

## Response to reviewer #1

MS no: ACP-2021-521

We thank the reviewer for the constructive comments and suggestions that helped improve the manuscript. This document outlines the reviewer's comments (in **bold-blue**), followed by the author's responses and changes made in the revised manuscript (in *italics*). A marked-up version of the manuscript showing the revisions is appended to this response file.

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### **Comment:**

**Single scattering albedo (SSA) is a very important parameter in assessing the radiative impact of aerosols and on which there is meager data globally. The paper is a welcome addition to the aerosol literature in this regard. The authors have made use of satellite data of CERES and MODIS in obtaining global maps of SSA based on the concept of critical optical depth. They have presented the maps for different seasons as well considering a four year period. It is hoped such maps will be generated on annual basis subject to sufficient data availability.**

**The authors made a very clear presentation of the method of analysis including the error estimates. On the whole the paper will be a very important contribution to the area of aerosol radiative impact.**

We greatly appreciate and thank the reviewer for the summary evaluation, positive recommendations, and valuable feedback. Yes, on completion of peer-review, we intend to publish this dataset online and generate datasets for extended periods for future studies.

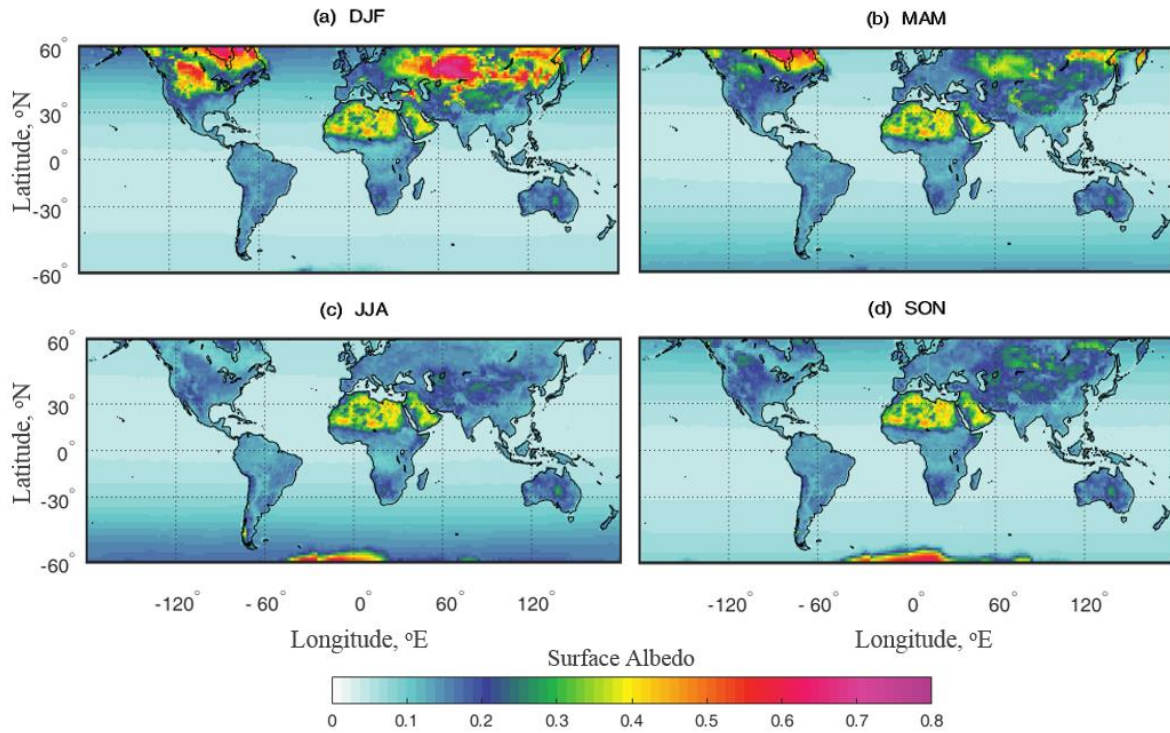
### **Comment:**

#### **Specific comments/suggestions:**

- 1. As described in the paper, the surface albedo is an important parameter in the SSA estimation. So the surface albedo maps for different seasons also should be given as in Fig 5 along with similar data for different seasons and regions in Table 1. This would greatly help in the discussion of the results.**

**Response:** This is indeed an insightful suggestion. We have added seasonal maps of surface albedo as Figure S2. And seasonal mean SSA values with standard deviation for each season have been listed in Table S2 for the various regions of interest.

**Additions to the supplementary file:**



**Figure S3.** Seasonal mean shortwave-integrated surface albedo from CERES

**Table S4.** Shortwave integrated seasonal mean surface albedo from CERES over regions of interest. Details of these regions are given in Table S1 and Fig. S1

Region	Surface Albedo			
	DJF	MAM	JJA	SON
Canadian Boreal Forest	0.36 ± 0.13	0.30 ± 0.12	0.12 ± 0.03	0.16 ± 0.05
Russian Boreal Forest	0.37 ± 0.10	0.27 ± 0.08	0.13 ± 0.02	0.20 ± 0.05
South African Forest	0.12 ± 0.01	0.13 ± 0.01	0.12 ± 0.02	0.13 ± 0.01
Amazon Forest	0.14 ± 0.01	0.14 ± 0.01	0.13 ± 0.02	0.14 ± 0.02
North East Atlantic	0.06 ± 0.01	0.05 ± 0.01	0.05 ± 0.01	0.05 ± 0.01
South East Atlantic	0.05 ± 0.01	0.05 ± 0.01	0.05 ± 0.01	0.05 ± 0.01
Eastern Pacific	0.05 ± 0.01	0.05 ± 0.00	0.05 ± 0.01	0.05 ± 0.00
Sahara	0.35 ± 0.06	0.34 ± 0.06	0.34 ± 0.06	0.34 ± 0.06
Indo Gangetic Plain	0.13 ± 0.02	0.13 ± 0.02	0.14 ± 0.02	0.13 ± 0.01
Eastern China	0.13 ± 0.04	0.13 ± 0.03	0.13 ± 0.03	0.13 ± 0.03

Arabian Sea	0.06 ± 0.01	0.05 ± 0.01	0.05 ± 0.02	0.05 ± 0.01
Bay of Bengal	0.05 ± 0.01	0.05 ± 0.01	0.05 ± 0.01	0.05 ± 0.01

**Comment:**

**Specific comments/suggestions:**

2. A brief description of the aerosol models used in the RT calculations should also be given.

Thank you for the comment. We have included the aerosol model specifications in the supplementary file.

**Additions to the supplementary file:**

**Table S5:** Normalized extinction coefficient of the aerosol model

$\lambda$ ( $\mu\text{m}$ )	$\text{Ext}_{\text{norm}}$	$\lambda$ ( $\mu\text{m}$ )	$\text{Ext}_{\text{norm}}$	$\lambda$ ( $\mu\text{m}$ )	$\text{Ext}_{\text{norm}}$
0.25	1.597	0.75	0.847	3.2	0.5075
0.3	1.467	0.8	0.8202	3.39	0.5047
0.35	1.334	0.9	0.7828	3.5	0.5062
0.4	1.224	1	0.7536	3.75	0.4828
0.45	1.135	1.25	0.7038	4	0.4629
0.5	1.061	1.5	0.6706	4.5	0.4395
0.55	1	1.75	0.6349	5	0.4103
0.6	0.9505	2	0.5883		
0.65	0.9106	2.5	0.4905		
0.7	0.8757	3	0.491		

**Table S6:** Phase function of the aerosol model (continued into Table S7)

$\lambda$ ( $\mu\text{m}$ )	Streams							
	1	2	3	4	5	6	7	8
0.25	0.754	0.606	0.473	0.397	0.342	0.307	0.283	0.265
0.3	0.738	0.589	0.452	0.379	0.325	0.293	0.270	0.254
0.35	0.738	0.592	0.456	0.386	0.333	0.303	0.279	0.264
0.4	0.741	0.598	0.463	0.395	0.343	0.313	0.290	0.275
0.45	0.743	0.602	0.468	0.403	0.351	0.323	0.299	0.284
0.5	0.746	0.607	0.474	0.411	0.359	0.331	0.308	0.292
0.55	0.748	0.611	0.478	0.416	0.364	0.337	0.313	0.297
0.6	0.749	0.615	0.481	0.421	0.368	0.342	0.317	0.301
0.65	0.750	0.618	0.485	0.426	0.373	0.347	0.321	0.305
0.7	0.751	0.620	0.487	0.429	0.376	0.350	0.323	0.306
0.75	0.752	0.623	0.490	0.433	0.378	0.352	0.325	0.308
0.8	0.755	0.628	0.494	0.437	0.382	0.355	0.327	0.310
0.9	0.756	0.631	0.496	0.440	0.383	0.356	0.326	0.308
1	0.756	0.632	0.496	0.440	0.382	0.354	0.323	0.304
1.25	0.766	0.643	0.505	0.442	0.380	0.346	0.314	0.291

1.5	0.777	0.651	0.512	0.441	0.376	0.337	0.302	0.276
1.75	0.798	0.673	0.536	0.455	0.385	0.339	0.300	0.271
2	0.826	0.707	0.577	0.491	0.415	0.362	0.316	0.282
2.5	0.858	0.750	0.636	0.552	0.476	0.418	0.365	0.323
3	0.871	0.765	0.662	0.578	0.505	0.444	0.391	0.346
3.2	0.836	0.708	0.584	0.491	0.414	0.354	0.304	0.264
3.39	0.818	0.682	0.548	0.453	0.375	0.317	0.270	0.233
3.5	0.808	0.670	0.530	0.434	0.356	0.299	0.253	0.217
3.75	0.805	0.667	0.524	0.429	0.349	0.292	0.246	0.210
4	0.797	0.660	0.513	0.421	0.340	0.284	0.238	0.202
4.5	0.795	0.655	0.507	0.413	0.331	0.275	0.228	0.192
5	0.808	0.663	0.520	0.420	0.338	0.278	0.230	0.192

**Table S7:** Phase function of aerosol model

$\lambda$ ( $\mu\text{m}$ )	Streams							
	9	10	11	12	13	14	15	16
0.25	0.252	0.241	0.233	0.226	0.219	0.214	0.209	0.204
0.3	0.241	0.232	0.224	0.217	0.211	0.205	0.200	0.196
0.35	0.251	0.242	0.233	0.226	0.219	0.214	0.208	0.203
0.4	0.262	0.252	0.243	0.235	0.228	0.222	0.216	0.210
0.45	0.270	0.260	0.251	0.242	0.235	0.228	0.221	0.215
0.5	0.278	0.267	0.257	0.248	0.240	0.233	0.226	0.219
0.55	0.283	0.271	0.261	0.251	0.243	0.235	0.227	0.220
0.6	0.286	0.274	0.263	0.253	0.244	0.235	0.228	0.220
0.65	0.289	0.277	0.265	0.255	0.245	0.236	0.228	0.220
0.7	0.290	0.277	0.265	0.254	0.244	0.235	0.226	0.218
0.75	0.291	0.277	0.265	0.253	0.243	0.233	0.225	0.216
0.8	0.292	0.278	0.265	0.253	0.242	0.232	0.223	0.214
0.9	0.289	0.274	0.261	0.248	0.237	0.226	0.217	0.208
1	0.284	0.269	0.254	0.241	0.230	0.219	0.209	0.200
1.25	0.271	0.253	0.238	0.224	0.212	0.200	0.190	0.180
1.5	0.255	0.236	0.220	0.205	0.193	0.181	0.171	0.161
1.75	0.246	0.226	0.208	0.193	0.180	0.168	0.157	0.147
2	0.252	0.229	0.208	0.191	0.176	0.163	0.151	0.141
2.5	0.287	0.257	0.231	0.208	0.189	0.172	0.157	0.144
3	0.307	0.274	0.245	0.220	0.198	0.179	0.162	0.148
3.2	0.231	0.203	0.180	0.161	0.144	0.130	0.117	0.107
3.39	0.203	0.178	0.157	0.140	0.125	0.113	0.102	0.093
3.5	0.188	0.165	0.146	0.130	0.116	0.104	0.094	0.085
3.75	0.181	0.157	0.139	0.122	0.109	0.097	0.088	0.079
4	0.174	0.150	0.132	0.116	0.103	0.091	0.082	0.073
4.5	0.164	0.141	0.122	0.107	0.094	0.083	0.073	0.066
5	0.163	0.139	0.120	0.103	0.090	0.079	0.070	0.062

**Comment:**

**Minor:** Page 13, line 8 from top: the value is 0.83 and not 0.81 (Table 1)

Thank you for pointing this out. The table has now been shifted to supplementary file and the ‘results and discussion’ section is now completely revised based on reviewer #2’s comments.

## Response to reviewer #2

MS no: ACP-2021-521

We thank the reviewer for the constructive comments and suggestions that helped improve the manuscript. We have considered each comment carefully and revised the manuscript accordingly. This document outlines the reviewer's comments (in **bold-blue**), followed by the author's responses and changes made in the revised manuscript. A marked-up version of the manuscript showing the revisions is appended to this response file.

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### **Comment:**

**This paper combines daily CERES flux retrievals and MODIS aerosol retrievals to estimate aerosol single scattering albedo (SSA) at 550 nm. SSA is, after aerosol optical depth (AOD), the key parameter determining aerosols' radiative effect, but is difficult to retrieve well from most spaceborne measurements. The authors expand the application of a technique called "critical optical depth" they have developed before to a global scale. There is a brief comparison to airborne data, and to similar SSA maps available from OMI.**

**The study is in scope for the journal, though is also a close fit for AMT. It is fairly clearly presented. My main criticism is that the numerous uncertainties in the technique are glossed over and the reader is instead presented (in the abstract and conclusions) with the claim that the global uncertainty is about 0.03. There is no real analysis to back up this number and it seems to be based on limited airborne measurements over India and surrounding oceans. The manuscript is not very long and I think that the paper would benefit from a much more thorough and honest discussion and quantification of uncertainty sources. Otherwise an inexperienced reader might believe the problem of determining SSA from space is essentially solved.**

**To that end, I recommend major revisions, and would like to review the revised version. I think the work is valuable but not yet at ACP quality.**

**Response:** We appreciate and thank the reviewer for the summary evaluation and valuable feedback. As suggested, we have included a more detailed uncertainty analysis and

comparisons with other datasets. These additional results, incorporated following the reviewer's comments, have vastly improved the manuscript.

**Comment:**

The paper is missing references to the existing literature. For example, a lot of similar work has been done framed in terms of “critical reflectance” or “albedo” rather than “optical depth”. The basic idea is the same, i.e. find a value of one parameter (surface/aerosol) where the top of atmosphere signal is invariant to changes in the other. Examples include Seidel and Popp (2012): <https://amt.copernicus.org/articles/5/1653/2012/> and Wells et al (2012): <https://doi.org/10.1029/2011JD016891> The authors should acknowledge and discuss the relative merits of other work using the same basic technique like this.

