of the water under solar irradiation (Chandrasekaran and Thomas, 1983; Goldstein et al., 2001; Shafirovich et al., 2001). An aqueous medium that attaches to particle surfaces offers an ideal environment for accumulating carbonate radicals. Consequently, concentrations of CO<sub>3</sub><sup>-</sup> and ·OH were set at the range from 1.0 × 10<sup>-10</sup> to 1 × 10<sup>-12</sup> mol L<sup>-1</sup> and from 1.0 × 10<sup>-12</sup> to 1 × 10<sup>-14</sup> mol L<sup>-1</sup> (Sulzberger et al., 1997) and *r* value could thus reach to 1.04 × 10<sup>4</sup> at most (Fig. S14). As a result, we speculated that the formation pathway of SO<sub>3</sub><sup>-</sup> via interaction between CO<sub>3</sub><sup>-</sup> and SO<sub>3</sub><sup>2</sup>- is a more effective route, corresponding well with experimental results.

In addition to the pathway launched by photo-generated holes, we also considered the sink of photo-generated electrons. In our reaction system,  $O_2$  is believed to be an electron trap and produce the superoxide radical ions  $(O_2^-)$ , which is reported to play a non-negligible role in sulfate formation (Shang et al., 2010) and should be taken into account to give a whole picture of reaction scheme in triggering sulfate formation on the surface of  $TiO_2$ -containing mineral dust particles. p-benzoquinone is a commonly-used  $O_2^-$  scavenger for trapping the  $O_2^-$  radical ions (Yan et al., 2018). Our supplementary data shows that adding an excess amount of p-benzoquinone into  $TiO_2$  particles reduces the sulfate yield by 32 % along with the appearance of sulfite ions over  $TiO_2$  particles upon exposure of  $SO_2$  (Fig.S12). Notably, the decrease in sulfate yield by around 30 % in the presence of  $O_2^-$  scavenger p-benzoquinone is almost complementary to that added with OH scavenger using TBA (70 %), pointing toward a minor sulfate formation pathway contributed by  $O_2^-$  relative to the major pathway by  $CO_3^-$  when carbonate ions are presented to efficiently capture OH ions. Following Shang's work (Shang et al., 2010),  $O_2^-$  involved  $SO_2$  oxidation can be given as Eqs. 11-13:

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$$e^-+O_2 \rightarrow O_2$$
 (Eq.11)

$$SO_2+O_2 \xrightarrow{\cdot} SO_3+O^{\cdot}$$
 (Eq.12)

$$350 \quad SO_3 + H_2O \rightarrow H_2SO_4 \tag{Eq.13}$$

Where intermediates  $SO_3$  formed via the interaction between  $SO_2$  and  $O_2$  subsequently couple with water molecules to produce sulfate species as a final product. pH is an important factor within aqueous chemical reaction processes and is likely to alter the dominated regime for sulfate production. Yet so far adjusting the pH of particle surfaces is quite tough, and exploring the role of dust surface pH in the reactivity of  $CO_3$  is not easily achieved. Notwithstanding, the increase of pH in  $TiO_2$  suspension was observed to promote the production of  $CO_3$  further strengthening the oxidation capability of dust particles (Fig. S11). In contrast, decreasing pH is expected to reduce the yield of  $CO_3$  since the reaction rate of  $CO_3$  with  $CO_3$  or orders of magnitude higher than that with  $CO_3$  or on this basis, the question arises whether the surface pH of mineral dust can be sustained to maintain fast  $SO_2$  oxidation triggered by  $CO_3$  in the typical lifespan of mineral dust.

Considering this, we thus plotted the heterogeneous sulfate production over TiO<sub>2</sub> and TiO<sub>2</sub>+CaCO<sub>3</sub> particles versus equivalent exposure time (Fig. S15). Clearly, the sulfate yield builds up steadily during the two-week equivalent exposure time (see a more detailed discussion on determining equivalent exposure time in supplementary text 6), suggesting that the regime of CO<sub>3</sub>- initiated SO<sub>2</sub> oxidation over TiO<sub>2</sub> and TiO<sub>2</sub>+CaCO<sub>3</sub> particles are slightly affected by the possible decrease of surface pH due to accumulation of sulfate production over entire reaction course. In the atmosphere, the lifetime of mineral dust particles ranges from several days to weeks (Bauer and Koch, 2005), and the equivalent exposure time considered in this study (nearly 2 weeks) falls right within the characteristic lifespan range of mineral dust particles. This leads us to deduce that persistent growth of sulfate shows a negligible effect on CO<sub>3</sub>- initiated SO<sub>2</sub> oxidation scheme proposed in this work.

