

1 **Mercury isotopic compositions in fine particles and offshore surface**
2 **seawater in a coastal area of East China: Implication for Hg sources**
3 **and atmospheric transformations**

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30 **Abstract.** Isotopic compositions of Hg in atmospheric particles (Hg_{PM}) are probably
31 the mixed results of emission sources and atmospheric processes. Here, we present Hg
32 isotopic compositions in daily fine particles ($PM_{2.5}$) collected from an industrial site
33 (CX) and a nearby mountain site (DMS) in a coastal area of East China, as well as in
34 surface seawater close to the industrial area, to reveal the influence of anthropogenic
35 emission sources and atmospheric transformations on Hg isotopes. The $PM_{2.5}$ samples
36 displayed a significant spatial difference in $\delta^{202}Hg$. For the CX, the negative $\delta^{202}Hg$
37 values are similar to those of source materials and the Hg_{PM} contents were well
38 correlated with chemical tracers, indicating the dominant contributions of local
39 industrial activities to $Hg_{PM_{2.5}}$. Whereas the observed positive $\delta^{202}Hg$ at the DMS was
40 likely associated with regional emissions and extended atmospheric processes during
41 transport. The $\Delta^{199}Hg$ values in $PM_{2.5}$ from the CX and DMS were comparable
42 positive. The unity slope of $\Delta^{199}Hg$ versus $\Delta^{201}Hg$ over all data suggests that the
43 odd-MIF of $Hg_{PM_{2.5}}$ was primarily induced by photo-reduction of Hg^{2+} in aerosols.
44 The positive $\Delta^{200}Hg$ values with a minor spatial difference were probably associated
45 with photo-oxidation of Hg^0 which is generally enhanced in the coastal environment.
46 Total Hg in offshore surface seawater was characterized by negative $\delta^{202}Hg$ and
47 near-zero $\Delta^{199}Hg$ and $\Delta^{200}Hg$ values, which are indistinguishable from Hg isotopes of
48 source materials. Overall, the $PM_{2.5}$ collected from industrial areas had comparable
49 $\delta^{202}Hg$ values but more positive $\Delta^{199}Hg$ and $\Delta^{200}Hg$ as compared to surface seawater.
50 The results indicate that atmospheric transformations would induce significant
51 fractionation of Hg isotopes and obscure the Hg isotopic signatures of anthropogenic
52 emissions.

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54 **Keywords:** Particle bound mercury; Surface seawater; Hg isotopes; Mercury sources;
55 Photo-chemical processes.

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60 **1. Introduction**

61 Mercury (Hg) is a genotoxic element and was ranked with the priority controlled
62 pollutants in many countries. Atmospheric Hg was operationally defined as three
63 forms: gaseous elemental mercury (GEM), gaseous oxidized mercury (GOM), and
64 particle bound mercury (PBM or Hg_{PM}) (Schroeder and Munthe, 1998). Previous
65 studies indicated that Hg_{PM} concentrations in urban and industrial areas could reach
66 up to hundreds even thousands of pg m⁻³, relative to tens of pg m⁻³ in uncontaminated
67 remote areas (Fu et al., 2015; Mao et al., 2016). In addition, Hg_{PM} can account for up
68 to 40% of atmospheric Hg in industrial areas, relative to < 5% in uncontaminated
69 areas (Guo et al., 2022; Schroeder and Munthe, 1998). Hence, particulate matter (PM)
70 can act as a vector of toxic Hg and inhalation of Hg-carrying particles is an important
71 pathway of human exposure to atmospheric Hg. Coal combustion, non-ferrous metal
72 smelting, and cement production were the three primary anthropogenic sources of
73 atmospheric Hg, which are responsible for 47% of total global Hg emissions (GMA,
74 2018). On the other hand, Hg_{PM} undergoes complex transport and transformation
75 processes in the atmosphere. Hg_{PM} can be formed by uptake of GOM in particles,
76 which made an important contribution to Hg_{PM} in the heavily particle polluted areas
77 (Xu et al., 2020). Whereas reduction of GOM binding with dissolved organic carbon
78 ligands in aqueous particles potentially converts it back to the gas phase (Horowitz et
79 al., 2017). In addition, Hg_{PM} has a residence time of several weeks as it can transport
80 and deposit at a regional scale (Selin, 2009). The research has suggested that
81 atmospheric Hg_{PM} is generally a combined result of anthropogenic emissions
82 and atmospheric processes, which plays a crucial role in the global cycling of Hg.

83 Analysis technique of Hg isotopes and mechanisms of Hg isotopic fractionation
84 have come a long way in the last decade (Blum and Johnson, 2017). Hg has seven
85 stable isotopes (including ¹⁹⁶Hg, ¹⁹⁸Hg, ¹⁹⁹Hg, ²⁰⁰Hg, ²⁰¹Hg, ²⁰²Hg, and ²⁰⁴Hg) and
86 exhibits mass dependent fractionation (MDF) and mass independent fractionation
87 (MIF) in various environmental samples and processes (Blum and Johnson, 2017;
88 Sonke and Blum, 2013; Yin et al., 2014a). MDF of Hg isotopes is often reported as
89 δ²⁰²Hg, while MIF of odd mass-numbered Hg isotopes (odd-MIF) is reported as

90 $\Delta^{199}\text{Hg}$ and $\Delta^{201}\text{Hg}$ and MIF of even Hg isotopes (even-MIF) as $\Delta^{200}\text{Hg}$ and $\Delta^{204}\text{Hg}$.
91 Previous laboratory and field studies have revealed that nearly all biogeochemical
92 processes induce MDF of Hg isotopes, whereas significant odd-MIF of Hg occurs
93 mainly in photochemical processes (Bergquist and Blum, 2007; Blum et al., 2014;
94 Malinovsky et al., 2010; Sun et al., 2016a). What's more, specific ratios of
95 $\Delta^{199}\text{Hg}/\Delta^{201}\text{Hg}$ have been reported for different transformation processes, i.e., ~ 1.0 for
96 photo-reduction of Hg^{2+} and ~ 1.6 for photo-oxidation of Hg^0 (Bergquist and Blum,
97 2007; Sun et al., 2016a). Even-MIF of Hg isotopes is observed mostly in atmosphere
98 related samples, which is suggested to associate with photo-oxidation of Hg^0 by UV
99 and oxidants (Blum and Johnson, 2017; Chen et al., 2012; Fu et al., 2019). Therefore,
100 Hg isotopes are capable of becoming useful tracers for the biogeochemical cycles of
101 Hg in the environment.

102 There is a large difficulty in sampling enough Hg mass for analyzing Hg isotopes
103 of atmospheric samples. Even so, a few public studies have proved that Hg isotopes
104 are useful tools to investigate potential sources and transformation processes of Hg in
105 the atmosphere. The PM from major cities, like in northern and western China, mostly
106 displayed significant negative MDF and near zero odd-MIF due to the dominant
107 impact of anthropogenic emissions (Huang et al., 2015, 2016, 2019, 2020; Xu et al.,
108 2017, 2019; Yu et al., 2016). A previous study conducted in remote areas of China has
109 reported that the Hg_{PM} exposed to air masses of regional and long-range sources had
110 distinct isotopic signatures (Fu et al., 2019). Recently, many studies have used Hg
111 isotopes to investigate the contribution of domestic emissions and transboundary Hg
112 transport to atmospheric Hg_{PM} . These studies have implied that the long-range
113 transboundary Hg transport from South Asia played a crucial role in the Himalayas
114 and the Tibetan Plateau, even in the southwestern and northwestern China (Fu et al.,
115 2019; Guo et al., 2021, 2022). East China is densely populated and one of the heaviest
116 industrialized area in China. The concentration of Hg_{PM} in this region has been well
117 characterized (Hong et al., 2016; Xu et al., 2020; Yu et al., 2015), but only two studies
118 conducted at the remote sites have referred to Hg_{PM} isotopes (Fu et al., 2019; Yu et al.,
119 2016). To the best of our knowledge, there is no report on the isotopic compositions of

120 Hg_{PM} from urban areas of East China. Likewise, the effect of atmospheric processes
121 on the fractionation of Hg isotopes in the coastal region has not been well elucidated.

122 This study determined Hg isotopic compositions in PM_{2.5} collected from an
123 industrial site and a mountain site in a coastal area of East China. Comparison of
124 Hg_{PM2.5} isotopes at the neighboring sites would eliminate the impacts of meteorology
125 and atmospheric Hg background which vary across space on Hg_{PM} isotopes.
126 Furthermore, this study measured the isotopic compositions of total mercury (THg) in
127 surface seawater close to the industrial area and distinguished Hg isotopes between
128 the atmospheric sample and surface medias. The objectives of this study are (1) to
129 differentiate the Hg isotopes in PM_{2.5} from the two neighboring industrial and
130 mountain sites; (2) use the Hg isotopes to explore the influence of anthropogenic
131 sources on the Hg_{PM}; (3) to reveal the role of atmospheric transformations in varying
132 Hg_{PM} isotopic compositions.