**Response:** Comparison with Kaufman's critical reflectance method, which has a similar basic technique, is a valuable discussion. Thank you for bringing this point. The details and references to the existing literature on the critical reflectance method have been added. Their relative merits are also discussed.

**Additions to the revised manuscript:**

Page 2 Line 5 to 11: Fraser and Kaufman., 1985 developed a critical surface reflectance method to retrieve SSA using satellite data. Their method is based on radiative transfer simulations, which showed a particular surface reflectance for which the top of atmosphere albedo is independent of AOD. Upward radiances between a clear and a hazy day over a varying surface reflectance region are used, along with radiative transfer simulations, to derive SSA. This method has been widely applied to data from various satellites to derive SSA over particular regions (Kaufman, 1987; Kaufman et al., 1990, 2001; Zhu et al., 2011; Wells et al., 2012). Seidel and Popp., 2012 have done extensive studies on the method's sensitivity to various parameters.

Page 2 Line 24 to Page 3 Line 2: The “critical optical depth” method developed in this research paper shares a similar concept to the critical surface reflectance method (Fraser and Kaufman., 1985). For a particular parameter (such as surface reflectance or optical depth), there exists a critical value at which the top of atmosphere albedo can be considered independent of variations in that parameter. Both the methods retrieve SSA by parameterizing the critical value as a function of SSA using radiative transfer

simulations. The critical reflectance method requires two-days data and large variations in surface reflectance over the region. It's suitable for retrieving daily SSA for a particular region. Whereas the critical optical method developed in this paper is suitable for retrieving monthly or seasonal global maps of SSA.

### References added

Kaufman, Y. J., Fraser, R. S., and Ferrare, R. A.: Satellite measurements of large-scale air pollution: methods, *J. Geophys. Res.*, 95, 9895–9909, <https://doi.org/10.1029/JD095iD07p09895>, 1990.

Kaufman, Y. J., Tanré, D., Dubovik, O., Karnieli, A., and Remer, L. A.: Absorption of sunlight by dust as inferred from satellite and ground-based remote sensing, *Geophys. Res. Lett.*, 28, 1479–1482, <https://doi.org/10.1029/2000GL012647>, 2001.

Seidel, F. C. and Popp, C.: Critical surface albedo and its implications to aerosol remote sensing, *Atmos. Meas. Tech.*, 5, 1653–1665, <https://doi.org/10.5194/AMT-5-1653-2012>, 2012.

Wells, K. C., Martins, J. V., Remer, L. A., Kreidenweis, S. M., and Stephens, G. L.: Critical reflectance derived from MODIS: Application for the retrieval of aerosol absorption over desert regions, *J. Geophys. Res. Atmos.*, 117, 3202, <https://doi.org/10.1029/2011JD016891>, 2012.

### Comment:

The uncertainty discussion really needs to be strengthened. There are claims of 0.03 throughout the paper but they are really not supported. The authors do not really acknowledge that e.g. aerosol vertical location matters a lot as well: you can get quite a different forcing if the aerosols change height due to interactions with Rayleigh scattering (which depends on pressure). This is well established by e.g. the OMI and combined UV-vis work the authors cite during the paper. Other key uncertainty sources are inconsistencies between the aerosol and surface properties assumed by the MODIS and CERES retrievals with each other and with the OPAC-based SBDART calculations. A further is the possibility of variation on scales of the regions used for the linear fitting process; the residuals on the fit (uncertainty on the intercept) would be one easy way to incorporate this effect. There are doubtless others as well. I think the authors need to list the potential uncertainty sources and try to quantify as many as possible – even if approximately – so it becomes clear which are the most important. The comparison

against airborne data is good to have but this is only a small part of the picture, and definitely not enough by itself.

**Response:** Thank you for this suggestion on strengthening the uncertainty discussion. We have included a new section on uncertainty analysis (section 5) where retrieval uncertainties due to possible perturbations in various parameters have been calculated and presented.

**Additions to the revised manuscript:**

**5 Uncertainty Analysis**

Table 1 identifies the major sources of error in the retrieval and summarizes their individual contribution. Uncertainty in the retrieved SSA was estimated by calculating retrieval sensitivities to perturbations in the possible error sources. The range of perturbation was based on published literature or reasonable assumptions for possible variations.

**Table 1.** Estimates of the uncertainty in retrieved SSA

Parameter	Input Uncertainty	Retrieval Uncertainty
Surface albedo	$\pm 0.01$	$\pm 0.03$
AOD	20% $\pm 0.05$ (land) 5% $\pm 0.03$ (ocean)	$\pm 0.02$
Angstrom exponent	$\pm 0.4$	$\pm 0.01$
Refractive index	$\pm 0.01$	$\pm 0.01$
Aerosol height	$\pm 1$ km	$\pm 0.01$
Aerosol type	Smoke vs dust	$\pm 0.01$
Residual of fit	$\pm 0.05$	$\pm 0.02$

Uncertainty in shortwave integrated surface albedo from CERES results in the maximum uncertainty in SSA of  $\pm 0.03$ . MODIS retrieved aerosol optical depth contains considerable uncertainties due to assumed aerosol models (Jeong et al., 2005). The MODIS aerosol optical depth uncertainty is 20%  $\pm 0.05$  over land (Chu et al., 2002) and 5%  $\pm 0.03$  over the ocean (Remer et al., 2002). The corresponding error in our retrieval is  $\pm 0.02$ . For a typical variation of angstrom exponent ( $\pm 0.4$ ) and refractive index



( $\pm 0.01$ ), the uncertainties vary depending on the surface albedo and are mostly around  $\pm 0.01$ .

Changes in aerosol height can vary the TOA radiances due to Rayleigh scattering interactions, which depend on pressure. Sensitivity to aerosol height was estimated by conducting a synthetic retrieval of SSA over a range of aerosol height values and perturbations from those heights. The average uncertainty observed for an aerosol height variation of  $\pm 1$  km was  $\pm 0.01$ . Many methods have been developed for detecting aerosol type, especially smoke vs. dust, to improve the uncertainties of various AOD and SSA retrievals.

Uncertainties due to possible variations on scales of the regions used for linear fitting were estimated as residuals of the fit. The uncertainty on the linear intercept is spatially dependent and is mostly around  $\pm 0.02$ , with higher values for those combinations having a slope close to zero during the regression. For highly correlated cases (i.e., correlation coefficient  $|r| > 0.5$ ), the probability of obtaining a slope close to zero is  $\sim 20\%$  over the ocean and  $< 5\%$  over land. These cases are mostly formed over regions where AOD variations are less. Regions having large variations in AOD values have lower uncertainty due to residual fit.

Overall, the algorithm is most sensitive to variations in surface albedo, followed by higher sensitivity towards AOD values used in the linear fit. The uncertainties are higher for scattering aerosols over bright surfaces and absorbing aerosols above dark surfaces. Sensitivity to water vapor is almost negligible, except in very few cases where the uncertainty is  $\pm 0.008$ . The CERES-MODIS algorithm is most effective over regions with large AOD variations and less surface albedo variations.

**Comment:**

**It is not clear to me if the derived SSA data are publicly available; I did not see a link in the paper. They ideally should be somewhere.**

**Response:** These datasets were generated as part of the author's ongoing Ph.D. research. All the datasets generated as part of the thesis work will be published online on the department's website on successful completion of the degree. For now, as mentioned in the data availability section, it will be available on request.

**Comment:**

**Specific comments:**

1. **Abstract:** As SSA is a spectrally varying quantity, the wavelength reported should be given here (550 nm). Additionally, the statement about uncertainty is basically unsupported and seems to come from the comparison against a small number of aircraft observations. I recommend that this statement is removed or made a lot weaker, e.g. “limited comparisons against airborne observations over India and surrounding oceans were generally in agreement within  $\pm 0.03$ ”. The abstract should be an honest summary of what is in the paper, not a place to hype up the work, as unfortunately many people only read abstracts and skim papers.

**Response:** We agree with the comment. The wavelength of retrieved SSA (550 nm) is now mentioned in the abstract Line 3. And the two sentences related to aircraft data and uncertainty have been replaced with the recommended statement, “Limited comparisons against airborne observations over India and surrounding oceans were generally within  $\pm 0.03$ .”

**Comment:**

2. **Page 6 line 15:** what exactly is the significance test done on here? This should be clearer. My guess is that it is on the linear correlation coefficient between the AOD and albedo difference, i.e. the authors are testing whether the probability of observing a correlation coefficient at least that large if there were truly no linear relationship between the two quantities is 0.05 or lower. Is that right?

**Response:** Yes, we are testing the significance of the correlation coefficient calculated during the linear regression between AOD and  $\Delta$ Albedo. The significance level is taken as 0.05, indicating that the risk of concluding that a correlation exists- when actually no correlation exists- is less than or equal to 5%. We have rephrased the sentence to improve the clarity.

**Additions to revised manuscript:**

Page 7 Line 15 to 17: A significance test on the correlation coefficient between AOD and  $\Delta$ Albedo is performed with a 0.05 significance level. Only those  $\tau_c$  values obtained through regressions that are statistically significant at 95% confidence level are utilized further to retrieve SSA.

**Comment:**

3. **Figure 2:** here and in the text, it is mentioned that small regression slopes mean that SSA cannot be determined well. Again, the linear model fit uncertainties (see general

comments) could be used to do this for every grid cell and associate an uncertainty. If this is not done, though, then the authors should show where and how frequently these conditions occur. It is not clear if, for example, if almost never happens or is common. In the latter case the seasonal maps shown later may have some additional sampling-related uncertainties.

**Response:** Following the reviewer's general comment #1, this point has been clarified in the revised manuscript's new Section 5 Uncertainty analysis.

**Additions to revised manuscript:**

Page 13 Line 12 to 17: Uncertainties due to possible variations on scales of the regions used for linear fitting were estimated as residuals of the fit. The uncertainty on the linear intercept is spatially dependent and is mostly around  $\pm 0.02$ , with higher values for those combinations having a slope close to zero during the regression. For highly correlated cases (i.e., correlation coefficient  $|r| > 0.5$ ), the probability of obtaining a slope close to zero is ~20% over the ocean and <5% over land. These cases are mostly formed over regions where AOD variations are less. Regions having large variations in AOD values have lower uncertainty due to residual fit.

**Comment:**

**4. Section 4: more information about the airborne measurement of SSA, including their uncertainty, is necessary. I encourage the authors to search for additional airborne data which may supplement their results from elsewhere in the world, which would strengthen the robustness of these comparisons. There have for example been NASA field campaigns through the US, south-eastern Atlantic, and Korea during this time frame, and these NASA data are publicly available (I am sure the investigators who spent considerable time collecting the data would be glad to see them used). Doubtless there are other resources as well.**

**Response:** Additional information about the aircraft measurements and their uncertainties has been included in the revised manuscript.

**Additions to the revised manuscript (shown in black font color):**

Page 14 Line 1 to 13: Babu et al. (2016), as part of RAWEX (Moorthy et al., 2016), derived SSA at 520 nm from aircraft measurements of scattering and absorption coefficients over the Indo-Gangetic Plain (IGP) and Central India during winter 2012

and spring/pre-monsoon 2013. Various measurements of aerosol properties were carried out in an instrumented Beechcraft B200 aircraft of the National Remote Sensing Centre, India. Manoj et al. (2019) estimated vertical profiles of SSA during the SWAAMI campaign conducted during monsoon (June - July) 2016 over IGP, Arabian Sea, and Bay of Bengal. Aerosol scattering coefficients were measured aboard the Facility for Airborne Atmospheric Measurements (FAAM) BAe-146 aircraft. Vaishya et al. (2018) estimated vertical profiles of SSA (520 nm) using an instrumented aircraft, Beechcraft B200, during SWAAMI-RAWEX campaign (June 2016). Instrument design and calibration were based on Anderson et al., 1996 and its application for Indian field experiments was as described by Nair et al., 2009. Uncertainties in the scattering coefficient measurement by nephelometer are  $\sim\pm 10\%$ , as reported by Anderson et al., 1996. As stated by Babu et al., 2016 uncertainties in the columnar SSA values estimated from RAWEX aircraft measurements depend mainly on instrumental uncertainties, sampling errors, and large spatial averaging.

**Response continued:**

Initially, while making the aircraft data comparisons, we had looked into other aircraft data available from various field campaigns such as ORACLES (South Eastern Atlantic), ACE-ENA (Northeastern Atlantic), DISCOVER-AQ (USA), etc. These flight datasets are available in the ASDC and ESPO data archives. They provide the raw data collected during the flights – such scattering coefficient measured by nephelometer at various latitude, longitude, and altitudes. These raw datasets need to be carefully processed considering the various instrument calibrations and experimental setup to obtain the scattering coefficient profiles over the flight track, from which the SSA profiles are computed. Further, these profiles need to be vertically integrated (also considering the flight's lat-lon variations) to obtain the columnar SSA required for comparison with the CERES-MODIS dataset. This entire work in itself would be an extensive experimental-data processing of the flight data. Carrying out these detailed computations would only provide a few datapoints for comparison with CERES-MODIS values over that region for the period.

The SWAAMI, RAWEX, and SWAAMI-RAWEX campaigns were organized and conducted by our research group. Hence the processed-datasets generated from the raw data by the experimentalists were readily available to us for comparison with the CERES-MODIS satellite data.

The suggestion provided by the reviewer to include other publicly available aircraft datasets is really a valuable point. It would surely add more points to the aircraft comparisons. But since this paper focuses more on satellite data and algorithms, performing extensive experimental calculations to obtain just a few data points would be tedious.

Instead, following reviewer's comment #6, we have included comparisons with POLDER and AERONET sites. AERONET sites were chosen based on the classification provided in Giles et al., 2012. The reviewer's suggestions to include these other datasets have significantly improved the manuscript. With these additional results now included in the revised manuscript, the addition of a few data points obtained from extensive flight data computations may not make significant improvements.

**Comment:**

**5. Section 4: I disagree with the framing of this section as a “validation” given the small extent of comparison and lack of detail or consideration of uncertainties. I suggest that it be renamed “Comparison with airborne observations” and the use of the word “validation” throughout be changed.**

**Response:** Done. Section 4 title has been rephrased as “Comparison with airborne observations.” All usage of ‘validation’ with aircraft data has been replaced with ‘comparison’ throughout the revised manuscript.

