A Novel Pathway of Atmospheric Sulfate Formation Through Carbonate Radical

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Abstract. Carbon dioxide is considered an inert gas that rarely participates in atmospheric chemical reactions. However, we

- 15 show here that CO_2 is involved in some important photo-oxidation reactions in the atmosphere through the formation of carbonate radicals (CO_3^{-}). This potentially active intermediate CO_3^{-} is routinely overlooked in atmospheric chemistry regarding its effect on sulfate formation. Present work demonstrates that SO_2 uptake coefficient is enhanced by 17 times on mineral dust particles driven by CO_3^{-} . It can be produced through two routes over mineral dust surfaces: i) hydroxyl radical + CO_3^{2-} ; ii) holes (h^+) + CO_3^{2-} . Importantly, upon irradiation mineral dust particles are able to produce gas-phase carbonate
- 20 radical ions when the atmospherically relevant concentration of CO_2 presents, therefore potentially promoting external sulfate aerosol formation and oxidative potential in the atmosphere. Employing a suite of laboratory investigations of sulfate formation in the presence of carbonate radical on the model and authentic dust particles, ground-based field measurements of sulfate and (bi)carbonate ions within ambient PM, together with density functional theory (DFT) calculations for single electron transfer processes in terms of CO_3 -initiated S(IV) oxidation, a novel role of carbonate radical in atmospheric chemistry is elucidated.

25 1. Introduction

Atmospheric composition changes are subjected to highly reactive light-induced radicals, such as hydroxyl (\cdot OH), hydroperoxyl (HO₂ \cdot), or nitrate radicals (NO₃ \cdot), which are able to alter not only compositions but also physical and chemical properties of particulate matter (Davis and Francisco, 2011; Platt et al., 1990; Prinn et al., 2001; Thompson, 1992). However, when atmospheric chemical reactions occur over nanometer-sized particles at ambient conditions, which creates a locally

30 enriched aqueous medium of unique chemical activity, other radicals might likewise gain importance. The carbonate radical (CO_3^{-}) is typically such an active radical. The lifetime of CO_3^{-} ranges from a microsecond to even a few milliseconds and its

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concentration can be two orders of magnitude higher than that of hydroxyl radicals over the water surface (Chandrasekaran and Thomas, 1983; Goldstein et al., 2001; Shafirovich et al., 2001; Sulzberger et al., 1997). In addition, the one-electron reduction potential of $E^0(CO_3^{-}/CO_3^{2-})$ couple is 1.78 V vs. NHE at neutral pH, leaving CO_3^{-} a strong oxidant in aquatic

- 35 chemistry (Bisby et al., 1998; Cope et al., 1973 ; Merouani et al., 2010). Previous studies concerning carbonate radical in aqueous media demonstrate that it reacts rapidly with some organic compounds with higher selectivity (Merouani et al., 2010), especially for those electron-rich compounds amines (Stenman et al., 2003; Yan et al., 2019). Also, it has been pointed out that the scavenging of hydroxyl radicals by (bi)carbonate species leads to the formation of CO_3^{-} ions (Graedel and Weschler, 1981), which promotes the degradation of phenol (Xiong et al., 2016). Besides, a higher second order of rate constant, lying at 10^9 M^-
- ¹ s⁻¹, has been reported for the reaction of CO_3^{-} with porphyrins (Ferrer-Sueta et al., 2003), indicating that this radical ion has great oxidation capability that may trigger atmospherically relevant chemical reactions. However, it is only regarded as a marginal intermediate in tropospheric anion chemistry so far (Beig and Brasseur, 2000; Dotan et al., 1977; Graedel and Weschler, 1981; Lehtipalo et al., 2016) and its underlying role as an active oxidant for heterogeneous reaction in the atmosphere is barely explored. Very recently, our group observed the promotional effect of CO_3^{-} on atmospheric nitrate
- 45 formation (Fang et al., 2021). Motivated by this finding, attempts were made to further explore its role in other important atmospherically-relevant reactions.

It is well documented that sulfate (SO_4^{2-}) is also a key constituent of aerosols in the atmosphere (Huang et al., 2015; Su et al., 2016). It is able to serve as the precursors of efficient cloud condensation nuclei, with optical properties leading to a cooling effect (Wang et al., 2011). As a consequence, the mechanism aspect of secondary sulfate formation was the focus of numerous

- 50 studies over the past decades (Hung et al., 2018; Stone, 2002; Zhang et al., 2015b). There is a consensus that high-valence sulfur (VI), produced from the oxidation of anthropogenic SO₂, is the dominant source for atmospheric secondary sulfate. However, a remarkable missing sulfate budget emerges for the atmospheric modeling, which underpredicts $SO_4^{2^2}$ by over 50 % (normalized mean bias) with respect to observational results when heterogeneous aerosol chemistry is not considered (Zheng et al., 2015). This indicates that the heterogeneous sulfate production pathway is a crucial process and exploring the
- 55 unconsidered heterogeneous mechanism is very likely to narrow the gap between observations in lab studies, field measurements, and numerical modelings. However, due to the missing chemical mechanism that initiated fast SO₂ oxidation, atmospheric models fail to capture the key feature of atmospheric observations of high sulfate production (Wang et al., 2020). Consequently, there are unknown heterogeneous reaction pathways of significance and previously unconsidered promoters that have great potential to accelerate sulfate formation.
- Due to the high stability of CO_2 under ambient conditions (Hossain et al., 2020), there are rare studies concerning the influence of CO_2 in atmospheric chemical processes (Deng et al., 2020; Liu et al., 2020a; Xia et al., 2021). CO_2 is demonstrated to form (bi)carbonate species over humidified dust particles (Baltrusaitis et al., 2011; Nanayakkara et al., 2014) and reduced to CO under solar illumination (Deng et al., 2020). However, its impact on atmospheric heterogeneous reactions remains poorly characterized. Our earlier laboratory study shows that CO_2 decreases the sulfate formation on aluminum oxide particles
- 65 in the dark (Liu et al., 2020b) while upon solar illumination its role in SO₂ oxidation over mineral dust surfaces is still an open

question. In addition, carbonate salt is enriched in authentic dust aerosol (Cao et al., 2005) and reported to reach over 10 % wt. of Asian dust particles (McNaughton et al., 2009). It is generally accepted that CO_3^{2-} affects atmospheric chemistry and aerosol characteristics mainly through its intrinsic alkalinity, which buffers aerosol acidity and favors the sulfate formation (Bao et al., 2010; Kerminen et al., 2001; Yu et al., 2018). In fact, either CO₂ or carbonate salt is able to produce the active CO_3^{-} under the

ambient circumstance and increase the oxidative capacity in the atmosphere. Combined with our previous investigation of CO_3^{-} (Fang et al., 2021), this radical ion is likely to be a driving force for fast SO₂ oxidation. However, to the best of our knowledge, no work has considered how and to what extent the carbonate radical influences SO₂ heterogeneous oxidation in the atmosphere.