133 **2. Experiment**

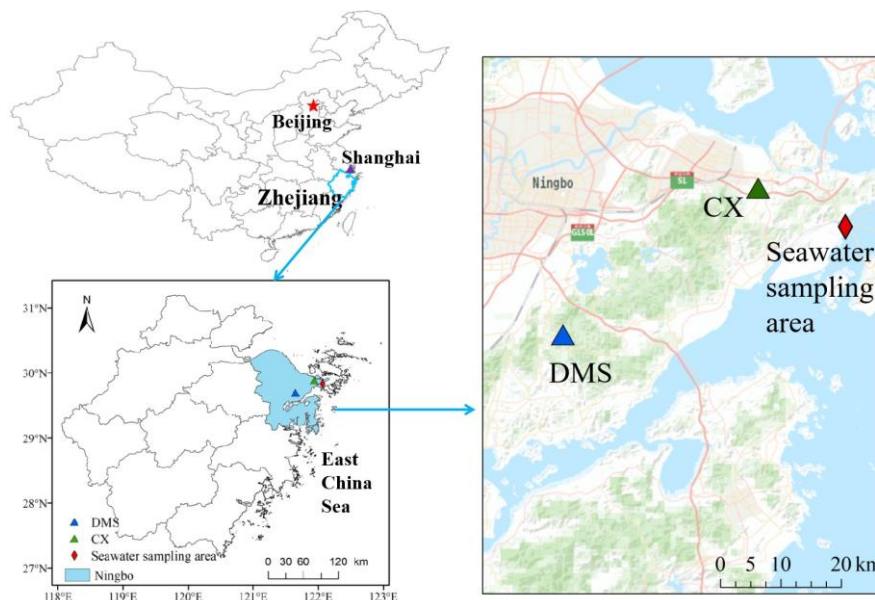
134 *2.1. Study area description*

135 PM_{2.5} sampling was conducted at an industrial site (Chunxiao, CX) and a nearby
136 mountain site (Daimeishan, DMS) on the east coast of Zhejiang province, East China
137 (Fig. 1). The study region experiences a typical subtropical monsoon climate, with sea
138 breeze in summer and continental breeze in winter. The average annual temperature,
139 precipitation, relative humidity, and sunshine hours were 18.1 °C, 1608 mm, 76.8%,
140 and 1797 h, respectively.

141 The CX (121.91 °E, 29.87 °N, 15 m a.g.l.) is located in the Urban Environment
142 Observation and Research Station, Chinese Academy of Sciences, Beilun District,
143 Ningbo. Ningbo is a highly industrial city and there is a high density of industrial
144 activities around the CX. Potential Hg point sources include a large coal-fired power
145 plant (5000 MW) approximately 20 km to the northwest, a Chlor-alkali plant 20 km to
146 the northeast, and an automobile assembly plant within 1 km of the site. The CX is in
147 close proximity to the East China Sea (ECS, ~ 0.6 km), thus clean air masses from the
148 sea in warm seasons would dilute atmospheric Hg at the CX. The DMS (121.62 °E,
149 29.68 ° N, 450 m a.s.l.) is located at the summit of Mountain Damei, which is

150 surrounded by trees. The site is 20 km to the coast of the ECS and approximately 22
151 km south of Ningbo. There are no significant Hg point sources within a radius of ~10
152 km from the DMS. However, an early study reported that intense regional emissions,
153 like industrial activities and coal combustion in the Yangtze River Delta and the
154 neighboring region of Anhui, Jiangsu, and Zhejiang Provinces, caused a high
155 atmospheric Hg concentration at the DMS (Yu et al., 2015).

156 Surface seawater samples were collected in the offshore area of Ningbo. The
157 seawater sampling area (about 122.04 ° E, 29.82 ° N, Fig. 1) is approximately 1 km
158 west of the Beilun District, Ningbo. The salinity of the seawater samples ranged from
159 21.2‰ to 29.5‰. The pH of the seawater samples was in the range of 5.7 ~ 8.5, with
160 the mean value of 7.5 ± 0.6 .



161
162 **Fig. 1** Locations of PM_{2.5} (CX: industrial site; DMS: mountain site) and surface
163 seawater sampling area.

164 2.2. Sample collection and analysis

165 2.2.1. Sampling of PM_{2.5}

166 The period of PM_{2.5} sampling was from Jul. 2017 to Jun. 2018. Daily PM_{2.5}
167 samples were collected 1~2 times a week at the CX (except Jan. and Feb. 2018) and
168 once a week at the DMS. Field blank samples were collected at each site. PM_{2.5}
169 samples were collected on a preheated quartz-fiber filter (500 °C, 4 h, 8 × 10 inches,
170 Whatman) using a high volume sampler (Tianhong TH1000H, China) with a flow rate

171 of $1.05 \text{ m}^3 \text{ min}^{-1}$. The filters were conditioned at $24 \pm 1 \text{ }^\circ\text{C}$ and $52 \pm 2\%$. The mass
172 loading of $\text{PM}_{2.5}$ on filters was determined by mass difference before and after
173 sampling. The filter samples were wrapped in aluminum foils and stored at $-20 \text{ }^\circ\text{C}$
174 until analysis.

175 *2.2.2. Concentration of Hg and other chemical species in $\text{PM}_{2.5}$*

176 Six punches (ca. 0.5 cm^2 per punch) of each sampling filter were digested by a
177 10 mL of 40% aqua regia (HNO_3 : $\text{HCl} = 1:3$, v/v) in a water bath at $95 \text{ }^\circ\text{C}$ for 5 min,
178 then the solution was oxidized by 1 mL of 0.2 M BrCl and bathed for another 30 min.
179 After cooling down, the extracted solution was diluted to 15 mL with ultra-pure water
180 and then analyzed by cold-vapor atomic fluorescence spectrometry (CVAFS, Brooks
181 Rand Model III, USA) following the EPA method 1631. The content of Hg on blank
182 filters can be negligible (42.5 pg at the CX and 27.0 pg at the DMS) relative to those
183 on sample filters.

184 $\text{PM}_{2.5}$ samples selected for Hg isotopes analysis were also measured for 8 water
185 soluble inorganic ions (K^+ , Ca^{2+} , Na^+ , Mg^{2+} , Cl^- , SO_4^{2-} , NO_3^- , and NH_4^+), elemental
186 carbon (EC), organic carbon (OC), and levoglucosan. The water soluble ions were
187 analyzed by ion chromatography (ICS-3000, Dionex, USA). EC and OC were
188 analyzed using a carbon analyzer (Model 4, Sunset Lab., USA) and NOISH protocol.
189 Analytical procedures and quality control procedures have been described by Xu et al.
190 (2018). Levoglucosan, an excellent indicator of biomass burning, was analyzed using
191 a gas chromatography – mass spectrometer detector (GC – MS, Agilent 7890A-5975C,
192 Agilent Tech. Inc., USA). Levoglucosan analytical procedures have been presented in
193 detail elsewhere (Hong et al., 2019).

194 *2.2.3. Sampling and analysis of Hg in seawater*

195 Seawater samples were collected from the surface of the offshore sampling area
196 twice a month during Jul. 2017 ~ Jun. 2018, except Feb. 2018. Each time, three ~50
197 mL of duplicate seawater samples were collected for THg content analysis. Final THg
198 content was determined by the average of three duplicate samples. In addition, ~2 L
199 surface seawater was sampled for Hg isotopes analysis each time. The seawater
200 samples were stored in brown glass bottles and preserved with 1% (v/v) HCl in the

201 laboratory. They were analyzed for Hg content and isotopic compositions in a month.

202 Total Hg content in seawater samples was measured by CVAFS (Brooks Rand
203 Model III, USA). A 25 mL of seawater sample was digested with 0.2 M BrCl at least
204 12 h in advance and then analyzed using the EPA method 1631. More details can be
205 found in a previous study (Xu et al., 2014). Method blank was processed by bottles
206 filling up with ultra-pure water instead of seawater. The blank was lower than 10 pg
207 (n = 15), which can be negligible compared to the samples.