**Comment:**

**6. Section 5: comparing against OMI is one good choice; the authors might also mention the POLDER archive, which is similar or higher quality for SSA, but ended in 2013 before the time period the authors used here. The results could also be compared to global aerosol model simulations or reanalyses. And, although the authors briefly mention AERONET, it would be worthwhile to add a comparison with AERONET for regions where there is a persistent repeatable high aerosol loading. The authors could take AERONET climatologies themselves or go to other analyses, e.g. Giles et al (2012) <https://doi.org/10.1029/2012JD018127> for various types of aerosol or Sayer et al (2014) <https://acp.copernicus.org/articles/14/11493/2014/> for smoke in various regions. POLDER results could also be used in a climatological sense. Finally here the authors should be clearer that the reason for less OMI coverage over oceans is not so much cloudiness but in fact that over ocean the OMI retrieval is only done if the UVAI is high (I believe 0.7 or above). So this introduces a sampling bias towards high-AOD, high**

absorption cases (as we know baseline sea spray is not very absorbing) which is likely the main reason that OMI SSA is patchier and has lower values over ocean. This could be tested by also subsampling the MODIS-CERES data to only examine those times when OMI also has a retrieval. POLDER and reanalyses do not have this issue. Expanding the comparisons would provide further evidence for where the authors' technique may be valuable or where there are issues with one or other data set.

**Response:** Thank you for these suggestions to include comparisons with other satellite and ground-based SSA datasets. These suggestions have helped improve the revised manuscript. Summary of work done based on this comment:

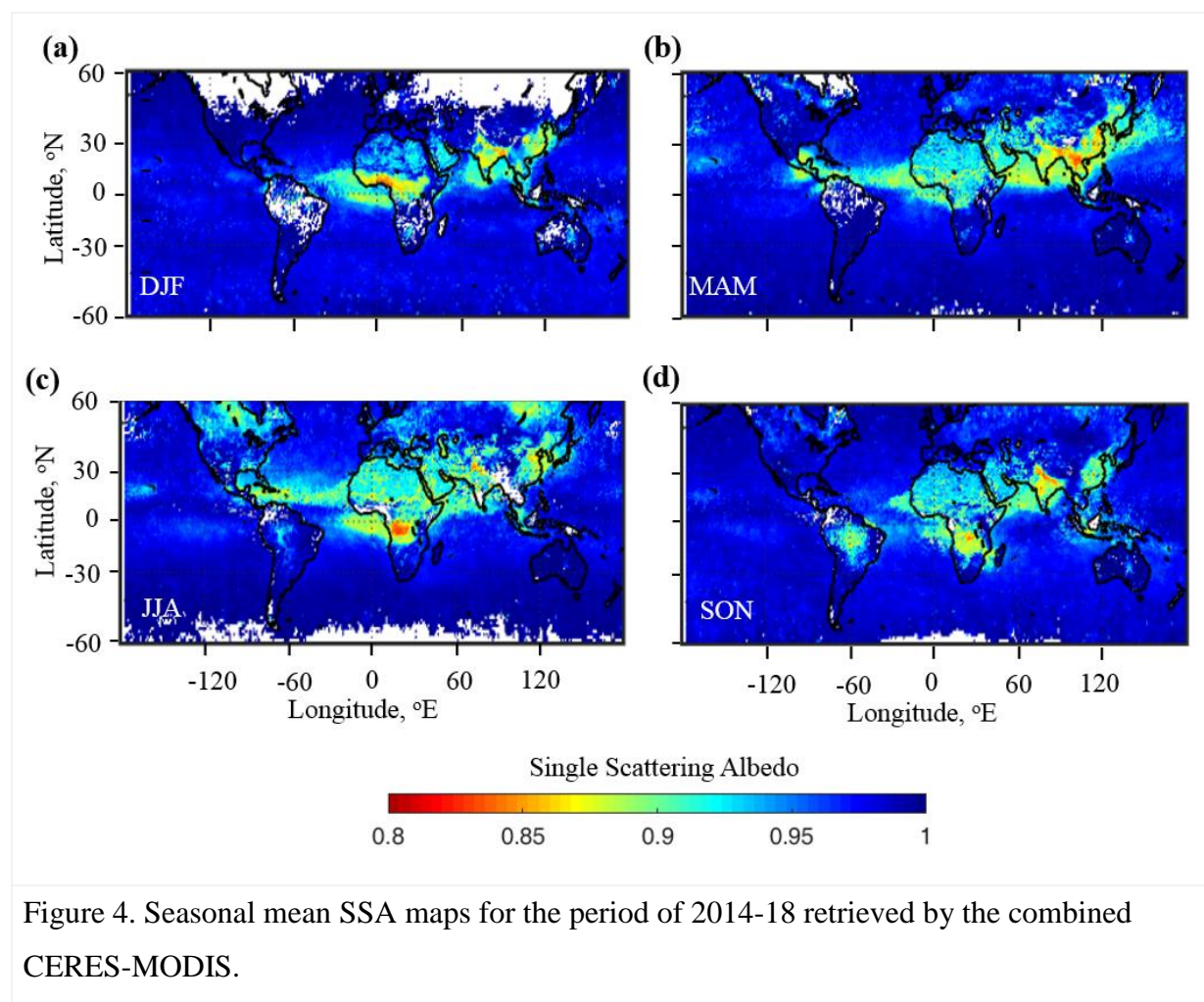
- Alongside OMI SSA (500 nm), we have compared the CERES-MODIS SSA dataset (550 nm) with POLDER climatological SSA (565 nm).
- The complete “results and discussion” section has been rewritten, emphasizing the advantages of each of the three datasets and the issues with one or the other data set. The table containing seasonal SSA values for different regions of interest has been shifted to the supplementary file.
- For the study period of 2014-18, the CERES-MODIS SSA has also been compared to monthly AERONET SSA data (440 nm) for the corresponding period for various AERONET sites. As suggested by the reviewer, we have chosen AERONET sites based on the type of aerosols as given by Giles et al., 2012. These results have been incorporated as the new Section 7 in the revised manuscript.
- Following the reviewer's comment, we intended to compare with the MERRAero reanalysis dataset. It would have been a valuable addition to the paper. Unfortunately, the OpenDap server was not accessible for downloading the data. Monthly SSA data files downloaded from the GEOS-5 data server were missing data values in it. We tried with the GrADS data server, but the download link was unavailable. We had also got our manuscript response deadline extended with ACP, hoping the data server would be up running by then. But we couldn't get to download the files.

#### **Additions to the revised manuscript:**

#### **4 Results and discussion**

Fig. 4 shows the seasonal-mean global maps of SSA (550 nm) retrieved by the combined CERES-MODIS algorithm for the five years of 2014-2018. Data are averaged for different

seasons: DJF (December-January-February), MAM (March-April-May), JJA (June-July-August), and SON (September-October-November).



The retrieved SSA dataset (500 nm) was compared with other widely used global SSA datasets – OMI SSA (500 nm) and climatological POLDER SSA (565 nm). OMAERUVd V3 (Torres et al., 2007; Torres et al., 2013; Ahn et al., 2014) for the corresponding period are shown in panels a, c, e, and g in Fig 5. And POLDER 1-2 Level 3 climatological seasonal mean SSA maps are shown in panels b, d, f, and h in Fig 5. For a generalized qualitative comparison, we can assume that SSA does not vary much for the small 50 nm spectral difference between CERES-MODIS and OMI SSA. (Zhu et al., 2011; Jethva et al., 2014).

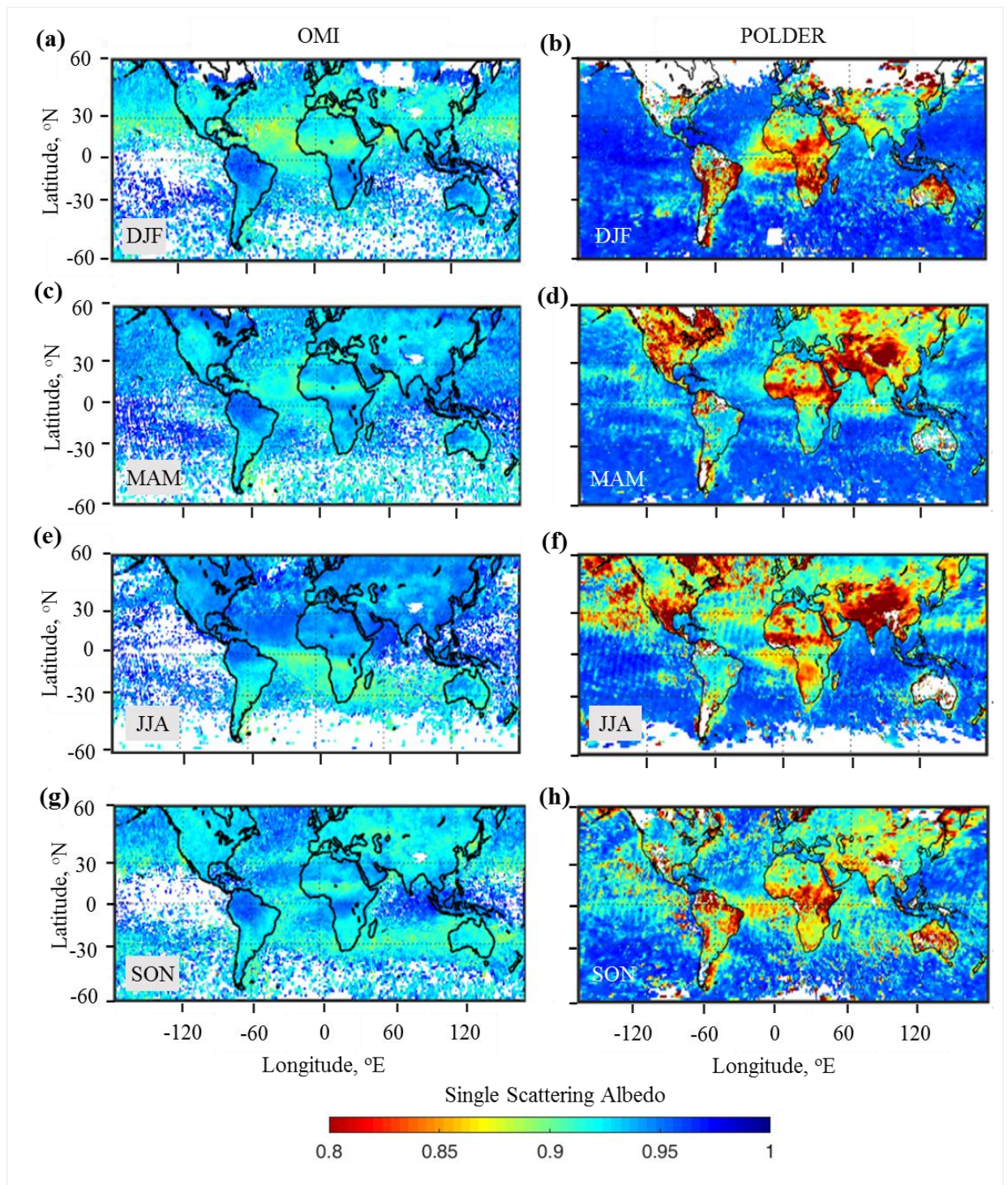


Figure 5. Seasonal mean SSA maps of OMI (500 nm) and POLDER (565 nm) in panels a,c,e,g and b,d,f,h respectively.

From a quick comparison between Fig 4 and Fig S2 SSA maps, the following points can be noted:

- Over the ocean, OMI retrieves SSA only for regions with high values of UVAI, leading to large data gaps. In comparison, we can notice that CERES-MODIS and



POLDER have better data coverage on a global scale. In the CERES-MODIS maps, the absence of data is mostly due to the unavailability of MODIS AOD.

- The Global Ocean, a relatively dark surface covering more than 70% of the Earth's surface, plays a significant role in determining global aerosol radiative forcing effects. Therefore, the better data coverage over oceans by the CERES-MODIS and POLDER provides better input for radiative forcing calculations.
- CERES-MODIS maps capture a wider range of SSA values. Regions with very low SSA can easily be identified as the sources of absorbing aerosols. OMI SSA values are mostly above 0.9 and do not clearly capture the sources and transport of absorbing aerosols.
- Both POLDER and OMI SSA values are more accurate in the UV wavelengths since SSA is primarily retrieved in the UV regions and extrapolated to visible wavelengths using aerosol models. Whereas CERES-MODIS retrieves SSA directly at 550 nm, hence is more accurate for SSA values in the visible wavelengths.
- Over the land, POLDER shows very low SSA values ( $< 0.85$ ), thus indicating the presence of highly absorbing aerosols even over less polluted regions. OMI values are around 0.9 over land and do not clearly identify the presence of absorbing aerosols. Whereas SSA values are within reasonable range over land as retrieved by the CERES-MODIS method – high SSA values over relatively pristine regions, lower SSA values over sources and transport of absorbing aerosols.
- Seasonal trends in forest fire can be noticed in POLDER maps and distinctly identifiable in CERES-MODIS SSA maps. Every year forest fires are common in specific seasons in Canadian and Russian Boreal forests (JJA), Amazon forest (SON) and South African forest (JJA and SON).
- The Indo-Gangetic plain (IGP) is a densely populated region spotted with several coal-based thermal power plants and seasonal stubble burning. Low SSA values are retrieved by both POLDER and CERES-MODIS over IGP. Whereas OMI shows values around 0.9 throughout the year. Similar pattern can be observed over Eastern China, one of the most highly polluted industrial region.

From the above points, we can draw conclusions about the advantages of each dataset. OMI, CERES, and MODIS instruments are still operational, whereas POLDER datasets are available only till 2013. OMI and POLDER SSA datasets are more suitable for UV wavelengths, whereas the CERES-MODIS SSA dataset provides more accurate SSA over visible wavelengths. OMI provides operational daily global SSA maps, whereas the CERES-MODIS algorithm is more suitable for obtaining monthly/seasonal global SSA maps. Over the ocean, the POLDER dataset has more coverage than OMI and identifies the transport of aerosols across the oceans. Hence, POLDER SSA and CERES-MODIS SSA can be used for studying SSA values over the ocean in the UV and visible wavelengths, respectively. Over the land, OMI retrieves high SSA values, whereas POLDER shows very low SSA values even over relatively pristine regions. Hence, the CERES-MODIS dataset retrieves reasonable SSA values over both polluted and less polluted regions for visible wavelengths.