Additionally, dust particles are reported to eject the radical ions from the surface under solar light irradiation, showing a non-negligible contribution to sulfate aerosol formation (Chen et al., 2021; Dupart et al., 2012), as described as:

370 Mineral Dust +  $hv \rightarrow ROS$  (g) (Eq.14)

ROS (g)+ humidified Air+  $SO_2 \rightarrow Sulfate(g)$  (Eq.15)

Where ROS (g) stands for the active intermediates in the gas phase. Over 400 ppm of  $CO_2$  is universal in the atmosphere, and it is expected to form (bi)carbonate ions once enters into the atmospheric aqueous media such as aerosol water, cloud droplets as well as fog environment. Bi(carbonate) ions are then prone to react with hydroxyl radical ions to form carbonate radicals. Following this line of reasoning, we attempt to monitor the plausible gas ROS species that are formed in the presence of  $CO_2$  (see a detailed discussion about the measurement approach and experimental setup in supplementary text 7 and Fig. S16).

When  $CO_2$  (atmospheric relevant concentration) is introduced into the homemade flow-cell chamber, with the intervening gap between  $TiO_2$ -coated film and probe molecule solution fixing at nearly 2 mm, and the short distance of which allows possible gaseous ROS to diffuse and react with aniline molecular (None, 2013). An increased degradation rate of aniline was seen, which can be attributed to the generation of active carbonate radical ions (Fig. S17). The maximum concentration of steady-state  $CO_3$ - radical ions supplied by partition processes between gas phase and solid-liquid phases (humified dust particles) was determined to be  $1.39 \times 10^{-13}$  M for  $TiO_2$ +Air+CO<sub>2</sub> system, which is over one order of magnitudes higher than that of ·OH for  $TiO_2$ +Air+system (2.15 ×10<sup>-15</sup> M). This observation matches with the earlier study where the concentration of carbonate radical can be two orders of magnitudes than ·OH over the water surface (Sulzberger et al., 1997).

The above results suggest that the photochemistry that involves carbonate ions, more precisely CO<sub>3</sub>- radical, increases sulfate formation. This finding broadens the prevailing view that acceleration of SO<sub>2</sub> oxidation over the carbonate salt is merely due to the favorable neutralization of H<sub>2</sub>SO<sub>4</sub> over an alkaline surface. To be important, upon irradiation active component TiO<sub>2</sub> in mineral dust produce carbonate radical in the gas phase when CO<sub>3</sub>- precursor CO<sub>2</sub> is presented, therefore potentially promoting sulfate aerosol formation in the atmosphere. Overall, it could be speculated that carbonate radical ions strengthen the oxidative capability of TiO<sub>2</sub>-containing mineral dust particles, and consequently accelerate SO<sub>2</sub> oxidation.

## 3.2.4 Field Measurements of Sulfate and (Bi)carbonate Ions.

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Complement field sampling and analysis were further conducted to examine our hypothesis that intermediates  $CO_3$  may play role in secondary sulfate formation in the atmosphere. We first considered the meteorological condition wind speed, which is an important parameter determining whether the local chemical process gains importance in affecting secondary sulfate formation. Meteorological information was collected from the open-access database (<a href="https://www.aqistudy.cn/">https://www.aqistudy.cn/</a>). During the sampling period, the wind scale mainly varies from 0 to 1, corresponding to the wind speed ranging from 0 to 1.5 m s<sup>-1</sup> (Fig. S18). All plots shown in Fig. S18 give rise to a statistical wind speed of  $0.76 \pm 0.73$ , which represents the weak dispersion of pollutants at low wind speed (not exceeding 2.5 m s<sup>-1</sup>)(Witkowska et al., 2016; Wu et al., 2020), indicating that local source is a dominant contributor to local air pollution.