208 *2.3. Analysis of Hg isotopic compositions*

209 *2.3.1. PM_{2.5} sample processing*

210 Due to the effects of precipitation and short sampling duration, the mass of Hg
211 on most of PM_{2.5} samples was not sufficient for isotopes detection. A total of 20 PM_{2.5}
212 samples, including 10 at the CX and 10 at the DMS, were chosen for Hg isotopes
213 analysis. Pre-concentration of Hg from PM_{2.5} samples was conducted following a
214 dual-stage combustion protocol (Huang et al., 2015). To be specific, a tube furnace
215 (OTF-1200X-II, Kejin, China) consisting of two combustion stages was used. A
216 sampling filter was embedded in a furnace quartz tube (50 mm OD, 46 mm ID, 80 cm
217 length). The tube was then placed in the furnace so that the filter was at the first
218 combustion stage. The second decomposition stage was heated up in advance and
219 maintained at 1000 °C, then the first combustion stage was heated up to 950 °C
220 through a temperature-programmed procedure. The combustion procedure was run
221 with no samples in the furnace quartz tube before PM_{2.5} sample treatment every day to
222 remove residual volatiles. The released Hg was transferred by O₂/Ar gas (30%/70%)
223 at a flow rate of 20 mL min⁻¹ and then trapped by a 10 mL of 40% inverse aqua regia
224 (2: 4: 9 ratio of 10 M HCl, 15 M HNO₃ and ultra-pure water) in a designed glass
225 bottle. In advance of PM_{2.5} sample analysis, the accuracy of dual-stage combustion
226 method was assessed by the analysis of the standard NIST SRM 3133 Hg (dripped on
227 blank filters) and the certified reference material GBW07434. The Hg recovery
228 efficiency of the dual-stage protocol was in the range of 87.6% ~ 103.3% (mean: 95.0
229 ± 5.1%, n = 6).

230 *2.3.2. Seawater sample processing*

231 A total of 20 seawater samples were analyzed for Hg isotopes. ~2 L seawater
232 sample was mixed with a 4 mL of 300 g L⁻¹ NH₂OH HCl solution for neutralizing
233 excess BrCl and then an 8 mL of 200 g L⁻¹ SnCl₂ solution for reducing the oxidized
234 Hg. The pre-treated seawater sample was stirred and bubbled for 1 h with Hg-free N₂
235 at a flow rate of 400 mL min⁻¹. The gaseous Hg purged off seawater samples was
236 collected by a series of three gold traps. The gold traps were heated and the released
237 Hg was transferred by Hg-free N₂ at 10~15 mL min⁻¹ and concentrated by a 10 mL of
238 40% inverse aqua regia.

239 2.3.3. Hg isotopes analysis

240 All trapping solutions were preserved with 1% (v/v) BrCl and stored at 4 °C in
241 the dark before Hg isotopes analysis. Hg isotopic compositions were measured by a
242 multi-collector inductively coupled plasma mass spectrometer (MC-ICP-MS, Nu
243 Instruments Ltd. UK) equipped with an introduction device following the protocols
244 presented in previous studies (Huang et al., 2015; Huang et al., 2021; Lin et al., 2015).
245 The introduction device includes a modified cold-vapor generator (CVG) and an
246 Aridus III nebulizer for respective Hg and Tl introduction. Between standard and
247 sample, the CVG was rinsed with 3% (v/v) HNO₃ solution to ensure the Hg signal
248 returned to the background level. Instrument mass bias was corrected using both an
249 internal standard (NIST 997 Tl) and a strict sample-standard bracketing method
250 (NIST 3133 Hg). A reference material NIST 8610 was measured repeatedly for
251 quality control. The pre-concentration solutions were diluted to about 1.5 ~ 3.0 ng
252 mL⁻¹ and the NIST 3133 and NIST 8610 were kept at 2.0 ng mL⁻¹ during the analysis
253 period. The MDF of Hg (represented by δ-value, ‰) is defined by the following
254 equation (Blum and Bergquist, 2007):

$$255 \quad \delta^{\text{xxx}}\text{Hg} (\text{‰}) = [(\text{xxxHg}/^{198}\text{Hg})_{\text{sample}}/(\text{xxxHg}/^{198}\text{Hg})_{\text{NIST 3133}} - 1] \times 1000 \quad (1)$$

256 where xxx = 199, 200, 201, and 202. The MIF of Hg (Δ-value, ‰) is calculated using
257 the theoretically predicted MDF as the following equation (Blum and Bergquist,
258 2007):

$$259 \quad \Delta^{\text{xxx}}\text{Hg} (\text{‰}) = \delta^{\text{xxx}}\text{Hg} - (\delta^{202}\text{Hg} \times \beta) \quad (2)$$

260 where the mass-dependent scaling factor β is 0.252 for ¹⁹⁹Hg, 0.502 for ²⁰⁰Hg, and

261 0.752 for ^{201}Hg . The repeated measurement of NIST 8610 during the analysis session
262 yielded $\delta^{202}\text{Hg}$ and $\Delta^{199}\text{Hg}$ values of $-0.60 \pm 0.15\text{‰}$ and $-0.02 \pm 0.06\text{‰}$ (2σ , $n = 7$). In
263 addition, a well-known reference material UM-Almaden showed a long-term average
264 of $\delta^{202}\text{Hg} = -0.59 \pm 0.10\text{‰}$ and $\Delta^{199}\text{Hg} = -0.03 \pm 0.07\text{‰}$ (2σ , $n = 25$), which are well
265 consistent with those in previous studies (Blum and Bergquist, 2007; Huang et al.,
266 2015). The samples of this study were measured only once, so the 2σ uncertainties
267 derived from repeated measurements of NIST 3133 standard during each analysis
268 section were applied to the samples.

269 **3. Results and discussion**

270 ***3.1. Concentrations and isotopic compositions of $\text{Hg}_{\text{PM}_{2.5}}$***

271 Concentrations and isotopes of $\text{Hg}_{\text{PM}_{2.5}}$ at industrial and mountain sites are
272 showed in Table 1. Mean volumetric concentrations of $\text{Hg}_{\text{PM}_{2.5}}$ were $16.3 \pm 17.8 \text{ pg}$
273 m^{-3} at the CX and $29.6 \pm 35.9 \text{ pg m}^{-3}$ at the DMS, which are comparable to those
274 observed at remote sites (Fu et al., 2019), but lower than those reported from urban
275 sites in China (Xu et al., 2019). The low volumetric concentrations of $\text{Hg}_{\text{PM}_{2.5}}$
276 observed in this study were likely associated with low $\text{PM}_{2.5}$ concentrations (mean:
277 $28.0 \text{ }\mu\text{g m}^{-3}$ at the CX and $34.8 \text{ }\mu\text{g m}^{-3}$ at the DMS) during the study period. Average
278 mass concentrations of $\text{Hg}_{\text{PM}_{2.5}}$ were $0.52 \pm 0.23 \text{ }\mu\text{g g}^{-1}$ (0.15 to 1.10, $n = 51$) at the
279 CX and $0.85 \pm 0.63 \text{ }\mu\text{g g}^{-1}$ (0.18 to 2.80, $n = 33$) at the DMS, respectively. A
280 relatively high concentration of $\text{Hg}_{\text{PM}_{2.5}}$ has been reported at the DMS before, which
281 was likely due to regional Hg emissions, mainly the industrial activities and coal
282 combustion in the Yangtze River Delta and the neighboring region of Anhui, Jiangsu,
283 and Zhejiang Provinces (Yu et al., 2015). The Hg contents in $\text{PM}_{2.5}$ of this study are
284 higher than those of natural sources (e.g., dust and topsoil; $0.056 \sim 0.30 \text{ }\mu\text{g g}^{-1}$;
285 Schleicher et al., 2015) and those of coals in China (mean: $0.22 \text{ }\mu\text{g g}^{-1}$; Yin et al.,
286 2014b), which implies a potential contribution of anthropogenic sources with high Hg
287 contents. The volumetric concentrations of $\text{Hg}_{\text{PM}_{2.5}}$ were closely correlated to
288 $\text{Hg}_{\text{PM}_{2.5}}/\text{PM}_{2.5}$ ratios both at the CX and DMS ($R^2 = 0.50$ and 0.60 , $p < 0.01$),
289 suggesting that atmospheric $\text{Hg}_{\text{PM}_{2.5}}$ concentrations were dependent on Hg contents.
290 Whereas, a weak correlation between the $\text{Hg}_{\text{PM}_{2.5}}$ volumetric concentrations and $\text{PM}_{2.5}$

291 masses was observed at the DMS ($R^2 = 0.25$, $p < 0.01$) in contrast to the CX ($R^2 =$
292 0.77 , $p < 0.01$). The result likely indicates that the DMS $\text{Hg}_{\text{PM}_{2.5}}$ was influenced by
293 diverse sources of $\text{PM}_{2.5}$ with different Hg levels and/or complex atmospheric Hg
294 transformations (Xu et al., 2019). This is supported by the higher variation coefficient
295 (VC = SD/Mean) of $\text{Hg}_{\text{PM}_{2.5}}$ mass concentrations at the DMS (74.1%) than the CX
296 (44.2%). Spatial differences of $\text{Hg}_{\text{PM}_{2.5}}$ were further examined by relationships of Hg
297 with other chemical species in $\text{PM}_{2.5}$ (Table S1). In contrast to DMS, the mass
298 concentrations of $\text{Hg}_{\text{PM}_{2.5}}$ at the CX were well correlated to chemical tracers, like Cl^- ,
299 NO_3^- , K^+ , and OC ($r = 0.40 \sim 0.57$, $p < 0.05$, Spearman correlation), implying the
300 contributions of steady anthropogenic sources to $\text{Hg}_{\text{PM}_{2.5}}$ in the industrial area.