In the current study, through laboratory studies, we presented that carbon dioxide and calcium carbonate, working as the 75 precursor of carbonate radicals, extend their ability to accelerate sulfate formation over authentic particles in the atmosphere. Together with quantum chemistry calculations, a detailed molecular mechanism regarding a single electron transfer (SET) process between carbonate radical and sulfite ions is elucidated. Furthermore, ground-based observations validate some findings from the laboratory-based simulations.

2. Experimental methods

80 2.1 Laboratory Studies

A series of characterizations were initially performed to investigate the mineral dust of concern by using X-ray diffraction (XRD) and Raman technique. The heterogeneous reaction of SO_2 on mineral dust particles in the presence of CO_2 and carbonate species were then investigated by the *in situ* Fourier transform spectrum (DRIFTS), ion chromatography (IC), Raman, electron spin resonance (ESR), and nanosecond transient absorption spectroscopy (NTAS). Furthermore, we employed three

85 types of authentic dust particles and four kinds of synthesized authentic simulants to probe the proposed scheme. Besides, the steady concentration of CO₃⁻⁻ ions in each reaction system was determined by High-Performance Liquid Chromatography (HPLC) using probe molecule aniline. All corresponding configuration setup, characterizations, and methodologies can be found in Supplement text 1, 2, 8-18, and Tables S1-S3, and Fig. S1-S12 and Fig. S15-S17. In terms of oxygen isotope experiments, more detailed procedures are available in Supplement, text 21.

90 2.2 Quantum Chemical Calculation

We employed density functional theory (DFT) calculations in the term of the single electron transfer (SET) process using Gaussian 09 package to investigate this novel route, detailed in Supplement text 19-20 and Fig. S13-S14.

2.3 Field Observations

Atmospheric aerosols were collected on the roof of our department using an 8-stage non-viable-cascade-impactor type sampler 95 (TISCH TE Inc., USA), with sampling details shown in supplement Text 22 and daily variations of wind scales shown in Fig. S18. The concentration of water-soluble sulfate ion was by measured IC analysis while that of (bi)carbonate ions was determined by the ionization balance approach (Supplement text 23-24). The relationship between (bi)carbonate ions and sulfate ions during the daytime and nighttime hours were then determined.

3. Results and discussion

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100 **3.2.1** Accelerated sulfate production in the presence of carbonate.

The physico-chemical properties of employed mineral dust proxies were first characterized (Fig. S1), consistent with earlier studies (Balachandran and Eror, 1982; Shang et al., 2010; Su et al., 2008), and spectral irradiance of the solar simulator applied in the present study is well covered by natural sunlight (Fig. S2), as much as possible having experimental results from the lab simulate the real atmosphere. Upon irradiation, the sulfate yield on TiO₂-CaCO₃ mixture particles (50 wt. % CaCO₃), measured by IC, is significantly enhanced by 7 times and 23 times compared to that of pristine TiO₂ and CaCO₃ (Fig. 1a), respectively. In stark contrast, there is a negligible increase of sulfate production detected on the TiO₂-CaCO₃ mixture relative to that of pristine CaCO₃ and TiO₂ in dark experiments (Fig. 1a). While great discrepancies in sulfate yield between dark and irradiation

experiments, it remains unclear for the role of carbonate salt in promoting sulfate formation. There is a prevailing view that neutralization of H₂SO₄ accounts for rapid SO₂ oxidation over carbonate salt particles, which needs careful consideration.

- 110 Following this speculation, two types of mixtures TiO₂-CaCO₃ and TiO₂-CaO were employed. In the dark condition (Fig. S3), both TiO₂-CaO and TiO₂-CaCO₃ almost yield an identical concentration of sulfite and sulfate as they are likely to present similar physical and chemical properties, e.g. surface pH and neutralization capability. Once irradiated, TiO₂-CaCO₃ particles produce nearly two times of sulfate than TiO₂-CaO particles, along with a sharp decrease of S(IV) species on the surface of TiO₂-CaCO₃ surfaces (see additional discussion in the supplementary text 2). These results allow us to assert that the carbonate-
- 115 containing system contains another important mechanism for sulfate generation beyond the production of an alkaline environment. Fig. 1b and 1c show the *in situ* diffuse reflectance infrared DRIFTS features of S(IV) and S(VI) species formed on theoretical and experimental TiO₂-CaCO₃ mixtures (wt./wt. = 50/50) upon irradiation for 90 min, respectively. The "theoretical" here is calculated based on the *in situ* DRIFTS of pristine TiO₂ and CaCO₃ through a simple linear superposition. These results suggest a synergistic effect presented in this mixture for sulfate formation under solar irradiation. In addition, a
- 120 more evident fingerprint SO_4^{2-} feature (Dong et al., 2009; Yann Batonneau et al., 2008) monitored by Raman spectroscopy appears over TiO₂-CaCO₃ particles compared to pristine TiO₂ particles (Fig. S4), in good agreement with the IC analyses and *in situ* DRIFTS measurements. Combining DRIFTS experiments with the obtained calibration curve (Fig. S5), we estimated that the uptake coefficient of TiO₂-CaCO₃ mixture (50 wt. % CaCO₃) is increased by about 17 times as compared to that of pure CaCO₃ or TiO₂ (Table S3). More importantly, upon irradiation SO₂ uptake coefficients for these dust proxies lie at the
- 125 order of magnitudes of 10^{-4} , indicating that the photochemical pathway associated with carbonate species is likely a potential driving force to trigger fast SO₂ oxidation in the atmosphere.



Figure 1. (a) Sulfate concentration quantified by IC on mineral dust particles after exposure to gaseous SO₂ under irradiation or dark for 30 min. *In situ* DRIFTS of S(IV) and S(VI) species yield on theoretical (b) and (c) experimental TiO₂-CaCO₃ mixtures (wt./wt. = 50/50) upon irradiation for 90 min. Reaction conditions: RH = 30 %, Light intensity (I) = 30 mW cm⁻², Total flow rate = 52.5 mL min⁻¹ and SO₂ = 2.21×10¹⁴ molecules cm⁻³. All spectra were processed by the Kubelka-Munk (K-M) algorithm. Noting that the production of sulfur species in theoretical TiO₂+CaCO₃ mixtures refer to 0.5 × K-M bands of sulfur species of TiO₂ + 0.5 × K-M bands of sulfur species of CaCO₃ while that for experimental TiO₂+CaCO₃ mixtures refer to 1 × K-M bands of sulfur species of TiO₂+CaCO₃ mixtures (wt./wt. = 50/50). (d) Energy Dispersive Spectroscopy (EDS) mapping of sulfur. (e) Selected HRTEM region containing a high density of sulfur for further observation and the red rectangle refers to the region shown in panel f. (f) The HRTEM image in high resolution with lattice fringes and (g) corresponding FFT power spectra, lattice indexing, and (1-6) inverse FFT analysis of lattice signal shown in panel g. In panel f, the term C-SO₄²⁻ stands for crystalline SO₄²⁻, i.e. CaSO₄ and Ti(SO₄)₂. Particles for the HRTEM measurement refer to TiO₂-CaCO₃ mixture particles upon exposure to the 4.42×10¹⁴ molecules cm⁻³ SO₂/N₂+O₂ for 60 min while other reaction conditions are as same as that of above sulfate quantification experiments.