It is generally accepted that under stagnant meteorological conditions (wind speed  $< 1.5~m~s^{-1}$ ), for the coarse-mode (2.5  $\mu$ m  $\sim 10~\mu$ m) of sulfate, the heterogeneous reaction of SO<sub>2</sub> on the dust surfaces is believed to be a major contributor (Liu

et al., 2017). This correlates to the fact that a large mass fraction of mineral dust is abundant in coarse-model particulate matter (PM) (Fang et al., 2017; Miller-Schulze et al., 2015), in which  $TiO_2$  was found at mass mixing ratios ranging from 0.1 to 10 % depending on the exact location where particles were uplifted (Chen et al., 2012; Hanisch and Crowley, 2003). Therefore, PM with relatively larger size dimensions is expected to contribute to secondary sulfate formation via heterogeneous reactions, which is supported by the recent field study where carbonate fraction of coarse PM is evidenced to promote secondary sulfate production (Song et al., 2018). Considering this, rather than determine the concentration of water-soluble ions in all stages, more attention was paid to PM collected in stages 1-4 (particles with their dimension  $\geq$  3.3  $\mu$ m). As (bi)carbonate ions are known as key precursors in producing  $CO_3$  and accelerate sulfate formation, quantifications of those relevant water-soluble ions were thus conducted (supplement text 23 and 24).

We further considered the relationships between sulfate ions and (bi)carbonate ions by means of linear regression analysis. However, under the low wind speed  $(0.76 \pm 0.73)$ , correlation coefficients  $R^2$  obtained for the relationship between bi(carbonate) and sulfate ions are not promising, 0.56 (sulfate vs carbonate) and 0.61 (sulfate vs bicarbonate) for  $PM_{3.3}$ - $PM_{9.0}$  during daytime hours. A plausible explanation is that although less significant, local primary emission source also brings bias and uncertainty to the correlation analysis. Shanghai is a coastal city, and sulfate species such as  $K_2SO_4$  and  $Na_2SO_4$  from the sea salt contribute to the local sulfate emission as well (Long et al., 2014). On the other hand, this novel  $SO_2$  oxidation channel is in the infant stage, and only active mineral dust components have been considered in this work whereas other components found in the coarse mode of PM such as organic matter, elemental carbon as well as sea salt (Cheung et al., 2011) are likely to involve this mechanism and alter the response of sulfate yield to  $SO_2$  heterogeneous uptake. In addition, the water-soluble ions determined in these samples (relatively small size) may not come from the net contribution of heterogenous reaction processes in absolute day-time and night-time periods. Some of the undesired processes that take place during day(nigh)-night(day) shifts may also contribute to the production of sulfate ions in separate sampling hours, thus reducing the correlation coefficients.

For those large particles (LP), that refer to the particles with a diameter large than 9  $\mu$ m in this work, sulfate ions show a rather weak or even no correlation to (bi)carbonate ions during the night-time and day-time hours (Fig. S19). This is likely due to the short lifetime of LP. Generally, the aerosol lifetime is on the order of less than an hour to days (Koelemeijer et al., 2006), highly depending on particle size. For example, the lifetime of PM<sub>10</sub> ranges from minutes to hours, and its travel distance, in general, is less than 10 km (Agustine et al., 2018). As a consequence, secondary sulfate formation through chemical reaction over LP is not significant with respect to *in situ* emissions. When PM downsizes to 2.5  $\mu$ m, PM<sub>2.5</sub> has a lifetime prolonged to nearly one day or longer (Liu et al., 2020b). Therefore, PM<sub>3.3</sub>-PM<sub>9.0</sub> are expected to have a relatively long lifetime, on the order of several hours, which enables the heterogeneous reaction process to become a more important contributor to overall sulfate ions measured in PM<sub>3.3</sub>-PM<sub>9.0</sub> than that in PM<sub>≥9.0</sub>. This is supported by our observations where during the daytime hours the correlation coefficients for PM<sub>3.3</sub>-PM<sub>9.0</sub>, i.e. 0.56 (sulfate vs carbonate) and 0.61 (sulfate vs bicarbonate), are higher than that of PM<sub>>9.0</sub>, i.e. 0.489 (sulfate vs carbonate) and 0.36 (sulfate vs bicarbonate), respectively.