301 $\delta^{202}\text{Hg}$ values for $\text{Hg}_{\text{PM}_{2.5}}$ at the CX were in the range of -1.11‰ to 0.08‰ (mean:
302 $-0.61 \pm 0.35\text{‰}$, $n = 10$), while $\delta^{202}\text{Hg}$ values at the DMS were significantly higher and
303 in a larger variation from -0.78‰ to 1.10‰ (mean: $0.12 \pm 0.63\text{‰}$, $n = 10$) ($p < 0.05$, T
304 Test; Table 1 and Table S2). Hg_{PM} isotopic compositions in multiple types of locations
305 are showed in Fig. 2 and Table S3. The $\delta^{202}\text{Hg}$ values at the CX basically overlap
306 those for PM in urban areas of China (mean: from -1.60‰ to -0.42‰), as well as
307 those for major source materials such as coal combustion, smelting, and cement plants
308 (mean: -1.10‰ , -0.87‰ , and -1.42‰ respectively, Huang et al., 2016) and those for
309 PM near anthropogenic emissions such as industry, landfill, traffic, and coal-fired
310 power plants (mean: from -2.41‰ to -0.58‰) (Fig. 2). The result likely indicates an
311 important contribution of anthropogenic sources to the CX $\text{Hg}_{\text{PM}_{2.5}}$. However, the
312 $\delta^{202}\text{Hg}$ values of above mentioned potential sources are not distinguishable, thus we
313 could not identify the specific sources of $\text{Hg}_{\text{PM}_{2.5}}$ solely based on Hg isotopes. On the
314 other hand, the slight positive $\delta^{202}\text{Hg}$ values at the DMS are quite different from those
315 observed at remote sites (mean: from -1.45‰ to -0.83‰ ; Fig. 2). Nevertheless, a less
316 negative MDF of $\text{Hg}_{\text{PM}_{2.5}}$ has also been reported at the DMS in a previous study
317 ($\delta^{202}\text{Hg} = -0.26\text{‰}$, Yu et al., 2016). Hg_{PM} from immediate anthropogenic emissions is
318 generally characterized by negative $\delta^{202}\text{Hg}$, which in turn suggests that the more
319 positive $\delta^{202}\text{Hg}$ of Hg_{PM} at the DMS might be affected by atmospheric processes, like
320 photo-reduction of Hg^{2+} (Bergquist and Blum, 2007; Zheng and Hintelmann, 2009).

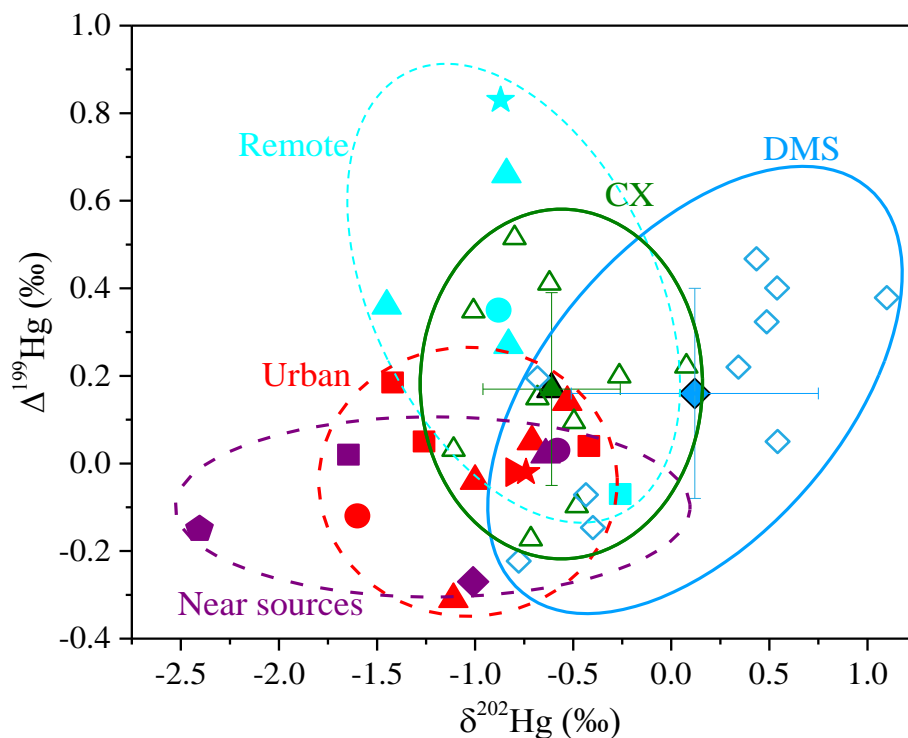
321 In contrast to MDF, the odd-MIF of $\text{Hg}_{\text{PM}_{2.5}}$ at the two sites were comparable
 322 ($p > 0.05$, T Test), with $\Delta^{199}\text{Hg}$ values of $0.17 \pm 0.22\text{‰}$ (from -0.17‰ to 0.52‰) at
 323 the CX and $0.16 \pm 0.24\text{‰}$ (from -0.22‰ to 0.47‰) at the DMS, respectively. The
 324 significant positive $\Delta^{199}\text{Hg}$ in this study are similar to those observed in coastal areas
 325 (Rolison et al., 2013; Yu et al., 2020) and in remote areas in China (Fu et al., 2019),
 326 but distinguishable from those in urban and industrial areas with near-zero values due
 327 to anthropogenic emissions (Das et al., 2016; Huang et al., 2016, 2018, 2020; Xu et al.,
 328 2019; Yu et al., 2016). A laboratory study has indicated that photo-reduction of Hg^{2+}
 329 restrains odd Hg in reactants (aerosols here) in priority, which shifts $\Delta^{199}\text{Hg}$ values
 330 positively (Bergquist and Blum, 2007). Thus, it's reasonably supposed that the
 331 positive odd-MIF of Hg_{PM} in the study region was associated with photo-reduction of
 332 Hg^{2+} in aerosols. As shown in Table S2 and Fig. S1, $\text{PM}_{2.5}$ samples affected by long
 333 range transport of air masses mostly had large positive $\Delta^{199}\text{Hg}$, like $\text{PM}_{2.5}$ collected on
 334 Apr. 4, 2018 from the CX and on Jan. 10, 2018 from the DMS. It's probably related to
 335 extensive photo-reduction of Hg^{2+} in aerosols during long-range transport, as previous
 336 studies suggested (Fu et al., 2019; Huang et al., 2016). In addition, the MIF of ^{200}Hg ,
 337 most probably relating to photo-reactions, was significant positive and displayed no
 338 spatial difference ($0.11 \pm 0.07\text{‰}$ at the CX and $0.14 \pm 0.07\text{‰}$ at the DMS; $p > 0.05$, T
 339 Test), which suggests enhanced and homogeneous photo-reactions in the study region.
 340 It is worth noting that a part of $\text{PM}_{2.5}$ samples collected from the DMS displayed
 341 negative $\delta^{202}\text{Hg}$ and near-zero $\Delta^{199}\text{Hg}$, similar to those from the CX (Fig. 2).
 342 Compared with the previous study (Yu et al., 2016), our results provide isotopes
 343 evidence that $\text{Hg}_{\text{PM}_{2.5}}$ at the DMS was affected by multiple sources and one of them
 344 might be regional anthropogenic emissions.

345 **Table 1** Concentrations and isotopic compositions of $\text{Hg}_{\text{PM}_{2.5}}$ at the industrial site (CX)
 346 and mountain site (DMS) in East China

Parameter ^a	CX		DMS	
	Mean \pm sd	Range	Mean \pm sd	Range
$\text{Hg}_{\text{PM}_{2.5}}$ ($\mu\text{g g}^{-1}$)	0.52 ± 0.23	0.15 ~ 1.10	0.85 ± 0.63	0.18 ~ 2.80
$\text{Hg}_{\text{PM}_{2.5}}$ (pg m^{-3})	16.3 ± 17.8	1.6 ~ 90.7	29.6 ± 35.9	2.9 ~ 181.3
$\delta^{202}\text{Hg}$ (‰)	-0.61 ± 0.35	-1.11 ~ 0.08	0.12 ± 0.63	-0.78 ~ 1.10

$\Delta^{199}\text{Hg}$ (‰)	0.17 ± 0.22	$-0.17 \sim 0.52$	0.16 ± 0.24	$-0.22 \sim 0.47$
$\Delta^{201}\text{Hg}$ (‰)	0.21 ± 0.18	$-0.07 \sim 0.48$	0.23 ± 0.36	$-0.29 \sim 0.66$
$\Delta^{200}\text{Hg}$ (‰)	0.11 ± 0.07	$-0.01 \sim 0.23$	0.14 ± 0.07	$0.06 \sim 0.28$
$\Delta^{204}\text{Hg}$ (‰)	0.19 ± 0.36	$-0.16 \sim 0.93$	3.58 ± 3.68	$0.26 \sim 11.38$

347 ^a 51 samples collected from CX and 32 samples from DMS were used for $\text{Hg}_{\text{PM}_{2.5}}$ concentration
348 analysis; 10 samples from each site for isotope analysis.