High-resolution transmission electron microscopy (HRTEM) analysis of TiO_2 -CaCO₃ particles after reaction, in combination with energy dispersive spectrometer mapping measurements of sulfur component, were conducted to investigate the synergistic effect between TiO_2 and carbonate ions (Fig. 1 d-f and Fig. S6). A region with a relatively high density of sulfur species was selected for further observation and the distribution of each component (Fig. 1g) was determined by fast Fourier

- 145 transformation (FFT) and inverse FFT analyses of the selected HRTEM image in high resolution with lattice fringes shown in Fig. 1f. Observation of crystalline Ti(SO₄)₂ and CaSO₄ on the interface of TiO₂ and CaCO₃ components imply that the synergistic effect on sulfate production likely originates from interplays of those two types of components under solar illumination. We further assessed the importance of interfacial contact between TiO₂ and CaCO₃ in sulfate production by two synthesis approaches in which the interface abundance is modulated for comparison. Typically, a "grinding" method was used
- 150 to make TiO_2 -CaCO₃ mixture with compact contact between those two components, thus leading to a strong interaction. Meanwhile, the "shaking" method is designed to create a TiO_2 -CaCO₃ mixture with weak interplay, leaving relatively fewer amounts of interfaces within the mixtures. The resulting mixing statuses of two samples meet our expectations, evidenced by the scanning electron microscope (SEM) technique (Fig. S7). IC quantification analysis suggests that particles with considerable junctions exhibit a more pronounced promotion for sulfate yield than those having relatively few junctions (Fig.
- 155 S8). These results emphasize the importance of an indispensable interface connection between TiO_2 and $CaCO_3$ in fast production upon irradiation.



Figure 2. Sulfate concentration quantified by IC. Sulfate concentration was measured by IC on mineral dust simulants after exposure to gaseous SO₂ (2.46×10¹⁴ molecule cm⁻³) under irradiation. Noting that SiO₂: Al₂O₃: CaCO₃: TiO₂ refers to the mass
160 fraction ratios of the components in simulants. Experiments were all conducted at RH of 30 % and Light intensity (*I*) of 30 mW cm⁻².

The rapid SO₂ oxidation pathway was further probed by employing mineral dust simulants where two dominant crust constituents SiO₂ and Al₂O₃ were introduced into TiO₂-CaCO₃ particles to mimic the authentic mineral dust particles in the

- 165 atmosphere, with specific component and corresponding ratio information shown in Table S1. It is worth mentioning that the determination of the ratio of each component in the simulants relies on the EDS mapping results of ATD particles. In Fig. 2, the introduction of TiO₂ components (≈ 1 % wt.) into SiO₂-Al₂O₃ leads to 81.6 % enhancement of sulfate production while merely 24.8 % wt. increase of sulfate yield was observed once ≈ 8 % wt. of CaCO₃ was incorporated into SiO₂-Al₂O₃ dust particles. Surprisingly, mixing of ≈ 1 % mass fraction of TiO₂ and ≈ 8 % wt. of CaCO₃ into SiO₂-Al₂O₃ gives rise to a 235 %
- 170 increase of sulfate formation relative to that of SiO₂-Al₂O₃. Hence, the observed synergistic effect on heterogeneous oxidation of SO₂ is likely to take effect in the atmosphere.

Fe₂O₃ is also one of the crucial components found in authentic mineral dust (El Zein et al., 2013), and it has been reported to produce ROS under solar irradiation (Li et al., 2019), thus likely involving the reaction mechanism proposed in this work. Similar to experiments using TiO₂+CaCO₃ mixture, alpha-Fe₂O₃+CaCO₃ are prepared by grinding alpha-Fe₂O₃ and CaCO₃. In

- 175 Fig. S9a, our results show that alpha-Fe₂O₃ can not trigger fast SO₂ oxidation in the presence of carbonate ions upon irradiation, which is distinguished from results we derived from TiO₂+CaCO₃ mixture. This can be explained by the fact that Fe₂O₃ shows a lower redox activity relative to TiO₂ (Fig. S9b), where its strong redox capability essentially enables photo-induced electrons and holes to produce O_2^{--} and \cdot OH radical ions. In stark contrast, the valence band and conduct band of Fe₂O₃ lie at -0.18 and at 1.68 V vs. NHE (pH = 7), lower than the redox potential required for generating O_2^{--} , \cdot OH as well as CO₃⁻⁻ (Li et al., 2016).
- 180 Hence, no promoted sulfate production is seen for Fe₂O₃+CaCO₃ particles under irradiation. More discussion on the inconsistency between our study and the previous results regarding the response of SO₂ oxidation to solar irradiation can be found in supplementary text 3.

Overall, we show that upon irradiation atmospherically relevant content of TiO_2 (nearly 1 %) found in authentic dust simulants is able to interact with carbonate ions to launch a fast SO_2 oxidation channel, which is beyond the conventional

185 regime of alkaline neutralization of H_2SO_4 . Unlike TiO₂, alpha- Fe₂O₃ lacks the ability to initiate fast SO₂ oxidation by generating CO₃⁻⁻ due to its limited photo-chemical activity although ferric chemistry is important in secondary sulfate formation in the atmosphere (Sullivan et al., 2007; Yermakov and Purmal, 2003).

3.2.2 Accelerated sulfate production in the presence of CO₂.

- Atmospheric CO₂ is also an important source of (bi)carbonate. Its influence on photochemical SO₂ uptake on mineral dust was thus studied. In the presence of atmospherically relevant CO₂ (9.83×10^{15} molecules cm⁻³), sulfate yield was increased under irradiation as compared to CO₂-free case (Fig. 3a and b). We cautiously examined the buffering effect of formed (bi)carbonate on sulfate production by time-resolved DRIFTS spectra (Fig. 3c and d). CO₂ suppresses both S(IV) and S(VI) products under the dark. This observation implies that (bi)carbonate ions that evolve active intermediates upon irradiation may be a plausible force to drive rapid sulfate formation rather than accumulating much more sulfur species through buffering effect. The reaction
- 195 kinetics of SO₂ on mineral dust particles follows the pseudo-first-order, as evidenced by the SO₂ concentration dependence

experiments (Fig. S10 a-f). Besides, a nearly 50 % increase of SO₂ uptake coefficient is observed for the mineral dust proxy TiO₂ after being exposed to 9.83×10^{15} molecules cm⁻³ (400 ppm) CO₂+SO₂/N₂+O₂ mixture.