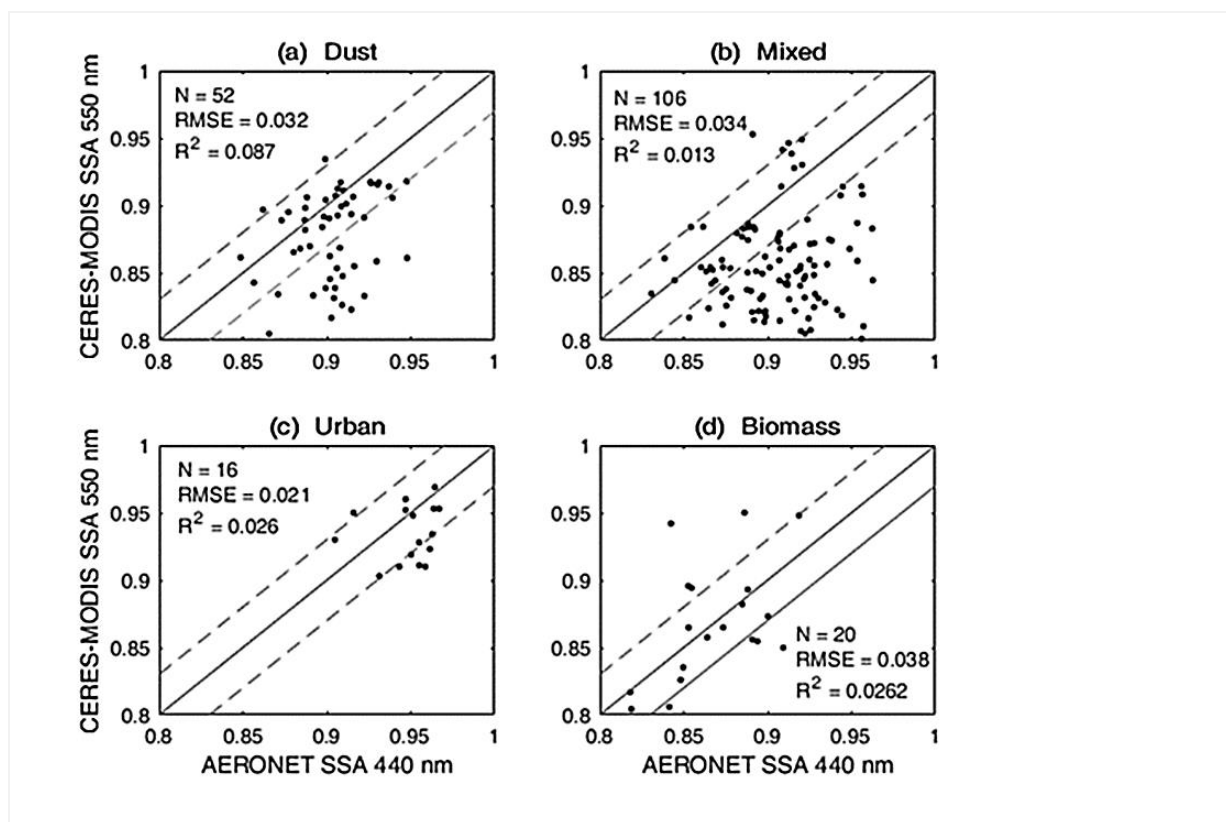
Global mean SSA retrieved by combined CERES-MODIS over land and ocean is 0.93 and 0.97, respectively (OMI: 0.94 and 0.94). Accurate SSA estimations are also required over regions of interest such as deserts, oceans, biomass-burning forests, and highly polluted industrial areas. Hence, seasonal mean SSA values retrieved by the combined CERES-MODIS algorithm, OMI, and POLDER are reported, in table S2, for major regions of interest as shown in Fig S1 and Table S1.

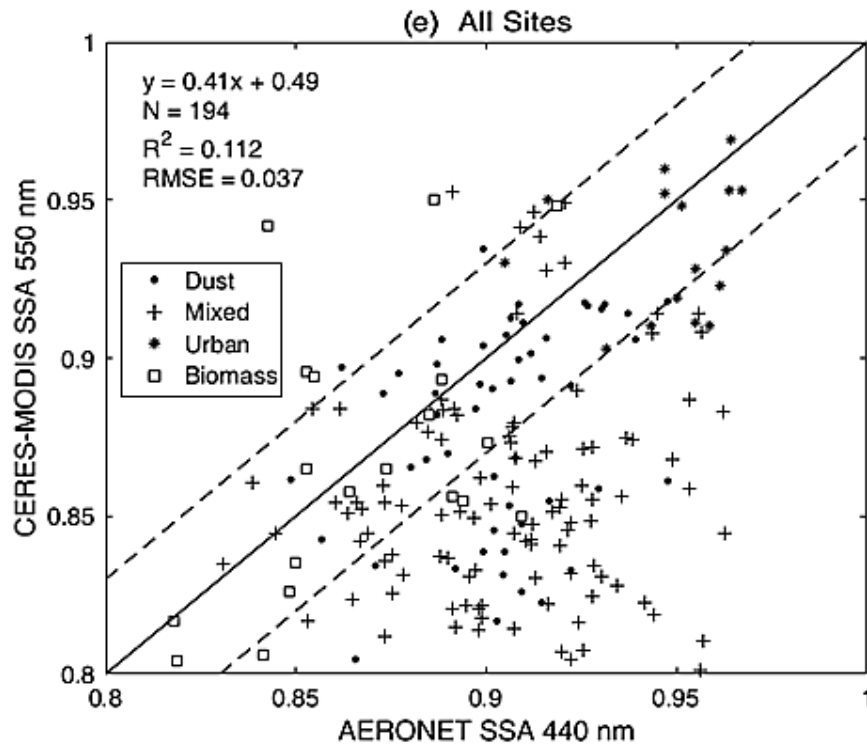
## **7 Comparison with AERONET data**

The Aerosol Robotic Network is a ground-based worldwide federated network of Cimel Sun photometers that measure extinction AOD from direct Sun measurements (Holben et al., 1998). The spectral diffuse sky radiations measured at different angles are inverted in conjunction with direct Sun measurements to derive the spectral SSA (440, 675, 870, and 1020 nm) and size distribution (Dubovik and King., 2000). The estimated uncertainty in retrieved SSA is largely attributed to the uncertainties in instrument calibration and is within 0.03 for AOD (440 nm) larger than 0.4. (Dubovik et al., 2000,2002).

AERONET version 3, level 2.0 monthly average values from selected sites were compared with corresponding CERES-MODIS SSA data. Sites were chosen to represent various types of aerosols following that of Giles et al., 2012. The location of the sites is shown in Fig S2 and Table S3. Scatter plots of comparison of AERONET SSA (440 nm) and CERES-MODIS SSA (550 nm) are shown in Fig 7.

Most AERONET SSA values are above 0.85, even in case of biomass burning aerosols. For dust type of aerosols (sites: Capo\_Verde, Dakar and Banizoumbaou), as shown in Fig 7a, AERONET and CERES-MODIS have better agreement. Whereas, in case of mixed type of aerosol (sites: SEDE\_BOKER, Kanpur, Xiang He and Illorin), most of the CERES-MODIS values are below 0.85, indicating highly absorbing type of aerosols (Fig 7b). For urban (sites: GSFC, Mexico\_city, Shirahama, Ispra and Moldova) and biomass (sites: Alta\_Floresta, Lake\_Argyle and Mongu) only very few data were available during the study period of 2014-18 as shown in Fig 7 panels c and d. Data points combined from all the sites are plotted together in Fig 7e showing a RMSE of 0.037. The large difference between SSA wavelengths of AERONET (440 nm) and CERES-MODIS (550 nm) could contribute to the variations observed between these two datasets. Overall, the resulting comparisons are agreeable within the uncertainties of both AERONET and CERES-MODIS datasets.





**Figure 7.** CERES-MODIS SSA (550 nm) vs AERONET SSA (440 nm) for various AERONET sites classified based on type of aerosols (Giles et al., 2012)

**Additions to the revised manuscript:** (underlined)

## 8. Summary and Conclusions

- Global maps of aerosol absorption have been generated following the concept of “critical optical depth”.
- The retrieved SSA values have been compared with available aircraft measurements. The limited comparison exercise shows that most of the retrieved SSA values are within  $\pm 0.03$ .
- We show that the combined CERES-MODIS algorithm better captures the spatial and seasonal variations in aerosol absorption and the resultant maps provide an improved global SSA database with fewer data gaps. Global mean SSA was estimated to be 0.93 and 0.97 over land and ocean, respectively
- The algorithm’s sensitivity to various parameters have been studied, which shows maximum sensitivity to changes in surface albedo. The algorithm is shown to be the most effective over regions with large AOD variations and less surface albedo variations.

- Comparison with SSA from 15 AERONET sites showed an acceptable agreement between AERONET and CERES-MODIS SSA, within their uncertainties.
- Overall, the combined CERES-MODIS algorithm provides global SSA maps with improved accuracy and better spatial coverage. These global maps provide valuable input for models to make assessment of aerosol-climate impacts on both regional and global scales.

**Comment:**

**In summary, this study has value but I think a much deeper treatment of uncertainty is needed. Otherwise it is not clear to what extent this technique improves our understanding of aerosol SSA, or where the largest challenges remain. We can't quantitatively move forward if we don't understand where we stand now.**

**Response:** Uncertainty analysis has been studied for various parameters and is now presented in a separate section 5. CERES-MODIS dataset has been compared both with OMI and POLDER. Their respective advantages and suitable area of usage are also discussed in the section 4, 'results and discussion'. Comparison were done with monthly AERONET data from 15 sites, chosen based on Giles et al., 2012 (section 7). We thank the reviewer for these detailed suggestions and comments, which greatly improved the manuscript.

# Global maps of aerosol single scattering albedo using combined CERES-MODIS retrieval

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**Abstract.** Single Scattering Albedo (SSA) is a leading contributor to the uncertainty in aerosol radiative impact assessments. Therefore accurate information on aerosol absorption is required on a global scale. In this study, we have applied a multi-satellite algorithm to retrieve SSA ([550 nm](#)) using the concept of ‘critical optical depth.’ Global maps of SSA were generated following this approach using spatially and temporally collocated data from Clouds and the Earth’s Radiant Energy System (CERES) and Moderate Resolution Imaging Spectroradiometer (MODIS) sensors on board Terra and Aqua satellites. [Limited comparisons against airborne observations over India and surrounding oceans were generally in agreement within  \$\pm 0.03\$ .](#) Global mean SSA estimated over land and ocean is 0.93 and 0.97, respectively. Seasonal and spatial distribution of SSA over various regions are also presented. The global maps of SSA, thus derived with improved accuracy, provide important input to climate models for assessing the climatic impact of aerosols on regional and global scales.

**Deleted:** The method has been validated using the data from aircraft-based measurements of various field campaigns. The retrieval uncertainty is  $\pm 0.03$  and depends on both the surface albedo and aerosol absorption.

## 1 Introduction

Atmospheric aerosols play a significant role in the Earth’s radiation budget (IPCC, 2013). The climatic impact of aerosols depends on their absorption and scattering properties, quantified by Single Scattering Albedo (SSA). Even a slight reduction in SSA can change the aerosol radiative forcing from cooling to warming, depending on the underlying surface albedo (Kaufman et al., 2001; Chand et al., 2009). However, the lack of an accurate global aerosol absorption database has led to SSA being the largest contributor to the total uncertainty in aerosol radiative impact assessment (IPCC, 2013).

The high spatio-temporal variability in aerosol properties entails the need for observations on a global scale (Dubovik et al., 2002; Levy et al., 2007; Remer et al., 2008; Hammer et al., 2018). Satellite data, despite its inherent limitation associated with an inverse problem, can provide the global perspective required in analysing

spatio-temporal aerosol characteristics (Torres et al., 2002; Lenoble et al., 2013). However, it is difficult to quantify the absorption over bright surfaces (Kaufman and Joseph, 1982; Ahn et al., 2014; Jethva et al., 2018). Hence, quantifying the aerosol absorption over land regions using satellite-based remote sensing remains a challenge even now (Torres et al., 2013; Jethva and Torres, 2019).

5 Fraser and Kaufman., 1985 developed a critical surface reflectance method to retrieve SSA using satellite data. Their method is based on radiative transfer simulations, which showed a particular surface reflectance for which the top of atmosphere albedo is independent of AOD. Upward radiances between a clear and a hazy day over a varying surface reflectance region are used, along with radiative transfer simulations, to derive SSA. This method has been widely applied to data from various satellites to derive SSA over particular regions (Kaufman, 1987; Kaufman et al., 1990, 2001; Zhu et al., 2011; Wells et al., 2012). Seidel and Popp., 2012 have done extensive studies on the method's sensitivity to various parameters.

10 Various studies have ascertained the inadequacy of single-sensor data in the accurate retrieval of aerosol absorption (Kaufman et al., 2001; Zhu et al., 2011). Dawn of the A-Train satellite constellation (Anderson et al., 2005) with spatially and temporally near-collocated observations facilitates multi-satellite retrieval of aerosol absorption (Eswaran et al., 2019; Hsu et al., 2000; Hu et al., 2007, 2009; Jeong and Hsu, 2008; Narasimhan and Sathesh, 2013; Sathesh et al., 2009) However, all these multi-sensor retrievals are in the Ultra Violet (UV) wavelengths, and SSA is extrapolated to visible wavelengths using spectral dependence of assumed particle size distribution. Sathesh and Srinivasan (2005) defined the concept of "critical optical depth" ( $\tau_c$ ) and introduced a method to retrieve SSA in the visible region by combining ground-based and satellite measurements. The method was validated/demonstrated over many locations, including the desert location of Solar Village in Saudi Arabia, using Aerosol Robotic Network (AERONET) data.

15 In this paper, we have utilized the concept of  $\tau_c$  and further extended the methodology to develop the combined CERES-MODIS retrieval algorithm to derive regional and global maps of aerosol absorption (550 nm) using multi-satellite data. The "critical optical depth" method developed in this research paper shares a similar concept to the critical surface reflectance method (Fraser and Kaufman., 1985). For a particular parameter (such as surface reflectance or optical depth), there exists a critical value at which the top of atmosphere albedo can be considered independent of variations in that parameter. Both the methods retrieve SSA by parameterizing the critical value as a function of SSA using radiative transfer simulations. The critical reflectance method requires two-days data and large variations in surface reflectance over the region. It's suitable for retrieving daily SSA for a particular

region. Whereas the critical optical method developed in this paper is suitable for retrieving monthly or seasonal global maps of SSA.

The concept of  $\tau_c$ , which forms the scientific basis for the development of this retrieval algorithm is illustrated in Section 2. The various steps involved in the retrieval algorithm are detailed in the Section 3, data and methodology.

5 Section 4 presents the results and comparison with other satellite datasets. Uncertainty analysis is studied in Section 5. Comparison with aircraft measurements from various field campaigns are shown in Section 6. Comparison with AERONET data from 15 sites are shown in section 7. Summary and conclusions are provided in Section 8.

## 2 Critical optical depth

10 Let  $\Delta\alpha$  be the difference between the top of the atmosphere (TOA) albedo and surface albedo. Then, for a particular location, with a given surface albedo,  $\Delta\alpha$  variations are only due to changes in TOA albedo. The presence of absorbing aerosols over a bright surface decreases the TOA albedo. In contrast, scattering aerosols over a dark surface increase the TOA albedo. Thus, the increase (decrease) in aerosol loading due to scattering (absorbing) type of aerosols leads to an increase (decrease) in  $\Delta\alpha$ . The rate of change in  $\Delta\alpha$  with aerosol loading is dependent on SSA.