349
350 **Fig. 2** Isotopic compositions of Hg_{PM} at the multiple types of sites
351 (This study: \blacktriangle \triangle mean and each values at the CX, \blacklozenge \lozenge mean and each values at the DMS;
352 Remote sites: \star coast, \blacksquare \blacktriangle mountain, \bullet island (Fu et al., 2019; Rolison et al., 2013; Yu et al.,
353 2016); Urban sites in China: \blacktriangle Beijing, \bullet Changchun, \star Chengdu, \blacksquare Guiyang, \blacktriangle Xi'an (Huang
354 et al., 2015, 2016, 2019, 2020; Xu et al., 2017, 2019; Yu et al., 2016); Sites near emission sources:
355 \blacktriangle \blacksquare industrial, \bullet landfill, \blacklozenge traffic, \star near CFPP (Das et al., 2016; Huang et al., 2018; Yu et al.,
356 2016)

357 3.2. Influence of anthropogenic emissions on MDF of $\text{Hg}_{\text{PM}_{2.5}}$

358 Prior studies have compiled Hg isotopic compositions of major source materials,
359 such as fossil fuels, non-ferrous metal ores, and crustal rocks, which generally display
360 large negative $\delta^{202}\text{Hg}$ and negative or near-zero $\Delta^{199}\text{Hg}$ values (Huang et al., 2016;
361 Sun et al., 2016b). Combustion or/and industrial processing induces limited MIF (Sun
362 et al., 2013; Sun et al., 2016b), so we assumed that emitted Hg conserves the odd
363 isotope MIF of source materials. The $\Delta^{199}\text{Hg}$ values for most of the $\text{Hg}_{\text{PM}_{2.5}}$ samples
364 are distinguishable from those of source materials, indicating that anthropogenic
365 emissions were not the drive factors for odd-MIF of $\text{Hg}_{\text{PM}_{2.5}}$ in the study region. As

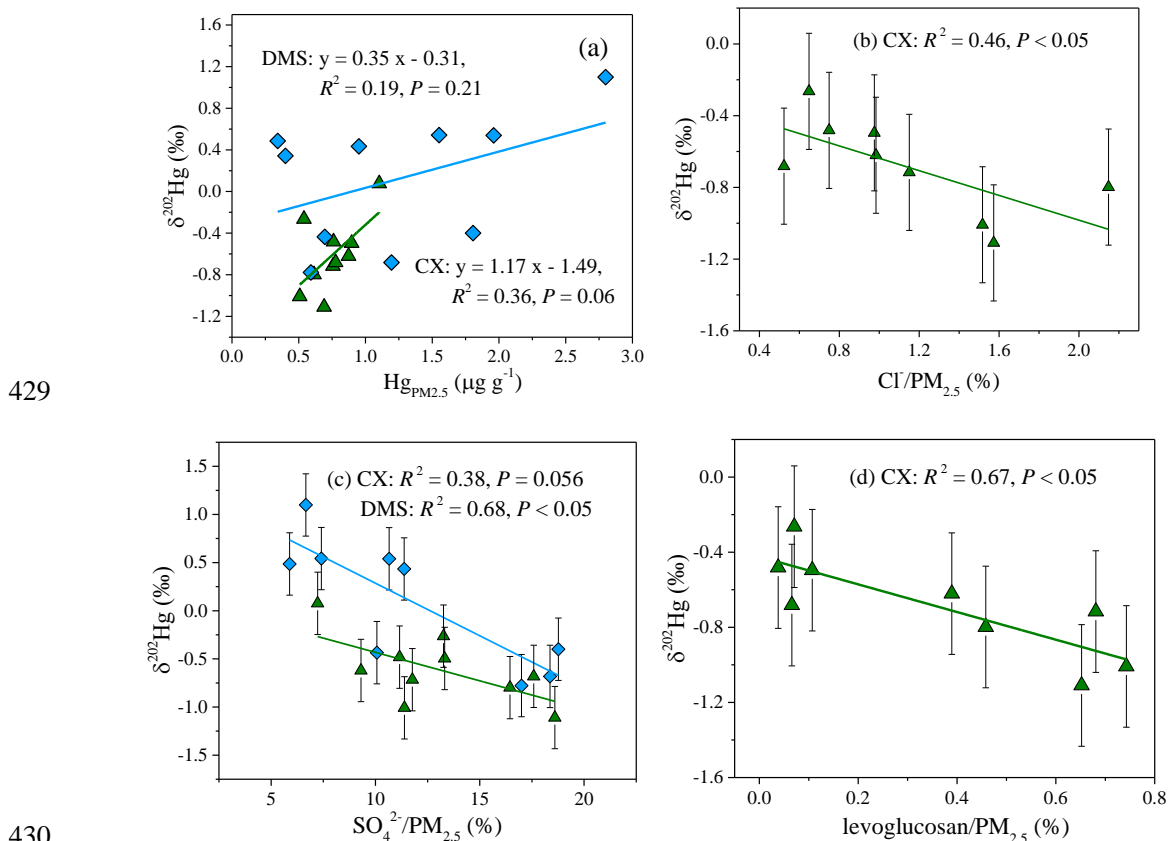
366 for MDF, the above analyses indicated that the MDF of $\text{Hg}_{\text{PM}_{2.5}}$ at the CX was
367 subjected to local anthropogenic sources, while the MDF at the DMS was probably
368 caused by the combination of atmospheric transformations and regional emissions.
369 The result was supported by the correlation between $\delta^{202}\text{Hg}$ values and $\text{Hg}_{\text{PM}_{2.5}}$
370 concentrations which was insignificant at the DMS, but significant at a loose level at
371 the CX (Fig. 3a). Pearson correlation between $\delta^{202}\text{Hg}$ and chemical components was
372 further conducted to explore the impacts of anthropogenic emissions on $\text{Hg}_{\text{PM}_{2.5}}$.

373 The $\delta^{202}\text{Hg}$ values at the CX were significantly correlated to Cl^- content ($R^2 =$
374 $0.46, P < 0.05$, Fig. 3b) and well associated with SO_4^{2-} content in $\text{PM}_{2.5}$ ($R^2 = 0.38, P$
375 $= 0.056$, Fig. 3c). In this study, Cl^- was mainly originated from coal combustion,
376 given that Cl^- content in $\text{PM}_{2.5}$ was not correlated to Na^+ . Besides, SO_4^{2-} was
377 primarily transformed from SO_2 which is mainly emitted from coal combustion. The
378 results imply that coal combustion played an important role in the MDF of $\text{Hg}_{\text{PM}_{2.5}}$ at
379 the CX. It has also been reported that coal combustion has a large contribution of
380 $\sim 50\%$ to total Hg emissions in Zhejiang province (Zhang et al., 2015). Differently
381 from the CX, the $\delta^{202}\text{Hg}$ values at the DMS were significantly correlated to SO_4^{2-} ($R^2 =$
382 $0.68, P < 0.05$, Fig. 3c), but not to Cl^- ($P > 0.05$, Table S4). It seems unlikely that
383 coal combustion was the predominant contributor to the positive MDF at the DMS.
384 Whereas under the influence of transport, the transformation of SO_2 to SO_4^{2-} usually
385 enhances and the photo-reduction of Hg^{2+} in aerosols tends to extensive which would
386 shift $\delta^{202}\text{Hg}$ to positive to a certain extent (Bergquist and Blum, 2007). The results
387 imply that coal combustion emissions at a regional scale or from long-range transport
388 had a potential impact on the MDF of $\text{Hg}_{\text{PM}_{2.5}}$ at the DMS, which is consistent with an
389 earlier study conducted at the same site based on Hg concentration and trajectory
390 analysis (Yu et al., 2015).