As another step toward a real scenario in the atmosphere, experimental trials employing authentic dust particles, i.e. Arizona test dust (ATD), clays IMt-2 (Illite, Mont., USA), and K-Ga-2 (Kaolin, Georgia, USA), were implemented (Table S2). In Fig.

- 4, K-Ga-2 clay exhibits the most marked promotional effect on sulfate yield (by nearly 100 % increased sulfate production in the CO₂-involved case under irradiation). This correlates with its considerable TiO₂ contents (3.43 %) in the K-Ga-2 clay, in which active intermediates are readily evolved from TiO₂ and (bi)carbonate species upon irradiation. However, the promotional effect of CO₂ on sulfate production under irradiation is weak for IMt-2 (the content of TiO₂ \approx 0.99 %) and ATD (the content of TiO₂ \approx 0.46 %) as compared to K-Ga-2 particles. This may correlate to their higher mass fraction of alkaline earth metal oxide (denoted as A.E.), which enables dust particles to possess a large number of (bi)carbonate species in the natural
- 205 Oxide (denoted as A.E.), which enables dust particles to possess a rarge number of (b)/carbonate species in the natural environment where they have experienced long-term exposure to atmospheric CO_2 during the regional transport. Therefore, the aforementioned synergetic effect takes effect over IMt-2 and ATD particles even without exposure to CO_2 due to the presence of abundant carbonate formed, and a less evident increase of sulfate yield was thus observed.



Figure 3. Time-resolved DRIFTS of S(IV) and S(VI) products over TiO₂ particles after exposure to SO₂/N₂+O₂ in the absence and presence of CO₂ upon irradiation (**a** and **b**) and those reactions under dark (**c** and **d**). Reaction conditions: RH = 30 %, Light intensity (I) = 30 mW cm⁻², total flow rate = 52.5 mL min⁻¹ and SO₂ = 7.37×10^{13} molecules cm⁻³.



Figure 4. Laboratory studies of sulfate production on authentic dust and clay membranes under the dark and irradiation (30 mW cm⁻²) upon exposure to 4.91×10^{14} molecules cm⁻³ SO₂/N₂+O₂ and 2.46×10^{18} molecules cm⁻³ CO₂+ 4.91×10^{14} molecules cm⁻³ SO₂/N₂+O₂ at RH of 30 % (total flow rate = 100 mL min⁻¹).

3.2.3 Reaction Mechanism.

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The heterogeneous reaction of SO₂ on dust particles in the atmosphere is a complicated process, covering a series of reactions taking place through both homogeneous and heterogeneous ways. At a sufficiently low RH condition (normally below 10 % RH), water readily dissociates on the surface of metal oxide under ambient atmospheric conditions, where metal oxide surface is terminated by hydroxyl groups that hydrogen bond to adsorbed water molecules (Cwiertny et al., 2008). In this case, SO₂ oxidation over dust particles is dominated by the heterogenous pathway, where the resulting hydroxyl groups capture SO₂ in the gas phase first and then stabilize it as adsorbed S(IV)_{ad}. Afterward, S(IV)_{ad} will be oxidized by oxidants in the atmosphere

- or photo-induced active intermediates produced from the dust surface upon irradiation. As the RH increases beyond 10 % -15 %, multilayer water coverage occurs, reaching approximately two monolayers at RH of 30 % (Mogili et al., 2006). Under these circumstances, the amount of water adsorbed onto the surface of the dust particles is believed to be sufficiently large that it is liquid-like in its physical and chemical properties (Cwiertny et al., 2008) (Peters and Ewing, 1997). In this work, heterogenous SO₂ oxidation over mineral dust proxies proceeds at the RH of 30 %, and two water layers absorb on dust
- 230 particles. Thus, radical ions are expected to play a key role in fast SO₂ oxidation and mechanism studies performed in solution phase are persuasive to some extent.

Our preliminary sulfate quantification results (Fig. 1-4) suggest that the presence of (bi)carbonate ions under solar light contributes to increased sulfate yield. In our carbonate-containing reaction system, a plausible intermediate is the active carbonate radical. They are readily produced via the following two pathways. First of all, carbonate anion can be directly oxidized by produced photo-induced holes (Eqs 2 and 3), as the redox potential of CO_3^{-7}/CO_3^{2-} is 1.78 V (vs NHE, at pH = 7).

which is lower than the TiO_2 valence band (VB) potential of 2.67 V (vs NHE, at pH = 7) (Li et al., 2016; Xiong et al., 2016):

Mineral dust $+hv \rightarrow h^++e^-$	(Eq. 1)
$h^++CO_3^2 \rightarrow CO_3^-$	(Eq. 2)

In the second pathway, carbonate radicals evolve through the reaction of (bi)carbonate anion with formed hydroxyl radicals •OH over mineral dust surfaces (Zhang et al., 2015a) (Eqs 3 and 4).

$$h^+ + H_2O_{ad} \rightarrow OH + H^+$$
 (Eq. 3)

$$\cdot OH + HCO_3 \cdot CO_3^2 \rightarrow CO_3^2 + H_2O/OH$$
(Eq. 4)

The above assumptions are supported by nanosecond transient absorption spectra (NTAS), in which signal (Δ OD) of carbonate radical CO₃⁻ at 600 nm (Bhattacharya et al., 1998) only emerged for dust suspension containing (bi)carbonate species

- 245 (Fig. 5a). A promoted degradation of aniline in TiO₂ suspension due to the presence of carbonate ions presents additional evidence of the formation of active CO₃⁻⁻ ions and strengthened oxidation capability of TiO₂ (Fig. S11, see additional discussion in supplementary text 4). The CO₃⁻⁻ induced chemistry was further evidenced by ·OH scavenging experiments using tertiary Butyl Alcohol (TBA) and isopropanol (i-PrOH) as they show lower reaction rates with CO₃⁻⁻ (k_{CO_3} ⁻⁻, i-PrOH < 4.0 ×10⁴ M⁻¹ s⁻¹ and k_{CO_3} ⁻⁻, t_{TBA} <1.6 ×10² M⁻¹ s⁻¹) relative to that with ·OH ($k_{i-PrOH--OH}$ < 1.9 ×10⁹ M⁻¹ s⁻¹ and ($k_{\cdot OH,TBA} = 6 \times 10^8$ M⁻¹ s⁻¹)
- 250 (Buxton et al., 2009; Liu et al., 2015) (Liu et al., 2015) (Li et al., 2020). Tertiary Butyl Alcohol (TBA) sharply decreases the yield of sulfate on TiO₂ surface by nearly 70 %, with sulfite ions being the dominant sulfur species (Fig. S12). Meanwhile, a significant loss of sulfate yield when TiO₂ suspension was added with i-PrOH (Fig. 5b). This is in strong contrast to the result of a carbonate-involved system where the reactivity is sustained, i.e., carbonate radicals offer an alternative reaction pathway for SO₂ oxidation. This is plausible since the carbonate ions are excellent 'OH scavenger, and CO₃⁻ becomes predominant
- 255 species at a relatively strong alkaline aqueous-like environment in the presence of carbonate salt. This is supported by the previous work (Sun et al., 2016), in which adding 0.1 M of NaHCO₃ into the UV/H₂O₂ system (H₂O₂ = 0.3 mM) were sufficient to suppress ·OH concentration to around 10^{-15} M, creating a carbonate radical dominated reaction system ([CO₃·⁻] = 8.64 × 10^{-12} M). In our supplementary experiments (Fig. S11), 0.2 M of carbonate salt was employed, and the reaction rate of CO₃²⁻ with ·OH is nearly two orders of magnitude higher than that of HCO₃⁻, thus giving rise to carbonate radical being the substitute
- 260 of hydroxyl radical in the reaction. The above results suggest that \cdot OH is a major contributor to sulfate yield on TiO₂ particles in the absence of carbonate ions while CO₃⁻⁻ ions dominate SO₂ oxidation over carbonate-containing dust particles upon irradiation. In addition to experimental investigations, the carbonate radical formation process is proved to be thermodynamically favorable, supported by density functional theory (DFT) calculations (Fig. S13).