Similarly, higher correlation coefficients are also observed for  $PM_{3.3}$ - $PM_{9.0}$  than  $PM_{\geq 9.0}$  in the sample collected during the nighttime periods.

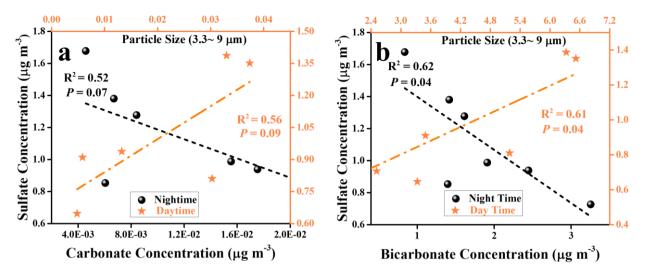


Figure 7. Field observation for the relationship between carbonate and sulfate ions during day-time and night-time hours. Linear relationship analyses for measured sulfate ions and estimated carbonate ions (a) and for measured sulfate ions and estimated bicarbonate ions (b) during the day-time and night-time hours, with particle sizes of PM ranging from 3.3 to 9 μm.

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While we note that the correlation coefficients between sulfate and (bi)carbonate are not promising in this work, ground-based field measurements of sulfate and (bi)carbonate ions shed light on their distinct correlations during the daytime and nighttime hours. In Fig. 7 and Fig.S19, the negative correlations between the mass concentrations of sulfate ions and (bi)carbonate ions are observed in the nighttime hours, consistent with the suppression of sulfate formation by CO<sub>2</sub> in the dark experiments. This is also supported by our previous lab study where CO<sub>2</sub>-derived (bi)carbonate species are demonstrated to block the active sites for yielding sulfate over mineral dust proxy aluminum oxide (Liu et al., 2020a). Instead, positive correlations were seen for those ions within PM sampled during the daytime hours regardless of size ranges and carbonate types (HCO<sub>3</sub>-/CO<sub>3</sub><sup>2</sup>-). This matches with the scenarios in which sulfate production upon irradiation in the presence of (bi) carbonate ions is increased over both model and authentic dust particles. Except the case (nighttime period, size larger than 9  $\mu$ m), most of the significance P values for their correlations were smaller than 0.1, with significance P values below 0.5 determined for bicarbonate vs sulfate, implying the plausible underlying connection between sulfate and (bi)carbonate ions. In fact, preceding ground-based observations of highly correlated relationship between Ca<sup>2+</sup> and SO<sub>4</sub><sup>2-</sup> water-soluble ions (Liu et al., 2020b) during the carbonate-enriched dust storm episodes, together with persistent reports on the significant role of photochemical channels in increasing the sulfate concentration during the daytime (Kim et al., 2017; Wei et al., 2019; Wu et al., 2017) indirectly reflects the possibility of accelerated SO<sub>2</sub> oxidation triggered by photo-generated active intermediates associated with carbonate species.

Overall, this is the first time that relationships between those ions are explored separately in these two periods. Taken together, carbonate radical is likely to promote sulfate production in the atmosphere during daytime hours. Detailed and systematic SO<sub>2</sub> oxidation channel triggered by CO<sub>3</sub>- needs further investigations to enable a better interpretation of correlations between these inorganic ions at the given meteorological conditions of sampling and physico-chemical properties of PM.