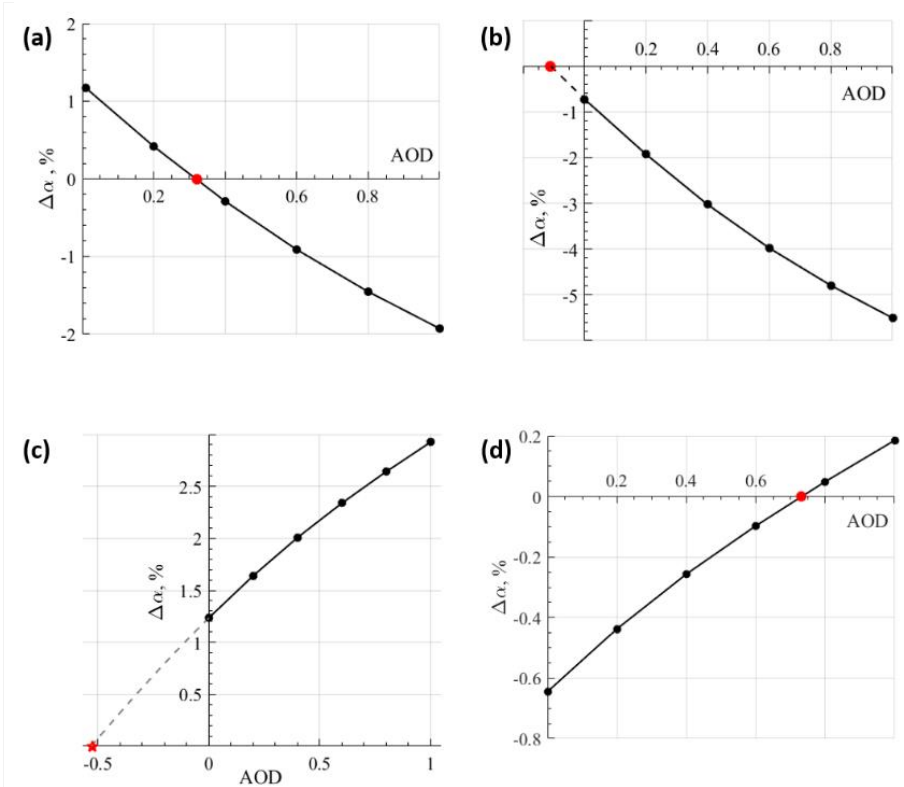
Satheesh and Srinivasan (2005) utilized this concept to retrieve SSA in the case of absorbing aerosols over a bright surface. In a pristine atmosphere (Aerosol Optical Depth = 0) over a bright surface, the  $\Delta\alpha$  is positive for solar zenith angle (SZA) = 0. Here, when absorbing aerosols become dominant,  $\Delta\alpha$  decreases with an increase in aerosol optical depth (AOD) and eventually turns negative. The AOD at which  $\Delta\alpha$  equals zero is defined as  $\tau_c$ . For a given surface albedo,  $\tau_c$  is the AOD at which the scattering and absorbing effects of the aerosol cancel each other. The rate of decrease in  $\Delta\alpha$  with the increase in AOD is higher when SSA is high and consequently lowers the resulting values of  $\tau_c$ . A radiative transfer (RT) model was then used to calculate the SSA that reproduces the same  $\tau_c$ , given atmospheric conditions.

**Deleted:** Section 4 presents the validation of SSA derived using this approach using aircraft measurements from various field campaigns.

**Deleted:** The global maps of SSA thus retrieved, its comparison with SSA from Ozone Monitoring Instrument (OMI), and the seasonal distribution of SSA over many regions are presented in Section 5.

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**Figure 1.** RT simulations (black dots) shows deriving  $\tau_c$  (red dot) for different cases of aerosols and surfaces. For pristine conditions (AOD = 0), diurnally-averaged  $\Delta\alpha$  is negative for bright surfaces and positive for dark surfaces. An increase in aerosol loading by absorbing (scattering) type of aerosol leads to decrease (increase) in TOA albedo. **(a)** Absorbing aerosols above a dark surface; **(b)** Absorbing aerosols above a bright surface; **(c)** Scattering aerosols above a dark surface; **(d)** Scattering aerosols above a bright surface.

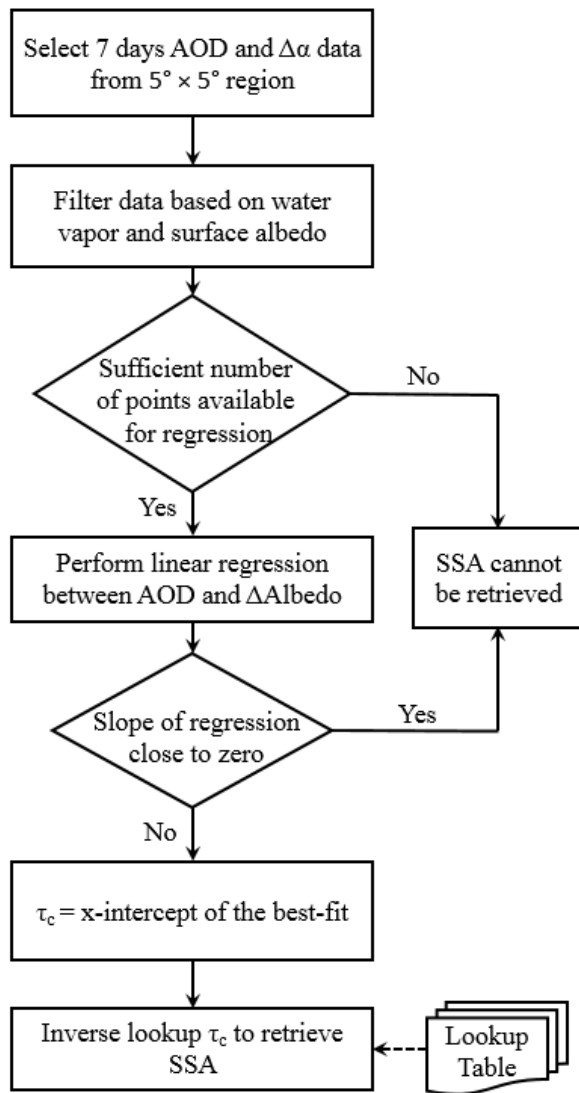
In this paper, the concept of  $\tau_c$  is extended to retrieve SSA for all scenarios of surfaces (dark and bright) and aerosols (absorbing and scattering). For AOD less than 1,  $\Delta\alpha$  is almost linearly dependent on AOD. Then  $\tau_c$  is mathematically the x-intercept when parameterizing the linear relationship.

- 5 Figure 1 shows the estimation of  $\tau_c$  for four different scenarios. Details of these RT simulations are given in Section 3.2. Unlike Satheesh and Srinivasan (2005), where simulations were carried out for  $SZA = 0$ , here the  $\Delta\alpha$  is diurnally averaged. Therefore, it is possible to have negative  $\Delta\alpha$  for AOD = 0 over relatively bright surfaces. It is difficult to retrieve SSA where the slope of regression line is close to zero.

### 3 Data and methodology

The Combined CERES-MODIS retrieval algorithm consists mainly of two steps: (1) determining  $\tau_c$  using MODIS and CERES data for a location, and (2) estimation of SSA that reproduces the same  $\tau_c$  for the associated atmospheric conditions and surface albedo of that particular location. Figure 2 shows the flowchart illustrating the combined CERES-MODIS retrieval algorithm.

TOA and surface fluxes, used to determine  $\Delta\alpha$ , are obtained from CERES SYN1deg-day (Edition 4.1) (Wielicki et al., 1996; Rutan et al., 2015). To avoid angular dependence of fluxes, the diurnally averaged flux data product from CERES is used, which is available only at 1° resolution. Hence, other satellite data sets in this study are also used at the same spatial resolution. AOD and total columnar water vapor are obtained from the MODIS Daily Global Product (MxD08\_D3 version 6.1). MODIS retrieves columnar AOD at 550 nm using two different types of algorithms – “Dark Target” (Levy et al., 2007, 2013) and “Deep Blue” (Hsu et al., 2004, 2006; Sayer et al., 2013). Dark target retrieves AOD over both land and ocean, whereas deep blue retrieves only over land. In this study, we have used a combined dark target and deep blue product.



**Figure 2.** Flowchart depicting the steps involved in combined CERES-MODIS retrieval of SSA for a particular location.

### 3.1 Determining the critical optical depth

The first step for retrieval is to determine  $\tau_c$  by linear regression analysis between  $\Delta\alpha$  vs. AOD as shown in Fig.

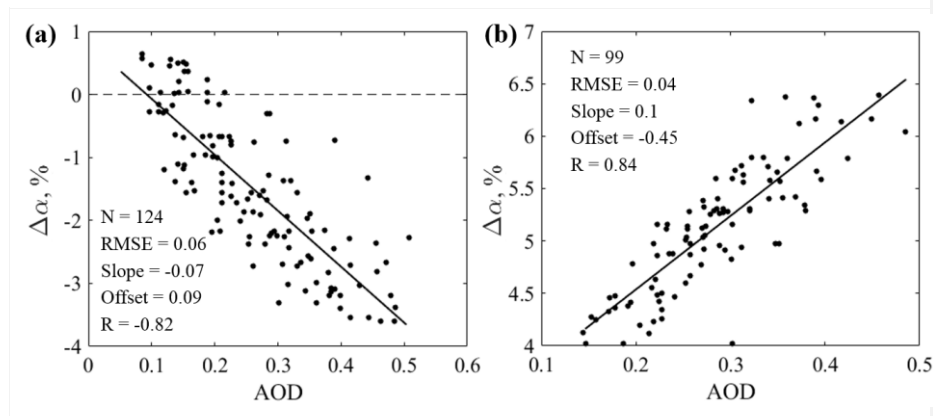
3. The x-intercept of the resultant line of best fit (i.e., the AOD at which  $\Delta\alpha = 0$ ) provides the value of  $\tau_c$ . CERES

and MODIS daily data are at  $1^\circ$  resolution, and SSA is retrieved for each  $1^\circ \times 1^\circ$  grid. In order to have adequate number of points for a meaningful regression analysis, it was required to use data over a larger interval (temporal and spatial) - whose extent is large enough to get a statistically significant fit but small enough to ensure insignificant variations in SSA. Thus, to determine  $\tau_c$  for a given pixel, seven days of data from its surrounding  $5^\circ \times 5^\circ$  region has been considered. This data is further constrained based on surface albedo and water vapor. Only those pixels in this region having surface albedo within  $\pm 0.025$  and water vapor within  $\pm 0.25$  cm of the given pixel are considered for regression analysis. These constraints ensure that the  $\tau_c$  determined from the best fit is dependent only on SSA and not affected by changes in surface albedo and water vapour. Figure 3a shows an example of regression with a positive correlation coefficient over the Arabian Sea. This can happen over regions of low surface albedo and the dominance of scattering aerosols. Figure 3b is an example of regression analysis with a negative correlation coefficient obtained over Sahara in the presence of dust aerosols.

The above procedure is repeated for all pixels, where data from the surrounding  $5^\circ \times 5^\circ$  region is used to determine  $\tau_c$  for each pixel. For the regression analysis, points which are outside one standard deviation are considered as outliers. Line of best fits with a slope close to zero yields extreme  $\tau_c$  values (very high positive/very low negative).

In such cases, we did not attempt a retrieval. A significance test on the correlation coefficient between AOD and  $\Delta\alpha$  is performed with a 0.05 significance level. Only those  $\tau_c$  values obtained through regressions that are statistically significant at 95% confidence level are utilized further to retrieve SSA.

**Deleted:** Only those  $\tau_c$  values that are statistically significant at 95% confidence level are utilised further for the retrieval of SSA.



**Figure 3.** Sample scatterplots between MODIS AOD and CERES  $\Delta\alpha$ . The solid lines represent the best-fits for (a) absorbing aerosols above the Sahara and (b) scattering aerosols above the Arabian Sea.  $\tau_c$  (AOD at which  $\Delta\alpha$  is zero) is the x-intercept of the best-fit line.

The final product of this step is a  $360 \times 180$  matrix that stores  $\tau_c$  value corresponding to each  $1^\circ$  pixel. In these matrices, not all points would have a  $\tau_c$  value owing to the insufficient number of points available for regression, either due to cloud-masking or large variations in surface albedo over the land. At least seven days of data is required to perform a statistically significant fit to compute  $\tau_c$  and retrieve SSA. The next step in the procedure is to estimate SSA from these  $\tau_c$  values using an inverse lookup table (LUT) approach.

### 3.2 Retrieval of SSA

Since the objective of this study is to retrieve SSA globally, look-up-tables (LUTs) were developed to reduce the computation time and avoid repeated RT simulations. The aerosol models from OPAC (Optical Properties of Aerosols and Clouds), developed by Hess et al., (1998), are given as input to SBDART (Santa Barbara DISORT Atmospheric Radiative Transfer) model (Ricchiuzzi et al., 1998) to simulate TOA fluxes. [Specifications of the model used are shown in Table S5, S6 and S7.](#)

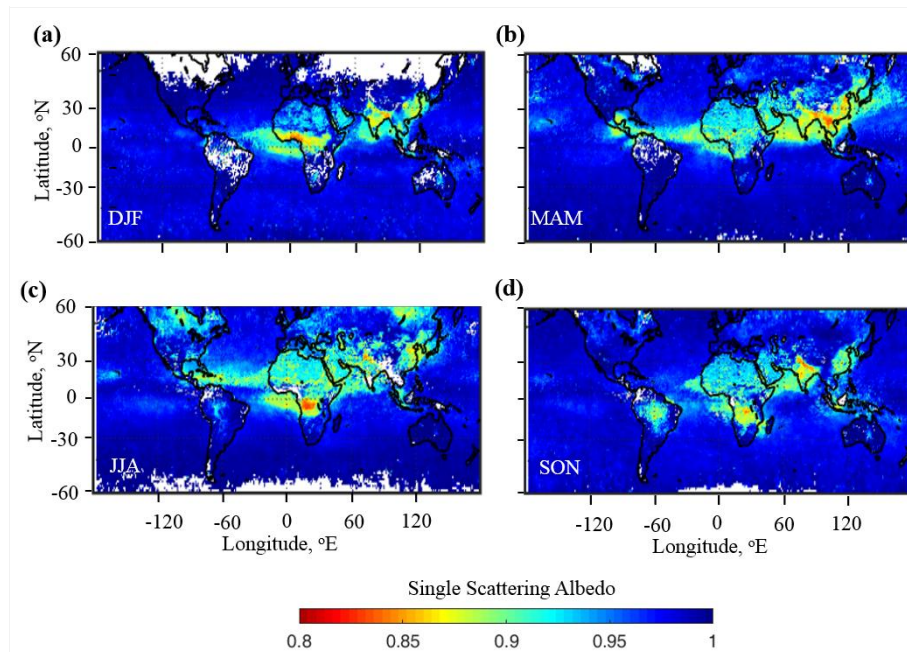
The RT computations were carried out to obtain the diurnally averaged (SZA:  $0^\circ$  to  $84^\circ$ ) TOA and surface fluxes using 16 radiation streams and spectrally integrated over the shortwave region (0.3 to  $5 \mu\text{m}$ ). For a particular case of surface albedo, water vapor, and SSA, AOD is varied from 0 to 1 in steps of 0.2 to generate its corresponding diurnally averaged  $\Delta\alpha$ . Then a linear fit is performed between AOD and simulated  $\Delta\alpha$  to determine  $\tau_c$ . A three-dimensional LUT that stores  $\tau_c$  for different combinations of surface albedo, water vapor, and SSA have been developed. The LUT is indexed by 11 values of surface albedo (0 to 0.5, increments of 0.05), 17 values of water vapour (0 to 8 cm, increments of 0.5 cm) and 10 values of SSA (0.8, 0.83, 0.85, 0.87, 0.9, 0.92, 0.95, 0.97, 0.99, and 1). A total of 89760 RT simulations were performed in the present study.