Figure 5. (a) Single-wavelength transient absorption spectra of various aqueous solutions. (b) Sulfate formation change Δ(SO4²⁻) determined by different sulfate concentrations with and without the addition of isopropanol as hydroxyl radical scavenger. (c) The difference in transient absorption kinetics of sulfite radical and carbonate radical at the various aqueous solutions and their corresponding growth-decay fit curves. ΔA-signal was recorded at 255 and 600 nm after pulsed 355 nm laser excitation. (d) ESR spectrometry of [DMPO–SO3⁻] intermediate formed in a solution of d TiO₂ (3mg~4 mL) + 0.1 M Na₂SO₃ and TiO₂ (3mg~4 mL) + 0.5 M Na₂CO₃ + 0.1 M Na₂SO₃. For clarity, the integrated areas of ESR profiles were also presented for direct comparison. Exp. and Sti. stand for experimental results and corresponding fitting results using software Isotropic Radicals.

On the other hand, the previous studies (Chameides and Davis, 1982; Das, 2001; Neta and Huie, 1985) agree with the key role of sulfite radical (SO₃⁻⁻) in rapid sulfate production in an aqueous medium, and the present reaction system creates a localized environment where SO₃⁻⁻ can be readily produced from the TiO₂ and S(IV) species upon solar illumination (Salama et al., 1995). Consequently, probe light of NTAS at wavelength 255 nm (ascribed to sulfite radical) and 600 nm (ascribed to carbonate radical) were simultaneously monitored (Ghalei et al., 2016; Goldstein et al., 2001; Hayon et al., 1972). A weak signal of sulfite radical was observed in the system of TiO₂+Na₂SO₃ suspension under irradiation (Fig. 5c). On the contrary, the sulfite radical signal is strengthened after the introduction of carbonate ions into the TiO₂+Na₂SO₃ suspension, along with

- a significant decrease of signal for carbonate radical. ESR data (Fig. 5d) further confirms the increase of SO_3^{-} after 2 min UV irradiation in the presence of carbonate ion. Based on the above results, one may deduce that the interplay between carbonate radical and sulfite ions is a crucial step giving rise to the increased SO_3^{-} which is responsible for rapid SO_2 oxidation through chain propagation reactions (Deng et al., 2017). Nevertheless, there are two possibilities that might explain the aforementioned interaction. One is the oxygen transfer and the other route is electron transfer, which needs further clarification.
- We first examined the oxygen transfer path through ¹⁸O isotope labeling experiments. TiO₂ particles were initially exposed to C¹⁶O₂/N₂ and C¹⁸O₂/N₂, followed by the exposure of SO₂/N₂+O₂ under irradiation (Fig. 6a). Bidentate carbonate band centered at 1573 cm⁻¹ appears after the introduction of C¹⁶O₂/N₂, while this band shifts to 1558 cm⁻¹ when C¹⁸O₂/N₂ is introduced, indicating the incorporation of ¹⁸O into bidentate carbonate species, in good agreement with the previous report (Liao et al., 2002). However, no shift of IR features at 1269, 1219, and 1159 cm⁻¹, assigned to (bi)sulfate species on TiO₂ particles, were observed throughout the reaction. This implies that the oxygen transfer path does not account for the rapid SO₂

oxidation on particles of concern.

In light of the above analysis, the electron transfer might be a plausible pathway to explain the fast oxidation within the reaction system. DFT calculations provide an accessible approach to study the electron transfer pathway. The result in Fig. 6b shows SO_3^{-1} formation is a SET process of CO_3^{-1} and SO_3^{2-1} , where O atom in SO_3^{2-1} transfers an electron to O atom in CO_3^{-1} to

- form SO_3^{-} and CO_3^{2-} . This SET reaction is a thermodynamically favorable process, with the difference of Gibbs free energy between reactant and product lying at -24.09 kcal mol⁻¹. However, we noted that the insufficient O_2 supply in aqueous media may be an underlying constraint to the proposed CO_3^{-} -initiated SO_2 oxidation pathway. Hence, we estimated both oxygen consumption and supply rates, and oxygen supply flux can be several orders of magnitude larger than corresponding consumption (see additional details in the supplementary text 5). Therefore, oxygen is sufficient in the reaction, allowing the
- 300 considered chain reactions to continually proceed. Taken above results and discussions together, the following reactions are proposed accordingly (Eqs. 5-8):

$$CO_3^- + SO_3^{2-} \rightarrow CO_3^{2+} SO_3^{--}$$
 (Eq. 5)

$$SO_3 + O_2 \rightarrow SO_5$$
 (Eq. 6)

$$\operatorname{SO}_5^- + \operatorname{SO}_3^2 \to \operatorname{SO}_4^- + \operatorname{SO}_4^{-2}$$
 (Eq. 7)

$$305 \quad SO_4^{-} + SO_3^{-} \rightarrow SO_4^{-} + SO_3^{-} \qquad (Eq. 8)$$

Another important issue needs to be addressed as well. SO_3^{-} can also be formed via the conventional reaction of $\cdot OH$ and SO_3^{2-} (Eqs. 9) and this process is also considered.

$$\cdot OH + SO_3^{2-} \rightarrow SO_3^{-+} OH^{-}$$
(Eq. 9)



Figure 6. (a) *In situ* DRIFTS of heterogeneous reaction of SO₂ on the TiO₂ particles for 2 and 60 min after being exposed to C¹⁶⁽¹⁸⁾O₂/N₂ for 20 min under irradiation. (b) Reaction pathway of interaction between hydroxyl radical (·OH) and sulfite (SO₃²⁻) and (c) Interaction between carbonate radical (CO₃⁻⁻) and sulfite (SO₃²⁻) through the SET process at the CCSD (T)-F12/cc-PVDZ-F12//M06-2X/6-311++G (3df, 3pd) level and ΔG_0^{SET} represents the difference in Gibbs free energy between reactant and product. The white, black, yellow, and red spheres represent H, C, S, and O atoms, respectively. In order to visualize the variation of surface products in oxygen isotope experiments (panel a),

315 DRIFTS features of these concerned species were plotted in dark colors. For interpretation of the references to color in the legends of panels b and c, the reader is referred to the Web version of this article.