391 It should be noted that the $\delta^{202}\text{Hg}$ values at the CX slightly shift to positive
392 compared to those for emitted Hg_{PM} from coal combustion. Hg isotopic compositions
393 of coals in China have large variations in MDF with $\delta^{202}\text{Hg}$ values from -2.36% to
394 -0.14% (Biswas et al., 2008; Yin et al., 2014b). A prior study estimated that emitted
395 Hg_{PM} has a shift of -0.5% relative to $\delta^{202}\text{Hg}$ of coal feeds based on the mass balance

396 model (Sun et al., 2014). Then $\delta^{202}\text{Hg}$ values for Hg_{PM} emitted from coal combustion
397 in China were estimated to be -2.86‰ to -0.64‰. There are many metal smelting
398 factories near the CX. Although we did not measure the tracers for smelting, the mean
399 $\delta^{202}\text{Hg}$ value for non-ferrous metal ores was reported to be $-0.47 \pm 0.77\%$ (Yin et al.,
400 2016). We assumed that Hg emitted from non-ferrous metal smelting conserves the
401 isotopes of source materials due to a lack of data for processing at the current stage
402 (Sun et al., 2016b). Then, less negative MDF of Hg from non-ferrous metal smelting
403 could explain the positive-shift MDF at the CX relative to coal combustion emissions.
404 Thus, it is reasonably inferred that the MDF of $\text{Hg}_{\text{PM}_{2.5}}$ at the CX is a result of
405 multiple anthropogenic sources such as coal combustion and non-ferrous metal
406 smelting. In addition, we found a close negative correlation between $\delta^{202}\text{Hg}$ and
407 levoglucosan content in $\text{PM}_{2.5}$ at the CX ($R^2 = 0.67$, $P < 0.05$, Fig. 3d) excluding a
408 $\text{PM}_{2.5}$ sample collected on Dec. 19, 2017. Levoglucosan is considered an excellent
409 indicative of biomass burning. Thus, we cannot rule out the possibility that the
410 contribution of biomass burning led to a negative deviation of $\delta^{202}\text{Hg}$ values at the CX
411 to some extent. Previous studies have reported that biological materials display
412 negative $\delta^{202}\text{Hg}$ and $\Delta^{199}\text{Hg}$ values, like foliage ($\delta^{202}\text{Hg}$: -2.67‰ to -1.79‰; $\Delta^{199}\text{Hg}$:
413 -0.47‰ to -0.06‰), litterfall samples ($\delta^{202}\text{Hg}$: -3.03‰ to -2.35‰; $\Delta^{199}\text{Hg}$: -0.44‰ to
414 -0.22‰), and lichen ($\delta^{202}\text{Hg}$: -2.32‰ to -1.83‰; $\Delta^{199}\text{Hg}$: -0.35‰ to -0.22‰)
415 (Demers et al., 2013; Jiskra et al., 2015; Yin et al., 2013; Yu et al., 2016; Zheng et al.,
416 2016). Such negative $\delta^{202}\text{Hg}$ and $\Delta^{199}\text{Hg}$ of biological materials could not explain the
417 isotopes of $\text{Hg}_{\text{PM}_{2.5}}$ in this study. Moreover, the contribution of biomass burning is
418 supposed to shift $\Delta^{199}\text{Hg}$ values negatively, but we found no significant negative
419 correlation between $\Delta^{199}\text{Hg}$ and K^+ or levoglucosan from the whole study period
420 (Table S4). This unexpected result might be due to the fact that the substantial
421 biomass burning often occurs in a short time (i.e., Mar. 2018, Fig. S2a,
422 <https://firms.modaps.eosdis.nasa.gov/>). In this study, the most negative odd-MIF was
423 observed for $\text{PM}_{2.5}$ samples collected on Mar. 21, 2018, with a $\Delta^{199}\text{Hg}$ value of
424 -0.17‰ at the CX and -0.22‰ at the DMS. Those $\text{PM}_{2.5}$ samples were likely related
425 to biomass burning, since they were associated with air masses originating from or

426 passing through the northeast of China with dense fire spots (Fig. S2b). The findings
 427 suggest the biomass burning was not the dominant contributor of $\text{Hg}_{\text{PM}_{2.5}}$ in the study
 428 region, but it would change the isotopes of $\text{Hg}_{\text{PM}_{2.5}}$ in some times.



430
 431 **Fig. 3** Relationships of $\delta^{202}\text{Hg}$ with (a) Hg , (b) Cl^- , (c) SO_4^{2-} , and (d) levoglucosan
 432 contents in $\text{PM}_{2.5}$ at CX (▲) or DMS (◆). Uncertainty (2σ) for $\delta^{202}\text{Hg}$ in $\text{PM}_{2.5}$ is
 433 0.25‰.

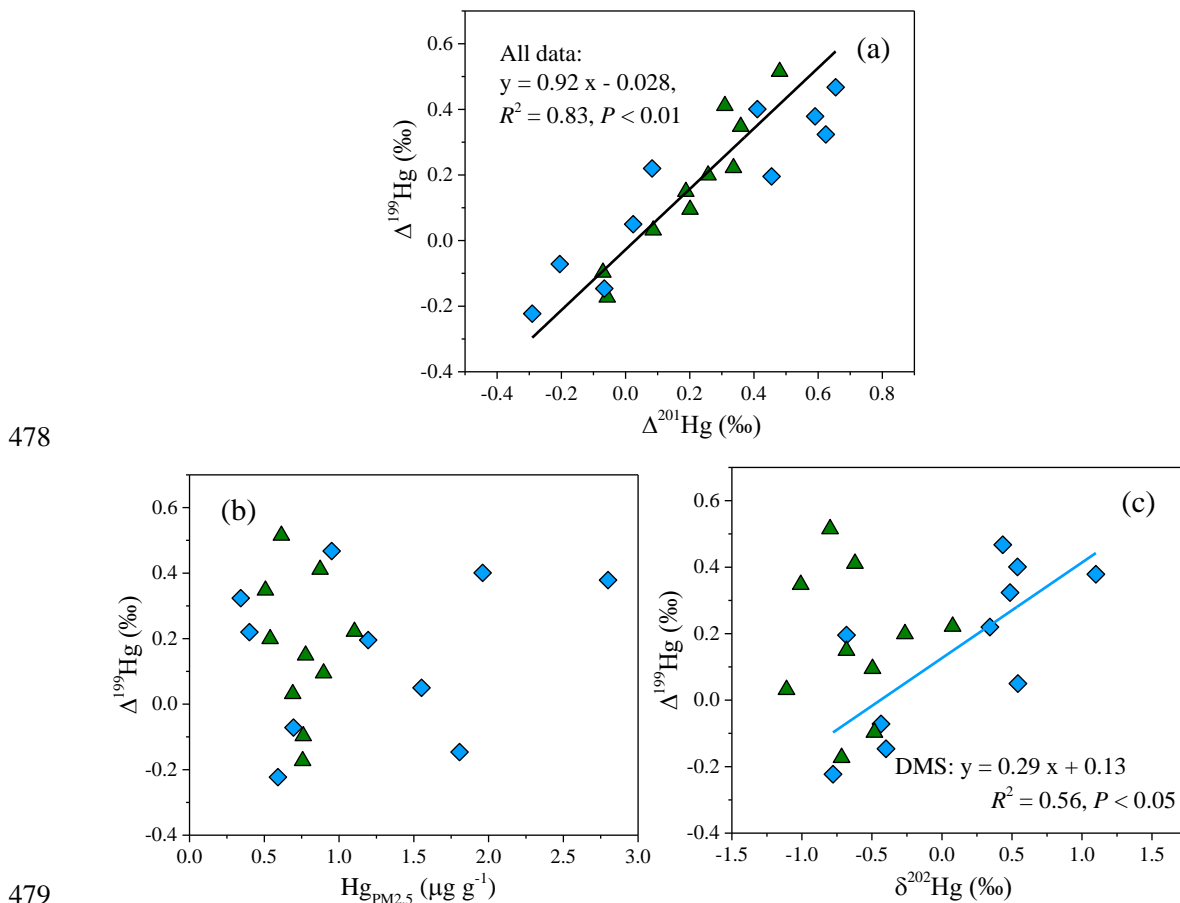
434 3.3. Influence of photo-chemical processes on isotopes of $\text{Hg}_{\text{PM}_{2.5}}$

435 Large odd-MIF of Hg isotopes in most $\text{PM}_{2.5}$ samples of this study was likely
 436 related to photo-chemical processes. An experiment study has found that the oxidation
 437 of Hg^0 by halogen atoms ($\text{Cl}\cdot$ or $\text{Br}\cdot$) results in a negative shift of $\Delta^{199}\text{Hg}$ in product
 438 Hg^{2+} (Sun et al., 2016a). Given that partitioning of Hg^{2+} between gas and particle
 439 phases leads to limited odd-MIF of Hg isotopes (Fu et al., 2019; Wiederhold et al.,
 440 2010), the formation of Hg_{PM} via oxidation of Hg^0 and following adsorption on
 441 particles could not explain the positive odd-MIF of $\text{Hg}_{\text{PM}_{2.5}}$ in this study. Previous
 442 experiments and field studies have reported that photo-reduction of inorganic Hg^{2+} in
 443 the aqueous solution induces odd-MIF of Hg isotopes and results in large positive
 444 $\Delta^{199}\text{Hg}$ values in reactants (aerosols here, Bergquist and Blum, 2007; Zheng and