The next step is to estimate SSA from  $\tau_c$  using the LUT. For a given surface albedo and water vapor of that pixel, we find the SSA associated with its determined  $\tau_c$ . An inverse lookup operation is performed on LUT by linear interpolation between the nearest two indices. SSA is estimated for each available  $\tau_c$  values of a pixel and then averaged to compute the seasonal mean SSA.

#### 4 Results and discussion

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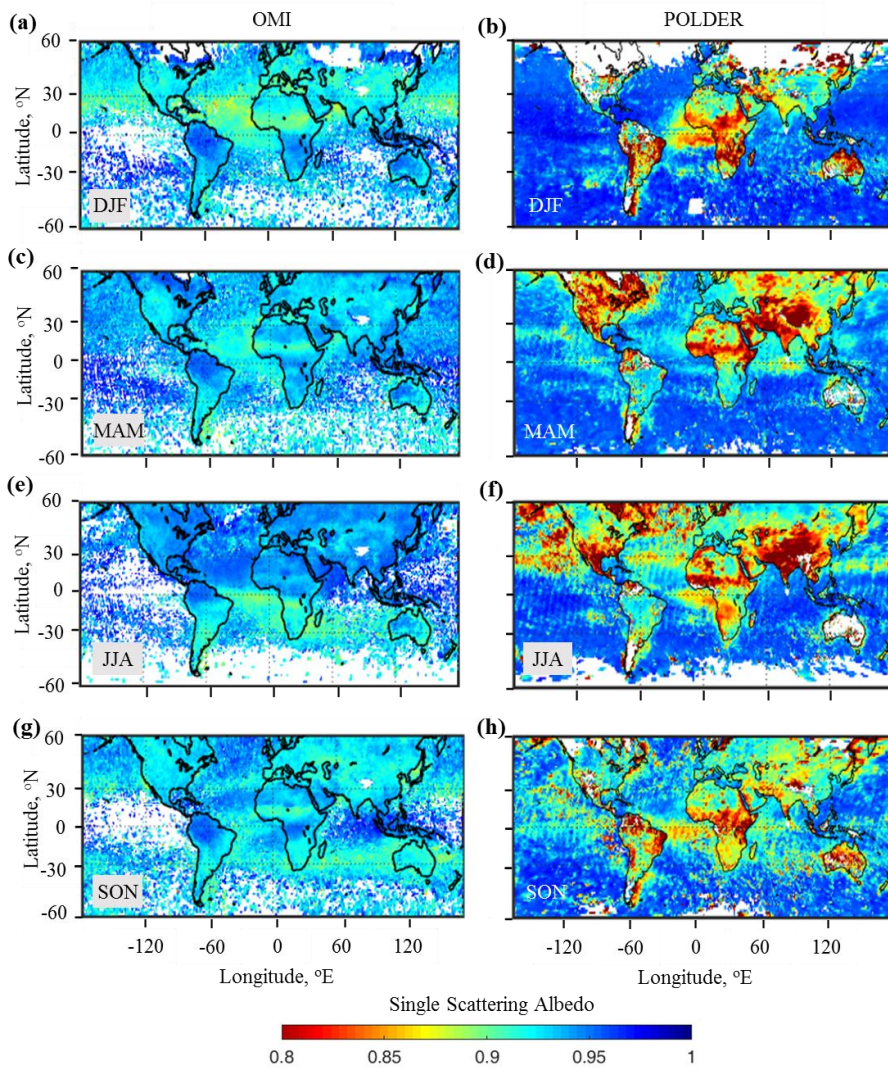
Fig. 4 shows the seasonal-mean global maps of SSA (550 nm) retrieved by the combined CERES-MODIS algorithm for the five years of 2014-2018. Data are averaged for different seasons: DJF (December-January-February), MAM (March-April-May), JJA (June-July-August), and SON (September-October-November).



**Figure 4.** Seasonal mean SSA maps for the period of 2014-18 retrieved by the combined CERES-MODIS.

5 The retrieved SSA dataset (500 nm) was compared with other widely used global SSA datasets – OMI SSA (500 nm) and climatological POLDER SSA (565 nm). OMAERUVd V3 (Torres et al., 2007; Torres et al., 2013; Ahn et al., 2014) for the corresponding period are shown in panels a, c, e, and g in Fig 5. And POLDER 1-2 Level 3 climatological seasonal mean SSA maps are shown in panels b, d, f, and h in Fig 5. For a generalized qualitative comparison, we can assume that SSA does not vary much for the small 50 nm spectral difference between CERES-

10 MODIS and OMI SSA. (Zhu et al., 2011; Jethva et al., 2014).



**Figure 5.** Seasonal mean SSA maps of OMI (500 nm) and POLDER (565 nm) in panels a,c,e,g and b,d,f,h respectively.

From a quick comparison between Fig 4 and Fig S2 SSA maps, the following points can be noted:

- Over the ocean, OMI retrieves SSA only for regions with high values of UVAI, leading to large data gaps. In comparison, we can notice that CERES-MODIS and POLDER have better data coverage on a global scale. In the CERES-MODIS maps, the absence of data is mostly due to the unavailability of MODIS AOD.

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- The Global Ocean, a relatively dark surface covering more than 70% of the Earth's surface, plays a significant role in determining global aerosol radiative forcing effects. Therefore, the better data coverage over oceans by the CERES-MODIS and POLDER provides better input for radiative forcing calculations.

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- CERES-MODIS maps capture a wider range of SSA values. Regions with very low SSA can easily be identified as the sources of absorbing aerosols. OMI SSA values are mostly above 0.9 and do not clearly capture the sources and transport of absorbing aerosols.

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- Both POLDER and OMI SSA values are more accurate in the UV wavelengths since SSA is primarily retrieved in the UV regions and extrapolated to visible wavelengths using aerosol models. Whereas CERES-MODIS retrieves SSA directly at 550 nm, hence is more accurate for SSA values in the visible wavelengths.

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- Over the land, POLDER shows very low SSA values (< 0.85), thus indicating the presence of highly absorbing aerosols even over less polluted regions. OMI values are around 0.9 over land and do not clearly identify the presence of absorbing aerosols. Whereas SSA values are within reasonable range over land as retrieved by the CERES-MODIS method – high SSA values over relatively pristine regions, lower SSA values over sources and transport of absorbing aerosols.

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- Seasonal trends in forest fire can be noticed in POLDER maps and distinctly identifiable in CERES-MODIS SSA maps. Every year forest fires are common in specific seasons in Canadian and Russian Boreal forests (JJA), Amazon forest (SON) and South African forest (JJA and SON).

- The Indo-Gangetic plain (IGP) is a densely populated region spotted with several coal-based thermal power plants and seasonal stubble burning. Low SSA values are retrieved by both POLDER and CERES-MODIS over IGP. Whereas OMI shows values around 0.9 throughout the year. Similar pattern can be observed over Eastern China, one of the most highly polluted industrial region.



From the above points, we can draw conclusions about the advantages of each dataset. OMI, CERES, and MODIS instruments are still operational, whereas POLDER datasets are available only till 2013. OMI and POLDER SSA datasets are more suitable for UV wavelengths, whereas the CERES-MODIS SSA dataset provides more accurate SSA over visible wavelengths. OMI provides operational daily global SSA maps, whereas the CERES-MODIS algorithm is more suitable for obtaining monthly/seasonal global SSA maps. Over the ocean, the POLDER dataset has more coverage than OMI and identifies the transport of aerosols across the oceans. Hence, POLDER SSA and CERES-MODIS SSA can be used for studying SSA values over the ocean in the UV and visible wavelengths, respectively. Over the land, OMI retrieves high SSA values, whereas POLDER shows very low SSA values even over relatively pristine regions. Hence, the CERES-MODIS dataset retrieves reasonable SSA values over both polluted and less polluted regions for visible wavelengths.

Global mean SSA retrieved by combined CERES-MODIS over land and ocean is 0.93 and 0.97, respectively (OMI: 0.94 and 0.94). Accurate SSA estimations are also required over regions of interest such as deserts, oceans, biomass-burning forests, and highly polluted industrial areas. Hence, seasonal mean SSA values retrieved by the combined CERES-MODIS algorithm, OMI, and POLDER are reported, in table S2, for major regions of interest as shown in Fig S1 and Table S1.

### 5 Uncertainty Analysis

Table 1 identifies the major sources of error in the retrieval and summarizes their individual contribution. Uncertainty in the retrieved SSA was estimated by calculating retrieval sensitivities to perturbations in the possible error sources. The range of perturbation was based on published literature or reasonable assumptions for possible variations.

**Table 1.** Estimates of the uncertainty in retrieved SSA

Parameter	Input Uncertainty	Retrieval Uncertainty
Surface albedo	$\pm 0.01$	$\pm 0.03$
AOD	20% $\pm 0.05$ (land) 5% $\pm 0.03$ (ocean)	$\pm 0.02$
Angstrom exponent	$\pm 0.4$	$\pm 0.01$
Refractive index	$\pm 0.01$	$\pm 0.01$
Aerosol height	$\pm 1$ km	$\pm 0.01$
Aerosol type	Smoke vs dust	$\pm 0.01$
Residual of fit	$\pm 0.05$	$\pm 0.02$

Uncertainty in shortwave integrated surface albedo from CERES results in the maximum uncertainty in SSA of  $\pm 0.03$ . MODIS retrieved aerosol optical depth contains considerable uncertainties due to assumed aerosol models (Jeong et al., 2005). The MODIS aerosol optical depth uncertainty is 20%  $\pm 0.05$  over land (Chu et al., 2002) and 5%  $\pm 0.03$  over the ocean (Remer et al., 2002). The corresponding error in our retrieval is  $\pm 0.02$ . For a typical variation of angstrom exponent ( $\pm 0.4$ ) and refractive index ( $\pm 0.01$ ), the uncertainties vary depending on the surface albedo and are mostly around  $\pm 0.01$ .

Changes in aerosol height can vary the TOA radiances due to Rayleigh scattering interactions, which depend on pressure. Sensitivity to aerosol height was estimated by conducting a synthetic retrieval of SSA over a range of aerosol height values and perturbations from those heights. The average uncertainty observed for an aerosol height variation of  $\pm 1$  km was  $\pm 0.01$ . Many methods have been developed for detecting aerosol type, especially smoke vs. dust, to improve the uncertainties of various AOD and SSA retrievals.

Uncertainties due to possible variations on scales of the regions used for linear fitting were estimated as residuals of the fit. The uncertainty on the linear intercept is spatially dependent and is mostly around  $\pm 0.02$ , with higher values for those combinations having a slope close to zero during the regression. For highly correlated cases (i.e., correlation coefficient  $|r| > 0.5$ ), the probability of obtaining a slope close to zero is  $\sim 20\%$  over the ocean and  $< 5\%$  over land. These cases are mostly formed over regions where AOD variations are less. Regions having large variations in AOD values have lower uncertainty due to residual fit.

Overall, the algorithm is most sensitive to variations in surface albedo, followed by higher sensitivity towards AOD values used in the linear fit. Seasonal mean maps of surface albedo are shown in Fig S3. The uncertainties are higher for scattering aerosols over bright surfaces and absorbing aerosols above dark surfaces. Sensitivity to water vapor is almost negligible, except in very few cases where the uncertainty is  $+ 0.008$ . The CERES-MODIS algorithm is most effective over regions with large AOD variations and less surface albedo variations.

#### 6 Comparison with airborne observations

For the comparison of columnar SSA values thus retrieved, we have used aircraft-based measurements of SSA from three campaigns: South West Asian Aerosol Monsoon Interactions (SWAAMI), Regional Aerosol Warming Experiment (RAWEX), and SWAAMI-RAWEX, to obtain column-integrated SSA. Available data points over India and adjoining oceanic regions (Arabian Sea and Bay of Bengal) from these field campaigns were compared with the retrieved SSA.

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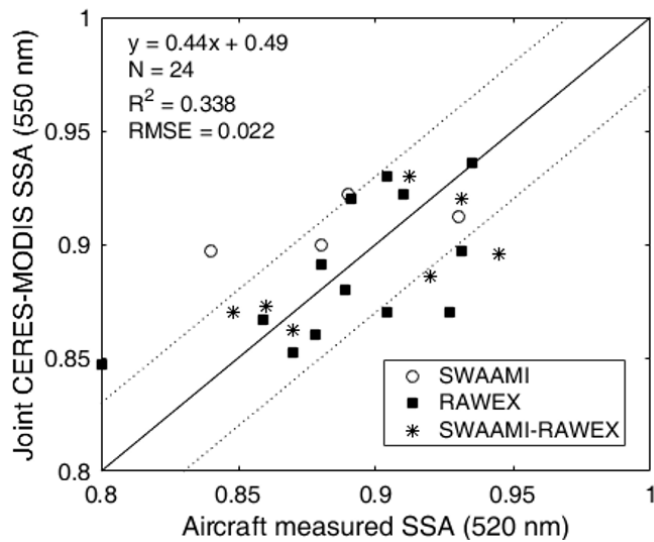
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Babu et al. (2016), as part of RAWEX (Moorthy et al., 2016), derived SSA at 520 nm from aircraft measurements of scattering and absorption coefficients over the Indo-Gangetic Plain (IGP) and Central India during winter 2012 and spring/pre-monsoon 2013. Various measurements of aerosol properties were carried out in an instrumented Beechcraft B200 aircraft of the National Remote Sensing Centre, India. Manoj et al. (2019) estimated vertical profiles of SSA during the SWAAMI campaign conducted during monsoon (June - July) 2016 over IGP, Arabian Sea, and Bay of Bengal. Aerosol scattering coefficients were measured aboard the Facility for Airborne Atmospheric Measurements (FAAM) BAe-146 aircraft. Vaishya et al. (2018) estimated vertical profiles of SSA (520 nm) using an instrumented aircraft, Beechcraft B200, during SWAAMI-RAWEX campaign (June 2016). Instrument design and calibration were based on Anderson et al., 1996 and its application for Indian field experiments was as described by Nair et al., 2009. Uncertainties in the scattering coefficient measurement by nephelometer are  $\sim\pm 10\%$ , as reported by Anderson et al., 1996. As stated by Babu et al., 2016 uncertainties in the columnar SSA values estimated from RAWEX aircraft measurements depend mainly on instrumental uncertainties, sampling errors, and large spatial averaging.