In this SET process, electron donor SO_3^{2-} reacts spontaneously with electron acceptor $\cdot OH$ (Fig. 6c) and the calculated activation free energy barrier ΔG^{\neq}_{SET} for this SET reaction is 2.50 kcal mol⁻¹. Hence, the reaction process of $\cdot OH$ with SO_3^{2-} is diffusion-controlled, and the total rate constant k_{SET-2} was calculated to be $7.12 \times 10^9 \text{ M}^{-1} \text{s}^{-1}$. In comparison, the rate constant k_{SET-1} of the diffusion-controlled SET process for CO_3^{--} and SO_3^{2-} was estimated to be $7.42 \times 10^9 \text{ M}^{-1} \text{s}^{-1}$. Despite a slight net

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increase of the rate, the distinguishable concentration of CO_3 and OH should also be taken into account for the rate comparison in varied reaction paths. To visualize the difference, relative rates were calculated according to Eq. 10:

$$r = \frac{v_{\text{CO}_3^{-+} + \text{SO}_3^{2^-}}}{v_{\text{OH}^+ + \text{SO}_3^{2^-}}} = \frac{k_{\text{SET}^{-1}}[\text{CO}_3^{--}][\text{SO}_3^{-2^-}]}{k_{\text{SET}^{-2}}[\text{OH}][\text{SO}_3^{-2^-}]}$$
(Eq.10)

- 325 Where *r* is the ratio of two reaction rates, $[CO_3^{--}]$, $[SO_3^{2-}]$, and $[\cdot OH]$ refer to the concentration of corresponding reactants. Previous literature suggests the concentration of carbonate radicals is able to show two orders of magnitude higher than that of hydroxyl radical at the surface of the water under solar irradiation (Chandrasekaran and Thomas, 1983; Goldstein et al., 2001; Shafirovich et al., 2001). An aqueous medium that attaches to particle surfaces offers an ideal environment for accumulating carbonate radicals. Consequently, concentrations of CO_3^{--} and $\cdot OH$ were set at the range from 1.0×10^{-10} to $1 \times$
- 330 10^{-12} mol L⁻¹ and from 1.0×10^{-12} to 1×10^{-14} mol L⁻¹ (Sulzberger et al., 1997) and *r* value could thus reach to 1.04×10^4 at most (Fig. S14). As a result, we speculated that the formation pathway of SO₃⁻⁻ via interaction between CO₃⁻⁻ and SO₃²⁻ is a more effective route, corresponding well with experimental results.

In addition to the pathway launched by photo-generated holes, we also considered the sink of photo-generated electrons. In our reaction system, O_2 is believed to be an electron trap and produce the superoxide radical ions (O_2^{-}) , which is reported to

- 335 play a non-negligible role in sulfate formation (Shang et al., 2010) and should be taken into account to give a whole picture of reaction scheme in triggering sulfate formation on the surface of TiO₂-containing mineral dust particles. *p*-benzoquinone is a commonly-used O_2^{--} scavenger for trapping the O_2^{--} radical ions (Yan et al., 2018). Our supplementary data shows that adding an excess amount of *p*-benzoquinone into TiO₂ particles reduces the sulfate yield by 32 % along with the appearance of sulfite ions over TiO₂ particles upon exposure of SO₂ (Fig.S12). Notably, the decrease in sulfate yield by around 30 % in the presence
- of O_2^{--} scavenger *p*-benzoquinone is almost complementary to that added with •OH scavenger using TBA (70 %), pointing toward a minor sulfate formation pathway contributed by O_2^{--} relative to the major pathway by CO_3^{--} when carbonate ions are presented to efficiently capture •OH ions. Following Shang's work (Shang et al., 2010), O_2^{--} involved SO₂ oxidation can be given as Eqs. 11-13:

$$e^++O_2 \rightarrow O_2^{-+}$$
 (Eq.11)

$$345 \quad SO_2 + O_2 \xrightarrow{\bullet} SO_3 + O^{-\bullet} \tag{Eq. 12}$$

$$SO_3+H_2O\rightarrow H_2SO_4$$
 (Eq.13)

Where intermediates SO_3 formed via the interaction between SO_2 and O_2^- subsequently couple with water molecules to produce sulfate species as a final product. pH is an important factor within aqueous chemical reaction processes and is likely to alter the dominated regime for sulfate production. Yet so far adjusting the pH of particle surfaces is quite tough, and

- 350 exploring the role of dust surface pH in the reactivity of CO_3^{-1} is not easily achieved. Notwithstanding, the increase of pH in TiO₂ suspension was observed to promote the production of CO_3^{-1} , further strengthening the oxidation capability of dust particles (Fig. S11). In contrast, decreasing pH is expected to reduce the yield of CO_3^{-1} since the reaction rate of CO_3^{-2} with \cdot OH is nearly two orders of magnitude higher than that with HCO_3^{-1} . On this basis, the question arises whether the surface pH of mineral dust can be sustained to maintain fast SO₂ oxidation triggered by CO_3^{-1} in the typical lifespan of mineral dust.
- 355 Considering this, we thus plotted the heterogeneous sulfate production over TiO₂ and TiO₂+CaCO₃ particles versus equivalent exposure time (Fig. S15). Clearly, the sulfate yield builds up steadily during the two-week equivalent exposure time (see a more detailed discussion on determining equivalent exposure time in supplementary text 6), suggesting that the regime of CO₃⁻⁻ initiated SO₂ oxidation over TiO₂ and TiO₂+CaCO₃ particles are slightly affected by the possible decrease of surface pH due to accumulation of sulfate production over entire reaction course. In the atmosphere, the lifetime of mineral dust particles ranges from several days to weeks (Bauer and Koch, 2005), and the equivalent exposure time considered in this study (nearly 2 weeks) falls right within the characteristic lifespan range of mineral dust particles. This leads us to deduce that

persistent growth of sulfate shows a negligible effect on CO₃⁻ initiated SO₂ oxidation scheme proposed in this work.

Additionally, dust particles are reported to eject the radical ions from the surface under solar light irradiation, showing a non-negligible contribution to sulfate aerosol formation (Chen et al., 2021; Dupart et al., 2012), as described as:

365 Mineral Dust + $hv \rightarrow ROS$ (g)

(Eq.14)

ROS (g)+ humidified Air+ SO₂ \rightarrow Sulfate(g)

Where ROS (g) stands for the active intermediates in the gas phase. Over 400 ppm of CO_2 is universal in the atmosphere, and it is expected to form (bi)carbonate ions once enters into the atmospheric aqueous media such as aerosol water, cloud droplets as well as fog environment. Bi(carbonate) ions are then prone to react with hydroxyl radical ions to form carbonate radicals.