445 Hintelmann, 2009). Hence, photo-reduction of Hg^{2+} in aerosols was invoked as a key
446 factor for the odd-MIF of $\text{Hg}_{\text{PM}_{2.5}}$ in the study region. The linear relationship between
447 $\Delta^{199}\text{Hg}$ and $\Delta^{201}\text{Hg}$ is often used to identify the odd-MIF processes of Hg isotopes.
448 The slope of $\Delta^{199}\text{Hg}$ versus $\Delta^{201}\text{Hg}$ yielded from the data of each site was 1.16 ($R^2 =$
449 0.92) at the CX and 0.63 ($R^2 = 0.85$) at the DMS, respectively. The data over the two
450 sites defined a straight line with a slope of 0.92 ($R^2 = 0.83$, $P < 0.01$; Fig. 4a). The
451 near-unity slope of $\Delta^{199}\text{Hg}$ versus $\Delta^{201}\text{Hg}$ was widely observed in particles from
452 coastal site and from other locations in Asia (Fu et al., 2019; Rolison et al., 2013;
453 Huang et al., 2016, 2019; Xu et al., 2019). The $\Delta^{199}\text{Hg}/\Delta^{201}\text{Hg}$ ratios of this study are
454 more consistent with the indicative ratio of aqueous photo-reduction of inorganic Hg^{2+}
455 (~ 1.0 , Bergquist and Blum, 2007; Zheng and Hintelmann, 2009), but different from
456 the ratios of other processes, like photo-oxidation (1.64 by $\text{Br}\cdot$ and 1.89 by $\text{Cl}\cdot$, Sun et
457 al., 2016) and photo-demethylation (1.36, Bergquist and Blum, 2007). Therefore, the
458 photo-reduction of Hg^{2+} in aerosols might be the critical factor for the observed
459 positive odd-MIF of $\text{Hg}_{\text{PM}_{2.5}}$ in the study region.

460 The similarity of odd-MIF anomaly between the CX and DMS suggests the
461 photo-reduction of Hg^{2+} in aerosols was homogeneous at a regional scale. However,
462 the relationships of $\Delta^{199}\text{Hg}$ with $\text{Hg}_{\text{PM}_{2.5}}$ content and $\delta^{202}\text{Hg}$ showed distinct spatial
463 differences. For the DMS, the $\Delta^{199}\text{Hg}$ values generally decreased with $\text{Hg}_{\text{PM}_{2.5}}$ content
464 increased (Fig. 4b) and the correlation between $\Delta^{199}\text{Hg}$ and $\delta^{202}\text{Hg}$ was significantly
465 positive ($R^2 = 0.56$, $P < 0.05$; Fig. 4c). Experimental studies indicated that
466 photo-reduction of Hg^{2+} releases Hg^0 and preferentially retains odd and heavier
467 isotopes in solutions (Bergquist and Blum, 2007; Zheng and Hintelmann, 2009),
468 which is expected to result in a positive relationship between $\Delta^{199}\text{Hg}$ and $\delta^{202}\text{Hg}$ and
469 an inverse relationship between $\Delta^{199}\text{Hg}$ and $\text{Hg}_{\text{PM}_{2.5}}$ content. In this study, the
470 consistent relationships of $\Delta^{199}\text{Hg}$ with $\delta^{202}\text{Hg}$ and $\text{Hg}_{\text{PM}_{2.5}}$ at the DMS strongly imply
471 a predominant role of photo-reduction of Hg^{2+} in isotopic fractionation of $\text{Hg}_{\text{PM}_{2.5}}$ at
472 this site. Meanwhile, the $\delta^{202}\text{Hg}$ signatures of anthropogenic emissions from regional
473 and long-range transport might be largely obscured by photo-reduction process, which
474 well explains the positive $\delta^{202}\text{Hg}$ at the DMS. In contrast, the variation of $\Delta^{199}\text{Hg}$ at

475 the CX was not associated with $\text{Hg}_{\text{PM}_{2.5}}$ contents or $\delta^{202}\text{Hg}$. The result suggests an
 476 insignificant impact of photo-reduction of Hg^{2+} relative to anthropogenic sources on
 477 the MDF and Hg content in $\text{PM}_{2.5}$ at the CX.



479
 480 **Fig. 4** Linear relationships between (a) $\Delta^{199}\text{Hg}$ and $\Delta^{201}\text{Hg}$, (b) $\Delta^{199}\text{Hg}$ and $\text{Hg}_{\text{PM}_{2.5}}$
 481 content, and (c) $\Delta^{199}\text{Hg}$ and $\delta^{202}\text{Hg}$ at the CX (\blacktriangle) and DMS (\blacklozenge). Uncertainty (2σ) for
 482 $\Delta^{199}\text{Hg}$ and $\delta^{202}\text{Hg}$ in $\text{PM}_{2.5}$ is 0.03‰ and 0.25‰, respectively.

483 **3.4. Potential mechanism of even-MIF**

484 A small but significant MIF of ^{200}Hg was observed in most $\text{PM}_{2.5}$ samples from
 485 this study, with mean $\Delta^{200}\text{Hg}$ values of $0.11 \pm 0.07\text{‰}$ at the CX and $0.14 \pm 0.07\text{‰}$ at
 486 the DMS. They are more positive than those in urban (mean = 0.01‰ to 0.09‰, Das
 487 et al., 2016; Huang et al., 2016; Xu et al., 2017) and remote areas (mean = 0.07‰ to
 488 0.10‰, Fu et al., 2019), but similar to those in coastal areas and island (Fu et al., 2019;
 489 Rolison et al., 2013). In general, Hg emitted from anthropogenic sources has $\Delta^{200}\text{Hg}$
 490 of near-zero (Sun et al., 2016b), while large $\Delta^{200}\text{Hg}$ values are mainly observed in
 491 atmospheric samples, i.e., precipitation, gaseous Hg^{2+} , and aerosols (Chen et al., 2012;
 492 Fu et al., 2019; Rolison et al., 2013). Significant even-MIF of Hg isotopes has been

493 suggested to associate with photo-oxidation of Hg^0 , from upper troposphere or/and
494 from in situ involving UV light and oxidants (Chen et al., 2012; Fu et al., 2019). This
495 could help explain significant $\Delta^{200}\text{Hg}$ values in coastal areas where halogen atoms are
496 expected to be abundant. The $\Delta^{200}\text{Hg}$ values in $\text{PM}_{2.5}$ were not different between sites,
497 similar to $\Delta^{199}\text{Hg}$ values, which supports that the observed $\Delta^{200}\text{Hg}$ were associated
498 with photo-chemical processes of minor spatial difference.