Retrieved SSA, for the same period as the campaign, over a  $2^\circ \times 2^\circ$  region around the campaign location was utilized for comparison. Figure 4 shows the comparison of collocated aircraft measurements and CERES-MODIS retrieved SSA. The ideal 1:1 case (solid line), the absolute difference of 0.03 (dotted lines), and regression coefficients are also provided.

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**Figure 6.** Comparison of combined CERES-MODIS SSA with aircraft measurements during SWAAMI, RAWEX, and SWAAMI-RAWEX campaigns. The solid line shows the ideal 1:1 case and dotted lines represent the absolute difference of 0.03.

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Most of the points were within the absolute difference of 0.03. However, there are few exceptions. SSA values over the Bay of Bengal during SWAAMI campaign were reported as  $0.84 \pm 0.07$  during June-July by Manoj et al. (2019), whereas CERES-MODIS retrieves a higher SSA of  $\sim 0.89$  for the same time period. This large variation could be due to frequent cloud cover during the monsoon season, leading to fewer SSA points retrieved over the ocean and land. SSA estimated over Nagpur in Central India during RAWEX is  $\sim 0.8$ , while CERES-MODIS retrieves  $\sim 0.85$ . This inconsistency is due to the large surface albedo variations (standard deviation  $> 0.05$ ) over Central India, which leads to fewer points available for retrieval. Except for few such cases, most of the other points lie within an absolute difference of 0.03.

10 For comparison purposes, many previous studies have used ground-level SSA data from AERONET obtained through inversion methods (Zhu et al., 2011; Jethva et al., 2014). Even in this study, only very few points were available for comparison due to the limited number of direct measurements of columnar SSA. Despite this limitation, this comparison exercise provided confidence to generate global maps of SSA following this method.

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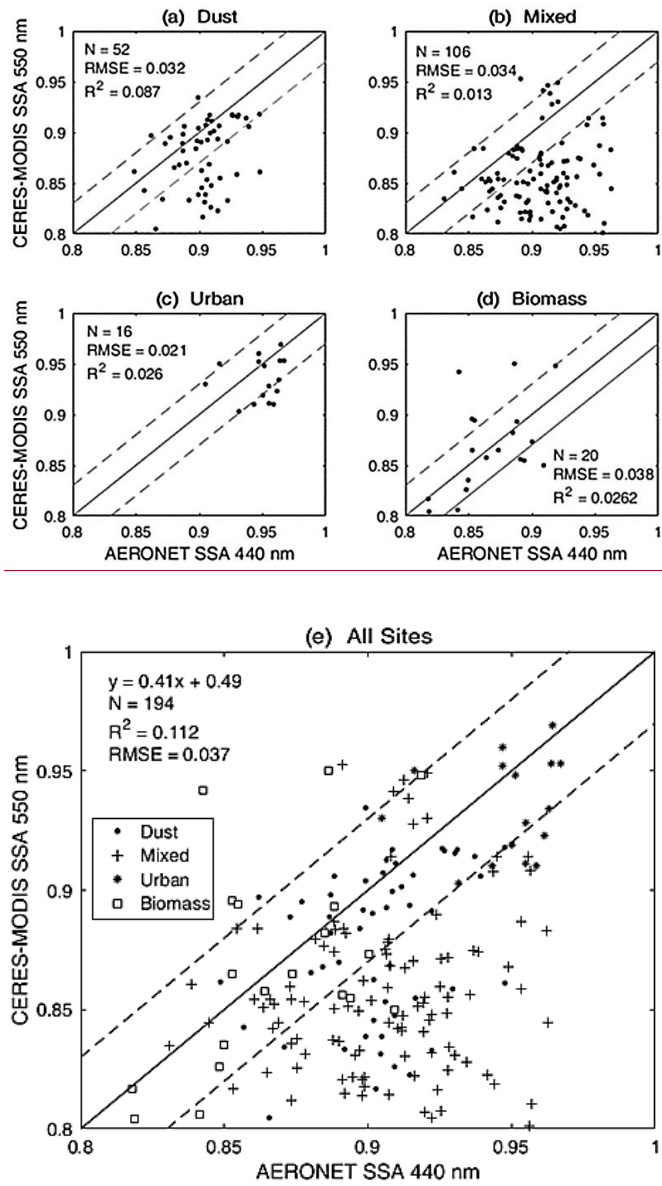
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## 7 Comparison with AERONET data

The Aerosol Robotic Network is a ground-based worldwide federated network of Cimel Sun photometers that measure extinction AOD from direct Sun measurements (Holben et al., 1998). The spectral diffuse sky radiations measured at different angles are inverted in conjunction with direct Sun measurements to derive the spectral SSA (440, 675, 870, and 1020 nm) and size distribution (Dubovik and King., 2000). The estimated uncertainty in retrieved SSA is largely attributed to the uncertainties in instrument calibration and is within 0.03 for AOD (440 nm) larger than 0.4. (Dubovik et al., 2000,2002).

AERONET version 3, level 2.0 monthly average values from selected sites were compared with corresponding CERES-MODIS SSA data. Sites were chosen to represent various types of aerosols following that of Giles et al., 2012. The location of the sites is shown in Fig S2 and Table S3. Scatter plots of comparison of AERONET SSA (440 nm) and CERES-MODIS SSA (550 nm) are shown in Fig 7.

Most AERONET SSA values are above 0.85, even in case of biomass burning aerosols. For dust type of aerosols (sites: Capo Verde, Dakar and Banizoumbaou), as shown in Fig 7a, AERONET and CERES-MODIS have better agreement. Whereas, in case of mixed type of aerosol (sites: SEDE BOKER, Kanpur, Xiang He and Illorin), most of the CERES-MODIS values are below 0.85, indicating highly absorbing type of aerosols (Fig 7b). For urban (sites: GSFC, Mexico city, Shirahama, Ispra and Moldova) and biomass (sites: Alta Floresta, Lake Argyle and Mongu) only very few data were available during the study period of 2014-18 as shown in Fig 7 panels c and d. Data points combined from all the sites are plotted together in Fig 7e showing a RMSE of 0.037. The large difference between SSA wavelengths of AERONET (440 nm) and CERES-MODIS (550 nm) could contribute to the variations observed between these two datasets. Overall, the resulting comparisons are agreeable within the uncertainties of both AERONET and CERES-MODIS datasets.



**Figure 7.** CERES-MODIS SSA (550 nm) vs AERONET SSA (440 nm) for various AERONET sites classified based on type of aerosols (Giles et al., 2012)

### 8. Summary and Conclusions

- Global maps of aerosol absorption have been generated following the concept of “critical optical depth”.
- The retrieved SSA values have been compared with available aircraft measurements. The limited comparison exercise shows that most of the retrieved SSA values are within  $\pm 0.03$ .
- We show that the combined CERES-MODIS algorithm better captures the spatial and seasonal variations in aerosol absorption and the resultant maps provide an improved global SSA database with fewer data gaps. Global mean SSA was estimated to be 0.93 and 0.97 over land and ocean, respectively
- The algorithm’s sensitivity to various parameters have been studied, which shows maximum sensitivity to changes in surface albedo. The algorithm is shown to be the most effective over regions with large AOD variations and less surface albedo variations.
- Comparison with SSA from 15 AERONET sites showed an acceptable agreement between AERONET and CERES-MODIS SSA, within their uncertainties.
- Overall, the combined CERES-MODIS algorithm provides global SSA maps with improved accuracy and better spatial coverage. These global maps provide valuable input for models to make assessment of aerosol-climate impacts on both regional and global scales.

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### Data Availability

MODIS and CERES data used in this study are available at <https://asdc.larc.nasa.gov/>. [The combined CERES-MODIS datasets are available upon request from the corresponding author.](#)

### Author Contributions

5 SKS conceptualized the method. AD developed the algorithm, carried out the simulations, and analyzed the data. AD wrote the manuscript with revisions from SKS.

### Competing interests

The authors declare they have no conflict of interest.

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## Supplementary Material

### Global maps of aerosol single scattering albedo using combined CERES-MODIS retrieval

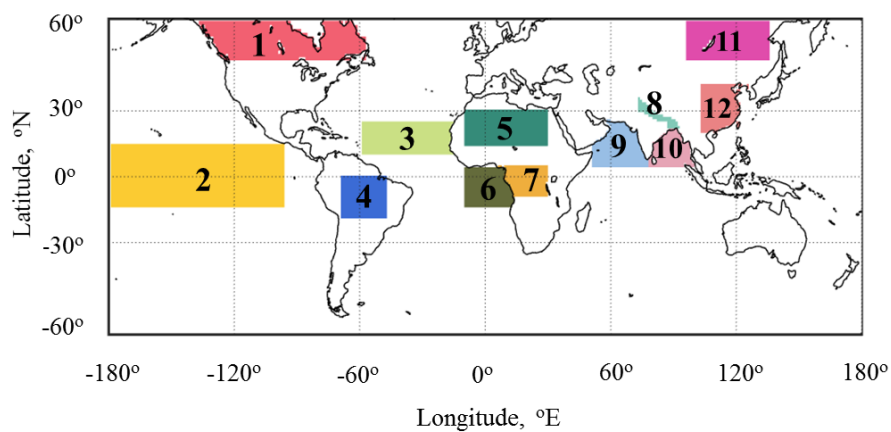
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**Figure S1.** Regions of interest (ROI). Details of each region are provided in Table S1

**Table S1.** Details of the regions shown in Fig. S1

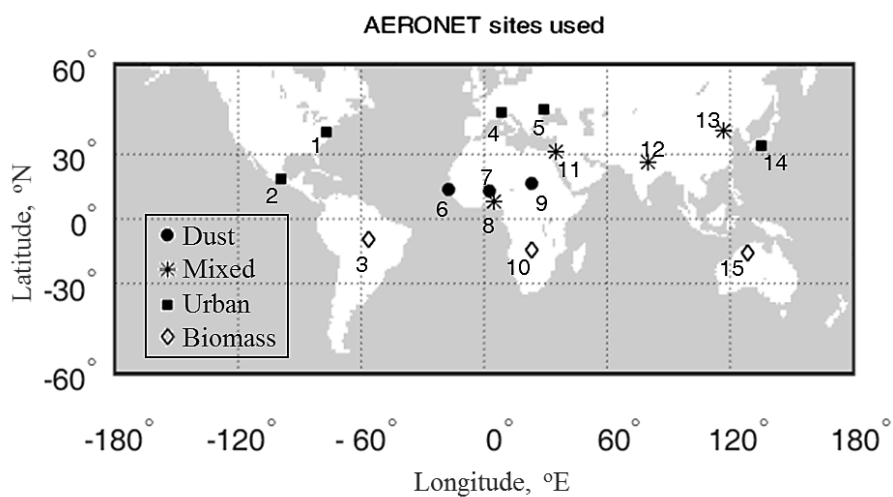
ROI No:	Region	General aerosol characteristics	Lat limit, °N	Lon limit, °E
1	Canadian Boreal Forest	Relatively pristine with seasonal biomass burning	48 to 60	-140 to -58
2	Eastern Pacific	Less polluted oceanic region	-15 to 15	-180 to -97
3	North East Atlantic	Highly polluted by dust transport and continental outflow from biomass burning	10 to 25	-60 to -10
4	Amazon	Relatively pristine with seasonal biomass burning	-20 to 0	-70 to -48
5	Sahara	Desert region with seasonal dust storms	14 to 30	-11 to 28
6	Southeast Atlantic	Highly polluted by dust transport and continental outflow from biomass burning	-15 to 4	-11 to 15
7	South African Forest	Relatively pristine with seasonal biomass burning	-10 to 5	3 to 29
8	Indo Gangetic Plain	A highly polluted industrial region with seasonal stubble burning and dust from the Thar desert	22 to 35	72 to 92
9	Arabian Sea	Continental outflow of pollution and dust	4 to 26	50 to 77
10	Bay of Bengal	Continental outflow of pollution	4 to 24	77 to 99
11	Russian Boreal Forest	Relatively pristine with seasonal biomass burning	48 to 60	95 to 135
12	Eastern China	A highly polluted industrial region	20 to 40	102 to 125

**Table S2.** Seasonal mean SSA over regions of interest from combined CERES-MODIS, OMI (given in round brackets) and POLDER (given in square brackets). Details of these regions are given in Table S1 and Fig. S1