370 Following this line of reasoning, we attempt to monitor the plausible gas ROS species that are formed in the presence of CO₂ (see a detailed discussion about the measurement approach and experimental setup in supplementary text 7 and Fig. S16).

When CO_2 (atmospheric relevant concentration) is introduced into the homemade flow-cell chamber, with the intervening gap between TiO_2 -coated film and probe molecule solution fixing at nearly 2 mm, and the short distance of which allows possible gaseous ROS to diffuse and react with aniline molecular (None, 2013). An increased degradation rate of aniline was

- 375 seen, which can be attributed to the generation of active carbonate radical ions (Fig. S17). The maximum concentration of steady-state CO_3^- radical ions supplied by partition processes between gas phase and solid-liquid phases (humified dust particles) was determined to be 1.39×10^{-13} M for TiO₂+Air+CO₂ system, which is over one order of magnitudes higher than that of \cdot OH for TiO₂+Air+system (2.15 ×10⁻¹⁵ M). This observation matches with the earlier study where the concentration of carbonate radical can be two orders of magnitudes than \cdot OH over the water surface (Sulzberger et al., 1997).
- 380 The above results suggest that the photochemistry that involves carbonate ions, more precisely CO_3^{--} radical, increases sulfate formation. This finding broadens the prevailing view that acceleration of SO₂ oxidation over the carbonate salt is merely due to the favorable neutralization of H₂SO₄ over an alkaline surface. To be important, upon irradiation active component TiO₂ in mineral dust produce carbonate radical in the gas phase when CO_3^{--} precursor CO_2 is presented, therefore potentially promoting sulfate aerosol formation in the atmosphere. Overall, it could be speculated that carbonate radical ions strengthen 385 the oxidative capability of TiO₂-containing mineral dust particles, and consequently accelerate SO₂ oxidation.

3.2.4 Field Measurements of Sulfate and (Bi)carbonate Ions.

Complement field sampling and analysis were further conducted to examine our hypothesis that intermediates CO_3^- may play role in secondary sulfate formation in the atmosphere. We first considered the meteorological condition wind speed, which is an important parameter determining whether the local chemical process gains importance in affecting secondary sulfate formation. Meteorological information was collected from the open-access database (<u>https://www.aqistudy.cn/</u>). During the sampling period, the wind scale mainly varies from 0 to 1, corresponding to the wind speed ranging from 0 to 1.5 m s⁻¹ (Fig. S18). All plots shown in Fig. S18 give rise to a statistical wind speed of 0.76 ± 0.73 , which represents the weak dispersion of pollutants at low wind speed (not exceeding 2.5 m s⁻¹)(Witkowska et al., 2016; Wu et al., 2020), indicating that local source is a dominant contributor to local air pollution.

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It is generally accepted that under stagnant meteorological conditions (wind speed < 1.5 m s^{-1}), for the coarse-mode (2.5 μ m ~ 10 μ m) of sulfate, the heterogeneous reaction of SO₂ on the dust surfaces is believed to be a major contributor (Liu et al., 2017). This correlates to the fact that a large mass fraction of mineral dust is abundant in coarse-model particulate matter (PM) (Fang et al., 2017; Miller-Schulze et al., 2015), in which TiO₂ was found at mass mixing ratios ranging from 0.1 to 10 %

depending on the exact location where particles were uplifted (Chen et al., 2012; Hanisch and Crowley, 2003). Therefore, PM

- 400 with relatively larger size dimensions is expected to contribute to secondary sulfate formation via heterogeneous reactions. which is supported by the recent field study where carbonate fraction of coarse PM is evidenced to promote secondary sulfate production (Song et al., 2018). Considering this, rather than determine the concentration of water-soluble ions in all stages, more attention was paid to PM collected in stages 1-4 (particles with their dimension $\ge 3.3 \ \mu m$). As (bi)carbonate ions are known as key precursors in producing CO_3^{-1} and accelerate sulfate formation, quantifications of those relevant water-soluble ions were thus conducted (supplement text 23 and 24).
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We further considered the relationships between sulfate ions and (bi)carbonate ions by means of linear regression analysis. However, under the low wind speed (0.76 \pm 0.73), correlation coefficients R² obtained for the relationship between bi(carbonate) and sulfate ions are not promising, 0.56 (sulfate vs carbonate) and 0.61 (sulfate vs bicarbonate) for PM_{3.3}-PM_{9.0} during daytime hours. A plausible explanation is that although less significant, local primary emission source also brings bias

- and uncertainty to the correlation analysis. Shanghai is a coastal city, and sulfate species such as K₂SO₄ and Na₂SO₄ from 410 the sea salt contribute to the local sulfate emission as well (Long et al., 2014). On the other hand, this novel SO_2 oxidation channel is in the infant stage, and only active mineral dust components have been considered in this work whereas other components found in the coarse mode of PM such as organic matter, elemental carbon as well as sea salt (Cheung et al., 2011) are likely to involve this mechanism and alter the response of sulfate yield to SO₂ heterogeneous uptake. In addition, the water-
- soluble ions determined in these samples (relatively small size) may not come from the net contribution of heterogenous 415 reaction processes in absolute day-time and night-time periods. Some of the undesired processes that take place during day(nigh)-night(day) shifts may also contribute to the prodcution of sulfate ions in separate sampling hours, thus reducing the correlation coefficients.
- For those large particles (LP), that refer to the particles with a diameter large than 9 µm in this work, sulfate ions show a rather weak or even no correlation to (bi)carbonate ions during the night-time and day-time hours (Fig. S19). This is likely due 420 to the short lifetime of LP. Generally, the aerosol lifetime is on the order of less than an hour to days (Koelemeijer et al., 2006), highly depending on particle size. For example, the lifetime of PM_{10} ranges from minutes to hours, and its travel distance, in general, is less than 10 km (Agustine et al., 2018). As a consequence, secondary sulfate formation through chemical reaction over LP is not significant with respect to in situ emissions. When PM downsizes to 2.5 µm, PM_{2.5} has a lifetime prolonged to
- nearly one day or longer (Liu et al., 2020b). Therefore, $PM_{3,3}$ -PM_{9,0} are expected to have a relatively long lifetime, on the order 425 of several hours, which enables the heterogeneous reaction process to become a more important contributor to overall sulfate ions measured in $PM_{3,3}$ -PM_{9,0} than that in PM_{>9,0}. This is supported by our observations where during the daytime hours the correlation coefficients for $PM_{3,3}$ - $PM_{9,0}$, i.e. 0.56 (sulfate vs carbonate) and 0.61 (sulfate vs bicarbonate), are higher than that of $PM_{\geq 9.0}$, i.e. 0.489 (sulfate vs carbonate) and 0.36 (sulfate vs bicarbonate), respectively. Similarly, higher correlation
- 430 coefficients are also observed for $PM_{3,3}$ -PM_{9,0} than PM>_{9,0} in the sample collected during the nighttime periods.