## 4. Conclusion

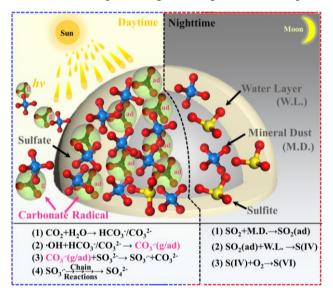
On the basis of the experimental and theoretical results derived from this work, we for the first time propose a novel reaction channel for fast SO<sub>2</sub> oxidation over mineral dust particles due to the formation of carbonate radical ions. A schematic chart for the sulfate formation in the presence of carbonate radicals upon solar irradiation or bi(carbonate) ions under dark conditions is summarized and elucidated in Fig. 8. During the night-time hours at 298 K (ambient temperature) CO<sub>2</sub>-derived (bi)carbonate species are prone to have a slightly negative effect on sulfate formation due to the competitive adsorption between CO<sub>2</sub> and SO<sub>2</sub>. For alkaline carbonate salt, it favors sulfate formation through the neutralization process. On the other hand, during the day-time hours, both CO<sub>2</sub>-derived (bi)carbonate species and carbonate salt work as the precursor of CO<sub>3</sub><sup>-</sup>, which promotes sulfate formation. Especially, uptake coefficients for carbonate salt containing mineral dust can be increased by 17 times, which is more pronounced than the increase due to the neutralization regime in the dark condition. Consistent with the findings reported in the earlier studies (Chen et al., 2021; Dupart et al., 2012), we observed the production of gas-phase CO<sub>3</sub><sup>-</sup> ions when mineral dust particles are irradiated in the presence of CO<sub>2</sub> (atmospherically-relevant concentration 400 ppm). This observation implies that the increased sulfate yield in part comes from promoted gas-phase secondary sulfate aerosol triggered by CO<sub>3</sub><sup>-</sup> (g).

By means of ROS scavenger experiments, direct observation of carbonate radical using NTAS analysis, oxygen isotope assay, ESR spectra as well as DFT calculations, CO<sub>3</sub><sup>-</sup>-initiated S(IV) oxidation involving single electron transfer process are elucidated. While carbonate radical ions are mainly responsible for rapid sulfate formation, superoxide radical ions are likely to serve as a minor pathway over TiO<sub>2</sub>-containing mineral dust particles. In addition, a weak correlation between sulfate ions and (bi)carbonate ions observed for PM<sub>3,3</sub>-PM<sub>9,0</sub> in this work correlates to non-chemical primary emission and complicated nature of CO<sub>3</sub><sup>-</sup> regime of sulfate production in the atmosphere. Nevertheless, complement field sampling of ambient PM and analysis of sulfate and (bi)carbonate ions in this study unfold their distinct correlations during the daytime and nighttime hours, these two tendencies of which agrees with the experimental observations.

In this work, only atmospheric secondary sulfate formation was considered, whereas the oxidation of primary organic species yet has not been investigated. In fact, carbonate radical ions are prone to rapidly react with electron-rich organics amines (Stenman et al., 2003; Yan et al., 2019) as well as phenol (Busset et al., 2007; Xiong et al., 2016), and it may potentially serve as the key oxidants that drive the fast formation of SOA in the atmosphere. Besides, observation of strengthened photochemistry launched by carbonate radicals suggests that such chemistry may be amplified on

atmospherically relevant reactions that occur in cloud droplets as well as fog water where they often contain hydroxyl radicals and water-soluble (bi)carbonate ions.

To be important, carbonate radical ions are observed to be formed in the gas phase in the atmospherically relevant CO<sub>2</sub> concentration (400 ppm) when mineral dust proxies are irradiated. This will help the formation of external sulfate aerosol formation. Since both sulfate aerosol and CO<sub>2</sub> are well known to affect the radiation budget and solar energy balance on the earth (Cheung et al., 2011; Möller, 1964), their overall influence on the global climate considering the increased yield of sulfate aerosol triggered by CO<sub>2</sub>, the precursor of carbonate radical, needs further investigation. Therefore, our study highlights the necessity for a comprehensive understanding of the CO<sub>3</sub><sup>-</sup> relevant chemistry in the underlying impacts of fine PM concentration, human health, and climate. All these assumptions need to be investigated in further detail. This study provides the first indication that carbonate radical not only plays a role as a marginal intermediate in tropospheric anion chemistry but also as a strong oxidant for surfacial processing of trace gas in the atmosphere.



**Figure 8.** Schematic of the sulfate formation in the presence and absence of carbonate radical. Noting that g and ad represent gas-phase and adsorbed carbonate radical ions, respectively.

Data availability. The data that support the results are available from the corresponding author upon request.

Author contributions. Y.L., Y.D. and L.Z. initially proposed the idea; Y.L. and Y.D. designed and performed most of the experiments; J.L. performed DFT calculations; Y.L., X.Z. and T.W. contributed to field samplings and data analysis; K.L., K.G., A.B., I.N., X.Z., C.G., and L.Z. provided suggestions on the experiments and paper writing; All authors wrote the manuscript.

Competing interests. The authors declare that they have no conflict of interest.

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