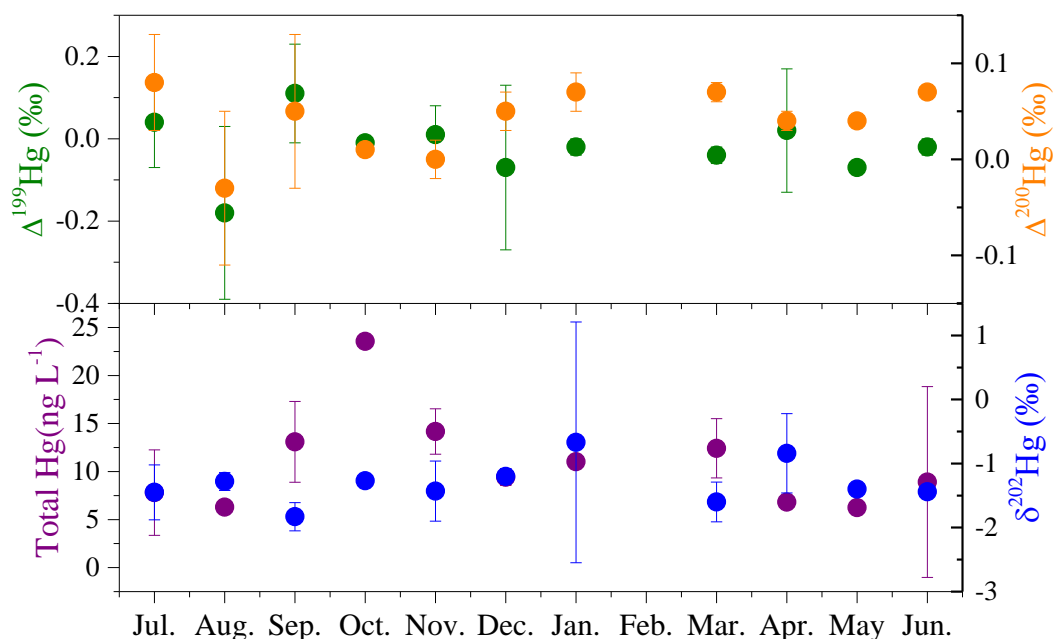
499 Gaseous elemental Hg is the predominant form of Hg in atmosphere, which in
500 China was generally characterized by slight negative even-MIF and odd-MIF (Fu et
501 al., 2018; Yu et al., 2020). Given that gas-particle partitioning of Hg^{2+} is strongly
502 temperature-dependent, this process unlikely produces the MIF of Hg isotopes (Fu et
503 al., 2019). Thus, comparing the MIF of Hg isotopes between Hg^0 and Hg_{PM} might
504 shed light on the effect of species conversion on Hg isotopes. The experimental study
505 showed that the oxidation of Hg^0 vapor by $\text{Cl}\cdot$ or $\text{Br}\cdot$ results in positive $\Delta^{200}\text{Hg}$ values
506 in products (Sun et al., 2016a). Thus, this process, which generally enhances in the
507 coastal environment with abundant halogen atoms (Wang et al., 2019), would well
508 explain for the detectable positive $\Delta^{200}\text{Hg}$ values in products. However, the oxidation
509 of Hg^0 by $\text{Cl}\cdot$ or $\text{Br}\cdot$ should produce a negative odd-MIF in products (Sun et al.,
510 2016a), which is inconsistent with the observed positive $\Delta^{199}\text{Hg}$ in $\text{PM}_{2.5}$. A recent
511 study has reported that oxidation of Hg^0 by oxidizers other than $\text{Cl}\cdot$ or $\text{Br}\cdot$ might
512 induce a positive odd-MIF in the Hg^{2+} , but this deduction lied on limit isotopic study
513 of Hg^0 oxidation (Yu et al., 2020). Alternatively, the magnitude of photo-reduction of
514 Hg^{2+} larger than the oxidation of Hg^0 by Cl and Br atoms could result in the observed
515 positive odd-MIF of Hg_{PM} . Since photo-reduction of Hg^{2+} most likely occurred in
516 aerosols as previously discussed, oxidation of Hg^0 and consequent photo-reduction of
517 Hg^{2+} in aerosols was the possible cause of the positive $\Delta^{199}\text{Hg}$ and $\Delta^{200}\text{Hg}$ values of
518 $\text{PM}_{2.5}$ in the study region.

519 ***3.5. Isotopes of Hg in adjacent surface seawater***

520 Hg isotopes are often used to track the transport and transformations of Hg in the
521 environment. The average concentration of THg in seawater was $10.5 \pm 5.0 \text{ ng L}^{-1}$,
522 with a range of $1.9 \sim 23.6 \text{ ng L}^{-1}$ (Table S1). As shown in Fig. 5, the concentrations of

523 seawater THg displayed distinct time variations, with higher levels during Sep. ~ Mar.
524 than during Apr. ~ Aug, which is likely related to the precipitation cycle. The average
525 $\delta^{202}\text{Hg}$ value of seawater samples was $-1.31 \pm 0.59\text{‰}$, with most samples fell in the
526 range of $-2.00\text{‰} \sim -1.00\text{‰}$. Whereas the MIF of Hg isotopes in seawater samples was
527 not significant, with mean $\Delta^{199}\text{Hg}$, $\Delta^{201}\text{Hg}$, and $\Delta^{200}\text{Hg}$ values of $-0.02 \pm 0.07\text{‰}$, 0.00
528 $\pm 0.05\text{‰}$, and $0.04 \pm 0.03\text{‰}$, respectively. The negative MDF and near-zero MIF of
529 surface seawater are well consistent with those of source materials (Huang et al., 2016;
530 Sun et al., 2016b), suggesting the dominant effect of anthropogenic emissions on Hg
531 in offshore surface seawater. A minor change in the intensity of industrial activities as
532 expected among the months also supports the above deduction.

533 Isotopic compositions of THg in surface seawater and Hg_{PM} at the adjacent
534 industrial site are consistent in MDF but not in MIF. Similar results were obtained
535 comparing to wet deposition which presented negative $\delta^{202}\text{Hg}$ and positive $\Delta^{199}\text{Hg}$
536 and $\Delta^{200}\text{Hg}$ values (Chen et al., 2012; Huang et al., 2018). The negative MDF of Hg in
537 industrial $\text{PM}_{2.5}$ and adjacent surface seawater implies an important contribution of
538 local anthropogenic sources. On the other hand, the unity slope of $\Delta^{199}\text{Hg}$ versus
539 $\Delta^{201}\text{Hg}$ ($\Delta^{199}\text{Hg} = 1.12 \times \Delta^{201}\text{Hg} - 0.02$, $R^2 = 0.68$, $n = 19$, Fig. S3) indicates that the
540 odd-MIF of Hg isotopes in surface seawater was mainly produced by photo-reduction
541 of Hg^{2+} . Whereas, the minor $\Delta^{199}\text{Hg}$ anomalies suggest that photo-reduction was not
542 evident for surface seawater. A big discrepancy in the MIF of Hg isotopes between
543 atmospheric samples and surface seawater further evidences that atmospheric
544 transformations would induce the significant MIF of Hg isotopes and obscure the Hg
545 isotopic signatures of anthropogenic emissions.



546

547 **Fig. 5** Monthly variations of total Hg concentration, $\delta^{202}\text{Hg}$, $\Delta^{199}\text{Hg}$ and $\Delta^{200}\text{Hg}$ of
 548 surface seawater during the sampling period from July 2017 to June 2018

549 **4. Conclusion**

550 This study investigated Hg isotopic compositions in $\text{PM}_{2.5}$ collected from the
 551 neighboring industrial and mountain sites in a coastal area and in surface seawater
 552 close to the industrial area. $\text{Hg}_{\text{PM}_{2.5}}$ displayed a significant spatial difference in MDF
 553 but not in odd-MIF. Negative $\delta^{202}\text{Hg}$ in $\text{PM}_{2.5}$ at the CX was primarily induced by
 554 local industrial activities like coal combustion. Whereas, the slight positive $\delta^{202}\text{Hg}$ at
 555 the DMS could not be fully explained by anthropogenic emissions. Other than the
 556 effect of regional transport, a close correlation between $\delta^{202}\text{Hg}$ and $\Delta^{199}\text{Hg}$ at the
 557 DMS implies that photo-chemical processes likely contributed to the MDF of $\text{Hg}_{\text{PM}_{2.5}}$.
 558 Significant positive odd-MIF of $\text{Hg}_{\text{PM}_{2.5}}$ and the unity slope of $\Delta^{199}\text{Hg}$ versus $\Delta^{201}\text{Hg}$
 559 indicate an important role of photo-reduction of Hg^{2+} in aerosols. The observed
 560 positive $\Delta^{200}\text{Hg}$ values in this study were probably associated with photo-oxidation of
 561 Hg^0 which is generally enhanced in the coastal environment. THg in surface seawater
 562 was characterized by negative MDF and near-zero odd-MIF, which is more consistent
 563 with isotopic signatures of source materials. The MIF anomalies of Hg isotopes were
 564 larger for atmospheric $\text{PM}_{2.5}$ than for surface seawater, suggesting that atmospheric
 565 transformations induce significant MIF of Hg isotopes and obscure Hg isotopic
 566 signatures of initial anthropogenic emissions.

567 **Novelty statement**

568 A comparison of isotopic compositions of Hg_{PM_{2.5}} was conducted between the
569 neighboring industrial and mountain sites, which effectively revealed the influence of
570 anthropogenic emission sources and transformation processes on Hg isotopes. Hg
571 isotopic compositions in industrial PM_{2.5} and offshore surface seawater were also
572 compared. The results indicate that atmospheric transformations would induce
573 significant fractionation of Hg isotopes and obscure specific Hg isotopic signatures of
574 initial emissions.

575

576 **Data availability.** HYSPLIT trajectory model and gridded meteorological data
577 (Global Data Assimilation System, GDAS1) are available from the US National
578 Oceanic and Atmospheric Administration (<http://ready.arl.noaa.gov>). Fire data are
579 available in the Fire Information for Resource Management System (FIRMS,
580 <https://firms2.modaps.eosdis.nasa.gov/map/#d:2021-04-26..2021-04-27;@6.7,2.0,3z>).
581 All data in this study are available upon request to the first author via email
582 (linglingxu@iue.ac.cn).

583

584 **Author contributions.** JSC, LLX, and YRZ designed this study. MRL, LQY, YTC,
585 LT and HX conducted the sampling. YRZ and LLX participated in sample treatment
586 and measurements. LLX wrote the paper. JYS and YPC helped the graphics
587 production. All authors reviewed the paper.

588

589 **Competing interests.** The authors declare that they have no conflict of interest.

590

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600

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