Figure 7. Field observation for the relationship between carbonate and sulfate ions during day-time and night-time hours. Linear relationship analyses for measured sulfate ions and estimated carbonate ions (**a**) and for measured sulfate ions and estimated bicarbonate ions (**b**) during the day-time and night-time hours, with particle sizes of PM ranging from 3.3 to 9 μ m.

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While we note that the correlation coefficients between sulfate and (bi)carbonate are not promising in this work, groundbased field measurements of sulfate and (bi)carbonate ions shed light on their distinct correlations during the daytime and nighttime hours. In Fig. 7 and Fig.S19, the negative correlations between the mass concentrations of sulfate ions and (bi)carbonate ions are observed in the nighttime hours, consistent with the suppression of sulfate formation by CO_2 in the dark experiments. This is also supported by our previous lab study where CO_2 -derived (bi)carbonate species are demonstrated to block the active sites for yielding sulfate over mineral dust proxy aluminum oxide (Liu et al., 2020a). Instead, positive correlations were seen for those ions within PM sampled during the daytime hours regardless of size ranges and carbonate types (HCO₃⁻/CO₃²⁻). This matches with the scenarios in which sulfate production upon irradiation in the presence of (bi) carbonate ions is increased over both model and authentic dust particles. Except the case (nighttime period, size larger than 9

- 445 μ m), most of the significance *P* values for their correlations were smaller than 0.1, with significance *P* values below 0.5 determined for bicarbonate vs sulfate, implying the plausible underlying connection between sulfate and (bi)carbonate ions. In fact, preceding ground-based observations of highly correlated relationship between Ca²⁺ and SO₄²⁻ water-soluble ions (Liu et al., 2020b) during the carbonate-enriched dust storm episodes, together with persistent reports on the significant role of photochemical channels in increasing the sulfate concentration during the daytime (Kim et al., 2017; Wei et al., 2019; Wu et al., 2019; W
- 450 al., 2017) indirectly reflects the possibility of accelerated SO₂ oxidation triggered by photo-generated active intermediates associated with carbonate species.

Overall, this is the first time that relationships between those ions are explored separately in these two periods. Taken together, carbonate radical is likely to promote sulfate production in the atmosphere during daytime hours. Detailed and

systematic SO₂ oxidation channel triggered by CO₃⁻ needs further investigations to enable a better interpretation of correlations

455 between these inorganic ions at the given meteorological conditions of sampling and physico-chemical properties of PM.

4. Conclusion

On the basis of the experimental and theoretical results derived from this work, we for the first time propose a novel reaction channel for fast SO₂ oxidation over mineral dust particles due to the formation of carbonate radical ions. A schematic chart for the sulfate formation in the presence of carbonate radicals upon solar irradiation or bi(carbonate) ions under dark conditions is summarized and elucidated in Fig. 8. During the night-time hours at 298 K (ambient temperature) CO₂-derived (bi)carbonate species are prone to have a slightly negative effect on sulfate formation due to the competitive adsorption between CO₂ and SO₂. For alkaline carbonate salt, it favors sulfate formation through the neutralization process. On the other hand, during the day-time hours, both CO₂-derived (bi)carbonate species and carbonate salt work as the precursor of CO₃⁻⁻, which promotes sulfate formation. Especially, uptake coefficients for carbonate salt containing mineral dust can be increased by 17 times, which is more pronounced than the increase due to the neutralization regime in the dark condition. Consistent with the findings reported in the earlier studies (Chen et al., 2021; Dupart et al., 2012), we observed the production of gas-phase CO₃⁻⁻ ions when mineral dust particles are irradiated in the presence of CO₂ (atmospherically-relevant concentration 400 ppm). This observation implies that the increased sulfate yield in part comes from promoted gas-phase secondary sulfate aerosol triggered by CO₃⁻⁻ (g).

470 By means of ROS scavenger experiments, direct observation of carbonate radical using NTAS analysis, oxygen isotope assay, ESR spectra as well as DFT calculations, CO₃⁻⁻-initiated S(IV) oxidation involving single electron transfer process are elucidated. While carbonate radical ions are mainly responsible for rapid sulfate formation, superoxide radical ions are likely to serve as a minor pathway over TiO₂-containing mineral dust particles. In addition, a weak correlation between sulfate ions and (bi)carbonate ions observed for PM_{3,3}-PM_{9,0} in this work correlates to non-chemical primary emission and complicated nature of CO₃⁻⁻ regime of sulfate production in the atmosphere. Nevertheless, complement field sampling of ambient PM and analysis of sulfate and (bi)carbonate ions in this study unfold their distinct correlations during the daytime and nighttime hours, these two tendencies of which agrees with the experimental observations.

In this work, only atmospheric secondary sulfate formation was considered, whereas the oxidation of primary organic species yet has not been investigated. In fact, carbonate radical ions are prone to rapidly react with electron-rich organics amines (Stenman et al., 2003; Yan et al., 2019) as well as phenol (Busset et al., 2007; Xiong et al., 2016), and it may potentially serve as the key oxidants that drive the fast formation of SOA in the atmosphere. Besides, observation of strengthened photochemistry launched by carbonate radicals suggests that such chemistry may be amplified on atmospherically relevant reactions that occur in cloud droplets as well as fog water where they often contain hydroxyl radicals and water-soluble (bi)carbonate ions.

485 To be important, carbonate radical ions are observed to be formed in the gas phase in the atmospherically relevant CO₂ concentration (400 ppm) when mineral dust proxies are irradiated. This will help the formation of external sulfate aerosol formation. Since both sulfate aerosol and CO₂ are well known to affect the radiation budget and solar energy balance on the earth (Cheung et al., 2011; Möller, 1964), their overall influence on the global climate considering the increased yield of sulfate aerosol triggered by CO₂, the precursor of carbonate radical, needs further investigation. Therefore, our study highlights the necessity for a comprehensive understanding of the CO₃⁻⁻ relevant chemistry in the underlying impacts of fine PM concentration, human health, and climate. All these assumptions need to be investigated in further detail. This study provides the first indication that carbonate radical not only plays a role as a marginal intermediate in tropospheric anion chemistry but also as a strong oxidant for surfacial processing of trace gas in the atmosphere.



495 **Figure 8.** Schematic of the sulfate formation in the presence and absence of carbonate radical. Noting that g and ad represent gas-phase and adsorbed carbonate radical ions, respectively.

Data availability. The data that support the results are available from the corresponding author upon request.

500 *Author contributions*. Y.L., Y.D. and L.Z. initially proposed the idea; Y.L. and Y.D. designed and performed most of the experiments; J.L. performed DFT calculations; Y.L., X.Z. and T.W. contributed to field samplings and data analysis; K.L., K.G., A.B., I.N., X.Z., C.G., and L.Z. provided suggestions on the experiments and paper writing; All authors wrote the manuscript.

Competing interests. The authors declare that they have no conflict of interest.

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