Region	CERES-MODIS SSA 550 nm (OMI SSA 500 nm) [POLDER SSA 565 nm]			
	DJF	MAM	JJA	SON
Canadian Boreal Forest	<u>NODATA</u> <u>(0.95 ± 0.02)</u> <u>[0.96 ± 0.04]</u>	<u>0.96 ± 0.02</u> <u>(0.94 ± 0.01)</u> <u>[0.84 ± 0.05]</u>	<u>0.91 ± 0.02</u> <u>(0.94 ± 0.01)</u> <u>[0.89 ± 0.04]</u>	<u>0.94 ± 0.02</u> <u>(0.93 ± 0.01)</u> <u>[0.90 ± 0.05]</u>
Russian Boreal Forest	<u>NO DATA</u> <u>(0.95 ± 0.02)</u> <u>[0.81 ± 0.08]</u>	<u>0.96 ± 0.02</u> <u>(0.94 ± 0.01)</u> <u>[0.89 ± 0.03]</u>	<u>0.90 ± 0.01</u> <u>(0.94 ± 0.01)</u> <u>[0.91 ± 0.03]</u>	<u>0.96 ± 0.01</u> <u>(0.93 ± 0.01)</u> <u>[0.89 ± 0.05]</u>
South African Forest	<u>0.91 ± 0.02</u> <u>(0.93 ± 0.01)</u> <u>[0.84 ± 0.03]</u>	<u>0.92 ± 0.01</u> <u>(0.94 ± 0.01)</u> <u>[0.90 ± 0.03]</u>	<u>0.83 ± 0.01</u> <u>(0.93 ± 0.02)</u> <u>[0.88 ± 0.03]</u>	<u>0.90 ± 0.01</u> <u>(0.94 ± 0.01)</u> <u>[0.85 ± 0.05]</u>
Amazon Forest	<u>0.96 ± 0.02</u> <u>(0.95 ± 0.01)</u> <u>[0.84 ± 0.07]</u>	<u>0.98 ± 0.01</u> <u>(0.95 ± 0.01)</u> <u>[0.91 ± 0.05]</u>	<u>0.97 ± 0.02</u> <u>(0.93 ± 0.01)</u> <u>[0.92 ± 0.02]</u>	<u>0.89 ± 0.02</u> <u>(0.94 ± 0.01)</u> <u>[0.87 ± 0.04]</u>
North East Atlantic	<u>0.96 ± 0.02</u> <u>(0.90 ± 0.01)</u> <u>[0.94 ± 0.03]</u>	<u>0.94 ± 0.02</u> <u>(0.92 ± 0.01)</u> <u>[0.93 ± 0.01]</u>	<u>0.92 ± 0.02</u> <u>(0.95 ± 0.01)</u> <u>[0.93 ± 0.02]</u>	<u>0.93 ± 0.03</u> <u>(0.94 ± 0.01)</u> <u>[0.94 ± 0.01]</u>
South East Atlantic	<u>0.92 ± 0.02</u> <u>(0.92 ± 0.01)</u> <u>[0.88 ± 0.04]</u>	<u>0.94 ± 0.02</u> <u>(0.92 ± 0.01)</u> <u>[0.94 ± 0.01]</u>	<u>0.89 ± 0.01</u> <u>(0.91 ± 0.01)</u> <u>[0.88 ± 0.03]</u>	<u>0.92 ± 0.02</u> <u>(0.94 ± 0.01)</u> <u>[0.89 ± 0.03]</u>
Eastern Pacific	<u>0.97 ± 0.01</u> <u>(0.94 ± 0.02)</u> <u>[0.97 ± 0.01]</u>	<u>0.97 ± 0.01</u> <u>(0.95 ± 0.02)</u> <u>[0.95 ± 0.02]</u>	<u>0.96 ± 0.01</u> <u>(0.95 ± 0.02)</u> <u>[0.95 ± 0.02]</u>	<u>0.97 ± 0.01</u> <u>(0.95 ± 0.02)</u> <u>[0.93 ± 0.03]</u>
Sahara	<u>0.93 ± 0.01</u> <u>(0.92 ± 0.01)</u> <u>[0.90 ± 0.03]</u>	<u>0.93 ± 0.01</u> <u>(0.93 ± 0.01)</u> <u>[0.88 ± 0.03]</u>	<u>0.91 ± 0.02</u> <u>(0.94 ± 0.01)</u> <u>[0.87 ± 0.04]</u>	<u>0.92 ± 0.02</u> <u>(0.93 ± 0.01)</u> <u>[0.90 ± 0.03]</u>
Indo Gangetic Plain	<u>0.88 ± 0.01</u> <u>(0.92 ± 0.01)</u> <u>[0.89 ± 0.01]</u>	<u>0.87 ± 0.01</u> <u>(0.92 ± 0.01)</u> <u>[0.83 ± 0.02]</u>	<u>0.85 ± 0.02</u> <u>(0.95 ± 0.01)</u> <u>[0.77 ± 0.03]</u>	<u>0.83 ± 0.01</u> <u>(0.92 ± 0.01)</u> <u>[0.89 ± 0.01]</u>
Eastern China	<u>0.92 ± 0.01</u> <u>(0.92 ± 0.01)</u> <u>[0.91 ± 0.01]</u>	<u>0.90 ± 0.01</u> <u>(0.94 ± 0.01)</u> <u>[0.87 ± 0.02]</u>	<u>0.87 ± 0.01</u> <u>(0.95 ± 0.01)</u> <u>[0.84 ± 0.04]</u>	<u>0.88 ± 0.02</u> <u>(0.93 ± 0.01)</u> <u>[0.91 ± 0.03]</u>
Arabian Sea	<u>0.92 ± 0.01</u> <u>(0.91 ± 0.02)</u> <u>[0.94 ± 0.02]</u>	<u>0.89 ± 0.01</u> <u>(0.93 ± 0.01)</u> <u>[0.92 ± 0.02]</u>	<u>0.91 ± 0.01</u> <u>(0.96 ± 0.01)</u> <u>[0.94 ± 0.02]</u>	<u>0.89 ± 0.01</u> <u>(0.93 ± 0.02)</u> <u>[0.93 ± 0.02]</u>
Bay of Bengal	<u>0.91 ± 0.01</u> <u>(0.92 ± 0.01)</u> <u>[0.93 ± 0.02]</u>	<u>0.90 ± 0.01</u> <u>(0.94 ± 0.01)</u> <u>[0.91 ± 0.02]</u>	<u>0.91 ± 0.02</u> <u>(0.95 ± 0.01)</u> <u>[0.95 ± 0.02]</u>	<u>0.91 ± 0.02</u> <u>(0.94 ± 0.02)</u> <u>[0.93 ± 0.03]</u>

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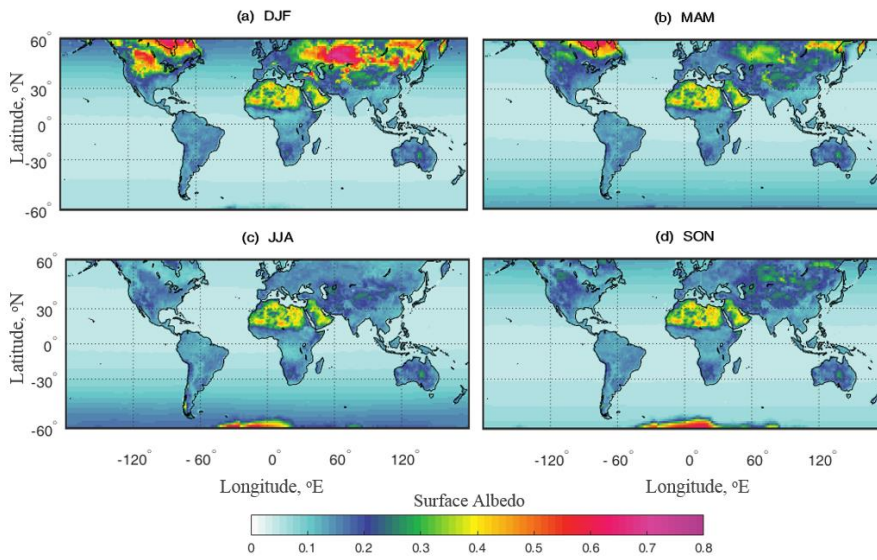


**Figure S2.** Map showing location of AERONET sites used in this study. The type of aerosols (dust, mixed, urban and biomass) were as defined in Giles et al., 2012

**Table S3:** Name of AERONET site as shown in Fig. S2

No.	Name	No.	Name	No.	Name
1	GSFC	6	Capo_Verde	11	SEDE_BOKER
2	Mexico_City	7	Dakar	12	Kanpur
3	Alta_Floresta	8	Illorin	13	XiangHe
4	Ispra	9	Banizoumbou	14	Shirahama
5	Moldova	10	Mongu	15	Lake_Argyle

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**Figure S3.** Seasonal mean shortwave-integrated surface albedo from CERES

**Table S4.** Shortwave integrated seasonal mean surface albedo from CERES over regions of interest. Details of these regions are given in Table S1 and Fig. S1

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Region	Surface Albedo			
	DJF	MAM	JJA	SON
Canadian Boreal Forest	$0.36 \pm 0.13$	$0.30 \pm 0.12$	$0.12 \pm 0.03$	$0.16 \pm 0.05$
Russian Boreal Forest	$0.37 \pm 0.10$	$0.27 \pm 0.08$	$0.13 \pm 0.02$	$0.20 \pm 0.05$
South African Forest	$0.12 \pm 0.01$	$0.13 \pm 0.01$	$0.12 \pm 0.02$	$0.13 \pm 0.01$
Amazon Forest	$0.14 \pm 0.01$	$0.14 \pm 0.01$	$0.13 \pm 0.02$	$0.14 \pm 0.02$
North East Atlantic	$0.06 \pm 0.01$	$0.05 \pm 0.01$	$0.05 \pm 0.01$	$0.05 \pm 0.01$
South East Atlantic	$0.05 \pm 0.01$	$0.05 \pm 0.01$	$0.05 \pm 0.01$	$0.05 \pm 0.01$
Eastern Pacific	$0.05 \pm 0.01$	$0.05 \pm 0.00$	$0.05 \pm 0.01$	$0.05 \pm 0.00$
Sahara	$0.35 \pm 0.06$	$0.34 \pm 0.06$	$0.34 \pm 0.06$	$0.34 \pm 0.06$
Indo Gangetic Plain	$0.13 \pm 0.02$	$0.13 \pm 0.02$	$0.14 \pm 0.02$	$0.13 \pm 0.01$
Eastern China	$0.13 \pm 0.04$	$0.13 \pm 0.03$	$0.13 \pm 0.03$	$0.13 \pm 0.03$
Arabian Sea	$0.06 \pm 0.01$	$0.05 \pm 0.01$	$0.05 \pm 0.02$	$0.05 \pm 0.01$
Bay of Bengal	$0.05 \pm 0.01$	$0.05 \pm 0.01$	$0.05 \pm 0.01$	$0.05 \pm 0.01$

**Table S5:** Normalized extinction coefficient of the aerosol model

$\lambda$ ( $\mu\text{m}$ )	$\text{Ext}_{\text{norm}}$	$\lambda$ ( $\mu\text{m}$ )	$\text{Ext}_{\text{norm}}$	$\lambda$ ( $\mu\text{m}$ )	$\text{Ext}_{\text{norm}}$
0.25	1.597	0.75	0.847	3.2	0.5075
0.3	1.467	0.8	0.8202	3.39	0.5047
0.35	1.334	0.9	0.7828	3.5	0.5062
0.4	1.224	1	0.7536	3.75	0.4828
0.45	1.135	1.25	0.7038	4	0.4629
0.5	1.061	1.5	0.6706	4.5	0.4395
0.55	1	1.75	0.6349	5	0.4103
0.6	0.9505	2	0.5883		
0.65	0.9106	2.5	0.4905		
0.7	0.8757	3	0.491		

**Table S6:** Phase function of the aerosol model (continued into Table S7)

$\lambda$ ( $\mu\text{m}$ )	Streams							
	1	2	3	4	5	6	7	8
0.25	0.754	0.606	0.473	0.397	0.342	0.307	0.283	0.265
0.3	0.738	0.589	0.452	0.379	0.325	0.293	0.270	0.254
0.35	0.738	0.592	0.456	0.386	0.333	0.303	0.279	0.264
0.4	0.741	0.598	0.463	0.395	0.343	0.313	0.290	0.275
0.45	0.743	0.602	0.468	0.403	0.351	0.323	0.299	0.284
0.5	0.746	0.607	0.474	0.411	0.359	0.331	0.308	0.292
0.55	0.748	0.611	0.478	0.416	0.364	0.337	0.313	0.297
0.6	0.749	0.615	0.481	0.421	0.368	0.342	0.317	0.301
0.65	0.750	0.618	0.485	0.426	0.373	0.347	0.321	0.305
0.7	0.751	0.620	0.487	0.429	0.376	0.350	0.323	0.306
0.75	0.752	0.623	0.490	0.433	0.378	0.352	0.325	0.308
0.8	0.755	0.628	0.494	0.437	0.382	0.355	0.327	0.310
0.9	0.756	0.631	0.496	0.440	0.383	0.356	0.326	0.308
1	0.756	0.632	0.496	0.440	0.382	0.354	0.323	0.304
1.25	0.766	0.643	0.505	0.442	0.380	0.346	0.314	0.291
1.5	0.777	0.651	0.512	0.441	0.376	0.337	0.302	0.276
1.75	0.798	0.673	0.536	0.455	0.385	0.339	0.300	0.271
2	0.826	0.707	0.577	0.491	0.415	0.362	0.316	0.282
2.5	0.858	0.750	0.636	0.552	0.476	0.418	0.365	0.323
3	0.871	0.765	0.662	0.578	0.505	0.444	0.391	0.346
3.2	0.836	0.708	0.584	0.491	0.414	0.354	0.304	0.264
3.39	0.818	0.682	0.548	0.453	0.375	0.317	0.270	0.233
3.5	0.808	0.670	0.530	0.434	0.356	0.299	0.253	0.217
3.75	0.805	0.667	0.524	0.429	0.349	0.292	0.246	0.210
4	0.797	0.660	0.513	0.421	0.340	0.284	0.238	0.202
4.5	0.795	0.655	0.507	0.413	0.331	0.275	0.228	0.192
5	0.808	0.663	0.520	0.420	0.338	0.278	0.230	0.192

**Table S7:** Phase function of aerosol model

$\lambda$ ( $\mu\text{m}$ )	Streams							
	9	10	11	12	13	14	15	16
0.25	0.252	0.241	0.233	0.226	0.219	0.214	0.209	0.204
0.3	0.241	0.232	0.224	0.217	0.211	0.205	0.200	0.196
0.35	0.251	0.242	0.233	0.226	0.219	0.214	0.208	0.203
0.4	0.262	0.252	0.243	0.235	0.228	0.222	0.216	0.210
0.45	0.270	0.260	0.251	0.242	0.235	0.228	0.221	0.215
0.5	0.278	0.267	0.257	0.248	0.240	0.233	0.226	0.219
0.55	0.283	0.271	0.261	0.251	0.243	0.235	0.227	0.220
0.6	0.286	0.274	0.263	0.253	0.244	0.235	0.228	0.220
0.65	0.289	0.277	0.265	0.255	0.245	0.236	0.228	0.220
0.7	0.290	0.277	0.265	0.254	0.244	0.235	0.226	0.218
0.75	0.291	0.277	0.265	0.253	0.243	0.233	0.225	0.216
0.8	0.292	0.278	0.265	0.253	0.242	0.232	0.223	0.214
0.9	0.289	0.274	0.261	0.248	0.237	0.226	0.217	0.208
1	0.284	0.269	0.254	0.241	0.230	0.219	0.209	0.200
1.25	0.271	0.253	0.238	0.224	0.212	0.200	0.190	0.180
1.5	0.255	0.236	0.220	0.205	0.193	0.181	0.171	0.161
1.75	0.246	0.226	0.208	0.193	0.180	0.168	0.157	0.147
2	0.252	0.229	0.208	0.191	0.176	0.163	0.151	0.141
2.5	0.287	0.257	0.231	0.208	0.189	0.172	0.157	0.144
3	0.307	0.274	0.245	0.220	0.198	0.179	0.162	0.148
3.2	0.231	0.203	0.180	0.161	0.144	0.130	0.117	0.107
3.39	0.203	0.178	0.157	0.140	0.125	0.113	0.102	0.093
3.5	0.188	0.165	0.146	0.130	0.116	0.104	0.094	0.085
3.75	0.181	0.157	0.139	0.122	0.109	0.097	0.088	0.079
4	0.174	0.150	0.132	0.116	0.103	0.091	0.082	0.073
4.5	0.164	0.141	0.122	0.107	0.094	0.083	0.073	0.066
5	0.163	0.139	0.120	0.103	0.090	0.079	0.070	0.062